Title: Retrieval of SO_2 columns from FY3F/OMS instrument observations

Authors: Huanhuan Yan, Andreas Richter, Xingying Zhang, Anja Schönhardt, and Thomas Visarius

We thank the AMT Editors and the anonymous Referees #1 and #2 for handling the review process, carefully reading our manuscript, and providing constructive comments. The paper was corrected according to the suggestions of the two reviewers. We hope that in dealing with the comments put forward by Anonymous Referees #1 and #2, the quality of the manuscript was improved. We addressed the comments below.

Answer to Referee #1

1. **Reviewer 1:** Treatment of missing retrieval data. It is understandable that the test retrievals can fail to produce valid data in certain circumstances for certain areas. However, I find the treatment of this missing data in the manuscript to be rather odd. The authors take great care to document the technical nature of missing data in retrieval output (e.g. they explicitly state how "nan values" and fill values are represented in color maps), which I do not believe to be relevant in a scientific publication. However, in most cases they do not explain why this data is missing, which would be highly relevant, and often essential. These explanations should be added.

   Authors: We thank the reviewer for this valuable comment. We agree that, in addition to describing how missing values are represented in the figures, it is important to explain the reasons for missing retrieval data. In the revised manuscript, we have added explanations for the main causes of missing data in the retrieval results, such as the gap between two orbits, Nan values in the downloaded TROPOMI data, and invalid DOAS fitting results.

The following changes were made in the revised manuscript:

1) P15, L302: For the Nan values in the TROPOMI COBRA data, we contacted one of the developers of the COBRA algorithm, Dr. Nicolas Theys, to determine their origin. The TROPOMI COBRA L3 grid product (0.022° × 0.022° equal latitude–longitude grid) was generated from L2 data using the HARP gridding tool with a Quality Assurance (QA) filter (QA > 0.5) to remove low-quality data.

   In Section 4, at the end of the second paragraph, we added the following sentence to explain the cause of Nan values in the TROPOMI COBRA data: "The quality filtering leads to some gaps in the COBRA data."

2) In Figure 8, revised "The missing pixels in the Figure c are caused by Nan values in TROPOMI COBRA data" to "The missing pixels in Figure c are due to quality filtering applied to TROPOMI COBRA data" to clarify the cause of the missing pixels.

   In Figure 9, revised "The missing pixels in the Figures b and c are caused by the gap between the two orbits and the Nan values in TROPOMI COBRA data" to "The missing pixels in Figure b are due to the gap between the two TROPOMI orbits, while those in Figure c are due to the quality filtering applied in the TROPOMI COBRA data" to more clearly explain the causes of the missing pixels in each subfigure.

   In addition, similar descriptions of missing data in TROPOMI COBRA data have been modified throughout the manuscript.

2. **Reviewer 1:** Lack of summary/conclusions of agreement between OMS and TROPOMI**.** I would have liked to see more clear statements and a summary regarding agreement of OMS and TROPOMI data. The authors provide quite a few possible reasons for differences in SO_2 column values, but do they generally believe that those reasons are sufficient to explain the differences? How do these differences compare to OMS and TROPOMI error estimates? Also, although error

analysis was discussed in detail, a clear representation and discussion of the resulting error values seems to be missing.

Authors: We thank the reviewer for this comment. We agree that the manuscript would benefit from a clearer summary of the agreement between OMS and TROPOMI SO_2 products. Our responses to these suggested revisions are provided below.

1) **Comment:** I would have liked to see more clear statements and a summary regarding agreement of OMS and TROPOMI data.

Answer: In the revised manuscript, the first paragraph of Section 6 has been reorganized into four paragraphs, with expanded descriptions that more clearly elaborate on both the agreement and the differences between the OMS and TROPOMI results. The revised content in the revised manuscript is as follows (in red font):

This study utilized TOA reflected radiance data from the Chinese FY3F/OMS-N instrument, launched in August 2023, to retrieve global SO_2 columns with a DOAS approach. Based on the characteristics of the OMS instrument and the performance of its L1 data, specific schemes, including solar spectrum selection, spectral soft calibration, and background offset correction, were developed to effectively reduce along-track stripes and across-track asymmetry in the initial OMS SO_2 retrievals.

The OMS SO_2 retrievals were compared with TROPOMI DOAS and TROPOMI COBRA SO_2 products in clean oceanic regions, under volcanic eruption conditions, and in anthropogenic emission regions. The comparison results indicate that OMS retrievals show reasonable agreement with TROPOMI products, have good stability in clean oceanic regions and can be used to monitor SO_2 emissions from volcanic eruptions and anthropogenic sources. In selected clean oceanic regions, the SO_2 values of both OMS and

TROPOMI follow approximately a normal distribution centered around 0, with most values concentrated between -2 DU and 2 DU. For the Sundhnúkur and Nyamuragira volcanic eruptions, FY3F/OMS SO_2 retrievals successfully capture the spatial distribution and high-concentration plumes of volcanic SO_2, similar to the TROPOMI DOAS and TROPOMI COBRA 7 km SO_2 results. Over the Sundhnúkur volcano, OMS and TROPOMI DOAS show a high correlation of ~0.87, and OMS and TROPOMI COBRA reach ~0.76, indicating good overall agreement. However, OMS tends to underestimate SO_2 at high columns (>50 DU) due to saturation in the 312–326 nm fitting window. In anthropogenic emission regions, OMS and TROPOMI SO_2 products show generally good consistency in detecting anthropogenic SO_2 emissions, with correlation coefficients ranging from about 0.5–0.6 over the Persian Gulf and up to 0.91–0.93 over Norilsk.

The differences between OMS and TROPOMI SO_2 results may be related to differences in local overpass times, spatial resolution, observation angles, and the L1 and L2 processing algorithms (e.g., differences in L1 radiometric and spectral calibration methods, SO_2 retrieval fitting windows, AMF strategies). Among these, the AMF used in the SO_2 column retrieval is a major contributor to the differences between OMS and TROPOMI SO_2 results. For example, in the case of the Sundhnúkur volcano, the lack of accurate information on the vertical SO_2 profile can lead to discrepancies of more than a factor of two when comparing the OMS and TROPOMI SO_2 results. Random noise and uncertainties from background correction are relevant for low SO_2 scenarios, such as over the Persian Gulf, and lead to scatter in the order of several DU. However, the results for Norilsk demonstrate that under relatively constant emission conditions, good agreement can be achieved with a simple AMF when the satellite overpass times are well matched.

In summary, the agreement between the OMS and TROPOMI measurements is within expectations, taking into account the differences in

satellite overpass times and the uncertainties associated with AMF assumptions. With its high spectral and spatial resolution, morning overpass time, daily global coverage, and reliable SO_2 retrieval results, OMS will provide effective data support for monitoring the continuous SO_2 changes from global volcanic eruptions and anthropogenic activities, helping to fill the spatial and temporal gaps in the existing global satellite network.

2) **Comment:** The authors provide quite a few possible reasons for differences in SO_2 column values, but do they generally believe that those reasons are sufficient to explain the differences?

Answer: Although several possible reasons for the differences in SO_2 column values have been listed, we are not fully confident that these explanations are sufficient to account for all observed discrepancies. There may also be factors beyond our current understanding. At present, we thought that the reasons listed can explain the observed differences, and that the observed differences are in the expected order of magnitude.

3) **Comment:** How do these differences compare to OMS and TROPOMI error estimates?

Answer: Some of the error sources of OMS and TROPOMI SO_2 retrievals overlap with the factors causing differences between the two products. For example, errors in L1 radiometric and spectral calibration are not only error sources that affect the accuracy of both OMS and TROPOMI SO_2 retrievals, but also contribute to the differences between them. Similarly, the choice of retrieval window is both an error source for product accuracy and a factor leading to differences between OMS and TROPOMI SO_2 retrievals.

Some factors contributing to the differences are not intrinsic error sources of the OMS and TROPOMI SO_2 retrievals. For instance, the difference in local overpass times and spatial resolution will inevitably lead to discrepancies, as

the instantaneous atmospheric states observed by OMS and TROPOMI are not the same.

4) **Comment:** Also, although error analysis was discussed in detail, a clear representation and discussion of the resulting error values seems to be missing.

Answer: Thank the reviewer for this valuable comment.

There are many factors that affect the accuracy of SO_2 retrievals. In principle, one should analyze how each individual error source contributes to the total retrieval uncertainty. In practice, however, these error sources are complex and difficult to separate. For example, instrument random noise and calibration errors are challenging to quantify precisely since the true SO_2 values are not known. Although the official OMS L1 product provides specified performance requirements, our in-orbit tests indicate that the actual OMS L1 errors deviate from the planned specifications for certain parameters. At present, it is therefore difficult to obtain precise quantitative estimates of how instrument noise contributes to the SO_2 retrieval accuracy.

In addition, to address striping and cross-track asymmetry in the SO_2 SCD retrievals, we applied a background offset correction scheme. However, since the true SO_2 values are unknown, it is not possible to fully verify the accuracy of these corrections. While values over clean regions and their scatter could be used as an indicator of the correction quality, this method does not account for possible overcorrections or systematic biases.

Taking the OMS 20240823_1036 orbit as an example, here we present two plots of the SO_2 SCD retrieval RMSE and spectral fitting error (SFE). These figures show the magnitude and distribution of errors in the retrieved SCDs. Both RMSE and SFE values remain relatively large, with pronounced striping features, and are particularly high on the right side of the orbit. This is most likely related to instrument radiometric and spectral calibration errors, which are difficult to separate and remain a major source of uncertainty. The

corrections applied in our product largely reduce the impact on the SO_2 columns, but proper accounting for this in the error estimates is difficult. Ongoing improvements to the OMS L1 data are expected to reduce striping and asymmetry in the calibration in order to improve SO_2 retrieval accuracy. Considering these aspects, we decided not to include the RMSE and SFE results in this revised manuscript.

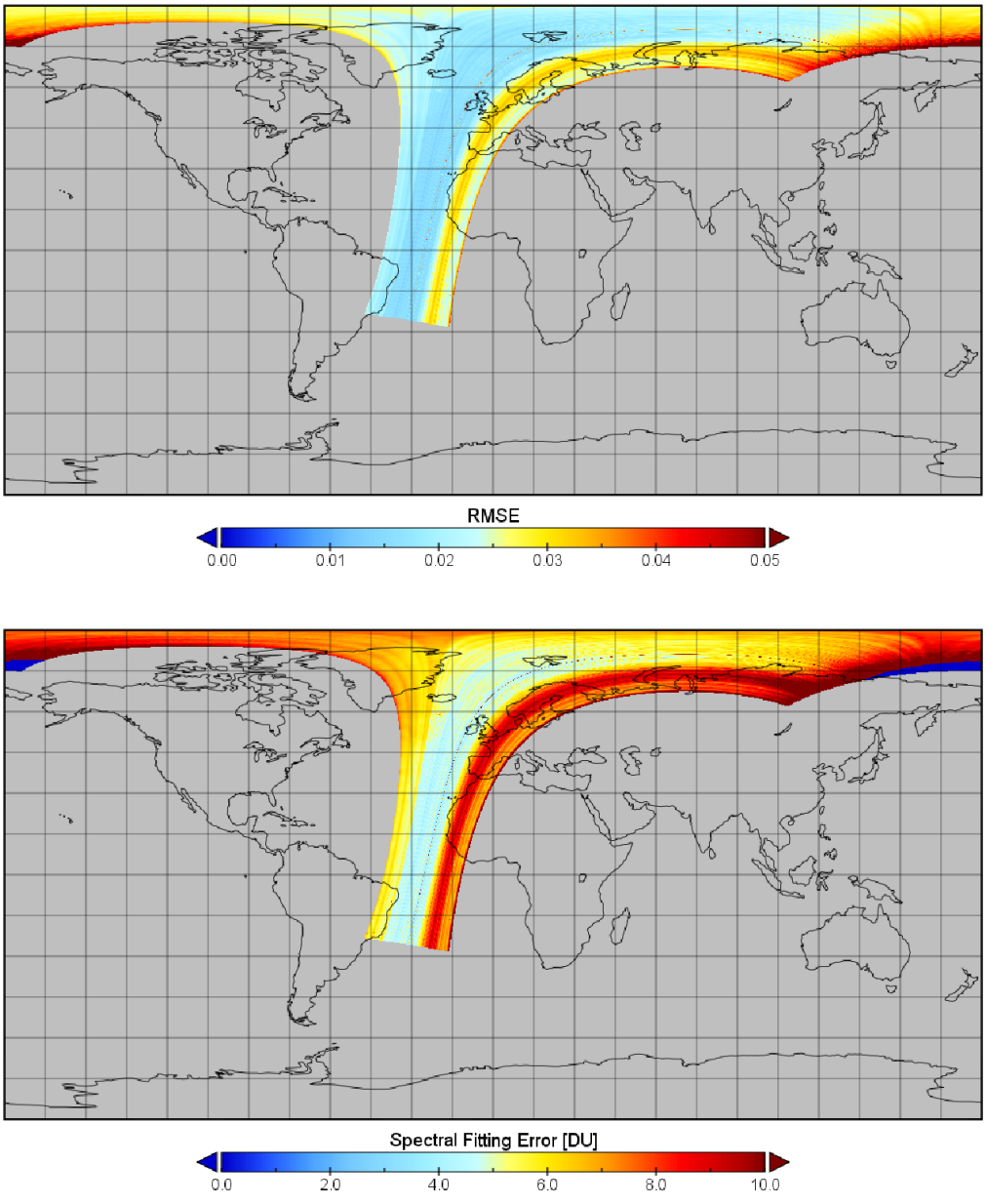


Figure: RMSE and Spectral Fitting Error of OMS SO_2 SCDs for OMS orbit 20240823_1036.

In response to Reviewer 2's comment, we have supplemented the AMF section (P35, L593-603) with additional calculation results to provide a rough estimate of the errors introduced by the simplified AMF strategy. The added contents are shown in red font below.

Neglecting forward model errors, Figure 22 shows the dependence of AMF on SZA, VZA, AS, wavelength, O_3 column, and SO_2 profiles. The AMFs were calculated with SCIATRAN Box-AMFs using six assumed SO_2 profiles. Here six SO_2 profiles were constructed, representing clean conditions, low, medium, and high anthropogenic SO_2 emissions, volcanic degassing with plume heights around 2 km, and volcanic eruptions with plume heights around 6 km. As shown in Fig. 22, although AMF values are generally close to 1 under typical atmospheric and surface conditions (non-ice/snow-covered), the magnitude of biases introduced by the simplified AMF approach (AMF=1 for clean regions and non-ice/snow-covered areas, while the other is AMF=2 for the ice/snow-covered areas) varies significantly with different conditions. Surface albedo is the major factor affecting AMF accuracy. For instance, AMFs can differ by up to a factor of three between AS = 0.05 and AS = 0.8. Furthermore, as shown in Eq. 3 and Fig. 22, the shape of the SO_2 vertical profile is critical for accurate AMF calculation. In extreme scenarios, such as volcanic eruptions with plume altitudes around 6 km and SO_2 columns of 120 DU, the use of a simplified AMF may lead to an overestimation of total SO_2 by a factor of 1.5–2.

3. **Reviewer 1:** Details mostly relevant to future work. I am not quite convinced that the rather long Section 5.2 should be included in the manuscript in its current form. As far as I understood, the retrievals presented in this paper were performed using the highly simplified AMF as described in Section 3.5. Section 5.2 introduces a much more advanced AMF treatment which was not used in the rest of this work (unless I missed something?). While these results do offer some insights into the uncertainties of the simple AMF implementation, the entire Section 5.2 seems to be focused on the Box-AMF which is envisaged to be much more than just means

of error analysis for the simple AMF used here. Therefore, I would suggest to either remove the whole discussion of Box-AMF, leaving it for future publications, or introduce it properly as one of the main methods, rather than just means of error analysis as the current manuscript structure seems to suggest.

Authors: Thank you for this comment.

We have moved the Box-AMF introduction from Section 5.2 to Section 3.5, introducing it properly as one of the main methods. Section 5.2 now only keeps the application to uncertainty estimates.

The Air Mass Factor (AMF) is a crucial component in SO_2 retrievals. In this study, due to the unavailability of accurate global SO_2 vertical profile data, we applied a simplified AMF strategy (AMF=1 for clean regions and non-ice/snow-covered areas, while the other is AMF=2 for the ice/snow-covered areas). While we understand the reviewer's concern regarding Section 5.2, removing the discussion of AMF errors entirely would leave readers without guidance on the OMS SO_2 AMF uncertainties and force them to consult external literature.

In the revised manuscript (P35, L593–L603 and Figure 22), to more clearly demonstrate the biases introduced by the simplified AMF strategy, we conducted AMF calculations using SCIATRAN Box-AMFs and assumed SO_2 profiles under a range of atmospheric and surface conditions. Specifically, six SO_2 profiles were constructed, representing clean conditions, low, medium, and high anthropogenic SO_2 emissions, volcanic degassing with plume heights around 2 km, and volcanic eruptions with plume heights around 6 km.

In the revised manuscript, additional text (shown in red font below) and the corresponding figure (Fig. 22 in the revised manuscript) have been added to further elaborate on these results. As shown in the Figure 22 below, we can see that the magnitude of biases introduced by the simplified AMF approach varies significantly under different conditions.

Neglecting forward model errors, Figure 22 shows the dependence of AMF on SZA,

VZA, AS, wavelength, O_3 column, and SO_2 profiles. The AMFs were calculated with SCIATRAN Box-AMFs using six assumed SO_2 profiles. Here six SO_2 profiles were constructed, representing clean conditions, low, medium, and high anthropogenic SO_2 emissions, volcanic degassing with plume heights around 2 km, and volcanic eruptions with plume heights around 6 km. As shown in Fig. 22, although AMF values are generally close to 1 under typical atmospheric and surface conditions (non-ice/snow-covered), the magnitude of biases introduced by the simplified AMF approach (AMF=1 for clean regions and non-ice/snow-covered areas, while the other is AMF=2 for the ice/snow-covered areas) varies significantly with different conditions. Surface albedo is the major factor affecting AMF accuracy. For instance, AMFs can differ by up to a factor of three between AS = 0.05 and AS = 0.8. Furthermore, as shown in Eq. 3 and Fig. 22, the shape of the SO_2 vertical profile is critical for accurate AMF calculation. In extreme scenarios, such as volcanic eruptions with plume altitudes around 6 km and SO_2 columns of 120 DU, the use of a simplified AMF may lead to an overestimation of total SO_2 by a factor of 1.5–2. Since the actual vertical distribution of atmospheric SO_2 is often difficult to get, a priori profiles from models are commonly used in AMF calculations. For regions with anthropogenic emissions, atmospheric chemistry models like GEOS-Chem and TM5 are often used to provide global SO_2 profiles for AMF calculation. The uncertainties in these profiles can also propagate into AMF calculations. In future work, we aim to incorporate high-resolution and satellite-synchronized SO_2 vertical profiles to improve the accuracy of AMF.

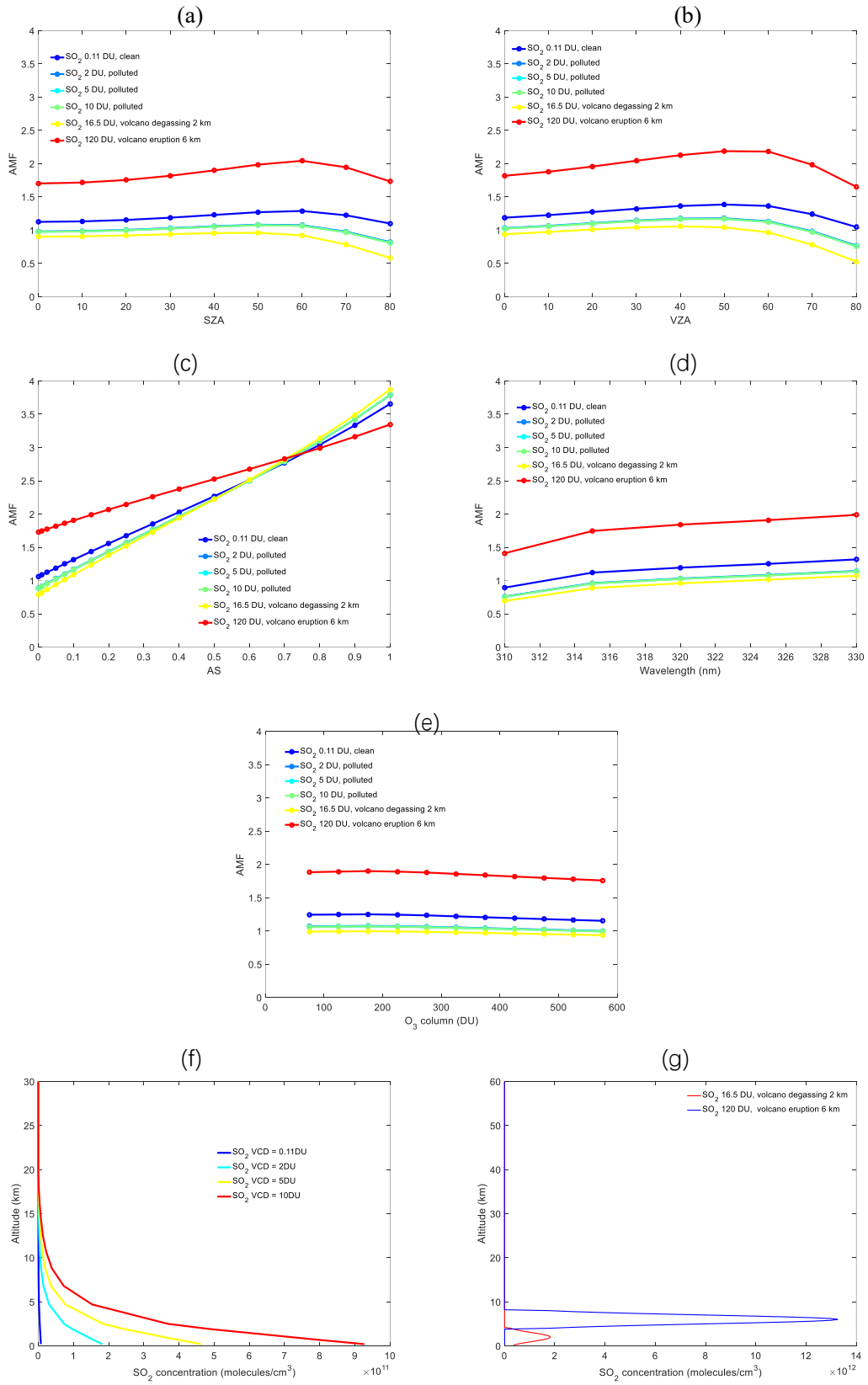


Figure: Dependence of AMF on SZA (a), VZA (b), AS (c), Wavelength (d), O_3 column (e), and SO_2 profiles. AMFs are calculated with SCIATRAN Box-AMFs using assumed SO_2 profiles. (f) Assumed SO_2 profiles corresponding to clean conditions, low, medium, and high anthropogenic SO_2 emissions; (g) Assumed SO_2

profiles corresponding to volcanic degassing with plume heights around 2 km, and volcanic eruption with plume heights around 6 km. The default SCIATRAN settings for Box-AMF calculation are as follows: wavelength=320 nm, clear sky, HS=0 km, O_3=275 DU, AS=0.05, SZA=32.9°, VZA=0°, RAA=0°.

4. **Reviewer 1:** Amount of figures. The manuscript contains a number of multi-panel figures that result from repeating the same analysis on different areas (or different spectral windows, etc.). This results in a large number of multi-panel figures, not all of which are adequately discussed in the main text (e.g. specific comment #1). I would suggest to remove some of the figures (or panels in the figures)

   Authors: Thanks for this comment. The intention of the authors was to demonstrate the reliability of the OMS SO_2 results across different times and locations, as we were concerned that a limited number of temporal or spatial cases might not fully reflect the reliability of the OMS SO_2 retrievals. We therefore prefer to keep the figures in the manuscript.

5. **Reviewer 1:** Some figures do not conform to journal standards. Some figures span multiple pages and have no labels for panels. As far as I know, this does not conform to journal standards.

   Authors: Thank the reviewer for pointing out this issue. In the revised manuscript, we have adjusted the figure layout to ensure, as much as possible, that each figure is presented on a single page rather than spanning multiple pages, although this could not be achieved for all figures. In addition, we have added panel labels to the relevant figures (e.g., Figure 21). Given the large number of figures, all corresponding modifications can be found in the revised manuscript. Any remaining nonconformities with journal standards will be discussed with Copernicus before publication.

6. **Reviewer 1:** Figure 2: Retrievals from the spectral windows of 325–335 nm and 360–390 nm appear to be complete failures. The authors should comment on why

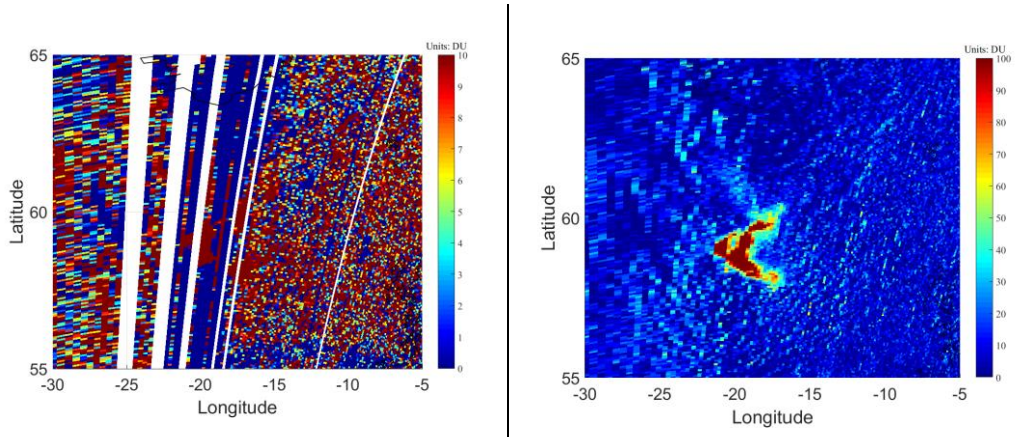
they failed, which would be far more useful than explicitly showing clearly unphysical retrieval results in (already overloaded) figures 2 and 3. The statement that a lot of NaN values were produced by the retrieval is very uninformative in this case.

Authors: Thank you for this comment and suggestion.

In the previous version of the manuscript, the SO_2 retrievals in these windows showed a large number of missing data. This was because the minimum value in the retrieval code had been set to −100 DU, and values lower than this were treated as outliers and assigned as Nan. After background offset correction, this led to a large number of missing values in the 325–335 nm and 360–390 nm retrievals over the Sundhnúkur volcano on August 23, 2024. In the revised manuscript, considering the broader variability of retrievals in the 325–335 nm and 360–390 nm windows, we reset the minimum value to −4000 DU. As a result, the missing data problem in Figure 2 has been largely eliminated. It should be noted that in the revised manuscript, we have also updated the 325–335 nm and 360–390 nm retrievals in Figure 2 (P8) by applying a different color scale range (0–100 DU and 0-400 DU) in order to better present the retrieval results (as shown in the figures below). In addition, we have updated the 325–335 nm and 360–390 nm retrievals in Figure 3 (P9) correspondingly.

We would like to emphasize that not only the SO_2 retrievals from the 325–335 nm and 360–390 nm windows in volcanic regions are generally higher than those from the 312–326 nm window, but also their standard deviations are quite large, even over clean and homogeneous oceanic regions.

325-335 nm



360–390 nm

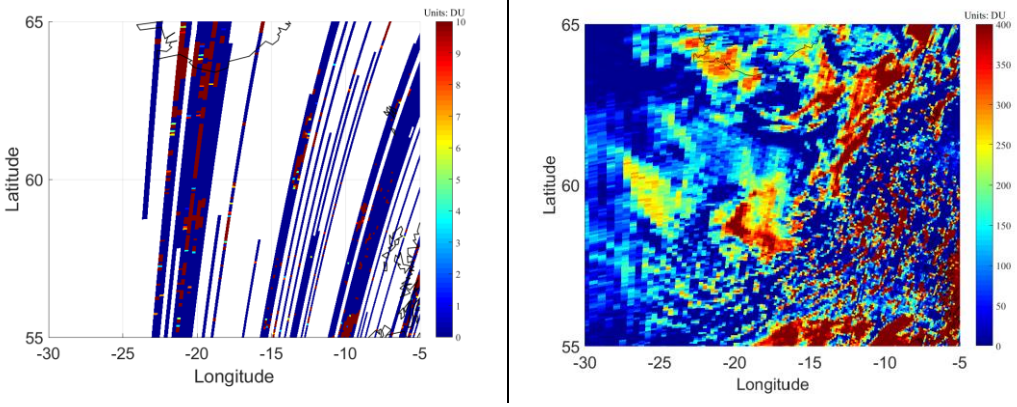


Figure: SO_2 retrievals from 325–335 nm and 360–390 nm fitting windows over the Sundhnúkur volcano on August 23, 2024 (OMS orbit 20240823_1036). DU=Dobson Units, 1 DU=2.69×10^{16} molecules/cm^2. The left panel shows the figures from the previous version of the manuscript, while the right panel shows the updated figures from the revised manuscript.

7. **Reviewer 1:** L114: "ISRF exhibits a flat top". Ideally, this should be accompanied by a figure or reference, so that the reader can get a sense of the nature and extent of the problem.

Authors: Thank you for this comment. In the paper "Preflight Spectral Calibration of the Ozone Monitoring Suite-Nadir on FengYun-3F Satellite" by

Wang Qian (2024), relevant figures and descriptions illustrating the issue of "ISRF exhibits a flat top" are provided. A screenshot of the figure from Wang's paper is shown below for reference. We have also added the corresponding citation in the revised manuscript.

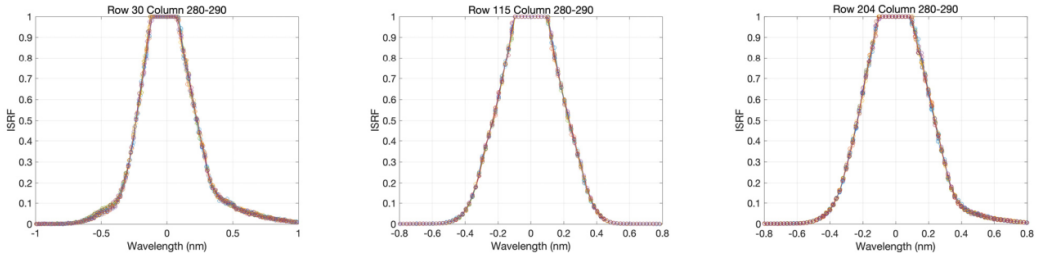


Figure 9. Samples from 11 spectral pixels in the VIS band are combined to give a single set of data, with a higher resolution shown by the colored circles. (**a**) Row 30; (**b**) row 119; (**c**) row 204.

8. **Reviewer 1:** L83-95: This paragraph contains a list of eight points. I think using numbers and/or special formatting would be more helpful to the reader than words like "seventhly".

   Authors: We agree with this suggestion. In the revised manuscript (L84-97), we have added numbering (e.g., (1), (2), …) to the eight points to make the paragraph clearer and more reader-friendly.

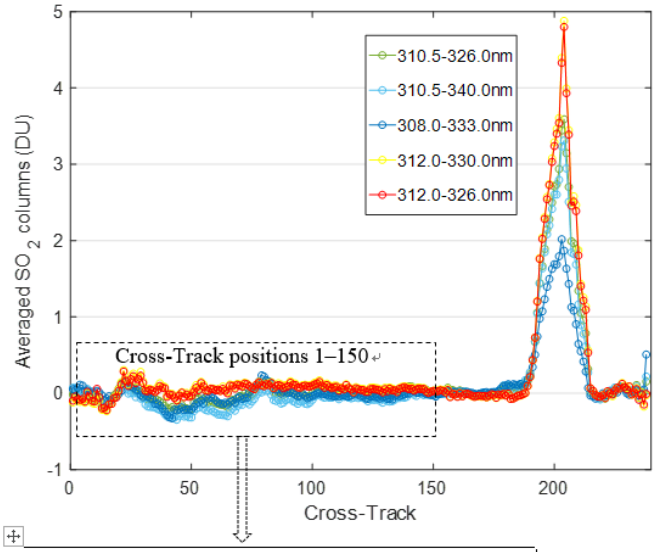
9. **Reviewer 1:** L133: Present tense here would be more consistent with the rest of the paragraph.

   Authors: This has been corrected in the revised manuscript (L134).

10. **Reviewer 1:** Figure 4 caption: I would suggest not to provide so many numerical values in a caption, but rather present them in a table, either as part of Figure 4 or separately.

    Authors: We agree with this suggestion. In the revised manuscript, we have presented the standard deviations and means of the different fitting windows as part of Figure 4, and removed the numerical descriptions from the caption. A screenshot

of the revised Figure 4 is attached here.



Fitting window (nm)	Standard deviation (DU)	Mean value (DU)
312–326	0.08	0.04
310.5–326	0.09	-0.03
310.5–340	0.1	-0.1
308–333	0.09	-0.04
312–330	0.09	0.06

Figure 4: Row-averaged SO_2 retrievals from OMS rows 700 to 870 of orbit 20240823_1036 by using different spectral fitting windows.

11. **Reviewer 1:** L343: The comparison of results from [. . .].

Authors: This has been corrected in the revised manuscript.

12. **Reviewer 1:** L503: Replace “difficult to be monitored and calibrated” with “difficult to monitor and calibrate”.

Authors: Thank the reviewer for pointing out this language issue. We have revised the sentence in the revised manuscript (L496), replacing “difficult to be monitored and calibrated” with “difficult to monitor and calibrate”.

Answer to Referee #2

1. **Reviewer 2:** The authors refer to this as the "first version" of the OMS SO_2. It would be useful to clearly state this in the title or abstract to manage reader expectations about its development stage and limitations.

   Authors: Thanks for this recommendation. We agree with this comment. The title of the manuscript has been revised to "First results of SO_2 columns from FY3F/OMS instrument observations" to reflect the limitations of the current version and to indicate the early stage of its development. In addition, at the end of the abstract, we have included the sentence "This study is the first to present SO_2 retrievals from FY3F/OMS observations" to further clarify this point.

2. **Reviewer 2:** The paper (L95-96) recommends filtering OMS data for SZA < 70 °, near-nadir, and cloud-free pixels. Were these recommended filters (including cloud-free) applied to the OMS data before comparison with TROPOMI? Please clarify the exact filtering used, especially given the statement about no cloud products.

   Authors: We are sorry for the unclear statement. When comparing with TROPOMI, we filtered OMS SO_2 retrievals greater than –10 DU, but did not apply filtering for pixels with cloud coverage or large SZA. In other words, OMS SO_2 retrievals with cloud coverage or large SZA were included in the comparison with TROPOMI SO_2 results as long as they were greater than –10 DU. To clarify this point, we have added the following description at the end of the first paragraph in Section 4 (L290-292) of the revised manuscript (in red font):

   "It should be noted that, in the comparison between OMS and TROPOMI SO_2 results, no filtering for cloud or SZA was applied to the OMS SO_2 data; instead, all retrievals greater than –10 DU were selected for the comparison."

3. **Reviewer 2:** The paper mentions the use of the TSIS HSRS hybrid solar

reference spectrum instead of OMS L1 solar irradiance due to degradation. Could this lead to systematic biases in the long term? What plans are in place to account for this degradation?

Authors: For FY-3F/OMS SO_2 retrievals, it is essential to obtain accurate solar irradiance data for calculating the TOA reflectance (radiance/irradiance). However, after one year in orbit, we found that the intensity of OMS irradiance at the shorter wavelength of 317 nm had decreased by about 8.83%, while at the longer wavelength of 331 nm it had decreased by about 6.07%. This degradation in OMS solar irradiance measurements is one of the main causes of the systematic low bias and striping observed in OMS trace gases retrievals, especially for ozone products.

In addition to the solar irradiance issue, OMS radiance measurements also suffer from certain calibration errors. The combined effect of these issues has led to systematic biases in the retrievals of OMS trace gases, as well as striping and left–right asymmetry problems. Given these multiple issues, we prioritized addressing the striping and left–right asymmetry problems by using the TSIS HSRS hybrid solar reference spectrum instead of the OMS L1 solar irradiance for total SO_2 retrievals. To date, the work on correcting the calibration errors in OMS L1 radiance and irradiance is still ongoing.

Using the TSIS HSRS hybrid solar reference spectrum can mitigate the effects of viewing-angle dependence and degradation on OMS SO_2 retrievals. However, because the solar reference spectrum does not include the instrument's specific characteristics, it may introduce systematic overestimation or underestimation in SO_2 column retrievals, and these errors could change over time. Such systematic biases introduced by the TSIS HSRS can be partially reduced through background offset correction, but they cannot be completely eliminated.

For the retrieval of atmospheric composition from OMS, using the OMS measured irradiance data can reduce some errors, and our preliminary plan for correcting the degradation of OMS irradiance in the future is to first correct the degradation in

the ultraviolet region, and then analyze the changes in the OMS diffuser based on the existing OMS irradiance observations for subsequent improvements.

4. **Reviewer 2:** The results in Figure 2 from using the 325–335 nm and 360–390 nm windows alone appear unconvincing. While these windows might be noisy in isolation, their value (as seen in TROPOMI's algorithm) lies in their combined use with stronger windows to avoid saturation at high SO_2 Since OMS shows saturation and underestimation for high SO_2, have the authors explored a multi-window fitting strategy to improve this?

   Authors: Thank you for this comment and suggestion. In general, it is well known that in volcanic eruption cases with high SO_2 columns, retrievals using strong SO_2 absorption bands are often subject to saturation effects. By contrast, retrievals using weaker SO_2 absorption bands in the longer UV range can effectively mitigate this saturation. The reviewer's suggestion is therefore very good and is indeed in line with our future plan for the OMS SO_2 product.

   Our preliminary strategy for future OMS SO_2 retrievals during volcanic eruptions is as follows. (1) Initial retrieval: Use the 312–326 nm window to perform a global SO_2 column retrieval. (2) Flagging potential saturation: Pixels with SO_2 columns above a certain threshold (e.g., >50 DU) are flagged as potentially affected by saturation at high SO_2. For these flagged pixels, additional retrievals are conducted using the 325–335 nm and 360–390 nm windows. (3) Assessment and replacement: The SO_2 values from the 325–335 nm and 360–390 nm windows are assessed to determine whether they fall within a reasonable range and exceed the corresponding 312–326 nm values. If both conditions are satisfied, the 312–326 nm retrievals are replaced by the maximum of the 325–335 nm or 360–390 nm values; otherwise, the original 312–326 nm retrievals are retained. In the revised manuscript, an abbreviated version of this plan has been added to Section 6 (L686-691), as indicated in red font below.

   Our preliminary strategy for such retrievals is as follows. First, a global SO_2

column retrieval is performed using the 312–326 nm window. Pixels exceeding a threshold (e.g., >50 DU) are flagged as potentially saturated. For these flagged pixels, additional retrievals are conducted using the 325–335 nm and 360–390 nm windows. If the new retrievals are within a reasonable range and exceed the 312–326 nm results, the maximum value from the 325–335 nm or 360–390 nm windows replaces the original retrieval; otherwise, the 312–326 nm retrieval is retained.

In addition, it should be noted that in the revised manuscript, we have also updated the 325–335 nm and 360–390 nm retrievals in Figure 2 (P8) by applying a different color scale range (0–100 DU and 0-400 DU) in order to better present the retrieval results (as shown in the figure below). In the previous version of the manuscript, the SO_2 retrievals in these windows showed a large amount of missing data. This was because the minimum value in the retrieval code had been set to −100 DU, and values lower than this were treated as outliers and assigned as Nan. After background offset correction, this led to a large number of missing values in the 325–335 nm and 360–390 nm retrievals over the Sundhnúkur volcano on August 23, 2024. In this revised manuscript, considering the broader variability of retrievals in the 325–335 nm and 360–390 nm windows, we reset the minimum value to −4000 DU. As a result, the missing data problem in Figure 2 has been largely eliminated. In addition, we have correspondingly updated the 325–335 nm and 360–390 nm retrievals in Figure 3 (P9).

We would like to emphasize that not only the SO_2 retrievals from the 325–335 nm and 360–390 nm windows in volcanic regions are generally higher than those from the 312–326 nm window, but also their standard deviations are quite large, even over clean and homogeneous oceanic regions.

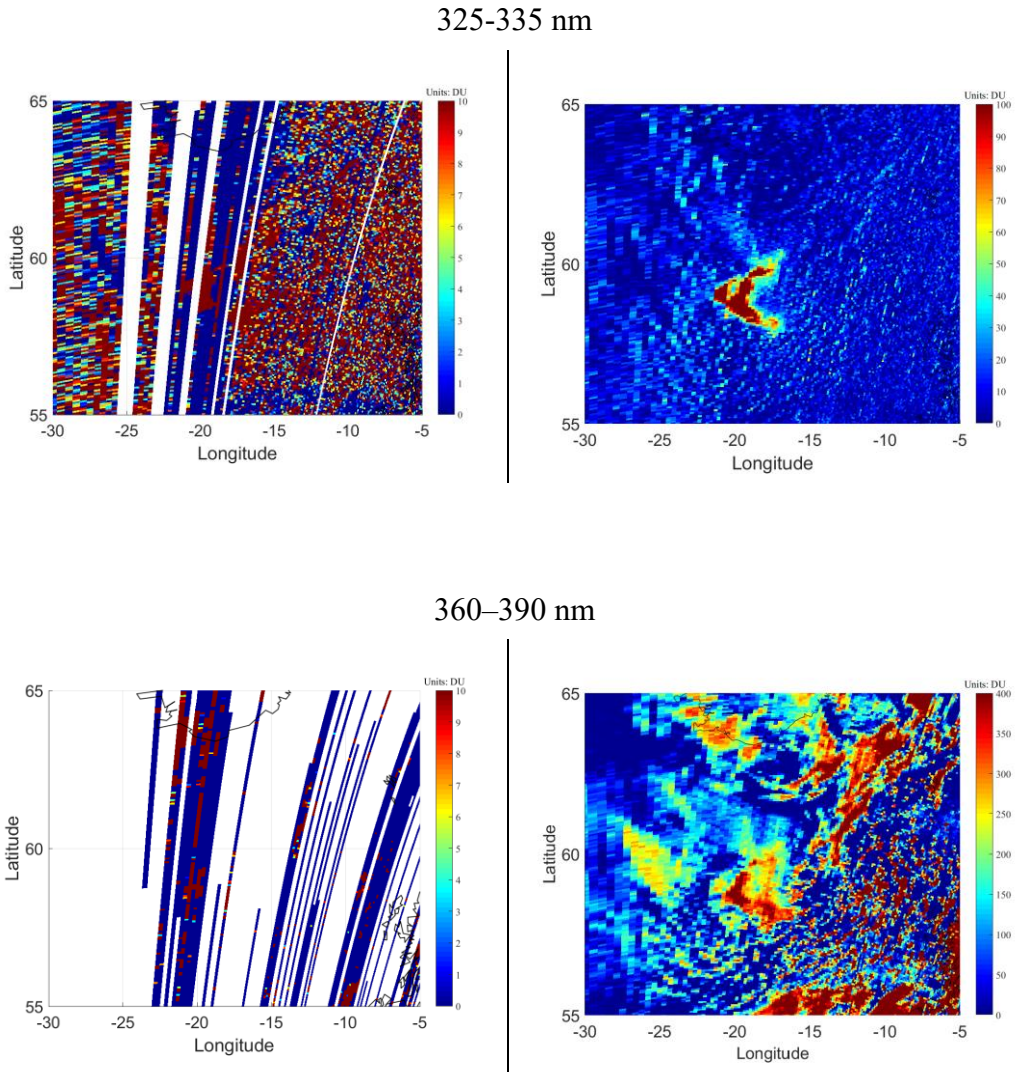


Figure: SO_2 retrievals from 325–335 nm and 360–390 nm fitting windows over the Sundhnúkur volcano on August 23, 2024 (OMS orbit 20240823_1036). DU=Dobson Units, 1 DU=2.69×10^{16} molecules/cm^2. The left panel shows the figure from the previous version of the manuscript, while the right panel shows the updated figure in the revised manuscript.

5. **Reviewer 2:** Section 3.3.2 states a single fixed Ring spectrum was used. However, Figure 19 clearly shows the Ring spectrum varies significantly with atmospheric and viewing conditions. Please clarify this apparent contradiction: was a single fixed Ring spectrum or a variable one used? If a single fixed spectrum was used, the significant error introduced by ignoring these variations should be quantitatively discussed and the actual Ring spectrum settings should be reflected

in Table 3.

Authors: Thanks for pointing this out, and sorry for the inaccurate description in the original manuscript. In this study, one single Ring spectrum simulated with the SCIATRAN model was used to retrieve OMS SO_2 for all pixels. As we know, the Ring effect in the UV wavelength range is a non-negligible component of the DOAS SCD fitting process. Initially, we generated a lookup table for the 310–330 nm wavelength range using the SCIATRAN model, considering variations in SZA, VZA, O_3 column, and AS to achieve more accurate retrievals. However, this lookup table is large and significantly increased computational cost for OMS SO_2 retrievals. Moreover, the OMS SO_2 retrievals indicate that the impact of the Ring spectrum variability on SO_2 retrievals is relatively small for high SO_2 concentrations, due to the weak correlation between the Ring spectrum and the satellite TOA reflectance. Therefore, in this study, we adopted the approach of using a single Ring spectrum for all OMS pixels.

To clarify this point, in the revised manuscript, the original text in Section 5.1 (blue font) has been modified and replaced with the following text (red font).

In the UV wavelength band, the Ring effect is a non-negligible part of the DOAS SCD fitting process. In this study, one single Ring spectrum simulated with the SCIATRAN model was used to retrieve OMS SO_2 for all pixels. As shown in Fig. 19, the Ring spectrum convolved with the OMS ISRF varies with SZA, VZA, O_3 column and AS within the 310–330 nm wavelength range. The variation of the Ring spectrum with RAA is negligible and is not presented in this study. The mean percentage change in the Ring spectrum is approximately 27.67% (absolute deviation: 0.0025) as the O_3 column varies from 175 DU to 575 DU, 47.34% (absolute deviation: 0.0048) as the AS varies from 0 to 1, 24.18% (absolute deviation: 0.0021) as the VZA varies from 0° to 75°, and 45.29% (absolute deviation: 0.0047) as the SZA varies from 0° to 80°. Although the Ring spectrum varies significantly with SZA, VZA, O_3 column, and AS within the 310–330 nm

wavelength range, its impact on SO_2 retrievals is relatively small for high SO_2 concentrations due to the weak correlation between the Ring spectrum and the satellite TOA reflectance. However, for low SO_2 concentrations, the influence is non-negligible.

In the UV wavelength band, the Ring effect is a non-negligible part of the DOAS SCD fitting process. As shown in Fig. 19, the Ring spectrum convolved with the OMS ISRF varies with SZA, VZA, O_3 column and AS within the 310–330 nm wavelength range. The variation of the Ring spectrum with RAA is negligible and is therefore not presented. The mean percentage change in the Ring spectrum is approximately 27.67% (absolute deviation: 0.0025) as the O_3 column varies from 175 DU to 575 DU, 47.34% (absolute deviation: 0.0048) as the AS varies from 0 to 1, 24.18% (absolute deviation: 0.0021) as the VZA varies from 0° to 75°, and 45.29% (absolute deviation: 0.0047) as the SZA varies from 0° to 80°. However, in this study, a single Ring spectrum simulated with the SCIATRAN model was used to retrieve OMS SO_2 for all pixels. The reason is that although the Ring spectrum varies significantly with SZA, VZA, O_3 column, and AS within the 310–330 nm wavelength range, the impact of Ring spectrum variability on SO_2 retrievals is relatively small due to the weak correlation between the Ring spectrum and the satellite TOA reflectance, especially in the case of volcanic eruptions with high SO_2 concentrations. Moreover, using Ring spectra that vary with SZA, VZA, O_3 column, and AS within the 310–330 nm wavelength range require constructing a large lookup table, which would significantly increase computational cost for OMS SO_2 column retrievals.

6. **Reviewer 2:** What improvements are considered at orbit edges in a future update?

Authors: At orbit edges, the retrieval accuracy is relatively low due to large viewing angles. This issue is also present in other satellite instruments such as TROPOMI and GOME-2. Ongoing improvements to the OMS L1 data are being made to reduce striping and asymmetry in the calibration to improve SO_2 retrieval

accuracy. Despite several attempted corrections, no effective strategy for the edge pixels is currently available and therefore we recommend that users take caution when using them. Nevertheless, the different overpass time and spatial coverage of OMS compared to other satellite instruments allow OMS to provide complementary observations, particularly for TROPOMI orbit edge pixels and in the tropical orbit gap regions.

7. **Reviewer 2:** For the Persian Gulf, the 0.5-0.6 correlations are not particularly high. What does this imply regarding the agreement between the OMS and TROPOMI products, and what factors contribute to these differences?

   Authors: We agree that in the Persian Gulf region, the correlation coefficient of 0.5–0.6 is not high. We consider that this may be related to the following factors: (1) the different overpass times of OMS (morning, about 10:00 am) and TROPOMI (afternoon, about 13:30 pm), during which atmospheric conditions may have changed, leading to discrepancies; (2) the differences in viewing geometries between the two instruments; (3) Random noise and uncertainties from background correction, which are relevant for low SO2 scenarios, such as over the Persian Gulf, in data from both instruments and lead to scatter in the order of several DU.

   In addition, the differences among the three datasets (OMS SO_2, TROPOMI DOAS SO_2, and TROPOMI COBRA PBL SO_2) vary over time: on August 23, 2024, OMS SO_2 retrievals were higher than those from both TROPOMI DOAS and TROPOMI COBRA, whereas on November 12, 2024, OMS SO_2 retrievals were lower than those of TROPOMI DOAS and TROPOMI COBRA.

8. **Reviewer 2:** How much does the difference of the overpass time between OMS and TROPOMI contribute to the observed differences in SO_2 retrievals, particularly in dynamic regions like volcanic or strong anthropogenic plumes?

   Authors: Thanks for this comment. The difference in overpass time between OMS and TROPOMI can contribute to the observed discrepancies in SO_2 retrievals,

particularly in regions with rapidly changing SO_2 concentrations, such as those affected by volcanic or strong anthropogenic plumes. However, isolating and quantifying the sole impact of overpass time requires accurate knowledge of the temporal evolution of SO_2 emissions and meteorological conditions within the time window between the OMS and TROPOMI overpasses, which is challenging for the current study. Our focus here is therefore on the overall differences between OMS and TROPOMI.

Nevertheless, in high-latitude regions, OMS and TROPOMI often provide multiple orbital coverages of the same area within a short period, which offers opportunities to see their SO_2 retrieval differences under relatively small overpass time differences. For example, as shown in Fig. 12, the Sundhnúkur volcano in the Northern Hemisphere high latitudes was captured by multiple OMS orbits, enabling near-simultaneous comparisons with TROPOMI. Compared with OMS orbit 20240823_1036, OMS orbit 20240823_1217 has a local overpass time closer to that of the TROPOMI orbit 20240823T125304 in the volcanic region, resulting in more consistent SO_2 retrievals between OMS and TROPOMI.

Future work employing high-temporal-resolution satellite observations (e.g., GEMS, TEMPO) or chemical transport model simulations could provide more detailed insights into the role of overpass time differences.

9. **Reviewer 2:** The paper highlights the simplified AMF approach as introducing significant errors and large uncertainties. Could the authors elaborate further on the estimated magnitude of biases introduced by these simplifications across a range of atmospheric and surface conditions? Furthermore, to enhance the accuracy and robustness of the product, particularly in complex regions like urban areas or near industrial sources, what are the plans for incorporating a more physically-based AMF calculation?

Authors: Thank you for this comment. Our responses are as follows:

1) **Comment:** The paper highlights the simplified AMF approach as introducing significant errors and large uncertainties. Could the authors elaborate further on the estimated magnitude of biases introduced by these simplifications across a range of atmospheric and surface conditions?

Answer: Thanks for this revision suggestion. We agree with the reviewer that using a simplified AMF (AMF=1 for clean regions and non-ice/snow-covered areas, while the other is AMF=2 for the ice/snow-covered areas) to convert slant columns to vertical columns introduces errors, particularly under complex atmospheric and surface conditions. The AMF value is influenced by multiple factors, including wavelength, solar zenith angle (SZA), viewing zenith angle (VZA), relative azimuth angle (RAA), surface albedo (AS), terrain height (HS), O_3 column, SO_2 vertical profile shape, and cloud fraction and altitude.

In the revised manuscript, we conducted AMF calculation by using SCIATRAN Box-AMFs and SO_2 profiles under a range of atmospheric and surface conditions to analyze the biases between simplified AMF and physically-based AMF. Specifically, we constructed six SO_2 profiles corresponding to clean conditions, low, medium, and high anthropogenic SO_2 emissions, volcanic degassing with plume heights around 2 km, and volcanic eruption with plume heights around 6 km.

Neglecting forward model errors, the figure below shows that the magnitude of biases introduced by the simplified AMF approach varies significantly under different conditions. In the revised manuscript, additional text (L593-603, shown in red font below) and the corresponding figure (Fig. 22 in the revised manuscript) have been added to further elaborate on these results.

Neglecting forward model errors, Figure 22 shows the dependence of AMF on SZA, VZA, AS, wavelength, HS, O_3 column, and SO_2 profiles. The AMFs were calculated with SCIATRAN Box-AMFs using assumed SO_2 profiles. As

shown in Fig. 22, although AMF values are generally close to 1 under typical atmospheric and surface conditions (non-ice/snow-covered), the magnitude of biases introduced by the simplified AMF approach (AMF=1 for clean regions and non-ice/snow-covered areas, while the other is AMF=2 for the ice/snow-covered areas) varies significantly with different conditions. Surface albedo is the major factor affecting AMF accuracy. For instance, AMFs can differ by up to a factor of three between AS = 0.05 and AS = 0.8. Furthermore, as shown in Eq. 3 and Fig. 22, the shape of the SO_2 vertical profile is critical for accurate AMF calculation. In extreme scenarios, such as volcanic eruptions with plume altitudes around 6 km and SO_2 columns of 120 DU, the use of a simplified AMF may lead to an overestimation of total SO_2 by a factor of 1.5–2. Since the actual vertical distribution of atmospheric SO_2 is often difficult to get, a priori profiles from models are commonly used in AMF calculations. For regions with anthropogenic emissions, atmospheric chemistry models like GEOS-Chem and TM5 are often used to provide global SO_2 profiles for AMF calculation. The uncertainties in these profiles can also propagate into AMF calculations. In future work, we aim to incorporate high-resolution and satellite-synchronized SO_2 vertical profiles to improve the accuracy of AMF.

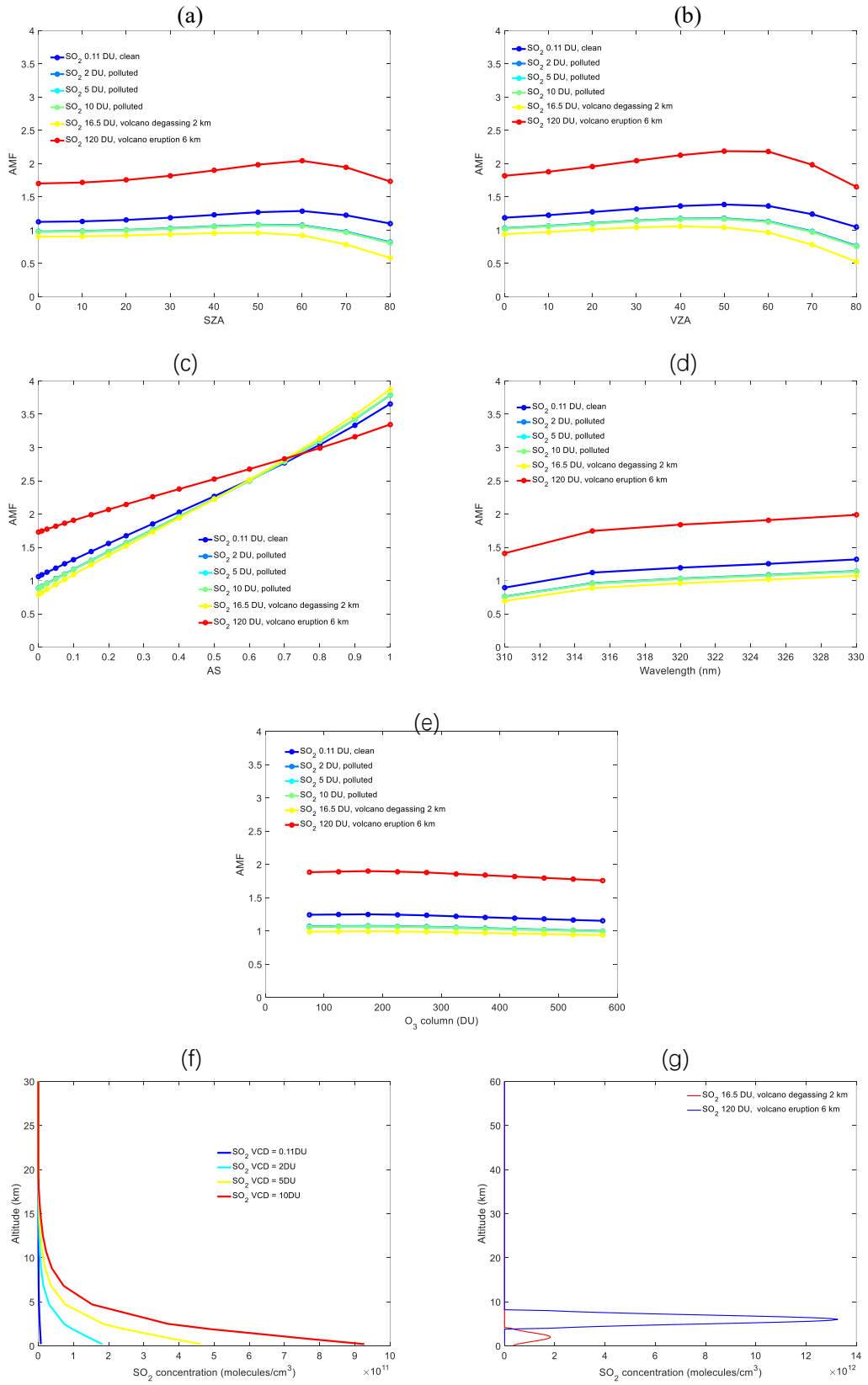


Figure: Dependence of AMF on SZA (a), VZA (b), AS (c), Wavelength (d), O_3 column (e), and SO_2 profiles. AMFs are calculated with SCIATRAN Box-AMFs using assumed SO_2 profiles. (f) Assumed SO_2 profiles corresponding to

clean conditions, low, medium, and high anthropogenic SO_2 emissions; (g) Assumed SO_2 profiles corresponding to volcanic degassing with plume heights around 2 km, and volcanic eruption with plume heights around 6 km. The default SCIATRAN settings for Box-AMF calculation are as follows: wavelength=320 nm, clear sky, HS=0 km, O_3=275 DU, AS=0.05, SZA=32.9°, VZA=0°, RAA=0°.

2) **Comment:** Furthermore, to enhance the accuracy and robustness of the product, particularly in complex regions like urban areas or near industrial sources, what are the plans for incorporating a more physically-based AMF calculation?

Answer: We thank the reviewer for their comment on the importance of a more physically-based AMF calculation, especially for improving retrievals in complex regions such as urban or industrial areas. Until now, we have been exploring collaboration with other research teams to obtain more accurate SO_2 vertical profiles representative of regions with anthropogenic emissions.

In future work, once high-resolution and accurate SO_2 profiles become available, we plan to develop an offline product. This offline product will employ more accurate AMFs to reprocess historical data and thereby improve the accuracy of the retrievals, particularly in regions with strong or anthropogenic emissions.

10. Reviewer 2: The current retrieval does not account for cloud and aerosol effects, which can introduce significant biases, particularly for boundary layer SO_2. What are the plans to integrate cloud and aerosol effects into future updates of the retrieval algorithm?

Authors: Thanks for this valuable comment. We agree that neglecting cloud and aerosol effects may introduce significant biases, especially for boundary layer SO_2 retrievals. Our plans for future updates are as follows:

(1) Short-term strategy: In the absence of reliable OMS cloud and aerosol products, we recommend users to primarily use clear-sky pixels. A simple radiance-based screening will be implemented to flag cloud-contaminated pixels in the SO_2 output file. In addition, cloud products from other instruments (e.g., TROPOMI, OMI, MODIS, or MERSI) may be incorporated as auxiliary inputs to reduce cloud-related uncertainties.

(2) Long-term strategy: Once mature OMS cloud and aerosol products become available, they will be integrated into the retrieval algorithm to enable accurate AMF calculations. This will significantly improve the accuracy of SO_2 retrievals, particularly for the boundary layer. Furthermore, quality control flags will be refined (e.g., pixels with cloud fraction > 0.3 or strong aerosol events will be filtered or flagged as low quality) to facilitate more reliable usage of the OMS SO_2 product.