

1 **Multi-Machine Learning Ensemble Regionalization of Hydrological
2 Parameters for Enhancing Flood Prediction in Ungauged
3 Mountainous Catchments**

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11

12 **Abstract:**

13 Machine learning-based parameter regionalization is an important method for
14 flood prediction in ungauged mountainous catchments. However, single machine
15 learning parameter regionalization often exhibits limitations in prediction accuracy and
16 robustness. Therefore, this study proposes a multi-machine learning ensemble
17 regionalization method that integrates Gradient Boosting Machine (GBM), K-Nearest
18 Neighbors (KNN), and Extremely Randomized Trees (ERT) methods (GBM-KNN-
19 ERT) to regionalize the sensitive parameters of the Topography-Based Subsurface
20 Storm Flow (Top-SSF) model. Validated across 80 mountainous catchments in
21 southwestern China, the GBM-KNN-ERT method demonstrates superior performance
22 with 90% of ungauged catchments achieving the Nash-Sutcliffe Efficiency (NSE)
23 above 0.9, representing a 67.44% improvement over the best single machine learning
24 parameter regionalization. Notably, the GBM-KNN-ERT method shows improved
25 robustness to climate change and changes in the number of donor catchments compared
26 to other regionalization methods. An optimal balance between accuracy and

27 computational efficiency was achieved using 20-40 high quality donor catchments
28 (NSE greater than 0.85). This study provides systematic evidence that multi-machine
29 learning ensemble can effectively address regionalization challenges in ungauged
30 mountainous regions, offering a reliable tool for water resource management and flood
31 disaster mitigation.

32 **Keywords:** Flood forecasting; Regionalization; Ungauged mountainous catchments;
33 Top-SSF model;

34

35 **Highlights:**

36 1. Proposes a novel multi-machine learning ensemble regionalization method
37 2. The GBM-KNN-ERT method increases the percentage of catchments with high-
38 accuracy flood predictions (NSE >0.9) to 90%, which is a 67.44% improvement
39 over the best single machine learning method.
40 3. The GBM-KNN-ERT method exhibits greater stability under climate change.

41

42 **1. Introduction**

43 Floods in mountainous catchments, encompassing both flash floods and general
44 larger-scale flood events which can be derived from mountainous upland catchments,
45 pose a significant threat to human safety and property, particularly in regions lacking
46 sufficient observational data (Luo et al., 2015; Zhai et al., 2018). While hydrological
47 models like the Topography-Based Subsurface Storm Flow (Top-SSF) model (Li et al.,
48 2024) offer promising simulation capabilities, their application in ungauged catchments
49 is severely limited by the absence of calibration data (Choi et al., 2023; Liu et al., 2018).
50 Effective parameter regionalization methods are therefore essential for transferring
51 hydrological knowledge from gauged to ungauged regions, enabling reliable flood
52 prediction in ungauged mountainous catchment (Garambois et al., 2015; Ragettli et al.,
53 2017; Xu et al., 2018).

54 Parameter regionalization is a crucial method for flood prediction in ungauged
55 catchments (Arsenault et al., 2022; Guo et al., 2021; Kratzert et al., 2019; Zhang et al.,
56 2020). Compared to purely data-driven methods, parameter regionalization offers
57 enhanced physical interpretability (Nearing et al., 2024; Tang et al., 2023; Zhang et al.,
58 2024). Existing parameter regionalization methods can be broadly classified into three
59 categories: similarity-based, hydrological signatures-based, and regression-based
60 (Arsenault et al., 2019; Wu et al., 2022). Similarity-based methods rely on the
61 assumption that catchments with similar characteristics exhibit similar hydrological
62 responses, considering spatial proximity (Arsenault et al., 2019; Pugliese et al., 2018;
63 Yang et al., 2018) and physical similarity (similar climatic and land cover conditions

64 have similar hydrological characteristics) (Kanishka et al., 2017; Papageorgaki et al.,
65 2016). Hydrological signature-based methods use hydrological signatures (quantitative
66 metrics that describe statistical or dynamic properties of streamflow) as an intermediate
67 link, establishing relationships first between model parameters and signatures, and then
68 between signatures and catchment descriptors to facilitate parameter transfer
69 (McMillan, 2021; Zhang et al., 2018). Regression-based methods, which directly link
70 hydrological model parameters to catchment descriptors, are widely used due to their
71 simplicity and computational efficiency (Guo et al., 2021; Kratzert et al., 2019; Song et
72 al., 2022; Wu et al., 2022). However, the performance of regression-based methods is
73 frequently constrained by the inherent nonlinearity in the relationships between model
74 parameters and catchment descriptors, coupled with the difficulty in adequately
75 capturing spatial heterogeneity, especially within complex mountainous terrain (Wu et
76 al., 2022).

77 Recent advances in machine learning offer potential solutions by capturing
78 nonlinear patterns in high-dimensional data. Methods such as Decision Tree (DT),
79 Extremely Randomized Trees (ERT), Gradient Boosting Machine (GBM), K-Nearest
80 Neighbor (KNN), Random Forest (RF), and Support Vector Machines (SVM) have
81 shown promise in parameter regionalization (Golian et al., 2021; Song et al., 2022).
82 However, existing machine learning-based parameter regionalization studies
83 predominantly focus on runoff prediction at coarser temporal scales (daily or monthly)
84 (Li et al., 2022; Wu et al., 2022), leaving a significant gap in high-resolution (hourly or
85 sub-hourly) flood prediction in ungauged mountainous catchments. Moreover, these

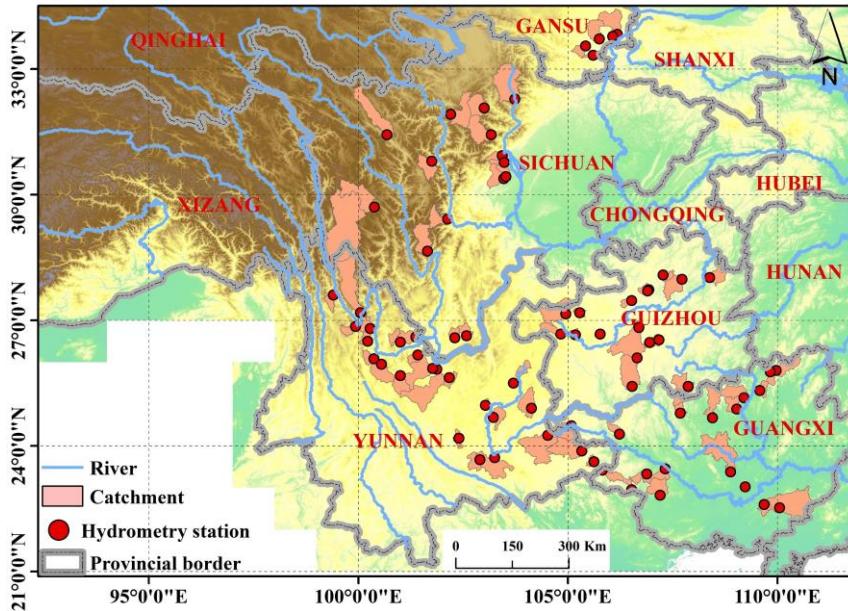
86 studies often rely on single machine learning methods to estimate all hydrological
87 model parameters (Golian et al., 2021; Song et al., 2022; Wu et al., 2022). Given that
88 different machine learning methods operate on distinct principles (Jordan et al., 2015;
89 Zounemat-Kermani et al., 2021) and hydrological model parameters represent diverse
90 hydrological processes (Li et al., 2024), a single machine learning method may not
91 adequately capture the complexity of model parameter estimation (Golian et al., 2021;
92 Wu et al., 2022). Therefore, exploring the multi-machine learning ensemble methods is
93 essential to improve the accuracy of high-resolution flood prediction in ungauged
94 mountainous catchments.

95 Southwest China's mountainous regions are particularly vulnerable to frequent
96 floods, leading to ecosystem degradation through habitat disruption and biodiversity
97 loss (Gan et al., 2018). The abundance of ungauged catchments in this region poses a
98 significant challenge to reliable flood prediction. To address this critical issue, we
99 systematically evaluate the performance of a novel multi-machine learning ensemble
100 method for regionalizing Top-SSF model parameters across 80 representative
101 catchments (mean area: 1,586 km²) in Southwest China. By assessing ensemble method
102 robustness under climate change and with varying donor catchment configurations, this
103 study aims to significantly enhance flood prediction accuracy in ungauged mountainous
104 catchments, contributing to improved ecosystem resilience, enhanced human safety,
105 and more effective water resource management in the face of escalating climatic
106 pressures.

107 **2. Study area and datasets**

108 **2.1. Study area**

109 This study investigated 80 mountainous catchments in Southwestern China,
110 encompassing Sichuan, Yunnan, Guangxi, Guizhou, and Chongqing provinces (Fig. 1).
111 This region exhibits diverse climatic zones, including subtropical monsoon, plateau
112 mountain, and tropical monsoon climates. The selected catchments have an average
113 area of 1,586 km² (ranging from 109 to 6,564 km²), with elevations ranging from 63 to
114 6,284 meters. Mean annual temperature varies from 15 to 20°C, and annual
115 precipitation ranges from 1,200 to 1,800 mm (Li et al., 2016), with approximately 80%
116 of the annual precipitation occurring during summer and autumn, contributing to
117 frequent flooding events (Cheng et al., 2019). These catchments are situated within a
118 heavily forested region, the second largest in China (Hua et al., 2018), with forest cover
119 ranging from 3% to 92% (mean: 51%), influencing evapotranspiration and runoff
120 generation. Dominant soil types, according to the Genetic Soil Classification of China
121 (Shi et al., 2004), include purple soil (12.20%), yellow soil (11.39%), and red soil
122 (9.52%), each with distinct hydrological properties.



123 **Fig.1.** Geographical distribution of the 80 gauged catchments used, with locations of
 124 hydrometry station (red points) and major rivers indicated.
 125

126 **2.2. Datasets**

127 Hourly flow data (2015–2018) for 80 mountainous catchments in China were
 128 sourced from the Hydrological Bureau of the Ministry of Water Resources, through
 129 China's hydrologic yearbooks, encompassing a spectrum of events from flash floods
 130 and general floods which can be derived from mountainous upland catchments. Hourly
 131 rainfall data (2015–2018) were obtained from ground meteorological stations across
 132 China (<http://en.weather.com.cn>), providing crucial input for hydrological modelling.
 133 Additional meteorological variables, including temperature, wind speed, dewpoint
 134 temperature, and surface net solar radiation, were obtained from the ERA5 hourly
 135 dataset (1940–present) (Hersbach et al., 2023), ensuring comprehensive atmospheric
 136 forcing. Relative humidity was estimated using dewpoint temperature. Historical

137 (1901–2021) and projected future (SSP585, 2022–2100) temperature and precipitation
138 data for China, averaged from the EC-Earth3, GFDL-ESM4, and MRI-ESM2-0 models
139 at 1 km resolution, were obtained from "A Big Earth Data Platform for Three Poles" to
140 assess the impact of climate change (Ding et al., 2020) (<http://poles.tpdc.ac.cn>).
141 Topographic data, including a 30m resolution Digital Elevation Model (DEM), used for
142 river network and topographic index derivation, were obtained from EARTHDATA
143 (<https://search.earthdata.nasa.gov/search>). Forest cover data (30m resolution) were
144 sourced from the Global Forest Cover and Forest Change Map
145 (<https://www.noda.ac.cn/>), providing information on vegetation characteristics. Bulk
146 density (BD) data were derived from the Soil Database of China for Land Surface
147 Modelling (Dai et al., 2013). Soil hydraulic parameters, specifically saturated hydraulic
148 conductivity (Ks_CH) for Clapp and Hornberger functions and the pore-connectivity
149 parameter (L) for van Genuchten and Mualem functions, were acquired from the China
150 Dataset of Soil Hydraulic Parameters Using Pedotransfer Functions for Land Surface
151 Modeling (Shangguan et al., 2013).
152

153 **Table 1.** Model forcing data and catchment descriptors information.

Data type	Name	Unit	Function
Hydro-meteorology	Rainfall	mm	Input for hydrological model
	Flood	m ³ /s	Used for model calibration (hourly resolution)
	Temperature	K	Input for hydrological model
	Surface pressure	Pa	
	Dewpoint temperature	K	
	wind speed	m/s	
	Surface net solar radiation	J/m ²	
	Relative humidity	%	
	1 km monthly precipitation (1901-2021)	mm	
	1 km monthly temperature (1901-2021)	°C	
Soil characteristics	1 km monthly temperature (2022-2100, SSP5-8.5, EC-Earth3, GFDL-ESM4, MRI-ESM2-0)	°C	Multi-year surface average as catchment descriptors
	1 km monthly precipitation (2022-2100, SSP5-8.5, EC-Earth3, GFDL-ESM4, MRI-ESM2-0)	mm	
	Soil bulk density (BD)	g/cm ³	
	Pore-connectivity parameter (L) for the van Genuchten and Mualem functions	-	
Topography	Saturated hydraulic conductivity (K _s _CH) of the Clapp and Hornberger Functions	cm d ⁻¹	Surface average as catchment descriptors
	Forest cover (FC)	%	
	DEM	m	
	Topographic index	-	
	Slope	mm ⁻¹	
	Catchment area	km ²	

154

3. Methodology

155

3.1. Hydrological model

156 Top-SSF is a semi-distributed hydrological model based on the well-established

157 TOPMODEL framework, which delineates sub-basins based on the topographic index.

158 It retains the key advantages of TOPMODEL, such as its parsimonious structure,

159 physical interpretability, and ease of parameter transfer (Beven et al., 2021; Gao et al.,

160 2018), consists of 15 parameters representing six key hydrological components: canopy

161 interception, infiltration, evapotranspiration, unsaturated zone moisture transport,

162 subsurface storm flow, and flow routing (Li et al., 2024). In the Top-SSF model, flood

163 can be comprised of four components: infiltration-excess overland flow, saturation-

164 excess overland flow, subsurface storm flow, and groundwater discharge.

165 Infiltration-excess overland flow occurs when the rainfall intensity exceeds the

166 infiltration capacity. In this study, infiltration is simulated using the Green-Ampt model.
 167 When surface ponding occurs, the infiltration rate is determined by solving the Green-
 168 Ampt equation iteratively, for which the Newton-Raphson method is employed. The
 169 infiltration rate (f_{in}) is given by:

$$170 \quad f_{in} = -\frac{Ks(CD+F_{satrt})}{Szm(1-e^{(F_{satrt}/Szm)})} \quad (1)$$

171 where, f_{in} is the infiltration rate (m/h); Ks is surface hydraulic conductivity (m/h);
 172 CD is capillary drive (m); F_{satrt} is the initial cumulative infiltration (m); Szm is the
 173 maximum water storage capacity in the unsaturated zone (m).

174 Saturation excess overland flow occurs at computational cell i when the
 175 groundwater table depth, S_i is less than or equal to zero (i.e., $S_i \leq 0$, indicating the
 176 water table has reached the surface). It is calculated as:

$$177 \quad r_{s,i} = \max\{Suz_i - \max(S_i, 0), 0\} \quad (2)$$

178 where, $r_{s,i}$ is the depth of saturation excess overland flow generated at cell i (m); Suz_i
 179 is the soil water storage in the unsaturated zone, at cell i (m); S_i is the groundwater table
 180 depth at cell i (m).

181 The depth of subsurface storm flow generated at computational cell i , $r_{sf,i}$ is
 182 given by:

$$183 \quad r_{sf,i} = q_{sf0}(1 - S_{sf,i}/S_{fmax}) \quad (3)$$

184 where, $r_{sf,i}$ is the depth of subsurface storm flow at cell i (m); q_{sf0} is initial subsurface
 185 storm flow (m); $S_{sf,i}$ is the water storage deficit in the subsurface storm flow zone at
 186 cell i (m).

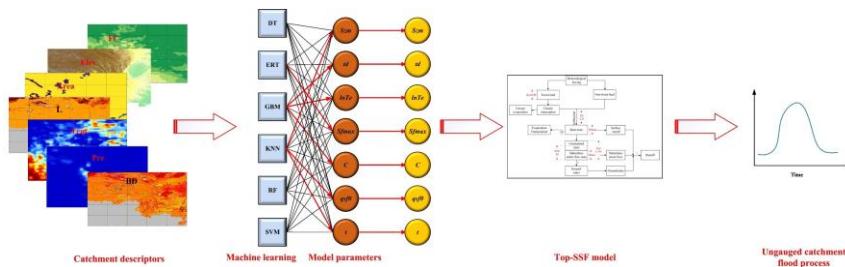
187 The depth of groundwater discharge is calculated as:

$$188 \quad r_b = e^{\ln Te - \lambda - \bar{S}_g/Szm} \quad (4)$$

189 where, r_b is depth of groundwater discharge (m); $\ln Te$ is the log of the areal average of
 190 $T0$ (m^2/h); is the catchment average topographic index; \bar{S}_g is the catchment average
 191 groundwater table depth (m). For the complete set of equations for the Top-SSF model,
 192 the reader is referred to the Supplementary Material and (Li et al., 2024).

193 **3.2. Multi-machine learning ensemble method**

194 To improve flood prediction accuracy in ungauged mountainous catchments, we
 195 proposed a multi-machine learning ensemble method for regionalizing sensitive
 196 parameters of the Top-SSF model. This method leverages the complementary strengths
 197 of multi-machine learning methods to estimate model parameters based on catchment
 198 descriptors (Fig. 2). The characteristics, strengths, and limitations of each machine
 199 learning method are summarized in Table 2. The ensemble method employs a cross-
 200 validation procedure to select the best-performing machine learning method for each
 201 sensitive parameter. These selections are then integrated into a unified regionalization
 202 scheme. By mitigating limitations inherent in single machine learning regionalization,
 203 such as model bias and overfitting, and by capturing complex hydrological processes
 204 in mountainous catchment, this ensemble method aims to achieve more accurate flood
 205 prediction in ungauged catchments.



206
 207 **Fig.2.** Multi-machine learning ensemble method for regionalization in ungauged mountainous
 208 catchments. The red line indicates the machine learning method that yielded the optimal
 209 parameter estimates.
 210

211 **Table 2.** Seven machine learning model characteristics, advantages and disadvantages.

Machine learning	Characteristic	Advantage	Disadvantages
DT	A single decision tree hierarchically partitions the data space using a tree structure, with internal nodes representing features, branches representing decision rules, and leaf nodes representing class labels.	High interpretability; Minimal data preprocessing.	Unstable; Tends to overfit.
ERT	Construct multiple decision trees with randomly selected feature values and randomly divided nodes (Geurts et al., 2006).	Low overfitting risk; Computational efficiency; Resilient to noise.	Possibility of increased bias; Limited interpretability.
GBM	Construct multiple decision trees. Multiple weak learners are trained iteratively and the loss function is optimised using gradient descent, progressively combined into a robust model through the learning rate (Friedman, 2002).	High accuracy for structured data; Robust to outliers; Minimal data preprocessing.	Limited interpretability; Complex adjustments.
KNN	It is a non-parametric, instance-based supervised learning algorithm. It operates by finding the K nearest data points in the training data to a given data point and making predictions based on these (Wani et al., 2017).	Simple and easy to implement. Learning process is quick.	Sensitivity to noisy and scale of data. Accuracy can be heavily impacted by the choice of K.
RF	A bagging algorithm proposed by Breiman (2001) that uses ensemble learning. Involves training numerous decision trees and aggregating predictions.	Simple and easy to implement; Low computational cost.	Prone to overfitting in noisy regression tasks.
SVM	Identifies hyperplanes in high-dimensional spaces to segregate data. The optimal hyperplane maximizes the margin between it and the nearest data points, termed support vectors (Sain, 1996).	Uses kernel functions to address nonlinear classification issues.	Sensitive to noise

212 **3.3. Parameter regionalization process**

213 The parameter regionalization process comprised four key steps: (1) Top-SSF
 214 model calibration and parameter sensitivity analysis; (2) selection of relevant catchment
 215 descriptors; (3) establishment of regionalization relationships between sensitive model
 216 parameters and catchment descriptors using multi-machine learning ensemble methods;
 217 and (4) evaluation of parameter regionalization performance.

218 **3.3.1. Top-SSF model calibration and parameter sensitivity analysis**

219 In this study, the Top-SSF model was employed to simulate hydrological processes.
 220 The model was driven by continuous hourly meteorological data, including rainfall,
 221 temperature, surface pressure, relative humidity, wind speed, and surface net solar
 222 radiation. For each catchment, model parameters were calibrated using two
 223 hydrologically independent and representative flood events. A third, distinct flood

224 event was then used for model validation. The Nash-Sutcliffe Efficiency (NSE) served
225 as the objective function during calibration, with parameter optimization achieved
226 using the Shuffled Complex Evolution (SCE-UA) algorithm (Duan et al., 1994), known
227 for its global convergence and robustness (Dakhlaoui et al., 2012; Qi et al., 2016).
228 Model performance was evaluated using the NSE, the relative error of flood peak flow
229 (Qp), and the absolute error in flood peak occurrence time (Tp), following China's
230 Specification for Hydrological Information Forecast (GB/T 22482-2008). These
231 metrics quantify the model's ability to predict flood dynamics, peak flow, and timing.
232 Following calibration, a sensitivity analysis was conducted to identify and exclude
233 insensitive model parameters (Lenhart et al., 2002), which were then used for
234 regionalization. This approach reduces the dimensionality of the regionalization
235 problem and improves the efficiency of the process.

236 The sensitivity index (Si) of each hydrological model parameter was determined
237 using the method of Lenhart et al. (2002), which assesses the influence of $\pm 10\%$
238 changes in parameter values (Eq. 5). Table 3 outlines the sensitivity analysis results for
239 the model parameters across the 80 mountainous catchments. The Si values are
240 categorized as follows (Guo et al., 2022): negligible sensitivity ($|Si| < 0.05$),
241 moderate sensitivity ($0.05 < |Si| < 0.2$), high sensitivity ($0.2 < |Si| < 1.00$), and
242 extremely high sensitivity ($|Si| \geq 1.00$). Based on the sensitivity analyses, seven
243 sensitive model parameters were identified: Szm , $lnTe$, $Sfmax$, C , $qsf0$, t (Table 3).

$$Si = \frac{1}{N} \sum_t^N \frac{(y_2(t) - y_1(t))/y_0(t)}{2\Delta x/x_0} \quad (5)$$

244 where $y_0(t)$ is the flood value of the calibrated parameter x_0 at time t ; Δx is the

246 adjusted parameter difference, $\Delta x/x_0=10\%$; $y_1(t)$ is the flood value of the calibrated
 247 parameter $x_0 - \Delta x$ at time t ; $y_2(t)$ is the flood value of the calibrated parameter $x_0 +$
 248 Δx at time t .

249 **Table 3.** Top-SSF model main modules and default range of parameters.

Module	Parameter	Definition	Unit	Default range	Sensitivity index
Canopy interception	Sc	Canopy storage capacity	m	0.00~0.01	<0.05
	St	Trunk storage capacity	m	0.00~0.01	<0.05
	Pt	Proportion of rain diverted into stemflow per cover	%	0.00~1.00	<0.05
Evapotranspiration	$Sr0$	Initial root zone storage deficit	m	0.00~0.02	<0.05
	$Srmax$	Maximum root zone storage deficit	m	0.00~2	<0.05
Infiltration	Ks	Surface hydraulic conductivity	m/h	0~0.01	<0.05
	CD	Capillary drive (Morel-Seytoux et al., 1974)	m	0~5	<0.05
Unsaturated zone	$Suz0$	Initial baseflow per unit area	m	0.00~ 10^{-4}	<0.05
	Szm	Soil maximum water storage capacity	m	0.00~1.00	0.19
	td	Unsaturated zone time delay per unit storage deficit	h/m	0~3	1.07
	$lnTe$	log of the areal average of T0	m ² /h	-2.00~1.00	3.4
Subsurface storm flow zone	$Sfmax$	Maximum subsurface storm flow zone deficit	m	0.00~0.01	0.16
	C	Transfer coefficient	m ⁻² /h	0.00~0.1	0.26
	$qsf0$	Initial subsurface storm flow per unit area	m	0.00~0.02	0.18
Routing	t	Flow routing correction coefficient	-	0.00~5.0	1.21

250 Note, the bolded values in the sensitivity index indicate sensitive model parameters (i.e.,
 251 $|Si|>0.05$).

252 3.3.2. Catchment descriptor selection

253 To mitigate the effects of multicollinearity on the accuracy and reliability of the
 254 parameter regionalization methods, catchment descriptors were screened using the
 255 variance inflation factor (VIF) and correlation coefficients. A VIF threshold of less than
 256 10 (VIF < 10) was used to indicate acceptably low multicollinearity (Salmeron et al.,
 257 2018). Initial screening identified strong correlations between several descriptor pairs,

258 notably L with Ks_CH, and Tem with Elev. Furthermore, the VIF values for Ks_CH
 259 and Slope were found to exceed 10. Consequently, Ks_CH and Slope were removed
 260 from the potential set of descriptors. Following their removal, a re-evaluation of the
 261 VIF for the remaining descriptors was conducted. Although a notable correlation exists
 262 between Tem and elevation (Elev), their VIF values in the reduced set were both below
 263 the threshold of 10. Given the importance of Tem for representing climate impacts and
 264 Elev as a key topographic driver, both were retained to preserve potentially valuable
 265 information. The final set of seven catchment descriptors selected for regionalization
 266 therefore comprised FC, Elev, Area, L, Tem, Pre, and BD. As illustrated in Fig. 3b, the
 267 correlations among these final descriptors and the sensitive model parameters are
 268 generally low (highest at 0.5), suggesting that the relationships are complex and
 269 nonlinear.



270
 271 **Fig.3.** Analysis of catchment descriptor relationships: (a) Correlation coefficients and variance
 272 inflation factors (VIF) among all descriptors; (b) Correlation coefficients between
 273 sensitive model parameters and descriptors with VIF values below 10.

274 3.3.3. Parameter regionalization

275 To simulate ungauged catchment conditions, each of the 80 catchments was
 276 iteratively treated as an ungauged catchment, with the remaining 79 catchments serving

277 as donor catchments. A parameter regionalization method was then constructed using
278 the catchment descriptors and sensitive model parameters of the donor catchments to
279 predict the seven sensitive model parameters for the ungauged catchment based on its
280 catchment descriptors. These predicted model parameters were then input into the Top-
281 SSF model to enable flood prediction in ungauged catchments. To ensure robust and
282 generalizable results, K-fold cross-validation (K = 10) was implemented. This involved
283 randomly partitioning the 79 donor catchments into K subsets, using one subset as a
284 test set and the remaining K-1 subsets for method training in each iteration (Jung, 2018).
285 This approach maximizes data utilization and minimizes bias associated with specific
286 data partitioning. Hyperparameter tuning for each machine learning method was
287 performed using RandomizedSearchCV (Bergstra et al., 2012), with the objective of
288 minimizing the difference between predicted and observed parameter values.

289 **3.3.4. Evaluated metrics**

290 The performance of the parameter regionalization methods was evaluated by
291 considering two key aspects. First, the accuracy of the methods in estimating sensitive
292 model parameters was assessed using three metrics: root mean square error (RMSE),
293 standard deviation (STD), and the coefficient of determination (R^2). The R^2 was used
294 to quantify the agreement between estimated and calibrated parameter sets. Second, to
295 evaluate the impact of parameter regionalization on flood prediction. The resulting
296 flood predictions were then evaluated using the NSE, Qp, and Tp metrics.

$$297 NSE = 1 - \frac{\sum_{j=1}^M (Q_{obs}(j) - Q_{sim}(j))^2}{\sum_{j=1}^M (Q_{obs}(j) - \bar{Q}_{obs})^2} \quad (6)$$

$$298 Q_p = \left| \frac{Q_{obs,p} - Q_{sim,p}}{Q_{obs,p}} \times 100\% \right| \quad (7)$$

299 $T_p = |T_{obs,p} - T_{sim,p}|$ (8)

300 where $Q_{obs}(j)$ is the observed flow rate (m^3/s); $Q_{sim}(j)$ is the simulated flow rate
301 (m^3/s); \bar{Q}_{obs} is the mean value of the observed flow rate (m^3/s); $Q_{obs,p}$ is the observed
302 flood peak flow (m^3/s); $Q_{sim,p}$ is the simulated flood peak flow (m^3/s); $T_{obs,p}$ is the
303 observed flood peak occurrence time (h); and $T_{sim,p}$ is the simulated flood peak
304 occurrence time (h).

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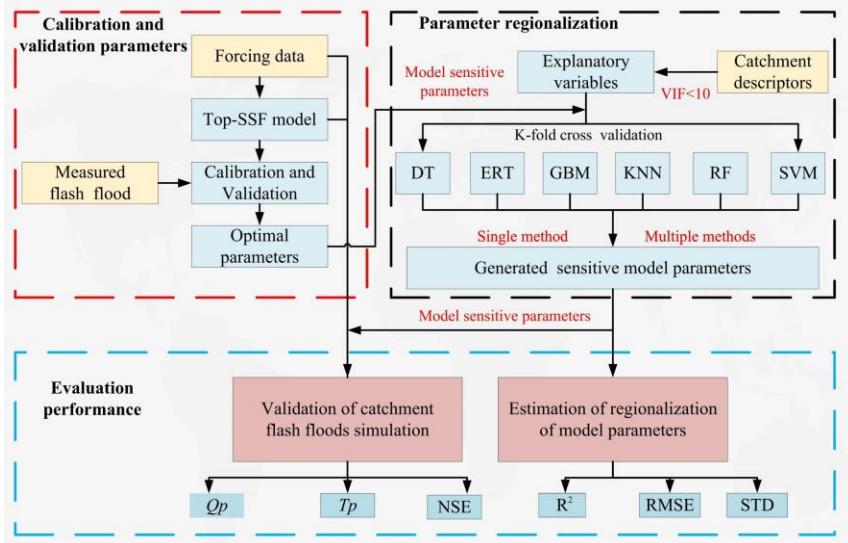
306 $RMSE = \sqrt{\frac{1}{N} \sum_{i=1}^N (X_i - Y_i)^2}$ (9)

307 $STD = \sqrt{\frac{1}{N-1} \sum_{i=1}^N (Y_i - \bar{Y})^2}$ (10)

308 $R^2 = \frac{[\sum_{i=1}^N (X_i - \bar{X})(Y_i - \bar{Y})]^2}{\sum_{i=1}^N (X_i - \bar{X})^2 \sum_{i=1}^N (Y_i - \bar{Y})^2}$ (11)

309 where X_i is the Top-SSF calibration model parameter value; Y_i is the model parameter
310 estimated value using the parameter regionalization method; \bar{X} and \bar{Y} are the mean
311 values of X_i and Y_i ; N is the sample size equal to 80.

312



313
314 **Fig.4.** Flowchart illustrating the parameter calibration, validation, and regionalization workflow.

315 Abbreviations: Top-SSF (Topography-Based Subsurface Storm Flow hydrological model),
316 DT (Decision Tree), ERT (Extremely Randomized Trees), GBM (Gradient Boosting
317 Machine), KNN (K-Nearest Neighbor), RF (Random Forest), SVM (Support Vector
318 Machine), NSE (Nash-Sutcliffe efficiency), R^2 (Coefficient of Determination), Qp (The
319 relative error of flood peak flow), Tp (The absolute error in flood peak occurrence time),
320 VIF (Variance inflation factor), RMSE (Root mean square error), STD (Standard
321 deviation).

322 **4. Result**

323 **4.1. Model performance**

324 The Top-SSF model demonstrated good flood simulation performance across the

325 80 gauged catchments, as quantified by NSE, Qp, and Tp. During the calibration period,

326 50% of the catchments achieved NSE values exceeding 0.78 (Fig. 5a), the median Qp

327 value was below 10% (Fig. 5b), and the median Tp value was within 2 hours (Fig. 5c).

328 The average NSE value was approximately 0.8, with a maximum of 0.96. The majority

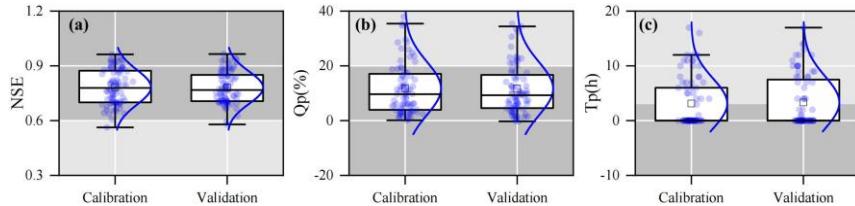
329 of Qp values were around 8%, and the majority of Tp values were below 2 hours.

330 During the validation period, the median NSE value was 0.76 (Fig. 5a), the median Qp

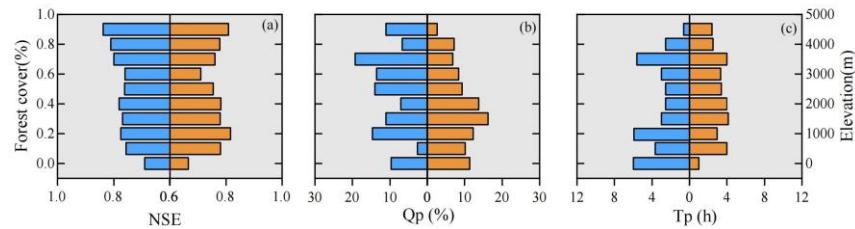
331 value was below 10% (Fig. 5b), and the median Tp value was within 4 hours (Fig.5c).
332 The hydrological response times for the 80 catchments were approximated as the time
333 from precipitation peak to flood peak. The estimated range is from 1 to 26 hours. This
334 diversity is indicative of the comprehensive nature of the study, which encompasses
335 both rapid flash floods in smaller basins and more general floods in larger, mountainous
336 catchments (mean area: 1,586 km²). For catchments with longer response times, a
337 median error of 2-4 hours remains operationally valuable for providing sufficient flood
338 warning lead time. It is noteworthy that the median Tp during the calibration period
339 (within 2 hours) satisfied China's Specification for Hydrological Information Forecast
340 (GB/T 22482-2008) stringent requirements for high-quality forecasts.

341 Model performance also exhibited some dependence on catchment characteristics.
342 For instance, NSE generally improved with increasing forest cover (Fig. 6a), potentially
343 due to the model's explicit representation of forest canopy interception and subsurface
344 storm flow generation mechanisms. The relationship between NSE, Qp, Tp and
345 elevation was more complex, suggesting a nonlinear influence of elevation on model
346 performance (Fig. 6 a-c). The demonstrated robust performance of the Top-SSF model
347 provides a strong foundation for its application in subsequent parameter regionalization
348 analyses.

349



350
351 **Fig. 5.** Boxplots of (a) NSE, (b) Qp, and (c) Tp during the calibration and validation periods
352 for 80 gauged catchments. The box represents the interquartile range, with the middle line
353 indicating the median (50th percentile). The whiskers represent the minimum and
354 maximum values. "□" represents the mean value. Dark grey indicates the range of flood
355 prediction criteria (i.e., NSE> 0.75, Qp<20%, and Tp < 2 hours).



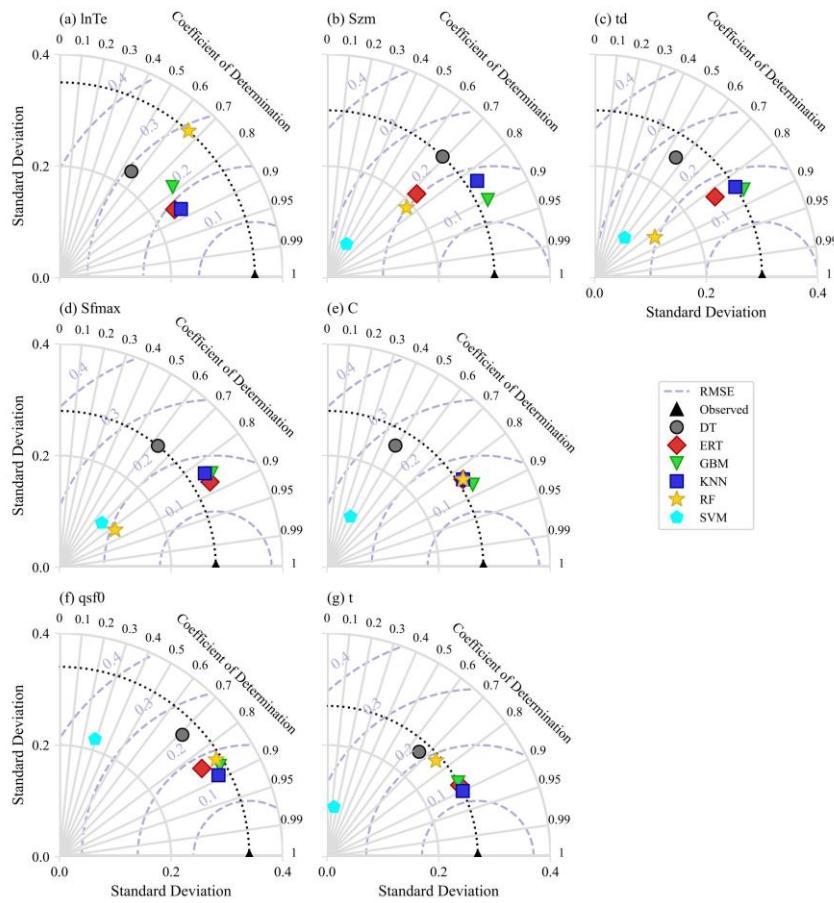
356
357 **Fig.6.** Influence of environmental factors on Top-SSF model performance in flood simulation.
358 The graphs illustrate the relationship between model evaluation metrics and forest cover
359 (left) and elevation (right).
360

361 **4.2. Results of parameter regionalization**

362 **4.2.1. Comparison of sensitive model parameter estimates**

363 The six single machine learning regionalization methods exhibited varying
364 performance in estimating sensitive model parameters (Fig. 7), likely due to differences
365 in catchment descriptor characteristics and the underlying principles of each method.
366 Their hyperparameter results are presented in Tables S1–S6 of the supplementary
367 material. The GBM demonstrated the highest accuracy in estimating Szm , td , and C
368 ($R^2 = 0.90$, 0.86, and 0.87, respectively,), with its estimates also exhibiting a STD that
369 closely matched the distribution of the calibrated parameter values. KNN provided the
370 most accurate estimates for $lnTe$, $qsf0$, and t ($R^2 = 0.87$, 0.89, and 0.90, respectively),
371 also with STD closely resembling the calibrated parameter distributions. ERT

372 performed best in estimating $Sfmax$ ($R^2 = 0.87$), but its performance was generally
 373 poorer for other parameters. DT, SVM, and RF methods generally showed lower
 374 performance across all sensitive model parameters. These differences in performance
 375 highlight the potential benefits of multi-machine learning ensemble methods for
 376 improving flood prediction in ungauged mountainous catchments.



377
 378 **Fig.7.** Performance of parameter regionalization methods assessed using Taylor diagrams. The
 379 diagrams show the accuracy of sensitive model parameter estimates, with the coefficient
 380 of determination (R^2) indicated by the radial axis, standard deviation (STD) by the
 381 horizontal and vertical axes, root mean square error (RMSE) by the grey-blue dotted lines,
 382 and the standard deviation of observations by the black dotted line."

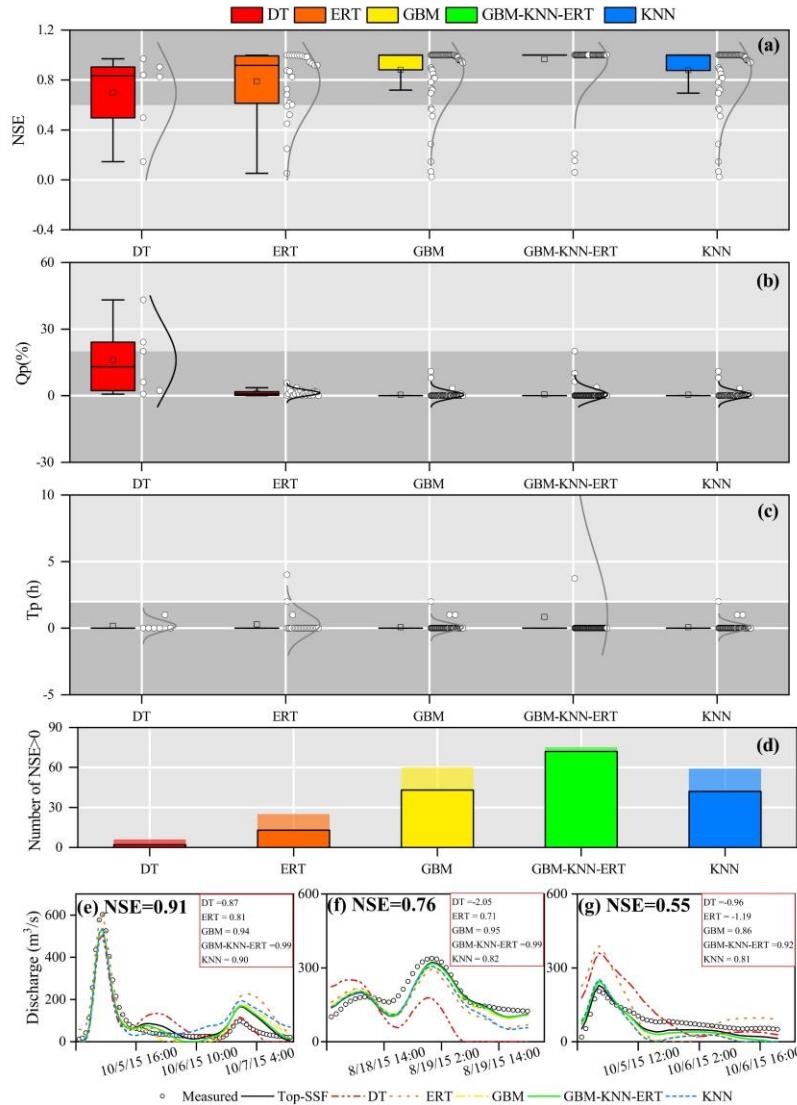
383 **4.2.2. Comparison of flood forecasting results**

384 The flood prediction performance of the Top-SSF model, integrated with different
385 parameter regionalization methods, was compared across 80 mountainous catchments
386 in southwestern China. The methods included single machine learning methods and a
387 multi-machine learning ensemble method (GBM-KNN-ERT), where GBM estimated
388 S_{zm} , td , and C ; KNN estimated $\ln Te$, $qsf0$, and t ; and ERT estimated $Sfmax$. The
389 performance of these parameter regionalization methods was then evaluated against the
390 performance of the Top-SSF model using calibrated parameters. Among the single
391 machine learning methods, GBM performed best, with 60 catchments achieving a
392 positive NSE ($NSE > 0$, Fig. 8d). Critically, for high-accuracy predictions ($NSE > 0.9$),
393 GBM succeeded in 43 catchments (54%), also showing strong performance with Q_p
394 less than 5% and T_p less than 1 hour in most cases (Fig. 8a-c). The GBM-KNN-ERT
395 ensemble method yielded even better results. It increased the number of catchments
396 with positive NSE to 75 (Fig. 8d). More impressively, the ensemble method achieved
397 exceptional performance ($NSE > 0.9$) in 72 catchments (90%). This represents a 67.44%
398 increase in the number of high-accuracy predictions compared to the best single method
399 (GBM). Furthermore, the ensemble method Q_p values were more concentrated around
400 zero, and 90% of catchments maintained near-zero T_p values. These results
401 demonstrate the superior potential of multi-machine learning ensembles for improving
402 flood prediction in ungauged catchments.

403 To further illustrate these performance differences visually, Fig. 8 (e, f, and g)
404 presents hydrographs from three randomly selected flood events. These events

405 represent cases where the calibrated Top-SSF model itself achieved high (NSE=0.91),
406 medium (NSE=0.76), and low (NSE=0.55) performance, respectively. A key insight
407 from these plots is that the Top-SSF simulation (solid black line) is the performance
408 benchmark for the regionalization methods. Although the models aim to approximate
409 measured floods, their performance is ultimately limited by the accuracy of the Top-
410 SSF model structure and its optimized parameters.

411 The hydrographs show how the GBM-KNN-ERT ensemble achieves superior
412 performance by leveraging the complementary strengths of its component methods. For
413 instance, in the high-performance case (Fig. 8e), the GBM and KNN methods capture
414 the overall shape well, but the ERT simulation provides a more precise estimation of
415 the primary flood peak. The final ensemble successfully integrates this peak accuracy,
416 resulting in the highest overall performance. Similarly, Fig. 8f shows that the ensemble
417 moderates the slow initial rise characteristic of the KNN method, leading to a more
418 realistic rising limb. The ensemble method ability to balance competing errors is most
419 evident in the low-performance case (Fig. 8g). During the recession phase, the ensemble
420 method averages the high bias of the ERT method with the low bias of the GBM and
421 KNN methods, producing a hydrograph that more closely resembles the benchmark
422 simulation than any single method could. This synergy demonstrates that the ensemble
423 method superior performance is a direct result of its ability to integrate the specific,
424 complementary strengths of each single method across different parts of the
425 hydrological process.



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Fig.8. Evaluation of flood prediction performance for different parameter regionalization methods. (a-c) show the distributions of Nash-Sutcliffe Efficiency (NSE), relative peak flow error (Qp), and peak time error (Tp) across all 80 catchments, with shaded regions indicating where flood prediction standards were met (NSE > 0.75, Qp < 20%, and Tp < 2 hours). (d) shows the number of catchments with NSE > 0 and the black border indicates the number of catchments with NSE > 0.9. (e-g) present example hydrographs comparing the simulated flood from each regionalization method against measured flood flow and the calibrated Top-SSF model benchmark for catchments where the benchmark model performance was (e) high (NSE=0.91), (f) medium (NSE=0.76), and (g) low (NSE=0.55).

436 **5. Discussion**

437 **5.1. Reliability of multi-machine learning ensemble in parameter regionalization**

438 In this study, the GBM-KNN-ERT method demonstrated superior regionalization
439 performance, highlighting the potential of ensemble methods for improving
440 hydrological predictions in ungauged mountainous catchments. The success of the
441 ensemble is rooted in the distinct learning mechanisms and behaviors of its individual
442 components, which were revealed during hyperparameter optimization.

443 The GBM method exhibited distinct parameter-specific sensitivities to
444 hyperparameters (Fig. 9a-c). For parameter C , the negative correlation between R^2 and
445 n_estimators (>300 trees) indicates overfitting risks when modeling complex rainfall-
446 runoff interactions in heterogeneous mountainous terrain (Fig. 9a). This aligns with
447 previous findings emphasizing the need for complexity control in hydrological
448 generalization (Schoups et al., 2008). Conversely, the improved R^2 for parameter td
449 with increased n_estimators highlights the capacity of ensemble learning to capture
450 complex, nonlinear relationships between catchment descriptors and hydrological
451 parameters (Hastie et al., 2009). The contrasting optimal max_depth of 10 layers for
452 parameter C , compared to shallower optimal depths (4 layers) for Szm and td , suggests
453 that parameters governing more complex hydrological processes in mountainous
454 catchments may require deeper decision trees to effectively capture the interactions
455 between climate, topography, and soil properties (Wainwright et al., 2013).

456 KNN performance exhibited pronounced sensitivity to neighbourhood size
457 (n_neighbors) and distance metric (p), highlighting the spatial heterogeneity of

458 catchment descriptors. For parameters *lnTe* and *qsf0*, optimal performance was
459 observed at *n_neighbors* =30 (Fig. 9d), aligns with the hypothesis that meaningful
460 hydrological similarities can emerge even in topographically complex mountainous
461 regions when considered at broader spatial scales (Li et al., 2022). Conversely,
462 parameter *t* achieved peak accuracy at *n_neighbors*=5, suggesting that localized, short-
463 term weather events and fine-scale topographic similarities in adjacent mountainous
464 areas can significantly influence local runoff processes (Garambois et al., 2015). The
465 Manhattan distance metric (*p*=1) outperformed Euclidean distance across all
466 parameters (Fig. 9e). This superiority stems from its ability to mitigate the curse of
467 dimensionality (Bellman, 1961) in high-dimensional datasets, a common characteristic
468 of mountainous catchments. In such datasets, sparse data distributions and the presence
469 of mixed variable types (e.g., topographic indices, land cover) can significantly degrade
470 the discriminative power of Euclidean distance (Rockström et al., 2023). The
471 robustness of the Manhattan distance arises from its axis-aligned sensitivity, which
472 provides a more effective means of handling feature scaling and integrating catchment
473 descriptors compared to the radial symmetry of Euclidean distance.

474 ERT performance was maximized at *max_features* = 0.1 (Fig. 9f). By restricting
475 the random sampling of features during node splits (using only 10% of the features),
476 both the diversity of the trees was enhanced and the effects of multicollinearity between
477 topographic and soil attributes were reduced. This finding aligns with the theory
478 proposed by Geurts et al. (2006), which suggests that random feature selection can
479 significantly improve model generalization, a particularly important consideration in

480 ungauged mountainous catchments characterized by high levels of inter-correlation
481 among predictor variables.

482 These distinct sensitivities and learning mechanisms form the scientific basis for
483 the superiority of the GBM-KNN-ERT method. As shown in Section 4.2, no single
484 machine learning method is universally optimal for all hydrological model parameters.
485 Instead, the ensemble method effectively allocates each parameter to the model best
486 suited for its regionalization. Specifically, GBM, with its capacity for modeling
487 complex interactions, proved optimal for integrated parameters like *Szm* and *td*. In
488 contrast, the instance-based KNN was superior for parameters like *lnTe*, which are
489 governed by physical similarity and spatial coherence. Finally, the highly randomized
490 nature of ERT provided the necessary robustness to model the noisy relationship
491 associated with the *Sfmax*. This synergistic combination, where each model
492 contributes its unique strength, results in a final regionalization method that is more
493 accurate and physically plausible than any individual method operating in isolation.

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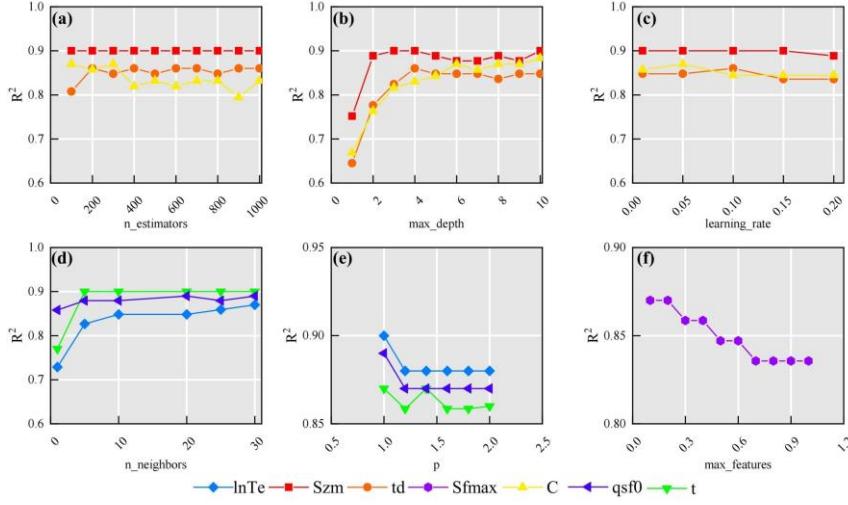


Fig.9. Sensitivity of parameter estimation performance to key hyperparameters in (a-c) GBM, (d-e) KNN method, and (f) ERT. (a) $n_estimators$ (number of decision trees in GBM), (b) max_depth (maximum depth of decision trees in GBM), (c) learning rate (GBM), (d) $n_neighbors$ (number of neighbors in KNN), (e) p -value of Minkowski distance (KNN; $p=1$: Manhattan distance, $p=2$: Euclidean distance), and (f) $max_features$ (ERT).

5.2. Combining multiple machine learning methods for parameter regionalization

Machine learning methods exhibit distinct strengths in hydrological parameter estimation due to fundamental differences in data processing mechanisms, pattern recognition strategies, and prediction generation (Bishop et al., 2006). This suggests that multi-machine learning ensemble methods have the potential to synergistically integrate advantages while effectively compensating for individual limitations, leading to more robust and accurate parameter estimates. As demonstrated in Fig. 10, the GBM-KNN-ERT method achieved notable improvements over any single machine learning method, particularly for sensitive parameters $lnTe$, $Sfmax$, $qsf0$ and t , with R^2 increases ranging from 0.02 to 0.03 compared to the best-performing GBM method (Fig.10e).

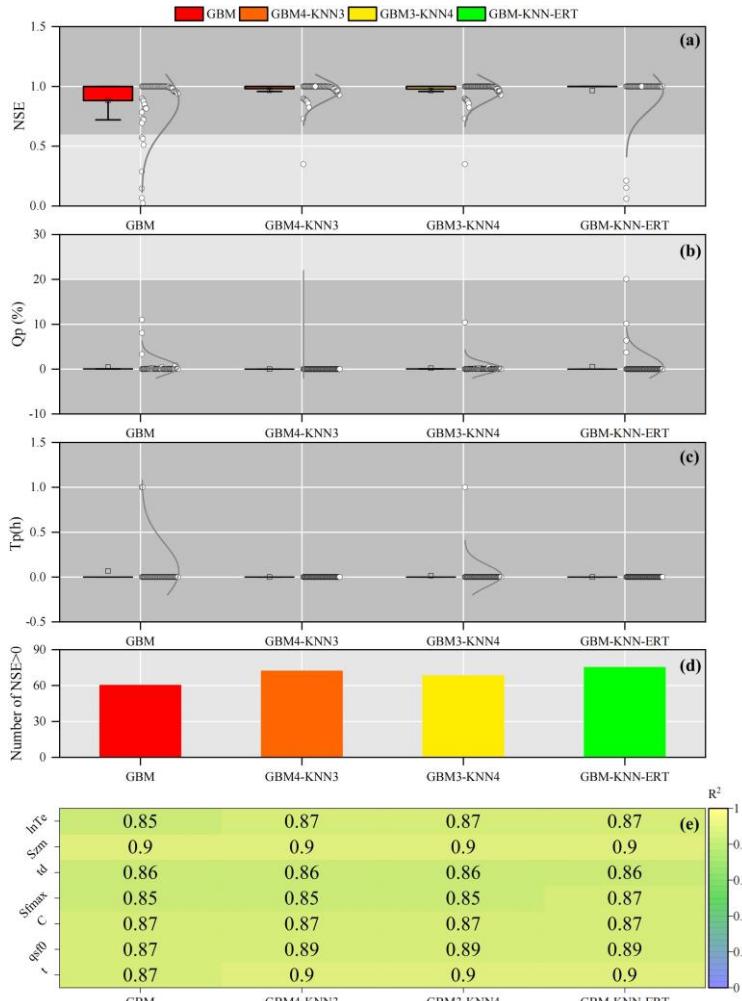
513 Interestingly, a comparison of GBM4-KNN3 (where $Sfmax$ is estimated by GBM)
514 and GBM3-KNN4 (where $Sfmax$ is estimated by KNN) revealed critical insights into
515 model parameter compatibility. Despite both achieving an identical R^2 of 0.85 for the
516 estimation of $Sfmax$, GBM4-KNN3 exhibited superior flood prediction performance,
517 with 72 catchments achieving $NSE > 0$ compared to only 68 catchments for GBM3-
518 KNN4. This suggests that GBM possesses an enhanced capability to resolve the
519 complex coupling between soil moisture dynamics and topography, leading to more
520 physically plausible representation of subsurface storm flow processes (Gupta et al.,
521 2023). The wider distribution of flood prediction performance observed for GBM3-
522 KNN4 (Fig. 10 a–c) further suggests that uncertainties introduced by KNN in the
523 estimation of $Sfmax$ may propagate nonlinearly during flood simulations, potentially
524 amplifying errors. This observation aligns with theoretical expectations that distance-
525 based methods may tend to oversmooth critical thresholds or sharp transitions in
526 heterogeneous environments, leading to a less accurate representation of hydrological
527 responses (Bellman, 1961).

528 Furthermore, an important consideration in adopting ensemble methods is the
529 trade-off between predictive accuracy and computational efficiency. To evaluate this
530 trade-off, the model training times for various parameter regionalization methods were
531 compared, and the results are summarized in Table 4. The analysis shows that the
532 proposed GBM-KNN-ERT ensemble, while providing the highest predictive accuracy,
533 required a total training time of 102.8 s. This is moderately higher than the best-
534 performing single model, GBM (57.6 s), and other simpler ensemble methods like

535 GBM4-KNN3 (36.1 s). The increased computational time for the GBM-KNN-ERT
536 method is primarily attributed to the inclusion of the ERT method for estimating the
537 S_{fmax} , which is inherently more computationally intensive than GBM or KNN.

538 However, it is crucial to contextualize this computational cost for operational use.
539 The process of training a regionalization method is an offline task, performed once to
540 establish the stable relationships between catchment descriptors and model parameters.
541 This one-time investment is not a constraint on real-time flood forecasting, as once the
542 method is trained, parameter estimation for a new ungauged catchment is nearly
543 instantaneous. For the reported computational times, all model training and simulations
544 were performed on a workstation equipped with an Intel(R) Core (TM) i9-10900K CPU
545 @ 3.70GHz, 32.0 GB of RAM, and an NVIDIA Quadro P1000 (4 GB) GPU, running
546 on a 64-bit Windows operating system with Python 3.9. Given this context, the modest
547 increase in one-time training cost is a justifiable investment for the significant
548 improvements achieved in flood prediction accuracy, model robustness, and stability.
549 Therefore, for applications in water resource management and flood risk assessment
550 where high accuracy is paramount, the GBM-KNN-ERT method strikes an optimal and
551 practical balance between computational efficiency and predictive performance.

批注 [L1]: Comment#1



552

553 **Fig.10.** Assessment of combined machine learning methods for improved parameter
554 regionalization in ungauged mountainous catchments. Performance is evaluated against
555 the GBM method, showing (a) NSE, (b) Qp, (c) Tp, (d) Number of catchments with NSE >
556 0, and (e) the difference in R^2 .

557 **Table 4.** Running time (s) for different parameter regionalization methods

	GBM	GBM4-KNN3	GBM3-KNN4	GBM-KNN-ERT	KNN	ERT
lnTe	11.3	3.4	3.4	3.7	3.6	74.4
Szm	7.8	7.5	7.7	7.8	0.6	76.7
td	8.2	8.1	8.0	8.5	0.6	74.7
Sfmax	7.7	8.2	0.6	73.6	0.5	74.9
C	7.8	7.7	7.7	8.0	0.6	74.9
qsfo	7.4	0.6	0.6	0.6	0.6	76.3
t	7.4	0.6	0.6	0.6	0.5	75.3
Sum	57.6	36.1	28.6	102.8	7.0	527.2

558 **5.3. The influence of donor catchment quantity on machine-learning parameter
559 regionalization**

560 The number of donor catchments used in machine learning-based parameter
561 regionalization methods is a critical factor influencing the accuracy and robustness of
562 hydrological predictions in ungauged catchments (Gauch et al., 2021; Song et al., 2022;
563 Zhang et al., 2022). This study investigated the influence of donor catchment quantity
564 (ranging from 20 to 80) on the flood prediction performance of the two best-performing
565 parameter regionalization methods (GBM4-KNN3 and GBM-KNN-ERT) across the 80
566 mountainous catchments (Fig.11). It is important to clarify that the following analysis
567 is not a method for selecting donor catchments based on physical similarity—a task
568 handled by the machine learning methods itself when it learns the relationships between
569 catchment descriptors and model parameters. Instead, this experiment serves as a
570 sensitivity analysis to understand how the regionalization performance is affected by
571 the overall quantity and quality of the available training data.

572 To systematically investigate the performance influence of donor catchment
573 quantity on parameter regionalization, two distinct sampling strategies were employed
574 across the 80 mountainous catchments. In Mode 1 (selection of donor catchments based
575 on decreasing NSE), which was designed to test the impact of data quality, a non-
576 monotonic relationship was observed. For both methods, regionalization performance
577 peaked with 20-40 donor catchments and then declined, particularly for the GBM4-
578 KNN3 method (Fig. 11a-c). This performance degradation is not due to increasing
579 catchment dissimilarity, but rather to the introduction of lower-quality training data. As
580 the donor pool expands beyond the best-performing catchments, it begins to include

581 catchments where the Top-SSF model calibration itself was less successful (i.e., lower
582 NSE values). These lower-performance samples may introduce noise and less reliable
583 parameter-descriptor relationships, which can mislead the training process (Gauch et
584 al., 2021; Zhang et al., 2022). Notably, the GBM-KNN-ERT method demonstrated
585 greater resilience to this degradation. Its performance, while also peaking early, did not
586 degrade as sharply and instead tended to stabilize after the inclusion of approximately
587 70 catchments. This suggests that the more complex ensemble structure has a superior
588 ability to suppress noise and generalize from a dataset containing a mix of high- and
589 low-quality examples, highlighting its enhanced robustness. In contrast, Mode 2
590 (random selection of donor catchments) demonstrated a consistent improvement in
591 regionalization performance for both NSE and Tp as the number of donor catchments
592 increased (Fig. 11d-f). However, while the average performance improves with data
593 quantity, it is important to acknowledge that this trend relies on the random samples
594 being generally representative; a poorly chosen random set could still reduce
595 generalizability. Notably, under both modes, the GBM-KNN-ERT method consistently
596 exhibited significantly greater performance stability compared to the alternative
597 ensemble, GBM4-KNN3. This enhanced robustness likely arises from its more
598 effective suppression of data heterogeneity and noise interference, indicating that more
599 complex ensemble methods possess a greater capacity to balance the benefits of
600 increased data quantity with the potential drawbacks of reduced data quality.

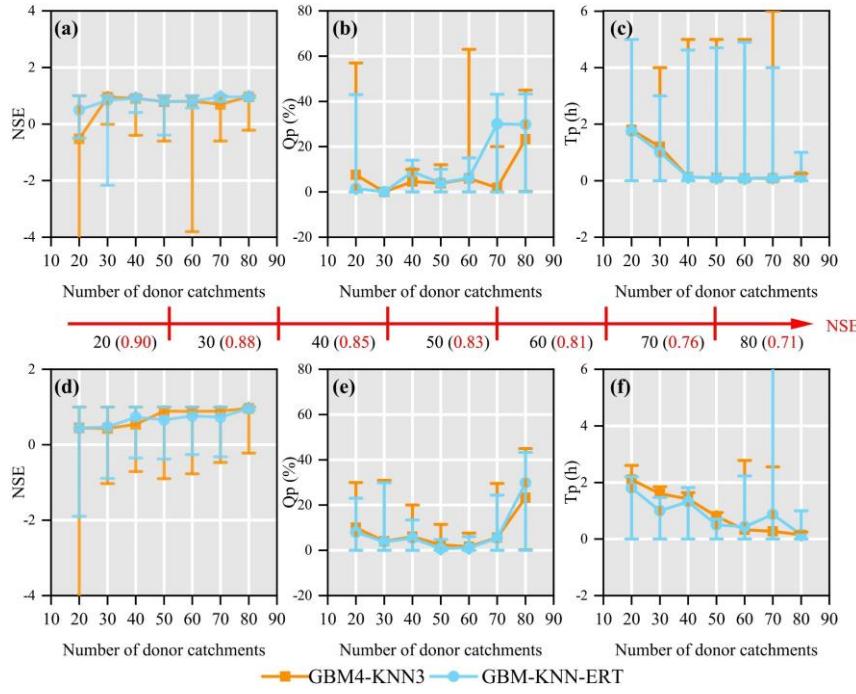


Fig. 11. Performance comparison of two donor catchment selection methods for parameter regionalization as a function of donor catchment quantity. Mode1 (a-c) selects donor catchments in order of decreasing NSE, while Mode 2 (d-f) selects them randomly. Flood prediction accuracy is assessed using NSE, Qp, and Tp. Error bars represent the full range (minimum to maximum) of the performance metrics.

5.4. The impact of climate change on parameter regionalization methods

The hydrological cycle within catchments is fundamentally governed by complex interactions between climate and environmental factors. The Intergovernmental Panel on Climate Change (IPCC) has consistently documented a continuous and accelerating transition in global climatic patterns, characterized by increased variability and extreme events (Pachauri et al., 2014). Consequently, future flood predictions derived from parameter regionalization methods are expected to exhibit increased uncertainty and variability, highlighting the substantial influence of climate change on the reliability and precision of flood predictions in ungauged mountainous catchments (Yang et al.,

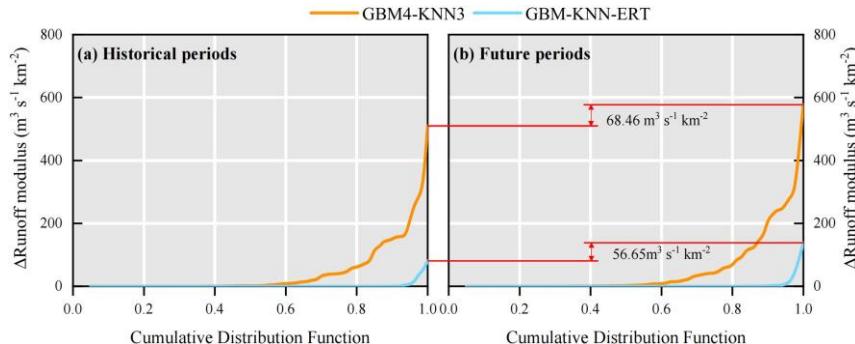
616 2019). Therefore, a sensitivity analysis was designed to evaluate the robustness of the
617 trained regionalization models when confronted with climatic conditions outside their
618 original training range.

619 To quantitatively assess the impact of climate change, an experiment was devised
620 where this impact was primarily reflected through changes in two key catchment
621 descriptors: Tem and Pre. For the historical period, these descriptors represent the multi-
622 year averages over 1901–2021, while for the future period, they represent the projected
623 multi-year averages over 2022–2100 under the SSP5-8.5 scenario. The regionalization
624 methods (GBM4-KNN3 and GBM-KNN-ERT), which were trained exclusively using
625 historical data, were then applied under these future conditions. Crucially, the method
626 structures and hyperparameters remained fixed, and no retraining was performed; only
627 the historical Tem and Pre values were replaced with their future projections. This
628 approach allows the response of the established historical relationships to new, out-of-
629 sample climatic inputs to be tested. The simulated peak discharges for this analysis were
630 derived from the same three flood events used in the calibration and validation of the
631 Top-SSF model. This experimental design is critical as it isolates the impact of the
632 changed model parameters from the compounding effect of a different future rainfall
633 event. Consequently, any observed change in the simulated flood peak is attributable
634 solely to the sensitivity of the regionalization method to the shift in climatic descriptors.
635 Cumulative distribution functions (CDFs) were then employed to illustrate the
636 discrepancies between the parameter regionalization simulations and the reference
637 simulations (derived from calibrated model parameters) across the historical and

638 projected future periods for the 80 catchments (Fig.12).

639 A comparative analysis of Fig. 12a and 12b reveals a clear amplification of the
640 absolute differences in predicted flood peaks (quantified as the error in runoff modulus)
641 between the two parameter regionalization methods and the reference Top-SSF model
642 simulations during the transition from the historical period to the projected future period.
643 Specifically, the maximum error in runoff modulus for the GBM4-KNN3 method
644 increased by $68.46 \text{ m}^3 \text{ s}^{-1} \text{ km}^{-2}$ from the historical period to the future period, while the
645 increase for the GBM-KNN-ERT method was a smaller $56.65 \text{ m}^3 \text{ s}^{-1} \text{ km}^{-2}$. These results
646 underscore that parameter regionalization methods are inherently sensitive to changing
647 climatic forcing. However, they also provide compelling evidence that the GBM-KNN-
648 ERT method exhibits superior stability and resilience under climate change,
649 demonstrating its potential for more reliable long-term flood risk assessment in
650 ungauged mountainous regions.

651 Exploring the effects of climate change on parameter regionalization methods
652 provides valuable insights for advancing flood prediction research in prediction in
653 ungauged basins. The enhanced stability demonstrated by the GBM-KNN-ERT
654 ensemble offers a promising direction for developing robust regionalization methods
655 capable of navigating the challenges of a non-stationary climate.



656
657 **Fig.12.** Comparison of flood peak runoff modulus between parameter regionalization and
658 calibrated Top-SSF model results, showing cumulative distribution functions (CDFs) of
659 absolute differences for 80 catchments during (a) historical and (b) future periods.

660 5.5. Uncertainty and limitation

661 The uncertainty in this study arises from several sources, including the
662 hydrological model, the regionalization methods, and the data itself. A critical
663 evaluation of these sources helps to contextualize our findings and assess the
664 generalizability of the ensemble method. Uncertainty from the hydrological model is
665 inherent in its structure and the calibrated parameters. Although the Top-SSF model
666 performed well, its parameters are effective values subject to equifinality. This
667 uncertainty in the true parameter values can be viewed as a form of calibration bias,
668 which serves as the target data for our regionalization. To mitigate this, we employed
669 the robust SCE-UA optimization algorithm and focused only on sensitive parameters.
670 Uncertainty is also introduced by the regionalization methods themselves, as the
671 training data derived from donor catchments are susceptible to errors that can impact
672 model performance (Mosavi et al., 2018; Xu et al., 2021).

673 A specific methodological choice was the exclusion of deep learning architectures,
674 such as Multilayer Perceptrons or Long Short-Term Memory (LSTM) networks. This

675 decision was guided by several factors. First, parameter regionalization is a static
676 regression problem, mapping time-invariant catchment descriptors to model parameters,
677 which does not align with the sequential data structure for which LSTM is designed.
678 Second, deep networks typically require large datasets to avoid overfitting; with a
679 dataset of 80 catchments, traditional machine learning methods like GBM and ERT are
680 often more robust and less prone to memorizing training data. **Third, a key advantage**
681 **of parameter regionalization is its potential for physical interpretability. Unlike DL**
682 **models, whose internal decision-making processes are often obscured within abstract**
683 **weight matrices, the ensemble methods employed here offer more accessible**
684 **transparency. The tree-based models (GBM and ERT) allow for the direct assessment**
685 **of feature importance, enabling the verification of physical consistency. Furthermore,**
686 **the KNN component provides instance-based interpretability by explicitly identifying**
687 **the specific donor catchments used for transfer. This preserves the traceable logic of**
688 **hydrological similarity, clearly indicating the geographical or physical source of the**
689 **transferred parameters, which may be crucial for building the trust of the method in the**
690 **water management of mountainous catchments.**

批注 [L2]: Comment#2

691 Furthermore, the primary contribution of this study is not the identification of a
692 single superior algorithm, but the demonstration of a data-driven framework for
693 constructing a locally optimal ensemble. The complementarity of the chosen models
694 was not assumed but empirically validated through a competitive evaluation process.
695 Each of the seven machine learning methods was independently trained and assessed
696 for its ability to estimate each sensitive parameter. The final GBM-KNN-ERT ensemble

697 was constructed by selecting only the empirically best-performing model for each
698 parameter based on objective metrics (R^2 , RMSE, STD). The very fact that different
699 methods were selected for different hydrological parameters provides direct empirical
700 evidence of their complementary strengths, thus validating the ensemble method.

701 Furthermore, the specific GBM-KNN-ERT ensemble identified is necessarily
702 data-dependent, raising questions about its transferability. However, this study primary
703 contribution is not the specific model combination itself, but rather the demonstration
704 of a data-driven method for constructing a locally optimal ensemble. This method is
705 designed to be generalizable; applying the same competitive evaluation process to a
706 new region would identify the best ensemble for that specific dataset. The key to
707 overcoming these limitations and ensuring robust generalization lies in genuine model
708 complementarity. The ensemble method's success is not an artifact of overfitting to
709 calibration bias or data quirks. Instead, it stems from a synergistic integration, where
710 different models are empirically shown to be better suited for regionalizing parameters
711 governed by distinct physical processes. The ensemble method's superior stability in
712 the out-of-sample climate change stress test further supports this conclusion, indicating
713 that it has captured robust underlying relationships, not just noise.

714 To manage methodological uncertainty, K-fold cross-validation was employed to
715 ensure robust performance evaluation, and RandomizedSearchCV was used for
716 hyperparameter tuning to minimize overfitting (Bergstra and Bengio, 2012). A key
717 methodological decision was to evaluate the regionalization methods against the
718 outputs of the calibrated Top-SSF model, rather than directly against observed flood

719 events. This approach was chosen for two primary reasons. First, it isolates the
720 performance of the parameter regionalization itself. The calibrated simulation
721 represents the theoretical upper bound of performance for the given hydrological model
722 structure; consequently, any deviation from this benchmark can be directly attributed
723 to imperfections in the regionalization method, rather than being confounded by the
724 inherent structural limitations of the Top-SSF model. Second, this strategy ensures that
725 the machine learning models learn the underlying physical relationships intended by
726 the hydrological model, not simply mimic data noise or measurement errors present in
727 the observations. If trained against raw observations, the machine learning methods
728 might derive spurious parameter sets that compensate for both the hydrological model's
729 structural flaws and observational errors. Such parameters could appear effective but
730 would lack physical meaning and generalizability. These measures, combined with the
731 evidence for model complementarity, provide a strong basis for the scientific validity
732 and potential for generalization of our proposed ensemble method.

733 **6. Conclusions**

734 This study introduces a novel multi-machine learning ensemble method (GBM-
735 KNN-ERT) to enhance model parameter transferability and improve flood prediction
736 in ungauged mountainous catchments. The proposed GBM-KNN-ERT method
737 demonstrated a substantial advancement in both flood prediction accuracy and model
738 robustness, achieving exceptional performance with 90% of ungauged catchments
739 exhibiting a NSE exceeding 0.9, a significant 67.44% improvement compared to the
740 best single machine learning method evaluated in this study. Importantly, the GBM-

741 KNN-ERT method exhibited remarkable stability under simulated climate change,
742 thereby highlighting its potential for reliable application in non-stationary hydrological
743 environments. Furthermore, the method demonstrated notable adaptability to varying
744 donor-catchment configurations, where an optimal balance between predictive
745 accuracy and computational efficiency with a relatively limited set of 20–40 high-
746 quality donor catchments ($\text{NSE} > 0.85$). By integrating the diverse strengths of multiple
747 machine learning with hydrological model, the proposed methodology significantly
748 advances the field of flood prediction in ungauged catchments, offering a reliable tool
749 for water resource management and flood disaster mitigation.

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755 **Competing interests**

756 The authors declare that they have no known competing financial interests or
757 personal relationships that could have appeared to influence the work reported in this
758 paper.

759 **Author contributions**

760 In this study, K L, G W, and J G were responsible for the conceptualization of the
761 research. Data curation was carried out by K L, L G, and X S, while formal analysis
762 was performed by K L, J G, and J M. The methodology was developed by K L, L G, P

763 H, and J L. Project administration was overseen by G W and J G. K L took the lead in
764 writing the original draft, and the writing, review, and editing process involved
765 contributions from K L, G W, J L, P H, J M, X Z, and J G.

766 **Code and data availability**

767 The code used in this study is available upon request from the authors. The
768 meteorological, soil characteristics, and topography datasets are publicly accessible
769 online, as detailed in Table 1. The hourly flood data for the 80 catchments were sourced
770 from China's Hydrological Yearbook. These data are not publicly available due to
771 governmental restrictions but can be accessed by contacting the corresponding author
772 for further information.

773 **References**

774 Arsenault, R., Breton-Dufour, M., Poulin, A., Dallaire, G. Romero-Lopez, R. (2019).
775 Streamflow prediction in ungauged basins: analysis of regionalization methods
776 in a hydrologically heterogeneous region of Mexico. *Hydrological Sciences
777 Journal*, 64(11): 1297-1311. <https://doi.org/10.1080/02626667.2019.1639716>

778 Arsenault, R., Martel, J., Mai, J. (2022). Continuous streamflow prediction in ungauged
779 basins: Long Short-Term Memory Neural Networks clearly outperform
780 hydrological models. *Hydrol. Earth Syst. Sci.* 1-29.
781 <https://doi.org/10.5194/hess-27-139-2023>

782 Bellman, R.E. (1961). On the reduction of dimensionality for classes of dynamic
783 programming processes. RAND Corp., Santa Monica, Calif., Paper P-2243.

784 Bergstra, J. Bengio, Y. (2012). Random search for hyper-parameter optimization.
785 *Journal of machine learning research*, 13(2).

786 Beven, K.J., Kirkby, M.J., Freer, J.E., Lamb, R. (2021). A history of TOPMODEL.
787 *Hydrology and Earth System Sciences*, 25(2): 527-549.
788 <https://doi.org/10.5194/hess-25-527-2021s>

789 Bishop, C.M. Nasrabadi, N.M., (2006). Pattern recognition and machine learning
790 (information science and statistics). New York: Springer - Verlag.

791 Breiman, L. (2001). Random forests. *Machine learning*, 45: 5-32.

792 Cheng, Q., Gao, L., Zuo, X. Zhong, F. (2019). Statistical analyses of spatial and
793 temporal variabilities in total, daytime, and nighttime precipitation indices and
794 of extreme dry/wet association with large-scale circulations of Southwest China,

795 1961–2016. Atmospheric research, 219: 166-182.
796 <https://doi.org/10.1109/ACCESS.2018.2886549>

797 Choi, J., Kim, U.Kim, S. (2023). Ecohydrologic model with satellite-based data for
798 predicting streamflow in ungauged basins. *Science of The Total Environment*,
799 903: 166617. <https://doi.org/10.1016/j.scitotenv.2023.166617>

800 Dai, Y., Shangguan, W., Duan, Q., Liu, B., Fu, S.Niu, G. (2013). Development of a
801 China dataset of soil hydraulic parameters using pedotransfer functions for land
802 surface modeling. *Journal of Hydrometeorology*, 14(3): 869-887.
803 <https://doi.org/10.1175/JHM-D-12-0149.1>

804 Dakhlaoui, H., Bargaoui, Z.Bárdossy, A. (2012). Toward a more efficient calibration
805 schema for HBV rainfall–runoff model. *Journal of Hydrology*, 444: 161-179.
806 <https://doi.org/10.1016/j.jhydrol.2012.04.015>

807 Ding, Y.Peng, S. (2020). Spatiotemporal trends and attribution of drought across China
808 from 1901–2100. *Sustainability*, 12(2): 477.
809 <https://doi.org/10.3390/su12020477>

810 Duan, Q., Sorooshian, S.Gupta, V.K. (1994). Optimal use of the SCE-UA global
811 optimization method for calibrating watershed models. *Journal of Hydrology*,
812 158(3): 265–284. [https://doi.org/10.1016/0022-1694\(94\)90057-4](https://doi.org/10.1016/0022-1694(94)90057-4)

813 Friedman, J.H. (2002). Stochastic gradient boosting. *Computational statistics & data
814 analysis*, 38(4): 367-378. [https://doi.org/10.1016/S0167-9473\(01\)00065-2](https://doi.org/10.1016/S0167-9473(01)00065-2)

815 Gan, B., Liu, X., Yang, X., Wang, X.Zhou, J. (2018). The impact of human activities on
816 the occurrence of mountain flood hazards: lessons from the 17 August 2015
817 flash flood/debris flow event in Xuyong County, south-western China.
818 *Geomatics, Natural Hazards and Risk*, 9(1): 816-840.
819 <https://doi.org/10.1080/19475705.2018.1480539>

820 Gao, J., Kirkby, M.Holden, J. (2018). The effect of interactions between rainfall
821 patterns and land-cover change on flood peaks in upland peatlands. *Journal of
822 Hydrology*, 567: 546-559. <https://doi.org/10.1016/j.jhydrol.2018.10.039>

823 Garambois, P.A., Roux, H., Larnier, K., Labat, D.Dartus, D. (2015). Parameter
824 regionalization for a process-oriented distributed model dedicated to flash
825 floods. *Journal of Hydrology*, 525: 383-399.
826 <https://doi.org/10.1016/j.jhydrol.2015.03.052>

827 Gauch, M., Mai, J.Lin, J. (2021). The proper care and feeding of CAMELS: How
828 limited training data affects streamflow prediction. *Environmental Modelling &
829 Software*, 135: 104926. <https://doi.org/10.1016/j.envsoft.2020.104926>

830 Geurts, P., Ernst, D.Wehenkel, L. (2006). Extremely randomized trees. *Machine
831 Learning*, 63(1): 3-42. <https://doi.org/10.1007/s10994-006-6226-1>

832 Golian, S., Murphy, C.Meresa, H. (2021). Regionalization of hydrological models for
833 flow estimation in ungauged catchments in Ireland. *Journal of Hydrology: Regional
834 Studies*, 36: 100859. <https://doi.org/10.1016/j.ejrh.2021.100859>

835 Guo, L., Huang, K., Wang, G.Lin, S. (2022). Development and evaluation of
836 temperature-induced variable source area runoff generation model. *Journal of
837 Hydrology*, 610: 127894. <https://doi.org/10.1016/j.jhydrol.2022.127894>

838 Guo, Y., Zhang, Y., Zhang, L. Wang, Z. (2021). Regionalization of hydrological
839 modeling for predicting streamflow in ungauged catchments: A comprehensive
840 review. Wiley Interdisciplinary Reviews: Water, 8(1): e1487.
841 <https://doi.org/10.1002/wat2.1487>

842 Gupta, A.K., Chakraborty, S., Ghosh, S.K. Ganguly, S. (2023). A machine learning
843 model for multi-class classification of quenched and partitioned steel
844 microstructure type by the k-nearest neighbor algorithm. Computational
845 Materials Science, 228: 112321.
846 <https://doi.org/10.1016/j.commatsci.2023.112321>

847 Hastie, T., Tibshirani, R. Friedman, J., (2009). The elements of statistical learning.
848 Citeseer.

849 Hersbach, H., Bell, B., Berrisford, P., Biavati, G., Horányi, A., Muñoz Sabater, J.,
850 Nicolas, J., Peubey, C., Radu, R., Rozum, I., Schepers, D., Simmons, A., Soci,
851 C., Dee, D. Thépaut, J.-N., (2023). ERA5 hourly data on single levels from 1940
852 to present, Copernicus Climate Change Service (C3S) Climate Data Store
853 (CDS)[Dataset]. <https://doi.org/10.24381/cds.adbb2d47> (Accessed on 08-06-
854 2023)

855 Hua, F., Wang, L., Fisher, B., Zheng, X., Wang, X., Douglas, W.Y., Tang, Y., Zhu,
856 J. Wilcove, D.S. (2018). Tree plantations displacing native forests: The nature
857 and drivers of apparent forest recovery on former croplands in Southwestern
858 China from 2000 to 2015. Biological Conservation, 222: 113-124.
859 <https://doi.org/10.1016/j.biocon.2018.03.034>

860 Jordan, M.I. Mitchell, T.M. (2015). Machine learning: Trends, perspectives, and
861 prospects. Science, 349(6245): 255-260. <https://doi.org/10.1126/science.aaa841>

862 Jung, Y. (2018). Multiple predicting K-fold cross-validation for model selection.
863 Journal of Nonparametric Statistics, 30(1): 197-215.
864 <https://doi.org/10.1080/10485252.2017.1404598>

865 Kanishka, G. Eldho, T. (2017). Watershed classification using isomap technique and
866 hydrometeorological attributes. Journal of Hydrologic Engineering, 22(10):
867 04017040. [https://doi.org/10.1061/\(ASCE\)HE.1943-5584.0001562](https://doi.org/10.1061/(ASCE)HE.1943-5584.0001562)

868 Kratzert, F., Klotz, D., Herrnegger, M., Sampson, A.K., Hochreiter, S. Nearing, G.S.
869 (2019). Toward improved predictions in ungauged basins: Exploiting the power
870 of machine learning. Water Resources Research, 55(12): 11344-11354.
871 <https://doi.org/10.1029/2019WR026065>

872 Lenhart, T., Eckhardt, K., Fohrer, N. Frede, H.G. (2002). Comparison of two different
873 approaches of sensitivity analysis. Physics and Chemistry of the Earth, Parts
874 A/B/C, 27(9): 645-654. [https://doi.org/10.1016/S1474-7065\(02\)00049-9](https://doi.org/10.1016/S1474-7065(02)00049-9)

875 Li, K., Wang, G., Gao, J., Guo, L., Li, J. Guan, M. (2024). The rainfall threshold of
876 forest cover for regulating extreme floods in mountainous catchments. Catena,
877 236: 107707. <https://doi.org/10.1016/j.catena.2023.107707>

878 Li, X., Khandelwal, A., Jia, X., Cutler, K., Ghosh, R., Renganathan, A., Xu, S., Tayal,
879 K., Nieber, J. Duffy, C. (2022). Regionalization in a global hydrologic deep
880 learning model: from physical descriptors to random vectors. Water Resources
881 Research, 58(8): e2021WR031794. <https://doi.org/10.1029/2021WR031794>

882 Li, Z., Xu, X., Yu, B., Xu, C., Liu, M. Wang, K. (2016). Quantifying the impacts of
883 climate and human activities on water and sediment discharge in a karst region
884 of southwest China. *Journal of Hydrology*, 542: 836-849.
885 <https://doi.org/10.1016/j.jhydrol.2016.09.049>

886 Liu, C., Guo, L., Ye, L., Zhang, S., Zhao, Y. Song, T. (2018). A review of advances in
887 China's flash flood early-warning system. *Natural hazards*, 92: 619-634.
888 <https://doi.org/10.1007/s11069-018-3173-7>

889 Luo, P., He, B., Takara, K., Xiong, Y.E., Nover, D., Duan, W. Fukushi, K. (2015).
890 Historical assessment of Chinese and Japanese flood management policies and
891 implications for managing future floods. *Environmental Science & Policy*, 48:
892 265-277. <https://doi.org/10.1016/j.envsci.2014.12.015>

893 McMillan, H.K. (2021). A review of hydrologic signatures and their applications. *Wiley
894 Interdisciplinary Reviews: Water*, 8(1): e1499.
895 <https://doi.org/10.1002/wat2.1499>

896 Morel-Seytoux, H.J. Khanji, J. (1974). Derivation of an equation of infiltration. *Water
897 Resources Research*, 10(4): 795-800.
898 <https://doi.org/10.1029/WR010i004p00795>

899 Mosavi, A., Ozturk, P. Chau, K.w. (2018). Flood prediction using machine learning
900 models: Literature review. *Water*, 10(11): 1536.
901 <https://doi.org/10.3390/w10111536>

902 Nearing, G., Cohen, D., Dube, V., Gauch, M., Gilon, O., Harrigan, S., Hassidim, A.,
903 Klotz, D., Kratzert, F., Metzger, A., Nevo, S., Pappenberger, F., Prudhomme, C.,
904 Shalev, G., Shenzis, S., Tekalign, T.Y., Weitzner, D. Matias, Y. (2024). Global
905 prediction of extreme floods in ungauged watersheds. *Nature*, 627(8004): 559-
906 563. <https://doi.org/10.1038/s41586-024-07145-1>

907 Pachauri, R.K., Allen, M.R., Barros, V.R., Broome, J., Cramer, W., Christ, R., Church,
908 J.A., Clarke, L., Dahe, Q. Dasgupta, P., (2014). Climate change 2014: synthesis
909 report. Contribution of Working Groups I, II and III to the fifth assessment
910 report of the Intergovernmental Panel on Climate Change.

911 Papageorgaki, I. Nalbantis, I. (2016). Classification of Drainage Basins Based on
912 Readily Available Information. *Water Resources Management*, 30(15): 5559-
913 5574. <https://doi.org/10.1007/s11269-016-1410-y>

914 Pugliese, A., Persiano, S., Bagli, S., Mazzoli, P., Parajka, J., Arheimer, B., Capell, R.,
915 Montanari, A., Blöschl, G. Castellarin, A. (2018). A geostatistical data-
916 assimilation technique for enhancing macro-scale rainfall-runoff simulations.
917 *Hydrology and Earth System Sciences*, 22(9): 4633-4648.
918 <https://doi.org/10.5194/hess-22-4633-2018>

919 Qi, W., Zhang, C., Fu, G. Zhou, H. (2016). Quantifying dynamic sensitivity of
920 optimization algorithm parameters to improve hydrological model calibration.
921 *Journal of Hydrology*, 533: 213-223.
922 <https://doi.org/10.1016/j.jhydrol.2015.11.052>

923 Ragettli, S., Zhou, J., Wang, H., Liu, C. Guo, L. (2017). Modeling flash floods in
924 ungauged mountain catchments of China: A decision tree learning approach for

925 parameter regionalization. *Journal of Hydrology*, 555: 330-346.
926 <https://doi.org/10.1016/j.jhydrol.2017.10.031>

927 Rockström, J., Gupta, J., Qin, D., Lade, S.J., Abrams, J.F., Andersen, L.S., Armstrong
928 McKay, D.I., Bai, X., Bala, G., Bunn, S.E., Ciobanu, D., DeClerck, F., Ebi, K.,
929 Gifford, L., Gordon, C., Hasan, S., Kanie, N., Lenton, T.M., Loriani, S.,
930 Liverman, D.M., Mohamed, A., Nakicenovic, N., Obura, D., Ospina, D.,
931 Prodani, K., Rammelt, C., Sakschewski, B., Scholtens, J., Stewart-Koster, B.,
932 Tharammal, T., van Vuuren, D., Verburg, P.H., Winkelmann, R., Zimm, C.,
933 Bennett, E.M., Bringezu, S., Broadgate, W., Green, P.A., Huang, L., Jacobson,
934 L., Ndehedehe, C., Pedde, S., Rocha, J., Scheffer, M., Schulte-Uebbing, L., de
935 Vries, W., Xiao, C., Xu, C., Xu, X., Zafra-Calvo, N.Zhang, X. (2023). Safe and
936 just Earth system boundaries. *Nature*, 619(7968): 102-111.
937 <https://doi.org/10.1038/s41586-023-06083-8>

938 Sain, S.R. (1996). The Nature of Statistical Learning Theory. *Technometrics*, 38(4):
939 409-409. <https://doi.org/10.1080/00401706.1996.10484565>

940 Salmeron, R., García, C.García, J. (2018). Variance inflation factor and condition
941 number in multiple linear regression. *Journal of statistical computation and*
942 *simulation*, 88(12): 2365-2384.
943 <https://doi.org/10.1080/00949655.2018.1463376>

944 Schoups, G., van de Giesen, N.C.Savenije, H.H.G. (2008). Model complexity control
945 for hydrologic prediction. *Water Resources Research*, 44(12).
946 <https://doi.org/10.1029/2008WR006836>

947 Shangguan, W., Dai, Y., Liu, B., Zhu, A., Duan, Q., Wu, L., Ji, D., Ye, A., Yuan,
948 H.Zhang, Q. (2013). A China data set of soil properties for land surface
949 modeling. *Journal of Advances in Modeling Earth Systems*, 5(2): 212-224.
950 <https://doi.org/10.1002/jame.20026>

951 Song, Z., Xia, J., Wang, G., She, D., Hu, C.Hong, S. (2022). Regionalization of
952 hydrological model parameters using gradient boosting machine. *Hydrology*
953 and *Earth System Sciences*, 26(2): 505-524. <https://doi.org/10.5194/hess-26-505-2022>

955 Tang, S., Sun, F., Liu, W., Wang, H., Feng, Y.Li, Z. (2023). Optimal Postprocessing
956 Strategies With LSTM for Global Streamflow Prediction in Ungauged Basins.
957 *Water Resources Research*, 59(7): e2022WR034352.
958 <https://doi.org/10.1029/2022WR034352>

959 Wainwright, J.Mulligan, M., (2013). Environmental modelling: finding simplicity in
960 complexity. John Wiley & Sons.

961 Wani, O., Beckers, J.V.L., Weerts, A.H.Solomatine, D.P. (2017). Residual uncertainty
962 estimation using instance-based learning with applications to hydrologic
963 forecasting. *Hydrol. Earth Syst. Sci.*, 21(8): 4021-4036.
964 <https://doi.org/10.5194/hess-21-4021-2017>

965 Wu, H., Zhang, J., Bao, Z., Wang, G., Wang, W., Yang, Y.Wang, J. (2022). Runoff
966 modeling in ungauged catchments using machine learning algorithm-based
967 model parameters regionalization methodology. *Engineering*.
968 <https://doi.org/10.1016/j.eng.2021.12.014>

969 Xu, Q., Chen, J., Peart, M.R., Ng, C.-N., Hau, B.C.H.Law, W.W.Y. (2018). Exploration
970 of severities of rainfall and runoff extremes in ungauged catchments: A case
971 study of Lai Chi Wo in Hong Kong, China. *Science of The Total Environment*,
972 634: 640-649. <https://doi.org/10.1016/j.scitotenv.2018.04.024>

973 Xu, T.Liang, F. (2021). Machine learning for hydrologic sciences: An introductory
974 overview. *Wiley Interdisciplinary Reviews: Water*, 8(5).
975 <https://doi.org/10.1002/wat2.1533>

976 Yang, X., Magnusson, J., Rizzi, J.Xu, C.-Y. (2018). Runoff prediction in ungauged
977 catchments in Norway: comparison of regionalization approaches. *Hydrology*
978 *Research*, 49(2): 487-505. <https://doi.org/10.2166/nh.2017.071>

979 Yang, X., Magnusson, J.Xu, C.Y. (2019). Transferability of regionalization methods
980 under changing climate. *Journal of Hydrology*, 568: 67-81.
981 <https://doi.org/10.1016/j.jhydrol.2018.10.030>

982 Zhai, X., Guo, L., Liu, R.Zhang, Y. (2018). Rainfall threshold determination for flash
983 flood warning in mountainous catchments with consideration of antecedent soil
984 moisture and rainfall pattern. *Natural Hazards*, 94: 605-625.
985 <https://doi.org/10.1007/s11069-018-3404-y>

986 Zhang, B., Ouyang, C., Cui, P., Xu, Q., Wang, D., Zhang, F., Li, Z., Fan, L., Lovati, M.,
987 Liu, Y.Zhang, Q. (2024). Deep learning for cross-region streamflow and flood
988 forecasting at a global scale. *The Innovation*, 5(3).
989 <https://doi.org/10.1016/j.xinn.2024.100617>

990 Zhang, Y., Chiew, F.H., Li, M.Post, D. (2018). Predicting runoff signatures using
991 regression and hydrological modeling approaches. *Water Resources Research*,
992 54(10): 7859-7878. <https://doi.org/10.1029/2018WR023325>

993 Zhang, Y., Chiew, F.H., Liu, C., Tang, Q., Xia, J., Tian, J., Kong, D.Li, C. (2020). Can
994 remotely sensed actual evapotranspiration facilitate hydrological prediction in
995 ungauged regions without runoff calibration? *Water Resources Research*, 56(1):
996 e2019WR026236. <https://doi.org/10.1029/2019WR026236>

997 Zhang, Y., Ragettli, S., Molnar, P., Fink, O.Peleg, N. (2022). Generalization of an
998 Encoder-Decoder LSTM model for flood prediction in ungauged catchments.
999 *Journal of Hydrology*, 614: 128577.
1000 <https://doi.org/10.1016/j.jhydrol.2022.128577>

1001 Zounemat-Kermani, M., Batelaan, O., Fadaee, M.Hinkelmann, R. (2021). Ensemble
1002 machine learning paradigms in hydrology: A review. *Journal of Hydrology*, 598:
1003 126266. <https://doi.org/10.1016/j.jhydrol.2021.126266>

1004