

EXSoDOS 1.0: downscaling of weather extremes shifts for ensemble climate projections using ground-based measurements, reanalysis and stochastic modelling

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Abstract. Accurately representing the changes of local extreme weather events in climate projections is crucial for climate impact assessment and adaptation services. Climate models ~~often struggle with capturing these events~~ do not explicitly represent these events as observed by weather stations due to their coarse spatial resolution. Existing downscale products successfully reduce overall biases of past or future ~~climateological-climatological~~ variables, but the representation of variability and extreme events including their past and future shifts under climate change are still not addressed. A new stochastic model, EXSoDOS, addresses this gap by the DOWnScaling of weather EXtremes Shifts for ensemble climate projections using ground-based measurements, reanalysis, and global climate models. This is done by using a stochastic model that correlates coarse-scale gridded historical climate records with the point-scale measurements. Therefore, EXSoDOS combines ground-based data (either from the Global Historical Climatological Network or user-specified), ERA5 ~~reanalysis, and~~ (ECMWF Re-Analysis 5), and GCM (global climate model(GCM)-projections-) projections from CMIP6 (Coupled Model Intercomparison Project Phase 6) to downscale past and future daily climate records. We demonstrate EXSoDOS for 5 use cases, resp. daily minimum temperature in Belgium, daily maximum temperature in Azerbaijan, heat stress in India, wind velocity in Germany and precipitation in Mali. It is found that EXSoDOS is able to represent annual cycle variability, density distributions, and extreme events of return periods of up to 10 years, while they are all underrepresented by the raw GCM outputs. Observed tendencies towards more extremes between two past periods 1961–1990 and 1991–2020 are also better represented. Projections under the SSP585 scenario suggest amplified extremes in maximum temperature, precipitation, and heat stress by 2071–2100. Furthermore, downscaling affects the outcomes of shifting extremes under future climate change, which is evident in terms of both absolute and relative changes, as well as changes in return periods. While limitations of statistical downscaling persist, it is concluded that EXSoDOS offers a novel method for estimating past and future shifts in weather extremes for weather stations with a sufficient daily record of data of multiple decades.

1 Introduction

Global (Eyring et al., 2016) and regional (Katrakou et al., 2024; Coppola et al., 2021) climate model initiatives have been successful in predicting many aspects of climate change, like increasing temperature and shifts in precipitation amounts. However, events like heavy precipitation, heavy wind, extreme heat (stress) and cold spells ~~are generally underrepresented as observed~~
25 ~~by (point-scale) weather stations are not explicitly represented~~ in climate projections. This results from ~~their~~ ~~a scale mismatch~~
~~between~~ coarse spatial resolution ~~of GCMs and point-scale observations~~ for which the scale of the extreme events are too small to be resolved (IPCC, 2021; Fischer and Knutti, 2015; Sillmann et al., 2013). Particularly, meteorological point-scale observations generally show much more erratic temporal variability than the coarser-scale gridded products. Especially, extreme precipitation and wind, but also temperature and heat stress, can be influenced by confounding local climate effects resulting
30 in local features like local convective precipitation events and urban heat islands. However, local extreme weather events have large impacts on society, infrastructure and ecosystems. Hence, their underrepresentation hampers extreme hazard assessment and their impacts under climate change. Not only the absolute representation but also the shifts of extreme weather distributions under past and future climate change are crucial for climate risk assessment and adaptation planning (int, 2023).

Several techniques have been developed to downscale global (Eyring et al., 2016) and regional (Katrakou et al., 2024;
35 Coppola et al., 2021) climate projections to represent their small-scale climate effects. On the one hand, they include mechanistic downscaling using convection permitting atmospheric numerical models with a resolution of 7km down to 1km (Kendon et al., 2021), high enough to resolve deep convection and associated extreme precipitation (Vanden Broucke et al., 2019; Brisson et al., 2016) and local land use like urban areas (Wouters et al., 2016; De Ridder et al., 2015). they have been effective to project extreme weather events (Fosser et al., 2024; Termonia et al., 2018; Wouters et al., 2017; Saeed et al.,
40 2017) ~~on~~. ~~On~~ the other hand, statistical downscaling methods have been developed (Maraun, 2016) to finer ~~resolution grids~~
~~(Lange, 2019; Karger et al., 2023)~~ ~~grids of 0.5° (eg., ISIMIP3BASD by Lange, 2019) and 1km resolution (eg., CHELSA-W5E5 by Karger~~
and point observation locations (Switanek et al., 2022), ~~whereas GCM resolutions are ~ 1–3°. Finally, a downscaling method~~
~~to estimate the probability parameters and return values was recently proposed by (Benestad et al., 2025), which can be used~~
~~as input to weather generators.~~

45 However, future climate assessment of local extremes is hampered by the requirements of computational resources and input data requirements. Particularly, mechanistic downscaling with atmospheric numerical models are computationally very demanding since higher-resolution climate also requires smaller integration time step because of the Courant-Friedrich-Lewy criterion (Prein et al., 2015). The most recent point-location stochastic downscaling models (Switanek et al., 2022; Volosciuk et al., 2017; Devis et al., 2013) and weather generators (Van De Velde et al., 2023; Brisson et al., 2015) have only applied to
50 specific areas or variables, or have not been applied to future climate projections. Gridded downscaling products may either have a low resolution (Lange, 2019) or do not cover ensemble future climate either (Karger et al., 2023). While these different downscaling products can reduce overall biases of coarse-scale ~~products~~ ~~time series~~, their representation of extremes in terms of variability, distributions and their return periods as observed from weather stations are not addressed, let stand the shifts under past and future climate change. ~~A downscaling method that represents past and future shifts in probability distributions and~~

55 return levels according to observations and that is simultaneously constrained to time series from model members of climate projections doesn't exist yet.

To fill this gap, we present EXSoDOS, a DOWnScaling method of weather EXtremes Shifts for climate projections using ground-based measurements, reanalysis, global climate models, and statistical downscaling. It downscales time series from an ensemble of global climate models to represent the variability and extremes as observed by weather stations. The model
60 implements a stochastic downscaling technique that is calibrated on point-scale observations and coarse-scale reanalysis, with over 84 years of climate historical climate data (from 1940 to present). Hereby, historical coarse-scale reanalysis data and observations are normalized and correlated with each other. This is done by mapping predictor and predictand values to standard normal space using their empirical CDFs (i.e., apply the probability integral transform and then the inverse normal CDF), acquiring correlation in that space, and finally mapping back sampled values through the inverse empirical CDF of observations.

65 While its stochastic downscaling approach builds on established perfect-prognosis concepts, the novelty of EXSoDOS lies in its ability to evaluate shifts in local weather extremes across past and future climatological periods for different variables. This is achieved by the combined use of (1) globally available long-term datasets including weather station observations, ERA5 (ECMWF Re-Analysis 5; Hersbach et al., 2020), and the GCM (global climate model) ensemble from CMIP6 (Coupled Model Intercomparison Project Phase 6), (2) a non-parametric treatment of distributions, and a (3) workflow that allows direct comparison of observed, reconstructed, and projected extremes. Its end-to-end design makes consistent assessment of extreme-event variability and return levels across multiple decades, and variables for any climate region possible.

The method can be used at locations where ~~climatological records~~ longterm records (60 years) of observations are available. It can either employ observational data from the Global Historical Climate Network (GHCN) hosted by the National Oceanic and Atmospheric Administration (NOAA; [https://www.ncei.noaa.gov/products/land-based-station/global-historical-climatology-network-](https://www.ncei.noaa.gov/products/land-based-station/global-historical-climatology-network)
75 [or ingest data from local observational sources, eg., measurements carried out by national meteorological offices, agro-meteorological or ~~envionmental~~ environmental institutes. In the case of the latter, one only requires the position \(in latitude longitude coordinates\) and the observed time series in a commonly used format \(csv or parquet\). The stochastic model is then applied on multi-member climate projections towards 2100 after bias-correcting them with reanalysis data. As such, the
80 method follows the downscaling strategy proposed by \(Switanek et al., 2022\), and is further designed to represent climate extremes regarding different weather variables including daily minimum and maximum temperature, heat stress temperature and wind speed. To evaluate the representation of extreme events with respect to observations and their shifts under past and future climate change, a number of statistics are employed, including the daily-to-annual variability, density distributions and overlaps, and return periods.](https://www.ncei.noaa.gov/products/land-based-station/global-historical-climatology-network-)

85 EXSoDOS runs quickly and automatically. For a single station, a full downscaling run for one scenario typically takes ~5–10 seconds, and a 10-member ensemble completes in <1 minute on a modern CPU, excluding one-time data download and gridded bias-adjustment preprocessing. Downloading of source data (ERA5, and CMIP6 climate models, station data) can take longer (network dependent) and quantile-delta mapping (QDM) bias-adjustment (incl. ERA5 grid upscaling) on continental grids can take hours, which is performed once per model/grid.

90 [EXSoDOS](#) can be applied at locations where a sufficiently extensive record of observations from weather stations are available, preferably [30-60](#) years of continuous data record or longer. It is also easily applicable by requiring only a few parameters, particularly the station coordinates, the weather time series under scope (either temperature, precipitation, wind, or heat stress). The synthesis of time series for a handful of stations and a handful of climate model members and climate scenarios takes less than an hour on a contemporary desktop computer, resulting in a negligible carbon footprint of computer resources. Our service is demonstrated for several specific use cases across the globe, in which the statistical model is calibrated and validated on historical data, and applied on climate projections. The demonstration is featured by metrics common to extremes climate assessment, including annual cycles, density distributions and return periods.

The structure of the remaining of this paper is as follows. In the methods section 2, the program interface with global data sources, the downscaling procedure, and bias correction are described. The statistics for validation and assessment as well as the case areas are described after. In the results section 3, we perform a validation of the model output for the historical period, and the assessment with climate projections for each of the use cases. We conclude the paper with perspectives and challenges for extreme weather assessment under climate change in section 4.

2 Method

To perform the downscaling, coarse-scale predictors (x) are correlated to the point-scale predictands (y). Calibration and validation of the statistical model is done on the basis of historical reanalysis reconstruction x_r and the measurements from the weather stations y_o . While ensemble models of climate projections offer the statistical climate properties on the multi-annual time series including anomalies, they are not in synchrony with observations or, in other words, their sequence of weather patterns generally differ from that of the observations. Consequently, they cannot be used to correlate the coarse-scale model output with local measurements. Therefore, historical reanalysis, which is constrained and in synchrony with the observations, is used for the calibration. For calibration, half of the days in the time series are chosen randomly for calibration (x_r^c, y_o^c) and the other half for validation (x_r^v, y_o^v). [We used a 50:50 split, because we require that the overall distribution and extremes to be represented on a climatological timescale in both calibration and validation. Over a 60-year period we have an equivalent of 30-year data for both, which is in line with climatology assessment standards of the World Meteorological Organization \(WMO, 2017\).](#) Finally, the downscaling is applied on the climate model projections x .

115 To ensure scale consistency of the predictors between calibration phase and application phase, the climate reanalysis is upscaled to the climate model grid for each individual model. Furthermore, the downscaling procedure follows the perfect prognosis assumption (Maraun, 2016), which means that the predictor is describing the reality free from biases. Therefore, the biased predictors from the climate models (x_b) are bias-adjusted with a quantile-delta mapping procedure (Cannon et al., 2015), see also section 2.3. The downscaling strategy above leads to the following subsequent steps, see also fig. 1:

- 120 1. Acquire climate reconstruction (reanalysis) and upscale to the climate model grid by aggregation and interpolation
2. Acquire historical predictand (y_o) from weather stations

3. Extract historical predictor (x_r) from upscaled reanalysis at point locations of the weather stations
4. Randomly select half of the historical data for model calibration (x_r^c and y_r^c) and the remaining half for model validation (x_r^c and y_r^c)
- 125 5. calibration of the statistical model using x^c and y^c (section 2.2)
6. Validation of the statistical model using x_r^v and comparing the output with y_r^v , see section 2.4)
7. Acquire (biased) climate projections for point locations (x_b)
8. Climate projections are bias-adjusted (indicated with x_m) with upscaled reanalysis using quantile-delta mapping, section 2.3)
- 130 9. Synthesis of downscaled climate projections (y_m) by applying the statistical model on bias-adjusted climate projections (section 2.5)

As such, the procedure performs bias-adjustment at the coarse scales and statistical downscaling towards the point scale in separate steps. This avoids inconsistencies arising from mixing up statistical properties (particularly, quantile values) across spatial scales. The latter could occur if one would apply the statistical downscaling directly on the climate model output without bias-adjustment on the grid level. Especially, trends under climate change including those of the (extreme) quantile values at the local scales can differ to those (spatially aggregated) at the coarse scales, so these should not be used for one another. As shown by (Maraun, 2013), mapping point-scale quantiles to coarse-scale quantiles, or vice versa, may lead to unrealistic trends of extremes leading to misinterpretation. The strategy of scale separation where scale transition is established with a statistical model only after bias-adjustment at the model grid level is in accordance to previous methods (Lange, 2019; Switanek et al., 2022; Volosciuk et al., 2017), and advised by previous analyses (Maraun, 2013).

140 The key assumptions of EXSoDOS include (see also Tab. S1):

1. Realism of bias-adjusted coarse scale predictors (perfect prognosis assumption),
2. Stationarity of predictor–predictand correlations,
3. Stochastic representation of subgrid variability,
- 145 4. Sufficient record length of observations to represent extremes up to ~10-year return periods,
5. QDM bias-adjustment stationarity (Cannon et al., 2015) and
6. Independence across stations and variables.

The different processing steps are implemented in Python and computationally optimized as vectorized numerical operations with Python Xarray, Pandas and NumPy.

150 In the next subsections, we elaborate the input data sets, the statistical downscaling and bias-adjustment in more detail.

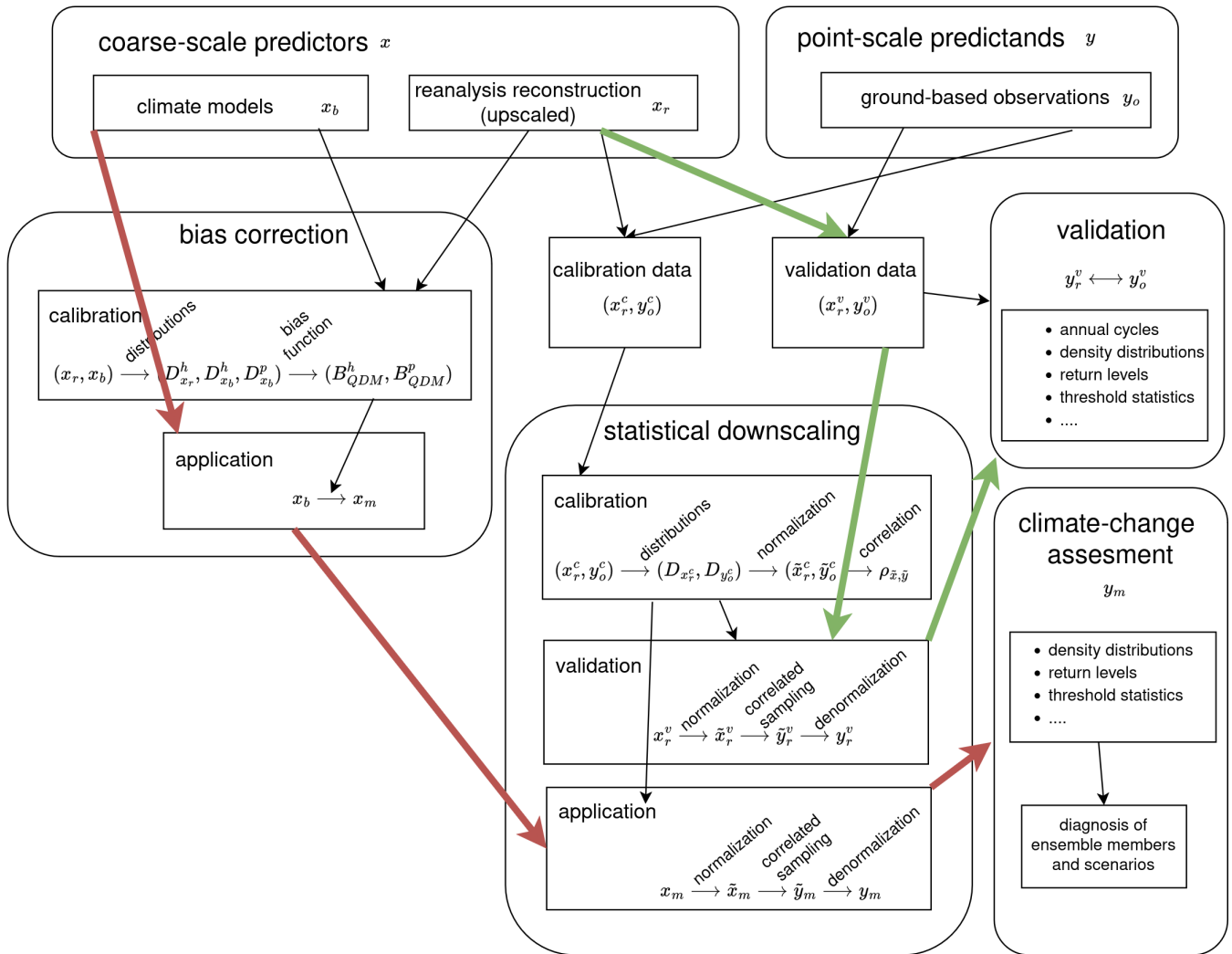


Figure 1. overview of EXSoDOS. The green arrows indicate the validation steps starting from the reanalysis reconstruction, whereas the red arrows indicate the assessment steps starting from the climate models.

2.1 Input

155 A key design feature of EXSoDOS is the consistent use of datasets with hyper-climatological temporal coverage (>60 years), which enables robust estimation of variability and extremes as well as their shifts between climatological periods. By relying exclusively on globally available datasets (ERA5, CMIP6, and station observations), the framework is fully transferable to any station location where a sufficiently long record of observations exists, including data-sparse regions such as Mali.

Observations collected from GHCN are used as predictand for calibration and validation. It includes 10010 stations for which it reports an available record of at least ~~30-years~~60 years, see Fig. 2. The minimum record length needs to be 60 years. On the one hand, the 60 year record is required to provide sufficient data for both calibration and validation. As a 50:50 split is applied, an equivalent of 30 years of data is available for both calibration and validation, which is in line climatology assessment standards of the World Meteorological Organization (WMO, 2017). On the other hand, one can address 2 x 30 years of records to evaluate the model performance to capture past shifts in weather extremes driven by global climate models. One should only use records with observational records covering >90% of the sample period. We do not infill gaps: missing values are treated as NaN and excluded from empirical CDF correlation construction. Alternatively, the service is able to read observational data provided by the user in a common format (csv, parquet, netcdf) for which one specifies their station coordinates (as elaborated 165 by one of the use cases, see Sect. 2.6.2).

Next, ERA5 (Hersbach et al., 2020) is used as reanalysis climate reconstruction for the coarse-scale predictor x_r . ERA5 reanalysis covers more than 80 years of data from 1940 until now that can be used for calibration and validation of the statistical model, hence suitable to build a robust statistical relation between predictor and predictand. Finally, the climate model ensemble members of the Coupled Model Intercomparison Project Phase 6 (CMIP6) see (Eyring et al., 2016) from 1931 up to 170 2100 are used as predictor input for EXSoDOS to generate and assess downscaled daily time series from past to future climate. These are downloaded from one of the data nodes of the Earth System Grid Federation (ESGF), see <https://esgf.llnl.gov/>. For demonstration of EXSoDOS, 9 models listed in the Tab. 1 are used, but the user can specify any other combination of model members available at ESGF.

2.2 Statistical downscaling

175 The stochastic downscaling strategy follows the general philosophy of perfect-prognosis and analog-based approaches previously proposed in the literature (eg., Volosciuk et al., 2017; Switanek et al., 2022). EXSoDOS does not introduce a fundamentally new class of stochastic models; instead, it adapts and extends these approaches within a unified framework that is explicitly designed for multi-decadal extreme-event assessment across historical and future climates.

The basic model strategy is to randomly sample point-scale predictands (y) from measurements in such a way that they are 180 correlating with coarse-scale predictors (x), as proposed by Switanek et al. (2022). As such, the predictors can be extracted from climate projection and then used as input for the model to generate future point-scale time series including future extreme values. The current method is now reformulated to work not only with precipitation but also with other climate parameters, including temperature, wind speed and heat stress. This is done by performing the normalizing and denormalization steps

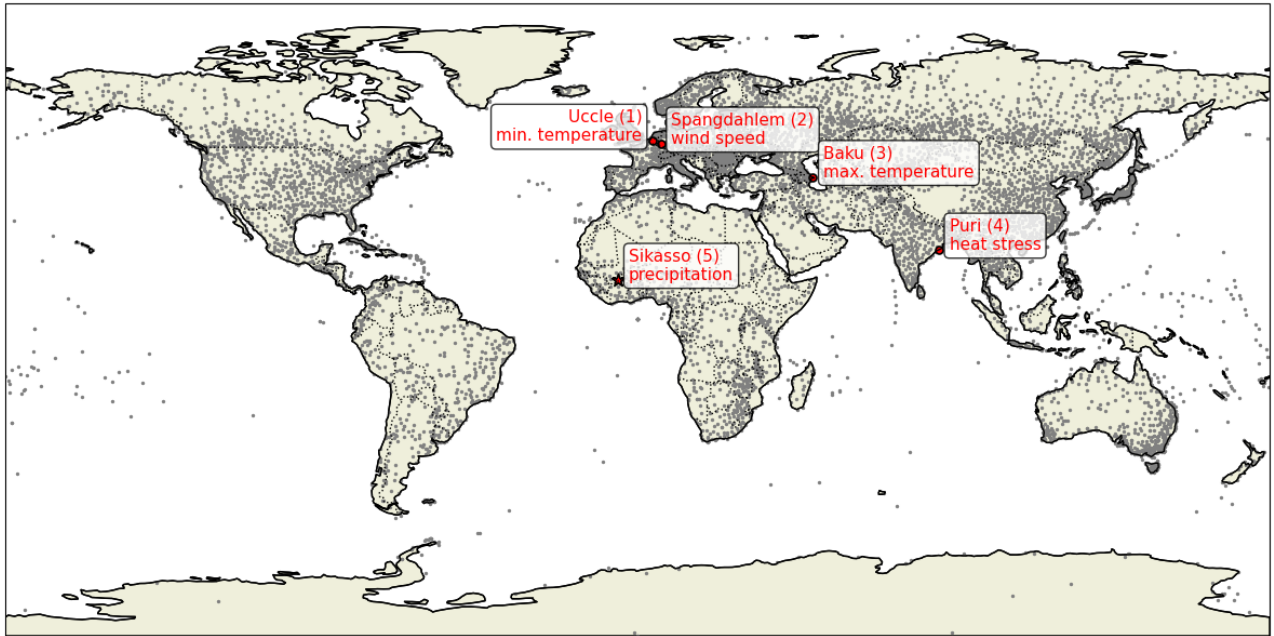


Figure 2. The 10010 stations from the global historical network (grey dots) of with it reports at least 30 years of data within the time span 1940–2023. EXSoDOS use cases for each of the variables provided by the global network are indicated as red circles and provided a user-input data as star for Sikasso (Mali).

Model	Member	Resolution	Reference
CanESM5	r1i1p1f1	$\sim 2.8^\circ \times 2.8^\circ$	Swart et al. (2019)
INM-CM4-8	r1i1p1f1	$\sim 1.5^\circ \times 2.0^\circ$	Volodin et al. (2018)
INM-CM5-0	r1i1p1f1	$\sim 1.5^\circ \times 2.0^\circ$	Volodin et al. (2017)
IPSL-CM6A-LR	r2i1p1f1	$\sim 2.5^\circ \times 2.5^\circ$	Boucher et al. (2020)
MIROC6	r1i1p1f1	$\sim 1.4^\circ \times 1.4^\circ$	Tatebe et al. (2019)
MPI-ESM1-2-LR	r10i1p1f1	$\sim 1.9^\circ \times 1.9^\circ$	Mauritsen et al. (2019)
MPI-ESM1-2-HR	r1i1p1f1	$\sim 0.9^\circ \times 0.9^\circ$	Gutjahr et al. (2019)
MRI-ESM2-0	r1i1p1f1	$\sim 1.1^\circ \times 1.1^\circ$	Yukimoto et al. (2019)
TaiESM1	r1i1p1f1	$\sim 1.25^\circ \times 1.25^\circ$	Lee et al. (2020)

Table 1. Climate models and references used in this study from the 6th phase of the Coupled Model Intercomparison Project (CMIP6), see Eyring et al. (2016).

with a non-parametric distributions. We further improve the model by implementing different predictor classes of different magnitude. Additionally, we implement a rescaling step of the point-scale predictand with the coarse-scale predictor before

applying the stochastic model, which leads to improved results on high extremes in case of precipitation and wind speed (see Sect. 3.1 for results). For the sake of clarity, we elaborate the details of the statistical model in the subsections below.

2.2.1 Stochastic model

190 A statistical relationship between x and y is established for each month separately. ~~To increase the sample size and to make the~~
~~Calibration is performed per calendar month. To increase sample size, and providing smooth~~ transition between months ~~more~~
~~smooth, calibration while preserving seasonal representativeness, we include~~ data from the ~~month before and the month after~~
~~is taken into account as well~~ adjacent months (± 1 month), yielding an effective 3-month seasonal window. A sensitivity test
using single-month windows showed slightly less robust tail estimates due to reduced calibration sample size per month (not
shown).

195 Afterwards, to allow the relation between the predictor and predictand to be dependent on the magnitude of the predictor, the sample data is subdivided into a number of categories ($n = 3$) of different magnitudes of the predictor percentiles. In the case of temperature and heat stress, the category bins of equal sample size are constructed, for which their boundaries are determined by equidistant values of the cumulative distribution function $p(x)$:

$$p_i = i/n \text{ for } 0 \leq i \leq n \quad (1)$$

200 In the case of precipitation and wind speed, daily distributions have large tails, which we want to represent in the stochastic model. We use an exponential profile for quantiles of category borders which provides more categories in the tails, hence to cover the large variation in the tails of the distributions:

$$p_i = \begin{cases} 1 - \exp(-l_i) & \text{for } 0 \leq i \leq n - 1 \\ 1 & \text{for } i = n \end{cases} \quad (2)$$

where

$$205 \quad l_i = i \ln(f_{\max}) / (n - 1) \text{ for } 0 \leq i \leq n - 1 \quad (3)$$

are equidistant values for i going from 0 to $n - 1$, and $\ln(f_{\max})$ is the logarithm of the (prescribed) chance (that is, frequency f_{\max}) for the observation to fall in the upper bin. As such, $p_{n-1} = 1 - f_{\max}$ represents the lower quantile border of that the upper bin. The exponential profile (note that $\ln(f_{\max})$ is negative hence also p_i) leads to more categories in the tails. Finally, the quantile categories q_i are calculated from the quantile function q_D of the calibration data x_r^c of the historical period (h):

$$210 \quad q_i = q_{D_{x_r^c}}^h(p_i) \quad (4)$$

Here above and throughout this manuscript, we follow the notation of Maraun (2016). For each of these categories, the predictand is correlated with the predictor variable. At first, the predictor and predictands are normalized within each category for the calibration data:

$$\tilde{x}_r^c = q_N(p_{D_{x_r^c}}(x_r^c)) \quad (5)$$

$$215 \quad \tilde{y}_o^c = q_N(p_{D_{y_o^c}}(y_o^c)) \quad (6)$$

where $p_{D_{x_r^c}}$ and $p_{D_{y_o^c}}$ are the cumulative distribution functions of the the calibration predictor x_r^c and predictand y_o^c , and q_N the quantile function of the normal distribution.

Subsequently, the correlation between the normalized predictor and normalized predictand is estimated from \tilde{x}_r^c and \tilde{y}_o^c :

$$\rho_{\tilde{x}\tilde{y}} \simeq \langle \tilde{x}_r^c \tilde{y}_o^c \rangle \quad (7)$$

220 where $\langle \cdot \rangle$ indicates the average.

The normalized predictand is now sampled as a linear combination of normalized predictor and an extra random variable r as follows:

$$\tilde{y} = \alpha \tilde{x} + \beta \tilde{r} \quad (8)$$

with

$$225 \quad \tilde{x} = q_N \left(p_{D_{x_r^c}}(x) \right) \quad (9)$$

the predictor normalized according to the estimated probability distribution function calculated above, and the predictor x can be either the upscaled reanalysis for validation x_r^v or the bias-adjusted climate models x_m for the application (predictor from climate projections). One can obtain \tilde{y} correlating with \tilde{x} with a correlation coefficient ρ by setting $\alpha = \rho$ and $\beta = \sqrt{1 - \alpha^2}$ (see appendix A), hence

$$230 \quad \tilde{y} = \rho \tilde{x} + \sqrt{1 - \rho^2} \tilde{r} \quad (10)$$

Finally, one can calculate the modeled predictand by inverting Eq. 6:

$$y = q_{D_{y_o^c}} \left(p_N(\tilde{y}) \right) \quad (11)$$

where p_N is the cumulative distribution function of the Gaussian normal distribution and $q_{D_{x^c}}$ is again the quantile function of the calibration dataset.

235 In contrast to Switanek et al. (2022), the current method for (de)normalizing the variables is non-parametric, hence doesn't take any assumptions on particular statistical distributions (eg., gamma distribution for precipitation or Weibull distribution for wind speed), hence any climate variable can be generated including precipitation, wind speed, temperature and heat stress, as long as the observation record is long enough to represent their density distributions.

2.2.2 Prescaling of predictand for precipitation and wind speed

240 ~~The statistical relation described above is improved for precipitation and wind speed by rescaling the predictand with the predictor. Therefore, the equation 6 is modified as follows:-~~

$$\underline{\tilde{y}_o^c = q_N \left(p_{D_{y_o^c}}(y_o^c) \right)}$$

with-

$$y_o'^c = y_o^c / (c + x_r)$$

245 where c is a scale parameter that is set to 30 mm/day for precipitation and 2 m/s for wind speed, which avoids division by zero for $x_r = 0$. The modelled predictand given by Eq. 11 is then modified as follows:-

$$\underline{y'} \equiv \frac{q_{D_{y_o'^c}}(p_N(\tilde{y}))}{\underline{y'} \cdot (c + x)}$$

250 For values of $x \gg c$, the correlated random sampling by Eqns. 10 and ?? is performed for the perturbation $y' \simeq y/x$ of the point-scale predictand from the coarse-scale predictor x instead of the point-scale predictand y itself. This leads to more robust results because it can be assumed that the expected value (ie., averaging over a large number of days for a particular value of x) of y approximately scales with the predictor x , or in other words, the expected value of y' is assumed to be a constant (and approximately 1). For values of $x \ll c$, the statistical model becomes equivalent to the original one by Eq. 11. This modification leads improved agreement of the modelled precipitation and wind speed with observed high extreme return
255 periods of 1 to 10 years, as will be shown in the results section.-

2.2.2 Detrending of climate projections

In the description above, a statistical relationship (ie., correlation through their normalized time series) is established between the predictor variable at the coarse scales and the predictand variable at the point scale. However especially in the case of temperature variables, distribution values in future climate at coarse scales exceed the range of the distribution of the historical
260 period because of global warming. This leads to a cutoff for highest extremes, since they are all mapped to the highest observed value from the historical record. To overcome this problem and still represent the high extremes in future global warming, the predictor from the projections are detrended by dividing by the 30-year mean of the given month and multiplying it by its mean of the reference period 1986–2015 (in order to avoid division by zero in this operation and the problems arising from negative values, temperature variables are expressed in Kelvin). This is done before normalizing the predictor in Eq. 9. Afterwards, the
265 sampled predictand from Eq. ??-11 is retrended by dividing again by the predictor mean of the reference period and multiplying with the 30-year mean. By performing these operations, the statistical relationship (correlation) between the coarse-scale and point-scale variables is now considered relative to their long-term averages, and it's assumed that it doesn't change under climate change. [Correlation stability and stationarity of the statistical relationship under climate change and its effect on model results are assessed for daily precipitation in Sikasso in Text S1, Tab. S2 and Fig. S3. It was found that the spread across
270 simulations using 6 different calibration sets \(even and odd years; 1961–2023, 1961–1990, 1991–2020\) is of the same order as observation sampling sets \(even vs odd years; 1961–1990 vs 1991–2020\). The underlying correlation coefficient where also found to be stable. This indicates that statistical relationships are robust under climate change. Nevertheless, uncertainties arising from different calibration sets needs to be kept in mind in climate change assessments.](#)

2.3 bias-adjustment of climate projections

275 In the previous section, a statistical model is constructed and calibrated on on historical reanalysis data used as predictor variables and the point-scale weather observations. To apply the statistical model to the climate projections as predictor, the consistency with the historical reanalysis needs to be maximized. Therefore, the raw CMIP6 ensemble climate model data is bias-adjusted against the historical reanalysis data with quantile delta-mapping. The latter is based on the quantile-delta mapping described in Cannon et al. (2015). For the sake of clarity, we elaborate the QDM correction below, in which we also
 280 follow the notation of Maraun (2016).

The model bias $B(p)$ is calculated as a function of the model's cumulative distribution function p (or 'p'robabilities corresponding to a given model quantile) with respect to the climate reconstruction x_r and the 'b'iasd climate model output x_b^h in the historical overlapping timeframe (h) from 1961-2022:

$$B(p) = q_{D_{x_b}^h}(p) - q_{D_{x_r}^h}(p) \quad (12)$$

285 This bias function determines the difference between the quantiles $q_{D_{x_b}^h}(p)$ of the climate model output and those of the climate reconstruction $q_{D_{x_r}^h}(p)$. For the historical period, the time series bias is then calculated as follows:

$$B_{QDM}^h(x_{b,i}) = B\left(p_{D_{x_b}^h}(x_{b,i})\right) \quad (13)$$

from which can find the final bias-adjusted time series by subtracting the bias in the historical timeframe:

$$x_{m,i} = x_{b,i} - B_{QDM}^h(x_{b,i}) \quad (14)$$

290 for i a day in the historical time frame h (1991-2020). To create future bias-adjusted time series, quantile delta mapping assumes that the bias associated to each percentile $B(p)$ (while changing the timeframe for calculating each percentile) is invariant under climate change. We get the following bias-adjustment function for the future timeframe p :

$$B_{QDM}^p(x_{b,i}) = B\left(p_{D_{x_b}^p}(x_{b,i})\right) \quad (15)$$

and:

295 $x_{m,i} = x_{b,i} - B_{QDM}^p(x_{b,i}) \quad (16)$

for i a day in the projected timeframe p with a 30-year timespan (eg., 2071-2100). For a given day i , the timeframe for projection (p) is chosen in such a way that the day i is in the center of the timeframe. In practice, timeframes and the corresponding bias-adjustment function B_{QDM}^p are created in steps of five years, which is a compromise between computational cost and a smooth transition of bias-adjustment between the timeframes.

300 It can be verified that the cumulative density distribution of the climate model $\left(p_{D_{x_b}^p}\right)$ is re-evaluated in the future period, which ensures that climate change signals of the quantile distribution is conserved. This is because the quantiles in the

respective historical and future cumulative distributions p undergo the same bias-adjustment $B(p)$, hence their difference remains conserved. This can be shown formerly by evaluating bias-adjustment (Eq. 16) on the quantiles $q_{D_{x_b}^p}(p)$ in the future timeframe:

$$\begin{aligned}
 305 \quad q_{D_{x_m}^p}(p) &= q_{D_{x_b}^p}(p) - B_{QDM}(q_{D_{x_b}^p}(p)) \\
 &= q_{D_{x_b}^p}(p) - B(p_{D_{x_b}^p}(q_{D_{x_b}^p}(p))) \\
 &= q_{D_{x_b}^p}(p) - B(p)
 \end{aligned} \tag{17}$$

and the bias-adjustment (Eq. 14) on the quantiles in the historical timeframe $q_{D_{x_b}^h}(p)$:

$$\begin{aligned}
 q_{D_{x_m}^h}(p) &= q_{D_x^h}(p) - B_{QDM}(q_{D_x^h}(p)) \\
 &= q_{D_x^h}(p) - B(p_{D_x^h}(q_{D_x^h}(p))) \\
 &= q_{D_x^h}(p) - B(p)
 \end{aligned} \tag{18}$$

Subtracting the two equations shows that the climate change signal on the quantiles are invariant for the QDM correction:

$$q_{D_{x_m}^p}(p) - q_{D_{x_m}^h}(p) = q_{D_{x_b}^p}(p) - q_{D_{x_b}^h}(p) \tag{19}$$

310 bias-adjustment of climate models are calculated over continental or global grids at once. To lower the computational cost and to obtain a more smooth bias-adjustment function, the bias is calculated for discrete values l of the cumulative distribution function p_l . As such, the bias function can be approximated as follows:

$$B(p) \simeq \alpha(p)B(p_l) + (1 - \alpha(p))B(p_{l+1}), \text{ for } l \text{ where } p_l \leq p < p_{l+1} \tag{20}$$

with

$$315 \quad \alpha = (q - q_l)/(q_{i+l} - q_l) \tag{21}$$

As for the statistical model above, we consider equidistant discrete probability levels (in this case $n = 15$) for temperature-like and heat-stress metrics (see Eq. 1), and an exponential profile for precipitation and wind speed (see Eq. 2) to represent their big tails. The time series $p_{D_{x_b}^p}(x_b)$ are calculated from the model time series for each 5-year window and for each month separately, from which the bias at each timestep $B(p_{D_{x_b}^p}(x_b))$ can be evaluated using Eq. 20 and subtracted from the model output in Eq. 16.

As mentioned earlier, the climate models are generally coarser than the reference historical data, so the climate reconstruction (reanalysis data) is spatially aggregated to the coarser climate model grid. The bias may depend on the time of the year, hence the bias-adjustment B is calculated for each month of the year separately. Hereby, the bias of each month is calculated by considering the month before, the month itself and the month after.

325 2.4 Validation of variability and extremes

Downscaled time series (y^v) are generated by applying the statistical model on the coarse-scale reanalysis predictor (x_r^v) over a climatological historical period (1961-2023), and validated with station observations y_o^v . The validation focusses on the

representation of the variability and extreme events. Therefore, their annual cycles, density distributions and return levels as a function of return period for the different datasets and timeframes are compared with the observations. For the annual cycles, the mean, median and 10th-to-90th percentile range of each day throughout the year is provided for highlighting the combined seasonality and variability of each variable throughout the year. Magnitude of extreme events are highlighted with density distributions, whereas occurrence of extreme events are highlighted with return levels as a function of return periods. Density distributions for all variables are weighted so that their integral sums to unity. This is except for precipitation for which the density is weighted with the average precipitation, in such a way that the integral sums to the average annual precipitation. The advantage of such a weighted density is that extreme precipitation values become more visible on the plots. Herewith, we report the overlap area between the different density distributions (Devis et al., 2013; Perkins et al., 2007):

$$S = \int_{-\text{inf}}^{+\text{inf}} \min(PDF(y_v), PDF^o(y_v)) dy_v \quad (22)$$

While the Perkins distribution overlap score provides an integrated measure of similarity between two probability density functions, it is inherently insensitive to compensating errors and offers limited insight into discrepancies in the distribution tails. Therefore, we complement the Perkins score with additional distribution diagnostics, including quantile-based loss metrics, Kolmogorov–Smirnov statistics, and explicit indicators of extremes such as high percentiles, return levels (value as a function of return period, see below), and dry-day frequencies. These metrics are represented in a table and provide a comprehensive and objective assessment of both the central tendencies and the extreme behavior of the downscaled variables.

The return period, also known as a recurrence interval or repeat interval, is a commonly used metric in extreme (weather) analysis that evaluates the average time between subsequent extreme events of a certain magnitude or return level. To show high extremes for daily maximum temperature, precipitation, wind speed and heat stress, return levels are determined that is exceeded as a function of the return period. To quantify low extremes for daily minimum temperature, we determine the return level that is subceeded (ie., becomes lower than the return level) as a function of the return period. In theory, the model can be validated for any return period, but this depends on the length of the observation record. To only retain ~~retain~~-statistically relevant results for the current validation period 1961-2023, we only show return periods of up to ~~5-10~~ years which lead to an averaging over least ~~6 validation samples per return~~ 3 validation samples for that maximum period over the 63 year period, ie., 63 years divided by ~~five-10~~ years, and divided by two since only half of the measurements are used for the validation.

All variables are extracted and processed directly from the respective data sources. This is except for heat stress. For the latter, we use the heat-stress temperature as proposed by Wouters et al. (2022):

$$T_s = T_w + 4.5 \left(1 - \left[\frac{RH}{100} \right]^2 \right) \quad (23)$$

which adjusts the wet-bulb temperature T_w to reflect the association with heatwave mortality under different levels of relative humidity (RH) and temperature.

2.5 Climate-change assessment of variability and extremes

For the climate change assessment, downscaled time series (y) are generated by applying the statistical model on multiple bias-corrected climate models as coarse-scale predictor (x_m). We show 10th-to-90th percentile ensemble spread of density distributions and return levels over the 9 different models member from CMIP6 (see Tab. 1). To demonstrate EXSoDOS, we show results from the Shared Socio-economic Pathway for fossil fueled development with 8.5Wm^{-1} radiative forcing (SSP585), see IPCC (2021). The user can change to other SSP scenarios to assess future climate uncertainties. Results are shown for three 30-year time frames, one for the future (2071-2100), and two past periods (1961-1990 and 1991-2020). For the two past periods, results from observations are also added to compare the past shifts of modeled weather extremes to the observations, and then assess whether and how these shifts would continue in the future.

2.6 Use cases

2.6.1 ~~Use cases based on NOAA data archive~~

To show the general utility of EXSoDOS according to particular local challenges of climate change, we apply and evaluate EXSoDOS for 5 cases across the globe. For each of these locations we show the results of one variable according to a particular local challenge of climate change (for their locations, see Fig. 2). The selected use cases are intended to demonstrate the methodology and validation workflow rather than to provide an exhaustive global evaluation. For any new application, local validation remains essential because data quality and predictor–predictand relationships are location dependent.

2.6.1 Use cases based on NOAA data archive

For Uccle (Belgium) in Europe, one of the challenges facing climate change is that many native vegetation (crop) species require the occurrence of freezing temperatures to ~~be able to blossom in spring~~ ensure proper dormancy release and phenological development of many perennial plant species (eg., Brunner et al., 2014; IPCC, 2021), hence minimum daily temperature is demonstrated for this location. For Spangdahlem (Germany) also in Europe, ~~the power yield of a cluster of wind turbines depends on the distribution of wind~~ wind speed variability and extremes directly affect wind energy yield and structural loads on turbines (Tobin et al., 2016; Devis et al., 2013; Pryor and Barthelmie, 2010), hence for this location we show results for daily wind speed. The Middle-East and the region of South-Causassus around the Caspian Sea contains one of the hot spots of high temperature of the world ~~, see for example Fig. 1D of Wouters et al. (2022)~~ (eg., Fig. 1D of Wouters et al., 2022; Perkins-Kirkpatrick and Le. Therefore, we assess daily maximum temperature of climate change on extremely high temperature For Baku (Azerbaijan) where COP29 took place in November '24. Finally, ~~Large parts of India are exposed~~ Coastal India is highly vulnerable to extreme heat stress due ~~a combination to the combined effects~~ of high temperature and humidity, ~~see for example Fig 1C of Wouters et al. (2022)~~ with documented impacts on mortality and labour productivity (eg., Fig. 1C of Wouters et al., 2022; Im et al., 2017; Raymond et al., 2020). We exemplify EXSoDOS by showing results of heat stress temperature Wouters et al. (2022) for Puri located along the east coast of India. For the use case locations above,

measurement data are extracted from the NOAA archive (see section 2.1), highlighting the applicability of EXSoDOS for any
390 location with measurements in this archive.

2.6.2 Use case based on local data by MALI-METEO

Mali, located in West Africa, is characterized by Sahelian and semi-arid climate. In Mali, rainfall is concentrated in the south
part, with more intense and regular rainfall, while it decreases as it moves northwards, where it becomes scarce. The data used
in this study are ground observations of daily rainfall times series from the MALI-METEO database for the period 1961 to
395 2023 from the Sikasso synoptic station (Latitude: 11°19'N, Longitude: 05°41'W, Altitude: 415.05m).

As rainfall is a dichotomous variable, its inter-annual variability in time and space is difficult to assess. Rainfall is a key
element in Mali, and particularly in Sikasso, playing a crucial role in agriculture, which is the main source of subsistence
for the local population. However, its variability has a direct impact on crop yields and the availability of water resources
and exacerbates the vulnerability of communities to climatic hazards. [Rainfall variability and the increasing relevance of
400 heavy-rainfall extremes across the Sahel have been widely documented, with important implications for agriculture and flood
risk \(Sanogo et al., 2015; Panthou et al., 2014\)](#). Daily times series of rainfall in Sikasso station clearly show that the rainfall
pattern in this locality is unimodal, with a more intense character in August. In fact, the absolute record for 24h cumulative
rainfall in Sikasso is 166.1mm, observed on August 12, 1963. Sikasso is a region where large quantities of rain are observed
every year. In 2023, flooding caused the loss of 123.5 hectares of agricultural land and affected 4,023 people, including 2
405 deaths (Source: Annual report of the Direction Générale de la Protection Civile, 2023). However, in 2024, the impact of
flooding increased considerably, with 7,017 hectares of agricultural land lost and 4,231 people affected (Source: Annual report
of the Direction Générale de la Protection Civile, 2024).

To demonstrate the applicability of EXSoDOS on locally supplied observations as an alternative to the NOAA database,
observations from MALI-METEO in Sikasso were used. Also, this use case was experimented during an interactive workshop
410 on weather and climate data (“ L’engagement des parties prenantes nationales dans la formulation de politiques agricoles
précises pour des stratégies d’adaptation au changement climatique au Mali”) held from October 21 to 25, 2024 in Bamako
(Mali) with an online Jupyter Notebook, see Carter (2024).

3 Results

3.1 Validation of downscaling reanalysis

415 We first validate the statistical downscaling algorithm by applying it on the validation predictor x_r^v (upscaled ERA5) and
compare its output with the observed predictand y_o^v . for the historical period 1961–2023. This is done for the different mete-
orological quantities in different study areas described in section 2.6. Annual cycles, distributions and return period for each
of the case studies are shown in Figs. 3, 4 and 5 [and Tab. 2](#). Overall, ERA5 matches well the average annual cycles of the
observations over multiple years for each variable (Fig. 3). However, the annual cycles of ERA5 underrepresent the variability

420 over years, as highlighted by the 10-to-90 percentile range of each day of the year in Fig. 3. ~~This is especially the case. This is also clear from underestimated values in standard deviation, 95th percentiles and 1-year return levels (Tab. 2), especially for precipitation in Sikasso (Mali), for which both dry days and extreme high precipitation are underrepresented where also the number of dry days is underestimated.~~ The underrepresentation is more subtle for the temperature variables, such as daily minimum temperature for Uccle in Belgium and the daily maximum temperature for Baku (Azerbaijan), and for heat stress
425 temperature for Puri (India). The underrepresentation of extremes is more pronounced when ERA5 is upscaled to a coarser resolution of global climate models (1° resolution), as extreme values are averaged ~~out~~ over larger grid sizes.

In contrast, all results for ERA5 downscaled to the station level match better the variability of the annual cycle (Fig. 3) and
distribution statistics (2) compared to the original ERA5 and distribution statistics. Herewith, also the overall bias is removed
for each variable. The improved representation of extremes by downscaling ~~becomes more clear when analyzing density
430 distributions. 4 is also clear from density distributions (Fig. 4) and return periods 5. (Fig. 5).~~ Distributions of precipitation
after downscaling are also better matching the observations (thin dashed line lines versus thick dashed linelines in Fig. 4) than
without downscaling of ERA5 (thick line versus thin dashed line), since they get shifted to more high extremes, while also the
number of dry days becomes larger and better matching observations (Fig. 4 Tab. 2). At the same time, the downscaling also
reduces the overall bias in annual average daily precipitation for Sikasso ~~where the downscaled value (1210 mm) is closer to
435 with a downscaled value of 3.3 mm matching the observed value (1205 mm) than, whereas it is underestimated by the~~ ERA5
value (842-2.9 mm). Other variables, particularly daily minimum and maximum temperature, wind speed and heat stress, also
show a better match with the observed distribution after downscaling with larger spread (Fig. 4). ~~The Perkins distribution
overlap scores (Eq. 22) confirms. Besides better statistics on mean and variability,~~ the better match of the distributions after
downscaling ~~, showing higher scores for each variable is confirmed with better score values for Perkins overlap score, the
440 quantile loss difference and the Kolmogorov-Smirnov statistic,~~ see Tab. 2.

All downscaled variables (including daily minimum temperature) show more extreme levels for the high return periods
compared to original ERA5 data, for which they approximate the return levels of the observations much better (Fig. 5). ~~We
also highlight that pre-scaling of the predictand with predictor (see Eq. ?? and ??) leads to more consistency of the model results
with extreme observations than without the pre-scaling (see Eq. 6 and 11) especially for precipitation, compare bottom panel of
445 Fig. 5 compared to Fig. 1 of supplementary material~~

Results and scores with standard quantile mapping (ERA5 QM) are added as comparison next to the full correlated sampling
(ERA5 DS), see Figs. 4 and 5, and Tab. 2. While the results from the standard quantile mapping (ERA5 QM) show similar
scores and distributions to the full correlated sampling (ERA5 DS), the latter one is preferred, since the former could lead to
inflation of trends in extremes (Maraun, 2013). Inflation for precipitation in Sikasso is also illustrated in the supporting text S2
450 with the help of Figs. S3, S4 and S5, and Tab. S3.

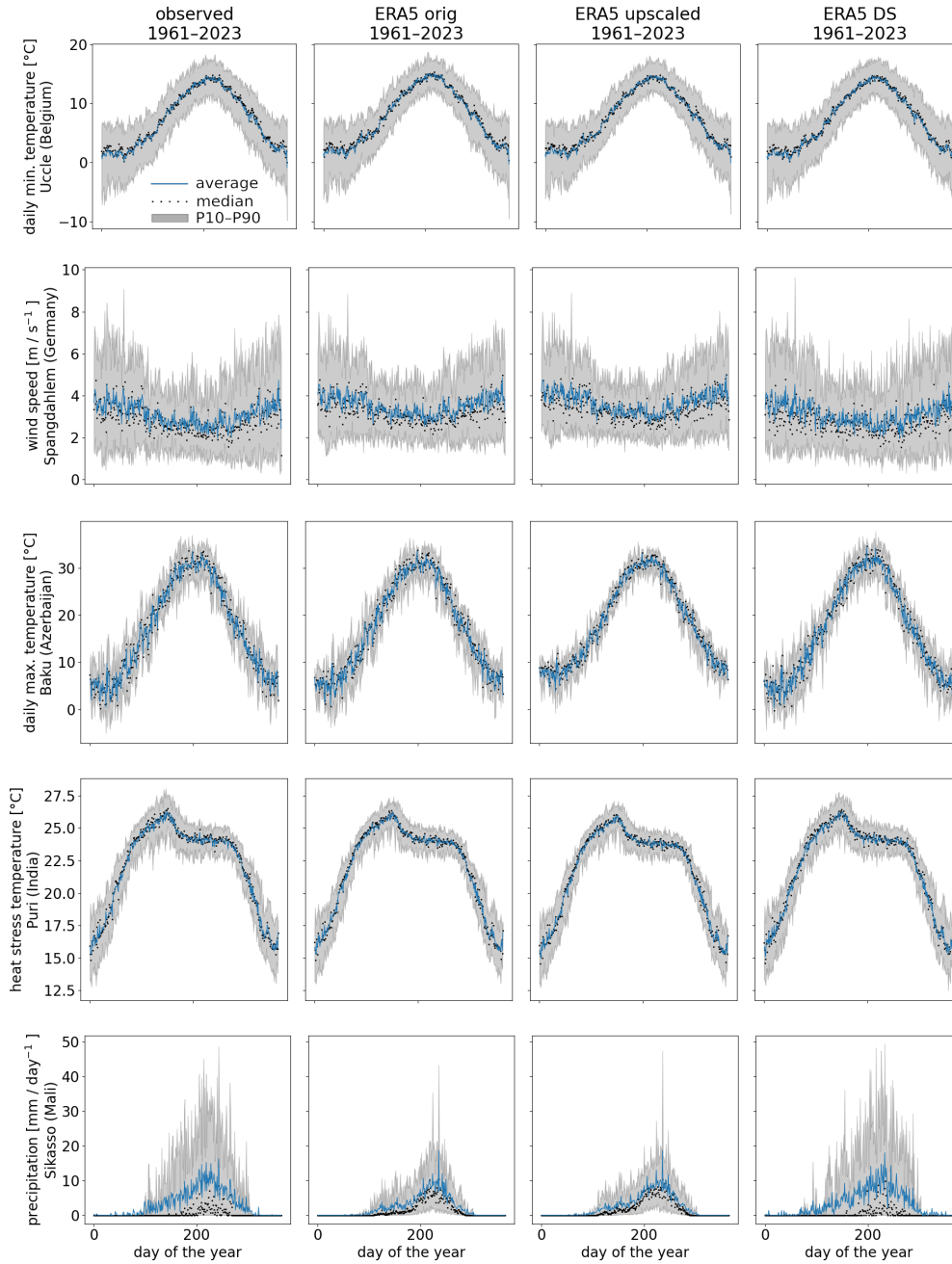


Figure 3. Annual cycles from observations for the period (1961-2023), ERA5 on its original grid (ERA5 orig), ERA5 upscaled to the grid of a climate model (ERA5 upscaled), and ERA5 downsampled with the statistical model (ERA5 DS). The blue lines indicate the average, whereas the black dots show the median of each day of the year over the multiple years. The grey bars indicate the range between the 10th and 90th percentile values over the different years. We show daily minimum temperature for Uccle in Belgium (first row), daily mean wind speed for Dahlem in Germany (second row), daily maximum temperature for Uccle in Belgium (third row), heat stress temperature for Puri in India (fourth row), and precipitation for Sikasso in Mali (last row).

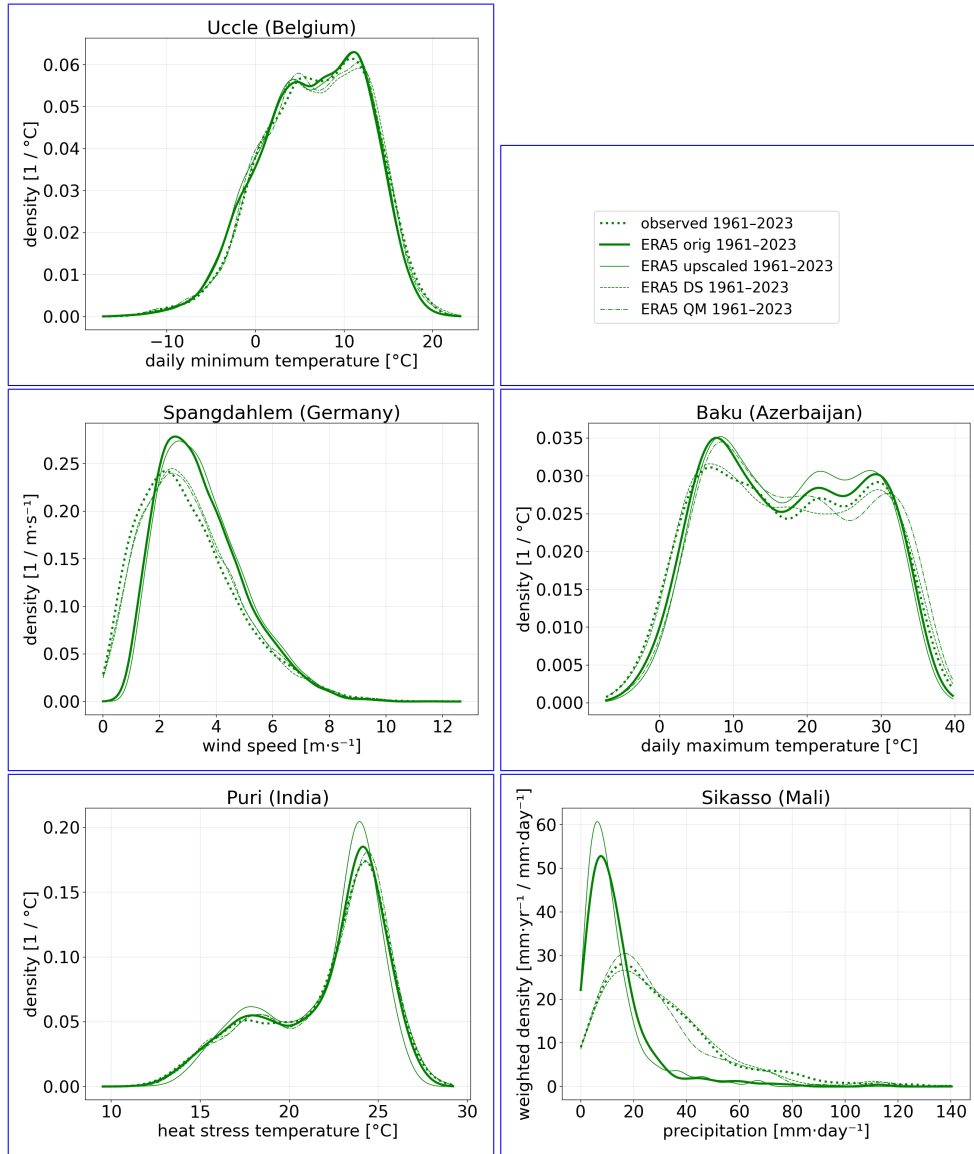


Figure 4. Density distribution from daily observations for the period (1961-2023), ERA5 on its original grid (ERA5 orig), ERA5 upscaled to the grid of a climate model (ERA5 upscaled), and ERA5 downscaled with the statistical model (ERA5 DS). We show daily minimum temperature for Uccle in Belgium (upper left panel), daily mean wind speed for Dahlem in Germany (upper right center left panel), daily maximum temperature for Uccle-Baku in Belgium-Azerbaijan (center left-right panel), heat stress temperature for Puri in India (center-right lower left panel), and precipitation for Sikasso in Mali (lower right panel). For precipitation, the annual number of dry days and the annual precipitation are provided, and for heat stress temperature the annual number of deadly heat days (heat stress temperature \rightarrow 27°C)

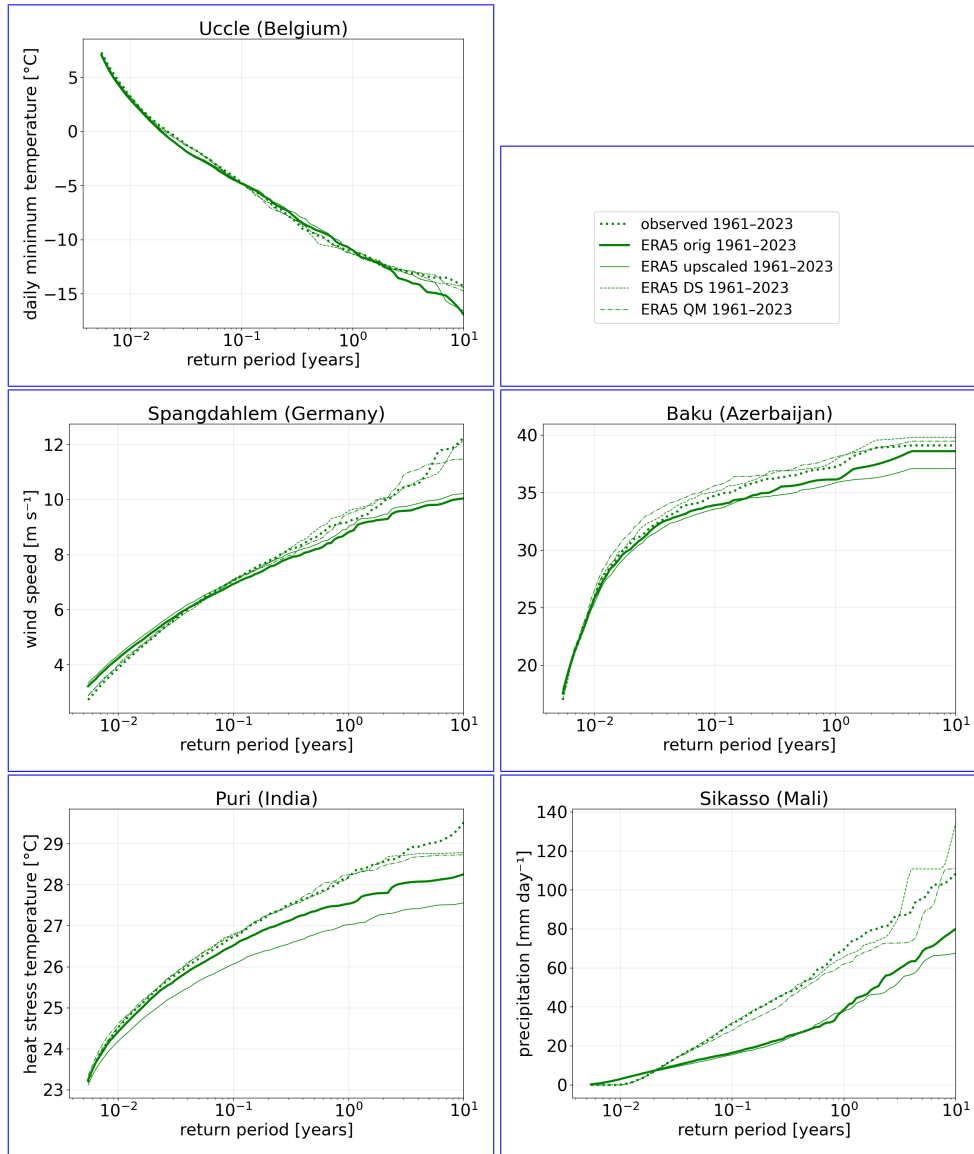


Figure 5. Idem as Fig. 4, but for return levels.

	Mean	Std	P95	1y return	Perkins	QLoss (ratio)	KS stat (p)		
<i>Daily minimum temperature (Uccle, Belgium)</i>									
observed	7.0 (+0.0)	4.9 (+0.0)	-2.8 (+0.0)	-11.0 (+0.0)	1.000	0.000 (1.000)	0.000 (1.000)		
ERA5 orig	6.6 (-0.4)	4.8 (-0.0)	-3.2 (-0.4)	-10.9 (+0.1)	0.970	0.006 (1.011)	0.027 (0.00)		
ERA5 upscaled	6.5 (-0.4)	4.8 (-0.0)	-3.1 (-0.3)	-10.8 (+0.2)	0.965	0.006 (1.012)	0.034 (0.00)		
ERA5 DS	6.9 (-0.0)	4.9 (+0.1)	-2.9 (+0.1)	-11.2 (-0.2)	0.980	0.000 (1.000)	0.018 (0.07)		
ERA5 QM	6.8 (-0.1)	4.9 (+0.0)	-2.8 (-0.0)	-11.3 (-0.2)	0.982	0.001 (1.002)	0.020 (0.04)		
<i>Wind speed (Spangdahlem, Germany)</i>									
observed	3.0 (+0.0)	1.4 (+0.0)	6.4 (+0.0)	9.2 (+0.0)	1.000	0.000 (1.000)	0.000 (1.000)		
ERA5 orig	3.5 (+0.4)	1.2 (-0.2)	6.4 (-0.0)	8.8 (-0.4)	0.846	0.000 (1.000)	0.159 (0.00)		
ERA5 upscaled	3.6 (+0.6)	1.2 (-0.2)	6.5 (+0.1)	8.9 (-0.2)	0.825	0.000 (1.001)	0.183 (0.00)		
ERA5 DS	3.1 (+0.1)	1.4 (-0.0)	6.3 (+0.1)	9.5 (+0.4)	0.959	0.000 (1.001)	0.043 (0.00)		
ERA5 QM	3.1 (+0.1)	1.4 (-0.0)	6.4 (-0.0)	9.5 (+0.3)	0.962	0.000 (1.000)	0.041 (0.00)		
<i>Daily maximum temperature (Baku, Azerbaijan)</i>									
observed	17.3 (+0.0)	9.3 (+0.0)	33.8 (+0.0)	37.1 (+0.0)	1.000	0.000 (1.000)	0.000 (1.000)		
ERA5 orig	17.6 (+0.2)	8.8 (-0.4)	33.1 (-0.7)	36.1 (-1.0)	0.962	0.004 (1.005)	0.041 (0.14)		
ERA5 upscaled	17.7 (+0.4)	8.5 (-0.8)	32.7 (-1.1)	35.6 (-1.5)	0.934	0.012 (1.013)	0.057 (0.01)		
ERA5 DS	17.4 (+0.1)	9.3 (+0.1)	34.1 (+0.3)	37.3 (+0.2)	0.981	0.002 (1.002)	0.021 (0.87)		
ERA5 QM	18.3 (+0.9)	9.0 (-0.3)	34.5 (+0.7)	37.8 (+0.7)	0.940	0.008 (1.009)	0.057 (0.01)		
<i>Heat stress temperature (Puri, India)</i>									
observed	22.1 (+0.0)	2.9 (+0.0)	26.3 (+0.0)	28.2 (+0.0)	Heat days / yr	7.1 (+0.0)	1.000	0.000 (1.000)	0.000 (1.000)
ERA5 orig	22.0 (-0.1)	2.9 (-0.0)	26.1 (-0.2)	27.5 (-0.6)	4.3 (-2.8)	0.974	0.001 (1.004)	0.022 (0.03)	
ERA5 upscaled	21.9 (-0.1)	2.7 (-0.2)	25.7 (-0.6)	27.0 (-1.1)	1.1 (-6.0)	0.920	0.012 (1.049)	0.072 (0.00)	
ERA5 DS	22.1 (+0.0)	2.9 (-0.0)	26.3 (+0.1)	28.1 (-0.0)	7.4 (+0.3)	0.984	0.000 (1.001)	0.011 (0.69)	
ERA5 QM	22.2 (+0.1)	2.8 (-0.0)	26.3 (+0.1)	28.2 (+0.1)	7.3 (+0.2)	0.972	0.000 (1.001)	0.027 (0.00)	
<i>Precipitation (Sikasso, Mali)</i>									
observed	3.3 (+0.0)	5.2 (+0.0)	21.4 (+0.0)	69.7 (+0.0)	Dry days / yr	281.6 (+0.0)	1.000	0.000 (1.000)	0.000 (1.000)
ERA5 orig	2.9 (-0.4)	3.5 (-1.7)	13.4 (-8.0)	38.4 (-31.3)	218.2 (-63.5)	0.530	0.146 (1.086)	0.355 (0.00)	
ERA5 upscaled	2.8 (-0.5)	3.3 (-1.9)	12.1 (-9.3)	37.2 (-32.5)	206.1 (-75.5)	0.487	0.200 (1.118)	0.457 (0.00)	
ERA5 DS	3.3 (-0.0)	5.2 (+0.0)	21.8 (+0.4)	64.8 (-4.8)	283.9 (+2.3)	0.953	0.000 (1.000)	0.010 (0.62)	
ERA5 QM	3.1 (-0.2)	4.9 (-0.3)	20.5 (-0.9)	61.8 (-7.8)	282.3 (+0.7)	0.905	0.001 (1.001)	0.007 (0.92)	

Table 2. Validation metrics for precipitation distributions from observations (Observed), original ERA5 (ERA5 orig), upscaled ERA5 (ERA5 upscaled), fully correlated stochastic downscaling (ERA5 DS), and quantile-mapping-only downscaling (ERA5 QM). For the mean, standard deviation (Std), 95th percentile (P95), 1-year return level (1y return), and annual number of dry days (in case of precipitation), absolute values are reported with deviations from observations in brackets. We further report the Perkins distribution-overlap scores score (Eq. 22) for different variables, the quantile loss difference (with ratio in brackets), and location the Kolmogorov–Smirnov statistic with the corresponding p-value.

Perkins distribution-overlap scores (Eq. 22) for the downscaled ERA5 (ERA5 DS), original ERA5 (ERA5 orig), and upscaled ERA5 (ERA5 upscaled).

3.2 Assessment of extremes under climate change

In this section, assessment is demonstrated for each of the variables for the different areas under the SSP585 scenario by comparing results of a 30-year timeframe by the end of the century (2071-2100) and of two 30-year timeframes of in the past (1961-1990 and 1991-2020) and one in the future under the SSP585 scenario (2071-2100), see Figs. 6 and 7, 7 and Tab. 3.

455 Overall, there is a tendency towards higher (extreme) values for the original climate projections (CMIP6) (Fig. 6) towards the future (green versus see Figs. 6 and 7, and Tab. 3) under past climate change (thin green line versus thin grey lines). This is except for daily wind speed and for minimum temperature, the low values become less extreme. For daily and maximum

temperature, precipitation and heat stress, the variables bias-corrected and downscaled from CMIP6 (CMIP6_BC_DS) could better capture the tendencies between the two past periods than the original CMIP6 (thin green line versus thin grey line). A shift towards more extremes is already present in the original, for which values remain stable. However, the CMIP6 data (thin green line versus thin grey line) but the distributions have a substantial bias (distributions have substantial biases (thin full lines versus thick dashed lines), just like the ERA5 values (both the original and upsealed to the climate model) as discussed in the previous section. Especially, the extremes in CMIP6 have an overall bias as discussed in the previous section, and this is also found for the shift under climate change. Thanks to the bias-correction and downscaling, the observed values and their shift under climate change are now captured by biascorrected and downscaled distribution (properties) and return levels are now better captured by CMIP6 (thick green line versus thick grey line BC_DS (thin dashed lines versus thick dashed line) for the two past period than for CMIP6 (thin full lines versus thick dashed lines), see Fig. 6. Particularly for precipitation (Sikasso), an overall underestimation (Figs. 6 and 7, and Tab 3. Hereby, CMIP6_BC_DS alleviates overall biases in the annual averages found for CMIP6 for each variable, see Tab. 3. Additionally, standard deviations, P95 values, 1-year return levels, and number of dry days and high extreme values are found for CMIP6 (see numbers in lower panel of Fig. 6), as (for Sikasso) and heat days (for Puri) generally tend to be higher after downscaling, except for 1-year return values for daily minimum temperature (Belgium) for which we get lower (hence also more extreme) return levels, see Tab. 3. They lead to a better representation of all these distribution properties, which was also found for ERA5 (original and upsealed) as shown in the previous section. The shift of precipitation distribution by CMIP6 also tends to be underestimated (thin green lines versus thin grey lines in Fig. 6), as well as their return values and their shifts (Fig. 7, except for standard deviation for daily maximum temperature (Baku). The underestimation in absolute values and their shifts is largely alleviated by employing the downscaling on bias-corrected CMIP6 (past tendencies of standard deviation, P95 and 1-year return values tend to be slightly lower for CMIP6_BC_DS); see Figs. 6 and 7. This is the case for all shifts in return periods, except for those of precipitation between two and six years. For the latter, there is a slight decrease in the observations while the downscaled values still show an increase. Daily minimum temperature (Uccle), daily maximum temperature (Baku) and heat stress temperature (Puri) all tend to be overestimated by variables compared to CMIP6. For daily precipitation in Sikasso, this lead to a better representation of the shifts: the tendency of the median 1-year return level between the two 30 year time frames (from 1961-1990 to 1991-2020) is lower for CMIP6, and these biases are all largely alleviated by BC_DS (from 65.58 to 67.75mm) than for the original CMIP6 (from 40.53 to 43.89mm) which better matches the observed tendency from (from 67.60mm/day to 69.70mm). The tendency of the P95 value of CMIP6_BC_DS, see Figs. 6 and 7. After this alleviation, it becomes clear that historical tendencies are well represented (from 21.02 to 22.22mm) is still higher than for the original CMIP6 (from 10.97 to 11.20mm) but that also leads to better match with the observed tendency (from 20.20mm to 22.49mm). For wind speed for Spangdahlem however, the overall distributions of the observations are captured well, but the shift in the observations towards higher wind speeds is not predicted, since large-scale wind speed distribution (CMIP6) doesn't show a shift either.

The tendencies of having higher extremes (except for wind speed) pull through to the end of the 21st century under the SSP585 scenario, see 6 and 7 Figs. 6, 7 and Tab. 3. While original CMIP6 ensemble precipitation already indicates a tendency towards more extremes in the future, the (red full lines versus green and grey full lines), the tendency of downscaled ensemble

(CMIP6_BS_DS) ~~substantially from the original CMIP6 ensemble (red dashed lines versus green and grey dashed lines)~~ differs substantially both in absolute and relative numbers. For ~~precipitation return periods between 0.1 and 1 year~~ 1-year return level of daily precipitation, for example, an ~~average increase from 17.1/day to 31.0/day~~ increase from 40.53mm (for 1961–1990) to 78.19mm (for 2071–2100) is found for the median ensemble, hence an increase of ~~13.9mm/day–37.6mm~~ ($\sim 8176\%$), is found for the raw CMIP6. At the same time, an average increase from ~~$\sim 43.3\text{mm/day}$ – 65.58mm~~ to ~~$\sim 57.2\text{mm/day}$ – 79.29mm~~ is found for CMIP6_BS_DS, hence an increase of ~~$\sim 14.4\text{mm/day}$ – 13.7mm~~ ($\sim 3320\%$). ~~A discrepancy also exists in terms of return periods: an event with a return period of 1 year according to CMIP6 ($\geq 32.0\text{mm/day}$) would have a return period 0.247 ($\sim 1/4$) year in a future climate, whereas an event with a return period of 1 year according to CMIP6_BC_DS ($\geq 69.1\text{mm/day}$) would have a return period of 0.349 ($\sim 1/3$) year.~~ Future temperature, heat stress and wind speed are also affected after downscaling. For heat stress, the (change in) number of days with extreme heat is higher after downscaling (~~from 5.73 to 92.87 days per year~~) than from the original CMIP6 ensemble ~~–see Fig. 7 median (from 4.50 to 81.80 days per year), see Tab. 3.~~ These results highlight that different types of climate-change assessments using downscaled GCM output ~~leads lead~~ to different results than those using original GCM output.

4 Conclusions

A downscaling method EXSoDOS is developed for the DOwnscaling of local weather EXtremes Shifts under global warming, including daily minimum and maximum temperature, precipitation, heat stress and wind speed. The model employs ground-based measurements, historical reanalysis climate reconstruction and ensemble climate projections as input. Stochastic modelling is performed by normalizing the local measurements and coarse climate reconstruction and determining the monthly correlation between them for 3 categories of magnitude. Afterwards, the correlation function is applied to different bias-corrected climate model members. The framework uses measurements from the data archive maintained by NOAA (National Oceanic and Atmospheric Administration; <https://www.ncei.noaa.gov/pub/data/>), the ERA5 reanalysis data from C3S (Copernicus Climate Change Service; <https://cds.climate.copernicus.eu>) and CMIP6 climate projections from ESGF (Earth System Grid Federation; <https://esgf-node.llnl.gov/esg-search>). One can also provide its custom measurement data as alternative observation input, as long as the data has climatological coverage of multiple tens of years.

Each of the extreme weather variables are tested for particular case areas around the world where records are available on climatological time scales. They include daily minimum temperature for Uccle (Belgium), daily maximum temperature for Baku (Azerbaijan), heat stress temperature for Kasungu (Malawi), wind speed for Dahlem (Germany), and daily precipitation in Sikasso (Mali). For these case studies, EXSoDOS was found to reproduce the annual cycle statistics (climatological mean, median and 10th-to-90th percentile range), the density distribution and return periods. The high skill of the model to reproduce extreme occurrences up to 10 year return periods is achieved by the two-step bias-correction and downscaling procedure in line with previous methods (Switanek et al., 2022; Volosciuk et al., 2017). ~~The results for extreme precipitation and wind speed, especially for occurrences of 1 to 10-year return periods, are further improved by the additional pre-sealing of the point-scale~~

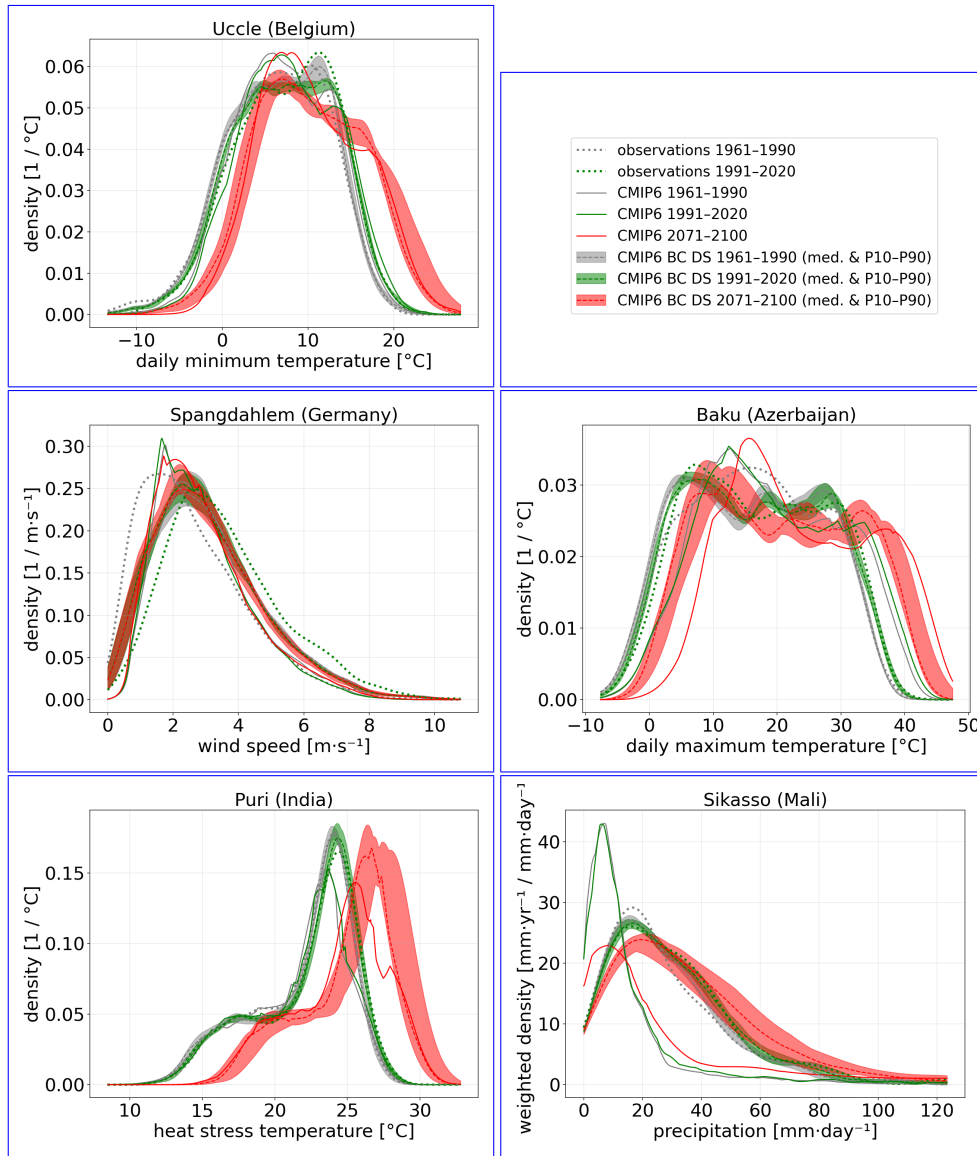


Figure 6. Idem as Fig. 4, but showing results for original CMIP6 climate projections (SSP585) including models listed in Tab. 1, bias-corrected and downscaled CMIP6 models (CMIP6_BC_DS). Observations and downscaled ERA5 are also included as comparison. Results are shown for two historical time frames 1961–1990 (in grey) and 1991–2020 (green), and for one future timeframe 2071–2100 (in red). For CMIP6_BS_DS, we show the 10-to-90 percentile spread over the different models.

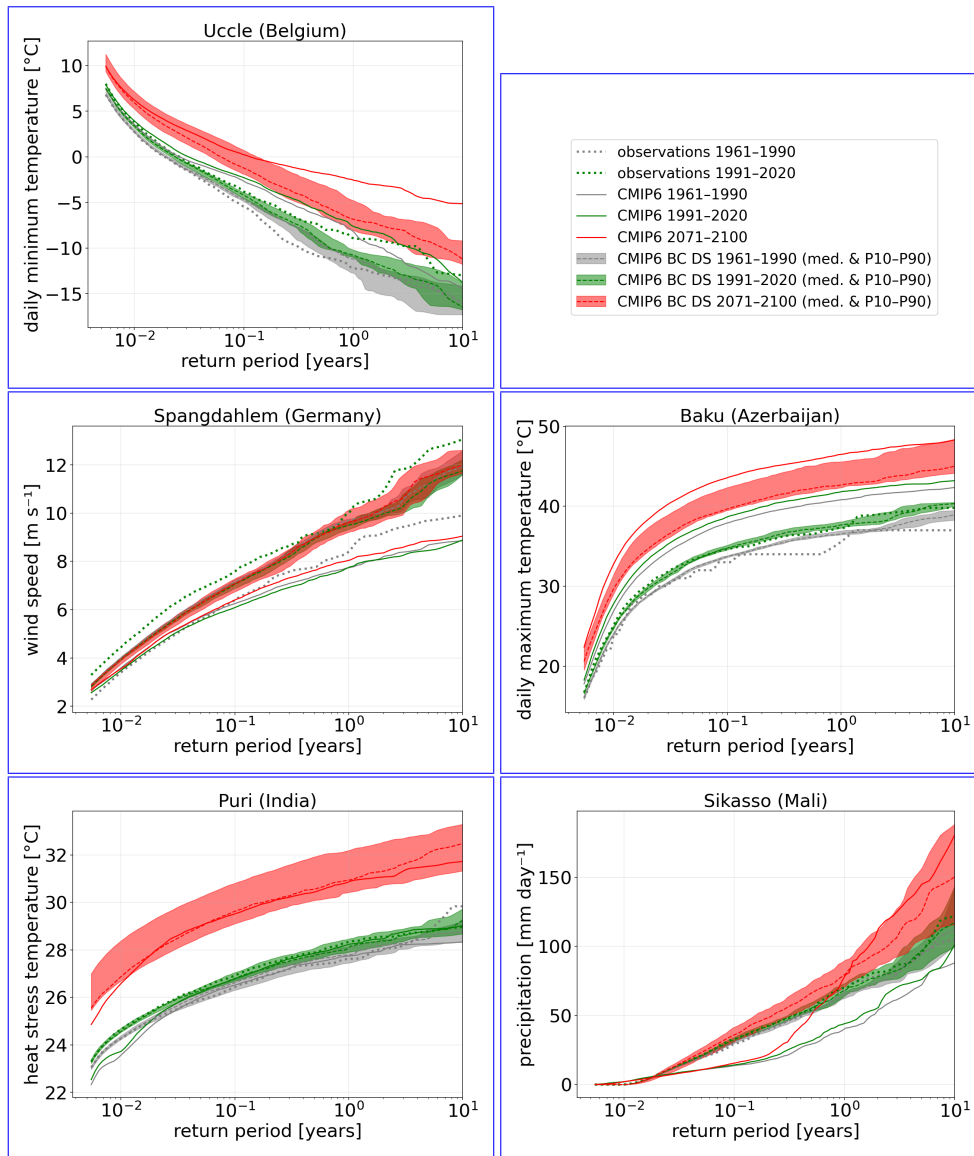


Figure 7. Idem as [Fig. 6](#), but for the return levels.

Dataset	Window	Average	Std	P95	1y return	
Daily minimum temperature (Uccle, Belgium)						
observations	1961–1990	6.41	5.92	15.31	-12.20	
observations	1991–2020	7.57	5.80	16.20	-8.90	
CMIP6	1961–1990	7.35 (5.48–8.20)	5.44 (5.08–6.10)	15.92 (13.93–17.06)	-8.11 (-13.27–-7.54)	
CMIP6	1991–2020	7.89 (6.11–9.07)	5.54 (5.17–6.31)	16.86 (14.82–18.36)	-7.57 (-10.69–-5.62)	
CMIP6	2071–2100	10.52 (8.90–13.32)	6.08 (5.56–6.46)	20.52 (18.89–23.90)	-2.57 (-4.96–-0.40)	
CMIP6_BC_DS	1961–1990	6.54 (6.47–6.81)	5.88 (5.79–5.98)	15.39 (15.14–15.64)	-11.21 (-12.72–-10.87)	
CMIP6_BC_DS	1991–2020	7.34 (7.12–7.37)	6.02 (5.92–6.06)	16.38 (16.20–16.43)	-10.76 (-11.01–-10.21)	
CMIP6_BC_DS	2071–2100	10.18 (9.65–11.65)	6.37 (6.06–6.58)	20.19 (19.73–22.08)	-6.86 (-7.99–-4.81)	
Daily mean wind speed (Spangdahlem, Germany)						
observations	1961–1990	2.59	1.63	5.68	8.39	
observations	1991–2020	3.58	1.79	6.94	10.05	
CMIP6	1961–1990	2.95 (2.06–3.90)	1.27 (0.95–1.74)	5.76 (3.93–6.80)	7.73 (5.45–9.26)	
CMIP6	1991–2020	2.85 (2.04–3.87)	1.26 (0.97–1.68)	5.62 (3.95–6.50)	7.77 (5.54–9.05)	
CMIP6	2071–2100	2.98 (1.98–3.83)	1.27 (0.97–1.66)	5.81 (3.90–6.40)	8.04 (5.66–8.59)	
CMIP6_BC_DS	1961–1990	3.15 (3.08–3.23)	1.77 (1.69–1.81)	6.48 (6.43–6.51)	9.71 (9.55–9.76)	
CMIP6_BC_DS	1991–2020	3.07 (3.03–3.20)	1.72 (1.67–1.79)	6.34 (6.23–6.43)	9.51 (9.32–9.64)	
CMIP6_BC_DS	2071–2100	3.13 (2.90–3.19)	1.74 (1.63–1.80)	6.35 (6.08–6.56)	9.62 (9.23–10.16)	
Daily maximum temperature (Baku, Azerbaijan)						
observations	1961–1990	16.31	9.94	32.00	37.00	
observations	1991–2020	17.07	10.47	33.40	37.30	
CMIP6	1961–1990	19.08 (16.34–20.68)	10.33 (9.49–11.50)	36.38 (32.27–37.38)	40.81 (36.73–41.93)	
CMIP6	1991–2020	19.76 (17.48–21.22)	10.46 (9.46–11.60)	37.20 (33.50–38.52)	41.84 (37.67–43.59)	
CMIP6	2071–2100	23.58 (23.12–25.25)	11.20 (10.01–12.64)	42.22 (40.19–44.10)	46.50 (44.58–49.01)	
CMIP6_BC_DS	1961–1990	15.86 (15.67–16.00)	10.62 (10.47–10.66)	32.47 (32.30–32.65)	36.55 (36.45–36.93)	
CMIP6_BC_DS	1991–2020	16.62 (16.47–16.80)	10.67 (10.58–10.74)	33.41 (33.36–33.86)	37.68 (37.32–38.06)	
CMIP6_BC_DS	2071–2100	20.44 (20.21–22.57)	11.38 (11.16–11.74)	38.47 (38.19–40.98)	42.66 (42.33–45.57)	
Heat stress temperature (Puri, India)						
observations	1961–1990	21.97	3.31	25.95	27.62	Heat days / yr
observations	1991–2020	22.15	3.55	26.47	28.35	4.40
CMIP6	1961–1990	21.12 (20.10–22.63)	3.51 (3.03–4.72)	26.08 (24.39–27.25)	27.74 (25.15–28.64)	4.50 (0.00–31.39)
CMIP6	1991–2020	21.38 (20.34–23.04)	3.54 (3.05–4.67)	26.31 (24.65–27.62)	28.24 (25.38–29.14)	7.47 (0.00–46.83)
CMIP6	2071–2100	23.77 (23.24–27.04)	3.25 (2.81–4.27)	29.10 (26.83–31.32)	30.85 (27.77–32.65)	81.80 (12.83–208.33)
CMIP6_BC_DS	1961–1990	21.84 (21.78–21.96)	3.48 (3.45–3.59)	26.18 (25.88–26.28)	27.78 (27.45–28.13)	5.73 (3.13–7.35)
CMIP6_BC_DS	1991–2020	22.14 (22.09–22.22)	3.49 (3.45–3.55)	26.35 (26.26–26.52)	28.06 (27.89–28.45)	7.77 (6.23–10.59)
CMIP6_BC_DS	2071–2100	24.66 (24.51–26.17)	3.25 (3.19–3.33)	29.07 (28.62–30.30)	30.93 (30.41–32.19)	92.87 (80.47–184.05)
Daily accumulated precipitation (Sikasso, Mali)						
observations	1961–1990	3.20	8.99	20.20	67.60	Dry days / yr
observations	1991–2020	3.40	9.62	22.49	69.70	281.84
CMIP6	1961–1990	2.18 (1.78–3.06)	5.14 (4.43–8.61)	9.94 (6.86–13.67)	42.39 (28.90–82.14)	229.97 (220.50–250.67)
CMIP6	1991–2020	2.35 (1.85–3.48)	5.72 (4.57–9.85)	10.86 (7.21–15.53)	47.24 (31.40–93.95)	230.83 (219.03–249.55)
CMIP6	2071–2100	2.52 (2.46–3.88)	7.89 (5.90–12.65)	11.25 (9.31–18.21)	69.75 (46.81–114.54)	235.57 (215.90–247.13)
CMIP6_BC_DS	1961–1990	3.20 (3.12–3.30)	9.24 (8.82–9.30)	21.10 (20.56–22.09)	64.36 (61.90–67.14)	284.97 (282.65–285.78)
CMIP6_BC_DS	1991–2020	3.40 (3.36–3.50)	9.72 (9.54–10.30)	22.54 (22.04–23.16)	69.43 (64.91–73.14)	284.48 (282.90–286.80)
CMIP6_BC_DS	2071–2100	3.92 (3.28–4.26)	11.68 (10.04–12.87)	25.67 (21.93–28.59)	82.20 (70.97–90.45)	285.28 (282.83–289.37)

Table 3. Annual mean, standard deviation, percentile 95 value and 1-year return value, annual dry days (for daily precipitation < 1mm), annual heat days (for heat stress temperature > 27°C) of observed and modelled time series. We include the original CMIP6 climate projections (SSP585) including models listed in Tab. 1, and the bias-corrected and downscaled CMIP6 models (CMIP6_BC_DS). Observations are also included as comparison. Results are shown for two historical time frames 1961–1990 and 1991–2020, and for one future timeframe 2071–2100. For the climate projections, we show median and percentile 10–90 ranges of the model ensemble.

525 ~~predictand with the coarse-scale predictor, hence for which high extremes are sampled as a perturbation of the grid-scale predictor.~~

EXSoDOS was also able to reproduce many of the observed shifts in extreme weather of the past climate (1991-2020 versus 1961-1990), including a tendency to more extreme precipitation and more extreme high temperature and heat stress, and a reduction in cold extremes. While the overall statistics of wind speed were found to be reproduced, the shift in more extreme wind in the observations was not found in CMIP6 or the downscaled time series. Applying the model on future climate projections towards the end of the 21st century (2071-2100), it was found that increasing extremes for daily maximum temperature, precipitation and heat stress are all exacerbated under future global warming. We further show that outcomes of shifting climate extremes are affected in different ways when applying a downscaling. Particularly for precipitation, relative and absolute climate-change signals are affected, as well as the changing return periods. As such, EXSoDOS downscaling is able to offer a baseline estimation of weather extremes and their past and future shifts under global warming for multiple weather variables for the first time, for which the statistical model is constrained with observations at any location on Earth where ~~climatological~~ hyper-climatological (>> 60 year) measurements are available. Since the ~~algorithm~~ algorithm has a low computational cost and uses globally available data, its application could be upscaled to perform assessments for the available data archive of weather stations across the globe (see Fig. 2), but also for single assessments with available data by local institutes.

The model is only evaluated for 5 sites, hence new applications require additional validation with local data. Nevertheless, the EXSoDOS, including its validation and application, is transferable to any station in the world. We further exemplify transferability by showing results for precipitation for 7 random stations across USA in the supplementary material, see resp. Figs. S1 (validation) and S2 (projection). EXSoDOS should be considered when long-term station weather station data is available and when representation extremes distribution (ie., tails) at point-scale is important to evaluate past and future climate change. In other cases, one should use existing state-of-the-art archives like CORDEX (Coppola et al., 2021) or CHELSA-W5E5 (Karger et al., 2023) providing grid-scale climate reconstruction and projections down to 1km resolution.

One should keep in mind the limitations of statistical downscaling when employing the model for future climate assessment. ~~Shifts of local extremes rely on the realism of the shiftthe climate variables at the coarse scale provided by the global climate models, and on the representativeness of the correlation function established between the~~ In this manuscript, the model is calibrated and validated with 63 year records, only providing assessments for return periods of up to 10 years. For higher return periods, one requires either substantially longer observational records or an explicit extreme-value extrapolation step requiring additional assumptions (e.g., employing Generalized Pareto Distribution or Generalized Extreme Value Distribution). One should interpret distribution shifts, especially those in far tails, conditional on the uncertainty related to calibration sets and representativeness/stability of correlation between coarse-scale ~~variable (predictor) predictor~~ and point-scale ~~variable (predictand) .~~ Particularly, the correlation is considered static, while in reality this correlation may change under global warming. The stochastic model ~~predictand under climate change (see sensitivity analysis in Text S1), methodological choices (e.g., detrending/retrending), and the realism of the shifts of the climate predictors at the coarse scale provided by the global climate models.~~

560 EXSoDOS also considers only one predictor ~~at a particular and one predictand for one~~ location at a time, ~~while local extremes~~
~~may be associated by the anomalies of multiple large-scale variables~~ which implies several limitations. At first, the model
~~doesn't explicitly capture its dependences on multiple local environmental (plant functional types, soil, urban environment...)~~
~~and climatic parameters (e.g., boundary-layer stability or circulation indicators)~~. As such, ~~the~~ complexity of physical processes
may be underrepresented by the model, ~~and considering compound events of different variables at neighbouring locations~~
565 ~~is not possible. In order to~~. To perform a more in-depth analysis of changing extremes and their underlying physics, one
should still ~~employ~~ rely on mechanistic high-resolution atmospheric numerical modelling. ~~At second, the simulation of only~~
~~one output variable (predictand) at a time ignores possible correlation among them, so hampers consistent representation of~~
~~compound hazards (e.g., heavy drought-heat, heavy rain-wind)~~. At third, modelling one location at a time does not preserve
~~spatial dependence between multiple sites (eg., a single convective rainstorm affecting multiple sites)~~.

570 The ~~accuracy of the~~ stochastic model could be improved ~~by considering multiple coarse-scale in~~ several ways. At first,
~~coherent time series for multiple sites and/or multiple variables can be achieved by correlating different~~ predictors and predic-
~~tands at different locations simultaneously. This would not only allow the improve the physics, but also the consideration of~~
~~compound events. This is possible by upgrading with correlation matrices among the predictands and predictors. Finally, in~~
~~order~~ for different variables and different sites. This can be done by normalizing predictands and predictors, subsequently transforming
575 ~~them to independent variables using their correlation matrix (or Gaussian copula)~~. After such calibration, one combines
~~predictor variables with random sampling and transforms the variables back to the correlated space, and finally one denormalizes~~
~~them again to their respective distributions. Multi-variable correlated sampling makes representation of compound hazards and~~
~~spatial dependence between different sites possible. Such a strategy that introduces a correlation matrix over different predictors~~
~~and predictands is conceptually similar to the multi-site approach proposed by (Switanek et al., 2022)~~. At second, extending the
580 ~~predictor set further with environmental and climatic parameters (mentioned above) is a promising avenue to better represent~~
~~regime-dependent land-atmosphere interactions. Finally~~ to improve the representation of underlying drivers and their shifts un-
der global warming, the ~~correlation function~~ correlated sampling could be upgraded ~~with artificial intelligence~~ using artificial
intelligence, in which one still represents both deterministic predictor-predictand relations and the remaining (unresolved)
variability and extremes with stochastic sampling.

585 *Code and data availability.* The code with documentation (README.md) and input data can be accessed on Zenodo at <https://doi.org/10.5281/zenodo.15387101> (Wouters, 2025).

Appendix A: Random sampling of a variable correlating with another

We look for α and β in such a way that \tilde{y} correlates with \tilde{x} with correlation coefficient ρ estimated from the calibration above.

$$\begin{aligned}\rho &= \langle \tilde{y}\tilde{x} \rangle \\ &= \langle \alpha x x \rangle + \langle \beta r x \rangle \\ &= \alpha \sigma_x^2 \\ \rho &= \alpha\end{aligned}\tag{A1}$$

590 where $\sigma_{\tilde{x}} = 1$

$$\begin{aligned}\sigma_y^2 &= \langle y y \rangle \\ &= \langle (\alpha x)^2 \rangle + \langle 2\alpha x \beta r \rangle + \langle (\beta r)^2 \rangle \\ &= \alpha^2 \sigma^2 + \beta^2 \sigma_r^2 \\ &= \alpha^2 + \beta^2\end{aligned}\tag{A2}$$

where $\sigma_y = \sigma_x = 1$. So

$$\beta = \sqrt{1 - \alpha^2}\tag{A3}$$

595 *Author contributions.* H.W. led the writing of the original draft, conceptualized the study, developed the methodology, performed formal analysis, and contributed to visualization. J.B. contributed to conceptualization and reviewed and edited the manuscript. F.P. acquired funding, administered the project, and performed validation. B.D. curated data, provided resources, contributed to validation, and reviewed and edited the manuscript. A.D. curated data, provided resources, contributed to validation, and reviewed and edited the manuscript. R.H. developed software and provided resources. D.L. contributed to conceptualization and supervised the research.

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References

- 605 Weather and Climate Extreme Events in a Changing Climate, in: *Climate Change 2021 – The Physical Science Basis: Working Group I Contribution to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change*, edited by Intergovernmental Panel on Climate Change (IPCC), pp. 1513–1766, Cambridge University Press, Cambridge, ISBN 978-1-009-15788-9, <https://doi.org/10.1017/9781009157896.013>, 2023.
- Benestad, R. E., Parding, K. M., and Dobler, A.: Downscaling the Probability of Heavy Rainfall over the Nordic Countries, *Hydrology and Earth System Sciences*, 29, 45–65, <https://doi.org/10.5194/hess-29-45-2025>, 2025.
- 610 Boucher, O., Servonnat, J., Albright, A. L., Aumont, O., Balkanski, Y., Bastrikov, V., Bekki, S., Bonnet, R., Bony, S., Bopp, L., Braconnot, P., Brockmann, P., Cadule, P., Caubel, A., Cheruy, F., Codron, F., Cozic, A., Cugnet, D., D’Andrea, F., Davini, P., De Lavergne, C., Denvil, S., Deshayes, J., Devilliers, M., Ducharne, A., Dufresne, J.-L., Dupont, E., Éthé, C., Fairhead, L., Falletti, L., Flavoni, S., Foujols, M.-A., Gardoll, S., Gastineau, G., Ghattas, J., Grandpeix, J.-Y., Guenet, B., Guez, E., L., Guilyardi, E., Guimberteau, M., Hauglustaine, D., Hourdin, F., Idelkadi, A., Joussaume, S., Kageyama, M., Khodri, M., Krinner, G., Lebas, N., Levvasseur, G., Lévy, C., Li, L., Lott, F., Lurton, T., Luysaert, S., Madec, G., Madeleine, J.-B., Maignan, F., Marchand, M., Marti, O., Mellul, L., Meurdesoif, Y., Mignot, J., Musat, I., Ottlé, C., Peylin, P., Planton, Y., Polcher, J., Rio, C., Rochetin, N., Rousset, C., Sepulchre, P., Sima, A., Swingedouw, D., Thiéblemont, R., Traore, A. K., Vancoppenolle, M., Vial, J., Vialard, J., Viovy, N., and Vuichard, N.: Presentation and Evaluation of the IPSL-CM6A-LR Climate Model, *Journal of Advances in Modeling Earth Systems*, 12, e2019MS002 010, <https://doi.org/10.1029/2019MS002010>, 2020.
- 620 Brisson, E., Demuzere, M., Willems, P., and Van Lipzig, N. P. M.: Assessment of Natural Climate Variability Using a Weather Generator, *Climate Dynamics*, 44, 495–508, <https://doi.org/10.1007/s00382-014-2122-8>, 2015.
- Brisson, E., Van Weverberg, K., Demuzere, M., Devis, A., Saeed, S., Stengel, M., and Van Lipzig, N. P. M.: How Well Can a Convection-Permitting Climate Model Reproduce Decadal Statistics of Precipitation, Temperature and Cloud Characteristics?, *Climate Dynamics*, 47, 3043–3061, <https://doi.org/10.1007/s00382-016-3012-z>, 2016.
- 625 Brunner, A. M., Evans, L. M., Hsu, C.-Y., and Sheng, X.: Vernalization and the Chilling Requirement to Exit Bud Dormancy: Shared or Separate Regulation?, *Frontiers in Plant Science*, 5, <https://doi.org/10.3389/fpls.2014.00732>, 2014.
- Cannon, A. J., Sobie, S. R., and Murdock, T. Q.: Bias Correction of GCM Precipitation by Quantile Mapping: How Well Do Methods Preserve Changes in Quantiles and Extremes?, *Journal of Climate*, 28, 6938–6959, <https://doi.org/10.1175/JCLI-D-14-00754.1>, 2015.
- Carter, T.: StratAdapt: Advancing Climate-Resilient Agriculture in Mali, [https://pressroom.icrisat.org/stratadapt-advancing-climate-resilient-](https://pressroom.icrisat.org/stratadapt-advancing-climate-resilient-agriculture-in-mali)
- 630 [agriculture-in-mali](https://pressroom.icrisat.org/stratadapt-advancing-climate-resilient-agriculture-in-mali), 2024.
- Coppola, E., Raffaele, F., Giorgi, F., Giuliani, G., Xuejie, G., Ciarlo, J. M., Sines, T. R., Torres-Alavez, J. A., Das, S., Di Sante, F., Pichelli, E., Glazer, R., Müller, S. K., Abba Omar, S., Ashfaq, M., Bukovsky, M., Im, E.-S., Jacob, D., Teichmann, C., Remedio, A., Remke, T., Kriegsmann, A., Bülow, K., Weber, T., Buntemeyer, L., Sieck, K., and Rechid, D.: Climate Hazard Indices Projections Based on CORDEX-CORE, CMIP5 and CMIP6 Ensemble, *Climate Dynamics*, 57, 1293–1383, <https://doi.org/10.1007/s00382-021-05640-z>, 2021.
- 635 De Ridder, K., Lauwaet, D., and Maiheu, B.: UrbClim – A Fast Urban Boundary Layer Climate Model, *Urban Climate*, 12, 21–48, <https://doi.org/10.1016/j.uclim.2015.01.001>, 2015.
- Devis, A., van Lipzig, N. P. M., and Demuzere, M.: A New Statistical Approach to Downscale Wind Speed Distributions at a Site in Northern Europe, *Journal of Geophysical Research: Atmospheres*, 118, 2272–2283, <https://doi.org/10.1002/jgrd.50245>, 2013.

- Eyring, V., Bony, S., Meehl, G. A., Senior, C. A., Stevens, B., Stouffer, R. J., and Taylor, K. E.: Overview of the Coupled Model Intercomparison Project Phase 6 (CMIP6) Experimental Design and Organization, *Geoscientific Model Development*, 9, 1937–1958, <https://doi.org/10.5194/gmd-9-1937-2016>, 2016.
- Fischer, E. M. and Knutti, R.: Anthropogenic Contribution to Global Occurrence of Heavy-Precipitation and High-Temperature Extremes, *Nature Climate Change*, 5, 560–564, <https://doi.org/10.1038/nclimate2617>, 2015.
- Fosser, G., Gaetani, M., Kendon, E. J., Adinolfi, M., Ban, N., Belušić, D., Caillaud, C., Careto, J. A. M., Coppola, E., Demory, M.-E., De Vries, H., Dobler, A., Feldmann, H., Goergen, K., Lenderink, G., Pichelli, E., Schär, C., Soares, P. M. M., Somot, S., and Tölle, M. H.: Convection-Permitting Climate Models Offer More Certain Extreme Rainfall Projections, *npj Climate and Atmospheric Science*, 7, 51, <https://doi.org/10.1038/s41612-024-00600-w>, 2024.
- Gutjahr, O., Putrasahan, D., Lohmann, K., Jungclaus, J. H., Von Storch, J.-S., Brüggemann, N., Haak, H., and Stössel, A.: Max Planck Institute Earth System Model (MPI-ESM1.2) for the High-Resolution Model Intercomparison Project (HighResMIP), *Geoscientific Model Development*, 12, 3241–3281, <https://doi.org/10.5194/gmd-12-3241-2019>, 2019.
- Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas, J., Peubey, C., Radu, R., Schepers, D., Simmons, A., Soci, C., Abdalla, S., Abellan, X., Balsamo, G., Bechtold, P., Biavati, G., Bidlot, J., Bonavita, M., De Chiara, G., Dahlgren, P., Dee, D., Diamantakis, M., Dragani, R., Flemming, J., Forbes, R., Fuentes, M., Geer, A., Haimberger, L., Healy, S., Hogan, R. J., Hólm, E., Janisková, M., Keeley, S., Laloyaux, P., Lopez, P., Lupu, C., Radnoti, G., De Rosnay, P., Rozum, I., Vamborg, F., Villaume, S., and Thépaut, J.-N.: The ERA5 Global Reanalysis, *Quarterly Journal of the Royal Meteorological Society*, 146, 1999–2049, <https://doi.org/10.1002/qj.3803>, 2020.
- Im, E.-S., Pal, J. S., and Eltahir, E. A. B.: Deadly Heat Waves Projected in the Densely Populated Agricultural Regions of South Asia, *Science Advances*, 3, e1603322, <https://doi.org/10.1126/sciadv.1603322>, 2017.
- IPCC: Climate Change 2021: The Physical Science Basis. Contribution of Working Group I to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change, Tech. rep., IPCC, 2021.
- Karger, D. N., Lange, S., Hari, C., Reyer, C. P. O., Conrad, O., Zimmermann, N. E., and Frieler, K.: CHELSA-W5E5: Daily 1 Km Meteorological Forcing Data for Climate Impact Studies, *Earth System Science Data*, 15, 2445–2464, <https://doi.org/10.5194/essd-15-2445-2023>, 2023.
- Katragkou, E., Sobolowski, S. P., Teichmann, C., Solmon, F., Pavlidis, V., Rechid, D., Hoffmann, P., Fernandez, J., Nikulin, G., and Jacob, D.: Delivering an Improved Framework for the New Generation of CMIP6-Driven EURO-CORDEX Regional Climate Simulations, *Bulletin of the American Meteorological Society*, 105, E962–E974, <https://doi.org/10.1175/BAMS-D-23-0131.1>, 2024.
- Kendon, E. J., Prein, A. F., Senior, C. A., and Stirling, A.: Challenges and Outlook for Convection-Permitting Climate Modelling, *Philosophical Transactions of the Royal Society A: Mathematical, Physical and Engineering Sciences*, 379, 20190547, <https://doi.org/10.1098/rsta.2019.0547>, 2021.
- Lange, S.: Trend-Preserving Bias Adjustment and Statistical Downscaling with ISIMIP3BASD (v1.0), *Geoscientific Model Development*, 12, 3055–3070, <https://doi.org/10.5194/gmd-12-3055-2019>, 2019.
- Lee, W.-L., Wang, Y.-C., Shiu, C.-J., Tsai, I.-c., Tu, C.-Y., Lan, Y.-Y., Chen, J.-P., Pan, H.-L., and Hsu, H.-H.: Taiwan Earth System Model Version 1: Description and Evaluation of Mean State, *Geoscientific Model Development*, 13, 3887–3904, <https://doi.org/10.5194/gmd-13-3887-2020>, 2020.
- Maraun, D.: Bias Correction, Quantile Mapping, and Downscaling: Revisiting the Inflation Issue, *Journal of Climate*, 26, 2137–2143, <https://doi.org/10.1175/JCLI-D-12-00821.1>, 2013.

- Maraun, D.: Bias Correcting Climate Change Simulations - a Critical Review, *Current Climate Change Reports*, 2, 211–220, <https://doi.org/10.1007/s40641-016-0050-x>, 2016.
- 680 Mauritsen, T., Bader, J., Becker, T., Behrens, J., Bittner, M., Brokopf, R., Brovkin, V., Claussen, M., Crueger, T., Esch, M., Fast, I., Fiedler, S., Fläschner, D., Gayler, V., Giorgetta, M., Goll, D. S., Haak, H., Hagemann, S., Hedemann, C., Hohenegger, C., Ilyina, T., Jahns, T., Jimenéz-de-la Cuesta, D., Jungclaus, J., Kleinen, T., Kloster, S., Kracher, D., Kinne, S., Kleberg, D., Lasslop, G., Kornblueh, L., Marotzke, J., Matei, D., Meraner, K., Mikolajewicz, U., Modali, K., Möbis, B., Müller, W. A., Nabel, J. E. M. S., Nam, C. C. W., Notz, D., Nyawira, S.-S., Paulsen, H., Peters, K., Pincus, R., Pohlmann, H., Pongratz, J., Popp, M., Raddatz, T. J., Rast, S., Redler, R., Reick, C. H., Rohrschneider, T., Schemann, V., Schmidt, H., Schnur, R., Schulzweida, U., Six, K. D., Stein, L., Stemmler, I., Stevens, B., Von Storch, 685 J.-S., Tian, F., Voigt, A., Vrese, P., Wieners, K.-H., Wilkenskjeld, S., Winkler, A., and Roeckner, E.: Developments in the MPI-M Earth System Model Version 1.2 (MPI-ESM1.2) and Its Response to Increasing CO₂, *Journal of Advances in Modeling Earth Systems*, 11, 998–1038, <https://doi.org/10.1029/2018MS001400>, 2019.
- Panthou, G., Vischel, T., and Lebel, T.: Recent Trends in the Regime of Extreme Rainfall in the Central Sahel, *International Journal of Climatology*, 34, 3998–4006, <https://doi.org/10.1002/joc.3984>, 2014.
- 690 Perkins, S. E., Pitman, A. J., Holbrook, N. J., and McAneney, J.: Evaluation of the AR4 Climate Models' Simulated Daily Maximum Temperature, Minimum Temperature, and Precipitation over Australia Using Probability Density Functions, *Journal of Climate*, 20, 4356–4376, <https://doi.org/10.1175/JCLI4253.1>, 2007.
- Perkins-Kirkpatrick, S. E. and Lewis, S. C.: Increasing Trends in Regional Heatwaves, *Nature Communications*, 11, 3357, <https://doi.org/10.1038/s41467-020-16970-7>, 2020.
- 695 Prein, A. F., Langhans, W., Fosser, G., Ferrone, A., Ban, N., Goergen, K., Keller, M., Tölle, M., Gutjahr, O., Feser, F., Brisson, E., Kollet, S., Schmidli, J., Van Lipzig, N. P. M., and Leung, R.: A Review on Regional Convection-permitting Climate Modeling: Demonstrations, Prospects, and Challenges, *Reviews of Geophysics*, 53, 323–361, <https://doi.org/10.1002/2014RG000475>, 2015.
- Pryor, S. C. and Barthelmie, R. J.: Climate Change Impacts on Wind Energy: A Review, *Renewable and Sustainable Energy Reviews*, 14, 430–437, <https://doi.org/10.1016/j.rser.2009.07.028>, 2010.
- 700 Raymond, C., Matthews, T., and Horton, R. M.: The Emergence of Heat and Humidity Too Severe for Human Tolerance, *Science Advances*, 6, eaaw1838, <https://doi.org/10.1126/sciadv.aaw1838>, 2020.
- Saeed, S., Brisson, E., Demuzere, M., Tabari, H., Willems, P., and Van Lipzig, N. P. M.: Multidecadal Convection Permitting Climate Simulations over Belgium: Sensitivity of Future Precipitation Extremes, *Atmospheric Science Letters*, 18, 29–36, <https://doi.org/10.1002/asl.720>, 2017.
- 705 Sanogo, S., Fink, A. H., Omotosho, J. A., Ba, A., Redl, R., and Ermert, V.: Spatio-Temporal Characteristics of the Recent Rainfall Recovery in West Africa, *International Journal of Climatology*, 35, 4589–4605, <https://doi.org/10.1002/joc.4309>, 2015.
- Sillmann, J., Kharin, V. V., Zhang, X., Zwiers, F. W., and Bronaugh, D.: Climate Extremes Indices in the CMIP5 Multimodel Ensemble: Part 1. Model Evaluation in the Present Climate, *Journal of Geophysical Research: Atmospheres*, 118, 1716–1733, <https://doi.org/10.1002/jgrd.50203>, 2013.
- 710 Swart, N. C., Cole, J. N. S., Kharin, V. V., Lazare, M., Scinocca, J. F., Gillett, N. P., Anstey, J., Arora, V., Christian, J. R., Hanna, S., Jiao, Y., Lee, W. G., Majaess, F., Saenko, O. A., Seiler, C., Seinen, C., Shao, A., Sigmond, M., Solheim, L., Von Salzen, K., Yang, D., and Winter, B.: The Canadian Earth System Model Version 5 (CanESM5.0.3), *Geoscientific Model Development*, 12, 4823–4873, <https://doi.org/10.5194/gmd-12-4823-2019>, 2019.

- Switanek, M., Maraun, D., and Bevacqua, E.: Stochastic Downscaling of Gridded Precipitation to Spatially Coherent Subgrid Precipitation
715 Fields Using a Transformed Gaussian Model, *International Journal of Climatology*, 42, 6126–6147, <https://doi.org/10.1002/joc.7581>, 2022.
- Tatebe, H., Ogura, T., Nitta, T., Komuro, Y., Ogochi, K., Takemura, T., Sudo, K., Sekiguchi, M., Abe, M., Saito, F., Chikira, M., Watanabe, S.,
Mori, M., Hirota, N., Kawatani, Y., Mochizuki, T., Yoshimura, K., Takata, K., O’ishi, R., Yamazaki, D., Suzuki, T., Kurogi, M., Kataoka,
T., Watanabe, M., and Kimoto, M.: Description and Basic Evaluation of Simulated Mean State, Internal Variability, and Climate Sensitivity
720 in MIROC6, *Geoscientific Model Development*, 12, 2727–2765, <https://doi.org/10.5194/gmd-12-2727-2019>, 2019.
- Termonia, P., Van Schaeybroeck, B., De Cruz, L., De Troch, R., Caluwaerts, S., Giot, O., Hamdi, R., Vannitsem, S., Duchêne, F., Willems,
P., Tabari, H., Van Uytven, E., Hosseinzadehtalaei, P., Van Lipzig, N., Wouters, H., Vanden Broucke, S., Van Ypersele, J.-P., Marbaix,
P., Villanueva-Birriel, C., Fettweis, X., Wyard, C., Scholzen, C., Doutreloup, S., De Ridder, K., Gobin, A., Lauwaet, D., Stavrakou, T.,
Bauwens, M., Müller, J.-F., Luyten, P., Ponsar, S., Van Den Eynde, D., and Pottiaux, E.: The CORDEX.Be Initiative as a Foundation for
725 Climate Services in Belgium, *Climate Services*, 11, 49–61, <https://doi.org/10.1016/j.cliser.2018.05.001>, 2018.
- Tobin, I., Jerez, S., Vautard, R., Thais, F., van Meijgaard, E., Prein, A., Déqué, M., Kotlarski, S., Maule, C. F., Nikulin, G., Noël, T.,
and Teichmann, C.: Climate Change Impacts on the Power Generation Potential of a European Mid-Century Wind Farms Scenario,
Environmental Research Letters, 11, 034 013, <https://doi.org/10.1088/1748-9326/11/3/034013>, 2016.
- Van De Velde, J., Demuzere, M., De Baets, B., and Verhoest, N.: Future Multivariate Weather Generation by Combining Bartlett-Lewis and
730 Vine Copula Models, *Hydrological Sciences Journal*, 68, 1–15, <https://doi.org/10.1080/02626667.2022.2144322>, 2023.
- Vanden Broucke, S., Wouters, H., Demuzere, M., and Van Lipzig, N. P. M.: The Influence of Convection-Permitting Regional Climate
Modeling on Future Projections of Extreme Precipitation: Dependency on Topography and Timescale, *Climate Dynamics*, 52, 5303–
5324, <https://doi.org/10.1007/s00382-018-4454-2>, 2019.
- Volodin, E. M., Mortikov, E. V., Kostykin, S. V., Galin, V. Ya., Lykossov, V. N., Gritsun, A. S., Diansky, N. A., Gusev, A. V., and
735 Iakovlev, N. G.: Simulation of the Present-Day Climate with the Climate Model INMCM5, *Climate Dynamics*, 49, 3715–3734,
<https://doi.org/10.1007/s00382-017-3539-7>, 2017.
- Volodin, E. M., Mortikov, E. V., Kostykin, S. V., Galin, V. Y., Lykossov, V. N., Gritsun, A. S., Diansky, N. A., Gusev, A. V., Iakovlev,
N. G., Shestakova, A. A., and Emelina, S. V.: Simulation of the Modern Climate Using the INM-CM48 Climate Model, *Russian Journal
of Numerical Analysis and Mathematical Modelling*, 33, 367–374, <https://doi.org/10.1515/rnam-2018-0032>, 2018.
- 740 Volosciuk, C., Maraun, D., Vrac, M., and Widmann, M.: A Combined Statistical Bias Correction and Stochastic Downscaling Method for
Precipitation, *Hydrology and Earth System Sciences*, 21, 1693–1719, <https://doi.org/10.5194/hess-21-1693-2017>, 2017.
- WMO: WMO Guidelines on the Calculation of Climate Normals (WMO-No. 1203). Authoritative Methods, Definitions, and Implementation
Details, Tech. rep., World Meteorological Organization, 2017.
- Wouters, H.: EXSoDOS: Downscaling of Changing Weather Extremes for Climate Projections: Code and Input Data for Demonstration,
745 Zenodo, <https://doi.org/10.5281/zenodo.15387101>, 2025.
- Wouters, H., Demuzere, M., Blahak, U., Fortuniak, K., Maiheu, B., Camps, J., Tielemans, D., and Van Lipzig, N. P. M.: The Efficient
Urban Canopy Dependency Parametrization (SURY) v1.0 Foratmospheric Modelling: Description and Application with the COSMO-
CLM Model for a Belgian Summer, *Geoscientific Model Development*, 9, 3027–3054, <https://doi.org/10.5194/gmd-9-3027-2016>, 2016.
- Wouters, H., De Ridder, K., Poelmans, L., Willems, P., Brouwers, J., Hosseinzadehtalaei, P., Tabari, H., Vanden Broucke, S., van Lipzig, N.
750 P. M., and Demuzere, M.: Heat Stress Increase under Climate Change Twice as Large in Cities as in Rural Areas: A Study for a Densely
Populated Midlatitude Maritime Region, *Geophysical Research Letters*, 44, 8997–9007, <https://doi.org/10.1002/2017GL074889>, 2017.

- Wouters, H., Keune, J., Petrova, I. Y., Van Heerwaarden, C. C., Teuling, A. J., Pal, J. S., Vilà-Guerau De Arellano, J., and Miralles, D. G.: Soil Drought Can Mitigate Deadly Heat Stress Thanks to a Reduction of Air Humidity, *Science Advances*, 8, eabe6653, <https://doi.org/10.1126/sciadv.abe6653>, 2022.
- 755 Yukimoto, S., Kawai, H., Koshiro, T., Oshima, N., Yoshida, K., Urakawa, S., Tsujino, H., Deushi, M., Tanaka, T., Hosaka, M., Yabu, S., Yoshimura, H., Shindo, E., Mizuta, R., Obata, A., Adachi, Y., and Ishii, M.: The Meteorological Research Institute Earth System Model Version 2.0, MRI-ESM2.0: Description and Basic Evaluation of the Physical Component, *Journal of the Meteorological Society of Japan*. Ser. II, 97, 931–965, <https://doi.org/10.2151/jmsj.2019-051>, 2019.