

# Criegee + HONO reaction: a bimolecular sink of Criegee, and the missing non-photolytic source of OH•

Vishva Jeet Anand<sup>1</sup>, Philips Kumar Rai<sup>1</sup>, and Pradeep Kumar<sup>1</sup>

<sup>1</sup>Department of Chemistry, Malaviya National Institute of Technology Jaipur, Jaipur, 302017, India

**Correspondence:** Pradeep Kumar (pradeep.chy@mnit.ac.in)

1 **Abstract.** One of the most important puzzles in atmospheric chemistry is a mismatch between observed and modelled con-  
2 centrations of OH•/HO<sub>2</sub>• in the presence of high concentration of volatile organic compounds. It is now well established that  
3 to fulfill this gap, one needs a reaction that is not only capable of producing OH• but also able to act as a sink of HO<sub>2</sub>•. In the  
4 present work, we are proposing the Criegee + HONO reaction as a possible solution of this puzzle. Our quantum chemical and  
5 kinetic calculations clearly suggest that this reaction can not only be an important source of OH radical but can also act as a sink  
6 of HO<sub>2</sub> radical. Our study also suggests that HONO has the potential to act as a bimolecular sink of Criegee intermediates, and  
7 for some Criegee intermediates under certain atmospheric condition it can even surpass the traditionally known bimolecular  
8 sinks such as SO<sub>2</sub> and water dimer, even in high humid conditions.

## 9 1 Introduction

10 It is well-known that the atmospheric chemistry is mainly dominated by the radicals (Anderson, 1987; Monks, 2005). Particu-  
11 larly in the troposphere, these radicals are key in degrading various pollutants, a phenomenon as important as the ozone layer  
12 for the existence of life (Weinstock, 1969; Lelieveld et al., 2004). The primary radicals responsible for the oxidative power of  
13 troposphere come from the HO<sub>X</sub> (OH•, HO<sub>2</sub>•, RO•, RO<sub>2</sub>• etc.) family (Prinn, 2003; Ehhalt, 1987; Khan et al., 2018). Among  
14 them, OH• is considered as the most important oxidant in the troposphere (Lelieveld et al., 2002, 2016). Although OH• is the  
15 most studied radical in the atmosphere, there are still open questions regarding its sources in the atmosphere (Heald and Kroll,  
16 2021; Yang et al., 2024). For a long time, it was believed that OH radicals are mainly formed in daytime via photolysis of  
17 tropospheric ozone (O<sub>3</sub>), and nitrous acid (HONO) (Calvert et al., 1994; Alicke et al., 2003; Griffith et al., 2016; Aumont et al.,  
18 2003). But now, with various on-field measurements (Geyer et al., 2003; Ren et al., 2003; Emmerson and Carslaw, 2009), it  
19 is well established that OH radicals are also present at night in sufficient amounts. In fact, average nighttime concentration of  
20 OH• ( $\sim 2.6 \times 10^5$  molecule cm<sup>-3</sup>) is only one order of magnitude lower than its average daytime concentration ( $\sim 1.9 \times 10^6$   
21 molecule cm<sup>-3</sup>) (Emmerson and Carslaw, 2009). As the lifetime of OH• is only  $\sim 1$  second, this much concentration of  
22 OH• during night indicates its *in situ* generation via non-photolytic sources. The major non-photolytic source of OH• is the  
23 recycling of HO<sub>2</sub>• radicals (Whalley et al., 2011; Stone et al., 2012; Hofzumahaus et al., 2009; Smith et al., 2006; Hens et al.,  
24 2013). Specifically, during the daytime, the primary reaction contributing to this recycling process is NO• + HO<sub>2</sub>•, whereas  
25 at night, the key reaction is NO<sub>3</sub>• + HO<sub>2</sub>• (Hall et al., 1988; Mellouki et al., 1988, 1993; Rai and Kumar, 2024). However,  
26 compared to photolytic sources, non-photolytic sources of OH• remain less understood in atmospheric chemistry (Brown and  
27 Stutz, 2012; Emmerson and Carslaw, 2009). This is evidenced by the fact that, in the atmosphere with a high concentration

28 of volatile organic compounds (VOCs), atmospheric models consistently under-predict the concentration of OH• compared to  
29 the observed value (Emmerson and Carslaw, 2009; Stone et al., 2012). This discrepancy is especially pronounced in winter  
30 (Harrison et al., 2006; Heard et al., 2004; Slater et al., 2020) and indoor environments (Østerstrøm et al., 2025; Gomez Alvarez  
31 et al., 2013; Reidy et al., 2023), where light plays a minimal role. In addition, the discrepancy between measured and observed  
32 value of OH• was also found to depend upon NO<sub>X</sub> concentration. Both under low NO<sub>X</sub> (Carslaw et al., 2001; Tan et al., 2001;  
33 Lelieveld et al., 2008; Tan et al., 2017) as well as high NO<sub>X</sub> (above 6 ppbv) (Slater et al., 2020), the discrepancy was found to  
34 be quite significant. As the primary recycling of HO<sub>2</sub>• to OH• occurs via NO<sub>X</sub>, the under-prediction of OH• by models under  
35 low NO<sub>X</sub> conditions suggests either the presence of another route for recycling or some new non-photolytic source of OH•.  
36 This hypothesis is further strengthened by a few combined experimental and modelling studies. For example, Lu et al. (Lu  
37 et al., 2012) have to introduce an artificial source of OH• ↔ HO<sub>2</sub>• inter-conversion (RO<sub>2</sub>• + X → HO<sub>2</sub>•, HO<sub>2</sub>• + X → OH•) in  
38 their atmospheric model to match the experimental concentration profile. In an another study, to match the experimental OH  
39 concentration with models, Whalley et al. (Whalley et al., 2011) increased the concentration of VOCs in their model. Although  
40 their computed OH• concentration becomes closer to experimental value, the mismatch between observed and measured con-  
41 centration of HO<sub>2</sub>• becomes worse. There have been various attempts to identify the missing source of OH• in the atmosphere  
42 (Paulot et al., 2009; Peeters et al., 2014; Sander et al., 2019). For example, Peeters et al. (Peeters et al., 2009; Peeters and  
43 Müller, 2010; Peeters et al., 2014) suggested that the oxidation of isoprene can regenerate HO<sub>X</sub> radicals in the presence of  
44 light via isoprene-peroxy radical interconversion and isomerisation pathways (Leuven Isoprene Mechanism (LIM)). Although  
45 the introduction of LIM into chemical models were found to improve the value of modelled OH• concentration, the modelled  
46 values still remain under-predicted (Crouse et al., 2011; Teng et al., 2017; Berndt et al., 2019; Novelli et al., 2020; J. Medeiros  
47 et al., 2022). Particularly, the LIM is more effective in regions where biogenic volatile organic compounds (BVOCs) dominate  
48 and NO<sub>X</sub> concentration is ultra low, e.g. rain forest regions (Whalley et al., 2011; Feiner et al., 2016; Lew et al., 2020). In  
49 contrast, in regions where sufficient anthropogenic sources of VOCs are present, e.g. in polluted areas, LIM is not effective.  
50 In addition, LIM is not fundamentally a HO<sub>2</sub>• to OH• interconversion process, rather it is the recycling of VOCs to OH•. In a  
51 recent study, Yang et al. (Yang et al., 2024) suggested that aldehyde could be an additional source of OH•. Authors proposed  
52 that the autoxidation of carbonyl organic peroxy radicals (R(CO)O<sub>2</sub>) derived from higher aldehydes, can produce OH• through  
53 photolysis (RAM mechanism). Though RAM mechanism efficiently predicts OH• production at low NO<sub>X</sub> concentrations, it  
54 still under-predicts the same at high NO<sub>X</sub> concentrations. Interestingly, when both LIM and RAM are incorporated into a base  
55 model in the presence of moderate concentration of NO<sub>X</sub>, OH• concentration improves significantly, but the discrepancy in the  
56 modelled and observed HO<sub>2</sub>• remains unresolved. It is also worth mentioning that photolysis is an important part of both, LIM  
57 and RAM, and hence, both of these mechanism do not offer any help in improving the model OH• concentration in nocturnal  
58 environment. Furthermore, both LIM and RAM are also not directly involved in recycling of HO<sub>2</sub>• to OH•. The discrepancy in  
59 the model occurs during both day and night (Faloona et al., 2001; Hens et al., 2013; Geyer et al., 2003), and is associated with  
60 HO<sub>2</sub>• to OH conversion (Whalley et al., 2011; Hofzumahaus et al., 2009). In light of these studies, we believe that the puzzle  
61 of missing OH• source is very much alive and the key to this puzzle may be a non-photolytic source capable of HO<sub>2</sub>• ↔ OH•  
62 recycling.

63 In the present work, we are proposing reaction of Criegee intermediate with HONO as a source of OH<sup>•</sup>. Criegee Intermediates  
64 (CIs) are formed during the ozonolysis of alkenes (Criegee, 1975; Johnson and Marston, 2008; Taatjes, 2017). In fact, alkene  
65 ozonolysis is a highly exothermic reaction produces energized CIs. Some of the energized CIs readily convert into OH<sup>•</sup> via  
66 unimolecular decomposition, while the remaining CIs get collisionally stabilized (sCI) (Horie and Moortgat, 1991; Donahue  
67 et al., 2011; Novelli et al., 2014; Alam et al., 2011). sCIs can undergo either a thermal unimolecular dissociation or a bimolec-  
68 ular reaction. Depending upon concentration of the co-reactant and rate constant of such bimolecular reaction, the bimolecular  
69 reaction paths can be the main sink of sCI (Osborn and Taatjes, 2015; Lin et al., 2015; Sheps et al., 2014; Vereecken and  
70 Francisco, 2012). There are several studies in the literature that suggest CI reacts rapidly with the trace gases present in the  
71 atmosphere (Cox et al., 2020; Mallick and Kumar, 2020; Vereecken et al., 2015; Long et al., 2016, 2021). In this work, we are  
72 suggesting HONO as a new partner for the bimolecular reaction of Criegee intermediates that is capable of producing OH rad-  
73 icals. The concentration of CI ( $\sim 10^4 - 10^5$  molecule  $\text{cm}^{-3}$ ) in the atmosphere is comparable with Cl<sup>•</sup> ( $\sim 5.0 \times 10^4 - 3.0 \times 10^5$   
74 molecule  $\text{cm}^{-3}$ ) and OH<sup>•</sup> ( $\sim 1.0 \times 10^5 - 4.0 \times 10^6$  molecule  $\text{cm}^{-3}$ ) (Khan et al., 2018; Novelli et al., 2017). Similarly, nitrous  
75 acid (HONO) is also an important trace gas present in the nighttime atmosphere in a considerable amount (Li et al., 2021; Song  
76 et al., 2023). The average concentration of HONO is  $\sim 8.9 \times 10^{10}$  molecule  $\text{cm}^{-3}$ , which can reach as high as  $\sim 6.9 \times 10^{11}$   
77 molecule  $\text{cm}^{-3}$  during the fog event (Pawar et al., 2024). Although a general wisdom about HONO is, its concentration builds  
78 up in nighttime, and in daytime, it decomposes via photolysis to give OH<sup>•</sup>, HONO itself is a highly reactive molecule and can  
79 participate in various bimolecular chemical reactions during night (Anglada and Sole, 2017; Lu et al., 2000; Wallington and  
80 Japar, 1989). Moreover, in indoor environments, high concentrations of OH<sup>•</sup> have been found to strongly correlate with high  
81 concentrations of HONO (Gomez Alvarez et al., 2013). It is important to mention that, the reaction of HONO with the simple  
82 Criegee intermediate (CH<sub>2</sub>OO) has already been investigated theoretically (Kumar et al., 2022). In that investigation, the major  
83 product was predicted to be hydroperoxymethyl nitrite (HPMN). We will show in the present work that the main product of  
84 this reaction is OH<sup>•</sup> and this newly found path is the dominant path of the title reaction.

## 85 2 Methodology

### 86 2.1 Electronic structure theory

87 There are two parts of electronic structure theory; optimization and subsequent single-point energy calculations. The criteria  
88 behind choosing a method for optimization is; it should be computationally not very demanding and at the same time, it should  
89 accurately predict the geometries and frequencies of the species involved in the reaction. Based on these criteria, in the present  
90 work, the CCSD(T)/CBS//M062X/aug-cc-pVTZ level of theory was chosen, which is known to give reasonable results in  
91 various previous studies (Kumar et al., 2022; Vereecken et al., 2017, 2014; Vereecken, 2017) for reactions involving Criegee  
92 intermediates. Gaussian16 software package (Frisch et al., 2016) has been used to carry out all the optimization and single-  
93 point energy calculations. To estimate energies at CCSD(T)/CBS level of theory, first, we calculated the single point energies  
94 at CCSD(T)/aug-cc-pVDZ, and CCSD(T)/aug-cc-pVTZ level of theory, and then extrapolated these energies to corresponding  
95 CBS limit using the method of Varandas and Pansini (Varandas and Pansini, 2014; Pansini et al., 2016) (see ESI for the details).

## 96 2.2 Kinetics

97 Energetics calculations shed light only on enthalpic requirement of the reaction, for a barrierless process, entropy is an equally  
98 important factor. Therefore, to account for both, enthalpy and entropy, we have estimated the rate constant for CH<sub>2</sub>OO +  
99 HONO reaction within a temperature range of 213–320 K.

100 The mechanism of CH<sub>2</sub>OO + HONO reaction can be represented by following reaction:



102 To calculate the overall rate constant of the title reaction, we have used the master equation approach as implemented in  
103 the MESMER software package. The reaction R1 proceeds in three steps. In the first step, the formation of RC occurs via  
104 a barrierless association of isolated reactants. MESMER uses the inverse-Laplace-transform (ILT) method to estimate the  
105 energy-dependent rate constant,  $k(E)$ , for this step. This, in turn, requires fitted Arrhenius parameters as input to MESMER,  
106 which are obtained using KTOOLS code as implemented in the MultiWell suite of programs (Barker et al., 2021). KTOOLS  
107 uses variational transition state theory (VTST) for the barrierless reaction. The inputs for KTOOLS are energies and frequen-  
108 cies calculated on potential energy surface (PES) scans along the coordinate describing the dissociation of RC into isolated  
109 reactants. Each point on the PES serves as a trial transition state; KTOOLS searches for the transition state for which the  
110 reaction flux is minimized. In the present work, we have obtained this PES scan at CCSD(T)/CBS//M062X/aug-cc-pVTZ level  
111 of theory (Table S9 of the ESI contains the energy as well as frequencies at each scan points). In the next step, RC under-  
112 goes unimolecular dissociation to PC via a transition state. MESMER uses Rice-Ramsperger-Kassel-Marcus (RRKM) theory,  
113 including tunneling contributions via an unsymmetrical Eckart barrier to compute the unimolecular reaction rate. In the final  
114 step, PC spontaneously dissociates to form isolated products. It is important to mention that we do not find any tight tran-  
115 sition state for product formation from PC; therefore, we have treated this step also using ILT assuming that rate constants  
116 are independent of temperature. The obtained rate constants within 213–320 K were then fitted with Arrhenius equation and  
117 supplied to the MESMER. It is worth noting that the reactant complex (RC) and the transition state (TS) exhibit hindered ro-  
118 tational motions, and multiple conformations may exist due to different torsional angles. To account for this, we have used the  
119 HinderedRotorQM1D model in MESMER to compute rate constants. Specifically, we performed a one-dimensional potential  
120 energy scan of OH torsion along the N–O bond in both RC and TS at CCSD(T)/CBS//M062X/aug-cc-pVTZ level of theory,  
121 that covers the full 0° to 360° range. The resulting energy profiles are used to calculate the hindered rotor partition functions.  
122 During this scan, we found local minima in both RC and TS, suggesting that our originally optimized structures correspond to  
123 the global minimum conformers. To verify this, we also manually searched for other possible minimum conformers and again  
124 found that our original structures are global minimum conformers. The Lennard-Jones (L-J) model is used to calculate the  
125 collision frequency between reactants and the bath gas. To obtain the L-J parameters for RC, we performed a PES scan along  
126 the reaction coordinate separating bath gas from RC, and fitted the obtained PES with the 12-6 L-J potential expression. **The**  
127 **fitted L-J parameters for RC turn out to be,  $\sigma = 3.1 \text{ \AA}$  and  $\epsilon = 895.5 \text{ K}$ .** A single-exponential down model is used to describe  
128 the collisional energy transfer probability with a maximum energy grain size of  $100 \text{ cm}^{-1}$  and  $\Delta E_{down} = 150 \text{ cm}^{-1}$ .

### 129 3 Results and discussion

130 In the present work, we have investigated the reactions of Criegee intermediates (CIs) with nitrous acid (HONO). It is known  
131 that the reactivity of CI is greatly influenced by the substitution group present on carbon center of the CI. Therefore, to account  
132 for it, we have studied two types of CIs; the simplest Criegee intermediate ( $\text{CH}_2\text{OO}$ ) and the dimethyl-substituted Criegee in-  
133 termediate ( $(\text{CH}_3)_2\text{COO}$ ). Another motivation for choosing  $(\text{CH}_3)_2\text{COO}$  comes from the fact that in contrast to simple Criegee  
134 which is formed only from the ozonolysis of ethene, the dimethyl-substituted Criegee intermediate can be generated from the  
135 ozonolysis of many highly abundant alkenes, such as terpenes and myrcene, and hence, the concentration of  $(\text{CH}_3)_2\text{COO}$  is  
136 significantly higher in the atmosphere. In this section, we will first discuss the energetics and kinetics of  $\text{CH}_2\text{OO} + \text{HONO}$   
137 reaction, followed by  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  reaction.

138 The potential energy surface for  $\text{CH}_2\text{OO} + \text{HONO}$  reaction is depicted in Figure 1. It is evident from Figure 1 that reaction  
139 occurs in three steps; in the first step,  $\text{CH}_2\text{OO}$  interacts with H atom of HONO via hydrogen bonding and forms a stable  
140 reactant-complex (RC1), which is  $\sim 10.1 \text{ kcal mol}^{-1}$  stable than isolated reactants. In the next step, RC1 undergoes a uni-  
141 molecular transformation to form product-complex (PC1) which has stabilization energy of  $\sim -44.7 \text{ kcal mol}^{-1}$  with respect to  
142 the isolated reactants. This happens via a transition-state (TS1) that is effectively  $\sim 8.0 \text{ kcal mol}^{-1}$  below the isolated reactants.  
143 In the last step, PC1 undergoes unimolecular dissociation to form final products, i.e.,  $\text{CH}_2\text{O}$ ,  $\text{OH}^\bullet$ , and  $\text{NO}_2$ . Gibbs free energy  
144 at 298 K associated with this conversion of PC1 to isolated products is  $\sim -2.5 \text{ kcal mol}^{-1}$  (shown in Figure S2 of ESI), which  
145 suggests that the formation of  $\text{OH}^\bullet$  via  $\text{CH}_2\text{OO} + \text{HONO}$  reaction is a spontaneous process. The overall reaction was found  
146 to be exothermic by  $\sim 17.3 \text{ kcal mol}^{-1}$  that lies close to the experimental value of  $\sim 16.9 \text{ kcal mol}^{-1}$  (Ruscic and Bross,  
147 2021), again confirming the adequacy of the methodology used. The computed bimolecular rate constant values ( $k_{bi}^{CH_2OO}$ ) for  
148  $\text{CH}_2\text{OO} + \text{HONO}$  reaction in the temperature range 213–320 K are given in Table 1. It is evident from Table 1 that the values  
149 of  $k_{bi}^{CH_2OO}$  slightly decrease with increasing temperature, a typical character of a barrierless process. For example, at 213 K,  
150 values of  $k_{bi}^{CH_2OO}$  is  $\sim 1.17 \times 10^{-11} \text{ cm}^3 \text{ molecule}^{-1} \text{ sec}^{-1}$  which becomes  $\sim 6.3 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ sec}^{-1}$  at 320 K.

151 Figure 2 depicts the potential energy surface of  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  reaction. It is evident from Figure 2 that  $(\text{CH}_3)_2\text{COO}$   
152 + HONO reaction also proceeds in three steps; in the first step,  $(\text{CH}_3)_2\text{COO}$  associates with HONO to form a stable reactant-  
153 complex (RC2) that is  $\sim 14.2 \text{ kcal mol}^{-1}$  more stable than isolated reactants. Next, RC2 transforms into product-complex  
154 (PC2) having stabilization energy of  $\sim -36.2 \text{ kcal mol}^{-1}$  with respect to isolated reactants. This transformation occurs through  
155 a transition state that lies  $\sim 10.1 \text{ kcal mol}^{-1}$  below the isolated reactants. At last, PC2 undergoes unimolecular dissociation to  
156 form final products, i.e.,  $(\text{CH}_3)_2\text{CO}$ ,  $\text{OH}^\bullet$ , and  $\text{NO}_2$ . Here also, the Gibbs free energy at 298 K associated with the conversion  
157 of PC2 to isolated products is  $\sim -6.3 \text{ kcal mol}^{-1}$  (Figure S2 of the ESI), making the overall product formation spontaneous.  
158 Using the energetics, we have also computed the rate constant for  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  reaction employing master equation  
159 in the same 213–320 K temperature range. The calculated bimolecular rate constants ( $k_{bi}^{(CH_3)_2COO}$ ) are listed in Table 1. It  
160 is evident from Table 1 that similar to  $\text{CH}_2\text{OO} + \text{HONO}$  reaction, here also the values of  $k_{bi}^{(CH_3)_2COO}$  slightly decrease with  
161 increasing temperature across the whole range of temperature. But the bimolecular rate constant of  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  re-  
162 action becomes  $\sim 2.6$  to  $3.6$  times higher compared to the same for  $\text{CH}_2\text{OO} + \text{HONO}$  reaction at all temperatures considered

163 in the present work. For example, at 298 K, the value of  $k_{bi}^{(CH_3)_2COO}$  is  $\sim 2.03 \times 10^{-11}$  cm<sup>3</sup> molecule<sup>-1</sup> sec<sup>-1</sup>, whereas the  
164 value of  $k_{bi}^{CH_2OO}$  is only  $\sim 7.2 \times 10^{-12}$  cm<sup>3</sup> molecule<sup>-1</sup> sec<sup>-1</sup>. It is worth noticing that, while computing the bimolecular rate  
165 constant, the capture rates of both the reactions are almost same (Table S3 of the ESI). The difference in the rate values of  
166 the two reactions depends on whether the reactant complex will proceed forward or backward, which further depends on the  
167 forward and backward Gibbs free energy barriers of the reactant complex. The Gibbs free energy profile at 298 K is shown  
168 in Figure S2 of the ESI. It is evident from Figure S2 that due to the higher stabilization of RC2 (corresponding to dimethyl-  
169 substituted CI), its reverse free energy barrier is high ( $\sim 2.9$  kcal mol<sup>-1</sup>), while the same is very low for RC1 (corresponding  
170 to simplest CI) ( $\sim -1.3$  kcal mol<sup>-1</sup>). Consequently, the relative yields of product are higher for the (CH<sub>3</sub>)<sub>2</sub>COO + HONO  
171 reaction compared to CH<sub>2</sub>COO + HONO reaction.

172 Lastly, it is important to discuss the uncertainties associated with the computed rate constant due to limitations in the method-  
173 ology (Fernández-Ramos et al., 2006). For example, a major source of uncertainty can originate from the fact that Criegee  
174 intermediates are known to possess moderate multireference character, and CCSD(T)/CBS sometimes fails in accurately pre-  
175 dicting the energetics of such reactions (Rai and Kumar, 2022; Mallick et al., 2019; Mallick and Kumar, 2018). It is worth  
176 mentioning that for multireference systems, incorporating higher-level excitations at the coupled-cluster level yield energet-  
177 ics within chemical accuracy (Tajti et al., 2004; Misiewicz et al., 2018; Nguyen et al., 2013; Anand and Kumar, 2023; Rai  
178 and Kumar, 2023). To assess the uncertainty in the energetics arising from the multireference character, we have carried out  
179 CCSDT(Q)/CBS calculations for the smaller Criegee intermediate reaction, i.e., CH<sub>2</sub>OO + HONO. We focused on key station-  
180 ary points; the reactant complex (RC) and the transition state (TS). The various components of the post-CCSD(T) corrections  
181 ( $\delta_T$  and  $\delta_{T(Q)}$ ) are provided in Table S7 of the ESI. It is evident from Table S7 that post-CCSD(T) corrections lead to only mi-  
182 nor changes in the calculated energetics of CH<sub>2</sub>OO + HONO reaction. Quantitatively, these corrections reduce the stabilization  
183 energy of RC by  $\sim 0.54$  kcal mol<sup>-1</sup>, while increasing the barrier height by a similar  $\sim 0.67$  kcal mol<sup>-1</sup>. Both variations fall  
184 well within the range of chemical accuracy. Furthermore, we have also estimated the capture and bimolecular rate constants  
185 using post-CCSD(T) energetics (see Table S8 of ESI), which suggest that at 298 K, the bimolecular rate constants calcu-  
186 lated at post-CCSD(T) and CCSD(T)/CBS levels are almost similar ( $5.53 \times 10^{-12}$  and  $7.21 \times 10^{-12}$  cm<sup>3</sup> molecule<sup>-1</sup> sec<sup>-1</sup>,  
187 respectively). This supports the reliability and computational efficiency of our chosen level of theory, CCSD(T)/CBS//M06-  
188 2X/aug-cc-pVTZ, for studying the title reaction. Another source of uncertainty in the computed rate constant may arise from  
189 the error in estimation of frequency. Such errors in frequency estimation may lead to  $2\sigma$  ( $\pm 2$  kcal mol<sup>-1</sup>) uncertainties in  
190 the computed barrier heights. To account for this, we have assumed an uncertainty of  $\pm 2$  kcal mol<sup>-1</sup> in both well depths and  
191 reaction barriers. Using this assumption, we estimated the resulting uncertainty in the rate constants at 213 K and 298 K for the  
192 model reaction CH<sub>2</sub>OO + HONO. Due to  $\pm 2$  kcal mol<sup>-1</sup> uncertainty in the reaction barriers and well depths, the deviation in  
193 the rate constant at 213 K is  $\sim 1.17_{-0.84}^{+1.8} \times 10^{-11}$  cm<sup>3</sup> molecule<sup>-1</sup> sec<sup>-1</sup> ( $\pm 2$  kcal mol<sup>-1</sup> reaction barriers) and  $\sim 1.17_{-0.08}^{+0.08} \times$   
194  $10^{-11}$  cm<sup>3</sup> molecule<sup>-1</sup> sec<sup>-1</sup> ( $\pm 2$  kcal mol<sup>-1</sup> well depths), respectively. At 298 K, the same becomes  $\sim 7.21_{-5.12}^{+11.04} \times 10^{-12}$   
195 cm<sup>3</sup> molecule<sup>-1</sup> sec<sup>-1</sup> ( $\pm 2$  kcal mol<sup>-1</sup> reaction barriers) and  $\sim 7.21_{-0.72}^{+0.72} \times 10^{-12}$  cm<sup>3</sup> molecule<sup>-1</sup> sec<sup>-1</sup> ( $\pm 2$  kcal mol<sup>-1</sup>  
196 well depths), respectively. This study suggests due to  $2\sigma$  error in the barrier height, there can be an error of a  $\sim$  factor-of-two  
197 in the estimated rate constant values. Our analysis also suggests that the uncertainty in the rate constant estimation is much

198 lower at low temperature region compare to high temperature regions. In addition, in the estimation of the partition function,  
199 the rigid rotor harmonic oscillator (RRHO) approximation is employed, which again can introduce some error in the final rate  
200 constant. For a typical  $2\sigma$  error, the uncertainty arising from the RRHO approximation can also contribute approximately a  
201 factor-of-two uncertainty in the evaluated partition function ratios.

## 202 4 Atmospheric implications

203 After estimating the energetics and kinetics of title reaction, it is important to discuss the impact of title reaction in the atmo-  
204 spheric chemistry. The importance of title reaction in the atmosphere critically depends on how it competes with other known  
205 sinks of Criegee intermediate, i.e.,  $\text{H}_2\text{O}$ ,  $(\text{H}_2\text{O})_2$ ,  $\text{NO}_2$ ,  $\text{NO}$ ,  $\text{CO}$ , and  $\text{SO}_2$ . The efficiency of a chemical reaction in the atmo-  
206 sphere depends upon two factors; rate of reaction and concentration of co-reactants. The effective rate constant ( $k_{eff}$ ) captures  
207 both of these factors as it is defined as the multiplication of bimolecular rate and concentration of co-reactants. Therefore,  
208 we have used  $k_{eff}$  to compare the effectiveness of title reaction compared to other sinks of Criegee intermediates. A list of  
209 effective rates for the reaction of CI with  $\text{H}_2\text{O}$ ,  $(\text{H}_2\text{O})_2$ ,  $\text{NO}_2$ ,  $\text{NO}$ ,  $\text{CO}$ , and  $\text{SO}_2$  at 298 K are provided in Table S4 of the  
210 ESI. To compute  $k_{eff}$ , the average concentrations of all the sinks have been taken from polluted urban environments. The  
211 corresponding rate coefficients of all the sinks are taken from experimental measurements. One can see from Table S4, the  
212 effective rate coefficients ( $k_{eff}$ ) of  $\text{CO}$ ,  $\text{NO}$ , and  $\text{NO}_2$  are lower compared to those of  $\text{SO}_2$ ,  $\text{H}_2\text{O}$ , and  $(\text{H}_2\text{O})_2$ . For example,  
213  $k_{eff}$  for the reaction of CI with  $\text{SO}_2$  is  $3.35 \text{ sec}^{-1}$ , while that for  $\text{NO}_2$  is only  $0.9 \text{ sec}^{-1}$ . Therefore, in the present work, we  
214 have focused our attention on a detailed comparison of the title reaction with  $\text{SO}_2$ ,  $\text{H}_2\text{O}$ , and  $(\text{H}_2\text{O})_2$ . As far as abundance of  
215 HONO is concerned, it is found in both regions; forested as well as polluted in significant amounts (Kim et al., 2015; Acker  
216 et al., 2006; Ren et al., 2010; Zhang et al., 2012; He et al., 2006; Su et al., 2008; Ren et al., 2006; Rondon and Sanhueza, 1989;  
217 Zhou et al., 2011; Pawar et al., 2024; Vereecken et al., 2012). Among the two, HONO concentrations are comparatively higher  
218 in polluted urban areas, such as megacities. Therefore, we expect HONO to play a more effective role as a sink for Criegee  
219 intermediates in such regions. Hence, we have used representative concentrations of HONO and  $\text{SO}_2$  in urban areas for the  
220 primary comparison. The concentration of water varies greatly in the atmosphere depending upon saturation vapour pressure  
221 and relative humidity (RH) (Anglada et al., 2013; Rai and Kumar, 2025). Therefore, in the case of  $\text{H}_2\text{O}$  and  $(\text{H}_2\text{O})_2$ , we have  
222 taken two concentrations; one calculated at 20% RH, and the other calculated at 100% RH. The former serves as lower limits  
223 of  $\text{H}_2\text{O}$  and  $(\text{H}_2\text{O})_2$  concentrations, whereas the latter serves as the upper limits of  $\text{H}_2\text{O}$  and  $(\text{H}_2\text{O})_2$  concentrations.  
224 For comparison, we have taken the rate constants reported by Lin et al. (Lin et al., 2016) for  $\text{H}_2\text{O}$  and  $(\text{H}_2\text{O})_2$ , and by Onel  
225 et al. (Onel et al., 2021) for  $\text{SO}_2$ . In Figure 3, we have compared the  $k_{eff}$  of  $\text{CH}_2\text{OO} + \text{HONO}$  with the  $k_{eff}$  of  $\text{CH}_2\text{OO}$   
226 +  $\text{H}_2\text{O}/(\text{H}_2\text{O})_2/\text{SO}_2$  reactions. Figure 3 shows that HONO is a minor sink of simplest Criegee intermediate ( $\text{CH}_2\text{OO}$ ) com-  
227 pare to  $\text{SO}_2$ ,  $\text{H}_2\text{O}$  and  $(\text{H}_2\text{O})_2$ . In fact, at 100% RH,  $k_{eff}$  of  $\text{CH}_2\text{OO} + (\text{H}_2\text{O})_2$  is the dominant reaction across the entire  
228 temperature range (213–320 K). At 20% RH,  $k_{eff}$  for  $\text{CH}_2\text{OO} + (\text{H}_2\text{O})_2$  and  $\text{CH}_2\text{OO} + \text{H}_2\text{O}$  remain dominant at higher  
229 temperatures, specifically within 235–320 K and 260–320 K, respectively. However, at lower temperatures,  $k_{eff}$  of  $\text{CH}_2\text{OO} +$   
230 HONO becomes dominant, surpassing both,  $\text{CH}_2\text{OO} + (\text{H}_2\text{O})_2$  and  $\text{CH}_2\text{OO} + \text{H}_2\text{O}$  in the range of 213–235 K and 213–260

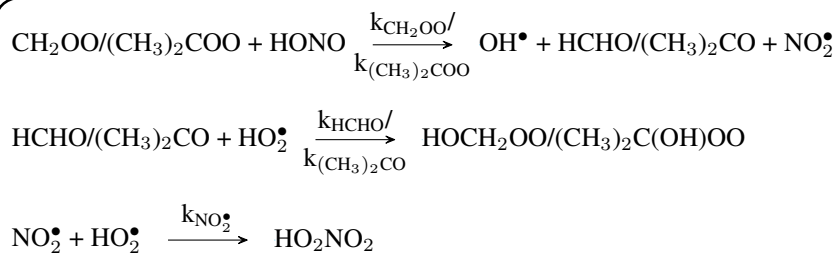
231 K, respectively. Although  $\text{CH}_2\text{OO} + \text{HONO}$  reaction dominates over  $\text{CH}_2\text{OO} + (\text{H}_2\text{O})_2$  and  $\text{CH}_2\text{OO} + \text{H}_2\text{O}$  at low tempera-  
232 ture and low humidity, it remains only a minor contributor compared to  $\text{CH}_2\text{OO} + \text{SO}_2$  reaction at the same conditions. For  
233 example,  $k_{eff}$  values of  $\text{CH}_2\text{OO} + \text{SO}_2$  reaction are  $\sim 5$  times higher than that of  $\text{CH}_2\text{OO} + \text{HONO}$  reaction within the whole  
234 temperature range, indicating that  $\text{CH}_2\text{OO} + \text{HONO}$  reaction is never a dominant sink of  $\text{CH}_2\text{OO}$  intermediate.

235 Similarly, we have compared our dimethyl substituted Criegee reaction ( $(\text{CH}_3)_2\text{COO} + \text{HONO}$ ) with other known bimolec-  
236 ular reactions of  $(\text{CH}_3)_2\text{COO}$ . Here also, we have computed  $k_{eff}$  for the comparison (see Figure 4). The rate constants of  
237  $(\text{CH}_3)_2\text{COO} + \text{SO}_2$  reaction (Smith et al., 2016) is known in the range of 283–303 K, and hence, we have compared its  $k_{eff}$   
238 in this temperature range with  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  reaction. Figure 4 shows that unlike  $\text{CH}_2\text{OO} + \text{HONO}$  reaction, here  
239  $k_{eff}$  of  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  is  $\sim 2$  times higher than the same for  $(\text{CH}_3)_2\text{COO} + \text{SO}_2$  reaction within 283–303 K. In ad-  
240 dition, it is worth mentioning that under certain atmospheric conditions, concentration of HONO can be quite high compared  
241 to  $\text{SO}_2$ . For example, during fog events, it is well known that concentration of  $\text{SO}_2$  drops significantly (Zhang et al., 2013)  
242 while concentration of HONO increases (Pawar et al., 2024), making HONO a potentially major bimolecular sink of Criegee  
243 intermediates in fog-like environments. In addition, as  $\text{SO}_2$  mainly comes from human activities, its concentrations are high  
244 in polluted areas and become quite very low in tropical forests and rural areas. In fact, its concentrations fall below detection  
245 limits in tropical forest regions (Vereecken et al., 2012). In contrast, although HONO concentration is high in polluted regions  
246 compared to a clean environment, due to the various *in situ* sources, HONO is present in reasonable amounts even in tropi-  
247 cal forest areas (Zhang et al., 2012). Therefore, in this region also, HONO is a more effective sink of CI compared to  $\text{SO}_2$ .  
248 Moreover, CI + HONO reaction is a hydrogen atom transfer (HAT) process, and hence, the presence of water can effectively  
249 catalyze this reaction (Buszek et al., 2012; Viegas and Varandas, 2012; Rai and Kumar, 2025). In contrast, the presence of  
250 water, particularly droplets and aerosols, can act as a sink for  $\text{SO}_2$  (Zhang et al., 2013), and hence, in the presence of wa-  
251 ter, Criegee +  $\text{SO}_2$  reaction should be less important compared to CI + HONO reaction. After establishing that compared to  
252  $\text{SO}_2$ , HONO is a more effective sink for  $(\text{CH}_3)_2\text{COO}$  under most of the conditions, at last, it is important to compare it with  
253  $(\text{CH}_3)_2\text{COO} + \text{H}_2\text{O}/(\text{H}_2\text{O})_2$  reactions (Vereecken et al., 2017). It can be seen from Figure 4 that even at 100% RH,  $k_{eff}$  of  
254  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  can dominate over  $k_{eff}$  of  $(\text{CH}_3)_2\text{COO} + \text{H}_2\text{O}$  and  $(\text{CH}_3)_2\text{COO} + (\text{H}_2\text{O})_2$  for a relatively wider range  
255 of temperatures. For example, the dominant temperature range of  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  is, 213–275 K for  $(\text{CH}_3)_2\text{COO} +$   
256  $(\text{H}_2\text{O})_2$  and 213–290 K for  $(\text{CH}_3)_2\text{COO} + \text{H}_2\text{O}$ . At 20% RH,  $k_{eff}$  of  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  becomes dominant over  $k_{eff}$  of  
257 both,  $(\text{CH}_3)_2\text{COO} + \text{H}_2\text{O}$  and  $(\text{CH}_3)_2\text{COO} + (\text{H}_2\text{O})_2$  in almost whole temperature range (213–310 K). For example, at 298 K,  
258  $k_{eff}$  of  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  is  $\sim 1.8 \text{ sec}^{-1}$ , which is 1.6 times and 2.2 times higher than the same for  $(\text{CH}_3)_2\text{COO} + \text{H}_2\text{O}$   
259 and  $(\text{CH}_3)_2\text{COO} + (\text{H}_2\text{O})_2$ , respectively. This suggests that the major bimolecular sink of substituted CI can be its reaction  
260 with HONO in the atmosphere even in the presence of high humidity and  $\text{SO}_2$ . At last, it is important to compare the  $k_{eff}$  of  
261  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  reaction with the unimolecular dissociation rate of  $(\text{CH}_3)_2\text{COO}$ . Figure 4 also contains the unimolec-  
262 ular dissociation rate of  $(\text{CH}_3)_2\text{COO}$ . It is evident from Figure 4 that unimolecular dissociation remains the dominant removal  
263 path of  $(\text{CH}_3)_2\text{COO}$  above 225 K temperature. Only below 225 K temperature, the bimolecular reaction of  $(\text{CH}_3)_2\text{COO} +$   
264 HONO becomes dominant. To conclude, although HONO is a dominant bimolecular sink for  $(\text{CH}_3)_2\text{COO}$ , it is still primarily  
265 removed by its unimolecular dissociation, particularly at room temperature. For example, the unimolecular dissociation rate of



266  $(\text{CH}_3)_2\text{COO}$  is  $\sim 276 \text{ sec}^{-1}$  at room temperature (Fang et al., 2017) whereas the  $k_{eff}$  of  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  is only  $\sim 1.8$   
267  $\text{sec}^{-1}$ . Interestingly, the unimolecular rate increases rapidly with temperature, while for the bimolecular reaction  $(\text{CH}_3)_2\text{COO}$   
268  $+ \text{HONO}$ ,  $k_{eff}$  decreases only slightly. As a result, at lower temperatures,  $k_{eff}$  may become comparable to the unimolecular  
269 dissociation rate of  $(\text{CH}_3)_2\text{COO}$ . For example, at 213 K,  $k_{eff}$  and the unimolecular rate constants are  $3.80 \text{ sec}^{-1}$  and  $1.82$   
270  $\text{sec}^{-1}$ , respectively. A comparison between  $k_{eff}$  and the unimolecular dissociation rate constant of  $(\text{CH}_3)_2\text{COO}$  within 213–  
271 320 K is provided in Table S6 of the ESI. It is evident from Table S6 that under conditions of high HONO concentration and  
272 low temperature, the bimolecular reaction of  $(\text{CH}_3)_2\text{COO}$  with HONO can compete with its unimolecular dissociation.  
273 Finally, it is important to assess the extent to which the title reaction can contribute in resolving the puzzle of mismatch be-  
274 tween measured and modelled  $\text{OH}^\bullet/\text{HO}_2^\bullet$  concentrations. It is important to mention that during daytime, HONO undergoes  
275 rapid photolysis; therefore, its concentration is higher in the absence of light, e.g. at night, indoors, in winter, etc. For example,  
276 the photolysis rate of HONO is known to be  $\sim 10^{-3} \text{ sec}^{-1}$ , which is several orders of magnitude higher than the effective rate  
277 constant of its reaction with Criegee intermediates ( $\sim 10^{-7} - 10^{-6} \text{ sec}^{-1}$ , computed using maximum Criegee concentration  
278 of  $\sim 10^5 \text{ molecule cm}^{-3}$ ) (Shabin et al., 2023). Therefore, during the peak of daytime, title reaction does not contribute much  
279 to  $\text{OH}^\bullet$  production; rather, it can play a key role in nocturnal atmospheric chemistry, specifically at times when both, concen-  
280 trations of HONO and CI are high, and, at the same time, the presence of light is minimal. To understand the efficiency of the  
281 title reaction in affecting  $\text{OH}^\bullet$  concentration in a nocturnal environment, we can compare it with  $\text{NO}_3^\bullet + \text{HO}_2^\bullet$  reaction, which  
282 is a well-known source of  $\text{OH}^\bullet$  at nighttime. The rate constants for  $\text{CH}_2\text{OO} + \text{HONO}$  reaction are  $\sim 2$  times higher compared  
283 to  $\text{NO}_3^\bullet + \text{HO}_2^\bullet$ . For example, at 298 K, the rate value for  $\text{CH}_2\text{OO} + \text{HONO}$  is  $\sim 7.21 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ sec}^{-1}$ , which is  
284 almost double compared to the rate value (Rai and Kumar, 2024) for  $\text{NO}_3^\bullet + \text{HO}_2^\bullet$ , i.e.,  $\sim 3.36 \times 10^{-12} \text{ cm}^3 \text{ molecule}^{-1} \text{ sec}^{-1}$ .  
285 In the atmosphere, average concentration of both  $\text{NO}_3^\bullet$  and  $\text{HO}_2^\bullet$  are  $\sim 10^8 \text{ molecule cm}^{-3}$  (Bottorff et al., 2023; Brown and  
286 Stutz, 2012), thus combined concentration turns out to be  $\sim 10^{16} \text{ molecule}^2 \text{ cm}^{-6}$ . Similarly, the combined concentration will  
287 be  $\sim 10^{15} \text{ molecule}^2 \text{ cm}^{-6}$  for  $\text{CH}_2\text{OO} + \text{HONO}$  under high concentrations of CI ( $\sim 10^5 \text{ molecule cm}^{-3}$ ) (Khan et al., 2018)  
288 and HONO ( $\sim 10^{10} \text{ molecule cm}^{-3}$ ) (Pawar et al., 2024). It suggests that  $\text{CH}_2\text{OO} + \text{HONO}$  reaction may be somewhat slower  
289 in producing  $\text{OH}^\bullet$ . However, since the rate of  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  reaction is one order of magnitude higher compared to  
290  $\text{NO}_3^\bullet + \text{HO}_2^\bullet$ , we believe both  $\text{NO}_3^\bullet + \text{HO}_2^\bullet$  and title reactions should be of similar importance as far as the production of night-  
291 time  $\text{OH}^\bullet$  is concerned. In other words, title reaction has the potential to serve as a significant contributor to  $\text{OH}^\bullet$  production  
292 in nighttime atmospheric chemistry.

293 Another factor worth noting is, besides  $\text{OH}^\bullet$ , the title reaction produces  $\text{HCHO}/(\text{CH}_3)_2\text{CO}$ , and  $\text{NO}_2^\bullet$  as products. It is well  
294 known that both  $\text{HCHO}/(\text{CH}_3)_2\text{CO}$  (Gao et al., 2024; Long et al., 2022; Hermans et al., 2004) and  $\text{NO}_2^\bullet$  (Christensen et al.,  
295 2004) can act as sinks for  $\text{HO}_2$  radicals (corresponding reactions are listed in the box below). It suggests that title reaction has  
296 the potential for recycling of  $\text{HO}_2^\bullet \leftrightarrow \text{OH}^\bullet$  process. To illustrate the ability of title reaction in recycling  $\text{HO}_2^\bullet \leftrightarrow \text{OH}^\bullet$  process,  
297 we have developed a kinetic model consisting of the following reactions (see ESI for the details):



298

299 This model requires two key components: first, the rate coefficients of the relevant reactions, which have been taken from  
 300 the recommended literature values (Gao et al., 2024; Hermans et al., 2004; Long et al., 2022; Christensen et al., 2004), and  
 301 second, a list of realistic initial concentrations of the reactive species involved in  $\text{HO}_2^\bullet \leftrightarrow \text{OH}^\bullet$  recycling process (Table S5  
 302 of the ESI). We first tracked the change in concentration of  $\text{OH}^\bullet$  and  $\text{HO}_2^\bullet$  using the first kinetic model consisting of  $\text{CH}_2\text{OO}$   
 303 + HONO reaction, followed by second model consisting of  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  reaction. Initial concentrations of relevant  
 304 species (HCHO, HONO,  $(\text{CH}_3)_2\text{CO}$ , and  $\text{HO}_2^\bullet$ ) are chosen based on literature values representing polluted urban conditions  
 305 (Vereecken et al., 2012; Pawar et al., 2024). Although the average concentration of  $\text{OH}^\bullet$  can vary within  $\sim 10^4$ – $10^6$  molecules  
 306  $\text{cm}^{-3}$  in the atmosphere, we have used a modelled value of it in the present work. In  $\text{CH}_2\text{OO} + \text{HONO}$  reaction model,  
 307 the initial  $\text{OH}^\bullet$  concentration was set to  $\sim 10^4$  molecules  $\text{cm}^{-3}$ , while in  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  model, it was set to  $\sim 10^5$   
 308 molecules  $\text{cm}^{-3}$ . This difference was chosen based on how much OH each reaction is expected to produce when no *in situ*  
 309 reactions are taking place from the byproducts of the title reaction. Since  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  reaction can generate more  
 310 OH, starting with a higher initial concentration helps one observe a noticeable change in  $\text{OH}^\bullet$  levels during the simulation.  
 311 This makes it easier to observe and compare the effect of  $\text{OH}^\bullet$  production between the two reactions. It is important to mention  
 312 that the maximum concentration of  $\text{OH}^\bullet$  can be taken as  $\sim 10^5$  molecules  $\text{cm}^{-3}$  in the kinetic model. This is because the  
 313 production of  $\text{OH}^\bullet$  is limited by the available concentration of CI which can be as high as  $\sim 10^5$  molecules  $\text{cm}^{-3}$ . Therefore,  
 314 taking  $\text{OH}^\bullet$  concentration more than  $\sim 10^5$  molecules  $\text{cm}^{-3}$  would produce no effect on the concentration of  $\text{OH}^\bullet$ . This also  
 315 reveals the fact that the title reaction is capable of producing  $\text{OH}^\bullet$  in regions where the concentration of  $\text{OH}^\bullet$  is already low.  
 316 Similarly, the concentration of  $\text{NO}_2$  can vary within  $\sim 10^{10}$ – $10^{12}$  molecules  $\text{cm}^{-3}$  in polluted urban regions. However, in the  
 317 present model, we have kept it at  $\sim 10^{10}$  molecules  $\text{cm}^{-3}$  in order to observe a clear numerical change in the values of  $\text{HO}_2^\bullet$ .  
 318 Taking a high concentration of  $\text{NO}_2$  ( $\sim 10^{12}$  molecules  $\text{cm}^{-3}$ ) would drastically consume  $\text{HO}_2^\bullet$ , and a gradual change would  
 319 not be observed.

320 We have divided the simulation results into two parts; first we will discuss  $\text{CH}_2\text{OO} + \text{HONO}$  reaction followed by  $(\text{CH}_3)_2\text{COO}$   
 321 + HONO. The model results have been shown in Figure 5. It is evident from Figure 5 that  $\text{CH}_2\text{OO} + \text{HONO}$  reaction increases  
 322  $\text{OH}^\bullet$  concentration while simultaneously reducing  $\text{HO}_2^\bullet$  concentration. Quantitatively, this reaction increases  $\text{OH}^\bullet$  production  
 323 by five times its initial value while decreasing  $\text{HO}_2^\bullet$  production by more than one order of magnitude. Furthermore, when  
 324 we consider dimethyl-substituted Criegee intermediate reaction ( $(\text{CH}_3)_2\text{COO} + \text{HONO}$ ),  $\text{OH}^\bullet$  production has been found to  
 325 increase by only a factor of two compared to its initial concentration, while  $\text{HO}_2^\bullet$  production again decreases by the same one  
 326 order of magnitude (Figure 5). The difference in  $\text{OH}^\bullet$  production can be attributed to the fact that, in case of  $(\text{CH}_3)_2\text{COO} +$

327 HONO, the initial OH• concentration was taken to be  $\sim 10^5$  molecules  $\text{cm}^{-3}$  compared to  $\sim 10^4$  molecules  $\text{cm}^{-3}$  in case of  
328  $\text{CH}_2\text{OO} + \text{HONO}$ . This further strengthens the fact that the effect of title reaction on OH• production will be more pronounced  
329 in the conditions where OH• concentration is lower in the atmosphere, e.g., at night. The overall simulation results suggest that  
330 incorporating title reaction into atmospheric models can improve their accuracy in predicting OH• and  $\text{HO}_2$  concentrations.  
331 It is important to note that the kinetics model used in the present work is preliminary. However, a more realistic impact of the  
332 title reaction on the budget of both OH• and  $\text{HO}_2$ , requires a more complete modeling. In order to do so, one needs accurate  
333 estimation of the rate constants for the reaction of HONO with various important Criegee intermediates. For bigger Criegee  
334 intermediates, computation will be more costly and require a separate study. In addition, being a HAT reaction, the effect of  
335 humidity on the title reaction is also important to build a complete model.

## 336 5 Conclusions

337 In this work, we have studied the energetics and kinetics of bimolecular reaction of simple and dimethyl-substituted Criegee  
338 with HONO using high-level electronic structure theory and chemical kinetics. Our quantum chemical calculations suggest that  
339 both of the reactions are barrierless and kinetic calculations reveal that reaction of substituted Criegee with HONO is  $\sim 2.6$ – $3.6$   
340 times faster than simple Criegee + HONO reaction. By comparing it with other known sinks of CI, we have shown that HONO  
341 can serve as a major bimolecular sink for bigger Criegee intermediate ( $(\text{CH}_3)_2\text{COO}$ ) and minor contributor at low humidity and  
342 low temperature for simple  $\text{CH}_2\text{OO}$ . In addition, we have also shown that title reaction can be an important source of OH• in  
343 nocturnal atmosphere. In addition, the products of CI + HONO reaction can be a sink for  $\text{HO}_2$  radicals, and hence this reaction  
344 is capable of  $\text{HO}_2 \leftrightarrow \text{OH}^\bullet$  recycling. Consequently, this reaction can be key in fulfilling the gap between the observed OH  
345 radicals and modelled values. Although in urban areas, HONO can be the dominant sink of certain CIs, it is important to notice  
346 that larger Criegee intermediates predominantly originate from biogenic volatile organic compounds (BVOCs). On the other  
347 hand, HONO concentrations in forested regions are also found to be moderate ( $\sim 10^8$  to  $10^{10}$  molecules  $\text{cm}^{-3}$ ). Therefore, we  
348 believe a separate study is required to understand the fate of larger Criegee intermediates in the presence of HONO. At last, we  
349 look forward to the experimental verification of our results.

350 *Author contributions.* VJA: Conducted the investigation, Writing–original draft, Formal analysis, curated the data. PKR: Contributed to partial  
351 formal analysis, writing, reviewing, and editing the manuscript. PK: Provided supervision, resources, and methodology; conceptualized  
352 the study; acquired funding; and contributed to the review and editing of the manuscript.

353 *Competing interests.* The authors declare that they have no conflict of interest.

354 *Acknowledgements.* V.J.A. and P.K.R acknowledge MNIT Jaipur for financial assistance. P.K. acknowledges DST, Govt. of India, for the  
355 financial support through the sanctioned project No. EEQ/2023/000351.

## 356 References

- 357 Acker, K., Möller, D., Wieprecht, W., Meixner, F. X., Bohn, B., Gilge, S., Plass-Dülmer, C., and Berresheim, H.: Strong daytime production  
358 of OH from HNO<sub>2</sub> at a rural mountain site, *Geophys. Res. Lett.*, 33, 2006.
- 359 Alam, M. S., Camredon, M., Rickard, A. R., Carr, T., Wyche, K. P., Hornsby, K. E., Monks, P. S., and Bloss, W. J.: Total radical yields from  
360 tropospheric ethene ozonolysis, *Phys. Chem. Chem. Phys.*, 13, 11 002–11 015, 2011.
- 361 Aliche, B., Geyer, A., Hofzumahaus, A., Holland, F., Konrad, S., Pätz, H., Schäfer, J., Stutz, J., Volz-Thomas, A., and Platt, U.: OH formation  
362 by HONO photolysis during the BERLIOZ experiment, *J. Geophys. Res.*, 108, PHO–3, 2003.
- 363 Anand, V. J. and Kumar, P.: Mechanistic insight into the N<sub>2</sub>O + O(<sup>1</sup>D, <sup>3</sup>P) reaction: role of post-CCSD (T) corrections and non-adiabatic  
364 effects, *Phys. Chem. Chem. Phys.*, 25, 33 119–33 129, 2023.
- 365 Anderson, J. G.: Free Radicals in the Earth's Atmosphere: Their Measurement and Interpretation, *Annu. Rev. Phys. Chem.*, 38, 489–520,  
366 1987.
- 367 Anglada, J. M. and Sole, A.: The atmospheric oxidation of HONO by OH, Cl, and ClO radicals, *J. Phys. Chem. A*, 121, 9698–9707, 2017.
- 368 Anglada, J. M., Hoffman, G. J., Slipchenko, L. V., M. Costa, M., Ruiz-Lopez, M. F., and Francisco, J. S.: Atmospheric significance of water  
369 clusters and ozone–water complexes, *J. Phys. Chem. A*, 117, 10 381–10 396, 2013.
- 370 Aumont, B., Chervier, F., and Laval, S.: Contribution of HONO sources to the NO<sub>x</sub>/HO<sub>x</sub>/O<sub>3</sub> chemistry in the polluted boundary layer,  
371 *Atmos. Environ.*, 37, 487–498, 2003.
- 372 Barker, J., Nguyen, T., Stanton, J., Aieta, C., Ceotto, M., Gabas, F., Kumar, T., Li, C., Lohr, L., Maranzana, A., et al.: MultiWell-2021 Software  
373 Suite; J. R. Barker, University of Michigan, Ann Arbor, Michigan, USA, <http://claspresearch.engin.umich.edu/multiwell/> (accessed march  
374 5, 2025), 2021.
- 375 Berndt, T., Hyttinen, N., Herrmann, H., and Hansel, A.: First oxidation products from the reaction of hydroxyl radicals with isoprene for  
376 pristine environmental conditions, *Commun. Chem.*, 2, 21, 2019.
- 377 Bottorff, B., Lew, M. M., Woo, Y., Rickly, P., Rollings, M. D., Deming, B., Anderson, D. C., Wood, E., Alwe, H. D., Millet, D. B., et al.: OH,  
378 HO<sub>2</sub>, and RO<sub>2</sub> radical chemistry in a rural forest environment: measurements, model comparisons, and evidence of a missing radical  
379 sink, *Atmos. Chem. Phys.*, 23, 10 287–10 311, 2023.
- 380 Brown, S. S. and Stutz, J.: Nighttime radical observations and chemistry, *Chem. Soc. Rev.*, 41, 6405–6447, 2012.
- 381 Buszek, R. J., Barker, J. R., and Francisco, J. S.: Water effect on the OH + HCl reaction, *J. Phys. Chem. A*, 116, 4712–4719, 2012.
- 382 Calvert, J., Yarwood, G., and Dunker, A.: An evaluation of the mechanism of nitrous acid formation in the urban atmosphere, *Res. Chem.*  
383 *Intermed.*, 20, 463–502, 1994.
- 384 Carslaw, N., Creasey, D., Harrison, D., Heard, D., Hunter, M., Jacobs, P., Jenkin, M., Lee, J., Lewis, A., Pilling, M., et al.: OH and HO<sub>2</sub>  
385 radical chemistry in a forested region of north-western Greece, *Atmos. Environ.*, 35, 4725–4737, 2001.
- 386 Christensen, L. E., Okumura, M., Sander, S. P., Friedl, R. R., Miller, C. E., and Sloan, J. J.: Measurements of the Rate Constant of HO<sub>2</sub> +  
387 NO<sub>2</sub> + N<sub>2</sub> → HO<sub>2</sub>NO<sub>2</sub> + N<sub>2</sub> Using Near-Infrared Wavelength-Modulation Spectroscopy and UV- Visible Absorption Spectroscopy, *J.*  
388 *Phys. Chem. A*, 108, 80–91, 2004.
- 389 Cox, R. A., Ammann, M., Crowley, J. N., Herrmann, H., Jenkin, M. E., McNeill, V. F., Mellouki, A., Troe, J., and Wallington, T. J.: Evaluated  
390 kinetic and photochemical data for atmospheric chemistry: Volume VII–Criegee intermediates, *Atmos. Chem. Phys.*, 20, 13 497–13 519,  
391 2020.
- 392 Criegee, R.: Mechanism of ozonolysis, *Angew. Chem. internat. Edit.*, 14, 745–752, 1975.

393 Crouse, J. D., Paulot, F., Kjaergaard, H. G., and Wennberg, P. O.: Peroxy radical isomerization in the oxidation of isoprene, *Phys. Chem.*  
394 *Chem. Phys.*, 13, 13 607–13 613, 2011.

395 Donahue, N. M., Drozd, G. T., Epstein, S. A., Presto, A. A., and Kroll, J. H.: Adventures in ozoneland: down the rabbit-hole, *Phys. Chem.*  
396 *Chem. Phys.*, 13, 10 848–10 857, 2011.

397 Ehhalt, D.: Free Radicals in the Atmosphere, *Free Radic. Res. Commun.*, 3, 153–164, 1987.

398 Emmerson, K. and Carslaw, N.: Night-time radical chemistry during the TORCH campaign, *Atmos. Environ.*, 43, 3220–3226, 2009.

399 Faloona, I., Tan, D., Brune, W., Hurst, J., Barket Jr, D., Couch, T. L., Shepson, P., Apel, E., Riemer, D., Thornberry, T., et al.: Nighttime  
400 observations of anomalously high levels of hydroxyl radicals above a deciduous forest canopy, *J. Geophys. Res. Atmos.*, 106, 24 315–  
401 24 333, 2001.

402 Fang, Y., Barber, V. P., Klippenstein, S. J., McCoy, A. B., and Lester, M. I.: Tunneling Effects in the Unimolecular Decay of (CH<sub>3</sub>)<sub>2</sub>COO  
403 Criegee Intermediates to OH Radical Products, *J. Chem. Phys.*, 146, 134 307, 2017.

404 Feiner, P. A., Brune, W. H., Miller, D. O., Zhang, L., Cohen, R. C., Romer, P. S., Goldstein, A. H., Keutsch, F. N., Skog, K. M., Wennberg,  
405 P. O., et al.: Testing atmospheric oxidation in an Alabama forest, *J. Atmos. Sci.*, 73, 4699–4710, 2016.

406 Fernández-Ramos, A., Miller, J. A., Klippenstein, S. J., and Truhlar, D. G.: Modeling the kinetics of bimolecular reactions, *Chem. Rev.*, 106,  
407 4518–4584, 2006.

408 Frisch, M. J., Trucks, G. W., Schlegel, H. B., Scuseria, G. E., Robb, M. A., Cheeseman, J. R., Scalmani, G., Barone, V., Petersson, G. A.,  
409 Nakatsuji, H., Li, X., Caricato, M., Marenich, A. V., Bloino, J., Janesko, B. G., Gomperts, R., Mennucci, B., Hratchian, H. P., Ortiz, J. V.,  
410 Izmaylov, A. F., Sonnenberg, J. L., Williams-Young, D., Ding, F., Lipparini, F., Egidi, F., Goings, J., Peng, B., Petrone, A., Henderson,  
411 T., Ranasinghe, D., Zakrzewski, V. G., Gao, J., Rega, N., Zheng, G., Liang, W., Hada, M., Ehara, M., Toyota, K., Fukuda, R., Hasegawa,  
412 J., Ishida, M., Nakajima, T., Honda, Y., Kitao, O., Nakai, H., Vreven, T., Throssell, K., Montgomery, Jr., J. A., Peralta, J. E., Ogliaro, F.,  
413 Bearpark, M. J., Heyd, J. J., Brothers, E. N., Kudin, K. N., Staroverov, V. N., Keith, T. A., Kobayashi, R., Normand, J., Raghavachari,  
414 K., Rendell, A. P., Burant, J. C., Iyengar, S. S., Tomasi, J., Cossi, M., Millam, J. M., Klene, M., Adamo, C., Cammi, R., Ochterski, J. W.,  
415 Martin, R. L., Morokuma, K., Farkas, O., Foresman, J. B., and Fox, D. J.: Gaussian 16 Revision C.01, gaussian Inc. Wallingford CT,  
416 2016.

417 Gao, Q., Shen, C., Zhang, H., Long, B., and Truhlar, D. G.: Quantitative kinetics reveal that reactions of HO<sub>2</sub> are a significant sink for  
418 aldehydes in the atmosphere and may initiate the formation of highly oxygenated molecules via autoxidation, *Phys. Chem. Chem. Phys.*,  
419 26, 16 160–16 174, 2024.

420 Geyer, A., Bächmann, K., Hofzumahaus, A., Holland, F., Konrad, S., Klüpfel, T., Pätz, H.-W., Perner, D., Mihelcic, D., Schäfer, H.-J., et al.:  
421 Nighttime formation of peroxy and hydroxyl radicals during the BERLIOZ campaign: Observations and modeling studies, *J. Geophys.*  
422 *Res. Atmos.*, 108, 2003.

423 Gomez Alvarez, E., Amedro, D., Afif, C., Gligorovski, S., Schoemaeker, C., Fittschen, C., Doussin, J.-F., and Wortham, H.: Unexpectedly  
424 high indoor hydroxyl radical concentrations associated with nitrous acid, *Proc. Natl. Acad. Sci.*, 110, 13 294–13 299, 2013.

425 Griffith, S. M., Hansen, R., Dusanter, S., Michoud, V., Gilman, J., Kuster, W., Veres, P., Graus, M., De Gouw, J., Roberts, J., et al.: Mea-  
426 surements of hydroxyl and hydroperoxy radicals during CalNex-LA: Model comparisons and radical budgets, *J. Geophys. Res.*, 121,  
427 4211–4232, 2016.

428 Hall, I. W., Wayne, R. P., Cox, R. A., Jenkin, M. E., and Hayman, G. D.: Kinetics of the reaction of nitrate radical with hydroperoxo, *J. Phys.*  
429 *Chem.*, 92, 5049–5054, 1988.

430 Harrison, R., Yin, J., Tilling, R., Cai, X., Seakins, P., Hopkins, J., Lansley, D., Lewis, A., Hunter, M., Heard, D., et al.: Measurement and  
431 modelling of air pollution and atmospheric chemistry in the UK West Midlands conurbation: Overview of the PUMA Consortium project,  
432 *Sci. Total Environ.*, 360, 5–25, 2006.

433 He, Y., Zhou, X., Hou, J., Gao, H., and Bertman, S. B.: Importance of dew in controlling the air-surface exchange of HONO in rural forested  
434 environments, *Geophys. Res. Lett.*, 33, 2006.

435 Heald, C. L. and Kroll, J. H.: A radical shift in air pollution, *Science*, 374, 688–689, 2021.

436 Heard, D., Carpenter, L., Creasey, D., Hopkins, J., Lee, J., Lewis, A., Pilling, M., Seakins, P., Carslaw, N., and Emmerson, K.: High levels of  
437 the hydroxyl radical in the winter urban troposphere, *Geophys. Res. Lett.*, 31, 2004.

438 Hens, K., Novelli, A., Martinez, M., Auld, J., Axinte, R., Bohn, B., Fischer, H., Keronen, P., Kubistin, D., Nölscher, A., et al.: Observation  
439 and modelling of HO<sub>x</sub> radicals in a boreal forest., *Atmos. Chem. Phys. Discuss.*, 13, 2013.

440 Hermans, I., Nguyen, T. L., Jacobs, P. A., and Peeters, J.: Tropopause chemistry revisited: HO<sub>2</sub><sup>\*</sup>-initiated oxidation as an efficient acetone  
441 sink, *J. Am. Chem. Soc.*, 126, 9908–9909, 2004.

442 Hofzumahaus, A., Rohrer, F., Lu, K., Bohn, B., Brauers, T., Chang, C.-C., Fuchs, H., Holland, F., Kita, K., Kondo, Y., et al.: Amplified trace  
443 gas removal in the troposphere, *science*, 324, 1702–1704, 2009.

444 Horie, O. and Moortgat, G.: Decomposition pathways of the excited Criegee intermediates in the ozonolysis of simple alkenes, *Atmos.*  
445 *Environ.*, 25A, 1881–1896, 1991.

446 J. Medeiros, D., Blitz, M. A., Seakins, P. W., and Whalley, L. K.: Direct measurements of isoprene autoxidation: Pinpointing atmospheric  
447 oxidation in tropical forests, *JACS Au*, 2, 809–818, 2022.

448 Johnson, D. and Marston, G.: The gas-phase ozonolysis of unsaturated volatile organic compounds in the troposphere, *Chem. Soc. Rev.*, 37,  
449 699–716, 2008.

450 Khan, M., Percival, C., Caravan, R., Taatjes, C., and Shallcross, D.: Criegee intermediates and their impacts on the troposphere, *Environ.*  
451 *Sci.: Process. Impacts*, 20, 437–453, 2018.

452 Kim, S., Kim, S.-Y., Lee, M., Shim, H., Wolfe, G., Guenther, A. B., He, A., Hong, Y., and Han, J.: Impact of isoprene and HONO chemistry  
453 on ozone and OVOC formation in a semirural South Korean forest, *Atmos. Chem. Phys.*, 15, 4357–4371, 2015.

454 Kumar, A., Mallick, S., and Kumar, P.: Nitrous acid (HONO) as a sink of the simplest Criegee intermediate in the atmosphere, *Phys. Chem.*  
455 *Chem. Phys.*, 24, 7458–7465, 2022.

456 Lelieveld, J., Peters, W., Dentener, F., and Krol, M.: Stability of tropospheric hydroxyl chemistry, *J. Geophys. Res.*, 107, ACH-17, 2002.

457 Lelieveld, J., Dentener, F., Peters, W., and Krol, M.: On the role of hydroxyl radicals in the self-cleansing capacity of the troposphere, *Atmos.*  
458 *Chem. Phys.*, 4, 2337–2344, 2004.

459 Lelieveld, J., Gromov, S., Pozzer, A., and Taraborrelli, D.: Global tropospheric hydroxyl distribution, budget and reactivity, *Atmos. Chem.*  
460 *Phys.*, 16, 12 477–12 493, 2016.

461 Lelieveld, J. a., Butler, T. M., Crowley, J. N., Dillon, T. J., Fischer, H., Ganzeveld, L., Harder, H., Lawrence, M. G., Martinez, M., Taraborrelli,  
462 D., et al.: Atmospheric oxidation capacity sustained by a tropical forest, *Nature*, 452, 737–740, 2008.

463 Lew, M. M., Rickly, P. S., Bottorff, B. P., Reidy, E., Sklaveniti, S., Léonardis, T., Locoge, N., Dusanter, S., Kundu, S., Wood, E., et al.: OH  
464 and HO<sub>2</sub> radical chemistry in a midlatitude forest: measurements and model comparisons, *Atmos. Chem. Phys.*, 20, 9209–9230, 2020.

465 Li, Y., Wang, X., Wu, Z., Li, L., Wang, C., Li, H., Zhang, X., Zhang, Y., Li, J., Gao, R., et al.: Atmospheric nitrous acid (HONO) in an alternate  
466 process of haze pollution and ozone pollution in urban Beijing in summertime: Variations, sources and contribution to atmospheric  
467 photochemistry, *Atmos. Res.*, 260, 105 689, 2021.

468 Lin, H.-Y., Huang, Y.-H., Wang, X., Bowman, J. M., Nishimura, Y., Witek, H. A., and Lee, Y.-P.: Infrared identification of the Criegee  
469 intermediates syn- and anti-CH<sub>3</sub>CHOO, and their distinct conformation-dependent reactivity, *Nat. Commun.*, 6, 7012, 2015.

470 Lin, L.-C., Chang, H.-T., Chang, C.-H., Chao, W., Smith, M. C., Chang, C.-H., Takahashi, K., et al.: Competition between H<sub>2</sub>O and (H<sub>2</sub>O)<sub>2</sub>  
471 reactions with CH<sub>2</sub>OO/CH<sub>3</sub>CHOO, *Phys. Chem. Chem. Phys.*, 18, 4557–4568, 2016.

472 Long, B., Bao, J. L., and Truhlar, D. G.: Atmospheric Chemistry of Criegee Intermediates: Unimolecular Reactions and Reactions with  
473 Water, *J. Am. Chem. Soc.*, 138, 14409–14422, 2016.

474 Long, B., Wang, Y., Xia, Y., He, X., Bao, J. L., and Truhlar, D. G.: Atmospheric Kinetics: Bimolecular Reactions of Carbonyl Oxide by a  
475 Triple-Level Strategy, *J. Am. Chem. Soc.*, 143, 8402–8413, 2021.

476 Long, B., Xia, Y., and Truhlar, D. G.: Quantitative kinetics of HO<sub>2</sub> reactions with aldehydes in the atmosphere: High-order dynamic corre-  
477 lation, anharmonicity, and falloff effects are all important, *J. Am. Chem. Soc.*, 144, 19910–19920, 2022.

478 Lu, K., Rohrer, F., Holland, F., Fuchs, H., Bohn, B., Brauers, T., Chang, C., Häseler, R., Hu, M., Kita, K., et al.: Observation and modelling  
479 of OH and HO<sub>2</sub> concentrations in the Pearl River Delta 2006: a missing OH source in a VOC rich atmosphere, *Atmos. Chem. Phys.*, 12,  
480 1541–1569, 2012.

481 Lu, X., Park, J., and Lin, M.-C.: Gas phase reactions of HONO with NO<sub>2</sub>, O<sub>3</sub>, and HCl: Ab initio and TST study, *J. Phys. Chem. A*, 104,  
482 8730–8738, 2000.

483 Mallick, S. and Kumar, P.: Impact of Post-CCSD(T) Corrections on Reaction Energetics and Rate Constants of the OH<sup>•</sup> + HCl Reaction, *J.*  
484 *Phys. Chem. A*, 122, 7151–7159, 2018.

485 Mallick, S. and Kumar, P.: The reaction of N<sub>2</sub>O with the Criegee intermediate: A theoretical study, *Comput. Theor. Chem.*, 1191, 113 023,  
486 2020.

487 Mallick, S., Kumar, A., and Kumar, P.: Revisiting the reaction energetics of the CH<sub>3</sub>O<sup>•</sup> + O<sub>2</sub> (<sup>3</sup>Σ<sup>-</sup>) reaction: the crucial role of post-  
488 CCSD(T) corrections, *Phys. Chem. Chem. Phys.*, 21, 6559–6565, 2019.

489 Mellouki, A., Le Bras, G., and Poulet, G.: Kinetics of the reactions of nitrate radical with hydroxyl and hydroperoxo, *J. Phys. Chem. A*, 92,  
490 2229–2234, 1988.

491 Mellouki, A., Talukdar, R., Bopeggedera, A., and Howard, C. J.: Study of the kinetics of the reactions of NO<sub>3</sub> with HO<sub>2</sub> and OH, *Int. J. Chem.*  
492 *Kinet.*, 25, 25–39, 1993.

493 Misiewicz, J. P., Elliott, S. N., Moore, K. B., and Schaefer, H. F.: Re-examining ammonia addition to the Criegee intermediate: converging  
494 to chemical accuracy, *Phys. Chem. Chem. Phys.*, 20, 7479–7491, 2018.

495 Monks, P. S.: Gas-phase radical chemistry in the troposphere, *Chem. Soc. Rev.*, 34, 376–395, 2005.

496 Nguyen, T. L., Li, J., Dawes, R., Stanton, J. F., and Guo, H.: Accurate determination of barrier height and kinetics for the F + H<sub>2</sub> → HF +  
497 OH reaction, *J. Phys. Chem. A*, 117, 8864–8872, 2013.

498 Novelli, A., Vereecken, L., Lelieveld, J., and Harder, H.: Direct observation of OH formation from stabilised Criegee intermediates, *Phys.*  
499 *Chem. Chem. Phys.*, 16, 19941–19951, 2014.

500 Novelli, A., Hens, K., Tatum Ernest, C., Martinez, M., Nölscher, A. C., Sinha, V., Paasonen, P., Petäjä, T., Sipilä, M., Elste, T., et al.:  
501 Estimating the atmospheric concentration of Criegee intermediates and their possible interference in a FAGE-LIF instrument, *Atmos.*  
502 *Chem. Phys.*, 17, 7807–7826, 2017.

503 Novelli, A., Vereecken, L., Bohn, B., Dorn, H.-P., Gkatzelis, G. I., Hofzumahaus, A., Holland, F., Reimer, D., Rohrer, F., Rosanka, S., et al.:  
504 Importance of isomerization reactions for OH radical regeneration from the photo-oxidation of isoprene investigated in the atmospheric  
505 simulation chamber SAPHIR, *Atmos. Chem. Phys.*, 20, 3333–3355, 2020.



506 Onel, L., Lade, R., Mortiboy, J., Blitz, M. A., Seakins, P. W., Heard, D. E., and Stone, D.: Kinetics of the gas phase reaction of the Criegee  
507 intermediate  $\text{CH}_2\text{OO}$  with  $\text{SO}_2$  as a function of temperature, *Phys. Chem. Chem. Phys.*, 23, 19 415–19 423, 2021.

508 Osborn, D. L. and Taatjes, C. A.: The physical chemistry of Criegee intermediates in the gas phase, *Int. Rev. Phys. Chem.*, 34, 309–360,  
509 2015.

510 Østerstrøm, F. F., Carter, T. J., Shaw, D. R., Abbatt, J. P., Abeleira, A., Arata, C., Bottorff, B. P., Cardoso-Saldaña, F. J., DeCarlo, P. F.,  
511 Farmer, D. K., et al.: Modelling indoor radical chemistry during the HOMEChem campaign, *Environ. Sci.: Process. Impacts*, 2025.

512 Pansini, F., Neto, A., and Varandas, A.: Extrapolation of Hartree–Fock and multiconfiguration self-consistent-field energies to the complete  
513 basis set limit, *Theor. Chem. Acc.*, 135, 1–6, 2016.

514 Paulot, F., Crounse, J. D., Kjaergaard, H. G., Kürten, A., St. Clair, J. M., Seinfeld, J. H., and Wennberg, P. O.: Unexpected epoxide formation  
515 in the gas-phase photooxidation of isoprene, *science*, 325, 730–733, 2009.

516 Pawar, P. V., Mahajan, A. S., and Ghude, S. D.: HONO chemistry and its impact on the atmospheric oxidizing capacity over the Indo-Gangetic  
517 Plain, *Sci. Total Environ.*, p. 174604, 2024.

518 Peeters, J. and Müller, J.-F.:  $\text{HO}_X$  radical regeneration in isoprene oxidation via peroxy radical isomerisations. II: experimental evidence  
519 and global impact, *Phys. Chem. Chem. Phys.*, 12, 14 227–14 235, 2010.

520 Peeters, J., Nguyen, T. L., and Vereecken, L.:  $\text{HO}_X$  radical regeneration in the oxidation of isoprene, *Phys. Chem. Chem. Phys.*, 11, 5935–  
521 5939, 2009.

522 Peeters, J., Müller, J.-F., Stavrou, T., and Nguyen, V. S.: Hydroxyl radical recycling in isoprene oxidation driven by hydrogen bonding and  
523 hydrogen tunneling: The upgraded LIM1 mechanism, *J. Phys. Chem. A*, 118, 8625–8643, 2014.

524 Prinn, R. G.: The Cleansing Capacity of the Atmosphere, *Annu. Rev. Environ. Resour.*, 28, 29–57, 2003.

525 Rai, P. K. and Kumar, P.: Role of post-CCSD (T) corrections in predicting the energetics and kinetics of the  $\text{OH}^\bullet + \text{O}_3$  reaction, *Phys. Chem.*  
526 *Chem. Phys.*, 24, 13 026–13 032, 2022.

527 Rai, P. K. and Kumar, P.: Accurate determination of reaction energetics and kinetics of  $\text{HO}_2^\bullet + \text{O}_3 \rightarrow \text{OH}^\bullet + 2\text{O}_2$  reaction, *Phys. Chem. Chem.*  
528 *Phys.*, 25, 8153–8160, 2023.

529 Rai, P. K. and Kumar, P.: Mechanistic Inside into the Gas-Phase  $\text{NO}_3 + \text{HO}_2$  Reaction, *J. Phys. Chem. A*, 128, 7907–7913, 2024.

530 Rai, P. K. and Kumar, P.: Influence of Water on the  $\text{NO}_3 + \text{HO}_2$  Reaction, *J. Phys. Chem. A*, 129, 2067–2076, 2025.

531 Reidy, E., Bottorff, B. P., Rosales, C. M. F., Cardoso-Saldaña, F. J., Arata, C., Zhou, S., Wang, C., Abeleira, A., Hildebrandt Ruiz, L.,  
532 Goldstein, A. H., et al.: Measurements of hydroxyl radical concentrations during indoor cooking events: Evidence of an unmeasured  
533 photolytic source of radicals, *Environ. Sci. Technol.*, 57, 896–908, 2023.

534 Ren, X., Harder, H., Martinez, M., Leshner, R. L., Oligier, A., Shirley, T., Adams, J., Simpas, J. B., and Brune, W. H.:  $\text{HO}_X$  concentrations  
535 and OH reactivity observations in New York City during PMTACS-NY2001, *Atmos. Environ.*, 37, 3627–3637, 2003.

536 Ren, X., Brune, W. H., Oligier, A., Metcalf, A. R., Simpas, J. B., Shirley, T., Schwab, J. J., Bai, C., Roychowdhury, U., Li, Y., et al.: OH,  
537  $\text{HO}_2$ , and OH reactivity during the PMTACS–NY Whiteface Mountain 2002 campaign: Observations and model comparison, *J. Geophys.*  
538 *Res. Atmos.*, 111, 2006.

539 Ren, X., Gao, H., Zhou, X., Crounse, J., Wennberg, P., Browne, E., LaFranchi, B., Cohen, R., McKay, M., Goldstein, A., et al.: Measurement  
540 of atmospheric nitrous acid at Bodgett Forest during BEARPEX2007, *Atmos. Chem. Phys.*, 10, 6283–6294, 2010.

541 Rondon, A. and Sanhueza, E.: High HONO atmospheric concentrations during vegetation burning in the tropical savannah, *Tellus B*, 41,  
542 474–477, 1989.

543 Ruscic, B. and Bross, D. H.: Active Thermochemical Tables (ATcT) Thermochemical Values ver. 1.122v,  
544 <https://doi.org/10.17038/CSE/1885921>, 2021.

545 Sander, R., Baumgaertner, A., Cabrera-Perez, D., Frank, F., Gromov, S., Grooß, J.-U., Harder, H., Huijnen, V., Jöckel, P., Karydis, V. A.,  
546 et al.: The community atmospheric chemistry box model CAABA/MECCA-4.0, *Geosci. Model Dev.*, 12, 1365–1385, 2019.

547 Shabin, M., Kumar, A., Hakkim, H., Rudich, Y., and Sinha, V.: Sources, sinks, and chemistry of stabilized Criegee intermediates in the  
548 indo-gangetic plain, *Sci. Total Environ.*, 896, 165 281, 2023.

549 Sheps, L., Scully, A. M., and Au, K.: UV absorption probing of the conformer-dependent reactivity of a Criegee intermediate  $\text{CH}_3\text{CHOO}$ ,  
550 *Phys. Chem. Chem. Phys.*, 16, 26 701–26 706, 2014.

551 Slater, E. J., Whalley, L. K., Woodward-Massey, R., Ye, C., Lee, J. D., Squires, F., Hopkins, J. R., Dunmore, R. E., Shaw, M., Hamilton, J. F.,  
552 et al.: Elevated levels of OH observed in haze events during wintertime in central Beijing, *Atmos. Chem. Phys.*, 20, 14 847–14 871, 2020.

553 Smith, M. C., Chao, W., Takahashi, K., Boering, K. A., and Lin, J. J.-M.: Unimolecular decomposition rate of the Criegee intermediate  
554  $(\text{CH}_3)_2\text{COO}$  measured directly with UV absorption spectroscopy, *J. Phys. Chem. A*, 120, 4789–4798, 2016.

555 Smith, S., Lee, J., Bloss, W., Johnson, G., Ingham, T., and Heard, D.: Concentrations of OH and  $\text{HO}_2$  radicals during NAMBLEX: measure-  
556 ments and steady state analysis, *Atmos. Chem. Phys.*, 6, 1435–1453, 2006.

557 Song, M., Zhao, X., Liu, P., Mu, J., He, G., Zhang, C., Tong, S., Xue, C., Zhao, X., Ge, M., et al.: Atmospheric  $\text{NO}_x$  oxidation as major  
558 sources for nitrous acid (HONO), *npj clim. atmos. sci.*, 6, 30, 2023.

559 Stone, D., Whalley, L. K., and Heard, D. E.: Tropospheric OH and  $\text{HO}_2$  radicals: field measurements and model comparisons, *Chem. Soc.*  
560 *Rev.*, 41, 6348–6404, 2012.

561 Su, H., Cheng, Y. F., Shao, M., Gao, D. F., Yu, Z. Y., Zeng, L. M., Slanina, J., Zhang, Y. H., and Wiedensohler, A.: Nitrous acid (HONO) and  
562 its daytime sources at a rural site during the 2004 PRIDE-PRD experiment in China, *J. Geophys. Res. Atmos.*, 113, 2008.

563 Taatjes, C. A.: Criegee intermediates: What direct production and detection can teach us about reactions of carbonyl oxides, *Annu. Rev.*  
564 *Phys. Chem.*, 68, 183–207, 2017.

565 Tajti, A., Szalay, P. G., Császár, A. G., Kállay, M., Gauss, J., Valeev, E. F., Flowers, B. A., Vázquez, J., and Stanton, J. F.: HEAT: High  
566 accuracy extrapolated ab initio thermochemistry, *J. Chem. Phys.*, 121, 11 599–11 613, 2004.

567 Tan, D., Faloon, I., Simpas, J., Brune, W., Shepson, P., Couch, T., Sumner, A., Carroll, M., Thornberry, T., Apel, E., et al.:  $\text{HO}_x$  budgets in  
568 a deciduous forest: Results from the PROPHET summer 1998 campaign, *J. Geophys. Res. Atmos.*, 106, 24 407–24 427, 2001.

569 Tan, Z., Fuchs, H., Lu, K., Hofzumahaus, A., Bohn, B., Broch, S., Dong, H., Gomm, S., Häseler, R., He, L., et al.: Radical chemistry at a  
570 rural site (Wangdu) in the North China Plain: observation and model calculations of OH,  $\text{HO}_2$  and  $\text{RO}_2$  radicals, *Atmos. Chem. Phys.*, 17,  
571 663–690, 2017.

572 Teng, A. P., Crouse, J. D., and Wennberg, P. O.: Isoprene peroxy radical dynamics, *J. Am. Chem. Soc.*, 139, 5367–5377, 2017.

573 Varandas, A. and Pansini, F.: Narrowing the error in electron correlation calculations by basis set re-hierarchization and use of the unified  
574 singlet and triplet electron-pair extrapolation scheme: Application to a test set of 106 systems, *J. Chem. Phys.*, 141, 224 113, 2014.

575 Vereecken, L.: The reaction of Criegee intermediates with acids and enols, *Phys. Chem. Chem. Phys.*, 19, 28 630–28 640, 2017.

576 Vereecken, L. and Francisco, J. S.: Theoretical studies of atmospheric reaction mechanisms in the troposphere, *Chem. Soc. Rev.*, 41, 6259–  
577 6293, 2012.

578 Vereecken, L., Harder, H., and Novelli, A.: The reaction of Criegee intermediates with NO,  $\text{RO}_2$ , and  $\text{SO}_2$ , and their fate in the atmosphere,  
579 *Phys. Chem. Chem. Phys.*, 14, 14 682–14 695, 2012.

580 Vereecken, L., Harder, H., and Novelli, A.: The reactions of Criegee intermediates with alkenes, ozone, and carbonyl oxides, *Phys. Chem.*  
581 *Chem. Phys.*, 16, 4039–4049, 2014.

582 Vereecken, L., Rickard, A., Newland, M., and Bloss, W.: Theoretical study of the reactions of Criegee intermediates with ozone, alkylhy-  
583 droperoxides, and carbon monoxide, *Phys. Chem. Chem. Phys.*, 17, 23 847–23 858, 2015.

584 Vereecken, L., Novelli, A., and Taraborrelli, D.: Unimolecular decay strongly limits the atmospheric impact of Criegee intermediates, *Phys.*  
585 *Chem. Chem. Phys.*, 19, 31 599–31 612, 2017.

586 Viegas, L. P. and Varandas, A. J.: Can water be a catalyst on the HO<sub>2</sub> + H<sub>2</sub>O + O<sub>3</sub> reactive cluster?, *Chem. Phys.*, 399, 17–22, 2012.

587 Wallington, T. J. and Japar, S. M.: Fourier transform infrared kinetic studies of the reaction of HONO with HNO<sub>3</sub>, NO<sub>3</sub> and N<sub>2</sub>O<sub>5</sub> at 295 K,  
588 *J. Atmos. Chem.*, 9, 399–409, 1989.

589 Weinstock, B.: Carbon monoxide: Residence time in the atmosphere, *Science*, 166, 224–225, 1969.

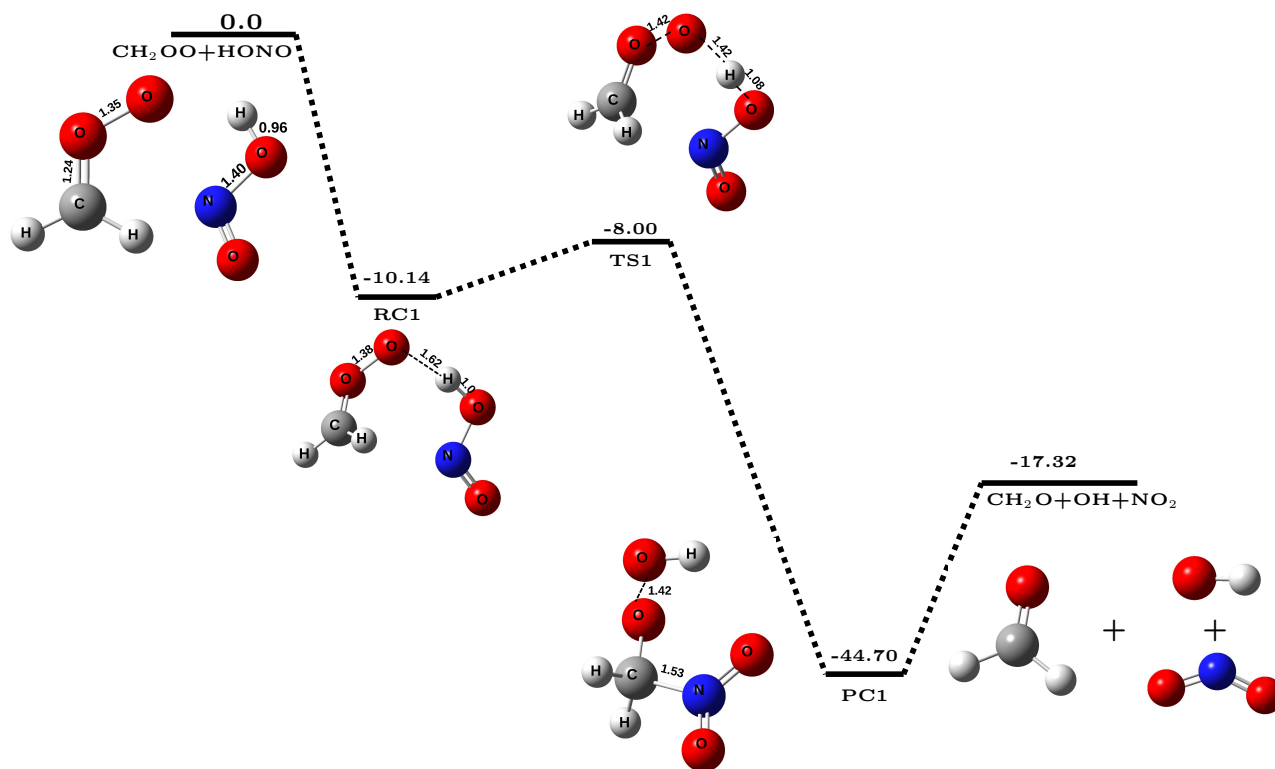
590 Whalley, L., Edwards, P., Furneaux, K., Goddard, A., Ingham, T., Evans, M. J., Stone, D., Hopkins, J., Jones, C. E., Karunaharan, A., et al.:  
591 Quantifying the magnitude of a missing hydroxyl radical source in a tropical rainforest, *Atmos. Chem. Phys.*, 11, 7223–7233, 2011.

592 Yang, X., Wang, H., Lu, K., Ma, X., Tan, Z., Long, B., Chen, X., Li, C., Zhai, T., Li, Y., et al.: Reactive aldehyde chemistry explains the  
593 missing source of hydroxyl radicals, *Nat. Commun.*, 15, 1648, 2024.

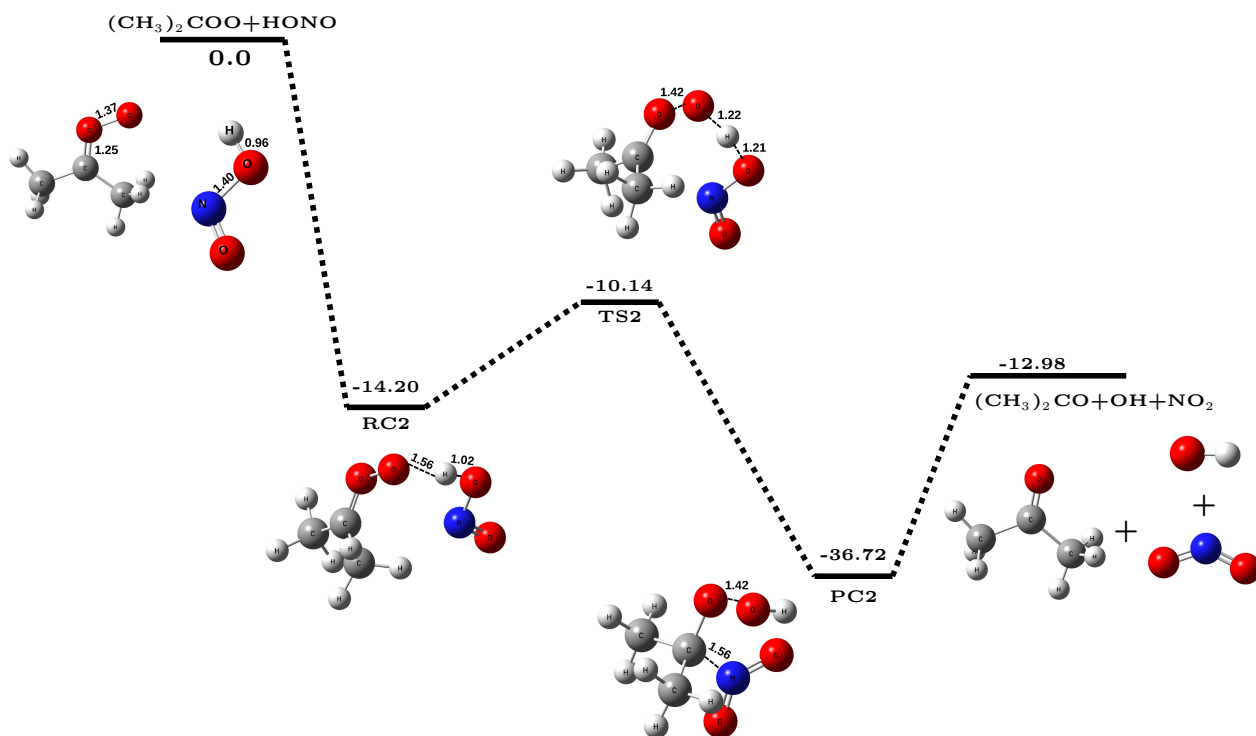
594 Zhang, N., Zhou, X., Bertman, S., Tang, D., Alaghmand, M., Shepson, P., and Carroll, M.: Measurements of ambient HONO concentrations  
595 and vertical HONO flux above a northern Michigan forest canopy, *Atmos. Chem. Phys.*, 12, 8285–8296, 2012.

596 Zhang, Q., Tie, X., Lin, W., Cao, J., Quan, J., Ran, L., and Xu, W.: Variability of SO<sub>2</sub> in an intensive fog in North China Plain: Evidence of  
597 high solubility of SO<sub>2</sub>, *Particuology*, 11, 41–47, 2013.

598 Zhou, X., Zhang, N., TerAvest, M., Tang, D., Hou, J., Bertman, S., Alaghmand, M., Shepson, P. B., Carroll, M. A., Griffith, S., et al.: Nitric  
599 acid photolysis on forest canopy surface as a source for tropospheric nitrous acid, *Nat. Geosci.*, 4, 440–443, 2011.



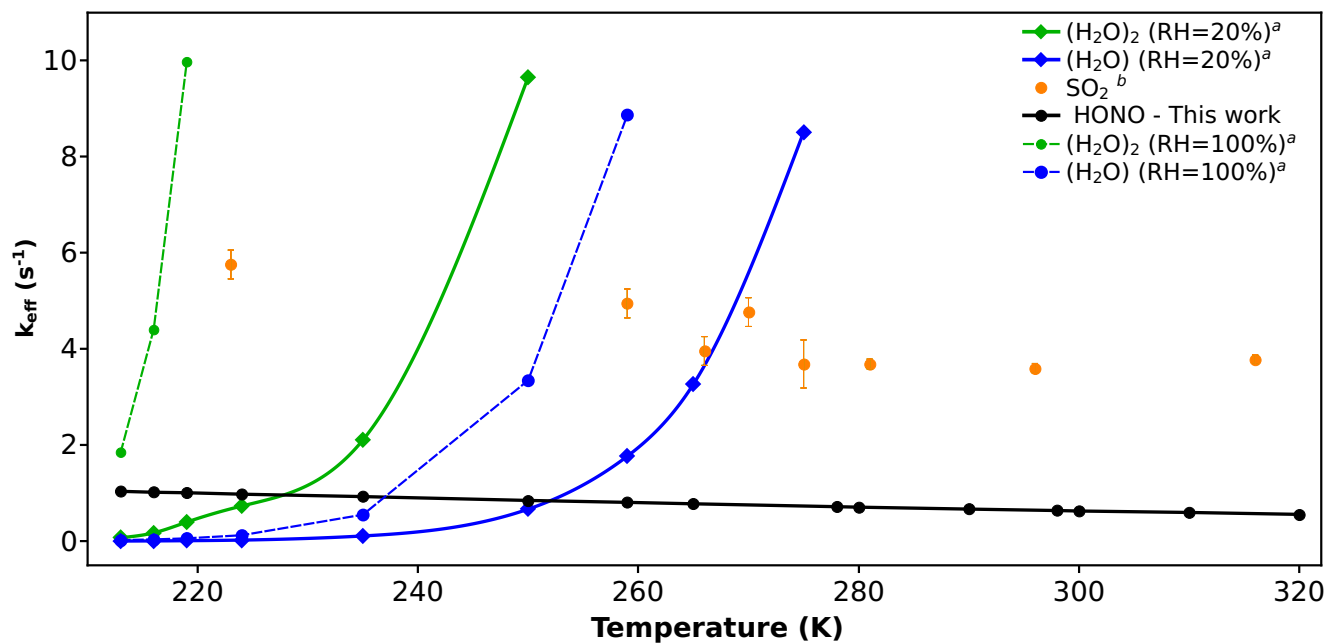
**Figure 1.** The potential energy surface for  $\text{CH}_2\text{OO} + \text{HONO}$  reaction (in  $\text{kcal mol}^{-1}$ ) obtained at CCSD(T)/CBS//M06-2X/aug-cc-pVTZ level of theory along with optimized geometries of species involved in the reaction.



**Figure 2.** The potential energy surface for  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  reaction (in kcal mol<sup>-1</sup>) obtained at CCSD(T)/CBS//M06-2X/aug-cc-pVTZ level of theory along with optimized geometries of species involved in the reaction.

**Table 1.** Bimolecular rate constants ( $k_{bi}$ , in  $\text{cm}^3 \text{ molecule}^{-1} \text{ sec}^{-1}$ ) for  $\text{CH}_2\text{OO}/(\text{CH}_3)_2\text{COO} + \text{HONO}$  reaction within the temperature range of 213–320 K.

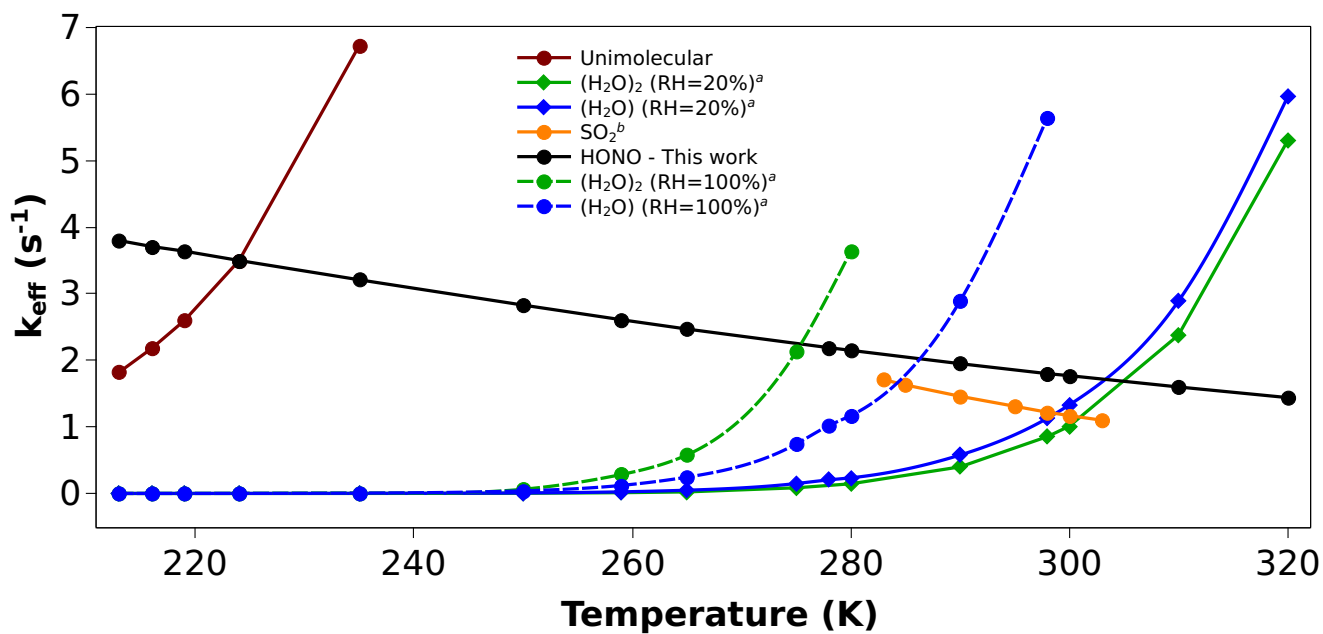
T (K)	$k_{bi}^{\text{CH}_2\text{OO}}$	$k_{bi}^{(\text{CH}_3)_2\text{COO}}$
213	$1.17 \times 10^{-11}$	$4.28 \times 10^{-11}$
216	$1.15 \times 10^{-11}$	$4.18 \times 10^{-11}$
219	$1.13 \times 10^{-11}$	$4.09 \times 10^{-11}$
224	$1.11 \times 10^{-11}$	$3.94 \times 10^{-11}$
235	$1.04 \times 10^{-11}$	$3.61 \times 10^{-11}$
250	$9.58 \times 10^{-12}$	$3.18 \times 10^{-11}$
259	$9.10 \times 10^{-12}$	$2.94 \times 10^{-11}$
265	$8.79 \times 10^{-12}$	$2.78 \times 10^{-11}$
278	$8.14 \times 10^{-12}$	$2.47 \times 10^{-11}$
280	$8.04 \times 10^{-12}$	$2.42 \times 10^{-11}$
290	$7.57 \times 10^{-12}$	$2.20 \times 10^{-11}$
298	$7.21 \times 10^{-12}$	$2.03 \times 10^{-11}$
300	$7.13 \times 10^{-12}$	$1.99 \times 10^{-11}$
310	$6.70 \times 10^{-12}$	$1.80 \times 10^{-11}$
320	$6.30 \times 10^{-12}$	$1.63 \times 10^{-11}$



**Figure 3.** Effective rate constant comparison ( $k_{eff}$ , in  $\text{sec}^{-1}$ ) of  $\text{CH}_2\text{OO} + \text{HONO}$  with the  $k_{eff}$  of previously known sinks of  $\text{CH}_2\text{OO}$ .

a. Values are taken from reference (Lin et al., 2016)

b. Values are taken from reference (Onel et al., 2021)

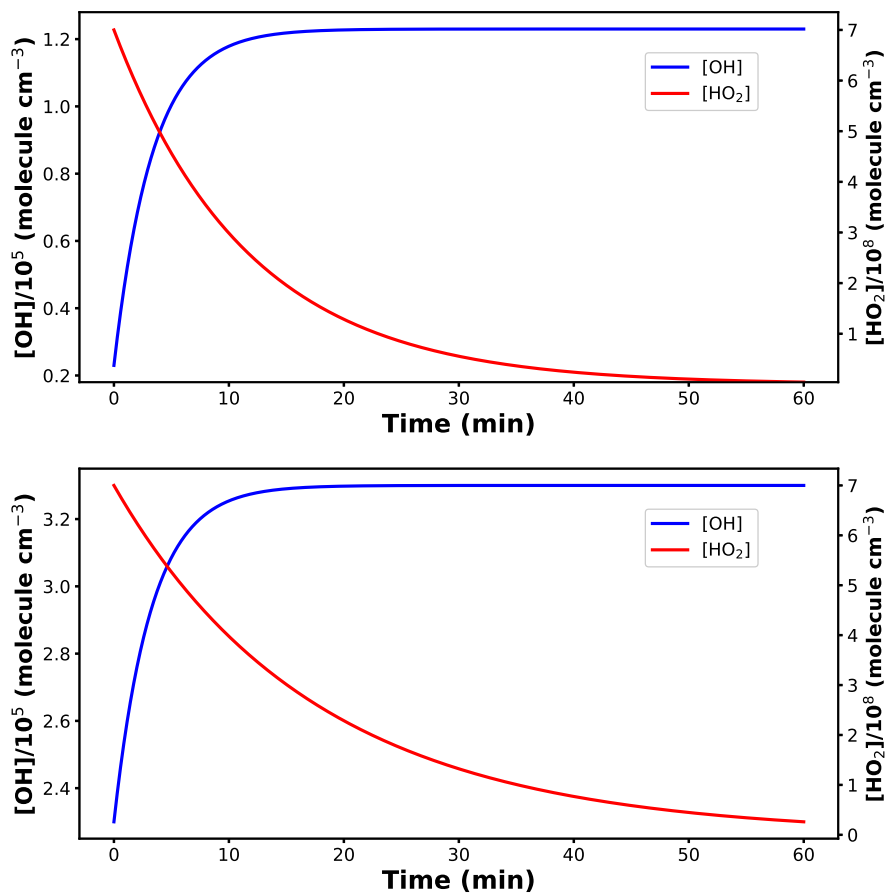


**Figure 4.** Effective rate constant comparison ( $k_{eff}$ , in  $\text{sec}^{-1}$ ) of  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  with the  $k_{eff}$  of previously known sinks of  $(\text{CH}_3)_2\text{COO}$ .

a. Values are taken from reference (Vereecken et al., 2017)

b. Values are taken from reference (Smith et al., 2016)





**Figure 5.** Top panel: Concentration profiles of  $\text{HO}_2^\bullet$  and  $\text{OH}^\bullet$  using  $\text{CH}_2\text{OO} + \text{HONO}$  reaction into the model. Bottom panel: Concentration profiles of  $\text{HO}_2^\bullet$  and  $\text{OH}^\bullet$  using  $(\text{CH}_3)_2\text{COO} + \text{HONO}$  reaction into the model.