

1 **BORIS-2 – a benthic ecosystem model based on allometry**

2 Adrian P. Martin¹, Anieke Brombacher¹, Noëlie Benoist¹, Brian J. Bett¹, Jennifer M. Durden¹, Sophy
3 Oliver¹, Andrew Yool¹

4 ¹National Oceanography Centre, Southampton, SO14 3ZH, UK

5 *Correspondence to:* Adrian P. Martin (adrian.martin@noc.ac.uk)

6 **Abstract.** We present a model describing the population dynamics of benthic biota, feeding from a common resource that is
7 supplied by a flux of sinking organic carbon arriving on the seafloor. By using allometric relationships for the physiological
8 processes of growth, mortality and respiration, and for food limitation, the model represents the population dynamics of
9 organisms ranging in size from bacteria (10^{-14} g wet weight C) to large metazoans (10^3 gwwt C). The effect of temperature on
10 physiological rates is also included. The only forcing information required is the ambient temperature and the rate of supply
11 of sinking organic carbon. The model can be used for, and tuned to, specific locations. However, a parameter set is provided
12 that is generally applicable. The ability of the model to simultaneously reproduce biomass size distributions at five contrasting
13 sites is demonstrated for this parameter set. Other examples of use are also shown, using the model to explore global patterns
14 of benthic biomass, and responding to a change in food supply.

15 **1 Introduction**

16 The surface ocean, or epipelagic, ecosystem has received considerable attention from modellers for a variety of reasons,
17 spanning from the magnitude of biogeochemical fluxes (e.g. Burd, 2024) and fundamental questions of ecosystem structure
18 (e.g. Woodson et al., 2018) and biodiversity (e.g. O’Dor et al., 2009) to more societal issues such as fisheries management
19 (e.g. Karp et al., 2023) and climate modelling (Kwiatkowski et al., 2020). However, the seafloor, or benthic, ecosystem has
20 received much less attention, particularly in the deeper regions away from the continental shelves. This is despite the regions
21 deeper than 1000 m constituting over half of the earth’s surface area (Ramirez-Llodra et al., 2010; Harris et al., 2014).

22
23 The benthic ecosystem of the deep ocean (aside from hydrothermal vents) is almost entirely dependent on external input for
24 food, with the majority in the form of organic material sinking down from the waters above. This means that the benthic
25 ecosystem is susceptible to changes in production of this organic material that may occur several kilometres above it (Ruhl et
26 al., 2008), such as in response to climate change (Yool et al., 2017). Benthic ecosystems are also subject to direct pressures
27 such as trawling, dredging, oil and gas activities, and seabed mining. To understand and to predict the future for benthic
28 ecosystems we therefore need models that adequately capture their response to such drivers, across the full ecosystem and over
29 appropriate timescales.

30
31 Building models that capture the key interactions within an ecosystem is of value for three reasons: construction of a model
32 forces us to identify the key processes and to articulate our understanding of them in a precise manner; the behaviour of the
33 model allows us to identify gaps and uncertainties in that knowledge; and by linking the model to forecasts for how
34 environmental drivers may change it allows us to make predictions for the fate of the ecosystem across different scenarios.
35 One modelling strategy is to represent an ecosystem as different functional groups, particularly those linked to particular fluxes
36 of interest into and out of the sediment e.g. deposit feeders and aerobic/anaerobic bacteria (e.g. Butenschön et al., 2016; Ersten
37 et al., 2018). This approach is valuable, for example, in studying the biogeochemistry of shelf systems where the interactions
38 between sediment, overlying water and benthic ecosystem may need to be captured because the feedback on the overlying
39 water column may be significant given the shallow depths. Shelf systems also benefit from a greater array of data to constrain

40 a model as they are more accessible for sampling than deeper waters. More generally, a paucity of data to constrain a model
41 or limited understanding of causal relationships are common hindrances, particularly for deep-sea ecosystems because of the
42 remote and challenging nature of the environment being studied. For situations where the ecosystem can be approximated as
43 unchanging in time, statistical methods have been used (e.g. Reiss et al., 2014), particularly for modelling distributions of
44 groups or individual species, but also for distributions of biomass (Wei et al., 2010; Jones et al., 2014). For deep-sea ecosystems
45 where data are sparser, an inverse approach has been used to estimate fluxes between functional and size category components
46 of the ecosystem at equilibrium (Soetaert & van Oevelen, 2009 ; Durden et al. 2017; de Jonge et al., 2020), with the size classes
47 mirroring those represented by typical benthic sampling techniques. However, behaviour such as switch-feeding (e.g.
48 alternating between suspension feeding and predation) in deep-sea fauna (Durden et al. 2015; Iken et al. 2001) complicates the
49 use of discrete functional groups based on feeding types (Durden et al. 2017).

50

51 Another approach, is to represent the community purely as a collection of different size classes of organisms (Kelly-Gerreyn
52 et al., 2014; Blanchard et al., 2011; Laguionie Marchais et al., 2020) rather than as functional groups or species. As described
53 below, this offers considerable simplification in model structure and parameterisation. Furthermore, by using allometric
54 relations to base the model on the representation of rates, rather than stocks, this approach also allows the response of
55 ecosystems with time to be tracked.

56

57 Considerable attention has been given to observations showing relationships which appear to scale in a consistent way with
58 body size, both at a population (e.g. abundance - White et al 2007) and an individual (e.g. physiological rates - Gillooly et al.
59 2001) level. This phenomenon has been widely observed, on land (e.g. Nagy 1987), in the air (e.g. Niven and Scharlemann,
60 2005) and in the sea (e.g. Molony and Field 1989) including the deep ocean (Durden et al. 2019; McClain et al. 2012; Mahaut
61 et al. 1995). That such behaviour has been observed across many habitats and orders of magnitude in size of organism
62 unsurprisingly led to a search for a “Universal” law explaining such behaviour. Metabolic rate controls ecological processes
63 at individual and ecosystem levels by determining resource uptake and allocation. The Metabolic Theory of Ecology (MTE;
64 West et al., 1997; Brown et al., 2004) asserts that, to first order, this rate is controlled by the size of organism and the ambient
65 temperature. This provides a potential explanation for the existence of a power-law relationship between physiological rates
66 and body size. However, there remains a discussion over the taxonomic or functional scale at which other features or processes
67 might disrupt any universal scaling (Seibel and Drazen, 2007), the precise value of the scaling (Isaac and Carbone 2010, Brey
68 2010; Glazier, 2022) and the extent to which such an approach applies to systems that are not in equilibrium (McCarthy et al.,
69 2019).

70

71 Notwithstanding these caveats, an allometric approach still has considerable value when applied at broad ecosystem scales.
72 To support use of an allometric approach, we give just a few examples for three key processes: growth, respiration and
73 mortality. Motivated by predictions of MTE, Ernest et al. (2003) successfully tested the predicted scaling exponent of -0.25
74 for growth rate, for organisms spanning 10^{-14} - 10^8 g in size. More specific to this study, with a focus on macrobenthos, Cusson
75 & Bourget (2005) brought together empirical relations from previous studies (their Table 8) that demonstrate similar evidence
76 of scaling of growth rate with size. For respiration, Mahaut et al. (1995) found a power-law scaling with size for deep-sea
77 organisms, spanning seven orders of magnitude. This was for a single location, however, so not suitable for testing a
78 temperature dependence. For mortality, McCoy and Gillooly (2008, 2009) brought together estimates of natural mortality
79 spanning 22 orders of magnitude including plants, fish, birds, mammals and invertebrates, finding that a power-law scaling
80 with size plus an exponential dependence on temperature captures the dominant pattern. A restriction of their data purely to
81 invertebrates found a similar result (McCoy and Gillooly, 2009), still spanning 11 orders of magnitude in size. Again, focussed
82 on marine benthic organisms, McClain et al., (2012) analysed data for growth, respiration and turnover (which can be a proxy

83 for mortality) and demonstrated a clear power-law scaling with size, with additional support for an exponential relationship
84 with temperature. There are, therefore, reasonable grounds for adopting an allometric approach.

85

86 Kelly-Gerreyn et al. (2014) constructed a dynamic model for benthic organisms based on allometry such that physiological
87 rates vary with body size (Benthic Organisms Resolved in Size – hereafter “BORIS-1”). BORIS-1 was capable of reproducing
88 the size distribution of organisms at three sites contrasting in depth between 150m and 1600m. This model assumed that all
89 organisms were detritivores, eating from a common pool of detritus supplied by organic material sinking to the seafloor, the
90 particulate organic carbon (POC) flux. BORIS-1 demonstrated that an allometric model with a small number of physiological
91 processes (ingestion, assimilation and respiration/mortality) that are common to all organisms but scale with body size can
92 capture the size-distribution of biomass seen in observations. However, the model has several limitations. The first is that it
93 only represents a limited range of sizes (8.9×10^{-7} to 2.9×10^{-2} g wet weight). It was therefore necessary to assume a specific
94 fraction of POC flux that was consumed and respired by organisms not represented by the model, and hence not available to
95 the modelled organisms. The omitted organisms included both smallest (e.g. bacteria) and largest (e.g. large sea cucumbers)
96 size ranges. The physiological rates were also not dependent on temperature even though there is evidence that physiological
97 rates typically increase as the environment warms (e.g. Gillooly et al., 2001). Additionally, the mortality rate had a dependency
98 on the POC flux in BORIS-1. This resulted in estimates of longevity that unrealistically varied across several orders of
99 magnitude for the same organism at different locations.

100

101 This paper presents an expanded and updated version of BORIS-1 that addresses these limitations. The resulting model,
102 BORIS-2, spans the full range of organism sizes, includes the physiological role of temperature, and is parameterised using a
103 larger dataset that includes observations from a greater range of sites with contrasting environments conditions, including the
104 abyssal ocean. A single parameter set that allows the model to capture the ecosystem structure across these sites is given and
105 examples are demonstrated for how the model may be used to study both local and global questions. It is worth stressing that
106 the aim of BORIS-2 is to capture broad ecosystem behaviour, i.e. macroecology, across the full range of body sizes, not to
107 capture the dynamics of specific species.

108 **2 BORIS-2 model description**

109 BORIS-2 represents benthic organisms spanning in size from bacteria to large metazoans (Figure 1). It does so by dividing
110 benthic organisms into size classes and using an allometric approach. The size classes are defined on the basis of individual
111 wet weight body mass (units grams wet weight – gwwt). From the smallest to the largest, each size class spans twice the range
112 of the former. More specifically, the mean body mass of organisms within a size class spans from 1.3×10^{-14} gwwt (13 fg) for
113 the smallest to 3.6×10^3 gwwt (3.6 kg) for the largest. The lower limit of the smallest class is 0.88×10^{-14} gwwt (9 fg) and the
114 upper limit of the largest is 5.1×10^3 gwwt (5.1 kg). The size classes are chosen to be consistent with those used for size-spectra
115 biomass data (e.g. Laguionie Marchais et al., 2020) and BORIS-1 (Kelly-Gerreyn et al., 2014). The smallest size class is based
116 on the smallest observed bacteria (Luef et al., 2015), using a conversion from gC to gwwt of 11.5 (Brey, 2010). The largest
117 size class is chosen to be broadly representative of benthic habitats. It is consistent with a detailed assessment of invertebrates
118 for a well-studied abyssal site, the Porcupine Abyssal Plain (one of the sites described in Section 2.4). These upper and lower
119 biomass limits, together with the factor of two scaling used between biomass size classes, sets the number of model size classes
120 to 59. For application of BORIS-2 to specific locations where larger organisms are known, the upper size limit is easily
121 changed. The BORIS-2 size range currently spans over 18 orders of magnitude.

122 2.1 Ecological interactions

123 BORIS-2 comprises a set of differential equations describing the time-varying behaviour of $N=59$ size classes ingesting a
 124 common resource, R , that represents the stock of detrital food available to the benthic community (i.e., in / on the seafloor or
 125 in the benthic boundary layer). The model does not capture any direct predation or cannibalism, and instead represents a
 126 community of detritivorous heterotrophs. This decision is based on the different nature of benthic and pelagic ecosystems.
 127 Total biomass in a size range increases with body mass in benthic ecosystems (e.g. Benoist, 2020; Kelly-Gerreyn et al., 2014).
 128 In the pelagic ocean, biomass is roughly equal across different sizes (Hatton et al., 2021). The greater accumulation of biomass
 129 in larger organisms in the benthic system is consistent with a greater transfer efficiency arising from a reduced role of predation
 130 and a greater role of feeding from a common detrital resource. We effectively assume that the dominant influence on benthic
 131 ecosystems is having neighbours competing for your food, rather than eating you. This assumption is discussed in more detail
 132 in Section 4.1.

133
 134 The total biomass represented by all organisms per square metre in each size class i of nominate mass M_i (units: grams wet
 135 weight - gwwt) is represented as B_i (units: gwwt/m²) which varies with time, t (units: d), according to the equation

$$136 \frac{dB_i}{dt} = \overbrace{g_i \cdot f(R, B_i) \cdot B_i}^{\text{gross growth}} - \overbrace{r_i \cdot B_i}^{\text{respiration}} - \overbrace{m_i \cdot B_i}^{\text{mortality}} \quad (1)$$

137 where g_i is the maximum specific growth rate and r_i and m_i are the specific rates of respiration and mortality, respectively. All
 138 of g_i , r_i and m_i have units of 1/d. The function $f(R, B_i)$ represents how growth is limited by increasing population size and/or
 139 decreasing resource availability (see Section 2.2). The associated equation controlling the amount of resource, R (units:
 140 gwwt/m²), is

$$141 \frac{dR}{dt} = F - \sum_{i=1}^N \left[\overbrace{g_i \cdot f(R, B_i) \cdot B_i}^{\text{gross growth}} - \overbrace{m_i \cdot B_i}^{\text{mortality}} \right] \quad (2)$$

142 where F is the POC flux to the seafloor through gravitational sinking of detritus (gwwt/m²/d). Note that at equilibrium the rate
 143 of supply of organic material, F , equals the respiration by the whole ecosystem, $\sum_{i=1}^N r_i \cdot B_i$. It is assumed that the long-term
 144 burial of organic material in sediment is negligible compared to POC and total respiration fluxes. (Section 4.1 discusses how
 145 this assumption might be relaxed.) A linear mortality term, $m_i \cdot B_i$, is used in BORIS-2. (The reason for using this rather than
 146 the quadratic $\mu_i \cdot B_i^2$ parameterisation used in BORIS-1 is given in the Appendix).

147 2.2 Growth limitation: food scarcity and interference

148 The growth limitation function, $f(R, B_i)$, reflects the impact on growth arising from competition for limited resources. This
 149 function is chosen to capture two effects. First, low availability of food, R , should lead to a reduced rate of intake and growth.
 150 Second, any increase in the number of organisms (for which B_i is a proxy) looking for food should reduce the likelihood of
 151 any individual finding it, a phenomenon known as interference (e.g. DeAngelis et al., 1975). More specifically, we assume the
 152 parameterisation

$$153 f(R, B_i) = 1/(1+a_i B_i/R) \quad (3)$$

154 The parameter a_i is present to account for how interference may scale with size. For example, larger organisms can search a
 155 given area more quickly than a smaller one, in general, either because motility generally increases with size or more simply
 156 because they occupy a greater area. Note that a_i is unitless. The function $f(R, B_i)$ varies between 0 and 1, with a value of zero
 157 entirely ceasing growth and a value of one leading to growth at the maximum rate, g_i . To demonstrate that the function has the
 158 required properties, first consider the case where food is very abundant such that $R \gg a_i B_i$. Then $f(R, B_i) \sim 1$ and there is no
 159 limitation of growth. If resource is scarce such that $R < a_i B_i$, then $f(R, B_i) \sim R/(a_i B_i)$ which is always a value less than one but
 160 increases and decreases linearly with both resource, R , and abundance of organisms, represented by B_i . Note that for simplicity

161 we currently only incorporate competition for resources within a size class. This is the simplest assumption given that different
 162 size classes seek food at different spatial scales. What is a meal for a bacterium is unlikely to be a meal for a holothurian. It
 163 would, however, be straightforward to include competition from other size classes simply by using a sum over those classes
 164 in the denominator. The form of the interference parameter, a_i , is discussed in the next section (2.3). Figure 2 and Section 2.4
 165 show further how the function $f(R, B_i)$ varies, in particular demonstrating that it affects all size classes equally i.e. that it does
 166 not lead to some size classes being food limited while others are growing near maximum rates.

167 2.3 Allometric and temperature influences

168 In BORIS-2 allometry is used to describe four physiological or physiologically affected processes across the range of body
 169 sizes. The physiological processes are growth (g_i), respiration (m_i) and mortality (r_i). The physiologically affected process is
 170 growth limitation, controlled by parameter a_i . All of these are assumed to be determined by size (body mass) and environmental
 171 temperature.

172

173 The effect of temperature is assumed to be identical for all four processes and represented by a function, $\theta(T)$, which is taken
 174 to be

$$175 \theta(T) = \exp[-E \cdot \tau/k] \quad (4)$$

176 with

$$177 \tau = 1/(T+T_{abs}) - 1/(T_{ref}+T_{abs}) \quad (5)$$

178 where T is temperature (units: °C), T_{ref} (units: °C) is a reference temperature, $T_{abs} = 273.15$ K converts T and T_{ref} to units of
 179 Kelvin (K), and k is Boltzmann's constant (8.62×10^{-5} eV/K). This is a widely used formulation applied both in empirical studies
 180 (e.g. Brey, 2010) and papers developing ideas around the Metabolic Theory of Ecology (e.g. Gillooly et al. 2001). E (units:
 181 eV) is often described as an activation energy. We discuss the value chosen for E in Section 2.4.2. T_{ref} is chosen to be 20°C.
 182 While this may seem an arbitrary choice of reference temperature, it has no impact on rates. Using a different T_{ref} simply
 183 requires a numerical change in parameters (g_0 , r_0 , m_0 and a_0) to compensate for the change.

184 It is assumed that the three physiological rates (g_i , m_i and r_i) scale with body size in an identical way. This is purely taking the
 185 option requiring fewest assumptions given the current uncertainty in how these rates vary with size. As the link between
 186 interference and physiology is more tentative, a_i is theoretically allowed to scale independently (but see Section 2.5). More
 187 specifically, growth, respiration and mortality have common scaling exponent β , whereas interference scales with exponent α :

$$188 g_i = \theta(T) \cdot g_0 \cdot M^\beta \quad (6)$$

$$189 r_i = \theta(T) \cdot r_0 \cdot M^\beta \quad (7)$$

$$190 m_i = \theta(T) \cdot m_0 \cdot M^\beta \quad (8)$$

$$191 a_i = \theta(T) \cdot a_0 \cdot M^\alpha \quad (9)$$

192 The values chosen for the seven parameters used in the model (g_0 , r_0 , m_0 , a_0 , α , β , E) are given in Section 2.5 (and Table 1),
 193 together with a description of the data used to constrain them. The performance of the model using this parameter set is then
 194 described in Section 2.6 and the uncertainties associated with their values are discussed in Section 2.7. Before then a steady
 195 state solution for the model is presented, both for its own use and as a source of useful information for constraining parameter
 196 values.

197

198 **2.4 Steady state solution**

199 The model has a steady state solution which can be written in a simple form. This provides a means to initialise simulations,
 200 to validate dynamical model runs (if run to equilibrium) or to accelerate model runs where time-scales are longer than organism
 201 response times.

202 Indicating steady state values with an asterisk, the steady state solution is

$$203 \quad R^* = \frac{F}{\sum_{i=1}^N \frac{r_i(g_i-r_i-m_i)}{a_i(r_i+m_i)}} \quad (10)$$

$$204 \quad B_i^* = \frac{(g_i-r_i-m_i)}{a_i(r_i+m_i)} \cdot R^* \quad (11)$$

205 This steady state solution provides a few insights into the behaviour of the model. First, both resource, R , and biomass in all
 206 size classes, B_i , increase linearly with F . This is not surprising as we would expect abundance of detritus and biomass to
 207 increase with increasing food supply. Second, B_i scales with size with the same exponent, $-\alpha$, as $1/a_i$ (i.e. $B_i^* \propto M^{-\alpha}$). This is
 208 because g_i , r_i and m_i all scale the same with size, as mentioned above, and so the scaling of $(g_i-r_i-m_i)$ in the numerator for B_i^*
 209 is cancelled by the identical scaling of (r_i+m_i) in the denominator. Hence, the biomass spectral slope is effectively set by
 210 interference. Although this might be unexpected it should be noted that the processes contributing to a_i are still very poorly
 211 known and its scaling is likely to be influenced by physiological processes, such as respiration associated with enhanced
 212 movement for example. The theoretical model of Damuth (2007) is potentially relevant here as it links competition for
 213 resources to allometric scaling and community wide energy use. Nevertheless, understanding the likely influences on
 214 interference is clearly a useful avenue for future research. A consequence of the inverse scaling of a_i and B_i^* is that $a_i B_i^*$ is the
 215 same for all size classes i.e. the growth limitation function f does not change with size at steady state.

216

217 Returning to the steady state solution, substituting Equations 6-9 into Equations 10 and 11 gives

$$218 \quad B_i^* = \left(\frac{1}{r_0 \theta(T)} \right) \cdot \left(\frac{M_i^{-\alpha}}{\sum_{j=1}^N M_j^{\beta-\alpha}} \right) \cdot F \quad (12)$$

$$219 \quad R^* = \left(\frac{\alpha_0 \cdot (r_0 + m_0)}{r_0 \cdot (g_0 - r_0 - m_0)} \right) \cdot \left(\frac{1}{\sum_{j=1}^N M_j^{\beta-\alpha}} \right) \cdot F \quad (13)$$

220 In addition to showing explicitly that B_i scales as $-\alpha$, as already mentioned, Equation 12 also reveals that the steady state
 221 biomass is independent of growth and mortality except for the scaling (β) and temperature dependence ($\theta(T)$) that they share
 222 with respiration. While this might seem at first surprising, it is because of the fundamental constraint that total respiration must
 223 match the POC flux, F , of arriving new organic material, i.e. $F = \sum_{i=1}^N r_i \cdot B_i^*$. At equilibrium, any change in growth or
 224 mortality arising from changing either g_0 or m_0 , respectively, is compensated by a change in food resource (R), rather than in
 225 B_i , to maintain this balance. This balance is also reflected in the influence of r_0 in Equation 12, with an increase in it
 226 corresponding to a compensating decrease in B_i^* . Similarly, if the specific respiration rate increases as a result of temperature
 227 increase (see Equation 7) then the higher physiological overhead means that a lower B_i^* is maintained.

228

229 The steady state solution is also of use in understanding the influence of interference in the growth limitation function. Equation
 230 11 implies that at steady state $a_i B_i^* = \delta \cdot R^*$ where

$$231 \quad \delta = \frac{(g_i-r_i-m_i)}{(r_i+m_i)}$$

232 Because the physiological rates scale identically, and have the same temperature dependence,

$$233 \quad \delta = \frac{(g_0 - r_0 - m_0)}{(r_0 + m_0)}$$

234 which is the same for all size classes and sites. (For the parameter set used here and described below, $\delta=1$.) Figure 2 shows
235 how the interference function varies with both R and B_i , as they vary either side of their steady state values.

236

237 **2.5 Observational constraints and choice of parameter values**

238 There are a range of observations that can be used to constrain parameter values but, as described below, none can be used to
239 set a parameter value in isolation. Relationships between parameters must be used to link the different observations together
240 as a collective constraint. It is also worth noting that there is no objective way to use these multiple constraints. As will be
241 seen below, the strength of some constraints is greater than others and trying to construct some overall cost function to optimise
242 all parameters simultaneously would require considerable subjectivity in how the constraints were translated into costs and
243 weighted relative to each other. For this reason, and because of the limited number of parameters and ability to calculate the
244 outcome of a given parameter set extremely quickly, values have instead been chosen by trial and error for the seven parameters
245 $g_0, r_0, m_0, a_0, \alpha, \beta, E$ to give an acceptable, if potentially not optimal, fit to observations. A summary of values can be found
246 in Table 1. While future users of BORIS-2 may choose to use a different approach to selecting parameter values, it will be
247 seen in Section 2.6 that the current set does a reasonable job and Section 2.7 describes the consequences associated with
248 adjusting these values.

249

250 Previous studies have highlighted observational evidence for allometric relationships, either in just one physiological rate (e.g.
251 Mahaut et al., 1995; Cusson & Bourget, 2005), or, in several (e.g. McClain et al., 2012). Because of the variations in scaling
252 reported in these studies (discussed in Section 2.7) we have used the observation of allometry as a starting point rather than
253 take a value for the scaling exponent from a specific study.

254

255 The following first describes the observational constraints. Based on these, the argument for the specific choices of parameter
256 values is then given. Table 1 has a summary of parameter values and Figures 3 and 4, and Table 2, summarise the observational
257 constraints and associated model diagnostics used to select them. The code needed to generate Figures 3 and 4 is also available
258 to allow the model to be re-tuned given additional data, different locations or different priorities (Martin et al., 2025).

259

260 Starting with the largest dataset, biomass can be used as a constraint for the interference parameter α . The steady state solution
261 (section 2.4) shows that a_i must scale in the opposite way to biomass distributions. Suitable observations from five sites are
262 available in selecting the value; a summary is found in Table 2. The Clarion Clipperton Zone (CCZ) is a vast abyssal plain in
263 the northeast Pacific. The data used here come from a site (17.2° N 122.6° W) of depth 4150 m with a low temperature (1.5
264 °C). Fladden Ground (FG) is in a shelf sea (153 m) and, unsurprisingly, with higher temperature (8 °C). The Faroe-Shetland
265 Channel (FSC) is a connection between the North Atlantic and the Arctic, with the lowest temperature (-1 °C) despite a depth
266 of only 1623 m. The Oman Margin (OM) is a slope site (507 m) and has the highest water temperature (13 °C). The final site
267 is the Porcupine Abyssal Plain (PAP), which is the deepest (4850 m), with reasonably cold temperature (2.6 °C). A general
268 decrease of temperature with depth is overlain with considerable variability due to local hydrography (notably FSC and OM).
269 In addition to spanning a range of contrasting temperatures and depths, the data from the five sites also covers complementary
270 size ranges of organisms. CCZ data are based on photographically surveyed megabenthos. FG, FSC, and OM data are based
271 on physically sampled meio- to macrobenthos. PAP is based on physically sampled macrobenthos and photographically
272 surveyed megabenthos. For CCZ, the data can be found in Benoist (2020), with sampling and methodology described in Simon-
273 Lledó et al. (2019) and Benoist et al., (2019). Details on data for FG, FSC and OM can be found in Kelly-Gerrey et al. (2014).
274 Additional information on the sampling and laboratory methodology can be found in Kaariainen et al. (2006). The benthic
275 ecosystem of the PAP site has been studied for decades (Hartman et al., 2021). The data used here, and presented in Benoist
276 (2020), combines analyses of macrobenthos and megabenthos. Descriptions of observational approach and the analysis

277 methodology for the megabenthos can be found in Morris et al. (2016) and Durden et al. (2020b). For macrobenthos, this
278 information can be found in Benoist (2020) and Ruhl et al. (2023). Further details on the treatment of size-resolved data, e.g.
279 to remove biases such as under-sampled size groups, can be found in Edwards et al., (2017, 2020) and Ruhl et al. (2023).
280 Observations of biomass versus size for each of the sites are shown in Figure 3. These are referred to as biomass spectra and
281 the gradient of the relationship (when plotted log-log as here) as the spectral slope, or scaling exponent. Despite some
282 variability, all sites exhibit an increase of biomass with body size, and in a manner that is consistent with a power-law
283 relationship. Figure 3 also shows the exponents found by fitting a power-law to the observations from each site individually.
284 There is no strong relationship between fitted exponent and environmental parameters, though the smaller magnitude scaling
285 exponents for the two shallow sites is something that has previously been seen in physiological rates rather than biomass
286 (Mahaut et al., 1995). The OM site additionally has a low oxygen concentration (Demopoulos et al., 2003) which has been
287 suggested to have a disproportionate impact on larger organisms (Quiroga et al., 2005) and which could therefore be
288 responsible for flattening the slope relative to other sites. Nevertheless, here we take the simplest assumption that all sites are
289 showing sufficiently similar behaviour in scaling to assume a common scaling exponent across the sites, leaving an
290 investigation of departures from this for other studies. Figure 3g shows the simultaneous fit to data from all sites. It is seen that
291 observations from the five sites cover different size ranges such that the composite dataset spans a substantially wider size
292 range than any individual site. Furthermore, the relationship of biomass with size appears consistent across the wider range. If
293 the data from all sites is combined then the scaling exponent for a power-law fit to all five sites simultaneously indicates a
294 scaling exponent of 0.26 (s.d. 0.016, $r^2=0.76$, $p<0.001$). The existence of similar scaling behaviour across multiple sites gives
295 more confidence that the model can be used globally.

296
297 The next observational constraint comes from the requirement that the supply of organic carbon to the ecosystem (the POC
298 flux, F) balances the total respiration of the organisms present at steady state. For the model, respiration is given by the sum
299 over size classes of the product of the specific respiration rate and biomass in each size class, $\sum r_i B_i$. Therefore, observations of
300 POC flux, co-located with the previously described size-resolved estimates of biomass, provide a useful joint constraint on r_0
301 and β . There are several sources of data for POC flux. First, and most directly, for the PAP site there is a long time-series of
302 sediment trap data (Lampitt & Pebody, 2023). Although there are sediment traps at both 3000 m and 4750 m, the latter is
303 thought to be biased by sediment resuspension as it is just 100m above the seafloor. The magnitude of this effect can vary with
304 time but previous work has shown that the flux near seabed at the PAP site is often in excess of that at 3000 m due to re-
305 suspension (Lampitt et al., 2000, 2001). For this reason, it is better to extrapolate the estimate from the 3000 m trap. For the
306 year 2012 (to best match the biomass observations) the annual carbon flux at 3000 m is 1.91 gC/m²/yr. The associated flux at
307 the seafloor can be roughly estimated using a widely used power-law scaling (Martin et al., 1987), such that the flux at the
308 seafloor at 4850 m equals $1.91 \text{ gC/m}^2/\text{yr} * (3000 \text{ m}/4850 \text{ m})^{0.858} = 1.3 \text{ gC/m}^2/\text{yr}$. An alternative way to estimate the sinking
309 flux at the seafloor is to use the Lutz et al. (2007) algorithm, which uses net primary production, sea surface temperature and
310 depth at a given location to estimate the flux. This allows estimates to be made for all five sites (not just PAP) – see Table 2.
311 Finally, Sediment Community Oxygen Consumption (SCOC) data (Stratmann et al., 2019) also allows the sinking flux to be
312 estimated, making the same assumption that the respiration (measured by oxygen consumption) must balance this flux, in this
313 case when averaged over the year. As SCOC is usually measured using chambers of ~50 cm across, the estimates exclude or
314 bias the contribution from larger organisms – not just those too large to fit but also those too scarce to be robustly sampled in
315 such an area - and care is needed in accounting for this (Laguionie Marchais et al., 2020). More generally, all of these sources
316 of POC flux data have significant associated uncertainties, which is why POC flux was not fixed when deriving parameter
317 values for the general use model configuration described here. Instead the model can be used to estimate the POC flux and
318 compared to these different observational estimates as a broader constraint. All observational estimates of POC flux used to
319 constrain model parameters, as well as the model values, are shown in Figure 4b. These values are also given in Table 2. Note

320 that even if there was no uncertainty in the observations for POC flux it would still not be possible to infer specific values for
321 β and r_0 using the data. This is because it is possible to simultaneously vary β and r_0 in a way that the total respiration remains
322 unchanged. Note that respiration also gives a link between the physiological (β) and interference (α) scaling exponents that
323 needs to be considered. Respiration by a size range of organisms is given by $r_i B_i$, which scales as $M_i^{\alpha-\beta}$. Hence, the difference
324 between α and β determines whether the energy use by a size range increases ($\alpha > \beta$), decreases ($\alpha < \beta$) or remains with same
325 ($\alpha = \beta$) with size. We are unaware of observations indicating an increase with size but there are observations suggesting $\alpha = \beta$
326 (Laguionie Marchais et al., 2020) and this is consistent with ‘energy equivalence’ which has been suggested theoretically
327 (Damuth, 2007).

328

329 The inherent relationships between the three physiological rates allow other observational constraints. Observations indicating
330 that physiological rates decrease with size were already described. Additionally, the maximum net growth rate, g_i , must equal
331 or exceed the sum of the respiration, r_i , and mortality, m_i , for all sizes for populations to be sustainable. Together, these two
332 statements imply that all rates must sit within a ‘window’ of parameter space (rate versus size, Figure 4c). We follow a similar
333 approach to that of Mahaut et al. (1995) and Kelly-Gerreyn et al. (2014), by using constraints at smallest and largest sizes of
334 organism. Having so many degrees of magnitude in size in the model means that care is needed for the parameter values to be
335 realistic at the two extremes of the sizes reproduced. Using constraints at intermediate sizes risks significant under- or over-
336 estimates for largest and smallest organisms through under-constrained extrapolation. Because rates are observed to decrease
337 with size, the top left corner of the window is set by the upper limit for the specific growth rate of the smallest size class of
338 organism. For benthic bacteria, we take a rough upper limit of 1 /d maximum specific growth rate. Though Dixon and Turley
339 (2001) find a rate of 0.1 /d, a wide range reported by Giovannelli et al. (2013) includes a value of 6 /d. An intermediate value
340 of 1 /d is marked in Figure 4c. The bottom right corner of the window is set by the lower limit for the specific mortality or
341 respiration rate (whichever is smallest) of the largest size class. The data collated by McClain et al. (2012) indicates lifetimes
342 for the largest organisms (few kg wwt) of order 50 years. Given this and assuming the same scaling, β , for all physiological
343 rates as stated earlier, comparison to POC fluxes described above indicates that respiration is a considerably larger rate than
344 mortality, so mortality defines the bottom right corner of the ‘window’. This is also marked in Figure 4c and lifetimes for the
345 largest organisms at each site are also given in Figure 4e. Note that both of these constraints need to be treated a little flexibly
346 as it is not realistic to set a precise limit in either case.

347

348 There are two further constraints, to inform the choice of values for β , g_0 , r_0 , m_0 and E . First, the ratio $(r_0+m_0)/g_0$ represents the
349 fraction of the maximum growth rate, g_0 , achieved by organisms when the system is in equilibrium. This should take a value
350 less than 1 for food to be limiting, as is expected for the seafloor (Smith et al., 2008). Second, decreasing E converges
351 physiological rates for the different sites, as the inter-site differences due to temperature are diminished. Doing so reduces the
352 inter-site differences in model POC flux (and R) for the same reason, because of the balance with total respiration.

353

354 The final observational constraint is Total Organic Carbon (TOC) in the benthic sediment which can be used to estimate the
355 amount of detritus available as food, R , which itself is a constraint on the parameter a_0 . This can be seen in the steady state
356 solution (Equation 13), where R increases linearly with a_0 . We make use of the compilation of Parameswaran et al., (2024)
357 who created an atlas of TOC at the seafloor surface, using a neural network approach applied to globally distributed estimates
358 calculated over the top 10 cm of sediment. This data source was chosen in preference to the alternative product of Atwood et
359 al., (2020) as the latter used estimates over the top 1m of sediment, which are an order of magnitude larger than those in
360 Parameswaran et al., (2024) and likely to represent carbon resources unavailable to the majority of detritivores on the seafloor.
361 The estimates for TOC using Parameswaran et al., (2024) for the five sites are shown in Figure 4e. These are not directly
362 comparable with R , however. There is considerable evidence that not all of sediment TOC is readily available as a food

363 resource, with typically 5% (e.g. De Jonge et al., 2020; Van Oevelen et al., 2011a, 2011b) regarded as ‘labile’ i.e. easily
364 consumed (see Discussion for more on this). We therefore multiply our model estimates for R at the five sites by a factor of
365 20 to give an estimate of TOC for comparison to the observational values. Estimates for TOC based on Parameswaran et al.,
366 (2024) are given in Figure 4e and Table 2, together with model estimates for R and TOC. There is considerable variability; in
367 observations, model values and their relative sizes. A weaker constraint for a_0 is the turnover time for R i.e. the time it would
368 take for R to be replaced by the POC flux (Figure 4d). A minimum turnover time of several years would be expected for the
369 system to be able to achieve steady state on an annual basis.

370

371 We now outline our choices of parameter values (Table 1) given the above constraints. Starting with α and β , taking $\alpha=\beta=-$
372 0.2 gives a scaling exponent consistent with observations for rates and energy equivalence but requires a biomass scaling that
373 is a little lower than the fit to observations. It is not possible to use the -0.26 scaling implied by the observed biomass spectra
374 without substantially worsening the fit between model and observational estimates of POC flux (Figure 4b). Given this scaling
375 for β , we take $g_0=0.001$ /d, $r_0=0.0045$ /d, $m_0=0.0005$ /d to fit rates within the window in Figure 4c. This also gives a value for
376 $(r_0+m_0)/g_0$ of 0.5, giving food limitation at steady state. We take a value of $E=0.35$ eV which is at the low end of observations
377 but using a larger value also worsens the match between model and observational estimates of POC flux. Finally, a value of
378 2000 is taken for a_0 as it gives the best agreement between model-estimated and observed TOC (Table 2) across the sites given
379 the other choices of parameter values.

380

381 The sensitivity of key metrics to the model parameters is shown in Figure 5. For each metric, its fractional change for a +/-
382 10% change in parameter value, from the values given above, is shown. This was carried out for each site assuming the
383 observed temperature and the POC flux that is consistent with a model fit to the biomass observations at that site with the
384 imposed 0.2 scaling exponent. The three parameters g_0 , α and β show the greatest sensitivity, with a 20-30% change for a 10%
385 change in parameter. The exceptions are: the largest organism lifetime, which is insensitive to g_0 and α ; and total biomass,
386 which is insensitive to g_0 . It should be noted that the sensitivities of α and β work in opposite directions, with a decreased
387 metric for one corresponding to an increase in the same metric for the other. This means that any simultaneous change to these
388 parameters while retaining energy equivalence (i.e. $\alpha=\beta$) would have a reinforcing rather than a compensating effect. Although
389 only the parameter g_0 shows a clear asymmetry about zero for the sensitivity analysis in Figure 5 it should be noted that the
390 metrics are generally non-linearly dependent on the parameters. This becomes more apparent if a larger fractional change to
391 parameters in the sensitivity analysis is used. There are some cases where a metric is linearly dependent on a parameter though
392 (e.g. R on a_0 – see Equation 13) and in these cases the effect of a +/- parameter change will always be symmetrical about zero.
393 An obvious feature of the sensitivity analyses is that only E has an influence which varies across sites. This is the combined
394 effect of two factors: temperature and POC flux are the only external influences that vary between sites; and the sensitivity
395 analysis shows fractional, not absolute, changes in the metrics i.e. the sensitivity is calculated as the metric evaluated at the
396 altered parameter value divided by that obtained using the standard value. Using the steady state solution (Equations 10 and
397 11) and taking each metric in turn we can see how this cross-site consistency arises. For R , as it varies linearly with the POC
398 flux, the flux cancels out in the sensitivity calculation. The temperature effect (Equation 4) varies across sites, but R is already
399 insensitive to this because growth, respiration, mortality and interference all share the same temperature dependence and their
400 effects cancel each other out (Equation 10). As TOC is calculated simply as a constant multiple of R , it shares the lack of
401 variation in sensitivity across sites. Similarly, the turnover time (τ) is calculated as the ratio of R to the POC flux, the latter of
402 which cancels when calculating the metric (as described above). Thus, τ also shows a sensitivity that is constant across sites.
403 The two metrics of total biomass (ΣB) and lifetime of largest organism (λ) do vary across sites for one of the parameters, the
404 activation energy E . The steady state solution for biomass (Equation 11, but perhaps clearer in Equation 12) shows that it
405 varies with temperature due to temperature’s influence (via $\theta(T)$) on physiological rates (specifically the respiration rate) for

406 which E is the key parameter. The sensitivity of total biomass to E is dictated by the ratio of $\theta(T)$ calculated at altered and
407 original values. Because $\theta(T)$ is a non-linear function of T and E , this ratio varies across sites. Similarly, as the lifetime of the
408 largest organisms is simply the reciprocal of their mortality, a physiological rate which varies with $\theta(T)$ across sites, then it
409 will too. For the other parameters, neither temperature nor POC flux give rise to variation in sensitivity across sites: none of
410 g_0 , r_0 , m_0 , a_0 , α , or β exert an influence through changes in temperature; biomass varies linearly with POC flux (Equation 11),
411 which therefore cancels out when calculating the sensitivity; and the lifetime of the largest organisms has no dependency on
412 the POC flux.

413

414 **2.6 Model performance**

415 The observations at all five sites show a power-law distribution of biomass with size, and the model performs well in capturing
416 this characteristic at each site (Figure 3). Note that the matching of magnitudes in biomass for each size class across sites in
417 Figure 3g arises from the normalisation for the simultaneous fit, not directly from model parameter value choices. Although
418 the choice of E influences the differences in biomass across sites, they are also influenced by inter-site differences in POC
419 flux. For POC flux (Figure 4b and Table 2), although there is some variability within the observations, the model and
420 observational estimates agree reasonably well and vary in a similar way across the sites. All have lowest fluxes at the deepest
421 sites (CCZ, PAP), highest at the shallowest (FG), and the fluxes at these two extremes differ by 1-2 orders of magnitude. For
422 the physiological rates, in Figure 4c the lines showing the rates should fall within ‘window’ marked as dotted black lines, and
423 described in Section 2.5. It is seen that although this is broadly the case, the chosen parameter values already lead to this
424 ‘window’ being stretched. Maximum bacterial (smallest size class) net growth rate is between 2-4 /d across the sites, higher
425 than the reference value of 1 /d. The net growth rate achieved at equilibrium to first order matches the respiration rate because
426 mortality is so much smaller, and varies between 1-2 /d. Life expectancies for the largest organisms are seen to span 40-82 yr
427 and straddle the reference value of 50 yr. Estimates for TOC using the model vary from 24 gC/m² to 6400 gC/m². The observed
428 range is smaller (from 320 to 2000 gC/m²) but given the large variability in observation and model estimates across sites, the
429 match is reasonable; at some places the model exceeds the observed value and at others it is lower. The turnover time (Figure
430 4d) is seen to be estimated at ~20 yr, using the model estimated POC flux. It is lower (3-15 yr) using the observations for POC
431 flux (Figure 4b) and model estimate of R because the model under-estimates the POC flux. Overall, using this set of parameters,
432 the model does a reasonable job of satisfying the constraints while simulating the biomass size distribution of the ecosystem
433 at strongly contrasting locations.

434

435 **2.7 Uncertainty in parameter values**

436 Like any other ecosystem model, there is no single objective choice of parameter values. The sensitivity analysis shown in
437 Figure 5 and discussed in Section 2.5 indicates the consequences of changing the current values. Additionally, it is insightful
438 to understand some of the restrictions on changing the values of the parameters.

439

440 A significant constraint arises from the mass scaling exponent for the biomass observations at the different sites. If the scaling
441 exponent is estimated using data from just a single site (Figure 3a-e), the estimates vary from 0.19 (for FG) to 0.33 (for CCZ).
442 The uncertainty is largest at OM (CI: [0.1,0.3]) and smallest at PAP (CI: [0.24,0.28]). If data from all sites are combined, as
443 described in Section 2.5, then the exponent is estimated at 0.26. A sensitivity analysis was done on how the estimate varies if
444 data from each site in turn is excluded. Excluding data from CCZ, FG, FSC, OM and PAP in turn gives a value that varies
445 between 0.25 and 0.27. Hence, none of the individual sites is having a significant effect on the estimate of the biomass scaling
446 exponent. As described in Section 2.5, the value used for the interference scaling exponent in the model is equivalent to having
447 a biomass scaling exponent of 0.2, below the values obtained from the observations described above. Also as described in
448 Section 2.5, using a greater magnitude value would either lead to a dominance of respiration by larger organisms (rather than

449 the 'energy equivalence' for which there is some observational support) or require an equivalent increase in the magnitude of
450 the scaling exponent for the physiological rates.

451

452 The chosen value of -0.2 for the exponent, β , is within the range of observational estimates. Most of these cluster around -0.2
453 to -0.25 (e.g. Mahout et al., 1995; Ernest et al., 2003). Lower values have been found though. McCoy & Gillooly (2009)
454 reported an exponent of -0.18 for invertebrates. Lower still, McClain et al. (2012) found an exponent of -0.11 for growth of
455 benthic organisms, though noted that this was inconsistent with the exponents they found for respiration and mortality (-0.2
456 and -0.24 respectively). Looking at Figure 4c it is apparent that a using larger magnitude (perhaps to allow a biomass scaling
457 of 0.25 and retain energy equivalence) would lead to a lower average respiration rate across the size classes, as the respiration
458 rate for the smallest size organism cannot get any bigger. An increase in the magnitude of β would, therefore, lead to a reduction
459 in the model estimate of the POC flux and the quality of its fit to observations. For example, if β were changed to -0.25, then
460 r_0 would need to be reduced by a factor of 5 to meet the constraint arising from the maximum bacterial growth rate and the
461 model POC flux estimates would therefore also decrease by a factor of 5, making them a further factor of 5 lower than
462 observations, which the model already underestimates.

463

464 The value chosen for E (0.35 eV) is at the lower end of values derived from observations. For a range of organisms not
465 restricted to marine ones, Savage et al. (2004) found values from 0.35eV to 0.84 eV. McClain et al. (2012) estimated E as
466 0.47eV for respiration and mortality but, with less confidence, 0.16 eV for growth. McCoy & Gillooly (2009) found a value
467 of 0.69 eV for respiration. For comparison, the canonical value for MTE is 0.63 eV (West et al., 1997; Brown et al. 2004). The
468 greatest influence of E is in accounting for significant differences in physiological rates between sites with strongly contrasting
469 temperatures. Changing temperature from 0°C to 10°C with $E=0.35$ eV increases rates by a factor of 1.7. Using $E=0.63$ eV
470 increases rates by a factor of 2.6. However, if $E=0.63$ eV then model-estimated POC fluxes reduce by 24-57% across the sites.
471 This is because the temperature effect $\theta(T)$ takes lower values with the larger E even though the difference between $\theta(T)$ at
472 different temperatures is greater.

473

474 The greatest subjectivity in choice of parameter value is in deciding the ratio $(r_0+m_0)/g_0$ which represents the fraction of the
475 maximum growth rate that an organism achieves at equilibrium. Another perspective on this is that it represents the degree of
476 food-limitation, with a value of 0 representative of total starvation and a value of 1 of a surfeit. We subjectively took this ratio
477 to be 0.5 to be consistent with a food-limited ecosystem which is dependent on material that is already the meagre remains of
478 food that was available to many other organisms as it sank down through the water column (Smith et al., 2008). It should also
479 be appreciated that two locations with high and low food supplies do not necessarily differ in the degree of food limitation, as
480 the population sizes at the two sites will reflect the supplies. It is unlikely that direct observations will become available to
481 constrain this ratio so changes to the current value are more likely to arise from a wish to change one of the three component
482 parameters, g_0 , r_0 or m_0 . Keeping β fixed, the first two of these, g_0 and r_0 , are already at the upper limit of observations whilst
483 m_0 is at the lower limit (Figure 4c). Figure 5 shows that decreasing g_0 by 10% would lead to a ~25% increase in R , TOC and
484 turnover time. It would also increase the time for the perturbed system shown in Figure 7 to recover to 95% of the steady state
485 value from 59 to 71 years. A 10% decrease in r_0 would result in the same decrease in estimated POC flux (or a compensating
486 10% increase in total biomass if it was wanted to preserve the total respiration, Figure 5). The same 10% decrease in r_0 would
487 give a ~10% decrease in R , TOC and turnover time (Figure 5). Finally a 10% increase in m_0 would give a 10% decrease in
488 lifetime for the largest organism. Returning to the ratio $(r_0+m_0)/g_0$, a 10% decrease in r_0 and m_0 would give a 10% decrease in
489 the ratio to 0.45. Decreasing g_0 by 10% would increase it by 11% to 0.56.

490

491 A change in a_0 results in a compensating change in R (Equation 11 shows that they are proportional at steady state) and
492 consequently for the model estimate for TOC. The current value of a_0 was chosen to minimise $\sum_{\text{sites}}[\log(\text{TOC}_{\text{est}}/\text{TOC}_{\text{obs}})]^2$.
493 Obviously, the choice of a_0 is sensitive to the subjective choice of this function. If a simple sum of squared differences was
494 used instead then a choice of $a_0=500$ would be more optimal, giving a lower range of estimates than those given in Figure 4e
495 and Table 2 (9.2 gC/m² CCZ; 2400 gC/m² FG; 560 gC/m² FSC; 960 gC/m² OM; 74 gC/m² PAP). We used the logged ratio to
496 choose a_0 here because of the orders of magnitude difference between observations across the sites (Figure 4e, Table 2). Using
497 a simple sum of squared differences biases the value to best fit observations with largest TOC concentrations. The turnover
498 time of R i.e. the time it takes the POC flux to replenish R if removed, has the same sensitivity to a_0 . This can affect the
499 recovery time to perturbations. Having slower recovery of R increases the recovery time for small organisms that would
500 otherwise recover much quicker than large ones because of higher physiological rates. The small organisms cannot fully
501 recover until R itself is recovered. This is apparent in the example of how BORIS2 may be used to study a response to a change
502 in POC flux given later (Section 3.2 and Figure 7). Using a value of 500 in this example would reduce the time taken for the
503 total biomass to recover to 95% of the steady state value from 59 to 27 years. Turnover times estimated using the model are
504 also affected by choice of a_0 as shown in Figure 5d. They are consistently ~20 years for the current choice of a_0 but reduce to
505 5 years if $a_0=500$.

506

507 **3. Use of BORIS-2**

508 BORIS-2 runs easily and quickly in Matlab (it was developed, tested and run in version 25.1.0.2973910 (R2025a) Update 1,
509 and a steady state solution (Section 2.4) is available for situations where equilibrium is the focus. The model requires only the
510 seafloor temperature and POC flux for a location as inputs. If BORIS-2 is to be used at a specific location then it may be
511 possible to estimate the local POC flux directly using in situ data from sediment traps (e.g. Durden et al. 2020a; Smith et al.,
512 2013), although the resuspension of material means that near seafloor data should be treated with care. Alternatively, if it is
513 intended to use BORIS-2 over larger areas, such as basin scales, then POC flux can be estimated less directly using algorithms
514 which estimate POC flux at any given depth using satellite remote sensing data (Lutz et al., 2007) or using global
515 biogeochemical model output (e.g. Yool et al., 2017; Figure 3.21 of Cooley et al., 2022). Alternatively, observations of
516 sediment community oxygen consumption (SCOC) rates (Smith et al., 2013; Stratmann et al., 2019), which would be expected
517 to roughly balance POC input on timescales for which the system could be viewed as in steady state, could be used. We now
518 give a few examples to illustrate the range of potential uses.

519

520 **3.1 Using the steady state solution**

521 If it is of interest to know how benthic biomass (i.e. the total amount of organisms that can be sustained) varies geographically,
522 it is useful to focus on the annual average biomass such that it can be assumed that the ecosystem is in steady state to first
523 order. This assumption allows Equations 10 and 11 to be used for quicker calculations. In Figure 6, use is made of data for
524 POC flux and temperature at the seafloor to produce a global map of benthic biomass. The POC flux data are generated using
525 the Lutz et al. (2007) algorithm while seafloor temperature data come from the World Ocean Atlas (Reagan et al., 2024).
526 Temperature is largely uniform, with little change across the abyssal plains, or even above seafloor ridges, due to the weak
527 vertical gradients in temperature in the deep ocean. The pattern of low values in subtropics with higher values in tropical,
528 subpolar, polar and coastal regions for the POC flux is similar to that seen in the export of organic material from the ocean
529 surface (e.g. Nowicki et al., 2022) but superimposed on this is the effect of depth. POC flux attenuates strongly with depth
530 (Martin et al., 1987), and a logarithmic scale is needed to capture the variation in seafloor POC flux from shelf to abyssal
531 regions. Because of the largely uniform distribution of seafloor temperature, that of seafloor biomass closely resembles that of
532 the POC flux for much of the ocean. Only in the Mediterranean and Red Sea are the impacts of much higher temperatures
533 visible with lower biomasses relative to the variations in POC fluxes because they need to balance greater physiological rates.

535 3.2 Running the model dynamically

536 The dynamic version of BORIS-2 runs easily on a standard laptop, taking just seconds for a thousand years. The forcing data
537 on POC flux, F , and the temperature, T , can also both vary with time if required. This allows a variety of transient responses
538 to be explored.

539

540 For temperature, long-term temporal change in deep-water temperatures has been detected, but is of a very small magnitude
541 (e.g., <0.002 °C/yr Garry et al., 2019). Stronger fluctuations at a site may arise near the boundaries of warm (e.g. Red Sea,
542 Mediterranean Sea), cool (e.g. Atlantic, Pacific, Indian Oceans), or cold (e.g. Arctic and Southern Oceans) deep waters if their
543 boundaries move in response to natural or climate-change related shifts. For example, at the Arctic-Atlantic transition in the
544 Greenland-Iceland-Faroe-Shetland region a 10°C shift in bottom water temperature can occur over a short spatial (bathymetric)
545 scale (e.g., Turrell et al., 1999) and so a near 10°C shift can occur on short time scales (hours, e.g., Bett, 2001). Generally,
546 though, because a 10°C change in temperature is required to create a roughly factor of 2 change in physiological rates, scenarios
547 where time-varying temperature has a significant impact on biomass are likely to be rare for deeper, off-shelf locations.

548

549 Significant changes in POC flux are more likely. For example, considerable uncertainty remains over the impact of climate
550 change on export of organic carbon from the ocean surface but future changes of up to 41% are possible (Henson et al., 2022).
551 Such changes in POC flux leaving the surface will impact the benthic ecosystem, which is dependent on the fraction of this
552 export that reaches the seafloor. One application of the dynamic version of BORIS-2 therefore is in exploring climate change
553 consequences for the benthos (e.g. Yool et al., 2017). A much simpler example of how the model can be used to study responses
554 to change in POC flux is shown in Figure 7. Here the ecosystem is initially in steady state but then the POC flux is doubled.
555 As is apparent in Figure 3, the different sizes of organisms will have very different biomasses. Hence, for ease of comparison
556 the biomasses and detritus are normalised in Figure 7 by dividing by their final value. Similarly, the variation of physiological
557 rates with size means that response times differ with size of organism. Using a log time scale allows this to be seen more
558 clearly. The smallest size class tracks the response of the detritus closely because faster physiological rates allow these
559 organisms to respond as quickly as the detritus changes. The larger organisms have slower rates and are seen to respond
560 significantly more slowly as a consequence.

561 4 Discussion

562 The BORIS-2 model has been presented. It allows simulations of the benthic community across the full size-range of
563 organisms. A single parameter set has been provided for general use that allows the model to reproduce observed biomass size
564 distributions at five sites contrasting strongly in location, depth and temperature, while meeting other constraints on POC flux,
565 expected physiological limitations of smallest and largest organisms and the amount of organic carbon available for food on
566 the seafloor. It is intended that BORIS-2 be used in a macroecological manner, not to represent the dynamics of specific
567 species. Physiological processes for organisms within the same size class can vary significantly so BORIS-2 is best suited for
568 examining questions related to the overall community or the relative behaviour between size classes. As with any model there
569 are aspects that represent limitations and, as a result, areas for further investigation.

570 4.1 Model assumptions

571 BORIS-2 assumes all organisms are detritivores feeding from the same common resource, detrital organic carbon on the
572 seafloor. In practice, a community will have organisms exhibiting a variety of feeding strategies of which detritivory is just
573 one. Predation, for example, is not captured explicitly by BORIS-2. However, on seafloors deeper than the euphotic zone and

574 outside of chemosynthetic systems, the benthic ecosystem is supported solely by the POC flux and predation is effectively a
575 secondary transformation of that carbon. One interpretation is that BORIS implicitly captures predation in the mortality term
576 but that the gains are distributed across all size ranges rather than received by specific ones. Even from that perspective,
577 BORIS-2 may under-estimate predation because mortality is parameterised based on natural mortality rate data. In the absence
578 of suitable data for predation rates and given the large uncertainties in natural mortality, the magnitude and significance of this
579 underestimate are uncertain. Separate population dynamics for detritivore and predator components of the benthic community
580 have been studied on the shelf (Blanchard et al., 2009) where it was found that predators might display a stronger increase of
581 biomass with size than detritivores. A size-based model presented in the same work to explore this interaction further found
582 that the presence of predators could cause a steepening of the biomass spectrum for detritivores where their size overlapped
583 with the prey range for predators. However, the predators of benthic organisms were assumed to be largely pelagic – a condition
584 that is not experienced in the deep ocean. While a similar coupled approach could be adopted in BORIS-2, the main difficulty
585 in incorporating carnivory into BORIS-2 is the requirement for data on the relative abundance of predators versus non-
586 predators across size classes. Such data are scarce even on the shelf (e.g. Blanchard et al., 2009). Without such information it
587 would be difficult to constrain sufficiently the model parameters.

588

589 An additional facet of the ecosystem that is simplified by BORIS-2 is how the organisms obtain their food. In reality, they
590 may be more or less mobile, allowing them to search for food. They may also be able to filter organic matter from seawater as
591 suspension feeders, intercepting food before it hits the seafloor or exploiting resuspended or advected material. In theory, the
592 parameter a_0 could be modified to reflect greater mobility while the growth parameter, g_0 , could be adjusted to capture the
593 effect of suspension feeding. Once again though, to incorporate such changes would require additional data on the relative
594 abundance of organisms with these different characteristics across size ranges.

595

596 A final assumption of BORIS-2 worth discussing is that no organic material is either refractory or buried. For burial, a fraction
597 of the POC flux and/or the mortality could alternatively be regarded as buried and removed from the system. Note that the
598 POC flux would then have to balance the sum of burial and respiration, so a lower respiration would be required to balance
599 the same POC flux. Estimates for burial vary from 37% of the POC flux arriving at the seafloor on the shelf to 4% in the deep
600 (>2000m) ocean (Dunne et al., 2007). The shelf value is in the absence of perturbations to the sediment such as fishing related
601 trawling. Based on the low deep-sea fraction, burial is omitted. It could easily be added as a (depth-dependent) ‘tax’ on seafloor
602 POC flux if needed subsequently. Regarding refractory organic carbon, it was described in Section 2.4 that it has been assumed
603 that only 5% of the total organic carbon in the surface sediment is readily available to the benthic ecosystem represented by
604 the model. The other 95% is regarded as refractory. Consider two scenarios. The first is that the POC flux arriving at the
605 seafloor is entirely labile and refractory carbon is created only by the seafloor ecosystem. In this case, it would be possible to
606 modify BORIS-2 such that a fraction of mortality passed into a refractory carbon pool rather than into R . In the second scenario,
607 the POC flux has a refractory component. This could be directed straight into a refractory pool. Reality is likely to be some
608 combination of these two scenarios. At steady state, all of the organic carbon entering the refractory pool must either be
609 respired or transformed to labile material and hence returned to R . In the hypothetical case of no respiration, then the flux of
610 organic material into and out of the refractory pool should balance, such that the net flux is zero. This is the implicit assumption
611 in BORIS2, such that this exchange is not modelled. In reality, some of the refractory carbon will be respired, and this could
612 be incorporated in BORIS-2 in the same way as burial, as a simple extra loss; applied to the POC flux or by adding creation
613 and respiration of refractory material by the benthic community as appropriate. In the absence of data from multiple sites for
614 the amount of refractory carbon arriving as POC flux or created by the benthic ecosystem, and the fraction of this that is
615 eventually respired or buried, the dynamics of the refractory pool are omitted. An additional aspect of the refractory carbon
616 dynamics is that there will be a bacterial population carrying out its respiration that is also not captured by BORIS-2. This

617 means that observational estimates of bacterial abundance in seafloor sediments are likely to be higher than those predicted by
618 BORIS-2. With a large pool of refractory carbon (Section 2.5) and associated bacterial doubling time up to thousands of years
619 (Jørgensen & Marshall, 2016), this additional population is likely to be much larger than represented by the smaller size classes
620 of the model that are feeding on *R*. For example, bringing together observations from a site in the abyssal Pacific to apply a
621 linear inverse model for the benthic system including refractory carbon, de Jonge et al. (2020) estimated the prokaryotic
622 biomass to be roughly equivalent to that for megafauna. In BORIS-2, the biomasses for the equivalent (smallest and largest)
623 size classes differ by a factor of 10^5 . This is not a straightforward comparison though as the prokaryotes in the de Jonge (2020)
624 study feed from both labile and refractory material. If future data suggest that bacteria need to be taken out of the allometric
625 framework and treated separately, the biomass estimates of remaining organisms are unlikely to change by more than a factor
626 of two (the extreme case of bacteria having total biomass equal to all other organisms present), with relative biomass of other
627 classes unchanged. To make such a change though would require information on the flux or fraction of organic carbon entering
628 the refractory pool, and the physiological rates of the bacteria ingesting and respiring it. Note that the very definition of
629 “refractory” is itself an uncertainty. The wide flexibility in the structure of molecules of organic carbon means that POC varies
630 widely in how “labile” or “refractory” it is. It is not a simple binary, so this adds a further layer of uncertainty.

631

632 Nevertheless, treating bacteria differently could offer one means of increasing the total respiration to give an improved match
633 to the observations for POC flux and oxygen consumption (Figure 4b). As an extreme example, if only bacteria were allowed
634 access to the POC flux arriving at the seafloor (rather than everything having equal access at present), with the detrital pool
635 being supplied instead by dying bacteria, then the observed biomass of other organisms would be supported by a fraction of
636 the POC flux equivalent to the ratio of mortality to respiration of the bacteria. This ratio is 0.1 for the current parameter set.
637 Without changing any parameter values this approach would therefore decrease the POC flux available for other organisms in
638 Figure 4b by an order of magnitude, with respiration of those organisms now exceeding estimates, as anticipated. It would also
639 increase bacterial biomass by two orders of magnitude (but still three orders of magnitude less than the total). Although there
640 is some evidence that bacteria respond quickly to POC deposition (Sweetman et al., 2018), other similar studies show a wider
641 response, including a rapid response by macrofaunal invertebrates (Witte et al., 2003). Note that the latter studies were
642 conducted at small physical scales that effectively exclude megafaunal invertebrates such that their relative influences and
643 responses are unknowns (Laguionie Marchais et al., 2020). The option to treat bacteria differently is therefore left for future
644 study.

645 **4.2 Other possible model extensions**

646 An aspect of BORIS-2 which may benefit future development is the restricted number of external influences. There are
647 currently only two: the supply of detrital material to the seafloor (POC flux) is the food source for all organisms and ambient
648 temperature is the only control other than size on metabolic rates.

649

650 The effect of oxygen concentration in seawater is not currently included in BORIS-2. Although it has been questioned whether
651 there is clear evidence for an oxygen effect on metabolism (Siebel and Drazen, 2007), a lack of clear response to low oxygen
652 by benthic communities might be as a result of a shift in community composition towards organisms more efficient at extracting
653 oxygen from waters with low concentrations (Childress and Seibel, 1998). That said, under reduced oxygen concentrations
654 there is evidence that macrobenthos shift to smaller body sizes (Pearson & Rosenberg, 1978), while meiobenthos may shift to
655 large body sizes (Moore & Bett, 1989). There may even be a tendency for megabenthos to be eliminated (Pearson & Rosenberg,
656 1978), though they may be enhanced at the peripheries of oxygen minimum zones (OMZs; Levin, 2003). Given the anticipated
657 expansion of oxygen minimum zones through climate change (Busecke et al., 2022), it is worth noting that commonly applied
658 thresholds for hypoxia range from 0.3-4 mgO₂/L, with a modal value of 2 mgO₂/L. However, the lethal and sublethal levels

659 for individual taxa vary greatly (Vaquer & Duarte, 2008). In a formal environmental monitoring context (e.g. EU Water
660 Framework Directive), oxygen concentrations below 4 mg O₂/L are considered to be of concern (Best et al., 2007). There is
661 therefore value in finding a way to incorporate an oxygen effect in BORIS-2 if sites <4 mgO₂/L are of interest, and particularly
662 if concentrations are likely to be below 2 mgO₂/L.

663

664 The impact of seafloor type is another area where BORIS-2 may benefit from further analysis and expansion. At present, for
665 simplicity, BORIS-2 makes no distinction in the nature of the seabed environment, other than bottom water temperature and
666 POC flux. The implicit assumption is that it is applied in a sedimentary environment. In practice, the seafloor represents a
667 range of environments varying on scales from a single manganese nodule to an ocean basin. Seafloor type can influence both
668 motility (with some suspension feeders favouring hard substrata) and the efficiency with which food can be obtained (such as
669 hills or trenches which can focus bottom currents carrying suspended POC). Whether BORIS-2 can be configured for different
670 seabed environments by suitably adjusting parameter values and/or by splitting the ecosystem into populations with different
671 feeding traits is left for future developers.

672 **5 Conclusions**

- 673 • Based on allometric scaling of metabolic processes, the BORIS-2 benthic ecosystem model is capable of simulating
674 population dynamics of organisms ranging in size from bacteria to large metazoans, over 18 orders of magnitude.
- 675 • The only external information required is the POC flux to the seafloor and the ambient temperature.
- 676 • It can be run dynamically but a steady state solution also exists and is given.
- 677 • A parameter set is provided suitable for general use globally and capable of simultaneously providing a good
678 reproduction of observed biomass size spectra at five locations contrasting in depth, food supply, and temperature.
- 679 • This model offers considerable flexibility in application, at a range of scales, from responses to regional perturbations
680 such as deep-sea mining, to studies of climate-driven global change in the benthos.

681

682 **Appendix: Differences between BORIS-1 and BORIS-2**

683 A brief description is given here of the differences between the BORIS-1 and BORIS-2 models. Full details of BORIS-1 can
684 be found in Kelly-Gerreyn (2014).

685 (1) Range of organism sizes reproduced

686 BORIS-2 has been designed to reproduce the full range of benthic organism sizes, whereas BORIS-1 focussed on a
687 limited range of sizes coincident with the data then available for comparison. BORIS-2 overlaps exactly with the 16
688 size classes of BORIS-1; size class 27 of BORIS-2 matches size class 1 of BORIS-1. BORIS-2 therefore extends for
689 26 smaller size classes and 17 larger size classes than BORIS-1, to provide more complete coverage of the range of
690 organism sizes.

691 (2) Choice and representation of physiological/ecological processes

692 Broadly, BORIS-1 and BORIS-2 are structurally similar, with dynamics arising from the three processes of growth,
693 respiration and mortality - but they differ a little in how they do this. First, in BORIS-1 growth is the net effect of
694 ingestion then assimilation. Ingestion was allowed to scale with body size in BORIS-1, but assimilation was just
695 assumed to be a fixed fraction of this. The difference between them was treated as waste and returned to *R*. To simplify
696 this in BORIS-2, a single net growth rate is used, effectively the combined product of ingestion and assimilation.
697 Also, in BORIS-1, growth rate increased linearly with the amount of food available. In practice an organism's ability
698 to ingest and assimilate food cannot increase indefinitely. In BORIS-2 the representation of growth is therefore
699 modified such that it saturates at high food abundance. It is also modified to include the effect of other organisms

700 competing for the limited food supply (Section 2.2). Second, respiration in BORIS-1 is represented as a fraction of
701 growth, and this fraction can vary independently with size. A consequence is that BORIS-1 does not specifically
702 capture basal metabolism, the ‘tax’ paid by any organism just to keep alive. In BORIS-2 respiration is represented as
703 a separate process independent of growth. This better represents basal metabolism. Additionally, an organism will
704 need to fuel active metabolism, the energy requirements above basic maintenance required for such things as
705 movement. In BORIS-2 it is implicitly assumed that this is included in the net growth rate. In BORIS-1 it was also
706 necessary to assume a fixed fraction of POC flux that was respired by organisms not captured by the model. By
707 expanding the size range to cover all organisms in BORIS-2, this assumption (and parameter) is no longer required.
708 Third, a linear mortality parameterisation was used in BORIS-2, in place of the quadratic one used in BORIS-1. This
709 choice is influenced by the impact of the mortality term on organism lifetimes. Temperature has been argued to be
710 the first order control on mortality (e.g. McCoy and Gillooly, 2008). There is found to be roughly a factor of two
711 increase in mortality rate per 10°C increase in environmental temperature. As context, the sites described in Section
712 2.5, which span from shelf to deep ocean, only encompass a 14°C range of temperatures, roughly consistent with a
713 factor less than 4 range in mortalities. The quadratic parameterisation, however, leads to a dependence of the specific
714 rate of mortality on biomass (i.e. with a quadratic mortality term the specific rate is $\mu_i B_i$), and hence on food supply
715 (the POC flux, F); this is because B_i increases with F (see Section 2.4) and F can vary significantly. Assuming the
716 ecosystem is in equilibrium, the respiration of organic carbon on the seafloor can be used to estimate F . Using the
717 large collection of globally distributed Sediment Community Oxygen Consumption (SCOC) data (Stratmann et al.,
718 2019), and limiting to sites less than 6000m in depth and accepting only in situ measurements, SCOC ranges from
719 0.14-110 mmol O₂ m⁻² d⁻¹ (2.5% and 97.5% confidence levels). Excluding shelf regions (<200m depth) reduces the
720 range to 0.08-12 mmol O₂ m⁻² d⁻¹, but it still spans several orders of magnitude. (Note that this result is unaffected by
721 the units as typically a constant factor is used to convert to carbon units.) Such large variations in POC flux would
722 induce similar variability in the longevity of organisms of the same size between different locations if a quadratic
723 mortality parameterisation was used. Therefore, using the linear form of mortality in BORIS-2 avoids a much larger
724 inter-site variability in mortality rate than is currently supported by observations.

725 (3) Observational constraints

726 BORIS-1 was compared to size-resolved data from 3 locations: FG, FSC and OM. While they do contrast in depth
727 and temperature (Table 2) they do not represent the deep seafloor that covers much of the globe. For BORIS-2,
728 additional data from CCZ and PAP are used giving data from two abyssal locations of contrasting food supply. For
729 BORIS-1, the assumption of a fraction respired by non-modelled organisms meant that a comparison to observations
730 of POC flux would be quite subjective. With this restriction removed in BORIS-2 the modelled POC flux is now
731 compared to 2 independent estimates of POC flux at 4 sites and 3 estimates at PAP. For BORIS-1 additional
732 constraints regarding the exponents for size scaling were imposed (see (4) below), as well as a range of expected
733 values for smallest (meiofauna) and largest (macrofauna) organisms. A similar thing to the latter is done for BORIS-
734 2 but, by necessity, for much smaller (bacteria) and larger (megafauna) organisms because of the expanded size range.
735 An additional constraint for BORIS-2 is provided by estimates of TOC in seafloor sediment, which provide a
736 constraint on R .

737 (4) Method of selecting parameter values

738 BORIS-2 has one fewer parameter (7) than BORIS-1 (8). This is despite BORIS-2 incorporating two new processes:
739 temperature sensitive physiology and interference. Without these additions BORIS-2 would have 5 parameters. To
740 find suitable parameter values for BORIS-1 an optimisation algorithm was used. With fewer parameters and a greater
741 set of constraints this was not necessary for BORIS-2. While it might be possible to construct a similar optimisation
742 routine for BORIS-2, it currently does not warrant the effort, particularly given the subjectivity in constructing the

743 necessary cost-function for the optimisation. The user can easily explore parameter space and make a decision on the
744 most suitable parameter values simply using Figure 3 and 4.

745

746 **Code and data availability**

747 All code and data for generating the figures in this paper and for using BORIS either at steady state or dynamically are
748 available on Zenodo at <https://doi.org/10.5281/zenodo.19235638> (Martin et al., 2026). A user manual can be found in the
749 Supplement to this manuscript.

750 **Author contribution**

751 AM was responsible for Writing the original draft and for Software, Validation and Visualisation. AB assisted in model
752 development, testing and sensitivity analysis. All authors were involved in the Conceptualization and Methodology,
753 contributing to the design of the model, and in the review and editing of the Writing. AY, AB, BJB, JD, and SO provided
754 advice related to the Validation of the model.

755 **Competing interests**

756 One author (AY) is a member of the editorial board of journal "Geoscientific Model Development".

757 **Acknowledgements**

758 This work was supported by the NERC National Capability funded CLASS project (NE/R015953/1), the NERC AtlantiS
759 (Atlantic Climate and Environment Impacts) programme ([https://noc.ac.uk/projects/atlantic-climate-environment-strategic-](https://noc.ac.uk/projects/atlantic-climate-environment-strategic-science)
760 [science](https://noc.ac.uk/projects/atlantic-climate-environment-strategic-science), NE/Y005589/1) and the NERC SMARTEX project (NE/T003537/1). We would also like to acknowledge Dr Jessica
761 Luo and an anonymous reviewer for their insightful and constructive reviews of this manuscript.

762

763 **References**

764 Atwood, T. B., Witt, A., Mayorga, J., Hammill, E. and Sala, E.: Global patterns in marine sediment carbon stocks, *Front. Mar.*
765 *Sci.*, 7, 165, doi: 10.3389/fmars.2020.00165, 2020.

766 Benoist, N.: Advances in the state-of-the-art in the quantitative ecology of the marine megabenthos, PhD thesis, University of
767 Southampton, 289 pages, 2020

768 Benoist, N. M. A., Bett, B.J., Morris, K.J. and Ruhl, H.A.: A generalised volumetric method to estimate the biomass of
769 photographically surveyed benthic megafauna, *Progress in Oceanography*, 178, 102188, doi: 10.1016/j.pocean.2019.102188,
770 2019

771 Best, M. A., Wither, A. W. and Coates, S.: Dissolved oxygen as a physico-chemical supporting element in the Water
772 Framework Directive, *Marine Pollution Bulletin*, 55, 1–6, doi: 10.1016/j.marpolbul.2006.08.037, 2007.

773 Bett, B.J.: UK Atlantic Margin Environmental Survey: Introduction and overview of bathyal benthic ecology, *Continental*
774 *Shelf Research*, 21, 917-956, doi: 10.1016/S0278-4343(00)00119-9, 2001.

775 Blanchard, J. L., et al.: How does abundance scale with body size in coupled size-structured food webs?, *Journal of Animal*
776 *Ecology* 78(1), 270-280, 2009.

777 Blanchard, J.L., Law, R., Castle, M.D. et al.: Coupled energy pathways and the resilience of size-structured food webs,
778 *Theoretical Ecology*, 4, 289–300, doi.org/10.1007/s12080-010-0078-9, 2011

779 Brey, T.: An empirical model for estimating aquatic invertebrate respiration, *Methods in Ecology and Evolution*, 1, 92-101,
780 doi:10.1111/j.2041-210X.2009.00008.x, 2010

781 Brey, T.: A multi-parameter artificial neural network model to estimate macrobenthic invertebrate productivity and production.
782 *Limnology and Oceanography: Methods*, 10, 581-589, doi.org/10.4319/lom.2012.10.581, 2012

783 Brown, J.H., Gillooly, J.F., Allen, A.P., Savage, V.M., and West, G.B.: Toward a metabolic theory of ecology. *Ecology*, 85,
784 1771-1789, doi: 10.1890/03-9000, 2004.

785 Burd, A. B.: Modeling the vertical flux of organic carbon in the global ocean, *Annual Review of Marine Science*, 16, 135-161.
786 doi.org/10.1146/annurev-marine-022123-102516, 2024

787 Busecke, J. J. M., Resplandy, L., Ditkovsky, S. J., and John, J. G.: Diverging fates of the Pacific Ocean Oxygen Minimum
788 Zone and its core in a warming world, *AGU Advances*, 3(6), e2021AV000470. doi: 10.1029/2021AV000470, 2022.

789 Butenschön, M., Clark, J., Aldridge, J.N., Allen, J.I., Artioli, Y., Blackford, J., Bruggeman, J., Cazenave, P., Ciavatta, S., Kay,
790 S., Lessin, G., van Leeuwen, S., van der Molen, J., de Mora, L., Polimene, L., Sailley, S., Stephens, N., Torres, R.: ERSEM
791 15.06: a generic model for marine biogeochemistry and the ecosystem dynamics of the lower trophic levels, *Geoscientific
792 Model Development*, 9(4), 1293–1339, doi: 10.5194/gmd-9-1293-2016, 2016

793 Childress, J. J. and B. A. Seibel: Life at stable low oxygen levels: adaptations of animals to oceanic oxygen minimum layers,
794 *Journal of Experimental Biology* 201(8), 1223-1232, 1998.

795 Cooley, S., D. Schoeman, L. Bopp, P. Boyd, S. Donner, D.Y. Ghebrehiwet, S.-I. Ito, W. Kiessling, P. Martinetto, E. Ojea, M.-
796 F. Racault, B. Rost, and M. Skern-Mauritzen: Ocean and Coastal Ecosystems and their Services. In: *Climate Change 2022:
797 Impacts, Adaptation, and Vulnerability. Contribution of Working Group II to the Sixth Assessment Report of the
798 Intergovernmental Panel on Climate Change* [H.-O. Pörtner, D.C. Roberts, M. Tignor, E.S. Poloczanska, K. Mintenbeck, A.
799 Alegría, M. Craig, S. Langsdorf, S. Löschke, V. Möller, A. Okem, B. Rama (eds.)]. Cambridge University Press, Cambridge,
800 UK and New York, NY, USA, pp. 379-550, doi:10.1017/9781009325844.005, 2022.

801 Cusson, M., & Bourget, E.: Global patterns of macroinvertebrate production in marine benthic habitats [Review]. *Marine
802 Ecology Progress Series*, 297, 1-14. doi.org/10.3354/meps297001, 2005.

803 Damuth J.: A macroevolutionary explanation for energy equivalence in the scaling of body size and population density, *Am.
804 Nat.*, 169, 621-31, doi: 10.1086/513495, 2007.

805 DeAngelis, D. L., Goldstein, R. A., and O'Neill, R. V.: A Model for trophic interaction, *Ecology*, 56(4), 881–892.
806 doi.org/10.2307/1936298, 1975.

807 de Jonge, D. S. W., Stratmann, T., Lins, L., Vanreusel, A., Purser, A., Marcon, Y., Rodrigues, C. F., Ravara, A., Esquete, P.,
808 Cunha, M. R., Simon-Lledó, E., van Breugel, P., Sweetman, A. K., Soetaert, K. and van Oevelen, D.: Abyssal food-web model
809 indicates faunal carbon flow recovery and impaired microbial loop 26 years after a sediment disturbance experiment, *Progress
810 in Oceanography*, 189, 102446, doi.org/10.1016/j.pocean.2020.102446, 2020.

811 Demopoulos, A. W. J., Smith, C. R., and Tyler, P. A.: The deep Indian Ocean floor, p. 219–237. In P. A. Tyler [ed.],
812 *Ecosystems of the world 28. Ecosystems of the deep oceans*. Elsevier, 2003.

813 Dixon, J. L. and Turley, C. M.: Measuring bacterial production in deep-sea sediments using ³H-Thymidine incorporation:
814 ecological significance, *Microbial Ecology*, 42(4), 549-561, 2001

815 Dunne, J. P., Sarmiento, J. L. and Gnanadesikan, A.: A synthesis of global particle export from the surface ocean and cycling
816 through the ocean interior and on the seafloor, *Global Biogeochemical Cycles*, 21(4), doi.org/10.1029/2006GB002907, 2007.

817 Durden, J.M., Bett, B.J., Ruhl, H.A.: The hemisessile lifestyle and feeding strategies of *Iosactis vagabunda* (Actiniaria,
818 Iosactiidae), a dominant megafaunal species of the Porcupine Abyssal Plain, *Deep-Sea Research Part I*, 102, 72-77.
819 10.1016/j.dsr.2015.04.010, 2015.

820 Durden, J.M., Ruhl, H.A., Pebody, C., Blackbird, S.J. and van Oevelen, D.: Differences in the carbon flows in the benthic food
821 webs of abyssal hill and plain habitats, *Limnology and Oceanography*, 62 (4). 1771-1782. doi.org/10.1002/lno.10532, 2017.

822 Durden, J.M., Bett, B.J., Huffard, C., Ruhl, H.A., and Smith Jr, K.L.: Abyssal deposit-feeding rates consistent with the
823 Metabolic Theory of Ecology, *Ecology*, 100, e02564. 10.1002/ecy.2564, 2019.

824 Durden, J.M., Bett, B.J., Huffard, C.L., Pebody, C., Ruhl, H.A., and Smith, K.L.: Response of deep-sea deposit-feeders to
825 detrital inputs: A comparison of two abyssal time-series sites. *Deep Sea Research Part II*, 173, 104677.
826 10.1016/j.dsr2.2019.104677, 2020a.

827 Durden, J. M., Bett, B. J. and Ruhl, H.A.: Subtle variation in abyssal terrain induces significant change in benthic megafaunal
828 abundance, diversity, and community structure, *Prog. Oceanogr.*, 186, 102395, doi: 10.1016/j.pocean.2020.102395, 2020b.

829 Edwards, A.M., Robinson, J.P.W., Plank, M.J., Baum, J.K. and Blanchard, J.L.: Testing and recommending methods for fitting
830 size spectra to data, *Methods Ecol. Evol.*, 8, 57-67, doi: 10.1111/2041-210X.12641, 2017.

831 Edwards, A.M., Robinson, J.P.W., Blanchard, J.L., Baum, J.K. and Plank M.J.: Accounting for the bin structure of data
832 removes bias when fitting size spectra, *Mar. Ecol. Prog. Ser.*, 636,19-33, doi: 10.3354/meps13230, 2020.

833 Ehrnsten, E., Norkko, A., Timmermann, K. and Gustafsson, B. G.: Benthic-pelagic coupling in coastal seas – Modelling
834 macrofaunal biomass and carbon processing in response to organic matter supply, *Journal of Marine Systems*, 196, 36-47,
835 2018.

836 Ernest, S. K. M., Enquist, B. J., Brown, J. H., Charnov, E. L., Gillooly, J. F., Savage, V. M., White, E. P., Smith, F. A., Hadly,
837 E. A., Haskell, J. P., Lyons, S. K., Maurer, B. A., Niklas, K. J. and Tiffney, B.: Thermodynamic and metabolic effects on the
838 scaling of production and population energy use, *Ecology Letters*, 6(11), 990-995, doi.org/10.1046/j.1461-0248.2003.00526.x,
839 2003

840 Garry, F. K., McDonagh, E. L., Blaker, A. T., Roberts, C. D., Desbruyères, D. G., Frajka-Williams, E., King, B. A.: Model-
841 derived uncertainties in deep ocean temperature trends between 1990 and 2010, *Journal of Geophysical Research: Oceans*,
842 124, 1155–1169. doi: 10.1029/2018JC014225, 2019.

843 Gillooly J.F., Brown, J.H., West, G.B., Savage, V.M. and Charno, E.L.: Effects of Size and Temperature on Metabolic Rate,
844 *Science*, 293, 2248-2251, doi: 10.1126/science.1061967, 2001

845 Giovannelli, D., Molari, M., d’Errico, G., Baldrighi, E., Pala, C., Manini, E.: Large-Scale Distribution and Activity of
846 Prokaryotes in Deep-Sea Surface Sediments of the Mediterranean Sea and the Adjacent Atlantic Ocean, *PLoS ONE*, 8,
847 10.1371/journal.pone.0072996, 2013

848 Glazier, D. S.: Variable metabolic scaling breaks the law: from ‘Newtonian’ to ‘Darwinian’ approaches, *Proc. R. Soc.*
849 *B.28920221605*, doi.org/10.1098/rspb.2022.1605, 2022.

850 Harris, P.T., Macmillan-Lawler, M., Rupp, J., and Baker, E.K. Geomorphology of the oceans. *Marine Geology*, 352, 4-24.
851 10.1016/j.margeo.2014.01.011, 2014.

852 Hartman, S.E., Bett, B.J., Durden, J.M., Henson, S.A., Iversen, M., Jeffrey, R.M., Horton, T., Lampitt, R., and Gates, A.R.:
853 Enduring science: Three decades of observing the Northeast Atlantic from the Porcupine Abyssal Plain Sustained Observatory
854 (PAP-SO), *Progress in Oceanography*, 191, 102508, doi: 10.1016/j.pocean.2020.102508, 2021.

855 Hatton, I. A., Heneghan, R. F., Bar-On, Y. M., & Galbraith, E. D.: The global ocean size spectrum from bacteria to whales.
856 *Science Advances*, 7, 10.1126/sciadv.abh3732, 2021

857 Henson, S.A., Laufkötter, C., Leung, S. et al.: Uncertain response of ocean biological carbon export in a changing world. *Nat.*
858 *Geosci.* 15, 248–254, doi.org/10.1038/s41561-022-00927-0, 2022

859 Iken, K., Brey, T., Wand, U., Voigt, J., and Junghans, P.: Food web structure of the benthic community at the Porcupine
860 Abyssal Plain (NE Atlantic): a stable isotope analysis, *Progress in Oceanography*, 50, 383-405, doi: 10.1016/S0079-
861 6611(01)00062-3, 2001.

862 Isaac, N.J.B., Carbone, C.: Why are metabolic scaling exponents so controversial? Quantifying variance and testing
863 hypotheses. *Ecology Letters*, 13, 728-735. 10.1111/j.1461-0248.2010.01461.x, 2010.

864 Jones, D.O.B., Yool, A., Wei, C.-L., Henson, S.A., Ruhl, H.A., Watson, R.A. and Gehlen, M.: Global reductions in seafloor
865 biomass in response to climate change, *Glob Change Biol*, 20: 1861-1872, doi.org/10.1111/gcb.12480, 2014.

866 Jørgensen, B. B., and Marshall, I. P. G.: Slow Microbial Life in the Seabed. *Annual Review of Marine Science*, 8, 311-332.
867 doi.org/10.1146/annurev-marine-010814-015535, 2016.

868 Kaariainen J. I. and Bett, B. J.: Evidence for benthic body size miniaturization in the deep sea, *Journal of the Marine Biological*
869 *Association of the United Kingdom*, 86, 1339-1345, doi: 10.1017/S0025315406014366, 2006.

870 Karp, M. A. et al.: Increasing the uptake of multispecies models in fisheries management, *ICES Journal of Marine Science*,
871 80(2), 243-257, doi.org/10.1093/icesjms/fsad001, 2023.

872 Kelly-Gerreyn, B. A., Martin, A. P., Bett, B. J., Anderson, T. R., Kaariainen, J. I., Main, C. E., Marcinko, C. J. and Yool, A.:
873 Benthic biomass size spectra in shelf and deep-sea sediments, *Biogeosciences*, 11, 6401-6416, doi:10.5194/bg-11-6401-2014,
874 2014.

875 Kwiatkowski, L., Torres, O., Bopp, L., Aumont, O., Chamberlain, M., Christian, J. R., Dunne, J. P., Gehlen, M., Ilyina, T.,
876 John, J. G., Lenton, A., Li, H., Lovenduski, N. S., Orr, J. C., Palmieri, J., Santana-Falcón, Y., Schwinger, J., Séférian, R.,
877 Stock, C. A., Tagliabue, A., Takano, Y., Tjiputra, J., Toyama, K., Tsujino, H., Watanabe, M., Yamamoto, A., Yool, A., and
878 Ziehn, T.: Twenty-first century ocean warming, acidification, deoxygenation, and upper-ocean nutrient and primary production
879 decline from CMIP6 model projections, *Biogeosciences*, 17, 3439–3470, <https://doi.org/10.5194/bg-17-3439-2020>, 2020.

880 Laguionie Marchais, C., Bett, B. J., Paterson, G. L. J., Smith, K. L., and Ruhl, H. A.: Using metabolic theory to assess structure
881 and function in the deep-sea benthos, including microbial and metazoan dominance, *Deep-Sea Research II*, 173, 104762. doi:
882 10.1016/j.dsr2.2020.104762, 2020.

883 Lampitt, R.S., Newton, P.P., Jickells, T.D., Thomson, J., King, P.: Near-bottom particle flux in the abyssal northeast Atlantic,
884 *Deep Sea Research Part II*, 47, Issues 9–11, 10.1016/S0967-0645(00)00016-3, 2000

885 Lampitt, R.S., Bett, B.J., Kiriakoulakis, K., Popova, E.E., Ragueneau, O., Vangriesheim, A., and Wolff, G.A.: Material supply
886 to the abyssal seafloor in the Northeast Atlantic, *Progress in Oceanography*, 50, 10.1016/S0079-6611(01)00047-7, 2001.

887 Lampitt R.S., and Pebody C.A.: Sediment Trap data from the Porcupine Abyssal Plain Sustained Observatory (PAP-SO) site
888 on PAP3 mooring at 3000 metres April 1989 - June 2019 Version 2. NERC EDS British Oceanographic Data Centre NOC.
889 doi:10.5285/06bd25d5-fcd3-0f63-e063-6c86abc0481e, 2023.

890 Lampitt, R.S., Briggs, N., Cael, B.B., Espinola, B., Hélaouët, P., Henson, S.A., Norrbin, F., Pebody, C.A. and Smeed, D.:
891 Deep ocean particle flux in the Northeast Atlantic over the past 30 years: carbon sequestration is controlled by ecosystem
892 structure in the upper ocean, *Front. Earth Sci.*, 11:1176196, doi: 10.3389/feart.2023.1176196, 2023.

893 Levin, L.A.: Oxygen minimum zone Benthos: Adaptation and community response to hypoxia, *Oceanography and Marine*
894 *Biology*, 41, 1-45, 2003.

895 Luef, B., Frischkorn, K. R., Wrighton, K. C., Holman, H.-Y. N., Birarda, G., Thomas, B. C., Singh, A., Williams, K. H.,
896 Siegerist, C. E., Tringe, S. G., Downing, K. H., Comolli, L. R. and Banfield, J. F.: Diverse uncultivated ultra-small bacterial
897 cells in groundwater, *Nature Communications*, 6(1), 6372, doi.org/10.1038/ncomms7372, 2015

898 Lutz, M. J., Caldeira, K., Dunbar, R.B., and Behrenfeld, M.J.: Seasonal rhythms of net primary production and
899 particulate organic carbon flux to depth describe the efficiency of biological pump in the global ocean, *J. Geophys. Res.*, 112,
900 C10011, doi:10.1029/2006JC003706, 2007

901 McCarthy J. K., Dwyer J. M. and Mokany, K.: A regional-scale assessment of using metabolic scaling theory to predict
902 ecosystem properties, *Proc. R. Soc. B*.28620192221, doi.org/10.1098/rspb.2019.2221, 2019.

903 McClain, C.R., Allen, A.P., Tittensor, D.P., and Rex, M.A.: Energetics of life on the deep seafloor, *Proceedings of the National*
904 *Academy of Sciences*, 109, 15366-15371. 10.1073/pnas.1208976109, 2012.

905 McCoy, M. W. and Gillooly, J. F.: Predicting natural mortality rates of plants and animals, *Ecology Letters*, 11(7), 710-716,
906 doi.org/10.1111/j.1461-0248.2008.01190.x, 2008.

907 McCoy, M. W. and Gillooly, J. F.: Predicting natural mortality rates of plants and animals, *Corrigendum, Ecology Letters*,
908 12(7), 731-733, doi.org/10.1111/j.1461-0248.2009.01338.x, 2009

909 Mahaut, M. L., Sibuet, M. and Shirayama, Y.: Weight-dependent respiration rates in deep-sea organisms, *Deep Sea Research*
910 Part I, 42(9), 1575-1582, 1995.

911 Martin, A. P., Benoist, N., Bett, B. J., Brombacher, A., Durden, J., Oliver, S., and Yool, A.: Matlab code associated with the
912 BORIS-2 benthic ecosystem model, Zenodo, <https://doi.org/10.5281/zenodo.19235638>, 2026.

913 Martin, J. H., Knauer, G. A., Karl, D. M., and Broenkow, W. W.: VERTEX: carbon cycling in the northeast Pacific, *Deep Sea*
914 *Research Part A.*, 34(2), 267-285, doi.org/10.1016/0198-0149(87)90086-0, 1987.

915 Molony, C.L., Field, J.G.: General allometric equations for rates of nutrient uptake, ingestion, and respiration in plankton
916 organisms. *Limnology Oceanography*, 34, 1290-1299, 1989.

917 Moore, C. G. and Bett, B. J.: The use of meiofauna in marine pollution impact assessment, *Zoological Journal of the Linnean*
918 *Society*, 96, 263–280, doi: 10.1111/j.1096-3642.1989.tb02260.x, 1989.

919 Morris, K. J., Bett, B. J., Durden, J. M., Benoist, N. M. A., Huvenne, V. A. I., Jones, D. O. B., Robert, K., Ichino, M. C., Wolff,
920 G. A., Ruhl, H. A.: Landscape-scale spatial heterogeneity in phytodetrital cover and megafauna biomass in the abyss links to
921 modest topographic variation. *Sci Rep* 6, 34080. Doi: 10.1038/srep34080, 2016.

922 Nagy, K.A.: Field Metabolic Rate and Food Requirement Scaling in Mammals and Birds. *Ecological Monographs*, 57, 112-
923 128. 10.2307/1942620, 1987.

924 Niven, J. E. and Scharlemann, J. P.W.: Do insect metabolic rates at rest and during flight scale with body mass? *Biol.*
925 *Lett.*1346–349, doi.org/10.1098/rsbl.2005.0311, 2005.

926 Nowicki, M., DeVries, T. and Siegel, D. A.: Quantifying the Carbon Export and Sequestration Pathways of the Ocean's
927 Biological Carbon Pump, *Global Biogeochemical Cycles*, 36(3), e2021GB007083, doi.org/10.1029/2021GB007083, 2022

928 Parameswaran, N., González, E., Burwicz-Galerie, E., Braack, M., and Wallmann, K.: NN-TOC v1: global prediction of total
929 organic carbon in marine sediments using deep neural networks, *EGU sphere [preprint]*, [https://doi.org/10.5194/egusphere-](https://doi.org/10.5194/egusphere-2024-1360)
930 2024-1360, 2024.

931 Pearson, T.H. and Rosenberg, R.: Macrobenthic Succession in Relation to Organic Enrichment and Pollution of the Marine
932 Environment. *Oceanography and Marine Biology—An Annual Review*, 16, 229-311, 1978.

933 O’Dor, R.K., Fennel, K. and Berghe, E. V.: A one ocean model of biodiversity, *Deep Sea Research Part II*, 56, 1816-1823,
934 doi.org/10.1016/j.dsr2.2009.05.023, 2009.

935 Quiroga, E., Quiñones, R., Palma, M., Sellanes, J., Gallardo, V.A., Gerdes, D., Rowe, G.: Biomass size-spectra of
936 macrobenthic communities in the oxygen minimum zone off Chile, *Estuarine, Coastal and Shelf Science*, 62(1), 217-231, 2005

937 Reagan, J. R., Boyer, T. P., García, H. E., Locarnini, R. A., Baranova, O. K., Bouchard, C., Cross, S. L., Mishonov, A. V.,
938 Paver, C. R., Seidov, D., Wang, Z., Dukhovskoy, D.: *World Ocean Atlas 2023*. NOAA National Centers for Environmental
939 Information. Dataset: NCEI Accession 0270533, 2024

940 Ramirez-Llodra, E., Brandt, A., Danovaro, R., De Mol, B., Escobar, E., German, C.R., Levin, L.A., Arbizu, P.M., Menot, L.,
941 Buhl-Mortensen, P., Narayanaswamy, B.E., Smith, C.R., Tittensor, D.P., Tyler, P.A., Vanreusel, A., Vecchione, M.: Deep,
942 diverse and definitely different: unique attributes of the world's largest ecosystem. *Biogeosciences*, 7, 2851-2899, 2010.
943 10.5194/bg-7-2851-2010

944 Reiss, H., et al.: Benthos distribution modelling and its relevance for marine ecosystem management, *Ices Journal of Marine*
945 *Science*, 72(2), 297-315, 2014.

946 Ruhl, H. A.: Community change in the variable resource habitat of the abyssal northeast Pacific, *Ecology*, 89(4), 991-1000,
947 2008.

948 Ruhl, H. A., Ellena, J. A., and Smith, Jr, K. J.: Connections between climate, food limitation, and carbon cycling in abyssal
949 sediment communities, *Proceedings of the National Academy of Sciences*, 105(44), 17006-17011, 2008.

950 Ruhl, H.A., Bett, B.J., Ingels, J., Martin, A., Gates, A.R., Yool, A., Benoist, N.M.A., Appeltans, W., Howell, K.L., Danovaro,
951 R.: Integrating ocean observations across body-size classes to deliver benthic invertebrate abundance and distribution
952 information. *Limnol. Oceanogr. Lett.*, 8, 692-706, doi: 10.1002/lol2.10332, 2023.

953 Savage, Van M., Gillooly, James F., Brown, James H., West, Geoffrey B. and Charnov, Eric L.: Effects of Body Size and
954 Temperature on Population Growth. *The American Naturalist*, 163(3), 429-441, doi.org/10.1086/381872, 2004

955 Seibel, B.A., and Drazen, J.C.: The rate of metabolism in marine animals: environmental constraints, ecological demands and
956 energetic opportunities. *Philosophical Transactions of the Royal Society B: Biological Sciences*, 362, 2061-2078.
957 10.1098/rstb.2007.2101, 2007.

958 Simon-Lledó, E., Bett, B. J., Huvenne, V. A. I., Schoening, T., Benoist, N. M. A., Jeffreys, R. M., Durden, J. M., Jones, D. O.
959 B.: Megafaunal variation in the abyssal landscape of the Clarion Clipperton Zone, *Progress in Oceanography*, 170, 119-133,
960 doi: 10.1016/j.pocean.2018.11.003, 2019

961 Smith, C. R., De Leo, F. C., Bernardino, A. F., Sweetman, A. K. and Arbizu, P. M.: Abyssal food limitation, ecosystem
962 structure and climate change, *Trends in Ecology & Evolution*, 23(9), 518-528, doi.org/10.1016/j.tree.2008.05.002, 2008

963 Smith, K.L., Ruhl, H.A., Kahru, M., Huffard, C.L., Sherman, A.D.: Deep ocean communities impacted by changing climate
964 over 24 y in the abyssal northeast Pacific Ocean, *Proceedings of the National Academy of Sciences*, 110.
965 10.1073/pnas.1315447110, 2013.

966 Soetaert, K., and van Oevelen, D.: Modeling food web interactions in benthic deep-sea ecosystems: A practical guide,
967 *Oceanography*, 22(1), 128–143, <https://doi.org/10.5670/oceanog.2009.13>, 2009.

968 Stratmann, T., et al.: The SCOC database, a large, open, and global database with sediment community oxygen consumption
969 rates, *Scientific Data* 6(1), 242, 2019.

970 Sweetman, A.K., Smith, C.R., Shulse, C.N., Maillot, B., Lindh, M., Church, M.J., Meyer, K.S., van Oevelen, D., Stratmann,
971 T. and Gooday, A.J.: Key role of bacteria in the short-term cycling of carbon at the abyssal seafloor in a low particulate organic
972 carbon flux region of the eastern Pacific Ocean. *Limnol Oceanogr*, 64, 10.1002/lno.11069, 2019.

973 Turrell, W. R., Slessor, G., Adams, R. D., Payne, R., and Gillibrand, P. A.: Decadal variability in the composition of Faroe
974 Shetland Channel bottom water. *Deep Sea Research I*, 46, 1-25. doi: 10.1016/S0967-0637(98)00067-3, 1999.

975 van Oevelen, D., Soetaert, K., Garcia, R., de Stigter, H. C., Cunha, M. R., Pusceddu, A. and Danovaro, R.: Canyon conditions
976 impact carbon flows in food webs of three sections of the Nazaré canyon, *Deep Sea Research Part II: Topical Studies in*
977 *Oceanography*, 58(23), 2461-2476, doi.org/10.1016/j.dsr2.2011.04.009, 2011a

978 van Oevelen, D., Bergmann, M., Soetaert, K., Bauerfeind, E., Hasemann, C., Klages, M., Schewe, I., Soltwedel, T. and
979 Budaeva, N. E.: Carbon flows in the benthic food web at the deep-sea observatory HAUSGARTEN (Fram Strait). *Deep Sea*
980 *Research Part I: Oceanographic Research Papers*, 58(11), 1069-1083, doi.org/10.1016/j.dsr.2011.08.002, 2011b.

981 Vaquer-Sunyer, R. and Carlos M. Duarte: Thresholds of hypoxia for marine biodiversity, *PNAS*, 105, 15452-15457. doi:
982 10.1073/pnas.0803833105, 2008.

983 Wei C-L, Rowe G.T., Escobar-Briones E., Boetius A., Soltwedel T., et al.: Global patterns and predictions of seafloor biomass
984 using random forests, *PLoS ONE* 5(12), e15323, doi:10.1371/journal.pone.0015323, 2010.

985 West, G. B., Brown, J. H. and Enquist, B. J.: A General Model for the Origin of Allometric Scaling Laws in Biology, *Science*,
986 276(5309), 122-126, 1997.

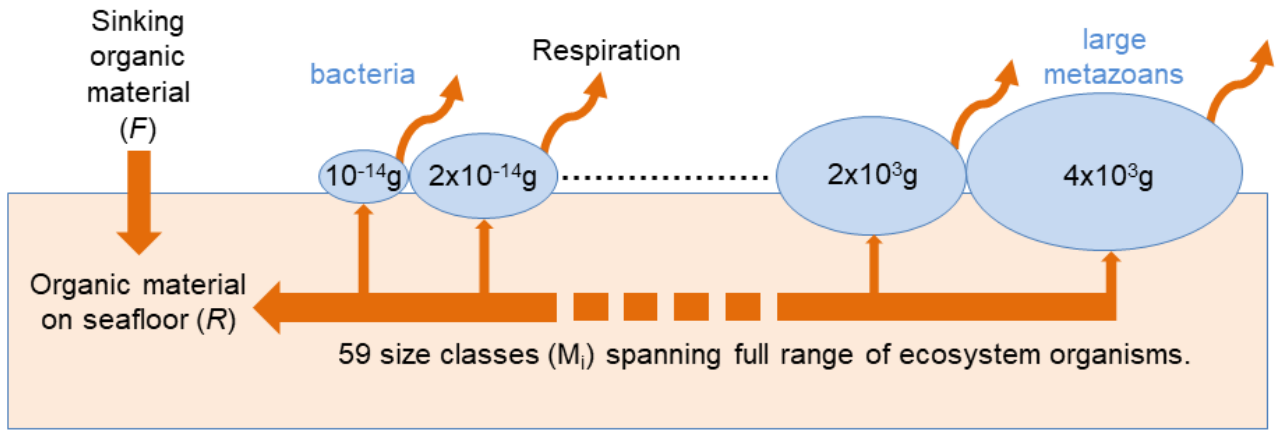
987 White, E.P., Ernest, S.K.M., Kerkhoff, A.J., Enquist, B.J.: Relationships between body size and abundance in ecology, *Trends*
988 *in Ecology & Evolution*, 22, 323-330. 10.1016/j.tree.2007.03.007, 2007

989 Witte, U., Wenzhöfer, F., Sommer, S., Boetius, A., Heinz, P., Aberle, N., Sand, M., Cremer, A., Abraham, W. R., Jørgensen,
990 B. B., and Pfannkuche, O.: In situ experimental evidence of the fate of a phytodetritus pulse at the abyssal sea floor, *Nature*,
991 424, 10.1038/nature01799, 2003.

992 Woodson, C., Schramski, J.R. & Joye, S.B.: A unifying theory for top-heavy ecosystem structure in the ocean. *Nat Commun*
993 **9**, 23 (2018). <https://doi.org/10.1038/s41467-017-02450-y>, 2018.

994 Yool, A, Martin, A.P., Anderson, T.R., Bett, B.J., Jones, D.O.B., Ruhl, H.A.: Big in the benthos: Future change of seafloor
995 community biomass in a global, body size-resolved model, *Global Change Biology*, **23**, 3554–3566.
996 doi.org/10.1111/gcb.13680, 2017

997



1000

1001

1002

1003

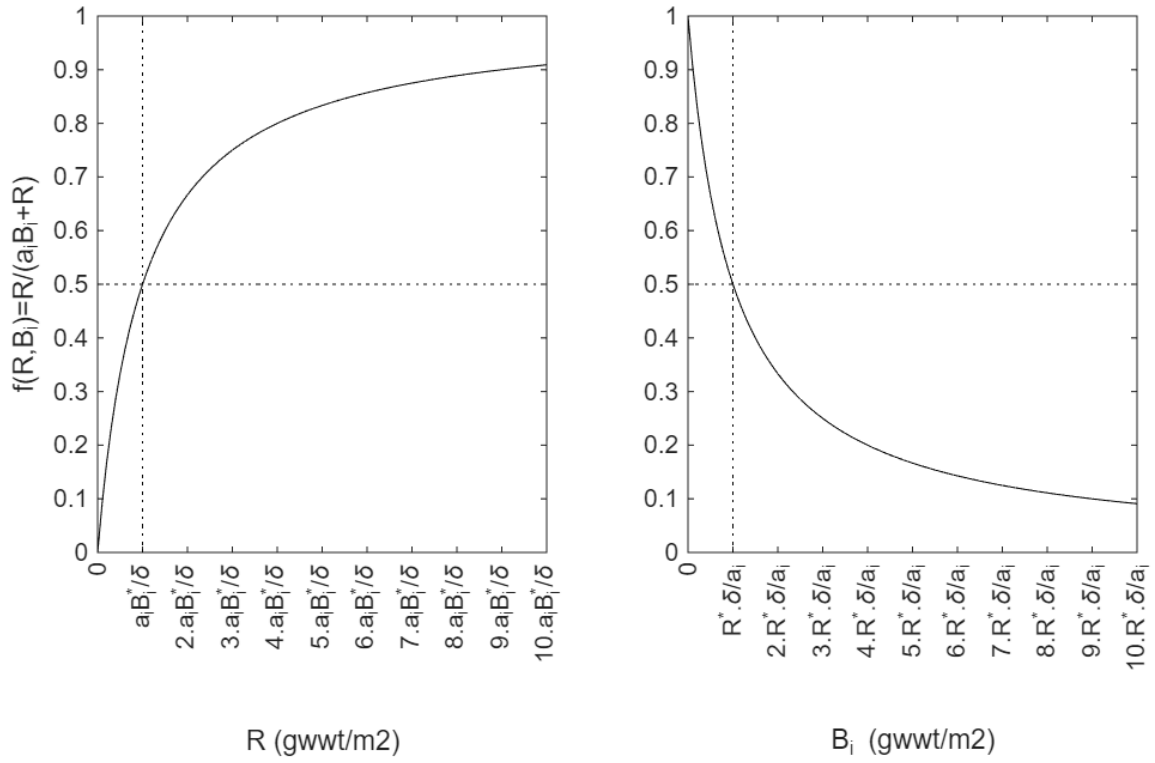
1004

1005

1006

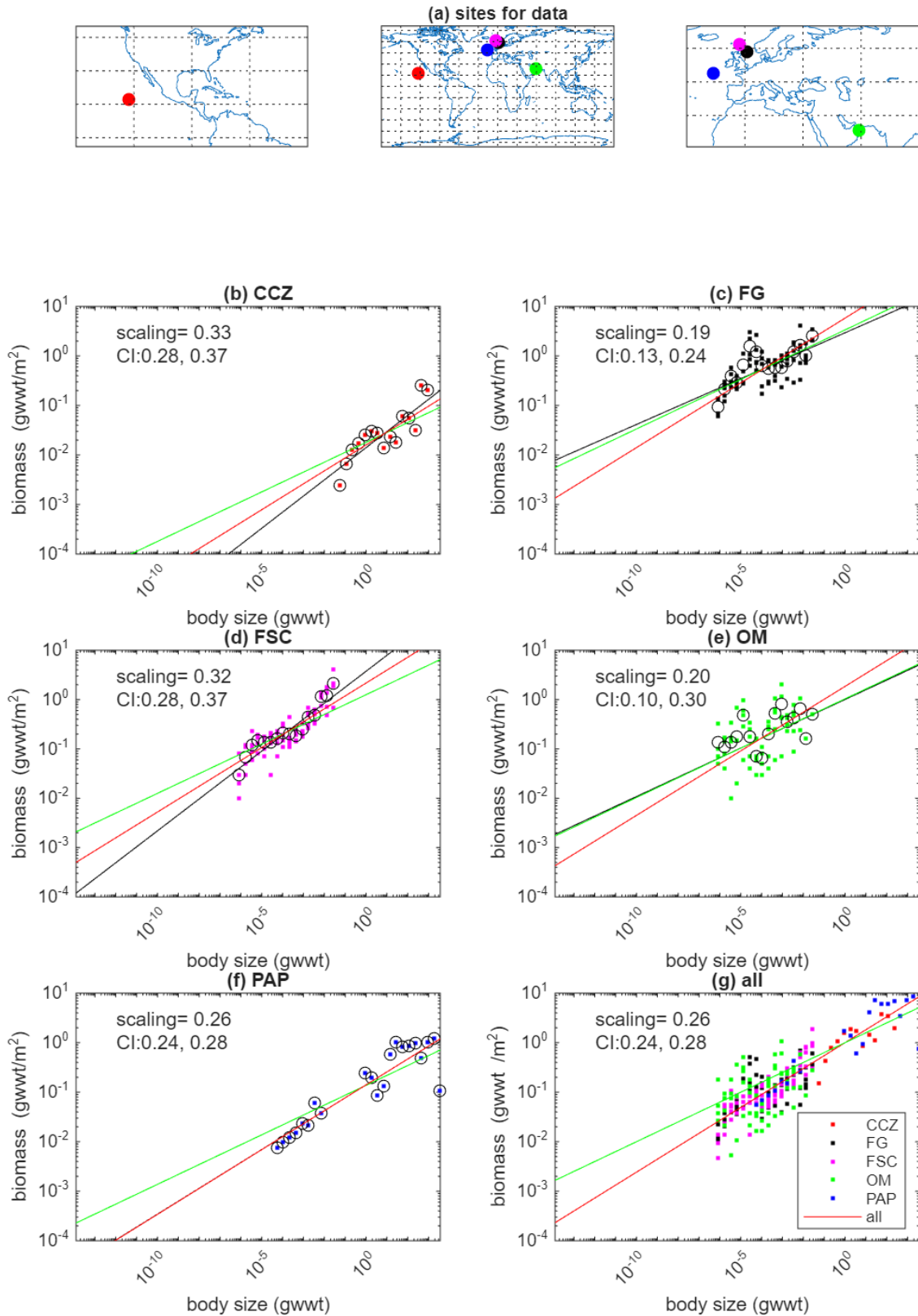
1007

Figure 1: Schematic of BORIS-2. Within each size class (M_i) the total biomass (B_i) is controlled by growth, using organic material on the seafloor, and losses to respiration and mortality. Organic material (R) accumulates on the seafloor from the deposition of sinking particulate organic carbon (F) and mortality of benthic organisms. Growth, respiration and mortality are all assumed to scale as a power-law with M_i and as an exponential function of ambient temperature (see Section 2). Numbers denote median mass (units: g wet weight) for each size class and example organisms of smallest and largest size classes are given in blue.



1008
1009
1010
1011
1012

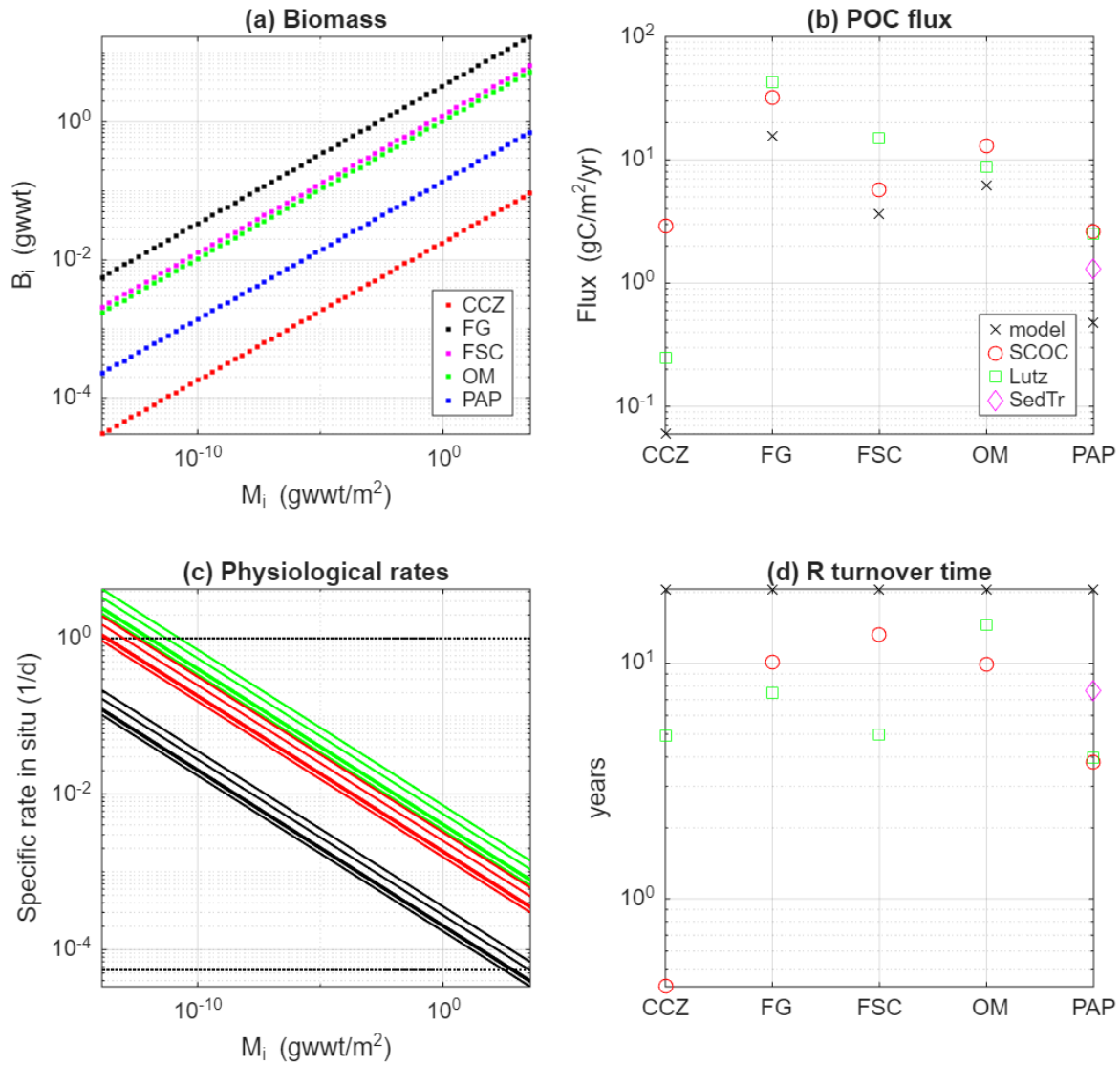
Figure 2: Plots showing how the growth limitation function, $f(R, B_i) = 1 / (1 + a_i B_i / R)$ varies as either R (left) or B_i (right) is varied. R^* and B_i^* denote values at steady state (vertical dotted line), where additionally $\delta R^* = a_i B_i$. Note that because of the scaling behaviour of a_i , $a_i B_i$ is the same for all i at any site.



1015

1016 **Figure 3** Locations (top) and observations of biomass as a function of body size for the five sites listed in Table 2. For each of the
 1017 sites the dots show observations, circles indicate means within size classes and the black line is a power-law fit to the observations.
 1018 The fitted values for the scaling exponent together with 95% confidence intervals are also shown. The bottom right panel shows the
 1019 fit to all sites simultaneously, assuming a common scaling exponent (fitted value shown with 95% CI) but allowing the pre-factor to
 1020 vary across sites. The dots show data from the sites with colours matching those used in the panels for individual sites. Note that in
 1021 the bottom right panel the data from each site has been normalised by dividing by the fitted pre-factor to allow the visual comparison.
 1022 The red line in all panels is the simultaneous fit to all sites. The green line in all plots is the relationship used in the model by fitting
 1023 a power-law with an imposed scaling of 0.2 to the same observations.

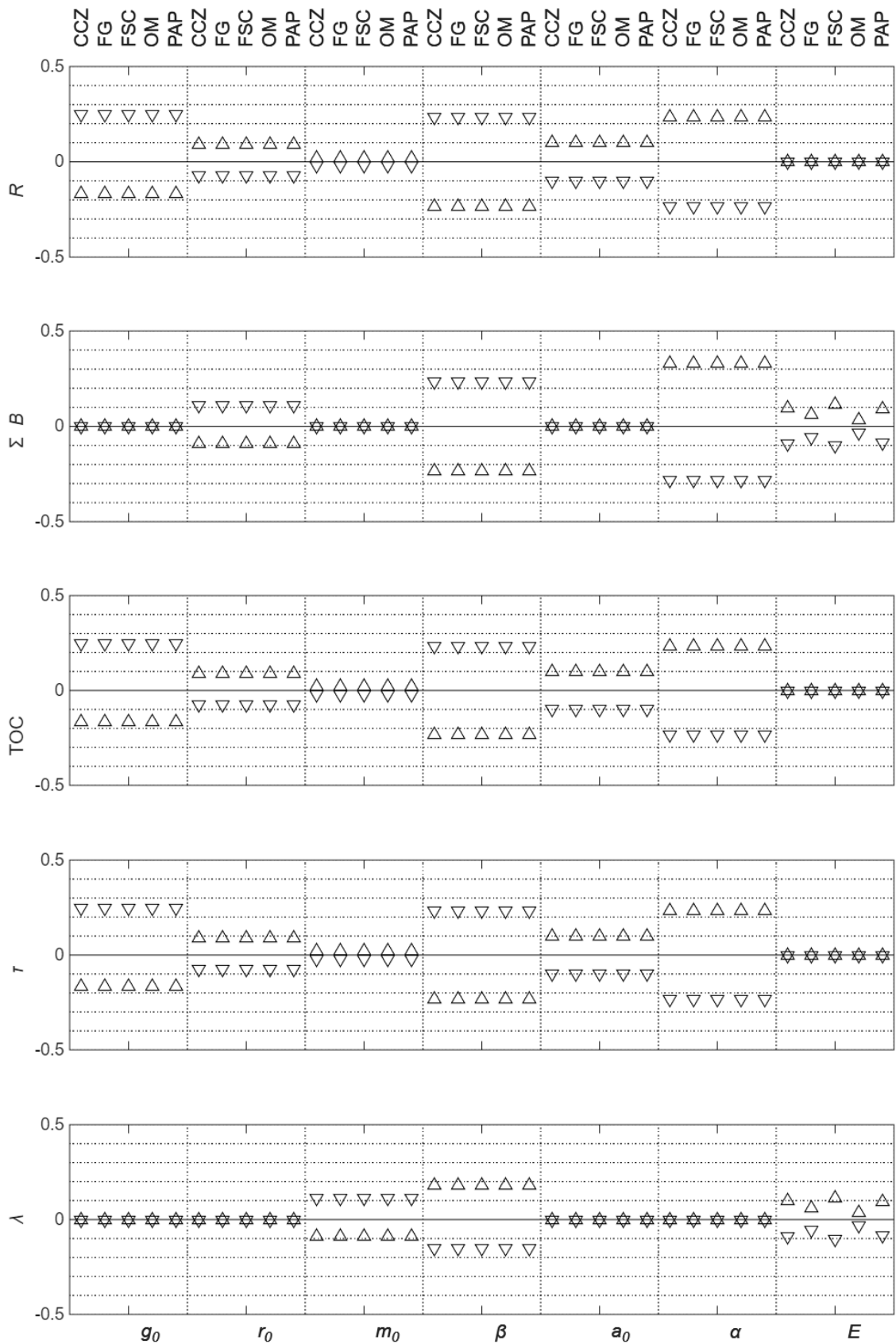
1024



(e) Table of diagnostics

Site	Total	R	TOC		Biggest organism lifetime
	biomass	gC/m2	Model	Obs	
	gC/m2	gC/m2	gC/m2	gC/m2	yr
CCZ	0.062	1.2	24	400	72
FG	12	320	6400	800	51
FSC	4.3	75	1500	550	82
OM	3.6	130	2600	2000	40
PAP	0.48	9.9	198	320	68

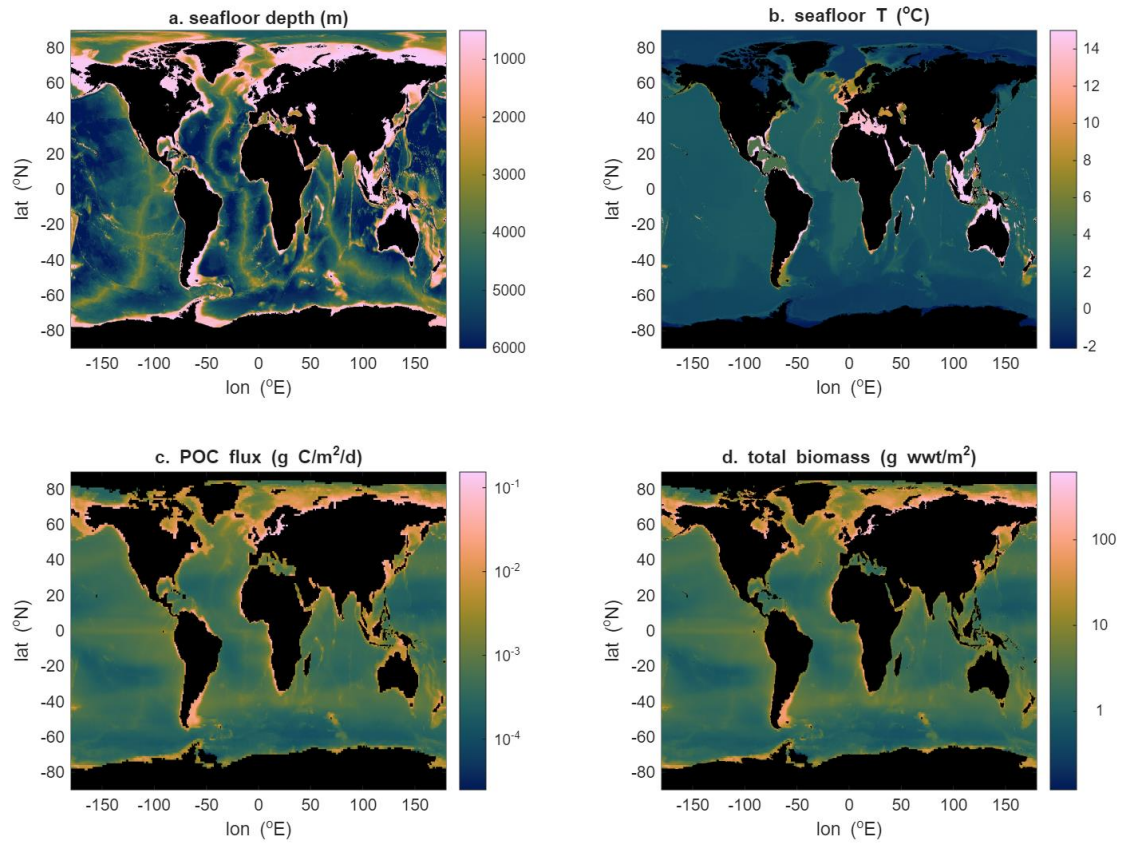
1025
1026 **Figure 4: Model diagnostics and observational constraints: (a) modelled biomass (gwwt/m²) size distributions for the five sites using;**
1027 **(b) POC flux estimates using the Lutz et al. (2007) algorithm (Lutz), SCOC data (SCOC), sediment trap observations (SedTr) and**
1028 **the model constrained by biomass data from each site (model). Note that sediment trap data are only available for PAP; (c) model**
1029 **specific rates for maximum net growth (green), respiration (red) and mortality (black) for each of the sites. Note that they differ**
1030 **between sites because of the temperature effect. The two dotted black lines correspond to the constraints of a maximum growth rate**
1031 **of 1 d⁻¹ for the smallest organisms and a lifetime of 50 yr for the largest organisms.; (d) turnover time (R/POC flux) for each of the**
1032 **sites, estimated using model estimates of R and all observation and model estimates for the POC flux from (b); (e) a summary table**
1033 **of diagnostic parameters including total biomass (model), R (model), TOC (model and observations) and biggest organisms' lifetime**
1034 **(model)**
1035



1036
 1037 **Figure 5** Sensitivity of the metrics detritus (R), total biomass (ΣB), TOC, turnover time of detritus (T) and lifetime of largest organism
 1038 (λ) to each of the model parameters across the five sites. In each case the sensitivity shows the fractional change resulting from a +/-
 1039 10% change in the parameter from the standard value. Symbols point in the direction in which the parameter was changed i.e.
 1040 upward-pointing pyramids indicate an increase in parameter value and inverted pyramids indicate a decrease.

1041

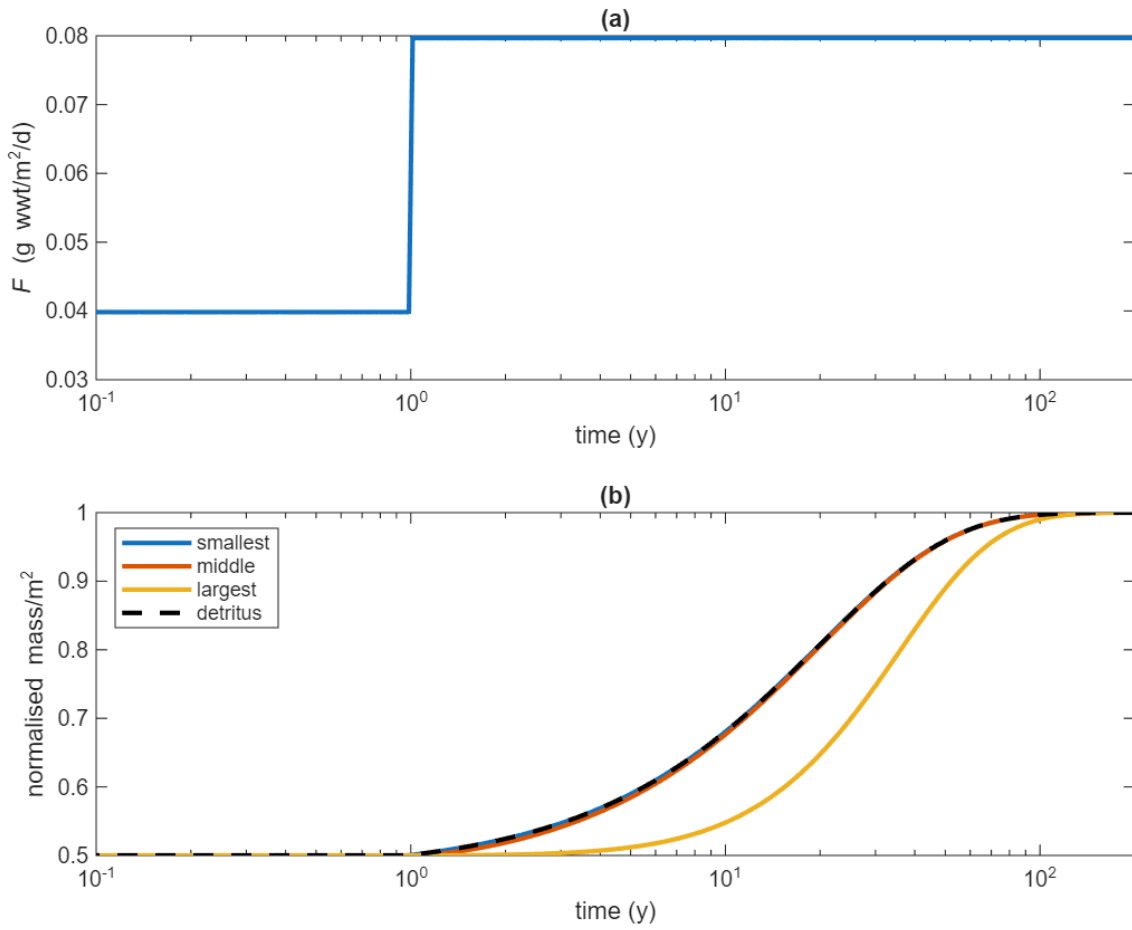
1042
1043
1044



1045

1046 **Figure 6** Example of using the steady state solution (Equations 10 and 11) to explore spatial variability in total benthic biomass:
1047 seafloor depth (a) and temperature (b) from the World Ocean Atlas, particulate organic carbon (POC) flux (estimated using Lutz
1048 et al., 2007) (c), and total biomass (d).

1049



1050

1051 **Figure 7: Modelled response to a perturbation in which the ecosystem is initially in steady state and then a doubling of the POC flux**
 1052 **takes place: (a) shows the POC flux, with doubling occurring after 1 year; (b) shows the response of organisms and detritus. For**
 1053 **clarity only 3 size classes are shown, the smallest, middle and largest ($i=1, 30$ and 59). For the same reason, the biomasses and**
 1054 **detritus are normalised by dividing by their final value. A log time scale is also used to highlight the different response timescales.**

1055

Parameter	Description	Value	Units
g_0	Max. net growth rate at 20°C for organism of size 1 gwwt	0.01	d ⁻¹
r_0	Respiration rate at 20°C for organism of size 1 gwwt	0.0045	d ⁻¹
m_0	Mortality rate at 20°C for organism of size 1 gwwt	0.0005	d ⁻¹
a_0	Interference pre-factor	2000	-
β	Scaling exponent for growth, respiration and mortality	-0.2	-
α	Scaling exponent for interference	-0.2	-
E	Activation energy	0.35	eV

1056

1057

1058 **Table 1: parameter set for BORIS-2**

1059

1060

1061

Site	Lat °N	Lon °E	Depth m	Temp °C	POC flux g C m ⁻² y ⁻¹				TOC g C m ⁻²		R g C m ⁻²
					Lutz	SCOC	Trap	Model	Obs	Model	Model
					CCZ	17.2	-122.6	4150	1.5	0.25	2.9
FG	58.3	0.9	153	8	43	32	-	16	800	6500	320
FSC	61.9	-2.8	1623	-1	15	5.7	-	3.6	550	1500	75
OM	23.4	59	507	13	8.8	13	-	6.2	2000	2600	130
PAP	48.8	-16.5	4850	2.6	2.5	2.6	1.3	0.48	320	200	9.9

1062

1063 **Table 2: Information on sites from which data were used to constrain parameter values for the general purpose**
1064 **parameter set given in Table 1, together with model diagnostics. POC flux is at the seafloor, estimated using the**
1065 **algorithm in Lutz et al. (2007), SCOC or sediment trap. TOC estimates come from Parameswaran (2024). The “Model”**
1066 **columns indicate model diagnostic values. Model TOC is calculated by assuming that R is 5% of TOC. Sources for data**
1067 **are given in Section 2.5**

1068

1069

1070