# Expanding Observational Capabilities of A Diode-Laser-Based Lidar Through Shot-To-Shot Modification of Laser Pulse Characteristics

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**Abstract.**

    A method for expanding the observational capabilities of diode-laser-based atmospheric lidar is discussed. A straightforward test, consisting of interleaved "long" and "short" laser pulses, was developed to demonstrate how shot-to-shot modification of laser pulse characteristics can enhance the performance of low-power, diode-laser-based lidar and could benefit atmospheric

5  observations. Two examples are given to demonstrate the technique. In the first, water vapor profiling is extended closer to the surface while simultaneously maintaining consistent far-range performance. These results are verified with collocated measurements. In the second example, clouds are resolved at high vertical spatial resolution with high signal-to-noise ratio. Details of the lidar instrument hardware and the method to combine the laser pulses of different durations are given.

## 1 Introduction

10  Lidar observations have been used for a number of atmospheric science applications, including, but not limited to, process studies, weather/climate monitoring, and assimilation into numerical weather prediction models. While scientific applications are typically the ultimate goal, scientific advancements are often fostered by technological advancements. As such, much of the focus of research on modern lidar systems centers around lidar hardware, in particular, the laser sources used. Different types laser sources can be employed in lidar, depending on the desired measurement requirements. Historically, diode lasers

15  have been primarily used in ceilometers but are expanding in their applicability in recent years to more quantitative lidar systems such as differential absorption lidar (DIAL) (Nehrir et al., 2011) and high spectral resolution lidar (HSRL) (Hayman and Spuler, 2017). Diode lasers are effective laser sources for these types of quantitative lidar because they are cost-effective, robust, available at a large number of wavelengths, have narrow bandwidth, and are flexible in their operational parameters. In contrast, the prime drawback of diode lasers is their low peak power.

20    Roughly a decade ago, the potential for quantitative atmospheric lidar systems based on diode lasers was first demonstrated (e.g. Nehrir et al. (2011); Repasky et al. (2013); Hayman and Spuler (2017)). The benefit of a diode-laser-based lidar architecture stems from higher reliability and lower-cost systems, not from their pure performance vis-à-vis output power. For high-power atmospheric lidar systems, solid-state lasers are typically used (e.g. Nd:YAG, Ti:Al$_2$O$_3$, or Alexandrite) (e.g. Thayer et al. (1997); Goldsmith et al. (1998); Eixmann et al. (2015); Wulfmeyer (1998); Strotkamp et al. (2019)). In contrast,

25  diode lasers have relatively low peak power, on the order of Watts – compared to 100s of Kilowatts or Megawatts from solid-state lasers. As part of a compensating mechanism, diode-laser-based lidar systems have typically used relatively long pulse lengths, on the order of 100s of nanoseconds to microseconds.

Generally, for a coaxial monostatic lidar system, the lidar system is blind when the laser pulse is outgoing. Detector recovery time can add to this blind zone. For typical high-powered solid-state lasers, the pulse durations are short (on the order of nanoseconds), and the blind period imposed by the pulse itself is practically irrelevant, leaving only detector recovery time to impact the dead time. However, for diode-laser-based lidar there is a direct trade-off. Longer pulses, having more average power, enable long-range observations but sacrifice near-range observational capability. Shorter pulses in contrast can be used to see at nearer ranges and at higher range resolution but lack the average power to see to greater distances.

Although diode lasers have relatively low peak power, they enable more flexibility, modularity, and agility than conventional

35  solid-state lasers. The concept to be explored in this work is to leverage that flexibility – to simply transmit multiple outgoing laser pulse lengths, altering laser parameters from shot-to-shot, to expand the observational range and resolution of diode-laser-based lidar systems. This has the benefit of allowing for rapid changes in observational sensitivity, with longer pulses used to see to further ranges and shorter pulses used to see to closer ranges and enable higher-resolution observations within dynamic targets, such as clouds. Although the concept of altering pulse characteristics from shot-to-shot as well as sub-pulse alterations

40  (typically referred to as pulse compression) is well explored for radar hardware (e.g. Ramp and Wingrove (1961); Mudukutore et al. (1998); O'Hora and Bech (2007); Richards et al. (2010)), to our knowledge these techniques have not been adopted for use in atmospheric lidar. This is likely due to the limited flexibility of high-power solid-state lasers where implementing such a capability would represent a significant engineering challenge. These lasers are best behaved in steady state operation due primarily to thermal transient effects (Hooker and Webb, 2010). Altering operational characteristics, such as pulse length, pulse

45  repetition frequency, or peak power, of most solid-state lasers significantly alters the thermal and refractive-index properties of the gain material. If not carefully mitigated, localized heating from the resultant thermal lensing can cause catastrophic damage to optical components. Although there are thermal effects for diode lasers, the relatively high wall-plug efficiency[1] and the comparatively lower peak power of diode-lasers allows most of the energy input into the laser to leave as light with localized heating minimized.

## 1.1 Contribution and Outline

50  The primary contributions of this work are: first, to implement a timing unit on an atmospheric observing lidar system to allow for this shot-to-shot control of laser pulse characteristics, to our knowledge for the first time, and second, to explore the enhanced observational capability of such a lidar system. The lidar hardware is briefly described in Section 2. The system used to perform the shot-to-shot alteration of laser characteristics is described in Section 3. The implementation of this system to

55  expand the observational capability of this lidar system to measure water vapor is described in Section 4. A discussion of the

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[1]We define wall plug efficiency to be the total output optical power from a laser divided by the total input electrical power it takes to run that laser. This may include energy for optical pumping, control systems, and thermal management.

ability to combine data from multiple pulse lengths is presented in Section 5. Finally, a discussion of the results, conclusions, and future work are given in Section 6.

## 2  MicroPulse Differential Absorption Lidar (MPD)

The MicroPulse Differential Absorption Lidar (MPD) used in this work uses a pair of DIAL systems. One DIAL operates near 828 nm and is used to measure water vapor absolute humidity while the other DIAL operates at 770 nm and is used, with an embedded HSRL, to measure temperature. The MPD instrument and most recent version of the laser system are described in Spuler et al. (2015) and Spuler et al. (2021), respectively, while the observations are described in Stillwell et al. (2020) and Hayman et al. (2024). There are two differences between the configuration presented in Spuler et al. (2015, 2021) and the one used for this work that should be noted. Additionally, while MPD consists of a pair of DIAL systems, we chose to focus on a single system, the water vapor system, for this work.

The first difference is in the optical hardware used to interleave the DIAL laser wavelengths. A trio of optical switches (a pair of 1x1 and a single 2x1 switch) are replaced with a pair of booster optical amplifiers (BOAs) that are operated in pulsed mode combined with a fiber optic coupler. When the pulse current is applied, the BOA acts to amplify a continuous wave seed laser. When off, the amplifier acts as a strong absorber (effectively a closed switch). This pair of amplifiers is coupled with a passive fiber coupler to a tapered semiconductor optical amplifier (TSOA). The pulses from the BOAs are interleaved so that only one wavelength from each DIAL may fire at a time. The overall performance of this laser design is similar to the performance described by Spuler et al. (2021) with the benefits of being more stable, robust, temperature insensitive, with better cross-talk mitigation between wavelengths, and more cost-effective. A block diagram of this design is provided in Appendix A, Figure A1, which can be compared to Spuler et al. (2021) Figure 1.

The second difference is the replacement of the main clock that controls the timing of the laser and receiver systems. The clock used for the system described by Spuler et al. (2015, 2021) was an 8-channel pulse generator (Quantum Composer 9530). This pulse generator served as the master clock for the lidar system. This pulse generator provided a consistent input to the lidar system, i.e. making the lasers operate at steady state. The replacement is a custom FPGA-based pulse generator, which is described in more detail in Section 3. This custom FPGA-based pulse generator provides all of the same electrical connections but allows for shot-to-shot changes in laser timing.

## 3  Primary Timing Unit

The FPGA based timing unit consists of a commercial off-the-shelf single-board computer module (control board) connected to a custom-built interface printed circuit board (interface board). The control board is based on a Xilinx Zynq System-on-Chip. This chip combines an FPGA and an ARM CPU capable of running embedded software. A combination of software and FPGA firmware implement the required control functions. The custom-built interface board converts electrical signals from the control board to provide TTL-like output to each of MPD's DIAL laser pairs (each consisting of 1 TSOA and 2 BOAs),

detectors (consisting of 3 identical avalanche photo diode gate signals), and signals to the MPD acquisitions system (1 trigger, 1 binary flag to indicate which laser, online or offline, is firing, and a binary flag indicating the pulse width).

The FPGA timing unit allows for the definition of multiple pulsing schemes and can switch from one scheme to the next in a single FPGA clock cycle (in this case 5 ns). The definition of each pulsing scheme includes individual definitions for the inputs provided by the original master clock, definitions of the pulse repetition time, the total number of pulses to fire, the split between the two DIAL wavelengths, and the amount of time to wait after completion of each block. Each of the inputs provided by the new master clock is defined in one of two ways: as a pulse or as a binary value. Each pulse is defined by 3 numbers: a pulse length, a delay, and an On-Off keying sequence. The intention of the pulse delay is to account for any rise times or any electrical delays that would be required to synchronize components, while the On-Off keying sequence is to allow for intrapulse modulations.

This timing unit was born out of a design used for the NSF NCAR HIAPER Cloud Radar (Vivekanandan et al., 2015). Some of the features are unused for this work, but are described and noted for future capability enhancement. Specifically, the amount of time to wait after completion of each block allows for different unambiguous ranges for each pulse block and provides block-to-block control of the pulse repetition frequency (used in Doppler radar for velocity unfolding). The On-Off keying is intended to allow intra-pulse modifications such as would be necessary to implement radar-style pulse compression by intra-pulse phase modulation.

For this work, we specified 3 blocks (of the 32 possible) as defined in Table 1. The first block we refer to as "Long", the second "Short", and the third "Very Short". Henceforth, quotation marks are ignored with names identified by their capitalization. The Long block defines the MPD settings needed to fire 1 $\mu$s laser pulses while the Short and Very Short block are the settings for 0.2 $\mu$s and 0.1 $\mu$s pulses, respectively. The Long block is chosen as it mimics the operational settings used for previous MPD observations (Spuler et al., 2015, 2021; Stillwell et al., 2020; Hayman et al., 2024). Each block fires 2 shots, one with the online laser and one with the offline laser for each DIAL pair, which are synchronized. This creates a series of 4 laser pulses before the system repeats. The laser output rate is 8 kHz, 2 kHz effectively for each pulse type (Online Long, Offline Long, Online (Very) Short, and Offline (Very) Short). This is shown in Figure 1.

## 4 Water Vapor Observations

NSF NCAR operates a network of five MPD systems. One of these instruments (MPD04) was equipped with the timing unit described in Section 3 and deployed at the NSF NCAR Marshall Field Test Site near Boulder, Colorado (39.95 N, 105.196 W) in fall 2024 (NCAR/EOL-MPD-Team, 2020). The MPD was run unattended from 16 October to 2 November 2024. A total of nine collocated radiosondes were launched at the Marshall field site. MPD was run with timing as shown in Figure 1 with an even split between short pulses and long pulses.

MPD observations (photon counts) were processed using the method of Stillwell et al. (2020), which itself is largely a restatement of the method originally defined by Spuler et al. (2015). Spectroscopic calculations were performed following the method of Hayman et al. (2019) trained with data from Hitran 2020 (Gordon et al., 2022). Error estimates are calculated

| Definition | Block 1 | Block 2 | Block 3 |
|---|---|---|---|
| Pulse Repetition Time | 125 $\mu$s | 125 $\mu$s | 125 $\mu$s |
| Number of Pulses | 2 | 2 | 2 |
| Number of Online Pulses | 1 | 1 | 1 |
| Number of Offline Pulses | 1 | 1 | 1 |
| Block Post Time | 0 $\mu$s | 0 $\mu$s | 0 $\mu$s |
| Channel Binary ID | 0 | 1 | 1 |
| TSOA | | | |
|     Timer Width | 1 $\mu$ | 0.2 $\mu$s | 0.1 $\mu$s |
|     Timer Delay | 0.1 $\mu$ | 0.1 $\mu$s | 0.1 $\mu$s |
| BOA | | | |
|     Timer Width | 1.1 $\mu$ | 0.3 $\mu$s | 0.2 $\mu$s |
|     Timer Delay | 0.1 $\mu$ | 0.1 $\mu$s | 0.1 $\mu$s |
| Detector Gate | | | |
|     Timer Width | 1.2 $\mu$ | 0.4 $\mu$s | 0.3 $\mu$s |
|     Timer Delay | 0 $\mu$ | 0 $\mu$s | 0 $\mu$s |
| Acquisition System Trigger | | | |
|     Timer Width | 1 $\mu$ | 1 $\mu$s | 1 $\mu$s |
|     Timer Delay | 0.1 $\mu$ | 0.1 $\mu$s | 0.1 $\mu$s |

**Table 1.** Clock settings used for this work. The timing unit is set to switch between blocks upon completion of pulsing defined by the Number of Pulses. Note that all On-Off keyed sequences are disabled for this work.
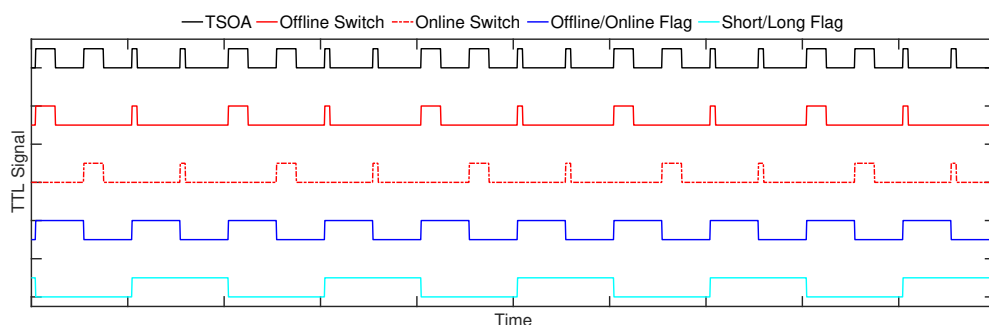
**Figure 1.** Timing diagram of the interleaved Long and Short modes defined in Table 1. The "Switch" signals are the drive current to the BOAs. Timing is to scale except pulses are intentionally enlarged by a factor of 40 for illustrative purposes. The bottom two signals are used as binary signals into the MPD acquisition system to parse out the data from the single water vapor detector to 1 of 4 acquisition channels where: 00: Offline Long, 01: Offline Short, 10:Online Long, 11: Online Short.

120     using the bootstrapping method defined in Spuler et al. (2021). Note that radiosondes measure relative humidity, whereas MPD measures absolute humidity. To compare the two observations, the relative humidity of the radiosondes was converted to absolute humidity using the saturation vapor pressure calculated using the approximation method of Lowe (1977). The saturation vapor pressure was then converted to number density using the ideal gas law.

    Absolute humidity and relative backscatter are shown in Figure 2. Radiosonde data are plotted on top of both MPD absolute
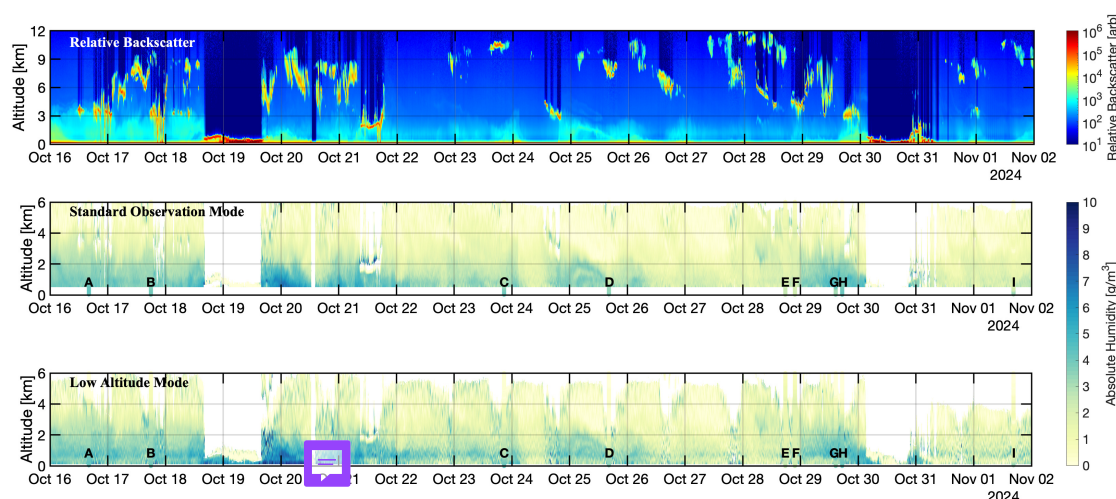
**Figure 2.** MPD data taken from October 16, 2024 through November 1, 2024. Water vapor measurements from 'Standard observation' (Long pulse) and 'Low Altitude' (Short/Very Short pulse) modes are taken per the settings in Table 1. Radiosonde data is overlaid on the water vapor panels with letters corresponding to panels in Figure 3. Relative backscatter is given for context – to identify where clouds are blocking the lidar signals and where high background levels from clouds are to be expected. All radiosonde labels are at the same altitude for context – to identify the the minimum observable altitude for Standard Observations.

125     humidity contours, each labeled with a letter corresponding to the panel in Figure 3. Relative backscatter originating from the Long laser pulses is given for context where the stronger signals are mostly clouds. The "Low Altitude Mode" shown in Figure 2 consists of the Short pulse block from October 16 to the middle of the day on October 30. It was changed for the last 2 days to the Very Short pulse block.

    There are a couple of things to note about Figure 2. First, it is generally difficult to differentiate the MPD absolute humidity

130     from those observed from the sonde; the easiest place to see a difference is in the lower 500 m of the Long pulse observations where MPD data is absent. This is expected given the extensive validation of the instrument and the quantitative nature of the DIAL method (Weckwerth et al., 2016; Spuler et al., 2021). What is noteworthy is that operating the system in two pulse modes has not proved detrimental to the original instrument performance, less the decline in performance that results from using effectively half the laser pulses in a given time period. Second, the Long pulses generally have data that extend further

135  in range than the short pulses. This is due to the higher average output power of the Long pulses increasing the observed SNR. This is especially true during daylight hours, which are approximately 13 UTC to 24 UTC and even more prominent under daytime clouds, which have the highest rate of background light observations. It is also clear that after the change from the Short to Very Short pulse blocks, the Low Mode water vapor maximum range drops.
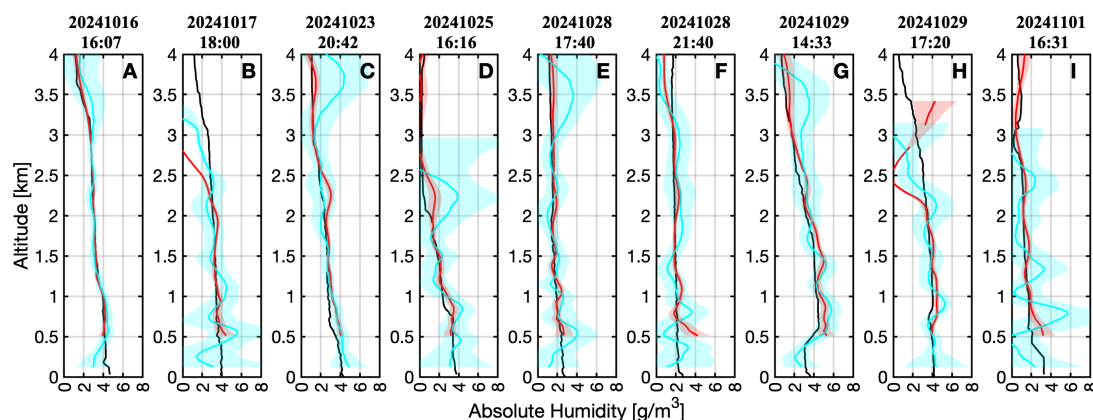
**Figure 3.** Comparisons of MPD WV data and collocated radiosondes, black, launched over the time period for Figure 2. Red is Long pulse mode and cyan is Short pulse mode for panels A-H and Very Short pulse mode for panel I. Solid lines are retrieved values with shaded bars giving error estimates.

Figure 3 shows comparisons of MPD data to radiosonde data. Radiosonde data is given in black and MPD measurements
140  are given in the solid lines, where red lines are Long pulses and blue are either Short or Very Short pulses. Shading of the same color is the measurement error estimate. There are three things that should be highlighted. First, the blue lines (either Short or Very Short pulses) extend significantly lower than the Long pulses; in this case the minimum range is reduced from approximately 500 meters to 90 meters. Second, of particular interest is the sonde shown in panel G, which shows a relatively strong surface inversion, which is well captured by the MPD with shorter pulses. This inversion, if observed with longer pulses is invisible. Using a method of inference, such as interpolation from the lowest range bin to the ground, would fail to see this
145  feature, which could have consequences related to interpretation of the data. Third, the error bounds are strongly linked to background count rate, and by extension daytime cloud cover. Panels A and G occur under clear conditions, panels C, D, E, and I occur under cirrus clouds, and panels B, F, and H under lower status clouds.

## 5  Sample Observations

150  So far, we have treated data from the MPD's different pulse modes independently. We used direct retrieval methodologies to convert the MPD observations (photon counts) to water vapor absolute humidity, yielding two independent estimates. This is analogous to more traditional lidar systems that use observing channels of different sensitivity for low- and high-altitude observations. Examples include using two or more optical receivers with duplicate detectors (e.g. von Zahn et al. (2000);

Reichardt et al. (2012); Newsom et al. (2020)) or splitting the measured detector signal electrically (e.g. Newsom et al. (2009);
155 Carroll et al. (2022)). This leaves a number of natural questions including: 1) Are the measurements necessarily equal? 2) If
not, which measurement is more representative of the atmospheric state? and 3) Can measurements be combined to yield a best
guess of the atmospheric state? It is important to note that this set of questions is not unique to the methodology presented here
and does not have a single universal solution.

Several methods have been proposed to answer these questions. Newsom et al. (2020) specify a fixed threshold whereby
160 the far range channel is used above a set attitude, the low range channel is used below a second set altitude, and a linearly
varying weighted average is used between those two altitude bounds. Carroll et al. (2022) specify something similar, where the
data merge locations are not specified by altitude but by the observed optical depth. Forward models have also been shown to
combine measurements of similar atmospheric variables with different sensitivity. For example, Löhnert et al. (2009) combine
observations from passive remote sensors in the infrared and microwave while Cadeddu et al. (2013) described combining
165 microwave remote sensors utilizing a pair of microwave water vapor emission lines with different zenith opacity. The method
of Marais and Hayman (2022) was used to forward model MPD water vapor absolute humidity data. It uses data from only one
sensitivity channel (in this case, our "Long" pulses), but it is theoretically extendable to two sensitivities. We elect to highlight
this method in particular because it has already been applied to MPD data, but applying it to this data is beyond the scope of this
work. Combining data in a robust manner should not be considered trivial. Presented here are two examples of measurements
170 of differing sensitivity, one straightforward to combine and one that is not.

## 5.1 Combined Water Vapor Retrievals

The method of data combination described by Newsom et al. (2020) is used with minor modifications to combine Long and
Short pulse data from October 29, 2024, presented in Figure 4. Data sets are merged with a linear weight beginning at 500
meters and ending at 750 meters, i.e. data below 500 meters come exclusively from Short pulses and data above 750 meters
175 come exclusively from Long pulses. Figure 4 also shows the error estimate of the presented data both as absolute humidity
and as a percent error. These data were selected because of a relatively rare moisture inversion for the first half of the day that
slowly dissipates. That feature is unique and stands out clearly on panel G of Figure 3 but largely disappears 3 hours later
shown in panel H of Figure 3.

Error estimates are provided in two ways because they highlight different components of DIAL-based moisture measure-
180 ments. First, the error for the Short pulses is naturally larger than the Long pulses, as shown in the middle panel of Figure 4.
This is first because Short pulses ($\approx 1 \mu J/pulse$) are approximately 5 times less powerful than Long pulses ($\approx 5 \mu J/pulse$)
and second, typical of most lidar systems, due to poor overlap at short ranges the MPD does not collect light as efficiently at
near ranges as it does at far ranges. Absolute error spikes near 18 UTC corresponding to background light enhancement caused
by a cloud. Near 21 UTC, the cloud layer passes. This is basically hidden in the Percent Error contour because the error is
185 rising due to background conditions as the absolute humidity is rising due to the air mass changing during the observation. It
is serendipitous that the magnitude of these two effects is almost the same. The percent error contour is helpful for identifying
clouds in this case. DIAL observations typically show large biases within cloud; as a result, most DIAL-based data near clouds
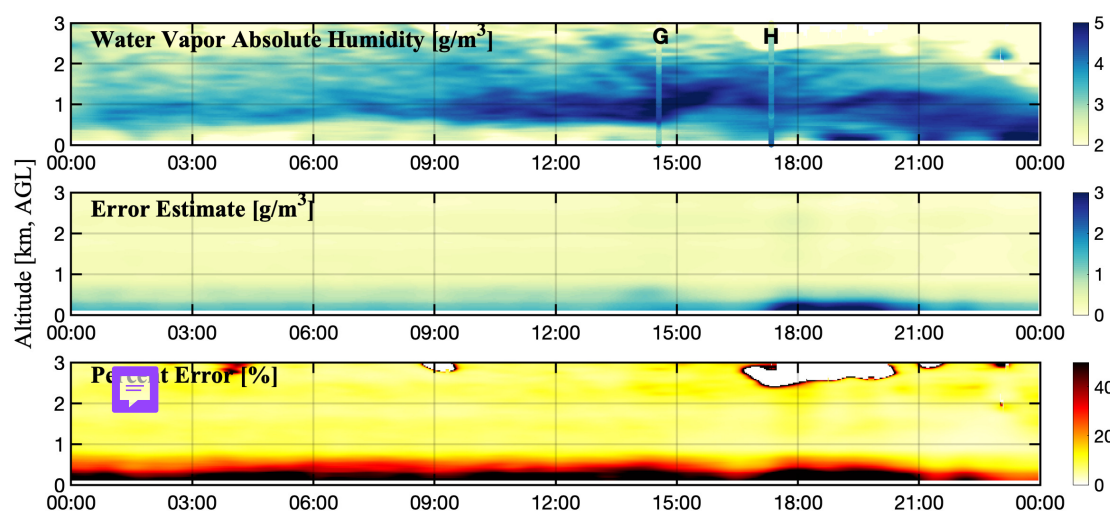
**Figure 4.** MPD data from October 29, 2024. Radiosondes overlaid on the top panel correspond to panels G and H from Figure 3.

and within clouds is masked. In this case, the bias is a strong dry bias. The absolute error estimate is quite low in cloud, not because the actual error is low, but because the observed variance during bootstrapping is low due to this dry bias. There is not an enhancement in absolute error from the cloud at 9 UTC near the ground because the Sun is not above the horizon, and thus the solar background observed is not enhanced by the cloud.

## 5.2 Cloud Retrievals

While combination of data using fixed altitude or optical depth thresholds may be possible in some instances, others scenes are more difficult. An example of this is given in Figure 5 showing observations from MPD04 in Albany, New York (42.68 N, 73.816 W) on May 7, 2024. This example contains a cloud scene that is best observed partially with Short pulses and partially with Long pulses.

This particular cloud scene contains distinct varieties of cloud with different signal dynamic ranges. From approximately 2.4 to 2.7 UTC, at approximately 7 km range, there is a strong scattering target that is relatively thin. The layer vanishes at approximately 3.4 UTC. It is possible that this is a thin supercooled liquid containing cloud based on the relative signal strength. That thin liquid layer is embedded within less strongly scattering targets extending from approximately 6.75 km to 12 km range. These are likely ice clouds given their relative scattering efficiency and the layer's temperature (the layer has a temperature of approximately -20° C at 7 km). For the sake of discussion, we will refer to these two cloud layers as liquid and ice for convenience of designation; it is practically irrelevant if this is correct. It is merely relevant that they have distinctly different optical depth per meter and, by extension, produce disparate photon count rate dynamic ranges.
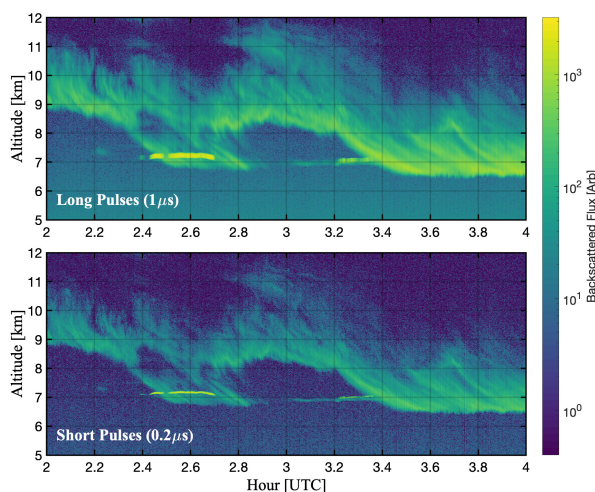
**Figure 5.** Cloud observations from May 7, 2024 in Albany, NY. Contours show observations of backscattered photon flux from the water vapor offline channel.

Examining the observation from the Long pulses, the ice clouds extend to at least 12 km; the same layer is not as clearly resolved when observed with Short pulses beyond 10.5 km. This is due to the higher overall power emitted from the Long pulses as compared to the Short pulses. In this case, the choice to use Short pulses causes errors in interpretation of the cloud scene; i.e. the observations are not completely representative of the atmosphere because they lack sensitivity. Considering that the tropopause tends to be lower at the poles and higher at the equator (maybe 7 km and 15 km above mean sea level, respectively), this could cause an observational bias of cloud occurrence and height.

In contrast to the observation of the ice clouds, the liquid cloud seems to have an oddly uniform thickness when observed with Long pulses. This is a result of the observations being a convolution of the scattering target and the laser pulse. In this case, the liquid cloud appears to be approximately 180 m thick when observed with a $1\,\mu s$ pulse (150 m range smearing) and 60 m thick when observed with a $0.2\,\mu s$ pulse (30 m range smearing). This indicates that the layer is likely not more than 30 m thick. The oddly uniform thickness when observed with the Long pulses is ostensibly an observation of the Long pulse itself. However, very thin cloud layers, such as those observed by McCullough et al. (2019), which were laminate structures on the order of 10 m thick, would be unobservable and effectively smeared together. The shorter the pulse, the more reasonable it would be to observe sharper features.

In this case, manual combination of the two layers would be challenging because they are not separated in space, one layer is embedded within the other. To combine this scene, and indeed many other cloud scenes in a general way, would require a method that was capable of balancing the sensitivity of each observation point by point. The previous example was possible because we assumed the nature of this balance of sensitivity, i.e. that the Short pulses would be worse at observing to greater range; that is untrue in this case and generally for cloud scenes. To our knowledge, concrete demonstration of such a method

is absent in the literature[2]. We suggest forward models might well be a method capable of balancing these sensitivities in a
general way without requiring heavy handed intervention, but we are aware of no published method capable of combining the
data from this scene.

## 6 Discussion and Conclusion

The shot-to-shot alteration of pulse characteristics in this work has effectively provided the MPD instrument with a low-altitude
observing channel. More traditional lidar systems have accomplished this in a number of ways, for example: 1) splitting the
measured signal optically and analyzing it with two or more optical receivers, 2) altering the electrical gain of different observa-
tional circuits used for record detector voltages, or 3) splitting detector signals to do both analog detection and photon-counting
detection simultaneously. It is worth noting that with a long pulse blinding the receiver, none of these other solutions are effec-
tive. The method introduced removes this issue, but also has the advantage that it introduces no additional optical components
and can work with purely photon-counting observations. The cost of this method is that the observational sensitivity comes
from altering the lidar transmitter, not the receiver. Traditional techniques focused on the receiver serve to attenuate both the
observed signal and the prime noise contribution, typically solar background and detector noise.

Given a single observational parameter (in this case, pulse width), maximum and minimum ranges have requirements that
are essentially diametrically opposed. Historically, MPD has been run with pulse widths between 0.65 $\mu$s and 1 $\mu$s, with the
vast majority being 1 $\mu$s. This allowed for observation typically between 300-500 meters and to approximately 4-6 km. This
was found to be a reasonable compromise between competing desire to simultaneously see both as high and as low as possible.
The timing unit demonstrated in this work has added a number of different parameters that can be used to better design the
pulse strategy. The pulse width has been extensively discussed, but the duty cycle of pulses can be adjusted as well.

In this case, we have maintained Long pulses as the MPD standard 1 $\mu$s and simply shortened pulses in an attempt to see
lower and at higher resolution. The 1 $\mu$s pulse was maintained because MPD performance in this mode is well understood and
has been extensively validated. However, given that the compromise to see to lower altitudes was primarily what prevented
the MPD pulse length from being longer originally, it is likely that longer pulses would increase the MPD maximum range
with virtually no penalty assuming it was paired with short pulses. Furthermore, we have also used a 50/50 split in Long
and Short pulses, respectively. This was chosen perforce by the MPD acquisition system requiring equal pulse numbers in its
preintegration. We are actively working to remove that limit. Assuming this hardware limitation is removed, it is also likely
that a 50/50 split is not ideal. Given that the loss of received signal is a function of range from the lidar system squared $(R^2)$,
it is expected that a more skewed split (e.g. 90/10 or 95/5) would yield better long-range performance while not providing a
major penalty for short-range observations.

The scope of this work has been to enhance the observational range of a diode-laser-based lidar system using shot-to-shot
changes to pulse parameters. The solution proposed and demonstrated is a new timing system that has allowed for pulse-to-

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[2]We note that the method of Hayman et al. (2020) might seem applicable here but was designed to balance the resolution of the smoothing kernel applied
to observations from a single sensitivity channel. In this case, the smoothing is done by the pulse length itself, not by the post-processing routines.

255   pulse changes in the laser shot duration. We have focused on a single parameter in an attempt to move the minimum range of observation to lower altitudes and to enhance the resolution of observations in clouds. In general, while it is beyond the scope of this work, we expect that extending observational range will be further accomplished by skewing the pulse split towards longer pulses and allowing for longer pulses than 1 $\mu$s. An example might be a 90/10 Long/Short split with 2 $\mu$s and 0.2 $\mu$s lengths, respectively. This would focus most of the laser power and observations at far range, where the signals are weakest,

260   while simultaneously measuring low altitudes with adequate power to fill in observations at lower altitudes. It has not been our goal to optimize this design space but rather demonstrate, to our knowledge for the first time, a lidar-based observational system capable of this behavior. Furthermore, we plan future work to update the forward model of Marais and Hayman (2022) to include observations from multiple sensitivity observations to better leverage the information content of the two signals.

## Appendix A: Appendix A
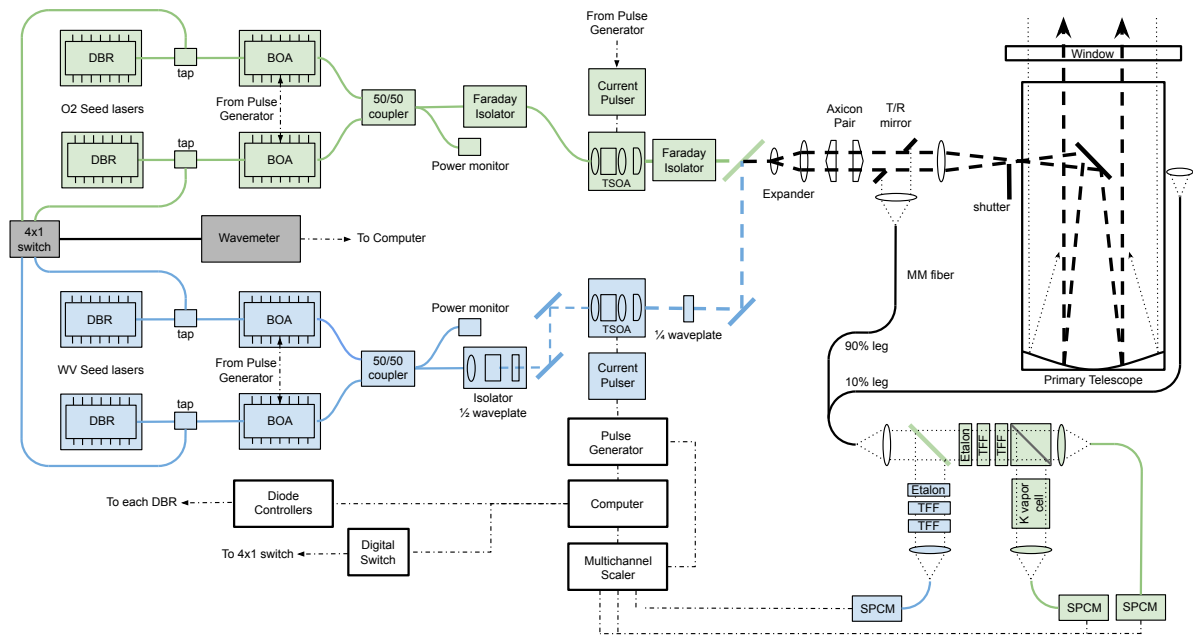
**Figure A1.** Block diagram of the MPD system. Blue components are for water vapor DIAL. Green components are needed for oxygen DIAL and potassium HSRL. Solid lines to the left of the TSOAs are polarization maintaining single mode fiber. Dashed lines are free space beam propogation. Acronyms: DBR: Distributed Bragg Reflector, BOA: Booster Optical Amplifier, TSOA: Tapered Semiconductor Optical Amplifier, TFF: Thin Film Filter, SPCM: Single Photon Counting Module, T/R: Transmit/Receiver, MM: Multi-Mode. The Pulse Generator shown is the master clock that has been altered for this work.

265   *Data availability.*   MPD data is cataloged in the NSF NCAR Earth Observing Laboratory's Field Data Archive with the following URL and digital object identifier: https://data.eol.ucar.edu/dataset/100.034 and https://doi.org/10.26023/MX0D-Z722-M406. (NCAR/EOL-MPD-Team, 2020)

*Author contributions.*   RAS, MH, and SS collectively conceived the presented study. AK designed and developed the FPGA-based timing unit. RAS preformed formal analysis with guidance from MH and SMS. RAS drafted the manuscript with input from all authors.

270   *Competing interests.*   The authors declare that they have no conflict of interest.

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# References

275    Cadeddu, M. P., Liljegren, J. C., and Turner, D. D.: The Atmospheric radiation measurement (ARM) program network of microwave radiometers: instrumentation, data, and retrievals, Atmospheric Measurement Techniques, 6, 2359–2372, https://doi.org/10.5194/amt-6-2359-2013, 2013.

Carroll, B. J., Nehrir, A. R., Kooi, S. A., Collins, J. E., Barton-Grimley, R. A., Notari, A., Harper, D. B., and Lee, J.: Differential absorption lidar measurements of water vapor by the High Altitude Lidar Observatory (HALO): retrieval framework and first results, Atmospheric

280    Measurement Techniques, 15, 605–626, https://doi.org/10.5194/amt-15-605-2022, 2022.

Eixmann, R., Gerding, M., Höffner, J., and Kopp, M.: Lidars With Narrow FOV for Daylight Measurements, IEEE Transactions on Geoscience and Remote Sensing, 53, 4548–4553, 2015.

Goldsmith, J. E. M., Blair, F. H., Bisson, S. E., and Turner, D. D.: Turn-Key Raman Lidar for Profiling Atmospheric Water Vapor, Clouds, and Aerosols, Appl. Opt., 37, 4979–4990, https://doi.org/10.1364/AO.37.004979, 1998.

285    Gordon, I. E., Rothman, L. S., Hargreaves, R. J., Hashemi, R., Karlovets, E. V., Skinner, F. M., Conway, E. K., Hill, C., Kochanov, R. V., Tan, Y., Wcisło, P., Finenko, A. A., Nelson, K., Bernath, P. F., Birk, M., Boudon, V., Campargue, A., Chance, K. V., Coustenis, A., Drouin, B. J., Flaud, J.-M., Gamache, R. R., Hodges, J. T., Jacquemart, D., Mlawer, E. J., Nikitin, A. V., Perevalov, V. I., Rotger, M., Tennyson, J., Toon, G. C., Tran, H., Tyuterev, V. G., Adkins, E. M., Baker, A., Barbe, A., Canè, E., Császár, A. G., Dudaryonok, A., Egorov, O., Fleisher, A. J., Fleurbaey, H., Foltynowicz, A., Furtenbacher, T., Harrison, J. J., Hartmann, J.-M., Horneman, V.-M., Huang, X., Karman, T., Karns,

290    J., Kassi, S., Kleiner, I., Kofman, V., Kwabia-Tchana, F., Lavrentieva, N. N., Lee, T. J., Long, D. A., Lukashevskaya, A. A., Lyulin, O. M., Makhnev, V. Y., Matt, W., Massie, S. T., Melosso, M., Mikhailenko, S. N., Mondelain, D., Müller, H. S. P., Naumenko, O. V., Perrin, A., Polyansky, O. L., Raddaoui, E., Raston, P. L., Reed, Z. D., Rey, M., Richard, C., Tóbiás, R., Sadiek, I., Schwenke, D. W., Starikova, E., Sung, K., Tamassia, F., Tashkun, S. A., Vander Auwera, J., Vasilenko, I. A., Vigasin, A. A., Villanueva, G. L., Vispoel, B., Wagner, G., Yachmenev, A., and Yurchenko, S. N.: The HITRAN2020 molecular spectroscopic database, Journal of Quantitative Spectroscopy and

295    Radiative Transfer, 277, 107 949, https://doi.org/https://doi.org/10.1016/j.jqsrt.2021.107949, 2022.

Hayman, M. and Spuler, S. M.: Demonstration of a diode-laser-based high spectral resolution lidar (HSRL) for quantitative profiling of clouds and aerosols, Optics Express, 25, A1096–A1110, https://doi.org/10.1364/OE.25.0A1096, 2017.

Hayman, M., Stillwell, R. A., and Spuler, S. M.: Fast computation of absorption spectra for lidar data processing using principal component analysis, Opt. Lett., 44, 1900–1903, https://doi.org/10.1364/OL.44.001900, 2019.

300    Hayman, M., Stillwell, R. A., and Spuler, S. M.: Optimization of linear signal processing in photon counting lidar using Poisson thinning, Optics Letters, 45, 5213–5216, https://doi.org/10.1364/OL.396498, 2020.

Hayman, M., Stillwell, R. A., Karboski, A., Marais, W. J., and Spuler, S. M.: Global estimation of range resolved thermodynamic profiles from micropulse differential absorption lidar, Opt. Express, 32, 14 442–14 460, https://doi.org/10.1364/OE.521178, 2024.

Hooker, S. and Webb, C.: Laser Physics, Oxford University Press, 2010.

305    Löhnert, U., Turner, D. D., and Crewell, S.: Ground-Based Temperature and Humidity Profiling Using Spectral Infrared and Microwave Observations. Part I: Simulated Retrieval Performance in Clear-Sky Conditions, Journal of Applied Meteorology and Climatology, 48, 1017 – 1032, https://doi.org/https://doi.org/10.1175/2008JAMC2060.1, 2009.

Lowe, P. R.: An Approximating Polynomial for Computation of Saturation Vapor Pressure, Journal of Applied Meteorology, 16, 100–103, 1977.

310   Marais, W. J. and Hayman, M.: Extending water vapor measurement capability of photon-limited differential absorption lidars through simultaneous denoising and inversion, Atmospheric Measurement Techniques, 15, 5159–5180, https://doi.org/10.5194/amt-15-5159-2022, 2022.

McCullough, E. M., Drummond, J. R., and Duck, T. J.: Lidar measurements of thin laminations within Arctic clouds, Atmospheric Chemistry and Physics, 19, 4595–4614, https://doi.org/10.5194/acp-19-4595-2019, 2019.

315   Mudukutore, A. S., Chandrasekar, V., and Keeler, R. J.: Pulse compression for weather radars, IEEE Transactions on Geoscience and Remote Sensing, 36, 125–142, https://doi.org/10.1109/36.655323, 1998.

NCAR/EOL-MPD-Team: NCAR MPD data. Version 1.0 (Version 1.0 ) [Data set]., https://doi.org/10.26023/MX0D-Z722-M406, 2020.

Nehrir, A., Repasky, K., and Carlsten, J.: Eye-Safe Diode-Laser-Based Micropulse Differential Absorption Lidar (DIAL) for Water Vapor Profiling in the Lower Troposphere, Journal of Atmospheric and Oceanic Technology - J ATMOS OCEAN TECHNOL, 28, 131–147,
320   https://doi.org/10.1175/2010JTECHA1452.1, 2011.

Newsom, R. K., Turner, D. D., Mielke, B., Clayton, M., Ferrare, R., and Sivaraman, C.: Simultaneous analog and photon counting detection for Raman lidar, Appl. Opt., 48, 3903–3914, https://doi.org/10.1364/AO.48.003903, 2009.

Newsom, R. K., Turner, D. D., Lehtinen, R., Münkel, C., Kallio, J., and Roininen, R.: Evaluation of a Compact Broadband Differential Absorption Lidar for Routine Water Vapor Profiling in the Atmospheric Boundary Layer, Journal of Atmospheric and Oceanic Technology,
325   37, 47–65, https://doi.org/10.1175/JTECH-D-18-0102.1, 2020.

O'Hora, F. and Bech, J.: Improving weather radar observations using pulse-compression techniques, Meteorological Applications, 14, 389–401, https://doi.org/https://doi.org/10.1002/met.38, 2007.

Ramp, H. O. and Wingrove, E. R.: Principles of pulse compression, IRE Transactions on Military Electronics, pp. 109–116, 1961.

Reichardt, J., Wandinger, U., Klein, V., Mattis, I., Hilber, B., and Begbie, R.: RAMSES: German Meteorological Service autonomous Raman
330   lidar for water vapor, temperature, aerosol, and cloud measurements, Appl. Opt., 51, 8111–8131, https://doi.org/10.1364/AO.51.008111, 2012.

Repasky, K. S., Moen, D., Spuler, S., Nehrir, A. R., and Carlsten, J. L.: Progress towards an Autonomous Field Deployable Diode-Laser-Based Differential Absorption Lidar (DIAL) for Profiling Water Vapor in the Lower Troposphere, Remote Sensing, 5, 6241–6259, https://doi.org/10.3390/rs5126241, 2013.

335   Richards, M. A., Scheer, J. A., and Holm, W. A.: Principles of Modern Radar Vol. 1: Basic Principles, SciTech Publishing, 2010.

Spuler, S. M., Repasky, K. S., Morley, B., Moen, D., Hayman, M., and Nehrir, A. R.: Field-deployable diode-laser-based differential absorption lidar (DIAL) for profiling water vapor, Atmospheric Measurement Techniques, 8, 1073–1087, 2015.

Spuler, S. M., Hayman, M., Stillwell, R. A., Carnes, J., Bernatsky, T., and Repasky, K. S.: MicroPulse DIAL (MPD) – a diode-laser-based lidar architecture for quantitative atmospheric profiling, Atmospheric Measurement Techniques, 14, 4593–4616, https://doi.org/10.5194/amt-
340   14-4593-2021, 2021.

Stillwell, R. A., Spuler, S. M., Hayman, M., Repasky, K. S., and Bunn, C. E.: Demonstration of a combined differential absorption and high spectral resolution lidar for profiling atmospheric temperature, Optics Express, 28, 71–93, https://doi.org/10.1364/OE.379804, 2020.

Strotkamp, M., Munk, A., Jungbluth, B., Hoffmann, H.-D., and Höffner, J.: Diode-pumped Alexandrite laser for next generation satellite-based earth observation lidar, CEAS Space Journal, 11, 413–422, https://doi.org/10.1007/s12567-019-00253-z, 2019.

345   Thayer, J. P., Nielson, N. B., Warren, R. E., and Heinselman, C. J.: Rayleigh lidar system for middle atmosphere research in the arctic, Optical Engineering, 36, 2045–2061, 1997.

Vivekanandan, J., Ellis, S., Tsai, P., Loew, E., Lee, W.-C., Emmett, J., Dixon, M., Burghart, C., and Rauenbuehler, S.: A wing pod-based millimeter wavelength airborne cloud radar, Geoscientific Instrumentation, Methods and Data Systems, 4, 161–176, https://doi.org/10.5194/gi-4-161-2015, 2015.

350  von Zahn, U., von Cossart, G., Fiedler, J., Fricke, K. H., Nelke, G., Baumgarten, G., Rees, D., Hauchecorne, A., and Adolfsen, K.: The ALOMAR Rayleigh/Mie/Raman lidar: objectives, configuration, and performance, Annales Geophysicae, pp. 815–833, 2000.

Weckwerth, T. M., Weber, K. J., Turner, D. D., and Spuler, S. M.: Validation of a Water Vapor Micropulse Differential Absorption Lidar (DIAL), Journal of Atmospheric and Oceanic Technology, 33, 2353–2372, https://doi.org/10.1175/JTECH-D-16-0119.1, 2016.

Wulfmeyer, V.: Ground-based differential absorption lidar for water-vapor and temperature profiling: development and specifications of a
355  high-performance laser transmitter, Appl. Opt., 37, 3804–3824, https://doi.org/10.1364/AO.37.003804, 1998.