



1 **Developing A Custom-Built Metal Cloud Chamber: Analysis of Aerosol**

2 **Coagulation at Low Humidities.**

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10

11 **Abstract**

12 We are developing an intermediate size (906 L) cloud chamber, and this paper reports on the
13 design and initial characterization of dry aerosol experiments. Specifically, we are determining
14 wall-loss and coagulation correction factors using the observed size distribution measurements
15 for surrogates of common aerosol classes: sodium chloride, sucrose, and soot. Results show
16 that, on average, sodium chloride, sucrose, and soot wall-loss rates converge to similar values
17 on relatively short time scales (<1 hour). The fitted coagulation correction factor, W_C^{-1} , for soot
18 particles (1.23 ± 0.312), indicates they adhere to each other more than sodium chloride



19 (0.969 ± 0.524) and sucrose (1.16 ± 1.38). This study lays the foundation for future experiments
20 at elevated humidity and supersaturation conditions to characterize the influence of particle
21 shape on coagulation and cloud parameters.

22 **1 Introduction**

23 Aerosol-cloud interactions remain one of the largest sources of uncertainty in the Earth's
24 radiation budget. By directly scattering, absorbing solar radiation and indirectly influencing
25 cloud formation, aerosols affect longwave and shortwave radiation in the Earth's atmosphere¹.
26 Despite sustained research efforts, these impacts still pose significant challenges to our
27 understanding of the aerosol cooling effect, estimated at $-0.86 \pm 0.56 \text{ W/m}^2$, and the effective
28 anthropogenic radiative forcing of Earth's climate (estimated at $-1.25 \pm 0.85 \text{ W/m}^2$)¹. The
29 complexity of aerosol sources, properties, and processing continues to hinder precise
30 quantification of these forcing estimates.

31 A critical source of aerosols is wildfire smoke, which can influence radiative budgets up
32 to a year depending on the transport and evolution of plumes²⁻⁴. Under extreme burning
33 conditions, wildfires can generate pyrocumulonimbus clouds, lofting large concentrations of
34 aerosol into the upper troposphere and lower stratosphere^{2,5,6}. These smoke particles can exert
35 prolonged effects on climate through chemical and physical processes such as condensation and
36 coagulation⁷⁻⁹. The fractal nature of soot particles further complicates our understanding of
37 their indirect effects on cloud formation and radiative properties¹⁰⁻¹². For instance, during the
38 Amazon biomass burning season, Koren et. al.¹³ reported a dramatic reduction in cumulus cloud



cover—from 38% under cleaner conditions to 0% during heavy smoke. However, Kaufman & Koren et. al.¹⁴ observed an increased cloud cover in regions with higher column aerosol concentrations. These discrepancies underscore the complexity of aerosol-cloud interactions, which depend on various factors such as aerosol composition, hygroscopicity, size distribution, supersaturation, and the prevailing atmospheric stability¹⁵. As wildfires increase in frequency and intensity due to climate change¹⁶, refining our knowledge of how these aerosols evolve and ultimately affect cloud development is crucial for improving climate models and future predictions.

Beyond large-scale aerosol effects, aging aerosols undergo microphysical transformations that can drastically alter their role in cloud processes. Condensation of organics and the mixing of sulfate with black carbon (BC) have both been shown to influence cloud dynamics¹⁷. Recent work indicates that larger BC agglomerates may form preferentially at cloud tops, while the heaviest-coated BC particles are most likely to be scavenged by cloud droplets^{18,19}. Modeling these highly dynamic processes remains challenging, as it requires accurately representing particle growth, mixing states, and cloud interactions^{20–24}.

Cloud chambers are valuable research tools for investigating microphysical mechanisms under well-controlled conditions^{25–28}. Existing cloud chambers are their own institutional facility in the case of CLOUD at CERN²⁹, AIDA Chamber EUROCHAMP³⁰, and PI-chamber at MTU²⁶. All chambers however, come with artifacts—most notably, the loss of particles to chamber walls through gravity, diffusion, convection, and electrostatic forces^{31–34}. Previous studies have



59 highlighted the importance of accounting for both size-dependent and time-dependent wall
60 losses^{35,36}.

61 In this paper, we introduce the development of a Los Alamos National Laboratory (LANL)
62 cloud chamber, which is specifically designed to investigate coagulation processes under
63 simulated conditions. We present initial data from experiments where aerosols were injected in
64 a dry environment to quantify losses to chamber walls, dilution, and coagulation effects.
65 Different types of aerosols were examined to validate known aerosol behaviors and characterize
66 coagulation. We further demonstrate the use of a python based aerosol package, Particula³⁷, to
67 model coagulation and wall-loss rates. Through these studies, we aim to refine experimental
68 design and advance understanding of how aerosols—particularly soot—undergo physical
69 transformations that shape their role in cloud formation and climate forcing.

70

71 **2 Chamber Development and Methods**

72 ***2.1 Setup of chamber and experiments***

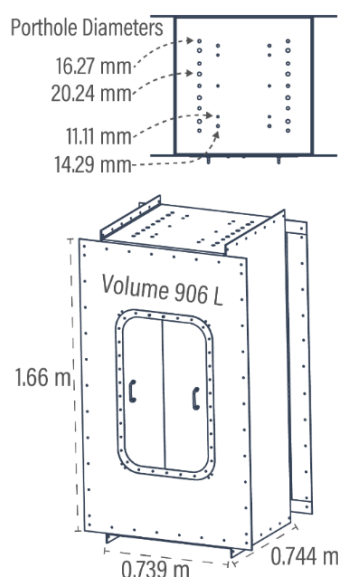
73 The LANL chamber is in the first phase of development with control of both temperature
74 and humidity to be added in future work. The 906 L chamber is made of 6 stainless steel walls
75 which are inert and reduce the effects of electrostatic charge. The rectangular body and
76 dimensions are shown in Figure 1. The chamber's joints are sealed with a fast cure marine
77 adhesive caulk (Sika, Sikaflex 291) and the outside junctions where the walls intersected were



78 sealed using ZIP System™ Stretch Tape (Huber Engineered Woods). Portholes were made for the
79 top and bottom plates where wires and probes be placed for measurements in the chamber and
80 for aerosols to flow in and out of the chamber. Unused portholes are sealed with Swagelok caps
81 and those used for probes and wires are sealed with a rubber gasket or a Teflon ferrule. A leak
82 test was performed by pressurizing the chamber by feeding clean air in and sealing every outlet
83 then seal any leak detected.

84 Copper tubing lines (3/8") are used to supply aerosols to the chamber and deliver
85 outflow sampling to instrumentation. Zero-air generators (T701 Teledyne Inc., USA) provide
86 clean dry air to push aerosol to the chamber and additional dilution air using Teflon tubing
87 (1/4"). Push flow enters at the bottom of the chamber, creating an upwards direction of flow.
88 Aerosols are sampled from an outlet at the top of the chamber. A dilution flow is connected to
89 the outlet line (88.9 mm from the outlet) to control aerosol concentrations and prevent
90 overwhelming the sampling instruments. A minimum sampling flow rate of 1.5 L/min was
91 needed to supply the instruments and we used a 1:5 ratio of push to dilution for the
92 experiments presented here. This infers a residence timescale within the chamber of 604
93 minutes (10 hours) and half-life of 418 minutes (6.9 hours). The flow rates are controlled with
94 mass flow controllers (MFC; Alicat). Prior to each experiment the chamber was flushed by
95 pushing clean air with a flow of ~10 L/min for at least 3 hours to reach background (~0-10 cm⁻³).

96



97

98 Figure 1. Schematic of LANL's 906 L chamber. The chamber has external dimensions of 1.66 m in
99 height, 0.739 m in width, and 0.744 m in depth. The design includes 56 portholes with
100 diameters ranging from 11.11 mm to 20.24 mm, shown across the top and mirrored on the
101 bottom.

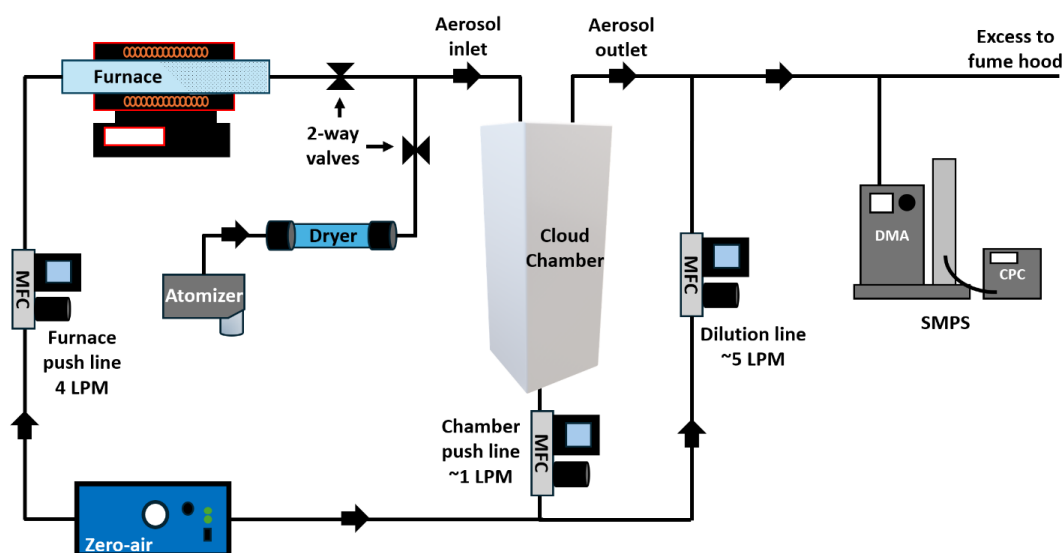
102 **2.2 Aerosol Generation and Instrumentation**

103 Two aqueous solutions and controlled combustion of dried biomaterial were used as the
104 sources of aerosols. Sodium chloride (NaCl; Sigma-Aldrich) was chosen because it is a well-
105 understood compound in aerosol studies. Sucrose (Sigma-Aldrich) was used to act as a
106 secondary organic aerosol surrogate, and it is also a well-studied aerosol. Each were dissolved in
107 deionized water (Milli-Q, 18.2 MΩ) in separate solutions and were put on an Atomizer Aerosol
108 Generator (3079, TSI Inc., USA). The particles coming out of the atomizer passed through a silica



109 gel diffusion drier at a generation flow rate of ~2.4 L/min. The duration of aerosol injection
110 varied based on the desired number concentration. To generate soot, 0.1 – 0.5 g samples of
111 dried biomaterial *Poa pratensis* (Kentucky bluegrass) were weighed out, placed on a quartz boat
112 and into a quartz-tube furnace (Carbolite Gero, TS1-1200, Verder Scientific, UK) that was set to
113 1000°C for a flaming combustion condition. Soot particles were pushed to the chamber by zero-
114 air at 4 L/min for 5 minutes, an estimated time for complete combustion of the sample.

115 Aerosol size and number distributions downstream of the cloud chamber were
116 measured with a scanning mobility particle sizer (SMPS) that consists of a Differential Mobility
117 Analyzer (3081 DMA, TSI Inc., USA) and a Condensation Particle Counter (3752, TSI Inc., USA).
118 Measurement settings were set to continuously scan for 3 minutes/scan; 160 seconds recording
119 with 20 seconds of purging, measuring sizes 15.7 – 764.5 nm. Our experimental matrix
120 consisted of 5 repeats of NaCl, 4 repeats of sucrose and 6 soot experiments with varying
121 biomaterial mass, they are outlined in Supplement Information Table 1. In all experiments the
122 first 6 hours of data were used to analyze results.



123

124 Figure 2. Schematic of experimental setup showing how aerosols are injected and sampled from
125 the chamber.

126

127 3 Theory on Chamber Processes

128 The processing of data from the LANL chamber experiments involved two key steps to
129 analyze the underlying aerosol processes of coagulation, wall loss, and dilution (chamber push
130 line). First, we determined the observed size-dependent particle rates: $dN(D_p)/dt$. The
131 measured size distributions were fitted to a two-mode lognormal distribution. The lognormal
132 distribution parameters were optimized using the Python library SciPy's optimization routines,
133 with the mean squared error as the cost function. We used multiple minimization methods and
134 selected the best fit for each timestep based on the highest Pearson R-squared value. The



135 methods included Nelder-Mead (Simplex algorithm), Powell's method (Powell's conjugate
136 direction method), L-BFGS-B (Limited-memory Broyden-Fletcher-Goldfarb-Shanno with Box
137 constraints), TNC (Truncated Newton Conjugate-Gradient method), SLSQP (Sequential Least
138 Squares Programming), and trust-constr (Trust Region Constrained method).

139 Second, we fitted these observed rates to theoretical rates calculated from Particula³⁷, a
140 python-based aerosol microphysics package. The first step was to generate a new time series at
141 a higher size resolution (log-spaced 250 bins), extrapolating to lower (20 nm) and upper (4 μm)
142 diameter limits. The size-dependent particle rate was then computed as the linear slope of 21
143 point moving window (10 before and 10 after). This final rate was subsequently used to fit the
144 underlying aerosol processes in Equation 1 where $N(D_p)$ represents the number concentration
145 of particles of diameter, D_p , K_{12} is the coagulation kernel, W_c^{-1} is the coagulation correction
146 factor, N_1 and N_2 are the concentrations of particles in the bins for K_{12} , k_{flow} is the dilution
147 rate, and β is the wall-loss rate.

$$148 \quad \frac{dN(D_p)}{dt} = W_c^{-1} K_{12} N_1 N_2 - k_{flow} N(D_p) - \beta N(D_p) \quad \text{Equation 1}$$

149 The coagulation term is governed by a Brownian Coagulation kernel, K_{12} , that captures
150 the collision frequency between bin number concentrations (N_1 and N_2). This kernel is
151 described in Seinfeld and Pandis³⁸ (Section 13; Fuchs form with alpha efficiency form of 13.56),
152 and calculated with Particula. Since K_{12} does not account for other interactions (e.g. Coulomb
153 interactions) that may lead to coagulation, W_c^{-1} , the coagulation correction factor, was



154 determined. In our analysis, W_c^{-1} is a free fit parameter to allow for un-modeled behaviors to
155 be represented. The dilution rate, $k_{flow} = Q/V$, characterizes how the clean air flow rate (Q) is
156 used to push sample flow out of the chamber volume (V). Finally, the wall-loss term, $\beta N(D_p)$,
157 accounts for the size-dependent removal of particles to the chamber walls.

158
$$\beta = \frac{1}{LWH} \left(\frac{4H(L+W)\sqrt{k_e D}}{\pi} + v_p LW \times \coth \left(\frac{\pi v_p}{4\sqrt{k_e D}} \right) \right) \quad \text{Equation 2}$$

159 Equation 2 shows the wall-loss rate (β) varies with particle size, derived from a rectangular-
160 chamber formulation adapted from Crump and Seinfeld³⁵ and Crump³⁶. It incorporates both
161 diffusion-driven transport and gravitational settling. In this formulation, L , W , and H denote the
162 chamber's length, width, and height, respectively; k_e is the eddy wall diffusivity (a free fit
163 parameter); D is the particle diffusion coefficient; and v_p is the particle gravitational settling
164 velocity.

165

166 **4 Results and Discussion**

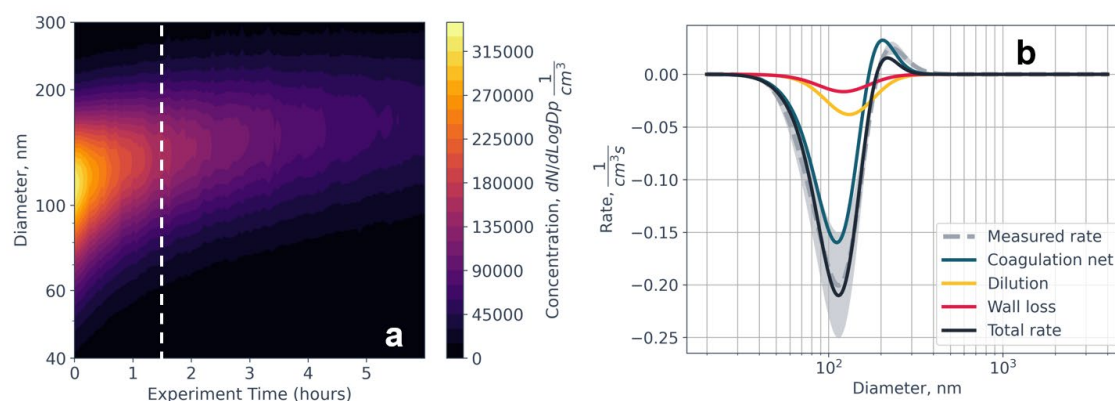
167 **4.1 Example Analysis**

168 We show, in Figure 3, the L-BFGS-B optimization routine that was used on Equation 1 for
169 experimental data from the smoke aerosol generated by combusting Kentucky bluegrass. Figure
170 3a shows the lognormal-fitted size distribution for the entire experiment, where particle growth
171 is evident as the mode diameter shifts to larger sizes over the six-hour period. Figure 3b breaks



down the observed rates after 1.5 hours into three calculated, time-varying, size-dependent components: coagulation, dilution, and wall-loss. At that time, coagulation dominates, reducing particles around 100 nm ($\sim 0.16 \text{ cm}^{-3} \text{ s}^{-1}$) and forming larger particles around 200 nm. From these fits we are specifically interested in the kernel correction factor to better understand the importance of agglomeration of freshly emitted BC fractal-like particles and how it changes in time.

178



179

Figure 3. **a)** Time series of the lognormal-fitted size distribution and concentration for a soot experiment. The dashed line marks a time slice at approximately 1.5 hours. **b)** At this time slice, particle loss rates are calculated, revealing both loss and gain of particles due to coagulation. (The time series of rates for individual aerosol species are provided in the Supplement.) In this panel, the dashed gray line represents the measured rate with uncertainty (shaded gray), while the blue, yellow, and pink lines correspond to the coagulation process, dilution, and wall loss, respectively.

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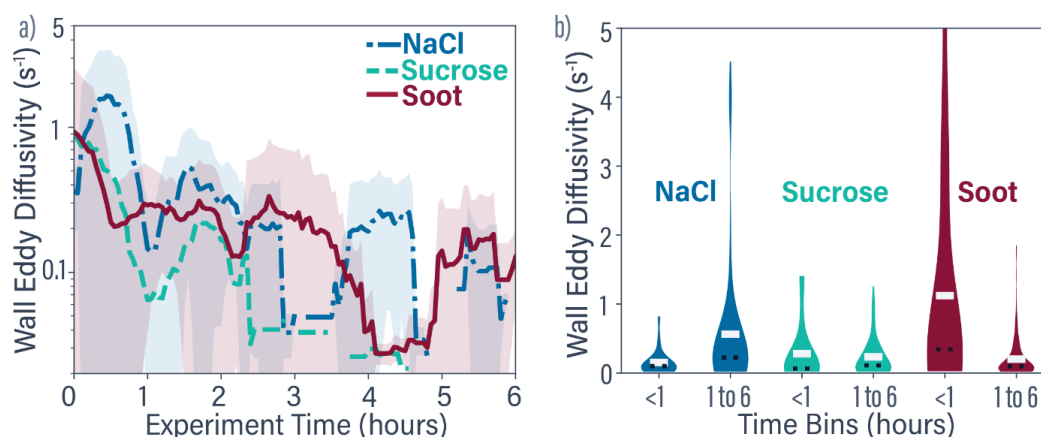
188 **4.2 Wall-loss Comparison**

189 In Figure 4a we show the average wall-loss rates for our three different aerosol types
190 based on 4–8 experiments each. Only results with valid optimizations and an R-squared above
191 0.85 were included. In the sucrose experiments, this filter led to data gaps during the later time
192 periods (2–6 hours) for inclusion in the analysis. To better compare with soot, we conducted
193 additional NaCl experiments to have a more complete timeseries for one of the comparisons.
194 The wall-loss rates during the first hour (< 1 hour) follow a similar trend apart from NaCl starting
195 at a low wall-loss rate then rising close to a rate of 2 s^{-1} . These initial wall-loss rates are
196 consistent with the general observation in chamber studies that early mixing processes and
197 injection conditions can dominate particle loss. Typical ranges reported in smog-chamber
198 experiments span from $< 1 \text{ s}^{-1}$ to tens of s^{-1} depending on injection flow and the use of a fan³⁹,
199 particle species^{40,41}, and chamber geometry⁴⁰. Over longer times (> 1 hours), all three aerosol
200 types converge toward similar wall-loss rates, in agreement with the literature indicating that
201 chamber turbulence and gravitational settling diminish over time as mixing subsides.

202 Figure 4b shows the statistical distribution of the wall-loss rates for each aerosol type
203 during the first hour and the subsequent five hours. NaCl and sucrose do not exhibit a large
204 variance in diffusivity for the first hour compared to soot which is $1.12 \pm 1.55 \text{ s}^{-1}$. NaCl, sucrose,
205 and soot show mean wall-loss rates of $0.562 \pm 0.975 \text{ s}^{-1}$, $0.233 \pm 0.286 \text{ s}^{-1}$, and $0.201 \pm 0.267 \text{ s}^{-1}$,
206 respectively. This convergence to relatively similar values is consistent with past observations in



207 smog-chamber experiments, where turbulent mixing dissipates, and the system approaches a
208 quasi-steady loss rate such as the CMU Teflon chamber^{33,34}, the CESAM chamber⁴⁰, and the AIR
209 chamber³⁹. However, NaCl and soot sucrose experiments display greater variability than soot,
210 likely due to residual chamber turbulence and differences in particle surface charge stemming
211 from their distinct generation methods (aerosolization vs. combustion).



212

213 Figure 4. a) Average time series of the calculated wall eddy diffusivity for NaCl (blue), sucrose
214 (green), and soot (red). Only fits with valid optimizations and r-squared greater than 0.85 are
215 included. b) Violin plots showing the mean (.), median (white bar) and overall distribution
216 range of wall eddy diffusivity values for each aerosol type in two-time bins (<1 hour and 1–6
217 hours). The width of each colored region represents the relative density of data points at that
218 value.

219



220 **4.3 Coagulation Corrections**

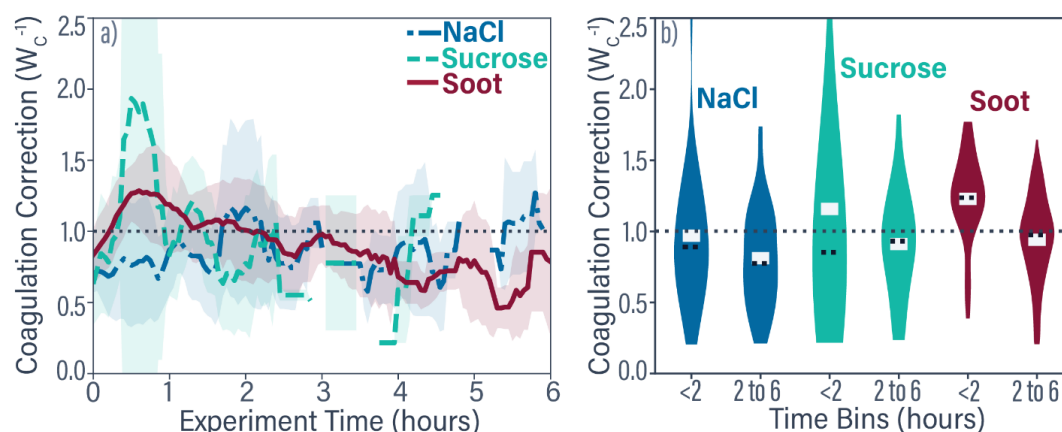
221 To investigate the influence of interparticle forces on aerosol coagulation, we fitted a
222 coagulation correction factor that would account for van der Waals forces, shape, and/or
223 Coulomb interactions in the coagulation rate. When $W_C^{-1} = 1$, collisions are effectively “elastic,”
224 with no net enhancement or inhibition. In contrast, $W_C^{-1} > 1$ indicates that coagulation is
225 enhanced (e.g. due to attractive forces, favorable particle morphology, or turbulence), whereas
226 $W_C^{-1} < 1$ implies reduced coagulation (e.g. electrostatic repulsion or other inhibiting effects).

227 In figure 5a., the soot experiments show an initial period where $W_C^{-1} > 1$, which may be
228 explained by the fractal nature of soot aggregates that can promote sticking or chain formation
229 upon collision. By the third hour in all experiments, accounting for the variation the average
230 coagulation corrections extend above and below 1. During this later phase, particle
231 concentrations ($< 3 \times 10^3 \text{ cm}^{-3}$) no longer sustain significant coagulation losses, consistent with
232 prior studies showing that coagulation becomes negligible under low concentration
233 conditions^{34,42,43}.

234 Figure 5b shows the distribution of coagulation corrections for these time periods. All
235 three aerosols show a mean W_C^{-1} value around 1 (0.969 ± 0.524 for NaCl, 1.16 ± 1.38 for
236 sucrose, and 1.23 ± 0.312 for soot), suggesting a slight repulsion or negligible net sticking
237 among particles. However, the standard deviations do encompass $W_C^{-1} = 1$. Soot exhibits a
238 slightly higher coagulation corrections initially followed by reduced values (0.941 ± 0.307) in
239 later periods. These observations align with the notion that both particle morphology (e.g.,



240 fractal soot structures) and injection-induced turbulence can transiently enhance coagulation,
241 but the effect diminishes as particles coagulate.



242
243 Figure 5. a) Average time series of the calculated coagulation correction for NaCl (blue), sucrose
244 (green), and soot (red). Only fits with valid optimizations and r-squared greater than 0.85 are
245 included. b) Violin plots showing the mean (..), median (white bar) and range of the calculated
246 coagulation correction for each aerosol type averaged across replicate experiments for the
247 indicated time bins (similar to Figure 4b).

248

249 5 Discussion

250 Our initial experiments in this new cloud chamber focused on dry conditions and a set of
251 aerosols to quantify how particles evolve in the absence of humidity (<10% relative humidity).
252 Despite the relatively simple setup—no temperature or humidity control—two key insights will



253 be used in future humidified experiments. First, the wall-loss rates converged to similar values
254 across all aerosol types after the first hour, indicating that early differences largely arose from
255 injection flow conditions and subsequent turbulence. Over time, these chamber conditions
256 stabilized, reinforcing the well-documented notion that particle wall losses approach a quasi-
257 steady state as mixing subsides.

258 A second important finding is that coagulation within the chamber is most pronounced
259 during the initial phase of each experiment. Though this is more uncertain due to larger relative
260 errors. Soot showed signs of coagulation enhancement, potentially attributable to its fractal
261 structure. Once total number concentrations fell below a few thousand particles per cubic
262 centimeter, coagulation slowed considerably, consistent with the literature. Collectively, these
263 observations highlight the dynamic interplay between wall loss, particle morphology, and
264 injection protocols in shaping the early stages of aerosol evolution in chamber studies.

265 Our results also shed light on the influence of particle composition and shape. While
266 aerosols like NaCl and sucrose exhibited expected behavior—initial collision enhancements near
267 unity—soot displayed additional complexity. Early-time coagulation factors for soot were
268 moderately elevated, suggesting that fractal aggregates can promote sticking or increased
269 collisional radius. Over longer times, the coagulation rates for all three aerosols converged to
270 near unity or below, indicating negligible net enhancement under steady-state conditions. These
271 observations set the stage for more detailed investigations of fractal-like particles under high
272 humidity environments (>90% relative humidity).



273 Although these initial experiments focused on low humidities, the chamber design
274 allows for temperature and humidity control to be integrated in future work. Extending to more
275 complex atmospherically relevant aerosol mixtures—such as soot mixed with organic vapors or
276 inorganic salts—will further elucidate aerosol aging pathways and cloud interactions.
277 Additionally, the use of more advanced aerosol instrumentation will improve the
278 characterization of particle morphologies and mixing states that evolve during cloud processing.

279

280 **6 Conclusion**

281 The custom-built 906 L stainless-steel chamber provided reproducible measurements of particle
282 size distributions under dry conditions, confirming its suitability for controlled aerosol research.
283 Although initial turbulence drove high wall-loss rates, these converged to stable values across
284 NaCl, sucrose, and soot—underscoring that injection protocols and mixing strongly influence
285 early aerosol behavior. The chamber’s intermediate size and flexible design for future
286 temperature and humidity controls make it a useful platform to investigate aerosol-cloud
287 interactions more comprehensively. Integrating additional measurements of particle shape,
288 chemical composition, and mixing state will further clarify the complexities of aerosol aging and
289 cloud formation. Building on these dry experiments, upcoming work at higher humidity will
290 reveal how aerosol coagulation and phase changes affect cloud processes such as droplet
291 activation and scavenging. By disentangling coagulation, dilution, and wall-loss mechanisms,
292 this chamber ultimately enables rigorous study of aerosol transformations—particularly for



293 fractal soot—in cloud-relevant environments, helping advance both scientific understanding
294 and climate prediction.

295

296 *Author contributions.* NAF performed the experiments, contributed to the analysis of the
297 results, and wrote the manuscript. KJG was responsible for the design of the chamber and the
298 study, development of the data analysis methods, and contributed to the writing of the
299 manuscript. KBB contributed to the design of the study, analysis of the results, and contributed
300 to the writing of the manuscript.

301

302 *Competing interests.* The authors declare no competing interests.

303

304 *Code/Data availability.* The python data analysis code is available from corresponding authors.

305

306 *Acknowledgments.* The authors would like to thank Hannah Brink, Ryan Farley, James E. Lee,
307 and Matt Nelson for their feedback on the manuscript. Release Number LA-UR-25-22994.

308

309 *Financial support.* Research presented in this article was supported by the Laboratory Directed
310 Research and Development program of Los Alamos National Laboratory under project number



20230248ER. This work was supported by the U.S. Department of Energy through the Los Alamos National Laboratory. Los Alamos National Laboratory is operated by Triad National Security, LLC, for the National Nuclear Security Administration of U.S. Department of Energy (Contract No. 89233218CNA000001).

References

1. Intergovernmental Panel On Climate Change (Ippc). *Climate Change 2021 – The Physical Science Basis: Working Group I Contribution to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change*. (Cambridge University Press, 2023).
doi:10.1017/9781009157896.
2. Yu, P. *et al.* Black carbon lofts wildfire smoke high into the stratosphere to form a persistent plume. *Science* **365**, 587–590 (2019).
3. D’Angelo, G., Guimond, S., Reisner, J., Peterson, D. A. & Dubey, M. Contrasting Stratospheric Smoke Mass and Lifetime From 2017 Canadian and 2019/2020 Australian Megafires: Global Simulations and Satellite Observations. *JGR Atmospheres* **127**, e2021JD036249 (2022).
4. Guimond, S. R., Reisner, J. & Dubey, M. The Dynamics of Megafire Smoke Plumes in Climate Models: Why a Converged Solution Matters for Physical Interpretations. *J Adv Model Earth Syst* **15**, e2022MS003432 (2023).



- 329 5. Rodriguez, B., Lareau, N. P., Kingsmill, D. E. & Clements, C. B. Extreme Pyroconvective
330 Updrafts During a Megafire. *Geophysical Research Letters* **47**, e2020GL089001 (2020).
- 331 6. Leach, R. N. & Gibson, C. V. Assessing the Potential for Pyroconvection and Wildfire Blow Ups.
332 *J. Operational Meteor.* 47–61 (2021) doi:10.15191/nwajom.2021.0904.
- 333 7. Fromm, M., Servranckx, R., Stocks, B. J. & Peterson, D. A. Understanding the critical elements
334 of the pyrocumulonimbus storm sparked by high-intensity wildland fire. *Commun Earth*
335 *Environ* **3**, 243 (2022).
- 336 8. Reisner, J. M. *et al.* Informed Multi-Scale Approach Applied to the British Columbia Fires of
337 Late Summer 2017. *JGR Atmospheres* **128**, e2022JD037238 (2023).
- 338 9. Gorkowski, K. *et al.* Insights into Pyrocumulus aerosol composition: black carbon content and
339 organic vapor condensation. *Environ. Sci.: Atmos.* **4**, 80–87 (2024).
- 340 10. Cotton, W. R. & Anthes, R. A. *Storm and Cloud Dynamics*. (Academic Press, San Diego,
341 2010).
- 342 11. Das, S., Colarco, P. R., Oman, L. D., Taha, G. & Torres, O. The long-term transport and
343 radiative impacts of the 2017 British Columbia pyrocumulonimbus smoke aerosols in the
344 stratosphere. *Atmos. Chem. Phys.* **21**, 12069–12090 (2021).
- 345 12. June, N. A. *et al.* Aerosol size distribution changes in FIREX-AQ biomass burning plumes:
346 the impact of plume concentration on coagulation and OA condensation/evaporation. *Atmos.*
347 *Chem. Phys.* **22**, 12803–12825 (2022).



- 348 13. Koren, I., Kaufman, Y. J., Remer, L. A. & Martins, J. V. Measurement of the Effect of
349 Amazon Smoke on Inhibition of Cloud Formation. *Science* **303**, 1342–1345 (2004).
- 350 14. Kaufman, Y. J. & Koren, I. Smoke and Pollution Aerosol Effect on Cloud Cover. *Science*
351 **313**, 655–658 (2006).
- 352 15. Feingold, G., Remer, L. A., Ramaprasad, J. & Kaufman, Y. J. Analysis of smoke impact on
353 clouds in Brazilian biomass burning regions: An extension of Twomey's approach. *J. Geophys.*
354 *Res.* **106**, 22907–22922 (2001).
- 355 16. Cunningham, C. X., Williamson, G. J. & Bowman, D. M. J. S. Increasing frequency and
356 intensity of the most extreme wildfires on Earth. *Nat Ecol Evol* **8**, 1420–1425 (2024).
- 357 17. Ching, J., West, M. & Riemer, N. Quantifying Impacts of Aerosol Mixing State on
358 Nucleation-Scavenging of Black Carbon Aerosol Particles. *Atmosphere* **9**, 17 (2018).
- 359 18. Taylor, J. W. *et al.* Size-dependent wet removal of black carbon in Canadian biomass
360 burning plumes. *Atmos. Chem. Phys.* **14**, 13755–13771 (2014).
- 361 19. Zanatta, M. *et al.* Airborne investigation of black carbon interaction with low-level,
362 persistent, mixed-phase clouds in the Arctic summer. *Atmos. Chem. Phys.* **23**, 7955–7973
363 (2023).
- 364 20. Riemer, N., West, M., Zaveri, R. A. & Easter, R. C. Simulating the evolution of soot mixing
365 state with a particle-resolved aerosol model. *J. Geophys. Res.* **114**, 2008JD011073 (2009).



- 366 21. Zaveri, R. A., Barnard, J. C., Easter, R. C., Riemer, N. & West, M. Particle-resolved
367 simulation of aerosol size, composition, mixing state, and the associated optical and cloud
368 condensation nuclei activation properties in an evolving urban plume. *J. Geophys. Res.* **115**,
369 2009JD013616 (2010).
- 370 22. Ching, J., Riemer, N. & West, M. Black carbon mixing state impacts on cloud
371 microphysical properties: Effects of aerosol plume and environmental conditions. *JGR*
372 *Atmospheres* **121**, 5990–6013 (2016).
- 373 23. Yao, Y., Dawson, M. L., Dabdub, D. & Riemer, N. Evaluating the Impacts of Cloud
374 Processing on Resuspended Aerosol Particles After Cloud Evaporation Using a Particle-
375 Resolved Model. *JGR Atmospheres* **126**, e2021JD034992 (2021).
- 376 24. Yang, F., Hoffmann, F., Shaw, R. A., Ovchinnikov, M. & Vogelmann, A. M. An
377 Intercomparison of Large-Eddy Simulations of a Convection Cloud Chamber Using Haze-
378 Capable Bin and Lagrangian Cloud Microphysics Schemes. *J Adv Model Earth Syst* **15**,
379 e2022MS003270 (2023).
- 380 25. Khlystou, A., Kos, G. P. A. & Ten Brink, H. M. A High-Flow Turbulent Cloud Chamber.
381 *Aerosol Science and Technology* **24**, 59–68 (1996).
- 382 26. Chang, K. *et al.* A Laboratory Facility to Study Gas–Aerosol–Cloud Interactions in a
383 Turbulent Environment: The Π Chamber. *Bulletin of the American Meteorological Society* **97**,
384 2343–2358 (2016).



- 385 27. Niedermeier, D. *et al.* Characterization and first results from LACIS-T: a moist-air wind
386 tunnel to study aerosol–cloud–turbulence interactions. *Atmos. Meas. Tech.* **13**, 2015–2033
387 (2020).
- 388 28. Shao, Y. *et al.* Characterisation of the Manchester Aerosol Chamber facility. *Atmos.*
389 *Meas. Tech.* **15**, 539–559 (2022).
- 390 29. The Cloud Collaboration. A study of the link between cosmic rays and clouds with a
391 cloud chamber at the CERN PS. Preprint at <https://doi.org/10.48550/ARXIV.PHYSICS/0104048>
392 (2001).
- 393 30. Wagner, R. *et al.* Chamber Simulations of Cloud Chemistry: The AIDA Chamber. in
394 *Environmental Simulation Chambers: Application to Atmospheric Chemical Processes* (eds.
395 Barnes, I. & Rudzinski, K. J.) vol. 62 67–82 (Kluwer Academic Publishers, Dordrecht, 2006).
- 396 31. Corner, J. & Pendlebury, E. D. The Coagulation and Deposition of a Stirred Aerosol. *Proc.*
397 *Phys. Soc. B* **64**, 645–654 (1951).
- 398 32. Fotou, G. P. & Pratsinis, S. E. A Correlation for Particle Wall Losses by Diffusion in Dilution
399 Chambers. *Aerosol Science and Technology* **18**, 213–218 (1993).
- 400 33. Wang, N., Jorga, S. D., Pierce, J. R., Donahue, N. M. & Pandis, S. N. Particle wall-loss
401 correction methods in smog chamber experiments. *Atmos. Meas. Tech.* **11**, 6577–6588
402 (2018).



- 403 34. Mahfouz, N. G. A. & Donahue, N. M. Primary ion diffusion charging and particle wall loss
404 in smog chamber experiments. *Aerosol Science and Technology* **54**, 1058–1069 (2020).
- 405 35. Crump, J. G. & Seinfeld, J. H. Turbulent deposition and gravitational sedimentation of an
406 aerosol in a vessel of arbitrary shape. *Journal of Aerosol Science* **12**, 405–415 (1981).
- 407 36. Crump, J. G., Flagan, R. C. & Seinfeld, J. H. Particle Wall Loss Rates in Vessels. *Aerosol*
408 *Science and Technology* **2**, 303–309 (1982).
- 409 37. Particula [computer software]. uncscode/particula: v0.1.3. Zenodo
410 <https://doi.org/10.5281/ZENODO.6634653> (2025).
- 411 38. Seinfeld, J. H. & Pandis, S. N. *Atmospheric Chemistry and Physics: From Air Pollution to*
412 *Climate Change*. (John Wiley & Sons, Incorporated, Newark, 2016).
- 413 39. Zong, T. *et al.* A new smog chamber system for atmospheric multiphase chemistry study:
414 design and characterization. *Atmos. Meas. Tech.* **16**, 3679–3692 (2023).
- 415 40. Wang, J. *et al.* Design of a new multi-phase experimental simulation chamber for
416 atmospheric photosmog, aerosol and cloud chemistry research. *Atmos. Meas. Tech.* **4**, 2465–
417 2494 (2011).
- 418 41. Li, K. *et al.* Smog chamber study on aging of combustion soot in isoprene/SO₂/NO_x
419 system: Changes of mass, size, effective density, morphology and mixing state. *Atmospheric*
420 *Research* **184**, 139–148 (2017).



- 421 42. Hussein, T. *et al.* Deposition rates on smooth surfaces and coagulation of aerosol
422 particles inside a test chamber. *Atmospheric Environment* **43**, 905–914 (2009).
- 423 43. Yu, H. *et al.* Experimental and Numerical Study on the Gravitational Deposition and
424 Coagulation of Aerosols. *Front. Energy Res.* **10**, 840503 (2022).
- 425