

Seismo-acoustic and GNSS observations of a record-breaking Black Sea storm: repurposing geophysical sensors for environmental monitoring

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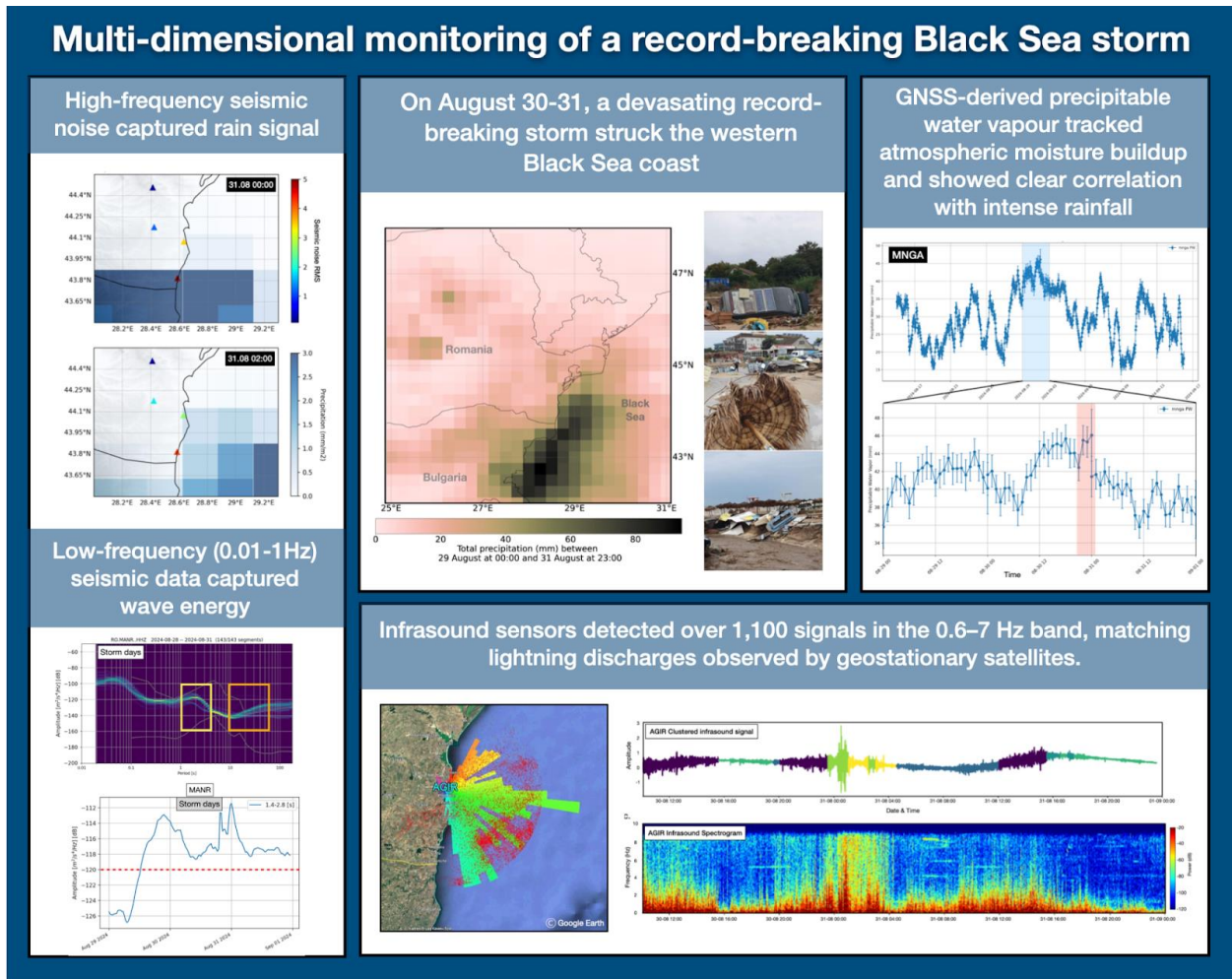
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Abstract

In August 2024, a devastating storm struck Romania's Black Sea coast, setting new precipitation records and marking an unusual change relative to historical climate observations highlighting the increasing frequency of extreme weather events. To investigate this extraordinary event, we integrated This study explores the integration of non-conventional sensors (seismic, GNSS, infrasound, and satellite data) with ERA5 meteorological reanalysis to monitor storm dynamics. High-frequency (>30 Hz) seismic signals captured precipitation, while microseismic bands (0.1-1Hz) reflected wave-induced ground motion. Analysis of infrasound data via unsupervised learning delineated periods of acoustic quiescence from storm-related activity. The temporal evolution of these infrasound states coincided with distinct patterns in seismic ground motion, suggesting a shared origin in the storm's atmospheric dynamics. The infrasound array also detected over 1,100 signals in the 0.6-7 Hz band, matching lightning discharges observed by geostationary satellites. GNSS data recorded a buildup of precipitable water vapor that peaked concurrently with intense rainfall, following a multi-day increase that preceded the main storm phase. This study highlights the value of integrating diverse, non-traditional datasets to enhance the resolution and depth of storm analysis. Their combined use offers a more holistic understanding of storm evolution and supports future research on the potential role of multi-sensor observations in improving early-warning systems in vulnerable coastal regions.

39 **Graphical abstract**

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49 **1. Introduction**

50 Climate change has become a critical global issue, with far-reaching effects on weather patterns
51 and the frequency and intensity of extreme events (Stott, 2016). These changes are not only
52 contributing to more severe weather events but also altering the timing, location, and duration of
53 storms, making them harder to predict and manage (Bengtsson et al., 2006). Understanding how
54 to effectively monitor and predict the behavior of storms, particularly extreme ones, is crucial for
55 improving forecasting models, enhancing early warning systems, and mitigating their impacts on
56 both natural and human systems.

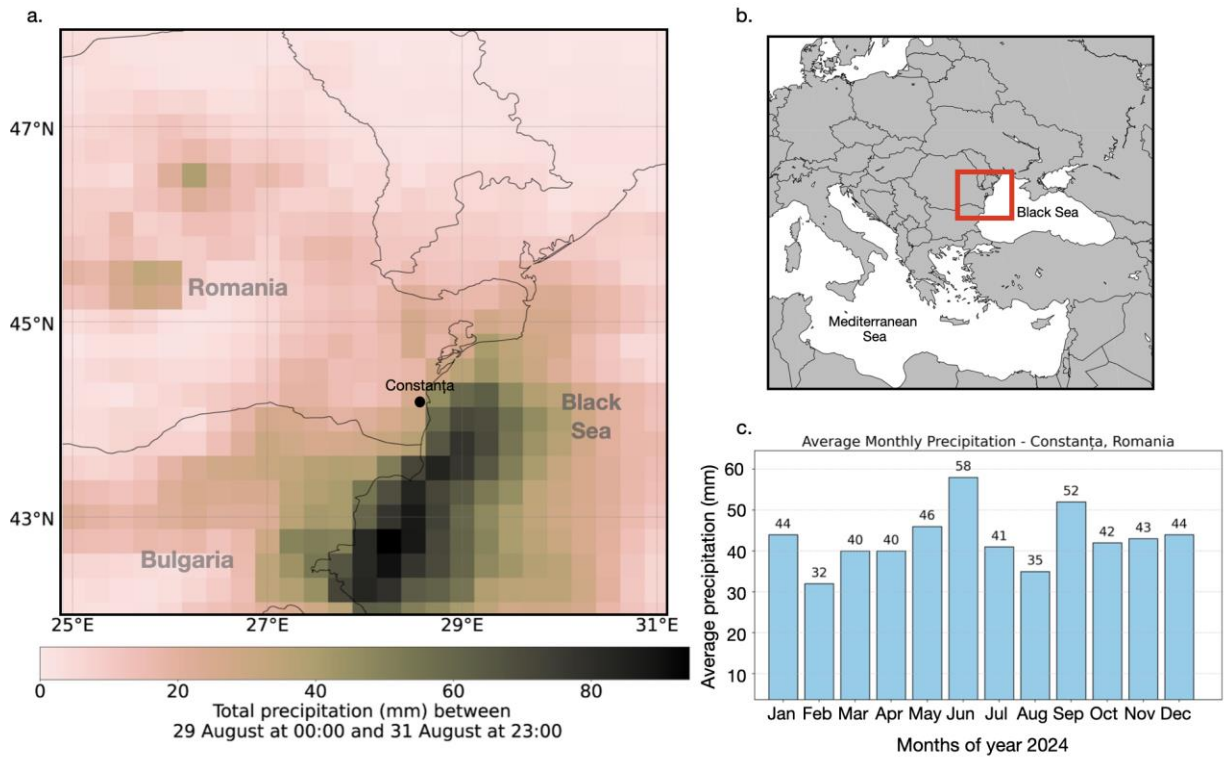
57 Traditional meteorological monitoring relies heavily on ground-based stations, weather radars, and
58 satellite observations to track and predict storm behavior (Kober and Tafferner, 2009). These
59 systems have been the backbone of weather forecasting for decades, providing valuable data on
60 temperature, pressure, wind speed, and precipitation. However, while these methods are effective,
61 they often have limitations in terms of spatial coverage (e.g. Sokol et al., 2021), particularly in
62 remote or hard-to-reach areas. Additionally, they may struggle to capture certain atmospheric
63 phenomena in real-time. As a result, non-conventional monitoring methods are increasingly being
64 integrated into storm tracking efforts to complement existing meteorological approaches (e.g. Bosy
65 et al., 2012; Burtin et al., 2016; Diaz et al., 2023; Coviello et al., 2024).

66 In this context, our study focuses on the integration of alternative environmental datasets, including
67 GNSS stations, infrasound sensors, and seismic data, to track the dynamics of an extreme storm
68 event, as part of a national climate change resilience strategy, implemented through the DTE
69 Climate project (<https://dteclimate.upb.ro/>). GNSS data provide valuable information on
70 atmospheric water vapor, helping to track moisture changes that influence storm formation and
71 intensity (Bosy et al, 2012; Marut et al., 2022). Infrasound sensors detect low-frequency acoustic
72 waves generated by storm activity, such as lightning or large-scale weather system movements
73 like microbaroms (e.g. Landès et al., 2012). Seismic data, though traditionally used for earthquake
74 monitoring, can also record vibrations caused by storm-induced pressure changes, making it useful
75 for detecting rainstorms, floods, or tropical cyclones (e.g. Retailleau and Gualtieri, 2021). Through
76 the integration of these diverse sensor networks, our work highlights their synergy in improving
77 storm detection, monitoring capabilities, and, potentially, early warning systems, contributing to
78 more robust climate resilience strategies.

79 **2. Study area and storm overview**

80
81 The Black Sea region (Figure 1) is characterized by a unique combination of geographic and
82 meteorological features that significantly influence its climate and weather patterns. Nestled
83 between Europe and Asia, the Black Sea is bordered by six countries with diverse landscapes, from
84 mountainous areas to flat plains. This geography, combined with the Black Sea's relatively shallow
85 waters compared to oceanic environments and its connection to the Mediterranean through the
86 Bosphorus Strait, creates an environment where rapid changes in weather are common.
87 Understanding the dynamics of these extreme weather events is crucial, as they can have a
88 profound impact on the environment, economy, and daily life in the region. Monitoring such events
89 is key to improving our ability to predict their occurrence and intensity. By studying the complex
90 atmospheric processes that govern these storms, we can enhance predictive models and refine early

91 warning systems, ultimately helping to mitigate the risks and protect the communities and
92 ecosystems most vulnerable to these extreme weather phenomena.



93
94 *Figure 1. a. Total precipitation accumulated (in mm, shaded according to the scale) between 29*
95 *August 00 UTC and 31 August 23:00 UTC extracted from ERA5 data. b. Map of Europe with red*
96 *square marking the study region; c. Average monthly precipitation rates in Constanța, Romania.*

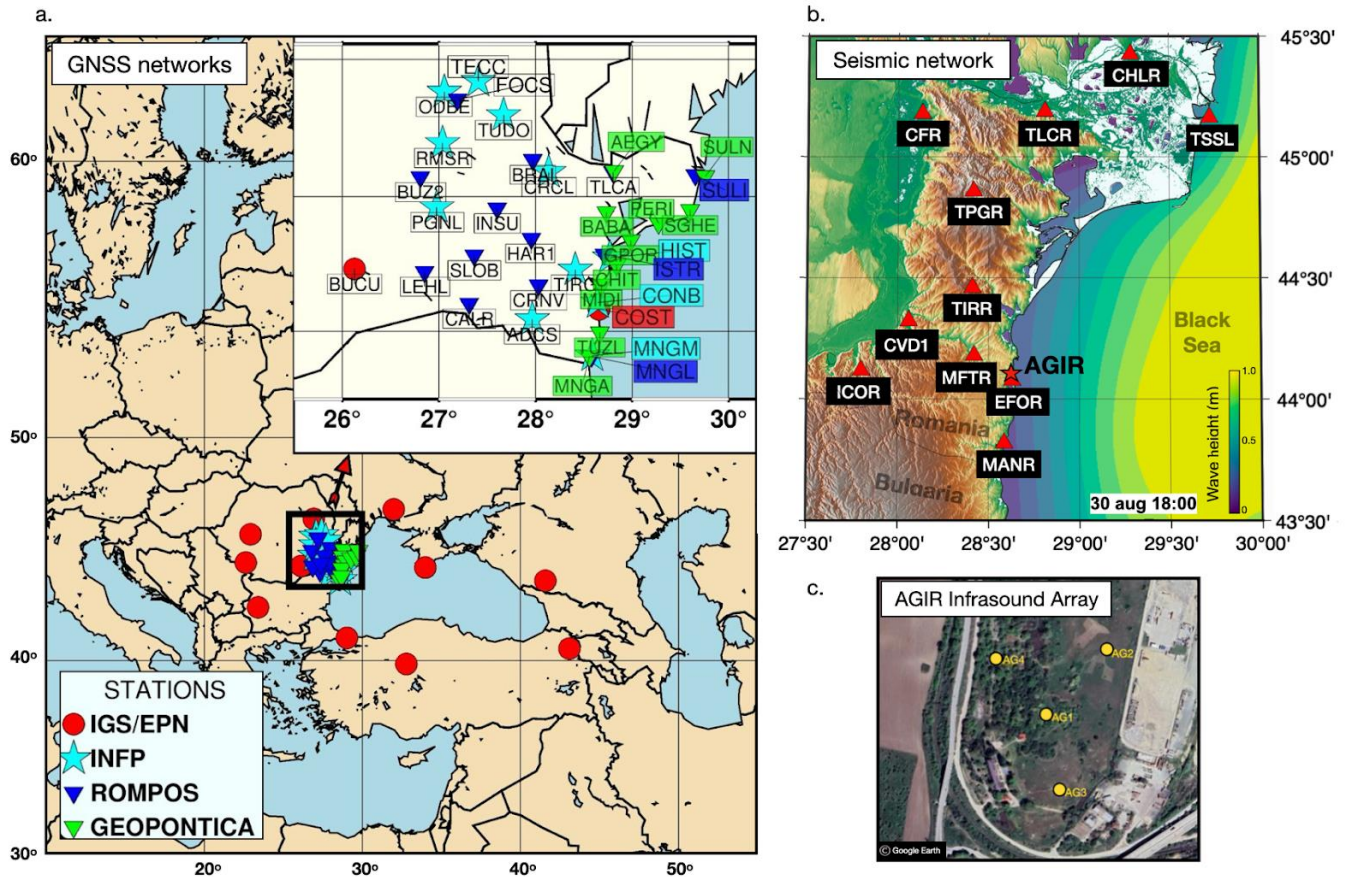
97 In August 2024, Romania experienced severe flooding, largely driven by a storm that brought
98 unusual precipitation patterns to the Black Sea coastal region. Exceptional precipitation totals were
99 recorded over south-eastern Romania in particular in Mangalia (234.7 mm), Agigea (145 mm),
100 and Tuzla (118 mm), leading to significant flooding in coastal towns (Figure 1). Over 800
101 emergency calls prompted large-scale intervention by ISU Dobrogea, focusing on evacuations,
102 debris clearance, and infrastructure restoration (Antonescu et al. 2024). According to the National
103 Meteorological Agency official records (https://www.meteoromania.ro/clim/caracterizare-lunara/cc_2024_08.html), one of the coastal stations at Mangalia, recorded a total of 343.6 mm of
104 precipitation in August 2024, breaking the previous record of 159.1 mm from 1947, and
105 significantly surpassing the average monthly precipitation values for this area (Figure 1c). A
106 remarkable 234.7 mm of this total fell in a single day on August 31, 2024, highlighting the event's
107 exceptional intensity.
108

109 An analysis conducted by ClimaMeter (www.climameter.org, Faranda et al. 2024, Antonescu et
110 al. 2024) immediately after the storm event, showed that low pressure systems similar to the one
111 that caused the floods typically produce 15% less rainfall (7 mm day⁻¹, or up to) in eastern
112 Romania compared to historical trends. However, this particular storm led to a significant local
113 increase in precipitation, particularly in Constanța, one of the coastal cities severely affected by

114 the flooding. In Constanța, where daily rainfall reached up to 5 mm day⁻¹, or up to 10% more than
115 usual. ClimaMeter’s analysis compares events of this type to historical analogues over the past
116 several decades, providing context for how unusual this storm was. marking a notable deviation
117 from the region’s typical weather behavior. The changes in precipitation that contributed to the
118 flooding are largely attributed. According to Antonescu et al. (2024), the local precipitation
119 anomaly in this event is mostly linked to human-induced climate change, with natural climate
120 variability likely playing a modest role. As climate change continues to influence weather patterns,
121 understanding the connection between changing precipitation levels and extreme weather events
122 like flooding is crucial for improving forecasting and resilience in the face of such disasters.
123 Studying this unusual storm and its atmospheric characteristics demonstrates the value of using all
124 available observational infrastructure, including geophysical sensors, to monitor extreme
125 precipitation events, gain insights into atmospheric processes, and support preparedness and
126 resilience in the context of ongoing climate variability.

127 **3. Data and Methods**

128 The analysis of the storm event integrates a variety of data sources and methodologies to provide
129 a comprehensive understanding of its dynamics. Seismic data, infrasound measurements, GNSS
130 water vapor data, and ERA5 reanalysis data are all utilized to capture different aspects of the
131 storm’s behavior (Figure 2). Seismic data offers insights into ground vibrations and atmospheric
132 disturbances, while infrasound monitoring detects low-frequency acoustic signals generated by
133 lightning and other meteorological phenomena. GNSS water vapor data provides valuable
134 information on atmospheric moisture. Additionally, ERA5 reanalysis data (Hersbach et al. 2020),
135 which provides detailed atmospheric and wave dynamics data, helps contextualize the storm’s
136 impact within broader weather patterns. Together, these diverse data sources enable a multifaceted
137 approach to studying the storm and its effects.



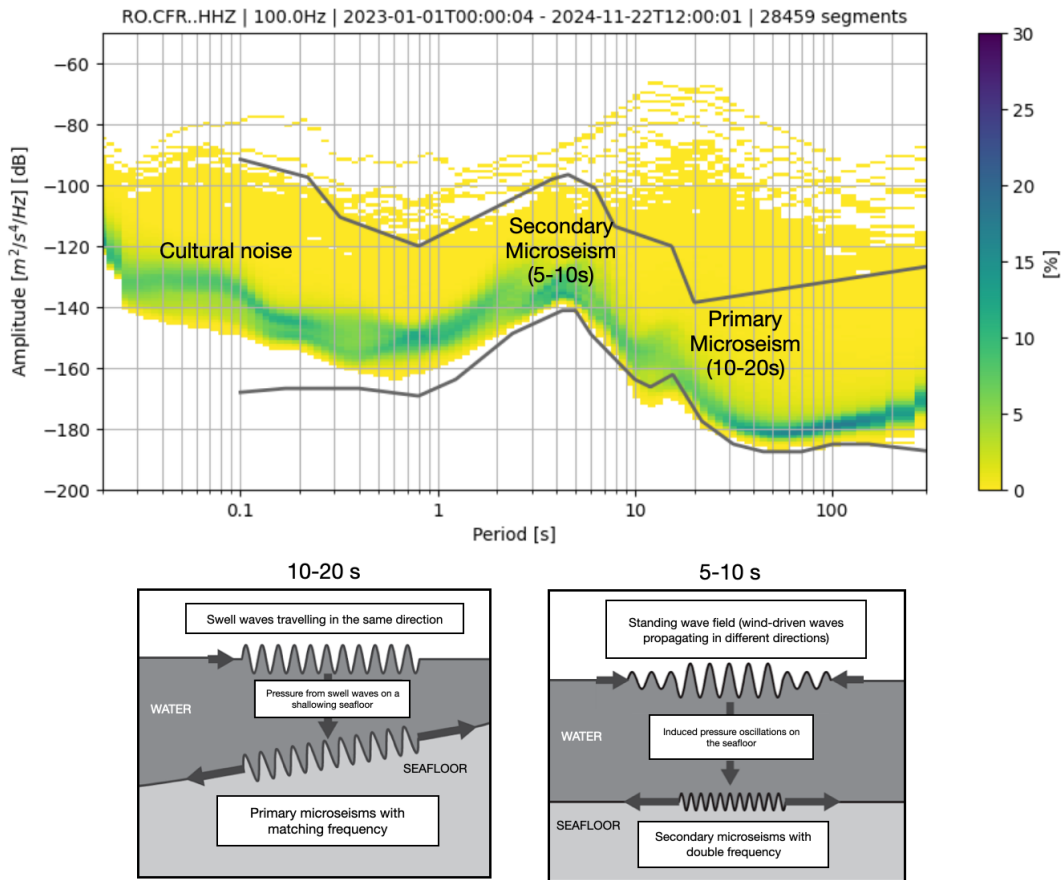
138
 139 *Figure 2. a. Map of Europe and the Black Sea coast showing GNSS stations and their belonging networks*
 140 *used for analysis in this study. b. Map of the Romanian sea coast showing seismic stations (red triangles)*
 141 *and the location of the AGIR infrasound array (star). Coloured contours represent total wave height at*
 142 *18:00 UTC on the 30th of August 2024 from ERA5 reanalysis data. c. The layout of the AGIR infrasound*
 143 *array.*

144
 145 **3.1 Seismic Data**

146 Seismic data represents vibrations of the Earth's surface, commonly referred to as seismic noise.
 147 These low-amplitude movements are recorded across the Earth's surface and are traditionally used
 148 to study the Earth's internal structure and detect earthquakes. Recently, it has increasingly found
 149 applications in meteorology and hydrology, particularly for monitoring weather events (e.g. Dias
 150 et al. 2023; Hua et al., 2023), destructive flood episodes (Burtin et al., 2016), ocean storms, and
 151 tropical cyclones (Gualtieri et al., 2018). Seismic noise can reveal the impact of atmospheric and
 152 oceanic conditions, providing valuable insights into weather events and climate changes (e.g.
 153 Bromirski et al., 2002; Aster et al, 2008; 2023). In particular, seismic data helps track variations
 154 in the Earth's surface caused by factors such as ocean waves, wind, and precipitation, offering a
 155 unique perspective on these phenomena (e.g. Grevemeyer et al., 2000; Borzi et al., 2022).

156 When the seismic noise is analyzed in the frequency domain, two clear peaks emerge in the
 157 spectrum (Figure 3), reflecting distinct types of ocean wave interactions (Koper et al., 2015;

158 Arduin et al., 2019; Tanimoto et al., 2023). The primary peak, observed in the range of 10-20
 159 seconds (0.05-1 Hz), is generated by the impact of "swell" waves traveling in the same direction,
 160 inducing pressure variations in the Earth's crust that match the period of the waves. The secondary
 161 peak, in the range of 5-10 seconds (0.1-0.5 Hz), is produced by wind-driven waves, which
 162 propagate in different directions and generate pressure oscillations on the ocean floor (Ebeling et
 163 al., 2012). These seismic signals directly link ocean conditions with seismic activity (Li et al.,
 164 2020), providing insights into large-scale weather phenomena like ocean storms.



165
 166 *Figure 3. Probabilistic Power Spectral Density (PPSD) of seismic noise for station CFR, over two years,*
 167 *showing key sources of primary and secondary microseisms. Below, sketches illustrate the generation*
 168 *mechanisms: primary microseisms are caused by unidirectional swell waves inducing pressure fluctuations*
 169 *on a shoaling seafloor, while secondary microseisms result from nonlinear interactions of wind-driven*
 170 *waves over deeper water (modified after Ebeling, 2012).*

171 Higher frequencies above 30 Hz are associated with the effects of precipitation and wind, as seen
 172 in studies like Rindraharisaona et al. (2022) or Diaz et al. (2023). These higher-frequency seismic
 173 signals help track more localized weather events, such as storms and heavy rainfall. Seismic data,
 174 when integrated with other meteorological tools, enhances the ability to monitor and predict
 175 weather events.

176 To analyse seismic data, the raw traces are first corrected for instrument response and converted
177 to units of velocity. These are then filtered with bandpass butterworth filters adapted to capture the
178 target signal: low pass filtering (<1 Hz) for wave-seafloor coupled interactions and high pass
179 filtering (>30 Hz) to identify possible signatures of precipitation, essentially induced pressure
180 fluctuations in the ground converted to weak seismic vibrations due to rain drops. Spectrograms
181 of these filtered seismic traces were computed using short-time Fourier transforms implemented
182 in the `scipy.signal` package, with the default 256-sample window length used for each segment,-to
183 visualise signatures of the hydro-meteorological phenomena in the frequency content of ground
184 vibrations.

185 Potential environmental signals in the seismic data were also investigated using power spectral
186 density (PSD) analysis. To account for variations over time, a Probabilistic Power Spectral Density
187 (PPSD) method was applied. The continuous waveform was divided into 1-hour time windows
188 with 50% overlap, and a PSD was computed for each window after instrument-response correction
189 and basic preprocessing. These estimates were combined into a probability distribution, providing
190 a statistical overview of typical and transient noise levels across frequencies. The PPSD was
191 produced using ObsPy (Beyreuther et al., 2010), which handles data gaps and ensures reliable
192 normalization.

193 Temporal variations in PSD amplitudes are also analyzed to track changes in seismic noise at
194 specific frequencies. By extracting PSD values at selected frequencies that are expected to capture
195 primary and secondary microseisms, time series of noise levels are generated. These temporal
196 PSDs allow for the identification of trends and correlations with environmental factors, such as
197 ocean wave activity or weather conditions.

198 **3.2 Acoustic Data**

199 Infrasound waves are low-frequency acoustic waves that are inaudible to the human ear, typically
200 below 20 Hz. These waves are generated by a variety of natural and anthropogenic sources,
201 including meteorological events, volcanic eruptions, earthquakes, and human activities such as
202 explosions and industrial processes (Campus and Christie, 2009; Bondár et al., 2022). In particular,
203 infrasound is often associated with phenomena like thunderstorms, ocean waves, and large-scale
204 atmospheric events, which generate pressure fluctuations that propagate through the atmosphere
205 (e.g. Stopa et al., 2012; Landès et al., 2012; Listowski et al., 2022). These waves provide valuable
206 information about the dynamics of weather systems (e.g. Hupe et al., 2019), making them an
207 essential tool for monitoring and understanding environmental processes (e.g. Brachet et al., 2009;
208 Hupe et al., 2022). Infrasound associated with thunderstorms, primarily generated by acoustic
209 waves from thunder, has been studied previously and shown to be detectable at distances ranging
210 from tens to hundreds of kilometers (e.g., Assink et al., 2008; Sindelarova et al., 2015; Šindelářová
211 et al., 2021). Nevertheless, infrasound arrays detect signals from multiple storm-related sources,
212 not just thunder (e.g., Waxler et al., 2024). In the present study, we build on this understanding by
213 integrating these signals with seismic, satellite, meteorological, and water vapor observations to
214 investigate what these complementary datasets reveal about storm evolution in a coastal
215 environment.

216 For the monitoring of infrasound signals, we use data from an infrasound array system located at
217 Eforie Nord-Agigea, Romania (AGIR, Figure 2). This array consists of multiple sensors, including

218 SIS-1 infrasonic sensors (Seismowave), equipped with global positioning systems (GPS) and noise
219 reduction technology.

220 To analyze the seismo-acoustic characteristics of the August 30-31 Black Sea storm, we used a
221 two-pronged approach: (1) single sensor signal analysis based on feature extraction and
222 unsupervised machine learning, and (2) array-based analysis using all the sensors of AGIR and
223 classic multi-channel correlation algorithms. Together, these methods provide complementary
224 insights into the acoustic behavior of the storm, capturing both local signal characteristics and
225 spatial coherence across sensors.

226 For the single-station analysis, infrasound data recorded at the AGIR sensor (Figure 2) was
227 segmented into 30-minute windows, and a set of time-frequency features was extracted to
228 characterize the signal dynamics (Supplementary Material). These features describe how energy
229 and frequency content evolve over time, providing insights into the structure of the infrasound
230 signal. Parameters such as spectral centroid and spectral rolloff are standard descriptors in acoustic
231 signal analysis and are suitable here because they effectively capture shifts in dominant frequency
232 produced by lightning-generated acoustic waves or the passage of pressure disturbances, while
233 spectral flux highlights changes in broadband acoustic energy (Pásztor et al., 2023). Spectral
234 entropy reflects the complexity of the frequency distribution, which increases during turbulent
235 atmospheric conditions, and the zero-crossing rate, mean, and variance of the power spectrum
236 summarize overall activity and variability. This feature set provides a compact representation of
237 the signal suitable for unsupervised machine-learning approaches such as clustering, techniques
238 widely used in data mining to identify patterns in multidimensional time-frequency data (e.g.,
239 Coates and Ng, 2012), and allows us to distinguish physically interpretable stages of storm-induced
240 changes in the infrasound wavefield.

241 The extracted features were used as input for K-Means clustering (MacQueen, 1967), an
242 unsupervised machine learning algorithm that partitions data into a predefined number of groups.
243 K-Means minimizes within-cluster variance by iteratively assigning feature vectors to the nearest
244 cluster centroid and updating the centroids based on the grouped data. This clustering method
245 enables the identification of distinct acoustic patterns in the signal (e.g. Pásztor et al., 2023),
246 offering a data-driven way to segment the storm’s infrasound profile without requiring prior labels
247 or assumptions. Prior to clustering, the features were standardized using z-score normalization, to
248 ensure comparable scaling across variables. The optimal number of clusters was determined using
249 the elbow method, which evaluates within-cluster variance as a function of cluster number
250 (Supplementary Material). To select the most informative features, we applied covariance pruning,
251 and the temporal evolution of the features was visualized to ensure meaningful representation. This
252 procedure resulted in six clusters, providing a balanced representation of the infrasound dynamics
253 while avoiding over-segmentation or overfitting. By combining multiple features in the clustering,
254 this method captures the evolving acoustic states of the storm in a compact, interpretable form.

255 In parallel with the single-station analysis, we also applied the Progressive Multi-Channel
256 Correlation (PMCC) method, as implemented in the DTK-PMCC software (Cansi and Le Pichon,
257 2008; Le Pichon et al., 2010) to detect and analyze coherent acoustic signals across an infrasound
258 array. The PMCC method targets signals generated by atmospheric sources such as lightning (i.e.,
259 associated thunders) or other pressure disturbances, operating in the low-frequency range of 0.7 to

260 7 Hz. It is specifically suited for mini-array configurations, where signal coherence between
261 closely spaced sensors can be exploited for precise signal detection and characterization.

262 The PMCC algorithm was implemented using a multi-resolution configuration following the
263 standardization proposed by Garcés (2013), with window lengths and frequency bands arranged
264 in third-octave bands. A total of 19 frequency bands were used, covering 0.1-7 Hz. Window
265 lengths decrease logarithmically with frequency, ranging from 258 s in the lowest band to 4 s in
266 the highest band. A 10% time step was applied (corresponding to 90% overlap between
267 consecutive windows), and this scheme repeats every decade. Within each time-frequency
268 segment, cross-correlations are computed between all sensor pairs to identify coherent wavefronts,
269 signals that exhibit consistent arrival times across the array. From these detections, PMCC
270 estimates several key propagation parameters, including backazimuth (the direction of arrival),
271 horizontal trace velocity, amplitude, duration, and dominant frequency. This approach is
272 particularly effective in noisy environments and enables the discrimination of storm-generated
273 infrasound from background signals or unrelated acoustic sources. The algorithm's output consists
274 of a time-frequency map of signal detections enriched with physical metadata, allowing for
275 detailed interpretation of the storm's acoustic footprint and its temporal evolution.

276 3.3 Satellite Observations

277 We also incorporated data from the Meteosat Third Generation (MTG) satellite system (Holmlund
278 et al., 2021), specifically from its Lightning Imager (LI) sensor (Viticchie et al., 2020). The MTG
279 satellites operate in geostationary orbit at approximately 36,000 km altitude, providing continuous
280 observations over Europe, Africa, and surrounding waters. The Lightning Imager detects cloud-
281 to-cloud, cloud-to-ground, and intra-cloud lightning flashes using four cameras that collectively
282 cover 86% of the Earth's visible disc from the satellite's perspective.

283 For this study, we used Level 2 group data, which includes the geographical coordinates and timing
284 of each detected flash. The MTG Lightning Imager detects total lightning (cloud-to-cloud and
285 cloud-to-ground) optically at 777 nm, with 4.5 km pixel resolution at the sub-satellite point and
286 1 ms frame rate (Holmlund et al., 2021; Kokou, 2023). Level-2 achieves detection efficiencies of
287 ~80-90%, capturing even weak flashes reliably, with false alarm rates <0.3 (Enno et al., 2025).
288 Flash geolocation uncertainty reaches 5-10 km near the edge of the instrument's field of view,
289 where off-nadir viewing geometry amplifies parallax effects (Bližňák & Sokol, 2026). By mapping
290 these detections, we were able to analyze the spatial distribution and temporal evolution of the
291 storm's lightning activity. The dataset also offered insights into the storm's intensity and structure,
292 complementing other meteorological observations.

293 Associations between infrasound detections and lightning flashes detected by MTG within 50 km
294 of the AGIR infrasound station were investigated by assuming direct-path acoustic propagation
295 and a correspondence between infrasound time-of-arrival and the MTG lightning discharge time
296 (after Assink et al., 2008):

$$297 t = t_{\text{MTG}} + d/c + \Delta t,$$

298 where d is the distance between the lightning discharge and the infrasound station, $c = 340\text{m/s}$,
299 and $\Delta t = \pm 10\text{s}$ accounts for timing uncertainty associated with the simplified propagation

300 assumption. In particular, infrasound travel time from thunder sources can vary due to atmospheric
301 temperature and wind variations along the propagation path, which affect the effective sound speed
302 and may introduce deviations from the assumed constant-velocity, straight-path propagation.
303 Additionally, a maximum angular deviation of 10° between the observed infrasound backazimuth
304 and the MTG-derived backazimuth is permitted for an association to be accepted.

305 **3.4 GNSS Data**

306 The use of GNSS technology for atmospheric monitoring provides a powerful tool for analyzing
307 extreme weather events. Beyond its well-known applications in navigation, timing, [positioning](#)
308 [and crustal dynamics](#) (Nistor et al., 2021a; Nistor et al., 2021b), GNSS has become a reliable
309 method for sensing tropospheric water vapour, an essential driver of weather systems and a key
310 variable in forecasting models (Guerova et al., 2016; Vaquero-Martínez and Antón, 2021). Over
311 the past two decades, ground-based GNSS networks in Europe have contributed significantly to
312 operational meteorology by providing near real-time estimates of atmospheric water vapour,
313 aiding in the detection and tracking of severe weather, including heavy rainfall and storms
314 (Karabatić et al., 2011; Priego et al., 2017; Jones et al., 2020). These high-resolution observations
315 have proven valuable for both nowcasting and validating numerical weather prediction models
316 (Wilgan et al., 2015; Bosy et al., 2012; Awange, 2012).

317 In this study, GNSS data were collected from several networks (Figure 2), including the
318 International GNSS Service (IGS, Johnston et al., 2017), the EUREF Permanent Network (EPN,
319 Bruyninx et al., 2012), the Romanian Position Determination System (ROMPOS, Iliescu et al.,
320 2019), and GEOPONTICA (Dimitriu et al., 2017). A total of 37 permanent GNSS stations were
321 analyzed over a 30-day period, with the rainiest interval selected at the midpoint of the study
322 period. These stations provide high-quality, continuous observations critical for atmospheric
323 monitoring.

324 The data were processed using a double-differenced, ionosphere-free combination of L1 and L2
325 carrier phases. This approach helps minimize errors such as ionospheric delays, satellite clock
326 biases, and other common atmospheric effects. The resulting Zenith Tropospheric Delay (ZTD)
327 values were then corrected using the Vienna Mapping Functions 3 (VMF3, Landskron et al., 2018),
328 which improves the accuracy of ZTD by accounting for variations in the troposphere's atmospheric
329 conditions. Once the ZTD was refined, it was converted into integrated precipitable water vapor
330 (PWV) using surface meteorological data (temperature and pressure) from co-located weather
331 stations, following the method outlined by Bosy et al. (2012). This process allowed for the
332 derivation of high-resolution atmospheric water vapor content, critical for analyzing the dynamics
333 of the extreme storm event over the Black Sea. By combining GNSS-derived PWV with data from
334 other observational sources, the study captured the temporal and spatial variations in atmospheric
335 moisture, offering valuable insights into the storm's development and intensity.

336 **3.5 Meteorological Data**

337 To compare the infrasound signals captured during the Black Sea extreme storm event, we
338 extracted meteorological data from the open-access ERA5 reanalysis dataset, produced by the
339 European Centre for Medium-Range Weather Forecasts (ECMWF). This dataset provides a
340 comprehensive record of global weather conditions from 1950 to the present (Hersbach et al.,

341 2023). ERA5 combines observational data and advanced numerical models to generate relatively
342 high-resolution atmospheric parameters compared with earlier global reanalyses, including
343 precipitation (Figure 1), wind speed, and wave height. ERA5 has been extensively validated (Jiao
344 et al., 2021; Wu et al., 2022; Soci et al., 2024) and is widely used in studies of storm evolution and
345 precipitation dynamics (e.g. Dullart et al., 2020; Tiberia et al., 2021; Price et al., 2025), making it
346 a suitable choice for the mesoscale processes examined here.

347 For our study, the ERA5 data was used to track the meteorological context of the storm, offering
348 insights into the intensity of precipitation, the evolution of wind patterns, and the development of
349 oceanic wave heights. With a temporal resolution of 1 hour and spatial resolution of $0.25^\circ \times 0.25^\circ$,
350 ERA5 allows for a mesoscale comparison of the storm's meteorological characteristics over time.
351 While its spatial averaging cannot resolve localized convective-scale precipitation, it provides a
352 vital benchmark for qualitative comparisons and for testing multi-sensor monitoring potential.
353 These comparisons help us understand the storm's dynamics and assess its impact, further
354 enhancing the interpretation of infrasound signals and aiding in future storm prediction and
355 monitoring efforts. The open-access nature of ERA5 ensures broad accessibility, contributing to
356 the transparency and reproducibility of our storm analysis (Copernicus Climate Change Service,
357 Climate Data Store, 2023).

358 **4. Results**

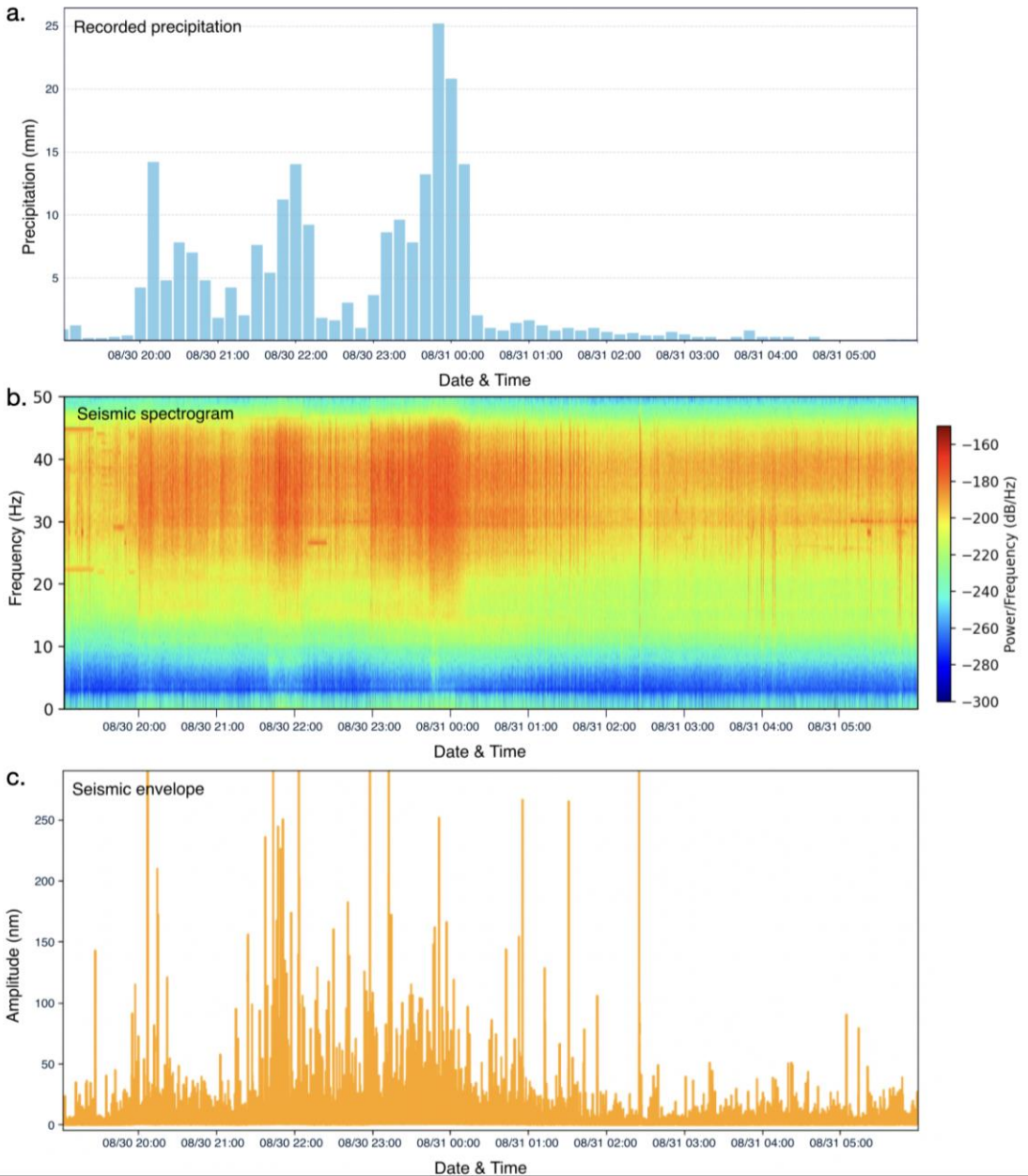
359 **4.1 Seismic signatures of storm evolution**

360 High frequency (>30 Hz) analysis of seismic noise reveals strong signals during periods of intense
361 rainfall (Figure 4). Specifically, the displacement envelope at station MANR and its spectrogram
362 for 30 August, 12:00 UTC to 31 August, 06:00 UTC (Figures 4b, c) reveal strong signal around
363 midnight, when recorded precipitation exceeded 20 mm per 10 minutes. Similar temporal patterns
364 in the seismic spectrogram were also visible when compared with hourly precipitation levels from
365 ERA5, indicating that the high amplitude of energy observed above 30 Hz is most plausibly
366 generated by raindrop impacts.

367 However, this correspondence is not uniform across all rainfall episodes. While the main
368 precipitation maximum on 30-31 August produces a clear and sustained seismic response, several
369 lower-intensity precipitation pulses show a much weaker or no recognizable signature in either the
370 seismic envelope or spectrogram. This behaviour is consistent with previous work (e.g.,
371 Rindraharisaona et al., 2022), which demonstrates that only rainfall above a certain intensity, or
372 involving sufficiently large drops, generates impact forces strong enough to be detected by
373 broadband seismometers. Our observations therefore reflect both strong positive correlations
374 during intense rainfall and the lack of seismic expression for weaker precipitation. This selective
375 sensitivity supports the interpretation that high-frequency seismic noise can reflect strong rainfall
376 peaks but is less responsive to light or moderate precipitation, an important nuance when
377 interpreting multi-sensor relationships in this study.

378 Anthropogenic seismic noise is typically strongest at low to mid frequencies (<25 Hz), where day-
379 night variations reflect traffic, human activity, and transient signals from machinery, while higher-
380 frequency bands (25-45 Hz) may include periodic contributions from rotating equipment (e.g.,
381 Gross & Ritter, 2008; Diaz et al., 2017). The bandwidth targeting rainfall in this case is between

382 30-50 Hz, which is above the dominant frequency content of most anthropogenic sources and
383 overlaps with raindrop-impact energy documented in recent rainfall-seismic studies.

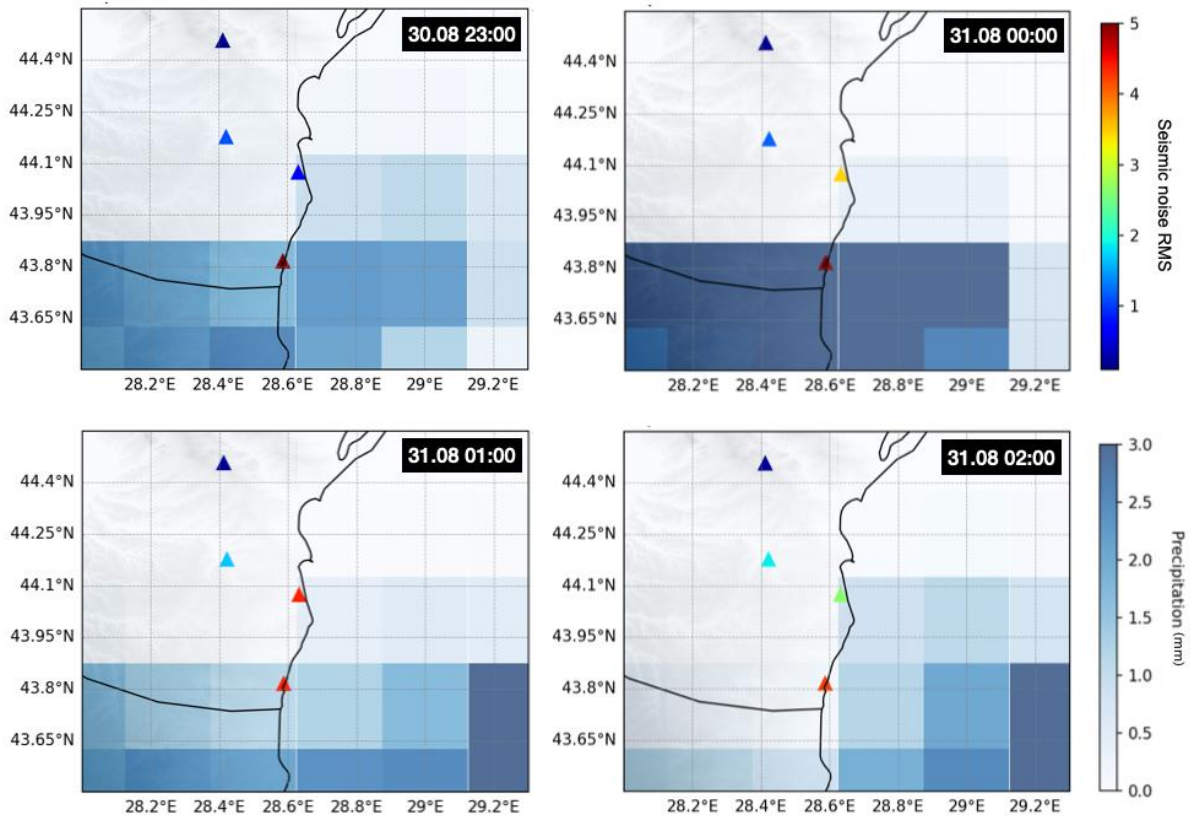


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385 *Figure 4. High frequency (30-50Hz) observations of the storm at station MANR. a. Time series of total*
386 *precipitation every 10 minutes from the ANM station at Mangalia b. Spectrogram of the seismic time series*
387 *for station MANR. c. Envelope of the displacement seismogram at station MANR.*

388 To visualise the signature of the storm passing over the network of broadband seismic stations in
389 the coastal area, we also plotted the hourly precipitation values with the hourly root-mean-square
390 amplitudes of the high-frequency (>30 Hz) seismic velocity envelopes recorded at seismic stations.
391 Figure 5 shows four snapshots of hourly plots of gridded precipitation data from ERA5, which
392 have a lower amplitude than point measurements at the Mangalia station, due to the averaging over

393 the grid block. This figure presents a temporal coincidence between changing precipitation patterns
 394 from ERA5 data and the amplitudes of high-frequency seismic noise. This observation further
 395 supports the likelihood of a causal relationship. These high-frequency seismic signals could
 396 potentially be explored as a near real time indicator of intense rainfall events, providing a
 397 conceptual basis for a simple streaming detection approach.



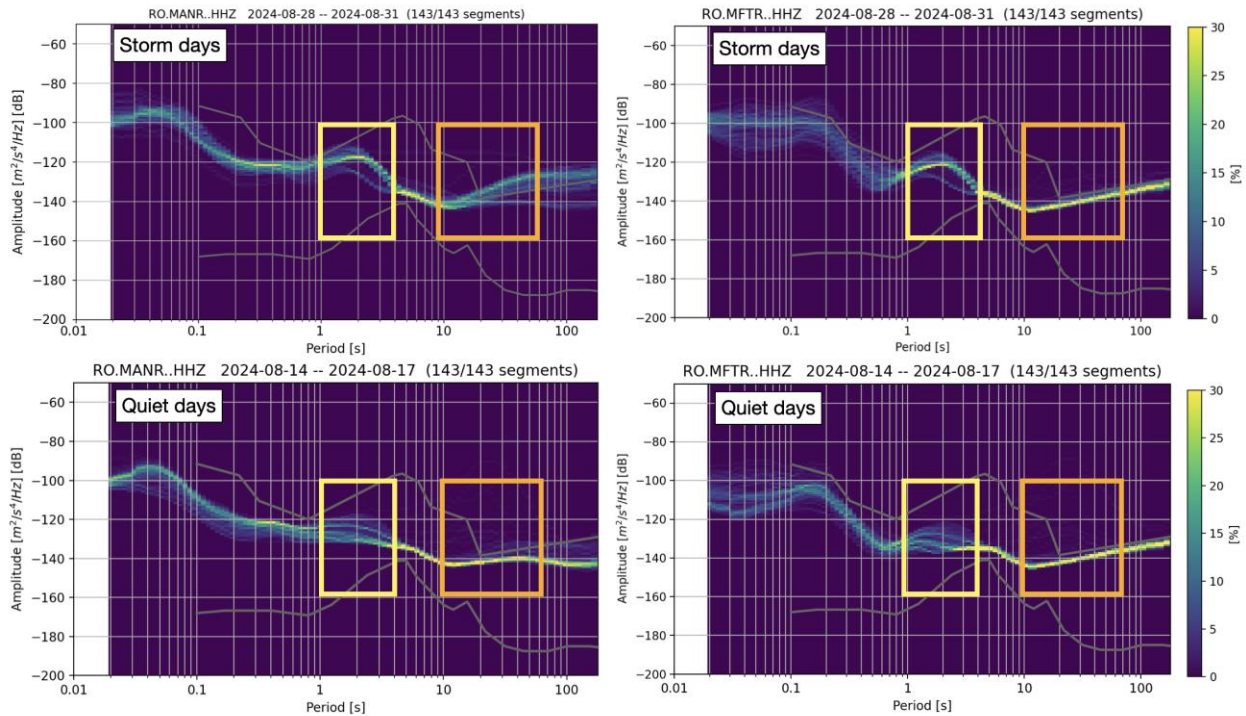
398
 399 *Figure 5. Distribution of hourly RMS of high pass filtered (>30 Hz) seismograms and precipitation data.*
 400 *Colours indicate hourly RMS amplitude of velocity envelopes filtered 30-50 Hz. Background coloured*
 401 *grid indicates the total precipitation (mm) from ERA5 data.*

402 The analysis of the microseismic noise frequency band is closely linked to the interaction between
 403 ocean waves and the seafloor, which is influenced by storm conditions. To assess the storm's
 404 impact, we analyze the PPSD (Probabilistic Power Spectral Density) of noise recorded at several
 405 stations during both storm and quiet days, using the latter as a baseline. Figure 6 shows examples
 406 of PPSD at stations MANR and MFTR (Figure 2), revealing differences in PSD amplitudes across
 407 the primary and secondary microseismic bands. These differences indicate the presence of high-
 408 intensity wind-driven waves and swell energy in the sea.

409 The secondary microseismic band, in particular, shows a significant rise in amplitude during
 410 storms, consistent with established mechanisms linking storm-driven wave activity and seafloor
 411 pressure fluctuations (Figure 3) to enhanced secondary microseism generation, while local factors
 412 such as bathymetry or wave direction may modulate the response (Bromirski et al., 2002; Ebeling

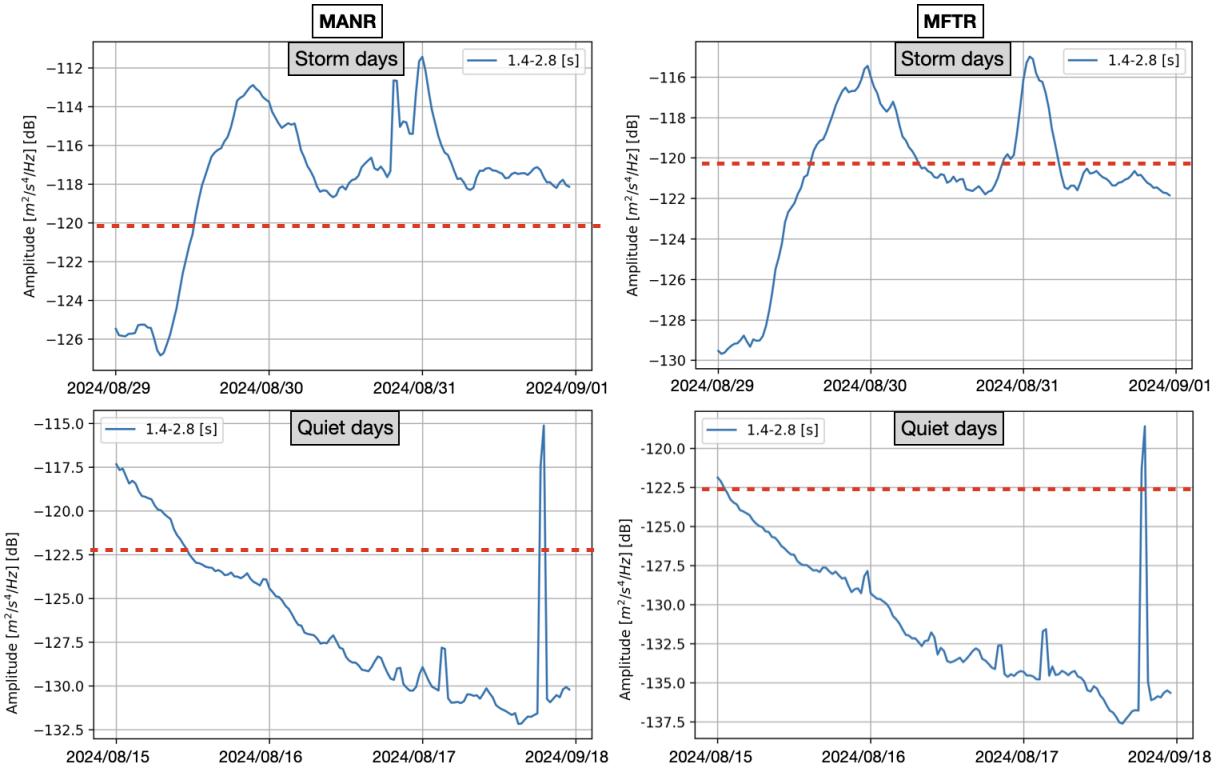
413 et al., 2012; Arduin et al., 2019). On quiet days, the PSD remains consistently lower, typically
 414 staying below the -120 dB threshold. This stark contrast emphasizes the role of atmospheric
 415 conditions in modulating seismic noise, with storms causing a notable increase in energy across
 416 both frequency bands. The temporal evolution of the PSD values (Figure 7) further highlights the
 417 storm's impact, with fluctuations corresponding to changes in environmental factors, reinforcing
 418 the connection between storm activity and the observed seismic signals.

419



420
 421 *Figure 6. Probabilistic Power Spectral Density (PPSD) plots for two seismic stations near the Black Sea*
 422 *coast capturing the target storm signal in the microseismic bandwidths (marked with rectangles). The top*
 423 *panels show the PPSD distributions across frequencies, indicating the probability of power spectral*
 424 *density values in percentage for days including the Black Sea storm. The bottom panels show PPSD for*
 425 *days with no recorded events.*

426 Anthropogenic seismic noise does not significantly affect the microseismic band (0.1-1 Hz).
 427 Human-generated vibrations predominantly occupy frequencies above 1 Hz, while long-period
 428 microseisms are produced by ocean wave interactions and are coherent over large distances. The
 429 temporal evolution of the microseismic energy observed in this study matches changes in wave
 430 state associated with the storm rather than any local activity. Similar to the findings of Gross &
 431 Ritter (2009), the sub-Hz frequency range is dominated by natural sources, with anthropogenic
 432 contributions being negligible.



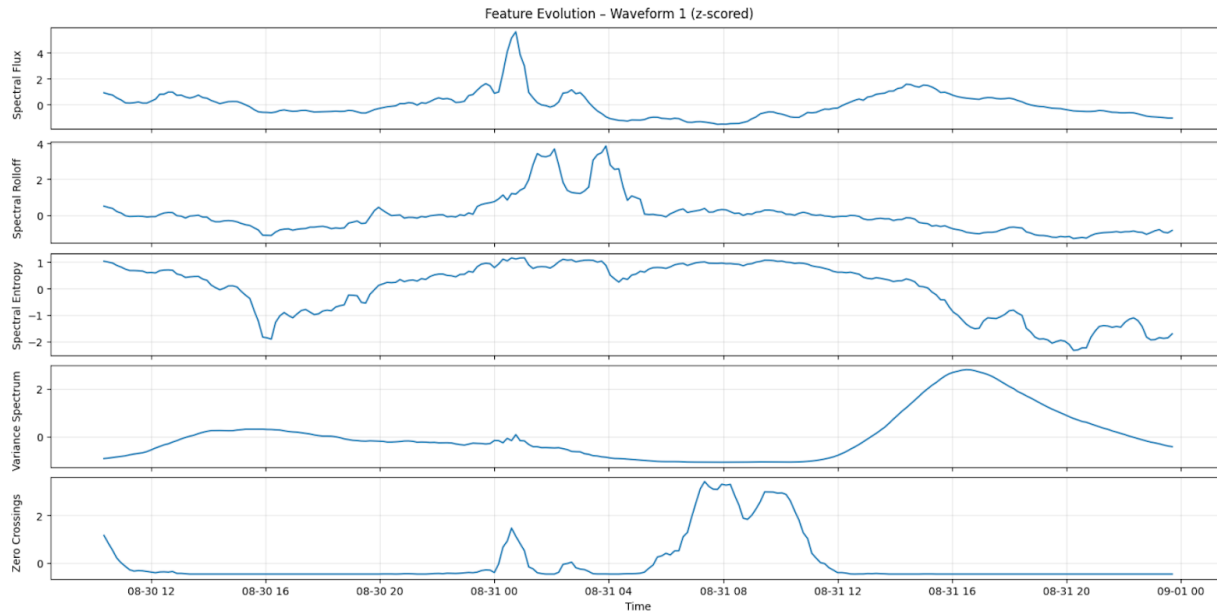
433

434 *Figure 7. Temporal PSDs for two seismic stations near the Black Sea Coast in the secondary microseismic*
 435 *band showing significantly higher values (>-120 dB) during stormy days compared to days without*
 436 *recorded precipitation.*

437 4.2 Infrasonic and satellite lightning observations

438 4.2.1 Single-station feature extraction

439 The evolution of time-frequency features over the duration of the Black Sea storm revealed distinct
 440 patterns in the infrasonic signal (Figures 8 and 9). Centroid and rolloff show parallel behavior
 441 because they are both frequency-domain descriptors tied to the distribution of spectral energy, and
 442 so both respond strongly to the same uplift in energy during the storm's peak. Spectral flux, by
 443 contrast, quantifies inter-frame spectral change, so its peak occurs where the spectrum transitions
 444 most rapidly, even when that does not coincide with the maximum absolute energy (e.g. Pásztor
 445 et al., 2023). Finally, the zero-crossing rate reflects time-domain volatility, not spectral shape,
 446 which explains its distinct pattern, such as the storm's later stages may introduce broadband
 447 turbulence or noise components that boost zero crossings independently of the spectral shifts
 448 visible in the first two panels. While the individual features varied over time, it is the combination
 449 of these features through K-means clustering that effectively identifies the time frame
 450 corresponding to the main precipitation episode. Zero-crossings exhibited more variable patterns,
 451 reflecting high-frequency fluctuations, but the joint clustering of all features robustly captures the
 452 timing of the storm's most intense phases.

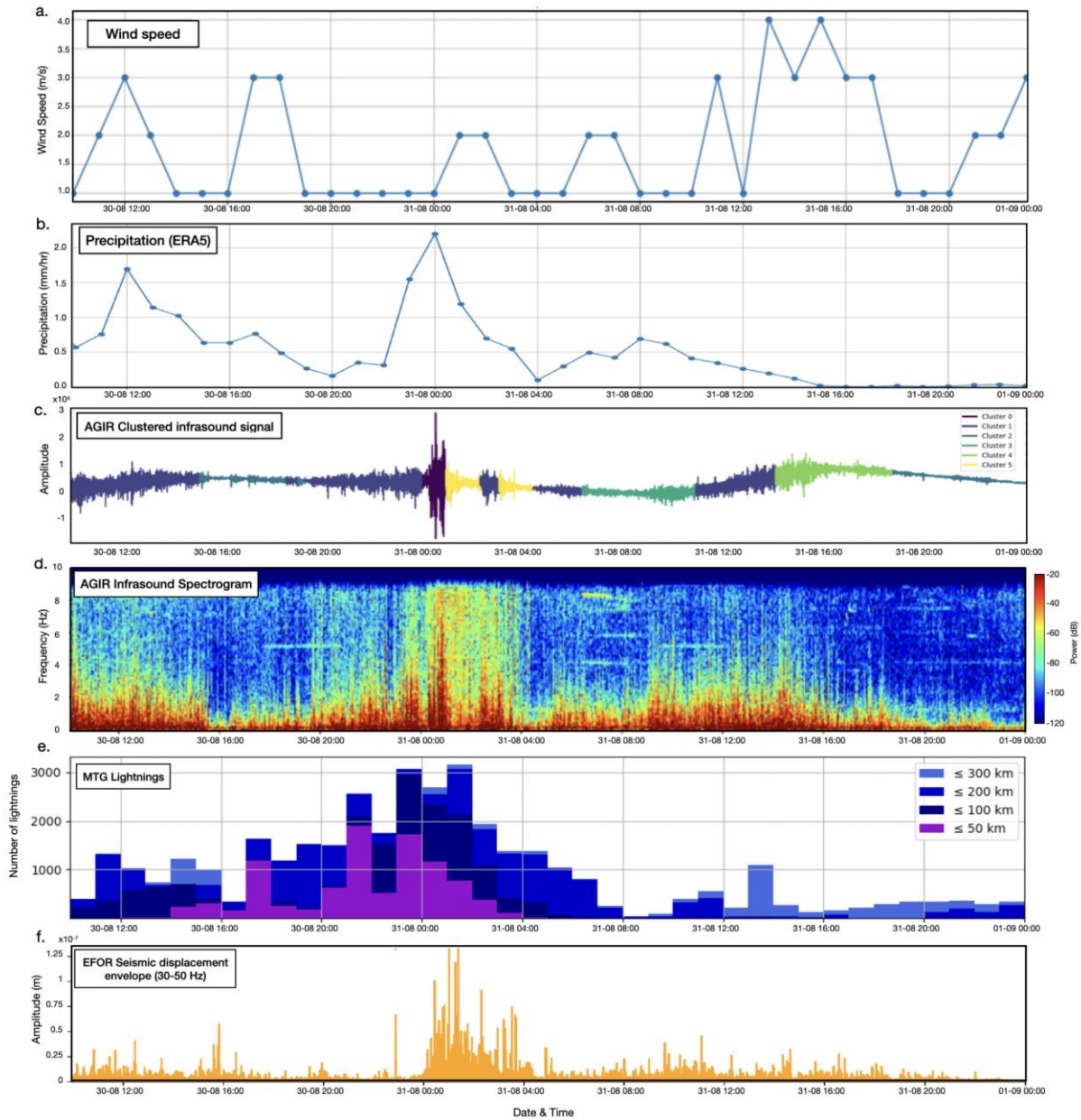


453

454 *Figure 8. Time-frequency feature analysis for the single-station infrasound signal recorded at AGIR during*
 455 *the Black Sea Storm.*

456 K-means clustering separated the acoustic data into six groups with distinct spectral and amplitude
 457 characteristics (Figure 9). These clusters highlight acoustic states that may relate to different
 458 environmental conditions during the monitoring period. For example, Cluster 0 coincides with
 459 periods of intense precipitation and stronger winds, and also aligns with enhanced lightning
 460 activity, suggesting a link with the most energetic phases of the convective system. Cluster 1
 461 captures intervals with moderate amplitudes but persistently elevated background acoustic levels,
 462 without corresponding rainfall or wind peaks, and with comparatively reduced lightning
 463 occurrence. Cluster 2 reflects calmer conditions with low amplitudes and little or no precipitation
 464 and minimal lightning activity. Transitional patterns also arise, such as Cluster 3, which appear
 465 before intervals grouped in Cluster 1 and mark intermediate acoustic activity. Overall, the
 466 clustering approach demonstrates that combining multiple features reveals consistent acoustic
 467 regimes and can help differentiate environmental conditions, without relying on any single
 468 parameter.

469 Interestingly, the spectral content of the infrasound signal showed similarities to seismic signal
 470 envelopes, particularly in the high frequency ranges (Figure 9), which may suggest a connection
 471 between the atmospheric pressure waves detected by infrasound and the ground vibrations
 472 captured by seismic instruments. When considered alongside the temporal evolution of lightning
 473 activity, this overlap implies that both seismic and infrasound signals could be complementary in
 474 capturing different aspects of storm dynamics, with seismic signals reflecting ground vibrations
 475 and infrasound capturing the atmospheric processes, and lightning indicating convective intensity.



476
 477 *Figure 9. Clustering results of infrasound signals recorded at AGIR from the August Black Sea storm event*
 478 *compared with precipitation data and seismic signal from EFOR station. a. Wind speed per hour measured*
 479 *at ANM meteorological station Constanta; b. Average precipitation data from $1^{\circ} \times 1^{\circ}$ around AGIR. c. Raw*
 480 *infrasound signal recorded at the AGIR sensor during the period of August 30–31, 2024, with different*
 481 *segments color-coded according to the cluster they belong to, based on K-means clustering of 30-minute*
 482 *time-frequency feature windows. d. The corresponding spectrogram generated using Blackman windowing*
 483 *with 128 samples and 70% overlap; e. Number of lightnings detected by the MTG satellite. f. Seismic*
 484 *displacement envelope at station EFOR, filtered between 30–50 Hz.*

485

486

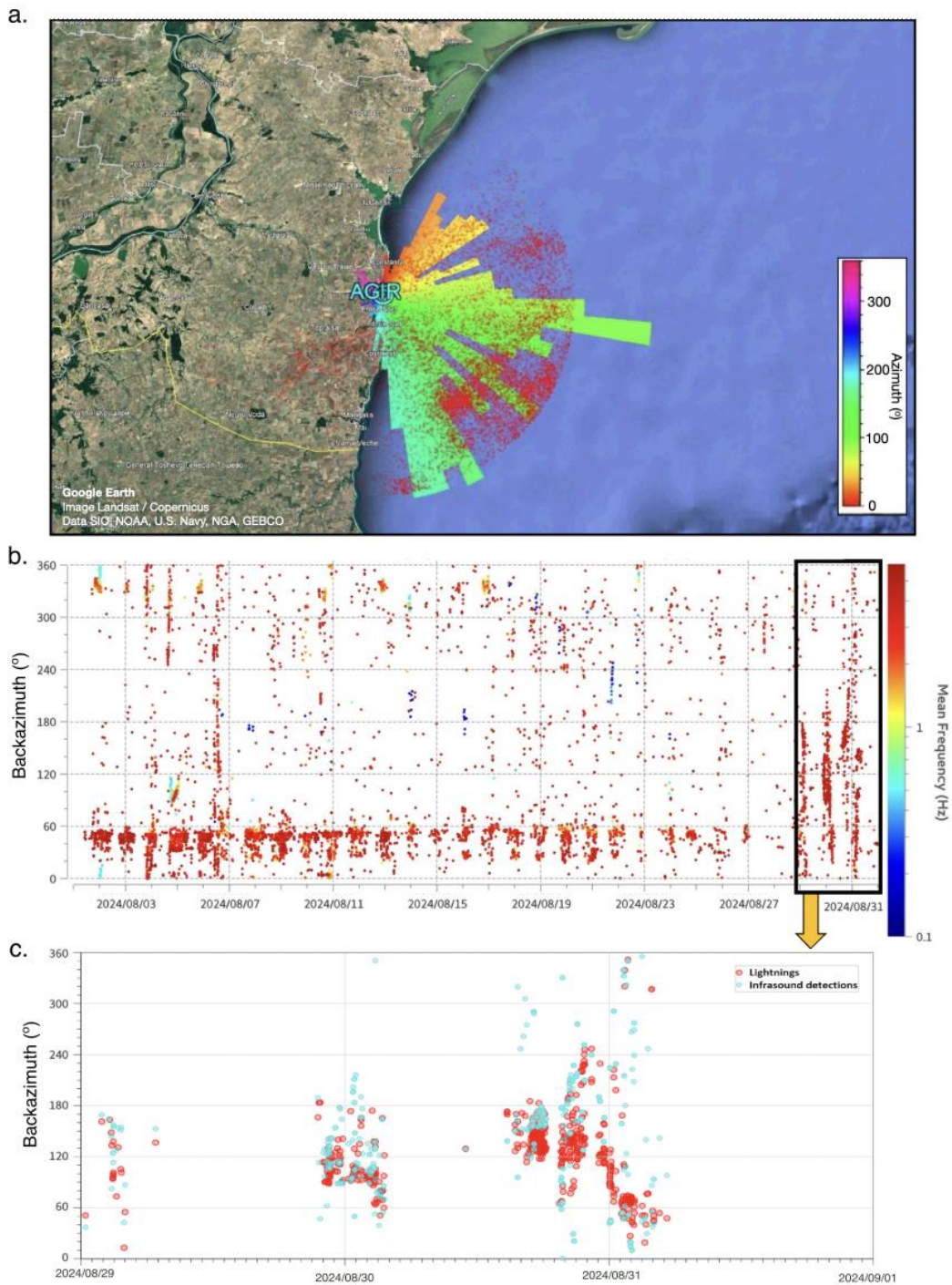
487

488 4.2.2 Array analysis and lightning detection

489 Using multiple sensors from the AGIR infrasound array with the PMCC algorithm allowed us to
490 isolate coherent infrasound signals and estimate their propagation parameters, such as backazimuth
491 and arrival times, across the sensor network. Nearly 11,000 lightning flashes were detected by the
492 MTG Lightning Imager within 50 km of the AGIR station during the Black Sea storm (Figure
493 10a), providing a dense reference set for comparison. Within the relevant 0.6-7 Hz frequency band,
494 approximately 1,100 infrasound detections were identified. These signals primarily consisted of
495 long-duration wave trains with frequent amplitude peaks and short-lived disturbances
496 characteristic of lightning-generated infrasound. Their dominant frequency was around 3 Hz, and
497 amplitudes ranged from 0.01 to 3.4 Pa (Figure 10b).

498 A subset of the infrasound detections could be confidently associated with individual lightning
499 discharges based on temporal alignment and backazimuth consistency. Overall, we were able to
500 associate 6.4% of lightning flashes with infrasound detections at AGIR (Figure 10c). This level of
501 correspondence is reasonable given that only a subset of lightning discharges radiate infrasound
502 efficiently and that propagation strongly depends on altitude, source geometry, and atmospheric
503 conditions. Nevertheless, these coincident detections demonstrate that lightning-generated
504 infrasound was clearly recorded during the storm and that the PMCC-derived acoustic signatures
505 provide valuable insight into the evolution of electrical activity and storm dynamics.

506 Anthropogenic noise sources, such as wind turbines (e.g., Jakobsen, 2005), industrial machinery
507 (Gastmeier and Howe, 2008), and road traffic (Grafkina et al., 2019), are well-documented
508 challenges for infrasound studies because they often generate persistent, periodic, or tonal signals
509 that can mask natural atmospheric phenomena. The AGIR infrasound array used here is located in
510 a semi-rural setting, distant from major roads and industrial facilities, which reduces the likelihood
511 of local anthropogenic contamination. Several independent lines of evidence indicate that such
512 contamination is negligible in this case study. First, the strongest infrasound signatures occurred
513 during night-time hours, when human activity is minimal. Second, both the clustering and PMCC
514 analyses identify transient signals with energy peaking around ~ 3 Hz, which contrasts sharply with
515 the more continuous or harmonic spectral patterns typically produced by anthropogenic sources.
516 Third, the temporal alignment of these acoustic signatures with independent observations of
517 lightning and precipitation provides ~~strong~~ confirmation that the detected infrasound variability is
518 storm-related rather than anthropogenic in origin.



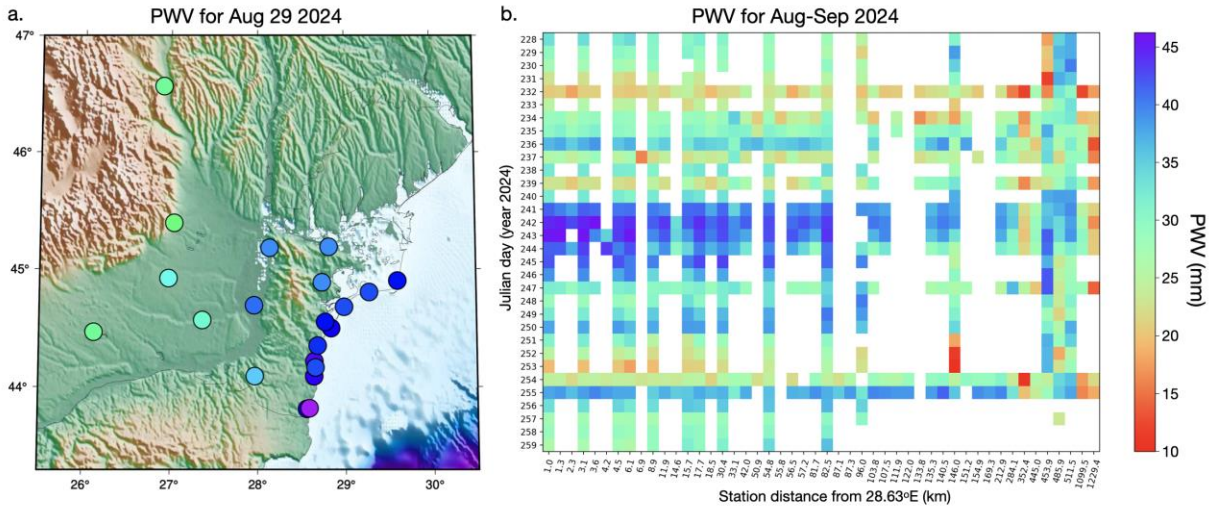
519

520 *Figure 10. a. Polar histogram of infrasound detections from the AGIR station, displayed with © Google*
 521 *Earth, along with the locations of lightning strikes detected by the MTG Lightning Imager system for the*
 522 *period from 29-31st of August 2024. The geographical position of the AGIR infrasound station is also*
 523 *shown on the map. b. diagram of high-frequency detections from the AGIR infrasound station between 1-*
 524 *31st of August 2024. c. Associations between events detected by the AGIR infrasound array and the MTG*
 525 *satellite database for 29-31st of August 2024.*

526 **4.3 GNSS-derived precipitable water vapor trends**

527 The analysis of daily GNSS-derived precipitable water vapor (Figure 11) reveals clear temporal
528 variations, with the highest PWV values consistently recorded on stormy days (>40 mm on DOY
529 240-243, i.e. August 27-30). Notably, the peak values occurred between DOY 241 and DOY 243
530 (Figure 11b), when the heaviest rainfall was observed (Figure 1). Coastal stations showed
531 extremely high PWV values (>40 mm) compared to inland stations (<30 mm), with a slight
532 decrease in PWV away from the coast (Figure 11a). This spatial distribution highlights the
533 geographical gradient of atmospheric moisture, with the highest PWV concentrations near coastal
534 areas, also decreasing gradually toward the north away from the storm peak. Interestingly, some
535 inland stations (BUCU, PGNL, RMSR) recorded their peak PWV on DOY 255, corresponding to
536 the onset of the Boris storm, another significant extreme rainfall event that swept through Central
537 and Eastern Europe (Athanasé et al., 2024).

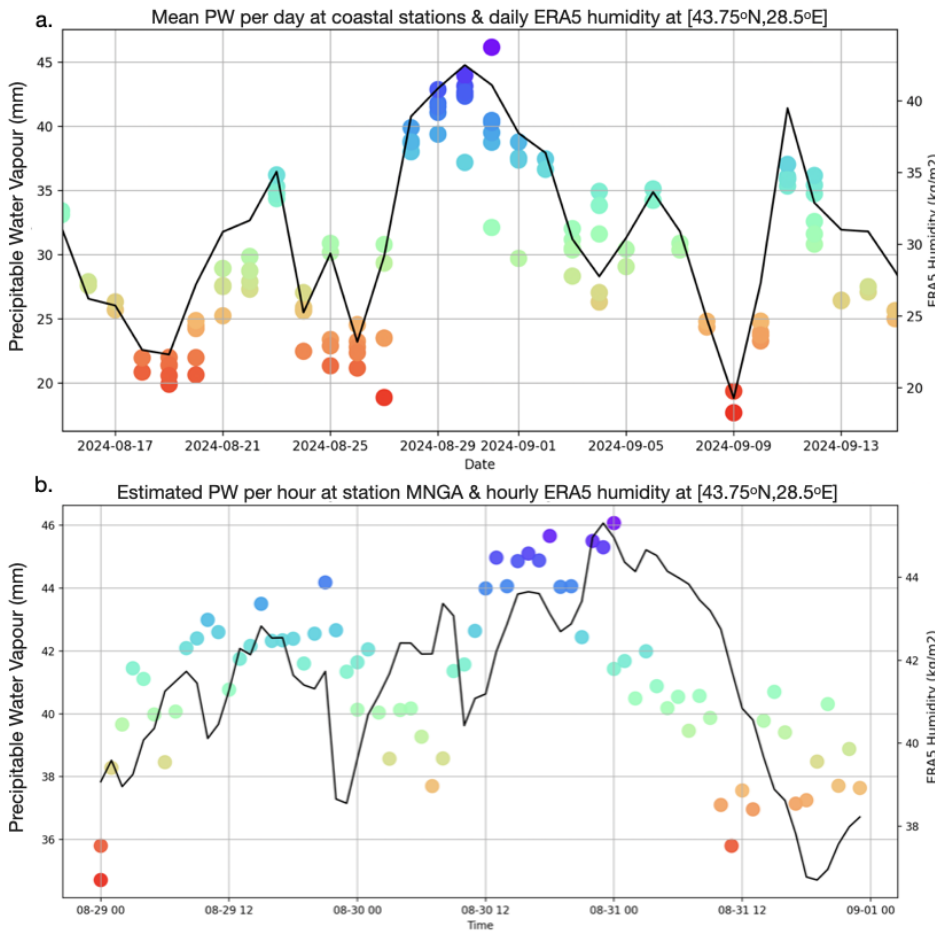
538 Elevated PWV was observed as early as August 27th (Figure 12a), suggesting that the tropospheric
539 moisture loading began to increase several days before the onset of the rainfall. This increase in
540 PWV may act as an indicator of a developing weather system. Remarkably, although HAR1,
541 located inland, did not directly experience the extreme rainfall, it exhibited similar PWV behavior
542 to coastal stations, suggesting that GNSS stations, even outside the immediate storm zones, can
543 capture atmospheric signals indicative of intense precipitation. This finding offers a valuable
544 precedent, showing that PWV measurements at GNSS stations not directly in the storm path can
545 still provide critical insights into moisture dynamics at the tropospheric level. The comparison
546 with ERA5’s total-column water vapour further supports this interpretation, as the broad temporal
547 evolution of ERA5 humidity mirrors the GNSS-derived daily PWV patterns, despite the inherently
548 coarser resolution of the reanalysis data.



549
550 *Figure 11. a: Map of GNSS stations coloured as a function of PWV estimated for the 29th of August. b.*
551 *Daily PWV values for each station, plotted as a function of their longitudinal distance relative to 28.63°E.*

552 Using the hourly PWV data, Figure 12b illustrates the evolution of water vapour at the MNGA
553 station, which recorded the heaviest rainfall in the study area. Notably, MNGA also showed a rapid

554 buildup of PWV, reaching values greater than 44 mm just a few hours before the storm event. This
 555 rapid increase in PWV suggests that the accumulation of atmospheric moisture may precede to
 556 extreme weather events, such as intense rainfall and storms. This observation aligns with known
 557 atmospheric dynamics, where a significant increase in water vapor content precedes heavy
 558 precipitation, highlighting the potential usefulness of GNSS-based PWV monitoring for studying
 559 pre-storm atmospheric moisture variability. The general rising trend toward the event is present in
 560 both GNSS-based and ERA5 reanalysis datasets, although some minor fluctuations are not
 561 matched. After the storm, the GNSS PWV drops sharply while ERA5 maintains elevated values
 562 for several hours. These differences show that GNSS can resolve rapid, real-time atmospheric
 563 changes that may be blurred in large-scale weather model products.



564

565 *Figure 12. Comparison of GNSS-derived precipitable water vapor (coloured circles) with independent*
 566 *humidity data (black line) from ERA5's total column vertically integrated water vapor parameter. a. Mean*
 567 *PWV values per day at selected GNSS stations (CONB, HARI, MNGA, MNGL, MNGM, TUZL) alongside*
 568 *daily averaged ERA5 humidity. b. Estimated PWV values per hour at the MNGA GNSS station compared*
 569 *with hourly ERA5 humidity at the 43.75°N, 28.5°E grid point. In both panels, GNSS data points are*
 570 *represented by colored circles where the color scale corresponds to the PWV magnitude, consistent with*
 571 *the vertical axis.*

572 **5. Discussion**

573 Storm evolution, in the meteorological sense, describes the sequence of processes from pre-storm
574 atmospheric moisture accumulation to convective initiation, peak rainfall, electrical activity, and
575 the associated marine response along coastal areas. The multi-sensor dataset used here captures
576 these different stages: GNSS-PWV documents the build-up of column water vapor before
577 convective onset, infrasound detects lightning-generated acoustic waves and pressure disturbances
578 during the mature convective phase, high-frequency seismic noise reflects the timing and spatial
579 progression of intense rainfall at the surface, microseisms respond to storm-driven changes in sea
580 state, and ERA5/MTG provide the mesoscale structure that ties these geophysical signals together.
581 By observing the same storm through these complementary physical pathways, we can outline a
582 more detailed picture of how the storm developed, intensified, and decayed than is possible from
583 individual datasets.

584 The integration of infrasound, seismic, and GNSS data in monitoring the extreme storm event over
585 the Black Sea provides valuable insights into the dynamics of storm behavior and illustrates the
586 potential of repurposing non-conventional sensors for meteorological analysis. Infrasound data,
587 for instance, revealed a clear acoustic signature of lightning activity, with signals detected in the
588 range of 0.6 to 7 Hz corresponding to electrical discharges. The high frequency of infrasound
589 detections (around 1,100) suggests that it may provide useful information for tracking storm-
590 related phenomena, particularly lightning, which is difficult to capture with traditional methods.
591 However, the signals did not always perfectly align with lightning strikes, indicating that other
592 factors, such as the movement of storm systems or variations in atmospheric conditions, may
593 influence infrasound signatures. This suggests that refining the correlation between infrasound
594 signals and lightning activity could be an avenue for future research, particularly in cases of sparse
595 lightning or in remote regions.

596 Seismic data alone also showed a strong connection between high-frequency seismic noise and
597 heavy rainfall, supporting previous studies that linked seismic signals to rainfall intensity. The
598 distinction between high-frequency and low-frequency seismic noise is particularly noteworthy.
599 Increases in high-frequency seismic noise occurred during intense precipitation, while low-
600 frequency signals were associated with wave height and storm-driven winds. This suggests that
601 different seismic frequencies capture distinct storm dynamics, with high-frequency signals
602 reflecting localized rainfall impacts and low-frequency signals tied to broader atmospheric and
603 oceanic interactions. This dual-frequency approach provides a more nuanced interpretation of
604 seismic data in storm monitoring, highlighting its complexity.

605 A key finding of this study is that K-means clustering of multiple acoustic features, including
606 spectral centroids, roll-off, flux, and zero-crossing rate, effectively segmented the infrasound
607 record into distinct storm phases. This approach proved more robust than relying on any single
608 feature, as the combination captured the complex, evolving nature of the storm's acoustic signature.
609 For instance, the identified clusters delineated periods of intense precipitation, elevated
610 background acoustic levels, and calm intervals, providing a data-driven overview of the storm's
611 progression. The fact that these acoustically defined phases align with independent meteorological
612 observations, such as rainfall peaks, confirms that the infrasound signal variability is a direct
613 response to the storm's atmospheric dynamics.

614 Furthermore, the observed spectral similarity between the infrasound signals and high-frequency
615 seismic envelopes suggests a coupled seismo-acoustic response to the storm. This implies that the

616 same atmospheric forcing, such as pressure fluctuations from rainfall and wind, generates
617 complementary signals in the atmosphere (infrasound) and the ground (seismic waves). Our
618 findings are consistent with other studies of intense weather systems, where coupled microbarom-
619 microseism signals have been shown to track storm structure and evolution (e.g., Butler & Aucan,
620 2018; Smirnov, 2021). The coherent acoustic and seismic responses to atmospheric-oceanic
621 pressures, as also documented in Distributed Acoustic Sensing studies (Taweessintananon et al.,
622 2023) and surf studies (Francoeur et al., 2025), reinforce the interpretation of a shared source
623 mechanism. Therefore, a major and logical next step is to move beyond analyzing these datasets
624 in parallel and to perform joint clustering of seismo-acoustic data (e.g. Floroiu et al., 2025). Such
625 an integrated approach could unlock a more comprehensive, multi-physics understanding of storm
626 dynamics by simultaneously characterizing the coupled atmospheric and ground-borne wavefields.

627 The temporal variations observed in GNSS-derived integrated precipitable water vapor provide
628 valuable insights into atmospheric moisture dynamics before extreme weather events. The
629 pronounced increase in PWV, particularly in the days leading up to and during the storm, supports
630 the link between elevated atmospheric water vapor and precipitation. Notably, the PWV buildup
631 starting roughly three days before the extreme rainfall suggests that rising moisture levels in the
632 troposphere were observed prior to the intense precipitation in this event. Even stations located up
633 to 130 km inland, such as HAR1, recorded similar PWV trends, indicating that GNSS stations
634 outside direct storm zones can still provide crucial atmospheric data. Hourly PWV trends further
635 revealed a rapid increase several hours before precipitation, with values exceeding 44 mm/hr,
636 highlighting the accumulation of moisture just before heavy rainfall. These findings align with the
637 notion that increasing atmospheric moisture acts as a precursor to intense precipitation,
638 highlighting the potential of GNSS-based PWV monitoring for tracking moisture and
639 understanding short-term atmospheric fluctuations.

640 The integration of GNSS, infrasound, and seismic data provides a more comprehensive
641 understanding of storm dynamics than any single data source alone. The synergy between these
642 diverse sensor types allows for the detection of atmospheric moisture, lightning activity, rainfall-
643 induced seismic signals, and storm-driven oceanic interactions. Future research should focus on
644 refining unsupervised learning algorithms for infrasound and seismic signal classification,
645 optimizing joint clustering techniques, and improving the integration of these data sources to
646 enhance storm forecasting and early-warning systems. This case study ~~highlights suggests that this~~
647 ~~the potential of~~ multi-sensor ~~observations approach may help improve our ability to predict to~~
648 ~~enhance our understanding of~~ extreme ~~storm behavior and support the development of future~~
649 ~~early-warning strategies.~~ weather events, understand their impacts, and mitigate associated risks.

650 **6. Conclusions**

651 This study presents a comprehensive analysis of a record-breaking storm over the Black Sea, using
652 a combination of GNSS, infrasound, and seismic data to capture the dynamics of extreme weather
653 events. Our findings underscore the power of multi-sensor networks in enhancing the
654 understanding of storm behavior, particularly in the context of atmospheric moisture, lightning
655 activity, and storm-induced seismic signals. GNSS-derived integrated precipitable water vapor
656 indicates a clear buildup of atmospheric moisture hours before the onset of heavy rainfall,
657 providing valuable insights into the lead-up to extreme precipitation events. Infrasound and
658 seismic data further complemented this analysis, with infrasound providing useful observations of

659 lightning activity and seismic data revealing the link between rainfall intensity and high-frequency
660 seismic noise.

661 The storm analysed here represents an exceptional meteorological event and one of the most
662 intense storms recorded in the region in recent years. ~~According to An attribution analysis using~~
663 ~~the ClimaMeter analysis of framework~~ (Antonescu et al. (2024), ~~the characteristics of this storm~~
664 ~~places~~ identified a detectable anthropogenic climate change signal associated with this event,
665 ~~placing~~ it within the broader context of extreme ~~events under in a changing~~ warming climate
666 ~~conditions~~. The integration of GNSS, infrasound, and seismic data provides a more nuanced and
667 holistic view of storm dynamics, highlighting the need for advanced monitoring systems to predict
668 and respond to such extreme events. Looking forward, such multi-sensor approaches may support
669 future developments in integrated environmental monitoring and research into early-warning
670 capabilities, ultimately contributing to improved understanding and characterization of high-
671 impact atmospheric events.

672 7. Code availability

673 Seismic data were processed with the open-source python framework for seismology Obspy
674 (Beyreuther et al., 2010). Infrasound data was processed with the WinPMCC software (Le Pichon
675 et al., 2010) developed by CEA/DASE (French Atomic Energy Commission, Environmental
676 Assessment and Monitoring Department) and open-source Python libraries for signal processing.
677 Some of the figures were made with GMT (Generic Mapping Tools, Wessel et al., 2019). The
678 GNSS data was processed using Gamit/Globk (Herring et al., 2020) developed by Massachusetts
679 Institute of Technology (<http://www-gpsg.mit.edu/gg/>).

680 8. Data availability

681 Processed infrasound, seismic, and GNSS-derived integrated water vapour data can be visualized
682 and accessed via the INFP monitoring platforms: <https://infp.ro/dashboard-reactive.php> and
683 <https://reactive.infp.ro/events/>. Seismic data are part of the Romanian National Seismic Network
684 maintained by the National institute for Earth Physics (NIEP, www.infp.ro) and are freely
685 available in the miniseed format via EIDA (European Integrated Data Archive,
686 <https://www.orfeus-eu.org/data/eida/>). GNSS data are available for download from NIEP
687 (<http://gps.infp.ro/#/download>) and are provided in the standardized RINEX v2 format, with 24-
688 hour files sampled at 30-second intervals. Infrasound data at AGIR are available to download from
689 NIEP via FDSN dataselect web service. Hourly hydro-meteorological data were obtained from the
690 Copernicus Climate Change Service, Climate Data Store (<https://doi.org/10.24381/cds.bd0915c6>),
691 ERA5 dataset (Hersbach et al., 2023). Limited wind and precipitation data were downloaded from
692 the National Meteorological Agency from <https://www.meteoromania.ro/grafice/> (accessed on
693 13.09.2024) and from https://www.meteoromania.ro/clim/caracterizare-lunara/cc_2024_08.html
694 (accessed on 10.11.2025), respectively. Lightning data came from Meteosat Third Generation
695 Lightning Imager operated by EUMETSAT (The European Organisation for the Exploitation of
696 Meteorological Satellites, <https://www.eumetsat.int/>).

697

698 9. Author contribution

699 **Laura Petrescu:** Conceptualization, Methodology, Software, Formal analysis, Data Curation,
700 Writing-Original Draft, Visualization; **Bogdan Antonescu:** Conceptualization, Writing-Review &
701 Editing, Visualization; **Sorin Nistor:** Software, Formal Analysis, Data curation, Visualisation,
702 Writing-Review & Editing; **Iustin Floroiu:** Methodology, Software, Formal analysis, Data
703 Curation, Writing-Original Draft, Visualization; **Dragoş Ene:** Software, Formal analysis, Data
704 Curation, Writing-Review & Editing; **Daniela Ghica:** Software, Formal analysis, Data Curation;
705 **Constantin Ionescu:** Funding Acquisition, Resources, Project administration; **Andrei Anghel:**
706 Methodology, Supervision; **Mihai Dăţcu:** Methodology, Supervision, Funding Acquisition,
707 Resources, Project administration.

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713 anonymous reviewers for their valuable feedback, which helped improve the manuscript.

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719 Digitalization of Romania.

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