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First insights into deep convection by the Doppler velocity measurements of the EarthCARE's Cloud Profiling Radar

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Abstract.

Convective updrafts and downdrafts play a vital role in Earth's energy and water cycles by modulating vertical energy and moisture transport and shaping precipitation patterns. Despite their importance, the characteristics of convective motions and their relationship to the near-storm environment remain poorly constrained by observations.

The payload of the recently launched EarthCARE satellite mission includes a 94-GHz Cloud Profiling Radar (CPR) with Doppler capability. In this study, we present first-light CPR Doppler velocity observations in deep convective clouds. These early examples offer a first glimpse into the dynamic nature of cloud systems. The narrow footprint of the CPR helps reduce the impact of multiple scattering and non-uniform beam filling (NUBF) on the Doppler velocity measurements. However, the instrument's low Nyquist velocity presents a significant challenge for recovering the true Doppler velocity profiles in deep convective systems.

The CPR Doppler velocity observations are expected to challenge traditional methodologies for identifying deep convective cores, which typically rely on reflectivity-based thresholds. We showcase examples that demonstrate the synergy between CPR Doppler velocity measurements and geostationary satellite observations, illustrating how their combined use can help capture the evolution of the convective lifecycle.

These results align with EarthCARE's broader mission objectives and highlight the potential of spaceborne Doppler radar to significantly advance our understanding of cloud dynamics and convection in the climate system.

1 Introduction

Deep convective clouds are responsible for the vertical transport of air and water, one of the most influential yet poorly constrained by measurements atmospheric process. Deep convection is crucial in balancing the Earth's heat budget and influencing large-scale weather patterns, including cloud formation and the development of storms and extreme weather (Hartmann et al. 1984). Deep convective events typically occur in tropical regions, but they affect the global atmospheric circulation beyond the tropics via anvil detrainment processes and latent heat release via precipitation (Gasparini et al. 2021; Hartmann et al. 2018).

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A number of microphysical processes are active during convective initiation and development that are not yet well understood or properly implemented in models (Prein et al. 2015; Arakawa 2004; Bony et al. 2015).

Despite the importance of deep convection, several aspects of deep convective clouds remain challenging to represent in high-resolution models and even observations (Fridlind et al. 2017; Ladino et al. 2017). Models also struggle to accurately represent convective updrafts, leaving significant observational gaps (Varble et al. 2014). Surface and airborne radar observations have provided valuable insight into the structure and magnitude of convective updrafts, but the observational record is very sparse and mostly available over land (Giangrande et al. 2013; Wang et al. 2020; Oue et al. 2019; J. Yang et al. 2016; Jeyaratnam et al. 2021; North et al. 2017). On the other hand, satellite observations can provide global coverage and sufficient sampling of deep convection and associated clouds and precipitation (Lee et al. 2021). In particular, the 3-D structure of deep convective clouds has been extensively studied using observations from spaceborne radars.

The Tropical Rainfall Measuring Mission (TRMM), developed by the National Aeronautics and Space Administration (NASA) and the National Space Development Agency of Japan (NASDA), introduced the first spaceborne radar in space, a 13.8 GHz Precipitation Radar (PR) (Kummerow et al. 1998; Kummerow et al. 2000). The TRMM PR was operational from 1997 to 2015 and advanced our understanding of tropical convection and associated rainfall (Yokoyama et al. 2014; Xu et al. 2012). Studies using the TRMM PR data analyzed convective system structures, including diurnal cycles and vertical profiles (Hamada et al. 2015). TRMM's success led to the Global Precipitation Measurement (GPM) mission launched in 2014 by NASA and the Japan Aerospace Exploration Agency (JAXA) which enhances TRMM's capabilities with improved resolution and higher latitude coverage (Skofronick-Jackson et al. 2017). The GPM mission features a Dual-frequency Precipitation Radar (DPR) that operates at Ku (35.5 GHz) and Ka (13.6 GHz) bands, providing 3D precipitation structures (Skofronick-Jackson et al. 2018). Studies using GPM PR data show deep convection reaching the tropopause predominantly over land, consistent with TRMM findings (Liu et al. 2020; Battaglia et al. 2020; Liu et al. 2016). Ni et al. 2019 analyzed ice microphysical properties, revealing larger ice particles and higher ice water content in land-based deep convective cores. The poor sensitivity of the PR and DPR limited their ability to capture the 3D structure of the upper-level tropospheric cloud structures.

The CloudSat-CALIPSO mission (Stephens et al. 2002), part of NASA's A-Train since 2004, provided detailed cloud vertical structures. Its Cloud Profiling Radar (CPR) with 240 m vertical resolution captured convective cloud features, aiding studies on convective cores and updrafts (Takahashi et al. 2017a). Findings indicate stronger convective cores and lower entrainment rates over land, enabling higher-altitude particle transport. However, CloudSat's narrow along-track sampling (1.4 km cross-track) limits representation of spatially heterogeneous deep convective cores (DCCs). To mitigate biases, CloudSat data have been integrated with passive sensors, such as MODIS cloud top temperature, for improved convective characterization (Luo et al. 2008; Luo et al. 2010; Luo et al. 2014; K. Yang et al. 2023).

Finally, in May 2024, the Earth, Cloud, Aerosol and Radiation Explorer (EarthCARE, Illingworth et al. 2015), a joint European Space Agency (ESA) and JAXA mission was successfully launched. The EarthCARE mission aims to improve cloud-aerosol-radiation interaction studies and enhance numerical weather prediction (NWP) models and climate simulations. EarthCARE carries a 94-GHz Doppler Cloud Profiling Radar (CPR), High-Spectral Resolution Lidar (ATLID), Multi-Spectral Imager (MSI), and Broad-Band Radiometer (BBR). Launched after CloudSat-CALIPSO ended operations in 2023, EarthCARE

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benefits from improved radar sensitivity due to its lower orbit and having all instruments on the same platform (Illingworth et al. 2015; Wehr et al. 2023). Most importantly, the EC mission features the first spaceborne radar with Doppler capability (Kollias et al. 2018b; Kollias et al. 2014a; Kollias et al. 2022a). The availability of Doppler measurements from space will offer a unique opportunity for the collection of a global dataset of vertical motions in clouds and precipitation. This global data set is expected to improve our understanding of convective motions in clouds and help evaluate current parameterizations of convective mass flux in cloud resolution models (Manabe et al. 1964; Tiedtke 1989; Bechtold et al. 2001).

Here, a first assessment of the performance of the EC CPR Doppler velocity measurements in deep convection is presented. The main objectives of this study are to describe and interpret convective cores as observed by the EC-CPR, leveraging joint Doppler velocity and reflectivity measurements, and to compare these observations with geostationary data. For the first time, Doppler velocity from a spaceborne radar is used to identify and characterize convective cores, providing insights into their internal dynamics and updraft structures (Kollias et al. 2023). Coincident MSI observations are compared with geostationary MSG (Meteosat Second Generation) imagery to assess the capability of passive sensors in detecting convection and tracking its evolution.

2 CPR Doppler velocity observations in deep convection

One of the most exciting new measurement capabilities of the EarthCARE mission is the CPR Doppler velocity measurements. Several factors are expected to impact the quality of the CPR Doppler velocity measurements (Tanelli et al. 2002; Tanelli et al. 2005; Kollias et al. 2014b; Kollias et al. 2018a; Kollias et al. 2022b). The EarthCARE satellite speed of 7.6 ms^{-1} introduces significant broadening (decorrelation) of the CPR phase measurements that results in significant uncertainty in the Doppler velocity estimates (Kollias et al. 2014b; Kollias et al. 2022b). Antenna mispointing is another source of uncertainty (Tanelli et al. 2005; Battaglia et al. 2014; Puigdomènech Treserras et al. 2025). In deep convection, additional factors such as attenuation, multiple scattering (Battaglia et al. 2008; Battaglia et al. 2010; Battaglia et al. 2011c), non-uniform beam filling (Tanelli et al. 2002; Kollias et al. 2022b), and aliasing (Sy et al. 2014) can have a significant impact on the observed Doppler velocities and introduce considerable uncertainty and biases.

An example of CPR observations of a deep convective system is shown in Fig. 1. The CPR observations were collected on September 18, 2024, over Western Africa on a descending (daytime) orbit. Here, CPR Level 2a (L2a) C-PRO data products are used (Kollias et al. 2023). These products are derived from the CPR Level 1b data plus auxiliary meteorological data. The L2a C-PRO data product was released available to the research community on March 2025 (Eisinger et al. 2023). The CPR reflective image (Fig. 1a) illustrates the vertical structure of a wide deep precipitating system. CloudSat-based studies of deep convection mainly use the reflectivity profile features near cloud top to identify deep convective cores (DCC, Takahashi et al. 2012; Luo et al. 2014; Takahashi et al. 2017b; Stephens et al. 2024). The underlying reasoning is that the overshooting of radar reflectivity is an indicator of the larger-size particles pushed higher up; this is only possible with the presence of strong rising updrafts. Three criteria are commonly adopted for convection identification (Takahashi et al. 2014):

- CPR cloud mask (2B-GEOPROF product) greater than 20;

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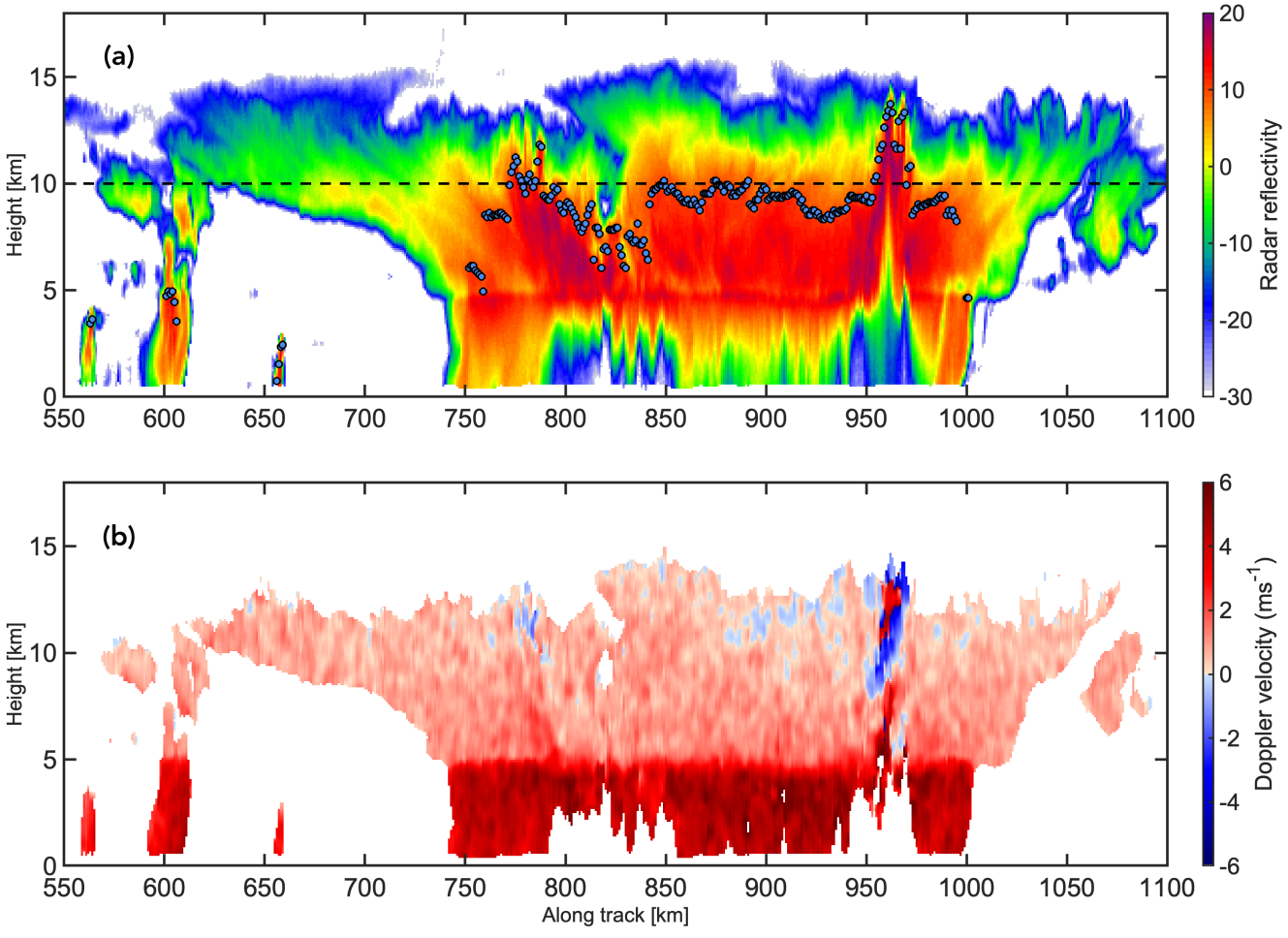

Figure 1. *(a) CPR reflectivity during a large-scale, deep precipitating system with embedded convection observed on September 19, 2024 over the Tropical Western Pacific (Frame 1760E). The horizontal line indicates the 10 km height, and the blue circles indicate the maximum height where a dBZ value above +10 dBZ is observed. (b) the CPR Doppler velocity measurements after a 4-km along-track integration (Kollias et al. 2023). Positive Doppler velocities indicate hydrometors' movement towards the ground.*

- A continuous radar echo from below 2 to above 10 km, thus a thick cloud deck;
- The 10 dBZ echo top height which is indicative of the level where large size particles are lofted by strong convection (Luo et al. 2008) above 10 km.

In Fig. 1a the 10 dBZ echo top height is very close to the 10 km height for a significant part of the deep precipitating system. In two areas (775-790 km and 950-975 km along track), the 10 dBZ echo top height is well above the 10 km height and closer

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to the cloud top height. Luo et al. 2014 introduced a fourth criterion for detecting DCCs, which requires that the 10 dBZ echo top height be within 2 km of the cloud top height determined by the CPR.

The CPR Doppler velocity measurements for the same event can assist into evaluating these different methodologies for identifying DCCs. Figure 1b shows the CPR Doppler velocity averaged over a 4-km along-track distance. The CPR Doppler velocity measurements are shown only in areas where there CPR reflectivity exceeds -21 dBZ. The native CPR along track resolution is 500 m, thus, a total of nine CPR Doppler velocity estimates (their respective real and imaginary parts of the lag-1 pulse pair estimator) have been averaged (Kollias et al. 2023). The averaging of the pulse-pair Doppler velocity estimator lag-1 real and imaginary parts is immune to velocity folding. Before the along track averaging, the CPR Doppler velocities have been corrected for antenna mispointing (Puigdomènech Treserras et al. 2025) and non-uniform beam filling (NUBF) Doppler velocity biases (Kollias et al. 2014b; Sy et al. 2014).

The nadir-pointing CPR Doppler velocity V_D represents the sum of the vertical air motion W_{AIR} and the reflectivity-weighted Doppler sedimentation velocity of the hydrometeors V_T^D:

$$V_D = W_{AIR} + V_T^D. \tag{1}$$

The V_T^D term can only take positive values (downward motion) while the W_{AIR} term can take both positive (downdraft) and negative (updraft) values. The majority of the observed V_D in Fig. 1b are positive. This implies that the V_T^D magnitude is higher than that of the embedded W_{AIR} updrafts. This suggests the presence of negligible vertical air motions ($|W_{AIR}| < 2$ ms^{-1}). A typical example profile of the CPR Doppler velocity and corresponding radar reflectivity is stratiform precipitation conditions is shown in Fig. 2. The most pronounced V_D feature is its melting layer signature just below 5 km height that indicates the phase change from the slowly falling solid ice/snow particles to the fast falling liquid raindrops around the 0°C isotherm (Fig. 2a). The 1-km CPR-averaged Doppler velocity profiles exhibit the same trend but exhibit considerable fluctuations (Kollias et al. 2014b). The noisiness of the CPR 1-km averaged Doppler velocities makes the estimation of the hydrometeors' size and/or density at the 1-km resolution challenging (Kollias et al. 2022b; Mroz et al. 2023). The melting layer signature is also evident in the CPR reflectivity profile with a pronounced increase around the 0 °C isotherm (Fig. 2b). The ice-to-rain Doppler velocity transition is a well-known feature of the Doppler velocity in cold-rain systems, routinely observed by ground-based and airborne Doppler radars (among others, Fabry et al. 1995; Heymsfield et al. 2010), but for the first time with EC it is possible to see it from space.

The CPR short wavelength ($\lambda = 3.2$ mm) and Pulse Repetition Frequency (PRF) determine the CPR Nyquist velocity (V_N) folding ($V_N = \lambda PRF/4$, yellow lines in Fig. 2a). The V_N is the maximum unambiguous velocity that can be detected by the CPR without aliasing (or velocity folding). If the V_D exceeds V_N then folding occurs. During stratiform conditions, in the ice layer, velocity folding is rare even for the 1-km CPR Doppler velocity estimates (Fig. 2a). Below the melting layer, V_T^D can reach values up to 6.5 ms^{-1} (Kollias et al. 2022c). Here, velocity folding can occur especially in the 1-km CPR Doppler velocity estimates, which are altogether noisier. In Fig. 2a the 1-km Doppler velocity estimates outside the V_N boundaries have already been corrected for velocity folding. The assumption used for the unfolding is that negative Doppler velocities below

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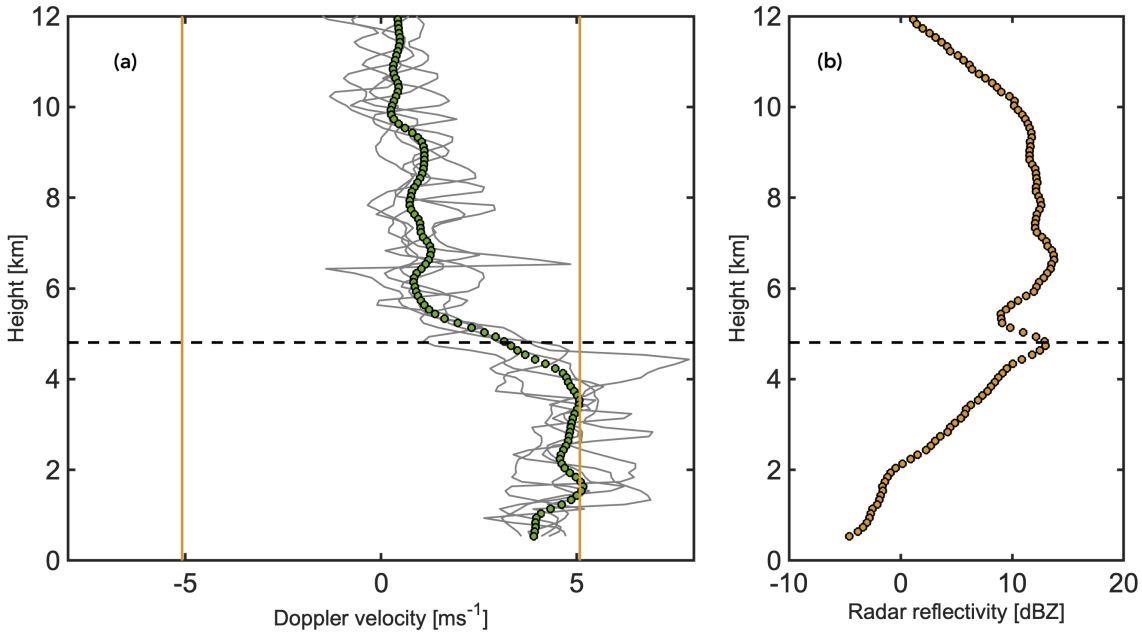

Figure 2. *(a) the CPR Doppler velocity profiles at along track distance of 875 km. The 4-km CPR Doppler velocity estimate is shown in green circles and the 1-km Doppler velocity estimates within a 2 km distance from 875 km are shown in gray lines. The yellow vertical lines indicate the CPR Nyquist velocity and the horizontal dashed line indicates the melting layer height. (b) the corresponding CPR reflectivity at along track distance of 875 km.*

the melting layer in a stratiform precipitation profile are the results of V_D exceeding $+V_N$. Subsequently, all negative V_D values below the 0°C isotherm are unfolded by adding 2 V_N to them.

The interpretation of the CPR Doppler velocity profile in deep stratiform layers serves as a baseline for understanding convective updrafts. In Fig. 1b, updrafts are depicted as regions with negative (upward) 4-km-averaged V_D estimates in cold temperatures (Fig. 1b). Not including the along-track interval 950-975 km, the clusters of negative V_D are located near the cloud top. Since ice particles are smaller at colder temperatures, it is plausible that near cold cloud tops, weak gravity waves and updrafts contribute to an overall negative (upward) CPR Doppler velocity signal. The estimation of air vertical velocity, W_{AIR}, requires knowledge of the Doppler terminal fall speed V_T^D. An estimate of the V_T^D can be provided by the sedimentation velocity best estimate variable in the L2A C-CD data product (Kollias et al. 2023). Interestingly, two regions with 10 dBZ echo top height well above the 10 km altitude exhibit such dynamical features. At 950-975 km along-track, a deep and coherent dynamical structure is observed, characterized by strong upward motions extending from 8 to 14 km. This vertically oriented feature represents a deep convective updraft and is collocated with the highest 10 dBZ echo top heights. The W_{AIR} within this convective updraft is strong enough to cause velocity folding, depicted as a red patch of Doppler velocities embedded within the negative Doppler velocity cluster.

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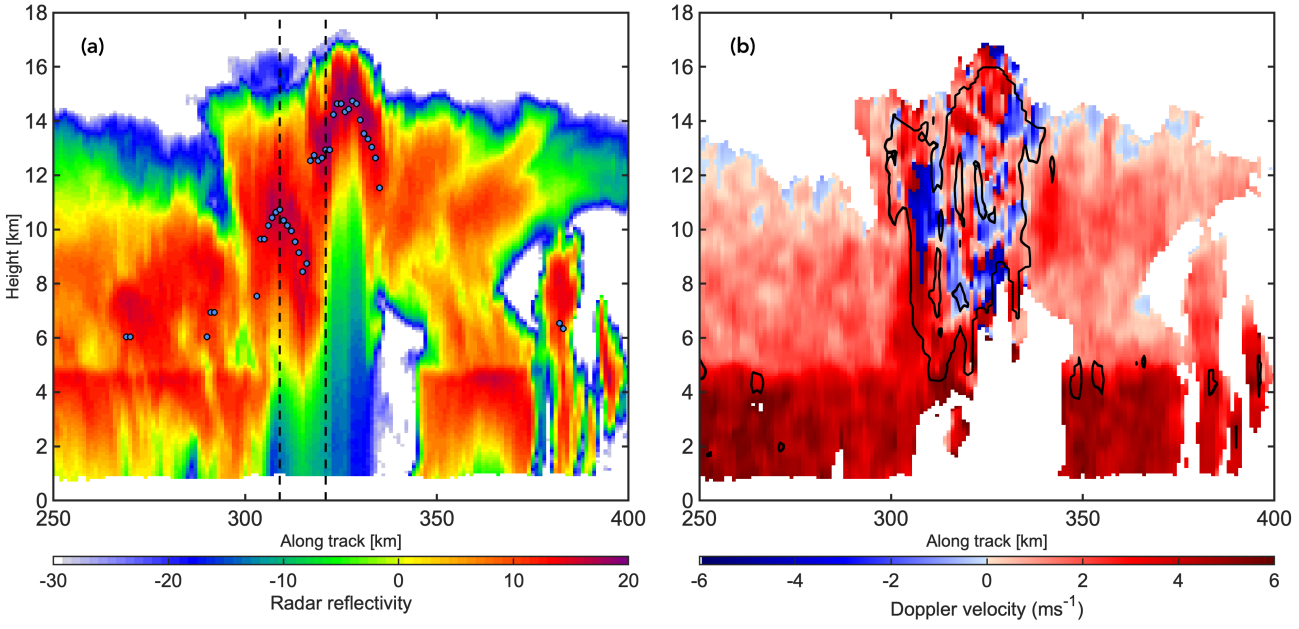

Figure 3. *(a) CPR reflectivity during a deep convective event system on September 18, 2024 over Western Africa (Frame 1752E). The blue circles indicate the height where multiple scattering effects become important. The vertical dashed lines indicate the locations where CPR profiles will be shown in later figures. (b) the CPR Doppler velocity measurements after a 4-km along-track integration (Kollias et al. 2023). Positive Doppler velocities indicate hydrometors' movement towards the ground. The black contour indicates the area where the 4-km CPR Doppler velocity standard deviation exceeds 2 ms^{-1}. A box of 3 km along-track by 2 km in range is used for the estimation of the standard deviation.*

The complexity of the V_D profiles in deep convection is examined using a sample deep convective cloud (DCC) observed by the CPR (Fig. 3). The DCC is located between 300 and 350 km along track and is characterized by overshooting cloud tops reaching up to 17 km in altitude. Strong attenuation is observed (Fig. 3a), and the smooth appearance of radar reflectivity echoes extending to and below the surface indicates the presence of moderate multiple scattering effects (Battaglia et al. 2010). Regions contaminated by multiple scattering are currently identified in the C-FMR product (Kollias et al. 2023) using a simple flagging approach based on the methodology proposed by Battaglia et al. 2011a. The blue-filled circles denote the height at which multiple scattering effects on radar reflectivity are expected to become significant. To correctly interpret Doppler velocities in deep convection, it is essential to investigate the influence of multiple scattering on the Doppler signal (Battaglia et al. 2011b). However, since this is not the focus of the current study, our interpretation will be limited to the portion of the V_D profiles above the height where multiple scattering effects begin to become significant.

The DCC V_D measurements are shown in Fig. 3b. The V_D profiles substantial variability, with regions of both positive and negative values. The range of V_D values span the full Nyquist interval [-V_N: +V_N]. The convective V_D profiles are characterized by frequent Doppler velocity aliasing. Fig. 3b presents the 4-km averaged V_D. Velocity aliasing is even more pronounced at the 1-km averaged V_D. The observed V_D variability serves as a strong indicator of the presence of convective updrafts and

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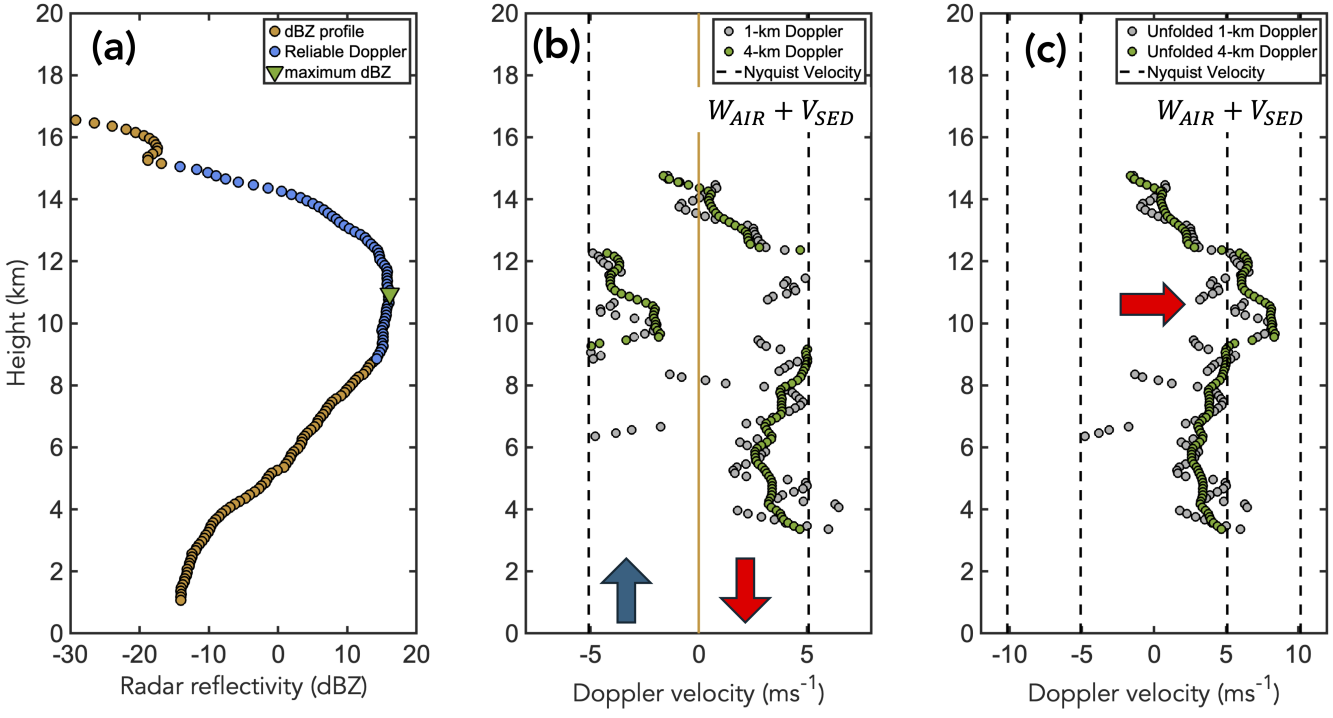

Figure 4. *(a) the CPR reflectivity profile at along-track distance of 309 km. The blue filled circles section of the CPR reflectivity profile indicate the CPR range gates where the Doppler velocity estimates are considered unaffected by multiple scattering. The green triangle indicates the height of the maximum radar reflectivity. (b) The 4-km CPR Doppler velocity profile (green filled circles) and the 1-km CPR Doppler velocity profile (gray filled circles). The black dashed vertical lines indicate the CPR Nyquist Doppler velocity. (c) The unfolded 4-km CPR Doppler velocity profile (green filled circles) and the unfolded 1-km CPR Doppler velocity profile (gray filled circles). The black dashed vertical lines indicate the CPR Nyquist Doppler velocity*

downdrafts. In Figure 3b, the black outline highlights regions where the standard deviation of Doppler velocity exceeds 2 m/s. The standard deviation is calculated within a moving window of 3 km horizontally and 2 km vertically, centered on each pixel, to capture Doppler velocity variations in both the along-track and across-track Doppler velocity directions.

Two example profiles corresponding to the along-track locations indicated by the black dashed lines in Fig. 3b are analyzed here to explore the complexity of the V_D in deep convective cores. The first profile is shown in Fig. 4. The CPR reflectivity profile is presented in Fig. 4a. The blue-filled circles mark the CPR range gates where Doppler velocity estimates are considered unaffected by multiple scattering. Additionally, V_D estimates near the cloud top are excluded if CPR reflectivity falls below –15 dBZ. The maximum reflectivity is observed at an altitude of 11 km, more than 5 km below the cloud top. The corresponding V_D profiles, averaged over 1-km and 4-km along-track intervals, are shown in Fig. 4b. The black dashed lines denote the CPR Nyquist velocity bounds, while the vertical yellow line indicates zero Doppler velocity. As expected, the 4-km-averaged V_D exhibits lower variability with height compared to the 1-km V_D estimates. This vertical correlation is expected, given that the CPR pulse length is 500 m and V_D is estimated every 100 m.

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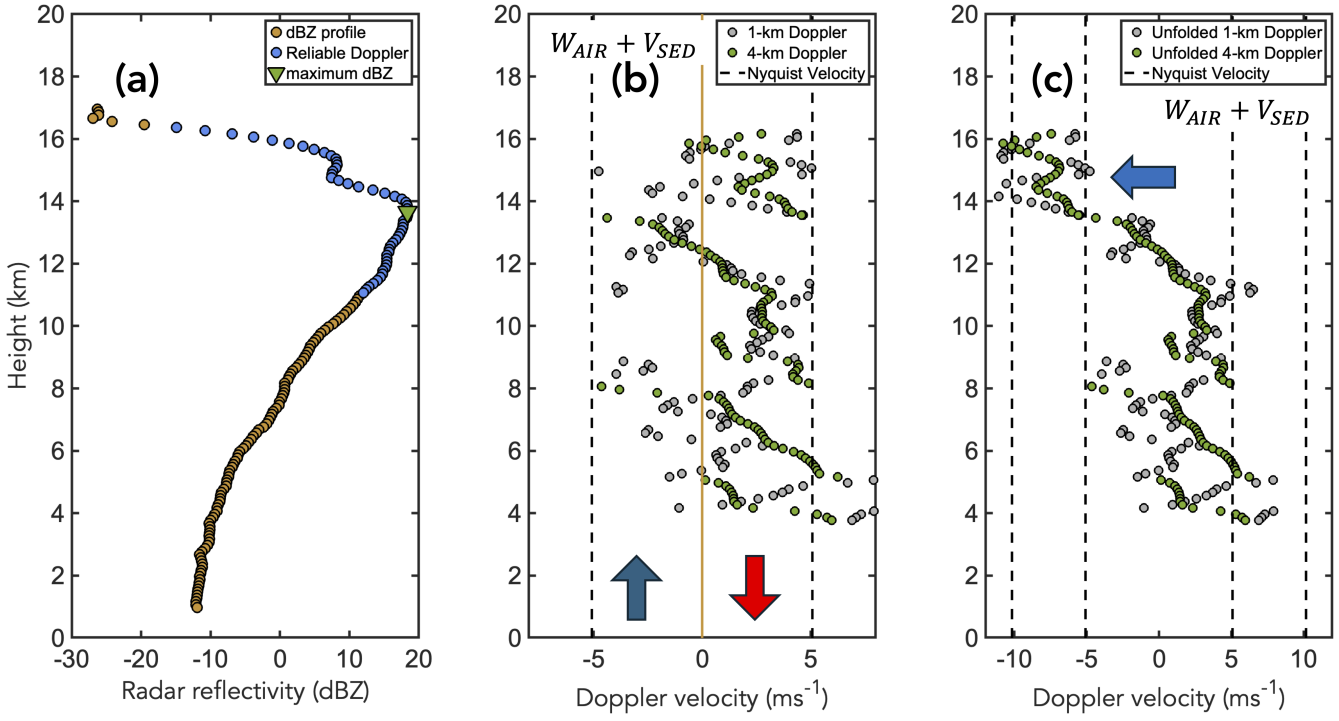

Figure 5. *(a) The CPR reflectivity profile at along track distance of 321 km. The blue filled circles section of the CPR reflectivity profile indicate the CPR range gates where the Doppler velocity estimates are considered unaffected by multiple scattering. The green triangle indicates the height of the maximum radar reflectivity. (b) The 4-km CPR Doppler velocity profile (green filled circles) and the 1-km CPR Doppler velocity profile (gray filled circles). The black dashed vertical lines indicate the CPR Nyquist Doppler velocity. (c) The unfolded 4-km CPR Doppler velocity profile (green filled circles) and the unfolded 1-km CPR Doppler velocity profile (gray filled circles). The black dashed vertical lines indicate the CPR Nyquist Doppler velocity.*

Here, we focus on interpreting the V_D estimates within the section identified as having reliable Doppler velocity retrievals. Beginning with the 4-km profile: near the cloud top, the V_D is positive, indicating the presence of an updraft. Below 14 km, the V_D turns negative, which may indicate the presence of large hydrometeors falling, a downdraft, or a combination of the two, resulting in an apparent downward motion. The abrupt jump of more than 10 m/s in the profile at 12.5 km is attributed to velocity aliasing. In general if the absolute value of the difference between two consecutive Doppler measurements exceed the Nyquist velocity, then adding $\pm 2\, V_N$ to one of the velocity produces a smoother profile. Due to the noisiness of the measurement the identification of a fold is not so straightforward and there will be some ambiguity for points with jumps in v_D close to v_N (e.g. for the 4-km integration values between $V_N - 1$ and $V_N + 1$ m/s). In this example the difference is much larger, so folding is identified unambiguously and unfolding is straightforward. All the segment of the profile between 9 and 12.5 km is therefore aliased; Fig. 4c shows the unfolded 1-km and 4-km V_D profiles. The aliased negative sections of the 4-km profile have been corrected by adding $2\, V_N$. The unfolded 4-km profile displays a smooth vertical structure. Except for a small region

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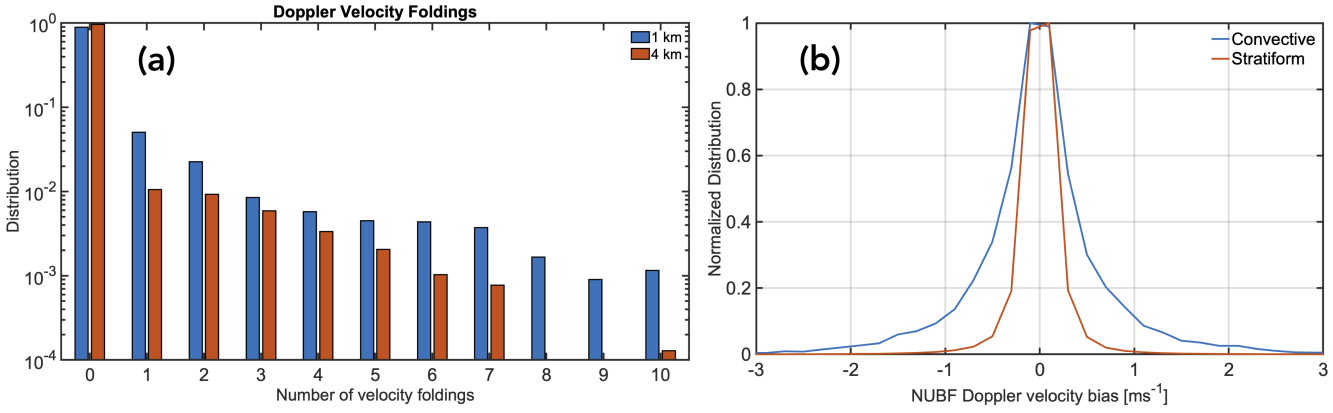

Figure 6. *(a) The normalized frequency of occurrence of CPR Doppler velocity folding, (b) The NUBF induced CPR Doppler velocity bias in* ms^{-1} *in convective and stratiform regions.*

near the cloud top, the V_D values remain negative, suggesting that in this upper part of the convective tower all hydrometeors are falling to the ground.

The second profile is shown in Fig. 5. The CPR reflectivity profile is presented in Fig. 5a. This profile is selected from the elevated cloud top region of the deep convective cloud. In this case, the maximum CPR reflectivity is detected higher in the profile, only 2.5 km below the cloud top. The corresponding V_D profiles, averaged over 1-km and 4-km along-track intervals, are shown in Fig. 5b. These V_D profiles appear more complex due to the presence of significant multiple scattering effects. Fig. 5c shows the unfolded 1-km and 4-km V_D profiles. In the case, the section of the 4-km averaged V_D from the cloud top to the height of 13.8 km is identified as aliased and corrected by subtracting 2 V_N. The unfolded 4-km profile displays a smooth vertical structure. A strong updraft is present above 12 km height, and its magnitude exceeds 10 ms^{-1} near the cloud top.

The analysis of the two convective V_D profiles underscores the challenges associated with unfolding CPR Doppler velocity profiles in deep convection. The low Nyquist velocity of the EarthCARE CPR ($V_N < 6$ ms^{-1} even for the highest PRF) combined with the presence of strong updrafts and downdrafts frequently results in complex V_D profiles. The normalized frequency of occurrence of CPR Doppler velocity folding is shown in Fig. 6a. In the ice portion of the deep precipitating layers, no velocity folding is observed in approximately 90% and 97% of V_D of the 1- and 4-km averaged V_D, respectively. Furthermore, NUBF conditions—expressed as the along-track gradient of CPR radar reflectivity—are more pronounced in deep convective cores. Based on comprehensive statistics from a large dataset of deep precipitating layers, the standard deviation of the along-track gradient of CPR radar reflectivity is 11.2 dB/km and 2.1 dB/km in stratiform conditions. Figure 6b shows the distribution of the corresponding NUBF V_D velocity biases in both convective and stratiform conditions. The standard deviations of the NUBF velocity bias distributions are 1.85 and 0.35 ms^{-1} for convective and stratiform conditions, respectively. A NUBF velocity bias correction algorithm is applied to the V_D data to reduce these biases (Kollias et al. 2023). On average, the residual error is about 20% of the observed NUBF-induced velocity bias (Kollias et al. 2022c). The resulting residual random errors are 0.37 ms^{-1} in convective conditions and 0.07 ms^{-1} in stratiform conditions. The residual random errors in convection are five

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times higher than those in stratiform conditions, yet they are practically negligible considering the expected magnitudes of the convective updrafts and downdrafts.

The analysis suggests that the primary cause of the frequent occurrence of velocity folding in deep convection is the presence of strong convective updrafts and downdrafts. The design of the EarthCARE CPR was finalized more than 20 years ago, at a time when incorporating multiple pulse repetition frequency (PRF) operating modes (Kollias et al. 2007) or polarization diversity techniques (Battaglia et al. 2013) was considered technologically risky. These techniques have since been recommended for future proposed missions, such as WIVERN (Illingworth et al. 2018; Battaglia et al. 2025). While there is considerable experience within the research community in applying radar Doppler velocity dealiasing techniques, most of these methods are designed for scenarios with slowly varying horizontal wind fields and focus on reconstructing spatial and temporal continuity in the velocity field (Eilts et al. 1990; Zhang et al. 2006; Feldmann et al. 2020; Louf et al. 2020). However, there is significantly less experience with Doppler velocity dealiasing in deep convective clouds, particularly under conditions characterized by narrow Nyquist velocity intervals (Maahn et al. 2012). In such environments, vertical continuity cannot be assumed—convective updrafts and downdrafts often have limited vertical extent, and different portions of the hydrometeor column may be in distinct stages of their lifecycle, experiencing markedly different dynamical conditions. For example, the reflectivity profile in Fig. 5a exhibits a double-peak structure above 12 km, indicating a possible vertical discontinuity near the local CPR reflectivity minimum around 15 km. This discontinuity is also reflected in the Doppler velocity profile shown in Fig. 5b. A top-down dealiasing technique—based on the assumption that the upper boundary of the deep convective Doppler velocity profile is free of aliasing—was applied; however, the resulting dealiased Doppler velocity profiles were found to be unphysical. As previously noted, horizontal continuity cannot be assumed in deep convection, as convective updrafts and downdrafts typically have finite horizontal scales. Moreover, the presence of multiple scattering and strong attenuation further limits the applicability of simplifying assumptions such as $W_{AIR} \approx 0$ below the melting layer height.

3 Synergy between the CPR Doppler, MSI and Geostationary Observations

Geostationary satellites today provide a quasi-global coverage in a wide, common set of wavelengths across different agencies, ranging from visible shortwaves to infrared (IR) (Fiolleau et al. 2024). Over the past decade, the capabilities of geostationary satellites have increased significantly in terms of spectral diversity in observations. A synergistic effort is currently underway to merge imagery from the different instruments, recognizing the unprecedented spatial, temporal, and spectral coverage of Earth observation as a crucial contribution for the scientific and meteorological community. EarthCARE plays a primary role in the creation of this homogenized global observation product. These include advanced cloud and aerosol detection and property retrievals, as well as the ability to characterize these properties as a function of time. This temporal dimension enables deeper insights into convective storm lifecycle processes and the corresponding environmental responses.

An example highlighting the possible synergies between the CPR Doppler velocity observations and MSI and Geostationary observations is presented here. The convective event occurred on 7 November 2024 over the southern Mediterranean and the Atlas Mountains in North Africa, during a descending (daytime) orbit (Fig. 7). Figure 7a show the visible band ($0.6\,\mu\mathrm{m}$)

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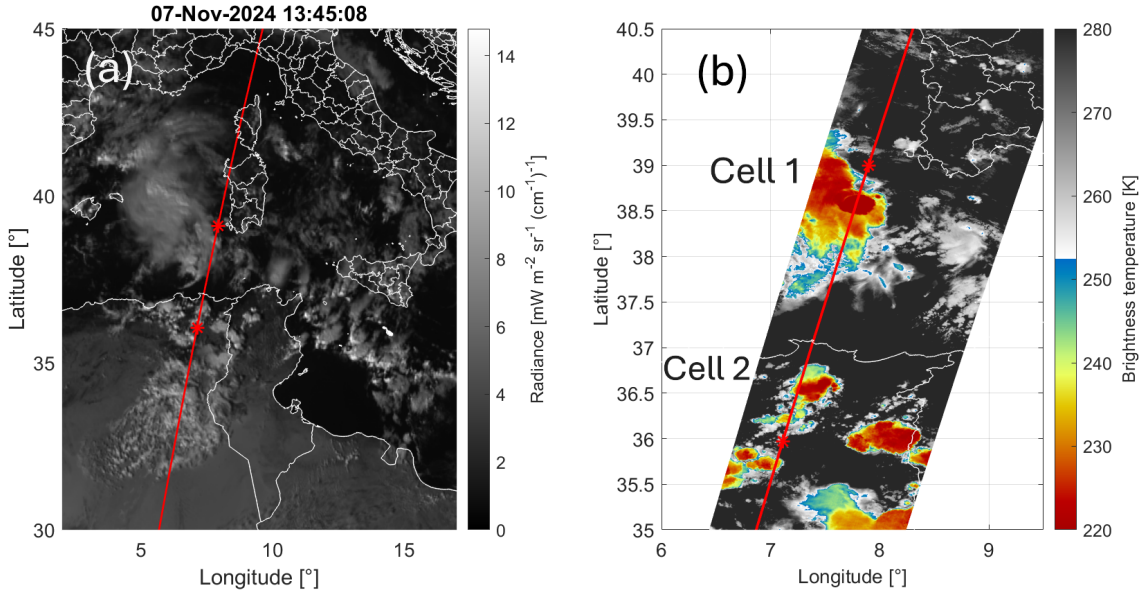

Figure 7. *(a) Radiance from channel 1 (0.6 µm) of MSG, on November 7th, 2024 at 13:45 UTC. The EarthCARE ground track, corrected for the parallax is shown by the red line. (b) the MSI IR channel data from the EarthCARE satellite. The overpass time is 13:43 UTC on November 7th, 2024, frame 2530D. The red line is the satellite ground track. The segment between the two stars is plotted in Fig. 8.*

radiance from the Meteosat Second Generation (MSG) satellite, with the EarthCARE satellite ground track overlaid in red. Figure 7b displays the MSI brightness temperatures in the clean infrared band (10.8 µm). Several convective cloud complexes are detected, and some deep convective clouds were sampled by the CPR. The analysis focuses on two of these cells —cell 1 and cell 2— which are highlighted in Fig. 7.

Figure 8 illustrates the vertical cross-section of the CPR reflectivity and Doppler velocity along the satellite overpass, specifically for the segment between the two red stars in Fig. 7. The along track resolution is 1 km. Two distinct convective complexes are apparent in Fig. 8. Cell 1 spans a broad area, with an anvil extending over 120 km along the EarthCARE orbit. This well-developed anvil reaches an altitude of 10 km and is primarily detrained southward from the main convective core. Between 38.4° and 38.5° latitude, a stratiform region is identified, characterized by a continuous reflectivity echo extending from the surface to cloud-top heights above 10 km. The melting layer is clearly marked by the presence of a bright band. North of this stratiform region, around 38.7° latitude, a stronger convective core is observed, featuring a thick reflectivity column exceeding 15 dBZ. This core reaches nearly 12 km in echo-top height, overshooting the anvil cloud top. Furthermore, the cluster of high-reflectivity profiles below approximately 8 km shows signs of strong attenuation, indicative of significant hydrometeor loading.

Cell 2, located around 36.5° latitude, is part of a cluster of high-reflectivity cells, all characterized by strong attenuation below the freezing level located at approximately 4 km. The CPR captures several distinct convective cores within this region. Between 36.5° and 36.75° latitude, high cloud-top echoes reaching 8–10 km, reflectivities exceeding 10 dBZ, and significant

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attenuation collectively indicate a well-developed convective system. Further south, between 36° and 36.3° latitude, the lower height of the reflectivity echo, combined with the presence of very high reflectivity between 4 km and 6 km, suggests a convective cell still in its developmental phase—likely in an earlier stage of its life cycle. In Fig. 8c, the orange line represents the number of Doppler velocity foldings per profile, while the blue line indicates the number of pixels per profile that exceed the stratiform range threshold.

Using the CloudSat methodology, a DCC would have been identified in Cell 1 between 38.5° and 38.7° latitude (green bar on the right in Fig. 8a) whereas only the central tower in Cell 2, located around 36.5° latitude, would be classified as deep convective (green bar on the left of Fig. 8a). The other convective cores do not meet the required criteria for cloud-top echo height and echo continuity.

To place the CPR convective cloud snapshots within the context of their life cycle, observations from EUMETSAT's Meteosat Second Generation (MSG) satellite are analyzed. Fig. 9 and 10 display the corresponding MSG SEVIRI 1.5 rapid scan frames from channel 9 (10.8 µm) captured at 5-minute intervals before and after the EarthCARE overpass for the two convective cells analyzed in this study. In these figures, the solid black line represents the EarthCARE ground track, corrected for parallax using cloud-top height derived from radar data, while the dashed line shows the original, uncorrected ground track position. The markers correspond to feature locations shown in Fig. 8a, with the black star indicating the position of the minimum brightness temperature tracked within the cell. Strong updrafts, including overshooting tops (Khlopenkov et al. 2021), are expected to be well captured by geostationary sensors. However, this assumption may not hold in cases where convection is embedded within a thick cloud deck or occurs beneath an extensive anvil cloud, where updrafts are obscured and not directly observable by spaceborne infrared and visible passive instruments.

The Tracking and Object-Based Analysis of Clouds *tobac* algorithm is a robust and well-supported algorithm for feature detection and tracking of convective clouds (Heikenfeld et al. 2019; Sokolowsky et al. 2024). In this study, *tobac* is applied to MSG imagery to track the minimum brightness temperature within Cells 1 and 2 (Fig. 11). As shown in Fig. 9, Cell 1 is present well before the EarthCARE overpass and is already in a mature stage of development. Although the cloud top is very cold, no significant cooling is detectable in association with the ongoing embedded convective updraft observed by the CPR. It is likely that, as the cell began detraining mass into the anvil, multiple sparse convective cores developed beneath it. In such cases, radar observations are essential for accurately identifying and characterizing convection. According to the geostationary tracking, at the time of the EarthCARE overpass (13:44 UTC), the minimum brightness temperature in Cell 1 is already below 215 K (Fig. 11a) and fluctuates only slightly—by a few kelvin—during the minutes surrounding the overpass. At this stage of the convective lifecycle, the evolution of the minimum brightness temperature is no longer representative of fine-scale structures such as overshooting tops or highly localized, intense updrafts—features that, in contrast, are well captured by the EC-CPR.

In contrast, Cell 2 is more isolated, which facilitates more effective tracking and allows its evolution to be observed from the early stages. Fig. 10 reveals a secondary cold spot on the southwest side of the cell, which later merges with the main convective core. The *tobac* tracking algorithm identifies a single feature, prioritizing the cooling associated with the main core while disregarding the cloud-top cooling of the smaller, secondary feature.

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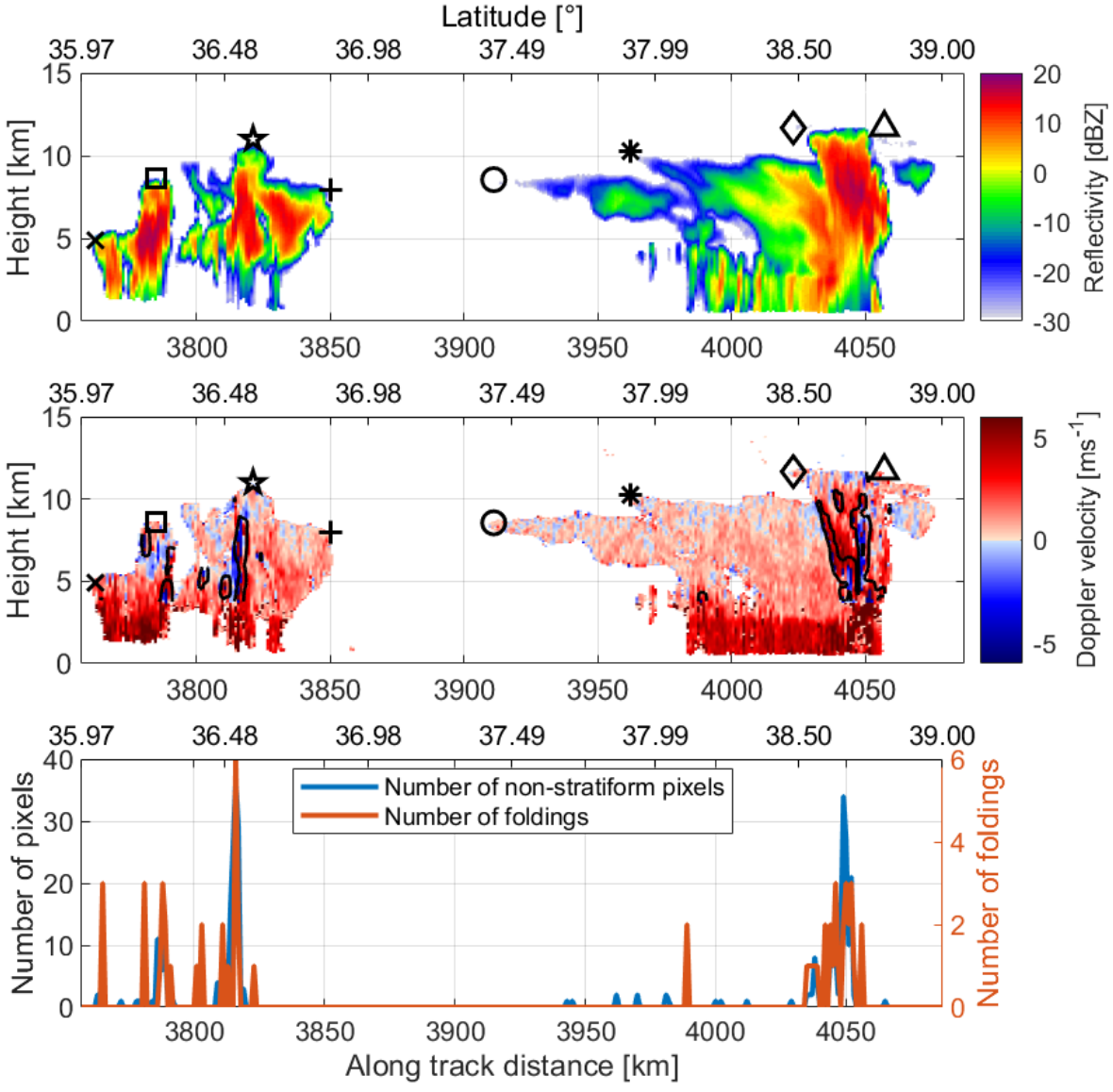

Figure 8. *EarthCARE overpass 2530D over South Mediterranean sea and Atlas mountains on November 7, 2024, at 13:43 UTC. (a) Reflectivity in dBZ, clutter removed. The green bars correspond to profiles labelled as convective from CloudSat methodology described in Takahashi et al. 2014. (b) Doppler velocity, corrected for antenna pointing and NUBF. Black contour is the standard deviation (calculated in a window 3 km horizontally and 1.1 km vertically) that exceeds 2 m/s. (c) Number of foldings per profile and number of pixels per profile that exceed the startiform interval [-2 3] m/s.*

A significant cooling phase is observed during the first 20 minutes of the cell's development. Following this initial phase, the *tobac*-tracked cloud-top temperature remains nearly constant, plateauing at approximately 220 K over the subsequent hour.

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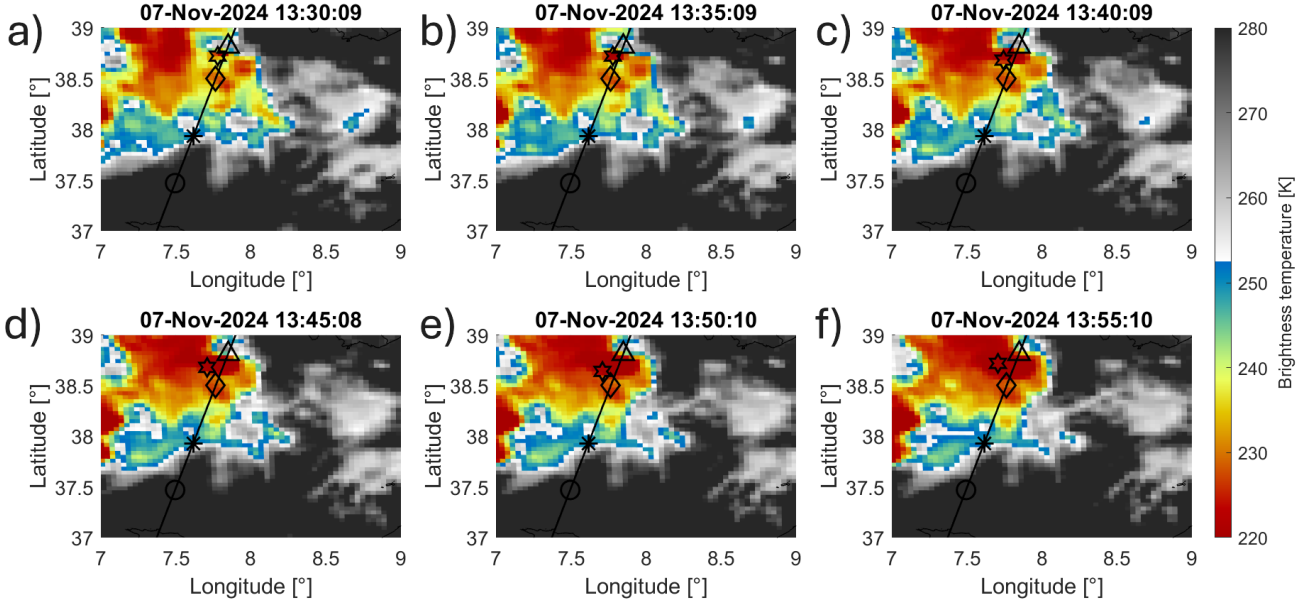

Figure 9. *Successive images depicting the time evolution 10 minutes before (a), 5 minutes before (b), closest in time (c), 5 minutes after (d) and 10 after (e) EarthCARE overpass 2530D on November 7th, 2024, zoom on cell 1. The colors represent the brightness temperature from channel 9 (*$10.8\,\mu m$*) , MSG rapid scans. Black solid line represent the ground track of EarthCARE, corrected for parallax (dashed line is the original ground track). The black markers correspond to Fig. 8a. The black star is the minimum of the brightness temperature that is tracked.*

The EarthCARE overpass (indicated by the red dashed line in the time series in Fig. 11a) occurs when the cell is already in its mature phase. Once again, it is challenging to directly correlate the CPR data—offering detailed vertical cross-sections of internal cloud structure—with the geostationary observations, which characterize the average behavior of the convective system based on cloud-top cooling rates. While *tobac* tracking provides valuable temporal context, it can not capture the fine-scale vertical variability and internal dynamics revealed by the EC-CPR.

This discussion reinforces the limitations of relying solely on geostationary infrared cooling rates for characterizing convection. While IR observations are effective at capturing relatively large and isolated updrafts near cloud tops, embedded convection and sub-kilometer-scale vertical motions largely go undetected. Resolving these features requires spaceborne radar observations, such as those provided by the EarthCARE CPR.

4 Conclusions

Spaceborne radar observations—such as those collected during NASA's CloudSat mission and the NASA-JAXA TRMM and GPM missions—have provided valuable global observations of storm and convective cloud reflectivity structures. However, direct observations of convective dynamics at the global scale have been lacking until now. The recently launched ESA-JAXA

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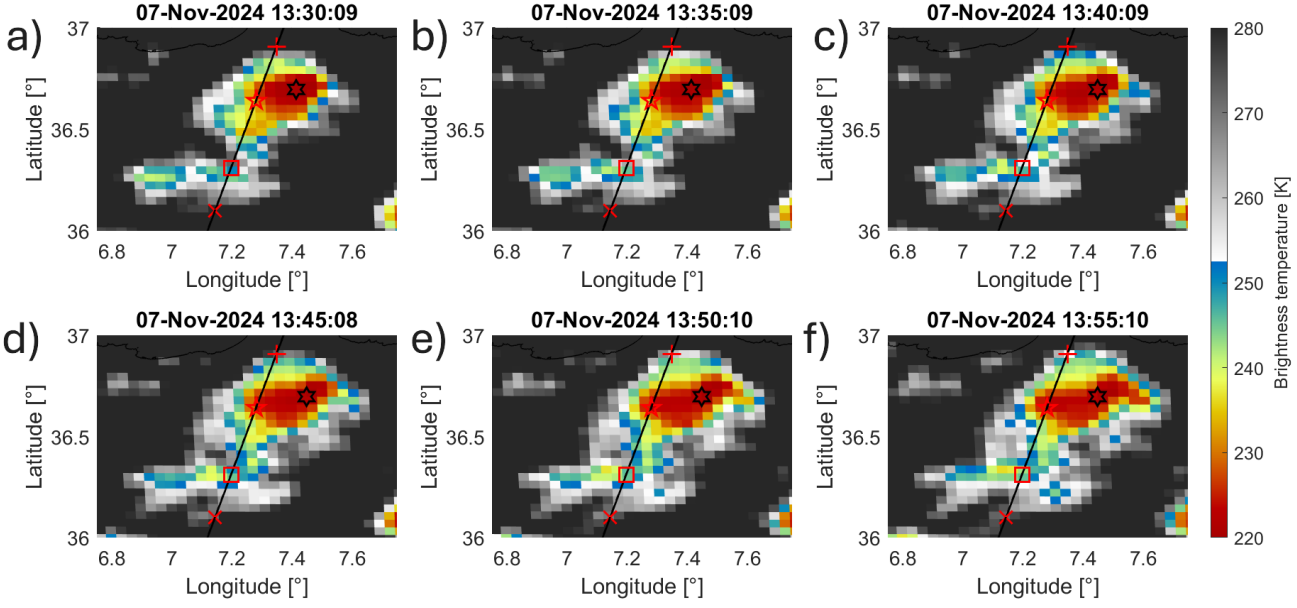

Figure 10. *Successive images depicting the time evolution 10 minutes before (a), 5 minutes before (b), closest in time (c), 5 minutes after (d) and 10 after (e) EarthCARE overpass 2530D on November 7th, 2024, zoom on cell 2. The colors represent the brightness temperature from channel 9 (10.8 µm), measured by MSG rapid scans. Black solid line represent the ground track of EarthCARE, corrected for parallax (dashed line is the original ground track). Red markers correspond in shape to Fig. 8a. The black star is the position of the minimum brightness temperature that is tracked.*

EarthCARE mission, equipped with a Doppler-capable radar, fills this critical observational gap and marks the beginning of a new era of satellite-based radar measurements to improve our understanding of convective dynamics.

Before launch, there were numerous questions regarding the quality of Doppler velocity measurements in deep convection, particularly due to anticipated challenges such as strong attenuation, multiple scattering, non-uniform beam filling (NUBF) effects, the limitations imposed by a narrow Nyquist velocity range, and the complexity introduced by the vertical and horizontal variability of convective cores. In this study, CPR transects across various convective systems are analyzed to assess and illustrate the impact of these challenges on the interpretation of Doppler velocity profiles.

The availability of Doppler velocity measurements from space provides valuable new insights into the presence, as well as the horizontal and vertical extent, of convective updrafts and downdrafts. Doppler velocity-based detection of convective cores is compared with traditional reflectivity-based methods. This comparison is expected to inform a revision of the detection criteria used in previous spaceborne radar studies.

Furthermore, when combined with co-located infrared observations from geostationary satellites, CPR Doppler measurements offer new perspectives on the use of cloud-top cooling rates—computed as time derivatives of brightness temperature—as proxies for convective intensity.

Some preliminary conclusions of this work are summarized in the following.

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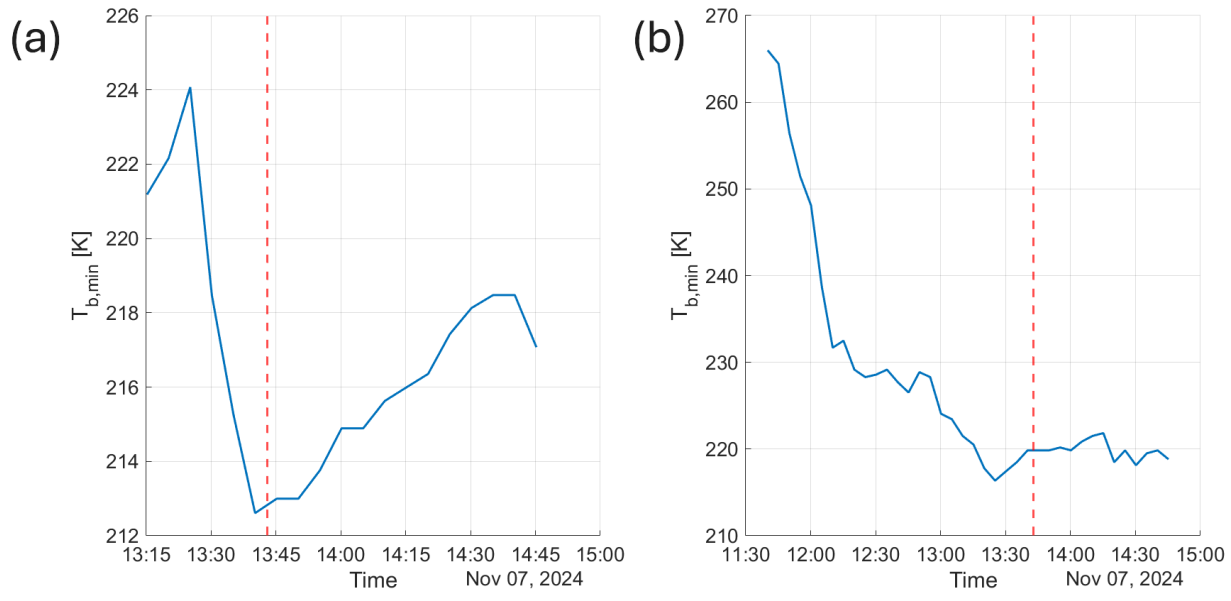

Figure 11. *Minimum brightness temperature (in K) within the cell, as detected and tracked with tobac. Red dashed line corresponds to the EC overpass time. (a) Cell 1. (b) Cell 2.*

1. The first images of Doppler velocities measured by the EarthCARE Cloud Profiling Radar (EC-CPR) offer an unprecedented view of convective motions on a global scale. While these images immediately reveal the presence of convection, the quantitative interpretation of the CPR signal—such as the estimation of updraft and downdraft velocities or convective mass fluxes—will require further analysis. This need arises from the inherent complexity of convective dynamics, compounded by signal noise and the limitations imposed by the narrow Nyquist velocity range.

   The CPR Doppler velocity measurements will serve as the foundation for a dynamics-based convection identification algorithm, designed to augment existing reflectivity-based detection methods. As demonstrated in the case study, parameters such as the standard deviation of Doppler velocity computed within a 3 km horizontal by 2 km vertical window, or the frequency of Nyquist velocity foldings, can serve as reliable indicators of convective activity.

2. The development of a robust algorithm for unfolding CPR Doppler velocity (V_D) measurements in deep convective clouds is currently underway. The first step is to characterize the complexity of the V_D field and to identify the primary sources of velocity discontinuities in deep convection. Initially, the focus will be limited to convective profiles exhibiting fewer than three Doppler velocity foldings at the 4-km along-track resolution—an approach expected to encompass more than 99% of the observed CPR V_D profiles. In cases where velocity aliasing is not observed in the 4-km averaged V_D, but is present in the 1-km averaged profile, the 4-km averaged V_D can be used as a weak constraint to unfold the 1-km averaged V_D by minimizing the difference between the two. In more complex cases, such as those shown in this study, the morphology of the CPR reflectivity profile will be used to determine the vertical continuity of the convective

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column. In addition, V_D estimates at 500 m (native CPR along track resolution), 1-km or 4-km will be combined for the estimation of the unfolded CPR V_D profile.

3. The CPR provides a unique capability for observing embedded convection and sub-kilometer-scale convective cells, thereby overcoming key limitations of convective observations derived from geostationary imagery. In particular, convective motion estimates based on cloud-top cooling rates are effective primarily for updrafts that are both comparable in size to the geostationary sensor's resolution (typically larger than 2 km at mid-latitudes) and located near the cloud top. As such, this method is generally limited to convective cells in the early stages of development or to those exhibiting overshooting tops..

4. Geostationary imagery, on the other hand, offers significant potential for providing the spatio-temporal context of convection—such as whether it is part of a mesoscale system or an isolated cell, and whether it is in the early, mature, or decaying stage of its lifecycle. Additionally, geostationary observations are well-suited for quantifying updraft strength in isolated convective cells, where the time series of minimum cloud-top brightness temperature is expected to be strongly correlated with the intensity of the updraft.

The Doppler capability of EarthCARE's Cloud Profiling Radar (CPR) represents a major innovation, enabling the direct observation of vertical air motions and the terminal fall speeds of hydrometeors. Nonetheless, substantial effort is still required to fully harness this capability and convert these measurements into actionable insights for atmospheric science and modeling.

As a next step, a new convection classification framework will be developed using Doppler velocity and radar-derived features. Once established, this classification—when integrated with synergistic geostationary observations—will support the systematic identification of convective regimes and their associated characteristics. This framework will then be applied to generate global-scale statistics.

These efforts will significantly enhance our understanding of convective dynamics at the global scale and are expected to inform and validate high-resolution weather and climate models.

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Author contributions. AG performed the analysis and provided the draft of the manuscript, AB supervised the study and reviewed the paper, PK provided the description of the Doppler velocity analysis and reviewed the paper, BPT provided the updated dataset and reviewed the paper.

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Competing interests. At least one of the (co-)authors is a member of the editorial board of Atmospheric Measurement Techniques.

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