

# 1 **BORIS-2 – a benthic ecosystem model based on allometry**

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6 **Abstract.** We present a model describing the population dynamics of benthic biota, feeding from a common resource that is  
7 supplied by a flux of sinking organic carbon arriving on the seafloor. By using allometric relationships for the physiological  
8 processes of growth, mortality and respiration, and for food limitation, the model represents the population dynamics of  
9 organisms ranging in size from bacteria ( $10^{-14}$  g wet weight C) to large metazoans ( $10^3$  gwwt C). The effect of temperature on  
10 physiological rates is also included. The only forcing information required is the ambient temperature and the rate of supply  
11 of sinking organic carbon. The model can be used for, and tuned to, specific locations. However, a parameter set is provided  
12 that is generally applicable. The ability of the model to simultaneously reproduce biomass size distributions at five contrasting  
13 sites is demonstrated for this parameter set. Other examples of use are also shown, using the model to explore global patterns  
14 of benthic biomass, and responding to a change in food supply.

## 15 **1 Introduction**

16 The surface ocean, or epipelagic, ecosystem has received considerable attention from modellers for a variety of reasons,  
17 spanning from the magnitude of biogeochemical fluxes (e.g. Burd, 2024) and fundamental questions of ecosystem structure  
18 (e.g. Woodson et al., 2018) and biodiversity (e.g. O’Dor et al., 2009) to more societal issues such as fisheries management  
19 (e.g. Karp et al., 2023) and climate modelling (Kwiatkowski et al., 2020). However, the seafloor, or benthic, ecosystem has  
20 received much less attention, particularly in the deeper regions away from the continental shelves. This is despite the regions  
21 deeper than 1000 m constituting over half of the earth’s surface area (Ramirez-Llodra et al., 2010; Harris et al., 2014).

22  
23 The benthic ecosystem of the deep ocean (aside from hydrothermal vents) is almost entirely dependent on external input for  
24 food, with the majority in the form of organic material sinking down from the waters above. This means that the benthic  
25 ecosystem is susceptible to changes in production of this organic material that may occur several kilometres above it (Ruhl et  
26 al., 2008), such as in response to climate change (Yool et al., 2017). Benthic ecosystems are also subject to direct pressures  
27 such as trawling, dredging, oil and gas activities, and seabed mining. To understand and to predict the future for benthic  
28 ecosystems we therefore need models that adequately capture their response to such drivers, across the full ecosystem and over  
29 appropriate timescales.

30  
31 Building models that capture the key interactions within an ecosystem is of value for three reasons: construction of a model  
32 forces us to identify the key processes and to articulate our understanding of them in a precise manner; the behaviour of the  
33 model allows us to identify gaps and uncertainties in that knowledge; and by linking the model to forecasts for how  
34 environmental drivers may change it allows us to make predictions for the fate of the ecosystem across different scenarios.  
35 One modelling strategy is to represent an ecosystem as different functional groups, particularly those linked to particular fluxes  
36 of interest into and out of the sediment e.g. deposit feeders and aerobic/anaerobic bacteria (e.g. Butenschön et al., 2016; Ersten  
37 et al., 2018). This approach is valuable, for example, in studying the biogeochemistry of shelf systems where the interactions  
38 between sediment, overlying water and benthic ecosystem may need to be captured because the feedback on the overlying  
39 water column may be significant given the shallow depths. Shelf systems also benefit from a greater array of data to constrain

40 a model as they are more accessible for sampling than deeper waters. More generally, a paucity of data to constrain a model  
41 or limited understanding of causal relationships are common hindrances, particularly for deep-sea ecosystems because of the  
42 remote and challenging nature of the environment being studied. For situations where the ecosystem can be approximated as  
43 unchanging in time, statistical methods have been used (e.g. Reiss et al., 2014), particularly for modelling distributions of  
44 groups or individual species, but also for distributions of biomass (Wei et al., 2010; Jones et al., 2014). For deep-sea ecosystems  
45 where data are sparser, an inverse approach has been used to estimate fluxes between functional and size category components  
46 of the ecosystem at equilibrium (Soetaert & van Oevelen, 2009 ; Durden et al. 2017; de Jonge et al., 2020), with the size classes  
47 mirroring those represented by typical benthic sampling techniques. However, behaviour such as switch-feeding (e.g.  
48 alternating between suspension feeding and predation) in deep-sea fauna (Durden et al. 2015; Iken et al. 2001) complicates the  
49 use of discrete functional groups based on feeding types (Durden et al. 2017).

50

51 Another approach, is to represent the community purely as a collection of different size classes of organisms (Kelly-Gerreyn  
52 et al., 2014; Blanchard et al., 2011; Laguionie Marchais et al., 2020) rather than as functional groups or species. As described  
53 below, this offers considerable simplification in model structure and parameterisation. Furthermore, by using allometric  
54 relations to base the model on the representation of rates, rather than stocks, this approach also allows the response of  
55 ecosystems with time to be tracked.

56

57 Considerable attention has been given to observations showing relationships which appear to scale in a consistent way with  
58 body size, both at a population (e.g. abundance - White et al 2007) and an individual (e.g. physiological rates - Gillooly et al.  
59 2001) level. This phenomenon has been widely observed, on land (e.g. Nagy 1987), in the air (e.g. Niven and Scharlemann,  
60 2005) and in the sea (e.g. Molony and Field 1989) including the deep ocean (Durden et al. 2019; McClain et al. 2012; Mahaut  
61 et al. 1995). That such behaviour has been observed across many habitats and orders of magnitude in size of organism  
62 unsurprisingly led to a search for a “Universal” law explaining such behaviour. Metabolic rate controls ecological processes  
63 at individual and ecosystem levels by determining resource uptake and allocation. The Metabolic Theory of Ecology (MTE;  
64 West et al., 1997; Brown et al., 2004) asserts that, to first order, this rate is controlled by the size of organism and the ambient  
65 temperature. This provides a potential explanation for the existence of a power-law relationship between physiological rates  
66 and body size. However, there remains a discussion over the taxonomic or functional scale at which other features or processes  
67 might disrupt any universal scaling (Seibel and Drazen, 2007), the precise value of the scaling (Isaac and Carbone 2010, Brey  
68 2010; Glazier, 2022) and the extent to which such an approach applies to systems that are not in equilibrium (McCarthy et al.,  
69 2019).

70

71 Notwithstanding these caveats, an allometric approach still has considerable value when applied at broad ecosystem scales.  
72 To support use of an allometric approach, we give just a few examples for three key processes: growth, respiration and  
73 mortality. Motivated by predictions of MTE, Ernest et al. (2003) successfully tested the predicted scaling exponent of -0.25  
74 for growth rate, for organisms spanning  $10^{-14}$ - $10^8$ g in size. More specific to this study, with a focus on macrobenthos, Cusson  
75 & Bourget (2005) brought together empirical relations from previous studies (their Table 8) that demonstrate similar evidence  
76 of scaling of growth rate with size. For respiration, Mahaut et al. (1995) found a power-law scaling with size for deep-sea  
77 organisms, spanning seven orders of magnitude. This was for a single location, however, so not suitable for testing a  
78 temperature dependence. For mortality, McCoy and Gillooly (2008, 2009) brought together estimates of natural mortality  
79 spanning 22 orders of magnitude including plants, fish, birds, mammals and invertebrates, finding that a power-law scaling  
80 with size plus an exponential dependence on temperature captures the dominant pattern. A restriction of their data purely to  
81 invertebrates found a similar result (McCoy and Gillooly, 2009), still spanning 11 orders of magnitude in size. Again, focussed  
82 on marine benthic organisms, McClain et al., (2012) analysed data for growth, respiration and turnover (which can be a proxy

83 for mortality) and demonstrated a clear power-law scaling with size, with additional support for an exponential relationship  
84 with temperature. There are, therefore, reasonable grounds for adopting an allometric approach.

85

86 Kelly-Gerreyn et al. (2014) constructed a dynamic model for benthic organisms based on allometry such that physiological  
87 rates vary with body size (Benthic Organisms Resolved in Size – hereafter “BORIS-1”). BORIS-1 was capable of reproducing  
88 the size distribution of organisms at three sites contrasting in depth between 150m and 1600m. This model assumed that all  
89 organisms were detritivores, eating from a common pool of detritus supplied by organic material sinking to the seafloor, the  
90 particulate organic carbon (POC) flux. BORIS-1 demonstrated that an allometric model with a small number of physiological  
91 processes (ingestion, assimilation and respiration/mortality) that are common to all organisms but scale with body size can  
92 capture the size-distribution of biomass seen in observations. However, the model has several limitations. The first is that it  
93 only represents a limited range of sizes ( $8.9 \times 10^{-7}$  to  $2.9 \times 10^{-2}$  g wet weight). It was therefore necessary to assume a specific  
94 fraction of POC flux that was consumed and respired by organisms not represented by the model, and hence not available to  
95 the modelled organisms. The omitted organisms included both smallest (e.g. bacteria) and largest (e.g. large sea cucumbers)  
96 size ranges. The physiological rates were also not dependent on temperature even though there is evidence that physiological  
97 rates typically increase as the environment warms (e.g. Gillooly et al., 2001). Additionally, the mortality rate had a dependency  
98 on the POC flux in BORIS-1. This resulted in estimates of longevity that unrealistically varied across several orders of  
99 magnitude for the same organism at different locations.

100

101 This paper presents an expanded and updated version of BORIS-1 that addresses these limitations. The resulting model,  
102 BORIS-2, spans the full range of organism sizes, includes the physiological role of temperature, and is parameterised using a  
103 larger dataset that includes observations from a greater range of sites with contrasting environments conditions, including the  
104 abyssal ocean. A single parameter set that allows the model to capture the ecosystem structure across these sites is given and  
105 examples are demonstrated for how the model may be used to study both local and global questions. It is worth stressing that  
106 the aim of BORIS-2 is to capture broad ecosystem behaviour, i.e. macroecology, across the full range of body sizes, not to  
107 capture the dynamics of specific species.

## 108 **2 BORIS-2 model description**

109 BORIS-2 represents benthic organisms spanning in size from bacteria to large metazoans (Figure 1). It does so by dividing  
110 benthic organisms into size classes and using an allometric approach. The size classes are defined on the basis of individual  
111 wet weight body mass (units grams wet weight – gwwt). From the smallest to the largest, each size class spans twice the range  
112 of the former. More specifically, the mean body mass of organisms within a size class spans from  $1.3 \times 10^{-14}$  gwwt (13 fg) for  
113 the smallest to  $3.6 \times 10^3$  gwwt (3.6 kg) for the largest. The lower limit of the smallest class is  $0.88 \times 10^{-14}$  gwwt (9 fg) and the  
114 upper limit of the largest is  $5.1 \times 10^3$  gwwt (5.1 kg). The size classes are chosen to be consistent with those used for size-spectra  
115 biomass data (e.g. Laguionie Marchais et al., 2020) and BORIS-1 (Kelly-Gerreyn et al., 2014). The smallest size class is based  
116 on the smallest observed bacteria (Luef et al., 2015), using a conversion from gC to gwwt of 11.5 (Brey, 2010). The largest  
117 size class is chosen to be broadly representative of benthic habitats. It is consistent with a detailed assessment of invertebrates  
118 for a well-studied abyssal site, the Porcupine Abyssal Plain (one of the sites described in Section 2.4). These upper and lower  
119 biomass limits, together with the factor of two scaling used between biomass size classes, sets the number of model size classes  
120 to 59. For application of BORIS-2 to specific locations where larger organisms are known, the upper size limit is easily  
121 changed. The BORIS-2 size range currently spans over 18 orders of magnitude.

## 122 2.1 Ecological interactions

123 BORIS-2 comprises a set of differential equations describing the time-varying behaviour of  $N=59$  size classes ingesting a  
 124 common resource,  $R$ , that represents the stock of detrital food available to the benthic community (i.e., in / on the seafloor or  
 125 in the benthic boundary layer). The model does not capture any direct predation or cannibalism, and instead represents a  
 126 community of detritivorous heterotrophs. This decision is based on the different nature of benthic and pelagic ecosystems.  
 127 Total biomass in a size range increases with body mass in benthic ecosystems (e.g. Benoist, 2020; Kelly-Gerreyn et al., 2014).  
 128 In the pelagic ocean, biomass is roughly equal across different sizes (Hatton et al., 2021). The greater accumulation of biomass  
 129 in larger organisms in the benthic system is consistent with a greater transfer efficiency arising from a reduced role of predation  
 130 and a greater role of feeding from a common detrital resource. We effectively assume that the dominant influence on benthic  
 131 ecosystems is having neighbours competing for your food, rather than eating you. This assumption is discussed in more detail  
 132 in Section 4.1.

133  
 134 The total biomass represented by all organisms per square metre in each size class  $i$  of nominate mass  $M_i$  (units: grams wet  
 135 weight - gwwt) is represented as  $B_i$  (units: gwwt/m<sup>2</sup>) which varies with time,  $t$  (units: d), according to the equation

$$136 \frac{dB_i}{dt} = \overbrace{g_i \cdot f(R, B_i) \cdot B_i}^{\text{gross growth}} - \overbrace{r_i \cdot B_i}^{\text{respiration}} - \overbrace{m_i \cdot B_i}^{\text{mortality}} \quad (1)$$

137 where  $g_i$  is the maximum specific growth rate and  $r_i$  and  $m_i$  are the specific rates of respiration and mortality, respectively. All  
 138 of  $g_i$ ,  $r_i$  and  $m_i$  have units of 1/d. The function  $f(R, B_i)$  represents how growth is limited by increasing population size and/or  
 139 decreasing resource availability (see Section 2.2). The associated equation controlling the amount of resource,  $R$  (units:  
 140 gwwt/m<sup>2</sup>), is

$$141 \frac{dR}{dt} = F - \sum_{i=1}^N \left[ \overbrace{g_i \cdot f(R, B_i) \cdot B_i}^{\text{gross growth}} - \overbrace{m_i \cdot B_i}^{\text{mortality}} \right] \quad (2)$$

142 where  $F$  is the POC flux to the seafloor through gravitational sinking of detritus (gwwt/m<sup>2</sup>/d). Note that at equilibrium the rate  
 143 of supply of organic material,  $F$ , equals the respiration by the whole ecosystem,  $\sum_{i=1}^N r_i \cdot B_i$ . It is assumed that the long-term  
 144 burial of organic material in sediment is negligible compared to POC and total respiration fluxes. (Section 4.1 discusses how  
 145 this assumption might be relaxed.) A linear mortality term,  $m_i \cdot B_i$ , is used in BORIS-2. (The reason for using this rather than  
 146 the quadratic  $\mu_i \cdot B_i^2$  parameterisation used in BORIS-1 is given in the Appendix).

## 147 2.2 Growth limitation: food scarcity and interference

148 The growth limitation function,  $f(R, B_i)$ , reflects the impact on growth arising from competition for limited resources. This  
 149 function is chosen to capture two effects. First, low availability of food,  $R$ , should lead to a reduced rate of intake and growth.  
 150 Second, any increase in the number of organisms (for which  $B_i$  is a proxy) looking for food should reduce the likelihood of  
 151 any individual finding it, a phenomenon known as interference (e.g. DeAngelis et al., 1975). More specifically, we assume the  
 152 parameterisation

$$153 f(R, B_i) = 1/(1+a_i B_i/R) \quad (3)$$

154 The parameter  $a_i$  is present to account for how interference may scale with size. For example, larger organisms can search a  
 155 given area more quickly than a smaller one, in general, either because motility generally increases with size or more simply  
 156 because they occupy a greater area. Note that  $a_i$  is unitless. The function  $f(R, B_i)$  varies between 0 and 1, with a value of zero  
 157 entirely ceasing growth and a value of one leading to growth at the maximum rate,  $g_i$ . To demonstrate that the function has the  
 158 required properties, first consider the case where food is very abundant such that  $R \gg a_i B_i$ . Then  $f(R, B_i) \sim 1$  and there is no  
 159 limitation of growth. If resource is scarce such that  $R < a_i B_i$ , then  $f(R, B_i) \sim R/(a_i B_i)$  which is always a value less than one but  
 160 increases and decreases linearly with both resource,  $R$ , and abundance of organisms, represented by  $B_i$ . Note that for simplicity

161 we currently only incorporate competition for resources within a size class. This is the simplest assumption given that different  
 162 size classes seek food at different spatial scales. What is a meal for a bacterium is unlikely to be a meal for a holothurian. It  
 163 would, however, be straightforward to include competition from other size classes simply by using a sum over those classes  
 164 in the denominator. The form of the interference parameter,  $a_i$ , is discussed in the next section (2.3). Figure 2 and Section 2.4  
 165 show further how the function  $f(R, B_i)$  varies, in particular demonstrating that it affects all size classes equally i.e. that it does  
 166 not lead to some size classes being food limited while others are growing near maximum rates.

### 167 2.3 Allometric and temperature influences

168 In BORIS-2 allometry is used to describe four physiological or physiologically affected processes across the range of body  
 169 sizes. The physiological processes are growth ( $g_i$ ), respiration ( $m_i$ ) and mortality ( $r_i$ ). The physiologically affected process is  
 170 growth limitation, controlled by parameter  $a_i$ . All of these are assumed to be determined by size (body mass) and environmental  
 171 temperature.

172

173 The effect of temperature is assumed to be identical for all four processes and represented by a function,  $\theta(T)$ , which is taken  
 174 to be

$$175 \theta(T) = \exp[-E \cdot \tau/k] \quad (4)$$

176 with

$$177 \tau = 1/(T+T_{abs}) - 1/(T_{ref}+T_{abs}) \quad (5)$$

178 where  $T$  is temperature (units: °C),  $T_{ref}$  (units: °C) is a reference temperature,  $T_{abs} = 273.15$  K converts  $T$  and  $T_{ref}$  to units of  
 179 Kelvin (K), and  $k$  is Boltzmann's constant ( $8.62 \times 10^{-5}$  eV/K). This is a widely used formulation applied both in empirical studies  
 180 (e.g. Brey, 2010) and papers developing ideas around the Metabolic Theory of Ecology (e.g. Gillooly et al. 2001).  $E$  (units:  
 181 eV) is often described as an activation energy. We discuss the value chosen for  $E$  in Section 2.4.2.  $T_{ref}$  is chosen to be 20°C.  
 182 While this may seem an arbitrary choice of reference temperature, it has no impact on rates. Using a different  $T_{ref}$  simply  
 183 requires a numerical change in parameters ( $g_0$ ,  $r_0$ ,  $m_0$  and  $a_0$ ) to compensate for the change.

184 It is assumed that the three physiological rates ( $g_i$ ,  $m_i$  and  $r_i$ ) scale with body size in an identical way. This is purely taking the  
 185 option requiring fewest assumptions given the current uncertainty in how these rates vary with size. As the link between  
 186 interference and physiology is more tentative,  $a_i$  is theoretically allowed to scale independently (but see Section 2.5). More  
 187 specifically, growth, respiration and mortality have common scaling exponent  $\beta$ , whereas interference scales with exponent  $\alpha$ :

$$188 g_i = \theta(T) \cdot g_0 \cdot M^\beta \quad (6)$$

$$189 r_i = \theta(T) \cdot r_0 \cdot M^\beta \quad (7)$$

$$190 m_i = \theta(T) \cdot m_0 \cdot M^\beta \quad (8)$$

$$191 a_i = \theta(T) \cdot a_0 \cdot M^\alpha \quad (9)$$

192 The values chosen for the seven parameters used in the model ( $g_0$ ,  $r_0$ ,  $m_0$ ,  $a_0$ ,  $\alpha$ ,  $\beta$ ,  $E$ ) are given in Section 2.5 (and Table 1),  
 193 together with a description of the data used to constrain them. The performance of the model using this parameter set is then  
 194 described in Section 2.6 and the uncertainties associated with their values are discussed in Section 2.7. Before then a steady  
 195 state solution for the model is presented, both for its own use and as a source of useful information for constraining parameter  
 196 values.

197

198 **2.4 Steady state solution**

199 The model has a steady state solution which can be written in a simple form. This provides a means to initialise simulations,  
 200 to validate dynamical model runs (if run to equilibrium) or to accelerate model runs where time-scales are longer than organism  
 201 response times.

202 Indicating steady state values with an asterisk, the steady state solution is

$$203 \quad R^* = \frac{F}{\sum_{i=1}^N \frac{r_i(g_i-r_i-m_i)}{a_i(r_i+m_i)}} \quad (10)$$

$$204 \quad B_i^* = \frac{(g_i-r_i-m_i)}{a_i(r_i+m_i)} \cdot R^* \quad (11)$$

205 This steady state solution provides a few insights into the behaviour of the model. First, both resource,  $R$ , and biomass in all  
 206 size classes,  $B_i$ , increase linearly with  $F$ . This is not surprising as we would expect abundance of detritus and biomass to  
 207 increase with increasing food supply. Second,  $B_i$  scales with size with the same exponent,  $-\alpha$ , as  $1/a_i$  (i.e.  $B_i^* \propto M^{-\alpha}$ ). This is  
 208 because  $g_i$ ,  $r_i$  and  $m_i$  all scale the same with size, as mentioned above, and so the scaling of  $(g_i-r_i-m_i)$  in the numerator for  $B_i^*$   
 209 is cancelled by the identical scaling of  $(r_i+m_i)$  in the denominator. Hence, the biomass spectral slope is effectively set by  
 210 interference. Although this might be unexpected it should be noted that the processes contributing to  $a_i$  are still very poorly  
 211 known and its scaling is likely to be influenced by physiological processes, such as respiration associated with enhanced  
 212 movement for example. The theoretical model of Damuth (2007) is potentially relevant here as it links competition for  
 213 resources to allometric scaling and community wide energy use. Nevertheless, understanding the likely influences on  
 214 interference is clearly a useful avenue for future research. A consequence of the inverse scaling of  $a_i$  and  $B_i^*$  is that  $a_i B_i^*$  is the  
 215 same for all size classes i.e. the growth limitation function  $f$  does not change with size at steady state.

216

217 Returning to the steady state solution, substituting Equations 6-9 into Equations 10 and 11 gives

$$218 \quad B_i^* = \left( \frac{1}{r_0 \theta(T)} \right) \cdot \left( \frac{M_i^{-\alpha}}{\sum_{j=1}^N M_j^{\beta-\alpha}} \right) \cdot F \quad (12)$$

$$219 \quad R^* = \left( \frac{\alpha_0 \cdot (r_0 + m_0)}{r_0 \cdot (g_0 - r_0 - m_0)} \right) \cdot \left( \frac{1}{\sum_{j=1}^N M_j^{\beta-\alpha}} \right) \cdot F \quad (13)$$

220 In addition to showing explicitly that  $B_i$  scales as  $-\alpha$ , as already mentioned, Equation 12 also reveals that the steady state  
 221 biomass is independent of growth and mortality except for the scaling ( $\beta$ ) and temperature dependence ( $\theta(T)$ ) that they share  
 222 with respiration. While this might seem at first surprising, it is because of the fundamental constraint that total respiration must  
 223 match the POC flux,  $F$ , of arriving new organic material, i.e.  $F = \sum_{i=1}^N r_i \cdot B_i^*$ . At equilibrium, any change in growth or  
 224 mortality arising from changing either  $g_0$  or  $m_0$ , respectively, is compensated by a change in food resource ( $R$ ), rather than in  
 225  $B_i$ , to maintain this balance. This balance is also reflected in the influence of  $r_0$  in Equation 12, with an increase in it  
 226 corresponding to a compensating decrease in  $B_i^*$ . Similarly, if the specific respiration rate increases as a result of temperature  
 227 increase (see Equation 7) then the higher physiological overhead means that a lower  $B_i^*$  is maintained.

228

229 The steady state solution is also of use in understanding the influence of interference in the growth limitation function. Equation  
 230 11 implies that at steady state  $a_i B_i^* = \delta \cdot R^*$  where

$$231 \quad \delta = \frac{(g_i-r_i-m_i)}{(r_i+m_i)}$$

232 Because the physiological rates scale identically, and have the same temperature dependence,

$$233 \quad \delta = \frac{(g_0 - r_0 - m_0)}{(r_0 + m_0)}$$

234 which is the same for all size classes and sites. (For the parameter set used here and described below,  $\delta=1$ .) Figure 2 shows  
235 how the interference function varies with both  $R$  and  $B_i$ , as they vary either side of their steady state values.

236

## 237 **2.5 Observational constraints and choice of parameter values**

238 There are a range of observations that can be used to constrain parameter values but, as described below, none can be used to  
239 set a parameter value in isolation. Relationships between parameters must be used to link the different observations together  
240 as a collective constraint. It is also worth noting that there is no objective way to use these multiple constraints. As will be  
241 seen below, the strength of some constraints is greater than others and trying to construct some overall cost function to optimise  
242 all parameters simultaneously would require considerable subjectivity in how the constraints were translated into costs and  
243 weighted relative to each other. For this reason, and because of the limited number of parameters and ability to calculate the  
244 outcome of a given parameter set extremely quickly, values have instead been chosen by trial and error for the seven parameters  
245  $g_0, r_0, m_0, a_0, \alpha, \beta, E$  to give an acceptable, if potentially not optimal, fit to observations. A summary of values can be found  
246 in Table 1. While future users of BORIS-2 may choose to use a different approach to selecting parameter values, it will be  
247 seen in Section 2.6 that the current set does a reasonable job and Section 2.7 describes the consequences associated with  
248 adjusting these values.

249

250 Previous studies have highlighted observational evidence for allometric relationships, either in just one physiological rate (e.g.  
251 Mahaut et al., 1995; Cusson & Bourget, 2005), or, in several (e.g. McClain et al., 2012). Because of the variations in scaling  
252 reported in these studies (discussed in Section 2.7) we have used the observation of allometry as a starting point rather than  
253 take a value for the scaling exponent from a specific study.

254

255 The following first describes the observational constraints. Based on these, the argument for the specific choices of parameter  
256 values is then given. Table 1 has a summary of parameter values and Figures 3 and 4, and Table 2, summarise the observational  
257 constraints and associated model diagnostics used to select them. The code needed to generate Figures 3 and 4 is also available  
258 to allow the model to be re-tuned given additional data, different locations or different priorities (Martin et al., 2025).

259

260 Starting with the largest dataset, biomass can be used as a constraint for the interference parameter  $\alpha$ . The steady state solution  
261 (section 2.4) shows that  $a_i$  must scale in the opposite way to biomass distributions. Suitable observations from five sites are  
262 available in selecting the value; a summary is found in Table 2. The Clarion Clipperton Zone (CCZ) is a vast abyssal plain in  
263 the northeast Pacific. The data used here come from a site (17.2° N 122.6° W) of depth 4150 m with a low temperature (1.5  
264 °C). Fladden Ground (FG) is in a shelf sea (153 m) and, unsurprisingly, with higher temperature (8 °C). The Faroe-Shetland  
265 Channel (FSC) is a connection between the North Atlantic and the Arctic, with the lowest temperature (-1 °C) despite a depth  
266 of only 1623 m. The Oman Margin (OM) is a slope site (507 m) and has the highest water temperature (13 °C). The final site  
267 is the Porcupine Abyssal Plain (PAP), which is the deepest (4850 m), with reasonably cold temperature (2.6 °C). A general  
268 decrease of temperature with depth is overlain with considerable variability due to local hydrography (notably FSC and OM).  
269 In addition to spanning a range of contrasting temperatures and depths, the data from the five sites also covers complementary  
270 size ranges of organisms. CCZ data are based on photographically surveyed megabenthos. FG, FSC, and OM data are based  
271 on physically sampled meio- to macrobenthos. PAP is based on physically sampled macrobenthos and photographically  
272 surveyed megabenthos. For CCZ, the data can be found in Benoist (2020), with sampling and methodology described in Simon-  
273 Lledó et al. (2019) and Benoist et al., (2019). Details on data for FG, FSC and OM can be found in Kelly-Gerrey et al. (2014).  
274 Additional information on the sampling and laboratory methodology can be found in Kaariainen et al. (2006). The benthic  
275 ecosystem of the PAP site has been studied for decades (Hartman et al., 2021). The data used here, and presented in Benoist  
276 (2020), combines analyses of macrobenthos and megabenthos. Descriptions of observational approach and the analysis

277 methodology for the megabenthos can be found in Morris et al. (2016) and Durden et al. (2020b). For macrobenthos, this  
278 information can be found in Benoist (2020) and Ruhl et al. (2023). Further details on the treatment of size-resolved data, e.g.  
279 to remove biases such as under-sampled size groups, can be found in Edwards et al., (2017, 2020) and Ruhl et al. (2023).  
280 Observations of biomass versus size for each of the sites are shown in Figure 3. These are referred to as biomass spectra and  
281 the gradient of the relationship (when plotted log-log as here) as the spectral slope, or scaling exponent. Despite some  
282 variability, all sites exhibit an increase of biomass with body size, and in a manner that is consistent with a power-law  
283 relationship. Figure 3 also shows the exponents found by fitting a power-law to the observations from each site individually.  
284 There is no strong relationship between fitted exponent and environmental parameters, though the smaller magnitude scaling  
285 exponents for the two shallow sites is something that has previously been seen in physiological rates rather than biomass  
286 (Mahaut et al., 1995). The OM site additionally has a low oxygen concentration (Demopoulos et al., 2003) which has been  
287 suggested to have a disproportionate impact on larger organisms (Quiroga et al., 2005) and which could therefore be  
288 responsible for flattening the slope relative to other sites. Nevertheless, here we take the simplest assumption that all sites are  
289 showing sufficiently similar behaviour in scaling to assume a common scaling exponent across the sites, leaving an  
290 investigation of departures from this for other studies. Figure 3g shows the simultaneous fit to data from all sites. It is seen that  
291 observations from the five sites cover different size ranges such that the composite dataset spans a substantially wider size  
292 range than any individual site. Furthermore, the relationship of biomass with size appears consistent across the wider range. If  
293 the data from all sites is combined then the scaling exponent for a power-law fit to all five sites simultaneously indicates a  
294 scaling exponent of 0.26 (s.d. 0.016,  $r^2=0.76$ ,  $p<0.001$ ). The existence of similar scaling behaviour across multiple sites gives  
295 more confidence that the model can be used globally.

296  
297 The next observational constraint comes from the requirement that the supply of organic carbon to the ecosystem (the POC  
298 flux,  $F$ ) balances the total respiration of the organisms present at steady state. For the model, respiration is given by the sum  
299 over size classes of the product of the specific respiration rate and biomass in each size class,  $\sum r_i B_i$ . Therefore, observations of  
300 POC flux, co-located with the previously described size-resolved estimates of biomass, provide a useful joint constraint on  $r_0$   
301 and  $\beta$ . There are several sources of data for POC flux. First, and most directly, for the PAP site there is a long time-series of  
302 sediment trap data (Lampitt & Pebody, 2023). Although there are sediment traps at both 3000 m and 4750 m, the latter is  
303 thought to be biased by sediment resuspension as it is just 100m above the seafloor. The magnitude of this effect can vary with  
304 time but previous work has shown that the flux near seabed at the PAP site is often in excess of that at 3000 m due to re-  
305 suspension (Lampitt et al., 2000, 2001). For this reason, it is better to extrapolate the estimate from the 3000 m trap. For the  
306 year 2012 (to best match the biomass observations) the annual carbon flux at 3000 m is 1.91 gC/m<sup>2</sup>/yr. The associated flux at  
307 the seafloor can be roughly estimated using a widely used power-law scaling (Martin et al., 1987), such that the flux at the  
308 seafloor at 4850 m equals  $1.91 \text{ gC/m}^2/\text{yr} * (3000 \text{ m}/4850 \text{ m})^{0.858} = 1.3 \text{ gC/m}^2/\text{yr}$ . An alternative way to estimate the sinking  
309 flux at the seafloor is to use the Lutz et al. (2007) algorithm, which uses net primary production, sea surface temperature and  
310 depth at a given location to estimate the flux. This allows estimates to be made for all five sites (not just PAP) – see Table 2.  
311 Finally, Sediment Community Oxygen Consumption (SCOC) data (Stratmann et al., 2019) also allows the sinking flux to be  
312 estimated, making the same assumption that the respiration (measured by oxygen consumption) must balance this flux, in this  
313 case when averaged over the year. As SCOC is usually measured using chambers of ~50 cm across, the estimates exclude or  
314 bias the contribution from larger organisms – not just those too large to fit but also those too scarce to be robustly sampled in  
315 such an area - and care is needed in accounting for this (Laguionie Marchais et al., 2020). More generally, all of these sources  
316 of POC flux data have significant associated uncertainties, which is why POC flux was not fixed when deriving parameter  
317 values for the general use model configuration described here. Instead the model can be used to estimate the POC flux and  
318 compared to these different observational estimates as a broader constraint. All observational estimates of POC flux used to  
319 constrain model parameters, as well as the model values, are shown in Figure 4b. These values are also given in Table 2. Note

320 that even if there was no uncertainty in the observations for POC flux it would still not be possible to infer specific values for  
321  $\beta$  and  $r_0$  using the data. This is because it is possible to simultaneously vary  $\beta$  and  $r_0$  in a way that the total respiration remains  
322 unchanged. Note that respiration also gives a link between the physiological ( $\beta$ ) and interference ( $\alpha$ ) scaling exponents that  
323 needs to be considered. Respiration by a size range of organisms is given by  $r_i B_i$ , which scales as  $M_i^{\alpha-\beta}$ . Hence, the difference  
324 between  $\alpha$  and  $\beta$  determines whether the energy use by a size range increases ( $\alpha > \beta$ ), decreases ( $\alpha < \beta$ ) or remains with same  
325 ( $\alpha = \beta$ ) with size. We are unaware of observations indicating an increase with size but there are observations suggesting  $\alpha = \beta$   
326 (Laguionie Marchais et al., 2020) and this is consistent with ‘energy equivalence’ which has been suggested theoretically  
327 (Damuth, 2007).

328

329 The inherent relationships between the three physiological rates allow other observational constraints. Observations indicating  
330 that physiological rates decrease with size were already described. Additionally, the maximum net growth rate,  $g_i$ , must equal  
331 or exceed the sum of the respiration,  $r_i$ , and mortality,  $m_i$ , for all sizes for populations to be sustainable. Together, these two  
332 statements imply that all rates must sit within a ‘window’ of parameter space (rate versus size, Figure 4c). We follow a similar  
333 approach to that of Mahaut et al. (1995) and Kelly-Gerreyn et al. (2014), by using constraints at smallest and largest sizes of  
334 organism. Having so many degrees of magnitude in size in the model means that care is needed for the parameter values to be  
335 realistic at the two extremes of the sizes reproduced. Using constraints at intermediate sizes risks significant under- or over-  
336 estimates for largest and smallest organisms through under-constrained extrapolation. Because rates are observed to decrease  
337 with size, the top left corner of the window is set by the upper limit for the specific growth rate of the smallest size class of  
338 organism. For benthic bacteria, we take a rough upper limit of 1 /d maximum specific growth rate. Though Dixon and Turley  
339 (2001) find a rate of 0.1 /d, a wide range reported by Giovannelli et al. (2013) includes a value of 6 /d. An intermediate value  
340 of 1 /d is marked in Figure 4c. The bottom right corner of the window is set by the lower limit for the specific mortality or  
341 respiration rate (whichever is smallest) of the largest size class. The data collated by McClain et al. (2012) indicates lifetimes  
342 for the largest organisms (few kg wwt) of order 50 years. Given this and assuming the same scaling,  $\beta$ , for all physiological  
343 rates as stated earlier, comparison to POC fluxes described above indicates that respiration is a considerably larger rate than  
344 mortality, so mortality defines the bottom right corner of the ‘window’. This is also marked in Figure 4c and lifetimes for the  
345 largest organisms at each site are also given in Figure 4e. Note that both of these constraints need to be treated a little flexibly  
346 as it is not realistic to set a precise limit in either case.

347

348 There are two further constraints, to inform the choice of values for  $\beta$ ,  $g_0$ ,  $r_0$ ,  $m_0$  and  $E$ . First, the ratio  $(r_0 + m_0)/g_0$  represents the  
349 fraction of the maximum growth rate,  $g_0$ , achieved by organisms when the system is in equilibrium. This should take a value  
350 less than 1 for food to be limiting, as is expected for the seafloor (Smith et al., 2008). Second, decreasing  $E$  converges  
351 physiological rates for the different sites, as the inter-site differences due to temperature are diminished. Doing so reduces the  
352 inter-site differences in model POC flux (and  $R$ ) for the same reason, because of the balance with total respiration.

353

354 The final observational constraint is Total Organic Carbon (TOC) in the benthic sediment which can be used to estimate the  
355 amount of detritus available as food,  $R$ , which itself is a constraint on the parameter  $a_0$ . This can be seen in the steady state  
356 solution (Equation 13), where  $R$  increases linearly with  $a_0$ . We make use of the compilation of Parameswaran et al., (2024)  
357 who created an atlas of TOC at the seafloor surface, using a neural network approach applied to globally distributed estimates  
358 calculated over the top 10 cm of sediment. This data source was chosen in preference to the alternative product of Atwood et  
359 al., (2020) as the latter used estimates over the top 1m of sediment, which are an order of magnitude larger than those in  
360 Parameswaran et al., (2024) and likely to represent carbon resources unavailable to the majority of detritivores on the seafloor.  
361 The estimates for TOC using Parameswaran et al., (2024) for the five sites are shown in Figure 4e. These are not directly  
362 comparable with  $R$ , however. There is considerable evidence that not all of sediment TOC is readily available as a food

363 resource, with typically 5% (e.g. De Jonge et al., 2020; Van Oevelen et al., 2011a, 2011b) regarded as ‘labile’ i.e. easily  
364 consumed (see Discussion for more on this). We therefore multiply our model estimates for  $R$  at the five sites by a factor of  
365 20 to give an estimate of TOC for comparison to the observational values. Estimates for TOC based on Parameswaran et al.,  
366 (2024) are given in Figure 4e and Table 2, together with model estimates for  $R$  and TOC. There is considerable variability; in  
367 observations, model values and their relative sizes. A weaker constraint for  $a_0$  is the turnover time for  $R$  i.e. the time it would  
368 take for  $R$  to be replaced by the POC flux (Figure 4d). A minimum turnover time of several years would be expected for the  
369 system to be able to achieve steady state on an annual basis.

370

371 We now outline our choices of parameter values (Table 1) given the above constraints. Starting with  $\alpha$  and  $\beta$ , taking  $\alpha=\beta=$   
372 0.2 gives a scaling exponent consistent with observations for rates and energy equivalence but requires a biomass scaling that  
373 is a little lower than the fit to observations. It is not possible to use the -0.26 scaling implied by the observed biomass spectra  
374 without substantially worsening the fit between model and observational estimates of POC flux (Figure 4b). Given this scaling  
375 for  $\beta$ , we take  $g_0=0.001$  /d,  $r_0=0.0045$  /d,  $m_0=0.0005$  /d to fit rates within the window in Figure 4c. This also gives a value for  
376  $(r_0+m_0)/g_0$  of 0.5, giving food limitation at steady state. We take a value of  $E=0.35$  eV which is at the low end of observations  
377 but using a larger value also worsens the match between model and observational estimates of POC flux. Finally, a value of  
378 2000 is taken for  $a_0$  as it gives the best agreement between model-estimated and observed TOC (Table 2) across the sites given  
379 the other choices of parameter values.

380

381 The sensitivity of key metrics to the model parameters is shown in Figure 5. For each metric, its fractional change for a +/-  
382 10% change in parameter value, from the values given above, is shown. This was carried out for each site assuming the  
383 observed temperature and the POC flux that is consistent with a model fit to the biomass observations at that site with the  
384 imposed 0.2 scaling exponent. The three parameters,  $g_0$ ,  $\alpha$  and  $\beta$  show the greatest sensitivity, with a 20-30% change for a  
385 10% change in parameter. The exceptions are: the largest organism lifetime, which is insensitive to  $g_0$  and  $\alpha$ , and total  
386 biomass, which is insensitive to  $g_0$ . It should be noted that the sensitivities of  $\alpha$  and  $\beta$  work in opposite directions, with a  
387 decreased metric for one corresponding to an increase in the same metric for the other. This means that any simultaneous  
388 change to these parameters while retaining energy equivalence (i.e.  $\alpha=\beta$ ) would have a reinforcing rather than a compensating  
389 effect. Although only the parameter  $g_0$  shows a clear asymmetry about zero for the sensitivity analysis in Figure 5 it should be  
390 noted that the metrics are generally non-linearly dependent on the parameters. This becomes more apparent if a larger fractional  
391 change to parameters in the sensitivity analysis is used. There are some cases where a metric is linearly dependent on a  
392 parameter though (e.g.  $R$  on  $a_0$  – see Equation 13) and in these cases the effect of a +/- parameter change will always be  
393 symmetrical about zero. An obvious feature of the sensitivity analyses is that only  $E$  has an influence which varies across sites.  
394 This is the combined effect of two factors: temperature and POC flux are the only external influences that vary between sites;  
395 and the sensitivity analysis shows fractional, not absolute, changes in the metrics i.e. (This can be understood by looking at the  
396 steady state solution (Equations 10 and 11) and remembering that the temperature and POC flux are the only things that vary  
397 between sites as external influences. The sensitivity is calculated as the metric evaluated at the altered parameter value divided  
398 by the value that obtained using at the standard value. Using the steady state solution (Equations 10 and 11) and taking each  
399 metric in turn we can see how this cross-site consistency arises. For  $R$ , as it varies linearly with the POC flux, the POC flux  
400 cancels out in the sensitivity calculation. The temperature effect (Equation 4) varies across sites, but  $R$  is already insensitive  
401 to this because growth, respiration, mortality and interference all share the same temperature dependence and their effects  
402 cancel each other out (Equation 10). As TOC is calculated simply as a constant multiple of  $R$ , it shares the lack of variation in  
403 sensitivity across sites. Similarly, As the turnover time ( $\tau$ ) is calculated as the ratio of  $R$  to the POC flux (the latter of which  
404 cancels when calculating the metric, (as described above), it also). Thus,  $\tau$  also shows a sensitivity that is constant across sites.  
405 The two metrics of total biomass ( $\Sigma B$ ) and lifetime of largest organism ( $\lambda$ ) do vary across sites for one of the parameters, the

406 ~~activation energy  $E$  do show variation in sensitivity across sites though. The steady state solution for biomass (Equation 11,~~  
407 ~~but perhaps clearer in Equation 12) shows that the temperature effects do not cancel out for the biomass metric; it varies with~~  
408 ~~temperature due to temperature's influence (via  $\theta(T)$ ) on physiological rates (specifically the respiration rate) for which  $E$  is~~  
409 ~~the key parameter. The sensitivity of total biomass to  $E$  is dictated by the ratio of  $\theta(T)$  calculated at altered and original values.~~  
410 ~~Because  $\theta(T)$  is a non-linear function of  $T$  and  $E$ , this ratio varies across sites, so the sensitivity to  $E$  varies with temperature~~  
411 ~~across the sites. Similarly, as the lifetime of the largest organisms is simply the reciprocal of their mortality, a physiological~~  
412 ~~rate which varies with  $\theta(T)$  temperature across sites, then it will too. For the other parameters, neither temperature nor POC~~  
413 ~~flux give rise to variation in sensitivity across sites: none of  $g_0$ ,  $r_0$ ,  $m_0$ ,  $a_0$ ,  $\alpha$ , or  $\beta$  exert an influence through changes in~~  
414 ~~temperature; biomass varies linearly with POC flux (Equation 11), which therefore cancels out when calculating the sensitivity;~~  
415 ~~and the lifetime of the largest organisms has no dependency on the POC flux.~~

416

## 417 2.6 Model performance

418 The observations at all five sites show a power-law distribution of biomass with size, and the model performs well in capturing  
419 this characteristic at each site (Figure 3). Note that the matching of magnitudes in biomass for each size class across sites in  
420 Figure 3g arises from the normalisation for the simultaneous fit, not directly from model parameter value choices. Although  
421 the choice of  $E$  influences the differences in biomass across sites, they are also influenced by inter-site differences in POC  
422 flux. For POC flux (Figure 4b and Table 2), although there is some variability within the observations, the model and  
423 observational estimates agree reasonably well and vary in a similar way across the sites. All have lowest fluxes at the deepest  
424 sites (CCZ, PAP), highest at the shallowest (FG), and the fluxes at these two extremes differ by 1-2 orders of magnitude. For  
425 the physiological rates, in Figure 4c the lines showing the rates should fall within 'window' marked as dotted black lines, and  
426 described in Section 2.5. It is seen that although this is broadly the case, the chosen parameter values already lead to this  
427 'window' being stretched. Maximum bacterial (smallest size class) net growth rate is between 2-4 /d across the sites, higher  
428 than the reference value of 1 /d. The net growth rate achieved at equilibrium to first order matches the respiration rate because  
429 mortality is so much smaller, and varies between 1-2 /d. Life expectancies for the largest organisms are seen to span 40-82 yr  
430 and straddle the reference value of 50 yr. Estimates for TOC using the model vary from 24 gC/m<sup>2</sup> to 6400 gC/m<sup>2</sup>. The observed  
431 range is smaller (from 320 to 2000 gC/m<sup>2</sup>) but given the large variability in observation and model estimates across sites, the  
432 match is reasonable; at some places the model exceeds the observed value and at others it is lower. The turnover time (Figure  
433 4d) is seen to be estimated at ~20 yr, using the model estimated POC flux. It is lower (3-15 yr) using the observations for POC  
434 flux (Figure 4b) and model estimate of  $R$  because the model under-estimates the POC flux. Overall, using this set of parameters,  
435 the model does a reasonable job of satisfying the constraints while simulating the biomass size distribution of the ecosystem  
436 at strongly contrasting locations.

437

## 438 2.7 Uncertainty in parameter values

439 Like any other ecosystem model, there is no single objective choice of parameter values. The sensitivity analysis shown in  
440 Figure 5 and discussed in Section 2.5 indicates the consequences of changing the current values. Additionally, it is insightful  
441 to understand some of the restrictions on changing the values of the parameters.

442

443 A significant constraint arises from the mass scaling exponent for the biomass observations at the different sites. If the scaling  
444 exponent is estimated using data from just a single site (Figure 3a-e), the estimates vary from 0.19 (for FG) to 0.33 (for CCZ).  
445 The uncertainty is largest at OM (CI: [0.1,0.3]) and smallest at PAP (CI: [0.24,0.28]). If data from all sites are combined, as  
446 described in Section 2.5, then the exponent is estimated at 0.26. A sensitivity analysis was done on how the estimate varies if  
447 data from each site in turn is excluded. Excluding data from CCZ, FG, FSC, OM and PAP in turn gives a value that varies  
448 between 0.25 and 0.27. Hence, none of the individual sites is having a significant effect on the estimate of the biomass scaling

449 exponent. As described in Section 2.5, the value used for the interference scaling exponent in the model is equivalent to having  
450 a biomass scaling exponent of 0.2, below the values obtained from the observations described above. Also as described in  
451 Section 2.5, using a greater magnitude value would either lead to a dominance of respiration by larger organisms (rather than  
452 the 'energy equivalence' for which there is some observational support) or require an equivalent increase in the magnitude of  
453 the scaling exponent for the physiological rates.

454

455 The chosen value of -0.2 for the exponent,  $\beta$ , is within the range of observational estimates. Most of these cluster around -0.2  
456 to -0.25 (e.g. Mahout et al., 1995; Ernest et al., 2003). Lower values have been found though. McCoy & Gillooly (2009)  
457 reported an exponent of -0.18 for invertebrates. Lower still, McClain et al. (2012) found an exponent of -0.11 for growth of  
458 benthic organisms, though noted that this was inconsistent with the exponents they found for respiration and mortality (-0.2  
459 and -0.24 respectively). Looking at Figure 4c it is apparent that a using larger magnitude (perhaps to allow a biomass scaling  
460 of 0.25 and retain energy equivalence) would lead to a lower average respiration rate across the size classes, as the respiration  
461 rate for the smallest size organism cannot get any bigger. An increase in the magnitude of  $\beta$  would, therefore, lead to a reduction  
462 in the model estimate of the POC flux and the quality of its fit to observations. For example, if  $\beta$  were changed to -0.25, then  
463  $r_0$  would need to be reduced by a factor of 5 to meet the constraint arising from the maximum bacterial growth rate and the  
464 model POC flux estimates would therefore also decrease by a factor of 5, making them a further factor of 5 lower than  
465 observations, which the model already underestimates.

466

467 The value chosen for  $E$  (0.35 eV) is at the lower end of values derived from observations. For a range of organisms not  
468 restricted to marine ones, Savage et al. (2004) found values from 0.35eV to 0.84 eV. McClain et al. (2012) estimated  $E$  as  
469 0.47eV for respiration and mortality but, with less confidence, 0.16 eV for growth. McCoy & Gillooly (2009) found a value  
470 of 0.69 eV for respiration. For comparison, the canonical value for MTE is 0.63 eV (West et al., 1997; Brown et al. 2004). The  
471 greatest influence of  $E$  is in accounting for significant differences in physiological rates between sites with strongly contrasting  
472 temperatures. Changing temperature from 0°C to 10°C with  $E=0.35$  eV increases rates by a factor of 1.7. Using  $E=0.63$  eV  
473 increases rates by a factor of 2.6. However, if  $E=0.63$  eV then model-estimated POC fluxes reduce by 24-57% across the sites.  
474 This is because the temperature effect  $\theta(T)$  takes lower values with the larger  $E$  even though the difference between  $\theta(T)$  at  
475 different temperatures is greater.

476

477 The greatest subjectivity in choice of parameter value is in deciding the ratio  $(r_0+m_0)/g_0$  which represents the fraction of the  
478 maximum growth rate that an organism achieves at equilibrium. Another perspective on this is that it represents the degree of  
479 food-limitation, with a value of 0 representative of total starvation and a value of 1 of a surfeit. We subjectively took this ratio  
480 to be 0.5 to be consistent with a food-limited ecosystem which is dependent on material that is already the meagre remains of  
481 food that was available to many other organisms as it sank down through the water column (Smith et al., 2008). It should also  
482 be appreciated that two locations with high and low food supplies do not necessarily differ in the degree of food limitation, as  
483 the population sizes at the two sites will reflect the supplies. It is unlikely that direct observations will become available to  
484 constrain this ratio so changes to the current value are more likely to arise from a wish to change one of the three component  
485 parameters,  $g_0$ ,  $r_0$  or  $m_0$ . Keeping  $\beta$  fixed, the first two of these,  $g_0$  and  $r_0$ , are already at the upper limit of observations whilst  
486  $m_0$  is at the lower limit (Figure 4c). Figure 5 shows that decreasing  $g_0$  by 10% would lead to a ~25% increase in  $R$ , TOC and  
487 turnover time. It would also increase the time for the perturbed system shown in Figure 7 to recover to 95% of the steady state  
488 value from 59 to 71 years. A 10% decrease in  $r_0$  would result in the same decrease in estimated POC flux (or a compensating  
489 10% increase in total biomass if it was wanted to preserve the total respiration, Figure 5). The same 10% decrease in  $r_0$  would  
490 give a ~10% decrease in  $R$ , TOC and turnover time (Figure 5). Finally a 10% increase in  $m_0$  would give a 10% decrease in

491 lifetime for the largest organism. Returning to the ratio  $(r_0+m_0)/g_0$ , a 10% decrease in  $r_0$  and  $m_0$  would give a 10% decrease in  
492 the ratio to 0.45. Decreasing  $g_0$  by 10% would increase it by 11% to 0.56.

493

494 A change in  $a_0$  results in a compensating change in  $R$  (Equation 11 shows that they are proportional at steady state ) and  
495 consequently for the model estimate for TOC. The current value of  $a_0$  was chosen to minimise  $\sum_{\text{sites}}[\log(\text{TOC}_{\text{est}}/\text{TOC}_{\text{obs}})]^2$ .  
496 Obviously, the choice of  $a_0$  is sensitive to the subjective choice of this function. If a simple sum of squared differences was  
497 used instead then a choice of  $a_0=500$  would be more optimal, giving a lower range of estimates than those given in Figure 4e  
498 and Table 2 (9.2 gC/m<sup>2</sup> CCZ; 2400 gC/m<sup>2</sup> FG; 560 gC/m<sup>2</sup> FSC; 960 gC/m<sup>2</sup> OM; 74 gC/m<sup>2</sup> PAP). We used the logged ratio to  
499 choose  $a_0$  here because of the orders of magnitude difference between observations across the sites (Figure 4e, Table 2). Using  
500 a simple sum of squared differences biases the value to best fit observations with largest TOC concentrations. The turnover  
501 time of  $R$  i.e. the time it takes the POC flux to replenish  $R$  if removed, has the same sensitivity to  $a_0$ . This can affect the  
502 recovery time to perturbations. Having slower recovery of  $R$  increases the recovery time for small organisms that would  
503 otherwise recover much quicker than large ones because of higher physiological rates. The small organisms cannot fully  
504 recover until  $R$  itself is recovered. This is apparent in the example of how BORIS2 may be used to study a response to a change  
505 in POC flux given later (Section 3.2 and Figure 7). Using a value of 500 in this example would reduce the time taken for the  
506 total biomass to recover to 95% of the steady state value from 59 to 27 years. Turnover times estimated using the model are  
507 also affected by choice of  $a_0$  as shown in Figure 5d. They are consistently ~20 years for the current choice of  $a_0$  but reduce to  
508 5 years if  $a_0=500$ .

509

### 510 3. Use of BORIS-2

511 BORIS-2 runs easily and quickly in Matlab (it was developed, tested and run in version 25.1.0.2973910 (R2025a) Update 1,  
512 and a steady state solution (Section 2.4) is available for situations where equilibrium is the focus. The model requires only the  
513 seafloor temperature and POC flux for a location as inputs. If BORIS-2 is to be used at a specific location then it may be  
514 possible to estimate the local POC flux directly using in situ data from sediment traps (e.g. Durden et al. 2020a; Smith et al.,  
515 2013), although the resuspension of material means that near seafloor data should be treated with care. Alternatively, if it is  
516 intended to use BORIS-2 over larger areas, such as basin scales, then POC flux can be estimated less directly using algorithms  
517 which estimate POC flux at any given depth using satellite remote sensing data (Lutz et al., 2007) or using global  
518 biogeochemical model output (e.g. Yool et al., 2017; Figure 3.21 of Cooley et al., 2022). Alternatively, observations of  
519 sediment community oxygen consumption (SCOC) rates (Smith et al., 2013; Stratmann et al., 2019), which would be expected  
520 to roughly balance POC input on timescales for which the system could be viewed as in steady state, could be used. We now  
521 give a few examples to illustrate the range of potential uses.

522

#### 523 3.1 Using the steady state solution

524 If it is of interest to know how benthic biomass (i.e. the total amount of organisms that can be sustained) varies geographically,  
525 it is useful to focus on the annual average biomass such that it can be assumed that the ecosystem is in steady state to first  
526 order. This assumption allows Equations 10 and 11 to be used for quicker calculations. In Figure 6, use is made of data for  
527 POC flux and temperature at the seafloor to produce a global map of benthic biomass. The POC flux data are generated using  
528 the Lutz et al. (2007) algorithm while seafloor temperature data come from the World Ocean Atlas (Reagan et al., 2024).  
529 Temperature is largely uniform, with little change across the abyssal plains, or even above seafloor ridges, due to the weak  
530 vertical gradients in temperature in the deep ocean. The pattern of low values in subtropics with higher values in tropical,  
531 subpolar, polar and coastal regions for the POC flux is similar to that seen in the export of organic material from the ocean  
532 surface (e.g. Nowicki et al., 2022) but superimposed on this is the effect of depth. POC flux attenuates strongly with depth  
533 (Martin et al., 1987), and a logarithmic scale is needed to capture the variation in seafloor POC flux from shelf to abyssal

534 regions. Because of the largely uniform distribution of seafloor temperature, that of seafloor biomass closely resembles that of  
535 the POC flux for much of the ocean. Only in the Mediterranean and Red Sea are the impacts of much higher temperatures  
536 visible with lower biomasses relative to the variations in POC fluxes because they need to balance greater physiological rates.

537

### 538 **3.2 Running the model dynamically**

539 The dynamic version of BORIS-2 runs easily on a standard laptop, taking just seconds for a thousand years. The forcing data  
540 on POC flux,  $F$ , and the temperature,  $T$ , can also both vary with time if required. This allows a variety of transient responses  
541 to be explored.

542

543 For temperature, long-term temporal change in deep-water temperatures has been detected, but is of a very small magnitude  
544 (e.g.,  $<0.002$  °C/yr Garry et al., 2019). Stronger fluctuations at a site may arise near the boundaries of warm (e.g. Red Sea,  
545 Mediterranean Sea), cool (e.g. Atlantic, Pacific, Indian Oceans), or cold (e.g. Arctic and Southern Oceans) deep waters if their  
546 boundaries move in response to natural or climate-change related shifts. For example, at the Arctic-Atlantic transition in the  
547 Greenland-Iceland-Faroe-Shetland region a  $10^{\circ}\text{C}$  shift in bottom water temperature can occur over a short spatial (bathymetric)  
548 scale (e.g., Turrell et al., 1999) and so a near  $10^{\circ}\text{C}$  shift can occur on short time scales (hours, e.g., Bett, 2001). Generally,  
549 though, because a  $10^{\circ}\text{C}$  change in temperature is required to create a roughly factor of 2 change in physiological rates, scenarios  
550 where time-varying temperature has a significant impact on biomass are likely to be rare for deeper, off-shelf locations.

551

552 Significant changes in POC flux are more likely. For example, considerable uncertainty remains over the impact of climate  
553 change on export of organic carbon from the ocean surface but future changes of up to 41% are possible (Henson et al., 2022).  
554 Such changes in POC flux leaving the surface will impact the benthic ecosystem, which is dependent on the fraction of this  
555 export that reaches the seafloor. One application of the dynamic version of BORIS-2 therefore is in exploring climate change  
556 consequences for the benthos (e.g. Yool et al., 2017). A much simpler example of how the model can be used to study responses  
557 to change in POC flux is shown in Figure 7. Here the ecosystem is initially in steady state but then the POC flux is doubled.  
558 As is apparent in Figure 3, the different sizes of organisms will have very different biomasses. Hence, for ease of comparison  
559 the biomasses and detritus are normalised in Figure 7 by dividing by their final value. Similarly, the variation of physiological  
560 rates with size means that response times differ with size of organism. Using a log time scale allows this to be seen more  
561 clearly. The smallest size class tracks the response of the detritus closely because faster physiological rates allow these  
562 organisms to respond as quickly as the detritus changes. The larger organisms have slower rates and are seen to respond  
563 significantly more slowly as a consequence.

## 564 **4 Discussion**

565 The BORIS-2 model has been presented. It allows simulations of the benthic community across the full size-range of  
566 organisms. A single parameter set has been provided for general use that allows the model to reproduce observed biomass size  
567 distributions at five sites contrasting strongly in location, depth and temperature, while meeting other constraints on POC flux,  
568 expected physiological limitations of smallest and largest organisms and the amount of organic carbon available for food on  
569 the seafloor. It is intended that BORIS-2 be used in a macroecological manner, not to represent the dynamics of specific  
570 species. Physiological processes for organisms within the same size class can vary significantly so BORIS-2 is best suited for  
571 examining questions related to the overall community or the relative behaviour between size classes. As with any model there  
572 are aspects that represent limitations and, as a result, areas for further investigation.

### 573 **4.1 Model assumptions**

574 BORIS-2 assumes all organisms are detritivores feeding from the same common resource, detrital organic carbon on the  
575 seafloor. In practice, a community will have organisms exhibiting a variety of feeding strategies of which detritivory is just  
576 one. Predation, for example, is not captured explicitly by BORIS-2. However, on seafloors deeper than the euphotic zone and  
577 outside of chemosynthetic systems, the benthic ecosystem is supported solely by the POC flux and predation is effectively a  
578 secondary transformation of that carbon. One interpretation is that BORIS implicitly captures predation in the mortality term  
579 but that the gains are distributed across all size ranges rather than received by specific ones. Even from that perspective,  
580 BORIS-2 may under-estimate predation because mortality is parameterised based on natural mortality rate data. In the absence  
581 of suitable data for predation rates and given the large uncertainties in natural mortality, the magnitude and significance of this  
582 underestimate are uncertain. Separate population dynamics for detritivore and predator components of the benthic community  
583 have been studied on the shelf (Blanchard et al., 2009) where it was found that predators might display a stronger increase of  
584 biomass with size than detritivores. A size-based model presented in the same work to explore this interaction further found  
585 that the presence of predators could cause a steepening of the biomass spectrum for detritivores where their size overlapped  
586 with the prey range for predators. However, the predators of benthic organisms were assumed to be largely pelagic – a condition  
587 that is not experienced in the deep ocean. While a similar coupled approach could be adopted in BORIS-2, the main difficulty  
588 in incorporating carnivory into BORIS-2 is the requirement for data on the relative abundance of predators versus non-  
589 predators across size classes. Such data are scarce even on the shelf (e.g. Blanchard et al., 2009). Without such information it  
590 would be difficult to constrain sufficiently the model parameters.

591

592 An additional facet of the ecosystem that is simplified by BORIS-2 is how the organisms obtain their food. In reality, they  
593 may be more or less mobile, allowing them to search for food. They may also be able to filter organic matter from seawater as  
594 suspension feeders, intercepting food before it hits the seafloor or exploiting resuspended or advected material. In theory, the  
595 parameter  $a_0$  could be modified to reflect greater mobility while the growth parameter,  $g_0$ , could be adjusted to capture the  
596 effect of suspension feeding. Once again though, to incorporate such changes would require additional data on the relative  
597 abundance of organisms with these different characteristics across size ranges.

598

599 A final assumption of BORIS-2 worth discussing is that no organic material is either refractory or buried. For burial, a fraction  
600 of the POC flux and/or the mortality could alternatively be regarded as buried and removed from the system. Note that the  
601 POC flux would then have to balance the sum of burial and respiration, so a lower respiration would be required to balance  
602 the same POC flux. Estimates for burial vary from 37% of the POC flux arriving at the seafloor on the shelf to 4% in the deep  
603 (>2000m) ocean (Dunne et al., 2007). The shelf value is in the absence of perturbations to the sediment such as fishing related  
604 trawling. Based on the low deep-sea fraction, burial is omitted. It could easily be added as a (depth-dependent) ‘tax’ on seafloor  
605 POC flux if needed subsequently. Regarding refractory organic carbon, it was described in Section 2.4 that it has been assumed  
606 that only 5% of the total organic carbon in the surface sediment is readily available to the benthic ecosystem represented by  
607 the model. The other 95% is regarded as refractory. Consider two scenarios. The first is that the POC flux arriving at the  
608 seafloor is entirely labile and refractory carbon is created only by the seafloor ecosystem. In this case, it would be possible to  
609 modify BORIS-2 such that a fraction of mortality passed into a refractory carbon pool rather than into  $R$ . In the second scenario,  
610 the POC flux has a refractory component. This could be directed straight into a refractory pool. Reality is likely to be some  
611 combination of these two scenarios. At steady state, all of the organic carbon entering the refractory pool must either be  
612 respired or transformed to labile material and hence returned to  $R$ . In the hypothetical case of no respiration, then the flux of  
613 organic material into and out of the refractory pool should balance, such that the net flux is zero. This is the implicit assumption  
614 in BORIS2, such that this exchange is not modelled. In reality, some of the refractory carbon will be respired, and this could  
615 be incorporated in BORIS-2 in the same way as burial, as a simple extra loss; applied to the POC flux or by adding creation  
616 and respiration of refractory material by the benthic community as appropriate. In the absence of data from multiple sites for

617 the amount of refractory carbon arriving as POC flux or created by the benthic ecosystem, and the fraction of this that is  
618 eventually respired or buried, the dynamics of the refractory pool are omitted. An additional aspect of the refractory carbon  
619 dynamics is that there will be a bacterial population carrying out its respiration that is also not captured by BORIS-2. This  
620 means that observational estimates of bacterial abundance in seafloor sediments are likely to be higher than those predicted by  
621 BORIS-2. With a large pool of refractory carbon (Section 2.5) and associated bacterial doubling time up to thousands of years  
622 (Jørgensen & Marshall, 2016), this additional population is likely to be much larger than represented by the smaller size classes  
623 of the model that are feeding on *R*. For example, bringing together observations from a site in the abyssal Pacific to apply a  
624 linear inverse model for the benthic system including refractory carbon, de Jonge et al. (2020) estimated the prokaryotic  
625 biomass to be roughly equivalent to that for megafauna. In BORIS-2, the biomasses for the equivalent (smallest and largest)  
626 size classes differ by a factor of  $10^5$ . This is not a straightforward comparison though as the prokaryotes in the de Jonge (2020)  
627 study feed from both labile and refractory material. If future data suggest that bacteria need to be taken out of the allometric  
628 framework and treated separately, the biomass estimates of remaining organisms are unlikely to change by more than a factor  
629 of two (the extreme case of bacteria having total biomass equal to all other organisms present), with relative biomass of other  
630 classes unchanged. To make such a change though would require information on the flux or fraction of organic carbon entering  
631 the refractory pool, and the physiological rates of the bacteria ingesting and respiring it. Note that the very definition of  
632 “refractory” is itself an uncertainty. The wide flexibility in the structure of molecules of organic carbon means that POC varies  
633 widely in how “labile” or “refractory” it is. It is not a simple binary, so this adds a further layer of uncertainty.

634

635 Nevertheless, treating bacteria differently could offer one means of increasing the total respiration to give an improved match  
636 to the observations for POC flux and oxygen consumption (Figure 4b). As an extreme example, if only bacteria were allowed  
637 access to the POC flux arriving at the seafloor (rather than everything having equal access at present), with the detrital pool  
638 being supplied instead by dying bacteria, then the observed biomass of other organisms would be supported by a fraction of  
639 the POC flux equivalent to the ratio of mortality to respiration of the bacteria. This ratio is 0.1 for the current parameter set.  
640 Without changing any parameter values this approach would therefore decrease the POC flux available for other organisms in  
641 Figure 4b by an order of magnitude, with respiration of those organisms now exceeding estimates, as anticipated. It would also  
642 increase bacterial biomass by two orders of magnitude (but still three orders of magnitude less than the total). Although there  
643 is some evidence that bacteria respond quickly to POC deposition (Sweetman et al., 2018), other similar studies show a wider  
644 response, including a rapid response by macrofaunal invertebrates (Witte et al., 2003). Note that the latter studies were  
645 conducted at small physical scales that effectively exclude megafaunal invertebrates such that their relative influences and  
646 responses are unknowns (Laguionie Marchais et al., 2020). The option to treat bacteria differently is therefore left for future  
647 study.

#### 648 **4.2 Other possible model extensions**

649 An aspect of BORIS-2 which may benefit future development is the restricted number of external influences. There are  
650 currently only two: the supply of detrital material to the seafloor (POC flux) is the food source for all organisms and ambient  
651 temperature is the only control other than size on metabolic rates.

652

653 The effect of oxygen concentration in seawater is not currently included in BORIS-2. Although it has been questioned whether  
654 there is clear evidence for an oxygen effect on metabolism (Siebel and Drazen, 2007), a lack of clear response to low oxygen  
655 by benthic communities might be as a result of a shift in community composition towards organisms more efficient at extracting  
656 oxygen from waters with low concentrations (Childress and Seibel, 1998). That said, under reduced oxygen concentrations  
657 there is evidence that macrobenthos shift to smaller body sizes (Pearson & Rosenberg, 1978), while meiobenthos may shift to  
658 large body sizes (Moore & Bett, 1989). There may even be a tendency for megabenthos to be eliminated (Pearson & Rosenberg,

659 1978), though they may be enhanced at the peripheries of oxygen minimum zones (OMZs; Levin, 2003). Given the anticipated  
660 expansion of oxygen minimum zones through climate change (Busecke et al., 2022), it is worth noting that commonly applied  
661 thresholds for hypoxia range from 0.3-4 mgO<sub>2</sub>/L, with a modal value of 2 mgO<sub>2</sub>/L. However, the lethal and sublethal levels  
662 for individual taxa vary greatly (Vaquer & Duarte, 2008). In a formal environmental monitoring context (e.g. EU Water  
663 Framework Directive), oxygen concentrations below 4 mg O<sub>2</sub>/L are considered to be of concern (Best et al., 2007). There is  
664 therefore value in finding a way to incorporate an oxygen effect in BORIS-2 if sites <4 mgO<sub>2</sub>/L are of interest, and particularly  
665 if concentrations are likely to be below 2 mgO<sub>2</sub>/L.

666

667 The impact of seafloor type is another area where BORIS-2 may benefit from further analysis and expansion. At present, for  
668 simplicity, BORIS-2 makes no distinction in the nature of the seabed environment, other than bottom water temperature and  
669 POC flux. The implicit assumption is that it is applied in a sedimentary environment. In practice, the seafloor represents a  
670 range of environments varying on scales from a single manganese nodule to an ocean basin. Seafloor type can influence both  
671 motility (with some suspension feeders favouring hard substrata) and the efficiency with which food can be obtained (such as  
672 hills or trenches which can focus bottom currents carrying suspended POC). Whether BORIS-2 can be configured for different  
673 seabed environments by suitably adjusting parameter values and/or by splitting the ecosystem into populations with different  
674 feeding traits is left for future developers.

## 675 5 Conclusions

- 676 • Based on allometric scaling of metabolic processes, the BORIS-2 benthic ecosystem model is capable of simulating  
677 population dynamics of organisms ranging in size from bacteria to large metazoans, over 18 orders of magnitude.
- 678 • The only external information required is the POC flux to the seafloor and the ambient temperature.
- 679 • It can be run dynamically but a steady state solution also exists and is given.
- 680 • A parameter set is provided suitable for general use globally and capable of simultaneously providing a good  
681 reproduction of observed biomass size spectra at five locations contrasting in depth, food supply, and temperature.
- 682 • This model offers considerable flexibility in application, at a range of scales, from responses to regional perturbations  
683 such as deep-sea mining, to studies of climate-driven global change in the benthos.

684

## 685 Appendix: Differences between BORIS-1 and BORIS-2

686 A brief description is given here of the differences between the BORIS-1 and BORIS-2 models. Full details of BORIS-1 can  
687 be found in Kelly-Gerreyn (2014).

### 688 (1) Range of organism sizes reproduced

689 BORIS-2 has been designed to reproduce the full range of benthic organism sizes, whereas BORIS-1 focussed on a  
690 limited range of sizes coincident with the data then available for comparison. BORIS-2 overlaps exactly with the 16  
691 size classes of BORIS-1; size class 27 of BORIS-2 matches size class 1 of BORIS-1. BORIS-2 therefore extends for  
692 26 smaller size classes and 17 larger size classes than BORIS-1, to provide more complete coverage of the range of  
693 organism sizes.

### 694 (2) Choice and representation of physiological/ecological processes

695 Broadly, BORIS-1 and BORIS-2 are structurally similar, with dynamics arising from the three processes of growth,  
696 respiration and mortality - but they differ a little in how they do this. First, in BORIS-1 growth is the net effect of  
697 ingestion then assimilation. Ingestion was allowed to scale with body size in BORIS-1, but assimilation was just  
698 assumed to be a fixed fraction of this. The difference between them was treated as waste and returned to *R*. To simplify  
699 this in BORIS-2, a single net growth rate is used, effectively the combined product of ingestion and assimilation.

700 Also, in BORIS-1, growth rate increased linearly with the amount of food available. In practice an organism's ability  
701 to ingest and assimilate food cannot increase indefinitely. In BORIS-2 the representation of growth is therefore  
702 modified such that it saturates at high food abundance. It is also modified to include the effect of other organisms  
703 competing for the limited food supply (Section 2.2). Second, respiration in BORIS-1 is represented as a fraction of  
704 growth, and this fraction can vary independently with size. A consequence is that BORIS-1 does not specifically  
705 capture basal metabolism, the 'tax' paid by any organism just to keep alive. In BORIS-2 respiration is represented as  
706 a separate process independent of growth. This better represents basal metabolism. Additionally, an organism will  
707 need to fuel active metabolism, the energy requirements above basic maintenance required for such things as  
708 movement. In BORIS-2 it is implicitly assumed that this is included in the net growth rate. In BORIS-1 it was also  
709 necessary to assume a fixed fraction of POC flux that was respired by organisms not captured by the model. By  
710 expanding the size range to cover all organisms in BORIS-2, this assumption (and parameter) is no longer required.  
711 Third, a linear mortality parameterisation was used in BORIS-2, in place of the quadratic one used in BORIS-1. This  
712 choice is influenced by the impact of the mortality term on organism lifetimes. Temperature has been argued to be  
713 the first order control on mortality (e.g. McCoy and Gillooly, 2008). There is found to be roughly a factor of two  
714 increase in mortality rate per 10°C increase in environmental temperature. As context, the sites described in Section  
715 2.5, which span from shelf to deep ocean, only encompass a 14°C range of temperatures, roughly consistent with a  
716 factor less than 4 range in mortalities. The quadratic parameterisation, however, leads to a dependence of the specific  
717 rate of mortality on biomass (i.e. with a quadratic mortality term the specific rate is  $\mu_i B_i$ ), and hence on food supply  
718 (the POC flux,  $F$ ); this is because  $B_i$  increases with  $F$  (see Section 2.4) and  $F$  can vary significantly. Assuming the  
719 ecosystem is in equilibrium, the respiration of organic carbon on the seafloor can be used to estimate  $F$ . Using the  
720 large collection of globally distributed Sediment Community Oxygen Consumption (SCOC) data (Stratmann et al.,  
721 2019), and limiting to sites less than 6000m in depth and accepting only in situ measurements, SCOC ranges from  
722 0.14-110 mmol O<sub>2</sub> m<sup>-2</sup> d<sup>-1</sup> (2.5% and 97.5% confidence levels). Excluding shelf regions (<200m depth) reduces the  
723 range to 0.08-12 mmol O<sub>2</sub> m<sup>-2</sup> d<sup>-1</sup>, but it still spans several orders of magnitude. (Note that this result is unaffected by  
724 the units as typically a constant factor is used to convert to carbon units.) Such large variations in POC flux would  
725 induce similar variability in the longevity of organisms of the same size between different locations if a quadratic  
726 mortality parameterisation was used. Therefore, using the linear form of mortality in BORIS-2 avoids a much larger  
727 inter-site variability in mortality rate than is currently supported by observations.

### 728 (3) Observational constraints

729 BORIS-1 was compared to size-resolved data from 3 locations: FG, FSC and OM. While they do contrast in depth  
730 and temperature (Table 2) they do not represent the deep seafloor that covers much of the globe. For BORIS-2,  
731 additional data from CCZ and PAP are used giving data from two abyssal locations of contrasting food supply. For  
732 BORIS-1, the assumption of a fraction respired by non-modelled organisms meant that a comparison to observations  
733 of POC flux would be quite subjective. With this restriction removed in BORIS-2 the modelled POC flux is now  
734 compared to 2 independent estimates of POC flux at 4 sites and 3 estimates at PAP. For BORIS-1 additional  
735 constraints regarding the exponents for size scaling were imposed (see (4) below), as well as a range of expected  
736 values for smallest (meiofauna) and largest (macrofauna) organisms. A similar thing to the latter is done for BORIS-  
737 2 but, by necessity, for much smaller (bacteria) and larger (megafauna) organisms because of the expanded size range.  
738 An additional constraint for BORIS-2 is provided by estimates of TOC in seafloor sediment, which provide a  
739 constraint on  $R$ .

### 740 (4) Method of selecting parameter values

741 BORIS-2 has one fewer parameter (7) than BORIS-1 (8). This is despite BORIS-2 incorporating two new processes:  
742 temperature sensitive physiology and interference. Without these additions BORIS-2 would have 5 parameters. To

743 find suitable parameter values for BORIS-1 an optimisation algorithm was used. With fewer parameters and a greater  
744 set of constraints this was not necessary for BORIS-2. While it might be possible to construct a similar optimisation  
745 routine for BORIS-2, it currently does not warrant the effort, particularly given the subjectivity in constructing the  
746 necessary cost-function for the optimisation. The user can easily explore parameter space and make a decision on the  
747 most suitable parameter values simply using Figure 3 and 4.

748

#### 749 **Code and data availability**

750 All code and data for generating the figures in this paper and for using BORIS either at steady state or dynamically are  
751 available on Zenodo at <https://doi.org/10.5281/zenodo.19235638> (Martin et al., 2026). A user manual can be found in the  
752 Supplement to this manuscript.

#### 753 **Author contribution**

754 AM was responsible for Writing the original draft and for Software, Validation and Visualisation. AB assisted in model  
755 development, testing and sensitivity analysis. All authors were involved in the Conceptualization and Methodology,  
756 contributing to the design of the model, and in the review and editing of the Writing. AY, AB, BJB, JD, and SO provided  
757 advice related to the Validation of the model.

#### 758 **Competing interests**

759 One author (AY) is a member of the editorial board of journal "Geoscientific Model Development".

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765

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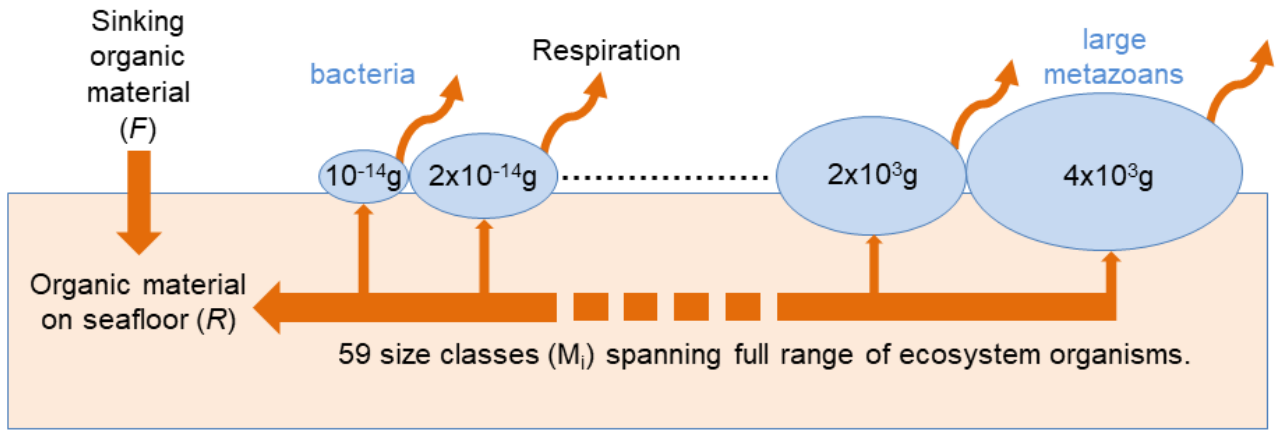
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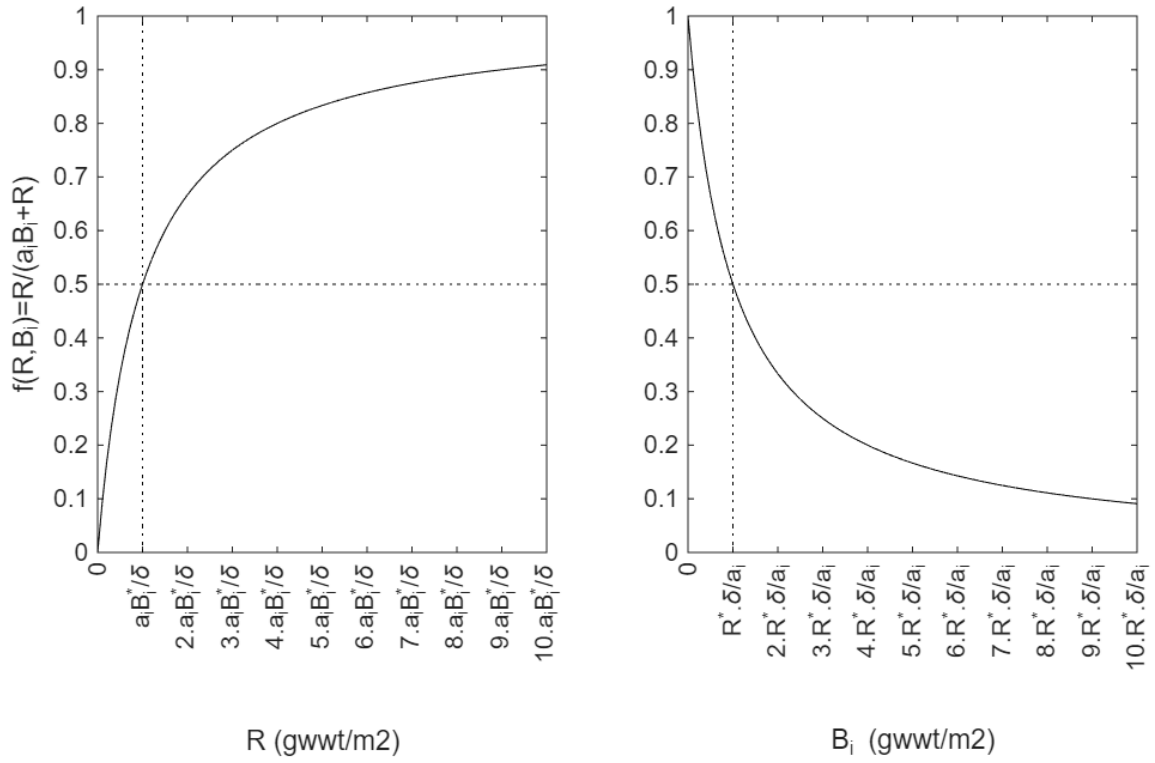
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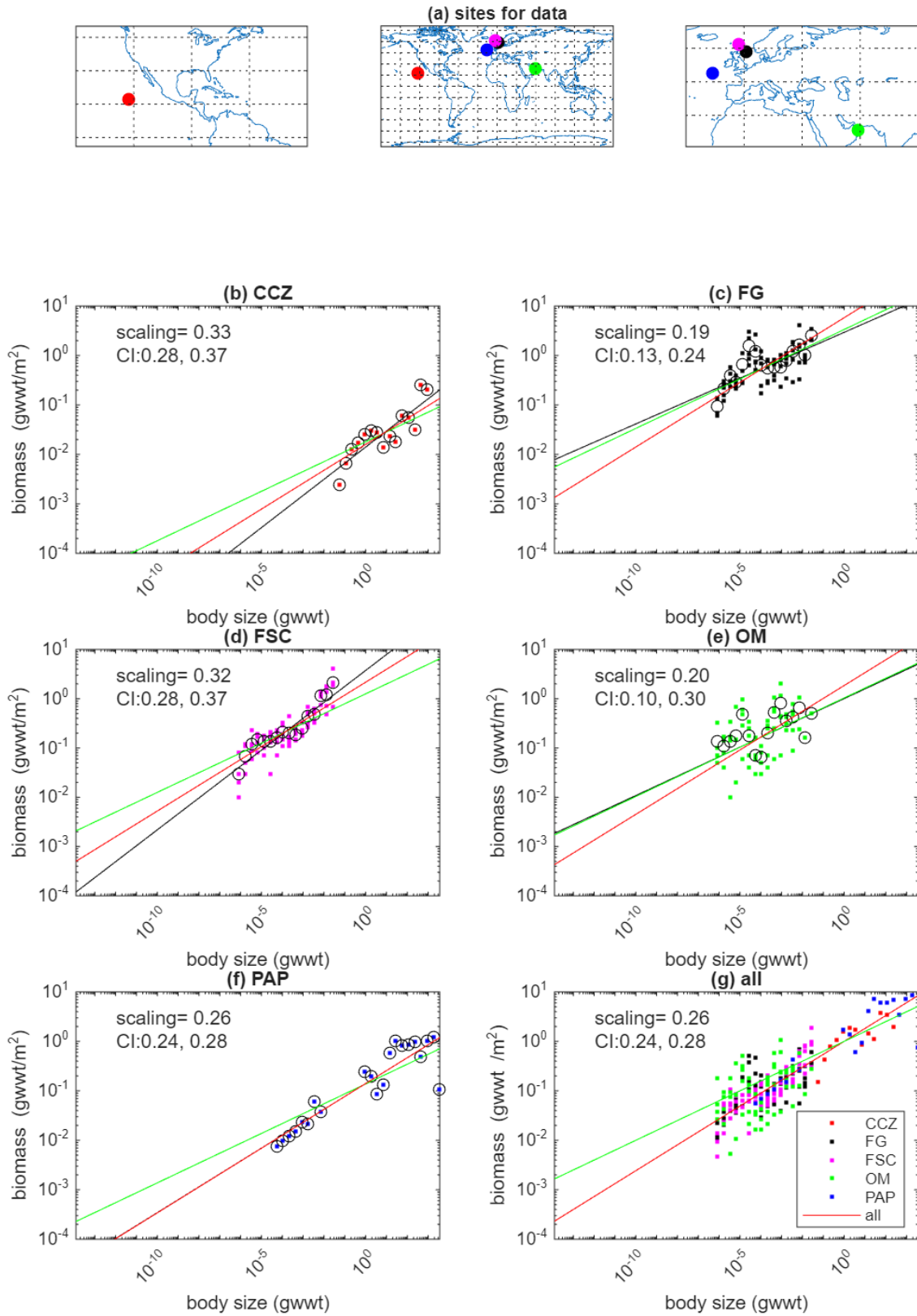
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Figure 1: Schematic of BORIS-2. Within each size class ( $M_i$ ) the total biomass ( $B_i$ ) is controlled by growth, using organic material on the seafloor, and losses to respiration and mortality. Organic material ( $R$ ) accumulates on the seafloor from the deposition of sinking particulate organic carbon ( $F$ ) and mortality of benthic organisms. Growth, respiration and mortality are all assumed to scale as a power-law with  $M_i$  and as an exponential function of ambient temperature (see Section 2). Numbers denote median mass (units: g wet weight) for each size class and example organisms of smallest and largest size classes are given in blue.



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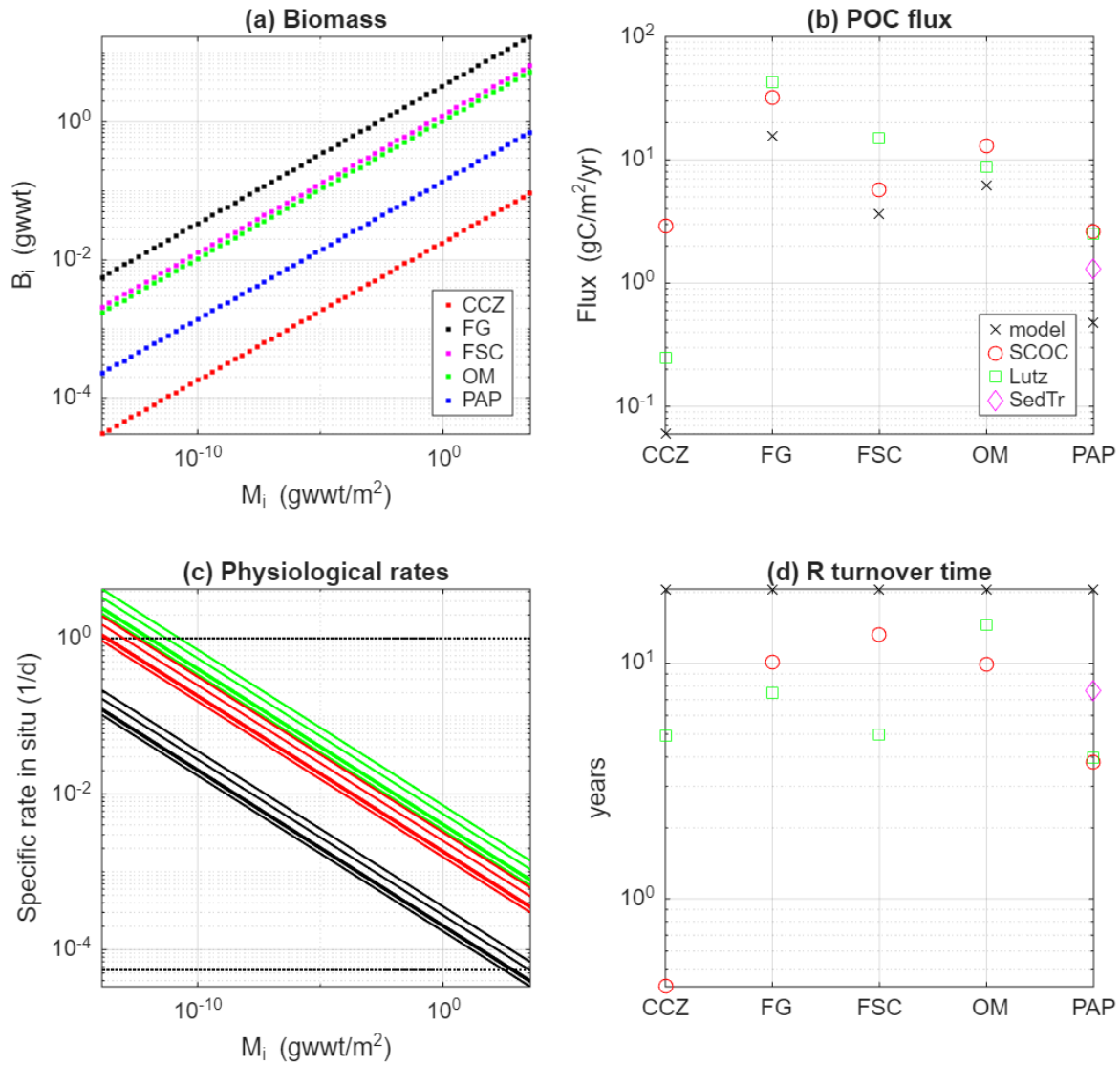
Figure 2: Plots showing how the growth limitation function,  $f(R, B_i) = 1 / (1 + a_i B_i / R)$  varies as either  $R$  (left) or  $B_i$  (right) is varied.  $R^*$  and  $B_i^*$  denote values at steady state (vertical dotted line), where additionally  $\delta R^* = a_i B_i$ . Note that because of the scaling behaviour of  $a_i$ ,  $a_i B_i$  is the same for all  $i$  at any site.



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1019 **Figure 3** Locations (top) and observations of biomass as a function of body size for the five sites listed in Table 2. For each of the  
 1020 sites the dots show observations, circles indicate means within size classes and the black line is a power-law fit to the observations.  
 1021 The fitted values for the scaling exponent together with 95% confidence intervals are also shown. The bottom right panel shows the  
 1022 fit to all sites simultaneously, assuming a common scaling exponent (fitted value shown with 95% CI) but allowing the pre-factor to  
 1023 vary across sites. The dots show data from the sites with colours matching those used in the panels for individual sites. Note that in  
 1024 the bottom right panel the data from each site has been normalised by dividing by the fitted pre-factor to allow the visual comparison.  
 1025 The red line in all panels is the simultaneous fit to all sites. The green line in all plots is the relationship used in the model by fitting  
 1026 a power-law with an imposed scaling of 0.2 to the same observations.

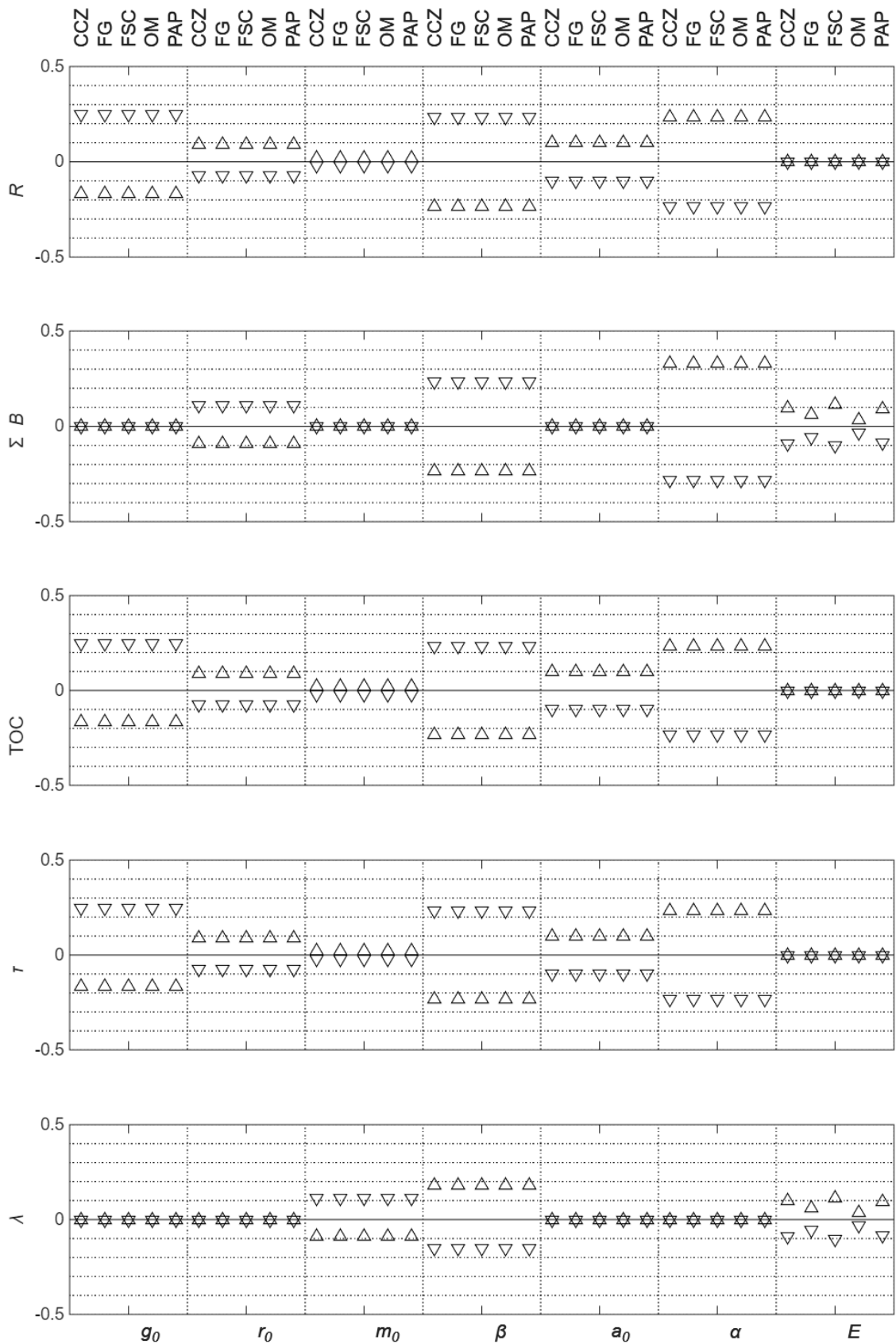
1027



(e) Table of diagnostics

Site	Total biomass gC/m2	R gC/m2	TOC Model gC/m2	TOC Obs gC/m2	Biggest organism lifetime yr
CCZ	0.062	1.2	24	400	72
FG	12	320	6400	800	51
FSC	4.3	75	1500	550	82
OM	3.6	130	2600	2000	40
PAP	0.48	9.9	198	320	68

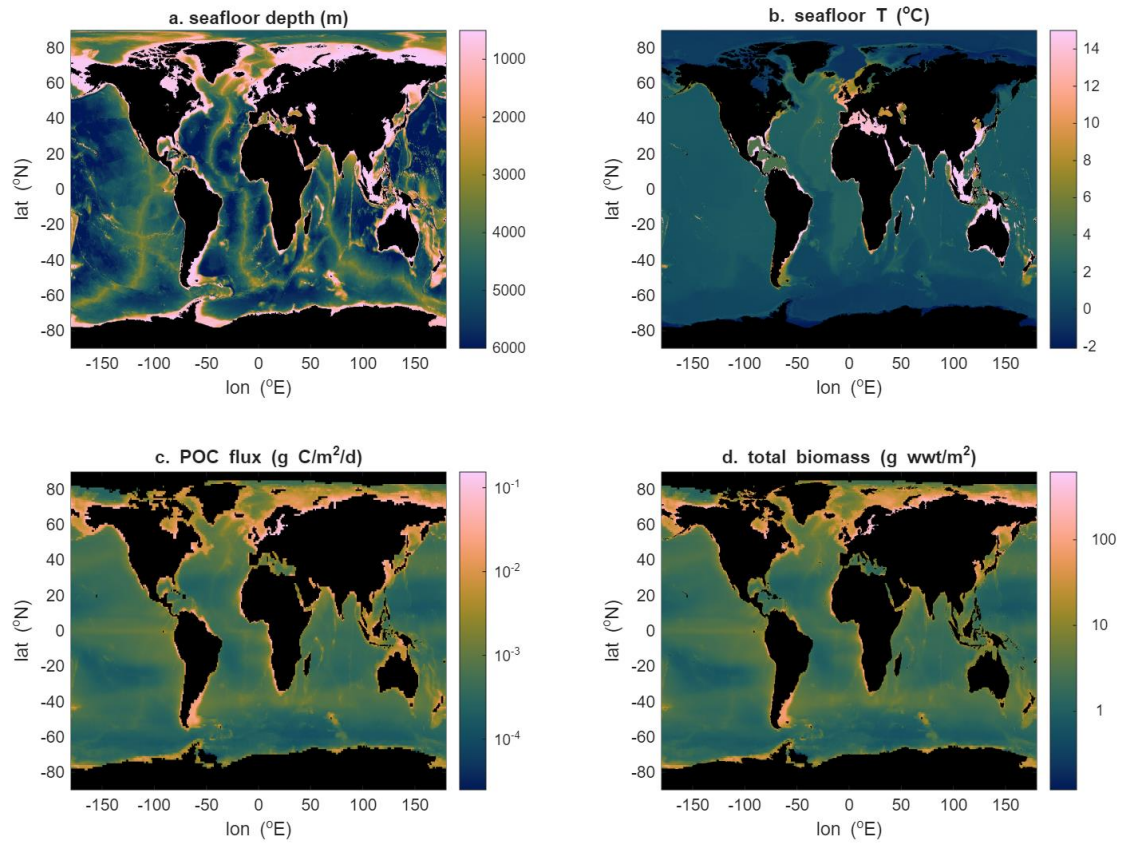
1028  
 1029 **Figure 4: Model diagnostics and observational constraints: (a) modelled biomass (gwwt/m<sup>2</sup>) size distributions for the five sites using;**  
 1030 **(b) POC flux estimates using the Lutz et al. (2007) algorithm (Lutz), SCOC data (SCOC), sediment trap observations (SedTr) and**  
 1031 **the model constrained by biomass data from each site (model). Note that sediment trap data are only available for PAP; (c) model**  
 1032 **specific rates for maximum net growth (green), respiration (red) and mortality (black) for each of the sites. Note that they differ**  
 1033 **between sites because of the temperature effect. The two dotted black lines correspond to the constraints of a maximum growth rate**  
 1034 **of 1 d<sup>-1</sup> for the smallest organisms and a lifetime of 50 yr for the largest organisms.; (d) turnover time (R/POC flux) for each of the**  
 1035 **sites, estimated using model estimates of R and all observation and model estimates for the POC flux from (b); (e) a summary table**  
 1036 **of diagnostic parameters including total biomass (model), R (model), TOC (model and observations) and biggest organisms' lifetime**  
 1037 **(model)**  
 1038



1039  
 1040 **Figure 5** Sensitivity of the metrics detritus ( $R$ ), total biomass ( $\Sigma B$ ), TOC, turnover time of detritus ( $T$ ) and lifetime of largest organism  
 1041 ( $\lambda$ ) to each of the model parameters across the five sites. In each case the sensitivity shows the fractional change resulting from a +/-  
 1042 10% change in the parameter from the standard value. Symbols point in the direction in which the parameter was changed i.e.  
 1043 upward-pointing pyramids indicate an increase in parameter value and inverted pyramids indicate a decrease.

1044

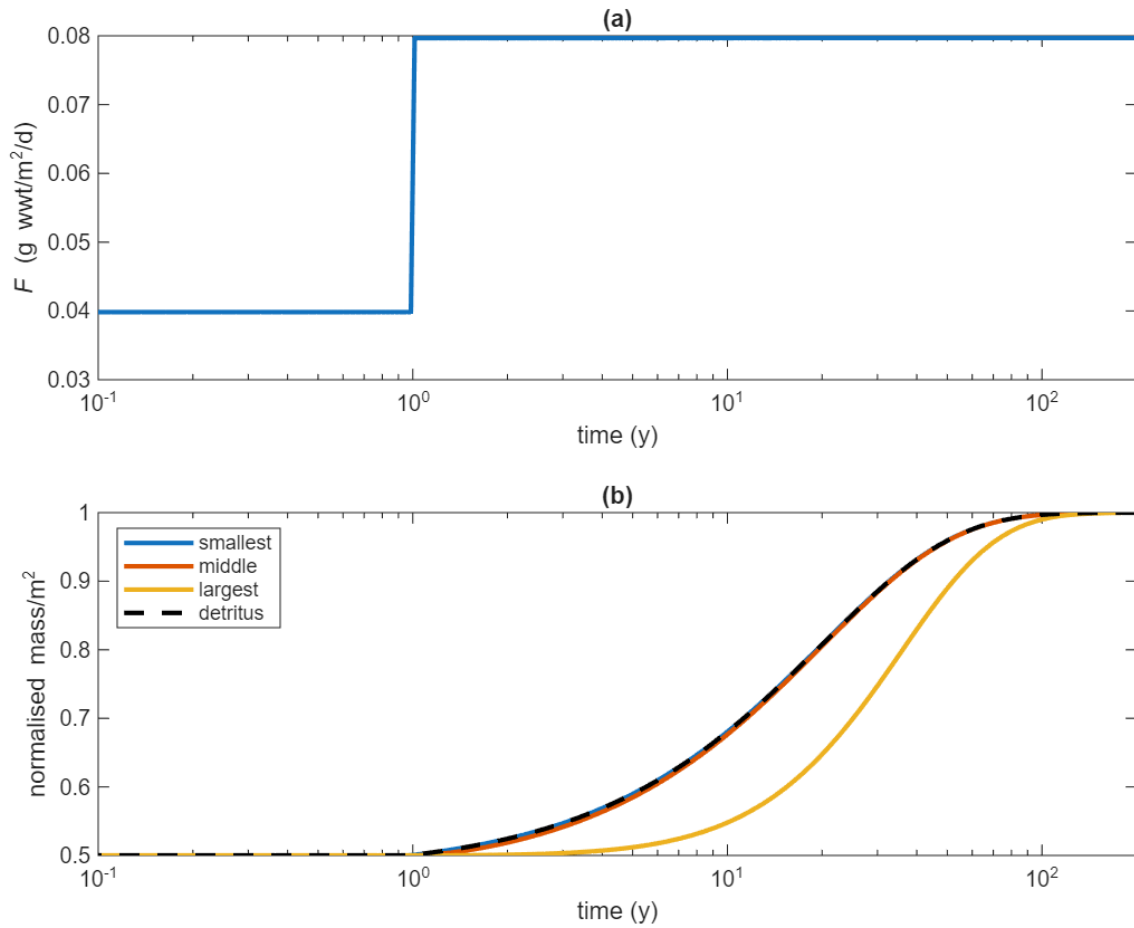
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1049 **Figure 6** Example of using the steady state solution (Equations 10 and 11) to explore spatial variability in total benthic biomass:  
1050 seafloor depth (a) and temperature (b) from the World Ocean Atlas, particulate organic carbon (POC) flux (estimated using Lutz  
1051 et al., 2007) (c), and total biomass (d).

1052



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1054 **Figure 7: Modelled response to a perturbation in which the ecosystem is initially in steady state and then a doubling of the POC flux**  
 1055 **takes place: (a) shows the POC flux, with doubling occurring after 1 year; (b) shows the response of organisms and detritus. For**  
 1056 **clarity only 3 size classes are shown, the smallest, middle and largest ( $i=1, 30$  and  $59$ ). For the same reason, the biomasses and**  
 1057 **detritus are normalised by dividing by their final value. A log time scale is also used to highlight the different response timescales.**

1058

Parameter	Description	Value	Units
$g_0$	Max. net growth rate at 20°C for organism of size 1 gwwt	0.01	d <sup>-1</sup>
$r_0$	Respiration rate at 20°C for organism of size 1 gwwt	0.0045	d <sup>-1</sup>
$m_0$	Mortality rate at 20°C for organism of size 1 gwwt	0.0005	d <sup>-1</sup>
$a_0$	Interference pre-factor	2000	-
$\beta$	Scaling exponent for growth, respiration and mortality	-0.2	-
$\alpha$	Scaling exponent for interference	-0.2	-
$E$	Activation energy	0.35	eV

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1061 **Table 1: parameter set for BORIS-2**

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Site	Lat °N	Lon °E	Depth m	Temp °C	POC flux g C m <sup>-2</sup> y <sup>-1</sup>				TOC g C m <sup>-2</sup>		R g C m <sup>-2</sup>
					Lutz	SCOC	Trap	Model	Obs	Model	Model
					CCZ	17.2	-122.6	4150	1.5	0.25	2.9
FG	58.3	0.9	153	8	43	32	-	16	800	6500	320
FSC	61.9	-2.8	1623	-1	15	5.7	-	3.6	550	1500	75
OM	23.4	59	507	13	8.8	13	-	6.2	2000	2600	130
PAP	48.8	-16.5	4850	2.6	2.5	2.6	1.3	0.48	320	200	9.9

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1066 **Table 2: Information on sites from which data were used to constrain parameter values for the general purpose**  
1067 **parameter set given in Table 1, together with model diagnostics. POC flux is at the seafloor, estimated using the**  
1068 **algorithm in Lutz et al. (2007), SCOC or sediment trap. TOC estimates come from Parameswaran (2024). The “Model”**  
1069 **columns indicate model diagnostic values. Model TOC is calculated by assuming that R is 5% of TOC. Sources for data**  
1070 **are given in Section 2.5**

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