



1 **Ideas and perspectives: Tipping into the unknown—the global consequences**
2 **of biogeochemical system collapse**

3 Salvador Sánchez-Carrillo^{1*}, Felipe García-Oliva², Javier Alcocer³, Rocío Alcántara-
4 Hernández⁴, Martín Merino-Ibarra⁵, & David G. Angeler^{1,6,7}

5 ¹Dep. Biogeochemistry and Microbial Ecology, Museo Nacional de Ciencias Naturales-
6 CSIC (MNCN-CSIC), Serrano 115 bis, E-28006-Madrid, Spain

7 ²Instituto de Investigaciones en Ecosistemas y Sustentabilidad, Universidad Nacional
8 Autónoma de México, Morelia, Michoacán, Mexico

9 ³Grupo de Investigación en Limnología Tropical, FES Iztacala, Universidad Nacional
10 Autónoma de México, Tlalnepantla, Estado de México, Mexico

11 ⁴Instituto de Geociencias, Universidad Nacional Autónoma de México, Ciudad de
12 México, Mexico

13 ⁵Instituto de Ciencias del Mar y Limnología, Universidad Nacional Autónoma de
14 México, Ciudad de México, Mexico

15 ⁶School of Natural Resources, University of Nebraska - Lincoln, Lincoln, NE, USA

16 ⁷IMPACT, Institute of Mental and Physical Health and Clinical Translations, Deakin
17 University, Melbourne, Australia

18 *Author for correspondence sanchez.carrillo@mncn.csic.es

19 **Abstract**

20 With Earth's tipping points being approached and already exceeded, human-driven
21 disruptions of carbon (C), nitrogen (N), and phosphorus (P) cycles are forcing Earth's
22 biogeochemical systems toward irreversible collapse. Potential functional extinction of
23 key regulatory processes—such as carbonate precipitation, denitrification, phosphorus
24 burial, and methane oxidation—could lead to cascading failures, spread beyond
25 biogeochemical domains and further contribute to destabilize planetary homeostasis.
26 This study examines scenarios how these disruptions may interact across
27 biogeochemical cycles, and how they might create reinforcing feedback loops that



28 amplify climate change, accelerate ecosystem degradation, and alter atmospheric and
29 oceanic chemistry.

30 Our scenarios envision carbon cycle disruptions (ocean acidification, soil carbon loss)
31 weaken CO₂/CH₄ sequestration and boost emissions. Eutrophication and oxygen
32 depletion threaten the nitrogen cycle, increasing nitrates and N₂O. Meanwhile,
33 phosphorus release from hypoxic sediments sustains eutrophication and intensifies
34 greenhouse gas emissions. Changing feedback loops may prevent the recovery of
35 desirable biogeochemical conditions. A shift to anaerobic metabolism would favor
36 sulfate reduction, methanogenesis, and ammonification, triggering biodiversity collapse,
37 expanding anoxic zones, and allowing microbial extremophiles to dominate—echoing
38 early anoxic Earth.

39 The loss of these regulatory functions would exacerbate global warming and push
40 Earth's ecosystems irreversibly into a "Hothouse Earth" regime. Immediate governance
41 action is necessary to mitigate these risks by managing nutrient cycles, protecting
42 carbon sinks, and incorporating biogeochemical feedbacks into climate policies.
43 Without intervention, the accelerating extinction of key biogeochemical processes may
44 render long-term climate protection unattainable.

45 **Keywords:** biogeochemical cycles, functional extinction, feedback loops, climate
46 change, tipping points, ecosystem collapse, planetary boundaries, resilience.

47

48 1. Introduction

49 Human domination of the biophysical environment in an accelerating Anthropocene is
50 moving Planet Earth dangerously towards and beyond global tipping elements
51 (Richardson et al. 2023). Global environmental change has surpassed planetary
52 boundaries, entering a high-risk period for humanity (IPCC, 2018). For instance, land
53 system alterations, ocean acidification and freshwater eutrophication are escalating and
54 nearing critical thresholds (Steffen et al. 20015).

55 Biogeochemical cycles are not exempt from rapid environmental change.
56 Unprecedented human alterations of the environment fundamentally alter the carbon
57 (C), nitrogen (N), and phosphorus (P) cycles, disrupting biogeochemical fluxes, and



58 threatening functional key regulatory processes (Galloway et al. 2008). The loss of
59 these processes could push ecosystems beyond critical thresholds, triggering cascading
60 effects (Rocha et al. 2018) and moving Earth into an irreversible Hothouse Earth regime
61 (Steffen et al. 2018). Once such a shift has occurred the current Holocene climate
62 regime will be elusive. Under such a globally novel Earth systems regime, extinction
63 and fundamental reorganization of microbial pathways, mineralization reactions, and
64 nutrient transformation processes may amplify positive feedback loops, and bolster run-
65 away climate change, ecosystem degradation, biodiversity loss and global
66 biogeochemical disruptions (Falkowski et al., 2000).

67 Among the most vulnerable processes are carbonate precipitation and burial (Ridgwell
68 and & Zeebe, 2005), which regulate long-term carbon sequestration in the oceans
69 (Berner, 2004); denitrification and anammox, essential for balancing reactive nitrogen
70 levels (Thamdrup & Dalsgaard, 2002); and phosphorus occlusion and burial, which
71 prevent eutrophication and limit primary production (Ruttenberg, 2003). Their
72 substantial alteration would destabilize local ecosystems and amplify planetary-scale
73 feedback loops that climate change. As these disruptions approach critical thresholds,
74 they may trigger tipping points beyond which recovery becomes improbable, leading to
75 irreversible shifts in the Earth system.

76 In this perspective, we project evidence-based state-of-the-art research into and build
77 speculative but potentially realistic worst-case scenarios about the consequences of the
78 potential collapse of biogeochemical processes. Such an assumption is partly motivated
79 by the current climate trajectory pointing to a global temperature increase exceeding
80 2°C which would bring severe, widespread impacts (IPCC, 2022). Specifically, we
81 evaluate the mechanisms driving the decline, interactions, and cascading effects on
82 climate regulation and ecosystem stability across altered C, N, and P cycles. Through a
83 global feedback loop perspective, we assess how these disruptions could reinforce one
84 another. Substantial local, regional, and global alterations of interrelated
85 biogeochemical processes may contribute unexpectedly to pushing our Earth toward a
86 fundamentally different and less hospitable regime. The scope and complexity of
87 potential biogeochemical collapse prevent our analysis from being exhaustive; however,
88 our paper's broad yet realistic tendencies justify an urgent call for governance and
89 management action.



90 We acknowledge upfront that the scenarios presented are alarmist, speculative and lack
91 quantifiable precision. However, the article's primary aim is to emphasize the urgency
92 of addressing critical biogeochemical disruptions, recognizing that quantifiable
93 evidence and a focus on human capacity to adapt and transform may prevent these
94 scenarios from becoming real. Recognizing the potential for system resilience and
95 human innovation is essential, as these should complement, not detract from, the urgent
96 need for a global, coordinated response to prevent irreversible damage (Rockström et
97 al., 2009). The emphasis on immediate action aligns with the principle that preparing
98 for dire possibilities is preferable to being caught unprepared (Taleb, 2007).

99 **2. Functional extinction of carbon processes: from sequestration to atmospheric**
100 **release**

101 The carbon cycle is experiencing irreversible shifts as sequestration pathways decline
102 and emission routes intensify. Ocean acidification, driven by CO₂ absorption,
103 suppresses carbonate precipitation, essential for long-term carbon burial (Table 1). As
104 pH declines, calcifying organisms (e.g., corals, foraminifera, coccolithophores) struggle
105 to form calcium carbonate, reducing marine carbon storage and weakening the ocean's
106 role as a CO₂ sink, accelerating atmospheric carbon accumulation in a self-reinforcing
107 loop (Feely et al., 2004; Orr et al., 2005).

108 Rising temperatures and microbial priming effects on land (the stimulation of soil
109 microbial activity by fresh organic inputs, which accelerates the decomposition of older,
110 stable carbon pools) are shifting soils from carbon sinks to sources. While CO₂
111 fertilization temporarily enhances photosynthesis, warming accelerates organic matter
112 decomposition, outpacing sequestration (Davidson & Janssens, 2006). The functional
113 extinction of slow humification pathways, which stabilize carbon in long-term pools,
114 would deplete soil carbon stocks, intensifying CO₂ emissions and increasing ecosystem
115 desertification (Cotrufo et al., 2019). Simultaneously, permafrost thaw is mobilizing vast
116 reservoirs of ancient organic carbon, fueling microbial decomposition and releasing
117 both CO₂ and methane (Schuur et al., 2015; Turetsky et al., 2020). The rapid
118 destabilization of these frozen carbon stocks represents a tipping point beyond which
119 permafrost ecosystems could shift from a long-term sink to a sustained source of
120 atmospheric carbon. In addition to the above, the forest C pool is threatened by the
121 increment of temperatures and reduction of precipitation, because the severe wildfires



122 have increased in forest ecosystems with low fire frequency as boreal and humid-
123 temperate forests (IPCC, 2019). These dynamics may be beyond human control and
124 must be mitigated by global human action.

125 The shift from methanotrophic mechanisms to unchecked methanogenesis in wetlands
126 could mark an approaching critical tipping point. As warming intensifies, prolonged
127 waterlogging and oxygen depletion lead to the collapse of aerobic methanotrophic
128 communities, reducing methane oxidation capacity (Bange et al., 2010). Although
129 anaerobic oxidation of methane (AOM) via sulfate or nitrate reduction can act as an
130 alternative sink, its efficiency depends on the availability of these electron acceptors,
131 which may be limited in freshwater wetlands and eutrophic systems (Raghoebarsing et
132 al., 2006; Ettwig et al., 2010). Competition for sulfates and nitrates can further favor
133 methanogenesis over AOM in highly anoxic conditions, exacerbating methane
134 accumulation. Since methane has 25 times the warming potential of CO₂ (Myhre et al.,
135 2013), this shift reinforces run-away climate warming and ecosystem degradation.

136 **3. Nitrogen cycle disruptions: the extinction of denitrification and the rise of** 137 **nitrous oxide**

138 Human-driven nitrogen inputs, particularly fertilizer overuse, are disrupting microbial
139 nitrogen transformations (Galloway et al., 2008). N fertilizer use has increased around
140 8-fold in 55 years (between 1961 and 2016) for global crop production (Martinez-
141 Dalmau et al., 2021), which has probably favored the increase in N₂O emissions
142 because agriculture has been reported to be responsible for two-thirds of the global
143 emissions of this GHG (Janssens-Maenhout et al., 2017). A critical risk is the
144 functional extinction of denitrification, which converts reactive nitrogen to inert N₂ gas.
145 Under eutrophic and anoxic conditions, denitrifying bacteria are replaced by microbial
146 communities favoring incomplete denitrification, leading to increased nitrous oxide
147 (N₂O) emissions, a greenhouse gas 300 times the warming potential of CO₂ (Davidson,
148 2009). This shift could exponentially increase N₂O accumulation, further accelerating
149 climate change.

150 Another vulnerable pathway is anaerobic ammonium oxidation (anammox), essential
151 for nitrogen removal in marine and freshwater systems (Kuypers et al., 2003). This
152 process is highly sensitive to thermal and chemical disruptions; its collapse could cause



153 ammonium (NH_4^+) to accumulate in aquatic ecosystems, fueling eutrophication, toxic
154 algal blooms, and oxygen depletion (Jetten et al., 2001). The extinction of nitrifying
155 bacteria due to acidification and eutrophication would create a nitrogen-saturated
156 environment, exacerbating imbalances between nitrogen fixation and removal.

157 **4. Phosphorus cycle breakdown: loss of sedimentary storage and self-reinforcing** 158 **eutrophication**

159 Phosphorus, a non-renewable nutrient critical for primary production, is undergoing
160 irreversible alterations due to agricultural runoff and climate change (Filippelli, 2008).
161 The extinction of phosphorus precipitation and burial pathways, driven by ocean
162 acidification and anoxic sediment conditions, could permanently alter phosphorus
163 availability (Conley et al., 2009). Under normal conditions, phosphorus binds with iron
164 and calcium in sediments, removing it from biological circulation. However, increasing
165 hypoxia releases stored phosphorus back into the water, perpetuating eutrophication in a
166 positive feedback loop (Elser et al., 2007).

167 Another collapse scenario is emerging in phosphorus-limited ecosystems: phosphorus
168 depletion in soils due to intensive agriculture and erosion threatens terrestrial primary
169 production (Sternberg et al., 2017). In addition, global reserves of phosphate rock may
170 be depleted within 10 years, as several authors estimate that the peak of P fertilizer
171 production will occur around 2030 (Cordell et al., 2009), declining after this year. As
172 phosphorus becomes scarce, plant growth declines, reducing carbon sequestration and
173 compounding climate change while threatening food security and eroding ecosystem
174 resilience.

175 **5. The domino effect: tipping points and cascading extinctions across** 176 **biogeochemical functions**

177 The simultaneous and interrelated disruption of the C, N, and P cycles initiates multi-
178 layered feedback loops, driving Earth's biogeochemical systems toward and beyond
179 irreversible collapse and novel configurations. The interdependence of these cycles
180 amplifies disturbances, creating a domino effect where the extinction of one
181 biogeochemical function accelerates the degradation of others (Falkowski et al., 2000).
182 Several routes of change can be envisioned:



183 **Collapse of methanotrophic mechanisms:** As warming intensifies, oxygen depletion
184 in wetlands, thawing permafrost, and aquatic systems increasingly suppress aerobic
185 methanotrophic processes, reducing methane oxidation capacity and leading to
186 unchecked methane accumulation (Davidson & Janssens, 2006; Bange et al., 2010).
187 While anaerobic methane oxidation (AOM) linked to sulfate and nitrate reduction can
188 provide an alternative sink, its effectiveness is constrained by electron acceptor
189 availability, particularly in freshwater wetlands and eutrophic environments
190 (Raghoebarsing et al., 2006; Ettwig et al., 2010). In highly anoxic conditions, sulfate
191 and nitrate limitation, coupled with increased organic matter decomposition, further
192 shifts microbial dominance toward methanogenesis, intensifying CH₄ emissions. This
193 feedback loop accelerates warming and enhances soil respiration and permafrost
194 thawing, releasing additional CH₄ and CO₂, reinforcing global climate destabilization.

195 **Denitrification failure:** Oxygen depletion disrupts nitrogen removal, causing nitrate
196 overaccumulation and increased N₂O emissions, exacerbating greenhouse warming
197 (IPCC, 2021). Excess nitrate intensifies eutrophication, driving algal blooms and further
198 oxygen loss, reinforcing anoxia and suppressing denitrification. This cycle locks
199 nitrogen in aquatic systems, promoting ammonification and worsening water quality,
200 biodiversity loss, and ecosystem instability (Galloway et al., 2008). As denitrification
201 declines, methanogenesis becomes more dominant, shifting microbial pathways toward
202 methane (CH₄) production, further amplifying global warming. If left unchecked,
203 nitrogen-saturated, hypoxic environments will become permanent greenhouse gas
204 sources, accelerating planetary instability (a state where cascading biogeochemical
205 disruptions drive irreversible climate shifts, ecosystem collapse, and feedback loops that
206 push Earth beyond safe operating conditions).

207 **Phosphorus burial loss:** Without effective P removal, eutrophication remains self-
208 sustaining, driving algal blooms and oxygen depletion, suppressing nitrification, and
209 amplifying CH₄ and N₂O emissions. Anoxic sediments release stored phosphorus,
210 perpetuating this cycle indefinitely (Conley et al., 2009; Elser et al., 2007; Spears et al.
211 2013). As eutrophication worsens, sulfate-reducing bacteria thrive, producing toxic
212 hydrogen sulfide (H₂S) and further degrading biodiversity. If phosphorus burial collapse
213 continues, aquatic systems risk becoming permanently hypoxic or anoxic, locked in a
214 cycle of nutrient saturation, greenhouse gas emissions, and ecosystem failure (i.e., the



215 loss of structure and function in the ecosystem; Díaz & Rosenberg, 2008; Rockström et
216 al., 2009; Barnosky et al., 2012), contributing to global warming.

217 As these feedback compounds, tipping points emerge, beyond which natural recovery
218 mechanisms fail (Scheffer et al., 2012, Hessen et al. 2024). The functional extinction of
219 microbial communities mediating C sequestration, N removal, and P stabilization may
220 drive large-scale regime shifts—from carbon-sequestering forests to grasslands, from
221 oxygen-rich oceans to anoxic dead zones, and from nutrient-cycling wetlands to
222 methane-producing landscapes (Lenton et al., 2008).

223 These cascading disruptions could irreversibly shift Earth's biogeochemical regime and
224 have cascading effects to impair global biodiversity and social-ecological resilience.
225 The accelerating breakdown of these interconnected processes supports the concern
226 about the emergence of a new planetary Hothouse Earth regime (Steffen et al., 2018).
227 Once flipped into such a regime, recovery to the previous conditions is no longer viable,
228 adaptation becomes obsolete and transformation is the only viable option for navigating
229 the accelerating Anthropocene (Angeler et al. 2020a).

230 **6. A Planetary shift: the consequences of biogeochemical extinction**

231 Suppose the regulatory processes of the carbon, nitrogen, and phosphorus cycles
232 collapse. Earth will undergo a fundamental transition toward a drastically altered
233 biosphere (Fig.1). The loss of carbonate precipitation, denitrification, phosphorus burial,
234 and methane oxidation would not only amplify greenhouse gas accumulation but also
235 trigger systemic shifts in atmospheric composition, ocean chemistry, and ecosystem
236 function. As a result, the planet would enter a new equilibrium dominated by anaerobic
237 metabolism, extreme biogeochemical imbalances, and widespread ecological
238 degradation (Fig. 1).

239 **6.1. The decline of habitability: from biosphere to chemosphere**

240 The extinction of biogeochemical processes responsible for climate regulation would
241 drive a significant decline in Earth's habitability (Fig. 1a). Without effective CO₂
242 sequestration via carbonate precipitation and terrestrial carbon storage, atmospheric C
243 concentrations would continue rising indefinitely, intensifying global warming beyond
244 current worst-case projections (Feely et al., 2004; Orr et al., 2005). Wetlands,



245 permafrost, and coastal sediments—historically net carbon sinks—would become
246 permanent methane sources, further compounding the greenhouse effect (Davidson &
247 Janssens, 2006; Bange et al., 2010). Rising temperatures and feedback-driven
248 permafrost thawing would release vast amounts of stored C (primarily as CH₄),
249 accelerating global aridification and heat stress (Lenton et al., 2008; Fig. 1a).

250 Simultaneously, oceanic oxygen depletion would expand beyond coastal dead zones,
251 reducing habitable marine environments (Rabalais et al., 2010). The loss of
252 denitrification and anammox would lock reactive nitrogen into aquatic systems, fueling
253 sustained eutrophication and reinforcing widespread anoxia (Conley et al., 2009). The
254 intensification of oxygen minimum zones (OMZs) in midwater layers, driven by
255 increased organic matter input and nutrient loading, coupled with excess nitrogen and
256 phosphorus, would create conditions favoring sulfate-reducing bacteria, leading to
257 periodic hydrogen sulfide (H₂S) releases—a known driver of mass extinctions in Earth’s
258 past (Canfield, 1998; Knoll et al., 2007).

259 On land, agricultural collapse would result from P depletion and N saturation. Primary
260 production would decline in P-limited regions, leading to soil degradation and
261 widespread desertification (Elser et al., 2007; Sternberg et al., 2017). In N-saturated
262 systems, toxic algal blooms and soil acidification would disrupt terrestrial food webs,
263 further reducing ecosystem resilience and increasing biodiversity loss (Galloway et al.,
264 2008). The combined failure of C sequestration, N removal, and P stabilization would
265 make large areas of the planet inhospitable for higher-order life forms, shifting Earth’s
266 surface toward a chemically unstable, hypoxic, and high-CO₂ state.

267 **6.2. A new biogeochemical regime: anaerobic metabolism and microbial** 268 **dominance**

269 The expansion of anoxic zones, rather than a complete loss of oxygen throughout the
270 ocean, would shift carbon processing toward anaerobic metabolic pathways—
271 particularly sulfate reduction, methanogenesis, and fermentation—reinforcing localized
272 anoxia and greenhouse gas emissions (Schlesinger, 1997; Fig. 1b). Sulfate-reducing
273 bacteria (SRB) would outcompete aerobic decomposers in these oxygen-depleted
274 regions, generating high concentrations of H₂S, which is not only toxic to most



275 metazoans but also exacerbates oxygen depletion through further microbial and
276 chemical interactions (Canfield, 1998; Fig. 1b).

277 Methanogenic archaea would proliferate in wetlands, anoxic oceans, and thawing
278 permafrost, shifting the C cycle balance toward CH₄ accumulation (Bange et al., 2010).
279 The extinction of aerobic methane-oxidizing bacteria (MOB) would eliminate a key
280 sink for CH₄, further exacerbating warming (Ettwig et al., 2010; Raghoebarsing et al.,
281 2006). In N cycling, the decline of denitrification would favor ammonification,
282 increasing ammonium accumulation (NH₄⁺). At the same time, losing anammox would
283 remove a major pathway for N removal from aquatic systems, leading to a chemically
284 reduced, eutrophic environment (Jetten et al., 2001).

285 As these shifts progress, macroorganisms dependent on oxygen and stable nutrient
286 cycles would face increasing extinction risks. Complex multicellular life, particularly
287 large vertebrates, would struggle to survive in a world dominated by water scarcity,
288 anoxia, GHG buildup, and disrupted primary production. Ecosystem productivity would
289 decline, leading to food web simplification and the collapse of biodiversity hotspots
290 (Steffen et al., 2015). In contrast, microbial extremophiles adapted to anoxic, high-
291 sulfide, and high-methane environments would thrive, forming the dominant
292 biosphere—a shift reminiscent of early Earth conditions before the Great Oxidation
293 Event (Fig. 1b).

294 **6.3. From an oxygen-rich earth to a methane and sulfide-dominated world**

295 Under current conditions, the biosphere is largely regulated by oxygenic photosynthesis,
296 aerobic respiration, and nutrient cycling that, despite local imbalances, maintains large-
297 scale biogeochemical stability (Fig. 1c). These processes maintain atmospheric O₂
298 levels, drive carbon sequestration, and stabilize N and P fluxes. However, with the
299 collapse of oxygenic biogeochemical functions, the planet would transition into a high-
300 CO₂, high-CH₄, and high-H₂S regime, chemically resembling Earth's Precambrian
301 epochs or situations of past mass extinction events (Knoll et al., 2007; Lenton et al.,
302 2008; Fig. 1c).

303 Due to nutrient imbalances, ocean acidification, and CO₂ oversaturation, photosynthetic
304 primary producers would face severe limitations, leading to a net decline in ecosystem
305 productivity. Instead, chemosynthetic and anaerobic microbial processes—such as



hydrogenotrophic methanogenesis, sulfur-based energy metabolism, and fermentative C cycling—would dominate biogeochemical fluxes (Falkowski et al., 2000). The terrestrial and marine C cycles would become increasingly driven by microbial respiration rather than plant sequestration, fundamentally altering ecosystem structure.

This transition would render large portions of the planet inhospitable to complex life (Fig. 1c). Marine dead zones, initially localized in eutrophic coastal regions, would expand into open oceans, reducing viable habitats for fish, corals, and other marine organisms. On land, desertification and altered precipitation patterns would lead to the contraction of forests and grasslands, further reducing carbon uptake (Davidson & Janssens, 2006). The biosphere would be dominated by microbial communities adapted to extreme environments, similar to the deep-sea hydrothermal vent ecosystems that persist under oxygen-deficient, high-sulfide conditions.

Conclusions: a planetary shift beyond recovery

The escalating disruption of Earth's biogeochemical cycles risks the irreversible extinction of critical regulatory processes, exacerbating global warming, ecosystem collapse, and nutrient imbalances (Fig. 1). The loss of microbial pathways governing methane oxidation, nitrogen transformation, and phosphorus burial threatens to amplify feedback loops that push planetary stability beyond recoverable limits.

The extinction of these known biogeochemical processes would push Earth beyond its current stable regime, leading to a chemically reduced, microbially dominated world with diminished habitability for complex life. The collapse of CO₂ sequestration, N removal, and P stabilization would accelerate global warming, disrupt nutrient cycles, and shift ecosystem function toward anaerobic metabolism. The resulting feedback loops would lock the planet into a high-CO₂, high-CH₄, and low-O₂ equilibrium, where microbial extremophiles outcompete oxygen-dependent life.

Governance and mitigation must go beyond emissions reduction to actively maintain biogeochemical stability. Prioritizing wetland conservation, soil carbon restoration, and nutrient pollution control can delay the collapse of key ecosystem functions, allowing us to gain more time for designing strategies to navigate the Anthropocene for best-possible sustainability. Advancing microbial ecology research, carbon sequestration modeling, and N and P cycles restoration is essential to integrate biogeochemical



337 feedbacks into climate mitigation policies. Without urgent action based on
338 interdisciplinary collaborations across sectors of society that embrace the known and
339 unknown consequences of rapid social-ecological change (Angeler et al. 2020b), the
340 irreversible extinction of biogeochemical processes may accelerate planetary instability,
341 rendering stabilization of the Holocene climate unattainable.

342 Critics may argue that the scenarios laid out are alarmist or that the evidence lacks
343 quantifiable precision. However, this article intends to highlight the urgency of
344 addressing these critical issues. The development of certain scenarios aims not to
345 overstate risks, but to prepare for potential outcomes that scientific research indicates as
346 plausible. The interconnectedness and complexity of biogeochemical cycles make
347 precise predictions challenging, but this does not diminish the validity of the concerns
348 raised.

349 Furthermore, we acknowledge the need for governance and policy development to
350 mitigate these risks. While there may be calls for more quantifiable evidence or a
351 stronger emphasis on social and ecological adaptive and transformative capacity
352 (Angeler et al., 2019; Michaels et al., 2025), the primary goal is to raise awareness and
353 prompt action. Recognizing the potential for natural system resilience and human
354 innovation is crucial, and these aspects should complement, not detract from, the urgent
355 call for a global, coordinated response to prevent irreversible damage. The article's
356 emphasis on immediate action and comprehensive understanding aligns with the
357 principle that it's better to prepare for dire possibilities than to be caught unprepared.

358 **Author contribution**

359 SSC: conceptualization, writing - original draft, and visualization. All authors:
360 investigation, supervision, writing - review and editing, and final approval.

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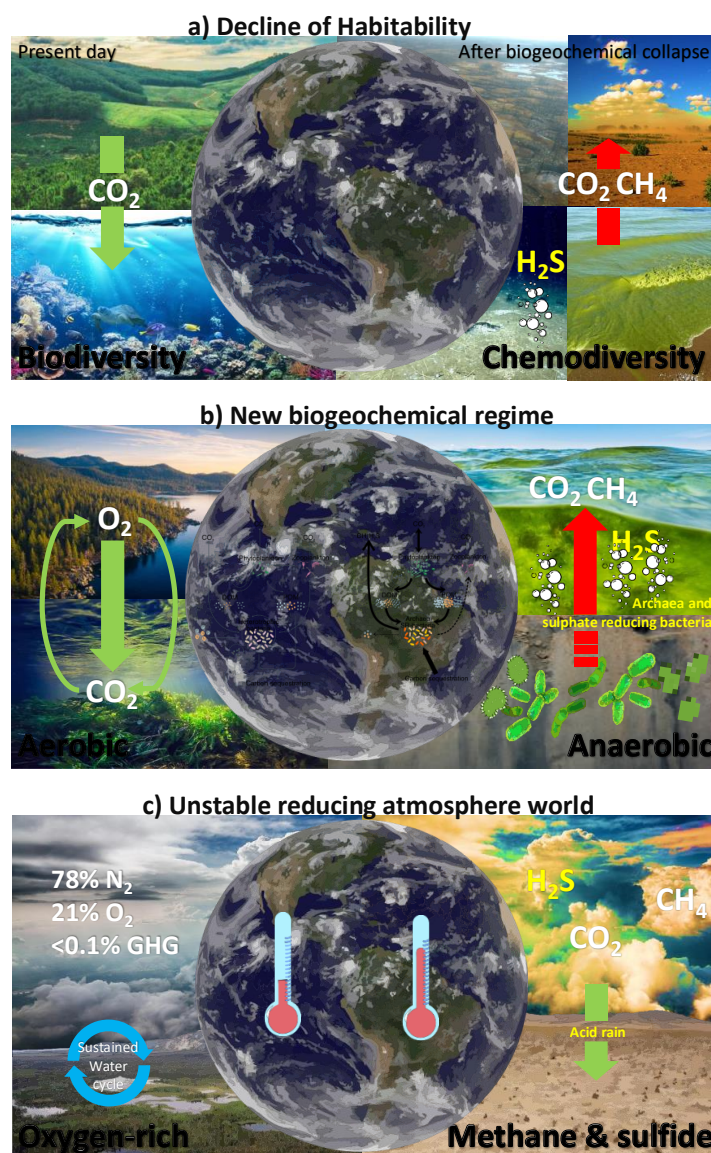


529 Table 1. Biogeochemical process disruptions under global environmental change: key
530 risks and consequences.

Biogeochemical process	Main disruptor	Potential risk	Consequence for global biogeochemistry
Carbon sequestration in oceans	Ocean acidification due to CO ₂ absorption	Reduced efficiency of biological carbon pump	Weakened oceanic carbon sink increasing atmospheric CO ₂
Soil carbon storage	Warming-induced microbial decomposition	Decline in humification, leading to faster carbon turnover	Soil carbon instability, amplifying climate feedbacks
Wetland methane cycling	Enhanced methanogenesis due to CO ₂ and temperature rise	Loss of aerobic methanotrophy, increasing methane emissions	Higher CH ₄ emissions, intensifying global warming
Carbonate precipitation & burial	Lower pH reducing carbonate ion availability	Decreased carbonate formation, reducing oceanic CO ₂ buffer	Loss of long-term marine CO ₂ sequestration, disrupting ecosystems
Anaerobic respiration pathways	Nutrient loading & climate warming favoring methanogenesis	Shift from denitrification to methanogenesis, enhancing warming	Greater CH ₄ and N ₂ O emissions, accelerating climate change
Methanotrophic processes	Oxygen depletion suppressing methane oxidation	Collapse of methane oxidation, increasing greenhouse gas flux	Methane accumulation in anoxic zones, worsening feedback loops
Nitrification & denitrification	Excess nitrogen inputs disrupting balance	Loss of anammox process, altering nitrogen cycle dynamics	Increased nitrous oxide emissions, intensifying greenhouse effect
Phosphorus precipitation	Eutrophication overwhelming phosphorus sinks	Reduced phosphorus retention, leading to long-term eutrophication	Enhanced eutrophication, creating hypoxic zones and biodiversity loss
Microbial phosphorus solubilization	Nutrient saturation shifting microbial communities	Decreased microbial P cycling, impairing ecosystem productivity	Reduced phosphorus bioavailability, limiting primary production

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534 Figure 1. The transition from an Earth regime with desirable biogeochemical
535 functioning to a degraded regime with impaired processes. (a) From a biodiversity-
536 dominated world with oxygenated oceans and stable carbon and nutrient cycles to a
537 chemosphere dominated by methane and sulfide pathways. (b) Future Earth under
538 biogeochemical collapse: anaerobic ocean zones, widespread anoxia, disrupted nutrient
539 cycles, and microbial dominance. (c) Atmospheric shift: from an oxygen-rich
540 atmosphere to a methane- and sulfide-dominated world, leading to extreme climatic
541 instability and ecosystem collapse.
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