# Supplementary Material (S1)

# S1: Detailed model design, structure and equations in CoupModel

### 1. Soil surface energy balance

Energy balance is important to land surface exchanges in CoupModel. In computing the energy fluxes, total net radiation ( $R_{n,tot}$ ) was calculated from net longwave radiation and shortwave incoming radiation Ris. LWin was estimated from meteorological measurements using Konzelmann et al. (1994) function, while  $R_{is}$  is one of the driving variables of the model. The fluxes were partitioned in CoupModel according to the surface energy balance (eq.A1),

$$\underbrace{\frac{\sigma(\varepsilon_{s}(T_{s}+273.15)^{4}-\underbrace{\varepsilon_{a}(T_{a}+273.15)^{4})}_{LW_{out}}-\underbrace{\varepsilon_{a}(T_{a}+273.15)^{4})}_{R_{n,iot}}}_{R_{n,iot}}+(1-a_{r})R_{is}=H+LE+q_{h}+\frac{dF}{dt} \quad (eq. A1)$$

where  $LW_{out}$  is the longwave radiation emitted from the ground calculated by using simulated temperature of soil surface, Ts with consideration of snow surface temperatures in winter. LWin represents the incoming longwave radiation from the atmosphere which is based on measured air temperature, Ta (forcing variable).  $\sigma$  is the Stefan-Boltzmann constant and is the  $\varepsilon_s$  emissivity of the ground. Atmospheric emissivity  $\varepsilon_a$  is estimated from Konzelmann et al (1994) function. The surface albedo of ground  $a_r$  was simulated by considering the dynamic area cover of peat soil and snow. Peat soil albedo was set to 7% as measured by Munro et al. (2010). The snow approach used in the study follows Gustafsson et al. (2001) where snow albedo was assumed to be a function of snow age with 90% for newly formed snow which decreased to 40% c.a. 1 month time. CoupModel partitioned  $R_{n,tot}$  into laten (LE), sensible (H) and soil heat flux ( $q_h$ ) fluxes respectively. Change energy storage (dF/dt) within the measured reference height ( $Z_{ref}$ = 1.5m) and soil surface was assumed to be zero in the study.

The model calculated H and LE between the soil surface and reference level of meteorological inputs from gradients of temperature and vapour pressure respectively

$$H_{s} = \rho_{a}c_{p}\frac{T_{s} - T_{a}}{r_{as}}$$
 (eq. A2)
$$LE = \frac{\rho_{a}c_{p}\left(e_{surf} - e_{a}\right)}{\gamma}$$
 (eq. A3)
$$q_{h} = k_{h}\frac{T_{s} - T_{l}}{\frac{\Delta z_{l}}{2}}$$
 (eq. A4)

Where  $\rho_a$  is the soil density, cp is he air heat capacity,  $\gamma$  is the psychrometric constant,  $e_{surf}$  is the vapour pressure of the soil surface,  $e_a$  is the actual vapour pressure of the atmosphere (forcing variable), aerodynamic resistance above the soil surface  $r_{as}$  which was assumed to be the same for sensible and latent heat flux. Under neutral conditions,  $r_{as}$  (eqs. 2 &3) was calculated as a function of wind speed and temperature gradients,

$$r_{as} = \frac{1}{k^2 u} \ln \left( \frac{z_{ref} - d}{z_{OM}} \right) \ln \left( \frac{z_{ref} - d}{z_{OH}} \right)$$
 (eq.A5)

Where measured wind speed, u (forcing variable) was given at the reference height, k (=0.4) is von Karman's constant, d is the displacement height (d=0),  $z_{OM}$  and  $z_{OH}$  are the roughness length for momentum and heat respectively. Under no-neutral conditions, eg. (A5) was further corrected with the Monin-Obukhov stability function (Beljaars and Holtslag 1991). This involves a dimensionless factor  $kB^{-1}$ . In this study,  $kB^{-1} = 2.3$  was obtained from the synthesized value from available measured peatland data (Humphreys et al., 2006)

Alvenas and Jansson (1997) scheme was used to account for the non-equilibrium effects of rapid moisture fluctuations close to the soil surface. This was done by modifying  $e_{surf}$  by an available factor that is driven by soil moisture availability (eq. A7) calculated by water balance at the first soil layer (eq. A8)

$$e_{surf} = e(T_s) *\times exp^{\left(\frac{\psi Mwge_c}{R(T_s + 273.15)}\right)}$$
 (eq. A6)

$$e_c = 10^{(-\delta\psi_g)} \tag{eq. A7}$$

$$\delta_{\mathcal{S}}(t) = \max\left(-2, \min(1, \delta_{\mathcal{S}}(t-1) + (P-E)\Delta t\right) \tag{eq. A8}$$

Where  $e(T_s)^*$  is the saturated vapor pressure at the surface temperature  $T_s$ ,  $\psi$  is the mean soil water potential in the topsoil layer, g is the gravity constant, and R is the gas constant,  $e_c$  is an empirical correction factor that compensates for the differences between the mean soil moisture potential in the top-soil layer and the soil moisture potential at the surface, defined as eq. (A7), eq. (A8) calculates the water balance and was assumed to vary from a deficient of -2mm to a surplus of +1mm at one timestep t, where  $\delta_s$  and  $\psi_g$  are parameters.

#### 2. Soil temperature

Soil heat flux  $q_h$  in eq. (A1) was further used to calculate the soil temperature profile by considering heat conduction and convection flow (eqs. A9-A11),

$$q_h = -k_h \frac{\partial T_s}{\partial z} + C_w T_s q_w \tag{eq. A9}$$

$$k_h = h_1 + h_2 \Theta \tag{eq. A10}$$

$$\frac{\partial (CT)}{\partial t} - L_f \rho \frac{\partial \theta_i}{\theta_t} = \frac{\partial}{\partial z} (-q_h)$$
 (eq. A11)

Where dT/dz is the gradient of soil temperature with depth,  $k_h$  is the thermal conductivity of peat soil calculated by an empirical approach of de Vries (1975), assuming proportional to the soil water content,  $\theta$  and  $h_1$  and  $h_2$  are parameters (eq. 10). The subscript h in eq. (A9) is heat, w is liquid water, i and f in eq. (A11) refer to ice and freezing respectively, t is time, C is heat capacity, t is latent heat of melting. In this study model default values were used for the snow and soil frost modules.

### 3. Soil water flow

The soil water process (eqs. A12-A13) was described based on the matrix flow and bypass flow,

$$q_w = -k_w \left(\frac{\partial \psi}{\partial z} - 1\right) + q_{\text{bypass}}$$
 (eq. A12)

$$\frac{\partial \theta}{\partial t} = -\frac{\partial q_w}{\partial z} \tag{eq. A13}$$

Where qw is the water flow flux, kw is the unsaturated hydraulic conductivity, is the water tension/potential and qbypass is the bypass flow.

The soil water retention curve was based on the van Genuchten (1980) approach

$$S_e = \frac{1}{(1 + (\alpha \psi)^n)^m} = \frac{\theta - \theta_r}{\theta_s - \theta_r}$$
 (eq. A14)

Where Se is the effective saturated water content,  $\alpha$ , n, m are van Genuchten parameters (m=1-1/n),  $\theta_r$  is the residual water content,  $\theta_s$  is the saturated water content.

The (un)saturated hydraulic conductivity was calculated by Mualem (1976) and modified to consider macropores (pp 63-68 in Jansson and Karlberg, 2011)

$$k_{w}^{*} = k_{\text{mat}} S_{e}^{\left(n+2+\frac{2}{\lambda}\right)}$$
 (eq. A15)  

$$k_{\text{mat}} = 10^{\left((\log k_{\text{sat}} - \log h_{\text{com}})h_{\text{sens}} + \log k_{\text{sat}}\right)}$$
 (eq. A16)  

$$k_{w}^{**} = 10^{\left(\log(k_{w} * (\theta_{s} - \theta_{m})) + \frac{\theta - \theta_{s} + \theta_{m}}{\theta_{m}} \log\left(\frac{k_{\text{sat}}}{k_{w} * (\theta_{s} - \theta_{m})}\right)\right)}$$
 (eq. A17)  

$$k_{w} = (r_{AOT} + r_{AlT}T_{s}) max(k_{w} * k_{minus})$$
 (eq. A18)

Where (eqs. A15-A16) accounts for the matric pores,  $k_w^*$  is the unsaturated hydraulic conductivity,  $k_{\text{mat}}$  is the saturated matrix conductivity,  $h_{\text{com}}$  and  $h_{\text{sens}}$  are parameters. The (eq. A17) additionally accounts for the conductivity in macropores,  $\theta_m$  is the water content at which the macropore water flow occurs,  $k_{\text{sat}}$  is the saturated total hydraulic conductivity and  $k_w^*$  ( $\theta_s - \theta_m$ ) is the unsaturated hydraulic conductivity calculated by eq. (A15), the effects of temperature on hydraulic conductivity were calculated by eq. (A18),  $r_{AOT}$  and  $r_{AlT}$  are scaled parameters for temperature,  $k_{minus}$  is the minimum hydraulic conductivity parameter.

The bypass flow (eqs. A19-A21) was calculated by an empirical approach (Janssonet al, 2005)

$$q_{\text{bypass}} = \begin{cases} 0 & 0 < q_{\text{in}} < S_{\text{mat}} \\ q_{\text{in}} - q_{\text{mat}} & q_{\text{in}} \ge S_{\text{mat}} \end{cases}$$

$$q_{\text{mat}} = \begin{cases} max \left( k_w(\theta) \left( \frac{\partial \psi}{\partial z} + 1 \right), q_{\text{in}} \right) & 0 < q_{\text{in}} < S_{\text{mat}} \\ S_{\text{mat}} & q_{\text{in}} \ge S_{\text{mat}} \end{cases}$$

$$(eq. A19)$$

$$S_{\text{mat}} = a_{\text{scale}} a_r k_{\text{mat}} pF$$
 (eq. A21)

Where qin is the infiltration rate,  $S_{\text{mat}}$  is the sorption capacity rate,  $k_{\text{mat}}$  is the maximum conductivity of the matric pores,  $a_r$  is the compartment thickness (for each layer), pF is  $10^{\log}$  of the water tension  $\psi$  and  $a_{\text{scale}}$  is the sorption scaling coefficient.

The lateral flow across the swamp was calculated by Darcy-type function (He et al., 2021),

$$q_{wp} = \int_{z_p}^{z_{sat}} k_{sat} \frac{(z_{sat} - z_p)}{d_u d_p} d_z \qquad (eq. A22)$$

Where  $q_{wp}$  is the horizontal flow rate,  $d_u$  is the unit length of the horizontal element,  $z_{sat}$  is the simulated depth of the groundwater level,  $k_{sat}$  is the saturated hydraulic conductivity of each soil layer including macropores,  $d_p$  is the distance between drainage pipes,  $z_p$  is the level of drainage pipes.

During over-saturated periods (e.g. snowmelt and heavy rainfall), the flow of water in the upper soil layer can be directed upwards and added to the surface runoff to drainage pipes. The model formed a pool of water on the soil surface when precipitation exceeded the infiltration capacity of the first soil layer. Water in the surface pool could either infiltrate with a delay into the soil profile or be lost as surface runoff, as overland flow. The flow was regulated by a first-order runoff coefficient (eq. A21)

## 4. Plant biotic processes

Canopy photosynthetic rate and C assimilation was estimated by light use efficiency approach,

$$C_{atm \to a} = \varepsilon_l \eta f(T_l) f(CN_l) f\left(\frac{E_{ta}}{E_{tn}}\right) R_{s,pl} \quad (eq. A23)$$

Where rate of photosynthesis (g C m<sup>-2</sup> day<sup>-1</sup>),  $C_{atm\to a}$  is a function of global radiation absorbed by the swamp's canopy  $R_{s,pl}$  but ultimately this is constrained by unfavorable temperature  $f(T_l)$ , nitrogen  $f(CN_l)$ , and water  $f\left(\frac{E_{ta}}{E_{tp}}\right)$  conditions.  $\varepsilon_l$  is the radiation use efficiency and  $\eta$  is the conversion factor from biomass to carbon

where  $\varepsilon_L$  is the radiation use efficiency and  $\eta$  is the conversion factor from biomass to carbon.  $R_{s,pt}$  is the global radiation absorbed by canopy and  $f(T_l)$ ,  $f(CN_l)$ , and  $f(E_{ta}/E_{tp})$  limitations due to unfavourable temperature, nitrogen, and water conditions.

### Response function for leaf temperature $f(T_1)$

$$f(T_{l} - p_{\text{man}})/(p_{o1} - p_{\text{mox}}) \quad p_{\text{mox}} \leq T_{l} \leq p_{o1}$$

$$f(T_{l}) = 1 \quad p_{o1} < T_{l} < p_{o2}$$

$$1 - (T_{l} - p_{o2})/(p_{\text{max}} - p_{o2}) \quad p_{o2} \leq T_{l} \leq p_{\text{max}}$$

$$0 \quad T_{l} > p_{\text{max}}$$

$$(eq. A24)$$

where  $p_{mu}$ ,  $p_{ol}$ ,  $p_{o2}$  and  $p_{mux}$  are parameters and  $T_l$  the leaf temperature.

### Response function for fixed leaf C:N ratio f(CN<sub>1</sub>)

$$f(CN_l) = p_{\text{fixedN}}$$
 (eq. A25) Where  $p_{\text{fixedN}}$  is a parameter.

# Response function for transpiration $f(E_{ta}/E_{tp})$

$$f(E_{ta}/E_{pp}) = \frac{E_{ta}}{E_{tp}}$$
 (eq. A26) where  $E_{ta}$  and  $E_{tp}$  are actual and potential transpiration

# Allocation of new assimilates to the leaves $C_{a\rightarrow Leaf}$

 $C_{a \to L \text{eaf}} = l_{c1} \cdot C_a$  (eq. A27) where  $l_{c1}$ , is a parameter and  $C_a$  is the new assimilated carbon

# Allocation of new assimilates to the roots $C_{a\rightarrow Root}$

 $C_{a \to \text{Root}} = (1 - l_{c1}) \cdot C_a$  (eq. A28) where  $l_{c1}$ , is a parameter and  $C_a$  the new assimilated carbon.

# The plant growth and maintenance respiration processes (g C m<sup>-2</sup> day<sup>-1</sup>) are estimate by

$$C_{\text{respleaf}} = k_{\text{mrespleaf}} \cdot f(T_a) \cdot C_{\text{leaf}} + k_{\text{gresp}} \cdot C_{a \to \text{Leaf}}$$
 (eq. A29)

Where  $k_{\rm mrespleaf}$  is the maintenance coefficient for leaves,  $k_{\rm gresp}$  is the growth respiration coefficient,  $f(T_a)$  is the temperature response,  $C_{a \to {\rm Leaf}}$  is the carbon allocation to leaf and  $C_{\rm leaf}$  is the carbon content of leaf. Similar methods were used for stem and root respiration estimates. In this case, the equation calculates respiration from stem and roots by substituting  $k_{\rm mrespleaf}$  with  $k_{\rm mrespstem}$ , &  $k_{\rm mresproot}$  and using the corresponding storage pools. Respiration from the old carbon pools is estimated with the same maintenance respiration coefficients that is used for respiration from new carbon pools.

### Temperature response for maintenance respiration

$$f(T) = t_{Q10}^{(T-t_{\text{Qlose}})/10}$$
 (eq. A30) where  $t_{\text{Q10}}$  and  $t_{\text{Q10 bas}}$  are parameters.

# Reallocation of C from leaf pool to stem pool. It is represented here as pool for senescent leaves

 $C_{\text{Leaf} \to \text{Sem}} = I_{LS} \cdot C_{\text{Leaf}}$  where  $l_{LS}$  is a parameter and  $C_{L \text{ eaf}}$  the carbon in the leaf pool.

### Leaf C entering the surface litter pool

$$C_{\text{Leaf} \to \text{LitterSurface}} = f(T_{\text{Sum}}) \cdot f(A_l) \cdot S_{\text{newleaf}} \cdot C_{\text{Leaf}}$$
 (eq. A31)

where  $s_{\text{newleaf}}$  is a scaling factor and stem C is calculated analogously with  $S_{\text{newstem}}$ 

### Leaf litter fall dependence of temperature sum

$$f(T_{Sum}) = l_{Lcl} + (l_{Lc2} - l_{Lcl}) \cdot \min\left(1, \frac{\max(0, T_{Sum} - t_{L1})}{\max(1, t_{L2} - t_{LI})}\right) \text{ (eq. A32)}$$

where  $t_{L1}$ ,  $t_{L2}$ ,  $l_{Lc1}$  and  $l_{Lc2}$  are parameters and  $T_{Sum}$  is the so called "dorming" temperature sum,  $T_{DormSum}$ .  $T_{DormSum}$  is calculated at the end to the growing season when air temperature is below the

threshold temperature  $T_{\text{DormTth}}$ , as the accumulated difference between  $T_{\text{DormTth}}$  and Ta.  $T_{\text{DormTth}}$  is a parameter. The stem litter rate is calculated analogously with the parameters  $t_{Sl}$ ,  $t_{S2}$ ,  $l_{Sl}$  and  $l_{Sc2}$ .

# Leaf litter fall dependence of LAI

$$f(A_l) = e^{l_{L_{aiEnh}} \cdot A_l}$$
 (eq. A33)

where  $I_{\text{LaiEnh}}$  is a parameter and  $A_l$  the leaf area index

### Root C entering the soil litter pool of the corresponding layer

$$C_{\text{Root} \to \text{Litter}} = f(l_{\text{Rc}}) \cdot C_{\text{Root}} \cdot S_{\text{newroot}}$$
 (eq. A34)

where  $s_{\text{newroot}}$  is a scaling factor. The root litter rate function,  $f(l_{Rc})$ , can be calculated with Eq. (A32) by exchanging the parameters  $t_{L1}$ ,  $t_{L2}$ , lLcI and  $l_{Lc2}$  to  $t_{Rl}$ ,  $t_{R2}$ ,  $l_{Rc1}$  and  $l_{Rc2}$ .

Litter fall from roots, leaves and stems in "old" biomass are calculated similarly to the "new" biomass but with the important exception that some of the old leaves may be retained

$$C_{\text{OldLeaf} \rightarrow \text{LitterSurface}} = f(l_{Lc}) \cdot (C_{\text{OldLeaf}} - C_{\text{RemainLeaf}}) s_{\text{old leaf}} \text{ (eq. A35)}$$

where or  $s_{\rm oldeaf}$  is a scaling factor. The litter fall for stems and roots is calculated analogously.

# Fraction of the whole $C_{OldLeaf}$ pool that will be excluded from the calculation of the litterfall from the old leaves

$$C_{\text{RemainLeaf}} = C_{\text{OldLeaf}} \left( 1 - \frac{1}{l_{life} - 1} \right)$$
 where  $l_{\text{life}}$  is a parameter (eq. A36)

# Allocation to the mobile C pool for developing new leaves during litter fall

$$C_{\text{Mobile}} = (C_{\text{Leaf} \rightarrow \text{LitterSurface}} + C_{\text{OldLeaf} \rightarrow \text{LitterSurface}}). m_{retain} \text{ (eq. A37)}$$

where  $m_{retain}$  is an allocation coefficient

# Allocation from mobile C pool at leafing as an additional supply. This process goes on as long there is C left in the mobile pool.

$$C_{\text{Mobile} \to \text{Leaf}} = C_{\text{Mobile}} \cdot m_{\text{shoot}}$$
 (eq. A38)

where  $m_{\rm shoot}$  is an allocation coefficient and  $C_{\rm Mobile}$  the carbon in the mobile pool.

### 5. Plant abiotic process

# Plant interception of global radiation (MJ m<sup>-2</sup> day<sup>-1</sup>)

$$R_{s,pl} = \left(1 - e^{-k_{rn}\frac{A_l}{f_{cc}}}\right) \cdot f_{cc} (1 - a_{pl}) R_{is} \text{ (eq. A39)}$$

where  $k_{rn}$  is the light use extinction coefficient given as a single parameter common for all plants,  $f_{cc}$  is the surface canopy cover,  $a_{pl}$  is the plant albedo and  $R_{is}$ , is the global radiation

# Surface canopy cover (m<sup>2</sup>m<sup>-2</sup>)

$$f_{cc} = p_{cmax}(1 - e^{-p_{ck}A_l})$$
 (eq. A40)

Where  $p_{cmax}$  is a parameter that determines the maximum surface cover and  $p_{ck}$  is a parameter that governs the speed at which the maximum surface cover is reached.  $A_l$  is the leaf area index of the plant.

# Leaf Area Index (m<sup>2</sup>m<sup>-2</sup>)

$$A_l = \frac{B_l}{p_{l,sp}}$$
 (eq. A41)

Where  $p_{l,sp}$  is a parameter estimated from specific leaf area and  $B_l$  is the total mass of leaf (i.e leaf C content in the leaves,  $C_{\text{Leaf}} + C_{\text{OldLeaf}}$ ).

# Potential transpiration (m<sup>2</sup>m<sup>-2</sup>)

$$L_v E_{tp} = \frac{\Delta R_n + \rho_a c_p \frac{(e_S - e_a)}{r_a}}{\Delta + \gamma \left(1 + \frac{r_S}{r_a}\right)}$$
 (eq. A42)

where  $R_n$  is net radiation available for transpiration,  $e_s$  is the vapour pressure at saturation,  $e_a$  is the actual vapour pressure,  $\rho_a$  is air density,  $c_p$  is the specific heat of air at constant pressure,  $L_v$  is the latent heat of vaporisation,  $\Delta$  is the slope of saturated vapour pressure versus temperature curve,  $\gamma$  is the psychrometer 'constant',  $r_s$  is 'effective' surface resistance and  $r_a$  is the aerodynamic resistance.

$$r_s = \frac{1}{\max(A_1 g_1, 0.001)}$$
 (eq. A43)

where  $g_l$  is the leaf conductance.

$$g_l = \frac{R_{ls}}{R_{ls} + g_{ris}} \frac{g_{\text{max}}}{1 + \frac{(e_s - e_a)}{g_{vpd}}}$$
 (eq. A43)

where  $g_{ris}$ ,  $g_{max}$  and  $g_{vpd}$  are parameter values,  $g_{maxwin}$  corresponds to  $g_{vpd}$  in winter.  $R_{is}$ , is the global radiation and  $(e_s - e_a)$  the vapour pressure deficit.

## 6. Soil respiration process

Decomposition of soil organic matter was simulated by first-order kinetics (eqs. A44-A45),

# Decomposition of the fast C pools (g C m<sup>-2</sup> day<sup>-1</sup>)

$$C_{\text{DecompL}} = k_l . f(T) . f(\theta) . C_{\text{labile}}$$
 (eq. A44)

where  $k_l$  is a parameter that describes the 1<sup>st</sup> order decomposition rates and f(T) and  $f(\theta)$  are response functions for soil temperature and moisture in the certain layer

# Decomposition of the slow C pools (g C m<sup>-2</sup> day<sup>-1</sup>)

$$C_{\text{DecompL}} = k_h . f(T) . f(\theta) . C_{\text{refractory}}$$
 (eq. A45)

where  $k_h$  is a parameter that describes the decomposition rates and f(T) and  $f(\theta)$  are response functions for soil temperature and moisture in the certain layer

### Response function for soil temperature

$$f(T) = t_{Q10}^{(T - t_{Q10bas})/10}$$
 (eq. A46)

Where  $t_{OIO}$  and  $t_{O10bas}$  are parameters and T is the soil temperature in the certain layer.

### Response function for soil moisture was calculated by (eq. A47)

$$p_{\theta \text{satact}} \qquad \qquad \theta = \theta_{s}$$
 
$$f(\theta) = \min \begin{pmatrix} \left(\frac{\theta_{s} - \theta}{p_{\theta U_{pp}}}\right)^{p_{\theta p}} (1 - p_{\theta \text{ satact}}) + p_{\theta \text{satact}}, \\ \left(\frac{\theta - \theta_{\text{wilt}}}{p_{\theta Low}}\right)^{p_{\theta p}} \end{pmatrix} \theta_{\text{wilt}} \leq \theta \leq \theta_{s}$$
 
$$0 \qquad \qquad \theta < \theta_{\text{wilt}}$$

where  $p_{\theta \text{Upp}}$ ,  $p_{\theta \text{Low}}$ ,  $p_{\theta \text{Satact}}$  and  $p_{\theta \rho}$  are parameters and the variables,  $\theta_{s}$ ,  $\theta_{\text{wilt}}$ , and  $\theta$ , are the soil moisture content at saturation, the soil moisture content at the wilting point, and the actual soil moisture content, respectively. Soil respiration was assumed to be optimal at 60% and decreased linearly with either water content above or below (Or et al., 2007). The produced CO<sub>2</sub> was assumed to emit directly into the atmosphere thus no transpiration of CO2 gas in the soil profile was simulated.

## Litter from inactive surface litter pool entering the fast SOC pool at continuous rate

$$C_{\text{LitterSurface} \rightarrow \text{Litter } l} = l_{l1} \cdot C_{\text{LitterSurface}}$$
 (eq. A48)

where  $I_{l1}$  is a parameter and  $C_{Litter\ Surface}$  the carbon in the surface litter pool.

# Amount of decomposition product from fast SOC pools being released as CO2

$$C_{\text{Litter}\rightarrow CO_2} = (1 - f_{e,l}) \cdot C_{\text{DecompL}} \text{ (eq. A49)}$$

# Amount of decomposition products from fast SOC pools entering the slow decomposition pools

$$C_{\text{Litter-> refractory}} = f_{e,l} \cdot f_{h,l} \cdot C_{\text{DecompL}}$$
 (eq. A50)

Amount of decomposition products from the slow SOC pools being returned to the fast decomposition pools

$$C_{\text{Litter} \rightarrow \text{Litter}} = f_{e,l} (1 - f_{h,l}) \cdot C_{\text{DecompL}}$$
 (eq. A51)

Amount of decomposition products from the slow SOC pools being released as CO<sub>2</sub>

$$C_{\text{refractory}} = f_{e,l} \cdot C_{\text{Demmpl}} \text{ (eq. A52)}$$

Where  $C_{\text{DecompL}}$  represents the decomposition rate of litter pool,  $f_{e,l}$  is the efficiency of decay of litter and  $f_{h,l}$  is the fraction of C and N in the labile pool that will enter the refractory C pool.

S2: Initial Carbon and Nitrogen layers used for Beverly Swamp in Coup

Depth	Init L1C (active)	Init HC (passive)	Init L1N (active)	Init HN (passive)
	(g/m2)	(g/m2)	(gN/m2)	(gN/m2)
0-5	1067.949	1220.513	30.51282	61.02564
5-10	1719.494	1965.136	49.12839	98.25678
10-15	1747.074	1996.656	49.9164	99.8328
15-20	311.4887	5161.813	8.899677	258.0906
20-25	207.8285	3444.015	5.937956	172.2007
25-30	358.1405	5934.9	10.23259	296.745
30-40	1020.338	5247.451	29.1525	262.3725
40-50	1056.745	5434.69	30.19272	271.7345
50-65	649.2692	3339.099	18.55055	166.9549
65-80	1091.221	5611.992	31.17774	280.5996
80-100	1111.299	18415.81	31.7514	920.7905
100-120	647.3291	10727.17	18.49512	536.3584
120-150	1058.263	17536.93	30.23609	876.8466

# S3: List of model parameters used in the model run that differ from the model default for Beverly swamp

Symbol	Parameters	Value	Units	References
Zo	Surface roughness length	2.7±0.5	m	Munro 1987 & Munro et al 2000
$P_{\rm ral}$	Air resistance inside canopy	6.7-10	s/m	Munro 1987 & Munro et al 2000
g <sub>max</sub>	Maximal conductance of fully stomata	0.01-0.02	m/s	Munro 1987 & Munro et al 2000
$\mathbf{g}_{\mathrm{ris}}$	Global radiation intensity that represents half-light saturation in light response	5300000	J/m²/day	Thambipillai, 2000
$g_{ m vpd}$	Vapour pressure deficit that corresponds to a 50% reduction of stomata conductance	125	Pa	Munro 1987 & Munro et al 2000
$lpha_{ m dry}$	Soil albedo when tension >104cm H <sub>2</sub> O	7	%	Munro et al 2000; Thambipillai, 2000
$lpha_{ m wet}$	Soil albedo when tension <10cm H <sub>2</sub> O	15	%	Munro et al 2000; Thambipillai, 2000
$lpha_{leaf}$	Albedo of vegetation canopy	15-16	%	Munro et al 2000; Thambipillai, 2000
kB <sup>-1</sup>	Difference between the natural logarithm of surface roughness length for momentum and heat	2.5	-	Munro et al 2000; Thambipillai, 2000
$ heta_{ m sat}$	Total porosity	80-85	Vol %	Czerneda, 1985 & Munro 1982
$\theta_{\mathrm{m}}$	Macroporosity	18-28	Vol %	Word et al., 2022

Ksat	Total saturated hydraulic conductivity <sup>1</sup>	125-0.08	mm d-1	Gupta et al. (2023); Czerneda, 1985
$lpha_{ m surf}$	First-order coefficient for surface runoff	0.4	_	assumed
$P_{\theta Low}$	Lower range for moisture response	12	Vol%	assumed
$P_{ heta Low}$	Upper range for moisture response	12	Vol%	assumed
$lpha_{ m scale}$	The first-order coefficient for the surface runoof	0.1	_	Assumed
$\theta_{\rm r}$	Residual water content	22	Vol %	Dimitrov and Lafleur (2021); Menberu et al
$ heta_{ m wilt}$	Wilting point	30	Vol %	(2021) ; Letts et al. (2000); Liu and Lennartz et al. (2018)
Wpmax	Maximum amount of water stored on soil surface without causing surface runoff	200	mm	Valverde, 1978; Woo and Valverde (1981)
$t_{WA}$	Temperature coefficient in the temperature response function	0.59		Assumed
twB	Temperature coefficient in the temperature response function	0.28		Assumed
$\Delta z_{humus}$	The thickness of peat layer	0.85	m	Munro et al. 2000
$k_{rn}$	Extinction coefficient	0.52		Thambipillai, 2000

 $<sup>^{1}</sup>$  K<sub>sat</sub> values were used for the simulated 12 layers which is arranged from top to bottom layer above 13

$l_{life}$	Max leaf lifetime	2	Years	Welch, 1985
<b>P</b> life	Max plant lifetime	400	years	Kanda et al (1996); Withington et al. (2006)
$arepsilon_L$	Radiation use efficiency for photosynthesis at optimum temperature, moisture and C-N ratio	2.8 - 4	gDw/MJ	Horn & Schulz, 2010; Wang et al., 2020
$k_{ m gresp}$	Growth respiration coefficient	0.17	Day-1	
$\mathbf{k}_{ ext{mrespleaf}}$	Maintenance respiration coefficient for leaves	0.002	Day-1	
kmrespstem	Maintenance respiration coefficient for stem	0.0001	Day-1	
kmresproot	Maintenance respiration coefficient for root	0.015	Day-1	Amthor & Baldochhi (2001); Amthor 1984; Amthor 1986;
kmrespcroot	Maintenance respiration coefficient for coarse root	0.0008	Day-1	Cannel & Thornley, 2000
mretain	coefficient for determining allocation to mobile internal storage pool	0.9	Day-1	assumed
mshoot	Coefficient for determining allocation from the mobile pool to the leaf at leafing.	0.5	Day-1	assumed
$l_{ m Lc1}$	rate coefficient for the leaf litter fall before the first threshold	0.001	Day-1	Thambipillai, 2000

	temperature sum tL1 is reached			
l <sub>Lc2</sub>	rate coefficient for the leaf litter fall after the second threshold temperature sum tL2 is reached	0.05	Day-1	Thambipillai, 2000
ISc1	rate coefficient for the litter fall from stems before the first threshold temperature sum tS1 is reached	1.73e-05	Day-1	Thambipillai, 2000
k <sub>l</sub>	First-order decomposition coefficient for labile C	0.001	Day	Field measurement (unpublished), (Literature) Kendall et al (2020)
k <sub>h</sub>	First-order decomposition coefficient for refractory C	3 x 10-6	day	Field measurement (unpublished), Middleton et al. (2020)
l <sub>h</sub>	Fraction of above ground residues that enter litter 1 pool	0.001	day	Kendall et al (2020
cn <sub>m</sub>	C:N of soil microbes	7	-	Wang et al. (2015)
C <sub>tot</sub>	Total soil C at 1.5 m profile	98083	gC/m <sup>2</sup>	Waddington et al. (unpublished data); Schmidt et al. (2024)
C <sub>tot, layer</sub>	Total soil C for each simulated layer* <sup>2</sup>	13623	gC/m <sup>2</sup>	
Q <sub>10</sub>	Q <sub>10</sub> value for decomposition	2-3	-	Byun et al. (2021)
T <sub>aamp</sub>	Amplitude of analytical air temperature	12	оС	Weather station data

<sup>-</sup>

 $<sup>^{\</sup>rm 2}$  This is an average as different values were used for different layers

T <sub>amean</sub>	Mean value in air	7.64	oC	Weather station data
	temperature			
	function (also for			
	estimating lower			
	boundary condition)			

# S4: Result of initial parameter screening using One-at-a-time

Parameter	Temperature (1)	Temperature (6)	Saturation Level	TotalWater Content(1)	TotalWater Content(6)	Soil respiration	Radiation Net Tot	TotalLatent Flow	TotalSensible Flow	LeafArea Index(1)	Snow Depth	Surface pool
TemQ10												
TempAirMean												
Theta Lower Range												
Theta upper Range												
Reference Height												
AlbedoLeaf												
TemQ10Bas												
Albedo Dry												
Albedo Wet												
CritThresholdDry												
DrainSpacing												
PrecA0Corr												
PrecA1Corr												
MeltCoefAirTemp												
MeltCoefGlobRad												
KBMinusOne												
RaIncreaseWithLAI												
WindLessExchange Canopy												
WithinCanopyRes												
TempDiffPrec_Air												
CN Ratio Microbe												
RateCoefHumus												
RateCoefLitter1												
RateCoefSurf L1												
SaturationActivity												

AScaleSorption						
SPMaxCover						
SurfCoef						
SPCoverTotal						
SurfPoolMax						
GWSourceFlow						
GWSourceLayer						
AirMinContent						
TempCoefA						
TempCoefB						
CFrozenSurfCorr						
Maximal Cover						
Specific LeafArea						
MaxLeafLifetime						
Max Plant Lifetime						
Shoot Coef						
MobileAlloCoef						
RadEfficiency						
Leafc1						
RootWaterc1						
RootCNc1						
RootMassc1						
GrowthCoef						
MCoefLeaf						
MCoefStem						
MCoefRoot						
MCoefCoarseRoot						
LeafRate1						
LeafRate2						
LeafTsum1						

stemRate1						
stemRate2						
StemTsum1						
RootRate1						
Rootrate2						
CoarseRootRate1						
CoarseRootRate2						
LAI enh Coef						
Roughness Min						
AirResisLAi Effect						
Conduct Ris						
Conduct VPD						
Conduct Max						
Roughness max						
n-value (1)						
n-value (6)						
n-value (9)						
n-value (11)						
Alpha(1)						
Alpha(6)						
Alpha (9)						
Alpha(11)						
Wilting point						
Residual water Total						
conductivity(1)						
Total conductivity(6)						
Total						
conductivity(9) Total						
conductivity(11) Total						
conductivity(12)						

n Tortuosity						
Macropore(1)						
Macropore(6)						
Macropore(9)						
Macropore(11)						
Macropore(12)						

	Sensitivity classes	
Class	Index	Sensitivity
I	$0.00 \le  I  \le 0.05$	small to negligible
II	$0.05 \le  I  < 0.20$	Medium
III	$0.20 \le  I  \le 1.00$	High
IV	$ \mathbf{I}  \ge 1.0$	Very High

# **S5:** Correlation matrix of calibrated parameters

-	k <sub>mre speo</sub> €	L	p <sub>0p</sub>	P <sub>esasa</sub> a	p <sub>oll so</sub>	<b>G</b> <sub>Amin</sub>	xhf	р	olow F	D <sub>zmot</sub>	k <sub>h</sub>	C <sub>rain</sub>	l <sub>et</sub>	wa.	I <sub>Lc2</sub>	c <sub>max</sub>	lsp	a <sub>eg</sub> t	<sub>pQ10bas</sub> k	, 4	, t	210	k <sub>gresp</sub>	q, <sub>of</sub>	g <sub>ed</sub>	g <sub>max</sub>	m <sub>retain</sub>	kB-1	m <sub>T</sub>	a <sub>wet</sub>	d, I	Al <sub>erico</sub> (	ļ <sub>i</sub> l	le1	P <sub>maxt</sub>	W <sub>pmax</sub>	q <sub>sol</sub> k	C <sub>sati</sub>	K <sub>sat6</sub>
k <sub>mesposa</sub>																																							
seroot		-0.3	-0.2	0	0	-0	.3	0	0.4	0.2	0	0.1	0.1	-0.4	0.1	0	0.3	0.3	-0.2	-0.1	0	-0.5	0	0.2	0.4	0.3	0.1	0.1	0.2	0	0	0.2	0.2	-0.2	-0.1	-0.4	0	0.3	0.1
eį			0	0.1	-0.1	l -0	.1	0	0	0	0.3	-0.1	0	0.2	0.1	0.2	0	-0.3	-0.1	0.2	-0.1	0.2	0	-0.4	0.1	-0.1	-0.1	0.2	0.4	0.3	-0.3	-0.2	0.1	0.5	-0.3	0.1	-0.3	-0.5	0.1
p <sub>0p</sub>				0	0.2	-0	.1 (	0.1	-0.3	0.3	0	0.1	-0.1	0.4	0.4	0	0	0.1	0.4	0.2	0.3	0.2	0	0	-0.3	-0.1	0	-0.2	-0.3	0.2	0.3	-0.2	-0.3	0	0	-0.2	0	-0.1	0.2
Pesatact					0	-0	.1 -0	0.3	0.3	0	-0.2	0.4	-0.1	0	-0.2	-0.1	0.2	-0.3	0	0.1	0.2	0.3	0.4	-0.2	-0.2	0.5	0.2	-0.2	0.2	0.4	-0.5	0	0.4	0	-0.2	0.3	-0.2	0	0.3
P <sub>qUpp</sub>						-0	.3	0	0.4	-0.2	0.1	0.4	-0.1	-0.2	0.1	0.1	0	-0.4	0.5	0.3	-0.4	0	-0.3	0	0.3	0.3	-0.4	0.1	-0.4	0.4	0.3	-0.2	-0.3	0.5	0.4	0.2	0.3	-0.1	-0.2
<b>G</b> <sub>kmin</sub>							-(	0.4	-0.1	0.2	-0.2	-0.2	0.3	0.2	-0.1	-0.3	0	0.3	-0.3	-0.1	0.2	0.1	0.2	0	-0.2	-0.2	0.4	0.2	-0.1	-0.6	0	0.4	-0.1	-0.1	-0.1	-0.1	0.3	0.1	0.2
xhf									-0.3	-0.2	0.6	-0.1	0.1	0	-0.1	0.1	-0.1	0	0.2	0	0	-0.1	0	0.1	-0.1	-0.4	-0.5	-0.2	0.1	-0.2	0.1	-0.2	0	-0.2	0.2	-0.2	0	0	-0.4
P <sub>qlow</sub>										-0.1	-0.1	0.6	-0.2	-0.7	0	-0.2	0.1	-0.3	0.1	0	-0.5	0.2	0.2	-0.3	0.4	0.6	0	0	0.2	0.2	-0.3	0	-0.1	0.1	0.1	0.2	0.2	0	0.2
P <sub>aroot</sub>											-0.2	0.1	-0.1	0	0.1	-0.4	0.1	0.4	-0.3	0.3	0.1	-0.2	0.4	-0.2	-0.1	0	0.2	0.4	0	-0.2	0	0	-0.1	-0.2	-0.1	-0.5	-0.2	0.3	0.3
k,												-0.1	0	0.2	0	0.1	-0.1	-0.2	0.3	0.2	0	0.1	0.1	-0.3	0.1	-0.3	-0.4	0.2	0.1	0.1	0.2	0.2	0.3	0.3	0.3	0	-0.1	-0.2	-0.3
Crain													-0.3	-0.3	-0.2	-0.5	-0.1	-0.3	0.4	0.1	-0.2	0.2	0.5	-0.5	-0.1	0.4	-0.1	-0.1	. 0	0.2	0	0	-0.1	-0.1	0.5	0.3	-0.1	0	0.3
l <sub>d</sub>														0.3	-0.1	0.2	0.5	0.7	-0.2	-0.6	0.4	-0.6	-0.2	0.6	-0.2	-0.2	-0.2	-0.3	-0.3	-0.2	0	0.1	-0.1	0.2	0	-0.4	0	-0.3	0
t <sub>WA</sub>															-0.1	0.3	0.1	0.2	0	0	0.6	0	0	0.2	-0.2	-0.4	0.3	-0.1	. 0	-0.1	0.2	0.3	0.1	0.1	-0.1	0	0	-0.2	0
l <sub>1/2</sub>																-0.1	-0.1	0.1	0.2	0.2	0.2	0.1	-0.5	-0.1	-0.1	-0.1	0	-0.2	-0.2	0.3	-0.1	-0.4	-0.2	0.2	-0.2	-0.3	-0.2	-0.2	0.2
C <sub>max</sub>																	0.4	-0.1	-0.1	0	0.2	-0.1	-0.5	0.5	0.2	0.1	-0.1	-0.2	-0.1	0.3	0	0	0.1	0.3	-0.1	0	0.3	-0.5	-0.4
Pi.sp																		0.6	-0.2	-0.2	0.3	-0.4	-0.1	0.5	0	0.5	0.1	-0.3	-0.1	0.1	-0.1	0.2	0.2	0	0	-0.4	0	-0.4	-0.1
a <sub>eg</sub>																			-0.3	-0.5	0.4	-0.6	0	0.5	-0.2	-0.1	0.1	-0.2	-0.2	-0.3	0.1	0.2	-0.2	-0.3	-0.1	-0.7	-0.2	0	0.1
t <sub>pQL0bas</sub>																				0	-0.3	0.4	-0.1	-0.2	-0.1	0	-0.4	-0.2	-0.4	0.3	0.5	-0.1	-0.4	0.3	0.4	0.3	0	-0.2	0.2
k,																					-0.2	0.4	0.1	-0.4	0.1	0.2	0.2	0.5	0.2	0.2	-0.1	-0.2	0.3	0.1	-0.1	0.1	0.1	0.2	-0.3
l <sub>tt</sub>																						-0.1	0	0.3	-0.4	-0.2	0.3	-0.5	0	0	-0.2	0.1	0.3	-0.2	-0.2	-0.3	-0.1	-0.2	0.1
tae																							0.1	-0.6	0.1	-0.1	0.2	0	0.2	0.1	-0.2	-0.1	0.1	0.1	0	0.6	0.2	-0.1	0.3
k <sub>gresp</sub>																								-0.5	-0.1	0.1	0.1	0.3	0.2	-0.2	-0.1	0.5	0.3	-0.4	0.1	0.1	-0.2	0.3	0.2
Q <sub>tof</sub>																									0	-0.1	0		-0.2	-0.2	0.1	-0.1	-0.1	-0.1	-0.1	-0.3	0.3	0	-0.3
Sept																										0					0	0.2	0.1	0			0.4	0.2	0
g <sub>rmax</sub>																											0.2	0				0	0.2	0			0	-0.1	-0.1
m <sub>retain</sub>																												0.2	0.4	-0.2	-0.2	0.4	0.4	-0.4	-0.5	0	0.2	0.2	0.2
kB <sup>-1</sup>																													0.1			0.4	0.1	0.1	-	-	0	0.5	-0.1
m <sub>T</sub>																														-0.2		-0.1	0.5	-0.4		0.1	0	0.2	0.1
a <sub>wet</sub>																															0		0	0.4			-0.1	-0.3	0
d <sub>p</sub>																																0.3	-0.5	0.1	0.4	0	0.1	0.1	-0.2
LAI <sub>enho</sub>																																							
oef																																	0.3	-0.1			0.2	0.2	0
<b>S</b> ris																																		-0.3			-0.1	0.1	-0.3
l <sub>te1</sub>																																			0.2		-0.2	-0.5	0.1
P <sub>maxt</sub>																																				0.2	0	-0.2	-0.1
W <sub>prrex</sub>																																					0.2	0.1	
Q <sub>tol</sub>																																						0.1	-0.3
k <sub>sat1</sub>																																							0

# S6: Result of parameter contribution to variable performance with LMG Approach

Category	Parameter	Soilresp(cal)	Soilresp(val)	Stemp5	Stemp30	SL (cal)	SL	VMC5	VMC30	Nrad	LH	SH	LAI	SnowD	SP (cal)	SP (val)
plant growth	MCoefCoarseRoot	2.121416	1.477459	8.237431	12.1217	1.187035	17.74539	1.958169	4.767888	6.843652	2.5908	3.96741	10.120669	0.7853312	7.533316	6.719465
water/abiotic	ThetaPowerCoef	7.31755	2.090769	7.294556	3.988291	5.82293	2.310807	9.194762	2.728837	1.641421	2.517004	13.08791	1.845788	9.9324373	2.711138	4.553528
temp and decom-abiotic	TemQ10	3.38563	3.591051	15.32889	5.032576	3.499937	4.164385	1.835659	1.754209	6.103535	1.248388	1.856529	7.562064	0.974746	11.13657	3.54019
water-abiotic	Saturation activity	19.032887	6.029802	3.726173	8.712364	2.779985	1.768566	6.311641	5.844128	13.70476	3.935608	5.333488	6.856898	2.1444549	1.165058	2.954025
soil thermal	ThScaleLog(1)	2.77317	5.571461	4.079623	5.135376	4.40426	3.929442	4.940231	5.749493	10.63224	4.292136	7.725473	8.078626	9.7169917	2.218961	3.019597
soil organic	RateCoefHumus	2.002456	4.926929	2.882412	1.993718	2.652958	6.900096	3.326349	2.480241	3.123797	6.954205	7.417158	8.072348	2.1838908	4.72409	2.614537
meteorological	PrecA0Corr	3.926877	6.32134	3.140031	1.147461	2.66219	1.225755	7.812461	5.822728	1.504695	3.863147	3.205525	1.508219	1.7608698	9.293432	8.184454
plant	AlbedoLeaf	3.637218	2.044274	7.800618	5.18149	3.673719	10.14206	7.674131	1.644069	8.028443	4.908113	5.794787	5.340018	2.5878091	7.580291	10.59164
soil organic	RateCoefLitter1	2.283223	4.270437	3.114548	5.069016	6.191558	1.756832	1.220147	3.971254	3.271041	7.293449	3.626779	8.778116	1.2721331	1.795226	1.437049
soil organic	RateCoefSurfL1	4.098293	2.965316	5.875754	3.252724	1.77197	12.68042	7.161687	5.493897	2.456879	5.269038	3.192613	2.106959	1.449582	2.74919	9.2759
plant growth	GrowthCoef	10.574973	21.048423	2.033687	2.959492	11.60148	2.642283	13.57133	11.97936	5.986806	10.26584	5.410165	7.162934	3.3258733	2.559314	10.2463
Drainage and deep percola	GWSourceFlow	7.342453	10.080261	1.7985	5.581834	4.069777	3.404689	7.94457	3.722327	11.74869	5.1041	11.87913	4.499984	4.1570726	7.660973	11.73978
transpiration (plant )	Conduct Max	3.597613	8.882092	3.841812	8.000887	2.412481	8.196215	4.953836	3.441733	7.459955	4.108898	3.651364	4.448836	2.4006713	3.50104	1.712176
plant	MobileAlloCoef	3.726829	7.44953	4.444536	5.736773	5.692012	4.188876	9.603209	12.94847	7.071648	3.708091	7.459666	2.539816	4.4515116	9.891208	8.456312
snowpack	MeltCoefAirTemp	7.115477	2.370553	14.26561	8.392565	15.5446	4.61058	4.321144	7.912447	2.800823	5.317808	5.136222	2.393013	40.1541276	16.568312	4.167339
radiation/ET	Albedo Wet	2.933628	3.97973	5.651772	5.521492	7.196675	2.909862	2.061132	5.983924	1.994771	18.91448	2.283797	1.521929	4.0062391	2.491929	1.494299
plant	LAI enh Coef	10.300507	2.741922	3.166416	8.150327	9.233269	3.870245	3.35022	8.524324	2.152355	3.188891	3.261322	11.081197	3.8636574	2.712412	2.13573
plant	LeafRate1	3.8298	4.158654	3.317631	4.021919	9.603164	7.553502	2.759323	5.230672	3.4745	6.520004	5.710656	6.082587	4.8326011	3.70754	7.157681