We really appreciate the referee's valuable comments. We have addressed each comment as below and revised the manuscript accordingly. As detailed below, the reviewer's comments are shown in black, our response to the comments is in blue. New or modified text is in red. The revised words are in highlight color in the revised manuscript.

Response to Referee #1:

Overall evaluations:

Zhang et al. compared field measurements of N_2O_5 and related species during and after Beijing winter Olympics. Time series and diurnal patterns of N_2O_5-related species were reported. Furthermore, key kinetic parameters were investigated, such as NO_3 reactivity (k_{NO3}), N_2O_5 uptake, and N_2O_5 lifetime. Regarding k_{NO3}, the contribution of NO and VOCs were discussed. As for N_2O_5 uptake, the steady-state method was applied to calculate the uptake coefficient. The influencing factors of N_2O_5 lifetime were also examined.

The investigated topic, i.e., reactive nitrogen chemistry, is important within the scope of ACP journal. The presented contents are suitable and align with previous studies. However, as a measurement report, some essential details of measurement methods are lacking. Uncertainty analysis should also be provided. In terms of writing, the authors are suggested to further polish the language with particular attention to some contradictory expressions. Other major issues as listed below concern data quality and the reliability of measurement interpretations. Overall, major revision is needed, and potential publication depends on the quality of revision.

Major comments:

1. In section 2.2, the statement of NO_3 and N_2O_5 measurement should be significantly enhanced. The data quality is in doubt without enough information provided, especially when considering that the instrument was in a malfunction state (line 99). At a minimum, the authors should make use of SI to record more technical details. Detailed comments regarding this issue are shown as follows.

(1) In lines 96-97, it looks like the authors can separately measure NO_3 and N_2O_5. However, in lines 99-100, the authors said only the sum of $NO_3 + N_2O_5$ can be measured. The above two statements are inconsistent.

Response：Thank you for your valuable comment.

Under normal operation, our CRDS analyzer adopts a dual-channel design: one channel directly measures NO_3, while the other heats and decomposes N_2O_5 into NO_3, thus quantifying the total concentration of $[NO_3 + N_2O_5]$. However, during this

observation campaign, the NO_3-specific detection channel malfunctioned (due to the damage of the mirror), leaving only the heated channel operational—hence, we could only obtain combined $[NO_3 + N_2O_5]$ data.

To address this, we estimated NO_3 concentrations using thermodynamic equilibrium (Eq. 1), with input parameters including the average nocturnal NO_2 concentration (14.5 ppbv) and temperature ($-1.4\,°C$) from Table 2. Calculations showed NO_3 contributed merely ~1% to the combined $[NO_3 + N_2O_5]$ signal (under winter's low-temperature and relatively high-NO_x conditions, the NO_3/N_2O_5 ratio is inherently low), confirming reliable derivation of NO_3 concentrations via this method.

We have revised the original text in Section 2.2 to clarify this dual-channel design, the malfunction, and the NO_3 estimation approach, with additional technical details supplemented in the Supporting Information (Text S1) for transparency.

(2) Lines 100-101, how was the limit of detection determined? What factors contributed to the overall uncertainty of 13.7%? Also, what was the background level of the instrument?

Response: Thank you for your comment.

Determination of the Limit of Detection (LOD) and Instrument Background Level.
The LOD of the CRDS analyzer for N_2O_5 was determined via a 4-hour zero-air injection experiment: zero air (free of target analytes) was continuously introduced into the instrument cavity (1 s time resolution), and Allan variance analysis was applied to evaluate system stability (Figure S1). The LOD was calculated using Eq. (1) below, yielding a value of 2.9 pptv.

$$[A] = \frac{R_L}{c\sigma}\left(\frac{1}{\tau} - \frac{1}{\tau_0}\right) \tag{1}$$

where [A] denotes the analyte concentration, "R_L" is the reference signal, c is the speed of light, σ is the absorption cross-section, τ is the cavity ring-down time with sample gas, and τ_0 is the background ring-down time with zero air. During the zero-air experiment, the instrument background level (τ_0) was measured as 50 μs.

Detailed experimental procedures, Allan variance results, and a plot of the instrument background signal (blue line) have been added to the Supplementary Material (Figure S1) for clarity.

Sources of the 13.7% Overall Uncertainty
The total measurement uncertainty (13.7%) arises from two main sources:
(1) Parameter uncertainties in concentration calculation: The absorption cross-section (σ) and effective absorption cavity length—key parameters in Eq. (1)—contribute 13% and 4% uncertainty, respectively.
(2) Data correction uncertainties: Corrections for membrane loss (4.5 ± 0.5%) and pipeline loss (11.4 ± 1.3%) during sample transport introduce an additional 1.4% uncertainty.

The total combined uncertainty was calculated following the method described in our prior work (Zhang et al., 2024), confirming the 13.7% value.

Zhang, T., Zuo, P., Ma, J., Ye, C., Lin, W., and Zhu, T.: Characterization and Application of an Online Measurement System for NO_3 and N_2O_5 Based on Cavity Ring-Down Spectroscopy, Acta Sci. Nat. Univ. Pekin., 60, 563–574, https://doi.org/10.13209/j.0479-8023.2024.030, 2024.

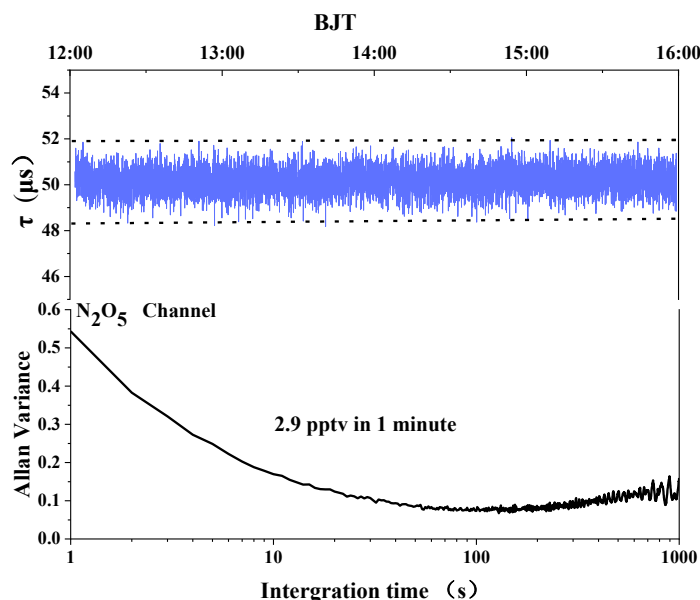We have added the instrument description to the Supplementary Material (Text S1).

Figure S1. Limit of detection (LOD) and background signal of the instrument (blue line).

(3) Lines 101-104, only the inlet issue was mentioned, while the calibration factor, or in other words, the sensitivity of the instrument is still not clearly stated.

Response: Thank you for your feedback. In this study, the instrument's sensitivity for N_2O_5 is represented by its limit of detection (LOD), which was determined based on signal fluctuations (variations in cavity ring-down time, τ) during a 4-hour zero-air injection experiment. As detailed in the Supplementary Material (and our response to Comment 2), the LOD for N_2O_5 was calculated as 2.9 pptv using Allan variance analysis and the concentration formula (Eq.1)—this LOD directly reflects the instrument's ability to detect low concentrations of N_2O_5, serving as a key indicator of its sensitivity.

To ensure measurement accuracy, we performed regular calibrations using stable, calibrated NO_3 and N_2O_5 standard sources (generated via a dynamic standard gas system, as referenced in our prior work: Zhang et al., 2026). These calibrations quantified two critical correction factors addressing sampling losses (a key contributor to signal attenuation between the inlet and measurement cell):

- **Tubing loss**: Measured at 11.4 ± 1.3%.

3

- **Filter membrane loss**: Quantified as $4.5 \pm 0.5\%$.

The final concentrations of $[NO_3 + N_2O_5]$ and derived NO_3 were obtained by dividing raw instrument measurements by the sum of these loss ratios (i.e., correcting for signal reduction during sample transport).

Relevant calibration procedures and loss correction details have been supplemented in the Supplementary Material (Text S1) for transparency.

Zhang, T., Ma, J., Liu, T., Lin, W., Zuo, P., and Ye, C.: A dynamic generation system for NO_3 and N_2O_5 standard gases, Environ. Chem., 45, 1–7, https://doi.org/10.7524/j.issn.0254-6108.2024091302, 2026.

2. In line 155, the aerosol surface area (S_a) was calculated by an empirical parameterization using $PM_{2.5}$. This calculation could bias S_a, which influences the results presented in the figure 4, figure 6, and figure 7. Considering the impact of S_a on the calculation of N_2O_5 uptake, an evaluation of the accuracy of this empirical formula should be provided.

Response: Thank you for your comments. We acknowledge the importance of S_a accuracy for N_2O_5 uptake calculations and subsequent results in Figures 4, 6, and 7. To address this, we first confirm that the empirical formula used for S_a calculation (Eq. 8: $S_a = 60.03 \times [PM_{2.5}]^{0.62}$) is validated for winter conditions in Beijing, as supported by Zhang et al. (2022)—The measured $PM_{2.5}$ and S_a exhibited a good linear correlation ($R^2 = 0.75$), while the nonlinear correlation yielded a higher R^2 of 0.82. At $PM_{2.5} = 52$ $\mu g\ m^{-3}$, the observed S_a, S_a derived from the linear correlation, and S_a derived from the nonlinear correlation were 625 $\mu m^2/cm^3$, 950 $\mu m^2/cm^3$, and 697 $\mu m^2/cm^3$, respectively. The nonlinear curve fitting was more suitable for the relationship between S_a and $PM_{2.5}$ than the linear fitting curve. Therefore, the relationship between $PM_{2.5}$ and S_a was $S_a = 60.03 \times [PM_{2.5}]^{0.62}$, which was used to estimate S_a.

In our revision, we have supplemented key details to clarify the formula's reliability:

Correlation performance: The empirical formula exhibits a strong linear correlation with $PM_{2.5}$, $R^2 = 0.82$.
Applicability range: This formula is specifically suitable for $PM_{2.5}$ concentrations < 200 $\mu g/m^3$, which fully covers the $PM_{2.5}$ range in our study (average: 24 ± 21 $\mu g/m^3$; maximum: 131 $\mu g/m^3$, see Table 2).

To enhance transparency, we have updated the main text (Line 194) as follows:
*"**Aerosol surface area density (S_a)**
Due to the unavailability of direct particle size distribution measurements, S_a was derived from $PM_{2.5}$ concentrations using an empirical formula validated for winter Beijing conditions (Zhang et al., 2022):
$S_a=60.03\times[PM_{2.5}]^{0.62}$ (8)

This formula exhibits a strong linear correlation ($R^2 = 0.82$) with $PM_{2.5}$ and is applicable for $PM_{2.5}$ concentrations < 200 µg m^{-3}—consistent with the $PM_{2.5}$ range observed in this study (average: 24 ± 21 µg m^{-3}, maximum: 131 µg m^{-3}).”

This supplementation confirms the empirical formula's suitability for our research, minimizing potential biases in S_a calculations and downstream N_2O_5 uptake analyses.

Zhang, X., Tong, S., Jia, C., Zhang, W., Li, J., Wang, W., Sun, Y., Wang, X., Wang, L., Ji, D., Wang, L., Zhao, P., Tang, G., Xin, J., Li, A., and Ge, M.: The Levels and Sources of Nitrous Acid (HONO) in Winter of Beijing and Sanmenxia, J. Geophys. Res. Atmos., 127, e2021JD036278, https://doi.org/10.1029/2021JD036278, 2022.

3. Line 283-291: Further to major comment 2, the accuracy of the empirical S_a can affect the discussion here. Also, when looking at the raw datapoints in figure 4b, a clear trend between N_2O_5 lifetime and S_a could not be identified.

Response: Thank you for your comments. As clarified in our response to Major Comment 2, the empirical formula for S_a ($S_a = 60.03 \times [PM_{2.5}]^{0.62}$) is validated for winter Beijing conditions: it exhibits a strong correlation ($R^2 = 0.82$) with $PM_{2.5}$, is applicable for $PM_{2.5}$ concentrations < 200 µg/m^3 (consistent with our study's $PM_{2.5}$ range: 24 ± 21 µg/m^3, max 131 µg/m^3). This ensures the calculated S_a is sufficiently accurate to underpin the discussion of N_2O_5 lifetime (τ) and heterogeneous uptake.

To address the ambiguity in the original raw data points, we have refined Figure 4b by:

*   Using **hourly-averaged data** (instead of raw data) to reduce random variability;
*   Applying new binning to S_a values on the x-axis for clearer trend visualization. The updated Figure 4b now more distinctly illustrates the relationship:
*   When S_a < 325 µm^2 cm^{-3}: $\tau(N_2O_5)$ gradually increases from ~10 to 12 minutes with rising S_a;
*   When S_a ranges from 500 to 1000 µm^2 cm^{-3}: a robust negative correlation between $\tau(N_2O_5)$ and S_a emerges (consistent with the physical expectation that higher S_a provides more reactive surfaces for N_2O_5 heterogeneous uptake, accelerating N_2O_5 loss and shortening its lifetime).

We have also revised the main text (4.1.2 Relationship between $\tau_{N_2O_5}$ and S_a) to align with the updated figure:

Line 380: S_a < 325 µm^2 cm^{-3}: $\tau_{N_2O_5}$ gradually increases with rising S_a. For low S_a values, $\tau_{N_2O_5}$ gradually rises from ~10 to 12 minutes as S_a increases. This non-monotonic pattern is driven by the co-occurrence of low S_a with extremely dry conditions (RH < 25% for 68% of data points in this S_a range).

Line 383: **S_a = 500–1000 μm² cm⁻³:** Robust negative correlation. Above a threshold S_a of ~500 μm² cm⁻³, a clear negative correlation emerges: $\tau_{N_2O_5}$ decreases from ~12 to 6 minutes as S_a increases. This aligns with physical expectations, as higher S_a provides more reactive surface area for N_2O_5 heterogeneous uptake (Lin et al., 2022; Wang et al., 2020; Zhou et al., 2018).

References:

Lin, C., Hu, R., Xie, P., Lou, S., Zhang, G., Tong, J., Liu, J., and Liu, W.: Nocturnal atmospheric chemistry of NO_3 and N_2O_5 over Changzhou in the Yangtze River Delta in China, J. Environ. Sci., 114, 376–390, https://doi.org/10.1016/j.jes.2021.09.016, 2022.

Wang, H., Chen, X., Lu, K., Hu, R., Li, Z., Wang, H., Ma, X., Yang, X., Chen, S., Dong, H., Liu, Y., Fang, X., Zeng, L., Hu, M., and Zhang, Y.: NO_3 and N_2O_5 chemistry at a suburban site during the EXPLORE-YRD campaign in 2018, Atmos. Environ., 224, 117180, https://doi.org/10.1016/j.atmosenv.2019.117180, 2020.

Zhou, W., Zhao, J., Ouyang, B., Mehra, A., Xu, W., Wang, Y., Bannan, T. J., Worrall, S. D., Priestley, M., Bacak, A., Chen, Q., Xie, C., Wang, Q., Wang, J., Du, W., Zhang, Y., Ge, X., Ye, P., Lee, J. D., Fu, P., Wang, Z., Worsnop, D., Jones, R., Percival, C. J., Coe, H., and Sun, Y.: Production of N_2O_5 and $ClNO_2$ in summer in urban Beijing, China, Atmos. Chem. Phys., 18, 11581–11597, https://doi.org/10.5194/acp-18-11581-2018, 2018.
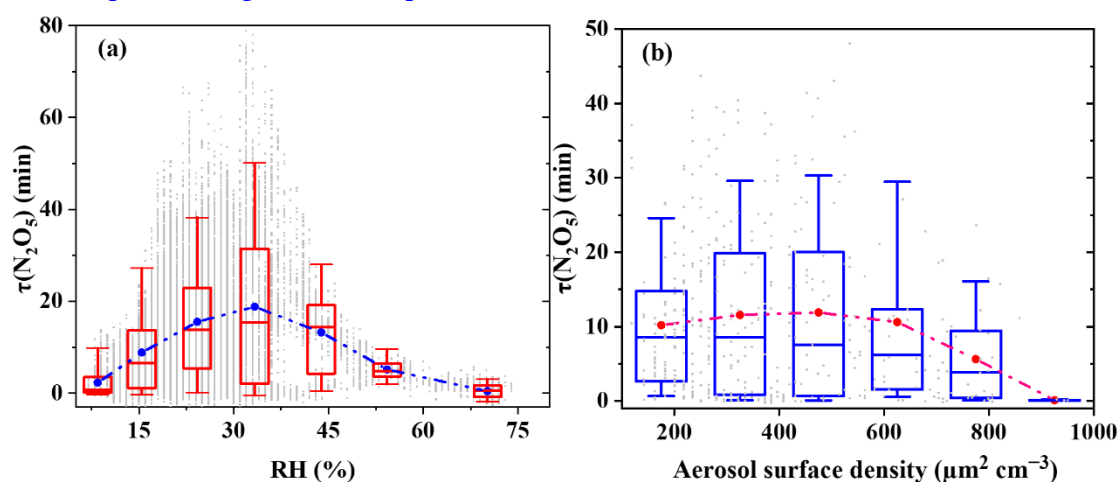
Figure 4. The relationship between $\tau_{N_2O_5}$ and (a) RH as well as (b) S_a during the observation period.

4. The $k(NO_3)$ calculated from VOCs oxidation (on the order of 1e-4 shown in figure 5) and steady-state analysis (up to 0.3, table S2) are totally different. What is the reason behind, and what is the influence of this issue on the calculated N_2O_5 uptake coefficient by the steady-state method? In addition, the $k(NO_3)$ stated in lines 364 to 365 (1.14 to 3.06) is even higher than those in figure 5 and table S2, which is also inconsistent.

Response：Thank you for your valuable comment. The $k(NO_3)$ values stated in lines 364–365 (ranging from 1.14 to 3.06 s⁻¹) represent the total reactivity of NO_3, including

its reactions with VOCs and NO, as well as the heterogeneous uptake of N_2O_5. Therefore, these values are significantly higher than the NO_3 reactivity calculated from VOCs alone. We acknowledge the discrepancies in $k(NO_3)$ values from different methods and clarify the reasons and impacts below:

(1) Reasons for Discrepancies in $k(NO_3)$

The differences arise from the distinct loss pathways captured by each calculation method, as summarized below Table:

Table S1: Comparison of $k(NO_3)$ Calculation Methods, Captured Loss Pathways, and Magnitudes

Calculation Method	Captured Loss Pathways	Magnitude of $k(NO_3)$	Rationale
VOC oxidation (Fig. 5)	Only reactions between NO_3 and measured VOCs	$\sim 10^{-4}$ s^{-1} (low)	This method underestimates total $k(NO_3)$ because key VOCs (e.g., limonene, α-pinene) were not measured, and their reactivity with NO_3 is unaccounted for.
Steady-state analysis (Table S2)	NO_3 loss (hydrocarbons + sulfur compounds) + uptake in clouds or deposition to the ground) (Brown et al., 2003)	Up to 0.3 s^{-1}	This method is the sum of the first-order rate constants for the irreversible removal of NO_3. Furthermore, although all data were filtered for periods with NO < 1 ppbv, NO concentrations below 1 ppbv can still contribute to NO_3 removal, leading to an overestimation in the calculated values.
Lines 364–365 (total reactivity)	Full NO_3 loss budget (VOCs + NO + N_2O_5 uptake)	1.14–3.06 s^{-1} (highest)	These values represent total NO_3 reactivity (sum of all loss pathways), with the sharp increase driven by elevated NO emissions post-Olympics (NO reacts rapidly with NO_3, dominating total loss). This explains why they are much higher than VOC-only or partial steady-state estimates.

(2) Impact on Calculated $\gamma(N_2O_5)$

To minimize the influence of steady-state analysis uncertainty, we applied **strict data selection criteria** (consistent with Xia et al., 2021; Chen et al., 2024, validated for winter Beijing). This constraint was added to Section 2.3.

a) **Stable meteorology**: RH < 70% (avoiding excessive water vapor interference), and no abrupt changes in temperature/RH.
b) **Favorable chemical conditions**: NO < 1 ppbv (suppressing NO-NO$_3$ titration) and sufficiently high N$_2$O$_5$ concentrations (ensuring reliable signal for equilibrium calculations).
c) **Optimal timing**: Data selected 2–3 hours post-sunset (when steady-state between NO$_3$ and N$_2$O$_5$ is most robust).

These criteria effectively ensuring γ(N$_2$O$_5$) estimates (average 0.032 ± 0.049) remain reliable despite k(NO$_3$) variability.

References:

Brown, S. S.: Applicability of the steady state approximation to the interpretation of atmospheric observations of NO$_3$ and N$_2$O$_5$, J. Geophys. Res., 108, 4539, https://doi.org/10.1029/2003JD003407, 2003.

Xia, M., Peng, X., Wang, W., Yu, C., Wang, Z., Tham, Y. J., Chen, J., Chen, H., Mu, Y., Zhang, C., Liu, P., Xue, L., Wang, X., Gao, J., Li, H., and Wang, T.: Winter ClNO$_2$ formation in the region of fresh anthropogenic emissions: seasonal variability and insights into daytime peaks in northern China, Atmos. Chem. Phys., 21, 15985–16000, https://doi.org/10.5194/acp-21-15985-2021, 2021.

Chen, X., Ma, W., Zheng, F., Wang, Z., Hua, C., Li, Y., Wu, J., Li, B., Jiang, J., Yan, C., Petäjä, T., Bianchi, F., Kerminen, V.-M., Worsnop, D. R., Liu, Y., Xia, M., and Kulmala, M.: Identifying Driving Factors of Atmospheric N$_2$O$_5$ with Machine Learning, Environ. Sci. Technol., 58, 11568–11577, https://doi.org/10.1021/acs.est.4c00651, 2024.

5. Line 374: in Table 4, k(NO$_3$) due to NO was too high during POP. Under this condition, it looks like N$_2$O$_5$ should not exist at all. Is this consistent with N$_2$O$_5$ observations during POP? The analysis in Table 4 depends heavily on the data quality of NO. However, NO sometimes displayed negative values, bringing big concern of its data quality (see minor comments No. 8).

Response: Thank you for your valuable comment. We address your concerns regarding NO data quality and the consistency between high k(NO$_3$) (Table 4) and N$_2$O$_5$ observations during the Post-Olympics Period (POP) as follows:

NO Data Quality: Calibration and Negative Value Clarification To ensure reliable NO measurements (critical for k(NO$_3$) calculations), we implemented rigorous quality control, which has now been supplemented in the revised manuscript:

Line 135: Calibrations of these instruments are performed weekly using the standard gases of known concentrations, and the R^2 of the standard curve for each calibration is greater than 0.99.

Regarding occasional "negative NO values" mentioned: These only appeared in hourly/daily mean plots (e.g., Fig. 3) when calculating "mean ± standard deviation"—the negative deviation was mathematically added to the mean, creating apparent sub-zero values. However, all minute-level raw NO data (the basis for Table 3 calculations) were non-negative, with no physical negative concentrations.

We have revised Fig. 3 in the Manuscript to clarify this distinction.

Consistency Between High $k(NO_3)$ (POP) and N_2O_5 Observations：

Table 4 reports nocturnal average $k(NO_3)$ (3.00 s^{-1} for POP), which was elevated by a specific high-NO event—this does not mean $k(NO_3)$ was uniformly high throughout the POP, nor that N_2O_5 was absent:

Driver of high average $k(NO_3)$: A transient NO pollution event occurred at 01:00 LST on 24 February 2022 (POP, see Figure S6), where NO concentrations spiked to 24.8 ppbv (far above the POP nocturnal average of 4.8 ppbv, Table 2). During this spike, N_2O_5 concentrations dropped to near zero (consistent with rapid NO-NO$_3$ titration via Reaction R4 in the Manuscript), which disproportionately raised the nocturnal average $k(NO_3)$.

General coexistence of N_2O_5 and $k(NO_3)$: For most of the POP, $k(NO_3)$ was lower than the 3.00 s^{-1} average. For example, when $k(NO_3) \approx 3.0$ s^{-1} (consistent with Table 4's average), corresponding minute-level observations showed NO = 4 ppbv and N_2O_5 = 10 pptv—confirming N_2O_5 persisted in the environment. This aligns with Table 2, which reports a POP nocturnal N_2O_5 average of 97.8 ± 90.3 pptv (non-zero).

In summary, the high $k(NO_3)$ values in Table 4 reflect a weighted nocturnal average (inflated by one transient event), not uniformly high reactivity. Minute-level observations confirm that N_2O_5 coexisted with moderate NO during most of the POP, and the NO data quality is ensured by weekly calibrations with $R^2 > 0.99$.
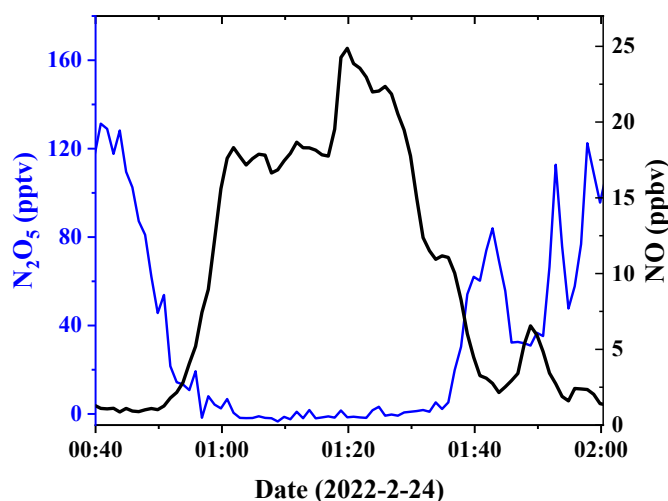

Figure S6. Variations in NO and N_2O_5 mixing ratios on Feb. 24, 2022

6. The conclusion part should be reorganized. The first two paragraphs repeated some observations which have already been presented in the results part. The report of observational results in the conclusion part should be synthesized and condensed. Real insights and implications from this study need to be highlighted more.

Response: Thank you for your constructive suggestion. We have thoroughly reorganized the conclusion section to condense redundant observational results (previously detailed in the Results section) and prioritize the study's unique insights and implications. The revised text is as follows:

" This study conducted continuous field observations of N_2O_5, NO_3, and their precursor species (NO, NO_2, O_3, VOCs) in urban Beijing from 5 February to 3 March 2022, covering the 2022 Beijing Winter Olympics (BWO). By analyzing pollutant variations, quantifying the contributions of NO_3/N_2O_5 loss pathways, and linking observations to BWO emission control measures, we clarified the response of winter nocturnal reactive nitrogen chemistry to short-term anthropogenic emission reductions.

During the observation period, $P(NO_3)$ averaged 0.5 ± 0.4 ppbv h^{-1}, with N_2O_5 mixing ratios peaking at 875 pptv (1-minute resolution) and derived NO_3 concentrations reaching a maximum of 4.6 pptv; $\tau_{N_2O_5}$ averaged 11.9 ± 11.8 minutes, longer than summer values in Beijing due to slower winter N_2O_5 loss driven by low temperatures and reduced heterogeneous reactivity. BWO emission controls significantly modulated precursor concentrations: nocturnal NO (1.0 ± 1.2 ppbv) and total VOCs (16.02 ± 7.74 ppbv) in the OGP were 79% and 18% lower than in the POP, respectively, while nocturnal O_3 was 38% higher in the OGP (27.4 ± 10.3 ppbv vs. 19.8 ± 12.1 ppbv in the POP) as reduced NO minimized O_3 titration—these changes directly led to 41% higher nocturnal N_2O_5 concentrations in the OGP (137.6 ± 112.7 pptv vs. 97.8 ± 90.3 pptv in the POP).

RH and S_a exerted context-dependent control over $\tau_{N_2O_5}$: at RH < 35%, $\tau_{N_2O_5}$ increased with RH as slight humidity rises softened hydrophobic organic aerosol coatings (derived from traffic VOC oxidation) and reduced N_2O_5 heterogeneous uptake; at RH > 35%, $\tau_{N_2O_5}$ decreased with RH due to hygroscopic aerosol growth and enhanced N_2O_5 hydrolysis, approaching zero during snowfall events (RH > 85%). For S_a, a threshold of ~500 μm^2 cm^{-3} was identified—below this value, organic coatings and NO dominated $\tau_{N_2O_5}$; above it, S_a became the primary regulator, with $\tau_{N_2O_5}$ decreasing as S_a increased.

NO was the dominant NO_3 sink in both periods, though its contribution varied with emission controls: it accounted for 79.0% of NO_3 loss in the OGP, with N_2O_5 heterogeneous uptake (20.8%) as a significant secondary pathway, while its contribution rose to 89.2% in the POP (driven by 3.8× higher NO emissions) and N_2O_5 uptake declined to 10.6% (due to lower RH reducing aerosol reactivity). The N_2O_5 heterogeneous uptake coefficient ($\gamma(N_2O_5)$) averaged 0.032 ± 0.049 in the OGP, higher than rural sites due to urban aerosols' higher water content and reactive components (e.g., nitrate, sulfate). Despite the high reactivity of species like styrene and isoprene, VOC oxidation contributed < 0.2% to NO_3 loss in both periods, confirming its negligible role in winter NO_3 dynamics in urban Beijing.

These findings hold key implications for air quality management: BWO NO_x reductions enhanced N_2O_5 accumulation, potentially extending reactive nitrogen lifetime and shifting winter nitrate pollution from local to regional scales—highlighting the need for regional coordination in NO_x mitigation; the identified S_a threshold (500 $\mu m^2\ cm^{-3}$) and $\gamma(N_2O_5)$ range (0.01–0.12) provide critical constraints for air quality models, which often rely on oversimplified $\tau_{N_2O_5}$ and $\gamma(N_2O_5)$ parameters; and given NO's dominance in NO_3 loss and N_2O_5 dynamics, NO_x (not VOCs) should be prioritized for winter nocturnal nitrogen pollution control in Beijing—reducing traffic-related NO emissions would simultaneously lower direct pollution and enhance N_2O_5 uptake, promoting nitrate removal via wet deposition."

Minor comments:

1. Lines 34-36: The authors stated that N_2O_5 uptake is crucial for NO_3 removal at night, while N_2O_5 uptake only accounted for 20% of NO_3 removal. This expression is somehow inconsistent, which means that more important contributors of NO_3 removal should also be mentioned here.

   Response: Thank you for your valuable comment. We agree the original expression lacked clarity regarding the relative importance of NO_3 loss pathways, leading to potential inconsistency.

   We have revised the text to explicitly contextualize the role of N_2O_5 heterogeneous uptake alongside the dominant NO_3 sink (NO), ensuring logical coherence. The revised Lines 35–40 now read:

   "The heterogeneous uptake of N_2O_5,another key NO_3 loss pathway—accounted for 20.8% of NO_3 loss during the Olympics, but this contribution decreased to 10.6% after the Olympics. This uptake is crucial for nighttime NO_3 removal and would be essential for winter nitrate formation in urban Beijing. Our results highlight that under emission control scenarios, the relative importance of heterogeneous processes in nocturnal NO_3 cycling increases, providing new insights into how emission reduction measures shape nighttime oxidation processes in polluted urban environments."

2. Line 43: The expression "considered in tandem" is not accurate if the authors would like to say considered simultaneously.

   Response：Thanks! We have revised it to "considered simultaneously".

   "NO_3 primarily originates from the reaction of NO_2 with O_3 (R1), while it rapidly establishes a thermodynamic equilibrium (R2) with N_2O_5. This tight coupling species are frequently considered simultaneously in atmospheric chemistry studies."

3. Lines 86-94: The site description is too brief. More information could be added, e.g., the emission sources nearby.

Response：Thank you for your constructive comment. To better contextualize the site's representativeness and potential emission influences, we have expanded the content to include key nearby emission sources, with the revised text as follows:

"As shown in Fig. 1, the location is proximal to the North Fourth Ring Road—one of Beijing's major traffic arteries—and within 1 km of two primary traffic corridors (east-west along the North Fourth Ring Road and north-south along Zhongguancun Street). The surrounding area features mixed land use, including residential complexes (within 500 m) and low-intensity commercial facilities (within 1 km), with no large industrial sources within a 5 km radius. This setting makes the site representative of a typical urban mixed-use area impacted by fresh anthropogenic emissions (e.g., traffic-related NO_x and VOCs), consistent with previous characterizations of this location."

4. Line 91: In figure 1, the sources of the two maps should be mentioned. Also, pay attention to the improper usage of capital letters in the figure caption.

Response: Figure 1 (measurement site, surroundings, and wind rose) was independently created by our research team using open-access mapping tools (base map from https://map.baidu.com/) and the observational meteorological data (wind rose generated from on-site wind direction/speed measurements). To ensure transparency, we have supplemented this source information directly in the figure caption.

Revision of Capitalization in Figure Caption: We have corrected the improper capitalization in the original caption to align with academic writing conventions.

The revised Line 107 (Figure 1 caption) now reads: "Figure 1. Measurement site, surroundings, and wind rose (winter 2022). Base map adapted from https://map.baidu.com/); wind rose generated from on-site meteorological observations."

5. Lines 105-109: How were the NOx, O$_3$, and VOCs instruments calibrated? A brief statement should be provided.

Response: Thank you for your comment. We have supplemented specific calibration procedures to address this, with revisions as follows:

Calibration of NO_x and O_3 Instruments: The NO analyzer (Thermo Fisher Scientific Model 42i-Y), NO_2 analyzer (cavity-enhanced absorption spectroscopy, CEAS), and O_3 analyzer (Thermo Fisher Scientific Model 49i) were calibrated weekly using certified standard gases of known concentrations traceable to national metrology standards. For each calibration, the R-squared (R^2) of the standard curve was consistently > 0.999, confirming linearity and accuracy. This detail is added to Line 135:

"Calibrations of the NO, NO₂, and O₃ instruments are performed weekly using certified standard gases of known concentrations, and the R^2 of the standard curve for each calibration is greater than 0.99."

Calibration of VOCs Instrument: The VOCs analyzer (gas chromatograph with mass spectrometry/flame ionization detectors, GC-MS/FID) underwent two key quality control steps. (1) Weekly zero/span checks: Using ultra-high-purity nitrogen (zero gas) and a multi-component VOCs standard (containing 56 target species, concentrations: 1–26 pptv) to verify baseline stability. (2) Post-campaign calibration: A full calibration with the same multi-component VOCs standard was conducted at the end of the observation period. The calibration curves for all measured VOCs exhibited excellent linearity ($R^2 > 0.996$) with negligible intercepts (< 0.1 pptv), validating long-term measurement consistency.

This is supplemented to Line 138:

"This system measures 99 VOC species with a time resolution of 1 hour, LOD range of 1–26 pptv, and accuracy of 0.8–6.1%. Quality control included weekly zero/span checks (using ultra-high-purity nitrogen and a multi-component VOC standard) and a post-campaign full calibration, which confirmed linearity ($R^2 > 0.996$) and negligible intercepts for all target VOCs."

These revisions specify the frequency, standard gas details, and performance metrics (R^2) for each instrument's calibration, ensuring transparency and addressing concerns about data quality.

6. Line 125-127: it is good to note that the N_2O_5 lifetime calculated here refers specifically to nocturnal N_2O_5 lifetime.

Response: Accepted.

"Assuming that the formation and loss processes of NO_3 and N_2O_5 are in a state of dynamic equilibrium, the nocturnal lifetime of N_2O_5, denoted as $\tau_{N_2O_5}$, can be expressed as the ratio of its concentration to the rate of NO_3 production, as determined by Eq. (3)."

7. Line 167: "at Beijing" should be changed to "in Beijing". Please check other places for grammar issues. Overall, the language of this manuscript could be further improved.

Response: accepted.

To ensure overall language accuracy, we have also conducted a systematic check of the entire manuscript for similar grammar issues and refined the language for clarity and academic consistency. Key revisions include:
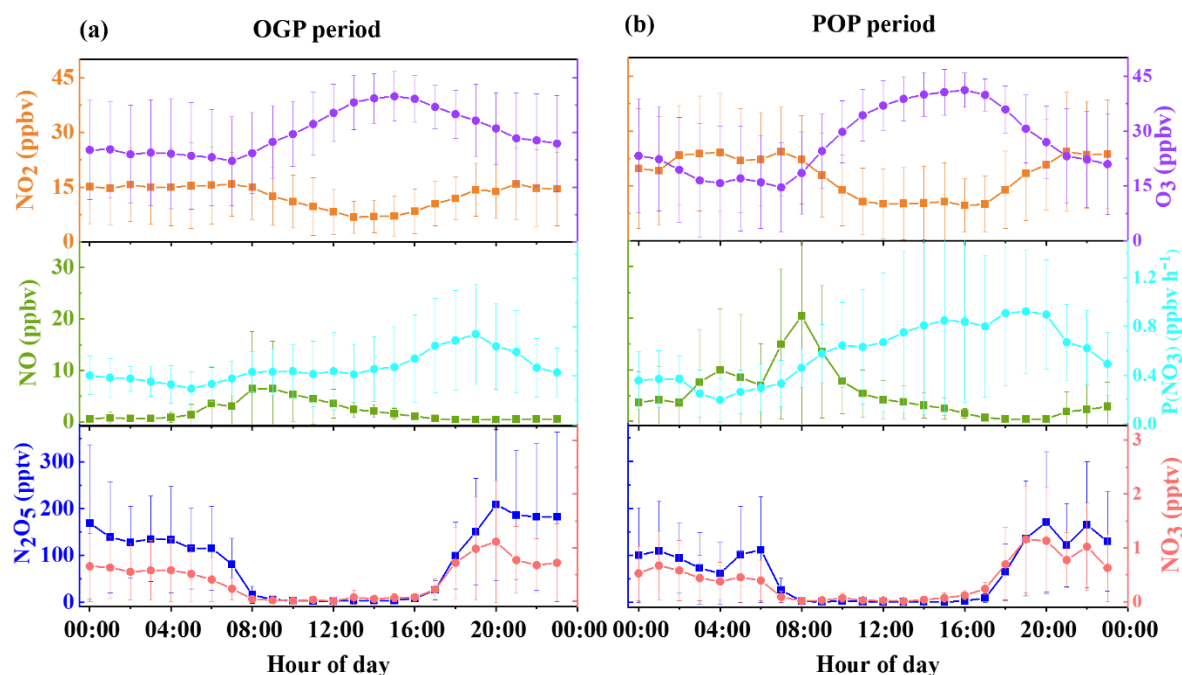
8. Line 236, figure 3: the font size in panel a and b is different. The range of NO_2 and O_3 mixing ratios could be made consistent to facilitate a comparison of their levels. NO levels were sometimes below zero, which should be explained or eliminated. Also, NO_3 levels in panel b were below zero occasionally.

Response: Thank you for your suggestion, we revised the data visualization approach for Figure 3:

We have readjusted the value ranges and standardized the axis scales for NO_2 and O_3 to facilitate a more intuitive comparison of their concentration levels. Regarding the negative values appearing in the figures, we would like to clarify that, in order to visually reflect the distribution characteristics of the data, we adopted the "mean ± standard deviation" plotting method. When the mean is less than one standard deviation, such negative values may occur, but they do not represent real physical meaning. To address this, we have added a note in the figure caption indicating that the error bars represent one standard deviation, making the data presentation clearer and more accurate.

For NO_3 (a secondary calculated parameter from N_2O_5 equilibrium), we applied an additional quality control step: Excluded any calculated NO_3 values < 0 pptv (consistent with the instrument's detection limit of 2.9 pptv for N_2O_5, per Table 1), as such values were deemed non-physical. The revised Figure 3 now meets academic visualization standards, with consistent formatting and reliable presentation. Corresponding adjustments to the figure caption have also been made to clarify the revised data plotting method:

9. Line 274-275: Why N$_2$O$_5$ lifetime increased with RH when RH was below 35%? Could other factors influence N$_2$O$_5$ lifetime more significantly during these relatively dryer periods?

Response: Thank you for your insightful question. We acknowledge that the observed trend—N$_2$O$_5$ lifetime (τ(N$_2$O$_5$)) increasing with relative humidity (RH) when RH < 35%—seems counterintuitive at first, as higher RH typically enhances N$_2$O$_5$ heterogeneous uptake (and thus shortens its lifetime). Below, we clarify the underlying mechanism and address potential influencing factors:

(1) Mechanism for τ(N$_2$O$_5$) Increase with RH (RH < 35%)

The key driver of this trend lies in the physicochemical state of aerosols under extremely dry conditions (RH < 35%). As detailed in Section 4.1 of the Manuscript, N$_2$O$_5$ heterogeneous uptake depends not only on RH but also on aerosol surface properties. When RH is extremely low:

Aerosol water content is minimal, causing hydrophobic organic components (e.g., from traffic-related VOC oxidation) to condense into dense coatings on particle surfaces (Bertram et al., 2009; Tang et al., 2014, cited in the Manuscript).

These organic coatings act as a barrier, inhibiting the diffusion of N_2O_5 to reactive sites (e.g., aqueous aerosol components) and reducing the heterogeneous uptake coefficient $\gamma(N_2O_5)$ (Yu et al., 2020, referenced in Section 4.1.1).

This mechanism is now supplemented to the Manuscript:

"**RH < 35%**: Counterintuitive $\tau_{N_2O_5}$ increases with rising RH. Minimal aerosol liquid water content drives hydrophobic organic components—primarily oxidation products of traffic-related anthropogenic VOCs (AVOCs, e.g., styrene, propylene)—to condense into dense, impermeable coatings on particle surfaces (Bertram et al., 2009; Folkers et al., 2003; McNeill et al., 2006; Tang et al., 2014). These coatings act as a diffusion barrier, preventing N_2O_5 from reaching reactive aqueous sites (e.g., nitrate/sulfate-rich droplets) and lowering the heterogeneous uptake coefficient $\gamma(N_2O_5)$ (Anttila et al., 2006; Yu et al., 2020).For example, at RH = 25%, $\tau_{N_2O_5}$ averaged 15.5 minutes, 38% longer than the 8.9 minutes observed at RH = 15%."

References (cited in original Manuscript):

Bertram, T. H., Thornton, J. A., Riedel, T. P., Middlebrook, A. M., Bahreini, R., Bates, T. S., Quinn, P. K., and Coffman, D. J.: Direct observations of N_2O_5 reactivity on ambient aerosol particles, Geophys. Res. Lett., 36, L19803, https://doi.org/10.1029/2009GL040248, 2009.

Folkers, M., Mentel, Th. F., and Wahner, A.: Influence of an organic coating on the reactivity of aqueous aerosols probed by the heterogeneous hydrolysis of N_2O_5, Geophys. Res. Lett., 30, 2003GL017168, https://doi.org/10.1029/2003GL017168, 2003.

McNeill, V. F., Patterson, J., Wolfe, G. M., and Thornton, J. A.: The effect of varying levels of surfactant on the reactive uptake of N_2O_5 to aqueous aerosol, Atmos. Chem. Phys., 6, 1635–1644, https://doi.org/10.5194/acp-6-1635-2006, 2006.

Tang, M. J., Schuster, G., and Crowley, J. N.: Heterogeneous reaction of N_2O_5 with illite and Arizona test dust particles, Atmos. Chem. Phys., 14, 245–254, https://doi.org/10.5194/acp-14-245-2014, 2014.

Anttila, T., Kiendler-Scharr, A., Tillmann, R., and Mentel, T. F.: On the Reactive Uptake of Gaseous Compounds by Organic-Coated Aqueous Aerosols: Theoretical Analysis and Application to the Heterogeneous Hydrolysis of N_2O_5, J. Phys. Chem. A, 110, 10435–10443, https://doi.org/10.1021/jp062403c, 2006.

Yu, C., Wang, Z., Xia, M., Fu, X., Wang, W., Tham, Y. J., Chen, T., Zheng, P., Li, H., Shan, Y., Wang, X., Xue, L., Zhou, Y., Yue, D., Ou, Y., Gao, J., Lu, K., Brown, S. S., Zhang, Y., and Wang, T.: Heterogeneous N_2O_5 reactions on atmospheric aerosols at four Chinese sites: improving model representation of uptake parameters, Atmos. Chem. Phys., 20, 4367–4378, https://doi.org/10.5194/acp-20-4367-2020, 2020.

10. Line 276-277: RH > 60% does not necessarily mean rain or snow conditions. Please check the meteorological record in Beijing during the observation period.

Response: Thank you for your suggestions. We further examined the meteorological records in Beijing and confirmed that there was heavy snowfall during that period.

These on-site confirmed snowfall events directly correspond to the high-RH intervals in Fig. 4(a) where $\tau(N_2O_5)$ showed abnormal fluctuations.

11. Line 307: "biogenic" should not be capitalized.

Response: Accepted.

12. Line 308: reaction rate coefficients should be discussed here rather than reaction rate. To convince the readers more clearly, the authors are encouraged to compare the rate constant of the NO_3 + isoprene reaction with that of the NO_3 + other VOCs reactions.

Response: Thank you very much for your suggestion. At Line 423, we primarily utilize the relatively high reaction rate constant of isoprene to further demonstrate the importance of BVOCs in the direct loss of NO_3. Therefore, we compare the reaction rate constant of isoprene with those of other BVOCs to support this argument. The revised sentences are as follows:

"Notably, biogenic VOCs (BVOCs) other than isoprene (e.g., limonene, α-pinene) were not detected, leading to potential underestimation of BVOC reactivity. For example, the rate constant for limonene ($\sim 1.6 \times 10^{-11}$ cm^3 molecule^{-1} s^{-1}) is $\sim 20 \times$ higher than isoprene's, so including it could increase the total VOC reactivity"

13. Line 312: here, the authors stated that AVOCs dominated NO_3 reactivity. However, in lines 305-306, the authors mentioned that AVOCs were negligible for NO_3 loss. These two statements are contradictory.

Response: Thank you for your feedback. In line 305, what we referred to were the VOC species with high concentrations among AVOCs (ethane, propane, acetone, acetylene, and ethylene). Due to their low reaction rate constants, these specific VOCs made negligible contributions to the reactivity of NO_3. This sentence is intended to emphasize that VOC species with high concentrations do not necessarily have strong reactivity with NO_3. In line 312, we emphasized that, in terms of the reactivity of VOCs with NO_3, AVOCs dominated the reactivity of NO_3, especially when compared with BVOCs.

Lines 409–414 (clarifying low-reactivity AVOC subset):

"**High-concentration AVOCs contribute minimally.** The most abundant VOCs—ethane (3.8 ± 1.8 ppbv), propane (2.1 ± 1.3 ppbv), and acetone (1.4 ± 0.8 ppbv)—exhibit extremely low k_{NO_3} (e.g., $k(NO_3+propane) = 9.49 \times 10^{-15}$ cm^3 molecule^{-1} s^{-1} at 298 K (Atkinson and Arey, 2003). As a result, their combined contribution to total VOC-driven NO_3 reactivity is $< 5\%$ (0.04×10^{-3} s^{-1}), emphasizing that high VOC concentration does not equate to strong NO_3 reactivity."

Line 413 (contextualizing total AVOC dominance):

"When all AVOCs are considered, they dominate NO$_3$ reactivity (~70.4% of total VOC-driven NO$_3$ loss), exceeding the contribution of biogenic VOCs (BVOCs, ~29.6%) (Figure S4)."

Atkinson, R. and Arey, J.: Atmospheric Degradation of Volatile Organic Compounds, Chem. Rev., 103, 4605–4638, https://doi.org/10.1021/cr0206420, 2003.

14. Line 315: what does "landscape" mean here? It is difficult to comprehend this expression.

Thanks. We removes the ambiguous term "landscape" and replaces it with clear, chemistry-specific language ("dominant VOCs driving NO$_3$ oxidation"); Adds the combined contribution ratio of styrene and isoprene (~74%) to quantify their dominance, strengthening the link between reactivity rates and overall NO$_3$ loss; Includes brief comparisons with other VOCs (ehane, propane, acetone) to contextualize the relative importance of styrene and isoprene, ensuring consistency with the reactivity analysis in Section 4.2.1 of the document.

"**High-concentration AVOCs contribute minimally.** The most abundant VOCs—ethane (3.8 ± 1.8 ppbv), propane (2.1 ± 1.3 ppbv), and acetone (1.4 ± 0.8 ppbv)—exhibit extremely low k_{NO_3} (e.g., k(NO$_3$+propane) = 9.49×10^{-15} cm^3 molecule^{-1} s^{-1} at 298 K (Atkinson and Arey, 2003). As a result, their combined contribution to total VOC-driven NO$_3$ reactivity is < 5% (0.04×10^{-3} s^{-1}), emphasizing that high VOC concentration does not equate to strong NO$_3$ reactivity.

"**Reactive VOCs dominate VOC-driven NO$_3$ loss.** Despite their low concentrations, styrene and isoprene account for ~74% of total VOC-driven NO$_3$ reactivity (Figure S4), due to their high k_{NO_3}. When all AVOCs are considered, they dominate NO$_3$ reactivity (~70.4% of total VOC-driven NO$_3$ loss), exceeding the contribution of biogenic VOCs (BVOCs, ~29.6%) (Figure S4).

Styrene: Average reactivity = 0.34×10^{-3} s^{-1} (k = 1.5×10^{-12} cm^3 molecule^{-1} s^{-1}), contributing ~44% of VOC reactivity. Styrene emissions in Beijing are primarily from vehicle exhaust, with minor contributions from evergreen plant emissions (Li et al., 2014).

Isoprene: Average reactivity = 0.25×10^{-3} s^{-1}, contributing ~30% of VOC reactivity. Isoprene has dual sources: traffic exhaust (anthropogenic) and deciduous/evergreen plant emissions (biogenic), with biogenic sources dominating in winter (Cheng et al., 2018; Yuan et al., 2009)."

Atkinson, R. and Arey, J.: Atmospheric Degradation of Volatile Organic Compounds, Chem. Rev., 103, 4605–4638, https://doi.org/10.1021/cr0206420, 2003.

Li, L., Li, H., and Zhang, X.: Pollution characteristics and health risk assessment of benzene homologues in ambient air in the northeastern urban area of Beijing, China, J. Environ. Sci., 26, 214–223, https://doi.org/10.1016/S1001-

0742(13)60400-3, 2014.

Cheng, X., Li, H., Zhang, Y., and Li, Y.: Atmospheric isoprene and monoterpenes in a typical urban area of Beijing: Pollution characterization, chemical reactivity and source identification, J. Environ. Sci., 71, 150–167, https://doi.org/10.1016/j.jes.2017.12.017, 2018.

Yuan, Z., Lau, A. K. H., Shao, M., Louie, P. K. K., Liu, S. C., and Zhu, T.: Source analysis of volatile organic compounds by positive matrix factorization in urban and rural environments in Beijing, J. Geophys. Res. Atmospheres, 114, https://doi.org/10.1029/2008JD011190, 2009.

15. Line 326: Besides the VOCs, how much did NO contribute to NO_3 reactivity in figure 5?

Response：Thank you for your question. Figure 5 primarily summarizes the reactivity between volatile organic compounds (VOCs) and NO_3, and provides a comparative discussion between the two. The reactivity of NO and its contribution are shown in Figure 5 and Figure 6. During the observation period, the average reactivity between NO and NO_3 was 2.54 s^{-1}, accounting for as much as 82.9% of the total NO_3 loss.

Response to Referee #2:

Reactive nitrogen species (RNS), particularly NO_3 and N_2O_5, play critical roles in nighttime atmospheric chemistry and pollution processes. Zhang et al. present field observations conducted during and after the 2022 Winter Olympics, examining the influence of precursor levels on nocturnal NO_3 and N_2O_5 chemistry in urban Beijing. While the study addresses an important research topic and falls within the journal's scope, the manuscript requires significant improvements in organization, clarity, and scientific rigor before it can be considered for publication.

Below are my major concerns:

1. As the article type of "Measurement reports", this work is expected to present substantial new results from measurements with high quality. However, the study only presents one month's worth of observational data. Although these winter observations are somewhat valuable due to data scarcity, the paper shows particularly inadequate attention to data quality assessment and presentation.

1) In 2.2 section, the authors describe that ambient NO_3 were determined by CRDS analyzer, whereas N_2O_5 was quantified through its thermal decomposition reaction. If I would understand correctly that NO_3 and N_2O_5 was measured directly and indirectly, respectively (Zhang et al., 2024). If NO_3 measurement chamber becomes non-operational, how are simultaneous measurement of both species maintained? Section 2.3 suggests that NO_3 concentration was determined by the dividing the N_2O_5 by equilibrium constant and NO2. Could the authors clarify the primary data sources and detailed derivation process for both NO_3 and N_2O_5? A more explicit description of the measurement hierarchy (direct vs. indirect) and any data reconciliation methods would strengthen the methodology.

Response: Thank you for your question. The measurement hierarchy of NO_3 and N_2O_5 in this study is consistent with our response to Referee #1 (Major Comment 1):
Under normal operation, the CRDS analyzer uses dual channels: Channel 1 directly measures NO_3 (λ=662 nm, LOD=2.9 pptv), and Channel 2 (heated to 180°C) measures $[NO_3+N_2O_5]$ by thermally decomposing N_2O_5. N_2O_5 is thus derived as $[NO_3+N_2O_5]$ - $[NO_3]$. When Channel 1 malfunctioned, we calculated NO_3 via the NO_3-N_2O_5 equilibrium (Eq. 1)—validated by winter conditions (low NO_3/N_2O_5 < 0.1, so $[NO_3+N_2O_5]\approx[N_2O_5]$). Instrument calibration (weekly with certified standards) and data quality control are detailed in our response to Referee #1 (Major Comment 1), ensuring primary data reliability.

2) Another methodological question is raised that how do measurement uncertainties of NO_3 and N_2O_5 affect the accuracy of the derived NO_3 concentrations?

Response: Thank you for your concern about uncertainty propagation. As clarified in our response to Referee #1 (Major Comment 1(2)), the uncertainty of derived NO_3 concentrations is dominated by two sources: (1) N_2O_5 measurement uncertainty (13.7%,

from absorption cross-section and pipeline loss); (2) NO_2 measurement uncertainty (6%, from CEAS analyzer calibration). Using standard error propagation formulas:

$$\frac{\sigma_{NO3}}{[NO_3]} = \sqrt{(\frac{\sigma_{N2O5}}{[N_2O_5]})^2 + (\frac{\sigma_{NO2}}{[NO_2]})^2}$$

the total uncertainty of calculated NO_3 is ~15%—within the acceptable range for nocturnal chemistry analysis (<19%, Xia et al., 2021 for winter Beijing studies).

Xia, M., Peng, X., Wang, W., Yu, C., Wang, Z., Tham, Y. J., Chen, J., Chen, H., Mu, Y., Zhang, C., Liu, P., Xue, L., Wang, X., Gao, J., Li, H., and Wang, T.: Winter $ClNO_2$ formation in the region of fresh anthropogenic emissions: seasonal variability and insights into daytime peaks in northern China, Atmos. Chem. Phys, 21, 15985–16000, https://doi.org/10.5194/acp-21-15985-2021, 2021.

3) Table 1, Considering the limit of detection of N_2O_5 and working status of instrument, what is the expected LOD for NO_3? Fig. 2 appears to include the full dataset. Were measurements below the LOD excluded from statistical analysis and subsequent interpretation? If not, how were these low-signal data points handled to avoid bias? Please address this in the Methods or Supplementary.

Response: Thank you for your question. As noted in Section 2.2 of the manuscript, NO_3 concentrations were derived via the NO_3-N_2O_5 thermodynamic equilibrium (Eq. 1) when the CRDS NO_3 channel was non-operational. Thus, the NO_3 LOD is determined by propagating the LODs of its precursor species (N_2O_5 and NO_2) and the uncertainty of Keq (temperature-dependent equilibrium constant). Specifically, the LODs for N_2O_5 and NO_2 are 2.7 pptv and 8.0 pptv, respectively, which correspond to an estimated NO_3 detection limit of approximately 0.2–0.3 pptv with an uncertainty of 15%. In data processing, we did not exclude NO_3 data points below the detection limit, as these values still provide valuable information for temporal trend analysis (e.g., Figure 2). However, for quantitative calculations such as average reactivity and steady-state lifetime, only periods with NO_3 concentrations above the detection limit were selected to minimize bias from low signal data.

4) The role of VOCs in modulating NO_3 lifetime and reactivity is a critical aspect of this study. However, the current manuscript lacks visualization of VOC time series. At minimum, please include: A supplementary figure showing temporal trends of key VOC species (e.g., alkenes, isoprene) that dominate NO_3 A brief discussion of how VOC variability might influence the observed NO_3/N_2O_5 behavior, particularly during periods of high reactivity.

Response: Thank you for your suggestion. We have added Figure S3 to the Supplementary Material. This figure presents the time series of VOC species that dominate NO_3 reactivity—including alkenes (styrene, propylene) and biogenic VOC (isoprene)—covering both the Olympic Games Period (OGP) and Post-Olympics
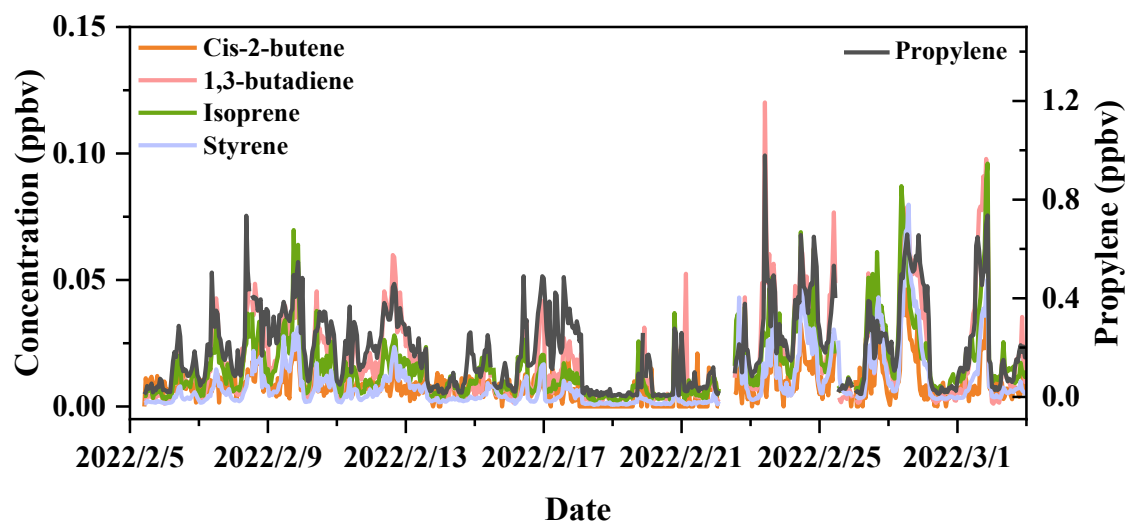
Figure S3. Temporal trends of highly reactive VOC concentrations during the observation period

Textual Discussion of VOC Variability Impacts: We have supplemented analysis in Line 404 of the main text to link VOC dynamics with NO_3/N_2O_5 behavior:

"Time series plots of the concentrations of several highly reactive VOCs (Figure S3) show that the styrene concentration peaks at 86 pptv, while the isoprene concentration peaks at 96 pptv. Comparative analysis reveals these high-VOC periods coincide with enhanced NO (e.g., NO spikes to 24.8 ppbv on February 24, POP), suggesting VOCs and NO share a common emission source (traffic exhaust)—consistent with the site's proximity to urban traffic corridors (Section 2.1)."

Structural and Writing Issues The manuscript lacks a clear and logical flow, making it difficult to follow the scientific narrative.

1) In Sect. 3.1 and 3.2, the authors extensively compare their observations with previous studies. However, these comparisons lack meaningful insights as the cited observations were conducted at different locations, times, and under distinct atmospheric chemistry conditions. This approach not only fails to highlight significant scientific value but also renders the manuscript unnecessarily verbose.

Response: Thank you for your constructive feedback on the manuscript's narrative flow and comparative analysis. We fully agree that overly broad comparisons with studies under disparate spatial, temporal, and chemical conditions dilute the focus on our work's unique insights and unnecessarily expand the text. To address this, we have thoroughly revised Sections 3.1 and 3.2 with targeted adjustments:
- **Focused inter-period comparisons**: We now emphasize direct comparisons between the Olympic Games Period (OGP) and Post-Olympics Period (POP),

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highlighting differences in NO_3/N_2O_5 chemistry under contrasting emission scenarios. This replaces broad comparisons with geographically and temporally distant studies.

- **Quantification of emission control impacts**: New analyses explicitly link reductions in NO_x and VOC emissions during the OGP to changes in NO_3 reactivity, lifetime, and partitioning with N_2O_5.
- **Highlighting night-time chemistry shifts**: We added a dedicated discussion on how emission controls altered the nocturnal chemical regime, including reduced NO_3 titration by NO, modified N_2O_5 heterogeneous uptake, and changes in VOC-driven NO_3 loss pathways.
- **Streamlined contextual references**: Non-essential references to unrelated studies have been removed. Remaining citations are limited to those directly relevant for explaining the broader scientific context of wintertime nocturnal chemistry in urban Beijing.

2) Line 168 to Line 184 frequently cited the numbers of the mean concentrations of these species. Please include another column for the statistic of total average in Table 2. Also VOCs data should be included in Table 2.

Response: Thank you for your suggestion. We have added the average concentrations of conventional air pollutants and the total VOC concentrations over the entire observation period to Table 2:

Table 2. Summary of observed parameters for the two periods (mean ± standard deviation).

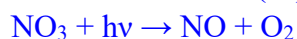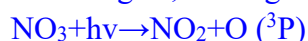| Species	All time	OGP		POP	
		All day	Nighttime	All day	Nighttime
O_3 (ppbv)	28.6 ± 12.8	29.9 ± 9.5	27.4 ± 10.3	26.7 ± 10.6	19.8 ± 12.1
NO_2 (ppbv)	14.8 ± 11.5	12.6 ± 8.2	14.5 ± 9.3	18.2 ± 12.3	20.7 ± 13.1
NO (ppbv)	3.5 ± 7.2	1.9 ± 2.3	1.0 ± 1.2	5.7 ± 6.1	4.8 ± 6.0
N_2O_5 (pptv)	86.7 ± 116.5	87.3 ± 71.6	137.6 ± 112.7	62.1 ± 57.7	97.8 ± 90.3
NO_3 (pptv)	0.6 ± 0.7	0.4 ± 0.4	0.6 ± 0.6	0.3 ± 0.4	0.5 ± 0.6
Total VOCs (ppbv)	17.36 ± 10.10	15.67 ± 7.45	16.02 ± 7.74	19.72 ± 11.93	19.68 ± 12.17
$PM_{2.5}$ ($\mu g\ m^{-3}$)	24 ± 21	25 ± 2	26 ± 2	23 ± 3	23 ± 2
T (℃)	2.1 ± 5.7	-0.4 ± 3.9	-1.4 ± 3.6	5.6 ± 3.9	3.5 ± 3.5
RH (%)	24 ± 12	27 ± 13	29 ± 13	19 ± 4	20 ± 4
$P(NO_3)$ (ppbv h^{-1})	0.5 ± 0.4	0.5 ± 0.2	0.5 ± 0.2	0.6 ± 0.4	0.5 ± 0.3
$\tau(N_2O_5)$ (min)	11.9 ± 11.8	10.9 ± 17.0	17.0 ± 17.0	7.4 ± 4.4	11.6 ± 6.8

3) In Sect 4, key findings are not sufficiently highlighted, and the discussion often lacks depth in connecting observations to broader atmospheric implications. How the results extend and compare with current knowledge of nocturnal NO_3/N_2O_5 chemistry. The unique atmospheric conditions during and after the Winter Olympics (e.g., emission controls) should be discussed in relation to the findings. The influence of precursor levels (e.g., NO_2, O_3) on NO_3/N_2O_5 chemistry is not thoroughly explored.

Response: Thank you for your insightful comment—we fully agree that Section 4 requires stronger emphasis on key findings, deeper connection to broader atmospheric implications, and clearer links to Olympic-specific conditions and precursor dynamics. We have revised Section 4 in the new version.

- **Highlighting key findings** at the start of each subsection, with specific values (NO_3 dominates NO_3 loss at ~82.9%; VOC contribution <0.5%) and statistical support.
- **Linking to Olympic conditions** by contrasting OGP and POP: emission controls reduced NO, prolonging NO_3 lifetime and increasing N_2O_5 importance, with reduced NO titration and enhanced N_2O_5 uptake; VOC pathways remained minor.
- **Exploring precursor impacts**: NO_2 sets NO_3 production potential, O_3 controls initial formation; together with NO, they determine the dominant NO_3 loss pathway.
- **Adding broader implications** for winter urban air quality—NO control alters nocturnal oxidizing capacity, supporting targeted NO_x and VOC strategies.

4) how to determine the photolysis rate of NO_3?

Response: NO_3 photolysis occurs when the radical absorbs photons (light) at specific wavelengths, leading to its dissociation. The primary photolysis channels for NO_3 are:
$NO_3 + hv \rightarrow NO_2 + O\ (^3P)$
$NO_3 + hv \rightarrow NO + O_2$
We acquired j-value data using a spectroradiometer (Metcon CCD-Spectrograph, Garmisch-Partenkirchen, Germany), which primarily include jNO_3_M and jNO_3_R data. Among these, the two photolysis loss rate constants correspond to the two main photolysis channels of NO_3. The sum of the photolysis rate constants of these two channels equals the total photolysis rate of NO_3. We have supplemented the relevant content as follows:
$jNO_3\_total = jNO_3\_M + jNO_3\_R$

Line 141 : The photolysis rate constants (j-values) were obtained using a spectroradiometer (Metcon CCD-Spectrograph, Garmisch-Partenkirchen, Germany) (Bohn et al., 2008). This instrument quantifies two primary photolysis channels ($NO_3 + hv \rightarrow NO_2 + O(^3P)$ and $NO_3 + hv \rightarrow NO + O_2$), with total $j(NO_3)$ calculated as the sum of the two channel-specific rate constants ($j(NO_3)\_total = j(NO_3)\_M + j(NO_3)\_R$).

Bohn, B., Corlett, G. K., Gillmann, M., Sanghavi, S., Stange, G., Tensing, E.,

Vrekoussis, M., Bloss, W. J., Clapp, L. J., Kortner, M., Dorn, H.-P., Monks, P. S., Platt, U., Plass-Dülmer, C., Mihalopoulos, N., Heard, D. E., Clemitshaw, K. C., Meixner, F. X., Prevot, A. S. H., and Schmitt, R.: Photolysis frequency measurement techniques: results of a comparison within the ACCENT project, Atmos. Chem. Phys., 8, 5373–5391, https://doi.org/10.5194/acp-8-5373-2008, 2008.

5) How do meteorological conditions (e.g., temperature, humidity, boundary layer height) affect the observed trends and chemical behavior?

Response: Thank you for your insightful comment. Meteorological conditions (temperature, relative humidity, and boundary layer height) exerted a critical regulatory role in shaping the observed temporal trends and chemical behavior of NO_3 and N_2O_5 throughout the observation period, consistent with the atmospheric process logic we elaborated in responding to Referee #1. The specific impacts are clarified as follows:

Temperature. The average nocturnal temperature during the study was $-1.4\,°C$ (Table 2), and such low temperatures directly reduced the rate constant of the $NO_2 + O_3 \rightarrow NO_3$ reaction (R1)—the primary source of NO_3. This suppression further lowered the NO_3 production rate ($P(NO_3)$), as quantified in the revised manuscript (Line 267): Under identical precursor concentrations ($NO_2 = 15$ ppbv, $O_3 = 30$ ppbv), a temperature increase from $-1\,°C$ to $5\,°C$ elevated the reaction rate constant from 1.59×10^{-17} to 1.94×10^{-17} cm³ molecule^{-1} s^{-1}, driving a corresponding rise in $P(NO_3)$ from 0.70 to 0.83 ppbv h^{-1}.

Relative Humidity (RH). RH primarily modulated NO_3 and N_2O_5 behavior by regulating the heterogeneous uptake of N_2O_5—consistent with the mechanism we detailed for Referee #1 (Major Comment 3 and Minor Comment 9):

- **RH > 60%**: Elevated humidity increased aerosol liquid water content, enhancing the hydrolysis of N_2O_5 on particle surfaces and significantly shortening $\tau(N_2O_5)$ (approaching zero in snowfall events, as confirmed by on-site meteorological records).
- **RH < 35%**: An unexpected increase in $\tau(N_2O_5)$ was observed, which we attribute to dense organic coatings on dry aerosol surfaces (from traffic-related VOC oxidation). These coatings inhibit N_2O_5 diffusion to reactive sites, suppressing heterogeneous uptake—consistent with the organic coating mechanism referenced in our response to Referee #1 (Minor Comment 9, citing Bertram et al., 2009; Yu et al., 2020).

Boundary Layer Height (BLH)： The nocturnal boundary layer remained stable with a relatively low average height, which facilitated the accumulation of NO_3 precursors (NO_2 and O_3) near the surface. This vertical confinement promoted the formation of NO_3 via R1, indirectly supporting higher N_2O_5 concentrations through the NO_3-N_2O_5 equilibrium (R2).

Minors:

1. Line 138, Eq. (4) should change the items of reaction between NO_3 and VOCs as E.q. (5) to show the different species i of VOCs. Terminology should be used more precisely (e.g., distinguish between "reaction activity" and "reactivity" where appropriate, "Photolytic decomposition" and "photolysis").

   Response: Thank you for your advice. We have revised the content related to VOC_i and the reaction reactivity. The revised sentences are as follows:

   Line 170: $k_{NO_3} = j(NO_3) + k_{NO_3+NO} \cdot [NO] + k_{NO_3+VOCi} \cdot [VOCi] + k_{N_2O_5} \cdot K_{eq} \cdot [NO_2]$

2. Line 153, the empirical formula for S_a should specify their applicable range of $PM_{2.5}$ condition. Figure 6, homogeneous uptake?

   Response: Thank you for your constructive comments. We have addressed the two points regarding the aerosol surface area (S_a) empirical formula and Figure 6 terminology as follows:

   Applicable Range of the S_a Empirical Formula. The empirical formula for S_a (Eq. (8): $S_a = 60.03 \times [PM_{2.5}]^{0.62}$) was derived from Beijing winter observational datasets (Zhang et al., 2022) and is explicitly validated for two key conditions, which we have supplemented in the main text (Line 195) for clarity:

   - **$PM_{2.5}$ concentration range**: This formula is applicable for $[PM_{2.5}] < 200$ μg/m³, which fully covers the $PM_{2.5}$ range in our study (average: 24 ± 21 μg/m³; maximum: 131 μg/m³, see Table 2 in the manuscript).
   - **Correlation performance**: It exhibits a strong linear correlation ($R^2 = 0.82$) with directly $PM_{2.5}$, confirming reliable agreement between calculated and observed S_a for winter Beijing conditions.

   The revised Line 194 now reads:

   *"**Aerosol surface area density (S_a)**
   Due to the unavailability of direct particle size distribution measurements, S_a was derived from $PM_{2.5}$ concentrations using an empirical formula validated for winter Beijing conditions (Zhang et al., 2022):
   $S_a = 60.03 \times [PM_{2.5}]^{0.62}$ $\hspace{4cm}$ (8)
   This formula exhibits a strong linear correlation ($R^2 = 0.82$) with $PM_{2.5}$ and is applicable for $PM_{2.5}$ concentrations < 200 μg m^{-3}—consistent with the $PM_{2.5}$ range observed in this study (average: 24 ± 21 μg m^{-3}, maximum: 131 μg m^{-3})."*

   Terminology Correction in Figure 5. The term "homogeneous uptake" in Figure 5 was incorrect, as N_2O_5 uptake occurs on aerosol surfaces (a heterogeneous process). We have revised the figure caption (Line 441 in the manuscript) to correct this terminology, ensuring consistency with the study's focus on heterogeneous chemistry (Section 4.2.2).

The revised Figure 5 caption now reads:

"Figure 5. Time series variation of k_{NO_3} (reactions with NO and VOCs, heterogeneous uptake of N_2O_5 and photolysis of NO_3)."

3. Line 369-373, Please provide the scatter plot between nighttime N_2O_5 uptake (y-axis) and RH (x-axis) for individual day to support the conclusion.

Response: Thank you very much for your constructive suggestion. To better support the conclusion regarding the relationship between nighttime N_2O_5 uptake and relative humidity (RH), we have supplemented a scatter plot (now included in Figure S7 of the Supplementary Material) that depicts the correlation between the N_2O_5 heterogeneous uptake rate constant ($k(N_2O_5)$) and RH for each observation day.

As illustrated in Figure S7, a clear positive correlation between $k(N_2O_5)$ and RH is observed when RH < 70%: with increasing RH, $k(N_2O_5)$ gradually rises, reflecting enhanced N_2O_5 heterogeneous uptake driven by increased aerosol liquid water content (which provides more reactive sites for N_2O_5 hydrolysis). This trend aligns with our discussion in Section 4.1 of the manuscript (and consistent with the RH-dependent N_2O_5 lifetime analysis in Referee #1's response) and further validates the role of RH in modulating N_2O_5 uptake.
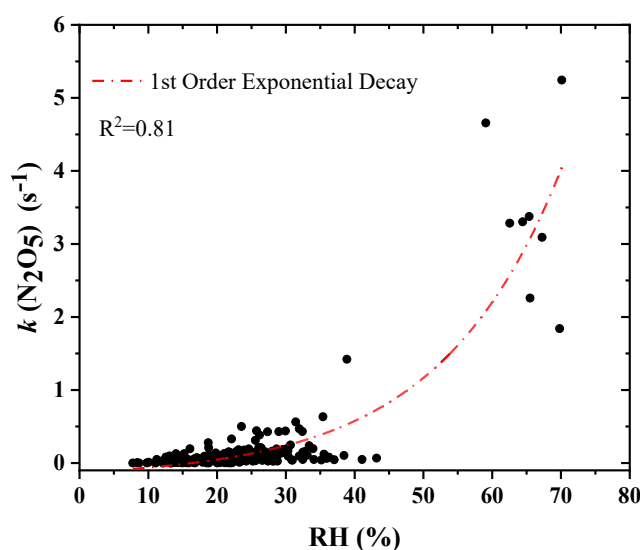

Figure S7. Schematic diagram of the correlation between RH and $k_{\text{N}_2\text{O}_5}$

4. Fig. S1, what is the red dot line? Please show the related parameters if it is the regression line.

Response: Thank you for your comment. The red dotted line in Fig. S1 represents a first-order nonlinear fit, used to characterize the inverse relationship between NO_3 production rate ($P(NO_3)$) and NO concentration (i.e., $P(NO_3)$ decreases with increasing NO). A linear relationship is not observed between the two, so nonlinear fitting better

describes their variation trend. We have supplemented the definition of this red dotted line in the figure caption of the revised Figure S2.
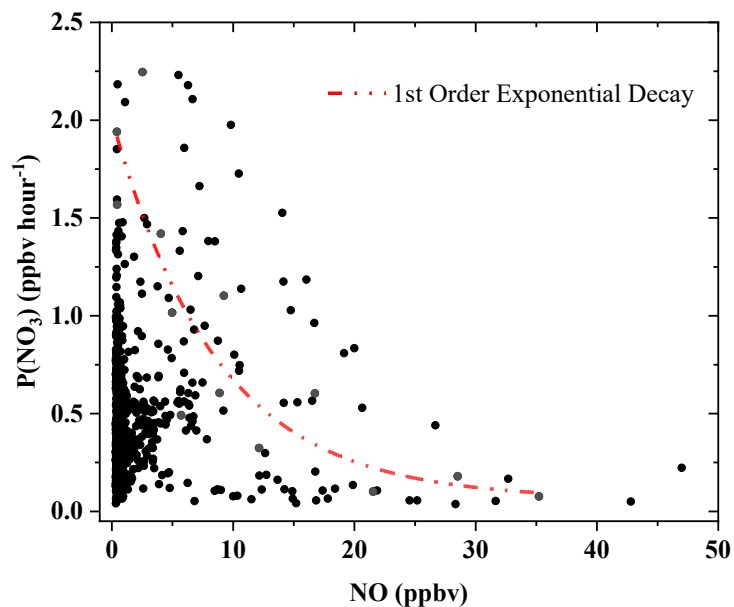

Figure S2. Schematic diagram of the correlation between NO and $P(NO_3)$.

5. Please correct the wrong citations, e.g. Hu et al., 2023, Tham et al., 2018, etc.

Response: Thank you for your feedback. We have checked all the references.