Dear Editors, dear reviewer,

We thank the reviewer for the insightful comments that greatly helped to rive our manuscript. Please find below the reviewer's comments in *black italics* and our answers in blue.

Sincerely,

Coline AriaGno (on behalf of all co-authors)

Summary: Ariano et al. explores how landslides influence landscapes evolution using numerical modelling in combination with topographic analyses. Specifically, they focus on three catchments in the western Alps that exhibit different morphologies (fluvial vs. glacial), to assess how landslide activity and erosion vary spatially and over time depending on the pre-existing landscape. Indeed, the different catchments represent a gradient in glacial imprint and deglaciation timing, allowing these regions to be used as a natural laboratory. The main objective is thus to predict and explore landslide activity and its role in transient landscape evolution during interglacial periods.

In general, this is a well-illustrated paper and well-written, with appropriate references. The scope of the study is well thought out, and I believe the results will be of interest to the wider geomorphological community. However, I do feel that the discussion could use some work (made easier to follow), to clearly communicate the implications of this work to the scientific community.

Particularly, I found that the main hypothesis of the paper is not defined consistently throughout the paper:

- 1: Line 143-145: "Our working hypothesis is that the different morphological signatures observed for Alpine catchments are evidencing both landslide activity and deglaciation timing."; So, the scope is to test that landscapes today are a result of both glacial erosion and landslide activity (as well as other processes). Very clear and feasible, and I believe this is indeed shown by the results.
- 2: Line 559: "Our landscape evolution model [...] has been designed to explore the hypothesis that landsliding represent a dominant geomorphological agent during postglacial periods."; Sort of similar to the first instance; landslides are important. Clear.
- 3: Line 665: "our initial hypothesis about the capacity of landslides to erase this glacial topographic inheritance over the last post-glacial period". So, here you state that the hypothesis is to test whether landslides can erase the glacial imprint over an interglacial cycle. This is very different from other two instances, and in my opinion much more difficult to test.
- 4: Line 704-706: "Here we discuss our initial hypothesis, that all the studied catchments had the same glacial topographic imprint, and show that the three catchments have a distinct erosion dynamics explained by diachronous landslide activity following different glacial retreat times". Again, this is different, and I am not sure I agree that you can be sure the three catchments all started out with an equal glacial imprint, give their differences e.g., in glacial duration (you do comment on this in the end).

In any case, my point would be that the hypothesis (or hypotheses) to be tested should be clear throughout the manuscript, and more care should be put in developing the rationale behind the 3. and 4. points listed here. These arguments are not trivial, and you are sort of cutting some corners by referring back to a hypothesis (that was stated differently initially).

Thank you for these insightful comments, also raised by the second reviewer. We have made some changes and clarified the hypotheses in the introduction and throughout the article (Lines 149). Specifically, we have retained the main hypotheses exposed in the introduction

and recalled at the beginning of the discussion, but have rephrased the points 3 and 4 above for clarity.

Can you really infer from this study that the initial glacial landscape of the 'fluvial' catchment has been erased by landslide activity, whereas the others are still ongoing? If the upper catchments will not transition into fluvial catchments over timescales of 100 kyr, how has the lower catchment managed to do the transition already? You come to this in the end of section 5.2, but I really think the argumentation for this point should be outlined much more clearly throughout. And it should be clearly stated where you land in terms of your initial hypothesis. You could also present these ideas with more caution: "If the three catchments had the same glacial topographic imprint initially...".

As you and the other reviewer suggested, we clarified our hypothesis (lines 149): Considering that the three catchments have a similar initial glacial inheritance, the fluvial (Pisse) catchment has already been strongly influenced by hillslopes processes, as testified by V-shaped valley in its lower section (fig 2). However, the 100-kyr simulation is not sufficient to induce similar changes (transition from U- to V-shaped valley) in the glacial (Pilatte) and intermediate (Etage) catchments, which means that a mechanistic approach, such as Hylands' model, cannot alone explain the fluvial topography of the fluvial (Pisse) catchment. That is why we argued in section 5.2 that the glacial shaping (erosion power) or/and the shorter glaciation duration in the fluvial (Pisse) catchment have an significant impact on the topographic evolution of the catchment during our simulations.

Thus, we do not infer that the initial glacial landscape of the 'fluvial' catchment has been erased by landslides activity. Instead, we argue that this actual topography is possibly a combination of a less intense glacial inheritance and longer/more active hillslope processes. We have modified the manuscript accordingly in section 5.2.

Another smaller comment I have relates to the time scale of the model simulations. The 100-kyr duration of the models seems a bit odd, i.e., to simulate landscape evolution over such a duration without considering glacial changes. It also seems a bit unnecessary given the overall scope, focused on interglacial timescales. Perhaps this could be justified (or the rationale behind could be explained), for instance by including in the discussion some reflections of how results would differ/be limited given a different choice. Right now, it is simply stated as a fact in the manuscript (line 615-616).

We agree with the reviewer that 100 kyr is not entirely realistic as a timescale for investigating the post-glacial period. However, this was raised during our modeling study as is an interesting and necessary duration which enables 1) to obtain clear temporal trends of erosion rates, smoothing out landslide variability at the onset of our simulations, and 2) to quantify the duration theoretically required to erase the glacial imprints and to reach stable hillslopes. We now mention this in the discussion section on line 638.

In addition, I think the discussion lacks some perspective related to the fact that the used (present-day) topography has already been influenced by these processes since glacial retreat (i.e., the DEM has already been affected by these processes to some extent – and given your conclusions potentially to a large extent!). Could this suggest that landslide erosion rates would have been even bigger between glacial retreat and now? Specifically, this study predicts a pulse in erosion rate by landslides over a few thousand years. Would this pulse already be done in the real world? Or has it simply been even bigger prior to today? Could such a trend be extrapolated back in time based on the presented results? Reflections on these questions could be added in the discussion.

Indeed, the initial topography we use in our models has already been influenced by post-glacial hillslope processes. Considering the results we have, landslide activity since the deglaciation should have been at least equal to the one we observe at the beginning of our simulation. It is however difficult to extrapolate further as the deglaciation time and the onset of post-glacial hillslope activity is asynchronous between our catchments but also within each catchment: lower areas having been deglaciated earlier than higher areas. Moreover, we do not consider the role of permafrost and its retreat, which is likely to have a large in controlling the intensity and timing of post-glacial landslide activity. It would, in turn, be interesting to investigate a possible delay between glacial retreat and landslide activity as modulated by the rate and timing of permafrost degradation.

We have added a few sentences in the discussion to mention this point and to add some perspectives on the influence of the chosen DEMs (modern topographies) used in our model (section 5.1.2, Line 645).

Finally, I have listed several comments and suggestions below that will hopefully be useful when making the final adjustments of the manuscript. In addition, I suggest going through the manuscript to do a final check of language (e.g., lines 152, 344-345, 444, 503, etc.) and consistency in reference style (e.g., line 51, 72, 351, 580 also using 'e.g.,' instead of 'e.g.', etc.).

We have checked the manuscript and language mistakes as suggested, and thank the reviewer for all the useful suggestions.

<u>Lines 28-32:</u> the mentioning of 'the glacial buzzsaw' might need a little more elaboration. It reads as if the glacial buzzsaw is usually attributed to a decrease in unstable slopes as well as a lowering of maximum topography. But is the mechanism in the glacial buzzsaw not that glaciers increase the steepness of their headwall slopes (i.e., increase in unstable slopes), such that hillslope processes are more active, and therefore by extension reducing the maximum elevation of a catchment? E.g., to quote one of the defining papers: Mitchell and Montgomery, 2006: "The summit altitudes are set by a combination of higher rates of glacial and paraglacial erosion above the ELA and enhanced hillslope processes due to the creation of steep topography."

It is true that this sentence is a bit ambiguous. To not overload the abstract, we decided to cut this mention of the glacial buzzsaw and glacial processes, as it is well developed and explained in the discussion of the manuscript (section 5.3.2).

Line 76: an uncovered landscape or uncovered landscapes. Corrected.

Lines 77-78: consider if this sentence should also be past tense. Changed as suggested.

Lines 84-85: "leading to a postglacial increase on both the frequency and intensity of hillslope events through time". I understand that the frequency and intensity go up as the regions deglaciates, but is the point not that it then decrease through time hereafter?

This sentence was not completely clear and the end part could be misleading. This has been rephrased for clarity.

Lines 98-100: maybe start with a 'while' Done.

Lines 134-138: I would mention the stochastic nature of the model here already. Done.

Line 152: 'Three'. Removed as suggested.

Line 181: I suggest consistency using 'three' versus '3'. I would suggest 'three'. Done.

Lines 205-211: you could consider also citing the new paper by Maxime Bernard here (see below). Reference added.

Line 222: why V+?. This was an error and has been corrected to "V-shaped".

<u>Line 263-264</u>: maybe this is the tradition when concerned with the used model. But I find it odd to refer to hillslope height, when talking about elevation/height change between two cells. Would maximum stable slope not be more appropriate?

Here, we have decided to retain the terminology used in the original framework of the Hylands model, which is widely referenced in the manuscript (Campforts et al., 2020).

Line 278: perhaps 'the erosion scar generates a failure plan' can be formulated more precisely. Is the erosion scar not generated by the failure and not vice versa? Modification done.

<u>Lines 282-284</u>: this was not completely clear to me - i.e., whether all DEM cells in the entire catchment above a certain plane would be considered unstable for one specific landslide event?

Starting from the trigger cell, the model identifies a failure plan upstream (following the Culmann angle) and all the cells above this plan are eroded. The roughness and the topographical irregularities of the terrain naturally limit the extent of the landslide.

<u>Line 286</u>: is it necessary to introduce Ff here, when not elaborated further? One could simply state that 'in this setup, all sediments are instantaneously evacuated.'

As suggested, we have been more synthetic in this part and we also have moved the details of the simulation (section 5.1.2), with a different *Ff value*, in the supplementary.

<u>Lines 292-324</u>: I would suggest simply to incorporate this section in the Model calibration section 3.3. I see no reason for dividing this into two distinct sections. For instance, the first part of section 3.3.1 gives info/context relevant to the text in 3.2 and right now you refer to back and forth between the sections several times. In addition, I found lines 314-324 difficult to follow. If parameters give rise to few landslides, how do you then generate a large amount of landslides? Maybe it is the 'Then' that leads to confusion. Should it be 'Either we compile multiple simulations ... or we reduced the return time...'. But still, it could be clarified how you calibrate t_{LS} while scaling this parameter.

The subsection 3.2 is not included in subsection 3.3 because it describes the procedure we used for the overall modelling, while subsection 3.3 details the specific calibration of the model for our simulations. We agree that these sections are closely related, which is why they are included in the same section 3.

We have modified the ambiguous part of this paragraph and made also some changes in subsection 3.1.1 for clarity in the presented parameters.

Line 344-345: *did* not display any clear rollover. Done.

<u>Line 346</u>: I guess this is not a matter of visualization but representation. Capitalize 'we' or perhaps add. 'However, we'

As the second reviewer suggested too, we added a connector at the beginning of the sentence.

Line 350-353: -2.3 is larger than -2.5 ;-). Done as suggested.

<u>Lines 397-401</u>: I would expect that the listed three combinations are part of a whole envelope of realistic parameter combinations, where the listed are just some examples. I would highlight that instead of listing specific values explicitly. Also, it would be nice to see an example of the spatial/temporal patterns in landslide activity for this selected catchment, perhaps for a few 'end-member simulations', showing the variability possible within reasonable values of parameter values (friction angle, cohesion, return time; e.g., supplementary figure, particularly if they are not very different – but that would be a point in itself).

We indeed chose the intermediate combination, among acceptable range of values for each parameter, to further develop our study and simulations. We modified the main text to be less specific about these examples of possible parameter combinations.

Moreover, as you rightly suggested, we ran new simulations with 'end-member parameters'; i.e. one with low parameter values (C=20 kPa, $t_{LS}=50$ kyr) and another with high parameters values (C=100 kPa, $t_{LS}=250$ kyr) and added the associated figures in the supplementary.

<u>Lines 425-427</u>: "landsliding results in homogeneous slopes which only slightly exceeds the internal angle of friction (i.e., 0.7, represented by white color in Fig. 6)." Seems like there are plenty of red colors still? Or do you mean only the regions associated with landslide activity? Again, would be nice with additional panels showing initial and/or changes in slope compared to initial values. Perhaps also comment on the high-slope regions that do not experience landslides.

Indeed, this sentence refers to the regions associated with landslide activity. The model does not erase all the steep slopes as it is stochastic. Please note that the color scale was different between the initial and final slope maps (Figs 1 & 6). We thought that it was more relevant to show the initial slope in the presentation of the study area (Fig. 1) but that it would be redundant to show it again in Figure 6.

<u>Lines 452-460</u>. For consistency, I would suggest referencing all figures when referring generally to all catchments, e.g., not only fig. 8 but also the corresponding supp. figs. Done as suggested.

<u>Line 483:</u> it seems a bit arbitrary with the selected elevation range of 2400-2800 m, to capture the minimum for both catchments. Why not simply specify a different elevation for each catchment? The large range makes it difficult to see in the left panels that there are fewer red dots in that interval (as it is so wide) – particularly for the glacial catchment.

The selected elevation range is linked to the fewer predicted landslides in the Etages (intermediate) catchment. Further, we relate this elevation range to glacial morphology. This elevation range refers to an area influenced by glaciation, which is indeed quite large due to multiple glacial cycles and ELA oscillations between glacial and interglacial periods. We think this is important to keep the same elevation range between the catchments since they experienced similar climate forcing. As it is not easy to show the quantity of landslides in the left panels, we have highlighted the density in the histogram in the right panels.

<u>Lines 487</u>: be careful using the word 'observations' in connection with model predictions. Thanks for this suggestion, this has been corrected.

<u>Line 493-495</u>: it is unclear how this is evident from Fig. S5. Should be S4, I assume. Also reference to Fig. 7G-H in the next sentence is unclear.

Same comment was raised by reviewer 2, we made the changes for clarity.

<u>Line 569</u>: I would suggest specifying rock uplift and sediment transport already in title and throughout. Done as suggested.

Lines 584-586: the language of this bit is unclear to me.

We rephrased the sentence for clarity: "However, despite these limitations, we believe that our modeling approach stays appropriate to assess the hillslope stability over 100-kyr timescales, which is largely dependent on climatically-shaped alpine topography and bedrock mechanical strength."

<u>Line 593</u>: return time of 150 kyr? Yes, this model parameterization is explained in the section 3.3.3.

<u>Lines 588-596</u>: I am not sure I understand the rationale behind the need for a model with an average erosion rate of 2-3 mm/yr to compare with the catchment-averaged erosion rate of 1 mm/yr. Do you want to imply that the **hillslope erosion rate needs to be higher that the average because other parts of the catchment have lower values?** Or is it because you are interested in the predicted longer-term erosion rate to be closer to 1 mm/yr? This is not clear from the text, and then why specifically 2-3 mm/yr was chosen? This would also rely on the assumptions you make about what has happened in the catchments since deglaciation until now (since you use present-day topography that have experienced many landslides already), which is what is reflected in the cosmo-derived rate.

This point has also been raised by reviewer 2, we rephrased the sentence for clarity and to argue for the supplementary test with higher erosion rates: "Considering effective sediment connectivity in the catchment (in our study area, main fluvial valleys are sediment bypass areas without significant incision but potential transient storage) and only landsliding to derive our catchment erosion rate, 1 mm/yr is likely to be an end member minimum value for our simulations."

<u>Lines 672-674:</u> I don't follow the argument here; can you be certain that the 'glacial' vs. 'fluvial' catchments are due to reshaping through hillslope processes? Given the much lower hillslope activity in the fluvial catchment, could this catchment simply have experiences less glacial modification in the first place? The intended message of this section in general is a bit difficult to follow (lines 668-682), could it be spelled out more clearly?

Indeed, the fluvial catchment has potentially experienced less glacial erosion and we come to this point in detail during the discussion (5.2.2), once all the arguments from the results have been gathered. However, the V-shaped valley we observed (Fig. 2) implies hillslope and fluvial activity in this catchment, which can be the result of postglacial reshaping of the topography. We slightly modified the paragraph and hope that it will be clearer for the reviewer.

<u>Line 736</u>: starting a new paragraph with 'this observation' is somewhat unclear. Please specify what 'this observation' refers to. Done we rephrased as suggested.

<u>Section 5.3.1</u>: this section is somewhat short and could potentially be included elsewhere. In addition, there is some discrepancy here related to other parts on the manuscript – arguing that U-Shaped valleys takes multiple glacial cycles to form, while other parts of the manuscript seem to suggest that the 'fluvial' catchment has transitioned from glacial to

fluvial during one deglaciation. This will likely be sorted out/become clear if the hypothesis of the paper will be clarified.

Section 5.3 reflects on the topographic evolution of mountains at the light of this study' outcomes. The short section 5.3.1 places the post-glacial period back in the Quaternary, i.e. with several successive glacial /interglacial periods. We are not certain that this reflection can be placed earlier, and section 5.3 would no longer be relevant if we moved section 5.3.1. Concerning the fluvial (Pisse) catchment, we argued that its lower part seems to have

achieved a V-shaped valley, characteristic of a fluvial erosive catchment. However, as explain in section 5.2.2, the glaciation was probably less intense in this area and the U-shaped valley not well marked.

<u>Lines 772-797</u>: as mentioned, I believe enhanced hillslope processes are already a recognized component of what has been presented as 'the glacial buzzsaw', which could be recognized in this section. This does not make the current study irrelevant in this context.

We agree and thank the reviewer for raising this point. The concept of glacial buzzsaw applies during glacial period, when, as correctly pointed by the reviewer, the mountainous reliefs created by glacier indirectly induce hillslope processes on steep slopes. Here we are referring rather to the interglacial period, when glaciers are no longer the main agent of erosion.

Moreover, the mechanisms associated with the buzzsaw are still not very clear in the literature. For example, Brozovic et al. (1997) and Egholm et al. (2009) illustrate the hypsometric distribution of elevation and slope, but landslides are not included in their studies. Only Mitchell and Montgomery (2011) addressed explicitly landslides as a potential mechanism of the buzzsaw. We think this is therefore an area that remains to be explored.

Comments on Figures:

Figure 3: please specify t_n , t_m , etc. in the caption. Also, for clarity there should be arrows from 'Trimline zone' to yellow circles in both sides. Done as suggested.

<u>Figure 4</u>, caption. 3. 10⁴ m² (the period .) should be fixed, here and throughout the paper. Corrected.

<u>Figure, 6</u>. I would suggest also to show panels with the change in slope (final slope versus initial slope). It is difficult to assess the changes not having the initial slopes at hand. The initial catchment slopes are presented in figure 1, and not to overload figure 6 we preferred to show only final slopes.

Figure 7: I would suggest adding thin horizontal lines at slope 0.7 to mark the internal angle of friction. Done as suggested.

<u>Figure 8</u>. *unclear why panel B is termed 'steepest slope'*. Modified to 'all slopes'.

Figure 9: y-axis label could simply be 'triggering point elevation (m)'. I think right panel C could be interpreted as bimodal, although I agree it is not as clear.

The y-axis label appears less ambiguous like that according to us. Concerning panel C, yes it could be but there is only one elevation interval which is lower than others, so the apparent bimodal distribution can result from our binning. We prefer to be careful and not to over interpret this result.

Cited References:

Bernard, M., van der Beek, P. A., Pedersen, V. K., & Colleps, C. (2025). Production and preservation of elevated low-relief surfaces in mountainous landscapes by Pliocene-Quaternary glaciations. AGU Advances, 6.