

A new method to retrieve relative humidity profiles from a synergy of Raman LiDAR, microwave radiometer and satellite

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Abstract Precise continuous measurements of relative humidity (RH) vertical profiles in the troposphere have emerged as a considerable scientific issue. In recent years, a combination of diverse ground-based remote sensing devices has effectively facilitated RH vertical profiling. This work introduces a newly developed approach for obtaining continuous RH profiles by integrating data from a Raman LiDAR, a microwave radiometer, and satellite sources. RH profiles obtained using synergistic approaches are subsequently compared with radiosonde data throughout a five-month observational study in China. Our suggested method for RH profiling demonstrates optimal concordance with the best correction coefficients R of 0.94 in Huhehaote (HHHT), 0.92 in Yibin (YB) and 0.93 in Qingyuan (QY), respectively. The mean value of RH

decreased with height and presented seasonal characteristics in QY. Finally, the RH height-time evolution was used to distinguish hail and heavy precipitation. This study firstly integrates satellite data into ground-based measurement to retrieve RH vertical profiles in China.

Keywords: relative humidity profiles, Raman LiDAR, microwave radiometer, satellite

1. Introduction

Relative humidity (RH) is a crucial parameter in characterizing aerosol-cloud interactions (Fan et al., 2007). It is also necessary as input for weather forecasting models (Petters and Kreidenweis, 2007; Wex et al., 2008; Mochida, 2014). The combination of these RH profiles with aerosol optical data allows us to obtain hygroscopic growth factors for different aerosol types (Zieger et al., 2013; Granados et al., 2015). However, the temporal resolution of routine observations performed by weather services is rather low, typically with one or two radiosonde launches per day (Schmetz et al., 2021). And significant mesoscale weather phenomena, including the frontal systems movement and the formation of convective boundary hygroscopic growth or clouds, transpire rapidly. Thus it is more challenging to adequately monitor the evolution of atmospheric profiles (Kang et al., 2019; Long et al., 2023; Chen et al., 2024). Consequently, precise information with great temporal resolution is essential for examining these events.

The current Raman LiDAR technology enables concurrent measurements of temperature and water vapor mixing ratio profiles to derive RH profiles (Reichardt et al., 2012; Brocard et al., 2013). But it requires calibration by the use of collocated and simultaneous observations from a radiosonde or microwave radiometer (MWR) (Mattis et al., 2002; Madonna et al., 2011; Foth et al., 2015). In addition, the average error of Raman LiDAR is relatively small within the effective height range but limited in the higher height detection.

MWR is another way to provide atmospheric RH observations with high temporal resolution (Hogg et al., 1983; Ware et al., 2003; Zhang et al., 2024). MWR has a certain penetration ability for harsh weather conditions such as clouds. But their vertical resolution and accuracy are not high, especially for RH which vary greatly (Xu et al., 2015). Thus it is challenging to deliver continuous high-resolution RH information with a single instrument. The synergy of information from both active and passive instruments can provide a more comprehensive understanding of atmospheric processes (Stankov, 1995; Furumoto et al., 2003; Delanoë and Hogan, 2008; Blumberg et al., 2015; Tuner et al., 2021). For example, when both Raman LiDAR and MWR are measuring simultaneously, continuous temperature, water vapor profiles

67 and thus RH profiles can be obtained operationally (Navas-Guzmán et al., 2014;
68 Barrera-Verdejo et al., 2016; Foth et al., 2017; Toporov et al., 2020). However, most of their
69 algorithms primarily utilize statistical methods. They perform data fusion between different
70 instruments based on long-term time-series data from individual locations. These approaches
71 are suitable for observations at single stations. Nonetheless, they lack universality when applied
72 to scenarios requiring data integration from multiple sites or broader geographical coverage.
73 Moreover, replacing instruments or equipment may also introduce additional inconsistencies.

74 For accurate RH profile retrieval at higher heights, satellites have global detection capabilities
75 and are highly effective for oceanic skies and remote land areas (Zhang et al., 2022; Wang et
76 al., 2023). For example, Wang et al. measured the subgrid-scale variability of critical relative
77 humidity (RH_c) to investigate the cloud parameterization from Cloud-Aerosol LiDAR and
78 Infrared Satellite Observation (CALIPSO) satellite data. Some deal with the retrieval of the
79 atmospheric layer averaged relative humidity profiles using data from the Microwave
80 Humidity Sounder (MHS) onboard the Metop satellite (Gangwar et al., 2014). Geostationary
81 Operational Environmental Satellite (GOES)-13 and the Moderate-Resolution Imaging
82 Spectroradiometer (MODIS) data are also be combined to estimate hourly relative humidity at
83 the surface level (Ramírez- Beltrán et al., 2019). Another sounder, Sondeur Atmospherique du
84 Profil d'Humidite Intertropical par Radiometrie (SAPHIR), onboard the Megha-Tropiques (MT)
85 satellite sounds the atmospheric humidity. Brogniez et al. (2013) and Gohil et al. (2013) have
86 shown the potential of SAPHIR sounder in retrieving the atmospheric humidity profile.

87 But the time resolution of polar orbit satellites is determined by the repeated coverage time of
88 the satellite orbit (Skou, et al., 2022). A single satellite generally only achieve repeated
89 observations twice a day. And the time resolution is also relatively low. Furthermore, few
90 observations are available from China's satellite Fenyun (FY). This study aims to introduce a
91 novel technique that integrates Raman LiDAR, MWR, and satellite data (FY4B) using an
92 optimum estimating methodology. It is given with a focus on two aspects: i) Evaluation of the
93 proposed synergetic method, and ii), investigation of the RH characteristics at different heights
94 and in different geographic regions. This paper is thus structured as follows. Descriptions of
95 the individual equipment is presented in Section 2. Section 3 illustrates the process of the new
96 synergetic algorithm combining the ground-based and satellite data. Section 4 presents the RH
97 statistic results and its time-height evolution in two convective cases. Finally, conclusions are
98 summarized in Section 5.

99 **2. Instrumentation**

100 *2.1 Raman LiDAR*

101 The Raman LiDAR method can assess the water vapor mixing ratio profiles through inelastic
102 backscattering signals from nitrogen at 387 nm and from water vapor at 407 nm (Whiteman,
103 1992; Mattis et al., 2002; Adam et al., 2010). At the lowest height, the intersection of the laser
104 beam with the receiver's field of view in the bistatic system is incomplete. Nevertheless, the
105 overlap of both Raman channels is presumed to be equivalent; thus, the overlap effect could be
106 minimal concerning water vapor measurements. The collected water vapor measurements, then
107 along with concurrent temperature profiles from a co-located MWR allow us to obtain RH
108 profiles. The vertical and temporal resolution of Raman LiDAR and other instruments are
109 listed in Table 1.

110 *2.2 Microwave Radiometer (MWR)*

111 The Microwave Radiometer (MWR) serves as a passive instrument designed to measure
112 atmospheric emissions across two frequency bands within the microwave spectrum (Cimini et
113 al., 2006; Crewell and Löhnert, 2007). There are seven channels set along the 22.235 GHz H₂O
114 absorption line. Humidity information can be extracted from these observations. The seven
115 channels from 51 to 58 GHz within the O₂ absorption complex encompass the vertical
116 temperature profile data. Consequently, the fully automatic MWR enables the derivation of
117 temperature and humidity profiles with a temporal resolution of up to 5 minutes. The method
118 for inverting temperature and humidity profiles is the neural network method in this study. It
119 uses statistical methods to optimize the long-term average radiosonde data and relies on
120 previous radiosonde data (Yang et al., 2023).

121 *2.3 Radiosonde data*

122 We use radiosonde data from the China Meteorological Administration (CMA) station for
123 reference analysis. It is located in the same place as the Raman LiDAR, and provides on-site
124 measurements of atmospheric pressure, temperature, and RH. During the observing campaign,
125 radiosondes were launched twice a day (08:00 and 20:00 Local Standard Time (LST)). The
126 height can be determined by the ascent time of the radiosonde balloon. The vertical resolution
127 of the raw data is 3 m/layer. To match other data, the vertical resolution of the raw data is
128 interpolated to 30 m (0-3000 m) and 250 m (3000-10000 m), respectively.

129 *2.3 Satellite*

130 In 2016 and 2021, China successfully deployed two second-generation geostationary
131 meteorological satellites, Fengyun-4A (FY4A) and Fengyun-4B (FY4B). They are both
132 equipped with the Geostationary Interferometric Infrared Sounder (GIIRS). The GIIRS
133 therefore became the first geostationary orbiting meteorological satellite (Yang et al., 2023).
134 This approach could achieve the detection of weather systems across China and its neighboring

regions with high temporal and spatial resolution. So it enables a more comprehensive understanding of the atmospheric vertical structure, including the retrieval of atmospheric temperature profiles for 1000 m layers and moisture profiles for 2000 m layers (Yang et al., 2017), respectively. In comparison to FY4A, the GIIRS on FY4B exhibits a broader spectral range, improved spectral resolution in the long-wave IR band, and superior detection sensitivity (Sufeng et al., 2022). Specifically, the temporal resolution of GIIRS has enhanced from 2.5 hours for FY4A to 2 hours for FY4B. In the meantime, the spatial resolution has progressed from 16000 m to 12000 m at nadir. The atmospheric humidity profiles utilized in this study, derived from GIIRS, are generated through the neural network algorithm created by the National Satellite Meteorological Centre (NSMC) (Bai et al., 2022). The data is available online: <http://fy4.nsmc.org.cn/nsmc/en/theme/FY4B.html> (accessed on 12 December 2024).

3. Methods and evaluation

3.1 LiDAR, MWR and satellite synergetic algorithm

This study aims to obtain a continuous time series of RH profiles by integrating ground-based remote sensing techniques, including Raman LiDAR, MWR, and satellite data, in a straightforward manner to facilitate a wide range of applications. The retrieval process involves a systematic four-step algorithm that integrates the Raman LiDAR water mixing ratio profile and MWR brightness temperatures along with satellite data. The retrieval framework is shown as in Fig. 1 and the retrieval process is detailed in the following paragraphs.

Step 1: Data quality control. Data with quality control codes of 0 and 1 for FY4B and 0 for ground-based remote sensing data is selected. The LiDAR only retains data with a SNR value greater than 3. The threshold value of the signal-to-noise ratio (SNR) is set as 3 based on our extensive comparisons with radiosonde data from CMA's long-term observations. The results indicate that selecting LiDAR signals with $\text{SNR} > 3$ can significantly improve the consistency between retrieved RH profiles and radiosonde measurements. So in the data selection period, the Raman signal starts with the first SNR greater than 3 and ends with five consecutive SNRs less than 3. The real-time observing data are designated as R_{radio} , R_{LiDAR} , R_{MWR} and $R_{\text{satellite}}$ in Fig. 2.

Step 2: Data spatial-temporal matching. This process aims to match the above quality-controlled data with the radiosonde data at a height of 0-10000 m in time and space. For the time matching, temperature from MWR and water vapor data from Raman LiDAR are selected corresponding to the radiosonde data time (00:00 LST and 20:00 LST). In terms of spatial matching, the FY4B data is selected from the nearest grid point to the ground observing

station for the horizontal scale. The data at vertical heights are interpolated to the resolution of 30 m (0-3000 m) and 250 m (3000-10000 m).

Step 3: Correction coefficient determination. The deviation between the temperature and humidity data of satellites and ground-based remote sensing data at each height is quantitatively calculated and analyzed to prepare for the optimal stitching process in the next step. Here the deviation of LiDAR, MWR and FY4B are designated as D_{LiDAR} , D_{MWR} and $D_{satellite}$, respectively.

$$D_{LiDAR} = R_{LiDAR} - R_{radio} \quad (1)$$

$$D_{MWR} = R_{MWR} - R_{radio} \quad (2)$$

$$D_{satellite} = R_{satellite} - R_{radio} \quad (3)$$

The correction coefficients C_{LiDAR} , C_{MWR} and $C_{satellite}$ are calculated as follows

$$C_{LiDAR} = (|D_{satellite}| + |D_{MWR}|) / [2 * (|D_{satellite}| + |D_{MWR}| + |D_{LiDAR}|)] \quad (4)$$

$$C_{MWR} = (|D_{satellite}| + |D_{LiDAR}|) / [2 * (|D_{satellite}| + |D_{MWR}| + |D_{LiDAR}|)] \quad (5)$$

$$C_{satellite} = (D_{MWR} + D_{LiDAR}) / [2 * (|D_{satellite}| + |D_{MWR}| + |D_{LiDAR}|)] \quad (6)$$

Step 4: Synergetic algorithm iteration and evaluation: Based on the above spatial-temporal data matching and correction coefficients calculation at different heights, a dynamic optimal stitching algorithm (Fig. 2) is conducted. To ensure the independence between the tested sample and the true value, the temperature and humidity profiles of the current time are fused using the correction coefficient of the previous time, and then compared with the radiosonde data at the same time for evaluation. The correlation coefficient (R), the root mean square error (RMSE), and mean bias (MB) are used as inspection indexes. Finally, the retrieved RH information S_{RH} could be obtained through the following formula.

$$S_{RH} = R_{satellite} * C_{satellite} + R_{MWR} * C_{MWR} + R_{LiDAR} * C_{LiDAR} \quad (7)$$

From the process we can see that compared to these existing techniques, our new method not only incorporates satellite data but also dynamically determines optimal fusion coefficients. Because the fusion coefficients are dynamically determined by comparing the deviations from other measurements with the reference of radiosonde, it highlights that this new algorithm is real-time calibrated. And it can guarantee the device model independence and geographical adaptability. Thus it eliminates constraints imposed by equipment specifications or observation locations, ensuring broad applicability across diverse scenarios.

3.2 Error analysis

To evaluate the performance of the synergetic algorithm for RH profiles, a comparative analysis was conducted between retrieved values and actual radiosonde measurements. Let N represent the total number of samples. The measured value is designated as O_i , with i representing the sample label. The value obtained through the new synergetic algorithm is designated as G_i . The evaluation indicators consist of MB, mean absolute bias (MAB) and RMSE are defined by the following formulas:

$$MB = \frac{\sum_{i=1}^N (G_i - O_i)}{N} \quad (8)$$

$$MAB = \frac{\sum_{i=1}^N |G_i - O_i|}{N} \quad (9)$$

$$RMSE = \sqrt{\frac{\sum_{i=1}^N (G_i - O_i)^2}{N}} \quad (10)$$

4. Results

4.1 General statistic information

A five-month data set has been chosen for a statistical analysis of RH profiles. The observation period spans from July 1 to November 30, 2024. The observing elements are RH data from 47 stations in China (yellow circles in Fig. 3) at the height of 0-10000 m. To investigate RH retrieval accuracy, we provide the comparison results of four methods (LiDAR, MWR, satellite, and synergetic algorithm) utilizing the radiosonde data as the reference at 47 sites in Table 2. Then Huhehaote (HHHT, northern China), Yibin (YB, middle China) and Qingyuan (QY, southern China) are selected as 3 representative sites (red stars in Fig. 3) for more detailed analysis, as shown in Fig. 4 and Table 3. The data samples for HHHT, YB, and QY are 3773, 7452 and 8110, respectively after quality control and elimination of the precipitation data.

Generally, the synergetic algorithm at 47 sites presents the maximum correlation coefficient R value of 0.98 with the minimum RMSE of 5.27% in Table 2. For three representative sites, the regression line from the synergetic algorithm at all heights similarly provides the best fitting results, with the largest correlation coefficients R of 0.94, 0.92 and 0.93 in HHHT, YB and QY respectively (Table 3). The correlation coefficient R for LiDAR measurement follows with marginally higher values of 0.83 in HHHT, 0.86 in YB and 0.86 in QY. It indicates its greater applicability compared to other single instruments. MWR presents the lowest R of 0.74 and 0.80 in HHHT and YB, while performing better ($R = 0.75$) than that from satellite ($R = 0.66$) in QY. In terms of RMSE, the LiDAR-, MWR- and satellite-derived RH all show values larger

229 than 18% at three sites. The synergistic use of a multi-source algorithm decreases the RMSE
 230 down to the lowest value of 10% in HHHT.

231 The regression line for LiDAR and MWR in HHHT, as illustrated in Fig. 4, exhibits a slope
 232 that is less than that of the one-to-one line. This implies that greater variations arise with
 233 increased RH in HHHT. Though the synergetic algorithm also presents similar trends, its
 234 RMSE decreased to 10% in HHHT. The regression line of LiDAR and MWR in YB and QY
 235 are larger than the one-to-one line, indicating the larger bias for less humid.

236 As RH vertical profiles are height-dependent, Fig. 5 presents the MB profiles observed at
 237 different heights in terms of four methods. Generally, the MB in the RH of LiDAR in the lower
 238 troposphere (below 3000 m) outperforms the other two single methods (MWR and satellite) at
 239 three sites. No significant biases between radiosonde and LiDAR are noticeable. Specifically,
 240 the lowest MB values (4.93% in HHHT, 2.63% in YB and 2.40% in QY) in the comprehensive
 241 region of the tropospheric region are achieved when LiDAR data is incorporated into the
 242 synergetic algorithm. This is because LiDAR is an active remote sensing technology with more
 243 accuracy compared to MWR and satellite. The LiDAR data's efficacy is enhanced at heights
 244 below 3000 m when integrated with data from other sources within the boundary layer.

245 However, the MB from LiDAR increased drastically above this height, up to the highest value
 246 28.67% in HHHT, 29.91% in YB and 20.09 % in QY. It is reasonable that the atmosphere
 247 changes so fast that radiosonde do not assess exactly the same air mass as LiDAR. In the
 248 meantime, LiDAR is increasingly constrained at elevated heights because of a decreased SNR.
 249 Hence LiDAR is more trustworthy in the lower layer, i.e. below 3000 m.

250 In contrast, the MB from satellite (FY4B) over 3000 m varied steadily within the range of
 251 approximately 15% at three sites. Therefore the satellite data in the far height range would be
 252 more reliable and could be employed in the synergetic algorithm at higher layers. Compared to
 253 LiDAR and satellite, the MB from MWR gives the largest uncertainty in HHHT at all heights.
 254 This may result from the discrepancy between the temperature recorded by the radiosonde and
 255 that obtained from the MWR in HHHT. However, it yields relatively less variation than
 256 LiDAR and satellite in YB and QY. Anyway, the synergetic method gives the best result for
 257 over three observing sites at almost all heights. And accurate measurements of RH vertical
 258 profiles provided here are highly beneficial for analyzing the hygroscopic growth of local
 259 aerosols.

260 The sources of the discrepancy can stem from several aspects. First, although all instruments
 261 are co-located in the ground, radiosondes deviate at higher heights. Signals can be disrupted if

clouds are present. Second, satellites provide gridded data, requiring the selection of ground observation points closest to its grid's latitude and longitude, which introduces uncertainty. Finally, both MWR and satellite are passive remote sensing technologies, which are inherently less precise than active remote sensing. Besides the inherent hardware difference, the errors during the retrieval process (e.g., neural networks for MWR) are also unavoidable.

4.2 Mean monthly analysis

RH mean monthly vertical profiles have been derived from the synergistic method illustrated in Fig. 6. Because RH profiles are retrieved from water ratio profiles and temperature profiles. For this property, the RH seasonal behavior may be more complicated. For example, no obvious seasonal behavior of RH profiles is found in HHHT or YB. However, QY still presents the most likely seasonal characteristic at most of the heights, with the highest mean values in summer at 1000-2000 m (80.65% in July) and lowest values at 7000-10000 m in late autumn (20.50% in November) in Fig. 6e-f. The elevated RH observed in QY's summer may be related to the sufficient water vapor and large transport volume as QY is located in coastal areas. So the characteristic of QY would be more dependent on water vapor.

For comparison, HHHT and YB are relatively random. Over 3000 m in HHHT (Fig. 6a-b), RH in August shows predominantly high values with the highest value of 65.37% at 5000-7000 m. Different from HHHT and QY, the RH profiles in November of YB interestingly show the highest values (83.95%) in the lower atmosphere (0-1000 m) in Fig. 6c-d. It suggests the reduced temperatures observed in autumn of YB promote proximity to saturation conditions, resulting in elevated RH values in November. It is also worth noting that RH above 3000 m in November of YB decreases dramatically as height increases, with the minimum RH of 13.91% at 7000-10000 m. That could be explained by more rapid fluctuations in the water vapor density and temperature in YB in the higher layer under the control of the subtropical monsoon climate zone. Anyway, this plot illustrates a clear decrease in the RH values with heights at three sites.

Though there is no obvious RH uncertainty caused by regional differences, we found that QY exhibits the predominant seasonal feature throughout most heights. In contrast, no discernible seasonal characteristics in RH profiles are observed in HHHT or YB. Thus we believe diverse atmospheric circulation patterns and geographical environments could result in regional variations in RH values.

4.3 Case analysis

We selected two different severe convective events in YB (one hailfall from 20:00 LST to 23:30 LST on 15 April and one heavy precipitation from 14:00 LST 25 May to 08:00 LST 26 May) for comparison in Fig. 7. At 23:00 LST on April 15, a thunderstorm with strong winds and hail occurred. The synergetic algorithm retrieved RH profile showed that before 22:00 LST, the RH was high (around 90%) at 3000 m height, low (20%-50%) between 3500 m and 8000 m, and above 80% between 8000 m and 9000 m (Fig. 7a). This indicates that before the severe convection, the upper and lower layers were relatively moist, while the middle layer (3500 m-8000 m) was dry (red arrow in Fig. 7a). Such a condition favors the evaporation and cooling of ice particles descending from the upper atmosphere, leading to refreezing and hail formation.

In contrast, the RH profile from 25 May to 26 May showed that the entire troposphere (0-10000 m) presented high RH values (>70%) starting at 19:00 LST, which was conducive to heavy precipitation (Fig. 7b). According to ground station observations, YB recorded an hourly rainfall of 52 mm at 21:00 LST, along with gale-force winds of 23 m/s (9th grade). Most areas in YB experienced precipitation, with localized heavy thunderstorms. From the above two cases, we can see that the RH in the middle troposphere can be used to distinguish between hail and heavy precipitation during severe convective events.

5. Conclusion

This study presents relative humidity (RH) measurements with a developed synergetic algorithm with the combination of Raman LiDAR, MWR, and satellite at three sites (northern China, middle of China and southern China) from 1 July to 31 November. The methodology for obtaining RH from the synergetic algorithm was introduced. The five-month field campaign was performed and linear regression between the LiDAR, MWR, satellite, synergetic algorithm and radiosonde data at the range 0-10000 m was presented to testify the accuracy.

Strong correlations of RH values over 0.9 were observed between radiosonde measurements and profiles derived from the synergetic approach at three representative sites in China. The lowest MB values (4.93% in HHHT, 2.63% in YB and 2.40% in QY) are observed when LiDAR data is integrated into the synergetic algorithm, which highlights the accuracy of the LiDAR data below 3000 m. However, the MB from LiDAR increased drastically above this height, which suggests the greater applicability of satellite or MWR in the middle and higher layers. In terms of the seasonal characteristic, QY exhibits the predominant seasonal feature throughout most heights, with peak mean values of 80.65% in July at 1000-2000 m and minimal values of 20.50% in November at 7000-10000 m. Finally, the RH evolution in two

convective events in YB suggests that the RH in the middle troposphere can be taken as a good indicator to distinguish hail and heavy precipitation.

These results validate the capabilities of the newly developed method to deliver accurate measurements of RH information throughout the troposphere. It also explores the potential of satellite data integration for RH profile retrieval for the first time. However, there are still problems with individual data at certain times during the fusing process. For example, there are few effective data filtered by quality control methods for FY4B data. Therefore, the matching accuracy and more high-quality FY4B data will be improved in future development.

Declaration of Competing Interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

Data availability

Raman LiDAR, MWR, satellite, radiosonde and other auxiliary data used to generate the results of this paper are available from the authors upon request (email: zychen@btbu.edu.cn).

Author contribution

Chengli Ji and Zhenyi Chen conceived the study. Jiajia Mao, Zhicheng Wang, and Xiaoyu Ren conducted the field measurements. Zhenyi Chen, Chengli Ji, Qiankai Jin, Feilong Li, Yuyang Liu and Wanlin Jian carried out the data analysis. Peitao Zhao and Yan Xiang offered academic help for this research. Zhenyi Chen and Chengli Ji wrote the paper with inputs from all co-authors.

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514

515 List of Tables

516 **Table 1** Instruments and monitoring parameters

Instrument	Parameters/units	Temporal-spatial Resolution
Raman LiDAR	Relative humidity (RH)	7.5 m, 3 minutes
Microwave radiometer (MWR)	Temperature (°C), Relative humidity (RH)	50 m, 3 minutes
FY4B	Relative humidity (RH)	1 hour

517

518 **Table 2** Assessment of the accuracy of four RH retrieval results (LiDAR, MWR, satellite and
519 synergetic algorithm) compared with radiosonde at 47 sites in China.

Comparison with radiosonde	Number of sample	R	MB (%)	MAB (%)	RMSE (%)
LiDAR	192111	0.91	0.56	6.7	10.67
MWR	192111	0.82	-1.49	10.79	14.31

satellite	192111	0.74	1.08	13.19	17.02
synergetic algorithm	192111	0.98	0.42	3.24	5.27

Table 3 The same as Table 2 but at three representative sites in China.

HHHT (northern China)	Comparison with radiosonde	Number of sample	R	RMSE (%)
	LiDAR	3771	0.83	20
	MWR	3771	0.74	25
	satellite	3771	0.76	24
	synergetic algorithm	3771	0.94	10
YB (middle China)	LiDAR	7452	0.86	19
	MWR	7452	0.80	26
	satellite	7452	0.83	29
	synergetic algorithm	7452	0.92	12
QY (southern China)	LiDAR	8110	0.86	18
	MWR	8110	0.75	19
	satellite	8110	0.66	21
	synergetic algorithm	8110	0.93	11

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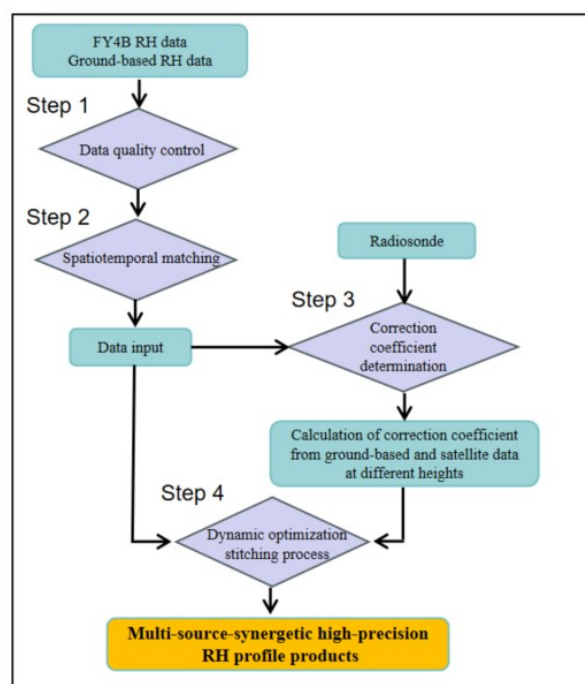


Fig. 1 Sketch of the retrieval scheme. Details are given in the text.

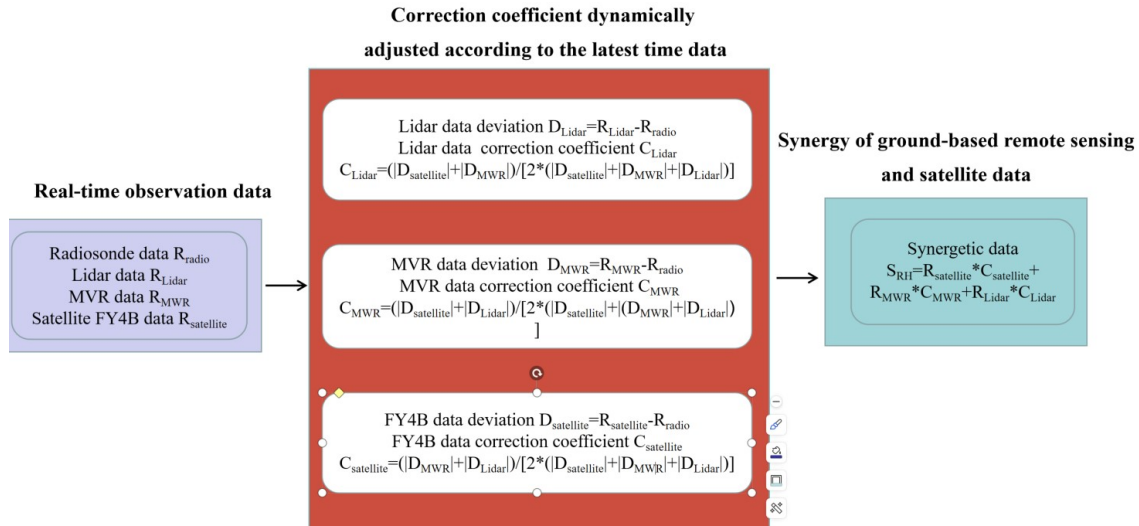


Fig. 2 The dynamic optimal stitching process

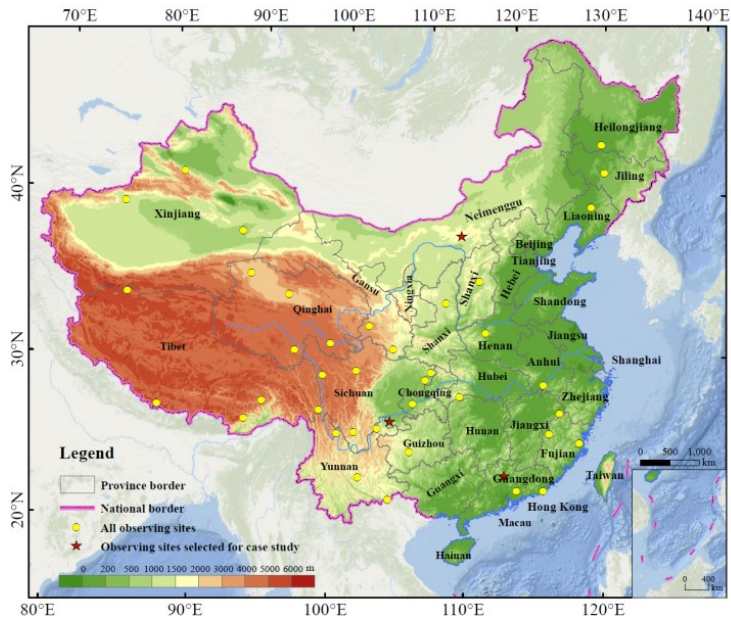
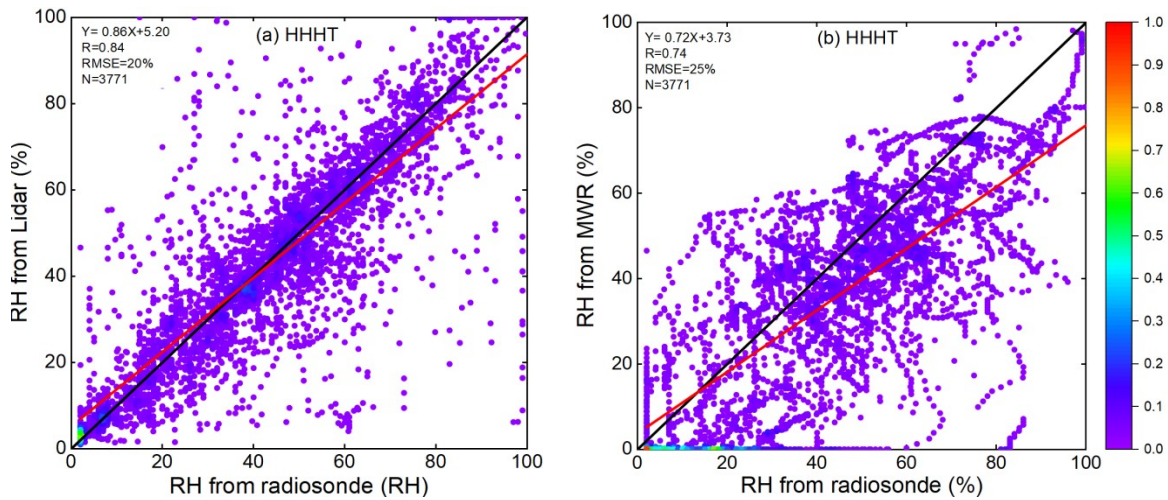
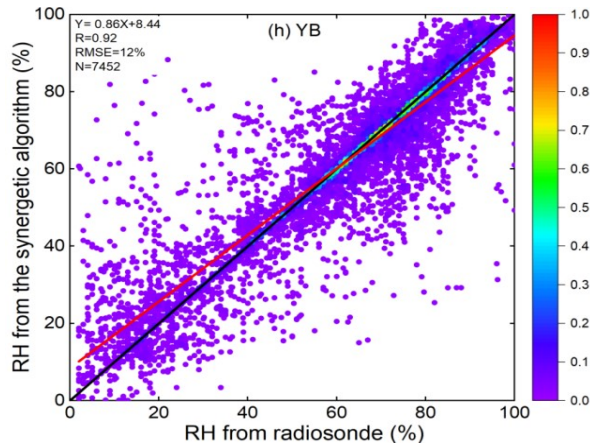
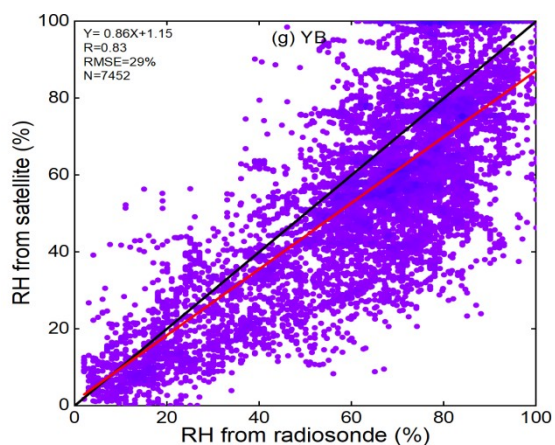
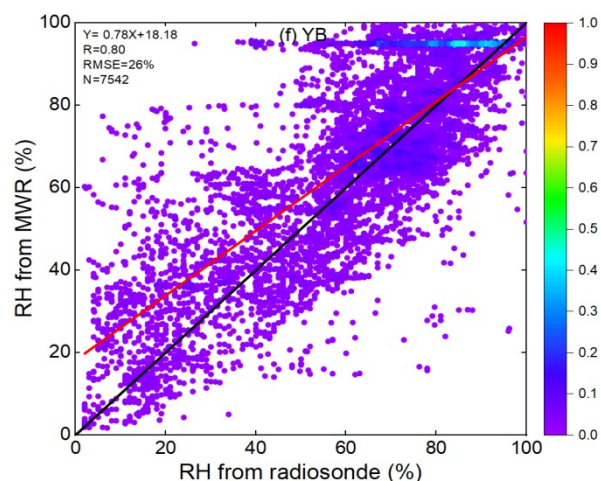
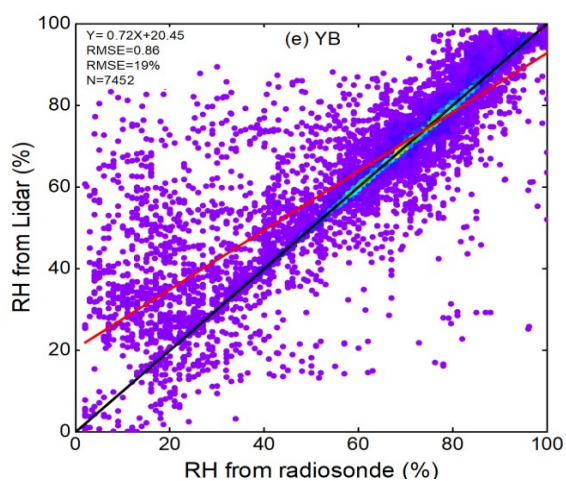
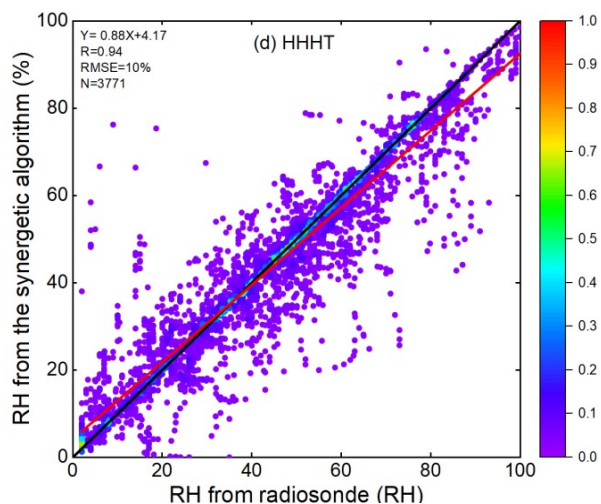
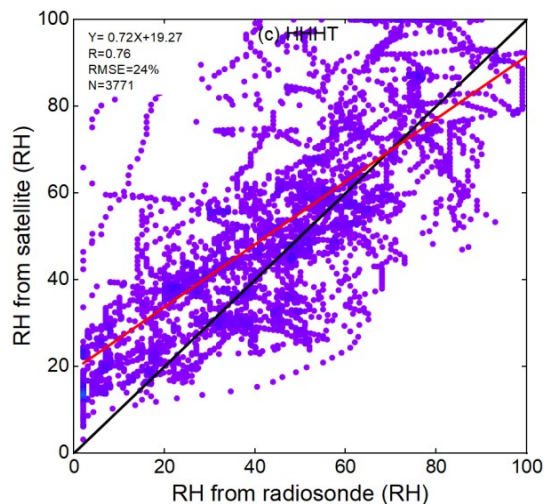


Fig. 3 The observing sites (yellow circles) and three selected sites (red stars) for statistics and case studies are marked in the





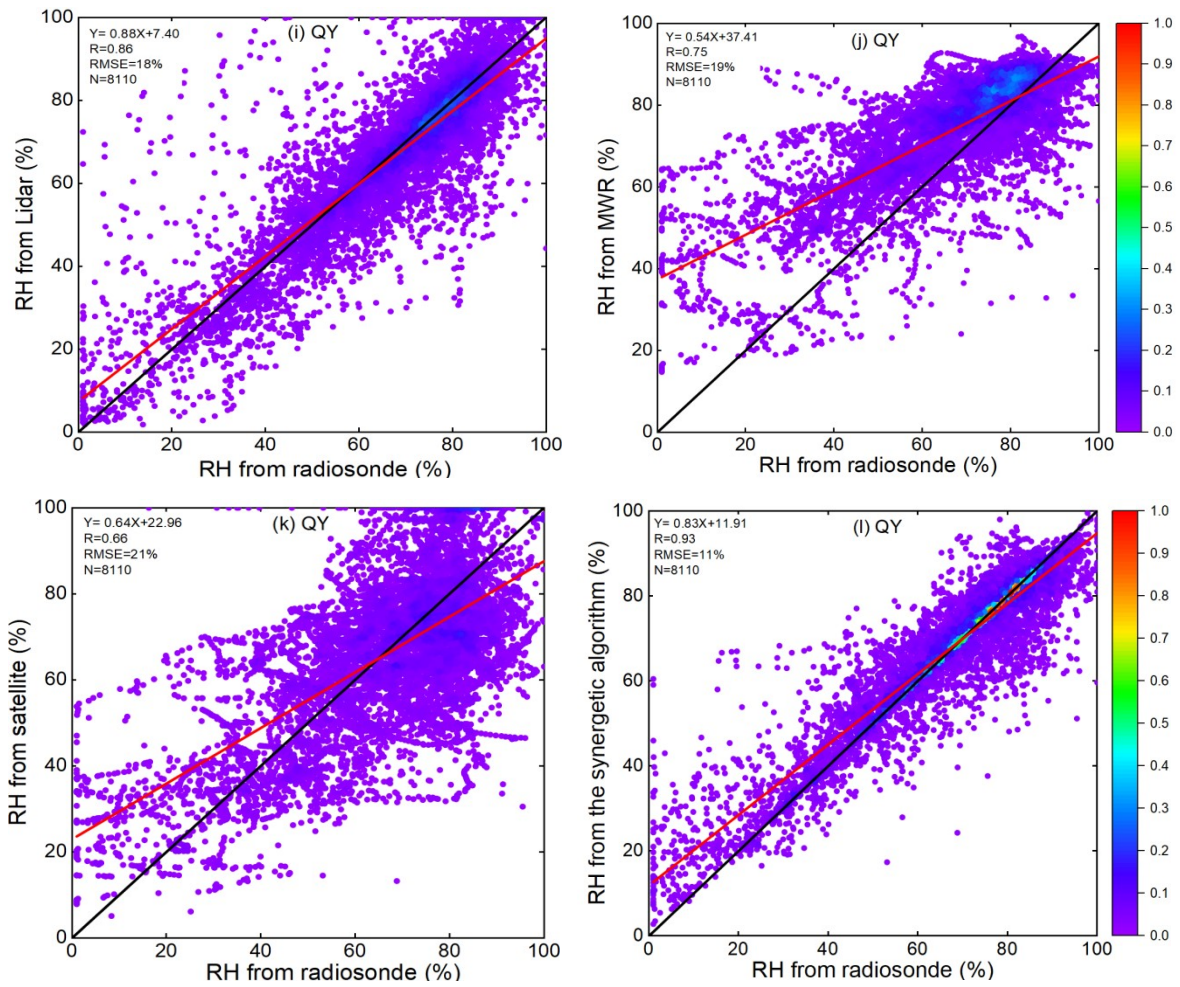
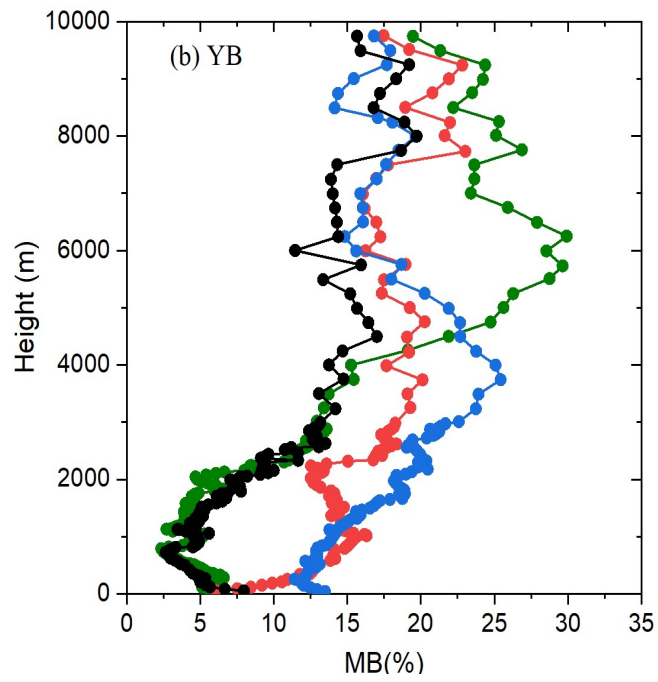
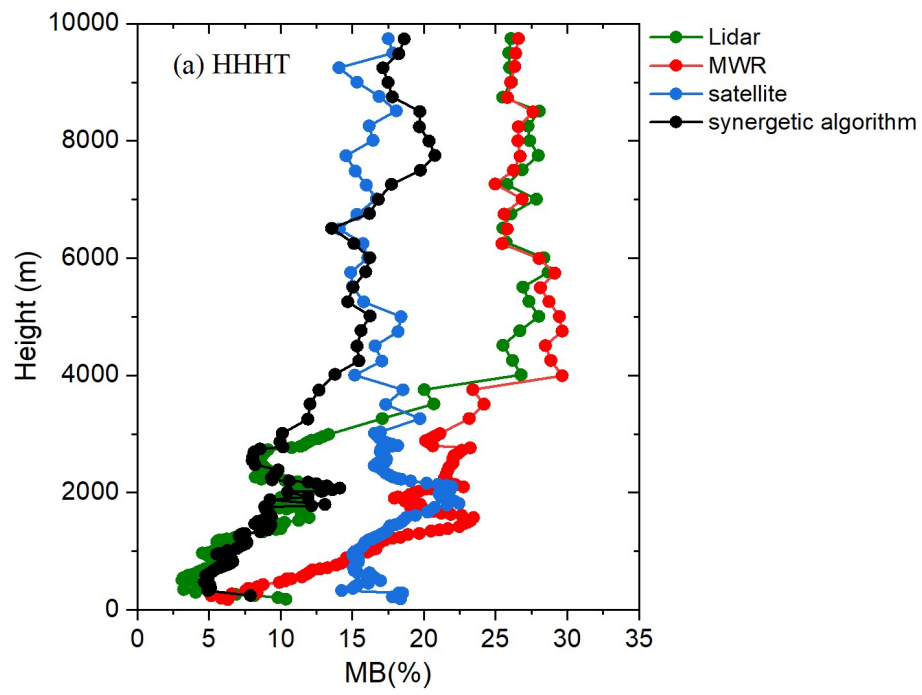


Fig. 4 Four-methods-retrieved RH results (LiDAR, MWR, satellite and synergetic algorithm) compared with radiosonde at three sites in China from 1 July to 31 November 2024. (a) Comparison between LiDAR and radiosonde in HHHT, (b) Comparison between MWR and radiosonde in HHHT, (c) Comparison between satellite and radiosonde HHHT, (d) Comparison between synergetic algorithm and radiosonde in HHHT; (e)-(h), the same as (a)-(d) but in YB. (i)-(l), the same as (a)-(d) but in QY. The red line shows the regression line. The black line is the one-to-one line.



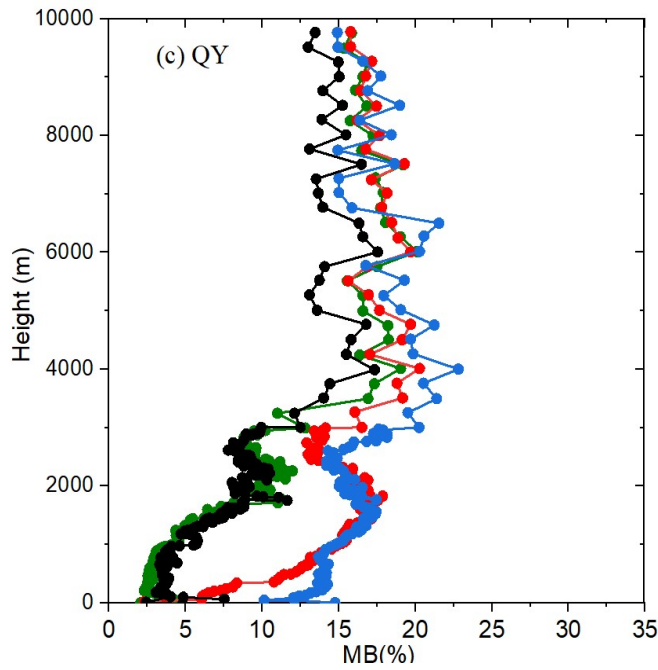
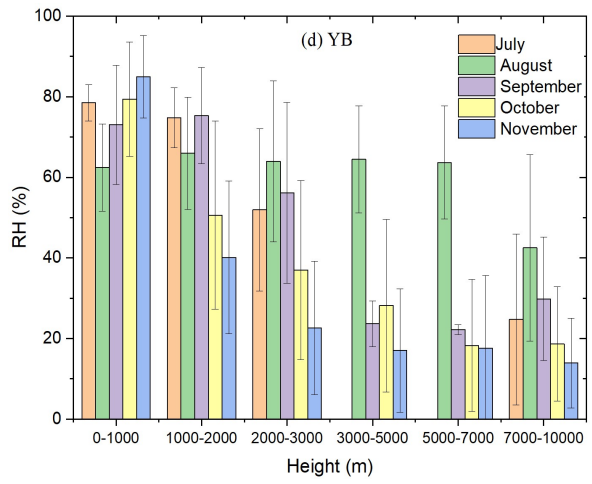
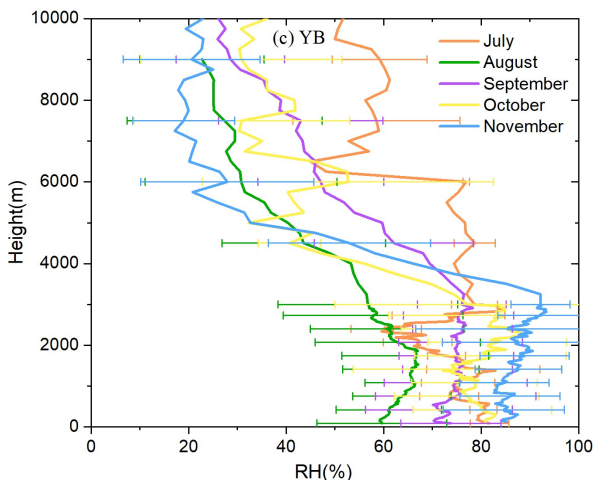
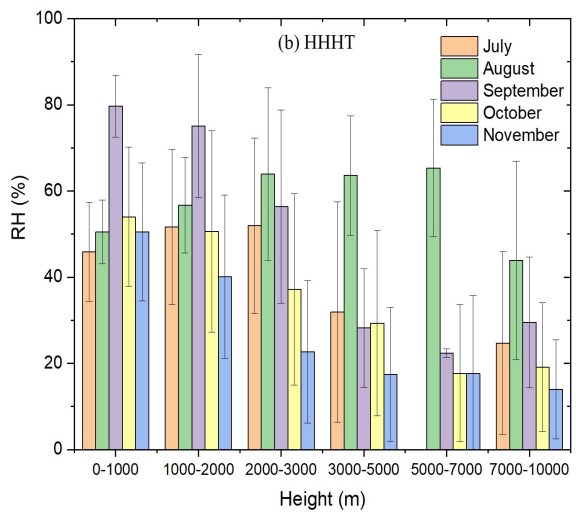
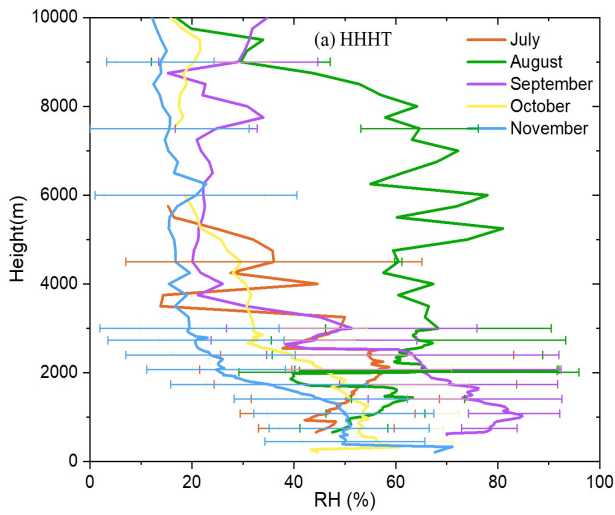


Fig. 5 RH vertical mean bias (MB) profiles retrieved from LiDAR, MWR, satellite and synergetic algorithm compared to the radiosonde data in (a) HHHT, (b) YB and (c) QY.



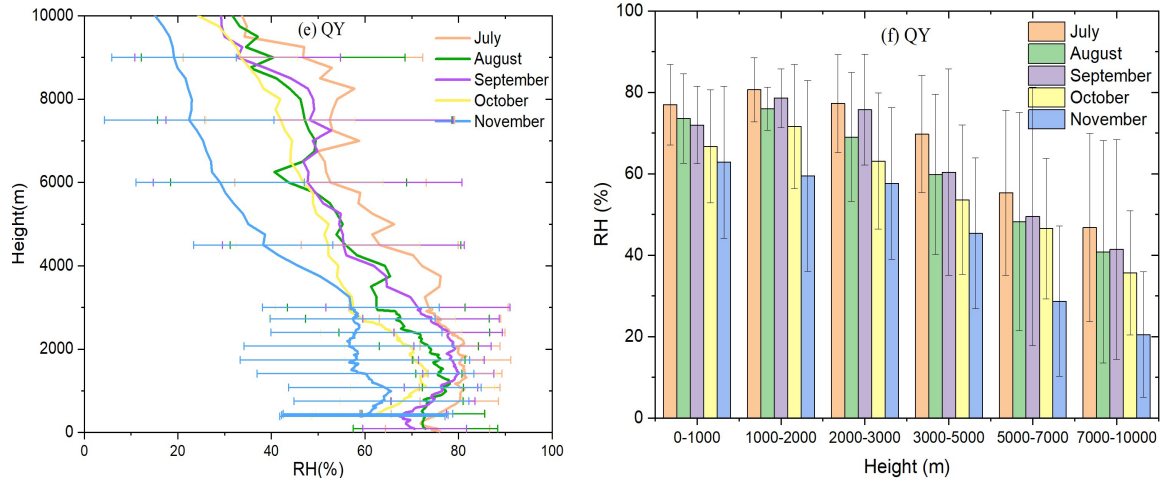


Fig. 6 RH Monthly vertical profiles (left) and monthly mean values for different heights (right) in (a)-(b) HHHT, (c)-(d) YB and (e)-(f) QY. The error bars indicate the standard deviation.

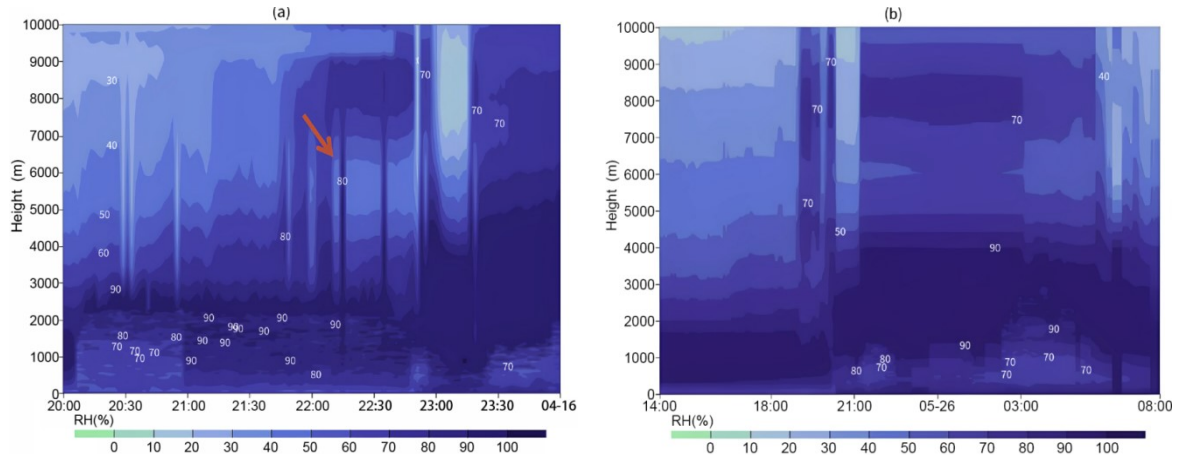


Fig. 7 Height-time display of RH from the synergetic retrieval during two convective cases (a) from 20:00 LST to 23:30 LST 15 April and (b) from 14:00 LST 25 May to 08:00 LST 26 May 2024 in YB. The red arrow indicates the less humidity in the layer when the hailfall occurred in the first convective case.