

Urban Weather Modeling using WRF: Linking Physical Assumptions, Code Implementation, and Observational Needs

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Abstract.

The Weather Research and Forecasting (WRF) model includes urban schemes that simulate the influence of urban surfaces on the atmosphere using parameterizations for flux, and radiative exchanges. Three core schemes—the Bulk urban parameterization, Single-Layer Urban Canopy Model (SLUCM), and Multi-Layer Urban Canopy Model (MLUCM)—represent increasing levels of complexity. Although the parameterizations within these urban schemes are described in the literature, their specific implementation remains poorly documented, thus slowing down model development efforts.

This manuscript provides a roadmap to the three urban schemes in WRF version 4.5.2, presenting equations using the same symbols as in the model code, along with references to code lines, and including graphics and explanations that connect the code to its physical foundations. Our thorough review of the urban parameterizations implemented in WRF version 4.5.2 highlighted a handful of parameters that may introduce discontinuities in simulations: (i) in the SLUCM, a 1 mm/hr rain rate threshold is employed to switch between two minimum moisture availability parameterizations, thus impacting latent heat flux calculations; (ii) in the SLUCM a threshold is used to partition shortwave radiation into direct and diffuse components; (iii) in all three urban schemes, the bulk Richardson number is employed to select the similarity function, which influences the vertical distribution of heat and momentum. We also identified a highly simplified treatment of the radiative balance on roof surfaces. The implications of these simplifications can be assessed through targeted observations across relevant conditions, including varying precipitation rates, cloud cover, and transitions between stability regimes. Furthermore, the widespread application of the Monin-Obukhov similarity theory in these urban schemes warrants model evaluation under highly stable and unstable conditions and in heterogeneous urban settings with variable land cover and building heights on scales finer than model resolution. To address these challenges, we offer guidance on observational strategies, emphasizing the need for multi-parametric measurements to capture potential compensating biases and multi-height measurements that align with the levels where quantities are diagnostic and prognosed in the model (i.e., the lowest atmospheric level of the WRF model). Finally, our inspection of the code revealed implementation bugs that have now been corrected in WRF versions 4.6.0 and 4.6.1. Sensitivity tests over the Atlanta urban area show that these corrections affect surface temperatures, underscoring the importance of performing rigorous documentation and verification of the implementation of parameterizations in model code.

1 Introduction

Weather Research and Forecasting (WRF) is a mesoscale atmospheric model widely used for research and short- to medium-range weather forecasting (Lynn and Yair (2010)) as well as regional long-term climate simulations Rasmussen et al. (2023). Over the past decade, development efforts have expanded the capabilities of the WRF model (Powers et al. (2017)) to include, among other things, chemistry (WRF-Chem; Grell et al. (2005)), hydrology (WRF-Hydro; Gochis et al. (2015)), and urban processes (WRF-Urban; Chen et al. (2011)).

Urban environments, now home to more than half of the global population (United Nations, Department of Economic and Social Affairs, Population Division, 2015), introduce significant heterogeneity in surface energy exchange due to impervious materials, complex geometries, and anthropogenic heat emissions. These complexities modify the atmospheric structure, especially in the lowest kilometer of the boundary layer (Lee et al., 2011), and require specialized modeling approaches. The community version of WRF includes three urban schemes—the Bulk scheme (also know as the Slab model), the Single-Layer Urban Canopy Model (SLUCM), and the Multi-Layer Urban Canopy Model (MLUCM)—that represent urban processes with varying levels of complexity. Despite their widespread use in urban climate research (e.g., Georgescu et al., 2024; Zhu and Ooka, 2023), the numerical implementations of these schemes are poorly documented, limiting model development, and lacking transparency regarding parameter manipulations needed to integrate urban parameterizations with other WRF modules. Improved documentation would help identify key parameters that influence model performance and guide targeted observations, better aligning modeling efforts with measurement strategies to enhance predictability.

This work seeks to complement the literature focused on the underlying physics of urban parameterizations in the WRF-Urban (e.g., Garuma, 2018; Tarasova et al., 2023; Grimmond et al., 2010; Masson, 2006; Barlow, 2014). It sets itself apart by providing a detailed roadmap of the parameterizations employed in WRF-Urban as implemented in the model version 4.5.2 using the exact mathematical symbols employed in the model code and referencing the code line numbers to accelerate future efforts to develop the code. We also provide new graphical schematics to guide the reader through the relevant physical concepts to improve users understanding of key model assumptions. We identified a handful of parameters that may introduce discontinuities in simulations, for which we offer guidance on strategies to assess their impact and/or tune them with observations. Additionally, our systematic code review uncovered several implementation bugs, which we document here. Sensitivity simulations are presented to illustrate the potential impact of these bugs on model output, emphasizing the importance of continued model code verification.

2 WRF-Urban modeling system

2.1 General description of the WRF Urban Canopy Models

55 The WRF-Urban code package contains three schemes to simulate the effects of urban surfaces and on turbulent fluxes and radiation: the Bulk urban parameterization, the SLUCM, and the MLUCM, which includes the Building Effect Parameterization (BEP) model, which is often coupled with the Building Energy Model (BEM).

The Bulk urban parameterization, originally presented in (Chen et al., 2011; Liu et al., 2006), simplifies urban areas by representing them as one flat, uniform surface with adjusted physical properties. These adjustments include increased heat capacity, 60 thermal conductivity, and surface roughness, along with reduced albedo. This simplified framework makes the Bulk urban Parameterization computationally efficient and well suited for studies where high-resolution urban details are unnecessary or where data availability and computational resources are constrained.

The WRF-Urban urban canopy models (UCMs) - the SLUCM (Kusaka et al. (2001)) and the BEP (Martilli et al. (2002)) - provide more detailed representations by simulating urban environments as idealized canyons. The SLUCM and BEP have 65 several surface types, each of which is treated independently, with separate calculations for turbulent and radiative processes. Both the SLUCM and BEP models can represent a mix of roof types, such as green roofs and traditional roofs, within the same grid cell. However, they are limited to representing a single type of wall and street per grid cell. The primary distinction between SLUCM and BEP lies in their vertical structure. The SLUCM represents buildings as an infinitely long 2D array of buildings of uniform height, and simulates the urban canopy as a single homogeneous layer beneath the first atmospheric layer (Hang 70 et al. (2024)). In contrast, the BEP model resolves the urban canopy layer into multiple vertical levels, which extend within but do not necessarily align with the atmospheric levels in the WRF model. Building heights are represented as a distribution function and, at each level, energy balance equations are solved taking into account the probability that buildings are equal to or exceeding that height. This approach allows BEP to provide a more detailed and realistic representation of interactions within and above urban canyons. However, the added complexity comes at the cost of higher computational demands compared 75 to single-layer models (Garuma (2018)). When coupled with the BEM model, BEP-BEM extends the dimensions of the BEP urban modeling framework by simulating, among other things, latent heat fluxes, windows on walls, building energy use, and anthropogenic heat emissions Salamanca and Martilli (2010). A single grid cell in a mesoscale model typically contains numerous buildings and streets that vary in physical characteristics such as height, width, and thermal properties. However, representing these heterogeneous features with a single, uniform type of building or street oversimplifies the complexity of real 80 urban morphology. Therefore, a more realistic depiction of urban areas requires advanced urban canopy models that incorporate multiple types of buildings and their distinct thermal attributes.

Fundamentally, each of these urban models prognoses the surface temperature T for each represented surface (e.g., roof, wall, and road) as well as turbulent fluxes (sensible heat SH and latent heat LH) between these urban surfaces and the ambient air. The parameters enter the energy balance calculation for each represented surface alongside the net radiation R_{net} , and the

85 ground heat flux G :

$$C_p \frac{\partial(T)}{\partial t} = R_{net} - G - SH - LH, \quad (1)$$

$$\text{where } R_{net} = SW + \epsilon(LW - UP_{flux}), \quad (2)$$

where C_p is the heat capacity of the surface per unit area, SW is the downward SWR flux absorbed by the surface, LW is the total downward longwave radiative heat flux, UP_{flux} is the outgoing LW radiative flux, and ϵ is the emissivity of the surface under consideration. Given their varying levels of complexity, each urban model in WRF employs distinct parameterization formulations and assumptions to compute fluxes. The following subsections delve into these differing approaches.

3 Turbulent fluxes

3.1 Sensible heat flux parameterizations

BOX 1 Underbraces under equations to inform the reader as to the location of the equation in the WRF v4.5.2 code. For example, Equation 3 for HFX is implemented on line 409 of file `module_sf_slab.F` in WRF v4.5.2.

95 Despite differences in the way they represent urban surfaces, the Bulk, SLUCM, and BEP parameterizations for sensible heat flux share several fundamental similarities. All employ a bulk aerodynamic formulation that depends on (i) the temperature gradient between the represented surface(s) and the adjacent air, (ii) wind speed, and (iii) heat/moisture transfer coefficients, which incorporate atmospheric stability effects through the Bulk Richardson number. Additionally, all rely on the Monin–Obukhov Similarity Theory (MOST) to estimate surface fluxes.

100 In the Bulk model, sensible heat flux is estimated based on the potential temperature difference between the slab surface (THG) and the temperature of the air at the first atmospheric level (THX), as well as the heat transfer coefficient of the air (FLHC) following the formulation of Zhang and Anthes (1982):

$$\underbrace{HFX}_{\substack{\text{Line 410,} \\ \text{module_sf_slab.F}}} = FLHC \times (THG - THX), \quad (3)$$

(FLHC) takes into account the specific heat of dry air at constant pressure CPM, the air density RHOX, and the potential temperature of the slab surface THGB as well as the turbulent velocity scale UST and the temperature scale MOL following:

$$\underbrace{FLHC}_{\substack{\text{Line 877} \\ \text{module_sf_sfclay.F}}} = CPM \times RHOX \times UST \times MOL / (THX - THGB), \quad (4)$$

The temperature scale MOL incorporates the potential temperature difference between the ground and the lowest atmospheric model level (DTG) and the integrated similarity function for heat PSIT following

$$\underbrace{MOL}_{\substack{\text{Line 815} \\ \text{module_sf_sfclay.F}}} = \frac{KARMAN \cdot DTG}{PSIT \cdot PRT}, \quad (5)$$

$$DGT = (THX - THGB) \quad (6)$$

where KARMAN is the von Karman constant and PRT is the Prandtl number with a constant value of 1.

The friction velocity UST incorporates the the wind speed at the lowest atmospheric level ($WSPD$) and the integrated similarity function for momentum $PSIX$, following:

$$UST = \frac{KARMAN * WSPD}{PSIX}, \quad (7)$$

Line 794
module_sf_sfclay.F

Under strongly unstable conditions, additional corrections can be applied to $WSPD$ to account for convective velocity scaling following Beljaars (1995) and subgrid-scale turbulence effects following Mahrt and Sun (1995).

The integrated similarity functions $PSIT$ and $PSIX$ were derived empirically using the MOST formulation (Beljaars and Holtslag (1991)), which also allows for the derivation of vertical profiles of meteorological variables, enabling the diagnostic estimation of quantities such as 2 m air temperature and 10 m wind speed at standard observational heights (see Box 2). In the Bulk scheme, the bulk Richardson number is used to select among four different integrated similarity formulations (Jiménez et al. (2012)). However, this reliance on bulk Richardson number thresholds to choose the appropriate integrated similarity functions introduces discontinuities and inconsistencies in flux calculations, particularly in transition regimes. Observational data could help inform the use of this parameter. Further discussion of this topic will be provided in Section 3.3.

Among the similarity function formulations implemented in the Bulk scheme are those of Fairall et al. (1996) for neutral to free convective conditions, and Chenge and Brutsaert (2005) for neutral to very stable conditions. These formulations extend the original similarity functions derived from the Kansas field program using a limited number of observations over flat, homogeneous surfaces Izumi and Barad (1970); Businger et al. (1971); Dyer (1974); Hicks (1976). Given the complexity and heterogeneity of urban environments, additional observational datasets that cover a wider range of urban settings could help refine these empirical similarity functions. A discussion of observational approaches for improving these formulations is provided in Section 3.3.

The integrated similarity functions ($PSIT$ and $PSIX$) are corrected for the excess aerodynamic resistance of urban areas following the proposed approach Chen and Zhang (2009). This effectively diminishes the efficiency of turbulent transfer in urban areas. This is achieved by adopting a modified formulation of the roughness length for heat, which takes into account the Reynolds number and the Zilitinkevich coefficient ($Czil$) - an empirical coefficient for bulk transfer. In the WRF version 4.5.2, $Czil$ is set to 0.1. Chen et al. (1997) and Chen and Zhang (2009) designed a calibration procedure for $Czil$ which relies on the comparison of observed and simulated fluxes collected over different land-cover type. They also present an empirical formula for $Czil$, which relates it to canopy height (Equation (4) in Chen and Zhang (2009)). The same procedure could be repeated over other urban environments to further tune $Czil$.

In the SLUCM a distinction is made during the sensible heat flux calculation for each surface ($H_{[surface]}$) to account for the fact that the roof surfaces are closer to the first atmospheric level, than building walls and ground. This distinction leads to using different air temperature and wind information when calculating heat flux for the roof (R) and green roof (GR) surfaces (sup eq. S1) compared to the wall (B) and ground (G) surfaces (eq. 8). Where canopy wind speed UC is used for the building walls and ground surfaces, wind speed at the lowest atmospheric level, denoted as UA , is used for the roof surfaces. Similarly, where the canopy temperature TCP is used for building walls and ground calculations, the air temperature at the lowest atmospheric level of the WRF mesoscale model TA is employed for roof surface

calculations.

$$\underbrace{H_{[B,G]}}_{\substack{\text{Lines: } \{1322,1323\} \\ \text{module_sf_urban.F}}} = RHO * CP * CH_{[B,G]} * UC * (T_{[B,G]}P - TCP), \quad (8)$$

where RHO denotes the air density at the lowest atmospheric layer, CP denotes the specific heat of dry air at constant pressure and $CH_{[GR,R,B,G]}$ represents the surface exchange coefficients.

Canopy-level wind speed UC is a diagnostic quantity representing wind speed at a level of 0.7 times the mean building height. It is estimated as the scaled version of the wind speed at roof level UR following the approach proposed by Inoue (1963):

$$\underbrace{UC}_{\substack{\text{Line } 857 \\ \text{module_sf_urban.F}}} = UR * EXP \left(-BB * \left(1 - \frac{ZC}{ZR} \right) \right), \quad (9)$$

$$\underbrace{UR}_{\substack{\text{Line } 852 \\ \text{module_sf_urban.F}}} = UA * LOG \left(\frac{ZR - ZDC}{Z0C} \right) / LOG \left(\frac{ZA - ZDC}{Z0C} \right), \quad (10)$$

$$\underbrace{BB}_{\substack{\text{Line } 856 \\ \text{module_sf_urban.F}}} = 0.4 * ZR / \left(0.4 * (ZR - ZDC) * alog \left(\frac{ZR - ZDC}{Z0C} \right) \right), \quad (11)$$

155 where ZC is the effective canyon depth taken to be 0.7 times the building height (ZR). The intermediate parameter BB encapsulates the effects of urban roughness (including the canyon roughness length ($Z0C$) and zero-plane displacement height (ZDC)) on the canopy wind speed as described in Macdonald et al. (1998). It is worth noting that a different approach is used to diagnose 10 m winds, which relies on MOST rather than the empirical damping of roof-level winds (see box 2).

Canopy temperature (TCP) is also a diagnostic quantity representing equivalent aerodynamic temperature at a level of 0.7 times the building height. It is computed as blend of the thermal influences from of the different elements in the urban canopy (the air, ground and walls) following:

$$\underbrace{TCP}_{\substack{\text{Line } 1440 \\ \text{module_sf_urban.F}}} = \frac{RW \cdot ALPHA_C \cdot TA + RW \cdot ALPHA_G \cdot TGP + W \cdot ALPHA_B \cdot TBP}{RW \cdot ALPHA_C + RW \cdot ALPHA_G + W \cdot ALPHA_B}, \quad (12)$$

165 where TA is the air temperature at the lowest atmospheric level, TGP is the ground surface temperature at previous time step, TBP is the wall surface temperature at previous time step. $ALPHA_C$, $ALPHA_G$, $ALPHA_B$ are the heat transfer coefficients associated with the canopy, ground, and wall surfaces, respectively (here acting as weighting coefficients). RW is the normalized canyon width, W is the doubled normalized building height (here used to normalize the contributions by the effective area over which each surface type impacts the local air temperature). It is worth noting that 2 m air temperature is not diagnosed using this approach, but rather using MOST (see box 2).

BOX 2 Diagnosing 2 m air temperature and 10 m wind speed 2 m air temperature and 10 m wind speed are frequently used as benchmarks to assess model performance; however, it is important to recognize that these quantities are not directly prognosed by WRF-Urban. Rather, they are diagnosed via scaling relationships based on similarity theory. As a result, uncertainties arise from both the underlying urban surface parameterizations and the diagnostic scaling procedures, complicating the attribution of errors. In the SLUCM, 2 m potential air temperature ($TH2$) is diagnosed using a logarithmic profile approach. Specifically, it is computed as a weighted combination of the effective urban surface temperature (TS) and the temperature at the lowest atmospheric level (TA), scaled by stability-corrected factors following:

$$\underbrace{TH2}_{\substack{\text{Line 1592} \\ \text{module_sf_urban.F}}} = TS + (TA - TS)(CHS/CHS2), \quad (13)$$

$$\underbrace{TS}_{\substack{\text{Line 1541} \\ \text{module_sf_urban.F}}} = TA + FLXTH/CHS, \quad (14)$$

where CHS is the turbulent transfer coefficient/aerodynamic conductance for heat at the lowest level of atmospheric model, $CHS2$ is the turbulent transfer coefficient for heat at the 2 m. The effective urban surface temperature (TS) takes into account the area-weighted sum of the sensible heat fluxes from the roof, ground, and wall surfaces ($FLXTH$).

Similarly, 10 m wind speed ($U10$) is diagnosed by scaling the wind speed at the lowest atmospheric level ($U1$) using the stability-corrected wind profile evaluated at 10 m ($PSIX10$).

$$\underbrace{U10}_{\substack{\text{Line 1586} \\ \text{module_sf_urban.F}}} = U1(PSIX10/PSIX), \quad (15)$$

where $PSIX$ is the integrated similarity function for momentum at lowest model level whereas $PSIX10$ is at 10 m height.

A similar approach is used for the Bulk scheme where $TH2$ and $U10$ are calculated in `module_sf_sfclay.F`.

In the BEP (or BEP-BEM) scheme, the urban surfaces are embedded into the mesoscale model through a multi-layer representation, such that the model's vertical resolution can be leveraged to more directly diagnose near-surface values. In practice, BEP (or BEP-BEM) assigns the temperature and wind speed of the lowest urban grid level to 2 m and 10 m levels for temperature and wind speed, respectively.

170 In WRF version 4.5.2, the surface exchange coefficients for heat and moisture transport (CH) used to estimate the sensible heat flux are not directly calculated. Rather, CH is estimated from solving for the heat transfer coefficients for each surface $ALPHA_{surface}$ - analogous to $FLHC$ in the Bulk scheme - following:

$$CH = VKRM/PSIH = KARMAN^2/PSIH/PSIM. \quad (16)$$

$$\underbrace{ALPHA_{[R,GR]}}_{\substack{\text{Lines [971,1078]} \\ \text{module_sf_urban.F}}} = RHO * CP * VKRM * USTAR/PSIH, \quad (17)$$

$$\underbrace{ALPHA_{[B,G]}}_{\substack{\text{Line 1684} \\ \text{module_sf_urban.F}}} = RHO * CP * KARMAN^2 * UC / PSIH / PSIM, \quad (18)$$

PSIH and PSIM represent the integrated similarity functions for heat and momentum - analogous to PSIT and PSIX in the Bulk scheme. *PSIH* and *PSIM* are estimated following MOST. In the SLUCM, the bulk Richardson number is again used to select among three different formulations. However, the Bulk scheme uses a threshold of 0.2, while the SLUCM uses 0.14, which may lead to inconsistencies between the bulk urban parameterization and urban canopy model outputs—particularly in transitional regimes. The SLUCM default includes the analytical formulations listed in Paulson (1970) and the option to use the formulation of Łobocki (1993) for the roof. A limited number of observation-based studies have examined the applicability of the similarity functions over highly heterogeneous urban surfaces, with most focusing on low-rise neighborhoods where buildings are less than 23-m tall (Rotach (1993); Zou et al. (2015); Moriwaki and Kanda (2006)). Two consistent trends have emerged from these studies. First, the flux-gradient relationship for the momentum approaches the classical similarity relations of Businger et al. (1971) at the transition between the roughness sublayer and the inertial sublayer. Second, within the roughness sublayer, momentum flux-gradient relations conform to classical MOST predictions under unstable (convective) conditions but deviate significantly under stable conditions. However, these trends require further validation through field observations, especially for urban regions with supertall buildings and with highly variable building heights for which we offer recommendations in Sec. 3.3.

In the SLUCM, the integrated similarity functions are also corrected for the excess aerodynamic resistance in urban areas following Kanda et al. (2007). In short, SLUCM corrects for excess aerodynamic resistance by directly modifying the resistance formulation via introducing an additional resistance term. The excess aerodynamic resistance for the wall and road is calculated based on user-defined urban class specific momentum and thermal roughness lengths. The thermal roughness length is dynamically adjusted based on the Reynolds number of roughness in a way similar to the empirical formula involving the factor *Czil*.

By embedding urban surfaces directly into the WRF model’s atmosphere, the BEP model eliminates the need for integrated similarity functions to estimate canopy-level temperature and winds. This shift makes the orientation of surfaces (vertical vs. horizontal) more relevant than their theoretical distance from the lowest atmospheric level. For the vertical surface (the walls (*w*)), the BEP (and BEP-BEM) model computes, at each level, the sensible heat flux (*sf_w*) using potential temperature gradient between the air and the wall surface (*pt – ptw*) and the heat transfer coefficient of the wall surface (*hc*) using the empirical relation developed by McAdams (1954) which is based on wind tunnel experiments performed by Jürges (1924):

$$\underbrace{sf_{[w]}}_{\substack{\text{Lines [4192]} \\ \text{module_sf_bep_bem.F}}} = hc * (pt - ptw), \quad (19)$$

$$\underbrace{hc}_{\substack{\text{Line 4168} \\ \text{module_sf_bep_bem.F}}} = 5.678 \times (1.09 + 0.23 \times (vett/0.3048)), \quad (vett < 4.88) \quad (20)$$

$$\underbrace{hc}_{\substack{\text{Line 4170} \\ \text{module_sf_bep_bem.F}}} = 5.678 * 0.53 * ((vett/0.3048)^{0.78}), \quad (vett \geq 4.88). \quad (21)$$

This empirical relation for *hc* relies solely on the horizontal wind speed (*vett*) and incorporates several empirical coefficients. However, its accuracy could be improved by validating it against a broader dataset, as further discussed in Section 3.3.

The BEP-BEM formulations for sensible heat for the horizontal surfaces (ground (*g*), and roof (*r*)) more explicitly accounts for aerodynamic resistance (through *rr* - the ratio of the drag coefficients for momentum and heat), stability effects (through *fh* - stability correction

function for heat flux), wind-driven turbulent exchange (through $utot$ - the wind speed) and temperature gradients (between surface and air).

$$210 \quad \underbrace{sf_{[g,r]}}_{\substack{\text{Lines: [g,r - 4710, 4702, 4698] \\ module_sf_bep_bem.F}}} = da_u \times cp_u \times \underbrace{aa^2 \times utot \times (pt - [ptg, ptr]) \times fh/rr}_{\substack{= thb_u \text{ in Line 2657 and } tha_u * pt_u = 0 \\ \text{of module_sf_bep.F}}}, \quad (22)$$

where da_u is air density, cp_u is the specific heat of the air at constant pressure, aa^2 is the surface drag coefficient for momentum, ptg is the potential temperature of the ground, ptr is the potential temperature of the roof, pt is the potential temperature of the air. rr is the ratio of the surface drag coefficients for momentum and heat and is set to a constant value of 0.74. fh is set to take a different value in stable, neutral, and unstable boundary layers as detailed in Louis (1979). Like in the Bulk scheme and SLUCM, the bulk Richardson number is used to determine the atmospheric stability regime.

3.2 Latent heat flux parameterizations

Many parallels also exist between the Bulk, SLUCM, and BEP parameterizations for latent heat flux. All employ a bulk aerodynamic formulation that depends on (i) the moisture gradient between the represented surface(s) and the adjacent air, (ii) wind speed, and (iii) heat/moisture transfer coefficients, which incorporate atmospheric stability effects through the Bulk Richardson number. Additionally, all parameterizations rely on MOST to estimate surface fluxes.

In the Bulk scheme, latent heat flux (LH) is estimated by calculating the moisture gradient in the form of the difference in moisture at the ground (QSG) and at the lowest atmospheric level (QX), and it factors a surface exchange coefficient for moisture (FLQC)

$$\underbrace{LH}_{\substack{\text{Line 412} \\ module_sf_slab.F}} = XLV \times QFX, \quad (23)$$

$$225 \quad \underbrace{QFX}_{\substack{\text{Line 411} \\ module_sf_slab.F}} = FLQC \times (QSG - QX) \quad (24)$$

$$\underbrace{FLQC}_{\substack{\text{Line 872} \\ module_sf_sfclay.F}} = RHOX \times MAVAIL \times UST \times KARMAN / DENOMQ, \quad (25)$$

where XLV is the latent heat of vaporization. $MAVAIL$ is surface moisture availability (ranging between 0 and 1), defined by the user in `LANDUSE.TBL`, where it is named as *SLMO*. Observations could be employed to inform parameterization of *MAVAIL* for improved latent heat flux estimation. More details on potential approaches are provided in Section 3.3. In the code version WRF V4.5.2, *DENOMQ* is also denoted as *PSIQ* and is set to *PSIT* which is the integrated similarity functions for heat presented in equation 5 of this manuscript.

When calculating the latent heat flux for each of its surfaces ($ELE_{[surface]}$), the SLUCM similarly considers the surface exchange coefficients for transport of moisture/heat (CH), the minimum moisture availability at the surfaces (BET), the wind speed, and the moisture gradient in the form of the difference in moisture at the surface ($QS0$) and in the atmosphere. Different sources of wind and surface moisture are used to calculate the latent heat flux for the wall (B) and ground (G) surfaces (eq. 26), in contrast to the roof (R), which is conceptually closer to the first atmospheric level (sup. eq. S2). For the wall and ground, canopy-level wind (UC) and canopy-level atmospheric moisture

(QCP) are used, whereas lower atmospheric level wind (UA) and moisture (QA) are used for the roof.

$$\underbrace{ELE}_{\text{Lines [1333, 1334] module_sf_urban.F}}_{[B,G]} = RHO * EL * CH_{[B,G]} * UC * BET_{[B,G]} * (QS0_{[B,G]} - QCP), \quad (26)$$

where EL represents the latent heat of vaporization.

240 Canopy-level atmospheric moisture (QCP) is estimated similarly as the canopy-level temperature by blending the moisture influences from the different elements of the urban canopy (see eq. 12 for a definition of the weight terms).

$$\underbrace{QCP}_{\text{Lines [1447] module_sf_urban.F}} = \frac{RW \cdot ALPHA_C \cdot QA + RW \cdot ALPHA_G \cdot BET_G \cdot QS0_G + W \cdot ALPHA_B \cdot BET_B \cdot QS0_B}{RW \cdot ALPHA_C + RW \cdot ALPHA_G \cdot BET_G + W \cdot ALPHA_B \cdot BET_B}. \quad (27)$$

BET - the minimum moisture availability at the surfaces - can take one of two forms depending on the value set to parameter IMP_SCHEME in the urban parameters table (i.e., $URBPARAM.TBL$, $URBPARAM_LCZ.TBL$, and $URBPARAM_UZE.TBL$). Option 2, considers more complex urban hydrology with water-holding scheme over impervious surface (Yang and Wang (2014)), while the default option (Option 1) sets BET to a value of 0.7 for rainfall events less than 1 mm/hr and for higher rainfall events sets it to:

$$\underbrace{BET}_{\text{Lines [991, 1241, 1251] module_sf_urban.F}}_{[R,B,G]} = Drel_{[R,B,G]} / dengimp(IMP_{[R,B,G]}) * porimp(IMP_{[R,B,G]}), \quad (28)$$

$$\underbrace{Drel}_{\text{Lines [984, 1232, 1234] module_sf_urban.F}}_{[R,B,G]} = Drel_{[RP,BP,GP]} + (RAIN1 - FLXHUM_{[RP,BP,GP]} * DELT / porimp(IMP_{[R,B,G]}), \quad (29)$$

250 where $RAIN1$ represents the rainfall velocity, and $DELT$ the time step, $porimp(IMP_{[R,B,G]})$ the porosity of the impervious surfaces, and $FLXHUM_{[RP,BP,GP]}$ the moisture transport rate from the previous time step. Additionally, $Drel_{[R,B,G]}$ and $Drel_{[RP,BP,GP]}$ refer to the water retention depths at the current and previous time steps, respectively, and $dengimp(IMP_{[R,B,G]})$ indicates the maximum water retention depth.

For the green roof (GR), the minimum moisture availability parameterization is replaced by one considering the sum of direct evaporation from soil ($EDIR$), evapotranspiration ($ETTR$), and canopy intercepted evaporation (ECR) divided by the potential evaporation:

$$\underbrace{ELE}_{\text{Lines [1132] module_sf_urban.F}}_{[GR]} = ETAR * RHO * EL / RHO, \quad \text{where} \quad (30)$$

$$\underbrace{ETAR}_{\text{Lines [1124] module_sf_urban.F}} = EDIR + ETTR + ECR, \quad (31)$$

where RHO is the air density.

260 A case study, UCAR WRF Support Forum (2024), reported a WRF simulation crash linked to an urban grid exhibiting an abnormally high negative sensible heat flux (-1450 W/m^2). The issue arose from excessive latent heat flux on impervious urban surfaces following rainfall, attributed to the minimum moisture availability parameter $BET_{[G]}$. Adjusting $BET_{[G]}$ from 0.7 to 0.1 for rainfall rates exceeding 1 mm/hr stabilized the simulation, highlighting $BET_{[G]}$'s critical role in urban energy balance modeling, yet observations of $BET_{[G]}$ remain scarce (further discussion of possible approaches is presented in Section 3.3).

265 The BEP model ignores the impact of precipitation on the production of latent heat and thus systematically set its flux to zero, which is a crude assumption, which is sure to bias its energy balance during and following rain as long as water remains on the surface. In the BEP-BEM model this assumption was relaxed and parameterizations for latent heat flux over the ground and roof were added, still neglecting latent heat flux for wall surfaces. The latent heat flux over the ground (lf_g), roof (lf_r) and green roof take into account the aerodynamic resistance to moisture transport ($rseg$), and water pool fraction (wf) parameters, following:

270
$$\underbrace{lf_{[g,r]}}_{\substack{\text{Lines [4711 \& 4586]} \\ \text{module_sf_bep_bem.F}}} = da_u \times latent \times wf_{[g,r]} \times (qv - qvsg) / (resg + rsveg). \quad (32)$$

where stomatal resistance to moisture transport ($rsveg$) is only nonzero for green roof. The water pool fraction ($wf_{[g,r]}$) is a function of water depth at the surface, a prognostic variable, (dg for the ground and dgr for the roof), as well as the maximum water holding capacity of the surface ($dgmax$ for the ground and $drmax$ for the roof) and considers the rainfall rate ($rainbl$), following:

$$\underbrace{wf_{[g,r]}}_{\substack{\text{Lines [2772 \& 2791]} \\ \text{module_sf_bep_bem.F}}} = [dg, dgr] / [dgmax, drmax] = \underbrace{([dg, dgr] + (rainbl + lf_{[g,r]} * dt / latent)) / [dgmax, drmax]}_{\substack{\text{Lines [1561 \& 1571]} \\ \text{module_sf_bep_bem.F}}}, \quad (33)$$

275 where da_u denotes the air density, $latent$ denotes the latent heat of vaporization, and dt is the time step. Mathematically, the minimum moisture availability (BET') in the SLUCM and the water pool fraction (wf) in the BEP-BEM model resemble each other and both represent the moisture availability making it an important parameter to evaluate using observations.

The aerodynamic resistance ($rseg$) to moisture transport incorporates the drag coefficient ratio for neutral stability ($rr = 0.74$), wind profile for the neutral stability, and a correction factor (fh) based on bulk Richardson number:

280
$$\underbrace{resg}_{\substack{\text{Line 4570} \\ \text{module_sf_bep_bem.F}}} = rr / (aa^2 \times utot \times fh), \quad (34)$$

Whereas the stomatal resistance to moisture transport ($rsveg$) follows the formulation of Chen and Dudhia (2001):

$$\underbrace{rsveg}_{\substack{\text{Line 4567} \\ \text{module_sf_bep_bem.F}}} = \min \left(\frac{rsmmin}{max(lai * \times f1 \times f2 \times f3 \times f4, 10^{-9}, rsmmax)} \right), \quad (35)$$

where $rsmmin$ is the minimum stomatal resistance, $rsmmax$ is the maximum resistance, and lai is the leaf area index. The fractional conductance ranges between 0 and 1: $f1$ represents the impact of photosynthetically active solar radiation on stomatal resistance which is a function of dimensionless active radiation and lai , $f2$ is the vapor pressure deficit which depends on the mixing ratio difference, $f3$ is a function of temperature difference and represent the air temperature effect on canopy resistance, and $f4$ is the water stress factor in computing the surface resistance to transpiration (Jacquemin and Noilhan (1990)). Stomatal resistance to moisture transport ($rsveg$) is typically measured indirectly by combining observations of latent heat flux with meteorological data and is put in the context of leaf area index. Allen et al. (2006) present recommendation to standardize the estimating of surface resistance from measurements. Their article highlights the need for

290 more observational data across diverse landscapes to refine canopy resistance estimates.

3.3 Observational approaches

A variety of observational techniques can be employed to evaluate the turbulent flux parameterizations in WRF-Urban. Table 1 summarizes what we believe are key variables, their links to model parameters, recommended instruments, and relevant literature on the techniques.

It is worth noting that multi-parametric observational datasets are especially useful for diagnosing compensating errors across coupled
295 model components and for evaluating model parameters that are challenging to measure. For example we note that by combining heat flux,
air temperature and land surface temperature measurements, the heat transfer coefficients can be evaluated. For example, Bruno et al. (2022)
evaluated heat transfer coefficients on building components using measurements collected at an outdoor experimental facility featuring seven
monitored building surfaces (walls, floor, ceiling, window, and door). The experiment involved deploying 62 temperature sensors and 13 heat
flux meters on internal and external surfaces to monitor thermal exchanges. A weather station was deployed to record outdoor temperature,
300 wind velocity, solar radiation, and sky infrared radiation. Their study showed that windows and doors are more sensitive to radiation and
conduction losses than opaque walls, thus suggesting that a single heat transfer coefficient may not be suitable to represent the entire wall
surface. They also noted that south-facing walls showed greater heat transfer coefficient fluctuations due to solar radiation suggesting that
adding consideration to solar angle may help improve the accuracy of wall heat flux calculations. Such experiments could be repeated for
various building types to help inform parameterizations of the heat transfer coefficients, especially the empirical formulation for the heat
305 transfer coefficient of the wall surface (h_c) employed in the BEP+BEM which strictly considers wind speed.

As for which urban regimes should be in focus: Given the challenges in estimating integrated similarity functions under very stable
or very unstable atmospheric conditions, we believe targeted observations in those regimes would be especially valuable. Also, since the
WRF UCMs apply thresholds based on the bulk Richardson number (Ri_b) to switch among similarity functions, observations collected near
regime transitions could help determine whether such thresholding is physically justified. Flux-gradient relationships, which form the basis
310 of similarity functions, in urban areas are challenging to evaluate not only due to complex surface conditions and limited observations, but
also because urban heterogeneity can lead to space-dependent flux-gradient behavior. Differences in building geometry, surface materials,
and vegetation across small spatial scales can alter local turbulence and vertical gradients, complicating the assumption of horizontally
homogeneous conditions underlying traditional similarity theory. As a result, evaluating existing flux-gradient formulations requires carefully
designed field campaigns that account for spatial variability—such as deploying dense sensor networks or mobile platforms to capture local
315 gradients. These considerations are critical for assessing the validity of parameterizations and improving model performance in heterogeneous
urban environments. Finally, we would argue that urban environments with high surface heterogeneity at subgrid scales warrant further study,
as they challenge the assumptions of uniform land cover or building morphology inherent to WRF-Urban. Observation-based approaches can
assist with estimating similarity relations for momentum by fitting the relationship between the dimensionless wind velocity gradient and
local stability parameters using observations spanning from the surface to the top of the roughness sublayer, which is estimated to extend
320 2 – 5 times the average building height (Zou et al. (2015); Rotach (1993); Moriwaki and Kanda (2006)). The dimensionless wind velocity
gradient can be determined by (i) fitting wind speed measurements taken at multiple heights to get the wind velocity gradient, (ii) estimating
the effective height above the zero-plane displacement height from field surveys, and (iii) calculating the friction velocity. The local stability
parameter can be computed following local similarity theory, incorporating key variables such as the effective height above the zero-plane
displacement height and the local Obukhov length. The latter derived from measurements of mean virtual potential temperature, kinematic
325 heat flux, and friction velocity obtainable from ultrasonic anemometers. The similarity relation for heat can be similarly estimated with
temperature measurement obtained from sonic anemometer or wire thermocouples.

Table 1: Observations for informing turbulent flux parameterizations in WRF-Urban.

Variable / Parameter	Measurement Methods	Key Applications	References
Sensible & latent heat flux	Eddy covariance (ECOR), small-/large-aperture scintillometers, dual-wavelength systems <i>Small-aperture for <250m paths; large-aperture for urban-scale fluxes; dual-wavelength estimates both H and LE.</i>	Validate simulated fluxes and assess surface transfer coefficients	Ward et al. (2015); Ward (2017); Salmond et al. (2012); Biraud et al. (2021)
Air temperature	Weather station, RTD, sondes	Validate air temperature, assess similarity functions (<i>PSIT</i> in <i>SLUCM</i> and <i>PSIH</i> in <i>BEP + BEM</i>) and estimate atmospheric stability (Ri_b)	Barlow (2014)
Surface temperature	IR cameras for individual surfaces and diurnal information, Landsat, MODIS and ECOSTress for global estimate at overpass time	Validate surface temperature	Jiménez-Muñoz et al. (2014); Hu et al. (2014)
Wind speed/direction	Anemometers, Doppler lidar, SODAR	Validate wind speed, assess similarity functions (<i>PSIX</i> in <i>SLUCM</i> and <i>PSIM</i> in <i>BEP + BEM</i>) and the surface transfer coefficients	Brotzge et al. (2020); Lamer et al. (2024)
Friction velocity	3D sonic anemometers, ECOR	Assess similarity functions	Raupach et al. (1991)
Rain rate	Disdrometers, rain gauges, radar	Trigger for different <i>BET</i> parameterizations in the <i>SLUCM</i>	Lamer et al. (2019); Uijlenhoet et al. (1997)
Soil/surface moisture	Neutron probes, dielectric sensors, SAR, pressure transducers	Assess <i>BET</i> parameterization in <i>SLUCM</i> and <i>wf</i> parameterization in <i>BEP + BEM</i>	S.U. et al. (2014); Rasheed et al. (2022); Mukhlisin et al. (2021); Mason et al. (2012); McDonald and Wang (2019)

Table 1 continued from previous page

Evapotranspiration (ET)	Indirect methods: Landsat-based Surface Energy Balance Algorithm for Land (SEBAL), improvements when including surface temperature information and, 3T-model	Assess <i>ETAR</i> parameterization in <i>SLUCM</i> , and <i>rsveg</i> parameterization in <i>BEP + BEM</i>	Faridatul et al. (2020); Qin et al. (2022); Qiu et al. (2023)
Leaf area index (LAI)	Photography, LAI sensors, lidar scans	In <i>rsveg</i> parameterization in <i>BEP + BEM</i> & in evapotranspiration (ETTR) in <i>SLUCM</i>	Chianucci et al. (2015)

4 Radiative Fluxes

4.1 Shortwave Radiation Parameterizations

BOX 3 The schematics shown in Figure 1 illustrate the shortwave radiation exchange between walls and the ground surface, as described by Equations 36-40. Likewise, Figure 2 depicts the radiation interactions between the ground and walls as formulated by the BEP scheme in Equations 41-43.

330 Parallels can be drawn between the SLUCM and BEP treatments of shortwave radiation (SWR) balance at surfaces. Both models employ the view factor formulation to represent the proportion of a surface visible from a particular point. Both models also consider SW reflections on the wall and ground surfaces, the SLUCM considers a single reflection, while the BEP model considers reflections at multiple levels. Both models have a simplified treatment of the roof surface, where shadows and reflections from walls are omitted.

In the SLUCM, equations for the SWR absorbed by each surface have a term for the direct (*SD*) and diffuse (*SQ*) components of
 335 SWR. The total solar SW radiation, provided by the WRF radiative transfer model, is divided into its direct and diffuse components using the SRATIO parameter. By default, SRATIO is set to 75%, assuming that 75% of the solar SW radiation is direct and 25% is diffuse. This formulation oversimplifies several factors influencing the direct-to-diffuse radiation ratio, such as cloud cover, aerosols, and atmospheric conditions, which exhibit significant diurnal and spatio-temporal variability. Observations could help provide guidance as to the true variability of SRATIO and could be used to formulate a parameterization for this parameter (see Section 4.3 for additional discussion).

340 In the SLUCM, the SWR absorbed by the rooftop surface (SR1) and green rooftop surface (SGR1) takes the form:

$$\underbrace{[SR1, SGR1]}_{\substack{\text{Line [921,922]} \\ \text{module_sf_urban.F}}} = \underbrace{SD(1 - [ALBR, ALBV])}_{\substack{\text{T1: Direct SWR from the sky} \\ \text{absorbed by the roof}}} + \underbrace{SQ(1 - [ALBR, ALBV])}_{\substack{\text{T2: Diffuse SWR from the sky} \\ \text{absorbed by the roof}}}, \quad (36)$$

where ALBR being the roof albedo, user-defined and specified in the *urb_param_table*, and ALBV the albedo of the green roof which is set to 0.2 in the *module_sf_urban.F*.

For surfaces within the urban canyon (ground and wall), the SLUCM model can account for shading effects - if this option is turned on
 345 using the *shadow_flag* parameter - and for a single reflection. This gives rise to more terms in equations describing the SWR absorbed

by these surfaces. For the ground:

$$\underbrace{SG1}_{\substack{\text{Line 923} \\ \text{module_sf_urban.F}}} = \underbrace{SD \frac{RW - SLX}{RW} (1 - ALBG)}_{\text{T1: Direct SWR from the sky on the lit portion of the ground absorbed by the ground}} + \underbrace{SQ \times VFGS (1 - ALBG)}_{\text{T2: Diffuse SWR from the sky seen \& absorbed by the ground}}, \quad (37)$$

$$\underbrace{SG2}_{\substack{\text{Line 925} \\ \text{module_sf_urban.F}}} = \underbrace{SD \frac{SLX}{W} \times ALBB \times VFGW (1 - ALBG)}_{\text{T1: Direct SWR from the sky incident on the lit portion of the wall reflected towards the ground and absorbed by the ground}} + \underbrace{SQ \times VFWS \times ALBB \times VFGW (1 - ALBG)}_{\text{T2: Diffuse SWR from the sky seen by the wall reflected towards the ground and absorbed by the ground}}, \quad (38)$$

350 with $ALBG$ being the ground albedo and $ALBB$ the wall albedo both user-defined and specified in the *urb_param_table*. The first terms in equation 37 and 38 depend on the ratio of the shadow length (SLX) to the normalized surface length or in other words the illuminated fraction of the surface. The normalized canyon width (RW) is employed in the estimation of SWR at the ground surface, while the normalized building height (W) is employed in the estimate of the SWR at the wall surface with the factor of two representing contributions from both walls ($2 \times$ normalized height). The shadow length (SLX) is influenced by several factors, including latitude, solar zenith angle, solar hour
355 angle, building height, street width, and street orientation, and is estimated between line 884 and 919 within the *module_sf_urban.F*. Most terms consider a view factor depending on the surface viewing the radiation. $VFGS$ represents how the ground views the sun, $VFGW$ represents how the ground views the wall and $VFWS$ represents how the wall views the sun.

For the wall surface, the SLUCM further considers one reflection between the East and West wall, which gives rise to two additional terms in the equation for SWR absorbed by the wall surface (SB), one for the direct and one for the diffuse component of SWR:

$$360 \quad \underbrace{SB1}_{\substack{\text{Line 924} \\ \text{module_sf_urban.F}}} = \underbrace{SD \frac{SLX}{W} (1 - ALBB)}_{\text{T1: Direct SWR from the sky incident on the lit portion of the wall absorbed by the wall}} + \underbrace{SQ \times VFWS (1 - ALBB)}_{\text{T2: Diffuse SWR from the sky, seen \& absorbed by the wall}}, \quad (39)$$

$$\begin{aligned} \underbrace{SB2}_{\substack{\text{Line 926} \\ \text{module_sf_urban.F}}} &= \underbrace{SD \frac{RW - SLX}{RW} \times ALBG \times VFWG (1 - ALBB)}_{\text{T1: Direct SWR from the sky incident on the lit portion of the ground reflected towards the wall \& absorbed by the wall}} + \underbrace{SQ \times VFGS \times ALBG \times VFWG (1 - ALBB)}_{\text{T2: Diffuse SWR from the sky seen by the ground reflected towards the wall \& absorbed by the wall}} + \\ &\quad \underbrace{SD \frac{SLX}{W} ALBB \times VFWW (1 - ALBB)}_{\text{T3: Direct SWR from the sky incident on the lit portion of the wall reflected towards the other wall absorbed by the wall: added to WRFv4.6.1}} + \underbrace{SQ \times VFWS \times ALBB \times VFWW (1 - ALBB)}_{\text{T4: Diffuse SWR from the sky seen by the wall reflected towards the other wall absorbed by that wall: added to WRFv4.6.1}}, \end{aligned} \quad (40)$$

where $VFWW$ represents the view factor from one wall to the other.

Unlike the SLUCM, the BEP model estimates the SWR incident on its surfaces at each urban grid level by solving a system of equations simultaneously, factoring in the probability of buildings being equal to height of a given level ($pb(z_{ju+1})$). This concept is implemented as
365 a system of equations in the form $[A]X = [B]$. Here, $[A]$ represents a $(2 \cdot nz_u + 1) \times (2 \cdot nz_u + 1)$ coefficient matrix for the SWR flux at the east and west walls and the ground (rsu^W , rsu^E , and rsu^g), nz_u is the number of urban vertical layers. The $[B]$ array with a size of $(2 \cdot nz_u + 1) \times 1$ includes the downward SWR from the sky incident on the walls and the ground. To facilitate explanation, this system of equations is decomposed into its components. The subscripts $s \rightarrow iu$, $s \rightarrow ju$, and $s \rightarrow g$ are used to indicate the direction of radiation
370 transport: from the sky to the iu^{th} -element of the wall, the sky to the ju^{th} -element of the opposite wall, the ju^{th} -element of the opposite wall to the ground, and the sky to the ground, respectively. Here, iu and ju denotes the iu^{th} -element (or level) of the current wall and the ju^{th} -element (or level) of the opposite wall, respectively.

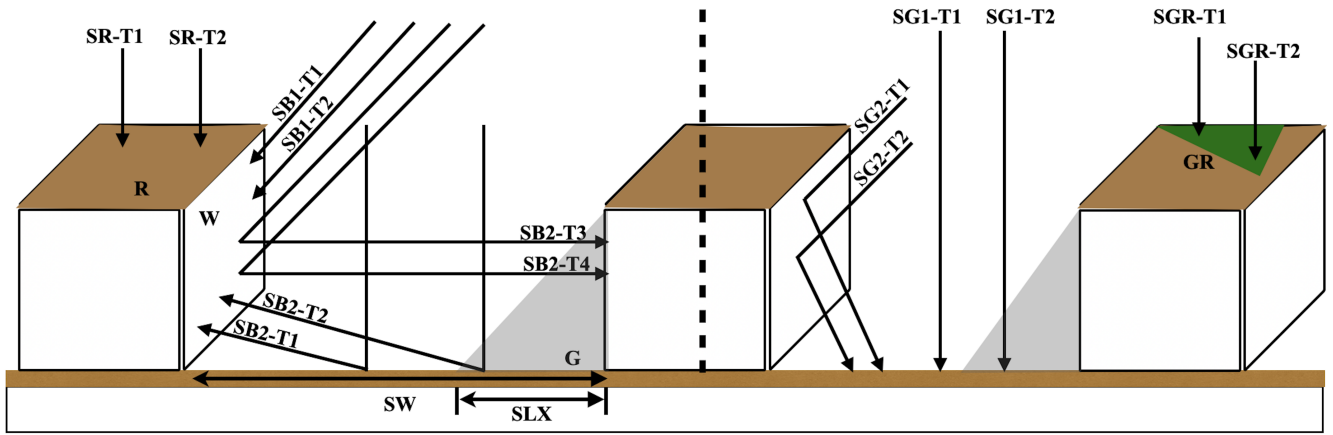


Figure 1. Schematic representation of SWR exchange in the SLUCM model. Depicted are the components of the SWR absorbed by the roof (*SR*), green roof (*SGR*), ground (*SG*), and wall (*SB*) surfaces in the SLUCM. The components are labeled to refer to the terms in the equations within this subsection; for instance, SG-T1 refers to the first term of the first equation describing the SWR absorbed by the ground surface in the SLUCM (i.e., equation 46). The gray triangle in the figure represent the shadow cast on the street and wall surfaces. The green triangle represents the green roof.

Another key distinction between the SLUCM and the BEP model is that BEP computes the incident (not the absorbed) SWR on each surface. The incident SWR is converted into the absorbed SW radiation for each surface by considering their albedo during the estimation of each surface's temperature. BEP accounts for the impact of shadows on the direct SWR in an independent step before estimating the incident SWR balance for each surface. The computation of direct SWR incident on the illuminated portion of the ground ($rs_{g \rightarrow g}$) and of the West wall ($rs_{s \rightarrow iu}^W$) of the East wall ($rs_{s \rightarrow iu}^E$) considers the canyon orientation:

$$\underbrace{rs_{s \rightarrow iu}^W}_{\text{Line 3301 module_sf_bep_bem.F}} = swddir \times rd \times ss/pb, \quad (41)$$

$$\underbrace{rs_{g \rightarrow g}}_{\text{Line 3311 module_sf_bep_bem.F}} = swddir * rd * ss/wsd, \quad (42)$$

where $swddir$ is the downward direct shortwave radiation, ss is the probability of the buildings with height equal to a given urban grid level, pb is the probability of the buildings with height equal to and taller than the given urban grid level, wsd represent the street width divided by the sine of the angle between direction of the sun and the face of the wall that accounts for the canyon orientation, rd is the ratio of portion of the wall lit to the horizontal surface illuminated area, that account for the shadow effect. Ultimately all orientations are averaged over the cell, considering the areas occupied by canyons with different orientations.

The SW radiation incident on the ground surface (rs_g) is comprised of three components (see sup eq. S3) while the SW radiation budget incorporates an additional term to account for SW radiation reflected from the opposite wall onto the wall of interest. Below, the formulation for the SW radiation budget of the West wall (rs_{iu}^W) is presented; the East wall follows an identical formulation with the indices [E, W]

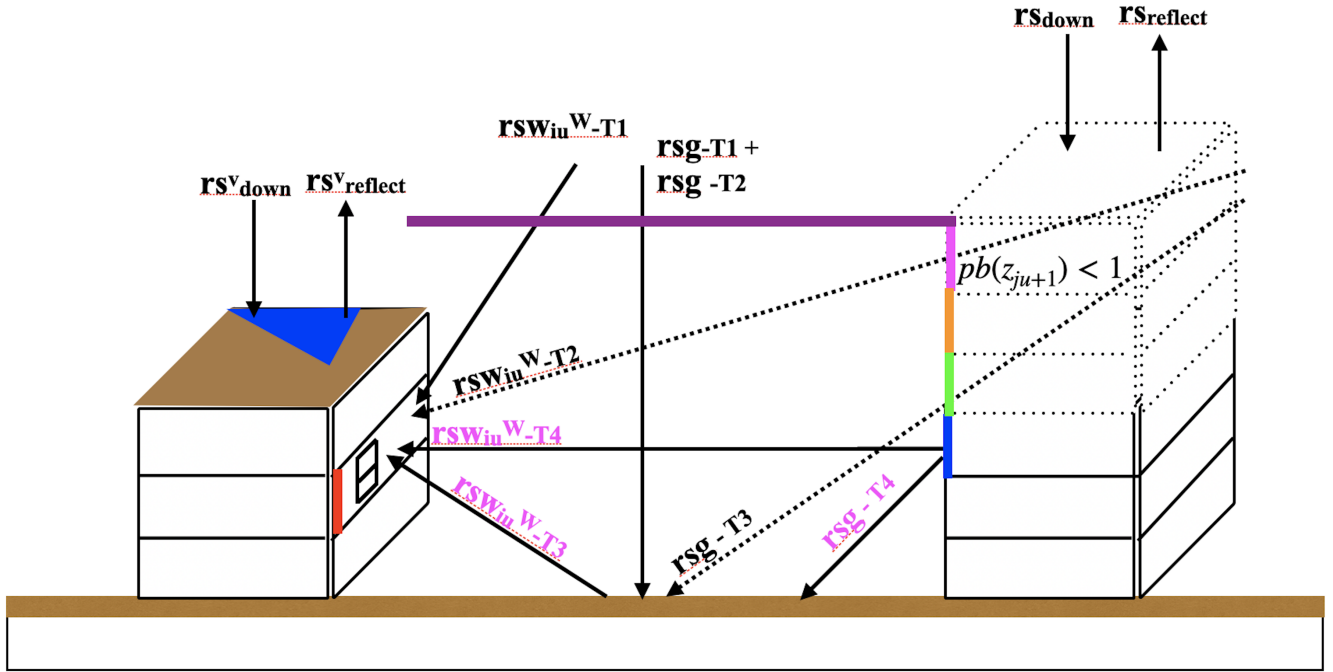


Figure 2. Schematic representation of SWR exchange in the BEP model. Depicted are the components of the incident SWR on the roof (rs), photovoltaic panel (rsv), ground (rsg) and wall (rsw) surfaces. The components are labeled to refer to the terms in the equations within this subsection; for instance, rsg -T1 refers to the first term of the first equation describing the SWR incident on the ground surface in the BEP (i.e., equation 43). Dashed building floors represent the probabilistic approach used to account for radiation not blocked by buildings (i.e., $(pb(z) < 1)$). Purple labels refer to the components that use the total surface SWR rather than the sky SWR

interchanged:

$$\begin{aligned}
 \underbrace{rsw_{iu}^W}_{\text{Line 3652, 3656 module_sf_bep_bem.F}} &= \underbrace{rsw_{s \rightarrow iu}^W}_{\text{T1: SWR incident on the illuminated portion of the west wall at level } iu \text{ considering the effect of shadows}} + \underbrace{rsdif_{s \rightarrow iu}^W \times fsw}_{\text{T2: Diffuse SWR from the sky to the } iu \text{ level of the west wall}} + \underbrace{rsg \times albg \times fgw}_{\text{T3: Total SWR at the ground reflected towards the wall}} + \\
 &\quad \underbrace{\sum_{ju=1, nu} rsw_{ju}^E \times pb(z_{ju+1}) albw \times fww}_{\text{T4: Total SWR at the east wall reflected towards the west wall by the portion of buildings at that level}}, \quad (43)
 \end{aligned}$$

with fgw denoting the the view factors of the wall looking at the ground, fww the view factor between the wall and the opposite wall, fsw is the view factor of the wall looking at the sky, fsg is the view factor of the ground looking at the sky, and $albg$ is the albedo of the ground surface. Notably, the BEP (or BEP-BEM) formulation uses fewer terms than the SLUCM approach because it employs the total SW radiation at the building wall to estimate the SW radiation reaching the ground, rather than recalculating each component of wall SW radiation (both direct and diffuse).

4.2 Longwave Radiation Parameterizations

BOX 4 Similarly to the shortwave radiation section, the schematics in Figure 3 correspond to Equations 44–46, illustrating the longwave radiation exchange between the walls and the ground surface, whereas Figure 4 depicts the longwave radiation interactions between the ground and walls represented by the BEP scheme, as formulated in Equation 47.

Similar to the formulation for the SWR balance on each surface, the formulation for the longwave radiation (LWR) balance on each surface in the SLUCM and BEP employs the view factor formulation to represent the proportion of a surface visible from a particular point. Both models also consider LWR reflections on the wall and ground surfaces, the SLUCM considers a single reflection, while the BEP model considers several reflections at multiple levels. Both models have a simplified treatment of the roof surface, where reflections from walls are omitted.

In the SLUCM, this approach leads to the following formulation of the net LWR for the roof surface (RR):

$$\underbrace{RR}_{\text{Line 1026 module_sf_urban.F}} = EPSR \left(\underbrace{RX}_{\text{T1: LWR from sky to roof}} - \underbrace{SIG \times TRP^4}_{\text{T2: LWR from roof to sky}} \right), \quad (44)$$

where RX represents the total downward LWR from the sky, which is supplied to the SLUCM model by the WRF atmospheric radiation scheme. The emissivity of the roof surface $EPSR$ and of the ground surface $EPSG$ and of the wall $EPSB$ is specified by the user in the *urb_param_table*. The outgoing LWR flux (RR - $T2$) is influenced by the roof surface temperature (TRP) from the previous time step and involves the Stefan-Boltzmann constant (SIG).

The net LWR at the ground surface (RG) has additional terms to account for LWR emitted from the wall surface towards the ground surface ($RG1$ - $T2$) as well as for reflected LWR reaching the ground surface ($RG2$) (see sup. eq. S4 and S5) while the net LWR at the wall (RB) has additional terms to account for LWR emitted by and reflected from the opposite wall. All LWR reaching the wall from reflection are contained in $RB2$:

$$\underbrace{RB1}_{\text{Line 1288 module_sf_urban.F}} = EPSB \left(\underbrace{RX \times VFWS}_{\text{T1: LWR from sky seen by wall}} + \underbrace{EPSG \times SIG \times TGP^4 \times VFWG}_{\text{T2: LWR from ground seen by wall}} + \underbrace{EPSB \times SIG \times TBP^4 \times VFWW}_{\text{T3: LWR from opposite wall seen by wall}} - \underbrace{SIG \times TBP^4}_{\text{T4: LWR from wall out}} \right). \quad (45)$$

$$\begin{aligned} \underbrace{RB2}_{\text{Line 1297 module_sf_urban.F}} = & EPSB \left(\underbrace{RX \times VFGS(1 - EPSG)VFWG}_{\text{T1: LWR from sky seen by ground reflected towards wall}} + \underbrace{EPSB \times SIG \times TBP^4 \times VFGW(1 - EPSG)VFWG}_{\text{T2: LWR from wall seen by ground reflected towards wall}} + \right. \\ & \underbrace{RX \times VFWS(1 - EPSB)VFWW}_{\text{T3: LWR from sky seen by opposite wall reflected towards wall}} + \underbrace{EPSG \times SIG \times TGP^4 \times VFWG(1 - EPSB)VFWW}_{\text{T4: LWR from ground seen by opposite wall reflected towards wall}} + \\ & \left. \underbrace{EPSB \times SIG \times TBP^4 \times VFWW(1 - EPSB)VFWW}_{\text{T5: LWR from opposite wall seen by wall reflected towards opposite wall}} \right), \end{aligned} \quad (46)$$

As it does to estimate the SWR incident on its various surface, the BEP (or BEP-BEM) model estimates the incident LWR at its various surfaces at its different levels using a system of equations, factoring in the probability of buildings being equal to height of a given level.

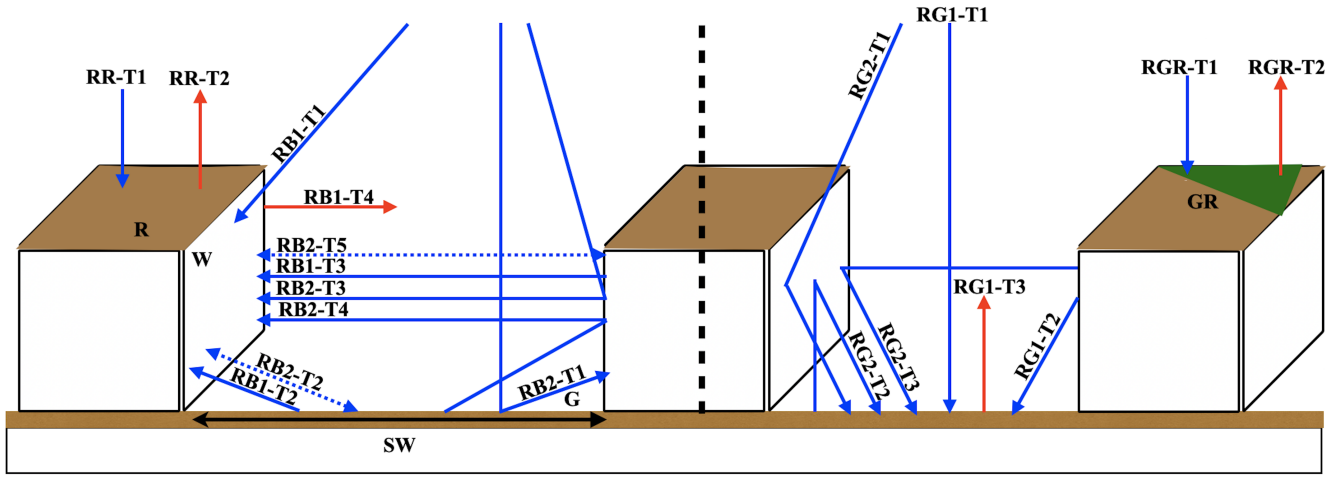


Figure 3. Schematic representation of LWR exchange in the SLUCM model. Depicted are the components of the LWR balance at the roof (*RR*), green roof (*RGR*), ground (*RG*) and wall (*RB*) surfaces. The components are labeled to refer to the terms in the equations within this subsection; for instance, *RG1-T1* refers to the first term of the first equation describing the net LWR at the ground surface in the SLUCM (i.e., equation 42). Blue arrows represent incoming LWR, while red arrows indicate outgoing LWR. Dashed double-sided arrows depict radiation leaving a surface and being reflected back to it. The green triangle represents a green roof.

Another key distinction between the SLUCM and the BEP model is that BEP computes the incident (not the net) LWR on each surface. The incident LWR is converted into the net LW radiation for each surface by considering their emissivity during the estimation of each surface's temperature.

The BEP-BEM module further considers the presence of windows on building walls. Windows are treated as another surface, with their own temperature tw_{lev} , and emissivity em_{win} and are set to cover a fraction of the wall p_{win} . The emissivity of the wall is denoted by em_w while the wall temperature is given by tw . The parameter sigma ($= 5.67 \times 10^{-8}$) is the Stefan-Boltzmann constant.

As a result, the incident LWR at the ground surface (rlg) is comprised of four components, including a weighted average of the contribution of the wall and window surface (see sup. eq. S6) and the incident LWR at the west wall surface (rlw^W) is comprised of two more components to account for the effect of the East wall:

$$\begin{aligned}
 \underbrace{rlw_{iu}^W}_{\text{Lines: 3535, 3539 module_sf_bep_bem.F}} &= \underbrace{rl \times f_{sw}}_{\text{T1: LWR from sky seen by west wall}} + \underbrace{\sum_{ju=1, nu} rl(1 - pb(z_{ju+1}))fw}_{\text{T2: Portion of sky LWR that passed through east walls that is seen by west wall}} + \underbrace{emg \times sigma \times tg_{av}^4 \times fgw}_{\text{T3: LWR emitted by ground seen by west wall}} + \underbrace{rlg(1 - emg)fgw}_{\text{T4: Total LWR at the ground reflected to west wall}} + \\
 &\quad \underbrace{\sum_{ju=1, nu} sigma \times (emw(1 - p_{win})tw^{4,E} + em_{win} \times p_{win} \times tw_{lev}^{4,E})pb(z_{ju+1})fw}_{\text{T5: LWR emitted by east wall that is seen by west wall}} + \\
 &\quad \underbrace{\sum_{ju=1, nu} rlw^E \times pb(z_{ju+1})(1 - emw(1 - p_{win}) - em_{win} * p_{win})fw}_{\text{T6: Total LWR at the east wall reflected to west wall}},
 \end{aligned} \tag{47}$$

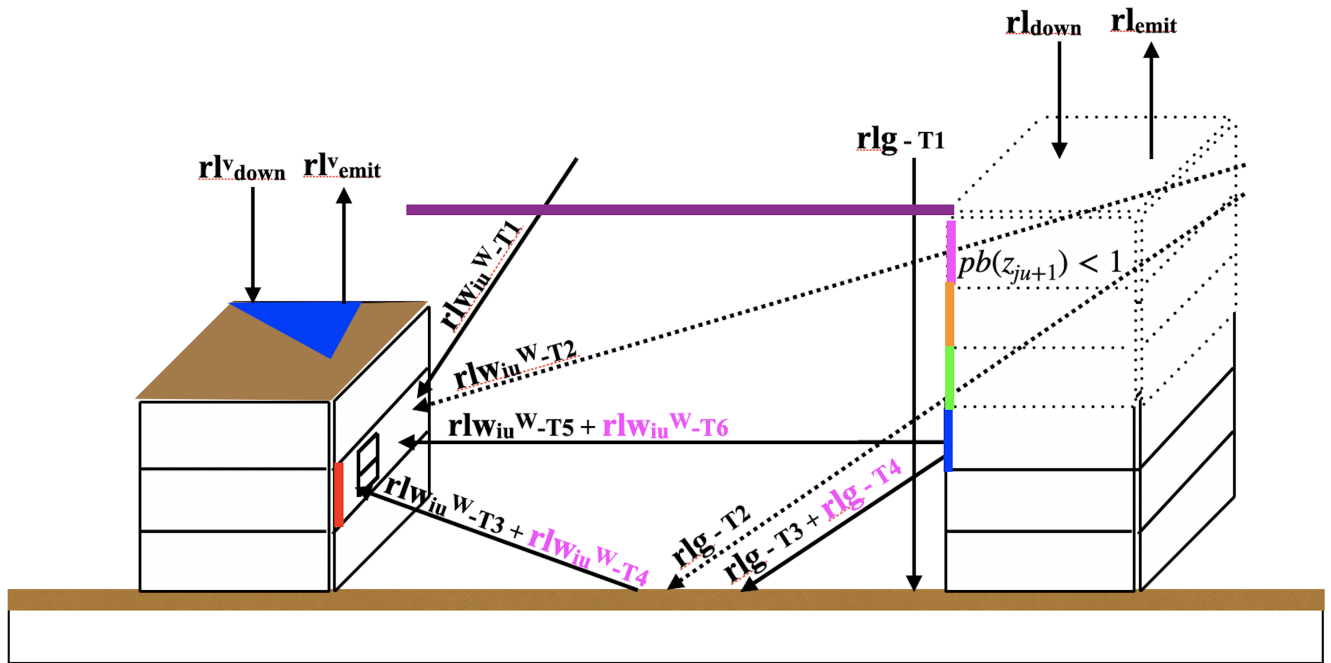


Figure 4. Schematic representation of LWR exchange in the BEP model. Depicted are the components of the incident SWR on the roof (rl), photovoltaic (PV) panel (rlv), ground (rlg) and wall (rlw) surfaces. The components are labeled to refer to the terms in the equations within this subsection; for instance, $rlg-T1$ refers to the first term of the first equation describing the LWR incident on the ground surface in the BEP (i.e., equation 46). Dashed building floors represent the probabilistic approach used to account for radiation not blocked by buildings (i.e., $pb(z) < 1$). The blue triangle represent a PV panel.

where t_g is the temperature of the ground surface in the previous time step. A similar equation exists for the east wall, where the "W" and "E" superscripts are interchanged. The discussions on radiative exchange between surfaces in SLUCM and BEP clearly demonstrate that the added complexity of BEP and MLUCM in modeling radiative transfer may offer greater benefits in urban environments with mixed building heights and complex geometries. However, the trade-off between model realism and computational cost should be carefully considered when setting up simulations for a specific city. A systematic evaluation of model performance versus complexity can help guide the choice of urban scheme.

4.3 Observational approaches

Observations can play a crucial role in evaluating the net SWR and LWR simulated for various surface types in WRF-Urban. Furthermore, they can be used to assess their parameterizations via measurements of shadow length and view factor. Observations could also be used to evaluate the impact of omitting reflections and shading on roof surfaces and the use of a fixed ratio (SRATIO) to separate solar radiation into its direct and diffuse components in the SLUCM.

Strategically positioned radiometers would allow us to measure incident and reflected SWR on various urban surfaces (Aoki and Mizutani (2015)). For wall surfaces, mounting a pyranometers facing outward from the wall, aligned with the wall's vertical plane, would provide a

measurement of direct sunlight plus diffuse sky radiation (Oliveira et al. (2024)). Another pyranometer, positioned at the same height but facing toward the wall, would measure radiation reflected from the wall. A shading structure can be employed to ensure that the pyranometer facing the wall captures only reflected radiation by blocking direct sunlight. To improve accuracy, the pyranometer should be positioned a short distance (10–20 cm) away from the wall to avoid obstruction effects. Similarly, using a pyrgeometer in this setup allows for the measurement of incident and outgoing LWR on the wall surface.

Radiometers are once again useful to measure the ratio of direct to diffuse solar radiation (Kassianov et al. (2024)). One common approach involves using a pyranometer equipped with a shadow-band to obstruct direct sunlight, thereby isolating diffuse irradiance measurements. However, this method requires correction factors to account for the portion of the sky occluded by the shadow-band. Another technique employs a pyrliometer in conjunction with a pyranometer (Wang et al. (2013)). The pyrliometer, aligned with the sun, measures direct normal irradiance, while the pyranometer captures global horizontal irradiance. By subtracting the direct component (adjusted for the solar zenith angle) from the global measurement, it is possible to derive diffuse horizontal irradiance. Additionally, rotating shadowband radiometers (RSRs) have been developed to automate the simultaneous measurement of global, direct, and diffuse irradiance components (Alexandrov et al. (2002)). These instruments utilize a rotating shadowband to intermittently shade the sensor, facilitating the differentiation between direct and diffuse radiation. The efficacy of RSRs in measuring diffuse solar irradiance has been evaluated in various studies, highlighting their utility in solar radiation research. Each of these methodologies offers distinct advantages and considerations, with the choice of technique often influenced by the specific requirements of the study and the available instrumentation.

Estimating shadow length in urban street canyons can be achieved through photographic methods with reference scaling. This approach involves capturing images of the area and using known dimensions within the image to scale and measure shadow lengths accurately.

Regarding the estimation of SVF, several methods have been developed and compared in the literature. Geometric methods involve calculating SVF based on the dimensions and arrangement of urban structures. Fish-eye photographic methods use hemispherical images to determine the proportion of visible sky. Global Positioning System (GPS) methods estimate SVF by analyzing positional data. Simulation methods utilize 3D city models or digital surface models to compute SVF. Big data approaches leverage street view images and deep learning algorithms to estimate SVF across large urban areas. A comprehensive review of these methods, including their principles, input data requirements, application contexts, accuracy, and efficiency, is provided by Miao et al. (2020).

5 Summary and Conclusions

Despite their broad application and widespread user base, the urban schemes within the WRF physics suite remain comparatively less documented, especially when it comes to the specific implementation of the parameterizations. This article serves two primary purposes.

First, it improves the transparency of the Bulk, Single Layer Urban Canopy Model (SLUCM) and Building Effect Parameterization coupled with Building Energy Model (BEM) model treatment of turbulent and radiative fluxes by creating a faithful roadmap to the code version v4.5.2 with matching symbols and code line number references. We bring the specific code implementation details back to the physics by providing descriptions and schematics to improve code accessibility to non-expert users. This systematic review of the code uncovered several implementation errors, which have since been addressed and corrected in subsequent versions of WRF (v4.6.0 and v4.6.1; see the appendix).

Second, this work narrowed down to the subset of parameters that can be evaluated using the state-of-the-art observation technique to improve the model performance. We identify several threshold-based parameterizations that can introduce discontinuities in simulations: (i) a rain rate threshold in SLUCM that controls minimum moisture availability and in turn affect latent heat flux, necessitating validation under

rainfall conditions above and below 1 mm/hr (Miao and Chen (2014)); (ii) an assumption that all surfaces are dry in BEP, thus leading to zero latent heat flux, which could lead to model errors especially during and after precipitation events, (iii) SWR partitioning in SLUCM using a fixed direct-to-diffuse ratio, which in turn affects surface air temperature, requiring validation under both clear and cloudy skies; and (iv) the use of bulk Richardson number in all three UCMs to select similarity functions, which impact fluxes and the vertical structure of temperature and wind, warranting model evaluation at transitions between atmospheric stability regimes. The wide use of Monin-Obukov similarity theory for flux calculations in the WRF urban models may also be a source of uncertainty particularly in highly stable and highly unstable atmospheric stability regimes and in environments with strong surface heterogeneity on the scale of the model resolution (i.e., highly variable land use land cover or building height and material at sub-km scale), warranting model evaluation in cities presenting dense building layouts with mixtures of high-rise and low-rise within the same neighborhood (e.g., New York City). We also identify a highly simplified treatment of the radiative balance at roof surfaces in all urban schemes, where the effects of shading and reflections from nearby walls are neglected. The implications of these simplifications can be assessed through targeted observations to evaluate their impact on simulation accuracy.

We offer guidance on how observational data can be used to evaluate and refine model assumptions, such as measurements of surface moisture availability, the direct-to-diffuse radiation ratio, and vertical profiles of air temperature and wind. We encourage the use of multi-parameter observations to identify compensating errors. We recommend collecting data at multiple heights, particularly at the model's lowest atmospheric level, to avoid confounding effects from diagnostic extrapolation schemes. For example, past campaigns employed a combination of tower-based and remote sensing measurements (e.g., the Basel UrBan Boundary Layer Experiment (BUBBLE) Rotach et al. (2005)). Given challenges with positioning instrumentation in cities, we are seeing the emergence of mobile platforms, including drones and truck-based remote sensing instrumentation in several modern field campaigns, as demonstrated in field campaigns such as the DOE Southwest Urban Field Laboratory (Southwest Integrated Field Laboratory (2024)) and *Urbisphere* (Fenner et al. (2024)).

Although this study has gone into details about model parameterization/schemes and touched on the topic of observational spatio-temporal coverage and comprehensiveness, other factors also impact model performance. These include boundary conditions (e.g., inflow winds), geographic input data (e.g., topography, land use), model setup choices (e.g., grid resolution), and uncertainties in measurement retrieval methods. These aspects warrant further investigation in future work.

Code availability. The Weather Research and Forecasting (WRF) model is freely available online and can be downloaded from: <https://github.com/wrf-model/WRF/releases>.

Data availability. ERA5 Reanalysis (0.25 Degree Latitude-Longitude Grid). Research Data Archive at the National Center for Atmospheric Research, Computational and Information Systems Laboratory. <https://doi.org/10.5065/BH6N-5N20>. The data produced in this study are available at: <https://doi.org/10.5281/zenodo.15191965> (Lin et al. (2025))

Appendix A

A1 Code Fixes and Model Sensitivity Outcomes

This study uncovered several inconsistencies between the urban parameterizations described in the literature and their implementation in WRF-Urban version 4.5.2. These inconsistencies - or also known as code bugs - are now documented in the WRF community forum as pull 2049, 2038, 2032, and 2101 (<https://github.com/wrf-model/WRF/pull/>) and were corrected in subsequent releases (beginning with v4.6.1). To assess the impact of these bug fixes on simulation outcomes, we conducted a one-week WRF sensitivity experiment centered over Atlanta, GA initialized from 2 August 2015, 0Z with the first 24-hour spin-up, capturing both clear-sky and precipitation conditions. Figure A1 illustrates the simulation setup, where panel (a) shows the outer and inner domains with their respective resolutions (5km and 1km), and panel (b) displays the spatial distribution of urban fraction.

Bug fixes were grouped in two sets, (i) in Building Energy Model (BEM) where we corrected the coefficients used in calculating the heat transfer coefficient for thermal exchange between the roof surface, PV panels, and the atmosphere using the empirical relation, (ii) in SLUCM where we corrected the stability parameter, Lech's surface functions for momentum and heat, added missing terms in short-wave radiation balance for wall, and used adequate surface temperature for outgoing LWR flux .

To demonstrate the impact of these bug fixes we run two sets of simulation, one without the bug fix and one with the bug fix, keeping all other physics and settings the same including the use of: Thompson microphysics scheme (Thompson and Eidhammer (2014)), Mellor-Yamada-Janjic (MYJ) planetary boundary layer scheme (Janjić (1994)), Rapid Radiative Transfer Model (RRTMG) shortwave and longwave radiation scheme (Iacono et al. (2008)), and the Noah-MP land surface model ((Yang et al., 2011; Niu et al., 2011; He et al., 2023)). For the BEP-BEM simulations, we employ a photovoltaic panels cover of 50 percent, while in the SLUCM simulations we employ a green roof cover of 50 percent.

The results demonstrate that even minor implementation issues can meaningfully influence model output. Fixes to the BEM led to changes in both 2 m air temperature (Figure A2(b)) and skin temperature on the order of 0.3 K (Figure A2(d)) both in urban grids and non-urban grids . Fixes to the SLUCM produced smaller but still notable effects, with temperature changes up to 0.15 K (Figure A3(b) and A3(d)).

Author contributions. KL: Conceptualization, funding acquisition, verification, supervision, writing - review and editing. PJ: Formal analysis, visualization, writing - original draft preparation, TL: Data curation, software, visualization, writing - review and editing. CH: Writing - verification, supervision, review and editing

Competing interests. One of the authors is a member of the editorial board of Geoscientific Model Development. The authors declare that they have no other competing interests.

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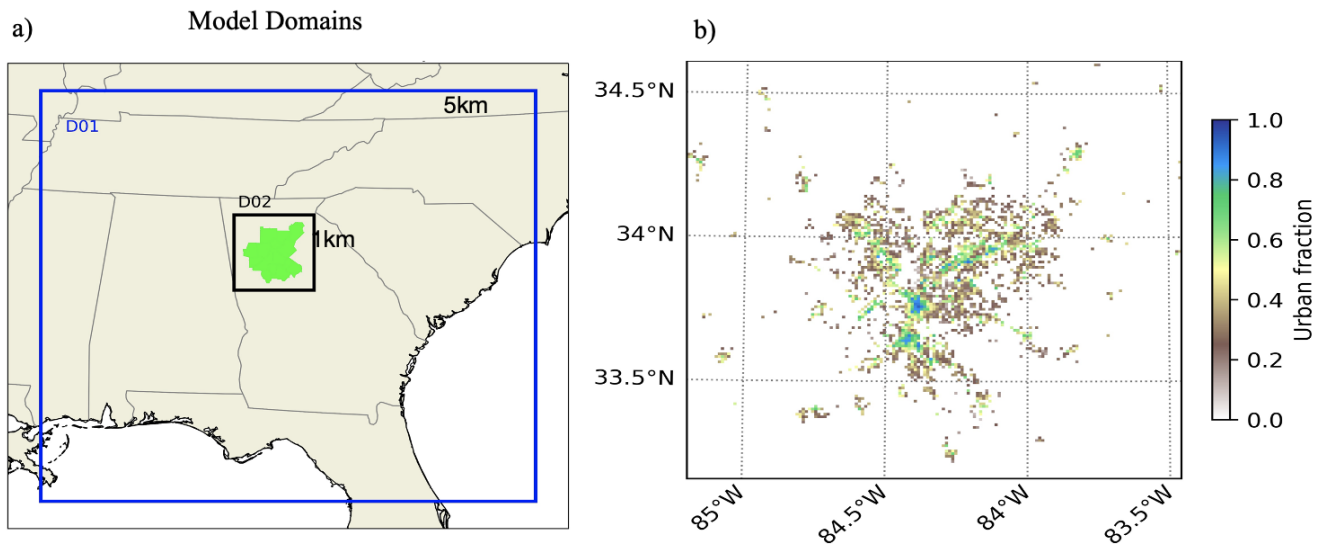


Figure A1. Panel (a) shows the outer (D01) and inner (D02) domains, with horizontal grid resolutions annotated for the WRF test simulations. Panel (b) displays the urban fraction contours within the D02 domain.

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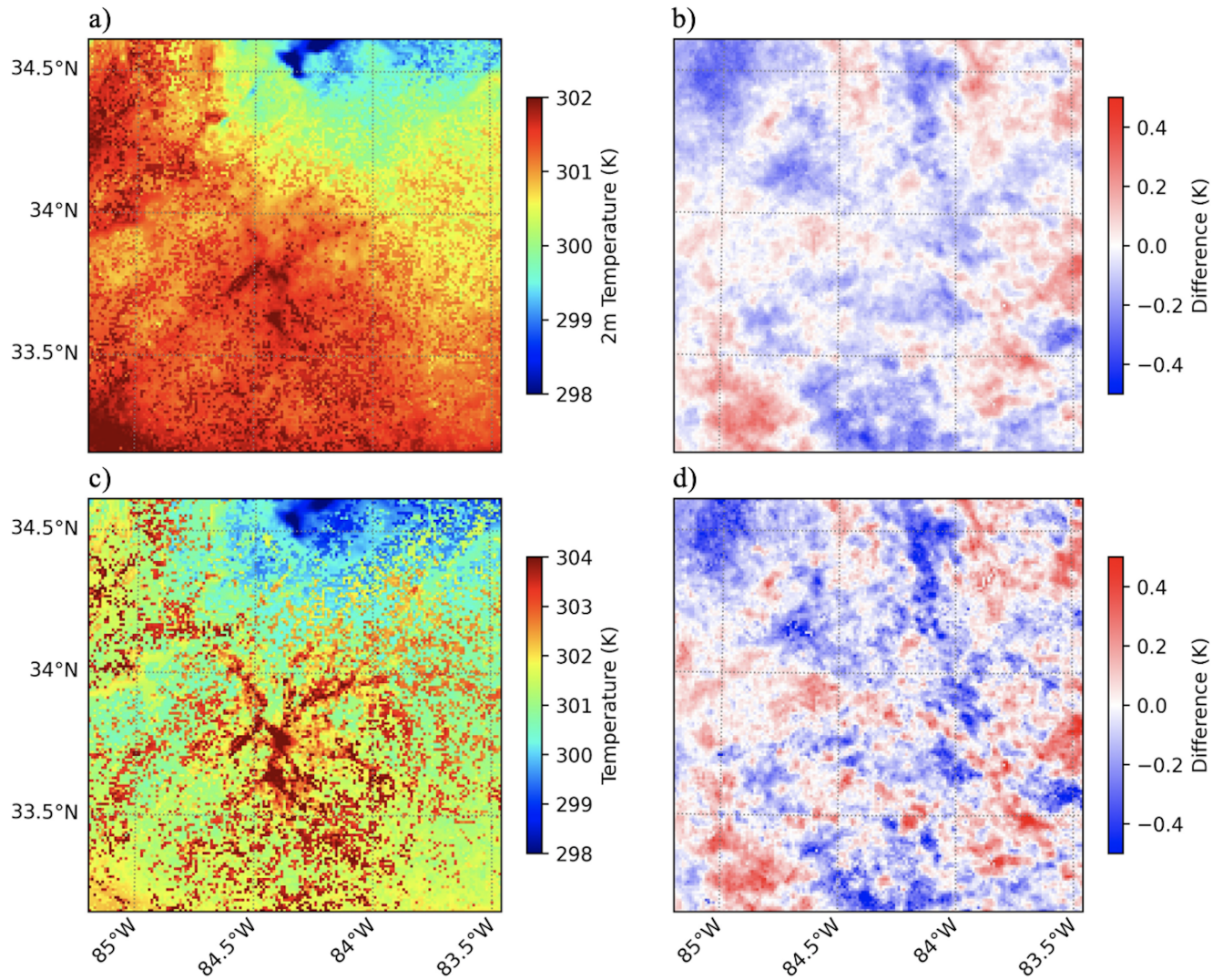


Figure A2. Panels (a) and (b) depict the average 2 m air temperature without bug fix and the 2 m air temperature difference between the cases studies with and without bug fixes implemented, respectively. Similarly, panels (c) and (d) represent the skin surface temperature output from WRF-BEP+BEM model without bug fixes implemented and the skin surface temperature difference between the case studies with and without bug fix. Panels (a)-(d) corresponds with the WRF model output for the BEP+BEM urban canopy model.

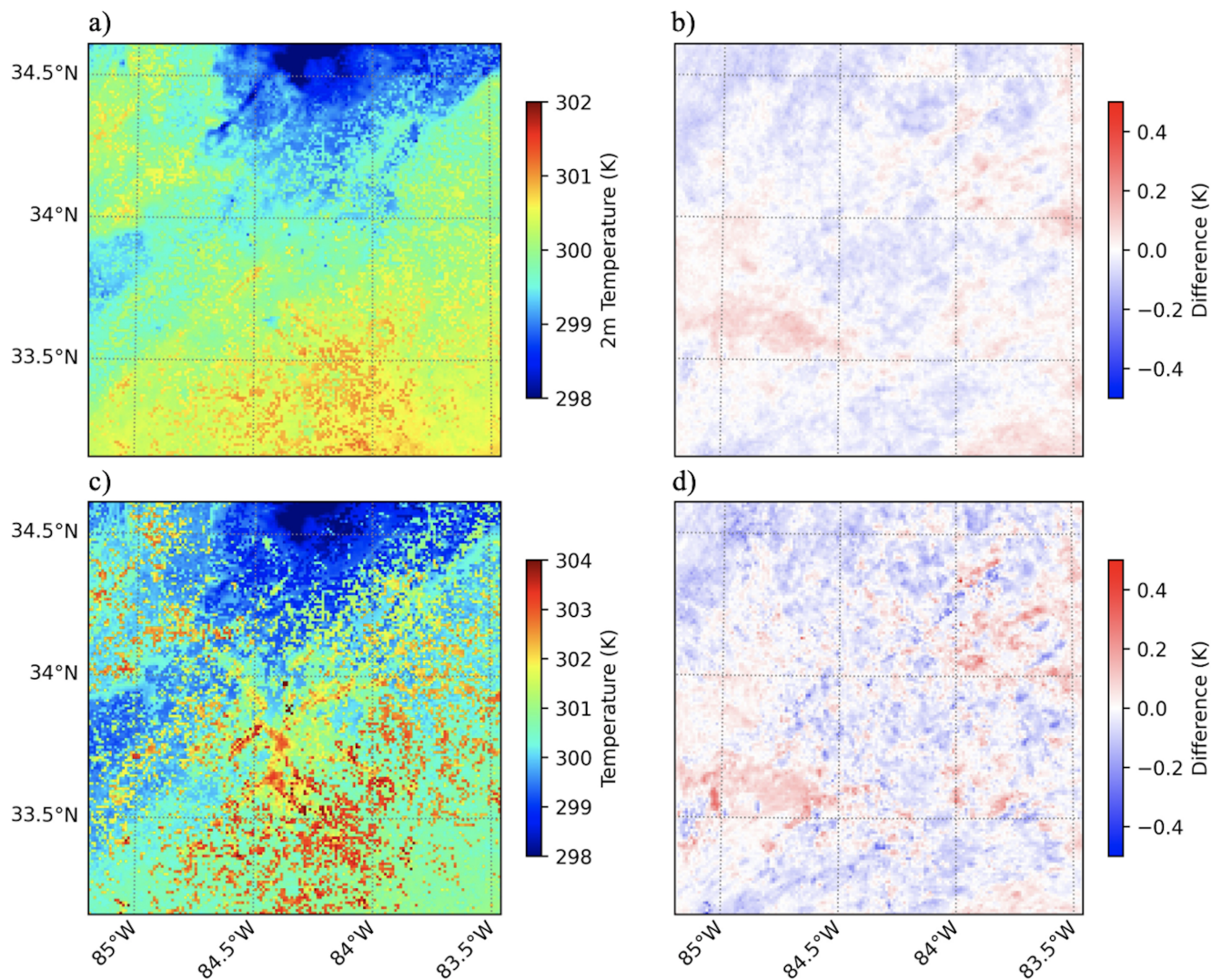


Figure A3. Panels (a) and (b) depict the average 2 m air temperature without bug fix and the 2 m air temperature difference between the cases studies with and without bug fixes implemented, respectively. Similarly, panels (c) and (d) represent the skin surface temperature output from WRF-SLUCM model without bug fixes implemented and the skin surface temperature difference between the case studies with and without bug fix. Panels (a)-(d) corresponds with the WRF model output for the SLUCM urban canopy model.

References

- Alexandrov, M. D., Lacis, A. A., Carlson, B. E., and Cairns, B.: Remote Sensing of Atmospheric Aerosols and Trace Gases by Means of Multifilter Rotating Shadowband Radiometer. Part I: Retrieval Algorithm, *Journal of the Atmospheric Sciences*, 59, 524 – 543, [https://doi.org/10.1175/1520-0469\(2002\)059<0524:RSOAAA>2.0.CO;2](https://doi.org/10.1175/1520-0469(2002)059<0524:RSOAAA>2.0.CO;2), 2002.
- 545 Allen, R. G., Pruitt, W. O., Wright, J. L., Howell, T. A., Ventura, F., Snyder, R., Itenfisu, D., Steduto, P., Berengena, J., Yrisarry, J. B., Smith, M., Pereira, L. S., Raes, D., Perrier, A., Alves, I., Walter, I., and Elliott, R.: A recommendation on standardized surface resistance for hourly calculation of reference ETo by the FAO56 Penman-Monteith method, *Agricultural Water Management*, 81, 1–22, <https://doi.org/10.1016/j.agwat.2005.03.007>, 2006.
- Aoki, T. and Mizutani, A.: Measurement of the Vertical Distribution of Reflected Solar Radiation, *Journal of Engineering and Technological Sciences*, 47, 160–169, <https://doi.org/10.5614/j.eng.technol.sci.2015.47.2.5>, 2015.
- 550 Barlow, J. F.: Progress in observing and modelling the urban boundary layer, *Urban Climate*, 10, 216–240, <https://doi.org/10.1016/j.uclim.2014.03.011>, iCUC8: The 8th International Conference on Urban Climate and the 10th Symposium on the Urban Environment, 2014.
- Beljaars, A. C. M.: The parametrization of surface fluxes in large-scale models under free convection, *Quarterly Journal of the Royal Meteorological Society*, 121, 255–270, <https://doi.org/10.1002/qj.49712152203>, 1995.
- 555 Beljaars, A. C. M. and Holtslag, A. A. M.: Flux Parameterization over Land Surfaces for Atmospheric Models, *Journal of Applied Meteorology and Climatology*, 30, 327 – 341, [https://doi.org/10.1175/1520-0450\(1991\)030<0327:FPOLSF>2.0.CO;2](https://doi.org/10.1175/1520-0450(1991)030<0327:FPOLSF>2.0.CO;2), 1991.
- Biraud, S., Chen, J., Christen, A., Davis, K., Lin, J., McFadden, J., Miller, C., Nemitz, E., Schade, G., Stagakis, S., Turnbull, J., and Vogt, R.: Eddy Covariance Measurements in Urban Environments: White Paper Prepared by the AmeriFlux Urban Fluxes Ad Hoc Committee, <https://ameriflux.lbl.gov/wp-content/uploads/2021/09/EC-in-Urban-Environment-2021-07-31-Final.pdf>, 2021.
- 560 Brotzge, J. A., Wang, J., Thorncroft, C. D., Joseph, E., Bain, N., Bassill, N., Farruggio, N., Freedman, J. M., Hemker, K., Johnston, D., Kane, E., McKim, S., Miller, S. D., Minder, J. R., Naple, P., Perez, S., Schwab, J. J., Schwab, M. J., and Sicker, J.: A Technical Overview of the New York State Mesonet Standard Network, *Journal of Atmospheric and Oceanic Technology*, 37, 1827 – 1845, <https://doi.org/10.1175/JTECH-D-19-0220.1>, 2020.
- 565 Bruno, R., Ferraro, V., Bevilacqua, P., and Arcuri, N.: On the assessment of the heat transfer coefficients on building components: A comparison between modeled and experimental data, *Building and Environment*, 216, 108 995, <https://doi.org/10.1016/j.buildenv.2022.108995>, 2022.
- Businger, J. A., Wyngaard, J. C., Izumi, Y., and Bradley, E. F.: Flux-Profile Relationships in the Atmospheric Surface Layer, *Journal of Atmospheric Sciences*, 28, 181 – 189, [https://doi.org/10.1175/1520-0469\(1971\)028<0181:FPRITA>2.0.CO;2](https://doi.org/10.1175/1520-0469(1971)028<0181:FPRITA>2.0.CO;2), 1971.
- 570 Chen, F. and Dudhia, J.: Coupling an Advanced Land Surface–Hydrology Model with the Penn State–NCAR MM5 Modeling System. Part I: Model Implementation and Sensitivity, *Monthly Weather Review*, 129, 569 – 585, [https://doi.org/10.1175/1520-0493\(2001\)129<0569:CAALSH>2.0.CO;2](https://doi.org/10.1175/1520-0493(2001)129<0569:CAALSH>2.0.CO;2), 2001.
- Chen, F. and Zhang, Y.: On the coupling strength between the land surface and the atmosphere: From viewpoint of surface exchange coefficients, *Geophysical Research Letters*, 36, <https://doi.org/10.1029/2009GL037980>, 2009.
- 575 Chen, F., Kusaka, H., Bornstein, R., Ching, J., Grimmond, C. S. B., Grossman-Clarke, S., Loridan, T., Manning, K. W., Martilli, A., Miao, S., Sailor, D., Salamanca, F. P., Taha, H., Tewari, M., Wang, X., Wyszogrodzki, A. A., and Zhang, C.: The integrated WRF/urban modelling

- system: development, evaluation, and applications to urban environmental problems, *International Journal of Climatology*, 31, 273–288, <https://doi.org/10.1002/joc.2158>, 2011.
- Chen, T. H., Henderson-Sellers, A., Milly, P. C. D., Pitman, A. J., Beljaars, A. C. M., Polcher, J., Abramopoulos, F., Boone, A., Chang, S.,
580 Chen, F., Dai, Y., Desborough, C. E., Dickinson, R. E., Dümenil, L., Ek, M., Garratt, J. R., Gedney, N., Gusev, Y. M., Kim, J., Koster, R., Kowalczyk, E. A., Laval, K., Lean, J., Lettenmaier, D., Liang, X., Mahfouf, J.-F., Mengelkamp, H.-T., Mitchell, K., Nasonova, O. N., Noilhan, J., Robock, A., Rosenzweig, C., Schaake, J., Schlosser, C. A., Schulz, J.-P., Shao, Y., Shmakin, A. B., Verseghy, D. L., Wetzel, P., Wood, E. F., Xue, Y., Yang, Z.-L., and Zeng, Q.: Cabauw Experimental Results from the Project for Intercomparison of Land-Surface Parameterization Schemes, *Journal of Climate*, 10, 1194 – 1215, [https://doi.org/10.1175/1520-0442\(1997\)010<1194:CERFTP>2.0.CO;2](https://doi.org/10.1175/1520-0442(1997)010<1194:CERFTP>2.0.CO;2),
585 1997.
- Chenge, Y. and Brutsaert, W.: Flux-profile Relationships for Wind Speed and Temperature in the Stable Atmospheric Boundary Layer, *Boundary-Layer Meteorology*, 114, 519–538, <https://doi.org/10.1007/s10546-004-1425-4>, 2005.
- Chianucci, F., Puletti, N., Giacomello, E., Cutini, A., and Corona, P.: Estimation of leaf area index in isolated trees with digital photography and its application to urban forestry, *Urban Forestry Urban Greening*, 14, 377–382, <https://doi.org/10.1016/j.ufug.2015.04.001>, 2015.
- 590 Dyer, A. J.: A review of flux-profile relationships, *Boundary-Layer Meteorology*, 7, 363–372, <https://doi.org/10.1007/BF00240838>, 1974.
- Fairall, C. W., Bradley, E. F., Rogers, D. P., Edson, J. B., and Young, G. S.: Bulk parameterization of air-sea fluxes for Tropical Ocean-Global Atmosphere Coupled-Ocean Atmosphere Response Experiment, *Journal of Geophysical Research: Oceans*, 101, 3747–3764, <https://doi.org/10.1029/95JC03205>, 1996.
- Faridatul, M. I., Wu, B., Zhu, X., and Wang, S.: Improving remote sensing based evapotranspiration modelling in a heterogeneous urban
595 environment, *Journal of Hydrology*, 581, 124 405, <https://doi.org/10.1016/j.jhydrol.2019.124405>, 2020.
- Fenner, D., Christen, A., Grimmond, S., Meier, F., Morrison, W., Zeeman, M., Barlow, J., Birkmann, J., Blunn, L., Chrysoulakis, N., Clements, M., Glazer, R., Hertwig, D., Kotthaus, S., König, K., Looschelders, D., Mitraka, Z., Poursanidis, D., Tsirantonakis, D., Bechtel, B., Benjamin, K., Beyrich, F., Briegel, F., Feigel, G., Gertsen, C., Iqbal, N., Kittner, J., Lean, H., Liu, Y., Luo, Z., McGrory, M., Metzger, S., Paskin, M., Ravan, M., Ruhtz, T., Saunders, B., Scherer, D., Smith, S. T., Stretton, M., Trachte, K., and Hove, M. V.: urbisphere-Berlin
600 Campaign: Investigating Multiscale Urban Impacts on the Atmospheric Boundary Layer, *Bulletin of the American Meteorological Society*, 105, E1929 – E1961, <https://doi.org/10.1175/BAMS-D-23-0030.1>, 2024.
- Garuma, G. F.: Review of urban surface parameterizations for numerical climate models, *Urban Climate*, 24, 830–851, <https://doi.org/10.1016/j.uclim.2017.10.006>, 2018.
- Georgescu, M., Broadbent, A. M., and Krayenhoff, E. S.: Quantifying the decrease in heat exposure through adaptation and mitigation in
605 twenty-first-century US cities, *Nature Cities*, 1, 42–50, <https://doi.org/10.1038/s44284-023-00001-9>, 2024.
- Gochis, D. J., Yu, W., and Yates, D. N.: The WRF-Hydro model technical description and user's guide, version 3.0., NCAR Technical Document, p. 120 pages, http://www.ral.ucar.edu/projects/wrf_hydro/, 2015.
- Grell, G. A., Peckham, S. E., Schmitz, R., McKeen, S. A., Frost, G., Skamarock, W. C., and Eder, B.: Fully coupled “online” chemistry within the WRF model, *Atmospheric Environment*, 39, 6957–6975, <https://doi.org/10.1016/j.atmosenv.2005.04.027>, 2005.
- 610 Grimmond, C. S. B., Blackett, M., Best, M. J., Barlow, J., Baik, J.-J., Belcher, S. E., Bohnenstengel, S. I., Calmet, I., Chen, F., Dandou, A., Fortuniak, K., Gouvea, M. L., Hamdi, R., Hendry, M., Kawai, T., Kawamoto, Y., Kondo, H., Krayenhoff, E. S., Lee, S.-H., Loridan, T., Martilli, A., Masson, V., Miao, S., Oleson, K., Pigeon, G., Porson, A., Ryu, Y.-H., Salamanca, F., Shashua-Bar, L., Steeneveld, G.-J., Tombrou, M., Voogt, J., Young, D., and Zhang, N.: The International Urban Energy Balance Models Comparison Project: First Results from Phase 1, *Journal of Applied Meteorology and Climatology*, 49, 1268 – 1292, <https://doi.org/10.1175/2010JAMC2354.1>, 2010.

- 615 Hang, J., Zeng, L., Li, X., and Wang, D.: Evaluation of a single-layer urban energy balance model using measured energy fluxes by scaled out-door experiments in humid subtropical climate, *Building and Environment*, 254, 111 364, <https://doi.org/10.1016/j.buildenv.2024.111364>, 2024.
- He, C., Valayamkunnath, P., Barlage, M., Chen, F., Gochis, D., Cabell, R., Schneider, T., Rasmussen, R., Niu, G.-Y., Yang, Z.-L., Niyogi, D., and Ek, M.: Modernizing the open-source community Noah-MP land surface model (version 5.0) with enhanced modularity, interoperability, and applicability, *EGUsphere*, 2023, 1–31, <https://doi.org/10.5194/egusphere-2023-675>, 2023.
- 620 Hicks, B. B.: Wind profile relationships from the ‘wangara’ experiment, *Quarterly Journal of the Royal Meteorological Society*, 102, 535–551, <https://doi.org/10.1002/qj.49710243304>, 1976.
- Hu, L., Brunsell, N. A., Monaghan, A. J., Barlage, M., and Wilhelmi, O. V.: How can we use MODIS land surface temperature to validate long-term urban model simulations?, *Journal of Geophysical Research: Atmospheres*, 119, 3185–3201, <https://doi.org/10.1002/2013JD021101>, 2014.
- 625 Iacono, M. J., Delamere, J. S., Mlawer, E. J., Shephard, M. W., Clough, S. A., and Collins, W. D.: Radiative forcing by long-lived greenhouse gases: Calculations with the AER radiative transfer models, *Journal of Geophysical Research: Atmospheres*, 113, <https://doi.org/10.1029/2008JD009944>, 2008.
- Inoue, E.: On the Turbulent Structure of Airflow within, *Journal of the Meteorological Society of Japan. Ser. II*, 41, 317–326, https://doi.org/10.2151/jmsj1923.41.6_317, 1963.
- 630 Izumi, Y. and Barad, M. L.: Wind Speeds as Measured by Cup and Sonic Anemometers and Influenced by Tower Structure, *Journal of Applied Meteorology and Climatology*, 9, 851 – 856, [https://doi.org/10.1175/1520-0450\(1970\)009<0851:WSAMBC>2.0.CO;2](https://doi.org/10.1175/1520-0450(1970)009<0851:WSAMBC>2.0.CO;2), 1970.
- Jacquemin, B. and Noilhan, J.: Sensitivity study and validation of a land surface parameterization using the HAPEX-MOBILHY data set, *Boundary-Layer Meteorology*, 52, 93–134, <https://doi.org/10.1007/BF00123180>, 1990.
- 635 Janjić, Z. I.: The Step-Mountain Eta Coordinate Model: Further Developments of the Convection, Viscous Sublayer, and Turbulence Closure Schemes, *Monthly Weather Review*, 122, 927 – 945, [https://doi.org/10.1175/1520-0493\(1994\)122<0927:TSMECM>2.0.CO;2](https://doi.org/10.1175/1520-0493(1994)122<0927:TSMECM>2.0.CO;2), 1994.
- Jiménez, P. A., Dudhia, J., González-Rouco, J. F., Navarro, J., Montávez, J. P., and García-Bustamante, E.: A Revised Scheme for the WRF Surface Layer Formulation, *Monthly Weather Review*, 140, 898 – 918, <https://doi.org/10.1175/MWR-D-11-00056.1>, 2012.
- Jiménez-Muñoz, J. C., Sobrino, J. A., Skoković, D., Mattar, C., and Cristóbal, J.: Land Surface Temperature Retrieval Methods From Landsat-8 Thermal Infrared Sensor Data, *IEEE Geoscience and Remote Sensing Letters*, 11, 1840–1843, <https://doi.org/10.1109/LGRS.2014.2312032>, 2014.
- 640 Jürges, W.: Der Wärmeübergang an einer ebenen Wand, Druck und Verlag von R. Oldenbourg, https://books.google.com/books/about/Der_W%C3%A4rme%C3%BCbergang_an_einer_ebenen_Wand.html?id=31cEvAEACAAJ, 1924.
- Kanda, M., Kanega, M., Kawai, T., Moriwaki, R., and Sugawara, H.: Roughness Lengths for Momentum and Heat Derived from Outdoor Urban Scale Models, *Journal of Applied Meteorology and Climatology*, 46, 1067 – 1079, <https://doi.org/10.1175/JAM2500.1>, 2007.
- 645 Kassianov, E., Flynn, C. J., Barnard, J. C., Ermold, B. D., and Comstock, J. M.: Shortwave Array Spectroradiometer-Hemispheric (SAS-He): design and evaluation, *Atmospheric Measurement Techniques*, 17, 4997–5013, <https://doi.org/10.5194/amt-17-4997-2024>, 2024.
- Kusaka, H., Kondo, H., Kikegawa, Y., and Kimura, F.: A Simple Single-Layer Urban Canopy Model For Atmospheric Models: Comparison With Multi-Layer And Slab Models, *Boundary-Layer Meteorology*, 101, 329–358, <https://doi.org/10.1023/A:1019207923078>, 2001.
- 650 Lamer, K., Puigdomènech Treserras, B., Zhu, Z., Isom, B., Bharadwaj, N., and Kollias, P.: Characterization of shallow oceanic precipitation using profiling and scanning radar observations at the Eastern North Atlantic ARM observatory, *Atmospheric Measurement Techniques*, 12, 4931–4947, <https://doi.org/10.5194/amt-12-4931-2019>, 2019.

- Lamer, K., Mages, Z., Treserras, B. P., Walter, P., Zhu, Z., Rapp, A. D., Nowotarski, C. J., Brooks, S. D., Flynn, J., Sharma, M., Klein, P., Spencer, M., Smith, E., Gebauer, J., Bell, T., Bunting, L., Griggs, T., Wagner, T. J., and McKeown, K.: Spatially distributed atmospheric boundary layer properties in Houston – A value-added observational dataset, *Scientific Data*, 11, 661, <https://doi.org/10.1038/s41597-024-03477-9>, 2024.
- Lee, S.-H., Kim, S.-W., Angevine, W. M., Bianco, L., McKeen, S. A., Senff, C. J., Trainer, M., Tucker, S. C., and Zamora, R. J.: Evaluation of urban surface parameterizations in the WRF model using measurements during the Texas Air Quality Study 2006 field campaign, *Atmospheric Chemistry and Physics*, 11, 2127–2143, <https://doi.org/10.5194/acp-11-2127-2011>, 2011.
- Lin, T.-S., Joshi, P., He, C., and Lamer, K.: WRF urban canopy models simulations dataset, <https://doi.org/10.5281/zenodo.15191965>, 2025.
- Liu, Y., Chen, F., Warner, T., and Basara, J.: Verification of a Mesoscale Data-Assimilation and Forecasting System for the Oklahoma City Area during the Joint Urban 2003 Field Project, *Journal of Applied Meteorology and Climatology*, 45, 912 – 929, <https://doi.org/10.1175/JAM2383.1>, 2006.
- Louis, J.-F.: A parametric model of vertical eddy fluxes in the atmosphere, *Boundary-Layer Meteorology*, 17, 187–202, <https://doi.org/10.1007/BF00117978>, 1979.
- Lynn, B. and Yair, Y.: Prediction of lightning flash density with the WRF model, *Advances in Geosciences*, 23, 11–16, <https://doi.org/10.5194/adgeo-23-11-2010>, 2010.
- Macdonald, R., Griffiths, R., and Hall, D.: An improved method for the estimation of surface roughness of obstacle arrays, *Atmospheric Environment*, 32, 1857–1864, [https://doi.org/10.1016/S1352-2310\(97\)00403-2](https://doi.org/10.1016/S1352-2310(97)00403-2), 1998.
- Mahrt, L. T. and Sun, J.: The Subgrid Velocity Scale in the Bulk Aerodynamic Relationship for Spatially Averaged Scalar Fluxes, *Monthly Weather Review*, 123, 3032 – 3041, [https://doi.org/10.1175/1520-0493\(1995\)123<3032:TSVSIT>2.0.CO;2](https://doi.org/10.1175/1520-0493(1995)123<3032:TSVSIT>2.0.CO;2), 1995.
- Martilli, A., Clappier, A., and Rotach, M. W.: An Urban Surface Exchange Parameterisation for Mesoscale Models, *Boundary-Layer Meteorology*, 104, 261–304, <https://doi.org/10.1023/A:1016099921195>, 2002.
- Mason, D. C., Davenport, I. J., Neal, J. C., Schumann, G. J.-P., and Bates, P. D.: Near Real-Time Flood Detection in Urban and Rural Areas Using High-Resolution Synthetic Aperture Radar Images, *IEEE Transactions on Geoscience and Remote Sensing*, 50, 3041–3052, <https://doi.org/10.1109/TGRS.2011.2178030>, 2012.
- Masson, V.: Urban surface modeling and the meso-scale impact of cities, *Theoretical and Applied Climatology*, 84, 35–45, <https://doi.org/10.1007/s00704-005-0142-3>, 2006.
- McAdams, W. H.: Heat Transmission, McGraw-Hill, https://books.google.com/books/about/Heat_Transmission.html?id=1ApRAAAAMAAJ, 1954.
- McDonald, S. and Wang, Z.: Real-Time Optimal Trajectory Generation for UAV to Rendezvous with an Aerial Orbit, *American Institute of Aeronautics and Astronautics*, <https://doi.org/10.2514/6.2019-3620>, 2019.
- Miao, C., Yu, S., Hu, Y., Zhang, H., He, X., and Chen, W.: Review of methods used to estimate the sky view factor in urban street canyons, *Building and Environment*, 168, 106497, <https://doi.org/10.1016/j.buildenv.2019.106497>, 2020.
- Miao, S. and Chen, F.: Enhanced modeling of latent heat flux from urban surfaces in the Noah/single-layer urban canopy coupled model, *Science China Earth Sciences*, 57, 2408–2416, <https://doi.org/10.1007/s11430-014-4829-0>, 2014.
- Moriwaki, R. and Kanda, M.: Flux-gradient profiles for momentum and heat over an urban surface, *Theoretical and Applied Climatology*, 84, 127–135, <https://doi.org/10.1007/s00704-005-0150-3>, 2006.

- Mukhlisin, M., Yunanto, B., Suharjono, A., Martono, M., Aprianoro, R., and Kusumawardani, R.: Monitoring rainfall intensity and moisture water content using soil column experiment, *IOP Conference Series: Materials Science and Engineering*, 1108, 012007, <https://doi.org/10.1088/1757-899X/1108/1/012007>, 2021.
- Niu, G.-Y., Yang, Z.-L., Mitchell, K. E., Chen, F., Ek, M. B., Barlage, M., Kumar, A., Manning, K., Niyogi, D., Rosero, E., Tewari, M., and Xia, Y.: The community Noah land surface model with multiparameterization options (Noah-MP): 1. Model description and evaluation with local-scale measurements, *Journal of Geophysical Research: Atmospheres*, 116, <https://doi.org/10.1029/2010JD015139>, 2011.
- Oliveira, M., Silva Lopes, H., Mendonça, P., Tenpierik, M., and Silva, L. T.: Solar Radiation Measurement Tools and Their Impact on In Situ Testing—A Portuguese Case Study, *Buildings*, 14, <https://doi.org/10.3390/buildings14072117>, 2024.
- Paulson, C. A.: The Mathematical Representation of Wind Speed and Temperature Profiles in the Unstable Atmospheric Surface Layer, *Journal of Applied Meteorology and Climatology*, 9, 857 – 861, [https://doi.org/10.1175/1520-0450\(1970\)009<0857:TMROWS>2.0.CO;2](https://doi.org/10.1175/1520-0450(1970)009<0857:TMROWS>2.0.CO;2), 1970.
- Powers, J. G., Klemp, J. B., Skamarock, W. C., Davis, C. A., Dudhia, J., Gill, D. O., Coen, J. L., Gochis, D. J., Ahmadov, R., Peckham, S. E., Grell, G. A., Michalakes, J., Trahan, S., Benjamin, S. G., Alexander, C. R., Dimego, G. J., Wang, W., Schwartz, C. S., Romine, G. S., Liu, Z., Snyder, C., Chen, F., Barlage, M. J., Yu, W., and Duda, M. G.: The Weather Research and Forecasting Model: Overview, System Efforts, and Future Directions, *Bulletin of the American Meteorological Society*, 98, 1717 – 1737, <https://doi.org/10.1175/BAMS-D-15-00308.1>, 2017.
- Qin, L., Yan, C., Yu, L., Chai, M., Wang, B., Hayat, M., Shi, Z., Gao, H., Jiang, X., Xiong, B., Mao, P., and Qiu, G. Y.: High-resolution spatio-temporal characteristics of urban evapotranspiration measured by unmanned aerial vehicle and infrared remote sensing, *Building and Environment*, 222, 109389, <https://doi.org/10.1016/j.buildenv.2022.109389>, 2022.
- Qiu, G. Y., Yan, C., and Liu, Y.: Urban evapotranspiration and its effects on water budget and energy balance: Review and perspectives, *Earth-Science Reviews*, 246, 104577, <https://doi.org/10.1016/j.earscirev.2023.104577>, 2023.
- Rasheed, M. W., Tang, J., Sarwar, A., Shah, S., Saddique, N., Khan, M. U., Imran Khan, M., Nawaz, S., Shamshiri, R. R., Aziz, M., and Sultan, M.: Soil Moisture Measuring Techniques and Factors Affecting the Moisture Dynamics: A Comprehensive Review, *Sustainability*, 14, <https://doi.org/10.3390/su141811538>, 2022.
- Rasmussen, R. M., Chen, F., Liu, C., Ikeda, K., Prein, A., Kim, J., Schneider, T., Dai, A., Gochis, D., Dugger, A., Zhang, Y., Jaye, A., Dudhia, J., He, C., Harrold, M., Xue, L., Chen, S., Newman, A., Dougherty, E., Abolafia-Rosenzweig, R., Lybarger, N. D., Viger, R., Lesmes, D., Skalak, K., Brakebill, J., Cline, D., Dunne, K., Rasmussen, K., and Miguez-Macho, G.: CONUS404: The NCAR-USGS 4-km Long-Term Regional Hydroclimate Reanalysis over the CONUS, *Bulletin of the American Meteorological Society*, 104, E1382 – E1408, <https://doi.org/10.1175/BAMS-D-21-0326.1>, 2023.
- Raupach, M. R., Antonia, R. A., and Rajagopalan, S.: Rough-Wall Turbulent Boundary Layers, *Applied Mechanics Reviews*, 44, 1–25, <https://doi.org/10.1115/1.3119492>, 1991.
- Rotach, M. W.: Turbulence close to a rough urban surface part I: Reynolds stress, *Boundary-Layer Meteorology*, 65, 1–28, <https://doi.org/10.1007/BF00708816>, 1993.
- Rotach, M. W., Vogt, R., Bernhofer, C., Batchvarova, E., Christen, A., Clappier, A., Feddersen, B., Gryning, S.-E., Martucci, G., Mayer, H., Mitev, V., Oke, T. R., Parlow, E., Richner, H., Roth, M., Roulet, Y.-A., Ruffieux, D., Salmond, J. A., Schatzmann, M., and Voogt, J. A.: BUBBLE – an Urban Boundary Layer Meteorology Project, *Theoretical and Applied Climatology*, 81, 231–261, <https://doi.org/10.1007/s00704-004-0117-9>, 2005.

- Salamanca, F. and Martilli, A.: A new Building Energy Model coupled with an Urban Canopy Parameterization for urban climate simulations—part II. Validation with one dimension off-line simulations, *Theoretical and Applied Climatology*, 99, 345–356, <https://doi.org/10.1007/s00704-009-0143-8>, 2010.
- Salmond, J. A., Roth, M., Oke, T. R., Christen, A., and Voogt, J. A.: Can Surface-Cover Tiles Be Summed to Give Neighborhood Fluxes in Cities?, *Journal of Applied Meteorology and Climatology*, 51, 133 – 149, <https://doi.org/10.1175/JAMC-D-11-078.1>, 2012.
- Southwest Integrated Field Laboratory: 2024 SW-IFL Annual Report, <https://sw-ifl.asu.edu/news>, 2024.
- S.U., S. L., Singh, D., and Shojaei Baghini, M.: A critical review of soil moisture measurement, *Measurement*, 54, 92–105, <https://doi.org/10.1016/j.measurement.2014.04.007>, 2014.
- Tarasova, M. A., Varentsov, M. I., and Stepanenko, V. M.: Parameterization of the Interaction between the Atmosphere and the Urban Surface: Current State and Prospects, *Izvestiya, Atmospheric and Oceanic Physics*, 59, 111–130, <https://doi.org/10.1134/S0001433823020068>, 2023.
- Thompson, G. and Eidhammer, T.: A Study of Aerosol Impacts on Clouds and Precipitation Development in a Large Winter Cyclone, *Journal of the Atmospheric Sciences*, 71, 3636 – 3658, <https://doi.org/10.1175/JAS-D-13-0305.1>, 2014.
- UCAR WRF Support Forum: Resolved: WRF Single-Layer UCM Crashed, <https://forum.mmm.ucar.edu/threads/resolved-wrf-single-layer-ucm-crashed.13453/>, accessed: 2024-07-16, 2024.
- Uijlenhoet, R., Stricker, J., and Russchenberg, H.: Application of X- and S-band radars for rain rate estimation over an urban area, *Physics and Chemistry of the Earth*, 22, 259–264, [https://doi.org/10.1016/S0079-1946\(97\)00141-9](https://doi.org/10.1016/S0079-1946(97)00141-9), hydrology, 1997.
- United Nations, Department of Economic and Social Affairs, Population Division: World Urbanization Prospects: The 2014 Revision, *ST/E-SA/SER.A/366*, 2015.
- Wang, K., Dickinson, R. E., Ma, Q., Augustine, J. A., and Wild, M.: Measurement Methods Affect the Observed Global Dimming and Brightening, *Journal of Climate*, 26, 4112 – 4120, <https://doi.org/10.1175/JCLI-D-12-00482.1>, 2013.
- Ward, H. C.: Scintillometry in urban and complex environments: a review, *Measurement Science and Technology*, 28, 064005, <https://doi.org/10.1088/1361-6501/aa5e85>, 2017.
- Ward, H. C., Evans, J. G., and Grimmond, C. S. B.: Infrared and millimetre-wave scintillometry in the suburban environment ndash; Part 2: Large-area sensible and latent heat fluxes, *Atmospheric Measurement Techniques*, 8, 1407–1424, <https://doi.org/10.5194/amt-8-1407-2015>, 2015.
- Yang, J. and Wang, Z.-H.: Physical parameterization and sensitivity of urban hydrological models: Application to green roof systems, *Building and Environment*, 75, 250–263, <https://doi.org/10.1016/j.buildenv.2014.02.006>, 2014.
- Yang, Z.-L., Niu, G.-Y., Mitchell, K. E., Chen, F., Ek, M. B., Barlage, M., Longuevergne, L., Manning, K., Niyogi, D., Tewari, M., and Xia, Y.: The community Noah land surface model with multiparameterization options (Noah-MP): 2. Evaluation over global river basins, *Journal of Geophysical Research: Atmospheres*, 116, <https://doi.org/10.1029/2010JD015140>, 2011.
- Zhang, D. and Anthes, R. A.: A High-Resolution Model of the Planetary Boundary Layer—Sensitivity Tests and Comparisons with SESAME-79 Data, *Journal of Applied Meteorology and Climatology*, 21, 1594 – 1609, [https://doi.org/10.1175/1520-0450\(1982\)021<1594:AHMOT>2.0.CO;2](https://doi.org/10.1175/1520-0450(1982)021<1594:AHMOT>2.0.CO;2), 1982.
- Zhu, D. and Ooka, R.: WRF-based scenario experiment research on urban heat island: A review, *Urban Climate*, 49, 101512, <https://doi.org/10.1016/j.uclim.2023.101512>, 2023.

- Zou, J., Liu, G., Sun, J., Zhang, H., and Yuan, R.: The momentum flux-gradient relations derived from field measurements in the urban roughness sublayer in three cities in China, *Journal of Geophysical Research: Atmospheres*, 120, 10,797–10,809, <https://doi.org/10.1002/2015JD023909>, 2015.
- 765 Łobocki, L.: A Procedure for the Derivation of Surface-Layer Bulk Relationships from Simplified Second-Order Closure Models, *Journal of Applied Meteorology and Climatology*, 32, 126 – 138, [https://doi.org/10.1175/1520-0450\(1993\)032<0126:APFTDO>2.0.CO;2](https://doi.org/10.1175/1520-0450(1993)032<0126:APFTDO>2.0.CO;2), 1993.