



Atmospheric Organosulfate Formation Regulated by Continental Outflows and Marine Emissions over East Asian Marginal Seas

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Abstract. Organosulfates (OSs) represent an unrecognized fraction and a potentially important source of marine organic aerosols. Based on shipboard observations over East Asian marginal seas, we characterized OSs in marine aerosols during
20 spring, summer, and autumn. The C₂–C₃ OSs and isoprene-/monoterpenes-derived OSs were quantified using synthesized standards. The total quantified OS concentrations ranged from 4.5 to 109.1 ng/m³, contributing 0.1%–3.2% of the mass concentration of marine organic aerosols. The highest OS concentrations, dominated by C₂–C₃ OSs and isoprene-OSs, were observed in summer, which surpassed the abundance of methane sulfonic acid, a key component in climate regulation by oceanic phytoplankton sulphur emissions. Abundant OS formation in summer was mainly attributed to the increased
25 isoprene emissions from the ocean. During the spring and autumn cruises, transported continental pollutants resulted in the higher fraction of monoterpene-derived (nitrooxy-)OSs, as well as the elevated OS concentrations over regions surrounded by the continent. This work highlights the joint effects of marine emissions and continental outflows on the formation and distribution of atmospheric OSs over marginal seas.

30 1 Introduction

Marine atmospheric aerosols play a vital role in climate change through influencing cloud formation and solar radiative balance (Li et al., 2022). Marine phytoplankton could produce dimethylsulfide (DMS), which further forms methane



35 sulphonic acid (MSA) or sulfate aerosols and then regulates the cloud condensation nuclei (CCN) formation and climate in the marine boundary layer (Andreae and Rosenfeld, 2008; Kettle and Andreae, 2000; Kloster et al., 2006). This is named CLAW hypothesis, proposed by Charlson et al. (1987) and Ayers et al. (1991). While following observation evidence and modelling studies indicated that CCN formation in marine atmospheres is far more complex than was recognized by the CLAW hypothesis (Quinn and Bates, 2011). This is mainly attributed to the unknown organic fractions in marine aerosols, including those primarily emitted by sea spray and secondarily formed organic aerosols (SOA) via the oxidation of volatile organic compounds (VOCs).

40 Traditional SOA tracers, including those from the oxidation of isoprene and monoterpene, etc., could explain only <10% of marine organic aerosols (Fu et al., 2011; Guo et al., 2020). Majority of the marine SOA components remain unknown till now. Abundant isoprene could be emitted from the ocean, and isoprene SOA has been proved to be one of the most important fractions in marine organic aerosols (Hu et al., 2013). Another important SOA formation pathway from isoprene and monoterpene oxidation is facilitated by acidic sulfate particles under high humidity conditions, with organosulfates (OSs) as the products (Brüggemann et al., 2020). Sulfate aerosols could be readily formed via DMS oxidation and are generally abundant over various marine environments (Andreae, 1990; Li et al., 2018; Yan et al., 2024). Recent studies have indicated the existence and importance of organic sulfur compounds, including OSs, in marine aerosols (Bao et al., 2018; Ye et al., 2021).

Atmospheric OSs constitute a large portion of organic aerosols (OA) in environments with substantial interactions of biogenic and anthropogenic emissions (Hettiyadura et al., 2019; Meade et al., 2016; Surratt et al., 2008; Wang et al., 2018). A recent cruise observation over Asian marginal seas suggested that OSs derived from isoprene and monoterpenes could contribute about 7% of the OA mass concentration (Wang et al., 2023b). Wang et al (2023) also indicated that isoprene/monoterpene-derived OSs could surpass the traditionally identified SOA tracers generated from isoprene or monoterpene oxidation (e.g., methylglyceric acid, alkene triols, hydroxyglutaric acid, pinic acid etc.). While these OS compounds derived from isoprene or monoterpenes were rarely detected in marine aerosols collected at a remote island site located in the Southern Ocean or in the southern Indian Ocean (Claeys et al., 2010; Cui et al., 2019). These could be attributed to the low biogenic VOCs emission/flux or the degradation of OSs over long-term storage (Claeys et al., 2010; Cui et al., 2019).

The existence or abundance of atmospheric OSs in marine aerosols have not been well evaluated or quantified till now (Hawkins et al., 2010; Wang et al., 2023b; Ye et al., 2021), which limited the understanding on their formation processes or their roles in the sulfur cycle and aerosol climate effects in marine atmospheres. In this study, atmospheric OSs over East Asian marginal seas were quantified using synthesized OS standards. We characterized the particulate OSs derived from



isoprene and monoterpenes, and investigated their spatial distributions, seasonal variations, as well as the dominant environmental factors of OS formation. Our results suggested that, over marginal seas, the spatiotemporal distribution of OS abundance and composition was dependent on the relative importance of marine emissions and continental outflows. This work highlights the vital roles of OSs in altering the sulfur cycle in marine boundary layer, and further studies in open ocean are needed to understand the influence of OSs on the climate effects of marine aerosols.

2 Methods

2.1 Cruise observation and sample collection

Marine aerosol samples were collected during summer (16 July–26 July) and autumn (21 October–2 November) in 2021, and during spring (14 April–25 April) in 2022 over the Yellow Sea and Bohai Sea (YBS). The fine particle ($PM_{2.5}$) and total suspended particulate (TSP) samples were collected using high-volume aerosol samplers (KB-1000, Qingdao Genstar Electronic Technology, China). The quartz fiber filters were pre-baked at 500 °C for 4.5 hours and wrapped in pre-baked aluminum foil after sampling. Aerosol samplers were placed on the top deck of the vessel “Lanhai 101”, approximately 8 m above the sea surface. Each aerosol sample was collected for 10–24 hrs, and a field blank sample was collected during each cruise.

During the observation, wind speed (WS), air temperature, and relative humidity (RH) were measured by a shipboard meteorological observatory. The *Chl-a* concentration in surface seawater was measured using a CE Turner Designs fluorometer. Concentrations of seawater isoprene were then estimated by empirical formulas based on previous studies (Ooki et al., 2015; Wang et al., 2023b). The 72 h backward trajectories of air masses from an altitude of 500 m above ground level were calculated using the HYSPLIT model (Version 5.2.1, NOAA), starting every 6 h (Fig. S1). Trajectories at the center site of the observation region were calculated to represent the air masses during each cruise over the YBS.

2.2 HPLC-MS analysis and OS quantification

An aliquot of each filter sample was extracted by methanol. The solutions were filtered using PTFE syringe filter (0.22 μ m), and evaporated to dryness under a gentle stream of N_2 gas. The dried residues were redissolved in methanol containing 0.1% formic acid (100 μ L). Organosulfate compounds were quantified using a QTRAP 4500 mass spectrometer (AB Sciex) coupled with an UHPLC system (Ultimate 3000, Thermo Scientific, DE) for the low-molecular-weight OSs, and an Exactive Plus-Orbitrap mass spectrometer (Thermo Scientific Inc.) with an UHPLC system (Ultimate 3000) for monoterpene-derived compounds. Mass spectrometry was operated using a negative-mode electrospray ionization. The monoterpene NOSs ($C_{10}H_{16}NO_7S^-$ and $C_9H_{14}NO_8S^-$) were identified in the extracted ion chromatogram mode, and other OS compounds were quantified in multiple-reaction monitoring (MRM) mode. In this work, C_2 – C_3 OSs (HAS, GAS, and LAS), isoprene-OSs



(IEPOX-OS, MAE-OS, and $C_5H_7O_7S^-$), monoterpene-OSs, and nitrooxy-OSs (NOSs) were quantified using synthesized OS standards (Table S1) (Wang et al., 2018).

Chromatographic separation of the low-molecular-weight OSs, including C_2 – C_3 OSs and isoprene-OSs/NOSs, was
95 optimized using a BEH Amide column (2.1 mm×100 mm, 1.7 μ m, Waters, USA) equipped with a pre-column. Hydrophilic interaction liquid chromatography separation is an accurate analytical method for quantifying the low-molecular-weight OSs (Hettiyadura et al., 2015). The injection volume was 2.0 μ L. The column was maintained at 35°C. Mobile eluents were solvent A: ammonium acetate buffer (10 mM, pH 9) in ultrapure water and solvent B: 10 mM ammonium acetate buffer (10 mM, pH 9) in acetonitrile/water (95:5). The flow rate was 0.4 mL/min at 0–2.5 min, then decreased to 0.35 mL/min from 2.5
100 to 11.5 min, and increased back to 0.4 mL/min from 11.5 to 18 min. The gradient elution was set as follows: 100% B at 0–0.4 min; reduced to 88% B at 0.4–2.4 min and maintained until 11 min; increased to 100% B at 11–11.5 min, and maintained at 100% B until 18 min to re-equilibrate the column.

Monoterpene OSs/NOSs were analyzed using an Acquity UPLC HSS T3 column (2.1 mm×100 mm, 1.8 μ m, Waters, USA) with a pre-column. The mobile eluents were solvent A (0.1% acetic acid in ultrapure water) and solvent B (0.1% acetic acid
105 in methanol) at a flow rate of 0.3 mL/min. The gradient elution procedure was performed as follows: 5% B at 0–1.5 min; increased to 54% B over 13.7 min and held for 1.0 min; then increased to 90% B over 1.8 min and held for 5 min; decreased to 5% B over 0.5 min and held for 1.5 min to re-equilibrate the column for next injection. The column temperature was maintained at 45°C, and the injection volume was 5.0 μ L.

2.3 Measurements of aerosol chemical composition

110 Elemental carbon (EC) and organic carbon (OC) in aerosol samples were quantified using a carbon analyzer (Model RT-3131, Sunset Laboratory, OR). The OA concentration was then calculated by multiplying OC by 1.6. Water-soluble cations (Na^+ , NH_4^+ , K^+ , Mg^{2+} , Ca^{2+}), anions (Cl^- , NO_3^- , SO_4^{2-}), and MSA were measured using ion chromatography systems (ICS-2100 and ICS-Aquion RFIC, Thermo Scientific). The concentrations of non-sea-salt potassium ion (nss- K^+) and non-sea-salt sulfate (nss- SO_4^{2-}) were calculated by $[K^+]-0.037\times[Na^+]$ and $[SO_4^{2-}]-0.2516\times[Na^+]$. The concentrations of $PM_{2.5}$ or TSP
115 were reconstructed by summing the concentrations of inorganic ions, OA, and EC in each aerosol sample.

3 Results and discussion

3.1 Concentration and composition of marine atmospheric OSs

The total quantified OSs and nitrooxy-OSs ranged from 4.5 to 109.1 ng/m³ in marine aerosols during the shipboard
120 observations over the YBS (Fig. 1, Table S1). The eleven quantified OS and NOS compounds contributed 0.1%–3.2% of the



OA mass concentrations over the YBS. The observed OS concentrations here were generally higher than the wintertime concentrations at inland sites, and lower than those in coastal regions (Kanellopoulos et al., 2022; Meade et al., 2016; Nguyen et al., 2014; Wang et al., 2020, 2021, 2022b) (Fig. 2). This was due to the active interactions between biogenic VOCs and sulfate aerosols under high RH conditions in coastal areas, which favored the aqueous-phase formation of OSs in the atmosphere. While the OS formation could be limited by the low biogenic VOC emissions or ambient RH in the wintertime inland environments (Wang et al., 2020). It is noted that, taking the autumn observation as an example, we compared the OS concentrations in the PM_{2.5} and the TSP samples simultaneously collected during the cruise (Fig. S2). The majority of the data points fall along the 1:1 line (Fig. S2). The presence of OSs is dominant in fine particles, and thus our further discussion is focused on the results of the PM_{2.5} samples.

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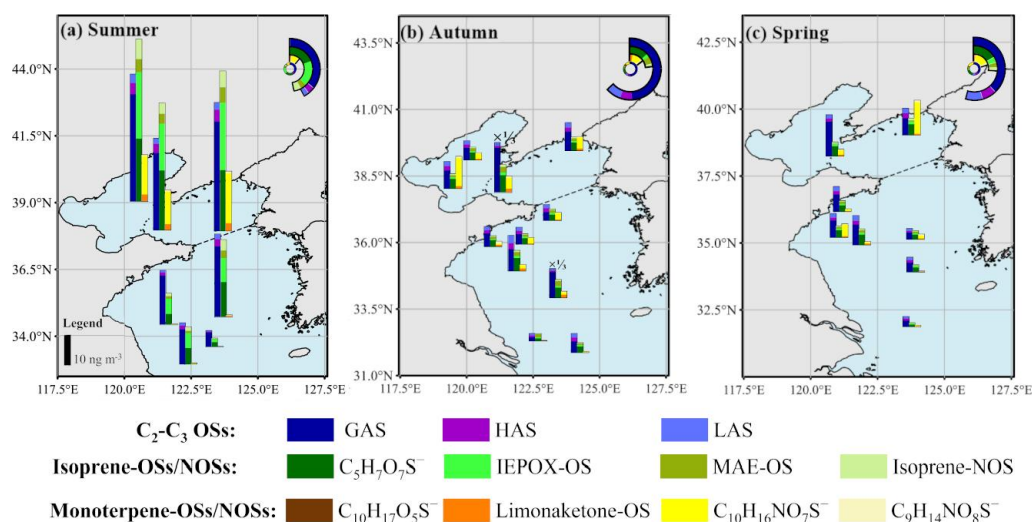


Figure 1. Spatial distributions of OSs in PM_{2.5} over the YBS during (a) summer, (b) autumn in 2021, and (c) spring in 2022. The inserted charts in panels (a, b, c) are the contribution of different OS compounds. The dotted lines in the panels are the dividing line of the North Yellow Sea (nYS) and the South Yellow Sea (sYS).

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The C₂–C₃ OSs, including glycolic acid sulfate (GAS), hydroxyacetone sulfate (HAS), and lactic acid sulfate (LAS), were the most abundant compound group across the observed seasons (Fig. 1, 2). The C₂–C₃ OSs concentrations were respectively 7.2±3.1, 24.2±12.4, and 12.8±14.4 ng/m³ in spring, summer, and autumn, comparable to the concentration levels at inland sites and lower than those in coastal areas (Fig. 2). In autumn and spring, the fraction of C₂–C₃ OSs, especially GAS, was much higher than other compound groups. The highest GAS concentration (47.8 ng/m³) over YBS was observed on 30 October during the autumn cruise. We noted that, in marine atmospheres, the contribution of C₂–C₃ OSs among the quantified OSs was much higher than those observed in various continental environments (Fig. 2). These low-molecular-



weight OSs could be formed via the oxidation of VOC precursors from both biogenic and anthropogenic origins (Wang et al., 2023a), and have been frequently observed as one of the most abundant OS groups in previous studies (Wang et al., 2018, 2020; Cai et al., 2020).

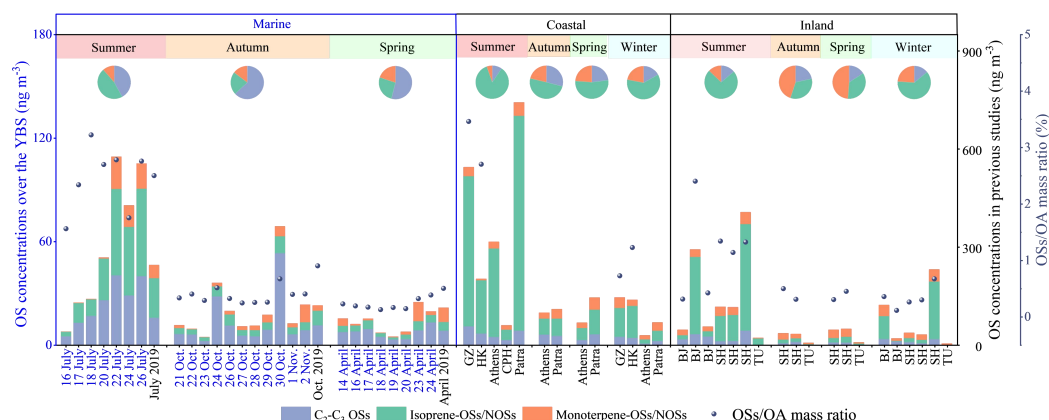


Figure 2. Atmospheric OS concentrations and mass ratios of (OSs+NOSs)/OA over the YBS in this study and in inland and coastal atmospheres reported in previous studies (Kanellopoulos et al., 2022; Meade et al., 2016; Nguyen et al., 2014; Wang et al., 2020, 2021, 2022b). The data labels in this work are denoted in blue, and those from previous studies are in black. The pie charts represent the average contribution of OS compound groups in each season. It is noted that the OS abundance over the YBS and at coastal or inland sites are represented by different y-axes concentration ranges.

The total concentration of quantified isoprene-OSs and NOSs ranged from 1.3 to 56.9 ng/m³, which were the most abundant group in summer over the YBS (Fig. 1). The predominance of isoprene OSs has been well documented at both coastal and inland sites (Fig. 2), which is attributed to the substantial biogenic isoprene emissions, especially during warmer seasons. Wu et al. (2021) reported abundant emission of isoprene from coastal and shelf seas (Wu et al., 2021), and isoprene OSs would then form via the interaction between sulfate aerosols and isoprene oxidation products (Surratt et al., 2010; Cooke et al., 2022). In the marine atmosphere over YBS, isoprene-derived OSs displayed a dominance by isoprene epoxydiol (IEPOX)-OS and C₅H₇O₇S⁻ during summer, while by C₅H₇O₇S⁻ during spring and autumn (Fig. 1). The C₅H₇O₇S⁻ compound has been suggested as a further oxidized or aged form of IEPOX-OSs (Armstrong et al., 2022; Chen et al., 2020). The abundant presence of C₅H₇O₇S⁻ in marine aerosols across seasons indicated the rapid oxidation and aging processes of isoprene SOA in marine atmospheres. The high contribution of C₅H₇O₇S⁻ molecule among isoprene-derived OSs has been reported in marine aerosols, as well as in coastal and inland atmospheres (Hettiyadura et al., 2015; Kanellopoulos et al., 2022; Surratt et al., 2008; Wang et al., 2018, 2022b, 2023b).

The IEPOX-OS, a typical low-NO oxidation product of isoprene, was one of the dominant compounds during the summer



cruise (Fig. 1). For the summertime samples, the contribution of IEPOX-OS among isoprene-OSs here is higher than that in a previous study conducted in 2019 over the YBS (Wang et al., 2023b). This could be due to the reduction of NO_x emissions in the North China Plain (Li et al., 2024), resulting in a lower NO condition in 2022 than in 2019. It is also noted that a BEH
170 Amide column was employed to separate the C₂-C₃ OSs and isoprene OSs in this work. The OS quantification here was more accurate than the study conducted in 2019, in which a reversed-phase column was used to separate the low-molecular-weight and highly polar OSs. This could be an additional reason for the different OS proportions between the two studies. The concentrations of methacrylic acid epoxide (MAE)-OS and isoprene-NOS, usually originated via NO/NO₂ pathway or under high-NO conditions (Worton et al., 2013), were much lower than those of IEPOX-OS and its aged product (C₅H₇O₇S⁻)
175 in the marine atmospheres (Fig. 1, Table S1).

The mass concentration and contribution of monoterpene-derived (nitrooxy-)OSs were lower than those of C₂-C₃ OSs and isoprene-derived OSs over the YBS (Fig. 1, 2). This compound group was dominated by monoterpene NOSs (C₁₀H₁₆NO₇S⁻), which were formed via the oxidation of monoterpenes in the presence of anthropogenic NO_x (Surratt et al., 2008; Wang et al., 2018). The formation of monoterpene OSs/NOSs in marine atmospheres was driven by the transported continental pollutants.
180 The concentration levels of monoterpene OSs/NOSs over the YBS were generally lower than those observed in continental atmospheres (Fig. 2) (He et al., 2014; Meade et al., 2016; Nguyen et al., 2014; Wang et al., 2020, 2021, 2022b).

3.2 Importance of OSs in marine atmospheres

The OS concentrations and mass contribution among marine OA were the highest in summer, followed by those in autumn and spring (Fig. 1, 2). The average OS concentration was 57.8±38.9, 20.4±19.7, and 13.3±8.3 ng/m³ in summer, autumn, and
185 spring, respectively. During the summer cruise, OSs occupied 1.6%–3.2% (2.5% on average) of the marine OA mass concentrations, which were comparable to those observed in coastal regions and higher than those at the inland sites (Fig. 2). The elevated concentration levels and contributions of biogenic OSs, especially isoprene OSs and C₂-C₃ OSs, in summer were attributed to the increased biogenic VOC emissions from marine phytoplankton or photochemical reactions in surface
190 microlayer (Conte et al., 2020). The higher seawater Chlorophyll-*a* (*Chl-a*) during summer indicated elevated isoprene production by phytoplankton, and higher temperature favored the sea-to-air transfer process of isoprene (Fig. S3). The vital importance of biogenic OSs to OA formation in summer has been highlighted in previous observations at both marine and continental sites (Hettiyadura et al., 2017; Kanellopoulos et al., 2022; Meade et al., 2016; Nguyen et al., 2014; Wang et al., 2020, 2021, 2022b). We cannot exclude the potential influence of terrestrial biogenic VOC emissions based on the
195 observational evidence. The air masses were dominantly from the open ocean. During spring or autumn, the lower *Chl-a* and air temperature resulted in the decreasing of biogenic OS formation (Fig. 2). Though the seawater *Chl-a* was at similar



concentration levels in spring and autumn (Fig. S3), the OS abundance was lower in spring. The ambient temperature was lower in spring, and the oceanic phytoplankton had not revived from the low temperature conditions throughout winter. Thus, the biological activity and biogenic VOC production were likely at low levels during the spring cruise.

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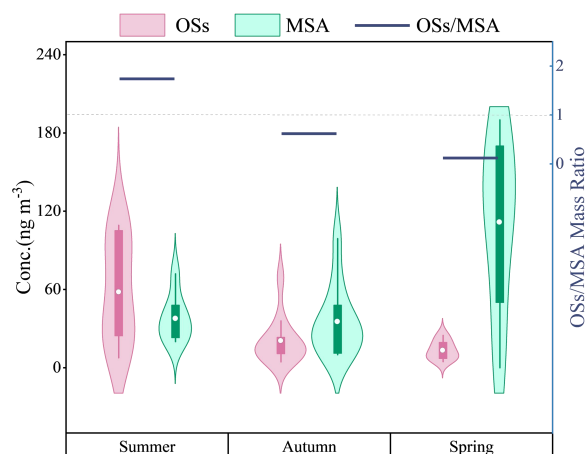


Figure 3. The OSs, MSA concentrations and OSs/MSA mass ratios in atmospheric aerosols over the YBS.

During summer, the active interactions between biogenic VOCs, especially isoprene, and acidic sulfate converted notable fractions of inorganic sulfate aerosols to OSs in marine atmospheres. The abundance of OSs was comparable to that of MSA in summer, and their mass ratios were higher than those in autumn and spring (Fig. 3). Atmospheric MSA contributes to the CCN formation in marine boundary layer, which is a vital species relevant to the CLAW hypothesis of oceanic phytoplankton-controlled climate regulation (Ayers et al., 1997; Charlson et al., 1987; Quinn and Bates, 2011). In addition, the OSs and MSA displayed strong correlations ($r = 0.86$, $p < 0.01$) in autumn (Fig. S4). This suggested that the atmospheric OSs and MSA formation was limited by the same environmental factors in autumn, which could be the lower biological activities indicated by the seawater *Chl-a* and temperatures (Fig. S3). The cruise observations indicated that organosulfate, besides MSA, should be taken into consideration when studying the sulfur cycle and its climate effects in marine atmospheres, especially over regions with high phytoplankton biomass and high temperature.

215 3.3 Seasonal variation of atmospheric OS composition

For the seasonal variations of OS composition, the chemical spaces of the autumn and spring samples are highly overlapped, which are different from that of the summer samples (Fig. 4a). The fraction of isoprene-derived (nitrooxy-)OSs was higher



during the summer cruise than those observed during the other two seasons. The autumn and spring samples generally showed a higher contribution by monoterpene-derived OS compounds. The seasonal variation was attributed to the relatively lower isoprene emissions, indicated by the lower seawater *Chl-a* (Fig. S3), and the more severe influence of anthropogenic pollutants transported from the continent in spring and autumn (Fig. S1). In addition to the air mass back trajectories, the more severe impacts of continental outflows in spring and autumn were also indicated by elevated elemental carbon (EC) concentrations ($0.5 \mu\text{gC}/\text{m}^3$ and $0.4 \mu\text{gC}/\text{m}^3$ compared to $0.2 \mu\text{gC}/\text{m}^3$ in summer). In marine atmospheres over the YBS, the relative contribution of monoterpene-derived (nitrooxy-)OSs was lower than that in continental atmospheres under more severe impacts of anthropogenic pollutants (Fig. 4a).

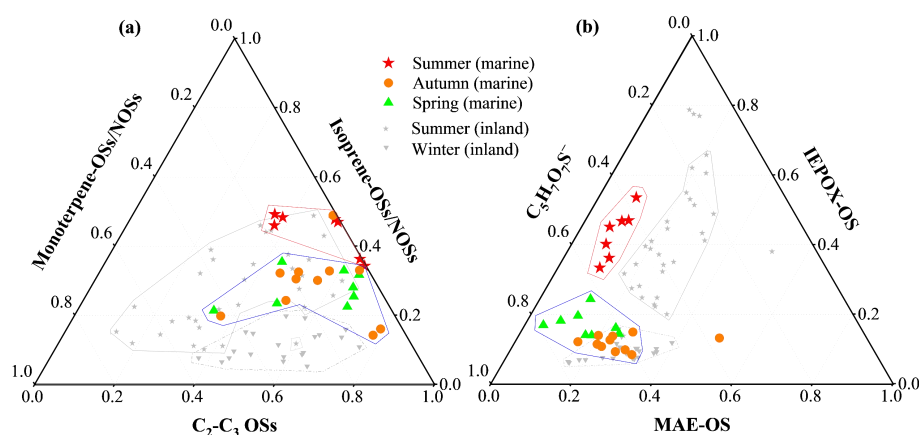


Figure 4. (a) Relative abundance of isoprene OSs, monoterpene OSs, and C₂-C₃ OSs, and (b) composition of isoprene OSs over the YBS in summer (red), autumn (orange) and spring (green). The results previously reported at the inland urban site (Wang et al., 2020) are colored gray.

In marine atmospheres over the YBS, different influences of marine emissions versus continental outflows across seasons resulted in the variation of C₂-C₃ OSs/isoprene-OSs mass ratios (Fig. S5). Strong correlations ($r=0.79-0.97$, $p \leq 0.05$) between isoprene-OSs and C₂-C₃ OSs suggested their consistent biogenic sources dominated by isoprene oxidation, which has been reported in previous studies (Schindelka et al., 2013; Surratt et al., 2008; Wang et al., 2020). In summer, the abundance of C₂-C₃ OSs was comparable to that of isoprene OSs. While, during autumn and spring, we observed higher mass ratios of C₂-C₃ OSs versus isoprene-OSs due to the additional sources of C₂-C₃ OSs contributed by anthropogenic sources (Fu, 2008; Huang et al., 2018; Liao et al., 2015).

The chemical space distributions of isoprene OSs also displayed obvious seasonal variations. The fraction of IEPOX-OSs was substantially higher, and that of MAE-OS was relatively lower in summer compared with those in spring and autumn



(Fig. 4b). The low-NO conditions in summer favored the IEPOX formation from isoprene oxidation via HO₂ pathway, while the formation of MAE via NO/NO₂ pathway would increase under the influence of continental pollutants in autumn and spring (Wang et al., 2020; Worton et al., 2013). During summer, the relative contribution of MAE-OS among isoprene-OSs in marine aerosols over the YBS was lower than those observed in continental atmospheres, indicated by the gray markers in Fig. 4b. This was due to the lower anthropogenic pollutants and NO conditions in marine atmospheres than in continental atmospheres. The proportions of C₅H₇O₇S⁻, a further oxidation or aged forms of IEPOX-OSs (Armstrong et al., 2022; Chen et al., 2020), were also higher in autumn and spring than in summer. The dominant presence of C₅H₇O₇S⁻ compared to IEPOX-OS indicated a highly oxidized state of marine SOA in spring and autumn.

3.4 Spatial distribution of OSs regulated by continental outflows

As shown in Fig. 5, atmospheric OSs concentrations over the Bohai Sea and the North Yellow Sea (nYBS, 51.3 ± 37.4 ng/m³) were notably higher than those over the South Yellow Sea (sYS, 16.1 ± 11.9 ng/m³). Surrounded by the continent, the nYBS region was under more severe impacts of transported anthropogenic pollutants compared with the relatively open sYS. This is also indicated by the variation of EC concentrations in atmospheric aerosols over nYBS and sYS areas (Fig. S6). Marine emissions dominated the biogenic OS formation over the YBS in summer, while we cannot exclude the potential influence of transported continental air masses, especially over the nYBS. This could be a reason for the higher OS concentrations over the nYBS than those over the sYS.

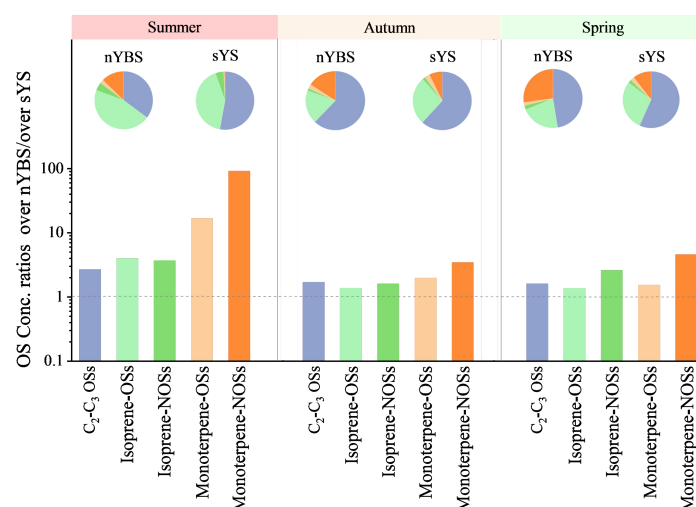


Figure 5. Concentration ratios of atmospheric OSs over the nYBS versus those over the sYS. The pie charts show the relative contribution of OS compound groups over the nYBS and the sYS during each season.



The concentration levels and compositions of OSs in atmospheric aerosols over the nYBS and the sYS are compared in Fig. 5. Among the quantified OSs derived from different VOC precursors, monoterpene-NOSs displayed the most obvious enhancement ratios over the nYBS compared to those over the sYS (Fig. 5). During the summer cruise, monoterpene-NOSs over the nYBS elevated to nearly two orders of magnitude higher than those over the sYS. The mass contributions of monoterpene-NOSs among the total OSs over nYBS were higher than those over the sYS, as shown in the pie charts of Fig. 5. Monoterpene-NOSs are usually formed via the interactions between anthropogenic NO_x , sulfate, and monoterpenes (Bryant et al., 2021, 2023; Wang et al., 2018). A recent study also suggested monoterpenes could be generated by biomass burning, besides the biogenic emissions (Wang et al., 2022a). The spatial difference of monoterpene-NOSs further indicated the more severe influence of anthropogenic pollutants over the nYBS.

The OS abundance displayed the most obvious enhancement over the nYBS in the summer samples, in which the concentrations of $\text{C}_2\text{--C}_3$ OSs and isoprene-OSs/NOSs over the nYBS elevated to 2.4 and 3.9 times of those over the sYS. The biogenic emissions from marine phytoplankton were more abundant in summer than in the other seasons. Transported anthropogenic pollutants over the nYBS would promote the formation of biogenic OSs via anthropogenic-biogenic interactions in marine atmospheres. Previous observation has suggested that the formation of biogenic SOA, including isoprene OSs, could be obviously mediated by anthropogenic sulfate and NO_x in regions with substantial anthropogenic-biogenic interactions (Xu et al., 2015). We noted that isoprene-OSs were not observed in remote marine aerosols over the Southern Ocean or the southern Indian Ocean, where the influence of transported anthropogenic pollutants was likely limited (Claeys et al., 2010; Cui et al., 2019). Our results suggested the universal existence of biogenic OSs in marine aerosols over regions with anthropogenic-marine interactions, while the presence of OSs in different marine environmental conditions needs further observation evidence.

3.5 Origins and influence factors of atmospheric OSs

Principal Component Analysis (PCA) was performed to further understand the sources of atmospheric OSs over the YBS (Fig. 6, Table S2). A total of 18 particulate components, including OSs, water-soluble ions, EC, and MSA, were chosen to carry out the statistics. Three factors could explain 83% of the measurements. Majority of the OS and NOS compounds showed high loadings in Factor 1, which explained 52% of the measurements. Characterized by high loadings of nss-sulfate, Cl^- , and low loadings of anthropogenic species (e.g., EC, nss- K^+), Factor 1 represented the sulfate-catalyzed reactions with VOCs dominated by marine emissions. Factor 2 shows high loadings of EC, nss- K^+ , and NO_3^- , suggesting the transported anthropogenic origins dominated by combustion emissions, which explained 21% of the measurements (Table S2). Factor 3, dominated by MSA, EC, and sea salts, was a mixed source of marine-anthropogenic interaction, which explained 10% of the



variance (Table S2).

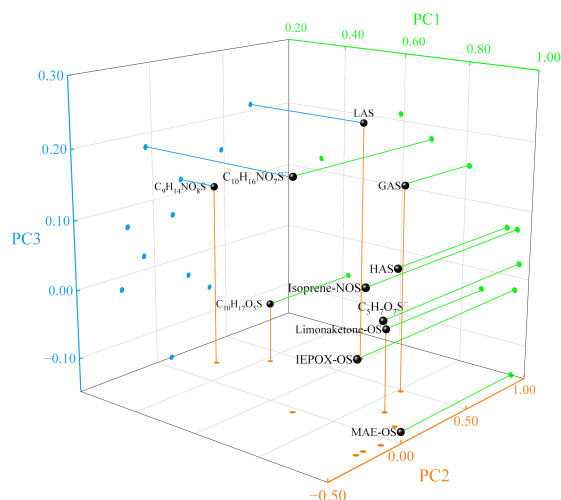


Figure 6. PCA statistics of the measured OSs and NOSs during the cruise observations. PC1, PC2, and PC3 represent the source of sulfate-catalyzed reactions with biogenic VOCs, transported anthropogenic origin, and a mixed source of marine-anthropogenic interaction, respectively.

The majority of quantified OS compounds, especially the isoprene-derived ones (IEPOX-OS, MAE-OS, $C_5H_7O_7S^-$, isoprene-NOS, GAS, and HAS), were dominated by the source of sulfate-catalyzed reactions with biogenic VOCs (Factor 1), as displayed in Fig. 6 and Table S2. The homogeneous origin of C_2 – C_3 OSs and isoprene-OSs/NOSs from the oxidation of isoprene has been approved in this work and previous observations (Surratt et al., 2008; Riva et al., 2016). This source factor was more related to the marine emissions, rather than anthropogenic pollutants, indicated by the low loadings of anthropogenic EC or $nss-K^+$. Isoprene could be largely emitted by phytoplankton and from photochemical processes in surface seawater, and then released into marine atmospheres (Brüggemann et al., 2018; Cui et al., 2023). The reactive uptake of isoprene by sulfate aerosols could be a vital reaction pathway for OS formation in marine aerosols (Wang et al., 2023b). Elevated OS concentrations were observed as the increasing of air temperature in summer or as the increasing of wind speed in spring (Fig. 7). Higher temperature or wind speed would promote the sea-to-air exchange of isoprene and favored the OS formation in marine atmospheres.

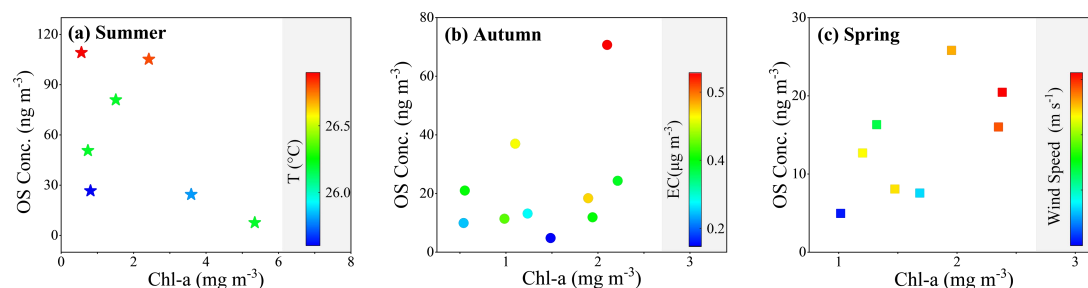


Figure 7. Variations of OS concentrations as a function of chlorophyll-a (Chl-a) in (a) summer, (b) autumn, and (c) spring. The markers are colored by air temperature, EC, and wind speed, respectively.

315 The loadings of monoterpene-OSs/NOSs in anthropogenic-related sources (Factor 2 and Factor 3) cannot be neglected, which was different from the main source of isoprene OSs from marine-dominated sulfate-biogenic VOC interaction (Factor 1). Lactic acid sulfate over the YBS showed comparable loadings in the transported anthropogenic origin (Factor 2, 0.70) and the marine-dominated sulfate-biogenic VOC interaction source (Factor 1, 0.59). The loadings of LAS in the mixed source of marine-anthropogenic interaction (Factor 3) were higher than other identified OS species (Fig. 6). A relatively high

320 loading of GAS (0.48) was also observed in Factor 2 (Table S2). The PCA result provided observational evidence on the additional sources of monoterpene-OSs/NOSs and C₂-C₃ OSs from transported anthropogenic pollutants over marginal seas. During the autumn cruise, higher OS concentration levels were observed when higher EC concentrations occurred, which also indicated the additional contribution of OSs by anthropogenic sources (Fig. 7b).

This work quantified and characterized the atmospheric OSs derived from isoprene and monoterpenes over the Asia marginal

325 seas. The chemical nature and distribution of OSs were modified by the joint influence of oceanic biological emissions and transported continental pollutants. The results highlight the abundant formation of airborne OSs in summer, which is promoted by the elevated biogenic VOC emissions from surface ocean. During high biological activity periods, atmospheric OS levels could surpass the MSA concentrations in marine aerosols, which is a vital species in the well-known climate regulation via oceanic phytoplankton sulphur emissions (CLAW hypothesis). Shipboard observations over open ocean areas

330 are needed to gain further understanding on the roles of OSs in modifying the sulfur cycle, biogenic VOC oxidation and regulating climate in marine boundary layer.

Author contribution

Y.W. designed and supervised the research. M.H. supervised and provided the instrumentations. S.L., Y.Z., Y.Y., Y.G., C.Y.

335 and Y.J. conducted the measurements. S.L. analyzed the data. Y.C.W synthesized the standards. S.L. and Y. W. wrote the



manuscript with contributions from all co-authors.

Declaration of competing interest

The authors declare no conflict of interest.

Data availability

340 The dataset is available upon request from the corresponding author.

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