



Sectoral attribution of greenhouse gas and pollutant emissions using multi-species eddy covariance on a tall tower in Zurich, Switzerland

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Abstract.

Eddy covariance measurement of species that are co-emitted with CO₂, such as carbon monoxide (CO) and nitrogen oxides NO and NO₂ (NO_x), provides an opportunity to attribute a total measured net flux to individual source or sink categories. This work presents eight months of continuous simultaneous measurements of fluxes (F) of CO₂, CO, NO_x, methane (CH₄), and nitrous oxide (N₂O) from an urban tall-tower (112 m agl) in Zurich, Switzerland. Median daily fluxes of F_{CO_2} were 1.47x larger in the winter (Nov-Mar) as opposed to summer (Aug-Oct) months (10.9 vs. 7.4 $\mu\text{mol m}^{-2} \text{s}^{-1}$); 1.08x greater for F_{CO} (30 vs. 28 $\text{nmol m}^{-2} \text{s}^{-1}$); 1.08x greater for F_{NO_x} (14 vs. 13 $\text{nmol m}^{-2} \text{s}^{-1}$); 1.01x greater for F_{CH_4} (13.5 vs. 13.3 $\text{nmol m}^{-2} \text{s}^{-1}$); and not statistically significantly different for $F_{\text{N}_2\text{O}}$. Flux ratios of $F_{\text{CO}}/F_{\text{CO}_2}$ and $F_{\text{NO}_x}/F_{\text{CO}_2}$ are well characterised by inventory molar emission ratios of stationary combustion and road transport in cold months. In warm months both $F_{\text{CO}}/F_{\text{CO}_2}$ and $F_{\text{NO}_x}/F_{\text{CO}_2}$ systematically exceed expected inventory ratios during the day, while no statistically significant seasonal difference is observed in $F_{\text{NO}_x}/F_{\text{CO}}$, indicating biospheric photosynthetic activity. A linear mixing model is proposed and applied to attribute half-hourly F_{CO_2} , F_{CO} , and F_{NO_x} to stationary combustion and road transport emission categories as well as determine the biospheric F_{CO_2} . Flux attribution is reasonable at certain times and from certain wind directions, but over-attributes CO and NO_x fluxes to road traffic and CO₂ fluxes to stationary combustion, and overestimates photosynthetic CO₂ uptake.

1 Introduction

The majority of the world's population now lives in urban areas, with the projected proportion to reach 68% by 2050. In the European Union, this proportion already exceeds 70% (United Nations, 2019). Cities are estimated to cause 70% of global energy-related GHG emissions (Lwasa et al., 2022) despite covering only 3% of the Earth's land surface (Liu et al., 2014). Cities are therefore crucial for emission mitigation initiatives, and cities in many parts of the world are leading emission reduction efforts, especially of carbon dioxide (CO₂, European Commission, 2021; Lwasa et al., 2022; Stadt Zürich, 2023).



To understand emissions at the city-scale, many cities have developed emission inventories based on bottom-up approaches. Estimates in the form of emission inventories are temporally coarse, and rely on a series of assumptions and scaling factors to downscale yearly bulk emissions to spatially and temporally resolved estimates, with large associated errors and uncertainties (Simon et al., 2008; Gately and Hutyra, 2017; Gurney et al., 2021).

Independent monitoring of total CO₂ emissions and the attribution of emissions to specific sectors is critical to determine the efficacy of emission reduction efforts and policy changes (Wu et al., 2016; Lauvaux et al., 2020). Eddy covariance (EC) flux measurements are currently the only method for providing direct and time-resolved in situ measurements of the vertical exchange of scalars in the atmosphere. Although urban areas violate the fundamental assumption of surface homogeneity that is inherent to estimating long-term exchange irrespective of wind direction (Aubinet et al., 2012), EC has been successfully applied in cities for over two decades to quantify dynamics of CO₂ emissions and monitor long- and short-term emission changes (e.g., Grimmond et al., 2002; Christen et al., 2011; Stagakis et al., 2019).

However, the EC method integrates all sources and sinks of CO₂ within its source area, making sectoral attribution of fossil CO₂ or separation of fossil (ff) and biospheric (bio) CO₂ difficult. Spatially and temporally resolved high-resolution emission inventories combined with footprint modelling have been used to attribute integral fluxes of CO₂ to emission sectors (e.g., Christen et al., 2011; Stagakis et al., 2023). Another promising approach is that of co-emitted species ratios (Nathan et al., 2018) in which the ratio between CO₂ and another species is used to attribute measured fluxes to specific emission processes or to separate bio and ffCO₂, e.g., by using the co-emitted species carbon monoxide (CO) and the CO/ffCO₂ ratio. Diurnal changes in concentrations of CO/ffCO₂ have been shown to correspond to diurnal changes in road traffic (Turnbull et al., 2015) in a city where CO emissions were predominantly traffic related. There is currently no method to measure ffCO₂ directly with the frequency needed to calculate EC fluxes. However, if the CO/ffCO₂ emission ratio is known then the measured flux ratios of CO and total CO₂ may be used to partition the CO₂ flux in to biospheric and fossil components (Wu et al., 2022). Further, if the sector-specific CO/ffCO₂ ratio is known, it may be possible to partition integral EC fluxes to specific emission sectors. City emission inventories therefore offer intriguing possibilities to estimate sector-specific emission ratios of CO₂ and co-emitted species, which may be used in turn to attribute integral EC fluxes to specific sectors at a half-hourly scale without the need for individual footprint modelling.

Urban EC has mostly focused on CO₂ measurements (Matthews and Schume, 2022). However, the EC method may be applied to any scalar measured with sufficient frequency. CO is a common urban pollutant and short-lived greenhouse gas produced by, among other processes, incomplete combustion of hydrocarbons, and is commonly co-emitted during combustion along with CO₂. For this reason, CO is commonly measured by air quality networks as lower-frequency (30 min to 1 h) average concentrations. Flux measurements of CO via EC are rare. A few studies have been conducted over agricultural sites (e.g., Cowan et al., 2018; Murphy et al., 2023) and CO is occasionally measured in urban environments as part of multi-species campaigns, e.g., alongside NO_x (Karl et al., 2017), CH₄ (Helfter et al., 2016), or N₂O (Famulari et al., 2010). Nitrogen oxides (NO_x) are other common urban pollutants created in combustion processes. Urban EC studies of NO_x are similarly scarce but have revealed significant underestimation of traffic emissions in selected inventories (Marr et al., 2013; Lee et al., 2015; Vaughan et al., 2016; Karl et al., 2017). Methane (CH₄) is a potent greenhouse gas also emitted through



human activity. While methane is released by combustion processes, it is less suitable as a tracer for fossil fuel emissions given its abundant non-combustion sources like leakage of natural-gas distribution networks, and non-fossil sources like landfills and agriculture (Nathan et al., 2018). Nevertheless, due to its strong global warming potential, urban methane emissions are receiving increasing attention, though only a few direct urban EC studies exist (Helfter et al., 2016; Pawlak and Fortuniak, 2016; Sticher et al., 2024). Nitrous oxide (N_2O) is another long-lived potent greenhouse gas emitted through anthropogenic activities. Similarly to CH_4 , it is less suited as an ffCO_2 tracer as there are many non-fossil production mechanisms of N_2O , and anthropogenic N_2O emissions are overwhelmingly agricultural on a global scale (Nathan et al., 2018; Tian et al., 2020). However, N_2O fluxes may be a useful tracer for fossil-fuel emissions within cities themselves where they have been shown to scale well with road traffic (Famulari et al., 2010) and have shown no seasonal variability in a semi-urban site (Järvi et al., 2014).

In this work we present the results of a study that measured simultaneous eddy covariance fluxes of CO_2 and co-emitted species CO , NO_x , CH_4 , and N_2O from a tall-tower near the city centre of Zurich, Switzerland, using a single instrument employing laser absorption spectroscopy. To our knowledge this is the first time that EC fluxes of these species have been measured together. We present aggregated diurnal and seasonal fluxes of each species in Sec. 3.1. In Sec. 3.2 we present ratios of the three most common and well-correlated species (CO_2 , CO , and NO_x) and compare these ratios to those from a city emission inventory. Finally in Sec. 3.3 we present a linear mixing model for partitioning total fluxes into source sectors and for estimating the biogenic CO_2 flux from observed species flux ratios and their expected ratios from the inventory.

2 Methods

2.1 Study Location and Site Description

The city of Zurich is the largest city in Switzerland with approximately 443,000 inhabitants in the city proper and approximately 1.6 million inhabitants in the wider metropolitan area. The city is situated at the northern edge of Lake Zurich at approximately 400 m asl in a valley surrounding the Limmat river running roughly SE-NW. The valley floor is relatively flat with a gentle slope towards the NW. The valley edges reach about 600 m asl to the N/NE and about 850 m asl to the S/SW (Fig. 1).

2.2 Instrumentation and Measurements

Eddy covariance (EC) measurements were carried out from August 2022 to March 2023 roughly 1.5 km NW of the city centre on a communications antenna on the top of a high-rise apartment building (Bullingerstrasse 73: $47^\circ 22' 52''$ N, $8^\circ 30' 26''$ E) built in the late 1970s. The building itself is 95 m tall and square in plan form with sides of 22.8 m. The instruments were installed 17 m above rooftop level on the antenna. The building is significantly taller than the surrounding buildings, the mean height of which is $13.3 \text{ m} \pm 8 \text{ m}$ in a 1.5 km radius, though the building is in a cluster of four nearly identical high-rise residential buildings set 43 to 68 m apart with heights 76 m (S), 85 m (SW), and 66 m (NNW). The displacement height within this 1.5 km radius is 23.4 m according to the parameterisation of Kanda et al. (2013).



Table 1. Variables and descriptions measured from the open-path (Campbell Scientific IRGASON) and closed-path (MIRO MGA7) EC systems.

OPEC (Campbell Scientific IRGASON)			CPEC (Miro MGA7)		
Variable	Description	Units	Variable	Description	Units
u_x	Longitudinal wind	m s^{-1}	χ_{CO_2}	CO ₂ molar fraction	mol mol^{-1}
u_y	Lateral wind	m s^{-1}	$\chi_{\text{H}_2\text{O}}$	H ₂ O molar fraction	mol mol^{-1}
u_z	Vertical wind	m s^{-1}	χ_{CO}	CO molar fraction	mol mol^{-1}
T_s	Sonic temperature	°C	χ_{NO}	NO molar fraction	mol mol^{-1}
T_a	Air temperature	°C	χ_{NO_2}	NO ₂ molar fraction	mol mol^{-1}
p	Air pressure	kPa	$\chi_{\text{N}_2\text{O}}$	N ₂ O molar fraction	mol mol^{-1}
ρ_{CO_2}	CO ₂ density	mg m^{-3}	χ_{CH_4}	CH ₄ molar fraction	mol mol^{-1}
$\rho_{\text{H}_2\text{O}}$	H ₂ O density	g m^{-3}			

The installation consisted of one integrated open-path eddy covariance system (OPEC) and a high-frequency closed-path (CPEC) multi-species gas analyser. The measured gases and other variables from these two instruments are summarised in Table 1. The open-path IRGASON (Campbell Scientific, Logan UT, USA) integrates a 3-dimensional sonic anemometer-thermometer with an open-path infrared gas analyser that measures CO₂ and H₂O densities. Both wind and gas measurements were acquired at 20 Hz. The IRGASON was mounted 111.8 m agl and oriented towards W (270° from N). Logging was performed on a Campbell Scientific CR6 datalogger (Campbell Scientific, Logan UT, USA).

The closed-path analyser was an MGA7 (Multi-species Gas Analyser, 7 species, MIRO Analytical, Wallisellen, Switzerland) located in a machine room on the roof-level of the high-rise with an inlet co-located with the IRGASON. The analyser was placed in a temperature-controlled enclosure for temperature stability. The inlet was connected to the analyser via approximately 33 m of Synflex tubing (8 mm outer diameter, 5.6 mm inner diameter). The MGA7 measured 7 different gas species (summarised in Table 1) simultaneously using mid-infrared quantum cascade laser spectroscopy (Li et al., 2013) at high frequencies suitable for EC measurements (nominally 10.3 Hz). The MGA7 measured all gases as a wet molar fraction (mol mol⁻¹). A NeoDry30E-2DBK (Kashiyama Industries, Saku City, Nagano Japan) dry vacuum pump downstream of the analyser drew sample air through the analyser at approximately 10 SL min⁻¹.

Fig. 1 shows the measurement location marked with a yellow triangle and land cover classification along with the full campaign 80% source area for the flux measurements derived with the FFP flux footprint model of Kljun et al. (2015). When weighting by the footprint, the dominant surface cover is vegetation (38%). Residential buildings comprise 10% while industrial and commercial buildings cover 3.5% and 4%, respectively. Roads are 9% of the footprint and rails 7.4%. The final 25% is categorised as *other*, namely bare or paved surfaces such as foot paths.

Wind characteristics at the measurement site and for the measurement period are shown in the wind roses in Fig. 2, separated by summer/autumn months (Aug-Oct, left) and winter/spring (Nov-Mar, right). In this work the terms summer or warm months

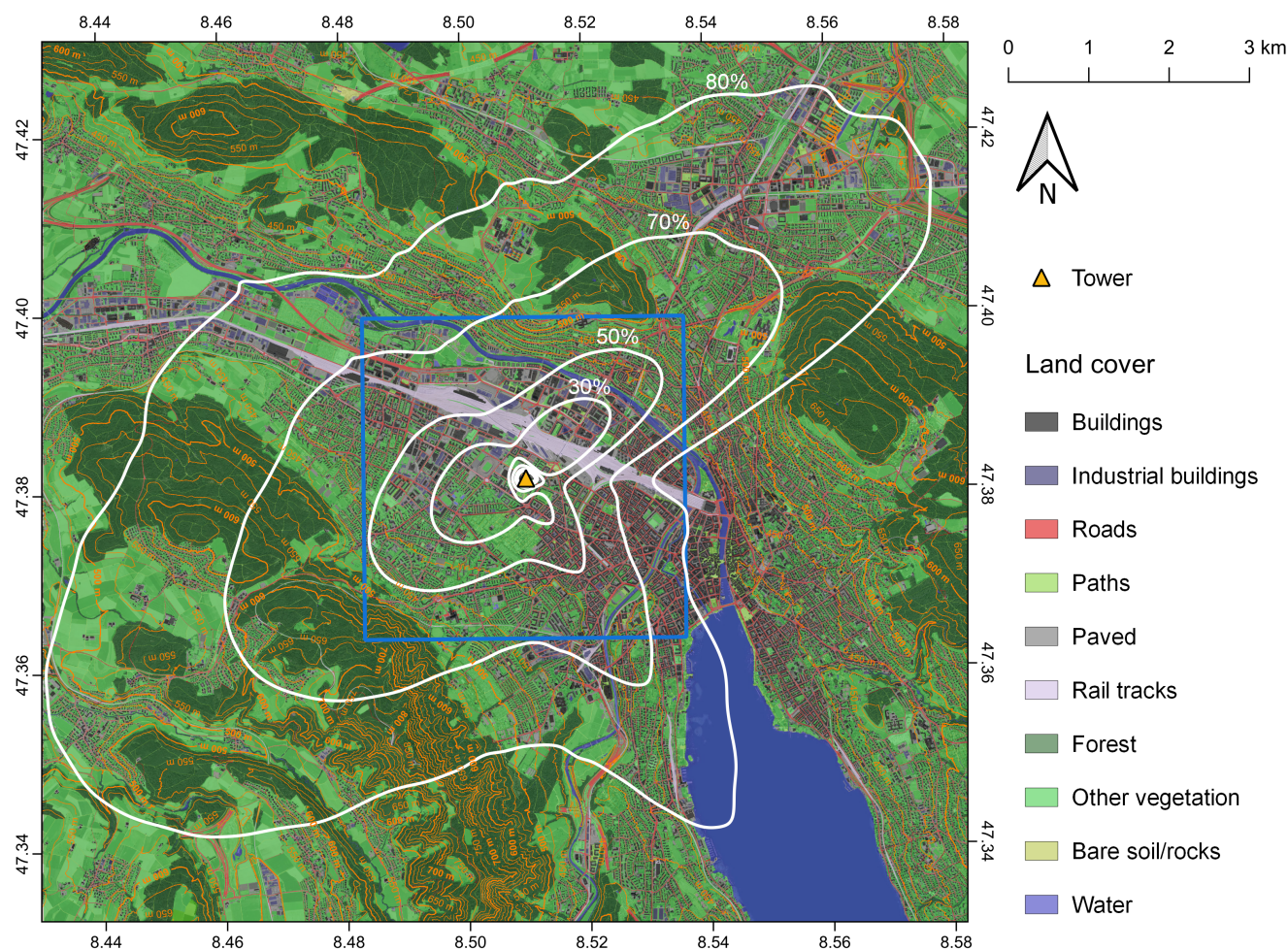


Figure 1. Location of measurement site (yellow triangle) over a land cover map of Zurich (Amtliche Vermessung). Isolines up to the 80% footprint according to Kljun et al. (2015) are overlaid in white. The blue box centred on the tower shows the 4 x 4 km² box centred on the tower from which reference inventory molar emission ratios are determined (Sec. 2.4).

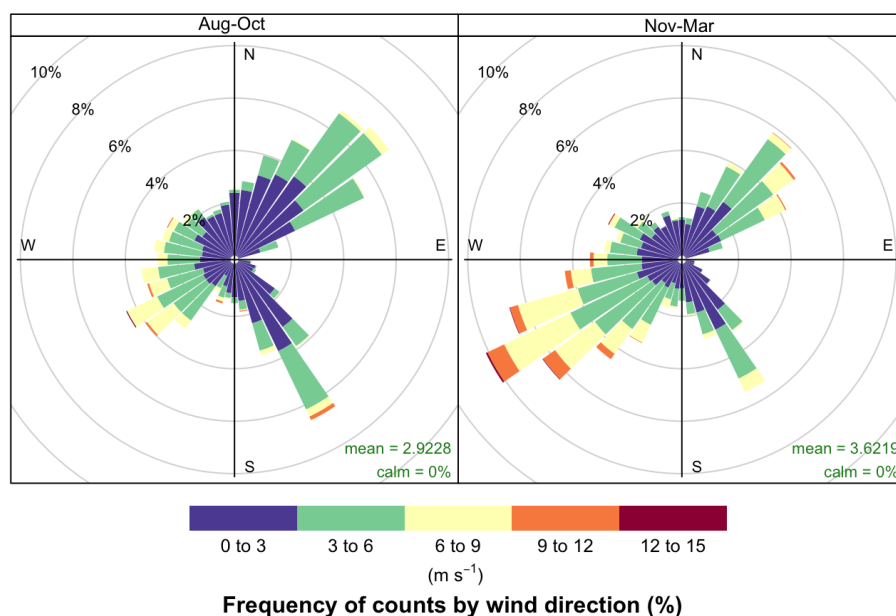


Figure 2. Wind regime for half-hour periods during the observation period which pass the CO₂ quality control tests.

will be used to refer to the Aug 1, 2022 - Oct 31, 2022 measurement period and the terms winter or cold months will be used
 110 for the Nov 1, 2022 - Mar 31, 2023 measurement period interchangeably. The chosen seasonal split broadly corresponds with
 changes in local emissions due to heating degree days as well as the change of local time from Central European Daylight
 Savings Time (CEDT, UTC+2) to Central European Standard Time (CET) on the morning of 30 October 2022. In the summer
 months (left) the wind is more frequently from the lake (SE) and calmer than the winter months, with a mean wind speed of
 2.9 m s⁻¹. In the winter months the wind is trimodal along an axis perpendicular to the valley, with one mode coming from the
 115 SW and another from the NE, and generally higher wind speeds with a mean of 3.6 m s⁻¹.

2.3 Processing and Quality Control

The OPEC (IRGASON) and CPEC (MGA7) systems were sampled at different rates which necessitated the merging of these
 two datasets. To create a harmonious dataset for EC calculations the gas measurements of the MGA7 were upsampled to 20
 Hz and merged with the IRGASON measurements using a nearest-neighbour approach with a 50 ms search window. Both
 120 systems were synchronised to UTC via an NTP server: The MGA7 once per week and the IRGASON logger every 24h.
 To account for unexpected clock drift and travel time of the sample gas through the inlet to the MGA7, the systems were
 synchronised by finding the time lag of maximum correlation between the 20 Hz CO₂ signals. Time lags for individual 30-
 minute averaging periods in which correlation was poor (correlation coefficient < 0.5) due to low IRGASON signal strength



were linearly interpolated to correct clock drift and then determined via covariance maximisation with the vertical wind using
125 a search window of ± 0.5 s. The median time lag was 4.15 s.

Open path measurements can be affected by erroneous spikes due to objects interfering (e.g. rain, dirt, insects along the
optical path) with the path. To remove problematic data but retain real increases in mixing ratios, a despiking algorithm was
run over the time series of vertical wind, CO_2 , and H_2O . Despiking of the open-path measurements of the IRGASON was
conducted using a modification of the median absolute deviation method described by Mauder et al. (2013), namely using the
130 upper and lower limits defined by Mauder et al. (2013) but keeping observations in which 3 or more consecutive outliers occur
to account for the large observed skewness of gas and temperature measurements typically observed in cities. No despiking
was performed on the closed path signals from the MGA7. Regular inspection of raw data did not reveal any data points that
would be considered spikes, and the environmental concerns that tend to cause spiky measurements in the open-path data do
not affect the closed-path system.

135 Calculation of turbulent fluxes F_χ was conducted via the EddyPro software (v7.0.9, Li-COR Biosystems) for 30-minute
blocks. Coordinate rotation was performed via double rotation and turbulent fluctuations defined by block average (Rebmann
et al., 2012). The Moncrieff et al. (1997) correction for high-pass filtering effects was used for both the OPEC and CPEC fluxes.
To account for high-frequency spectral attenuation in the CPEC, the low-pass correction according to Fratini et al. (2012) was
applied. The low-pass spectral correction according to Moncrieff et al. (2004) was applied to the OPEC.

140 Initial quality control was performed using the 0-1-2 flag system of Mauder and Foken (2004). This system incorporates two
tests: the steady state test and the integral turbulence characteristics (ITC) test, both according to Foken and Wichura (1996).
Each gas species is tested separately. Data flagged either 0 or 1 were included in the present analysis; data flagged 2 were
excluded.

The quality flag for flux ratios (e.g., $F_{\text{CO}}/F_{\text{CO}_2}$) was set to the highest flag of either component species, i.e., a period in
145 which the F_{CO} flag is 0 and the F_{CO_2} flag is 1 will result in $F_{\text{CO}}/F_{\text{CO}_2}$ flag 1. Winds from E were distorted by the antenna,
mounting equipment, and IRGASON instrument body. Therefore periods in which the mean vector wind direction over 30
minutes was between 70 and 100 degrees from N were excluded. A friction velocity (u_*) filter was also implemented. Friction
velocity filters are an accepted way to remove periods in which measured EC fluxes tend to decouple from surface fluxes
(Aubinet et al., 2012) though their usage in urban EC, where friction velocities tend to be greater, is debated (Matthews and
150 Schume, 2022). An examination of CO_2 fluxes binned by u_* showed significant and systematic decrease in flux magnitude at
 $u_* < 0.2 \text{ m s}^{-1}$ however, and therefore these observations were removed. Finally, periods in which there were known issues
with the MGA7 instrument due to software or hardware errors were removed. Fluxes were aggregated based on hourly binning
and therefore no attempt at gap-filling or quantification of error or uncertainty for individual flux periods was performed. The
storage flux was calculated using the default single-point estimation of EddyPro as no profile measurements below the tower
155 were available.

The study period of 1 August 2022 to 31 March 2023 contains 11 663 half-hour flux averaging periods. The distribution of
data quality is summarised in Table 2. Retention for the MGA7 species ranges from 53% (N_2O and H_2O) to 61% (CO_2 , CO ,
 NO_2) of the full measurement period. This is similar to the average retention value found in other urban EC studies (Matthews



Table 2. Gas fluxes and QC flags overview indicating the number of half-hour periods in each category. The final two columns indicate how many periods are kept for analysis as a number and as a percentage of the total number of half-hour periods during the study. Periods excluded by the u_* and other additional filtering are assigned flag 2.

Species	0-flag	1-flag	2-flag	NA	Incl. (n)	Incl. (pct)
F_{CO_2}	3436	3667	3292	1268	7103	61
$F_{\text{H}_2\text{O}}$	1940	4258	4101	1364	6198	53
F_{CO}	3463	3654	3278	1268	7117	61
F_{NO_x}	3633	3636	2921	1473	7269	62
$F_{\text{N}_2\text{O}}$	2694	3534	4167	1268	6228	53
F_{CH_4}	2334	3978	4083	1268	6312	54
$F_{\text{CO}}/F_{\text{CO}_2}$	2859	3920	3616	1268	6779	58
$F_{\text{NO}_x}/F_{\text{CO}_2}$	2971	3768	3451	1473	6739	58
$F_{\text{NO}_x}/F_{\text{CO}}$	2950	3793	3447	1473	6743	58
F_{CO_2} (IRGA)	2818	3082	4999	764	5900	51

and Schume, 2022, Table 1), though the wide range of different quality control regimes creates a large spread in retention rates. The large number of missing values are instrument downtime, both planned and unplanned, and reduce the data retention by 7-8%. The wind direction filter subsequently removed 253 periods (2.2%). The u_* filter removed a significant number of periods: between 7 and 10% depending on the species.

Initially calculated NO_x fluxes were significantly lower than would be expected from the city emissions inventory. A comparison of the average concentration of NO and NO_2 against measurements from local air quality measurement stations revealed a systematic underestimation of these species by the MGA7. While photochemical interactions may affect concentrations of NO and NO_2 between emission sources and the measurement tower, total NO_x should be conserved on the short (few minutes maximum) time scales between emission and measurement considered here (Lee et al., 2015). Therefore the surface NO_x concentrations were calculated for a local air quality measurement station (Zurich Kaserne, 47° 22' 39.23", 8° 31' 49.54") located in a park away from heavy emission sources such as traffic, and the total NO_x concentration from the MGA7 was calculated by summing NO and NO_2 molar fractions. The raw MGA7 measurements were scaled against the Kaserne measurements using a linear best fit model ($r^2 = 0.65$, RMSE 6.1 ppb) of hourly-averaged concentrations under day-time unstable conditions with wind speeds $> 3 \text{ m s}^{-1}$, assuming that vertical and horizontal concentration gradients are small under these conditions. The slope correction of 2.33 was applied to the raw MGA7 concentrations and a total NO_x flux calculated.

2.4 Emission Inventory

An emission inventory for 2022 was created by the city of Zurich on a GIS format which provides yearly total emissions of various species and pollutants for the city of Zurich (Brunner et al., 2025). The Swiss Laboratory for Air Pollution / Environ-



Table 3. Emissions inventory overview for a 4 x 4 km² box centred on the measurement tower. Categories given with their GNFR letter. Minor categories have been omitted.

Category	CO ₂ (kt yr ⁻¹)	CO (t yr ⁻¹)	NO _x (t yr ⁻¹)	CH ₄ (t yr ⁻¹)	N ₂ O (t yr ⁻¹)
Public Power (A)	15.6	1.5	9.7	1.7	0.03
Industry (B)	6.6	7.2	8.6	0.6	0.06
Other Stationary Combustion (C)	206.5	145.1	72.3	20.4	1.01
Road Transport (F)	71.6	158.9	130.6	3.5	2.60
Offroad (I)	5.1	34.7	18.5	0.1	0.21
Waste (J)	0.3	13.4	0.7	1.6	<0.01

Table 4. Molar ratios of select species and categories for the 4 x 4 km² area around the measurement tower: spatial average weighted by total mass of emissions and separated by major wind direction.

Species	Category	Ratio					Units
		All	NE	SE	SW	NW	
CO/CO ₂	Road transport	3.49	3.04	3.57	4.38	3.70	mmol mol ⁻¹
	Stationary combustion	1.10	1.28	0.95	1.40	0.94	mmol mol ⁻¹
NO _x /CO ₂	Road transport	1.74	1.62	1.77	1.81	1.88	mmol mol ⁻¹
	Stationary combustion	0.33	0.35	0.31	0.36	0.34	mmol mol ⁻¹
NO _x /CO	Road transport	0.50	0.54	0.49	0.41	0.51	mol mol ⁻¹
	Stationary combustion	0.30	0.27	0.33	0.26	0.36	mol mol ⁻¹

mental Technology of EMPA has processed this inventory according to standard Gridded Nomenclature for Reporting (GNFR) categories at 100 x 100 m² spatial resolution.

180 Within a 4 x 4 km² square centred on the tower (≈60-70% of the flux footprint for the study period) sectoral contributions as derived from the emission inventory are dominated by road transport and other stationary combustion (Table 3), accounting for 81% of CO emissions, 91% of non-respiration CO₂ emissions, 84% of NO_x emissions, 86% of CH₄ emissions, and 92% of N₂O emissions. Human respiration likely contributes an additional 10% of CO₂ and may be non-negligible for CH₄ (Brunner et al., 2025). For the purposes of sectoral attribution, characteristic molar emission ratios of the road transport and stationary combustion sectors were calculated from the inventory for the three species ratios CO/CO₂, NO_x/CO₂, and NO_x/CO. The
185 characteristic molar ratios are spatial averages weighted by the total mass of emissions for each inventory grid cell. The ratios are summarised in Table 4 for the 4 x 4 km² area surrounding the tower as well as separate ratios per cardinal direction.



2.5 Linear Mixing Model

Using simultaneous measurements of F_{CO_2} , F_{CO} , and F_{NO_x} , along with expected reference flux ratios per species pair and source sector (Table 4) a linear mixing model is proposed to attribute measured (total, $F_{x,\text{tot}}$) fluxes to biospheric ($F_{x,\text{bio}}$) and combustion origin, with combustion fluxes attributed to either road transport ($F_{x,\text{rt}}$) or stationary combustion ($F_{x,\text{sc}}$) source categories. Here $F_{x,\text{bio}}$ refers specifically to the net fluxes, i.e. biospheric sources (human, soil and plant respiration) minus biospheric sinks (photosynthetic uptake). Biogenic fuel sources such as the portion of biofuel in vehicle fuel mixture is attributed to the combusive $F_{x,\text{rt}}$.

We assume that for all three species all sources and sinks within the total flux-tower source area can be attributed to either the biospheric, road transport, or stationary combustion categories:

$$F_{\text{CO}_2,\text{tot}} = F_{\text{CO}_2,\text{bio}} + F_{\text{CO}_2,\text{rt}} + F_{\text{CO}_2,\text{sc}}. \quad (1)$$

Secondly, we assume no biospheric sources or sinks of CO or NO_x such that:

$$F_{\text{CO},\text{tot}} = F_{\text{CO},\text{rt}} + F_{\text{CO},\text{sc}}, \quad (2)$$

and

$$F_{\text{NO}_x,\text{tot}} = F_{\text{NO}_x,\text{rt}} + F_{\text{NO}_x,\text{sc}}. \quad (3)$$

From the inventory, we determine four independent ratios as described in Table 4:

$$a_{\text{rt}} = \frac{F_{\text{CO},\text{rt}}}{F_{\text{CO}_2,\text{rt}}}, \quad (4)$$

$$a_{\text{sc}} = \frac{F_{\text{CO},\text{sc}}}{F_{\text{CO}_2,\text{sc}}}, \quad (5)$$

$$b_{\text{rt}} = \frac{F_{\text{NO}_x,\text{rt}}}{F_{\text{CO}_2,\text{rt}}}, \quad (6)$$

and

$$b_{\text{sc}} = \frac{F_{\text{NO}_x,\text{sc}}}{F_{\text{CO}_2,\text{sc}}}. \quad (7)$$

Two additional ratios are defined dependent on Eqs. 4 to 7:

$$c_{\text{rt}} = \frac{a_{\text{rt}}}{b_{\text{rt}}} = \frac{F_{\text{CO},\text{rt}}}{F_{\text{NO}_x,\text{rt}}}, \quad (8)$$

and

$$c_{\text{sc}} = \frac{a_{\text{sc}}}{b_{\text{sc}}} = \frac{F_{\text{CO},\text{sc}}}{F_{\text{NO}_x,\text{sc}}}. \quad (9)$$



Using the budgets of Eqs. 1 to 3 and inventory ratios from 4 to 9 the total fluxes of NO_x and CO may be partitioned. Using Eqs. 8 and 9 in Eq. 2:

$$F_{\text{CO,tot}} = c_{\text{rt}} F_{\text{NO}_x,\text{rt}} + c_{\text{sc}} F_{\text{NO}_x,\text{sc}}, \quad (10)$$

and using Eq. 3 to eliminate $F_{\text{NO}_x,\text{rt}}$:

$$215 \quad F_{\text{CO,tot}} = c_{\text{rt}} F_{\text{NO}_x,\text{tot}} + F_{\text{NO}_x,\text{sc}}(c_{\text{sc}} - c_{\text{rt}}). \quad (11)$$

Solving for $F_{\text{NO}_x,\text{sc}}$:

$$F_{\text{NO}_x,\text{sc}} = \frac{F_{\text{CO,tot}} - c_{\text{rt}} F_{\text{NO}_x,\text{tot}}}{(c_{\text{sc}} - c_{\text{rt}})}. \quad (12)$$

Similarly

$$F_{\text{NO}_x,\text{rt}} = \frac{F_{\text{CO,tot}} - c_{\text{sc}} F_{\text{NO}_x,\text{tot}}}{(c_{\text{rt}} - c_{\text{sc}})}, \quad (13)$$

$$220 \quad F_{\text{CO,sc}} = \frac{F_{\text{NO}_x,\text{tot}} - \frac{1}{c_{\text{rt}}} F_{\text{CO,tot}}}{\left(\frac{1}{c_{\text{sc}}} - \frac{1}{c_{\text{rt}}}\right)}, \quad (14)$$

and

$$F_{\text{CO,rt}} = \frac{F_{\text{NO}_x,\text{tot}} - \frac{1}{c_{\text{sc}}} F_{\text{CO,tot}}}{\left(\frac{1}{c_{\text{rt}}} - \frac{1}{c_{\text{sc}}}\right)}. \quad (15)$$

Equations 12 to 15 can be directly solved using simultaneous measurement of F_{CO} and F_{NO_x} and thus total fluxes of CO and NO_x may be partitioned to these source sectors. These partitioned fluxes can then be used further to partition observed

$$225 \quad F_{\text{CO}_2,\text{tot}}. \text{ The } F_{\text{CO}_2,\text{tot}} \text{ budget from Eq. 1 may be rewritten as:}$$

$$F_{\text{CO}_2,\text{tot}} = F_{\text{CO}_2,\text{bio}} + \frac{F_{\text{CO,rt}}}{a_{\text{rt}}} + \frac{F_{\text{CO,sc}}}{a_{\text{sc}}} \quad (16)$$

and therefore

$$F_{\text{CO}_2,\text{bio}} = F_{\text{CO}_2,\text{tot}} - \frac{F_{\text{CO,rt}}}{a_{\text{rt}}} - \frac{F_{\text{CO,sc}}}{a_{\text{sc}}}. \quad (17)$$

Following a similar approach as above the non-biospheric CO_2 can be partitioned in two component categories:

$$230 \quad F_{\text{CO}_2,\text{sc}} = \frac{F_{\text{CO,tot}} - a_{\text{rt}}(F_{\text{CO}_2,\text{tot}} - F_{\text{CO}_2,\text{bio}})}{(a_{\text{sc}} - a_{\text{rt}})} \quad (18)$$

and

$$F_{\text{CO}_2,\text{rt}} = \frac{F_{\text{CO,tot}} - a_{\text{sc}}(F_{\text{CO}_2,\text{tot}} - F_{\text{CO}_2,\text{bio}})}{(a_{\text{rt}} - a_{\text{sc}})}. \quad (19)$$

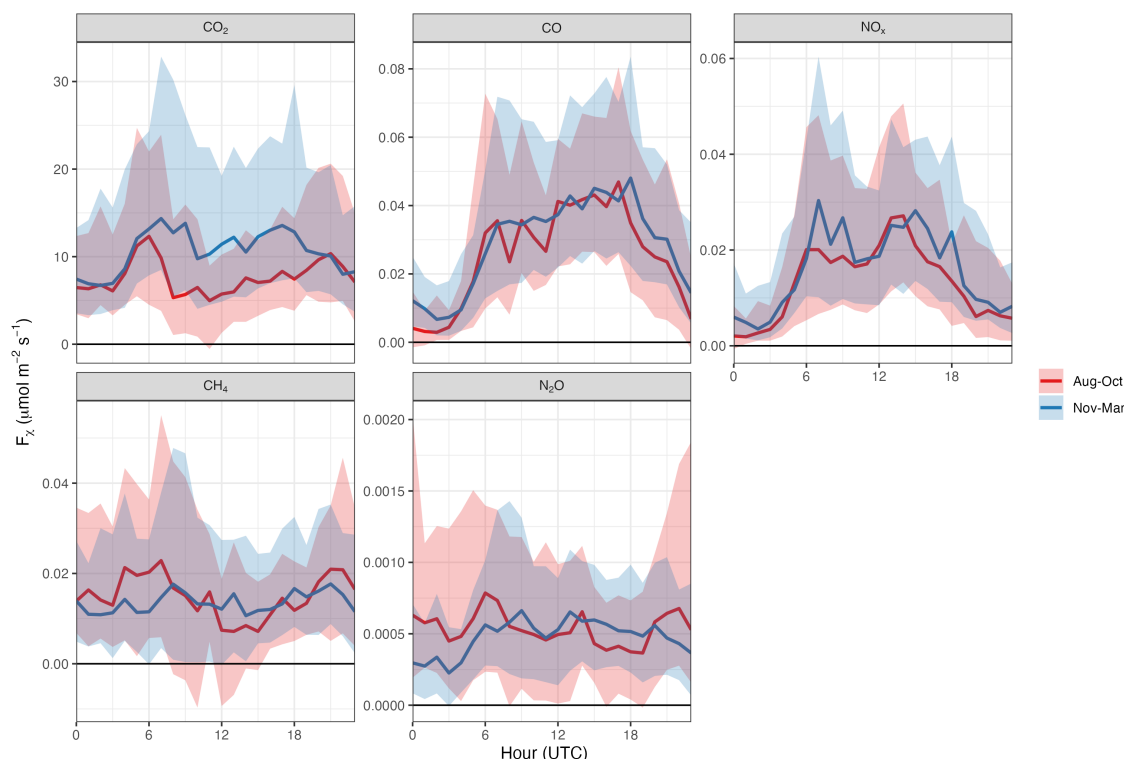


Figure 3. Diurnal and seasonal pattern of fluxes of CO_2 , CO , NO_x , CH_4 and N_2O measured by the CPEC. The solid line indicates the hourly median and the shaded band the inter-quartile (P_{25} to P_{75}) range. Fluxes are divided by season, where red is August to October 2022 (inclusive) and blue is November 2022 to March 2023 (inclusive).

3 Results and Discussion

3.1 Fluxes of Individual Species and their Correlation

235 Measured diurnal and seasonal patterns are shown for each species in Fig. 3. The median 24-hr F_{CO_2} in Aug-Oct was $7.4 \mu\text{mol m}^{-2} \text{s}^{-1}$ compared to $10.9 \mu\text{mol m}^{-2} \text{s}^{-1}$ in Nov-Mar with pronounced diurnal variability. F_{CO_2} for both seasons begins to peak in the morning hours of 5-6 UTC (7-8 local time in the summer and 6-7 local time in the winter). Median nocturnal F_{CO_2} is similar between seasons, varying by $< 0.5 \mu\text{mol m}^{-2} \text{s}^{-1}$ between 22 and 4 UTC. During the day F_{CO_2} is on average higher in the winter and the median winter flux is 1.5 times greater than the median summer flux. Fluxes are more skewed in the

240 winter where the mean F_{CO_2} is 1.6 times the median, compared to the summer in which the mean is 1.4 times greater than the median. The largest median CO_2 fluxes come from the direction of the city centre to the SE ($9.42 \mu\text{mol m}^{-2} \text{s}^{-1}$) while the smallest median fluxes ($5.50 \mu\text{mol m}^{-2} \text{s}^{-1}$) come from the more vegetated SW (not shown).

In contrast, there is much less seasonal variation in the CO fluxes, with median (mean) F_{CO} of 30 (43) $\text{nmol m}^{-2} \text{s}^{-1}$ in the winter compared to 28 (37) $\text{nmol m}^{-2} \text{s}^{-1}$ in the summer, a median winter enhancement of 8%. A Welch's t -test indicates



245 significant seasonal difference at $p < 0.01$. The clear one hour UTC offset in diurnal profiles between the seasons indicates that CO fluxes are dominated by anthropogenic rhythms, such as the work day and commuting hours, rather than responses to environmental conditions such as heating degree days. Nocturnal F_{CO} is on average 2 times greater in the winter compared to the summer. In the winter, average CO fluxes were greatest from the SE ($44 \text{ nmol m}^{-2} \text{ s}^{-1}$), and in summer approximately even between the SE and NE (34 and $35 \text{ nmol m}^{-2} \text{ s}^{-1}$, respectively).

250 Median hourly NO_x fluxes exhibit a bimodal diurnal course. While the average magnitudes are similar between seasons in the afternoon mode, they are about 1.25x greater in the winter compared to the summer during the morning mode. The median (mean) F_{NO_x} was 14 (24) $\text{nmol m}^{-2} \text{ s}^{-1}$ in winter and 13 (20) $\text{nmol m}^{-2} \text{ s}^{-1}$ in the summer, a winter enhancement of 8% (20%). The seasons are significantly different at $p < 0.01$. Some of the seasonal difference may be accounted for by seasonal changes in the wind regime. For example, there is little seasonal difference observed in the diurnal course of NO_x from the NE
255 or SW alone, but a larger sampling of SW directions during the morning rush hours in the summer months contributes to the appearance of large seasonal variation during these hours.

F_{CH_4} and $F_{\text{N}_2\text{O}}$ exhibit less seasonal difference and less clear diurnal patterns. The highest CH_4 fluxes are measured in the summer months between 4-7 UTC and 21-22 UTC when the median flux is between 20 and $23 \text{ nmol m}^{-2} \text{ s}^{-1}$. The summer also sees the lowest median fluxes between 12 and 15 UTC when the median flux is between 7 and $8 \text{ nmol m}^{-2} \text{ s}^{-1}$. In contrast,
260 the median range is smaller in the winter, between 11 and $19 \text{ nmol m}^{-2} \text{ s}^{-1}$. CH_4 is the only species for which nocturnal fluxes are on average higher than daytime fluxes. The median winter flux is only 1% greater compared to summer: 13.5 vs. $13.3 \text{ nmol m}^{-2} \text{ s}^{-1}$, and the mean winter flux is 18% greater: 23.9 vs. $20.3 \text{ nmol m}^{-2} \text{ s}^{-1}$. Nevertheless the seasons are statistically significantly different at $p < 0.01$. N_2O is the only species for which winter fluxes were observed to be smaller on average than summer fluxes, the median being 5% lower (0.49 to $0.52 \text{ nmol m}^{-2} \text{ s}^{-1}$) and the mean 7% lower (0.83 to $0.89 \text{ nmol m}^{-2} \text{ s}^{-1}$).
265 N_2O is the only species for which winter fluxes were observed to be smaller on average than summer fluxes, the median being 5% lower (0.49 to $0.52 \text{ nmol m}^{-2} \text{ s}^{-1}$) and the mean 7% lower (0.83 to $0.89 \text{ nmol m}^{-2} \text{ s}^{-1}$). However, a Welch's t -test indicates no statistically significant difference in mean $F_{\text{N}_2\text{O}}$ between the seasons ($p = 0.34$). Similarly to CH_4 , the nocturnal $F_{\text{N}_2\text{O}}$ was higher in the summer compared to the winter whereas the daytime $F_{\text{N}_2\text{O}}$ was lower, a ratio of 1.4 compared to 0.9 on average.

The hourly mean and median fluxes were integrated and the results are summarised in Table 5. For greenhouse gases CH_4 and N_2O , the CO_2 equivalent mass flux is also given using the 100-year global warming potential according to Forster et al.
270 (2021). Based on the mean day, the greenhouse gas emissions within the tower flux footprint consist of 95.8% CO_2 , 1.8% CH_4 , and 2.4% N_2O in the summer, and 97.2% CO_2 , 1.4% CH_4 , and 1.4% N_2O in the winter.

The correlation (Pearson's correlation coefficient r) between the fluxes of all species pairs on a 30-min scale is shown in Fig. 4. In the summer months (red, lower left) the highest correlation is seen between F_{NO_x} and F_{CO} at 0.72. Correlations at or above 0.6 were also observed between F_{CO_2} and F_{CO} , as well as between F_{CO_2} and F_{CH_4} . The lowest correlations (≤ 0.3)
275 were observed between $F_{\text{N}_2\text{O}}$ and F_{CO} , $F_{\text{N}_2\text{O}}$ and F_{NO_x} , as well as F_{NO_x} and F_{CH_4} . In the winter months (blue, upper right), correlations between nearly all species pairs increase, with the exception of F_{CO_2} and F_{CH_4} which decreases from 0.60 to 0.48. Correlations between F_{CO} and F_{NO_x} as well as F_{CO_2} and $F_{\text{N}_2\text{O}}$ remain essentially unchanged (0.01 difference between seasons). Co-emitted species CO and NO_x are selected as tracers for ffCO_2 in the following sections on the basis of their larger median fluxes and strong correlations in the winter.



Table 5. Summary of species fluxes. Results are split by season (Aug-Oct and Nov-Mar) and given for the median day (based on hourly medians) and mean day (based on hourly means). Results are presented as molar fluxes in $\text{mol m}^{-2} \text{yr}^{-1}$ and mass fluxes in $\text{t km}^{-2} \text{yr}^{-1}$. For mass fluxes of CH_4 and N_2O the CO_2 equivalent (CO_2e) mass flux is also given using 100-yr global warming potential factors of 28 (CH_4) and 273 (N_2O) (Forster et al., 2021).

		Summer (Aug-Oct)		Winter (Nov-Mar)	
		Molar flux ($\text{mol m}^{-2} \text{yr}^{-1}$)	Mass flux ($\text{t km}^{-2} \text{yr}^{-1}$)	Molar flux ($\text{mol m}^{-2} \text{yr}^{-1}$)	Mass flux ($\text{t km}^{-2} \text{yr}^{-1}$)
Mean	CO_2	357.0	15 711.7	501.24	22 059.8
	CO	1.06	29.81	1.30	36.44
	NO_x	0.61	28.18	0.71	32.78
	CH_4	0.72	10.82 (302.8 CO_2e)	0.76	11.43 (320.0 CO_2e)
	N_2O	0.032	1.43 (390.8 CO_2e)	0.026	1.17 (318.1 CO_2e)
Median	CO_2	240.83	10 599.0	355.13	15 629.2
	CO	0.80	22.56	0.93	26.19
	NO_x	0.45	20.66	0.52	23.70
	CH_4	0.48	7.23 (202.6 CO_2e)	0.45	6.83 (191.2 CO_2e)
	N_2O	0.018	0.81 (221.5 CO_2e)	0.015	0.66 (181.2 CO_2e)

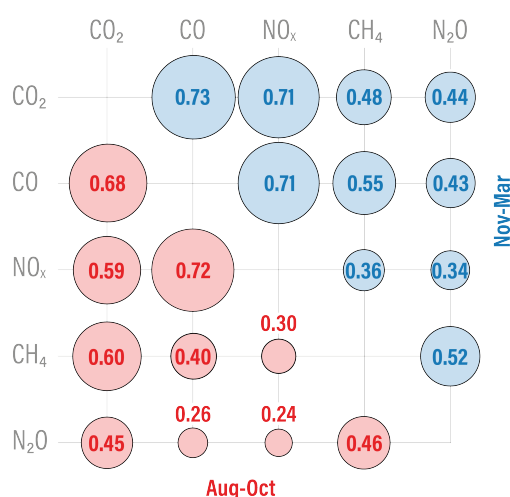


Figure 4. Correlation matrix of 30-min fluxes of measured species. The lower-left half is the summer months (Aug-Oct) and upper-right half is the winter months (Nov-Mar).



280 3.2 Ratios of Species Fluxes

Flux ratios of NO_x and CO to CO_2 were calculated along with the emission ratios for the largest emission sectors from the city emission inventory. In Fig. 5 the diurnal and seasonal ratios are shown for $F_{\text{NO}_x}/F_{\text{CO}_2}$ (top) and $F_{\text{CO}}/F_{\text{CO}_2}$ (bottom). The emission ratios determined from the inventory are shown as shaded bands using the range of values (wind-sector dependent) described in Table 4. Also shown as reference is the biospheric flux of human respiration which in both cases is 0.

285 As has been noted in other studies (Matthews and Schume, 2022; Helfter et al., 2016), tall-tower EC is more susceptible to large storage fluxes induced by boundary-layer dynamics that may be decoupled from actual surface emissions within a given 30-minute period. Emissions that accumulate within the nocturnal boundary layer and below the height of the EC system can register as anomalously large surface fluxes during the morning thermal mixing and breakup of the nocturnal boundary layer. The single-point storage correction employed here is rudimentary and may underestimate the true storage flux by over 50%
290 (Finnigan, 2006). However, even in the case of significantly underestimated storage fluxes, flux ratios should be unaffected and remain representative of the stored emission sources, which simplifies sectoral attribution. The steep increases in morning fluxes seen in CO_2 , CO, and NO_x in Fig. 3, e.g., do not correspond to similarly steep rises in their emission ratios in Fig. 5.

In both seasons and for both ratios, the diurnal minima are observed during the night. The observed $F_{\text{NO}_x}/F_{\text{CO}_2}$ nocturnal ratios are higher in winter than summer with median (mean) ratios of 0.89 (1.51) mmol mol^{-1} in the winter and 0.69 (0.76)
295 mmol mol^{-1} in the summer. Between 0 and 2 UTC in the summer, the median ratios are within 10% of the expected range of stationary combustion ratios. Outside these times, and for all hours during the winter, the median observed ratios are above the range expected for pure stationary combustion. The diurnal maxima in $F_{\text{NO}_x}/F_{\text{CO}_2}$ in both seasons are observed in the afternoon between 13 and 14 UTC. In the winter, the measured $F_{\text{NO}_x}/F_{\text{CO}_2}$ is $< 0.1 \text{ mmol mol}^{-1}$ from the inventory reference ratio for road transport between 13 and 15 UTC before gradually reducing through the afternoon and evening. These measure-
300 ments show very good agreement between measured $F_{\text{NO}_x}/F_{\text{CO}_2}$ and expected inventory NO_x/CO_2 molar emission ratios through the winter and a clear diurnal cycle that suggests observed ratios are dominated by stationary combustion through the night, dominated by road transport in the afternoon, and a mixture of the two during other times of day. This agreement is strong evidence for both the accuracy of the emission inventory as well as the possibility for emission ratios to reduce the artifacts of storage fluxes from single-point EC measurement towers. Observed nocturnal ratios are systematically higher in
305 winter than summer. This may be attributed to vegetation respiration of CO_2 , which is higher in the summer and produces $F_{\text{NO}_x}/F_{\text{CO}_2}$ of 0.

$F_{\text{NO}_x}/F_{\text{CO}_2}$ ratios observed through the summer follow a similar diurnal pattern, but with ratios that exceed the expected upper ratio from road transport between 11 and 16 UTC. Between these hours, the median measured ratio is between 0.5 and 1.9 mmol mol^{-1} greater than the expected ratio for road transport. This seasonal difference is a strong indication of photosynthetic
310 activity within the footprint causing a decrease in the absolute F_{CO_2} , and therefore a systematic increase in $F_{\text{NO}_x}/F_{\text{CO}_2}$. The median summer ratio is 1.5 mmol mol^{-1} compared to 1.2 mmol mol^{-1} in the winter and the mean ratios are 2.0 to 1.4 mmol mol^{-1} , respectively. Another seasonal difference is seen in the morning hours before 11 UTC where the summer $F_{\text{NO}_x}/F_{\text{CO}_2}$ exceeds the winter and tends on average towards the road transport ratio. This may be attributed to the use of heating during

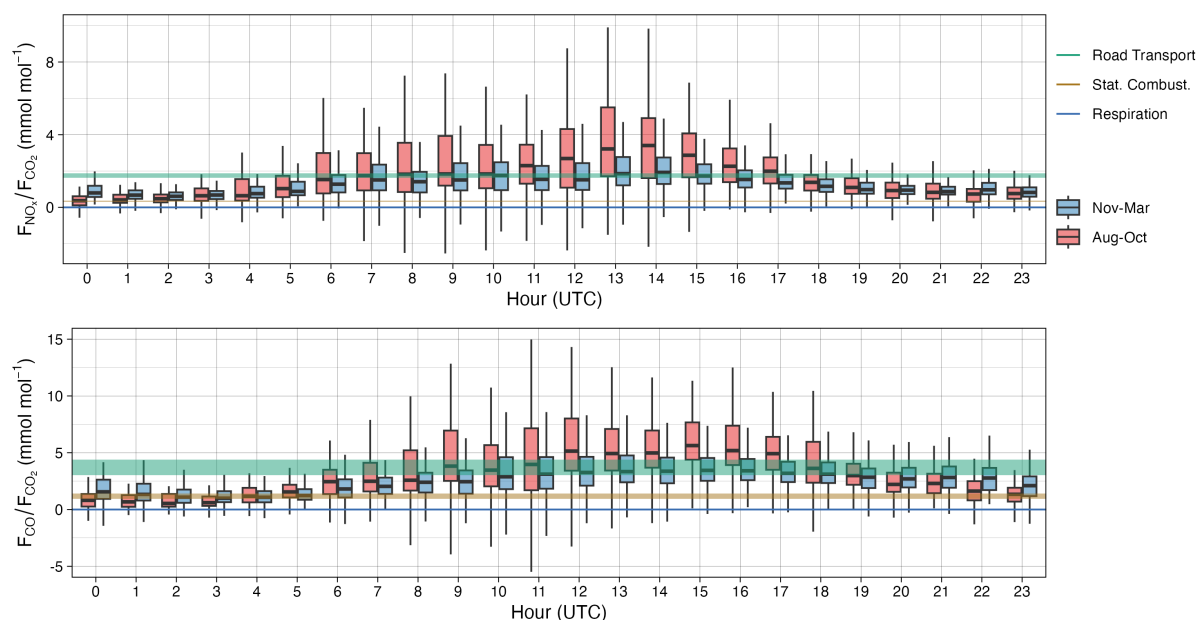


Figure 5. Diurnal ratios of F_{NO_x} to F_{CO_2} (top) and F_{CO} to F_{CO_2} (bottom) separated by season. The range of inventory emission ratios from Table 4 is shaded.

the colder winter months causing a more mixed signal, whereas in the summer months, in the absence of significant heating,
road transport is the more dominant signal.

Similar patterns are observed in the bottom panel of Fig. 5 for the $F_{\text{CO}}/F_{\text{CO}_2}$ ratios. Nocturnal ratios in the summer from 22 to 4 UTC agree well with the inventory stationary combustion ratio, on average $< 0.3 \text{ mmol mol}^{-1}$ from the median expected value. In contrast to the $F_{\text{NO}_x}/F_{\text{CO}_2}$ ratios, the $F_{\text{CO}}/F_{\text{CO}_2}$ remain elevated well above this ratio from 18 to 21 in both seasons and in the winter remain outside the expected range for stationary combustion through 23 UTC. Median nocturnal ratios in the winter are 35% greater than the summer (1.9 to $1.4 \text{ mmol mol}^{-1}$). During the winter daytime, a similar increase towards the expected road transport ratios is observed and the median flux ratio falls within this range from 12 to 18 UTC, with a mixture between the two source categories during other hours. Again in the summer there is a systematic exceedance of the upper road transport ratios during the afternoon. From 12 to 17 UTC, the median ratio is 0.7 to $1.8 \text{ mmol mol}^{-1}$ greater than the upper road transport reference ratio, and the median summer daytime ratio is 42% greater (4.4 to $3.1 \text{ mmol mol}^{-1}$) than the winter. Again, the observed $F_{\text{CO}}/F_{\text{CO}_2}$ in the winter are well bounded by the expected values determined from the emissions inventory whereas a systematic exceedance in the summer afternoon suggests biospheric influence.

Both panels in Fig. 5 suggest that reported inventory ratios are reasonable and the emissions within the tower footprint may be well characterised by a combination of stationary combustion and road transport during the winter. Nevertheless, there remain significant differences in the diurnal courses between the $F_{\text{NO}_x}/F_{\text{CO}_2}$ and $F_{\text{CO}}/F_{\text{CO}_2}$ ratios. Where the summer $F_{\text{NO}_x}/F_{\text{CO}_2}$ afternoon exceedance has one mode and peaks at 14 UTC, the $F_{\text{CO}}/F_{\text{CO}_2}$ arguably has two modes whose valley



occurs during 13 and 14 UTC, contradicting the trend of the NO_x ratios. The evening (18 to 23 UTC) winter $F_{\text{CO}}/F_{\text{CO}_2}$ ratios remain elevated towards the expected road transport ratio, suggesting a CO source in addition to road transport and stationary combustion. One possible explanation is wood combustion, which has a very high CO/CO₂ molar ratio of, e.g., 50 to 110 mmol mol⁻¹ depending on appliance and wood species (Evtyugina et al., 2014), and even a small number of wood-fired heating systems could therefore elevate the observed $F_{\text{CO}}/F_{\text{CO}_2}$.

The diurnal and seasonal flux ratios for NO_x/CO are shown in Fig. 6. In contrast to Fig. 5, there is no immediately clear seasonal variability in $F_{\text{NO}_x}/F_{\text{CO}}$ and the median winter and summer fluxes vary by 0.014 mol mol⁻¹ (0.485 mol mol⁻¹ in the winter, 0.471 mol mol⁻¹ in summer). While the mean summer ratio is double the mean winter ratio (0.852 against 0.425 mol mol⁻¹), this may be attributed to outliers when absolute flux magnitudes of both species become very small. A Welch's *t*-test indicates no statistically significant difference in means between the seasons ($p = 0.49$). The small absolute flux magnitudes contribute to considerable variability through much of the morning hours though in winter the median ratios from 0 to 3 UTC and 19 to 22 UTC vary by < 0.05 mol mol⁻¹. From 3 to 9 UTC and again 13 to 14 UTC the median ratios in both seasons exceed the upper expected ratio for road transport. These measured ratios overwhelmingly come from the NE wind sector where the median winter ratio is 1.4 to 1.9x greater than from the other directions (median 0.73 mol mol⁻¹ vs. 0.46 (SE), 0.39 (SW), and 0.51 (NW) mol mol⁻¹). This enhancement from the NE is also seen in summer where it is 1.2 to 1.6x greater than other directions and is driven by a similar pattern in enhanced NO_x fluxes (rather than lower CO fluxes). This may indicate the influence of a source sector with much higher NO_x to CO emission ratios to the NE. Emission ratios from the NW are also higher during rush hours than the highest expected road transport ratio and the expected ratios shown in Table 4 only characterise the SW and SE wind directions well.

Despite this directional variability, measured $F_{\text{NO}_x}/F_{\text{CO}}$ ratios broadly show higher measured ratios around 6-7 UTC and 13-14 UTC and lower ratios through the evening and night, though this pattern is muted by strong dependence on wind direction. The lower ratios during afternoon rush hours compared to morning rush hours may be due to higher CO emissions from vehicle cold-starts which can lead to larger CO emissions from similar traffic counts (Jayaratne et al., 2021). The lack of significant seasonal variability in the $F_{\text{NO}_x}/F_{\text{CO}}$ ratio suggests that the daytime seasonal variability in $F_{\text{NO}_x}/F_{\text{CO}_2}$ and $F_{\text{CO}}/F_{\text{CO}_2}$ is due to decrease in F_{CO_2} .

In Fig. 7 the flux ratios are directly compared. In Fig. 7-A a 2d histogram of measured $F_{\text{NO}_x}/F_{\text{CO}_2}$ and $F_{\text{CO}}/F_{\text{CO}_2}$ for the full campaign and the area bounded by the area-averaged inventory emission ratios is shown. While the general shape of the 2d histogram of the measured fluxes is described by the inventory ratios, a large number of observations fall outside this boundary. Notably, while the peak of the distribution falls within the expected range for stationary combustion for $F_{\text{CO}}/F_{\text{CO}_2}$, the $F_{\text{NO}_x}/F_{\text{CO}_2}$ peak falls outside the expected range.

Histograms of seasonal fluxes are shown in Fig. 7-B along with the range expected from the emissions inventory. Here again the seasonal variability in $F_{\text{NO}_x}/F_{\text{CO}_2}$ and $F_{\text{NO}_x}/F_{\text{CO}}$ is shown and the lack of significant seasonal variability in $F_{\text{NO}_x}/F_{\text{CO}}$ is made clear. The dotted horizontal lines indicating the inventory ranges show that for each species ratio the majority of fluxes are well captured by the inventory and the sectors of road transport and stationary combustion in both summer and winter.

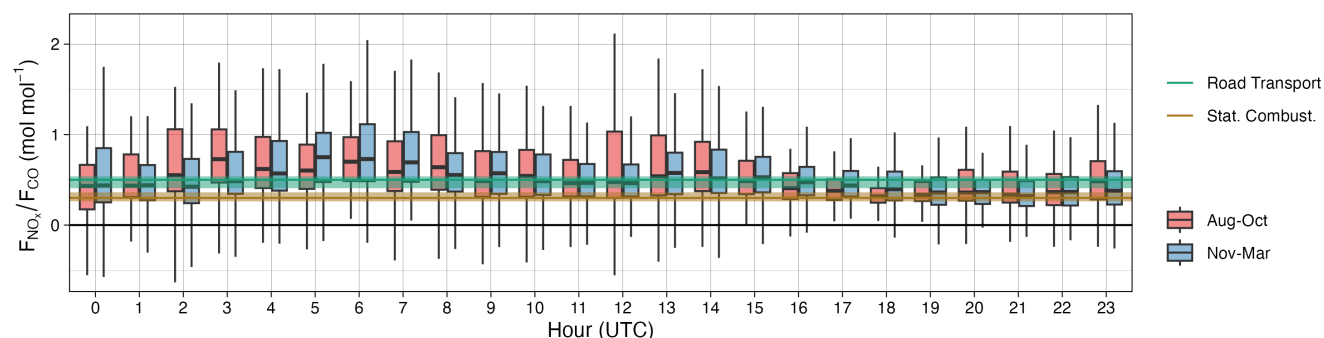


Figure 6. Diurnal and seasonal ratios of F_{NO_x} to F_{CO} . Inventory emission ratios from Table 4 are shaded.

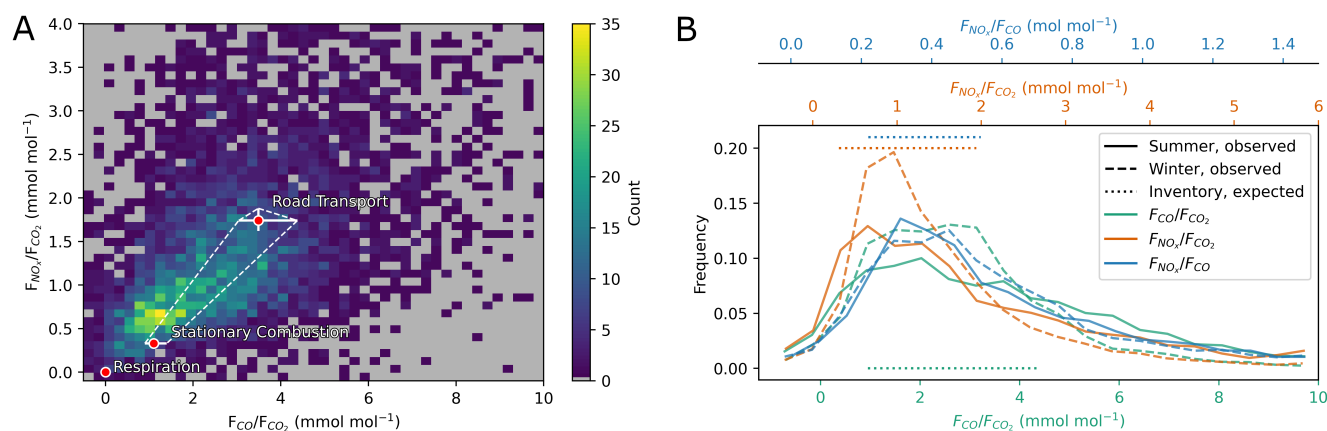


Figure 7. A - Heatmap of measured F_{NO_x}/F_{CO_2} and F_{CO}/F_{CO_2} ratios. Emission ratios from Table 4 are labelled and the area between is bounded by the dashed lines. B - Histograms of measured F_{NO_x}/F_{CO_2} , F_{CO}/F_{CO_2} , and F_{NO_x}/F_{CO} split by season. Inventory ratio ranges are shown as horizontal dotted lines.

365 In the winter, however, the peaks of the distributions for F_{NO_x}/F_{CO_2} and F_{CO}/F_{CO_2} are significantly higher, indicating that in the winter these sectors are better able to characterise the measured flux ratios. In the summer, the distribution peaks decrease and a greater portion of measured fluxes occur outside the inventory bounds. That this difference is not repeated in the NO_x to CO ratios is a strong indication that biospheric activity is being measured.

A non-negligible amount of winter observations occur outside the range predicted by the inventory. These may be attributed to a few reasons: Firstly, the selected inventory ratio ranges are comprised of a small number of spatially-averaged values with the intention of determining a characteristic ratio range for each source category. While these are weighted to the total mass of yearly emissions, variations in individual half-hour flux footprints may still sample areas outside these ratios. When the full range of ratios for each grid cell surrounding the tower is considered, the range of plausible values is greatly expanded: For CO/CO_2 the range is 0.13 to 89.17 $mmol\ mol^{-1}$. For NO_x/CO_2 0.19 to 3.78 $mmol\ mol^{-1}$. For NO_x/CO 0.016 to 4.58



375 mol mol⁻¹. While the extremes of the ranges represent cells with small contributions to total emissions, values outside the spatially-averaged mean may still be plausible depending on the individual 30-minute footprint. The overlap of the chosen inventory ratios with the histogram modes of Fig. 7-B indicates the values from Table 4 are acceptable characteristic ranges for the measured ratios despite not describing all measurements. Nevertheless, from Fig. 7-A it seems that nocturnal F_{NO_x} measurements especially are higher than expected.

380 Secondly, depending on the footprint, the contribution from sectors other than road transport and stationary combustion may impact measured ratios. While these two sectors comprise 91% of CO₂ from combustion in the 4 x 4 km⁻² area surrounding the tower according to the emission inventory, they comprise 81% of CO and 84% of NO_x. Some sources that may also contribute to the measured flux fall within the same ratio range. E.g., 7% of SW F_{CO} may be expected from industry where the expected CO/CO₂ ratio of 2.48 mmol mol⁻¹ falls between the expected stationary combustion and road transport values.
385 In contrast, 13% of F_{NO_x} from the NE is expected from public power, with a characteristic NO_x/CO ratio of 3.82 mol mol⁻¹, well outside the NO_x/CO ratios of Table 4.

Finally, in the case of $F_{\text{NO}_x}/F_{\text{CO}_2}$ and $F_{\text{CO}}/F_{\text{CO}_2}$ there may be considerable photosynthetic CO₂ uptake even through the winter (Wu et al., 2022) which produces higher than expected ratios during day, or enhanced CO₂ fluxes caused by respiration which may produce lower than expected ratios during night.

390 3.3 Flux Partitioning and Sectoral Attribution

The linear mixing model described in Section 2.5 was applied and the decomposed fluxes of CO, NO_x, and CO₂ are shown in Figure 8. Model input parameters for a and b were chosen from the full 4 x 4 km² footprint area of Table 4 and were kept constant across times and seasons.

Figure 8-A shows the source sector decomposition of F_{CO} . In summer and winter the majority of the total flux (57 and
395 59% respectively) is attributed to road transport, with the stationary combustion portion comprising 43% of the median flux in summer and 41% of the median flux in winter. In both seasons a bimodal distribution in $F_{\text{CO},\text{rt}}$ captures local rush hours. In the summer, the morning hours 0-6 UTC are attributed entirely to road transport, while in winter a higher median observed flux from 0-1 UTC is partitioned towards a small positive contribution by stationary combustion ($F_{\text{CO},\text{sc}}$). During the peak of the morning rush hour in both seasons the median modelled $F_{\text{CO},\text{rt}}$ exceeds the measured $F_{\text{CO},\text{tot}}$ (by 1.24 times at 6 UTC in
400 summer and 1.12 times at 7 UTC in winter), resulting in unrealistic negative $F_{\text{CO},\text{sc}}$ at these times as $F_{\text{CO},\text{sc}}$ and $F_{\text{CO},\text{rt}}$ must sum to $F_{\text{CO},\text{tot}}$ for each half-hour period. Through the course of the day, $F_{\text{CO},\text{sc}}$ generally increases and in both seasons reaches a daily maximum after the afternoon road transport mode. In the summer, $F_{\text{CO},\text{tot}}$ remains partitioned approximately equally between $F_{\text{CO},\text{rt}}$ and $F_{\text{CO},\text{sc}}$ until 23 UTC, while in winter the higher $F_{\text{CO},\text{tot}}$ in the evening hours remains largely attributed to $F_{\text{CO},\text{sc}}$ as $F_{\text{CO},\text{rt}}$ declines following the afternoon rush-hour maximum. The highest proportion of $F_{\text{CO},\text{sc}}$ is found in the
405 evening hours in the winter (19-23 UTC) when it comprises 50-70% of $F_{\text{CO},\text{tot}}$.

Similar patterns are observed for the sector decomposition of F_{NO_x} (Fig. 8-B). The $F_{\text{NO}_x,\text{tot}}$ is overwhelmingly (91%) partitioned to $F_{\text{NO}_x,\text{rt}}$, with $F_{\text{NO}_x,\text{sc}}$ comprising only 9% of the total flux in both seasons. The median hourly $F_{\text{NO}_x,\text{sc}}$ never exceeds the median hourly $F_{\text{NO}_x,\text{rt}}$, reaching a maximum of 40% to 60% of $F_{\text{NO}_x,\text{rt}}$ in the evening winter hours.

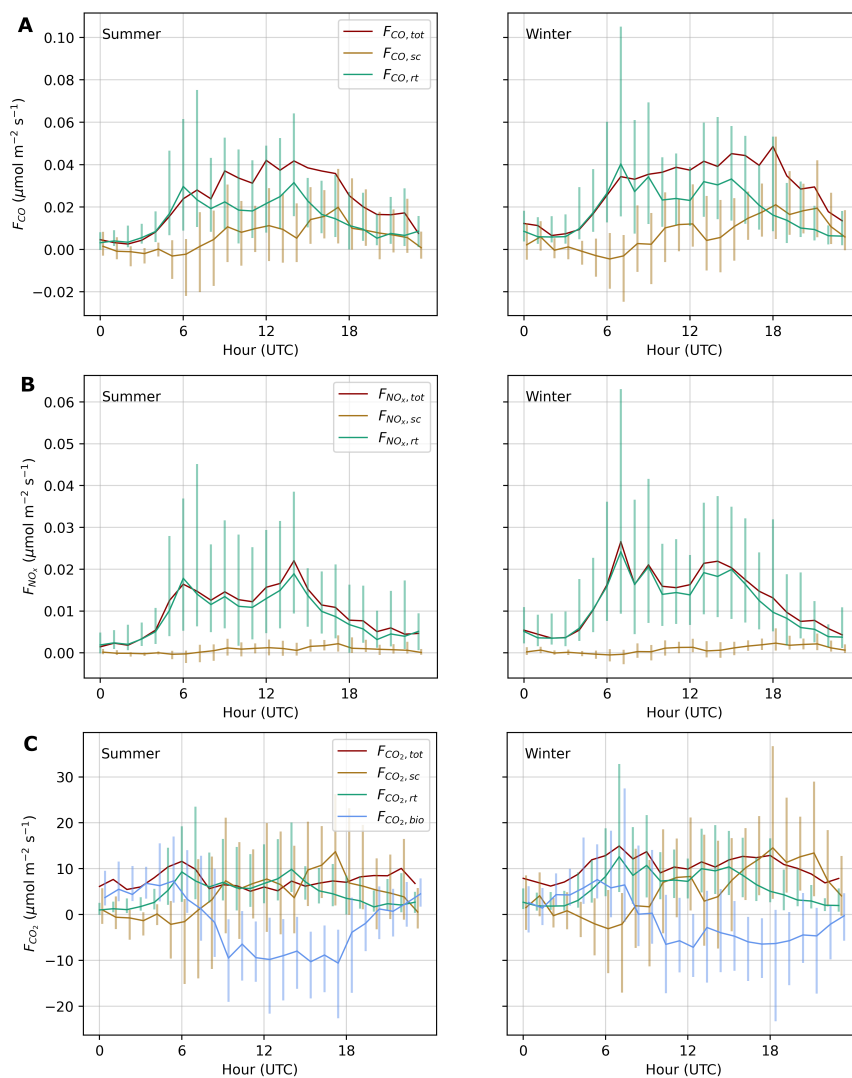


Figure 8. Modelled attribution of measured fluxes to source sector via the linear mixing model described in Section 2.5. A – F_{CO} , B – F_{NO_x} , C – F_{CO_2} . Median hourly course in UTC is shown, and the vertical bars indicate the inter-quartile range (P₂₅ to P₇₅) of the hourly bins. Differences in the diurnal course of measured fluxes from Fig. 3 may be attributed to the exclusion of the NE wind directions.



In Fig. 8-C the attribution of $F_{\text{CO}_2,\text{tot}}$ to $F_{\text{CO}_2,\text{sc}}$, $F_{\text{CO}_2,\text{rt}}$, and the biospheric component $F_{\text{CO}_2,\text{bio}}$ is shown. The compar-
 410 tively large nocturnal $F_{\text{CO}_2,\text{tot}}$ in the early morning hours is attributed largely to biospheric activity in both summer and winter
 with a larger summertime $F_{\text{CO}_2,\text{bio}}$ ($5.5 \mu\text{mol m}^{-2} \text{s}^{-1}$ average from 0 to 6 UTC) than winter ($4.2 \mu\text{mol m}^{-2} \text{s}^{-1}$), both in
 terms of absolute flux and as compared to $F_{\text{CO}_2,\text{rt}}$ and $F_{\text{CO}_2,\text{sc}}$. In both seasons there is a large daytime negative $F_{\text{CO}_2,\text{bio}}$,
 starting from 8 UTC in summer and from 10 UTC in winter. In the summer, median hourly $F_{\text{CO}_2,\text{bio}}$ remains negative and
 between -7 to $-10.5 \mu\text{mol m}^{-2} \text{s}^{-1}$ until 20 UTC when it again becomes positive and ranges from 1.3 to $6.9 \mu\text{mol m}^{-2} \text{s}^{-1}$
 415 between 20 and 7 UTC. In the winter, the daytime negative flux is smaller: between -0.2 and $-7 \mu\text{mol m}^{-2} \text{s}^{-1}$. However,
 it remains negative through 22 UTC. Modelled CO_2 exchange of parks within the tower footprint shows daytime magnitudes
 around $-10 \mu\text{mol m}^{-2} \text{s}^{-1}$ during the summer months but little exchange in the winter ($0 \pm 1 \mu\text{mol m}^{-2} \text{s}^{-1}$) (Stagakis et al.,
 2024). The negative nocturnal $F_{\text{CO}_2,\text{bio}}$ is unrealistic, and stems from the unrealistically large $F_{\text{CO}_2,\text{sc}}$ partitioning, which from
 17 to 22 is 100% to 155% of the measured $F_{\text{CO}_2,\text{tot}}$.

420 While certain reasonable patterns may be discerned from these results, it is also clear that the model is sensitive to the
 reference inventory ratios and can produce improbable and impossible results. The ratios chosen here produce a partitioning
 that significantly underestimates the contribution from stationary combustion in winter. From the inventory, annual mean
 $F_{\text{CO},\text{rt}}$ to $F_{\text{CO},\text{sc}}$ should be nearly 1:1 (38% vs. 42% of total) rather than the approximately 2:1 shown here (66% vs. 30% of
 measured CO). Similarly, $F_{\text{NO}_x,\text{rt}}$ to $F_{\text{NO}_x,\text{sc}}$ is expected to be approximately 2:1 (54% vs. 30% of total) rather than the 10:1
 425 (89% vs. 9% of measured NO_x) model partitioning. $F_{\text{CO}_2,\text{sc}}$ is expected to be 3 times $F_{\text{CO}_2,\text{rt}}$ (67% vs. 23% of total) but in the
 model they are partitioned roughly 1:1 (54% vs. 50% of measured CO_2). The mismatch in the observed stationary combustion
 ratios shown in Fig. 7-A is a likely source of poor partitioning of F_{NO_x} into stationary combustion. There are several implicit
 and explicit assumptions made in this model, violations of which could produce erroneous attribution:

A1. The emissions inventory is correct,

430 A2. The combustion emission ratios are unchanging in time,

A3. The combustion emission ratios are spatially homogeneous,

A4. The total flux of NO_x and the total flux of CO contain only non-biospheric components, and

A5. Measured combustion fluxes are solely due to either road transport, stationary combustion, or a mixture of the two.

A1 means that we are assuming no errors with regards to spatial upscaling or distribution of aggregated emissions within the
 435 city as well as perfect characterisation of the vehicle fleet and stationary combustion processes. A2 assumes that the emission
 ratio for a given category, e.g., road transport, is static in time and changes neither with fleet composition nor time of day nor
 time of year nor with changing environmental conditions. The fleet composition may be assumed to be unchanging on average
 for the short duration covered by this campaign, though trends in fleet composition in Zurich are resulting in a long-term
 decrease in NO_x from road transport. While the fleet composition on average may be unchanging, the different rhythms of
 440 traffic types (e.g., light-duty delivery vehicles vs. personal commuter cars) may lead to variation in the actual emission ratios



in any half-hour. Further, the influence of engine cold starts may produce diurnal variation in emission ratios. Jayaratne et al. (2021) showed that similar traffic counts during morning and afternoon rush hours in a city produce larger CO concentrations in the afternoon due to high CO emissions during cold starts which are concentrated within the urban core. Further, cold start enhancement of CO₂, CO, and NO_x exhibits a non-linear dependence on ambient temperature (Weilenmann et al., 2009; 445 Bielaczyc et al., 2011). A3 assumes that the emission ratio of a given category does not change with direction/source area. As discussed in Sec. 2.4 and 3.1 there is spatial variation in the per-grid cell inventory ratios surrounding the tower, and spatial heterogeneity is a perennial characteristic of urban emissions. Together A2 and A3 allow each emission category and species ratio to be characterised by a single value for the extent of the campaign. A4 is unlikely to be significantly violated, yet there could be biogenic production of CO due to VOCs (Griffin et al., 2007). A5 is untrue for the city as a whole, but a reasonable 450 simplification for the 4 x 4 km² area surrounding the tower. Though as discussed, the exceptions to A5 from certain sectors and directions could drive ratios well outside of the predicted range: enhanced F_{NO_x} from the NE especially contributes to a greater attribution of the flux to road transport. Finally, it must be noted that the linear correction applied to the NO_x measurements may cause an overestimation F_{NO_x} in some or all periods. This could lead to an over-attribution of net fluxes to road transport if the lower $F_{\text{NO}_x}/F_{\text{CO}_2}$ ratio of stationary combustion is not met, even if all other assumptions are met. An excessive correction 455 would explain this over-attribution as well as the apparent elevated measurements of Fig. 7.

To test the sensitivity of the linear mixing model, the calculations were re-run to consider all combinations of a and b from Table 4 at 0.1 mmol mol⁻¹ resolution and the partitioning of median fluxes considered for each cardinal direction separately. The results are summarised in Table 6. There are distinct directional differences regarding the partitioning of observed fluxes into source categories. While the range of model outputs is reasonable to the SW, there is no combination of inputs from Table 460 4 that produce reasonable partitioning of fluxes from the NE. The NW is also poorly partitioned while the SE produces a more reasonable partitioning.

It is not immediately clear from the inventory why this is so as the SW is similar to the other directions, with road transport and stationary combustion accounting for 80% of CO, 79% of NO_x, and 92% of combusive CO₂, with minority contributions from industry, offroad, and waste categories. To the NE the previously discussed contribution from public power may be 465 causing both the higher observed fluxes from this direction, and subsequently poorer partitioning of said fluxes.

Although the linear mixing model produces expected results in many situations, exceptions to A2 and A5 in the complex and heterogenous urban environment may ultimately pose too great a challenge for the application of such a model with fixed emission factors over long periods of time and large flux footprints. In future work, there should be a determination of where and when the model fails using footprint modelling of individual flux averaging periods as well as a spatially and temporally 470 resolved high-resolution emission inventory.

4 Conclusions

This work showcased eight months (August 2022 to March 2023) of continuous urban tall-tower eddy covariance measurements of CO₂ and four co-emitted species: CO, NO_x, CH₄, and N₂O. To our knowledge it is the first work to demonstrate



Table 6. Partitioning of fluxes compared to inventory for all combinations of input reference ratios. Inventory partitioning is given as a percentage of the total combustive flux of a species (% of total), i.e., considering all contributing source categories, as well as relative to each other (i.e., assuming the full flux may be attributed to one of the two categories). Model outputs are given as P_{25} - P_{75} (P_{50}) using the distribution of all model outputs.

Species	Inventory Reference		Modelled Partitioning				
	% of total	relative	All	SE	SW	NW	NE
CO _{sc}	38%	48%	-58 - 9 (-16%)	-23 - 26 (6%)	27 - 64 (45%)	-130 - -25 (-66%)	-286 - -95 (-163%)
CO _{rt}	42%	52%	78 - 146 (105%)	55 - 106 (76%)	25 - 62 (43%)	118 - 227 (159%)	192 - 388 (264%)
NO _{x,sc}	30%	36%	-35 - 5 (-9%)	-15 - 17 (4%)	19 - 55 (34%)	-69 - -12 (-31%)	-125 - -40 (-70%)
NO _{x,rt}	54%	64%	83 - 126 (99%)	67 - 101 (81%)	34 - 71 (55%)	104 - 162 (122%)	135 - 221 (164%)
CO _{2,sc}	67%	74%	-157 - 23 (-44%)	-45 - 51 (12%)	70 - 184 (122%)	-300 - -57 (-148%)	-790 - -250 (430%)
CO _{2,rt}	23%	26%	70 - 109 (85%)	37 - 60 (47%)	23 - 49 (37%)	89 - 144 (109%)	173 - 291 (215%)
CO _{2,bio}	-	-	30 - -122 (-70%)	-6 - 95 (35%)	-45 - -21 (-37%)	-84 - -206 (-137%)	-217 - -434 (-307%)

simultaneous flux measurements of these species over an urban area. While the EC flux footprint does not cover the complete administrative boundary of the city of Zurich, it is clear from these measurements that the city is a net source of CO₂, CO, NO_x, CH₄, and N₂O.

Considering the 100-year global warming potential of non-CO₂ greenhouse gases, the overall CO₂-equivalent emissions within the tower footprint consist of 95.8% CO₂, 1.8% CH₄, and 2.4% N₂O in summer, and 97.2% CO₂, 1.4% CH₄, and 1.4% N₂O in winter. Median wintertime enhancement of F_{CO_2} was largest at 1.47x (10.9 to 7.4 $\mu\text{mol m}^{-2} \text{s}^{-1}$) while seasonal enhancement of other species was smaller: 1.08x for F_{CO} ; 1.08x for F_{NO_x} ; 1.01x for F_{CH_4} ; and 0.95x for $F_{\text{N}_2\text{O}}$. Correlation between F_{CO_2} was highest with F_{CO} in the summer and with F_{CO} and F_{NO_x} in the winter.

Observed flux ratios of the three most-emitted species (CO₂, CO, and NO_x) were calculated and compared to characteristic spatially-averaged values from a city emission inventory within a 4 x 4 km² area centred on the measurement tower. Ratios of $F_{\text{NO}_x}/F_{\text{CO}_2}$ and $F_{\text{CO}}/F_{\text{CO}_2}$ in winter were generally well characterised by the stationary combustion and road transport categories of the emission inventory. Significant and systematic exceedance of the inventory ratios was observed in the summer afternoon hours, which was attributed to decreased net F_{CO_2} due to photosynthetic uptake. Measured ratios of $F_{\text{NO}_x}/F_{\text{CO}}$ were less well characterised by the inventory, with significant exceedance of the expected ratios coming from the NE. Nevertheless road traffic peaks seem well characterised by the measured $F_{\text{NO}_x}/F_{\text{CO}}$ ratios. No statistically significant seasonal difference was observed in $F_{\text{NO}_x}/F_{\text{CO}}$ ratios.

A linear mixing model for partitioning observed F_{CO_2} , F_{CO} , and F_{NO_x} into stationary combustion, road transport, and biospheric components was proposed and tested. Using characteristic spatially-averaged reference ratios for each category from the city emissions inventory, positive $F_{\text{CO}_2,\text{bio}}$ was found during the night and negative $F_{\text{CO}_2,\text{bio}}$ during the day in the warm months. In the colder months, positive $F_{\text{CO}_2,\text{bio}}$ was found during the morning hours and a reduced, but still negative,



$F_{\text{CO}_2, \text{bio}}$ was found from mid-day through midnight. $F_{\text{CO}_2, \text{rt}}$ and $F_{\text{CO}_2, \text{sc}}$ were partitioned approximately evenly: 50% and
495 54% of $F_{\text{CO}_2, \text{tot}}$, respectively. $F_{\text{CO}, \text{rt}}$ was overestimated compared to the inventory, with 66% of $F_{\text{CO}, \text{tot}}$ attributed to $F_{\text{CO}, \text{rt}}$
and 30% to $F_{\text{CO}, \text{sc}}$. $F_{\text{NO}_x, \text{rt}}$ was highly overestimated, with 89% of $F_{\text{NO}_x, \text{tot}}$ attributed to $F_{\text{NO}_x, \text{rt}}$ and 9% attributed to
 $F_{\text{NO}_x, \text{sc}}$. Total partitioning of species was very sensitive to inventory reference inputs and did not perform equally well for all
cardinal directions.

This work demonstrates the potential to partition CO_2 fluxes into different source and sink sectors using CO and NO_x as co-
500 emitted species. Measurement of co-emitted species provides important additional information on the probable source sector
contribution to individual 30-minute fluxes, however in the complex and heterogenous urban environment this information is
difficult to exploit on its own, without the use of a spatio-temporally-resolved emission inventory and individual flux footprint
modelling. It also highlights some challenges of simultaneous measurement of fluxes of co-emitted species via tall-tower urban
eddy covariance. In the next steps of this work, these measurements will be combined with modelled 30-minute flux footprints
505 as well as a temporally resolved city emission inventory to gain a more detailed understanding of flux ratios and their variability,
as well as further validation of the inventory itself.

Data availability. The raw data and processed fluxes used in this analysis are available from the ICOS Cities carbon portal <https://citydata.icos-cp.eu/portal/>.



Author contributions. RH oversaw the CPEC measurement system, wrote the manuscript and created the figures, and performed the analysis.

510 JH oversaw the CPEC installation, initial operation, and early analysis. SS installed and oversaw the OPEC measurements. BM produced the footprint and surface-cover analysis based on the model of NK. LC and DB created the emissions inventory. AC conceived the flux partitioning model. All authors reviewed and contributed to the manuscript and contributed intellectually to the work.

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