

Supporting Information for  
***Seismicity and Thermal Structure of the St. Paul Transform System, equatorial Atlantic:  
Insights from Focal Depth Analysis***

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**Contents of this file**

In this document, we provide material with supporting information to the manuscript. We include one additional text (Text S01) and five figures (Figure S01-05), and three tables (Table S01-03). Table S1 lists the focal depths and output information obtained with the surface waveform model of all 35 earthquakes. Table S02 shows the list of stations used in analysis, together their geographical coordinates. Table S03 shows the parameters of the geodynamic modeling described in Text S01. Figure S01/02 shows the lithospheric age/low-density layer of the St. Paul transform system mapped by Maia et al (2016), parallelly. Figure S03 shows all earthquakes with  $M_w > 4.7$  cataloged since 1990 by the Global Centroid Moment Tensor catalog (Ekström et al. 2012). Figure S04 shows a map with the locations of all seismic stations listed on Table S02. Figure S05 shows the results of geodynamic thermal models described on Text S01.

## Text S1

### A geodynamic hydrothermal cooling model (HTCM)

HTCM obtains a steady-state temperature field from geodynamic thermal flow simulations of a 3D ridge-transform system using the geodynamic finite element code ASPECT (Bangerth et al., 2022). This model is built upon the methodology of Liu et al (2023) and involves a nonlinear viscoplastic mantle rheology. We implement the process of hydrothermal cooling over the top 6 km and up to 600 °C by adopting the methodology proposed by Morgan and Chen (1993). This process can efficiently cool the new oceanic crust by promoting more efficient heat loss from the crust (e.g., Morton and Sleep, 1985; Theissen-Krah et al., 2011). Hydrothermal cooling in the shallow brittle lithosphere is parameterized in terms of an effective thermal conductivity enhanced by a non-dimensional Nusselt number, which quantifies the relative efficiency of convective hydrothermal cooling in comparison to conductive heat loss.

In the geodynamic thermal model, the geometry of the transform system is extracted from the bathymetric observations of STPS including the lengths of transform segments A-D. Mantle flow is driven passively by imposing horizontal transform slipping rate of 27.8 mm/yr on the top boundary. All lateral boundaries and the bottom boundary are open so that materials can freely flow in and out of the model domain to maintain mass balance. The models are incompressible and therefore do not resolve thermal contraction and also do not account for viscous dissipation. The model setup is shown in Figure S4 and model parameters are listed in Table S3.

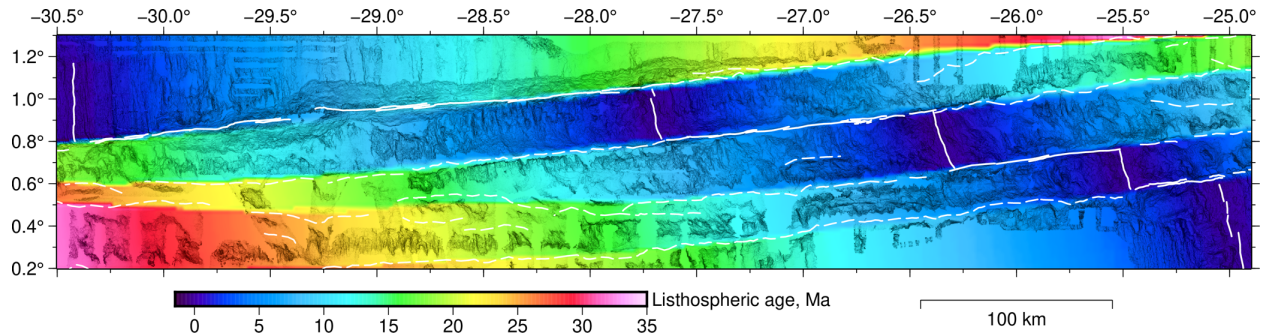


Figure S01 – Lithospheric age mapping of the St. Paul transform system (Maia et al 2016). Solid white lines represent the tectonic plate boundaries, with the dotted while lines indicating the fracture zones.

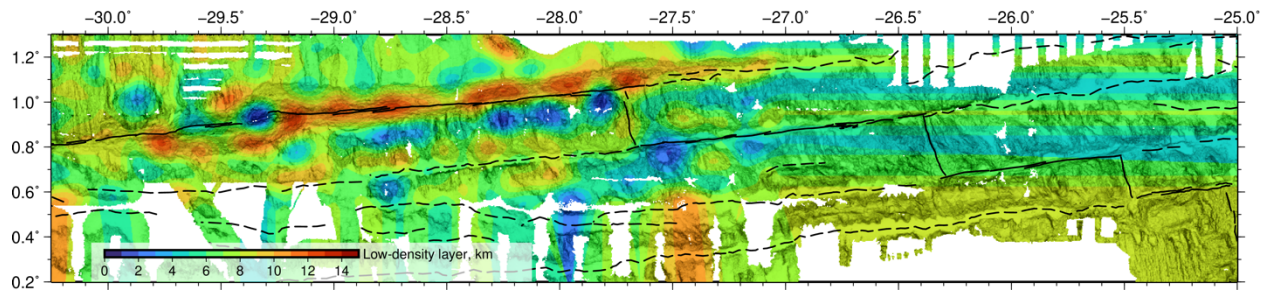


Figure S02 – Low-Density Layer mapping of the St. Paul transform system (Maia et al 2016). Solid black lines represent the tectonic plate boundaries, with the dotted while lines indicating the fracture zones.

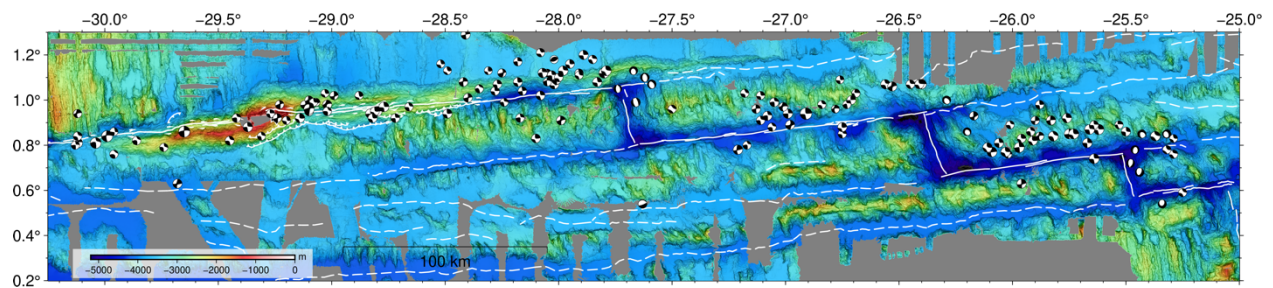
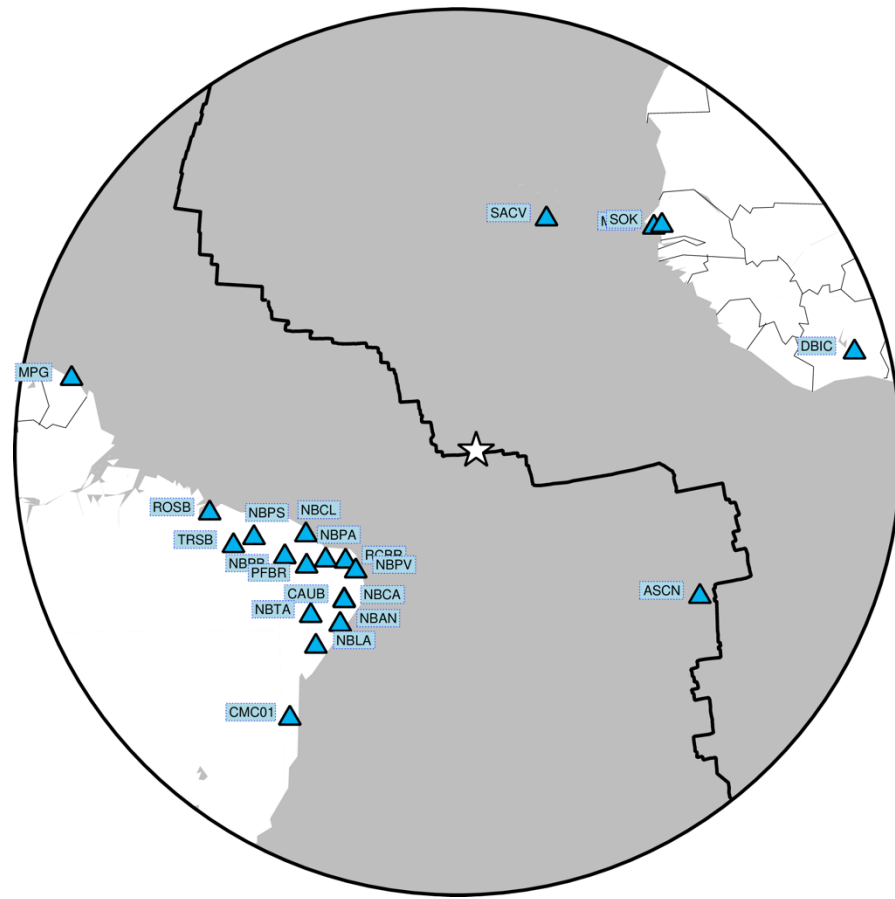
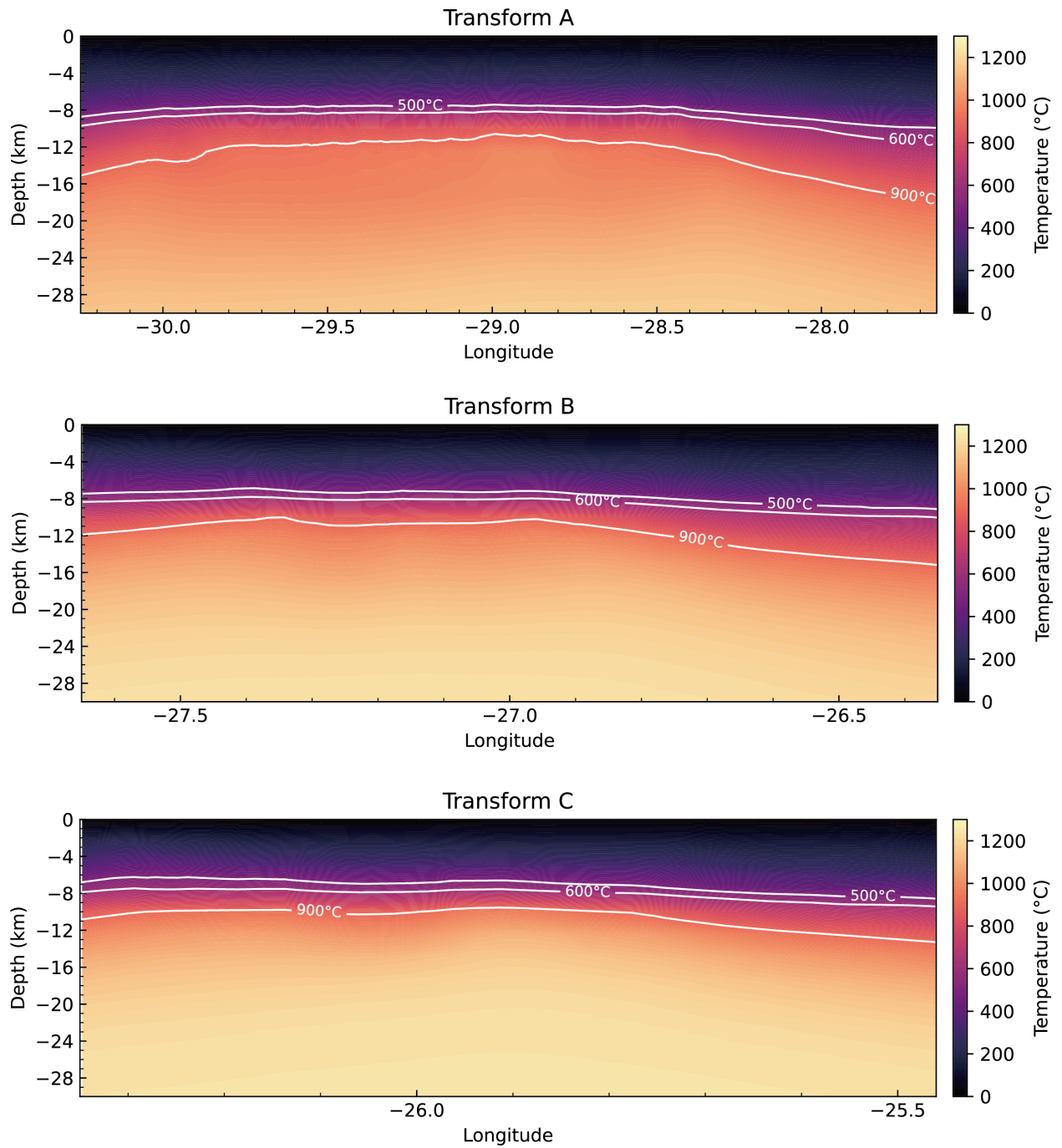


Figure S03 – Historical seismicity at St. Paul Transform System. Black beach balls refers to the historical earthquakes of the Global Centroid Moment Tensor catalog (Ekström et al. 2012) since 1990 with moment magnitude > 4.7.



*Figure S04 – Map with the locations of the 21 earthquakes used in surface waveform modeling.*



*Figure S05 – Results of geodynamic thermal models with a nonlinear viscoplastic mantle rheology referent to Transform A, B, and C. Hydrothermal circulation is parameterized to activate in regions where the temperature falls below 600 °C and the depth is less than 6 km.*

## References

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