



Technical note: Apportionment of Southeast Asian Biomass Burning and Urban Influence via In Situ Trace Gas Enhancement Ratios

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15 **Abstract.** Biomass burning and urban emissions are major contributors to poor air quality throughout Southeast Asia. Understanding these emissions is critical for predicting and mitigating their health impacts. The Cloud, Aerosol and Monsoon Processes Philippines Experiment (CAMP²Ex) field project in 2019 focused on regional sources in Southeast Asia and their effects on aerosol/cloud interactions using a combination of airborne, shipboard, and ground-based platforms. These measurements sampled a variety of pollution sources over the Sulu, Philippine, and South China seas during both the southwest monsoon and monsoon transition periods. CAMP²Ex provided a unique opportunity to examine how local and transported emissions affected air mass chemical composition and air quality. Correlations in airborne in situ gas enhancement ratios of CH₄ to CO were used to separate air masses with predominantly biomass burning, urban, or mixed influence, and isolating contributions from differing urban sources. HYSPLIT backtrajectory analysis was used to identify air mass sources, and resulting source regimes were examined for differences in ozone, reactive nitrogen, and aerosol chemical composition. 25 $\Delta\text{O}_3/\Delta\text{CO}$ enhancement ratios were observed to be constant between urban source regimes, yet $\Delta\text{NO}_y/\Delta\text{CO}$ enhancement ratios differed across these regimes. For biomass burning sources, enhancement ratios in $\Delta\text{O}_3/\Delta\text{CO}$ were lower than those reported by previous studies in this region. Organic aerosol mass fractions were 2-3 times higher in biomass burning influenced regimes compared to urban regimes. This technique represents a simple yet novel approach for separating emission influences in a complex environment that enable isolation of emission sources requiring relatively simple measurements.



1 Introduction

Biomass burning and urban fossil fuel emissions are both strong contributors to poor air quality. In particular, Southeast Asia is home to strong emissions of both types, and combined with many of the world's largest megacities, this leads to widespread increased mortality due to respiratory diseases (Lee et al., 2018). Uncertainties in emissions and emission sources can exacerbate the difficulty in predicting these mortality effects (Marvin et al., 2024), thus improving the ability to accurately identify air mass emission sources is a crucial need that will assist policymakers crafting policies to reduce these premature deaths.

Backtrajectory models have been utilized to great effect in identifying physical air mass histories (Fleming et al., 2012; Hilario et al., 2021, 2024; Robinson et al., 2011), but are limited by the resolution and accuracy of the assimilated meteorological information driving the model which can poorly represent sub-grid emissions and overestimate long-range transport (Harris et al., 2005). Gas tracer enhancement ratios, which exploit differing emission ratios, can provide a chemical ground truth at the measurement site (Halliday et al., 2019; Smith et al., 2015). While not providing an air mass physical history, the enhancement ratio can provide insight into the chemical history of the air mass, making it an excellent complement to backtrajectory models. The choice of gas tracers depends on the application and time scale of the emissions. Gases must either have similar chemical lifetimes to be viable or have a stable, known emission profile so that their differing lifetimes can be used as a chemical clock (Barnett et al., 2012; Gelencsér et al., 1997). For evaluating larger-scale regional or mesoscale emissions, relative enhancements in more stable tracers have been used to great effect (Plant et al., 2022). Commonly, enhancement ratios of carbon dioxide (CO₂) and carbon monoxide (CO) have been used to this end due to the direct relationship between their emission ratio and combustion efficiency (DiGangi et al., 2021; Halliday et al., 2019). However, this can be complicated in regions or seasons where biogenic or oceanic uptake of CO₂ can complicate the necessary stable tracer background (Silva et al., 2013).

In contrast, CO and methane (CH₄) are inefficiently scavenged by precipitation and cloud processes due to their low solubility, leaving dominant photochemical loss pathways typically on the order of many days-weeks for CO (Holloway et al., 2000; Khalil and Rasmussen, 1984) and years for CH₄ (Cicerone and Oremland, 1988; Naik et al., 2013; Voulgarakis et al., 2013), slow on the scale of regional or mesoscale transport. Thus, correlated enhancements between CO and CH₄ ($\Delta\text{CH}_4/\Delta\text{CO}$) can be primarily related to the air mass source. Biomass burning has been widely reported in the literature to have low $\Delta\text{CH}_4/\Delta\text{CO}$ enhancement ratios or emission ratios compared to the higher values reported from fossil fuel combustion. For example, based on shipborne measurements in the South China Sea while heavily influenced by biomass burning, Nara et al. (2017) reported a typical $\Delta\text{CH}_4/\Delta\text{CO}$ enhancement ratio of $3.71 \pm 0.88\%$. Worden et al. (2017) report biomass burning $\Delta\text{CH}_4/\Delta\text{CO}$ emission ratios between 0-20% over a variety of global domains based on Aura TES retrievals and GFED-driven GEOS-Chem modeling. In contrast, Helfter et al. (2016) observed $\Delta\text{CH}_4/\Delta\text{CO}$ emission ratios between 20-160% during urban eddy covariance flux studies over central London, UK, and Plant et al. (2019) observed mean $\Delta\text{CH}_4/\Delta\text{CO}$ enhancement ratios between 48-130% from airborne studies over six Eastern US cities.



We present a new approach that combines multiple methods for calculating enhancement ratios of $\Delta\text{CH}_4/\Delta\text{CO}$ to isolate contributions from differing chemical regimes observed during the 2019 Clouds, Aerosol and Monsoon Processes-Philippines Experiment (CAMP²Ex) field campaign (<https://espo.nasa.gov/camp2ex>). These chemical regimes were then compared with those determined through backtrajectory analysis to corroborate the expected source regimes and predict their emission origin. Finally, the regimes were used to examine contributors to poor air quality (ozone, reactive nitrogen, aerosol) by attributed source.

2 Methods

2.1 CAMP²Ex

CAMP²Ex was a joint field campaign sponsored by NASA, Naval Research Lab, and the Manila Observatory in 2019 based in the Philippines (Reid et al., 2023), and predominantly sampled airmasses over the Sulu, Philippine, and South China seas. The overarching science questions were focused on improving understanding of the coupling of aerosol, cloud, and radiation processes in a complex chemical and meteorological environment. The measurement platforms included two aircraft, the NASA P-3 and the SPEC Learjet 35 based at Clark International Airport in Luzon, the Scripps R/V Sally Ride stationed in the Philippine Sea, and a ground station at Manila Observatory inside the Ateneo de Manila University in Quezon City, Metro Manila. Flight profiles ranged over the South China, Philippine, and Sulu seas, capturing a variety of compositional and cloud environments (Hilario et al., 2021). The P-3 aircraft fielded a comprehensive payload, with a mixture of remote sensing and in situ cloud, aerosol, and composition instrumentation, data from which are the focus of this study. All flight data presented in this work are from contemporarily-averaged P-3 in situ data on a 5 s time base (i.e., a 5 s data merge), and all data are publicly available (NASA/LARC/SD/ASDC, 2020).

2.2 Measurements

Gas phase CO_2 , CH_4 , and CO mole fractions were collected at 0.4 Hz resolution via a commercial near-infrared cavity ringdown spectrometer (G2401-m: Picarro, Inc.) fitted with a custom airborne sampling system to increase stability and reduce sampling artifacts (DiGangi et al., 2021). Two-point calibrations were performed hourly during flight using standards prepared by NOAA ESRL and traceable to the WMO X2007 (CO_2), X2004A (CH_4), and X2014A (CO) scales. Precisions for CO_2 , CH_4 , and CO during CAMP²Ex were respectively 0.1 ppm, 1 ppb, and 5 ppb, and accuracies were respectively 0.1 ppm, 1 ppb, and 2%. Ozone (O_3) was measured using a pair of commercial dual-beam UV absorption sensors (Model 205: 2B Technologies, Inc.) fitted with a custom sampling manifold to enhance high altitude performance (Wei et al., 2021). Calibrations were performed in the laboratory prior to the campaign using a commercial NIST-traceable ozone generator (Model 306: 2B Technologies, Inc.). Due to an altitude-dependent discrepancy between the two sensors, final ozone data were reported as the average of the two measurements, with a combined uncertainty of 6% from both the intercomparison and the native uncertainty of the instrument.



95 Water vapor mixing ratios were measured via open-path near infrared absorption spectroscopy by the NASA Diode Laser Hygrometer (Diskin et al., 2002). Nitrogen oxide (NO), summed total NO + nitrogen dioxide (NO₂) (NO + NO₂ = NO_x), and total reactive nitrogen (NO_y) were measured using a two-channel chemiluminescence nitrogen oxide analyzer (Air Quality Design Inc.) sampling through an unheated PFA-sheathed inlet. NO and NO_x were sampled serially on one channel using a 395 nm UV LED photolytic NO₂ converter (Air Quality Design Inc.). A molybdenum converter operating at 300°C mounted
100 near the sample entry point to the aircraft cabin was used to convert NO_y species to NO for detection. The NO, NO₂, and NO_y limit of detection was 0.30, 0.27, 0.22 ppb with an uncertainty of 12%, 14%, and 12%, respectively. Submicron non-refractory aerosol mass concentrations were measured using a High-Resolution Time-of-Flight Aerosol Mass Spectrometer (HR-ToF-AMS, Aerodyne Research, LLC) using standard data processing methods (Canagaratna et al., 2007; DeCarlo et al., 2006). HR-ToF-AMS measurements were obtained behind a pressure-controlled inlet for stability and were recorded in FastMS mode (25
105 s open, 5 s closed) and averaged to approximately 30 s for this analysis. Mass concentrations are presented at standard temperature and pressure (273 K and 1013.25 hPa), and sulfate, nitrate, ammonium, chlorine, and organic mass concentrations were determined using standard fragmentation tables (Allan et al., 2004). In this analysis, total AMS aerosol mass is defined as the sum of each of these five AMS species contributions.

3 Emission Flag Derivation

110 To separate out different chemical regimes observed in the CAMP²Ex dataset, specifically biomass burning versus urban emissions, this work leverages differences in the populations of enhancement ratios of CH₄ and CO. A combination of two different approaches to calculating $\Delta\text{CH}_4/\Delta\text{CO}$ enhancement ratios was used to isolate the biomass burning and urban regimes: one using a single campaign background to determine enhancement ratios (Sect. 3.1) and a second using dynamic linear fitting on a rolling time window throughout each flight to find correlated enhancement ratios (Sect. 3.2). The final flag merging these
115 approaches was calculated for each timestamp in the 5 s data merge (Sect. 3.4).

3.1 Global Background Approach

Single concentrations each for CO and CH₄, 65 ppb and 1.85 ppm respectively, were chosen as the global regional background for the full dataset by fitting various segments of the urban and biomass burning lobes of the CH₄ vs CO distribution. However, much of the data at small CH₄ and/or CO enhancements above these background levels were still not separable into urban and
120 biomass burning regimes due to the variability in the true background concentration of each species. Thus, it was necessary to find the minimum ΔCH_4 and ΔCO enhancement above background at which the two regimes begin to be separable but also as low as possible to maximize the data capable of being assigned to each regime. To accomplish this, the frequency distribution of $\Delta\text{CH}_4/\Delta\text{CO}$ enhancement ratios were examined at various ΔCH_4 and ΔCO enhancement levels above background concentrations (Fig. 1a-b). In these plots, the high data frequency peaks of both the biomass burning regime (at low $\Delta\text{CH}_4/\Delta\text{CO}$
125 enhancement ratios) and the urban regime (at high $\Delta\text{CH}_4/\Delta\text{CO}$ enhancement ratios) appear as high peaks off scale. As the



cutoff enhancement concentrations were increased, a minimum in the frequency distribution emerges between 35-70% $\Delta\text{CH}_4/\Delta\text{CO}$ (Fig. 1a-b, grey shaded area). The cutoff concentration enhancements at which this minimum emerges were determined to be the minimum enhancement at which the urban and BB data populations could be isolated: 40 ppb for ΔCH_4 and 55 ppb for ΔCO . Hysteresis of ± 5 ppb CO and ± 5 ppb CH_4 along the data timeseries was added to avoid rapid regime switching influenced by instrument precision limitations, an example of which is shown in Supplemental Figure S1. Flag categories were then assigned by enhancement ratio and intensity of enhancement with respect to these cutoff concentrations (Fig. 2a).

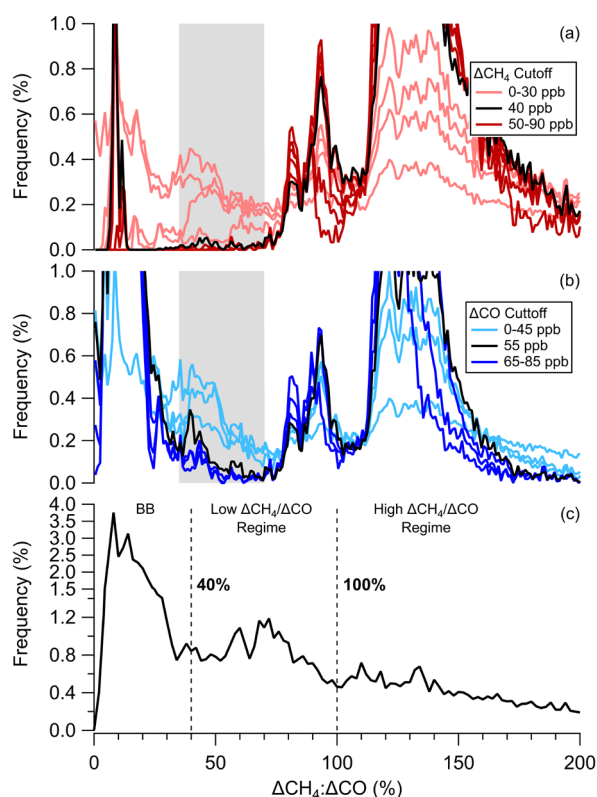


Figure 1: Sensitivity test of $\Delta\text{CH}_4/\Delta\text{CO}$ emission ratio distributions to concentration cutoffs in (a) CH_4 and (b) CO . Shaded area shows the minimum in $\Delta\text{CH}_4/\Delta\text{CO}$ chosen to separate the regimes above the selected cutoffs at $\Delta\text{CH}_4 = 40$ ppb and $\Delta\text{CO} = 55$ ppb. (c) Frequency distribution of $\Delta\text{CH}_4/\Delta\text{CO}$ enhancement ratios showing rolling slope regimes, with dashed lines representing the slope cutoffs between regimes at 40% and 100% $\Delta\text{CH}_4/\Delta\text{CO}$.

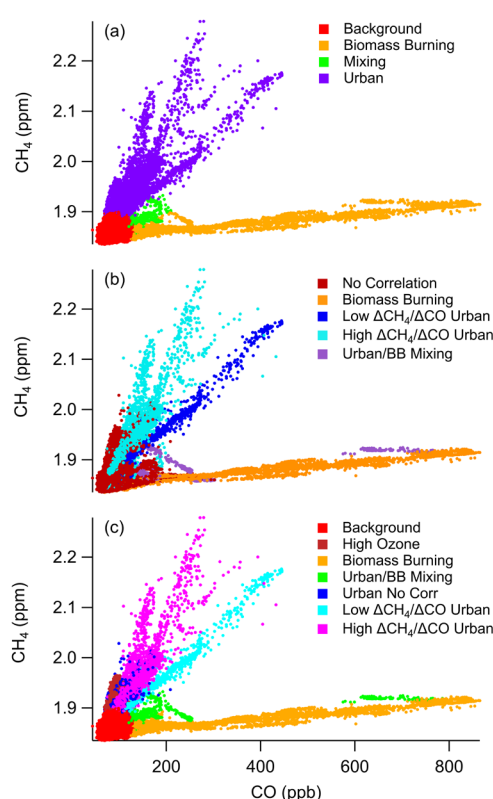


Figure 2: Tracer correlation plot of CH_4 vs CO mole fractions colored by regime excluding Clark-influenced data. (a) global background method, (b) rolling slope method, (c) final combined method.



3.2 Rolling Slope Approach

As the single background technique is limited due to the coarse assumption of a constant chemical background over the entirety of the campaign, a different approach was used to extract more fine detailed information from the dataset. With this approach, the linear fits of CH_4 vs CO were calculated using weighted orthogonal distance regression (ODRPACK95 - IGOR Pro v7; Wu and Yu, 2018) for 4 min rolling windows through each flight, similar to the technique used by Halliday et al. (2019) and DiGangi et al. (2021) for $\Delta\text{CO}/\Delta\text{CO}_2$ enhancement ratios (Fig. 2b). Rolling windows with lower goodness of fit ($r^2 < 0.5$) were filtered as uncorrelated. Slopes from the remaining fits represent a $\Delta\text{CH}_4/\Delta\text{CO}$ enhancement ratio product that is independent of changes in larger scale shifts in background concentration and more sensitive to mixing from recent emissions. In the frequency distribution of the $\Delta\text{CH}_4/\Delta\text{CO}$ slopes, three relative minima in the frequency distribution (Table 1) were chosen to separate the dataset into four regimes (Fig. 1c). One is a negative-sloped regime that corresponded to mixing between urban-influenced and biomass burning-influenced air. A second regime was between 0-40% $\Delta\text{CH}_4/\Delta\text{CO}$, peaking near $< 10\%$ $\Delta\text{CH}_4/\Delta\text{CO}$, with low slopes corresponding to biomass burning. The last two regimes had higher slopes consistent with urban influence: a lower slope regime between 40-100% $\Delta\text{CH}_4/\Delta\text{CO}$, peaking around 60-80% $\Delta\text{CH}_4/\Delta\text{CO}$, and a higher sloped regime containing slopes above 100% $\Delta\text{CH}_4/\Delta\text{CO}$ with a broad, multifeatured peak near 110-140% and a long tail toward higher $\Delta\text{CH}_4/\Delta\text{CO}$ slopes. Finally, the resulting flag was manually filtered to merge clearly temporally contiguous airmasses that may only have had correlation over a portion of the sampling region.

Regime	Global Background		Rolling Slope
	Enhancement Cutoff	Enhancement Ratio Cutoff	$\Delta\text{CH}_4/\Delta\text{CO}$ Limits
Urban/Biomass Burning Mixing	$\Delta\text{CO} < 55$ ppb $\Delta\text{CH}_4 < 40$ ppb	-	$\Delta\text{CH}_4/\Delta\text{CO} < 0\%$
Biomass Burning	$\Delta\text{CO} > 55$ ppb	$\Delta\text{CH}_4/\Delta\text{CO} < 35\%$	$\Delta\text{CH}_4/\Delta\text{CO} > 0\%$ $\Delta\text{CH}_4/\Delta\text{CO} < 40\%$
Urban Low $\Delta\text{CH}_4/\Delta\text{CO}$ Regime	$\Delta\text{CO} > 55$ ppb $\Delta\text{CH}_4 > 40$ ppb	$\Delta\text{CH}_4/\Delta\text{CO} > 35\%$ $\Delta\text{CH}_4/\Delta\text{CO} < 70\%$	$\Delta\text{CH}_4/\Delta\text{CO} > 40\%$ $\Delta\text{CH}_4/\Delta\text{CO} < 100\%$
Urban High $\Delta\text{CH}_4/\Delta\text{CO}$ Regime	$\Delta\text{CH}_4 > 40$ ppb	$\Delta\text{CH}_4/\Delta\text{CO} > 70\%$	$\Delta\text{CH}_4/\Delta\text{CO} > 100\%$

Table 1: Cutoff concentration enhancements for CO and CH_4 for both global and rolling slope background methods.

3.3 Special Cases

Data influenced by strong local sources during takeoff at Clark International Airport were flagged separately until the first sharp gradient in trace gas concentration/humidity or, if no gradient was discernible, until 1 km above ground level (AGL). Similarly, Clark influence prior to landing was flagged as either the last sharp trace gas concentration/humidity gradient or 1 km AGL through the end of the flight. A separate population of data was flagged as another special case shown in Supplemental Figure S2, a lobe in the $\Delta\text{CH}_4/\Delta\text{CO}$ scatterplot (Fig. S2a) identified due to a strong anticorrelation between ozone and water vapor data (Fig. S2b) at low water mixing ratios (< 8000 ppm_v). This may be indicative of stratospheric



mixing (Pan et al., 2014), which would normally also result in anti-correlations between ozone and both CH₄ and CO (Collins Jr. et al., 1993; Pan et al., 2004). In contrast, these ozone observations are uncorrelated with CO and appear positively correlated with CH₄ (Figs. S2c-d, respectively), analysis of which is deemed beyond the scope of this work. As these two cases are distinct from the rest of the dataset, both the Clark-influenced and potential stratospherically-influenced data were flagged as separate values in the data product and neglected in further analysis.

3.4 Final Combined Flag

The final flag background and biomass burning regimes were equivalent to those assigned by the single background method. The urban regime from the single background method was split into three separate regimes based on the correlated slope (or lack thereof) from the rolling slope method. The negative-sloped correlated regime was combined with the urban/BB mixing regime from the single background method to form the final urban/biomass burning mixing regime. Figure 2c shows the final regime apportionments with respect to CH₄ and CO, while Table 1 shows the final cutoffs in both tracer enhancement and enhancement ratios for each method.

4 Discussion

4.1 HYSPLIT Backtrajectory Comparison

To evaluate the utility of the emission flag and assign potential origins of the empirically-derived populations, hourly backtrajectories were calculated using the NOAA Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPLIT) model (Rolph et al., 2017; Stein et al., 2015), similar to those described in Hilario et al. (2021). Aircraft latitude, longitude, and altitude from the P-3 data merge were used to initialize the model for each 5 s point along the flight track for all flights. The Global Forecast System (GFS) 0.25°

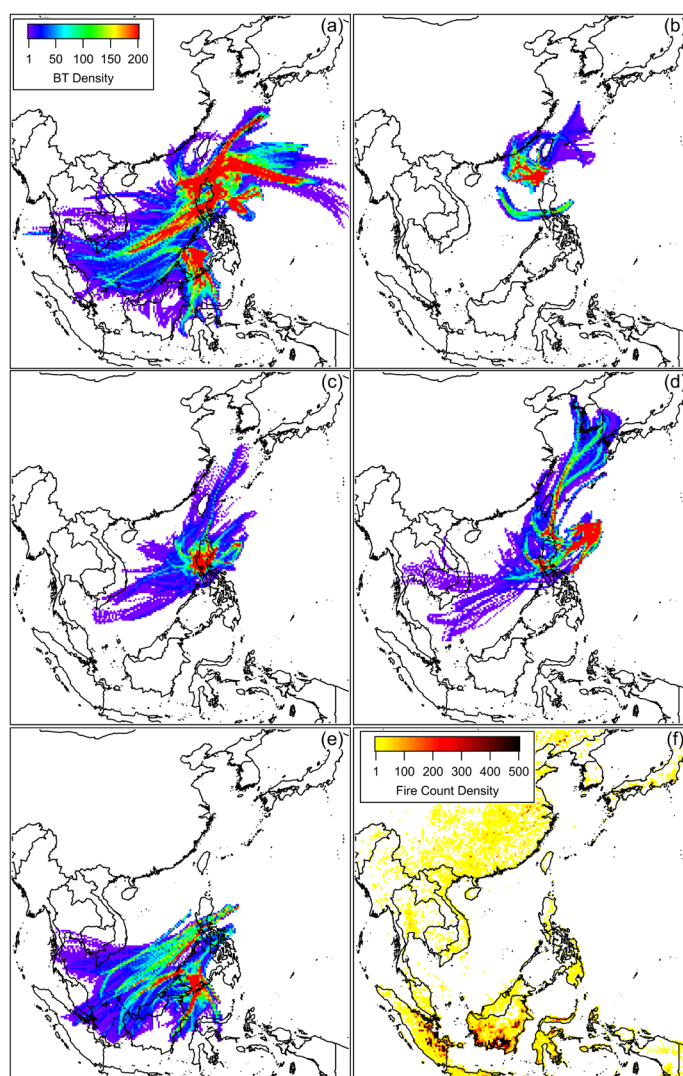


Figure 3: Heat map of HYSPLIT 48 h backtrajectories at 1 h resolution initialized every 5 s from flight track < 2 km, separated by regime: (a) background, (b) low $\Delta\text{CH}_4/\Delta\text{CO}$ urban, (c) high $\Delta\text{CH}_4/\Delta\text{CO}$ urban, (d) uncorrelated urban, and (e) biomass burning. (f) Heat map of VIIRS 375 m fire counts during September 2019.

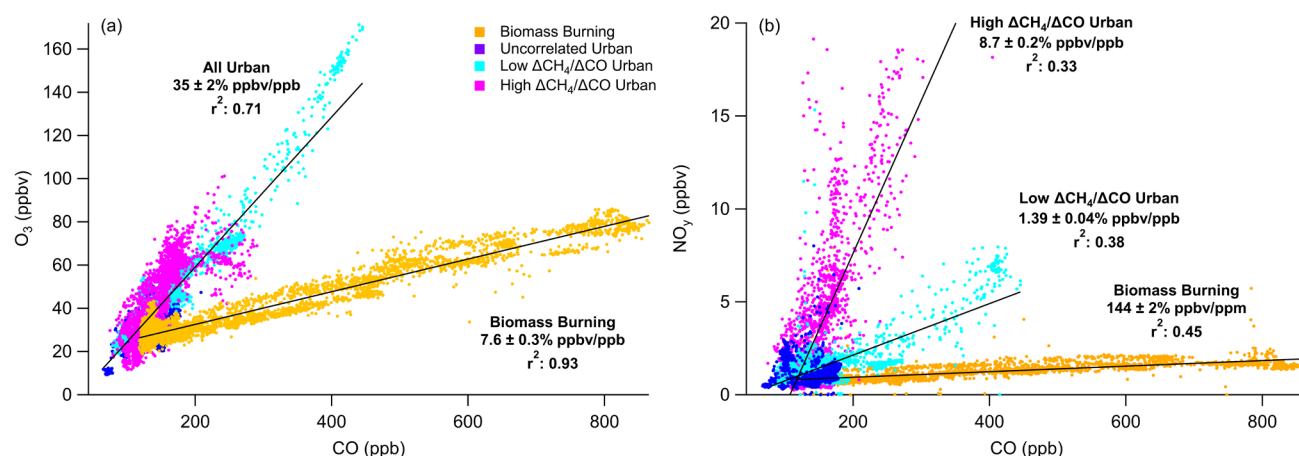


190 resolution reanalysis product from the National Centers for Environmental Protection (NCEP) was used to drive the modelled meteorology. Fig. 3 shows a summary of heat maps of the resulting hourly backtrajectory points under 2 km above mean sea level, an altitude cutoff intended to isolate contributions to the backtrajectories originating from the planetary boundary layer. Data were binned to 0.25° resolution where the bin color denotes the frequency of a contribution. Each heat map represents backtrajectories initialized during a different observed chemical regime.

195 Trajectories observed from the background chemical regime (Fig. 3a) were well dispersed around the region, though few backtrajectories appear over land in the boundary layer within the 48 h time window. In contrast, trajectories observed during the biomass burning chemical regime (Fig. 3b) were focused toward the southwest of the flight domain. The bulk of these were focused from the direction of Borneo and Sulawesi, though contributions appear from as far as Sumatra, the Malay peninsula, and the rest of mainland Southeast Asia. These results were consistent with the larger clusters in September 2019 observed

200 fire counts (Fig. 3f) from the VIIRS 375m product (NASA FIRMS, 2020). Of the correlated urban regimes, the low $\Delta\text{CH}_4/\Delta\text{CO}$ urban regime backtrajectories (Fig. 3c) were predominantly either from over continental China or from that direction. The high $\Delta\text{CH}_4/\Delta\text{CO}$ urban regime was observed more frequently and appears from the backtrajectories (Fig. 3d) to be focused locally in and around Luzon. The most frequent contributions in the uncorrelated urban regime (Fig. 3e) seemed to originate from Korea and Kyushu (Japan), the distance from which could be an explanation for existence of consistent CO and CH₄

205 enhancements with no short-term correlations between the species. The partitioning of the backtrajectories was based only on the $\Delta\text{CH}_4/\Delta\text{CO}$ -derived chemical regimes, yet results in distinct separation in the backtrajectory regimes, providing evidence toward the validity of this technique.



210 **Figure 4: Tracer-tracer plots of (a) O₃ and (b) NO_y versus CO separated by chemical regime. Trend lines denote linear fits for their respective regimes.**

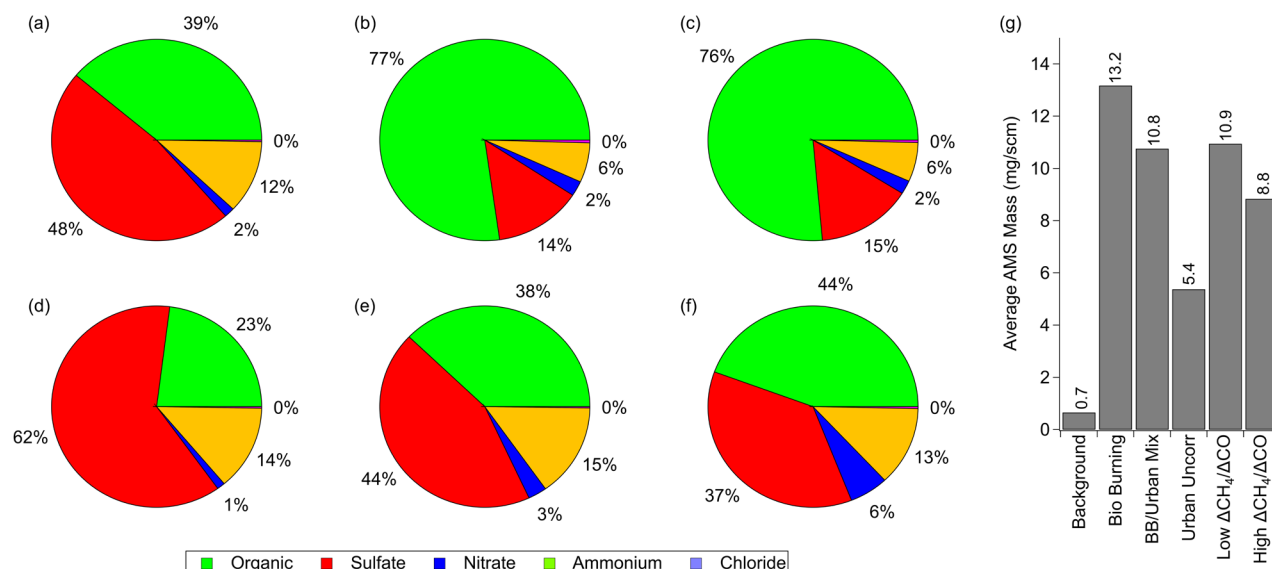


Figure 5: Relative aerosol chemical composition with respect to chemical regime: (a) background, (b) biomass burning, (c) biomass burning/urban mixing, (d) uncorrelated urban, (e) low $\Delta\text{CH}_4/\Delta\text{CO}$ urban, (f) high $\Delta\text{CH}_4/\Delta\text{CO}$ urban. (g) Average total AMS mass concentration by regime.

4.2 Ozone and Reactive Nitrogen Regime Dependencies

The ozone and reactive nitrogen measurements can give insight into the relative contributions within the different chemical regimes observed during the campaign. Fig. 4a shows enhancements in ozone with respect to CO for each regime. Biomass burning $\Delta\text{O}_3/\Delta\text{CO}$ is very strongly correlated (r^2 0.93) with a slope of $7.6 \pm 0.3\%$ ppbv/ppb, which is lower than both the $18 \pm 8\%$ ppbv/ppb reported by Lin et al. (2013) and the $20 \pm 1\%$ ppbv/ppbv reported by Kondo et al. (2004) in roughly the same region. Seasonality may affect these results somewhat, as the Lin et al. (2013) measurement values represent an annual range and the Kondo et al. (2004) values were only during the February-April, both contrasting with the measurements exclusively during August-October in this work. Otherwise, this may be indicative of a change in emission patterns over the previous decade, with lower O_3 production implying different NO_x or VOC loading. Urban $\Delta\text{O}_3/\Delta\text{CO}$ enhancements from all categories were not readily separable, but taken as one, they were still well correlated (r^2 0.71) with a slope of $35 \pm 2\%$ ppbv/ppb.

This consistency in $\Delta\text{O}_3/\Delta\text{CO}$ correlations among the urban regimes did not persist for the reactive nitrogen/CO ratios. $\Delta\text{NO}_y/\Delta\text{CO}$ enhancements were used to represent reactive nitrogen enhancements in the different regimes, as shown in Fig 4b. Biomass burning $\Delta\text{NO}_y/\Delta\text{CO}$ was not as correlated as $\Delta\text{O}_3/\Delta\text{CO}$ with an r^2 of 0.45 and a slope of 0.144 ± 0.002 pptv/ppb. $\Delta\text{NO}_y/\Delta\text{CO}$ enhancements were similarly correlated in the high and low $\Delta\text{CH}_4/\Delta\text{CO}$ urban regimes with r^2 values of 0.33 and 0.38, respectively. However, in contrast with the $\Delta\text{O}_3/\Delta\text{CO}$ urban correlations, the high and low $\Delta\text{CH}_4/\Delta\text{CO}$ urban regimes exhibited distinctive $\Delta\text{NO}_y/\Delta\text{CO}$ slopes, with the high $\Delta\text{CH}_4/\Delta\text{CO}$ regime slope (8.7 ± 0.2 pptv/ppb) a factor of six greater than the low $\Delta\text{CH}_4/\Delta\text{CO}$ urban regime slope (1.38 ± 0.02 pptv/ppb). This range of ratios is comparable to those reported by Kondo



et al. (2004), who observed $\Delta\text{NO}_y/\Delta\text{CO}$ slopes of 2.2 ± 0.6 pptv/ppbv in the planetary boundary layer and 7.9 ± 0.3 pptv/ppbv in the lower troposphere over Southeast Asia in primarily urban-influenced air during TRACE-P in 2001. The common ozone dependency under these drastically different reactive nitrogen environments suggests that ozone production in the low and high $\Delta\text{CH}_4/\Delta\text{CO}$ urban regimes may have been VOC-limited.

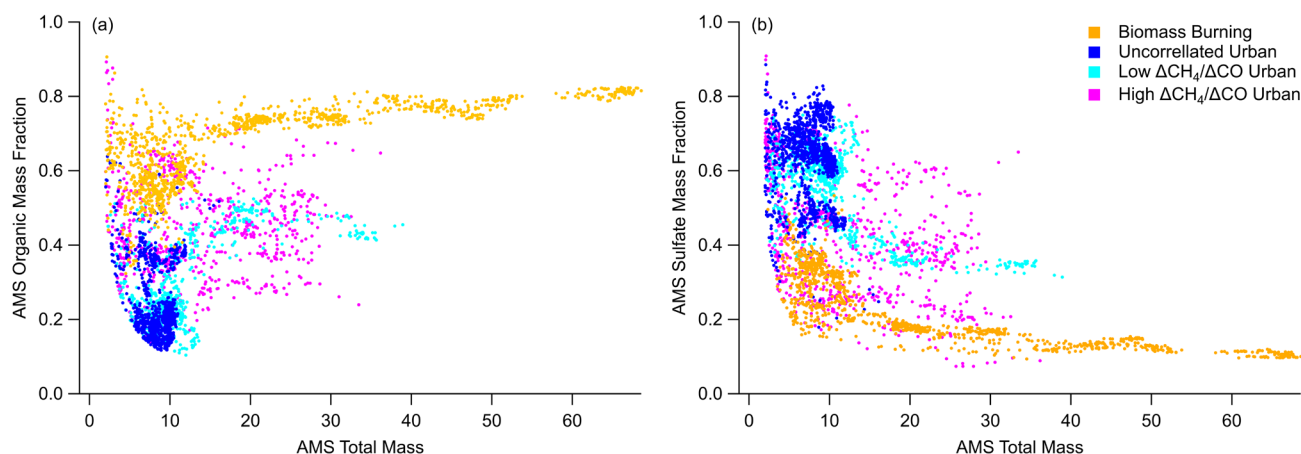


Figure 6: Relative AMS mass fractions of (a) organics and (b) sulfates versus total AMS mass concentration by chemical regime.

240 4.3 Aerosol Composition Regime Dependencies

To evaluate the chemical regime effects on aerosol composition, Fig. 5a-f shows the relative speciated contribution to aerosol mass observed by the AMS in each regime integrated over the entire campaign, while Fig. 5g shows the total aerosol mass observed in each regime, also over the entire campaign. The biomass burning regime (Fig. 5b) exhibited the highest average aerosol mass concentration ($13 \mu\text{g}/\text{m}^3$) and the largest relative mass contribution from organic aerosols (77%). The biomass/urban mixing regime (Fig. 5c) exhibited nearly identical relative aerosol contributions to the biomass burning regime, likely related to the observed larger mass loading from biomass burning compared to urban emissions. Among the urban regimes, the highest average mass concentration was observed in the low $\Delta\text{CH}_4/\Delta\text{CO}$ urban regime ($11 \mu\text{g}/\text{m}^3$), with relative aerosol contributions very similar to those observed in the background regime: $\sim 40\%$ each organic and sulfate aerosol mass. The high $\Delta\text{CH}_4/\Delta\text{CO}$ urban regime exhibited less aerosol on average ($8.8 \mu\text{g}/\text{m}^3$) than the low $\Delta\text{CH}_4/\Delta\text{CO}$ urban regime and had very similar composition with also $\sim 40\%$ each organic and sulfate aerosol mass. In the background regime (Fig 5a), average aerosol loading was much lower ($\sim 0.7 \mu\text{g}/\text{m}^3$) than for other regimes, and the largest relative contribution was from sulfate aerosols, with 49% of the loading, followed by organic aerosols for another 39%. In contrast, the biomass burning (Fig. 5b) and biomass/urban (Fig. 5c) mixing regimes had similar relative mass contributions with respectively 77% and 76% of the observed aerosol mass contribution from organic aerosol and only about 14% and 15% respectively from sulfate aerosol. Between the urban regimes, the uncorrelated regime had the lowest average aerosol mass but a larger relative sulfate aerosol contribution (62%) compared to the low and high $\Delta\text{CH}_4/\Delta\text{CO}$ urban regimes (44% and 37% respectively).



Fig. 6 shows the trend in relative mass fraction contributions for the two dominant aerosol mass contributors, organic and sulfate, as a function of total aerosol mass in the biomass burning and urban chemical regimes. Within the urban regimes, there was large variability in the mass fractions for both organic and sulfate. One exception seems to be that the mass fractions asymptote at nearly 50% (organic) and 35% (sulfate) of total aerosol mass within the low $\Delta\text{CH}_4/\Delta\text{CO}$ urban regime, though this could be an example of the more limited observations of this regime. In the biomass burning regime, the mass fractions have high total mass asymptotes near 80% (organic) and 10% (sulfate).

5 Summary

Enhancement ratios in airborne in situ CH_4 and CO were used to isolate predominant sources of air mass emission influence during the CAMP²Ex campaign over the western Pacific near the Philippines. Due to a more variable background in CO_2 during CAMP²Ex, $\Delta\text{CH}_4/\Delta\text{CO}$ was observed to exhibit greater utility in discerning combustion emission sources in comparison to the more commonly used $\Delta\text{CO}/\Delta\text{CO}_2$, a metric of combustion efficiency. A 5 s data product was derived for all flights assigning regimes of biomass burning and various urban sources based on natural groupings of ranges of $\Delta\text{CH}_4/\Delta\text{CO}$ enhancement ratios. HYSPLIT backtrajectory modeling was used to predict the likely emission origins of the differing regimes and to corroborate the method, with a higher sloped $\Delta\text{CH}_4/\Delta\text{CO}$ corresponding to local emissions and a lower sloped $\Delta\text{CH}_4/\Delta\text{CO}$ corresponding to continental outflow. $\Delta\text{O}_3/\Delta\text{CO}$ were individually strongly correlated in both biomass burning and urban regimes, but with a much lower slope for biomass burning ($7.6 \pm 0.3\%$ ppbv/ppb) compared to the slope for combined urban regimes ($35 \pm 2\%$ ppbv/ppb). $\Delta\text{NO}_y/\Delta\text{CO}$ were also highly correlated but had a sixfold higher slope for the local high $\Delta\text{CH}_4/\Delta\text{CO}$ urban regime compared to the continental outflow low $\Delta\text{CH}_4/\Delta\text{CO}$ urban regime. Organic aerosols were observed to dominate aerosol mass fractions for biomass burning influenced regimes at $\sim 75\%$ of the total mass fraction, whereas background as well as low and high $\Delta\text{CH}_4/\Delta\text{CO}$ urban regimes were more evenly balanced between sulfate and organic aerosols.

The multi-faceted method developed in this analysis provides a novel approach to assessing regional source contributions. One caveat is that this method appears to work well at receptor sites away from direct sources as in CAMP²Ex but may not work as well very close to emission sources or in air masses simultaneously influenced by both biomass burning and fossil fuel combustion. The use of these simple tracer relationships perhaps achieved using common, commercially-available instrumentation, may provide a feasible approach for constraining certain emission sources without requiring the cost and complexity of a larger-scale experimental effort. Thus, this method can more broadly assist scientists and policymakers in understanding the impact of emission type on air quality metrics.



285 Data Availability

All CAMP²Ex airborne data is available from NASA on the project data archive at <https://doi.org/10.5067/SUBORBITAL/CAMP2EX2018/DATA001> (NASA/LARC/SD/ASDC, 2020). The HYSPLIT model is a product of NOAA and available at <https://www.ready.noaa.gov/HYSPLIT.php>. VIIRS fire count data is available from NASA at <https://firms.modaps.eosdis.nasa.gov/download/> (NASA FIRMS, 2020).

290 Author Contributions

JPD, GSD, SY, SLA, JHF, CER, MAS, KLT, ELW, and LDZ participated in the data collection. JPD, GSD, LDZ, MOLC, JBS, MRAH, and AS performed the data analysis and provided feedback. JPD prepared the manuscript with contributions from all coauthors.

Competing Interests

295 AS is an editor of Atmospheric Chemistry and Physics.

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