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Long-term trends in daytime cirrus cloud radiative effects: Analyzing twenty years of Micropulse Lidar Network measurements at Greenbelt, Maryland in eastern North America

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Abstract. This pioneering study elucidates the long-term trends and intricate variability of the radiative impacts and optical characteristics of cirrus clouds over two decades, from 2003 to 2022 at the NASA Goddard Space Flight Center in Greenbelt, Maryland, USA, headquarters of the Micropulse Lidar Network (MPLNET) project. Over twenty years, analysis of the net cloud radiative effects (CREs) at both the top-of-the-atmosphere (TOA) and surface (SFC) reveals decreases in radiative flux 5 by -0.0017 and $-0.0035 \text{ W m}^{-2} \text{ yr}^{-1}$ and -0.0027 and $-0.048 \text{ W m}^{-2} \text{ yr}^{-1}$, respectively (based on the constrained solutions for lidar-derived 523/527/532 nm extinction coefficient (m^{-1}) solved for lidar ratios bounded by both 20 and 30 sr). Concurrently, pivotal attributes such as cloud boundary temperature and altitude and integrated optical depth exhibit noteworthy stability, punctuated only by minor seasonal shifts. This study also uncovers a persistent decline in surface albedo, with a derived trend of -0.00036 yr^{-1} . We further find that the interrelationship between CRE and surface albedo variation intensifies notably 10 during winter months. This leads to speculation that a decrease in the number of days of snow and ice is the main driver of the decrease in surface albedo. The decline in radiative flux at both the TOA and SFC can be perceived as a positive feedback loop that leads to increased atmospheric warming. The unveiled trends underscore the intricate synergy between albedo, radiative flux, and climate dynamics, pressing the need for vigilant monitoring of these shifts, given their profound implications for future climatic and circulatory phenomena.

15 1 Introduction

The Intergovernmental Panel on Climate Change (IPCC) was founded in 1988 by the World Meteorological Organization (WMO) with the support of the United Nations Environment Program. Its purpose is to examine scientific, technological, and socioeconomic data relevant to understanding climate change. More than three decades have passed since the first meeting in 1990, and the importance of clouds and aerosols in Earth's climate system remains a focal point in both the fifth and sixth 20 assessment reports (AR5 and AR6;(Stocker, 2014; Landi et al., 2021)). These reports highlight that clouds and aerosols sig-



nificantly influence the regulation of Earth's temperature and energy balance by reflecting and absorbing sunlight. The reports also discuss the challenges associated with accurately representing the behavior and interconnection of clouds and aerosols in climate models, as these processes can vary greatly depending on factors such as location and atmospheric conditions. The report emphasizes the need for a deeper understanding of the roles of clouds and aerosols in the climate system to improve 25 the accuracy of climate predictions. It stresses the importance of further research on the relationships among clouds, aerosols, and climate variations. AR6 also points out the importance of variations in cloud cover, as clouds are vital for controlling the global exchange of energy. Alterations in cloud patterns, whether due to interactions with aerosols or feedback mechanisms within clouds, can greatly influence how energy is distributed in the climate system.

Although numerous studies have evaluated short-term or geographically limited radiative impacts of cirrus clouds (Campbell 30 et al., 2016; Lolli et al., 2017a; Campbell et al., 2021), a comprehensive long-term evaluation spanning multiple decades at a single site has not yet been systematically carried out. To bridge this gap, this study presents the first detailed analysis of the two-decade-long record (2003-2022) at the NASA Micropulse Lidar Network (MPLNET) observational site in Greenbelt, Maryland, USA. The uniqueness and novelty of this analysis reside in the use of a continuous and homogeneous dataset in conjunction with rigorous statistical methodologies (Mann-Kendall and Sen slope analyses) to identify statistically significant 35 trends in the radiative effects of cirrus clouds at both the top of the atmosphere (TOA) and the surface (SFC). This broad temporal span not only enables the detection of subtle climatic patterns that emerge from intrinsic variability (Weatherhead et al., 1998b), but also provides an essential understanding of feedback processes specific to certain regions and their repercussions on global climate behavior. An essential part of understanding changes within a highly dynamic system such as the Earth's atmosphere involves a quantitative evaluation of trend significance, which is determined in relation to the precision of the 40 fundamental measurements (Weatherhead et al., 1998a). Consequently, careful consideration must be given to the duration and quality of the observations necessary to capture atmospheric trends of varying magnitudes (Weatherhead et al., 2002). Despite the importance of attributing changing conditions to the effects of climate variability, it can unfortunately be insufficient to simply establish long-term remote sensing platforms in the absence of a practical understanding of the challenges in identifying significant trends amid complex datasets (Weatherhead et al., 2018). Regarding the radiative and macrophysical characteristics 45 of clouds, the advancement of satellite technology, combined with the emergence of reliable ground-based profiling instruments, has led to the collection of more than two decades of measurements to evaluate the effects of climate change (Cermak et al., 2006).

This study aims to evaluate the long-term trends and variability of the radiative effects of cirrus clouds as well as their macrophysical properties. We focus on daytime conditions, since cirrus clouds can exert a net cooling or warming effect 50 depending on a host of internal and external factors (Stephens and Webster, 1981). Three previous studies by the same group provide snapshots of equivalent forcing and macrophysical properties at a series of different latitudes: Singapore (equatorial, (Lolli et al., 2017a)), Greenbelt, Maryland, USA (mid-latitude; (Campbell et al., 2016)) and Fairbanks, Alaska, USA (polar; (Campbell et al., 2021)). Therein, they describe how the daytime forcing of cirrus clouds exhibits a meridional gradient from warming near the equator to neutral or slight cooling near the poles.



55 The relative contribution of radiative effects of cirrus clouds is in direct contrast to the net cooling contribution overall (that is, day and night) of all clouds to the climate (near -20 W m^{-2} , for example (Ramanathan et al., 1989; Yi et al., 2017)). Better knowledge of the radiative behavior of cirrus clouds, their global variability and global-scale gradients, and composite physical properties, including dominant macrophysical properties, will improve numerical weather modeling on all scales, from day-to-day simulations Berry and Mace (2013) to climate-scale simulations Kärcher et al. (2006).

60 Nadir-pointing lidars have been routinely profiling the microphysical, optical, and geometric properties of cirrus clouds at different sites around the world for over 40 years (see, e.g., (Platt et al., 1987; Imasu and Iwasaka, 1991; Sassen et al., 2001; Sassen and Campbell, 2001)). These active optical remote sensing devices, developed in the first half of the last century, became more popular after the invention of CO_2 lasers (Ciofini et al., 2003), and are highly suitable for studying aerosols, clouds, and light precipitation Lolli et al. (2017b); Lolli (2021) due to their high spatial and temporal resolution and unique 65 spectral sensitivity. Cirrus clouds, with irregular particle sizes and shapes, have proven ideal targets for polarized lidar profiling ((Sassen, 2005)). Still, their early deployment and application was sporadic and widely distributed given their novelty, their highly-specialized technical nature, and the hazards of operating with an active laser source. The development and emergence of the eye-safe Micropulse Lidar system ((Spinhirne, 1993)) initiated the transition of lidar technology from an episodic remote sensing tool to continuous full-time measurements. At the turn of the century, NASA established the Micropulse Lidar Network 70 (MPLNET; Welton et al. (2001); Campbell et al. (2002); Welton et al. (2018)), its Earth Observing System (EOS) program ((Wielicki et al., 1995)). Now, more than 25 years into its existence, MPLNET maintains the longest continuous database of ground-based cirrus cloud observations on record.

75 The MPLNET database allows for a unique assessment of the long-term time evolution of cirrus cloud presence, their optical properties, and, through radiative transfer modeling, their radiative effects. This research aims to identify temporal trends related to global climate forcing. This study presents the most extensive long-term analysis of daytime cirrus cloud radiative effects using a continuous 20-year dataset from NASA MPLNET at GSFC, Maryland. Unlike previous studies that focused on shorter temporal scales or limited geographic regions, this research utilizes a unique combination of high-resolution ground-based lidar measurements and radiative transfer modeling to detect statistically significant trends in cirrus cloud radiative effects at both the top of the atmosphere (TOA) and the surface (SFC). The study also provides new information on the 80 interplay between cirrus clouds, surface albedo, and climate feedback mechanisms, revealing a persistent decline in radiative fluxes suggesting an evolving atmospheric energy balance. Using advanced statistical techniques, including the Mann-Kendall test and the Sen slope estimator, this work ensures robust trend detection despite natural variability. The findings have direct implications for improving climate models, as they highlight the importance of continued ground-based lidar monitoring for assessing long-term changes in cloud-radiation interactions.

85 Twenty years (2003-2022) of measurements were collected at the NASA Goddard Space Flight Center (GSFC) in Greenbelt, Maryland, USA (38.99 N, 76.85 W). The data record is normalized, when necessary, to avoid breaks in observations due to instrument failures and unrepresented months caused by adverse operating conditions (e.g., strong presence of fog and / or rain and / or water clouds). The significance of the temporal trends derived herein are characterized using both the non-parametric seasonal Mann-Kendall test in conjunction with the slope calculated using Sen's. Moreover, the significance of temporal trends



90 is determined by employing statistical methods that account for natural variability and measurement uncertainties (Weatherhead et al., 1998a). Our goal is to understand changes in cirrus clouds over the two-decade record collected at GSFC. However, we also want to understand the nature of the MPLNET record with respect to the significance of the trends derived therein. We seek to better understand the breadth of the MPLNET archive for climate study and wish to establish a baseline for the collection of lidar datasets supporting trend-based studies overall.

95 2 Methodology

We begin by describing the instruments and datasets used in the analysis. All input products are available through the MPLNET online data portal.

2.1 Instruments

The NASA MPLNET lidar network

100 The NASA MPLNET network has been operational since 1999 and consists of a collection of commercially available micropulse lidar (MPL) systems produced by LEICA Geosystems, Lanham, MD, USA (formerly SigmaSpace) (Welton et al., 2001). These MPL instruments (Spinhirne et al., 1995) are single-wavelength elastic lidars and are globally deployed to support the NASA Earth Observing System (EOS) program (Wielicki et al., 1995). MPLNET continuously monitors the atmosphere every 60 seconds from the surface up to 30 km. The adjustable vertical resolution of the profile software can be set in the range 105 of 0.030 to 0.075 km, depending on the station. The network operates under any meteorological conditions and to the limit of laser signal attenuation. Temporary and permanent observation sites are operational in various regions such as the polar, mid-latitude, tropical, and equatorial regions, to collect information on aerosol and cloud optical and geometric properties. The MPL system is usually co-located with NASA Aerosol Robotic Network (AERONET; Holben et al. (1998)) sunphotometers to reduce retrieval errors. The MPLNET products are publicly available website and follow the modified EOS convention, which 110 has three levels: Level 1, Level 1.5, and Level 2, all of which are available in near-real time (NRT). The primary difference between L2 and L1/15 files is that L2 may have additional post-calibrations applied as well as corrections to instrument temperatures. The only difference between the L1 and L1.5 products is that the data do not meet the L1.5 Quality Assurance (QA) criteria are screened and replaced with No Number (NaN) in the files.

Measurement Site

115 Twenty years of lidar-derived cirrus cloud optical property profiles were obtained from the MPLNET GSFC permanent observation site. The station is located in Greenbelt, MD, USA (38.99°N and 76.84°W) at an altitude of 50 m above sea level. The site is classified as urban, being part of the Washington-Baltimore metropolitan area.



2.2 Data Processing

This study presents for the first time the findings derived from analyzing such an extensive data set for TOA cloud forcing from 120 ground-based measurements. The instantaneous TOA CRE was standardized by multiplying its values by the daytime fraction obtained from the Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) onboard CALIPSO that operated from 2006-2023 (Vaughan et al., 2009; Winker et al., 2009)). The methodology previously reported in (Campbell et al., 2016; Lolli et al., 2017a; Campbell et al., 2021) was used to retrieve the cirrus cloud extinction profile and cloud optical depth (COD). In those 125 previous studies, lidar observations were used to assess a yearly and biannual analysis of CRE at three different observational sites deployed at three different latitudes: GSFC, MD, USA ((Campbell et al., 2016), Singapore (Lolli et al., 2017a) and Fairbanks, AK, USA (Campbell et al., 2021). These yearly analyses put into evidence that the daytime TOA net CRE shows latitudinal variability (South-North direction in the boreal hemisphere, North-South direction in the austral hemisphere) due to the different average solar zenith angles at the three different sites. The TOA net CRE is positive for equatorial and tropical regions, almost neutral at mid-latitudes, and slightly negative for polar regions.

130 **2.3 Retrieval methods**

2.3.1 Cirrus cloud optical depth

Single-layer cirrus clouds (that is, no aerosols, other cloud forms, or precipitation) are extracted from unconstrained MPLNET retrievals Campbell et al. (2008); Lewis et al. (2016, 2020) using lidar ratios of 20 and 30 sr to determine the upper and lower limits of the daytime cirrus CRE Yorks et al. (2011). These lidar ratios are then used to derive the profile of the cloud 135 extinction coefficient. The cloud layer height retrievals are carried out using the uncertainty-based cloud detection method (UCDM), which uses the uncertainty in the lidar signal, as explained by (Campbell et al., 2008). The presence of cirrus clouds is determined using the cloud top temperature threshold below -37°C and cloud base temperature below -25°C. The meteorological profiles required for radiative transfer, including temperature, specific humidity, and ozone, are derived from the Goddard Earth Observing System Model version 5 (GEOS-5) atmospheric general circulation model Rienecker et al. (2008); 140 Molod et al. (2012). Only cirrus cloud layers retrieved using the UCDM are reported for COD and extinction variables. For non-cirrus clouds and all layers identified using the gradient-based cloud detection method (GCDM), the COD and extinction are denoted by an IEEE floating-point value of Infinity. For this study, we focus only on optically thin cirrus clouds, whose column integrated optical depth is less than 3.0.

2.3.2 Cirrus cloud radiative effect

145 *The Fu-Liou-Gu Radiative Transfer Model*

The Fu-Liou-Gu (FLG) radiative transfer model was specifically chosen for this study because of its established accuracy and flexibility in simulating radiative processes involving clouds and aerosols. Compared to simpler single-scattering or two-stream



models, the FLG model effectively accounts for multiple scattering events and spectral variation of cloud and atmospheric constituents by employing the discrete-ordinate approach. This method significantly improves the representation of radiative interactions within optically thin cirrus clouds, a critical factor given their complex optical characteristics. Moreover, previous validation studies demonstrated that FLG reliably captures both direct and diffuse radiative components, essential for accurately quantifying the subtle radiative effects typical of cirrus clouds. These advantages make the FLG model particularly suitable for examining long-term trends in cloud radiative effects and assessing their implications for climate studies, thus motivating its selection for this analysis.

FLG model is a widely used tool for simulating the transport of broadband solar and infrared radiation through the atmosphere. It was developed by Fu and Liou (1992, 1993), and subsequently extended by Gu et al. (2003). The FLG model solves the radiative transfer equation using a discrete ordinate approach to represent the angular distribution of incoming and outgoing radiation. This method allows for both direct sun irradiance and diffuse sky radiation to be considered. In addition, scattering between different levels in the atmosphere can be considered, which greatly improves the accuracy of simulations compared to other simpler single-scattering models.

The FLG model works by assuming that incoming solar radiation is divided into 18 total spectral bands (6 in the solar (0.2 - 4 μm) and 12 in the infrared (4.5 - 1000 μm)) and the outgoing terrestrial radiation is divided into four flux streams (visible, short-wave infrared, long-wave infrared, and total reflected solar radiation). These fluxes are then calculated on the basis of their reflectivity or absorptivity depending on the properties of different layers within the atmosphere such as aerosols, clouds, or other components like water vapor or ozone. To accurately simulate atmospheric processes, the FLG model requires detailed knowledge of atmospheric properties such as temperature, humidity, and ozone profiles, aerosol concentrations, and optical thicknesses in each layer of air. These parameters are then used to compute various quantities such as transmittance across each layer, extinction coefficients due to scattering or absorption by aerosols and cloud droplets/ice crystals, surface reflectance, etc., all necessary components that influence how solar energy is partitioned within an atmosphere (Fu and Liou (1992, 1993); Gu et al. (2003)). The FLG model has been successfully applied in numerous studies to assess the radiative effect of clouds and aerosols Tosca et al. (2017); Lolli et al. (2019); Landi et al. (2021); Dolinar et al. (2024). In all of those cases, instantaneous radiative effects are computed by directly injecting the 1-minute aerosol/cloud profiles at 1-minute temporal resolution and 0.075 km spatial resolution into the FLG model. FLG then outputs the instantaneous upward and downward solar and infrared fluxes from which the radiative effects can be calculated at the TOA and surface.

The absolute daytime TOA net CRE is found as:

$$\text{Absolute daytime TOA net CRE} (W \cdot m^{-2}) = \sum_{\text{COD}=0}^3 \text{COD_RF}_i \times \overline{\text{CRF}}_i \times \text{CF} \times \text{DF} \quad (1)$$

where

- COD_RF – cloud optical depth relative frequency



– $\overline{\text{CRF}}$ – average cloud radiative forcing

180 – CF – cloud fraction (15-year CALIPSO record)

– DF – daytime fraction

2.4 Daytime cirrus cloud dataset statistical analysis

The optical properties of the daytime cirrus clouds (for example, the vertical extinction coefficient and the optical depth) were obtained from the MPLNET cloud product version 3 (CLD; (Lewis et al., 2016)). Consistently with (Campbell et al., 2016),

185 we define daytime clouds as those when the incoming solar flux is greater than the total (solar + IR) outgoing clear-sky flux at the TOA. In addition, to achieve a uniform database with balanced monthly representations, we visually inspect and verify the data for any breaks that might be attributed to various factors, including inaccuracies in the initial data collection, modifications in measurement techniques, or malfunctions in equipment. The CRE is not a continuous variable, such as temperature, due to its dependence on the existence of cirrus clouds. In this study, we apply the seasonal Mann-Kendall nonparametric test along 190 with Sen's slope to the monthly averaged daytime cirrus CRE and optical depth from 2003 to 2022. Given that cirrus clouds do not appear consistently, the level of autocorrelation remains minimal, as illustrated in Figure 1.

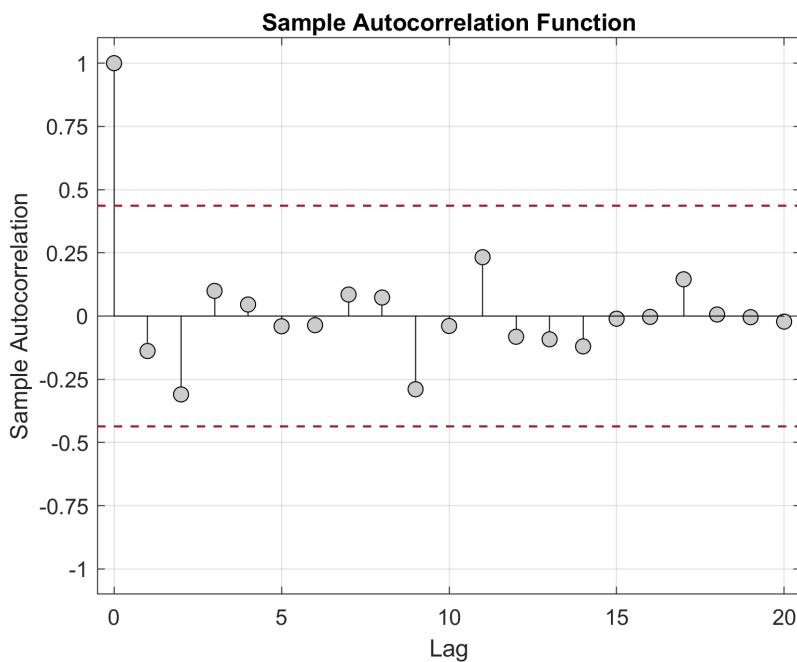


Figure 1: Averaged CRE showing no autocorrelation

Again, cirrus clouds are defined here using a cloud top temperature threshold (-37°C) based on the NASA GEOS-5 meteorological product (Campbell et al., 2015) and unconstrained extinction coefficient retrievals (for cirrus clouds) using bookend



lidar ratios of 20 sr and 30 sr. The COD is computed as the product of the lidar spatial resolution and the atmospheric extinction profile. If corresponding to a daytime cirrus cloud, the extinction profile is then input into FLG to compute the TOA and SFC CREs. The monthly average optical properties of the cloud (and the surface albedo) are calculated 85% over at least 1000 samples, as shown by the histogram of Figure 2.

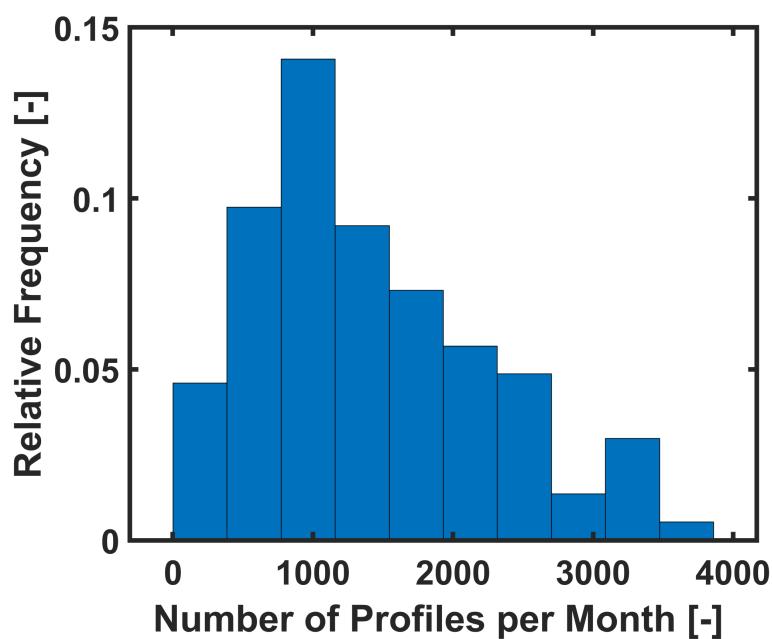


Figure 2: Month-by-month bin average histogram. Most of the months are the result of an average of at least 1000 cloud samples (85%)

Figure 3 illustrates a comprehensive statistical analysis of three key cloud-related parameters, cloud top temperature, cloud top altitude, and cloud geometric depth averaged over different temporal windows for the 20-year database estimated from the 30sr solution. The histograms display the relative frequency distributions of these parameters for the annual average and for each season: DJF (December, January, February), MAM (March, April, May), JJA (June, July, August), and SON (September, October, November). The first row of histograms shows the distribution of cloud top temperature, which ranges from approximately -80°C to -30°C. The distribution is roughly Gaussian in shape, with a peak around -55°C. There are subtle variations across seasons, indicating that while cloud top temperatures are relatively stable on an annual basis, there are minor seasonal fluctuations. The second row represents the distribution of the cloud top altitude, spanning from around 6 km to 18 km. These histograms also follow a Gaussian-like distribution, reaching a peak at approximately 13-14 km. Similarly to temperature, seasonal differences in altitude are minimal, suggesting a consistent vertical structure of clouds throughout the year. The third row focuses on the depth of the cloud, showing a different distribution compared with the other parameters. The distribution of cloud depth is skewed towards lower values, with a rapid decline in frequency as depth increases from 0 to 7 km. This indicates



210 that shallow clouds dominate the dataset. Although the proportion of shallow versus deep clouds shows slight seasonal variation, the overall pattern remains consistent. In summary, the data suggest that cloud characteristics such as top temperature and altitude are relatively stable across different temporal windows, with only minor seasonal variations, whereas cloud depth exhibits a skewed distribution that consistently favors shallower clouds throughout the 20-year period.

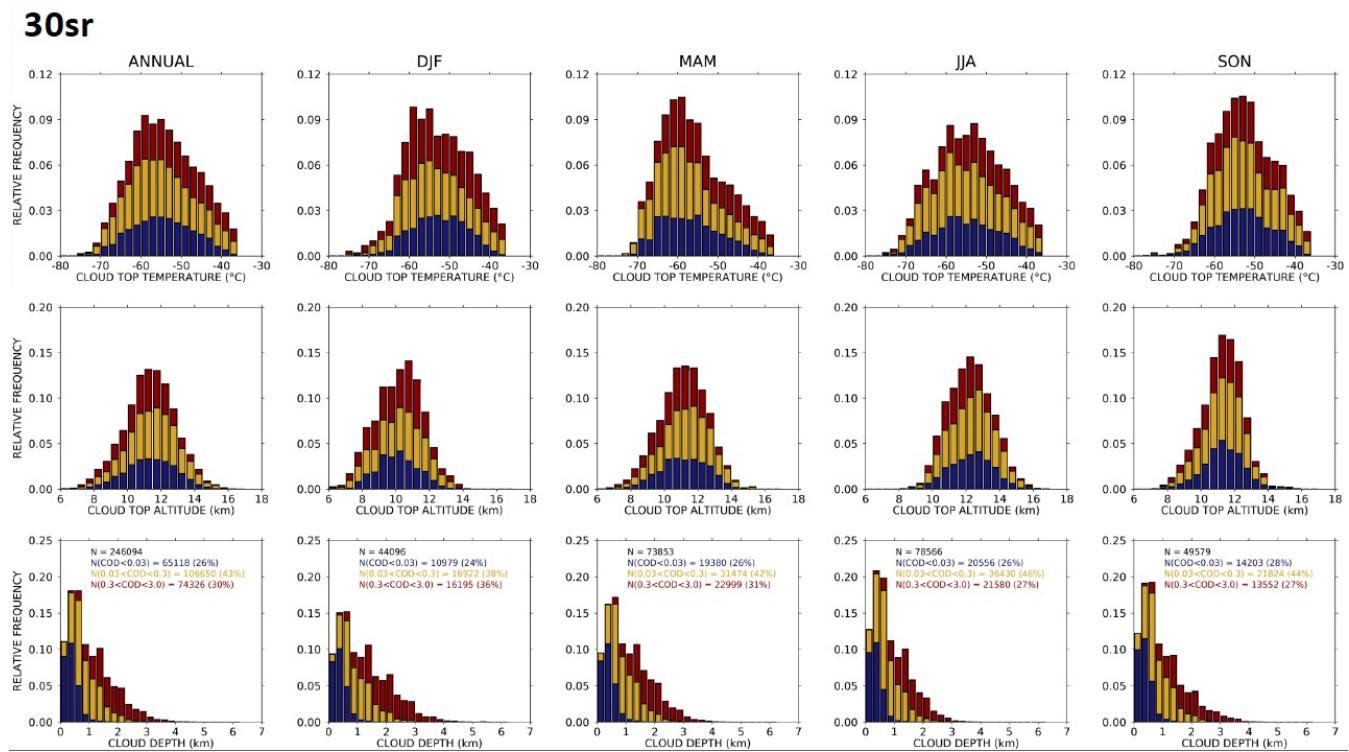


Figure 3: Seasonal distributions of cloud top temperature, altitude, and depth, estimated using a 30 sr lidar ratio, over a 20-year period. Cloud top temperature peaks around -55 °C, cloud top altitude around 13-14 km, and cloud depth is predominantly less than 2 km, showing minimal seasonal variations across all metrics.

2.5 Surface albedo

215 The albedo is a measure of surface brightness, and its significance in radiative transfer models stems from the fact that it influences how much incoming solar radiation is reflected by the surface. Albedo is defined as the percentage of incoming solar energy reflected by the Earth's surface (Ineichen et al., 1990). The quantity of energy absorbed by a medium is determined by its albedo; for example, surfaces with higher albedos, such as snow-covered areas, absorb less energy than surfaces with lower albedos, such as darker soils. The FLG radiative transfer model requires surface albedo as input to compute cloud heating rates and radiative effects. Changes in surface albedo produce significant effects on the Earth-atmosphere radiative budget. In 220 this study, for broadband solar reflection, the monthly mean gridded surface albedos (with 1 ° x 1 ° resolution) of the Clouds



and Earth Radiant Energy System (Wielicki et al., 1995) SYN1deg, version 4 product are utilized. The surface IR emission is considered to be a constant of 0.98.

3 Long-Term Trend Analysis

225 The optimal method for identifying temporal trends is a non-parametric, rank-based approach, well suited for managing missing data, negative values, and values below detection limits without relying on a specific statistical distribution. In this study, we employ the Mann-Kendall (MK) test to uncover long-term monotonic trends, either upward or downward, and use the Sen's slope estimator to determine the slope and its 95% confidence limits from the median of slopes calculated from all possible data pairs. Due to distinct seasonal patterns in many time series, the modified seasonal MK test (?) is consistently
230 used to assess trends over a 12-month period. The MK test is suitable for small samples that are often found in environmental studies. Furthermore, it is robust against heteroskedasticity, which refers to the condition in which the variance of errors, or the variability of a variable, is unequal across the range of values of a second variable that predicts it and lacks the assumption of normality (Mann, 1945; Kendall, 1948). Originally proposed as an extension of the Spearman rank correlation test, it can identify both linear and non-linear time-based changes unlike Spearman's rho, which only detects linear relationships (Yue
235 et al., 2002). Although the MK test has greater statistical power than other trend tests, it is sensitive to autocorrelation, which can increase the probability of false positives by incorrectly rejecting the null hypothesis of no trend (Collaud Coen et al., 2020). In our analysis, no data correction was applied as there was no evidence of autocorrelation (that is, Figure 1).

3.1 Minimum Detectable Trend

240 The minimum detectable trend analysis was performed to assess the detectability of trends in various atmospheric and surface parameters from the provided data set. The Weatherhead method (Weatherhead et al., 1998b) considers the noise level and autocorrelation structure of the time series to determine the smallest trend that can be reliably detected during the observation period. The following steps were carried out for each parameter:

1. **Calculation of Lag-1 autocorrelation (ϕ):** This measures the correlation of the time series with a one-month delay, indicating the persistence of the signal over time.
- 245 2. **Standard Deviation of Residuals (σ):** The residuals were calculated by subtracting the mean from each data point, and the standard deviation of these residuals was calculated. This provides an estimate of the variability of the data around its mean.
3. **Number of Data Points (n):** The total number of observations was counted, which in this dataset is 212, representing monthly measurements over approximately 17.67 years.
- 250 4. **Time Span (T):** The total time span of the data, calculated by dividing the number of data points by 12, gives the duration of the dataset in years.



5. Minimum Detectable Trend (MTD): Using the formula

$$MTD = \frac{\sigma\sqrt{1+\phi}}{\sqrt{n} \times T} \quad (2)$$

where σ is the standard deviation, ϕ is the lag-1 autocorrelation, n is the number of data points, and T is the time interval in which the minimum detectable trend was calculated. This value represents the smallest trend that can be detected with the given data, taking into account its variability and autocorrelation.
255

4 Results

This analysis explores the trends of daytime cirrus cloud CREs, optical depth, surface albedo, and solar zenith angle (SZA) over the past two decades from 2003 to 2022 at NASA GSFC in Greenbelt, MD, USA. The results of the MK analysis are
260 shown in the next section. We did not implement any other trend analysis techniques as the MK test performs better, as shown by Collaud Coen et al. (2020).

The observed negative trends, particularly pronounced during the winter and summer months, indicate an evolving atmospheric energy balance that could contribute to regional warming through positive feedback mechanisms. Concurrently, a significant decline in surface albedo, driven likely by reductions in snow and ice coverage, further amplifies these implications,
265 underscoring the climatological significance of the observed radiative trends. Detailed statistical findings from the Mann-Kendall trend test and Sen slope values are described in the following subsections.

4.1 Cirrus cloud atmospheric radiative effects

In this section, we perform a long-term trend analysis of the atmospheric radiative effects of cirrus clouds during the day (see, e.g., Figure 4). A positive (negative) TOA CRE indicates warming (cooling) of the Earth-atmosphere system. The seasonal
270 MK test is performed on monthly-averaged daytime CRE observations. We consider trends statistically significant at the 95% confidence level ($p < 0.05$). The results presented in Table 1 provide a comprehensive summary of the temporal trends of various atmospheric and surface parameters using the Mann-Kendall trend test. The variables analyzed include TOA and SFC net CREs and COD estimated from 20sr and 30sr lidar ratios. The results of the Mann-Kendall trend analysis indicate a statistically significant decreasing trend over 20 years in TOA20sr, TOA30sr, SFC20sr and SFC30sr ($p < 0.05$), confirming the presence of
275 a negative trend in these radiation components. This implies that the net radiative effect during the day, both at the TOA and surface, is decreasing over time, which could have implications for changes in atmospheric and surface energy balances. The associated Sen's slopes are then negative, indicating the annual rate of reduction. For example, the TOA30sr exhibits a Sen slope of $-0.031 \text{ W m}^{-2} \text{ yr}^{-1}$.

4.2 Other parameters

280 Parameters such as cloud top temperature (CTT), cloud base temperature (CBT), cloud top height (CTH) and cloud base height (CBH) do not show statistically significant trends, as evidenced by p-values that exceed the threshold of 0.05. This suggests

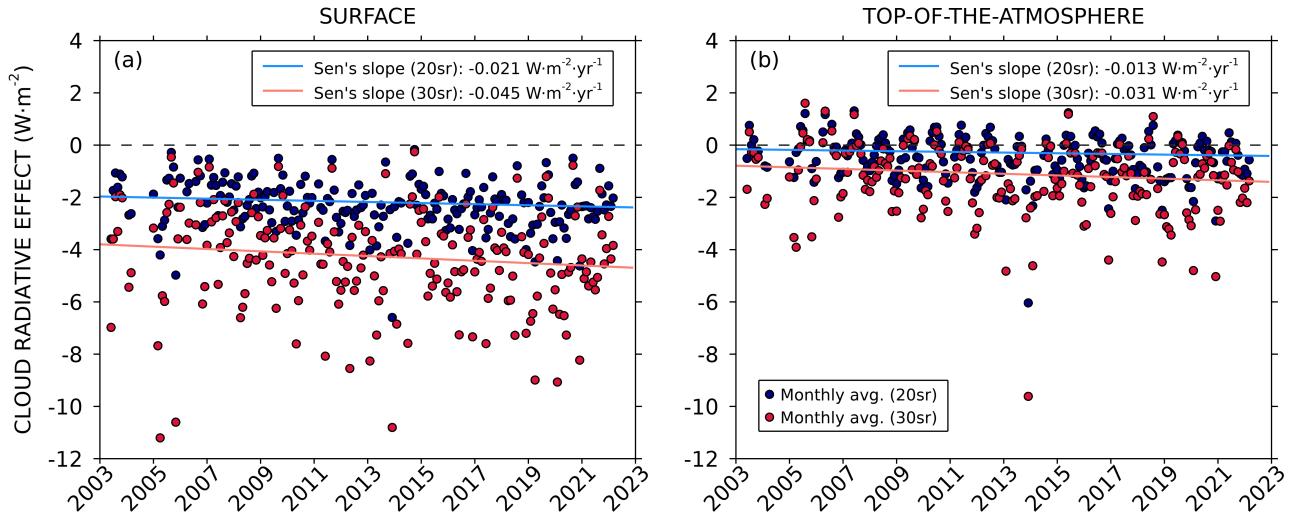


Figure 4: Temporal trends in CRE at the (a) SFC and (b) TOA from 2003 to 2022 using (blue) 20sr and (red) 30sr. Each point is the monthly average value with the Sen's slope overlaid.

the absence of notable alterations in cloud characteristics, including both temperature and altitude. However, SZA and cloud occurrence (i.e. count) show an increasing trend with a Mann-Kendall p-value of 0.021 and 0.048, respectively. The observed increasing trend in SZA, characterized by a Sen's slope of 0.23 yr^{-1} , may indicate, together with the count data, a significant

285 improvement in the MPL cloud detection capabilities. In particular, in 2010, the instrument was upgraded to a newer version that incorporated an enhanced signal-to-noise ratio. An increased SZA correlates with a higher solar background, which lowers the signal-to-noise ratio.

Overall, the decreasing trends in shortwave radiation components suggest changes in atmospheric transparency or cloud cover that might affect Earth's radiative balance. The unchanged optical properties of the clouds together with their altitude

290 suggest that these factors may not be the primary drivers of the observed radiative trends.

4.3 Cirrus cloud optical depth

We performed the seasonal MK test with Sen's slope estimation to assess long-term trends in cloud optical depth at both 20sr and 30sr (not shown). The results show that no discernible long-term monotonic global trend could be identified with confidence 95%, as evidenced by the p-values of 0.22 and 0.19, respectively. This implies that the optical properties of the

295 cirrus clouds remained relatively constant throughout the 20-year period.



Table 1: Mann-Kendall Test and Sen's Slope Results for 20-Year Trends Across Variables

Variable	MK Trend	MK P-value	Sen's Slope	Z-value	S-value
TOA20sr	decreasing	0.012	$-0.017 \text{ W m}^{-2} \text{ yr}^{-1}$	-2.44	-118
TOA30sr	decreasing	0.020	$-0.035 \text{ W m}^{-2} \text{ yr}^{-1}$	-2.31	-111
SFC20sr	decreasing	0.031	$-0.027 \text{ W m}^{-2} \text{ yr}^{-1}$	-2.15	-104
SFC30sr	decreasing	0.042	$-0.048 \text{ W m}^{-2} \text{ yr}^{-1}$	-2.05	-99
COD20sr	<i>no trend</i>	0.22	N/A	-1.24	-60
COD30sr	<i>no trend</i>	0.19	N/A	-1.31	-63
Albedo	decreasing	0.035	-0.00036 yr^{-1}	-2.15	-104
CTT	<i>no trend</i>	0.55	N/A	0.61	29
CBT	<i>no trend</i>	0.35	N/A	-0.94	-45
CTH	<i>no trend</i>	0.43	N/A	0.80	38
CBH	<i>no trend</i>	0.39	N/A	0.86	41
SZA	increasing	0.021	0.23 yr^{-1}	2.26	109
Counts	increasing	0.048	30.56 yr^{-1}	1.90	92

4.4 Surface albedo

Analysis of the monthly mean CERES surface albedos at GSFC showed a 20-year global negative trend at a confidence level 95%. The results of the global trend are shown in Figure 5. Surface albedo directly influences the daytime radiative transfer calculations, where a decreasing albedo translates into more absorbed radiation by the surface, reducing the reflected shortwave 300 flux at both the surface and at the TOA.

4.5 Climatological significance of the trends

We evaluated the minimum detectable trend for statistically significant MK trends. The results are depicted in Table 2

4.6 Seasonal Analysis

The variability of statistically significant temporal trends over 20 years of measurements is also analyzed by season. The 305 complete database is divided into four main seasons, December-January-February (DJF), March-April-May (MAM), June-July-August (JJA) and September-October-November (SON).

Seasonal analyses at both the SFC and TOA (Table 3) indicate consistent negative trends throughout all seasons, suggesting a reduction in radiative flux over the years. For SFC variables, during the winter months (DJF), SFC20sr shows a decreasing trend with a Sen's slope of $-0.053 \text{ W m}^{-2} \text{ yr}^{-1}$, while SFC30sr exhibits a stronger decline with a Sen's slope of $-0.12 \text{ W m}^{-2} \text{ yr}^{-1}$, 310 indicating a more pronounced reduction. In spring (MAM), SFC20sr shows a slight negative trend with a Sen's slope of $-0.018 \text{ W m}^{-2} \text{ yr}^{-1}$, while SFC30sr has a more substantial decrease of $-0.056 \text{ W m}^{-2} \text{ yr}^{-1}$. The summer season (JJA) shows

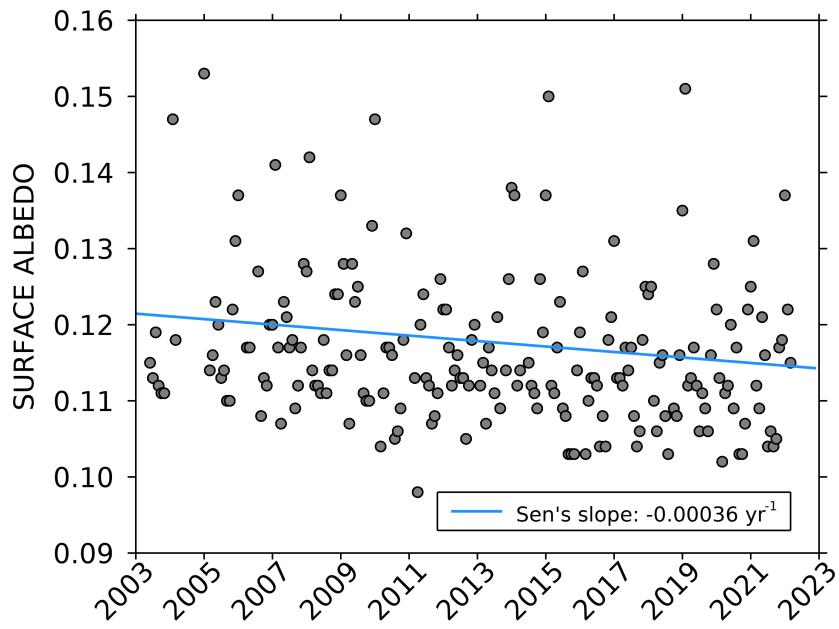


Figure 5: Surface albedo 20-years trend analysis is statistically significant with a negative Sen's slope of -0.00036 yr^{-1}

Table 2: The values show the lag-1 autocorrelation, standard deviation, and minimum detectable trend for the selected parameters in the dataset.

Parameter	Lag-1 Autocorrelation (φ)	Standard Deviation (σ)	Minimum Detectable Trend
TOA20sr	0.57	0.85 W m^{-2}	$0.0041 \text{ W m}^{-2} \text{ yr}^{-1}$
TOA30sr	0.49	1.32 W m^{-2}	$0.0063 \text{ W m}^{-2} \text{ yr}^{-1}$
SFC20sr	0.21	0.92 W m^{-2}	$0.0039 \text{ W m}^{-2} \text{ yr}^{-1}$
SFC30sr	0.22	1.89 W m^{-2}	$0.0081 \text{ W m}^{-2} \text{ yr}^{-1}$
Albedo	0.25	0.01	0.00012 yr^{-1}
SZA	0.70	9.27 deg	0.05 deg yr^{-1}
Count	0.49	802.77 counts	$3.80 \text{ counts yr}^{-1}$

a moderate negative trend for SFC20sr with a Sen's slope of $-0.035 \text{ W m}^{-2} \text{ yr}^{-1}$, while SFC30sr continues to demonstrate a stronger decline of $-0.0630 \text{ W m}^{-2} \text{ yr}^{-1}$. In the fall months (SON), SFC20sr exhibits a small negative trend with a Sen's slope of $-0.0169 \text{ W m}^{-2} \text{ yr}^{-1}$, and SFC30sr has a comparable downward trend with a slope of $-0.0360 \text{ W m}^{-2} \text{ yr}^{-1}$. For the TOA variables, during DJF, TOA20sr shows a negative trend with a Sen's slope of $-0.0276 \text{ W m}^{-2} \text{ yr}^{-1}$, while TOA30sr exhibits a steeper decline with a slope of $-0.1054 \text{ W m}^{-2} \text{ yr}^{-1}$.



In MAM, both TOA20sr and TOA30sr show similar decreasing trends with slopes of $-0.0435 \text{ W m}^{-2} \text{ yr}^{-1}$ and $-0.053 \text{ W m}^{-2} \text{ yr}^{-1}$, respectively. Summer (JJA) shows negative trends with TOA20sr at $-0.0365 \text{ W m}^{-2} \text{ yr}^{-1}$ and TOA30sr at $-0.0645 \text{ W m}^{-2} \text{ yr}^{-1}$. In SON, the trends continue with TOA20sr at $-0.0354 \text{ W m}^{-2} \text{ yr}^{-1}$ and TOA30sr at $-0.047 \text{ W m}^{-2} \text{ yr}^{-1}$.
 320 The observed decreasing trends in daytime net radiative fluxes at both the TOA and the surface emphasize potential shifts in regional climate dynamics, particularly during winter and summer seasons. The pronounced decrease in surface albedo, especially during the winter months, suggests a decrease in snow cover, likely driven by regional warming, which could amplify surface warming through positive feedback mechanisms. The significant decrease in surface albedo, especially during winter months, further underscores potential shifts in local climate systems, possibly accelerating snow and ice melt and influencing
 325 broader environmental conditions.

Overall, the analysis indicates that both surface and TOA cirrus CREs are decreasing consistently across all seasons, with the reductions generally being more pronounced at 30 sr compared with the 20 sr solution.

The SFC variables generally show stronger seasonal trends than the TOA, especially in winter (DJF), where the slope of SFC30sr is much steeper than the slope of TOA30sr. This could indicate a greater variability or sensitivity in surface radiative
 330 flux compared to the TOA flux in the winter season. In other seasons like MAM and SON, the trends for both SFC and TOA variables are more aligned in magnitude, suggesting more uniform changes in radiative fluxes across different atmospheric layers. Overall, while both SFC and TOA show consistent negative trends, the magnitude of these trends is more variable at the surface level, particularly for the SFC30sr variable, which shows the steepest seasonal declines across the board.

Table 3: Mann-Kendall Test and Sen's Slope Results for 20-Year seasonal trends Across CRE at TOA and SFC.

Season	SFC20sr [$\text{W m}^{-2} \text{ yr}^{-1}$]	SFC30sr [$\text{W m}^{-2} \text{ yr}^{-1}$]	TOA20sr [$\text{W m}^{-2} \text{ yr}^{-1}$]	TOA30sr [$\text{W m}^{-2} \text{ yr}^{-1}$]	SZA [deg yr^{-1}]	Albedo [yr^{-1}]
DJF	-0.053	-0.120	-0.033	-0.065	0.049	-0.00064
MAM	-0.018	-0.056	-0.012	-0.026	0.155	-0.00024
JJA	-0.035	-0.063	-0.023	-0.047	0.260	-0.00040
SON	-0.017	-0.036	-0.013	-0.022	0.210	-0.00035

Seasonal analyses of surface albedo and SFC/TOA cirrus CREs reveal similarities and differences in their trends between
 335 seasons. Albedo shows a consistent negative trend in all seasons, indicating a decrease in reflectivity over time. This aligns with the general reduction in the surface and TOA radiative fluxes. In winter (DJF), Albedo has a negative trend of -0.00064 yr^{-1} , which corresponds to significant reductions in SFC20sr and SFC30sr (-0.05 and $-0.12 \text{ W m}^{-2} \text{ yr}^{-1}$, respectively) and substantial decline in TOA30sr ($-0.07 \text{ W m}^{-2} \text{ yr}^{-1}$). This suggests that a decrease in reflectivity could be contributing to the reduction in radiative flux at both surface and atmospheric levels. In spring (MAM), the Albedo trend is less steep at
 340 -0.00025 yr^{-1} , while SFC and TOA also show less pronounced decreases, with SFC20sr and SFC30sr at -0.02 and $-0.06 \text{ W m}^{-2} \text{ yr}^{-1}$, respectively, and TOA20sr and TOA30sr at -0.01 and $-0.03 \text{ W m}^{-2} \text{ yr}^{-1}$. This partial agreement indicates a weaker correlation between changes in Albedo and radiative flux during spring. In summer (JJA), the Albedo trend remains



negative at -0.00040 yr^{-1} , with SFC and TOA showing similar decreasing trends, particularly for SFC30sr and TOA30sr (-0.06 and $-0.05 \text{ W m}^{-2} \text{ yr}^{-1}$, respectively), suggesting a more direct relationship between reduced reflectivity and declining radiative fluxes. Finally, in autumn (SON), the Albedo trend is -0.00036 yr^{-1} , which is consistent with the milder decreases in the SFC20sr and SFC30sr variables (-0.02 and $-0.04 \text{ W m}^{-2} \text{ yr}^{-1}$) and the TOA variables (-0.01 and $-0.02 \text{ W m}^{-2} \text{ yr}^{-1}$). This indicates a more coherent trend among all variables during the fall. Overall, while there is general agreement between Albedo, SFC, and TOA trends in each season, the strength of their correlation varies, with a stronger alignment in winter and summer and a partial agreement in spring and fall.

The solar zenith angle has a large influence on the radiative effects of the cirrus cloud, so understanding how this angle changes in reference to the long-term trend is crucial. We mentioned previously that MPLNET instrument designs changed in 2010, thus improving the signal-to-noise ratio and increasing cloud detection. Table 3 supports this technical advancement, where the seasonal trends in the SZA are all increasing. We would be remiss to not mention that the increasing trends in SZA may also be due to a systematic behavior change in cloud presence over the diurnal cycle. Quantifying this distinction is beyond the scope of this study. Trends in SZA are largest during the summer months (JJA) when there are more available daylight hours and the sun is higher in the sky.

4.7 Discussion

Seasonal analyses of surface albedo and cirrus cloud radiative effects (CRE) at the surface (SFC) and top-of-atmosphere (TOA) reveal consistent negative trends across all variables, suggesting a general reduction in net radiative flux over the last 20 years. Albedo, which represents Earth's surface reflectivity, shows a persistent decrease across all seasons, with the most significant reductions observed during winter (DJF) and summer (JJA). This declining trend in albedo indicates that less solar radiation is being reflected back into the atmosphere, leading to increased absorption by the surface and contributing to the observed decreases in cirrus CRE at both the SFC and TOA. During winter, the reduction in albedo, particularly at -0.00064 yr^{-1} , is accompanied by substantial decreases in SFC and TOA CRE. The SFC20sr and SFC30sr variables show decreases of -0.05 and $-0.12 \text{ W m}^{-2} \text{ yr}^{-1}$, respectively, while the TOA20sr and TOA30sr variables show reductions of -0.03 and $-0.07 \text{ W m}^{-2} \text{ yr}^{-1}$.

This alignment of trends suggests that the decreased reflectivity, likely due to less snow and ice cover, plays a significant role in the observed radiative flux changes. As snow and ice are highly reflective, their reduction in winter decreases albedo, allowing more solar energy to be absorbed by the surface, subsequently lowering the outgoing radiative flux at both SFC and TOA. The drop in the radiative effect at the surface and TOA has important implications for weather and climate alike. A decrease in albedo and subsequent reduction in radiative flux can lead to surface warming, which may accelerate snow and ice melt, creating a feedback loop that further reduces albedo and increases surface temperatures. This process can alter local and regional weather patterns, potentially leading to warmer winters with less snow cover and more variability in temperature and precipitation. Furthermore, changes in radiative flux at the TOA can influence large-scale atmospheric circulation patterns, potentially impacting weather systems and contributing to more extreme weather events. The identified trends underscore the intrinsic link between albedo, radiative flux, and climate dynamics, highlighting the importance of attentive observation of these alterations due to their significant impact on future climate and weather systems.



Although our findings provide robust evidence for statistically significant radiative trends, the analysis is constrained by certain methodological limitations. Specifically, this study uses observations from a single site in Greenbelt, Maryland, which may not fully represent broader global trends or variability. To mitigate these limitations, future research should incorporate 380 additional MPLNET sites in varying geographic and climatic regions and employ multilayer cloud radiative transfer models to better capture the complexity of cloud-radiation interactions. To further address the methodological limitations identified in this analysis, future studies should analyze data from additional MPLNET sites that cover diverse geographical and climatic regions. This broader spatial perspective would provide a clearer picture of global variability and trends in radiative effects of 385 the cirrus cloud. Moreover, incorporating multilayer cloud and aerosol radiative transfer modeling could significantly enhance the accuracy and realism of cloud-radiation interactions in climatological assessments. Finally, targeted in situ measurements and remote sensing studies that better characterize the microphysical properties of cirrus clouds, particularly ice crystal shape and size distributions, would improve the reliability of radiative forcing estimates and further refine their representation in climate models.

5 Conclusions and future perspectives

390 In this paper, the results of a 20-year trend analysis for the radiative effects and optical properties of daytime optically thin cirrus clouds were reported. This analysis provides robust evidence of significant declining trends in daytime cirrus cloud radiative effects and surface albedo over a 20-year observational period, underscoring an evolving atmospheric energy balance with substantial climatological implications. The persistent decrease in radiative flux at both the top of atmosphere (TOA) and 395 surface (SFC), coupled with the significant reduction in surface albedo, likely driven by decreasing snow and ice coverage, suggests the presence of positive climate feedback mechanisms, potentially intensifying regional warming and altering atmospheric circulation patterns. These findings highlight the crucial role of continuous ground-based lidar observations, such as those provided by MPLNET, in the accurate monitoring and quantification of long-term atmospheric changes. In addition, the insights gained through this analysis underline the importance of improving cloud representation in climate models, emphasizing the need for improved observational constraints to enhance future climate projections and inform climate adaptation 400 strategies.

The results discussed in this study are based on the longest record of ground-based cirrus cloud measurements, from 2003 to 405 2022, taken from the NASA Micropulse Lidar Network at Goddard Space Flight Center, Greenbelt, MD, USA. Herein, cirrus clouds are defined with a cloud top temperature threshold ($< -37^{\circ}\text{C}$), an optical depth threshold (< 3.0), and lidar-derived extinction coefficients are constrained using 20sr and 30 sr lidar ratios. Each 1-min average single layer daytime cirrus cloud extinction coefficient profile was injected into the Fu-Liou-Gu radiative transfer model to compute its cloud radiative effect (CRE, W/m^2) both at the top of the atmosphere (TOA) and at the surface (SFC). The instantaneous radiative effects were then normalized by the daytime fraction and cloud fraction obtained from the Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) instrument onboard of Cloud-Aerosol Lidar and Infrared Pathfinder Satellite Observations (CALIPSO) satellite observations. To perform a long-term trend analysis and reduce noise, the instantaneous CRE values were averaged monthly,



410 from which the seasonal Mann-Kendall statistical test with Sen's slope were performed. The trends in CRE are statistically significant at the confidence level 95% in both TOA and SFC and at 20 sr and 30 sr. At the TOA there is a net decrease of flux of -0.017 (20sr) and -0.035 (30sr) W/m^2 per year and -0.027(20sr) and -0.048 (30sr) W/m^2 per year at the SFC. Moreover, the surface albedo exhibits a statistically significant seasonal Mann-Kendall test over the two decades. The surface albedo decreased at a rate of -0.00036 $year^{-1}$. Lower albedos increase the absorbed energy of the surface, which translates to reduced
415 reflected shortwave radiative at both TOA and SFC.

420 Although crucial for assessing a climatologically significant trend in the net radiative effect of cirrus clouds during the day, this study shows some limitations due to the proposed methodology. For one, the radiative effects of aerosols (and other clouds) are not taken into account. The Fu-Liou-Gu radiative transfer model computations are instantaneous and only run on single-layer cirrus clouds, which, of course, limits the accuracy in assessing their true radiative effects within the entire climate
425 system. Furthermore, there are many unknowns regarding the microphysical and optical properties of cirrus clouds, such as the size and shape of ice crystals, which can influence their radiative behavior (see, e.g., (Dolinar et al., 2024)). Cirrus clouds are the only cloud type that, during the daytime, might cool or warm the Earth atmosphere system depending on their location, optical thickness, and underlying surface properties. For this reason, the scientific community recommends further research and continued monitoring to better understand the role of cirrus clouds in the broader context of climate change and to reduce uncertainty in our understanding of their impact.

430 Our results align closely with the themes highlighted in recent IPCC assessment reports, highlighting the intricate interplay between clouds, aerosols, and radiative forcing. These reports underscore the need for improved clouds representation in weather and climate models due to their substantial influence on the global energy balance. Given this, ongoing investment in long-term observational networks such as MPLNET remains critical, providing essential datasets to validate and improve weather forecasts and climate predictions. Continued efforts to refine cloud parameterization and sustained observational efforts are vital to improving our understanding and predictions of climate change and guiding effective policy interventions.

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