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Measurement report: Impact of domestic heating on dust
deposition sources in hyper-arid Qaidam Basin, northern
Qinghai-Xizang Plateau

Haixia Zhu^{abc}, Lufei Zhen^{abc}, Suping Zhao^{d*}, Xiying Zhang^{ab*}

^a Key Laboratory of Green and High-end Utilization of Salt Lake Resources, Qinghai Institute of Salt Lakes, Chinese Academy of Sciences, Xining, 810008, China.
^b Qinghai Provincial Key Laboratory of Geology and Environment of Salt Lakes, Qinghai Institute of Salt Lakes, Chinese Academy of Sciences, Xining, 810008, China.
^c University of the Chinese Academy of Sciences, Beijing, 100049, China
^d Key Laboratory of Cryospheric Science and Frozen Soil Engineering, Northwest Institute of Eco-Environment and Resources, Chinese Academy of Sciences, Lanzhou, 730000, China
** Corresponding author. E-mail address: xyzhchina@isl.ac.cn (Xiying Zhang) and zhaosp@lzb.ac.cn (Suping Zhao)*

23 **Highlights**

24 1. The temporal and spatial distribution of carbonaceous aerosols was analyzed using various carbon
25 indicators.

26 2. Domestic heating significantly increased atmospheric pollutants in rural areas.

27 3. The unique energy structure in Qaidam Basin significantly influenced the glaciers of the
28 Qinghai-Xizang Plateau and should not be overlooked.

29

30 Abstract

31 Given the unique energy profile of the Qaidam Basin (QDB), it is crucial to examine the impacts of
32 domestic heating on the Qinghai-Xizang Plateau (QXP) and global atmospheric systems. This
33 study collected monthly dust deposition at six sites in the southern QDB between 2020 and 2023.
34 We identified the sources of dust-fall during domestic heating (HP) and non-heating periods (NHP)
35 in urban and rural areas and its environmental effects. The results demonstrated that domestic
36 heating increased the concentration of water-soluble ions in rural areas, trace elements in urban
37 areas, and carbon emissions in both. Among various carbon indicators, organic carbon (OC) and
38 element carbon (EC) levels rose during the HP, with Char-EC as the primary component of EC
39 (80.44%). Char-EC concentrations were higher in urban areas (85.00%), while secondary organic
40 carbon (68.17%), the main contributor to OC, was more prevalent in rural (73.92%). The OC/EC
41 ratio in urban areas remained stable with an average of 2.16. In contrast, the rural OC/EC ratio was
42 significantly higher during the NHP (7.27 ± 4.66) than during the HP (4.57 ± 3.02). Additionally,
43 the char/soot ratio was elevated in the HP (5.06 ± 4.08) compared to the NHP (4.42 ± 3.09). The
44 OC/EC and char-EC/soot-EC ratios, along with PMF results, indicated that coal combustion
45 (17.28%) and biomass burning (32.50%) were the main contributors to dust deposition in rural areas,
46 strongly influenced by domestic heating, whereas urban dust predominantly originated from vehicle
47 traffic (44.43%) and industrial emissions (16.41%). Coal consumption in QDB was greater during
48 the HP than that of other dust sources in the QXP. This increased consumption leads to higher
49 emissions of atmospheric pollutants, which may accelerate glacier melting in the region.
50 Consequently, integrating QDB carbon aerosols into future environmental policies and climate
51 models for the QXP is essential. This study provides a reference for investigating carbonaceous
52 aerosols in climatically similar hyper-arid basins with intensive human activity and salt lake regions.

53 **Keywords:** Qinghai-Xizang Plateau; Qaidam Basin; Biomass burning; carbonaceous elements;
54 atmospheric dust deposition.

55

56 **Short summary**

57 This study collected dust samples from six sites in the Qaidam Basin, over three years to investigate
58 the impact of domestic heating on atmospheric dust in hyper-arid region. Our results indicate that
59 rural dust is significantly influenced by heating, particularly from coal_ and biomass burning,
60 which accounts for over 70% of total sources. The unique energy structure here has resulted in
61 distinct environmental effects from the emitted carbonaceous aerosols and useful for similar dry
62 areas.

63 1. Introduction

64 Atmospheric dust, a critical component of particulate matter (PM), serves as both an air quality
65 indicator and environmental stressor, influencing hydrological cycles and soil ecosystems (Feng et
66 al., 2019). Recent advancements in understanding PM characteristics—particularly chemical
67 composition (e.g. water-soluble ions, organic carbon, and elemental carbon) and source
68 apportionment—have been achieved through principal component analysis (PCA), chemical mass
69 balance (CMB), and positive matrix factorization (PMF) models (Lai et al., 2016; Yao et al., 2016a;
70 Zhang et al., 2015ba). PMF analysis of atmospheric dust in urban areas such as Lanzhou, Taiyuan,
71 and Jinan have identified diverse sources, including coal combustion, industrial emissions,
72 construction dust, windblown dust, vehicle emissions, and resuspended road dust. Seasonal
73 variations indicate that coal combustion during the domestic heating period and regional
74 meteorological conditions significantly influence dust deposition (Hu and Liu, 2022; Chen et al.,
75 2024; Yang et al., 2024; Zhang et al., 2022). These findings underscore the urgency of region-
76 specific pollution control strategies.

77 The Qinghai-Xizang Plateau (QXP) is a key regulator of Northern Hemisphere climate
78 variability and plays a vital role in global ecological and climatic stability, often referred to as the
79 “Asian Water Tower” (Liu et al., 2019; Liu et al., 2020b). –However, rapid glacier retreat on the
80 plateau poses risks to the Asian hydrological cycle and the monsoon system, with potential adverse
81 impacts if unchecked (Luo et al., 2020). –Beyond climate warming and increased humidity, black
82 carbon (BC) significantly accelerates glacial melt by inducing atmospheric warming and enhancing
83 radiative absorption at the glacier surface (Bond and Bergstrom, 2006; Chen et al., 2015). –Notably,
84 biomass burning in South and Central Asia during winter serves as a major source of BC, further
85 exacerbating glacier decline on the plateau (Zhang et al., 2015e2015b; Zheng et al., 2017; Xu et al.,
86 2018b). However, local sources within the QTP, particularly the Qaidam Basin (QDB) in the
87 northeastern region, should not be underestimated, as QDB is a key dust source for the plateau (Wei
88 et al., 2017; Zheng et al., 2021).

89 The QDB, known as the “Treasure Bowl” of the QXP, is rich in minerals, coal, oil, and gas,
90 positioning it as a key economic hub in northwest China. It has a high population density, and intense
91 human activity, yet it is highly sensitive to climate change. Extensive resource extraction has

92 rendered its ecosystem fragile (Li and Sha, 2022), exacerbating atmospheric pollution. Unlike South
93 Asia, Central Asia, and Xizang—where biomass fuels dominate—QDB relies primarily on a
94 mixture of coal, yak dung, and firewood for winter domestic heating, reflecting a unique energy
95 structure (Liu et al., 2008; Xiao et al., 2015; Behera et al., 2015; Kerimray et al., 2018; Jiang et al.,
96 2020; Shen et al., 2021). The combustion of coal releases significant pollutants, light-absorbing
97 organic compounds like BC and brown carbon, and hazardous gases such as ~~such as~~ carbon dioxide
98 (CO₂), nitrogen oxides (NO_x) and carbon monoxide (CO). These emissions impact human health
99 and exacerbate climate warming, thereby influencing regional and global climate systems (Munawer,
100 2018; Ye et al., 2020; Zhou et al., 2025). Consequently, we posit that seasonal carbon emissions in
101 QDB, particularly during winter domestic heating, could exert a unique influence on the climate
102 and ecological stability of the QXP.

103 Additionally, the QDB as a representative arid region with intensive human activity, exhibits
104 climatic and environmental conditions comparable to the Tarim and Junggar Basins in Xinjiang, the
105 Great Basin in the United States, and other hyper-arid areas. These regions are characterized by low
106 precipitation, rich mineral resources subject to significant anthropogenic impact, and abundant salt
107 lakes. Similarly, salt lakes such as the Uyuni in Bolivia, Atacama in Chile, and Ombre-Muerto in
108 Argentina are located on high plateaus averaging 3,000 m in elevation, with surrounding climates
109 and environments comparable to those of the QDB. Research in the Tarim and Junggar Basins has
110 predominantly focused on dust events, their sources, and associated gas emissions (Gao and
111 Washington, 2010; Liu et al., 2016b; Filonchik et al., 2018; Yu et al., 2019; Zhou et al., 2023). In
112 the Great Basin, studies largely address ozone and dust sources (Hahnenberger and Nicoll, 2012;
113 Vancuren and Gustin, 2015; Miller et al., 2015). Research on salt lake atmospheres has
114 predominantly focused on high-salinity dust emissions resulting from lakebed desiccation due to
115 resource extraction (Löw et al., 2013; Gholampour et al., 2015; Moravek et al., 2019; Christie et al.,
116 2025), with limited research on atmospheric carbon components, their sources, and environmental
117 impacts. Therefore, this research aims to investigate atmospheric carbonaceous aerosols in arid
118 basins with intensive human activity and climates comparable to the QDB, as well as in salt lakes
119 environments.

120 This study, conducted from January 2020 to March 2023, involved monthly dust deposition
121 sampling at six urban and rural monitoring sites located in the southern QDB. Samples were

122 categorized into two seasonal periods: the domestic heating (HP) and the non-domestic heating
123 (NHP) periods. The HYSPLIT model, and PMF receptor modeling using analyses of dust flux,
124 soluble ions, trace elements, and carbonaceous components, alongside OC/EC and char/soot
125 ratios—the primary sources of dust deposition were identified, with particular emphasis on
126 contributions from domestic heating. The study further evaluated the environmental impacts on the
127 QXP, considering its distinctive energy structure. Furthermore, ~~These~~these findings offer a
128 scientific basis and reference for examining atmospheric carbonaceous aerosols in arid basins with
129 similar climates and human activities to the QDB, as well as in salt lake regions.

130

131 **2. Materials and methods**

132 **2.1 Sampling**

133 The QDB, situated in the northeastern part of the QXP, is bordered by the Altyn-Tagh, Kunlun,
134 and Qilian mountains, making it one of China's largest intermontane basins (Zhang, 1987). With an
135 average elevation of 3,000 m, the basin features an extremely arid climate characterized by less than
136 20 mm of annual precipitation in the northwestern region, while evaporation rates exceed 2,000 mm.
137 The QDB is rich in salt lakes, non-ferrous metals, and hydrocarbon resources, with significant coal
138 deposits. It leads China in reserves of halite, potash, magnesium, lithium, strontium, asbestos,
139 earning it the nickname “Treasure Basin”. As a major salt lake resource area, it hosts 33 lakes—
140 including Qarhan, ~~Daqingkan~~Dachaidan, and Caka Salt Lake—and faces notable conflicts between
141 resource extraction and ecological preservation. Agriculturally, it cultivates crops like goji berries,
142 quinoa, and forage grass, and hosts China’s largest resource-rich circular economy pilot zone. The
143 permanent population of the basin is approximately 400,000, primarily using coal, yak dung, and
144 firewood for domestic heating (Jiang et al., 2020). Additionally, annual tourism peaks from May to
145 September, attracting around 17 million visitors, which likely amplifies atmospheric pollutant
146 emissions.

147 From January 2020 to March 2023, monthly dust deposition samples were collected from six
148 monitoring stations in the southern QDB. The stations included Xiao Zaohuo (XZH), Golmud
149 (GEM), Da Gele (LTC), Nuo Muhong (NMH), Balong (BLX), and Dulan station (DLX). Dry
150 deposition collection employed the glass ball method using Marble dust collector (MDCO) designed

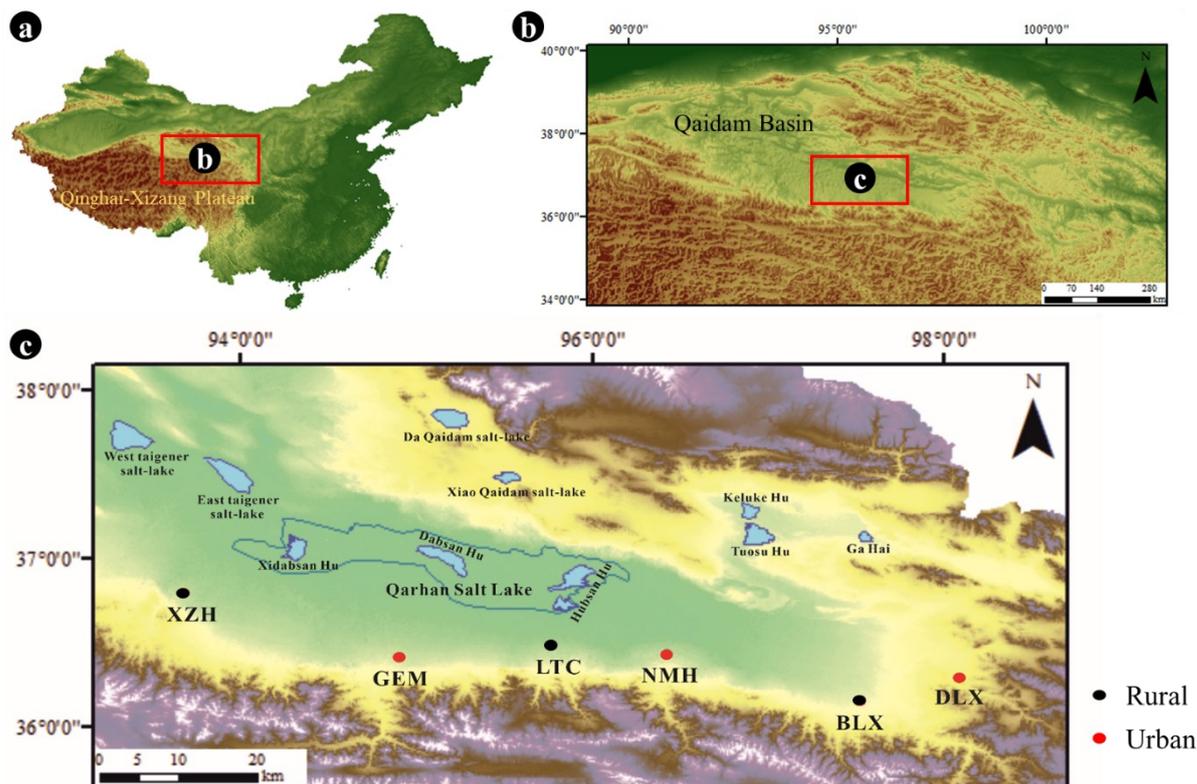
151 dust collection cylinders (Sow et al., 2006). The stainless steel collection device (50×30×30 cm)
152 contained a plastic sieve container of identical dimensions, with the sieve base positioned 10 cm
153 above the opening and perforated with 0.5 cm diameter holes (Figure S1). To minimize dust
154 resuspension during high wind events (Qian and Dong, 2004), two layers of 16 mm glass balls were
155 placed within the sieve container. A high-density polyethylene bag was attached to the base for
156 sample collection. According to Sow et al (2006), the collection efficiency of the MDCO decreases
157 with increasing wind speed, dropping below 20% when wind speed exceeds 3 m/s, and it
158 preferentially collects fine dust particles ranging from 10 to 31 μm in size (Chow, 1995). This type
159 of collector has been widely used in numerousmany studies to evaluate local dust conditions
160 (Abdollahi et al., 2021; Barjoe et al., 2021; Alzahrani et al., 2024). Therefore, we consider this
161 collector effective for capturing fine dust particles, and the actual dust deposition flux can be
162 estimated by accounting for its approximately 20% collection efficiency.

163 In this study, dust samples were collected monthly, with each sampling period lasting 30 or 31
164 days. The installation height and environment of the samplers are provided in Table S1. To ensure
165 only dry dust was collected, collection devices were covered during rain or snowfall. A total of 37,
166 39, 23, 30, 16, and 29 samples were obtained from XZH, GEM, LTC, NMH, BLX, and DLX stations,
167 respectively. Laboratory protocols incorporated biannual analyses with negative controls and
168 appropriate control samples. As continuous dust monitoring commenced in 2020, site blanks were
169 evaluated during initial sampling. Stations were classified as Urban (GEM, NMH, DLX) and Rural
170 (XZH, LTC, BLX) based on location characteristics. Consistent with the cold-arid climate in QDB,
171 the HP was defined as October-April, while the NHP spanned May-September.

172 Materials such as plant remnants, microfauna, and bird droppings were removed from the
173 sample bags with tweezers. The samples were then measured on a balance (0.0001 g) to determine
174 the dust deposition flux (Eq. 1) (Yu et al., 2016):

$$175 \quad M = \frac{m \times 30}{S \times K}, \quad (1)$$

176 where M is dust deposition [g/(m²·30d)]; m is the sample mass (g); S is the area of the dust collection
177 device (m²); and K is the actual number of sampling days per month (d).



178

179 **Figure 1.** Spatial distribution of monitoring stations in the southern Qaidam Basin. Urban stations
 180 (red) and rural stations (black) are labeled as follows: XZH (Xiao Zaohuo), GEM (Golmud), LTC
 181 (Da Gele), NMH (Nuo Muhong), BLX (Balong), DLX (Dulan).

182

183 2.2 Water-soluble inorganic ions

184 A 100 mg sample was weighed and transferred into a 250 mL bottle. The mixture underwent
 185 ultrasonic extraction for 20 minutes to ensure complete solubilization. The resulting supernatant
 186 was then filtered through a 0.45 μm filter for analysis. Based on preliminary experimental results,
 187 the concentrations of major ions (K^+ , Na^+ , Mg^{2+} , and Ca^{2+}) were measured using Inductively
 188 Coupled Plasma Optical Emission Spectrometer (ICP-OES, NexIon 2000). Anions (Cl^- , SO_4^{2-} , and
 189 NO_3^-) were quantified using ion chromatography (IC). To ensure measurement accuracy, samples
 190 were organized in sets of twenty, with one randomly selected sample from each group serving as a
 191 replicate, achieving an error margin of less than 10%. The detection limits for the various
 192 components were as follows: K^+ (0.0560 mg/L), Na^+ (0.0100 mg/L), Ca^{2+} (0.0037 mg/L), Mg^{2+}
 193 (0.0390 mg/L), SO_4^{2-} (0.0090 mg/L), NO_3^- (0.0125 mg/L), Cl^- (0.0100 mg/L). All standard solutions
 194 employed in the analysis were sourced from the National Standard Material Center.

195

196 **2.3 Trace element analysis**

197 According to the Chinese State Standard “Ambient air and waste gas from stationary sources
198 emission--determination of metal elements in ambient particles” (HJ 777-2015), the concentrations
199 of elements such as iron (Fe), aluminum (Al), silicon (Si), titanium (Ti), copper (Cu), cadmium (Cd),
200 chromium (Cr), manganese (Mn), nickel (Ni), zinc (Zn), lead (Pb), and vanadium (V) were
201 quantified using Inductively Coupled Plasma Mass Spectrometry (ICP-MS) and ICP-OES. A dust
202 sample weighing 0.100 g was placed in a Teflon cup, to which 20.0 mL of a nitric acid-hydrochloric
203 acid digestion solution was added. The sample was heated to reflux at $100 \pm 5^\circ\text{C}$ for 2 h under a
204 watch glass, then cooled. Following this, the inner walls of the cup were rinsed with water, and
205 approximately 10 mL of water was added, allowing the mixture to stand for 30 minutes for
206 extraction. The extract was then filtered into a 100 mL volumetric flask and diluted to volume with
207 distilled water for analysis. In cases where organic matter content was high, an appropriate amount
208 of hydrogen peroxide was introduced during digestion to decompose the organic materials. Prior to
209 sample analysis, the system was flushed with a rinse solution until the blank intensity value reached
210 a minimum, and samples were analyzed only after the signal stabilized. If the concentration of any
211 element exceeded the calibration range, the sample was diluted and reanalyzed.

212

213 **2.4 Carbon analysis**

214 This study utilized a combination of wet chemical treatment and thermal/optical reflection
215 (TOR) to analyze ~~trace-carbon~~ elements in dust deposition (Han et al., 2007b; Han et al., 2007a;
216 Han et al., 2016). Dust samples were treated with hydrochloric and hydrofluoric acids to remove
217 inorganic materials. The residual solution was then filtered through a pre-combusted quartz fiber
218 filter (Whatman, 450°C for 4 h, diameter 47 mm). This method has been widely applied to measure
219 OC and EC contents in lake sediments and urban soils (Han et al., 2009; Khan et al., 2009; Han et
220 al., 2011). Studies have shown that the EC collection efficiency of this method is approximately
221 99.6% (Zhan et al., 2013); however, its OC collection efficiency remains unclear. The filtered
222 samples were air-dried and analyzed for carbon content using a DRI 2001 thermal/optical carbon
223 analyzer (Atmoslytic Inc., Calabasas, CA) at the Institute of Earth Environment, Chinese Academy

224 of Sciences, adhering to the Interagency Monitoring of Protected Visual Environments (IMPROVE)
225 protocol.

226 A 0.544 cm² disc was extracted from the filter and placed in a quartz boat for analysis. During
227 the carbon analysis, the samples were initially heated in a 100% helium atmosphere, resulting in the
228 production of four organic carbon (OC) fractions (OC1, OC2, OC3, and OC4) at four different
229 temperature levels (140, 280, 480, and 580-°C). The atmosphere was subsequently changed to 2%
230 O₂/98% He, generating three elemental carbon (EC) fractions (EC1, EC2, and EC3) at three
231 temperatures (580, 740, and 840 °C). Volatile carbon underwent carbonization in an anaerobic
232 environment, indicated by a decrease in laser reflectance, and is referred to as "pyrolyzed organic
233 carbon" (OPC). In the oxidative atmosphere, OPC was emitted along with the original EC from the
234 filter. The amount of OPC is defined as the carbon evolved until the laser reflectance returned to its
235 baseline value (Han et al., 2007b). According to the IMPROVE protocol, EC is calculated as the
236 total of the three EC subfractions minus OPC (i.e., EC is defined as the sum of EC1, EC2, and EC3,
237 with OPC subtracted $EC = EC1 + EC2 + EC3 - OPC$). The method enables differentiation between soot
238 and ashchar, as determined by the gradual oxidation of these black carbon subtypes in standard
239 reference materials during the EC1 and the EC2 plus EC3 $EC2 + EC3$ steps, where soot-char is
240 defined as EC1 minus OPC $EC1 - OPC$ and ash-soot as the sum of EC2 and EC3 $EC2 + EC3$ (Han et
241 al., 2007a; Han et al., 2016).

242 Please note that in this manuscript, we interchangeably use the terms "EC" and "BC." While
243 these terms do not strictly refer to the same component, they serve as an adequate approximation
244 within the scope of this study (Seinfeld et al., 1998; Bond et al., 2004). We use "EC" when discussing
245 emissions and modeling components, reserving "BC" for climate-related discussions.

246

247 **2.5 Statistical analysis**

248 (1) Estimation of Secondary Organic Carbon

249 OC consists of primary organic carbon (POC) and secondary organic carbon (SOC). ~~An OC/EC~~
250 ~~ratio exceeding 2.0 indicates the possible presence of secondary organic aerosol (SOA) (Castro et~~
251 ~~al., 1999).~~ Due to the intricate physical and chemical processes involved, SOC in urban atmospheres
252 cannot be directly measured. Therefore, an indirect estimation method, known as the EC tracer

253 method, has been developed (Turpin and Huntzicker, 1991). If the concentrations of OC and EC are
254 available and primary OC from non-combustion sources ($OC_{\text{non-comb}}$) can be disregarded, EC can be
255 utilized as a tracer for POC from combustion sources, facilitating the estimation of SOC (Turpin
256 and Huntzicker, 1995):

$$257 \quad \text{POC} = \text{EC} \times (\text{OC}/\text{EC})_{\text{pri}}, \quad (2)$$

$$258 \quad \text{SOC} = \text{OC}_{\text{total}} - \text{POC}, \quad (3)$$

259 where OC_{total} represents total organic carbon.

260 Traditional methods for determining $(\text{OC}/\text{EC})_{\text{pri}}$ involve regressing OC and EC within a fixed
261 percentile range of the lowest (OC/EC) ratio data (typically 5-20%) or relying on sampling days
262 characterized by low photochemical activity and local emissions (Castro et al., 1999; Lim and
263 Turpin, 2002). However, these approaches are limited by their empirical nature, lacking clear
264 quantitative criteria for selecting the data subsets used to establish $(\text{OC}/\text{EC})_{\text{pri}}$. In this study, we
265 employed the minimum R squared (MRS) method (Millet et al., 2005; Wu and Yu, 2016; Wu et al.,
266 2018a) to determine $(\text{OC}/\text{EC})_{\text{pri}}$. This method calculates a set of hypothetical $(\text{OC}/\text{EC})_{\text{pri}}$ and SOC
267 values to identify the minimum correlation coefficient (R^2) between SOC and EC, allowing for the
268 accurate derivation of $(\text{OC}/\text{EC})_{\text{pri}}$. The computational procedure followed the algorithm developed
269 by Wu and Yu (2016) (available at: <https://sites.google.com/site/wuchengust>), implemented within
270 the Igor Pro environment (WaveMetrics, Inc., Lake Oswego, OR, USA). Due to the limited dataset
271 size and low temporal resolution, the MRS analysis was performed collectively across all sampling
272 sites. In this approach, the R^2 between SOC and EC was calculated iteratively for a range of
273 $(\text{OC}/\text{EC})_{\text{pri}}$ values spanning 0 to 10. The minimum R^2 value of 1.33 (Figure S2) identified the
274 optimal $(\text{OC}/\text{EC})_{\text{pri}}$ representative of the true primary ratio.

275
276 (2) Playa salt (ps) and non-playa salt (nps)

277 To differentiate the contributions of salt lake sources to water-soluble ions in atmospheric
278 deposition, we adopted a methodology similar to that used for marine aerosols. This approach relies
279 on the ratio of water-soluble ions (SO_4^{2-} , Ca^{2+} , K^+ , Mg^{2+} , Cl^-) to Na^+ in the salt lakes of the QDB,
280 enabling us to assess the contribution of ps- Na^+ components to nps (Zhang, 1987); details in Zhu
281 (2025).

$$282 \quad \text{nps-SO}_4^{2-} = [\text{SO}_4^{2-}] - 0.333 \cdot \text{ps-Na}^+, \quad (4)$$

283
$$\text{nps-Ca}^{2+} = [\text{Ca}^{2+}] - 0.062 \cdot \text{ps-Na}^+, \quad (5)$$

284
$$\text{nps-K}^+ = [\text{K}^+] - 0.087 \cdot \text{ps-Na}^+, \quad (6)$$

285
$$\text{nps-Mg}^{2+} = [\text{Mg}^{2+}] - 0.051 \cdot \text{ps-Na}^+, \quad (7)$$

286
$$\text{nps-Cl}^- = [\text{Cl}^-] - 2.287 \cdot \text{ps-Na}^+, \quad (8)$$

287 This was accomplished using equations that incorporate total Na^+ , total Ca^{2+} , the average
 288 crustal ratio $(\text{Na}^+/\text{Ca}^{2+})_{\text{crust}} = 0.56$ w/w (Bowen, 1979), and the average $(\text{Ca}^{2+}/\text{Na}^+)$ ratio for Qaidam
 289 salt lakes, $(\text{Ca}^{2+}/\text{Na}^+)_{\text{salt lake}} = 0.06$ w/w (Zhang, 1987). Among these, the mass concentration of
 290 $[\text{SO}_4^{2-}]$, $[\text{Ca}^{2+}]$, $[\text{K}^+]$, $[\text{Mg}^{2+}]$, $[\text{Na}^+]$ and $[\text{Cl}^-]$ were identified as constituents of dust-fall.

291
$$\text{ps-Na}^+ = [\text{Na}^+] - \text{nps-Na}^+, \quad (9)$$

292
$$\text{nps-Na}^+ = \text{nps-Ca}^{2+} \cdot (\text{Na}^+ / \text{Ca}^{2+})_{\text{crust}}, \quad (10)$$

293
$$\text{nps-Ca}^{2+} = [\text{Ca}^{2+}] - \text{ps-Ca}^{2+}, \quad (11)$$

294
$$\text{ps-Ca}^{2+} = \text{ps-Na}^+ \cdot (\text{Ca}^{2+} / \text{Na}^+)_{\text{salt lake}}, \quad (12)$$

295

296 (3) HYSPLIT backward trajectory model

297 Backward trajectory clustering analysis was conducted on sampling points using the TrajStat
 298 plugin within Meteoinfo software. Daily backward trajectories for 48 hours were calculated from
 299 January 2020 to March 2023 and classified monthly based on differences in the horizontal
 300 movement direction and velocity of air masses. The trajectories were initiated at Universal Time
 301 (UTC) 00:00, with a 6-hour increment, originating from 500 m above sea level (Yang et al., 2014).
 302 Meteorological data utilized in this research were obtained from the Global Data Assimilation
 303 System (GDAS) provided by the U.S. National Centers for Environmental Prediction (NCEP),
 304 covering the period from December 2019 to February 2023 (Meteoinfo software website:
 305 <http://meteothink.org>).

306

307 (4) PMF model

308 PMF is a multivariate factor analysis tool that decomposes a matrix of speciated sample data
 309 into two matrices: factor contributions (G) and factor profiles (F). The goal of the PMF model is to
 310 solve the chemical mass balance between measured species concentrations and the respective source
 311 profiles, with the purpose of minimizing the objective function Q (Eq. 13) based upon the
 312 uncertainties (uij) of measured species (Paatero and Tapper, 1994).

313
$$e_{ij} = x_{ij} - \sum_{h=1}^p g_{ih} f_{hj}; Q = \sum_{i=1}^m \sum_{j=1}^n (e_{ij}/h_{ij} s_{ij})^2, \quad (13)$$

314 where x_{ij} is the measured concentration of the j_{th} species in the i_{th} sample at receptor sites. f_{kj} is the
 315 source profile of the j_{th} species in the k_{th} factor and g_{ik} is the mass contribution of the k_{th} factor in
 316 the i_{th} sample. e_{ij} is the difference between modeled concentrations and measured concentrations.

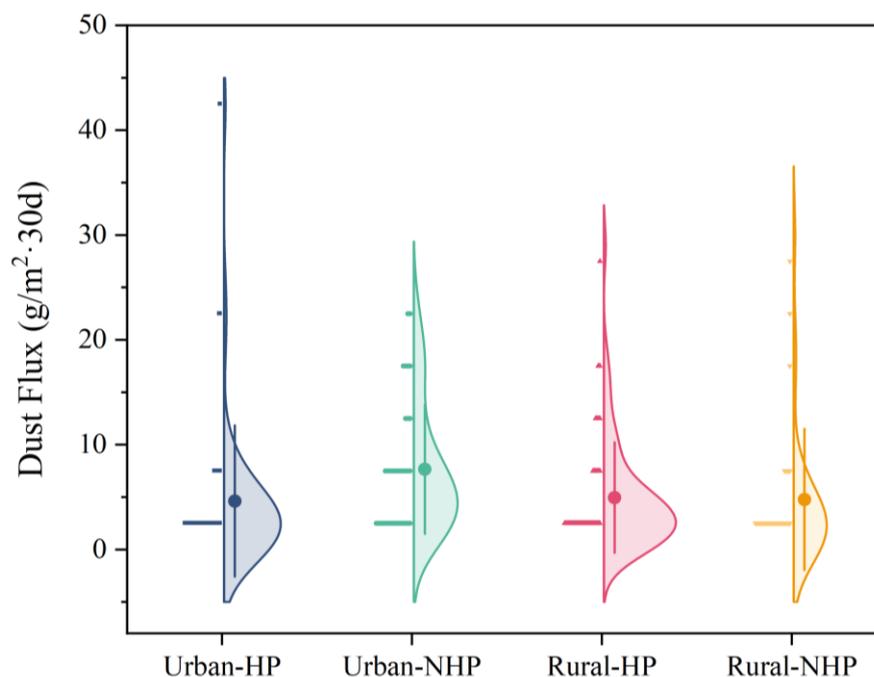
317 The uncertainty for individual species (u_{ij}) was defined as the sum of two components: the x_{ij}
 318 multiplied by an error fraction, and one-third of the method detection limit (MDL), calculated to be
 319 $x_{ij} \times \text{error fraction} + 1/3 \text{ MDL}$, where MDL is the method detection limit. For data below the MDL,
 320 concentrations were replaced by 1/2 MDL and the corresponding uncertainty was set to 5/6 MDL
 321 (Reff et al., 2007). An extended description of the PMF parameters used in this study and error
 322 estimate based on the model's Q value, displacement (DISP), and bootstrapping (BS) tests (DISP-
 323 BS) are provided in the Supplementary Information. The error assessment and uncertainty data for
 324 the PMF source apportionment can be found in Tables S2 and S3, respectively.

325

326 **3. Results and discussion**

327 **3.1 Atmospheric dust deposition flux and water ions concentration**

328 The total deposition flux (DF) in the southern QDB is $5.41 \pm 5.33 \text{ g/m}^2 \cdot 30\text{d}$, slightly lower
 329 than that of the Lake Aibi Basin ($3.32-23.410.77 \text{ g/m}^2 \cdot 30\text{d}$) (Li et al., 2022), but higher than the
 330 surrounding areas of Akatama Salt Lake ($2.93 \text{ g/m}^2 \cdot 30\text{d}$) (Wang et al., 2014). Specifically, DF was
 331 $4.67 \pm 4.96 \text{ g/m}^2 \cdot 30\text{d}$ in rural and $5.97 \pm 5.73 \text{ g/m}^2 \cdot 30\text{d}$ in urban areas. During the HP, DF in rural
 332 and urban areas were $4.62 \pm 4.15 \text{ g/m}^2 \cdot 30\text{d}$ and $4.95 \pm 5.25 \text{ g/m}^2 \cdot 30\text{d}$, respectively. In contrast, NHP
 333 showed increased DF values of $4.77 \pm 4.42 \text{ g/m}^2 \cdot 30\text{d}$ (rural) and $7.66 \pm 6.09 \text{ g/m}^2 \cdot 30\text{d}$ (urban)
 334 (Figure 2). Notably, Urban DF during NHP demonstrated a 54.6% increase relative to HP, while
 335 rural DF rose by 7.1%—contrasting previous findings that associated elevated DF with HP coal
 336 combustion (Cheng et al., 2009; Gao et al., 2013; Guo et al., 2010; Qi et al., 2018). We hypothesize
 337 that the increase in DF during the NHP is attributed to heightened tourism activity (May to
 338 September peaks season) in the QDB (Zhang et al., 2011), attracting approximately 17 million
 339 tourists annually, the number continues to grow. This influx likely leads to increased urban
 340 emissions, particularly in densely populated areas such as DLX and GEM (Figure S2S3),
 341 contributing to the elevated DF levels.

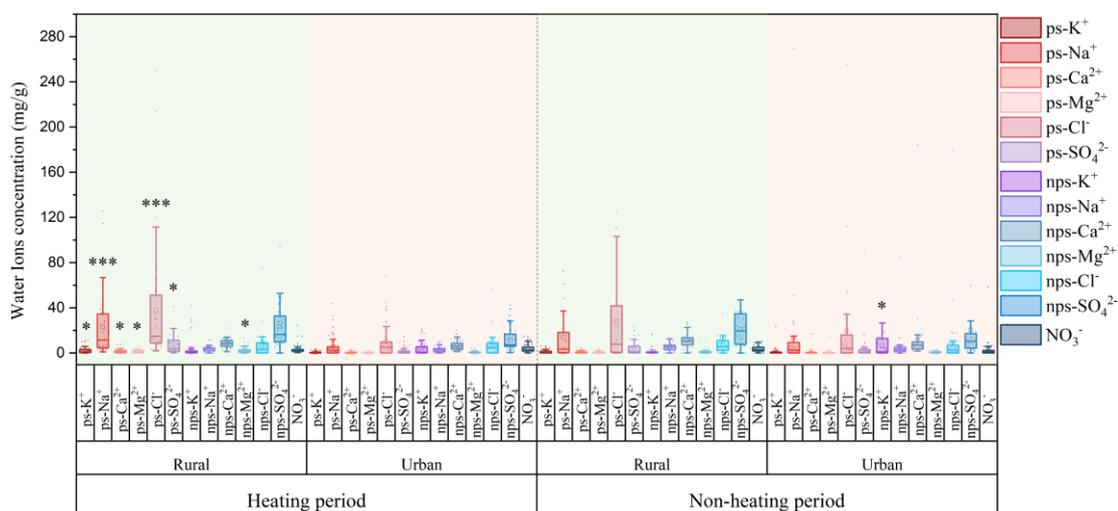


342
 343 **Figure 2** Dust flux distribution in urban and rural. The distribution of dust flux in four contexts:
 344 Urban with domestic heating period (Urban-HP), Urban with non-HP (Urban-NHP), Rural with
 345 domestic heating period (Rural-HP), and Rural with non-HP (Rural-NHP). Each violin plot
 346 illustrates the density distribution of dust flux values, with the central dot representing the mean,
 347 and the vertical lines indicating the interquartile range.

348

349 Water-soluble ion concentrations differed significantly between rural (115.31 mg/g) and urban
 350 (72.81 mg/g) areas. In rural, water-soluble ion content was greater during the HP than in the NHP,
 351 while the opposite trend was observed in urban areas (Figure 3, and S3S4). We categorized the
 352 water-soluble ions in dust deposits into ~~playa salt (ps)~~ and ~~non-playa salt (nps)~~ based on their sources,
 353 following the model of marine aerosols (Zhu et al., 2025). Playa salt content consistently surpassed
 354 nps in rural areas across both periods, while urban areas showed the opposite trend. Notably, during
 355 the NHP, ps content in urban and rural increased by 54.46 and 36.86% respectively. Backward
 356 trajectory analysis indicated that airflow in both rural and urban areas primarily originated from the
 357 northwest QDB and the eastern Tarim Basin during the HP, while during the NHP, it was influenced
 358 more broadly by the salt lake of the QDB (Figure S4S5, and S5S6), aligning with the observed
 359 variations in ps content. A similar increase in summer sea salt and non-sea salt ions has been reported
 360 in Rajkot, India, attributed to ocean wind direction (Gupta and Dhir, 2022).

361 The ratio of $\text{nps-SO}_4^{2-}/\text{NO}_3^-$ in soluble ions is used to differentiate between coal combustion
 362 (fixed sources) and vehicular emissions (mobile sources) (Arimoto et al., 1996; Shen et al., 2008).
 363 The higher $\text{nps-SO}_4^{2-}/\text{NO}_3^-$ ratio in urban areas compared to rural areas (Figure S7) indicates that
 364 during the study period, stationary sources (e.g., industrial emissions, coal, and biomass combustion)
 365 contributed more significantly to the ionic composition of urban dust-fall (Pipal et al., 2019). In
 366 contrast, the ratio was considerably lower during the NHP than during HP, suggesting that mobile
 367 sources likely played a more important role in the dust ionic composition during the NHP. Generally,
 368 emissions from coal combustion and biomass burning combustion was intensified in northern
 369 China during the HP (Liu et al., 2016), while vehicle emissions dominate during the NHP (Xu et al.,
 370 2012), supporting the reliability of this analysis. Nevertheless, further investigation integrating
 371 atmospheric emission inventories, source tracers, aerosol physicochemical processes, and
 372 meteorological conditions is warranted. Higher ratios in urban compared to rural areas (Figure S6)
 373 suggest a greater influence of fixed sources on urban dust deposition. Additionally, the lower nps-
 374 $\text{SO}_4^{2-}/\text{NO}_3^-$ ratio during the NHP indicates a predominant role of mobile sources in NHP dust fall.
 375 Typically, coal and biomass burning emissions intensify during the HP in northern China (Liu et al.,
 376 2016a), while vehicle emissions dominate in the NHP (Xu et al., 2012). These findings support the
 377 validity of the analysis, however, further validation using additional indicators is recommended.



378 **Figure 3** Concentrations of water ions in rural and urban across different seasonal periods (domestic
 379 heating and non-domestic heating periods). Asterisks indicate statistically significant differences
 380 between sites, with *, $P < 0.05$; **, $P < 0.01$; ***, $P < 0.001$.
 381
 382

3.2 Organic carbon and element carbon compositions

The average total carbon (TC) concentration in the southern QDB was 3.83 ± 3.2 mg/g, significantly lower than that of Huangshi, Hubei province (25.15 ± 11.79 mg/g) and Washington (157 ± 3.2 mg/g), Kumasi in West Africa (28 mg/g) and Xi'an (14.6 ± 5.8 mg/g) (Han et al., 2007a; Han et al., 2009b; Zhan et al., 2016; Bandowe et al., 2019), suggesting relatively low carbon emissions in the southern QDB. Average OC and EC levels in QDB are markedly lower than those in Xi'an (7.4 and 7.2 mg/g, respectively), Wuhu (33.26 and 22.49 mg/g, respectively), and Nanchang (25.15 and 11.46 mg/g, respectively) (Han et al., 2009a; Deng et al., 2014; Zhang, 2014), but significantly higher, particularly for EC, than in Nam Co (0.35 mg/g) (Chen et al., 2015).

In urban, TC content (3.05 ± 2.46 mg/g) were was marginally lower than the rural level (3.55 ± 3.56 mg/g), although this difference was not statistically significant. Contrasting spatial patterns emerged for carbon components: EC dominated urban areas (1.46 ± 1.60 mg/g), while OC prevailed in rural (2.25 ± 1.92 mg/g) (Figure 4, and S87). This disparity may be attributed to the long-term combustion of biomass, coal, and wood in rural settings (Na et al., 2004). It may also be associated with meteorological conditions, particularly heightened solar radiation resulting from reduced primary emissions in rural areas, which facilitates the formation of SOC (Xu et al., 2018a; Wang et al., 2019).

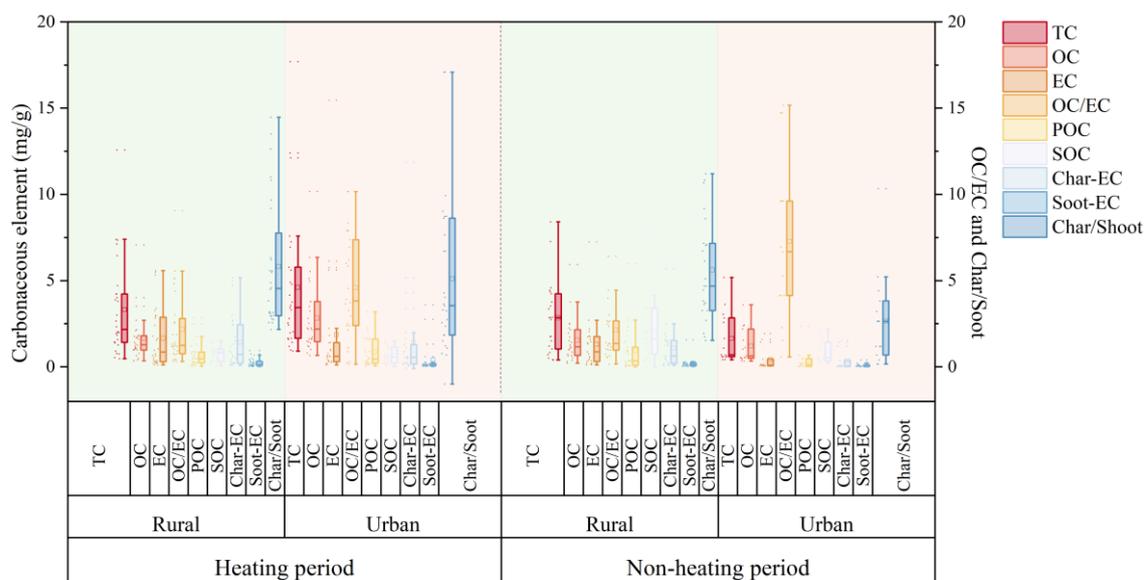


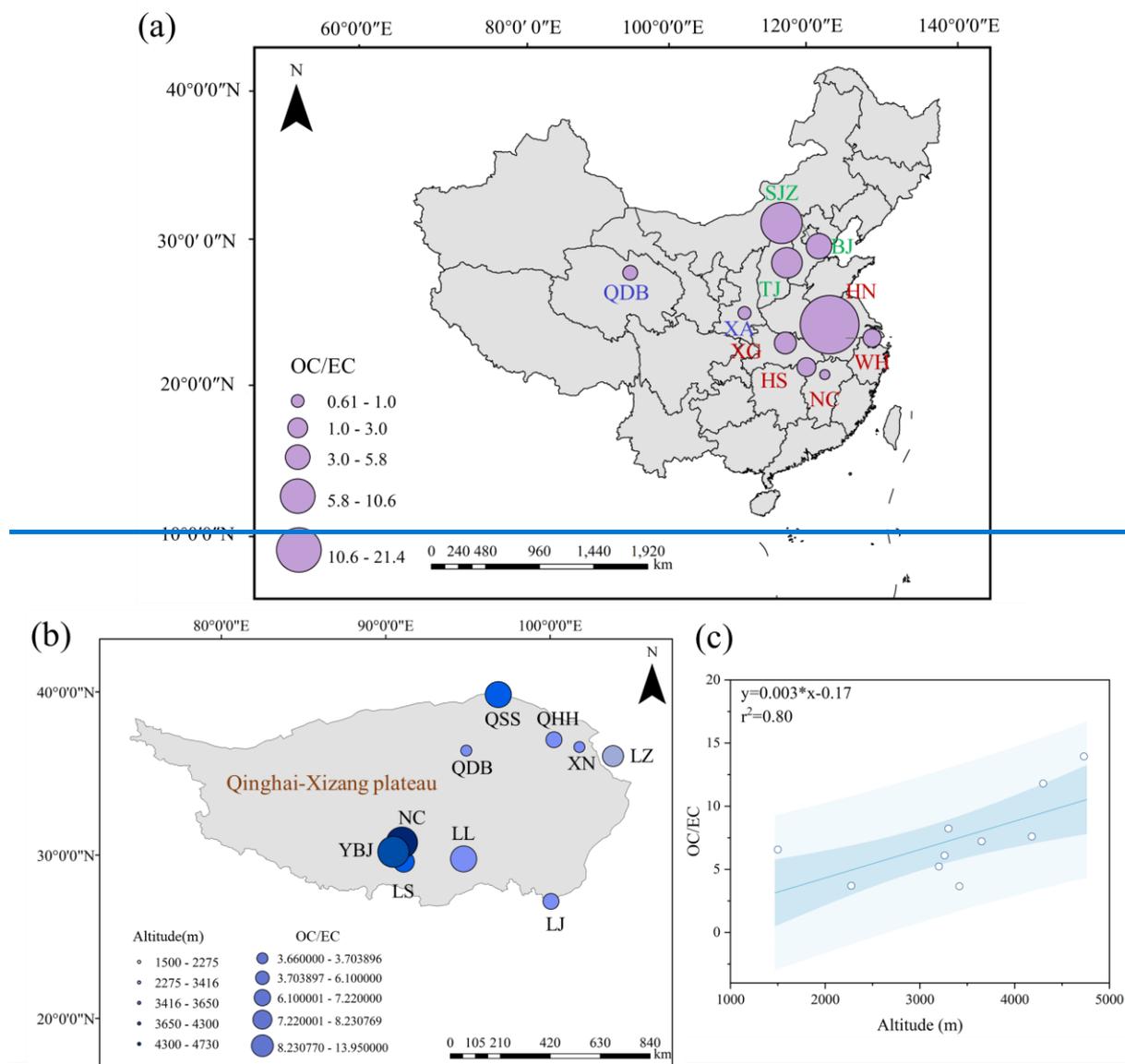
Figure 4 Concentrations of organic carbon (OC), elements carbon (EC), secondary organic carbon (OC), primary organic carbon (POC), char-EC, soot-EC and OC/EC, char/soot ratios in different

404 sites (Rural, Urban) and seasonal variations (domestic heating and non-domestic heating period).

405

406 Seasonal analysis revealed elevated carbonaceous compound concentrations, specifically OC
407 and EC, during the HP. This increase is primarily due to local domestic heating activities coupled
408 with adverse meteorological conditions, such as low temperature, weak winds (Oliveira et al., 2007;
409 Gong et al., 2017), weak atmospheric turbulence, and frequent atmospheric inversions (Guo et al.,
410 2016). Rural emissions primarily stem from coal and biomass burning for heating and cooking-
411 (Zhang et al., 2000; He et al., 2004), contributing to reduced OC and EC content in the NHP, whereas
412 elevated EC levels in urban areas are primarily linked to vehicular and industrial sources.
413 Spatiotemporal transport dynamics show EC depletion during rural ward pollutant migration due to
414 atmospheric dispersion, particularly affecting coarse particulate fractions.

415 Notably, rural carbon emissions exceed urban levels in the southern basin, potentially
416 attributable to extended HP duration (7 months) compelling low-grade fuel (crop residues, wood,
417 raw coal and yak dung) utilization (Na et al., 2004). In contrast, urban areas benefit from solar/wind
418 energy infrastructure and government-led clean heating initiatives (“suitable electricity for
419 electricity” policy), achieving 66.63% clean heating penetration (Statistical Yearbook of Haixi
420 Xizang Autonomous Prefecture of Qinghai Province), leading to a comparatively lower TC content.
421 The OC/EC ratio is a valuable indicator of carbonaceous aerosol sources. In this study, Urban areas
422 exhibited stable OC/EC ratios ranging from 0.15 to 15.16 (mean = 2.16), whereas rural areas showed
423 significantly higher ratios during NHP (7.27 ± 4.66) compared to HP (4.57 ± 3.02) (Figure 4 [and](#)
424 [S9S10](#)). Typically, the OC/EC ratio for vehicle emissions ranges from 0.7 to 2.4, for coal combustion
425 emissions from 0.3 to 7.6, and for biomass burning from 3.8 to 14.5 (Schmidl et al., 2008; Gonçalves
426 et al., 2010; Pio et al., 2011; Popovicheva et al., 2016).



427

428 **Figure 5** Distribution of OC/EC ratios across various regions of China (a), at different altitudes on
 429 the Qinghai Xizang Plateau (b), and the relationship between OC/EC ratios and altitude (c). Blue
 430 designations represent the Northwest region, red indicates the Central region, and green denotes the
 431 Eastern region. The color of the circles corresponds to altitude, while circle size reflects the
 432 magnitude of the OC/EC ratios.

433 These findings suggest that urban OC/EC ratios (0.15-9.05) are primarily associated with
 434 vehicle and coal combustion, while rural ratios (0.14-15.16) are predominantly linked to coal and
 435 biomass burning. A higher OC/EC ratio typically indicates a greater contribution from biomass
 436 combustion; in this study, the OC/EC ratio of rural was 5.56 ± 3.93 , which is significantly lower
 437 than values recorded in India (8.47) and Nam Co in Xizang (16.3 ± 4.4) (Saud et al., 2013; Chen et

438 al., 2015), yet higher than those observed in Shanxi (0.7-1.6), Beijing (1.9-2.7), and Tianjin rural
439 (2.66) (Zhang et al., 2007; Cheng et al., 2015; Wang et al., 2021). This finding indicates that
440 carbonaceous aerosols in rural QDB derive from both fossil fuel combustion and biomass burning,
441 suggesting source specificity.

442 ~~Figure 5a presents spatial variations in urban OC/EC ratios across China. The findings reveal~~
443 ~~that the Northwest region, represented by QDB urban and Xi'an (XA) (Han et al., 2009b), exhibits~~
444 ~~a significantly lower ratio (1.59 ± 0.56) compared to central regions, including Nanchang (NC)~~
445 ~~(Zhang, 2014), Huangshi (HS) (Zhan et al., 2016), Wuhu (WH) (Deng et al. 2014), Xiaogan (XG)~~
446 ~~(Zhan et al., 2022), and Huainan (HN) (Liu et al., 2020), where the ratio is 5.86 ± 7.81 . This ratio is~~
447 ~~also lower than that observed in eastern cities such as Beijing (BJ) (Tang et al., 2013), Tianjin (TJ)~~
448 ~~(Ma et al., 2019), and Shijiazhuang (SJZ) (Guo et al., 2018), which have a ratio of 6.83 ± 2.77 . This~~
449 ~~pattern is consistent with the trends in atmospheric PM OC/EC ratios (Xie et al., 2023), suggesting~~
450 ~~that the carbon in the dust of the QDB urban primarily results from coal combustion and industrial~~
451 ~~emissions, leading to elevated EC concentrations and lower OC/EC ratios (Liu et al., 2022).~~
452 ~~Conversely, cities with higher economic development, such as Beijing and Tianjin, characterized by~~
453 ~~greater population density and income levels, typically experience secondary pollution, resulting in~~
454 ~~higher OC/EC ratios. Furthermore, as the pretreatment used in this study removes impurities such~~
455 ~~as silicates from the carbonaceous components of dust, and the OC collection efficiency of this~~
456 ~~treatment is currently unknown, the lower OC/EC ratios observed may also be attributed to a lower~~
457 ~~OC collection efficiency.~~

458 ~~We analyzed the impact of altitude on the OC/EC ratio across various regions of the QXP by~~
459 ~~comparing aerosol emissions ($PM_{2.5}$, PM_{10} , TSP) from several sites: QDB (this study, altitude 3416~~
460 ~~m), Lhasa (LS, altitude 3650 m) (Wei et al., 2019), Nam Co (NC, altitude 4730 m) (Chen et al.,~~
461 ~~2015; Wu et al., 2018b), Xining (XN, altitude 2275 m) (Hu et al., 2020), Qilian Shan Station of~~
462 ~~Glaciology and Ecologic Environment (QSS, altitude 4180 m) (Xu et al., 2015), Lijiang (LJ, altitude~~
463 ~~3260 m) (Zhao et al., 2019), Qinghai Lake (QHH, altitude 3200 m) (Jun Li et al., 2013; Zhang et~~
464 ~~al., 2014), Lanzhou (LZ, altitude 1500 m) (Qi et al., 2024), Yangbajing (YBJ, altitude 4300 m)~~
465 ~~(Xiang et al., 2024), and Lulang (LL, altitude 3330 m) (Li et al., 2016). The findings (Figure 5b, c)~~
466 ~~demonstrated a significant positive correlation between altitude and the OC/EC ratio ($r^2 = 0.80$),~~
467 ~~indicating an altitude-dependent relationship (Zhao et al., 2022). This relationship may be attributed~~

468 ~~to the limited sources of EC at higher altitudes and the predominance of aged organic aerosols,~~
469 ~~which are rich in secondary OC, leading to increased OC/EC ratios (Sandrini et al., 2014; Wu et al.,~~
470 ~~2018b). The widespread biomass burning in the QXP further contributes to the presence of semi-~~
471 ~~volatile organic carbon in aerosols, which has a lower light absorption capacity (Koebach Bølling~~
472 ~~et al., 2009), thereby enhancing the OC/EC ratio.~~

473 ~~Both increased altitude and economic development elevate OC/EC ratio, though eastern~~
474 ~~regions exhibit lower values than high altitude zones. High altitude areas show depressed OC and~~
475 ~~EC concentrations overall, with ratio elevation potentially stemming from EC reduction via~~
476 ~~intensive biomass burning. Conversely, the higher ratios observed in economically developed urban~~
477 ~~are primarily driven by anthropogenic activities that facilitate the formation of SOA. These~~
478 ~~scenarios can be differentiated using the WSOC/OC ratio (Patel et al., 2022), though such analyses~~
479 ~~await experimental validation. Additionally, since this research primarily focused on PM larger than~~
480 ~~10 μm , the OC/EC results may be misleading. Further investigation of atmospheric particles smaller~~
481 ~~than 10 μm is necessary to elucidate the impact of altitude on carbonaceous aerosol emissions.~~

482 ~~The elevation of QDB (3416 m) is below the average for QXP (4,000 m), and it exhibits higher~~
483 ~~economic development and population density. Despite this, some herders still burn yak dung for~~
484 ~~heating. The OC, EC emissions, and OC/EC ratios indicate that the carbon sources—comprising~~
485 ~~coal, biomass, and yak dung—are distinct from those in both central and eastern of China~~
486 ~~economically developed regions and other high-altitude QXP areas, reflecting its unique regional~~
487 ~~characteristics.~~

489 **3.3 Char-EC and Soot-EC compositions**

490 EC is classified into soot and char (Han et al., 2009b), with char-EC and soot-EC defined as
491 EC1 minus OPC and the sum of EC2 and EC3 (EC1-OPC) and (EC2+EC3), respectively (Han et
492 al., 2007). Char-EC is typically produced from biomass burning at relatively low temperatures,
493 whereas soot-EC originates from high-temperature coal combustion and automotive emissions (Zhu
494 et al., 2010; Cao et al., 2013). The char-EC/soot-EC ratio, like the OC/EC ratio, serves as an
495 indicator of carbon aerosol sources. Since char and soot are mainly generated through combustion
496 processes, their ratio is typically influenced by two key factors: the primary emission source and the

497 deposition removal efficiency. For localized PM, such as in urban areas, the removal rate is generally
498 negligible (Han et al., 2009a).

499 Char-EC constitutes 75.88% of rural EC (74.71% HP; 78.84% NHP) and 85.00% of urban EC
500 (85.58% HP; 84.11% NHP) (Figure 4, [and S8S9](#)), demonstrating its dominance across spatial and
501 temporal scales. Research suggests that char-EC constitutes a larger proportion of coarse PM, while
502 soot-EC is more predominant in fine particles, resulting in extended atmospheric residence times
503 for soot-EC due to reduced deposition velocities (Han et al., 2009b). The increased levels of char-
504 EC during the urban HP are linked to complex sources, including biomass fuel usage and
505 transportation emissions, resulting in elevated char concentrations in urban areas and along busy
506 roadways (Kim et al., 2011).

507 The char/soot ratios for automobile emissions, coal combustion, and biomass burning are 0.60,
508 1.9, and 11.6, respectively (Chow et al., 2004; Chuang et al., 2014). Generally, high-temperature
509 combustion (e.g., vehicle and industrial processes) yields lower char and soot concentrations, while
510 low-temperature combustion (e.g., household cooking and biomass burning) results in higher ratios
511 (Han et al., 2016; Han et al., 2012; Han et al., 2010; Han et al., 2009a). Differences in char/soot
512 ratios between urban and rural areas across seasons may be linked to wheat straw burning,
513 contrasting with minimal vegetation combustion impacts in cities like Xi'an (Cao et al., 2005). The
514 char/soot ratio for dust-fall observed in this study (4.97; Figure 4, [and S9S10](#)) is slightly higher than
515 those recorded in Jinchang (3.84) (Han et al., 2009a) and Daheihe, Inner Mongolia (3.2) (Han et al.,
516 2008). The relatively stable concentration of soot-EC in this study, along with the elevated char/soot
517 ratio, suggests a correlation with higher coal consumption among local residents and industries. This
518 indicates that, in comparison to other regions, carbon emissions in the remote QDB are
519 predominantly sourced from coal and biomass burning, supporting [previous-OC/EC ratio](#) findings.
520 Furthermore, the char/soot ratio is elevated during HP, highlighting the predominant influence of
521 coal and biomass burning in rural areas during HP, while fossil fuel impacts are more pronounced
522 in NHP.

523

524 **3.4 POC and SOC compositions**

525 Aerosol samples with low OC/EC ratios typically exhibit low concentrations of POC, which

526 mainly comprises primary carbonaceous compounds. ~~Conversely, OC/EC ratios exceeding 2.0~~
527 ~~indicate substantial SOA formation (Chow et al., 1996; Gray et al., 1986). The MSRRS method~~
528 ~~enables discrimination of OC into POC and SOC (Method 2.7) (Yoo et al., 2022; Liu et al., 2023).~~

529 SOC constitutes a dominant fraction of OC in atmospheric aerosols. Research on carbonaceous
530 aerosols in various Chinese cities indicates that SOC contributes 67% (53-83%) and 57% (48-62%)
531 to rural and urban OC, respectively (Zhang et al., 2008)—marginally higher than the 62.62%
532 observed in this study (Figure 4, ~~and S4S11~~). Although SOC formation relies on solar radiation,
533 the QDB experiences high levels of solar energy (Liu et al., 2017), which may facilitate
534 photochemical oxidation of VOCs into SOC (Hama et al., 2022). Nevertheless, the consistently low
535 SOC concentrations indicate that VOCs emissions in QDB are significantly lower than the regional
536 averages observed across China, reflecting relatively low pollution levels. This finding is consistent
537 with the previously observed low concentrations of atmospheric carbon emission in the region.

538 During the NHP, rural areas exhibit the highest SOC/OC ratio of 86.70% (Figure ~~S4S12~~),
539 while urban areas record the lowest ratio of 44.32% during the HP. This trend reflects elevated
540 potential for photochemical activity and reduced contributions from POC, likely due to local
541 emission sources, such as traffic and coal combustion (Mbengue et al., 2018). The high SOC/OC
542 ratio suggests that SOC largely displaces OC. Our findings indicate that SOC levels are greater in
543 rural areas (66.52%) compared to urban regions (58.74%), likely attributable to significant coal
544 consumption for domestic heating, which enhances emissions of semi-volatile organic compounds
545 and organic gases (Dan, 2004). As for seasonal variations, studies in California show an increase in
546 SOC levels during warmer months, which is consistent with our results (Na et al., 2004). This
547 contrasts with the broader observation that higher temperatures typically lead to lower SOC
548 concentrations (Strader et al., 1999; Sheehan and Bowman, 2001). This discrepancy may stem from
549 varying sources of SOC emissions throughout the seasons, necessitating further investigation in
550 conjunction with other carbonaceous indicators.

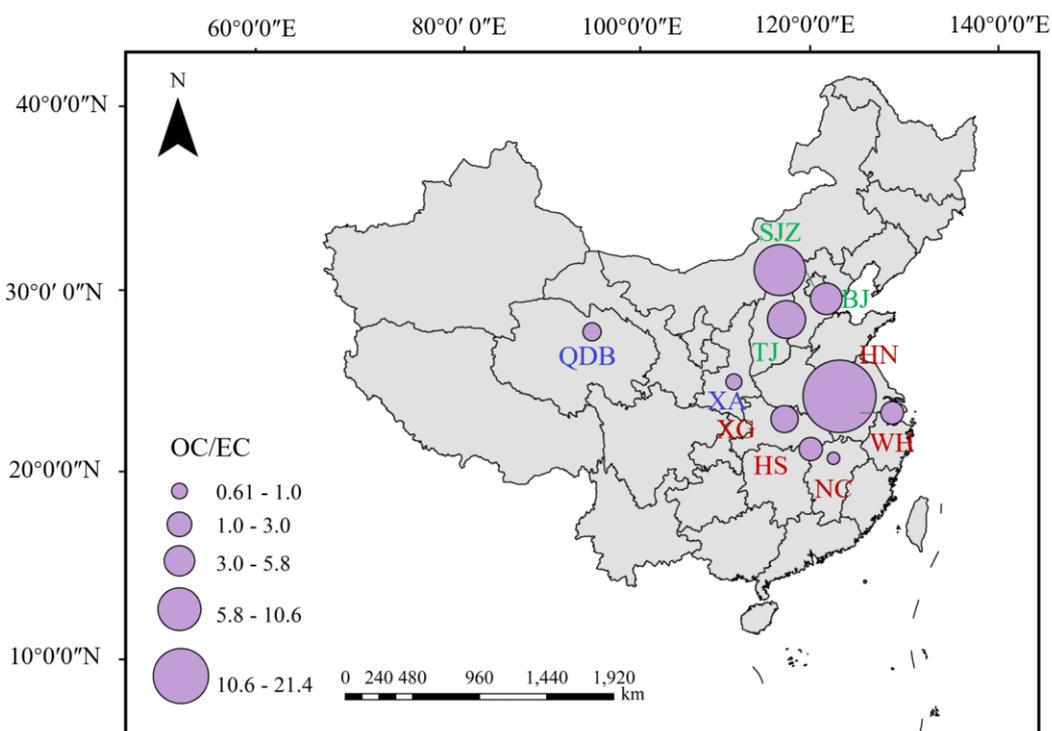
551 This study conducted a comparative analysis of carbonaceous element concentrations in
552 atmospheric dust-fall and road dust between the QDB and other global regions (Table S4). To ensure
553 data comparability, the selected road dust samples consisted of directly collected in-situ dust without
554 resuspension treatment. The results revealed that the concentrations of TC, OC, and EC in QDB
555 (3.27, 1.88, and 1.41 mg/g, respectively) were significantly lower than those in industrial or urban

556 areas such as Bolu, Turkey; New Delhi, India; and Ezhou, China, and were even lower than many
557 other Chinese cities (Table S4). The low concentrations of OC and EC in QDB indicate minimal
558 anthropogenic pollution influence in this region, and the data can represent the regional background
559 values of carbonaceous components in atmospheric dust-fall in the arid inland areas of East Asia
560 (Chen et al., 2019a). This is crucial for global models assessing the emission fluxes of carbonaceous
561 aerosols from dust source regions. In contrast, extremely high values of carbonaceous elements were
562 found primarily in urban road dust from locations like Bolu, Turkey (TC: 605.2 mg/g), Gwangju,
563 South Korea (TC: 31.97 mg/g), and Xi'an, China (TC: 36.53 mg/g), indicating strong influences
564 from traffic emissions (mainly non-exhaust emissions) (Wei et al., 2015; Lee et al., 2018; Demir et
565 al., 2022). For atmospheric dust-fall in major cities like New Delhi, India, and Wuhan, China, the
566 carbonaceous components are affected by the combination of traffic emissions (diesel vehicle
567 emissions being a major source of EC), industrial activities, and emissions from dense populations
568 (Deng et al., 2014; Zhang, 2014; Zhan et al., 2016; Mishra and Kulshrestha, 2017).

569 The OC/EC ratio in QDB (3.66) is at an intermediate level. It is much lower than that in regions
570 dominated by biomass burning, such as Kumasi, West Africa (17.07) and Huainan, China (21.4),
571 but is relatively close to ratios found in cities like Gwangju, South Korea (5.63) and Ulaanbaatar,
572 Mongolia (5.69), albeit with significantly lower concentrations (Lee et al., 2018; Bandowe et al.,
573 2019; Liu et al., 2020). We primarily analyzed the OC/EC ratios in cities across different regions of
574 China to reveal the influence of varying economic development levels.

575 Figure 5 presents spatial variations in urban OC/EC ratios across China. The findings reveal
576 that the Northwest region, represented by QDB urban and Xi'an (XA) (Han et al., 2009b), exhibits
577 a significantly lower ratio (1.59 ± 0.56) compared to central regions, including Nanchang (NC)
578 (Zhang, 2014), Huangshi (HS) (Zhan et al., 2016), Wuhu (WH) (Deng et al. 2014), Xiaogan (XG)
579 (Zhan et al., 2022), and Huainan (HN) (Liu et al., 2020), where the ratio is 5.86 ± 7.81 . This ratio is
580 also lower than that observed in eastern cities such as Beijing (BJ) (Tang et al., 2013), Tianjin (TJ)
581 (Ma et al., 2019), and Shijiazhuang (SJZ) (Guo et al., 2018), which have a ratio of 6.83 ± 2.77 . This
582 pattern is consistent with the trends in atmospheric PM OC/EC ratios (Xie et al., 2023), suggesting
583 that the carbon in the dust of the QDB urban primarily results from coal combustion and industrial
584 emissions, leading to elevated EC concentrations and lower OC/EC ratios (Liu et al., 2022).
585 Conversely, cities with higher economic development, such as Beijing and Tianjin, characterized by

586 greater population density and income levels, typically experience secondary pollution, resulting in
 587 higher OC/EC ratios. Furthermore, as the pretreatment used in this study removes impurities such
 588 as silicates from the carbonaceous components of dust, and the OC collection efficiency of this
 589 treatment is currently unknown, the lower OC/EC ratios observed may also be attributed to a lower
 590 OC collection efficiency.



591
 592 **Figure 5** Distribution of OC/EC ratios across various regions of China. (a), at different altitudes on
 593 the Qinghai Xizang Plateau (b), and the relationship between OC/EC ratios and altitude (c). Blue
 594 designations represent the Northwest region, red indicates the Central region, and green denotes the
 595 Eastern region. The color of the circles corresponds to altitude, while circle size reflects the
 596 magnitude of the OC/EC ratios.

598 The Char/Soot ratio in QDB is notably high at 5.04, significantly exceeding that of other
 599 regions such as Xi'an (0.99) and Wuhan (0.09) (Wei et al., 2015; Liu et al., 2021). Char-EC primarily
 600 originates from incomplete combustion processes like biomass burning and coal combustion. Soot-
 601 EC mainly derives from high-temperature combustion sources such as fuel oil and diesel vehicle
 602 exhaust (Han et al., 2009). The exceptionally high Char/Soot ratio in QDB strongly indicates that
 603 its limited carbonaceous components predominantly originated from relatively inefficient

604 combustion sources. These potentially included coal or small-scale biomass burning for local
605 residential/expedition activities (e.g., heating, cooking) and possibly long-range transported
606 biomass burning products (e.g., from forest/agricultural fires in South or Southeast Asia) (Han et al.,
607 2009; Han et al., 2006; Han et al., 2016). In contrast, the very low Char/Soot ratios observed in
608 cities like Wuhan and Xi'an clearly point to traffic source emissions as their primary contributor, a
609 finding likely influenced by the specific focus of those studies on road dust.

610 However, we fully recognize the fundamental differences in sources and composition between
611 road dust and atmospheric dust-fall. Road dust is primarily secondary dust formed from traffic
612 activities, construction dust, soil particles, and resuspended deposited atmospheric particles, with
613 its carbonaceous composition strongly reflecting intense local anthropogenic emissions (Casotti
614 Rienda and Alves, 2021). In contrast, atmospheric dust-fall integrates contributions from local
615 sources, regional transport, and even long-range transport. Therefore, direct comparison between
616 these two may introduce bias when interpreting regional pollution characteristics and the degree of
617 anthropogenic influence, which cannot be overlooked. Building on this analysis, the next phase of
618 this research will focus on the sampling and analysis of fine atmospheric particulate matter (PM_{2.5}
619 and PM₁) to more accurately elucidate the emission levels and environmental and climatic impacts
620 of carbonaceous aerosols in the QDB.

622 **3.5 Trace elements concentration**

623 The total concentration of major (Fe, Si, Al) and trace elements (Ti, Cr, Cd, Cu, Mn, Ni, Pb, V,
624 Zn) was determined to be 8.74 ± 5.82 mg/g, while arsenic (As) remained below the detection limit
625 in all analyzed samples. Crustally derived elements—Fe, Al, Si, and Ti—dominated the elemental
626 profile, aligning with dust composition patterns reported in Ira, Singapore, and Beijing-Hebei
627 regions, China (Joshi et al., 2009; Qiao et al., 2013; Eivazzadeh et al., 2021). In comparison to cities
628 such as Beijing, Shanghai, Xi'an, and Lanzhou, and Junggar Basin the levels of heavy metals in dust
629 from the QDB are relatively low (Jiang et al., 2018) (Supplementary Table [S4S5](#)).

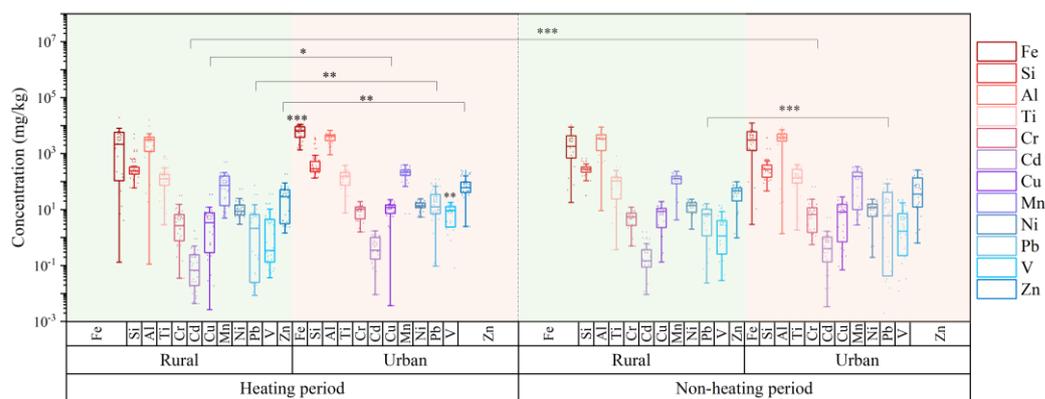
630 The low heavy metal content in dust deposition within the QDB can be attributed to the
631 following factors. The region has sparse human activity, lacks heavy industrial zones and dense
632 urban clusters, resulting in low total anthropogenic emissions of heavy metals. Furthermore, the

633 surface soil in the QDB itself has a low background level of heavy metals, primarily derived from
634 natural sources with relatively weak influence from human activities (Nuralykyzy et al., 2021; Chen
635 et al., 2021). From the meteorological perspectives, Meteorologically, the basin's high altitude,
636 strong winds, and arid conditions with minimal precipitation favor the dispersion of atmospheric
637 pollutants. This makes the formation of prolonged stagnant weather conditions unlikely, thereby
638 preventing the accumulation of pollutants and the occurrence of high concentrations near the ground.
639 A particularly unique aspect of the QDB is its role as a significant source of salt dust. The recent
640 studyResearch indicates that salt dust emitted from the playa lakes within the basin contributes
641 substantially to atmospheric dust deposition (Zhu et al., 2025). These salt dust particles, composed
642 mainly of soluble salts like NaCl and gypsum, may dilute the relative concentration of non-salt
643 components, such as heavy metals, when released into the atmosphere in large quantities. The
644 combined effect of these factors leads to the observed low heavy metal content in dust deposition
645 in this region.

646 Throughout both the HP and NHP, trace elements concentrations in urban areas were
647 consistently higher than in rural areas, with the exception of Ti (Figure 6; and S12S13). During the
648 HP, urban levels of Zn, Pb, and Cu were significantly elevated compared to rural areas, and Pb also
649 demonstrated a significant increase in urban during the NHP. In rural, the differences in metal
650 concentrations between HP and NHP were minimal. In contrast, urban areas exhibited higher
651 concentrations of all elements except for Ti, Cd, and Cr during the HP, with Fe and V showing
652 notably elevated levels compared to other regions. These variations in average concentrations
653 indicate that coal combustion for domestic heating in urban areas contributes to increased
654 atmospheric heavy metal levels (Duan and Tan, 2013; Meng et al., 2017). In contrast, Cd and Cr
655 exhibited mixed anthropogenic sources with limited coal combustion contributions, while Ti
656 concentrations remained stable across seasons, reflecting minimal anthropogenic influence.

657 Analysis of the carbonaceous components in QDB dust deposition reveals more intensive coal
658 combustion in rural areas, yet the heavy metal concentrations in atmospheric deposition are lower
659 than in urban areas. This observation can be explained by the following factors. First, pollution
660 sources in rural areas are relatively singular, whereas urban areas are influenced by more complex
661 heavy metal sources. During the heating period, heavy metals in the rural atmosphere of the QDB
662 mainly originate from coal and biomass combustion. In contrast, urban areas are affected by a wider

663 [range of sources, including industrial activities, traffic emissions, and others \(Liu et al., 2021; Huang](#)
 664 [et al., 2021a\). Additionally, the dense building layout in urban areas hinders pollutant dispersion,](#)
 665 [leading to accumulation, while the open terrain in rural areas facilitates dilution and diffusion. This](#)
 666 [pattern, where rural heavy metal concentrations \(particularly Pb, Cr, Cd, As, and other elements](#)
 667 [associated with coal combustion\) are lower than those in urban areas during the heating season,](#)
 668 [has also been observed in studies conducted in Northeast China, Shanghai, Taiyuan, the Yangtze River](#)
 669 [Delta, and Southern Nigeria \(Shi et al., 2012; Liu et al., 2021; Huang et al., 2021a; Liu et al., 2023b;](#)
 670 [Hilary et al., 2025\).](#)



671
 672 **Figure 6** Concentration of heavy metals by rural and urban settings during domestic heating and
 673 non-domestic heating periods. Significant differences are indicated by asterisks (* $p < 0.05$; ** $p <$
 674 0.01 ; *** $p < 0.001$).

676 3.6 Source apportionment

677 We conducted a PMF source apportionment analysis on soluble ions, trace ~~metals~~, and
 678 carbonaceous elements present in dust, specifically focusing on $ps-SO_4^{2-}$, $ps-Ca^{2+}$, $ps-K^+$, $ps-Mg^{2+}$,
 679 $ps-Cl^-$, $ps-Na^+$, $nps-SO_4^{2-}$, $nps-Ca^{2+}$, $nps-K^+$, $nps-Mg^{2+}$, $nps-Cl^-$, $nps-Na^+$, Fe, Si, Al, Ti, Cr, Cd, Cu,
 680 Mn, Ni, Pb, V, Zn, SOC, POC and Char-EC, Soot-EC. Seven source factors were identified based
 681 on prior research and an understanding of potential local sources: salt lakes, soil, vehicular
 682 emissions, secondary sources, biomass and coal burning, and industrial activities (Figure 7, [and](#)
 683 [S13S14](#)). [A plot of the time series is provided in Figure S15. The generation of Figure 7a involved](#)
 684 [extracting factors identified as the same source from the PMF factor profiles of each site \(Urban](#)
 685 [and Rural\) and each heating season \(HP and NHP\) shown in Figure S14. The arithmetic mean of](#)
 686 [the contributions from characteristic species \(elements and ions\) corresponding to each factor was](#)

687 calculated. Species with average contributions exceeding 20% were defined as characteristic species
688 of that source in atmospheric dust over the QDB.

689 The factor profiles for each element in these source categories represent the arithmetic mean
690 of profiles from six stations, with detailed operational methods provided in Supplementary Text S1.
691 The uncertainty of the source contributions was calculated directly from the standard error of the
692 multiple regression coefficients between the deposition flux (independent variable) and the source
693 contribution (dependent variable) at different monitoring sites (Belis et al., 2015; Manousakas et al.,
694 2017). The regression method assumes that all factors explaining the mass have been identified;
695 however, if a significant portion of the mass not directly related to the species in the PMF analysis
696 is omitted, the source contributions may be overestimated, which could be an important additional
697 source of uncertainty. The results are shown in Table S6. It must be noted that this method captures
698 only a portion of the total uncertainty, as it does not include errors from profile uncertainty or
699 rotational ambiguity. The low errors calculated by this method indicate a good model fit.

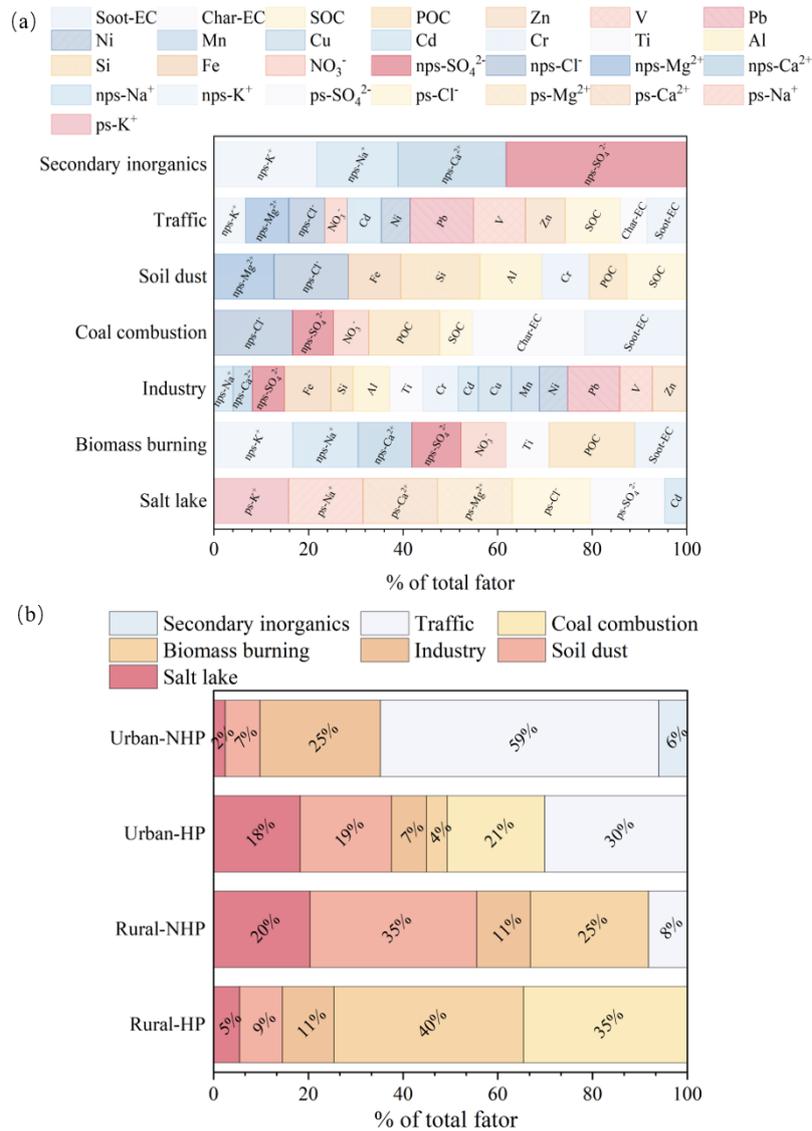
700 The ions ss-Na⁺, ss-Cl⁻, ss-SO₄²⁻, ss-Ca²⁺, ss-K⁺, and ss-Mg²⁺ are widely acknowledged as
701 indicators of sea salt (Ambade et al., 2022; Aswini et al., 2022; Gluscic et al., 2023). In this study,
702 we identified ps-Cl⁻ (82.71%), ps-Mg²⁺ (79.03%), ps-K⁺ (79.02%), ps-Na⁺ (78.69%), ps-Ca²⁺
703 (78.70%), and ps-SO₄²⁻ (78.69%) as key markers of salt lake sources. Furthermore, Cd (29.70%)
704 was detected at multiple sampling sites, indicating contributions from both salt lakes and industrial
705 emissions. Salt lake emissions were most pronounced in rural areas during the HP of 2023 at site
706 LTC and during the NHP of 2020 at site XZH. In urban areas, elevated contributions occurred during
707 the NHP and HP of 2022 at site GEM, and during the HP of 2023 at site DLX. The contribution of
708 salt lake sources in rural (12.93%) was significantly higher than in urban (10.33%). During the HP,
709 the proportion of salt lake sources in rural areas was 5.49%, compared to 20.37% in the NHP; urban
710 showed contributions of 18.24% during the HP and 2.42% in the NHP, showing opposite seasonal
711 trends. Backward trajectory simulations indicated that during the HP, airflows mass in both urban
712 and rural areas mainly originated from the northwestern part of the basin and the eastern Tarim
713 Basin, whereas during the NHP, they were broadly influenced by the salt lake regions within the
714 basin (Figure S6). The minor wind direction differences and the inter-distributed sampling points
715 (Figure 1), suggests no substantial geographical disparity between urban and rural sites. Additionally,
716 the ion content derived from playa salts in dust deposits increased during the NHP in both areas.

717 Therefore, we propose that the anomalous increase in salt lake contribution during the urban HP
718 may be closely related to human activities. The enhanced Urban Heat Island (UHI) effect~~urban heat~~
719 ~~island effect~~ and temperature inversion structures during the HP can alter boundary layer height,
720 turbulence, and deposition conditions, thereby increasing the residence time of externally
721 transported particles within the urban boundary layer and elevating their measured contribution
722 (Cichowicz and Bochenek, 2024). Urban heat sources and heating emissions may also modify local
723 transport pathways, leading to more concentrated deposition of dust originating from playa regions
724 over urban areas. Furthermore, dry road surfaces, increased traffic, and construction activities during
725 the HP can promote the ~~repeated~~ resuspension of previously deposited playa dust. The use and
726 subsequent resuspension of road de-icing salts (e.g., NaCl, CaCl₂) may further amplify the
727 contribution of tracer ions indicative of playa salts (Gertler et al., 2006; Casotti Rienda and Alves,
728 2021).~~This inverse trend suggests that seasonal variations differentially influence the contribution~~
729 ~~of salt lake sources in urban and rural, necessitating further research to elucidate the underlying~~
730 ~~driving factors.~~

731 The second factor pertains to soil dust, characterized by trace elements such as Si (37.17%)
732 and Al (29.18%), along with ions such as nps-Cl⁻ (34.90%), nps-Mg²⁺ (28.21%), and Fe (24.43%)
733 (Pervez et al., 2018; Tian et al., 2021). Additionally, the proportions of elements such as SOC
734 (28.20%) and POC (17.70%) suggest that the dust is likely mixed with fossil fuel emissions. Mg,
735 Al, Si, and Fe are typical tracers for soil dust (Liu et al., 2003; Heo et al., 2009). The temporal
736 variation of soil dust was largely consistent with that of the salt lake source, indicating the fact that
737 they may be influenced by similar factors. Notably, the contribution of soil dust in rural areas
738 (22.11%) exceeded that in urban areas (13.33%), indicating that soil dust is a major source of
739 atmospheric deposition. In urban areas, the contribution during the NHP was relatively low (7.36%),
740 likely due to higher wind speeds and the effectiveness of frequent summer precipitation (Zhang et
741 al., 2015a).

742 The third factor is ~~vehicular traffic~~ emissions, which are particularly pronounced in urban areas.
743 Key characteristic elements include Pb (59.52%), V (48.73%), nps-Mg²⁺ (40.78%), Zn (37.47%),
744 nps-Cl⁻(33.83%), Cd (32.43%), nps-K⁺ (29.64%), Ni (27.01%), -NO₃⁻ (20.99%) and SOC (51.71%),
745 Soot-EC (37.56%), Char-EC (24.94%) (Adeniran et al., 2017). ~~Furthermore, notable concentrations~~
746 ~~of nps-Mg²⁺ (40.78%), nps-K⁺ (29.64%), and NO₃⁻ (20.99%) were identified.~~ In rural areas,

747 vehicular emissions contributed 8.17% to atmospheric deposition during the NHP, whereas in urban
748 areas, the contribution was significantly higher at 45.13%, with 30.07% during the HP and 58.78%
749 in the NHP. These findings correlate with previous studies on OC/EC and char/soot ratios,
750 suggesting that carbonaceous elements in the NHP primarily derive from vehicular emissions. The
751 traffic emission factor in the QDB represents a mixture of vehicle exhaust and non-exhaust sources
752 (e.g., tire and brake wear, and resuspended road dust). Elements and ions including V, NO₃⁻, Ni, and
753 carbonaceous components primarily associated with vehicle exhaust (Cong et al., 2011; Zhang et
754 al., 2012). For instance, Ni can be emitted from fuel combustion and vehicle exhaust (Pacyna and
755 Pacyna, 2001). In contrast, elements such as Cu, Zn, nps-Mg²⁺, and nps-K⁺ originate from non-
756 exhaust vehicle emissions, including brake and tire wear, as well as the resuspension of road dust
757 (Amato et al., 2014). For example, Zn may derive from the wear of rubber tires on roads (Rogge et
758 al., 1993), Pb emissions may be related to wear (tires/brakes) (Smichowski et al., 2007), and Cu is
759 associated with brake wear (Lin et al., 2015). Furthermore, the presence of crustal elements and ions
760 such as Fe, Si, and nps-Mg²⁺ in the traffic emission factor for Urban-NHP, Urban-HP, and Rural-
761 NHP suggests an additional contribution from resuspended road dust (Chen et al., 2019). For traffic
762 emissions during the NHP, peaks were observed in both rural and urban areas in 2022, generally
763 concentrated in July, August, and September, coinciding with the tourism high season. During the
764 HP, traffic emissions primarily occurred in 2021 at GEM and in 2023 at DLX. Due to the impact of
765 the COVID-19 pandemic, tourist numbers in the QDB sharply declined in 2020 and 2021. The peak
766 in tourism activity in 2022 (Qinghai Statistical Yearbook, 2022) corresponded with the highest level
767 of traffic emissions during the three-year period, indicating a direct impact of tourism flux on
768 emission levels. However, given the relatively short sampling duration of this study (three years),
769 longer-term data and further research are needed to substantiate this hypothesis. The increase in
770 vehicular emissions during the NHP may be linked to the expanding tourism industry in the QDB,
771 particularly in cities like Golmud, which experience a rise in tourist numbers from May to
772 September, subsequently leading to a surge in population and vehicles, and thus elevating vehicular
773 emissions.



774

775 **Figure 7** Factor profile and contributions in urban and rural ~~area~~ area. (a) presents the factor profiles,
 776 represented as the arithmetic mean of individual elements across various locations, highlighting
 777 only those elements that constitute more than 20% of each profile. (b) illustrates the contributions
 778 of different sources at each location. [HP, domestic heating period; NHP, non-domestic heating
 779 period].

780

781 The fourth factor is coal combustion, characterized by high concentrations of nps-Cl⁻ (50.93%),
 782 nps-SO₄²⁻ (26.61%), NO₃⁻ (22.82%), and Char-EC (73.01%), Soot-EC (65.98%), POC (46.04%),
 783 SOC (21.12%) (Kundu and Stone, 2014; Contini et al., 2016). Zhang et al (2023): found that coal
 784 combustion emits particles rich in Cl⁻ (Zhang et al., 2023). Coal combustion was more intense at
 785 site LTC in rural areas and at site DLX in urban areas. Coal combustion occurs exclusively during

786 the HP, contributing 34.57% in rural areas and 20.63% in urban areas. These results align with earlier
787 studies on carbonaceous aerosols, indicating that the carbon content from coal combustion
788 emissions is higher in rural regions than in urban environments. Consistent with northern China, air
789 pollution in QDB urban has declined due to the adoption of clean heating technologies (Zhang et
790 al., 2021; Xue et al., 2023). However, rural coal combustion remains a major source of carbonaceous
791 aerosols during the HP.

792 The fifth factor, biomass burning, is characterized by significant concentrations of non-
793 precipitating species, including nps-K⁺ (42.76%), nps-Na⁺ (35.49%), nps-Ca²⁺ (29.23%), nps-SO₄²⁻
794 (26.56%), NO₃⁻(24.46%), Ti (23.41%) and POC (46.48%), Soot-EC (28.33%) (Simoneit, 2002;
795 Sulong et al., 2019). [K⁺ serves as an important tracer for biomass burning \(Cachier and Ducret,
796 1991\).](#) Biomass burning contributes 32.82% to rural atmospheric dust deposition, with a higher
797 proportion during the HP (39.21%) compared to NHP (26.44%). [Biomass burning made significant
798 contributions in urban areas in 2021 and 2022. In rural areas, biomass burning emissions were
799 particularly strong, especially during the HP at LTC and the NHP at XZH.](#) In urban, biomass burning
800 is primarily observed during the HP, contributing only 2.19%. These findings underscore biomass
801 burning as a major source of carbon emissions in rural settings, aligning with the prevalent use of
802 biomass fuels for cooking and domestic heating in Northern China's rural areas (Meng et al., 2019),
803 where 70% to 80% of energy demand are fulfilled by materials such as dung cakes, firewood,
804 charcoal, and crop residues (Tao et al., 2018; Shi et al., 2019). Furthermore, increased biomass
805 burning is also associated with the autumn harvest period (Chen et al., 2017; Li et al., 2021).

806 The sixth factor pertains to industrial emissions, which are characterized by high
807 concentrations of Pb (55.18%), Fe (48.91%), Cr (37.05%), Zn (36.42%), Ti (35.10%), Cu (34.91%),
808 V (34.44%), Ni (29.69%), Mn (29.59%) and Cd (21.29%) (Almeida et al., 2015; Yao et al., 2016b).
809 These elements were consistently detected across all sampling sites, alongside Al (38.84%), nps-
810 SO₄²⁻ (34.15%), Si (23.39%) and nps-Ca²⁺ (20.48%), nps-Na⁺ (20.02%), which indicate potential
811 contributions from soil dust. [Zn, Cu, Fe, and Mn are major chemical components in industrial
812 emission profiles; Cd is a trace element found in metallurgical industries \(Xu et al., 2022\), while Zn
813 and Mn emitte from oil combustion, metallurgy, and steel manufacturing processes \(You et al., 2017\).
814 Pb and Cd are associated with metal smelting and processing \(Fang et al., 2021\). Industrial
815 emissions showed greater contributions during the HP at XZH and at DLX.](#) In rural areas, industrial

816 emissions constitute 11.10% of carbon output, with contributions of 10.86% during the HP and
817 11.33% during the NHP. In urban areas, industrial emissions account for 16.41% overall, with 7.37%
818 during the HP and 25.44% in the NHP. ~~Moreover, industrial emissions are a primary source of~~
819 ~~atmospheric dust in urban regions, particularly in the QDB, which is rich in mineral resources and~~
820 ~~hosts numerous mining enterprises, thereby significantly contributing to regional air pollution. Due~~
821 ~~to the abundance of non-ferrous metal resources (e.g., lead-zinc ores), oil, natural gas, and saline~~
822 ~~minerals (e.g., potassium, lithium) in the QDB, the primary industrial activities are mining and~~
823 ~~associated chemical industries. Particularly around the GEM and DLX sites, the presence of~~
824 ~~numerous lead-zinc, iron, and copper mining enterprises leads to significant contributions from~~
825 ~~industrial emissions to urban dust fallout, making it one of the major sources of air pollution in the~~
826 ~~basin.~~

827 –The seventh factor is secondary inorganic aerosols, primarily composed of NO_3^- (68.54%),
828 nps-Mg^{2+} (41.04%), nps-Na^+ (39.01%) and nps-Ca^{2+} (30.79%) (Liu et al., 2015; Liu et al., 2016a).
829 High mass loadings of NO_3^- and SO_4^{2-} are characteristic of typical secondary inorganic aerosols
830 (Huang et al., 2021). Research indicates that NO_3^- and SO_4^{2-} primarily result from the conversion
831 of gaseous precursors, such as SO_2 and NO_x , through photochemical reactions, predominantly
832 sourced from local and regional emission (Liu et al., 2015; Tao et al., 2013). Secondary inorganic
833 aerosols are predominantly observed in urban areas during the NHP, where they contribute 6.00%
834 to total aerosol sources. This increase is likely due to elevated temperatures and enhanced solar
835 radiation during this period, which promote photochemical activity (Pandolfi et al., 2010). The
836 secondary inorganic aerosol source was identified only in Urban-NHP, peaking in August 2022 at
837 GEM and in June, July, and August 2022 at NMH. This trend closely followed the distribution of
838 traffic emissions, suggesting that the formation of secondary aerosols is linked to increased traffic
839 activity. Traffic emissions, particularly vehicle exhaust, are a significant source of secondary
840 inorganic aerosols (especially NO_3^-) in urban atmospheres (Ma et al., 2017). Source apportionment
841 studies in Beijing have similarly found that the contribution of secondary inorganic aerosols
842 increases in summer, closely associated with traffic emissions (Zhang et al., 2013).

843 Dust deposition sources exhibit significant seasonal and spatial variations. In the QDB, coal
844 combustion (27.60%) and biomass burning (22.21%) dominate during HP, transitioning to vehicular
845 traffic emissions (33.48%), soil dust (21.27%) and industry emission (18.39%) as the primary

846 contributor in NHP. Rural areas predominantly contribute to dust through biomass [burning](#) and coal
847 [burningcombustion](#), as well as natural sources like windblown dust and salt lake emissions. This
848 pattern aligns with increased coal usage for winter domestic heating and heightened biomass
849 burning for cooking and heating in rural [areas](#). In urban, dust deposition is briefly influenced by
850 anthropogenic activities, including [vehicletraffic](#) and industrial emissions, with minimal
851 contributions from domestic heating. Such differences can be attributed to varying economic
852 development models, industrial and energy structures, and levels of human activity (Kataki et al.,
853 2016).

854 This study observed that the contribution of biomass burning to atmospheric dust deposition in
855 rural areas of the QDB during the HP was higher than that of soil dust. Given that the collected dust
856 samples had particle sizes >10 μm , while biomass burning typically emits aerosols in the submicron
857 range, we propose several potential explanations. Firstly, during the HP, factors such as increased
858 soil moisture and snow cover significantly suppress soil dust emission, resulting in a lower intensity
859 than in other seasons (An et al., 2018; Yang et al., 2019). Simultaneously, biomass and coal burning
860 for heating increases substantially, leading to intense, short-term emissions of fine particles.
861 Although it was fine initially, these high-concentration ultrafine particles can undergo coagulation
862 or coalescence, aggregating with each other or onto pre-existing coarse particles, thereby increasing
863 their size (Butler and Mulholland, 2004; Kulmala et al., 2004; Li et al., 2020). Furthermore, fine
864 particles from biomass burning (e.g., carbonaceous materials) may mix internally with coarse
865 particles like soil dust or salt dust from the QDB, forming internally mixed particles (Li et al., 2003;
866 Hand et al., 2010). During source apportionment, such coarse particles are more likely to be
867 attributed to the biomass burning source. Additionally, the QDB is a significant source of salt dust
868 (Zhu et al., 2025). Salt dust particles (e.g., halite, gypsum) provide excellent condensation nuclei
869 for soluble substances emitted from biomass burning, greatly promoting hygroscopic growth (Li et
870 al., 2003; Kumar, 2010; Wang, 2013). The basin's topography also favors stable inversion layers,
871 inhibiting pollutant dispersion and allowing particles more time to grow, mix, and age in the
872 atmosphere. [Prevailing wintertime winds](#)~~Winter prevailing winds~~ may also transport pollutants from
873 surrounding regions into the basin.

874 Moreover, the dust in this study was collected using a passive sampler. Over 90% of the
875 collected dust particles were smaller than 100 μm , with approximately 25% less than 10 μm (Figure

876 S16), indicating the presence of fine particles (<10 µm), albeit in a relatively small proportion. The
877 particle size distribution of atmospheric dust deposition is similar to that of TSP, with both primarily
878 consisting of particles smaller than 100 µm. Using PMF source apportionment, this study identified
879 a notably high contribution from biomass burning in rural areas, particularly during the heating
880 period. Similarly, studies on atmospheric TSP in Iran (Ashrafi et al., 2018), the Qinghai-Xizang
881 Plateau (Lulang) (Zhao et al., 2013), Northeast China (Jia et al., 2024), and Qingdao (Liu et al.,
882 2022) have also reported significant contributions from biomass and coal combustion. This suggests
883 that contributions from biomass and coal combustion can indeed be observed in particles larger than
884 10 µm. Finally, the PMF model may have uncertainties in resolving sources with similar chemical
885 profiles. If the chemical compositions of local soil dust and biomass burning particles overlap after
886 long-range transport and complex atmospheric reactions, the model might not fully separate them
887 (Cesari et al., 2016).

888

889 **3.7 Environmental implication**

890 The source apportionment analysis using the PMF model indicates that in the QDB, rural dust-
891 fall predominantly originates from the combustion of solid fuels, —including firewood, yak dung,
892 and coal, —accounting for approximately 74.61% of the total contribution. This proportion
893 significantly exceeds contributions reported for rural areas in Beijing (41%) (Hua et al., 2018), Agra
894 (54.3%) (Agarwal et al., 2020), and Beihai, Guangxi Province (66.7%) (Zhang et al., 2019).

895 The higher contribution in this study likely reflects the local energy profile, as the sampling
896 site in Haixi Mongol and Xizang Autonomous Prefecture, Qinghai Province, primarily relies on coal,
897 yak dung, and firewood, constituting 58%, 23.5%, and 13% of rural energy consumption,
898 respectively (Jiang et al., 2020; Shen et al., 2021). In contrast, solid biomass fuels, including wood
899 and yak dung, account for over 70% of rural household energy consumption in Xizang, with yak
900 dung alone representing 53% (Liu et al., 2008; Xiao et al., 2015). Similar patterns emerge in South
901 and Central Asia, where biomass fuels dominate residential heating (firewood: 39%; dung: 29%)
902 (Amacher et al., 1999; Heltberg et al., 2000; Hoeck et al., 2007; Foysal et al., 2012; Behera et al.,
903 2015; Kerimray et al., 2018). In northern China, rural domestic heating primarily relies on coal
904 (46%), firewood (23.8%), and electricity (15.1%) (Tao et al., 2018), further highlighting the unique

905 energy composition of QDB.

906 Recent studies have shown that South Asia, Central Asia, and Xizang contribute significantly
907 to high concentrations of atmospheric PM, particularly BC, which accelerates glacier melting in the
908 QXP (Ming et al., 2010; Xia et al., 2011; Chen et al., 2015). The QDB is recognized as a significant
909 dust source affecting the glacier surfaces on the QXP, although it is often overlooked (Dong et al.,
910 2014; Wei et al., 2017; Zheng et al., 2021). Compared to other dust sources, the QDB exhibits higher
911 emissions from coal combustion, giving it a unique influence on the QXP. The organic matter and
912 pollutants, such as polycyclic aromatic hydrocarbons (PAHs), released from household solid fuel
913 combustion, particularly coal (98%) and dung (94%), are substantially higher than those from
914 firewood (Leavey et al., 2017; Secret et al., 2017; Ye et al., 2020). Consequently, the impact of PM
915 from coal combustion in the QDB on the QXP is significant. Specifically, the presence of BC in PM
916 increases glacier albedo, accelerating the melting of glaciers and snow in the region (Kang et al.,
917 2020), and impacting global freshwater resources (Huss and Hock, 2018). Additionally, BC
918 enhances cloud condensation nuclei (CCN), ice number concentration, and cloud cover (Zhou et al.,
919 2025), thereby influencing global climate change. Furthermore, coal combustion releases harmful
920 emissions, including CO₂, NO_x, CO, SO₂ and sulfur trioxide (SO₃) (Munawer, 2018), adversely
921 affecting local human health and exacerbating climate warming on the QXP (Liu et al., 2006; Li et
922 al., 2023), with broader implications for global climate. Therefore, the atmospheric pollutants
923 emission of the QDB deserves considerable attention. However, this study focuses primarily on
924 larger particles, indicating a need for further research on the environmental impacts of carbonaceous
925 aerosols in atmospheric PM within the QDB.

926 In addition to the distinctive energy consumption structure in the rural QDB, which leads to
927 significant contributions from coal and biomass burning during HP, atmospheric dust deposition in
928 the QDB during the NHP primarily originates from traffic and industrial emissions. The contribution
929 from traffic emissions during the NHP was twice that during the HP. Considering the larger particle
930 size of the dust samples collected in this study, the traffic-related dust is likely derived mainly from
931 vehicle non-exhaust emissions, such as road dust (Gondwal and Mandal, 2021). Beyond the unique
932 energy structure of QDB where coal and biomass dominate during the HP NHP atmospheric
933 dust primarily stems from vehicle and industrial emissions, with vehicle contributions being twice
934 as high as in the HP. This indicates that NHP atmospheric conditions are significantly influenced by

935 resource development and tourism. Sampling sites, such as Qarhan Salt Lake, along with GEM and
936 LTC stations within approximately 100 km (Figure 1), suggest that salt lake resource extraction has
937 a lower impact on regional aerosols than [vehicletraffic](#) emissions, despite salt lakes being the
938 primary resource. This is likely because salt lake development mainly involves solar evaporation
939 and chemical processes like extraction and adsorption, which emit fewer pollutants compared to
940 other mining methods (Zhen, 2010). Consequently, salt lake resource exploitation exerts a relatively
941 minor effect on local atmospheric carbonaceous aerosol.

942 Similar salt lakes with comparable environments to QDB—, such as Salar de Uyuni in Bolivia,
943 the Atacama Salt Lake in Chile, and Ombre Muerto in Argentina—, are rich in lithium resources (Li
944 et al., 2014), making them focal points for resource development. Additionally, Salar de Uyuni, the
945 Atacama Salt Lake, Junggar Basin, and the Great Salt Lake are renowned tourist destinations. This
946 suggests that, in arid basin salt lakes with similar climates and intensive human activity, atmospheric
947 carbonaceous aerosols are likely influenced by resource exploitation and tourism, especially tourism.
948 The study's findings can inform policy decisions regarding unexploited salt lakes in South America,
949 such as Ombre Muerto and Salar de Uyuni. However, while QDB also hosts mineral resources such
950 as copper, iron, and tin, this research focused on larger particles ($>100\ \mu\text{m}$), which are more
951 indicative of local sources. Given that sampling was conducted around the salt lakes, potential
952 impacts from other mineral resource developments may have been underestimated. Further research
953 is necessary to fully assess the environmental effects of carbonaceous aerosols in QDB atmospheric
954 particles.

955

956 **Conclusion**

957 This study analyzed the composition of dust deposition at six sampling sites in the southern
958 Qaidam Basin from January 2020 to March 2023 and examined DF, soluble ions, [metaltrace](#), and
959 carbonaceous element content in urban and rural samples during both domestic heating and non-
960 domestic heating periods. Through integrated application of backward trajectory modeling, PMF,
961 and carbon speciation indices, we identified dominant dust sources and evaluated domestic heating
962 impacts on atmospheric processes in remote regions.

963 The findings revealed that DF and carbon emissions were significantly higher in rural than in

964 urban areas. Among carbon indicators, urban ~~areas~~ exhibited elevated EC levels (1.46 ± 1.60 mg/g),
965 while OC levels were higher in rural (2.25 ± 1.92 mg/g). ~~Both altitude and e~~conomic development
966 increase OC/EC ratios, but they are driven by different intrinsic factors. Char-EC was the dominant
967 contributor to EC (80.44%), with urban Char-EC levels (85.00%) showing a notable increase
968 compared to rural levels (75.88%). SOC was the principal contributor to OC (68.17%), with rural
969 SOC levels surpassing (67%) those in urban areas (57%). The OC/EC and char-EC/soot-EC ratios,
970 along with PMF results, indicated that during HP, dust deposition in the QDB was primarily derived
971 from coal ~~combustion~~ (28.44%) and biomass burning (22.14%), while ~~vehiele~~traffic emissions
972 accounted for 34.19% of dust during NHP. Coal and biomass burning were the main contributors to
973 rural dust, strongly influenced by domestic heating, whereas urban dust predominantly originated
974 from ~~vehiele~~traffic (45.13%) and industrial emissions (16.41%). Compared to other dust sources in
975 the Qinghai-Xizang Plateau, coal consumption in the QDB is higher during the domestic heating
976 period. The resulting emissions of black carbon and greenhouse gases may exacerbate glacier
977 melting in the region, warranting increased attention. Given the distinctive carbonaceous aerosol
978 signatures identified in the Qaidam Basin, we recommend prioritizing their radiative forcing effects
979 in regional environmental policymaking and climate modeling frameworks. Furthermore, findings
980 of this study offer a valuable scientific basis for understanding atmospheric carbonaceous aerosols
981 in arid basins and salt lake regions with climates similar to QDB. They can particularly inform
982 policy decisions regarding unexploited salt lakes in South America, such as Ombre Muerto and Salar
983 de Uyuni.

984 However, this study primarily focused on larger-scale PM and examined the effects of heating
985 on carbonaceous aerosols in the QDB. It lacks an investigation of aerosols with smaller particle
986 sizes (e.g., PM_{10} , $PM_{2.5}$, PM_1), which is essential for a comprehensive understanding of
987 carbonaceous aerosol characteristics in this unique region. Furthermore, in addition to offline
988 observations, future research should incorporate online observations with high spatiotemporal
989 resolution and utilize numerical air quality models such as CMAQ, CAMx, WRF-CHEM, and
990 NAQPMS to analyze the spatiotemporal distribution and future trends of carbon aerosols in the
991 QDB.

992 993 **Author contribution**

994 HZ: Conceptualization, data curation, formal analysis, funding acquisition, investigation,
995 methodology, project administration, validation, writing – original draft.

996 LZ: Data curation, formal analysis, methodology.

997 SZ: funding acquisition, validation, writing – review & edited.

998 XZ: Supervision, conceptualization, funding acquisition, writing – review & edited.

999

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1006

1007 **Declaration of competing interest**

1008 The authors declare that there is no conflict of interest.

1009

1010 **Data availability**

1011 Datasets for this research has been uploaded in Zenodo and is available at
1012 <https://doi.org/10.5281/zenodo.14382853> (Zhu, 2024).

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