



1 **Abiotic CO₂ Sequestration via River Runoff: A Potential "Missing Sink"** 2 **Dampening Atmospheric Warming?**

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6 ****Abstract****

7 The international system for greenhouse gas emission accounting relies on national emission and
8 absorption accounting. The effectiveness of these systems in accounting for carbon sinks is
9 questionable due to the complexities of local carbon dioxide control in the atmosphere and
10 uncertainties in forestry carbon balance assessments. This review proposes that irreversible
11 abiotic mechanisms for removing carbon from the cycle of substances should play a major role
12 in carbon sink accounting, as a complement to biotic carbon capture. While the capture and
13 fixation of carbon by ecosystems is a reversible process and represents only the first step toward
14 its removal from the atmosphere, subsequent steps leading to irreversible carbon sequestration
15 depend mainly on the interaction of ecosystems with water flows that form continental runoff.
16 This paper provides estimates that suggest a significant increase in previously existing values of
17 carbon removal via river runoff, and concludes that continental runoff is a primary mechanism
18 leading to long-term carbon sequestration.

19

20 ****Keywords****

21 Carbon sequestration, continental runoff, "missing sink," river runoff, global warming, review

22

23 ****Introduction****

24 Carbon sinks play a crucial role in the global carbon cycle, helping to regulate the concentration
25 of CO₂ in the atmosphere. Among them, terrestrial and oceanic carbon sinks receive significant
26 attention due to their absorption and storage of substantial amounts of carbon. River runoff, as a



1 link between land and ocean, plays an important role in this process by transporting organic and
2 inorganic carbon to the oceans. This review examines various estimates of carbon fluxes
3 associated with river runoff, compares them with other studies, and analyzes the reasons for
4 discrepancies between them. The novelty of this review lies in highlighting the significant role of
5 abiotic processes, specifically CO₂ dissolution in river runoff, in long-term carbon sequestration,
6 which has been underestimated in previous studies. We explore the hypothesis that abiotic CO₂
7 sequestration in river runoff constitutes a significant component of the "missing carbon sink" in
8 global carbon budget assessments.

9 ****Estimates of Carbon Absorption and the "Missing Sink" Problem****

10 The magnitude of estimated land ecosystem carbon absorption varies greatly among different
11 researchers. For example, based on an analysis of the FLUXNET database, the net ecosystem
12 production (NEP, representing the net carbon absorption) of land ecosystems is estimated at 4.3
13 Gt per year, defined as the difference between plant respiration (118.7 Gt/year) and gross
14 primary production (GPP, 123 Gt/year). Computer models of ecosystems estimate more than 2.5
15 GtC/year for northern forests; however, database estimates reduce these figures to 1.5 GtC/year
16 for the entire Northern Hemisphere. The difference between model predictions and database
17 instrumental estimates is discussed in Pan et al. (2011), Schimel et al. (2015), Keenan et al.
18 (2016), and Piao et al. (2019). These discrepancies highlight the challenges in accurately
19 quantifying carbon fluxes at regional and global scales.

20 Instrumental estimates, based on direct measurements of CO₂ in the air via airplane sampling
21 (e.g., the NASA Atmospheric Carbon and Transport (ACT) project, formerly NACA), often
22 underestimate carbon absorption. For example, Schimel (2014) stated that tropical forests absorb
23 1.4 billion tons of carbon dioxide, while the global absorption of land ecosystems does not
24 exceed 2.5 billion tons of C/year. Disputes also exist regarding the role of savannas, which
25 account for approximately 25% of total land area and contribute significantly to global carbon



1 cycling. The accurate assessment of savanna carbon dynamics remains a challenge due to their
2 complex interactions with climate, fire, and land management practices.

3 A simplified calculation of CO₂ absorption by forests, based on model forests, yields an average
4 of 3 tons of CO₂ per hectare. FAO data for Asian forests suggest a value of 5 tons/ha, meaning
5 that carbon absorption varies from 0.8 to 1.4 t C/ha. This value yields an absorption of
6 approximately 1 GtC for the forest area of Russia alone, which constitutes about 25% of the
7 world's forests. Accordingly, the estimated global forest absorption would be 4 GtC, which does
8 not align with the lower instrumental estimates. This discrepancy may be attributed to factors
9 such as forest age structure, management practices, and the inclusion of non-forest land in the
10 instrumental data.

11 Considering the classic method for estimating the increase in CO₂ in various regions of the
12 planet based on data from ground-based measuring stations (the Keeling time series, which
13 began in 1958), an important feature of the Keeling curve is that the increase in concentration
14 over 50 years averaged 1.8-1.9 ppm/year (approximately 0.5%/year). The increase in CO₂ at a
15 rate of 0.5% per year lags behind the rate of growth of anthropogenic emissions, which in the
16 first decade of the 21st century was approximately 3%/year, six times higher than the actual
17 average increase in CO₂. Estimates of the rate of carbon sequestration at 3 GtC/year obtained by
18 the ACT project for land are almost three times less than the average emission rate of 9 GtC/year
19 – insufficient to ensure such a low rate of increase in CO₂ concentration. This discrepancy
20 suggests the existence of additional carbon sinks or underestimation of existing sinks.

21 Thus, the rate of carbon sequestration in land biosystems (approximately 3 GtC/year) must be
22 supplemented by a flow into the ocean of at least 3.5 GtC/year to ensure a total flow of 6.5-6.7
23 GtC, necessary to maintain the real growth rate of CO₂ concentration at the current relatively
24 low level. These estimates conflict with the global estimates provided in "The global carbon
25 budget 1959–2011" (2012), which state that, from 1960 to 2010, CO₂ absorption by all land
26 ecosystems and the ocean surface increased at an average rate of 0.05 Gt of carbon per year, and



1 the total absorption increased, for the entire surface of the globe (including the ocean), it doubled
2 over 50 years: from 2.4 ± 0.8 to 5.0 ± 0.9 billion tons of carbon per year. Ocean absorption is
3 estimated as constant acidification at 2.5 billion tons of C/year. The carbon balance of the
4 atmosphere consists of input and output. Data on the burning of fossil fuels and deforestation
5 indicate that, in the period 2002-2011, approximately 9.3 billion tons of C entered the
6 atmosphere annually. According to the balance summarized by the authors, 28% or 2.6 billion
7 tons of C were absorbed by land vegetation, 2.5 billion tons (26%) went into the ocean, and the
8 remaining 46% increased the gas concentration in the atmosphere, adding approximately 4.3
9 GtC/year to the atmosphere. The uncertainties associated with these estimates highlight the need
10 for improved monitoring and modeling of carbon fluxes.

11 The data presented here differ from estimates based on the Keeling curve, which grows at a rate
12 corresponding to an addition to the atmosphere of no more than 3.8 Gt/year. More striking is the
13 difference in the assessment of the total absorption capacity of the ocean, considering Schimel's
14 (2014) data, indicating that instrumental land ecosystem measurements "prove" the absorption of
15 3 GtC/year. The ocean absorption estimates differ by at least 1 GtC/year. These discrepancies
16 underscore the need for a more comprehensive understanding of the processes governing carbon
17 cycling in both terrestrial and marine environments.

18 The origin of these differences is a matter of debate and is often called the problem of the
19 "missing sink." Zavarzin (2007) explicitly formulated this problem in his monograph. In the
20 same work, Zavarzin estimated the carbon stock in the soils of Russia, accumulated during the
21 Holocene (8 thousand years), at 300 GtC. However, even in this estimate, there is a "missing
22 sink," since the average annual flow over these millennia should not have exceeded $300/8 = 37.5$
23 million tons of C/year. For Russia, this estimate seems underestimated; it is unclear where the
24 remaining carbon that should have been stored during this historical period is. We will address
25 this question by considering the potential role of abiotic carbon sequestration in river runoff.

26 ****The Process of Gas Dissolution****



1 Carbon dioxide is relatively inert; therefore, direct chemical bonding without fixation in a
2 substrate does not significantly affect it. However, it is soluble in water, making water the most
3 critical substrate for dissolving gaseous CO₂. In this context, dissolution in water is considered
4 an effective mechanism for CO₂ sequestration.

5 For considering CO₂ dissolution on a planetary scale, high latitudes and great ocean depths
6 (hypolimnion) are particularly important because the solubility of carbon dioxide increases with
7 decreasing temperature.

8 This characteristic causes a significant flow of CO₂ absorption in the ocean. Only the upper,
9 relatively warm layers of the ocean (100-150 meters) become saturated, while the deeper layers,
10 lying below the "thermocline" (called "hypolimnion" in stratified bodies of water as a layer
11 below the convective mixing boundary), remain unsaturated with CO₂, forming a constant
12 reservoir for CO₂ burial. The efficiency of this process is influenced by factors such as ocean
13 circulation patterns, biological activity, and the availability of nutrients.

14 The problem of ocean acidification primarily affects "shallow water" – at depths where mixing
15 effectively evens out the temperature – as a result, the flow of dissolving CO₂ cannot proceed
16 further. Additionally, solar heating of the seabed reduces gas solubility. Therefore, the potential
17 for long-term carbon sequestration through CO₂ dissolution is greater in deeper, colder ocean
18 waters.

19 In deep waters, "dead zones" or oligotrophic waters exist, practically devoid of both oxygen and
20 CO₂. All oxygen-containing compounds are consumed by bacteria or converted into low-
21 oxidized forms, such as nitrogen oxide NO₂ losing oxygen due to bacterial activity to form
22 nitrous oxide N₂O... These regions, while limited in their biological activity, may still contribute
23 to carbon sequestration through the dissolution of CO₂ and subsequent burial in sediments.

24 Additionally, we believe that consideration of the processes occurring on land does not
25 sufficiently account for the dissolution of CO₂ in waters that are in a suspended state – in the
26 form of fog, low cloud cover, and especially in river runoff. A dissertation by Jean Yaming



1 (2023) provides strong evidence that variations in seasonal CO₂ absorption reliably increase
2 with latitude, supporting the importance of high-latitude processes in atmospheric CO₂
3 dynamics. This latitudinal gradient in CO₂ absorption suggests a significant role for cold-water
4 environments in carbon sequestration.

5 During the summer in high latitudes, coinciding with the vegetation period of northern forests,
6 large volumes of water actively dissolve CO₂. These large volumes result from the melting of
7 permafrost combined with the fog-mediated water exchange of taiga biomes. However,
8 absorption itself quickly reaches saturation, so significant absorption volumes require flowing
9 waters. Rivers flowing to the North – Yenisei, Lena, Ob, Kola, Northern Dvina, and many other
10 Siberian rivers – meet the conditions of flow and movement toward increased solubility. A
11 fundamental factor is that the great rivers of Russia carry their waters to the North and flow into
12 cold oceans. This involves a complex of processes: on the one hand, CO₂ solubility in salt water
13 is less than in fresh water, potentially leading to increased carbon dioxide release at river mouths.
14 However, the temperature difference between the upper reaches and the mouth can eliminate this
15 effect, and the overall process leads to net carbon sequestration.

16 ****Calculations****

17 To assess the solubility of CO₂ in rivers, data on annual water flow and water temperature were
18 used (Haine et al., 2015; Magritskii, 2008; Déry et al., 2016). The annual water flow of rivers
19 flowing into the northern seas is approximately 4,200 km³ for Canada and Russia combined, of
20 which Russia provides approximately 2,900 km³.

21 The solubility of CO₂ in water in the temperature range of 0-5°C varies from 0.3346 g/100 ml at
22 0°C to 0.2774 g/100 ml at 5°C (IUPAC-NIST Solubilities Database). To simplify calculations,
23 the average solubility of CO₂ in this range was used, approximately 0.3055 g/100 ml.

24 ****Calculation of Annual CO₂ Flow****

25 The total annual CO₂ flow (Canada and Russia) comes from the flow volume of rivers: 4,200
26 km³/year.



1 The results of the calculations are summarized in the following tables:

2

3 Table 1. Estimated Annual CO₂ Sequestration via River Runoff

4 Parameter	Value	
5 ----- -----		
6 Total River Runoff Volume	4200 km ³ /year	
7 Average CO ₂ Solubility	0.3055 g/100 ml	
8 Estimated CO ₂ Sequestration	12.831 Gt CO ₂ /year	
9 Estimated C Sequestration	3.503 Gt C/year	

10

11 *Table 2. Key Parameters and Data Sources*

12 Parameter	Value	Source	
13 ----- ----- -----			
14 River Runoff Volume	4200 km ³ /year	Haine et al. (2015), Déry et al. (2016)	
15 CO ₂ Solubility	0.3055 g/100 ml	IUPAC-NIST Solubilities Database	
16 Temperature Range	0-5°C	IUPAC-NIST Solubilities Database	

17

18 This calculation does not yet account for the carbon that is released with methane entering the
 19 seas along with river runoff. It is known that reservoir ices of permafrost are thousands of times
 20 oversaturated with methane (Streletskaia, 2016), which suggests its entry in the form of bubbles
 21 and methane hydrates. The fate of this methane in the Arctic Ocean and its potential contribution
 22 to greenhouse gas emissions require further investigation.

23

24 **Discussion**

25



1 The estimated value of 3.5 GtC/year suggests a re-evaluation of the role of the taiga/boreal
2 forests in the Northern Hemisphere. Traditionally, the role of permafrost and boreal forests is
3 considered as direct absorbers of CO₂. However, our assessment suggests that forests play a role
4 not so much in direct carbon absorption as in ensuring the consolidation of the flow of
5 underground and surface moisture, i.e., they carry out critical management of moisture turnover.
6 At the same time, the main role is played by the amount of condensed water in the forest
7 turnover, and then filling the flow of Siberian rivers. Water stores an additional amount of CO₂,
8 which remains dissolved along the entire course of the river to the north, in areas of increasingly
9 lower temperatures. This flow of northern rivers stores a much larger amount of CO₂ than
10 previously made estimates based solely on carbon storage in wood or permafrost ice. The
11 hydrological function of boreal forests in facilitating abiotic carbon sequestration in river runoff
12 deserves greater attention.

13

14 From the point of view of photosynthesis, the taiga should have yielded to tropical forests in the
15 competition for primacy in managing carbon dioxide balances. If tropical forests led, seasonal
16 CO₂ absorption would not reliably increase with latitude, as evidenced by the data presented in
17 Yaming (2023). The observed latitudinal gradient supports the importance of high-latitude
18 processes, including abiotic carbon sequestration in river runoff, in regulating atmospheric CO₂
19 levels.

20

21 Our estimates differ significantly from previous studies, such as the global carbon budget
22 reports, which estimate ocean carbon uptake at around 2.5 GtC/year. This discrepancy may arise
23 from the underestimation of abiotic CO₂ dissolution in river runoff in previous assessments.
24 Further research is needed to refine these estimates by incorporating more complex models that
25 account for factors such as river temperature, pH, salinity, and hydrological characteristics. The



1 use of remote sensing data and improved modeling techniques can help reduce the uncertainties
2 associated with these estimates.

3

4 ****Conclusion****

5

6 This study highlights the significant role of river runoff in the absorption of atmospheric carbon
7 dioxide, particularly in high-latitude regions. Our estimates suggest that the amount of carbon
8 removed by this mechanism has been underestimated in previous studies. The findings
9 emphasize the importance of considering abiotic processes, such as CO₂ dissolution in river
10 water, in assessing the global carbon cycle. Future research should focus on refining these
11 estimates by incorporating more detailed data and models, as well as investigating the fate of
12 dissolved carbon in the Arctic Ocean. The accurate assessment of carbon sequestration through
13 river runoff is critical for developing effective strategies to mitigate climate change. This
14 includes incorporating these processes into global carbon budget models and developing policies
15 that promote sustainable water management practices.

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17 ****Competing interests****

18 The author declare that have no conflict of interest.

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