

rsfun v5.0.1: A model-data integration framework for simulating ecosystem processes

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Abstract.

Mechanistic vegetation models serve to estimate terrestrial carbon fluxes and climate impacts on ecosystems across diverse biotic and abiotic conditions. Systematically informing them with data is key for enhancing their predictive accuracy and estimating uncertainty. Here ~~we present~~, we demonstrate and evaluate the Simulating Optimal FUNctioning ~~{rsfun}~~ (rsfun) R package, providing which provides a computationally efficient ~~and parallelizable~~ implementation of the P-model for site-scale simulations of ecosystem photosynthesis and the acclimation of photosynthetic traits, complemented with functionalities for Bayesian model-data integration and ~~estimation of parameters and the estimation of model parameters and prediction~~ uncertainty. We ~~describe a use case to demonstrate the package functionalities for modelling ecosystem gross uptake at one flux measurement site, including model sensitivity analysis, Bayesian parameter calibration, and prediction uncertainty estimation.~~ {rsfun} lowers the bar of entry to ecosystem modelling and model-data integration and serves as an open-access resource for model development and dissemination. estimated model parameters simultaneously from observed time series of ecosystem gross primary productivity (GPP), and from globally distributed data on leaf carbon-13 isotopic discrimination ($\Delta^{13}\text{C}$) and the ratio of the maximum biochemical rates of carboxylation to electron transport ($V_{\text{cmax}}/J_{\text{max}}$). The multi-target calibration yielded unbiased predictions for all variables simultaneously and produced similar distributions of prediction-observation residuals for both calibration and out-of-sample test data, indicating that the model generalises robustly across diverse environments. We found that a step-wise approach to successive model integration and calibration yielded best results, and that correlations among parameters related to representing water stress effects underpinned non-robust parameter estimations. This likely indicates a dominant source of model structural uncertainty related to the representation of the response of photosynthesis to dry conditions in soil and air.

20 1 Introduction

The modelling of land ecosystem processes and structure, water, and carbon fluxes relies on both mechanistic and statistical approaches (~~Dietze et al., 2018; Hartig et al., 2012; Van Oijen et al., 2005~~)(Dietze et al., 2018; Hartig et al., 2012; Van Oijen et al., 2005). Mechanistic models are formulated as mathematical descriptions of functional relationships between the abiotic environment

and ecosystem states, rates, and dynamics. These descriptions reflect available theory and general empirical patterns and provide a means for "translating" hypotheses about governing principles and causal relationships into testable predictions ([Marquet et al., 2014](#)), and for upscaling model-based estimates in geographical space and to novel environmental conditions. However, mechanistic models rely also on ~~statistical (empirical) empirical~~ descriptions of processes at varying levels of abstraction.

Mechanistic models have model parameters that are either specified ~~directly~~ or fitted to data. A great advantage of mechanistic models is that they explicitly link known physical constants with process representations (e.g., molecular mass of CO₂ for diffusion and assimilation, or the gravitational constant and viscosity of water for its transport and transpiration). Other parameters may be specified based on independent measurements under controlled conditions (e.g., the activation energy of Arrhenius-type metabolic rates), or represent measurable plant functional traits, taken as constant over time and within plant functional types (PFTs - the basic unit in mechanistic vegetation models). Both types of parameters have traditionally been directly specified in models ~~directly~~ ('direct parameterization', ([Hartig et al., 2012](#))([Hartig et al., 2012](#))). Yet other model parameters may not be directly observable and describe processes that are not explicitly resolved but can be described at a higher level of abstraction. Such parameters are often fitted to observational data such that the agreement between one (or several) related model predictions and observations is optimised. Parameter estimation for mechanistic vegetation models typically employs generic optimization algorithms or Bayesian statistical approaches and is often used ~~for specifying to specify~~ diverse types of parameters (except for universal physical constants). Bayesian methods have the advantage that they enable a systematic assessment of the correlation structure among multiple fitted parameters, provide a means ~~for considering uncertainty in observations to consider uncertainty in inputs, observations, models,~~ and available prior information, generate probabilistic parameter estimations and model predictions, and provide a basis to ~~quantifying quantify~~ the constraints by various ~~of~~ calibration target data ~~or to identifying and to identify~~ errors arising from model structural choices ([Bagnara et al., 2015](#); [Dietze et al., 2018](#); [Hartig et al., 2012](#); [van Oijen et al., 2013](#); [Raj et al., 2018](#); [Van Oijen et al., 2005](#); [Xiao et al., 2019](#)) ([Bagnara et al., 2015](#); [Dietze et al., 2018](#); [Hartig et al., 2012](#); [van Oijen, 2017](#); [Raj et al., 2016](#); [Van Oijen et al., 2005](#); [Xiao et al., 2019](#)).

As the number of parameters increases in state-of-the-art mechanistic vegetation models, taking into account multiple PFTs and ecosystem components (e.g. soil, microbes, hydrology), larger amounts of data and computing resources are required to fully explore the parameter space ([Hartig et al., 2012](#))([Hartig et al., 2012](#)). This poses a limitation for systematic model-data integration and Bayesian parameter estimation. Eco-evolutionary optimality (EEO) principles have been proposed for reducing model complexity and for a robust grounding of models in governing principles ([Franklin et al., 2020](#); [Harrison et al., 2021](#)) ([Franklin et al., 2020](#); [Harrison et al., 2021](#)). They enable parameter-sparse representations, limit the distinction of separate PFTs, and may enable better model generalisations to novel environmental regimes. As such EEO principles make predictions of plant functional traits that would otherwise have to be prescribed – typically as temporally fixed model parameters. ~~However, although parameter-sparse, EEO-based vegetation models are not devoid of model parameters~~ Parameters in EEO models are considered to be universally valid, e.g., across different PFTs. Ideally, ~~remaining parameters they~~ represent known physical constants or quantities that can be measured independently. ~~But remaining parameters in optimality models typically also~~

represent quantities that are not directly measurable. However, not all can be measured directly – e.g., the marginal cost of water
60 in Cowan and Farquhar (1977) Cowan and Farquhar (1977), or the unit cost ratio in Prentice et al. (2014). These are considered
to be more universally valid (e.g., without distinctions between PFTs), but still must be Prentice et al. (2014) and need to be
fitted to data.

The P-model (Prentice et al., 2014; Stocker et al., 2020; Wang et al., 2017) (Prentice et al., 2014; Wang et al., 2017; Stocker et al., 2020
is an example of an EEO-guided model for terrestrial photosynthesis and its acclimation. It avoids the requirement for prescrib-
65 ing PFT-specific parameters of photosynthesis and stomatal regulation but instead predicts them from universal EEO principles
for the full range of environmental conditions across the Earth's (C_3 photosynthesis-dominated) biomes. However, not directly
observable parameters of the underlying Through its representation of the vegetation as a *big leaf* (ref), it represents the scaling
between leaf traits and ecosystem-level photosynthetic CO_2 uptake (gross primary productivity, GPP). However, despite its
foundation on EEO theory and of additional empirical parameterizations employed in the P-model (Stocker et al., 2020) remain
70 resulting parameter-sparseness, a small set of model parameters remains (Stocker et al., 2020) (Table 2) and must be specified
or fitted to data (Tab. 2).

Here, we provide a solution for this challenge, acknowledging that the data is an integral part of the modelling process
(Dietze et al., 2013) – even in theory-based models of ecosystem processes. We show how the most important parameters
contributing to uncertainty in the explore the respective constraints provided by ecosystem-level fluxes and leaf-level traits for
75 a probabilistic (Bayesian) multi-target estimation of the P-model parameters. The selection of observational target data types is
motivated by their known effectiveness in model calibration and parameter estimation from previous work (Prentice et al., 2014; Wang et al.
. Specifically, we use observations-based GPP time series from multiple eddy covariance measurement sites, and compilations
of globally distributed measurements of leaf traits, including the leaf carbon isotopic fractionation relative to the atmosphere
($\Delta^{13}C$, hereafter shortened to Δ), and the ratio of the maximum biochemical rate of carboxylation to electron transport
80 ($V_{cmax}/J_{max} \equiv VJ$). Data from these two leaf traits have previously been used for independently estimating separate model
parameters in the P-model (Wang et al., 2017). In Stocker et al. (2020), these independently estimated model parameters were
then specified for model simulations (direct parametrization), and were not subject to model parameter calibration. Here,
we demonstrate how the combined consideration of diverse data types within a Bayesian model-data integration framework
– combining ecosystem flux data and leaf traits data – enables the simultaneous estimation of a comprehensive set model
85 parameters that control functional dependencies of processes at multiple organisational levels – from the leaf to the canopy.
This enables a better understanding of interdependencies between model parameters and a more reliable estimation of model
prediction uncertainty.

Unbalanced observations of multiple calibration targets can lead to parameter estimates that compensate structural errors in
the (nonlinear) model, as shown previously with synthetic data (MacBean et al., 2016; Oberpriller et al., 2021; Cameron et al., 2022)
90 . We are therefore interested in the questions i) if the P-model can be estimated using observations of ecosystem-level
photosynthetic uptake (gross primary productivity, GPP). We provide calibrated in a consistent manner to these targets resulting
in unbiased parameter estimates (relative to expected ranges based on our process-based interpretation of each parameter), and
ii) if the calibrated model can be validated on an independent test data set.

We start by providing a brief description of the ~~theory embodied in the~~ P-model and introduce the P-model implementation in the Simulating Optimal FUNctioning ~~{rsofun}~~ *rsofun* version v5.0-1 modelling framework, made available as an R package (Stocker et al., 2025). A more comprehensive ~~P-model description~~ description of the P-model can be found in ~~Stocker et al. (2020).~~ Stocker et al. (2020). *rsofun* implements the P-model and its connection to data through Bayesian parameter optimisation and analysis. We demonstrate the functionalities ~~of {rsofun} through a case study for simulating GPP at a single flux measurement site. The case study includes sensitivity analysis~~ implemented in and accessible through *rsofun* with different calibration setups – i.e., different combinations of model parameters subjected to Bayesian calibration and different observational data types. Alternative calibration setups serve to elucidate the role of different observations in constraining estimates for different model parameters. *rsofun* functionalities demonstrated here include the flexible specification of likelihood functions for connecting model predictions to specific data types, Bayesian model calibration using Monte Carlo Markov Chain sampling, the analysis of posterior distributions of estimated parameters, and the estimation of prediction uncertainty, probabilistic predictions of GPP, and inference – the prediction of GPP with an estimation of its uncertainty. This paper presents the calibration to GPP observations only, but the package allows calibration to multiple targets simultaneously, including fluxes and leaf traits V_J , and Δ , and their evaluation against out-of-sample test data.

2 P-model description ~~Methods~~

2.1 P-model description

The P-model predicts the acclimation of leaf-level photosynthesis to a (slowly varying) environment based on EEO principles. It thereby yields a parameter-sparse representation of ecosystem-level quantities, generalising across (C_3 photosynthesis-dominated) vegetation types and biomes. The P-model combines established theory for C_3 photosynthesis following the Farquhar-von Caemmerer-Berry (FvCB, ~~(Farquhar et al., 1980)) model~~) model (Farquhar et al., 1980) with the Least-Cost hypothesis for the optimal balancing of water loss and carbon gain (Prentice et al., 2014)(Prentice et al., 2014), and the coordination hypothesis (Wang et al., 2017)(Wang et al., 2017), which states that the light and Rubisco-limited assimilation rates (as described by the FvCB model) are equal for representative daytime environmental conditions. ~~Based on these theoretical foundations, gross primary productivity (GPP) can be modelled as the product of absorbed photosynthetically active radiation, specified by the locally measured photosynthetically active radiation (PAR) and remotely sensed fAPAR,~~

The theory results in a prediction of the ratio of leaf-internal to ambient CO_2 concentration ($c_i : c_a \equiv \chi$) as a function of the atmospheric environment, characterised by the following meteorological variables: daytime mean air temperature T ($^{\circ}C$), daytime mean vapor pressure deficit D (Pa), and the atmospheric CO_2 partial pressure (c_a , Pa).

$$\chi = \frac{\Gamma^*}{c_a} + \left(1 - \frac{\Gamma^*}{c_a}\right) \frac{\xi}{\xi + \sqrt{D}} \quad (1)$$

$$\xi = \sqrt{\frac{\beta(K + \Gamma^*)}{1.6 \eta^*}} \quad (2)$$

Γ^* is the photorespiratory compensation point (Pa). η^* is the ratio the (temperature-dependent) viscosity of water, relative to its value at 25°C, and the theory-based prediction of the ecosystem-level light use efficiency (LUE) (Stocker et al., 2020; Wang et al., 2017). LUE acclimates to preceding environmental conditions with a characteristic, empirically determined (calibrated), time scale τ . Two latent (not directly observable) parameters govern the optimality-guided water-carbon trade-off: the K is the Michaelis-Menten coefficient for photosynthesis (Pa). β (unitless) is the unit cost ratio β (governing the balancing of maintaining the carboxylation capacity versus the transpiration stream) and c^* (the marginal cost of maintaining the electron transport rate). These have previously been calibrated separately to data and specified as fixed model parameters in the P-model (Stocker et al., 2020; Wang et al., 2017). The theory for predicting acclimated LUE then requires only the atmospheric environment to be specified (meteorological variables and), of carboxylation to transpiration in the EEO framework of Prentice et al. (2014), and is calibrated to data here (see Tab. 1). The functional dependency of Γ^* on temperature and atmospheric pressure, the dependency of η^* on temperature, and the dependency of K on temperature and atmospheric pressure are described in detail in Stocker et al. (2020) based on published work (Farquhar et al., 1980; Bernacchi et al., 2001; Huber et al., 2009). Involved parameters are held fixed here and are not calibrated.

A set of corollary predictions, physically and physiologically consistent with the simulated LUE predicted χ , follows. These include the acclimated base rates of photosynthetic capacities in the FvCB-model ($V_{\text{cmax}25}$ and $J_{\text{max}25}$), the acclimated average ratio of leaf internal to ambient. The following predicted quantities are used for model-data integration here. A complete description of the mathematical derivation of these quantities from first principles is given in (Stocker et al., 2020).

2.1.1 Isotope fractionation by photosynthesis

χ directly controls isotopic discrimination of carbon assimilates (Δ) relative to the atmospheric signature ($\delta^{13}\text{C}_a$) (Farquhar et al., 1989, 1991).

$$\Delta = a_{\Delta} + (b_{\Delta} - a_{\Delta}) \chi - f_{\Delta} \frac{\Gamma^*}{c_a}, \quad (3)$$

Here, parameters represent the isotope fractionation from CO_2 concentration ($c_i : c_a$), acclimated average daytime stomatal conductance (g_s diffusion in air ($a_{\Delta} = 4.4\%$, Craig (1953))), from Rubisco carboxylation ($b_{\Delta} = 27\%$), and the acclimated base rate of leaf dark respiration (R_{d25}). Physical constants and additional parameters that determine the instantaneous temperature dependence of V_{cmax} , J_{max} , R_d , and parameters in the FvCB-model are prescribed and held fixed in the P-model (Tab. A2 in (Stocker et al., 2020)). from photorespiration ($f_{\Delta} = 8.0\%$, Ubierna and Farquhar (2014)). Also these parameters were held fixed here and not subjected to calibration.

For simulating GPP, the P-model is conceived as a single-big-leaf model (Fig. 1). While describing leaf-level quantities at relatively high mechanistic detail, the link between the leaf and the canopy scale is not explicitly resolved. Instead, an empirical approach for leaf-to-canopy scaling of uptake is employed by treating the quantum yield parameter φ_0 to be representative for

2.1.2 Maximum rates of carboxylation and electron transport

155 ~~For daytime conditions, averaged over multiple days, the canopy-scale and allowing it to be calibrated to ecosystem-level flux data. The implementation of the P-model in the {rsfun} package (version v5.0) further includes an empirical parameterization of the temperature dependency of assumes the Rubisco carboxylation-limited and the quantum yield φ_0 , generalising the approach taken in Stoeker et al. (2020), electron transport-limited rates of photosynthesis to be equal:~~

$$A_C = A_J \quad (4)$$

160 ~~Following the Farquhar-von Caemmerer-Berry (FvCB) model for C_3 photosynthesis (Farquhar et al., 1980; von Caemmerer and Farquhar, the maximum rate of carboxylation V_{cmax} can thus be expressed as~~

$$V_{\text{cmax}} = \varphi_0 I_{\text{abs}} \frac{m'}{m_C}, \quad (5)$$

~~with~~

$$m_C = \frac{c_i - \Gamma^*}{c_i + K}, \quad (6)$$

165 ~~and with~~

$$m' = m \sqrt{1 - \left(\frac{c^*}{m}\right)^{2/3}} \quad (7)$$

~~and an empirical~~

$$m = \frac{c_i - \Gamma^*}{c_i + 2\Gamma^*}. \quad (8)$$

~~Here, φ_0 is the intrinsic quantum yield of photosystem II (mol mol^{-1}) which depends on the leaf temperature T (see Eq. 14) – here taken as identical to air temperature. c^* (unitless) is the unit cost of electron transport and is treated as a calibratable model parameter (see Table 2 for an overview of calibrated model parameters). I_{abs} is the photosynthetic photon flux density absorbed by the leaf.~~

~~Eq. 7 accounts for a limited electron transport capacity (J_{max}) such that m' can also be written as~~

$$m' = m \frac{1}{\sqrt{1 + \left(\frac{4 \varphi_0 I_{\text{abs}}}{J_{\text{max}}}\right)^2}} \quad (9)$$

175 ~~Again using $A_C = A_J$, J_{max} can be solved for and can be expressed as~~

$$J_{\text{max}} = \frac{4 \varphi_0 I_{\text{abs}}}{\sqrt{L^{-2} - 1}}, \quad (10)$$

~~with~~

$$L = V_{\text{cmax}} \frac{c_i + 2\Gamma^*}{\varphi_0 I_{\text{abs}}(c_i + K)}. \quad (11)$$

~~The ratio $VJ = V_{\text{cmax}}/J_{\text{max}}$ is finally calculated by dividing respective values obtained with Eqs. 5 and 10.~~

180 2.1.3 Gross primary productivity

Gross primary productivity (GPP) can be expressed in the form of a light use efficiency model (Prentice et al., 2024; Bao et al., 2022; Monte

$$\text{GPP} = \text{fAPAR} \cdot \text{PPFD} \cdot \text{LUE}, \quad (12)$$

185 with fAPAR being the fraction of absorbed photosynthetically active radiation (unitless), PPFD being the photosynthetic photon flux density, PPFD ($\text{mol s}^{-1} \text{ m}^{-2}$), and LUE (g C mol^{-1}) being the light use efficiency, calculated as

$$\text{LUE} = \varphi_0(T) f_\beta(\theta) M_C m'. \quad (13)$$

Here, f_β is the unitless soil moisture stress function, ~~generalising the approach taken in Stocker et al. (2020) (Appendix A):~~
~~Acclimating quantities are derived by~~ varying between 0 and 1 (see Eq. 15), with θ representing the plant-available soil water
 content in mm. M_C is the molar mass of C ($12.0107 \text{ g mol}^{-1}$). Note that the application of Eq. 12 assumes GPP to scale
 190 linearly with absorbed light. This functional relationship is assumed here to describe the relationship between multi-day sums
 of GPP and PAR and emerges from the assumption of the Coordination Hypothesis (Wang et al., 2017). However, it cannot be
 expected to describe the functional dependencies at shorter time scales, where the limitation by the electron transport capacity
 (J_{\max}) becomes effective (Mengoli et al., 2022; Farquhar et al., 1980).

Such acclimation is considered by employing the P-model theory to gradually varying environmental conditions (P, D, CO_2 ,
 195 and PAR) where variations are damped ~~and lagged by a characteristic (calibrated) by a low-pass filter with a characteristic,~~
~~empirically determined~~ time scale τ . ~~The continuous treatment of the acclimation time scale is different from Stocker et al. (2020)~~
~~, where monthly mean values of environmental variables were considered for the acclimation, (days) (see equation B1 for the~~
~~definition of the low-pass filter).~~

2.1.4 Quantum yield efficiency

200 The temperature dependency of the quantum yield efficiency φ_0 ~~can be turned off by setting a (see Tab. 2 and Appendix A) to~~
~~0 ('ORG' model setup in (Stocker et al., 2020)) is empirically parametrized as:~~

$$\varphi_0(T) = \varphi_0^* (1 + a_\varphi (T - b_\varphi)^2), \quad \text{bounded to the range } [0, \varphi_0^*]. \quad (14)$$

This is in contrast to the formulation used in Stocker et al. (2020), where a_φ and b_φ were effectively prescribed and not
 subjected to calibration.

205 2.1.5 Soil moisture stress

Soil moisture stress is computed as

$$f_\beta(\theta) = \begin{cases} 1 - \frac{(\theta - \theta^*)^2}{\theta^{*2}} & \text{if } x \geq 0, \\ 1 & \text{if } x < 0. \end{cases} \quad (15)$$

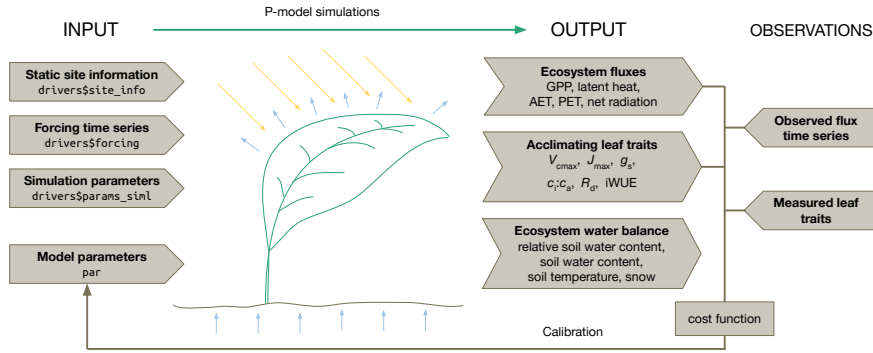


Figure 1. P-model inputs, outputs, and target observations for the calibration parameter estimation. The model takes as inputs static site information (longitude, latitude, elevation and root zone water storage capacity), time series of meteorological forcings (listed in Tab. 1), simulation parameters (spinup years, recycle period length, vegetation type (evergreen, deciduous, grass (/) with/without N-fixation)) and ; common model parameters for all sites, model parameters (listed in Tab. 2). The simulation returns a time series of several ecosystem fluxes, acclimating leaf traits and ecosystem water balance quantities states. By comparing these outputs to field measurements with measured traits and flux data in a Bayesian calibration routine, model parameters can be estimated.

where θ^* (mm) represents the threshold below which GPP is reduced. The soil moisture stress function can be turned off by setting β_0 (see Tab. 2 and Appendix A) to 1 ('BRC' model setup in (Stocker et al., 2020)). Taken together, the P-model approach is guided by EEO and thus yields predictions for quantities that otherwise must be prescribed and fitted to data. Yet, a small set of model parameters, related to empirical parametrizations and latent quantities remain and must be calibrated to data. See Tab. 2 for a list of all (calibratable) model parameters of the P-model implementation in the {rsfun} package water balance is simulated with a bucket model of soil water content (uniform total water storage capacity), whereby the dynamics of θ (mm) considers daily precipitation, the snow melt, a Priestley-Taylor-based evapotranspiration estimate, and runoff when the bucket is full. Except for implicitly enforcing $f_\beta(\theta = 0) = 0$ here, this formulation follows the description in Stocker et al. (2020) and is based on the SPLASH model (Davis et al., 2017).

2.2 Calibration setups

Different calibration setups were used to illustrate the constraints imposed by the different calibration targets for the same task of simultaneously fitting all parameters listed in Tab. 1.

Calibratable parameters in the {rsfun} The choice of calibration setups was guided by our expectation of constraints contained in the different target data sets and their combinations, given the model structure. Using common priors across (most) setups, the resulting parameter estimates can illustrate the constraints imposed by data alone. This approach can be seen as an extreme form of the approach suggested by Cameron et al. (2022), where they increased the number of observations of certain target variables to identify model structural errors. Here, the setups systematically tested each trait target individually:

Table 1. Forcing time series, Daily time series Overview of targets used and parameters estimatable in the following meteorological variables are required for simulations six calibration setups. If a spin-up period is specified, For parameter description and units see Table 2. Columns with c) mean the corresponding years are recycled for the spin-up posterior of that setup had been used as (truncated) prior.

Parameter name					
	Description	Data	Units	Model parameters	
temp-Setup	Daytime average air temperature Δ	$^{\circ}\text{C}$ tmin-VJ	Daily minimum air temperature-GPP	$^{\circ}\text{C}$ tmax- φ_0^*	Daily m
a)	$\text{mol m}^{-2} \text{s}^{-1}$				
netrad-b)	Net radiation	W m^{-2}			
patm-c)	Atmospheric pressure	Pa			
rain-d)	Rainfall as precipitation in liquid form	mm s^{-1}			
snow-e)	Snowfall in water equivalents	mm s^{-1}			
fapar-h)	Fraction of photosynthetic active radiation (fAPAR)				

225 Δ in setup a), VJ in setup b), and in combination in setup c); before then considering GPP alone setup d), and combined with the traits setup e). Finally, setup h) tested a stepwise approach using the posteriors from setup c) as prior specification, to inclusion of trait information in the final estimates in spite of potential model misspecifications.

We use the implementation of P-model **implementation** in the *rsosun* framework (Stocker et al., 2025) to model daily gross primary productivity, GPP, and leaf-level traits, namely the ratio of the rates of photosynthetic capacities in the FvCB model, 230 VJ, and the isotopic fractionation of assimilated carbon, Δ . Implementation details of the *rsosun* framework are provided in Appendix C.

Three latent (not directly observable) parameters govern the optimality-guided water-carbon trade-off: the unit cost ratio, β , (governing the balancing of maintaining the carboxylation capacity versus the transpiration stream), the marginal cost of maintaining the electron transport rate, c^* , and the quantum yield efficiency φ_0 (parametrized with φ_0^* , a_{φ} , b_{φ}).

235 The parameters τ , β , c^* have previously been calibrated separately to data and have been been specified as fixed model parameters in the P-model (Stocker et al., 2020; Wang et al., 2017). Here they were instead calibrated simultaneously to multiple calibration targets, together with additional parameters.

For simplicity, the same θ^* , the soil water volume (mm) below which plants are stressed, was used across all sites.

Three different calibration targets were used in this study. Δ represents accumulated assimilates and is influenced by stomatal 240 opening through the leaf-internal to ambient carbon dioxide ratio. Eq. 2 indicates its dependency on the parameter β . VJ is assumed constant throughout the season. Observations of this ratio (Eq. 5 and 10) are expected to inform the model parameters β and c^* . The maximum-a-posteriori GPP(*t*) observations represent daily values of ecosystem-level photosynthetic CO₂ uptake fluxes (with acclimation of LUE(*t*)). Eqs. 12 to 3 illustrate the dependence of these observations to model parameters β , c^* , as well as θ^* , τ , φ_0^* , a_{φ} , and b_{φ} .

245 Note that V_{cmax} and J_{max} both scale linearly with φ_0 (Eq. 5 and 10). This dependency to φ_0 cancels out when considering the ratio of the two VJ.

3 The {rsofun} model framework

2.1 Calibration target and test data

{rsofun} implements the P-model (Stocker et al., 2020) and provides off-the-shelf methods for Bayesian (probabilistic) parameter and prediction uncertainty estimation. {rsofun} is distributed as an R package on R's central and public package repository. {rsofun} also implements the BiomeE vegetation demography model (Weng et al., 2015). The latter is not further described here and is implemented at an experimental stage in {rsofun} version v5.0. The P-model implementation in {rsofun} is designed for time series simulations by accounting for temporal dependencies in the acclimation to a continuously varying environment (Tab. 1). Function wrappers in R make the simulation workflow user-friendly and all functions and input forcing data structures are comprehensively documented (<https://geco-bern.github.io/rsofun/>). The data set consisted of 50 sites with GPP flux time series (172'055 site-dates in total), 49 sites with a total of 597 individual VJ observations (multiple individual plants and/or species per site sampled), and 325 sites with a total of 2357 Δ observations. Data for all variables were split by sites for model calibration (training) and out-of-sample testing (Fig. 2 and listed in Tab. S2 in the supplemental material). The split into training and test data sets was performed in a stratified manner according to vegetation type and land cover class (Beck et al., 2018; Copernicus Climate Change Service, 2019; Hufkens and Stocker, 2025) to ensure balanced representations of each stratum in the test and training data set. For GPP sites, we additionally required that a site contained more than 12 years of good-quality gap-free data to be available as a training site. We used GPP data from 12 sites for training and data from the remainder 38 sites for testing. The VJ and the Δ data were split such that data from roughly 50% of all sites in each stratum were used for training and testing, respectively.

In {rsofun}, model parameters can be calibrated using a calibration function `calib_sofun()`, providing two modes of calibration, one based on generalised simulated annealing ({GenSA} R package) for global optimization (Xiang et al., 2013) and one based on Markov chain Monte Carlo (MCMC) implemented by the {BayesianTools} R package, giving access to a wide variety of Bayesian methods (Hartig et al., 2023). The former being fast, while the latter provides more informed parameter optimization statistics (Clark, 2004; Dietze et al., 2013). This gives the option for both exploratory and more in-depth analysis of estimated parameters. A set of standard cost functions are provided for the calibration, facilitating the exploration of various metrics or target variables and the specification of calibrated model parameters. Furthermore, the vignettes accompanying the package (<https://geco-bern.github.io/rsofun/articles/>) explain how to customise the calibration cost functions and interpret the calibration results.

3 Case study

We use {rsofun} to model GPP as estimated from ecosystem flux measurements taken at one site using the GPP observations were taken from FluxDataKit (FDK v3.4.2) (Hufkens and Stocker, 2025; Pastorello et al., 2020), which combines published releases of consistently processed eddy-covariance technique. The site, selected here for demonstration purposes, is Puéchabon (Rambal et al., 2004), an evergreen Mediterranean forest, dominated by *Quercus ilex*, growing on

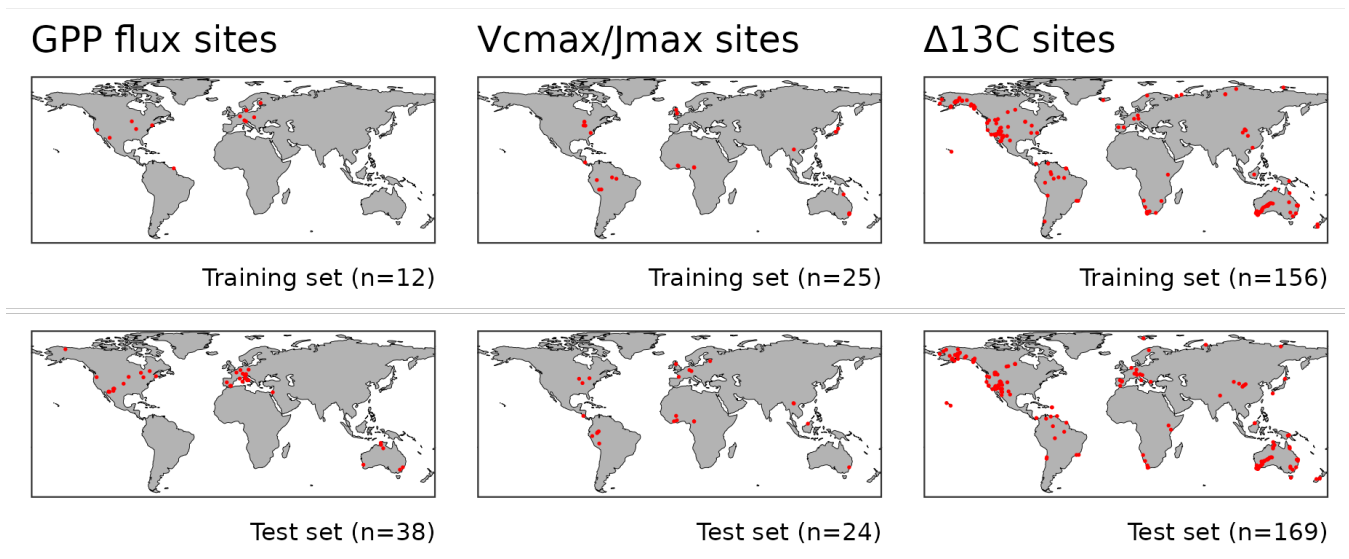


Figure 2. [Global maps of site with observations of the three target types in the training and test data sets.](#)

relatively shallow soil on karstic bedrock, and located in southern France. The climate is governed by a distinct seasonality in solar radiation and temperature, peaking in summer, and a dry period during summer months with recurrent ecosystem water limitation in late summer (Rambal et al., 2004). Data used here cover years 2007–2012. Both the model target data (GPP) and model forcing data (see Fig. 1) were measured at the site and obtained from the FLUXNET2015 dataset (Pastorello et al., 2020). Data for years 2013 and 2014 were available from FLUXNET2015 but removed here due to unexpectedly low GPP measurements (*pers. comm.* Jean-Marc Limousin). GPP data is taken as the FLUXNET standard variable (data from multiple regional networks (Ukkola, 2020; Warm Winter 2020 Team et al., 2022; Drought 2018 Team et al., 2020). GPP estimated through the nighttime flux partitioning was used (variable 'GPP_NT_VUT_REF') derived from the nighttime flux partitioning method (Reichstein et al., 2005) and was filtered (Reichstein et al., 2005). For model calibration and evaluation, we filtered data to retain only (daily) values that were derived from computed with at least 80% measured or good-quality, gap-filled (half-hourly) values. Sites, classified as croplands or wetlands were removed, as well as any site with five or less complete years of data. Based on visual inspection, the following site-year combinations exhibited spurious patterns and were additionally removed: ES-LJu, year 2006; US-Ho2, year 2007; CH-Dav, year 2010; US-Whs years 2016 and 2017.

2.1 Sensitivity analysis

Observations of VJ were obtained from the data compilation of Smith et al. (2019), containing data reported for top canopy from multiple sources (De Kauwe et al., 2016; Keenan and Niinemets, 2016; Smith and Dukes, 2017; Kattge et al., 2011; Wang et al., 201

The Observations of carbon-13 isotope discrimination in leaf material Δ were taken from a global data set (Cornwell et al., 2018; Cornwell et al., 2019), subsetting only observations that were marked as C3 plants. We used Δ values that were derived from the isotopic signature of leaf material in relation to the atmospheric signature at the date and latitude measurements were made.

2.1 Forcing data

For simulations of GPP time series, daily meteorological measurements, obtained in parallel with GPP observations, were used as model forcing. Daily forcing data was taken from FluxDataKit (FDK v3.4.2) (Hufkens and Stocker, 2025; Pastorello et al., 2020). The required daily variables are listed in Tab. S1 in the supplemental material.

For predictions of the two leaf traits (Δ , VJ), the P-model implementation in {rsfun} has a total of nine model parameters that are available for calibration (Tab. 2). Due to the relatively high computational cost of simultaneously calibrating all nine model parameters, we start by performing a sensitivity analysis to determine the most influential parameters on the model fit and exclude the least influential ones for the subsequent calibration step. Here, we apply the Morris method for global sensitivity analysis (Morris, 1991) from the {sensitivity} R package (Looss et al., 2023) and compute the sensitivity metrics μ^* , indicating the magnitude of the overall influence of was forced with average climate conditions during the growing season, derived from the global WorldClim data set (Fick and Hijmans, 2017) (comprising monthly averages of daily minimum, maximum and average temperature, vapor pressure, and solar radiation) and considering geographic positions of the sites.

The monthly WorldClim data were temporally disaggregated to daily values through polynomial interpolation (daily minimum, maximum, and average temperatures, cloud cover fraction, solar radiation, and vapor pressure). Interpolated daily maximum and minimum temperatures were then combined to an average daytime temperature using location-based day length assuming a sinusoidal temperature profile (Davis et al., 2017; Peng et al., 2023). The average daytime vapor pressure deficit (D) was derived from the average vapor pressure and daily maximum and minimum temperature.

Daily values were averaged to conditions representing the growing season. Growing season was defined as the period with daily average temperature above 0 °C. Then, daytime temperature, vapor pressure deficit D and solar radiation were averaged (mean) across all days of the growing season and used as model forcing (T , D , PAR) for a given parameter on the prediction, non-temporally resolved single prediction of VJ or Δ for each site. Atmospheric pressure P was derived from the ETOPO-1 digital elevation model (NOAA National Geophysical Data Center, 2009), using site positions and assuming standard atmospheric pressure. CO₂ was taken from yearly mean values from the Mauna Loa record (Keeling et al., 2017), using the corresponding observation year (or the year 2000 if observation year was unavailable).

2.2 Bayesian calibration

We estimated model parameters β , c^* , τ , φ_0^* , a_φ , b_φ , and θ^* (See Tab.2 for a description of parameters) in multiple combinations of parameters and target data (Sec. 2.1 and Tab. 1). Parameter estimation was done through a Bayesian calibration approach, using Markov chain Monte Carlo (MCMC) sampling (Clark, 2004; Dietze et al., 2013), using the DREAMzs sampling algorithm (Vrugt et al., 2009) as implemented in *BayesianTools* (Hartig et al., 2023). Eight independent chains were run, each for 100'000 iterations split among three internal chains, burn-in period was set to 30'000 and convergence was checked visually with trace

plots and Gelman-Rubin statistics (Gelman and Rubin, 1992). Parameters were calibrated to all sites' data simultaneously and are thus assumed to be universal across space and environmental conditions.

2.2.1 Likelihood

The choice of likelihood summarized our assumptions about different sources of uncertainties. Uncertainties in model inputs (parameters p and forcings x), in model structure f , and in the measured observations y of all target types (van Oijen, 2017) combine as

$$y + \varepsilon_y = f(p, x + \varepsilon_x) + \varepsilon_f$$

where ε_y and σ , a measure of the heterogeneity of a parameter's influence on the prediction across the parameter space. We assume ε_f represent (unknown) observational errors and model structural errors, respectively. ε_x is the error in the forcing data.

For all target variables, we assumed an additive and normally distributed model error term for the GPP prediction by the P-model (Trotsiuk et al., 2020) and express around the model prediction (Trotsiuk et al., 2020) and expressed the fit to observed data via the Gaussian log-likelihood. The sensitivity analysis result (Fig. 2) indicates that the quantum yield intercept parameter φ_0 is the most influential parameter, followed by the unit cost of electron transport e^* and the optimal temperature for the quantum yield b . The remaining parameters have relatively little influence on the evaluation of daily GPP predictions from the single site considered here and b_0 , the ratio of dark respiration to the temperature-normalised maximum carboxylation rate, has no influence on GPP predictions, which follows from the model structure. Additional analyses (not shown here) indicated that the convergence of the parameter calibration (shown in the next section) is undermined when calibrating β and e^* simultaneously with other model parameters. Therefore, and based on the sensitivity analysis (Fig. 2), we chose to hold e^*

$$\mathcal{L}(p | y) = \frac{1}{\sqrt{2\pi\sigma^2}} \exp\left(-\frac{1}{2\sigma^2}(y - f(p, x))^2\right) \quad (16)$$

with target-specific standard deviations σ_{GPP} , β , τ , σ_{Δ} , and b_0 constant and thereby excluded them from the Bayesian calibration procedure described in the next section. σ_{VI} . These standard deviations of the error model were estimated together with the model parameters (Tab. 2). Individual observations were considered independent from each other, thus the total likelihood for a dataset simply multiplied the likelihoods of each observation. With this likelihood, we neglected input error $\varepsilon_x = 0$ and lumped together observational uncertainties and model structural uncertainties into a single 'mismatch' or 'residual' uncertainty (van Oijen, 2017; Dietze et al., 2013).

Results from the Morris sensitivity analysis. The y-axis represents all nine model parameters and the Gaussian error standard deviation (error_gpp) and the x-axis the values of the statistics μ^* and σ (unitless, like the log-likelihood). μ^* indicates the magnitude of the overall influence of a certain parameter on the P-model output, while σ measures the heterogeneity of such influence across the parameter space. The parameter names are described and identified with corresponding symbols in Tab. 2. Since the P-model is conceived as a single-big-leaf model (Fig. 1), it represents average properties and fluxes for the whole canopy. The estimated residual uncertainty thus contains also a potential uncertainty due to the scale mismatch

between observation and model. Moreover, across-tree and across-species variabilities are also included since the likelihood was computed for each VJ and Δ observations of individual trees.

2.3 Bayesian calibration

365 While describing leaf-level quantities at relatively high mechanistic detail, the link between the leaf and the canopy-scale was not explicitly resolved. Instead, an empirical approach for leaf-to-canopy scaling of GPP was employed by treating the quantum yield parameter φ_0 to be representative for the canopy-scale and allowing it to be calibrated to ecosystem-level GPP flux data.

~~We simultaneously calibrate a subset~~

2.2.1 Priors

370 Prior distributions were defined based on prior knowledge and kept the same across all calibration setups, except for setup h), where the posteriors of the model parameters ~~that have been identified as particularly influential~~ (Sec. 4.1. β and c^* estimated from setup c), were used as priors (Tab. 2 and shown in grey in Fig. 3). These ~~include the model error term, φ_0~~ posteriors from setup c) were characterized as uni-variant normal distribution. For β , it was additionally truncated to the mean \pm three times the standard deviation.

The prior knowledge on the acclimation time scale τ was approximated by a normal distribution ($\mathcal{N}(14, 64)$) based on prior
375 findings (Mäkelä et al., 2004; Liu et al., 2024; Mengoli et al., 2022) and truncated to the range from 0 to 40 days). For all other parameters, uniform priors with distinct ranges were used. Ranges for β , ~~$a c^*$~~ , and φ_0^* were specified to range between 10% and 300% of published estimates of 146 (unitless), ~~0.41~~ (unitless), and 0.05 mol mol⁻¹, respectively, (Stocker et al., 2020; Wang et al., 2017). We chose wide uniform priors with the aim that posteriors would solely be informed by the used observations. The optimal temperature b_φ and shape parameter a_φ were specified to range between 10 and 30 °C and -0.004 and -0.001 °C⁻². The
380 soil moisture limitation threshold θ^* was specified to range between 1 and β_0 . ~~We use the Differential-Evolution-MCMC-zs (DEzs) sampler (Ter Braak and Vrugt, 2008), implemented in {BayesianTools} (Hartig et al., 2023), to estimate the posterior distribution of the calibrated parameters (Fig. 3). All parameters are given a uniform prior with bounds informed by their physical interpretations (Tab. 2). The prediction error is assumed to be normally distributed, as in the Morris analysis. On a 12th Gen Intel Core i7-1270P processor, it took 1100 sec. to run 3 independent MCMC chains of 24000 iterations (of which 12000 are discarded as burn-in period). The algorithm converged with a scale reduction factor (Gelman and Rubin, 1992) of 1.05 (≤ 1.1). More detail, examples, and explanations of calibration diagnostics are provided through the example vignettes in the package documentation or can be inferred directly from archived the scripts used to create the results (for both, see Sections on Code availability and Data availability). The calculation of the log-likelihood is implemented in the function `cost_likelihood_pmodel()`, enabling custom calibrations, also for multiple target variables that are considered simultaneously~~
390 during model calibration (examples also provided in the package documentation) 250 mm. Given the lack of prior knowledge on the error parameters characterizing the combined structural and observational uncertainty large, uninformative priors were assumed.

Prior and posterior distributions of the calibrated model parameters and error term. The maximum a posteriori (MAP) estimate for the gaussian error standard deviation is 1.06 and the estimate for each calibrated parameter are given in Tab. 2.

2.3 Prediction uncertainty

2.4 **Inference and prediction uncertainty estimation**

The parameter sets generated by the MCMC chains provide the basis for ~~inference (model prediction) and prediction uncertainty estimation~~, allowing us to get insights into the sources of uncertainty. We consider a simple representation of the uncertainty, split between the parameter uncertainty and the model error (Dietze, 2017) model prediction including an estimation of the uncertainty in the predictions on the train and test data set. Here, we propagated both characterized uncertainties: the parametric and the residual (structural and observational) uncertainty.

~~Predicted and observed daily mean GPP. The comparison is provided for the first year of GPP observations (black dots) at the site FR-Pue against GPP predictions (red line), calculated as the median of the posterior distribution. A light green band indicates the 90% credible interval for GPP predictions, which captures parameter uncertainty, while the 90% predictive interval for GPP predictions (in orange) captures model uncertainty.~~

~~Retaining 600~~ Retaining 20 samples from the combined Markov chains, ~~statistically~~ representative of the joint parameter posterior distribution (~~estimated during the calibration~~ including parameter correlations), we ran the P-model for each set of parameters to predict GPP. The credible interval was computed for each time step from the posterior distribution of predicted GPP. The prediction interval for GPP was computed by adding the Gaussian error standard deviation error, to the target variables ~~for~~ the predicted GPP. Fig. 4 shows that the uncertainty ascribable to the parameters (in green) is much smaller than the uncertainty due to the model error (orange area). A point estimate of GPP at each time step is calculated as the median of the posterior distribution of GPP (dark orange line) ~~test and training data set. These 20 sets of predictions represented the parametric uncertainty of the model.~~

3 **Discussion**

The {rsofun} R package provides a user-friendly and fast implementation of the P-model. It implements off-the-shelf model-data assimilation routines with simple function calls, while maintaining flexibility for future experiments and further development of model uncertainty estimation. The {rsofun} vegetation modelling framework is designed to strike a balance between integration and flexibility, enabled by extensible and user-defined calibration specifications and the parallelizable and fast low-level code implementation in Fortran. Its high computational efficiency offers the potential for effective model parameter and uncertainty estimation using Bayesian statistical methods—as demonstrated here.

With the function `calib_sofun()`, {rsofun} provides a blueprint for model-data assimilation and an implementation of the likelihood function with flexibility in selecting among a predefined list of model parameters and target observations. We have demonstrated here how ecosystem flux measurements of GPP from a single site can be used to estimate model parameters

of the P-model and generate estimations of prediction uncertainty. The approach taken here (as in previous studies with the P-model) was to specify latent parameters directly, based on independent observations. In Wang et al. (2017), β and c^* were determined from observations of V_{cmax} , J_{max} . Additionally, for each prediction (i.e. target, site, and date combination) three observational errors were drawn from the residual error model (characterized by σ_{GPP} , σ_{Δ} , and $c_i : c_a$). These were then used as constants for modelling GPP, while additional parameters related to empirical photosynthesis stress parameterizations were calibrated to GPP observations. Exploratory analyses (not shown) indicated that complementary observational constraints are necessary when calibrating β and c^* simultaneously with other model parameters. We have thus followed a simplified setup—used here for demonstration purposes. Future applications may use V_{cmax} , J_{max} , or $c_i : c_a$ data directly as additional calibration targets to provide such complementary constraints. Note also that the FvCB photosynthesis model contains additional parameters (see Tab. A2 in (Stocker et al., 2020)) that are treated as constants here—as in previous publications (Bloomfield et al., 2023; ?; Stocker et al., 2023) and added to the prediction. These 60 sets of predictions represented the combined parametric and residual uncertainty of the model. A third comparison with observations was based on one set of predictions using the Maximum A Posteriori estimate as a single set of parameters and without considering the residual uncertainty.

Our model implementation as an R package takes inspiration from the {r3PG} forest model (Trotsiuk et al., 2020), and our implementation of model-data integration on the basis of ecosystem data serves similar, yet reduced, aims and functionalities compared to PEcAn (<https://pecanproject.github.io/index.html>) (LeBauer et al., 2013). {rsofun} is designed to be minimally reliant on package dependencies and connections to specific data, while limiting the scope to a predefined set of process models (currently P-model, BiomeE at an experimentation stage). Note that the {rpmmodel} R package (available on CRAN) also provides an implementation of the P-model but is written fully in R and in the form of a function of a given environment—without a treatment of temporal dependencies and without the functionalities for model-data integration.

3 Results

We have demonstrated how important parameters contributing to uncertainty in GPP predictions by the P-model can be estimated using observations of ecosystem-level photosynthetic uptake flux time series. The Bayesian approach to model-data integration enables a probabilistic prediction of GPP and estimation of model parameters. The {rsofun} model implementation as an R package makes it possible to leverage a set of methods and complementary libraries for parameter estimation, sensitivity analysis, and calibration diagnostics—as demonstrated here. Its low-level code in Fortran is geared towards computational efficiency. Provision of {rsofun} as an open-access library aims at lowering the bar of entry to vegetation modelling for both field ecologists and computational ecologists and serves as an Open Science resource for future model development and experimentation and the further development of model uncertainty estimation.

3.1 Calibrated parameters

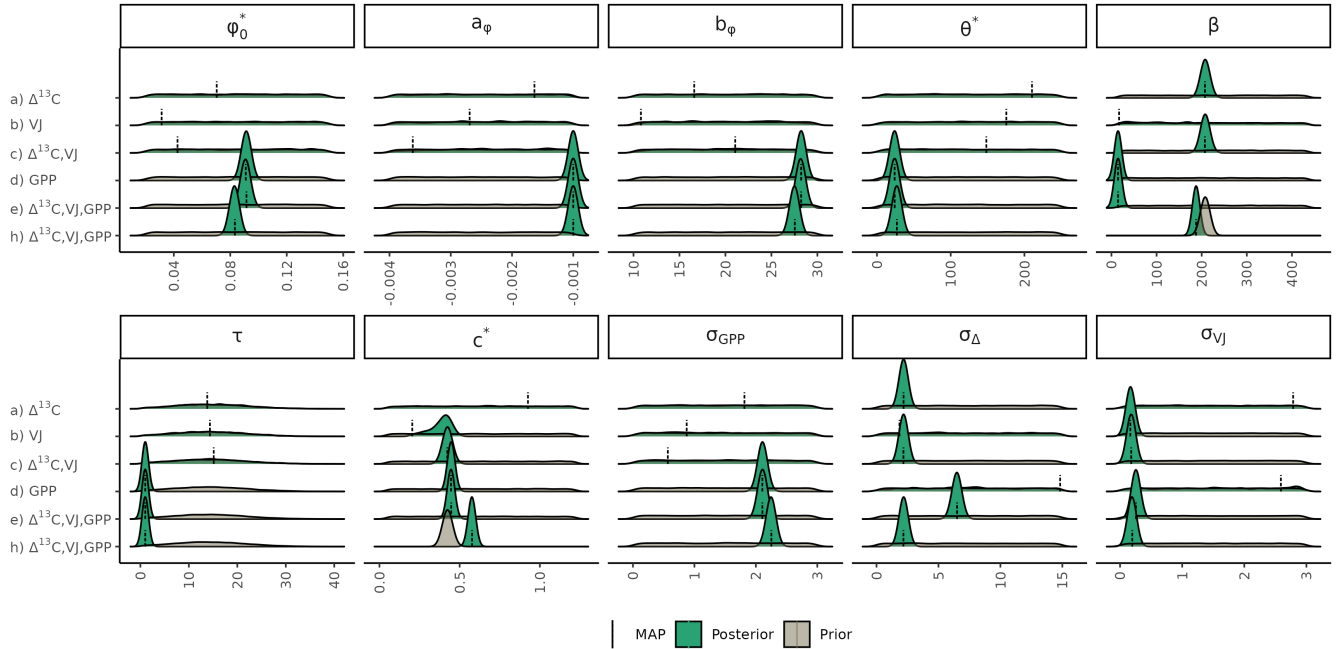


Figure 3. Prior and posterior distributions of the calibrated model parameters and error model terms in the setups. The maximum a posteriori (MAP) estimates are indicated with a dashed line and fixed parameters with solid lines. For parameter description and units refer to Table 2.

MCMC sampling with DREAMzs of 8 parallel, independent chains took between 2.5 (b) and 68 (h) hours to reach 100'000 iterations for the setups a) to h). The trace plot of the chain of setup h), including computations of the scale reduction factors (Gelman and Rubin, 1992) are reported in the Appendix (Fig. D1).

Posterior distributions of the estimated parameters varied across the different setups (Fig. 3). Setup a) (using only Δ as observational target) constrained β to a maximum a posteriori (MAP) of 207 (unitless), a median of 208 and with an inter-quartile range (IQR) from 203 to 213 and the residual prediction error σ_Δ to a MAP of 2.18 (unitless), which corresponds to a 10% of the mean of predicted Δ . Other parameters were not informed by the calibration and their posteriors remained largely identical to their prior distributions. This reflects model structural dependencies of parameters and predicted quantities (Δ is independent of the other model parameters, see Eq. 3).

4 Calibrated empirical functions

Setup b) (using only VJ as observational target) constrained c^* to a MAP of 0.214 (unitless), (median of 0.397, IQR from 0.347 to 0.428). As revealed by the posterior correlation analysis, these two parameters showed a strong correlation ($r = 0.86$, Fig. D2). These compensating effects were disentangled when simultaneously calibrating to Δ and VJ in setup c). This constrained β to a MAP of 207.1 (unitless), (median of 207.9, IQR from 203.5 to 212.3) and c^* to MAP of 0.419 (unitless) (median of

0.425, IQR from 0.410 to 0.439), slightly higher than in setup b), and avoided the correlation of posteriors (Figs. 3 and D3). The error model parameters associated with the two targets Δ and VI were estimated to MAPs of 2.18‰ and 0.178 (unitless).

470 ~~A new formulation of the temperature dependency of the quantum yield efficiency~~ Setups that use GPP as observational target and uninformed priors – setups d), and e) – yield estimates of β that are at the lower bound of the uniform prior range (i.e. 14.6 (unitless)), while τ is estimated to be exactly 1 day. This indicates that GPP observations "push" estimates of β towards extremely high unit costs of transpiration in relation to carboxylation, and that no smoothing of the daily meteorological conditions (Eq. B1) was necessary to optimize the likelihood of observing the GPP data. However, while improving the

475 likelihood of GPP, the fit with Δ observations was deteriorated in these setups, as indicated by an offset between model predictions and observations (Figs. 4 and 5). Only setup h), using GPP in combination with a truncated and prior for β , informed by the reduced setup of c), mitigates this offset. Also here, the posterior estimate of β came to lie at the border of the truncated region (14.6 (unitless)). The error model parameter associated with the GPP target was estimated to a MAP of 2.04 $\text{gC m}^{-2} \text{s}^{-1}$ in setup h), which was slightly smaller than the error in setups d) and e). In the posterior parameters of setup h),

480 correlations of $r = 0.89$ remain between φ_0^* and c^* and of $r = 0.81$ between φ_0 ~~is introduced in the {rsosfun} package, allowing more flexibility than Eq. 18 in Stocker et al. (2020). It is expressed as follows:~~ and b_φ (Fig. D4).

$$\varphi_0(T) = c(1 + a(T - b)^2) \quad \text{if } 0 < c(1 + a(T - b)^2) < 1,$$

$$\varphi_0(T) = 0 \quad \text{if } c(1 + a(T - b)^2) \leq 0,$$

$$\varphi_0(T) = 1 \quad \text{if } c(1 + a(T - b)^2) \geq 1.$$

485 3.1 Prediction uncertainty

Where T stands for temperature, c is the quantum yield efficiency at optimal temperature (in mol mol^{-1}), a is a unitless shape parameter and b is the optimal temperature. Whenever $a = 0$, the quantum yield efficiency is kept constant at $\varphi_0 = c$. A possible improvement for the model would be to use a peaked Arrhenius function instead of a parabola (Medlyn et al., 2002).

490 ~~Furthermore, the soil moisture stress function follows Stocker et al. (2020), but the parameters considered for calibration there differ from the calibratable parameters in the package:~~ Model predictions were unbiased and residuals were of similar magnitude when evaluated on the test and on the training data sets (Fig. 4), which indicates a good generalizability of the parametrized model. Including structural and observational uncertainties on top of parametric uncertainties only slightly increased the deviations between predicted and observed targets in setup h), with strongest relative increases for GPP.

$$\beta(\theta) = \frac{\beta_0 - 1}{\theta^{*2}} (\theta - \theta^*)^2 + 1 \quad \text{if } 0 \leq \theta \leq \theta^*,$$

495 $\beta(\theta) = 1 \quad \text{if } \theta > \theta^*.$

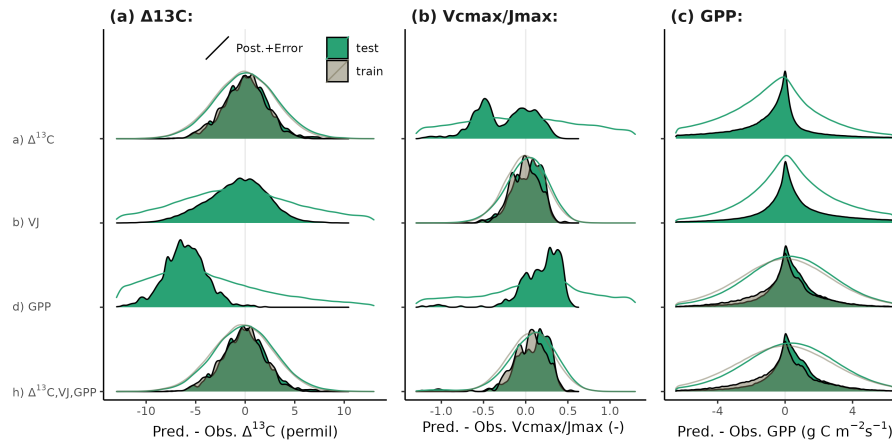


Figure 4. Density plot of residuals between predicted and observed values from the training and test sets for calibration setups a, b, d, and h. Model output is computed with parametric uncertainty (filled area, based on 20 samples from the posterior distribution) and residual uncertainty (solid line, based on 3 samples from error model). Model outputs are compared against individual observations (dates from all sites pooled for GPP, and individual observations of each site for Δ and VJ).

Δ and GPP predictions based on MAP parameter sets of setup h) showed no magnitude-dependent bias (linear slopes of regressions between predicted and observed values were close to 1), whereas VJ showed a prediction range that appears too large when compared to the observed range (slope close to 0, Fig. 5). Setup h) showed a slightly worse root mean squared error (RMSE) than setup e) for GPP, but clearly reduced RMSE for Δ and VJ.

500 In the equation above, θ stands for the plant-available soil water (in mm). Time series of GPP of a few select years on training site data showed that the model successfully reproduced seasonal patterns and differences across site in GPP for most sites (Fig. D5), with some shortcomings in accurate simulations of GPP during under dry conditions, as seen for site US-Var, and a general low bias for the moist tropical site of GF-Guy. The model also tends to systematically overestimate GPP in the early growing season at US-MMS and θ^* for the threshold indicating when the plants start being water stressed. The intercept β_0 is the β reduction at low soil moisture, that is $\beta_0 = \beta(0)$. This intercept is now calibrated directly, rather than expressed as a function of mean aridity (see eq. 20 in (Stocker et al., 2020)). US-PFa – a known bias (Luo et al., 2023).

505

4 Discussion

This study showed that a model for ecosystem photosynthesis and its acclimation to the environment can be robustly parameterised and that its predictions of multiple variables generalise well across a wide range of environmental conditions. Multiple model parameters can be estimated simultaneously by using diverse calibration target data types, combining ecosystem flux time series and static, species-specific traits data. This demonstrates how the explicit representation of connections between traits and process rates enable model-data integration on the basis of diverse observations, obtained at multiple organisational levels

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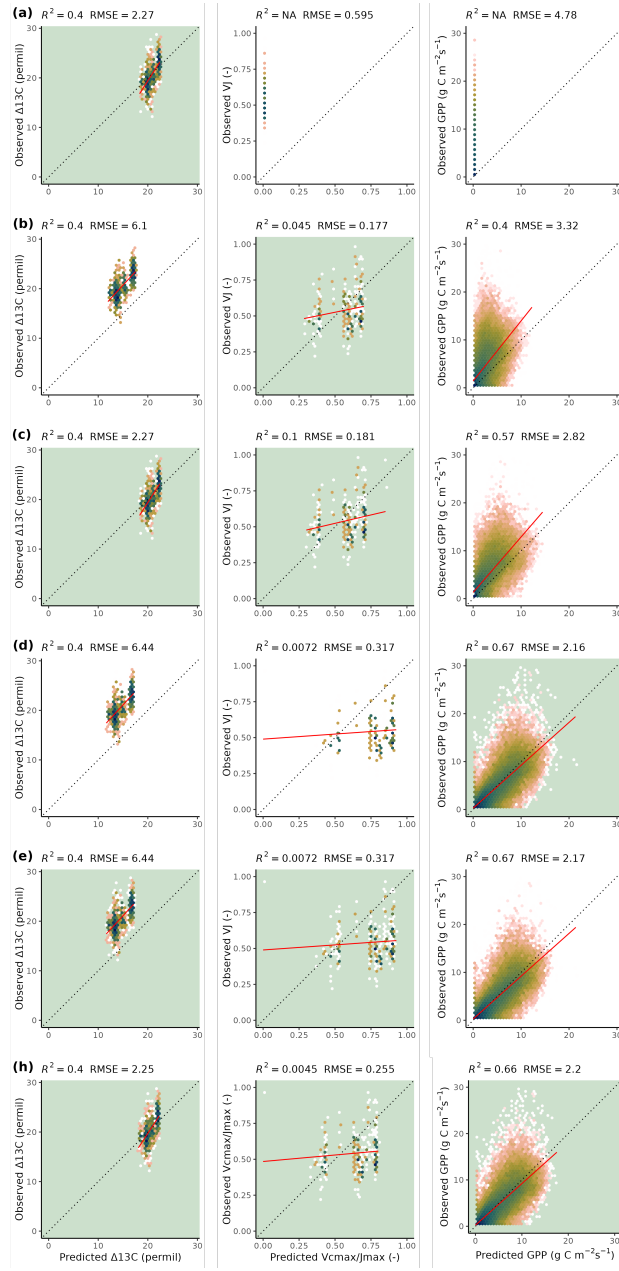


Figure 5. Predicted versus observed values for different target variables (Δ , VJ, and GPP along columns) and calibration setups (a, b, c, d, e, and h, along rows), evaluated on the test set. Model output is computed with Maximum A Posteriori parameter values (MAP) for each calibration setup. Note that the MAP parameters from setup a) result in null-predictions of VJ and GPP. Color indicates density, red line indicates a linear regression. Green panel backgrounds indicate which variables were used as targets for model calibration in the corresponding setup. Model outputs are compared against individual observations (predicted and observed daily GPP values from all sites pooled and site-specific predictions against all observations from the respective site for Δ and VJ).

– from the leaf to the canopy. The robustness of the parameter estimates is indicated by the convergence of the MCMC chains (Fig. D1) and the resulting narrow posterior distributions of Bayesian parameter estimates (Fig. 3). Despite being calibrated to only a relatively small set of sites with GPP data ($N=12$), the calibrated model generalizes well, as validated against an independent and much larger set of test data for GPP from 38 sites (Figs. 4 and 5).

Our results confirmed the expectation that the use of multiple observational targets yields more robust parameter estimates compared to a calibration setup that relies on a single data source, and that specific observation types imposed constraints for specific model parameters. Leaf carbon fractionation observations, Δ , allowed to constrain the unit cost ratio of carboxylation to transpiration, β (Fig. 3). Observed ratios of the biochemical rates of carboxylation to electron transport, VJ, constrained β and c^* , albeit with strong correlations between them, indicating compensating effects and a lack of robustness in resulting parameter estimates (Figs. 3 and D2). The combination of both these observation types allowed to constrain both parameters simultaneously, avoiding correlations between parameter estimates. This indicates that the use of two observational targets simultaneously made use of their complementary information content for parameter estimation in our model (Fig. D3).

~~We thank Florian Hartig for advice on uncertainty modelling and~~ Despite the general robustness of parameter estimates, we found several limitations and aspects that indicate challenges for model calibration in our case. When observations of GPP were included in calibration setups, parameter estimates of β differed strongly from results obtained from setups that used only observations of Δ and tended towards the lower margin of the uniform prior range – substantially lower than the value used for direct parameterisation of the model in previous work (Wang et al., 2017). β represents the ratio between the unit cost of carboxylation to transpiration within the EEO modelling framework applied here (Prentice et al., 2014). An extremely low value of β implies relatively high costs associated with transpiration, which is driven by VPD. The calibration tending towards low values of β potentially reflects a compensating effect for the ~~maintenance of the BayesianTools package, Volodymyr Trotsiuk for an initial template of the R package,~~ lack of GPP reductions under conditions of dry soils, e.g., during the dry summer periods at the site US-Var (Fig. D5). In other words, this apparent lack of robustness of parameter estimate may indicate a misspecification of the model structure. This interpretation could be tested with targeted setups (e.g., removing dry sites from the calibration data set) or by alternative specifications of the soil moisture stress that better accounts for its limiting effect on GPP.

To address this challenge and avoid unexpectedly low estimates of the unit cost ratio parameter β , we resorted to a step-wise Bayesian calibration (MacBean et al., 2016) and used the posterior distribution of setup c) as a prior in setup h). This resulted in the disappearance of offsets in Δ observation, a c^* closer to, but not at, the upper limit of the prior (MAP = 0.58), and ~~Colin Prentice for comments on the manuscript. B.D.S. was funded by the Swiss National Science Foundation grant PCEFP2_181115. This work is a contribution to the LEMONTREE (Land Ecosystem Models based On New Theory, obseRvations and ExperimEnts) project, funded through the generosity of Eric and Wendy Schmidt by recommendation of the Schmidt Futures program (K. H., B.D.S.).~~ a slightly lower φ_0^* than in GPP-only setup (d) or in the unconstrained GPP- Δ -VJ setup (e). The step-wise posterior estimation and prior specification of P-model parameters in setup h) yield estimates of all parameters that compare favourably with previous estimates (Stocker et al., 2020; Wang et al., 2017). This study estimated (MAP): $\beta=208$ (unitless, compared with 146 and the range from 200 to 240), $c^*=0.58$ (unitless, compared with 0.41), and $\varphi_0^*=0.05$ mol mol⁻¹

(compared to 0.05). However, the MAP estimate of $\tau = 1$, indicated an instantaneous acclimation appeared to yield better agreement of model output with daily GPP observations, than a delayed acclimation using smoothed versions of the environmental conditions. This contrasts with previous estimations of the acclimation time scale being on the order of 14 to 15 days (Mäkelä et al., 2004; L

Remaining correlations in posterior parameters of setup h) between φ_0^* and c^* indicate some equifinality. These compensating effects indicate another potential source of misinterpretation of what the parameter c^* represents in setup b) versus setup h). The Bayesian approach allows propagating the effect of this into predictive uncertainties by using posterior *distributions* instead of point-estimates of parameters. Still, this entanglement might be resolved in future calibrations by fixing one of these two parameters or by using additional observations. Future work is needed to look into causes of this posterior variability of c^* , identify potential observational constraints, and potentially to revise related model structures.

The calibrated model of setup h) showed unbiased predictions against observations in an independent test data set, indicating the model's generalizability. Based on the universal validity of the EEO parameters across plant functional types and biomes, the calibrated P-model can may be scaled to new locations sites and environmental conditions. However, it should be noted that our estimates of prediction uncertainty and the finding of robust generalisability only applies to environmental conditions that are within the domain of (or similarly distanced to) sites used for training and testing here (Ludwig et al., 2023). Further caveats apply. The choice of the specification of soil moisture stress, which uses a single global parameter θ^* , may be overly simplistic for describing diverse physiological responses to dry-downs across sites characterised by different soil texture (Fu et al., 2022; Wankmüller et al., 2024) and the model neglects the highly variable rooting zone water storage capacities across space globally (Stocker et al., 2023). Although generalization to the test data set did globally show unbiased predictions across all sites (Fig. 4), shortcomings were visible for certain sites (e.g., US-Var, Fig. 5) and warrant a re-consideration and potential revision of the related process representations. An improved representation of soil moisture stress effects and across-site variations of critical soil moisture thresholds and storage capacities may potentially also mitigate the need for targeted interventions into the calibration procedure, e.g., through the truncation of the prior for β or the step-wise approach to model calibration. More specifically, θ^* could potentially be expressed by additionally considering soil structural information (Wankmüller et al., 2024), which could be provided as additional information for each site. Alternatively, the Bayesian approach allows for the use of hierarchical models (van Oijen, 2017) that could make use of potential, globally available predictors (or covariates) for the site-specific parameters.

Generally, design choices related to the Bayesian likelihood specification and parameter calibration were adequate to showcase the sensitivity of parameter estimates across setups and retrieve an unbiased, generalizable model. While simple, the additive Gaussian error model chosen for the likelihood allowed to showcase the constraints, it could use improved functional forms more suited e.g. to the positivity of GPP, or to disentangle the observation uncertainty from model structural uncertainty (van Oijen, 2017). Independent estimates of observation uncertainty could be included as fixed parameters or as informed priors. This would potentially lead to a reduced residual error representing more closely the structural model uncertainty. Potential estimates for these could be the error on GPP measurements or errors of a trait measurement to represent the model scale (i.e. the ecosystem-level averages across species instead of individual observations). Alternatively, fitting to average

observations instead of individual observations might reduce the (independent) observational errors. This would mean fitting to averaged traits over different species at a given site or weekly cumulative fluxes instead of daily values.

585 A strength of the Bayesian approach is to update model parameters iteratively with additional data by using previous posterior as new prior distributions. Our illustration of this approach in setup h) relied on manual extraction and specification of these distributions. A future development step for the *rsofun* framework would thus be to support easier specifications of these type of stepwise calibration (MacBean et al., 2016) directly through the model interface.

Bagnara, M., Harrison, S. P., Prentice, I. C., Bartlein, P. J., Kelley, D. I., Daniau, A. L., Krawchuk, M. A., and Moritz, M. A.: A model-based approach to predicting fire weather and carbon dynamics in Mediterranean ecosystems, *Biogeosciences*, 12, 6173–6196, <https://doi.org/10.5194/bg-12-6173-2015>, 2015. P-model is limited in scope to the to canopy. We found even for models with this degree of complexity, parameter interactions and structural model errors can to lead parameter biases. The stepwise calibration approach allowed to rectify the perceived bias in β . This violated the initial aim of simultaneously estimating all parameters, but is interpreted here as indications of where to find structural deficiencies and target model improvements (Oberpriller et al., 2021). More comprehensive vegetation models of larger scope will likely face similar challenges, considering that structural deficiencies are unavoidable and more parameter interactions may arise through feedbacks in the soil-plant system. Our findings indicate that stepwise approaches to model calibration can be potential solutions to this challenge.

Bloomfield, K. J., van Hoolst, R., Balzarolo, M., Janssens, I. A., Veeva, S., Ghent, D., and Prentice, I. C.: Towards a General Monitoring System for Terrestrial Primary Production: A Test Spanning the European Drought of 2018, *Remote Sensing*, 15, 1693, <https://doi.org/10.3390/rs15061693>, 2023. While this study looked at seven model parameters simultaneously, several other parameters were held fixed, namely those involved in describing temperature and pressure dependencies of physical and physiological quantities, isotope fractionation, etc. This simplified the calibration and parameter estimation task, but assumed that the uncertainty stemming from these is negligible for the prediction of target variables. This is a strong assumption and should be relaxed in future attempts at calibrating P-model.

Clark, J. S.: Why environmental scientists are becoming Bayesians: Modelling with Bayes, *Ecology Letters*, 8, 2–14, <https://doi.org/10.1111/j.1461-0248.2004.00702.x>, 2004. Lastly, the chosen likelihood ignored input data uncertainty ε_x . For GPP sites, we expect small uncertainty in meteorological variables thanks to local measurements, but fAPAR data is obtained from satellite remote sensing data, which is likely to include uncertainties. For Δ and VJ sites, input forcing was based on a global dataset of high spatial resolution (~ 1 km around the equator). However, topographical effects and related micro-climata that deviate from the larger climate could be remaining sources of errors.

Cowan, I. R. and Farquhar, G. D.: Stomatal function in relation to leaf metabolism and environment. In: *Integration of Activity in The rsofun* R package provides a user-friendly and efficient implementation of the Higher Plant, 471–505, 1977. P-model and off-the-shelf model-data assimilation functionalities through its connection to the *BayesianTools* R package (Hartig et al., 2023), while maintaining flexibility for altered calibration setups and likelihood definitions. P-model's computational

[efficiency offered the potential for effective model parameter and uncertainty estimation using Bayesian statistical methods in combination with flux and traits data here.](#)

Dietze, M. C.: *Ecological Forecasting*, Princeton University Press, <https://doi.org/10.2307/j.ctvc7796h>, 2017. Our model implementation as an R package takes inspiration from the *r3PG* forest model (Trotsiuk et al., 2020), and our implementation of model-data integration on the basis of ecosystem data serves similar, yet reduced, aims and functionalities compared to PécAn (<https://pecanproject.github.io/index.html>) (LeBauer et al., 2013). *rsofun* is designed to be minimally reliant on package dependencies and connections to specific data, while limiting the scope to a predefined set of process models (currently P-model, BiomeE at an experimentation stage). *rsofun* and the implementation of our simulations and analyses analyses shown here in accompanying code (see code availability statement) and package vignettes provide a blueprint for model-data assimilation.

Dietze, M. C., Lebauer, D. S., and Kooper, R.: On improving the communication between models and data, *Plant, Cell & Environment*, 36, 1575–1585, <https://doi.org/10.1111/pee.12043>, 2013.

5 Conclusions

This study used *rsofun* (available as an R package) to calibrate latent P-model parameters to a set of flux and traits data, obtained from 193 training sites and 231 test sites. Predictions across all test sites showed that the calibrated model generalized well, not showing any biases and similar prediction residuals between train and test data. The Bayesian calibration also exhibited challenges. Structural uncertainty, the imbalanced data set, and a potentially too simplistic likelihood lead to biased parameters (β) and predictions (Δ) in the first calibration attempt. An alternative calibration setup made use of a stepwise calibration and the hierarchical design of the model structure and predictions, enabling a successive model integration and parameter calibration. The necessity for this approach to obtain robust and reliable parameter estimations may indicate model structural deficiencies, which we identified here as primarily being related to the representation of effects by dryness in soil and air.

Dietze, M. C., Fox, A., Beck-Johnson, L. M., Betancourt, J. L., Hooten, M. B., Jarnevich, C. S., Keitt, T. H., Kenney, M. A., Laney, C. M., Larsen, L. G., Loescher, H. W., Lunt, C. K., Pijanowski, B. C., Randerson, J. T., Read, E. K., Tredennick, A. T., Vargas, R., Weathers, K. C., and White, E. P.: Iterative near-term ecological forecasting: Needs, opportunities, and challenges, *Proc. Natl. Acad. Sci. U.S.A.*, 115, 1424–1432, <https://doi.org/10.1073/pnas.1710231115>, 2018.

Farquhar, G. D., Caemmerer, S., and Berry, J.: A biochemical model of photosynthetic CO_2 assimilation in leaves of C_3 species, *Planta*, 149, 78–90, 1980.

Code and data availability. The *rsofun* R package can be installed from CRAN (<https://cran.r-project.org/package=rsofun>) or directly from its source code on GitHub (publicly available at <https://github.com/geco-bern/rsofun> under an AGPLv3 licence). Versioned releases of the GitHub repository are deposited on Zenodo (<https://doi.org/10.5281/zenodo.15189864>). Code to reproduce the analysis and plots presented here is contained in the repository at https://github.com/geco-bern/rsofun_doc and archived on Zenodo (Bernhard and Stocker, 2025). The model forcing and evaluation data for GPP sites are based on the publicly available FLUXNET2015 data, prepared by FluxDataKit v3.4.2

(10.5281/zenodo.14808331). Scripts for generating these data files from open access sources are contained in the repository in the subdirectory 'data-raw/'. Outputs of the analysis presented here are archived in the subdirectory 'analysis-output/'. The model forcing for Δ and VJ sites are based on the publicly available WorldClim (Fick and Hijmans, 2017), ETOPO1 (NOAA National Geophysical Data Center, 2009), and Mauna Loa CO₂ (Keeling et al., 2017) data. The model evaluation data for Δ sites are based on data associated with (Cornwell et al., 2018) available at (Cornwell, 2025). The model evaluation data for VJ sites are based on the freely available subset of the data from (Smith et al., 2019). Scripts for generating these data are contained in the repository in subdirectory 'data-raw/'. Outputs of those scripts and of the analysis presented here are archived in the subdirectory 'data/' and 'analysis-output/'.

Franklin, O., Harrison, S. P., Dewar, R., Farrior, C. E., Brännström, Å., Dieckmann, U., Pietsch, S., Falster, D., Cramer, W., Loreau, M., Wang, H., Mäkelä, A., Rebel, K. T., Meron, E., Schymanski, S. J., Rovenskaya, E., Stocker, B. D., Zachle, S., Manzoni, S., Van Oijen, M., Wright, I. J., Ciais, P., Van Bodegom, P. M., Peñuelas, J., Hofhansl, F., Terrer, C., Soudzilovskaia, N. A., Midgley, G., and Prentice, I. C.: Organizing principles for vegetation dynamics, *Nat. Plants*, 6, 444–453, <https://doi.org/10.1038/s41467-020-1992->

Gelman, A. and Rubin, D. B.: Inference from Iterative Simulation Using Multiple Sequences, *Statist. Sci.*, 7, <https://doi.org/10.1214/ss/1177044492>.

Harrison, S. P., Cramer, W., Franklin, O., Prentice, I. C., Wang, H., Brännström, Å., de Boer, H., Dieckmann, U., Joshi, J., Keenan, T. F., Lavergne, A., Manzoni, S., Mengoli, G., Morfopoulos, C., Peñuelas, J., Pietsch, S., Rebel, K. T., Ryu, Y., Smith, N. G., Stocker, B. D., and Wright, I. J.: Eco-evolutionary optimality as a means to improve vegetation and land-surface models, *New Phytologist*, 231, 2125–2141, <https://doi.org/10.1111/nph.17558>, 2021.

665 Appendix B: [Acclimation](#)

Hartig, F., Dyke, J., Hickler, T., Higgins, S. I., O'Hara, R. B., Scheiter, S., and Huth, A.: Connecting dynamic vegetation models to data—an inverse perspective: Dynamic vegetation models—an inverse perspective, *J. Biogeogr.*, 39, 2240–2252, <https://doi.org/10.1111/j.1365-2699.2012.02745.x>, 2012. [Damped acclimation to daily environmental conditions was considered for the GPP prediction, while for \$\Delta\$ and VJ predictions used growing season average conditions.](#)

670 Hartig, F., Minunno, F., Paul, S., Cameron, D., Ott, T., and Pichler, M.: BayesianTools: General-Purpose MCMC and SMC Samplers and Tools for Bayesian Statistics, 2023. [The the low-pass filter of characteristic time scale \$\tau\$ \(days\) is defined as](#)

$$\text{The low-pass filter of characteristic time scale } \tau \text{ (days) is defined as } T(t+1) = T(t) + \frac{1}{\tau}(T'(t+1) - T(t)) \quad \text{initialized as } T(0) = T'(0), \quad (\text{B1})$$

[where \$T\$ is the low-pass filtered quantity \(here temperature\) and \$T'\$ the daily observations, resulting in a daily time series of the damped quantity. Equivalent expressions with the same \$\tau\$ are used for \$P\$, \$D\$, CO₂, and PAR.](#)

675 LeBauer, D. S., Wang, D., Richter, K. T., Davidson, C. C., and Dietze, M. C.: Facilitating feedbacks between field measurements and ecosystem models, *Ecological Monographs*, 83, 133–154, 2013.

Appendix C: The *rsofun* model framework

Looss, B., Iooss, B., Prieur, C., Veiga, S. de, Koehler, J. R., Pujol, G., and Saltelli, A.: sensitivity: Global Sensitivity Analysis of Model Outputs, <https://cran.r-project.org/web/packages/sensitivity/index.html>, *rsofun* implements the P-model (Stocker et al., 2020) and provides off-the-shelf methods for Bayesian (probabilistic) parameter and prediction uncertainty estimation. *rsofun* is distributed as an R package on R's central and public package repository. *rsofun* also implements the BiomeE vegetation demography model (Weng et al., 2017, 2015). The latter is not further described here and is implemented at an experimental stage in *rsofun* version v5.0. The P-model implementation in *rsofun* is designed for time series simulations by accounting for temporal dependencies in the acclimation to a continuously varying environment. Function wrappers in R make the simulation workflow user-friendly and all functions and input forcing data structures are comprehensively documented (<https://cran.r-project.org/web/packages/rsofun/index.html>) (https://cran.r-project.org/web/packages/rsofun/index.html, 2023).

Marquet, P. A., Allen, A. P., Brown, J. H., Dunne, J. A., Enquist, B. J., Gillooly, J. F., Gowaty, P. A., Green, J. L., Harte, J., Hubbell, S. P., O'Dwyer, J., Okie, J. G., Ostling, A., Ritchie, M., Storeh, D., and West, G. B.: On Theory in Ecology, *BioScience*, 64, 701–710, <https://doi.org/10.1093/biosci/biu098>, 2014.

Medlyn, B. E., Dreyer, E., Ellsworth, D., Forstreuter, M., Harley, P. C., Kirschbaum, M. U. F., Le Roux, X., Montpied, P., Strassmeyer, J., Walcroft, A., Wang, K., and Loustau, D.: Temperature response of parameters of a biochemically-based model of photosynthesis. II. A review of experimental data, *Plant, Cell & Environment*, 25, 1167–1179, <https://doi.org/10.1046/j.1365-3040.2002.00891.x>, 2002.

In *rsofun*, model parameters can be calibrated using a calibration function `calib_sofun()`, providing two modes of calibration, one based on generalised simulated annealing (`{GenSA}` R package) for global optimization (Xiang et al., 2013) and one based on Markov chain Monte Carlo (MCMC) implemented by the *BayesianTools* R package, giving access to a wide variety of Bayesian methods (Hartig et al., 2023). The former being fast, while the latter provides more informed parameter optimization statistics (Clark, 2004; Dietze et al., 2013). This gives the option for both exploratory and more in-depth analysis of estimated parameters. A set of standard cost functions are provided for the calibration, facilitating the exploration of various metrics or target variables and the specification of calibrated model parameters. Furthermore, the vignettes accompanying the package (<https://cran.r-project.org/web/packages/rsofun/index.html>) explain how to customise the calibration cost functions and interpret the calibration results.

Mengoli, G., Agustí-Panareda, A., Boussetta, S., Harrison, S. P., Trotta, C., and Prentice, I. C.: Ecosystem Photosynthesis in Land-Surface Models: A First-Principles Approach Incorporating Acclimation, *Journal of Advances in Modeling Earth Systems*, 14, e2021MS002767, <https://doi.org/10.1029/2021MS002767>, 2022.

Morris, M. D.: Factorial Sampling Plans for Preliminary Computational Experiments, *Technometrics*, 33, 161–174, <https://doi.org/10.1080/00401706.1991.10484804>, 1991.

Appendix D: Posterior parameter estimates

Pastorello, G., Trotta, C., Canfora, E., Chu, H., Christianson, D., Cheah, Y.-W., Poindexter, C., Chen, J., Elbashandy, A., Humphrey, M., Isaac, P., Polidori, D., Reichstein, M., Ribeca, A., van Ingen, C., Vuichard, N., Zhang, L., Amiro, B., Ammann,

out_calib_scen231_DREAMzs-100000-0iter_8x3chains_on_CPU8x1_continued.rds

GelmanDiagnostics: mpsrf=1.2; psrf: $\phi_0^*=1.09$, $a_\phi=1.13$, $b_\phi=1.13$, $\theta^*=1.05$, $\beta=1.09$, $\tau=1.11$, $c^*=1.10$, $\sigma_{GPP}=1.03$, $\sigma_\Delta=1.06$, $\sigma_{vj}=1.07$

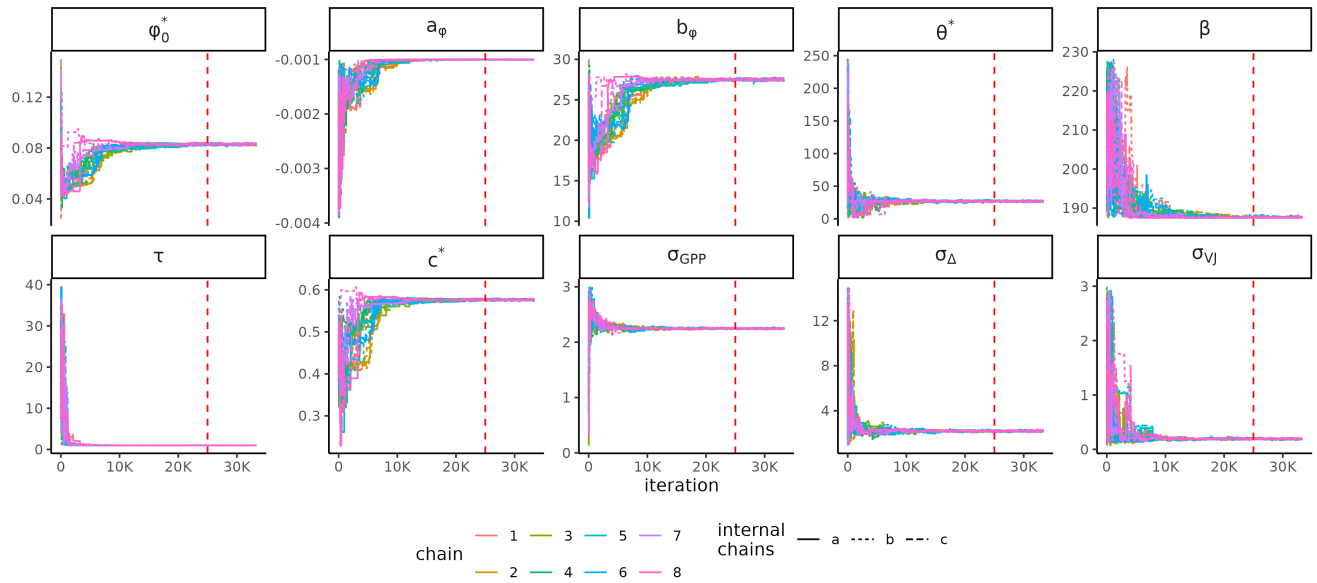


Figure D1. Trace plot of MCMC sampling of all parameters in setup h).

710 [C., Arain, M. A., and others: The FLUXNET2015 dataset and the ONEFlux processing pipeline for eddy covariance data, Sci Data, 7, 225, https://doi.org/10.1038/s41597-020-0534-3, 2020.](https://doi.org/10.1038/s41597-020-0534-3)

D1 [Posterior parameter sampling](#)

Table 2. Parameter listing with description, including Maximum A Posteriori (MAP) estimates and prior specifications, obtained from the Bayesian model different calibration setups (Sec. 4.2) are listed in column 'Values' columns). The values given in square brackets represent the lower and upper bounds of uniform prior distributions are given in parentheses or truncated normal prior distributions. Parameters that were held fixed for the calibration are marked with a single number in brackets and an asterisk (*) in the corresponding setup columns. Fixed site information was: longitude = 3.6°E, latitude = 43.7°N, elevation = 270 m a.s.l., and total root zone water storage capacity = 432 mm.

Symbol	Parameter	Description	Units	Values	Setup c)	Setup d)	Setup e)	Setup h)
	name		Setup a)	Setup b)				
	Units							
φ_0	kphio	Quantum yield at optimal temperature	mol mol^{-1}	0.041		0.09	0.09	0.08
φ_0^*	mol mol ⁻¹			(0.02, 0.15)		[0.02, 0.15]	[0.02, 0.15]	[0.02, 0.15]
α_a	kphio_par_a	Shape parameter for the temperature dependence of the quantum yield	$^{\circ}\text{C}^{-2}$	-0.0023		-0.001	-0.001	-0.001
	$^{\circ}\text{C}^{-2}$			(-0.004, -0.001)		[-0.004, -0.001]	[-0.004, -0.001]	[-0.004, -0.001]
b_e	kphio_par_b	Optimal temperature for the quantum yield	$^{\circ}\text{C}$	15.3		28.2	28.2	27.5
θ^*	$^{\circ}\text{C}$			(10, 30)		[10, 30]	[10, 30]	[10, 30]
	soilm_thetas0	Threshold plant-available soil water content in the soil moisture stress function	mm	158	soilm_beta0	Stress factor at low soil moisture; intercept for the soil moisture stress function	unitless	0.0001
	mm			(4.3, 432)			24.5	(0, 26.7)
		<u>Soil moisture limitation threshold (eq. 15)</u>					[1, 250]	[1, 250]
β	beta_uniteo	Unit cost ratio of carboxylation (maintenance of V_{cmax}) to transpiration	unitless	146	rd_to_vemax	Ratio of temperature-normalised dark respiration to the temperature-normalised maximum carboxylation rate	unitless	0.014
	unitless			(14.6, 438.0]			14.6	187.5
							[14.6, 438.0]	$\mathcal{N}(207.9, 6.8^2)^b$

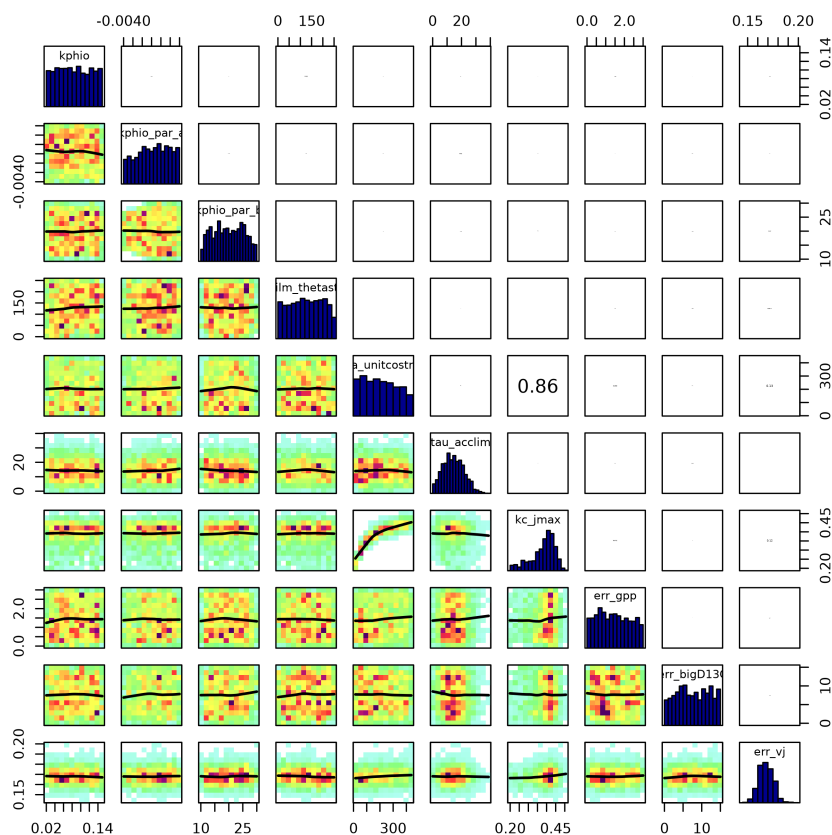


Figure D2. Parameter correlation of posterior parameters in setup b) with 100'000 iterations and 30'000 burnin.

Prentice, I. C., Dong, N., Gleason, S. M., Maire, V., and Wright, I. J.: Balancing the costs of carbon gain and water transport: testing a new theoretical framework for plant functional ecology, *Ecol Lett*, 17, 82–91, <https://doi.org/10.1111/ele.12211>, 2014.

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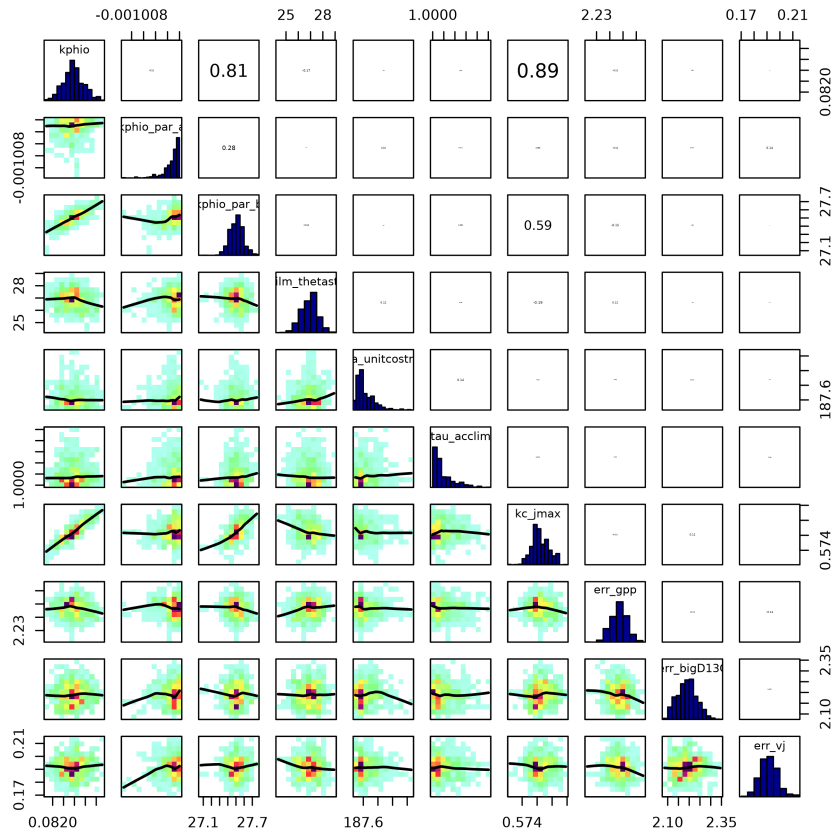


Figure D4. Parameter correlation of posterior parameters in setup h) with 100'000 iterations and 30'000 burnin.

Raj, R., Hamm, N. A. S., van der Tol, C., and Stein, A.: Uncertainty analysis of gross primary production partitioned from net ecosystem exchange measurements, *Biogeosciences*, 15, 1545–1557, <https://doi.org/10.5194/bg-15-1545-2018>, 2018.

D3 Posterior parameter predictions

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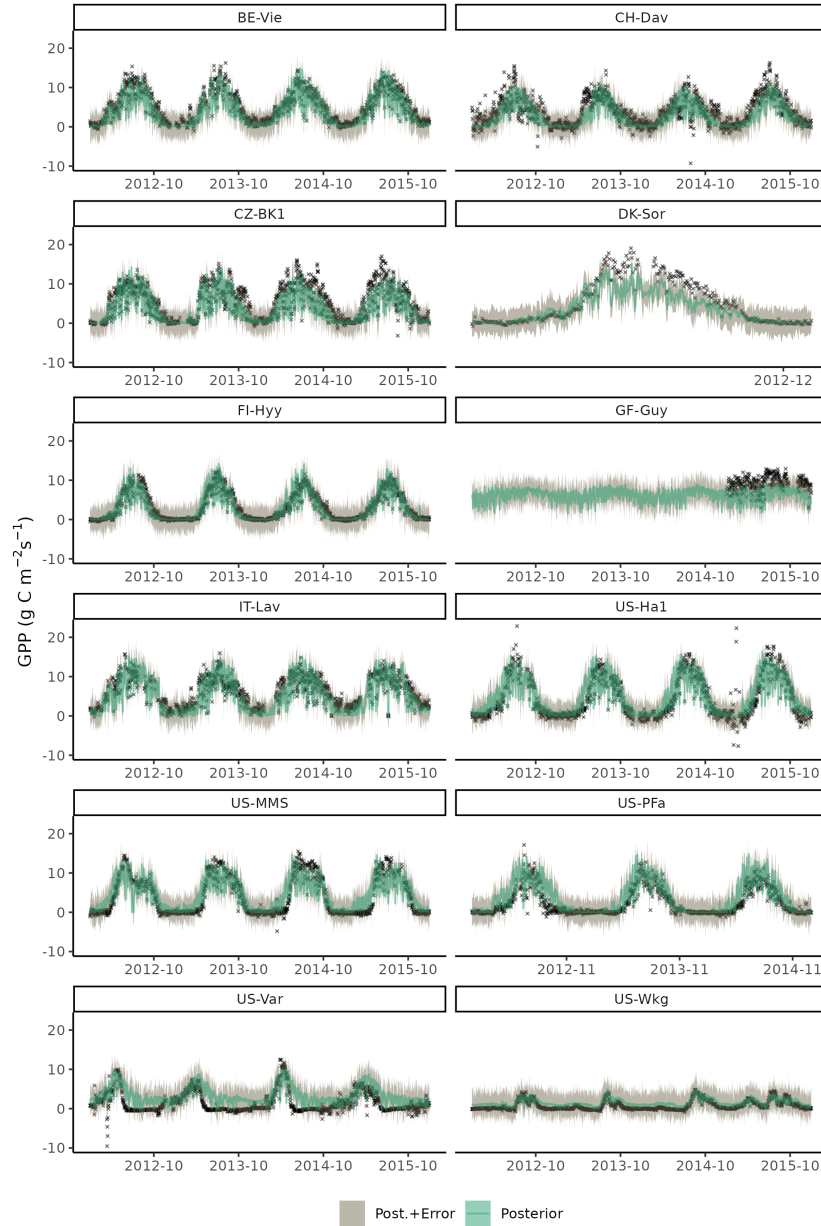


Figure D5. Time series plot of observed (black crosses) and modelled daily gross primary production (GPP) for setup h for years 2012 to 2015 of the training data set. Model output is computed with parametric uncertainty (green shaded area, based on 20 samples from the posterior) and structural uncertainty (grey shaded area, based on 3 samples from error model).

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- Author contributions.* B.D.S. and F.B. designed the study. B.D.S. wrote the Fortran code, J.A.P. wrote the calibration implementation in the R package, sensitivity analysis and uncertainty estimation. M.M. and F.B. prepared the package for publication to CRAN. K.H. and J.A.P. wrote the initial draft of the manuscript. F.B. processed the observational data, ran the analysis, and finalized the manuscript. All authors contributed to manuscript writing.
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[Volodymyr Trotsiuk for an initial template of the R package, and Colin Prentice for comments on the manuscript. We thank Nick Smith](#)
[for providing data for VJ. B.D.S. , Farrior, C. E., Dybzinski, R., and Pacala, S. W.: Predicting vegetation type through physiological](#)
[and environmental interactions with leaf traits: evergreen and deciduous forests in an earth system modeling framework, Global Change](#)
[Biology, 23, 2482–2498, was funded by the Swiss National Science Foundation grant PCEFP2_181115. This work is a contribution to the](#)
760 [LEMONTREE \(Land Ecosystem Models based On New Theory, obseRvations and ExperimEnts\) project, funded through the generosity](#)
[of Eric and Wendy Schmidt by recommendation of the Schmidt Futures program \(K.H., B.D.S.\). Calculations were performed on AMD](#)
[EPYC 7742 \(2.25, 3.4 GHz\) processors on UBELIX \(<https://doi.org/www.id.unibe.ch/10.1111/geb.13542, 2019>\), the HPC cluster at the](#)
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[Xiang, Y., Gubian, S., Suomela, B., and Hoeng, J. : Generalized Simulated Annealing for Global Optimization: The GenSA](#)
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