

Response to reviewer R1 comments

August 22, 2025

Thank you for your detailed and helpful review. In this document, reviewer comments are in **black** and our comments are in **red**. New text added to the manuscript is in **blue**.

This is a clearly written paper with nice figures describing nice analysis of an extraordinarily rare and hard to obtain dataset. The manuscript should be published.

Thank you for your positive assessment of our paper, we will address your individual points below

I do have a number of comments, questions and morsels for thought that I list below in the order in which I read. The majority are (very) minor, amounting to text and grammar nits, but some are more substantive. In particular I would like to see

- more supporting evidence behind the claim that mixing is weak (for the reasons given in the final comment below),

We show that the median TKE dissipation rate is 10^{-11} to 10^{-10} , which are very low values, comparable to the background TKE dissipation rate in the ocean (see Figures 6 and 7 in Waterhouse et al. (2014)) Waterhouse et al. (2014) does not give average values for epsilon, instead we can compare values for kappa. In our study median values of kappa

range between $0.2 \times 10^{-4} m^2 s^{-2} - 1.1 \times 10^{-4} m^2 s^{-2}$. Waterhouse et al. (2014) gives average deep ocean (depth between 1000 m and the bottom) values of kappa as $4.3(0.4 - 11.5) \times 10^{-4} m^2 s^{-1}$, with the values in parenthesis the 95-th percentile bootstrap confidence range. This indicates that our values of kappa lie within the lower range or just below the global distribution for the deep ocean. We will add a reference to global average values of mixing in Waterhouse et al. (2014). Additionally, we will include a new figure showing the distribution of measured epsilon in different ice shelf cavities. Our values lie within the range of previous observations as we have thus rephrased our abstract to remove the reference to “low mixing”, the sentence in question now reads “Rates of background mixing are $\varepsilon \approx 10^{-10} W kg^{-1}$ with patches of higher mixing of $\varepsilon \approx 10^{-8} W kg^{-1}$.”

- better figure 3 and 4, which currently mixes aspect ratios, has the reader going back and forth and does not allow direct comparisons of the most relevant quantities - specifically epsilon and the different instability indicators

We have combined Figures 3 and 4 to the new Figure 3. All panels now have the same aspect ratio.

- quantification of the ADCP vertical wavenumber response and hence justification of the numerical values of Ri presented (or alternately toning down the reference to specific values such as $Ri = 1/4$ given the estimates are noisy and not fully resolved),

We are unsure what the reviewer is asking about here. Do you refer to the vertical wave number response that Polzin et al. (2002) refer to when estimating turbulent mixing processes from vertical shear in the ADCP? We do not use the ADCP to calculate mixing, we only use it to get information on the horizontal velocity in the vicinity of our microstructure shear measurements. The Richardson number is calculated from the vertical shear between successive 8 m tall (in the vertical) ADCP bins, but this is not used for the turbulent shear calculations.

- 54 • justification for use of median versus mean

55 We use the median as it is less impacted by outliers or non-normal
56 distribution of values. If the data is normally distributed the median
57 and mean are identical, so there is no negative effect of using the median
58 as the default method for averaging values.

- 59 • and finally and perhaps most substantively, an explanation for why the
60 turbulent heat fluxes just above the bottom are important to measure.
61 Ie, is that the water that will eventually meet the grounding line, or
62 should the study have been done nearer the top of the mCDW water-
63 mass where the gradients and heat losses are much stronger?

64 Have added information to the text to clarify that we measure heat
65 fluxes close to the bottom to capture the effect of topography roughness
66 on the flow, to capture the mixing where the bottom intensified warm
67 inflow interacts with the seabed and due to practical constraints (the
68 ALR needs to stay within 100 m of the seabed to allow for accurate
69 dead-reckoning and bottom tracking). The schematic we added to the
70 Introduction (see your comment below) should also make the reason for
71 our interest in the lower mCDW clearer. We have added the following
72 sentence to the introduction: “our study targets the current of warm
73 mCDW flowing into the ice shelf cavity and maintains a dive track
74 close to the seabed. We investigate the circulation and mixing in the
75 mCDW inflow close to the bed of the cavity to understand the effect
76 of bathymetry on mixing and circulation. We quantify the upward
77 heat transport that cools the mCDW in the deepest part of the cavity
78 whilst warming the overlying mCDW (which can access the grounding
79 line and the ice shelf base; Figure 1), and investigate drivers for the
80 observed mixing. ”

81 Good luck. I enjoyed reading the paper and hope that these comments are
82 useful.

83 11: topography, turbulent or both not resolved?

84 We have clarified this sentence to confer that turbulent mixing is not re-
85 solved in models and topography is not resolved in bathymetry products or
86 models. The sentence now reads: “We show a highly complex spatial pat-
87 tern of turbulent mixing and of bottom topography. The bottom topography
88 is currently not resolved in bathymetry products and both the topography
89 and turbulent mixing are currently not resolved in models of ice-shelf–ocean
90 interactions.”

91 26: awkward

92 The sentence now reads “ The mCDW can cause melting at the grounding
93 line, leading to basal mass loss and grounding line retreat.”

94 35, 53: “this” is a weak reference. Please reword; see Strunk and White if
95 needed.

96 the sentences now read “ The depth at which meltwater enters the ocean is
97 influenced by where melt predominantly occurs. ”and “Due to the remote
98 location and difficult access, measuring turbulent kinetic energy dissipation
99 rate in ice shelf cavities is only now starting to become feasible.”, respectively.

100 48: Melt rates two words?

101 We have corrected this

102 52: This statement is actually not true: epsilon is the dissipation rate and
103 further assumptions must be invoked to infer the mixing. This needs to be
104 corrected and expanded upon.

105 We have clarified that ε is only a measure of turbulence if the turbulence is
106 isotropic. The sentence now reads “The turbulent kinetic energy dissipation
107 rate, ε , is the rate at which molecular viscosity dampens isotropic turbulence
108 generated at large scales by e.g. vertical or lateral shear, and is used to
109 quantify turbulent mixing.”

110 55: This would be a good place to distinguish what is different about this
111 study from the other two.

112 Thank you for your comment, the paragraph in question now reads: “To
113 our knowledge, there exist two published studies of mixing in an ice-shelf
114 cavity measured by an underwater vehicle, one under Pine Island Glacier
115 (Kimura et al., 2016), and one under the Filchner Ronne Ice Shelf (Davis
116 et al., 2022). We present a third such study, targeting DIS. DIS and Pine
117 Island Ice Shelf experience low tidal flows, whereas Filchner Ronne Ice Shelf
118 experiences strong tidal flows. Unlike Davis et al. (2022) and Kimura et al.
119 (2016), our study targets the current of warm mCDW flowing into the ice
120 shelf cavity and maintains a dive track close to the seabed. We investigate the
121 circulation and mixing in the mCDW inflow close to the bed of the cavity to
122 understand the effect of bathymetry on mixing and circulation. We quantify
123 the upward heat transport that cools the mCDW in the deepest part of the
124 cavity whilst warming the overlying mCDW (which can access the grounding
125 line and the ice shelf base; Figure 1), and investigate drivers for the observed
126 mixing. ”

127 56: which \rightarrow that. Also, is this the only reason mixing is important to
128 know for these situations?

129 Thank you, we have corrected that. Mixing at the seabed – ocean interface
130 is also important for nutrient transport, such as the transport of iron from
131 sedimentary sources to the euphotic zone. We refer to such processes in
132 the paragraph above: “The input of meltwater to the Amundsen Sea is also
133 important for biological activity in the region. The flow of mCDW along the
134 seafloor on its way into the DIS cavity enriches the mCDW in dissolved iron
135 and manganese while the meltwater from the ice shelf itself is a source of
136 particulate iron and manganese (van Manen et al., 2022). The addition of
137 glacial meltwater makes the outflowing mCDW more buoyant than the dense
138 mCDW inflow, transporting iron and manganese to the surface ocean (van
139 Manen et al., 2022) where they are important micronutrients for primary
140 producers (Twining & Baines, 2013).”

141 66-68: Please give order of magnitude of the clock offsets before correction
142 and the precision of the alignment afterwards.

143 We have added this information in the revised manuscript. The paragraph
144 now reads: “A clock offset of approximately 2 minutes between the ALR
145 CTD and the MicroRider was resolved by calculating lagged correlations
146 between the MicroRider pressure sensor and the CTD pressure sensor to find
147 the offset, then correcting for the identified clock offset and drift. ”.

148 74: Please explain why you used median instead of mean?

149 See our explanation above.

150 95: Could indicate this is likely because of $F = ma$; ie the same force on the
151 huge autos produces much smaller accelerations.

152 Thank you for this prompt, the revised sentence now reads: “ Unlike mi-
153 crostructure measurements performed with a small, light-weight AUV (e.g.
154 Kolås et al., 2022), the shear microstructure recorded on AutoSub Long
155 Range was not critically impacted by vehicle vibrations, possibly due to its
156 greater mass.”

157 105: on which this study focuses.

158 Thank you, we have made the correction.

159 105 general: is this the first paper that presents the details of shear mi-
160 crostructure from Autosub? Surprising if so but if true, you might consider
161 showing a few spectra and additional details, possibly in an appendix, so
162 that future work can cite this paper.

163 This is not the first such paper, we refer the reader to Davis et al. (2022) for
164 information of the spectral response of the shear probes on ALR. We have
165 added the sentence “The shear power spectra from a MicroRider mounted
166 on an ALR have been described in detail in Davis et al. (2022).” to the
167 manuscript.

168 111: Shih et al is a very bad reference for this! They find a Re_b -dependent
169 Gamma. Suggest just citing Osborn (1980). There are also now a handful of
170 observational references supporting the assertion that $\Gamma = 0.2$.

171 Thank you for pointing this out, we have removed the reference to Shih.
172 The sentence now reads “ $T = 0.2$ is the mixing efficiency, a measure of the
173 amount of available turbulent kinetic energy that is permanently converted
174 to potential energy by turbulent mixing, which is generally set to 0.2 (Osborn,
175 1980)”

176 113: How close to the bottom of the ice is the shallowest CTD measurement
177 shown? The very strong gradients at the very top of the cavity CTD casts
178 (Fig 2 black) are interesting.

179 The CTD cast goes right to the ice – ocean interface. We refer the reader
180 to A. Wåhlin et al. (2024) for a discussion of the CTD measurements at the
181 interface.

182 123 and throughout: I believe units should be in roman, not italicized, font.

183 We have corrected this where we found such instances, all remaining format-
184 ting will be finalized in the copy editing process.

185 136: Suggest reformatting the equation.

186 We have reformatted the equation.

187 140: Please make it very clear that Ri (under the ice at least) is based on a
188 single N^2 profile whereas the shear is a function of location and time. This
189 is OK, but appropriate caveats as to its governing local instabilities without
190 in-situ N^2 should be given.

191 Thank you for this comment, we have added the following words to the
192 paragraph describing Ri : “Thus, Ri is calculated from a constant value of N^2 ,
193 based on a single profile in the cavity, and shear is a function of space and
194 time along the track of the ALR. Variations of Ri due to variations in N^2
195 are not captured. For constant N^2 , Ri is low in areas of high shear.”

196 173: Generally, avoid “there is” in favor of more active language such as
197 “flow is to the ...”

198 We will change some of our wording where we deem appropriate in the revised
199 manuscript. The sentence in question here has been reworded to “A bottom
200 intensified southward current flows into the cavity in the east, between the
201 400 m and 900 m isobaths, and a shallower, bottom intensified northward
202 current flows out of the cavity in the west (Figure 3c).”

203 177: High compared to what?

204 This sentence has been rephrased to read “Below 500 m depth, turbulent ki-
205 netic energy dissipation is elevated (compared to other areas below 500 m
206 along the ice front) in the inflow. Turbulent kinetic energy dissipation is
207 $\approx 10^{-8} \text{ W kg}^{-1}$ in the inflow over an area approximately 7 km wide and 200 m
208 high (Figure 3d; turbulent kinetic energy dissipation rate is elevated between
209 38 km and 45 km of the ice front and ~ 200 m above the seabed).”

210 177: runon sentence.

211 In addition to adding context (see above), this sentence has been split into
212 shorter sentences.

213 Figure 3, lines 2 and 4 of caption: runon sentences. Also, the dots are said
214 to indicate the starting locations - but they are a continuous line. I’d have
215 thought there would just be two starting locations, one for center and one
216 for east? Please clarify.

217 The new Figure 3 has shorter sentences in the caption and the dots are
218 described as: “10-minute medians of the values measured by the ALR are
219 shown as coloured dots in panels a-d. The two dots with bold outlines show
220 the starting locations of the ALR east and centre short dive tracks into the
221 cavity.”

222 Figure 4: Personally I think it would be better to keep the aspect ratio
223 constant between Fig 3 and 4. Also, since you already plotted velocity in
224 Figure 3, suggest including a panel of N2. The aspect ratio is all the more
225 a problem later when the authors are comparing epsilon to the different
226 instability indicators - but the reader must go back and forth between figure

227 3 and 4. Suggest standardizing the aspect ratio and including an epsilon
228 panel in Figure 4. Possibly even adding Ri contours to the epsilon panel or
229 epsilon contours to the Ri panel since the authors are trying to demonstrate
230 correspondence between the two quantities.

231 Thank you for this feedback, we have combined Figure 3 and 4 into the
232 new Figure 3, in which all panels have the same aspect ratio. We have also
233 included a panel of N2 at the ice front. We have not plotted Ri contours on
234 the epsilon panel, as that proved to be confusing (switching between density
235 contours and Ri contours).

236 Also, the Ri panel is just a big sea of red. Consider plotting something else
237 to highlight the unstable regions such as Ri^{-1} or $Fr = Uz/N$.

238 The Ri panel is mainly red due to the choice of colourbar. We chose to plot
239 $Ri < 1/4$, $1 > Ri > 1/4$, and $Ri > 1$ as three different colours in keeping
240 with established practice (e.g. Dotto et al., 2025) to distinguish along criteria
241 for instability. Plotting $1/Ri$ would make it less obvious where $Ri < 1/4$. We
242 would like to avoid plotting additional instability metrics such as the Froude
243 number to avoid confusion.

244 182: Doesn't negative PV mean unstable? The whole water column is unsta-
245 ble? Is it backwards in the southern hemisphere? Some statements to clarify
246 would be useful.

247 We have clarified this in the text by adding the sentence: "Instabilities may
248 develop when potential vorticity and f have opposite signs, as f is nega-
249 tive in the southern hemisphere, potential vorticity > 0 indicates conditions
250 favourable to instability. "

251 188: I don't agree with this statement - the high dissipation does not appear
252 to me to line up at all with for Ri. Furthermore, given the ADCP's finite
253 vertical resolution and noise, some additional detail needs to be given on
254 how seriously we are to take the numerical value of Ri. I think that either
255 some wavenumber spectra and transfer functions a la Polzin 2002 need to be

256 included, or Ri used as a qualitative indicator.

257 As far as we understand Polzin et al., 2002 the vertical wavenumber re-
258 sponse of the ADCP is relevant when calculating turbulent dissipation from
259 the ADCP. We are not using the ADCP for turbulence. We use the VMP or
260 microrider for shear microstructure and the LADCP and ADCP on the ALR
261 to get an idea of the vertical and horizontal structure of the water column
262 at much larger scales, a background value if you will. Ri and other stability
263 criteria are frequently calculated from LADCP output with bin sizes of 8 m
264 and used in comparisons with microstructure data (e.g. Dotto et al., 2025;
265 Naveira Garabato et al., 2017; Naveira Garabato et al., 2019). We have clar-
266 ified our reference to Ri and mixing, the paragraph now reads: “The region of
267 high turbulent kinetic energy dissipation rate ε in the inflow (Figure 3d) co-
268 incides with instances of $R_i < 1/4$ captured at 40 km (Figure 3h), indicating
269 conditions favourable to turbulent mixing. Turbulent kinetic energy dissi-
270 pation rate is larger than 10^{-8} here, one to two orders of magnitude higher
271 than the background value (Figure 3d). Dotto et al. (2025) found similar
272 results for the outflow of DIS. Although areas of high ε extend beyond areas
273 of $Ri < 1/4$, ε is higher and Ri is lower in the upper watercolumn and close to
274 the seabed. We observe areas of low Ri and $Ri < 1/4$ that are not associated
275 with high values of ε , e.g. at 25 km along the transect. ”

276 191: I disagree; elevated mixing is much broader than the regions of $Ri < 1/4$
277 - augmenting my previous point.

278 We will clarify that the high epsilon includes, but extends beyond, the region
279 of low Ri . The relevant sentence now reads: “Although areas of high ε extend
280 beyond areas of $Ri < 1/4$, ε is higher and Ri is lower in the upper watercolumn
281 and close to the seabed.”

282 193: This statement is not justified. Epsilon appears surface intensified as
283 well. And while it is bottom intensified, I do not think the statement that
284 it is heightened over rough topography, shear or high currents (of which
285 you generally must choose either high current or high shear, not both...) is
286 supported. And as before, I don't think that high epsilon lines up with low

287 Ri either. Either way, if this statement is retained, more analysis needs to
288 be shown - scatter plots, binned averages, etc.

289 We have clarified that we are only considering epsilon below the Winter Water
290 layer, thus we do not discuss high epsilon at the surface. We have included
291 the following: “Below 500 m depth, turbulent kinetic energy dissipation is
292 elevated in the inflow (compared to other areas below 500 m along the ice
293 front). ” With regards to the ice shelf front, we have changed our statement
294 to read: “Our observations show turbulent mixing to be patchy, bottom
295 intensified and to coincide with high velocities (Figure 3).” We maintain that
296 in the cavity high epsilon is associated with high shear and low Ri and have
297 added correlations plots that show this to the manuscript. See more below.

298 197: runon sentence. And seemingly unrelated sentences. Ri governs shear
299 instability, not symmetric instability. . . (I understand they are highly corre-
300 lated here, but they are different, so clarification is needed).

301 We will insert a paragraph break before ”at the nearby Pine Island Ice
302 Shelf...”. The start of the new paragraph now reads: “ At the nearby Pine
303 Island Ice Shelf (PIIS) Naveira Garabato et al. (2017) conducted ADCP and
304 VMP transects along the calving front. Naveira Garabato et al. (2017) do
305 not detect a fast, narrow, turbulent inflow current, unlike what we observed
306 at DIS (Figure 3). High rates of turbulent kinetic energy dissipation below
307 the WW were mostly confined to the PIIS outflow. The PIIS is connected to
308 another ice shelf cavity to the north and may receive some of its inflow from
309 under this neighbouring ice shelf, which may decrease the inflow across the
310 PIIS front and possibly the turbulent mixing there. Additionally, the ice shelf
311 draft of the PIIS is deeper (≈ 400 m) than the DIS (≈ 350 m). The ice shelf
312 draft induces a barotropic jump (an abrupt change in water column thick-
313 ness, blocking flow along constant lines of water column thickness) and limits
314 barotropic inflow to the cavity (A. K. Wåhlin et al., 2020), thus decreasing
315 inflow current velocities and possibly turbulent mixing. ”. The sentence re-
316 garding Ri has been removed, the relevant paragraph now reads “Symmetric
317 instability is driven by high vertical current shear (Figure 3j). The region of

318 high turbulent kinetic energy dissipation rate ε in the inflow (Figure 3d) co-
319 incides with instances of $R_i < 1/4$ captured at 40 km (Figure 3h), indicating
320 conditions favourable to turbulent mixing.”

321 202: What is a barotropic jump?

322 It is an oceanographic term for an abrupt change in water column thickness.
323 This occurs at the ice shelf front, since ocean currents want to flow along lines
324 of uniform water column thickness, the ice shelf draft poses a barrier to flow,
325 even at depths deeper than its draft. We have added a parenthetical
326 “The ice shelf draft induces a barotropic jump (an abrupt change in water
327 column thickness, blocking flow along constant lines of water column thick-
328 ness) and limits barotropic inflow to the cavity (A. K. Wåhlin et al., 2020),
329 thus decreasing inflow current velocities and possibly turbulent mixing. ”to
330 the sentence in question.

331 207: Please rewrite this passive and vague sentence.

332 We have rephrased this sentence, it now reads “Because the ALR measure-
333 ments were not coincident in time with the LADCP section, the ALR may
334 have failed to capture transient patches of high turbulent kinetic energy dis-
335 sipation rate present in the LADCP section.”

336 204-210: Suggest moving this speculative bit to the discussion.

337 We originally had results and discussion split, but chose to integrate them to
338 avoid duplicating information and to limit jumping back and forth between
339 topics. We will retain this structure.

340 216: I think it would be nice to compare this to open ocean values at a
341 similar depth and/or abyssal values, for context. Otherwise “weakly stable”
342 doesn’t have meaning.

343 We have added typical open ocean values for N^2 in the Southern Ocean.
344 The sentence now reads: “We estimate N^2 below a depth of 900 m to be
345 $6 \times 10^{-7} \text{s}^{-1}$. This is about three orders of magnitude lower than typical open

346 ocean values for the southern ocean (King et al., 2012), indicating weakly
347 stable stratification in the cavity. ”

348 218: Style guides such as Strunk and White suggest avoiding “Figure x
349 shows...” in favor of “statement x is true (Figure y).”

350 We have changed this sentence to read: “In the cavity, the ALR detected
351 currents that flow predominantly southeastward with low vertical shear in
352 the east dive track, and a more mixed pattern in the two centre dive tracks
353 (Figures 5 and 6).”

354 223: Figure 6 and 5 -> Figures 5 and 6

355 Thank you for pointing out this typo, it has been corrected.

356 236: redundant. Suggest “Maximum values were” or “Values reached.”

357 We have changed the sentence to read: “Maximum values were $10^{-7} \text{ W kg}^{-1}$
358 (ϵ) and $10^{-2} \text{ m}^2 \text{ s}^{-1}$ (κ), respectively (Table 2).”

359 238: Again, I’m afraid I don’t see this. There are counter examples where
360 epsilon is high over flat bottoms. Please include plots that allow direct com-
361 parison such as plotting epsilon with Ri, current speed or bathymetric slope
362 over plotted, or scatter plots or binned averages (e.g. epsilon(Ri) etc) if you
363 want to make this claim.

364 We have included scatter plots in our revised manuscript. We stress that
365 our data are extremely noisy and thus correlation coefficients are low even if
366 relationships are statistically significant to the 0.1% level. Additionally, the
367 bathymetry gradient deeper in the cavity is affected by the low resolution
368 of BedMachine, preventing us from fully resolving the relationship between
369 bathymetry and epsilon.

370 241: Please remind reader that it’s Ri computed from in-situ where and

371 The remainder of this sentence seems to be missing, please clarify

372 245: Again, please include transfer function and instrument response infor-

373 mation if you wish to quantify the numerical value of Ri versus using it as a
374 qualitative indication. Note as well that these transfer functions and hence
375 the mapping of true to measured Ri will be different for the Autosub and the
376 LADCP.

377 Can you clarify what you mean by the mapping of true to measured Ri and
378 how that is influenced by the ADCP? It is common practice to calculate Ri
379 from vertical shear from the 8 m binned ADCP and we do not use the ADCP
380 to calculate fine scale microstructure.

381 257: Is it really necessary to use a package like this to compute a spatial
382 gradient? More fundamentally I do not see a relationship between RMS
383 bathymetric slope and dissipation rate.

384 We do not use a package to calculate the gradient. The bathymetry from ALR
385 is only 1D, to get a 2D gradient we use BedMachine to get the bathymetry
386 normal to and along the ALR dive track. Can you clarify what you mean by
387 "RMS bathymetric slope"? We do not calculate RMS of the slope and to our
388 eyes there is a clear relationship between the bathymetry and epsilon close to
389 0 km on the east dive track. We have included scatter plots, linear fit lines,
390 correlation coefficients and p-values for the relationship between bathymetric
391 slope and epsilon which clearly show a strong connection.

392 264 onwards: consider moving all of this comparison to past work to the
393 discussion, so that the results section just has your results?

394 A previous draft had results and discussion separated and the feedback from
395 several readers was that this caused unnecessary confusion, duplication and
396 jumping back and forth. We will keep the results and discussion merged.

397 270: I'm confused here, sorry. Weren't the ALR measurements entirely in
398 the warm inflow, since they were so deep?

399 We will clarify this sentence, you are correct that all our ALR measurements
400 are in the warm layer of in the cavity, but we define the inflow as the narrow
401 bottom intensified current along the 700 m isobath. The sentence now reads:

402 “We observed our highest mixing values in the bottom intensified inflow to
403 the cavity, whereas Kimura et al. (2016) observed the highest levels of mixing
404 close to the grounding line. Our ALR dive tracks did not reach the grounding
405 line, and the dive tracks of Kimura et al. (2016) did not cover the inflow of
406 the PIIS, making comparison difficult. Naveira Garabato et al. (2017) did
407 not find enhanced mixing in the PIIS inflow. ”

408 272: runon sentence.

409 The sentence in line 272 is not that long, are you sure that this is the line
410 number you meant?

411 273: Due to what mechanism? This sentence has been modified to read:
412 “Kimura et al. (2016) hypothesised that high (horizontal) density gradients
413 driven by temperature differences and a bathymetric ridge can drive a baro-
414 clinic current with strong vertical current shear. This high shear in turn
415 drives high levels of turbulence at the ridge under PIIS. Our study shows
416 that high density gradients are not a requirement for high levels of turbu-
417 lence.”

418 281: Please change “this” to “their” to avoid confusing with your study.

419 This change has been made.

420 285: If you are going to state dissipation rates this low, I think you do need
421 to demonstrate your minimum detectability threshold. Earlier you said it
422 was $1\text{e-}10$. So how then do you get a median lower than this.

423 The detection limit is between 10^{-11} and 10^{-10} depending on the dive track.
424 We never state that the detection limit is 10^{-10} and have clarified the sen-
425 tence you refer to. The paragraph now reads: “Smaller, narrower peaks at
426 frequencies below 10Hz in the accelerometer spectra are successfully removed
427 by the Goodman method for dissipation rates above $1 \times 10^{-8} \text{ W kg}^{-1}$. De-
428 viations from the fitted Nasmyth spectra remain for dissipation rates below
429 1×10^{-9} , arguing that quantitative estimates of dissipation rate in very qui-
430 escent regimes are not as reliable as estimates of high dissipation rates. Indi-

vidual dive tracks show good agreement between shear spectra and Nasmyth spectra for dissipation rates lower than $1 \times 10^{-10} \text{ W kg}^{-1}$. Where dissipation rates calculated from two orthogonal shear probes show good agreement, we are confident in reporting dissipation rates down to $1 \times 10^{-11} \text{ W kg}^{-1}$. Additionally, any signal in the shear spectra caused by the AUV motion, and not removed by the Goodman filter, will have minimal effects on the spatio-temporal pattern of high and low ε observed by the ALR or the qualitative assessment of these patterns, on which this study focuses.”

Again, I think median should be avoided for all quantities unless there is a good reason. Why not just use the mean?

We do not want to cause confusion by switching between mean and median for data with and without outliers or non-normal distributions, since $\text{median}(x) = \text{mean}(x)$ when x is normally distributed we think median is a better choice.

333: The reason for these calculations is revealed here - suggest giving it earlier to make the reader understand why they are being told all of this. More fundamentally, is that the only reason turbulence is important to measure under ice shelves? Ie, as a possible mitigator of the advective heat flux by these warm flows?

We have added the following paragraph to the introduction together with a schematic of the ice shelf cavity: “Basal melt under Dotson is highest close to the grounding line of the Kohler East (often referred to as Smith West) and Kohler West glaciers (Khazendar et al., 2016; Gourmelen et al., 2017). The Kohler West grounding line lies at the southern end of the dashed path shown in Figure 1a. A cross-section of the cavity along the path (Figure 1b) shows an idealized view of the cavity circulation under the Dotson Ice Shelf. Warm water entering the cavity in the east, and traveling along a path shallower than the 830 m deep sill (Jordan et al., 2020), can reach the grounding line. Warm water that reaches the grounding line causes high basal melt and grounding line retreat (Khazendar et al., 2016; Gourmelen et al., 2017). The sill may limit direct access of the deepest and warmest mCDW to the

462 grounding line (Jordan et al., 2020; Khazendar et al., 2016). The addition of
 463 meltwater to the warm, salty mCDW forms a buoyant plume which travels
 464 along the underside of the ice before exiting the cavity in the west. Along its
 465 path, the water experiences turbulent mixing which can transport heat and
 466 salt upward, modifying the properties of the water which ultimately interacts
 467 with the grounding line, and the properties of the buoyant plume exiting the
 468 cavity.” As far as the ice melt rate and modelling efforts in the cavity are
 469 concerned the heat flux is the major concern. As we discuss above, mixing
 470 is also important for the trace metal and nutrient transport, however we do
 471 not have measurements of concentration gradients in the cavity and can not
 472 make a statement as to how the mixing influences them.

473 I, at least as a non ice sheet person, would like to see a cartoon (words
 474 or actual graphic) showing a cross section of the hypothesized warm water
 475 flow to the grounding line. The reason for this is that I don’t currently
 476 understand why the study focused so much on the near-bottom mixing. I’d
 477 think that the heat loss out of the mCDW would be better quantified near
 478 its upper edge. As the authors point out, the water near the bottom is
 479 very weakly stratified so the heat fluxes are expected to be small. Aloft
 480 nearer the interface, the gradients would be stronger, but also the distance
 481 from the topography which is presumably generating most of the turbulence
 482 (my comments above about that not having been adequately demonstrated
 483 notwithstanding). So, statements that mixing is weak such as on lines 356-
 484 258 should be tempered somewhat. And I think the cartoon or written
 485 description of the flow giving readers the sense of which depths are thought
 486 the most likely to eventually contact the ice would help inform this discussion,
 487 at least for me.

488 Have added a schematic to the revised manuscript, thank you for the sug-
 489 gestion. We have added the following text to the introduction: “Basal melt
 490 under Dotson is highest close to the grounding line of the Kohler East (often
 491 referred to as Smith West) and Kohler West glaciers (Khazendar et al., 2016;
 492 Gourmelen et al., 2017). The Kohler West grounding line lies at the south-
 493 ern end of the dashed path shown in Figure 1a. A cross-section of the cavity

494 along the path (Figure 1b) shows an idealized view of the cavity circulation
 495 under the Dotson Ice Shelf. Warm water entering the cavity in the east,
 496 and traveling along a path shallower than the 830 m deep sill (Jordan et al.,
 497 2020), can reach the grounding line. Warm water that reaches the ground-
 498 ing line causes high basal melt and grounding line retreat (Khazendar et al.,
 499 2016; Gourmelen et al., 2017). The sill may limit direct access of the deepest
 500 and warmest mCDW to the grounding line (Jordan et al., 2020; Khazendar
 501 et al., 2016). The addition of meltwater to the warm, salty mCDW forms a
 502 buoyant plume which travels along the underside of the ice before exiting the
 503 cavity in the west. Along its path, the water experiences turbulent mixing
 504 which can transport heat and salt upward, modifying the properties of the
 505 water which ultimately interacts with the grounding line, and the properties
 506 of the buoyant plume exiting the cavity.”

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