



1 **Unveiling Sulfate Aerosol Persistence as the Dominant Control** 2 **of the Systematic Cooling Bias in CMIP6 Models: Quantification** 3 **and Corrective Strategies**

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17 **Abstract**

18 Including sophisticated aerosol schemes in the models of the sixth Coupled Model
19 Inter-comparison Project (CMIP6) has not improved historical climate simulations. In
20 particular, the models underestimate the surface air temperature anomaly (SATa) when
21 anthropogenic sulfur emissions increased in 1960~1990, making the reliability of the
22 CMIP6 projections questionable. Biases in cooling among the models are correlated
23 with sulfate burden and the deposition of sulfur is the process responsible. We show
24 that the lifetime of atmospheric sulfur, defined by a new global index for sulfur
25 deposition (an “Effective Sulfur Retention Timescale” (ESRT)), determines the cooling
26 biases. Reducing the biases to within the observational uncertainty is consistent with a
27 physically plausible ESRT of around one day, whereas the CMIP6 models overestimate
28 this timescale. Based on targeting a reduction of ESRT, post-CMIP6 improvements to
29 two models are shown to greatly improve SATa reproduction.

30



1. Introduction

Atmospheric aerosols have rapidly increased since the Industrial Revolution. Over this time period, the total aerosol effective radiative forcing (ERF) was dominated by the sulfate cooling effect, and offset a substantial portion of global-mean forcing from well-mixed greenhouse gases (IPCC, 2023). Without this historical aerosol ERF, the Paris Agreement's target of limiting global warming to 1.5°C above pre-industrial levels would have already been missed in 2015 (Hienola et al., 2018). Similarly, stopping all present-day anthropogenic aerosol emissions is estimated to induce a global-mean surface heating of 0.5~1.1°C (Samset et al., 2018). The year 2024 has been confirmed as the hottest year in human history, and was the first year to breach the 1.5°C warming limit (Bevacqua et al., 2025). Moreover, recent years have seen temperature trends accelerate, which may be due to reductions in atmospheric aerosols, especially aerosols produced by commercial shipping (Hansen et al., 2025). Hence, it has been suggested that even small emissions in relatively pristine air have substantial effects, and better constraining the ability of global-climate models to predict aerosol effects may be crucial to obtaining reliable projections.

The observed temporal evolution of historical surface air temperature (SAT) is one of the major metrics used for evaluating the performance of climate models. However, the SAT anomalies in the CMIP6 models are systematically lower than was observed for the 1960~1990 period, whereas the CMIP5 models on average track the instrumental record quite well (e.g., Flynn and Mauritsen, 2020). The 1960~1990 period is referred to as the “pothole cooling period” (PHC) in our previous study (Zhang et al., 2021a), due to the ‘pot-hole’ shaped dip in SAT at that time, and in this study hereafter. The PHC period is coincident with the so-called Great Acceleration period, in which the human enterprise was boosted remarkably and led to global-scale impacts on Earth System functioning (Steffen et al., 2007). Recent studies hypothesized that aerosol forcing in CMIP6 is stronger than in CMIP5 and is responsible for the suppressed late 20th-century warming (Dittus et al., 2020; Smith and Forster, 2021).

The anomalous cooling points towards a problem with the sulfur cycle in recent



ESMs or the emissions data (Hardacre et al., 2021; Wang et al., 2021). Considering the importance of the sulfur cycle in historical aerosol ERF, we examine the sulfur related processes in eleven CMIP6 models with aerosol schemes in this study. All the models are forced with CMIP6 historical anthropogenic aerosol emissions (Hoesly et al., 2018), and therefore differences in their sulfate burdens are mainly due to different representations of the sulfur cycle in the models.

We will identify the key processes that determine sulfate-burden in these models, and introduce a simple metric for measuring the level of activity in the sulfur cycle in the models on the global scale. This metric (an effective timescale for global cycling of atmospheric sulfate) can be easily calculated from time series of global means only, without the need for complex diagnostics of the sulfur-cycle processes. We show that our timescale is strongly correlated with sulfate burden and anomalous cooling and has a clear physical interpretation that allows each model's sulfur cycle to be calibrated using historical temperature biases.

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2. Results

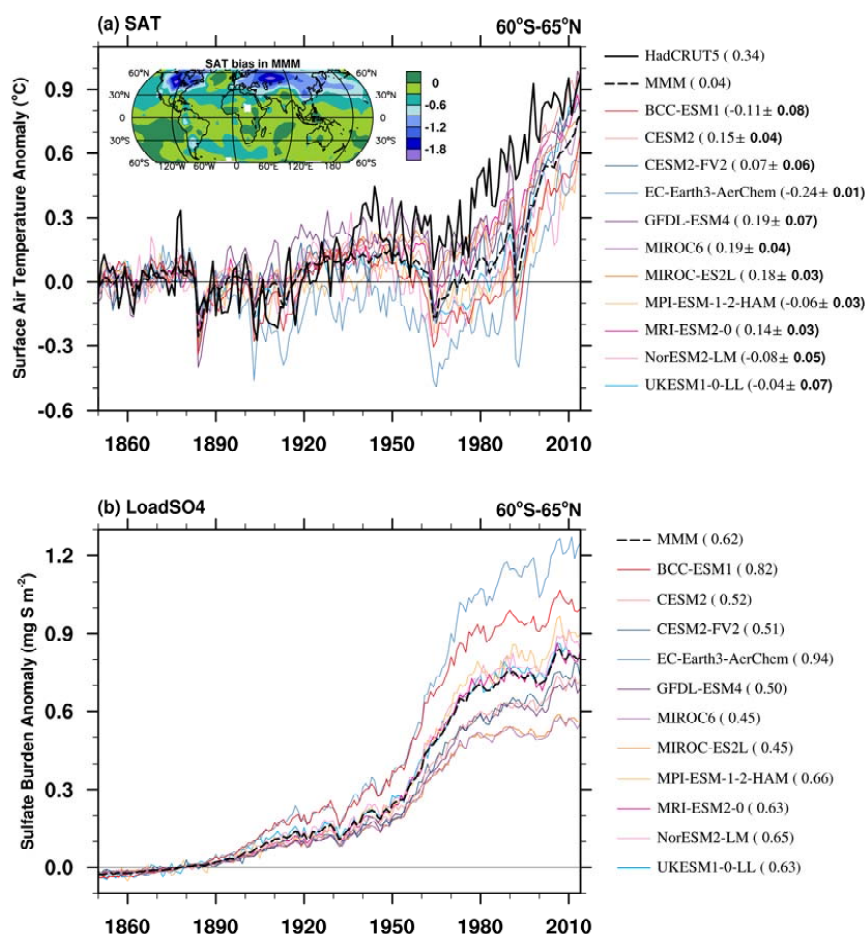
2.1 SAT and sulfate burden

The historical evolutions of near-global mean (60°S to 65°N) SATa in the eleven CMIP6 models with interactive aerosol schemes are shown in Fig. 1a. All the models tend to underestimate SATa since the 1950s. The cooling bias is most pronounced from 1960 to 1990, i.e., the PHC period. The SATa in PHC is about 0.34°C in the observations. However, the multi-model mean (MMM) SATa in the models is about 0.3°C lower with a large model spread. The SATa ranges from -0.24°C in EC-Earth3-AerChem to 0.19°C in GFDL-ESM4 and MIROC6. The cooling is noticeable at the mid to high latitude in the Northern Hemisphere (NH, as shown in the attached SATa map in Fig.1a). The sudden drop in SATa in the early 1960s and 1990s may be due to the stronger model responses to large volcanic eruptions, Mount Agung in 1963 and Mount Pinatubo in 1991, than in the observations (Chylek et al., 2020). The anomalous



88 cooling biases gets smaller later in the simulations, which is related to the generally
89 high sensitivity of the models to GHG forcing (Smith and Forster, 2021).

90



91

92 **Figure 1.** (a) Historical surface air temperature anomalies (SATa) relative to 1850–1900 mean for
93 HadCRUT5 (thick black line), the ensemble mean for each CMIP6 model (solid color lines), and
94 multi-model mean (MMM, dashed black line). The numbers in brackets are the mean results during
95 the pothole period (1960–1990) together with the inter-member spread for each model. Panel (b) is
96 the same as panel (a) but for sulfate burden anomalies for the first realization from each CMIP6
97 model (solid color lines) and MMM (dashed black line).

98

99 Global emissions of SO₂ grew steadily after the 1950s and peaked in the 1970s at



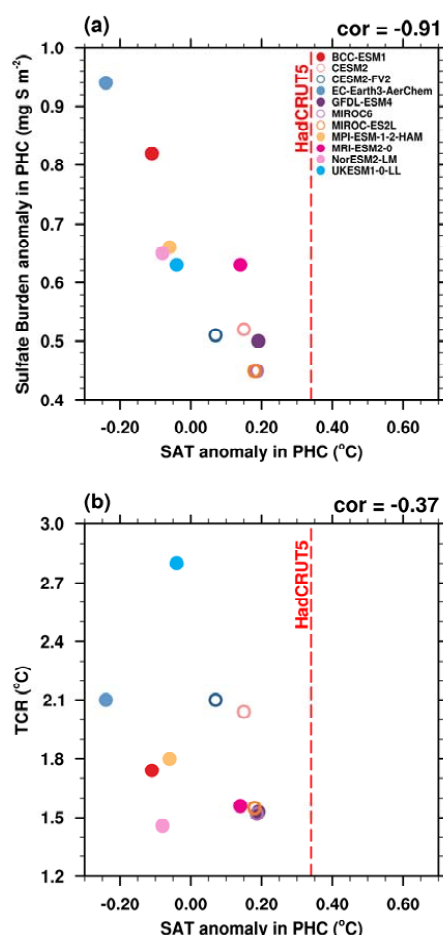
180Tg yr⁻¹, which is about 3.6 times the 1950s' emissions (Hoesly et al., 2018). As the main precursor of sulfate, the growing emission of SO₂ led to the accumulation of sulfate in the atmosphere, which interrupted a decades-long warming trend via the cooling effects of sulfate aerosols on climate, even though carbon-dioxide emissions continued to rise (Wilcox et al., 2013). Because of the emission control policies in Europe and North America (Hand et al., 2012; Vestreng et al., 2007), such as the Gothenburg Protocol (Eb, 1999) and the 1990 Clean Air Act Amendments in the U.S. (Likens et al., 2001), global anthropogenic SO₂ emissions were suppressed after the 1980s and SAT started to increase rapidly in the observations (Aas et al., 2019). However, anthropogenic SO₂ emission continued to increase over Asia due to industrial developments, although they have also decreased since 2006 in East Asia (Wang et al., 2021). Some of this decrease in SO₂ emissions at the beginning of the 21st century is not well represented in the CMIP6 emission inventory. But it is outside of the PHC period and the impact on SAT reproduction is beyond the scope of this paper.

In the 11 CMIP6 models, sulfate concentrations increased rapidly during the PHC period (Fig.1b). The intensified emission of anthropogenic SO₂ mainly comes from industries and the energy-transformation sector (e.g., Ohara et al., 2007; Vestreng et al., 2007). The SAT anomalies simulated by CMIP6 models are systematically lower than observations during the PHC period, indicating an excessively strong sulfate-induced cooling effect in CMIP6 models. The sulfate burden is the lowest in MIROC models (0.45 mg S m⁻²) with the smallest cooling bias (-0.15°C lower than HadCRUT5), and is doubled in EC-Earth3-AerChem (0.94mg S m⁻²) with the largest cooling (-0.58°C lower than HadCRUT5). Generally, the models with higher sulfate burdens anomalies also tend to underestimate SAT anomalies the most. As shown in Fig. 2a, the correlation coefficient between anomalous sulfate burden and SAT during the PHC is -0.91, significant at the 1% level using a Student's *t*-test. Interactive chemistry may have an impact on sulfate formation and affect the sulfate aerosol burdens in the atmosphere (Mulcahy et al., 2020). As shown in Fig.2a, models with interactive chemistry (color dots) seem to have higher sulfate burden anomaly and lower SATa than models without



129 (color circles). However, the relationship between sulfate burden and SATa is
130 consistent among models with and without interactive chemistry (i.e., there is no
131 obvious difference in relationship between loadSO₄ and SATa for models with and
132 without interactive chemistry).

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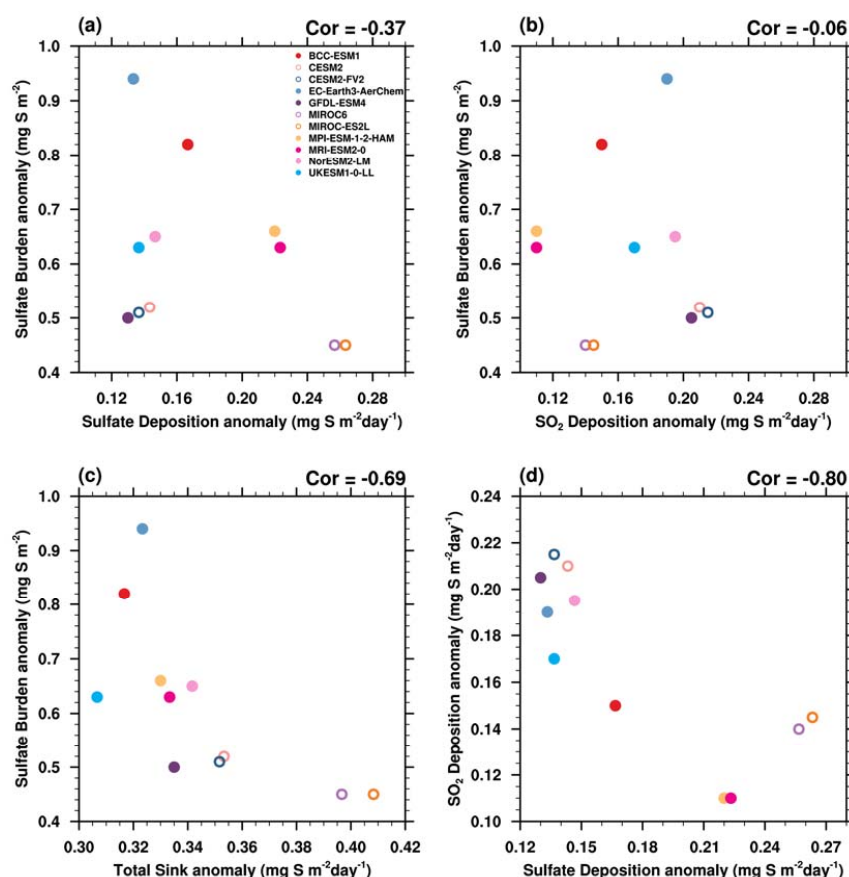
135 **Figure 2.** Scatter plots of SATa in PHC period (x-axis, °C) versus (a) sulfate burden anomaly in
136 PHC period (y-axis, mg S m⁻²) in historical experiments, and (b) the transient climate response (TCR,
137 °C) for each model calculated by 1pctCO₂ experiments. The corresponding correlation coefficient
138 (COR) is shown at the top-right corner of each panel. The anomalies are relative to 1850~1900
139 mean. Models with and without interactive chemistry are marked by color dots and color circles,
140 respectively.



Greenhouse gases (GHGs) also show a rapidly increasing trend in the PHC period. However, TCR, which can generally indicate the impact of GHGs, is insignificantly correlated with SAT anomalies in CMIP6 models and the correlation coefficient is even slightly negative (Fig.2b). Therefore, the biases of atmospheric sulfate burden and the associated sulfate aerosol cooling effects play an essential role in the anomalous-cooling biases in the CMIP6 models.

147

2.2 Sulfur deposition and a metric for the activity of the global sulfur cycle (ESRT)



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Figure 3. Sulfate burden anomaly in PHC period (mg S m^{-2} , y-axis) versus (a) sulfate deposition anomaly, (b) SO_2 deposition anomaly, and (c) total sulfur sink (sulfate and SO_2 deposition) anomaly in PHC period. (d) sulfate deposition anomaly (x-axis) versus SO_2 deposition anomaly (y-axis) in PHC period. Unit for deposition anomaly is $\text{mg S m}^{-2}\text{day}^{-1}$.



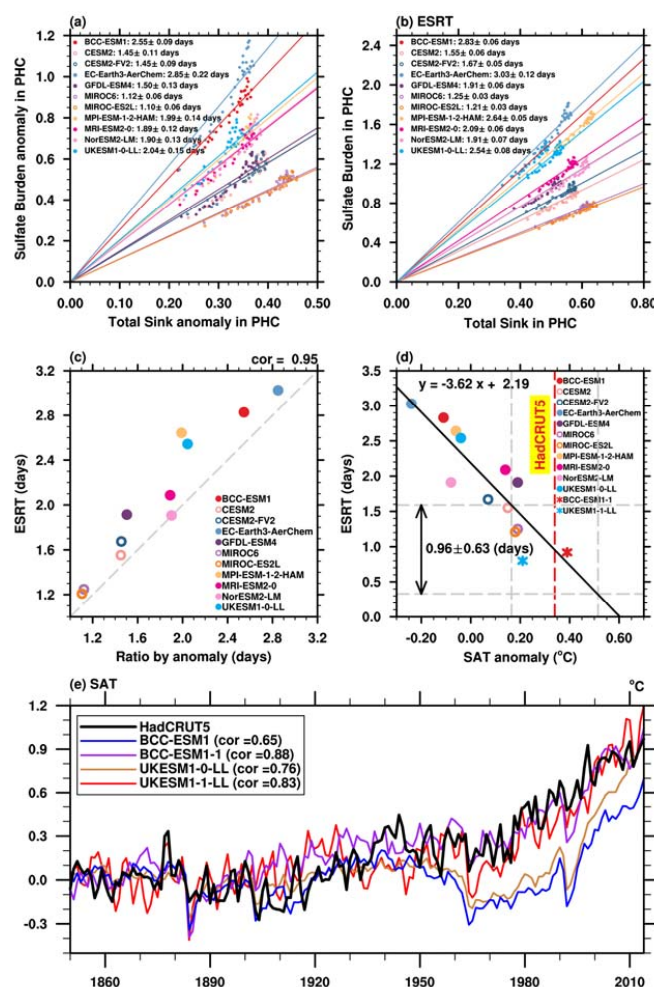
154 Fig. 3 shows comparisons of the global mean total sulfate burdens versus sinks of
155 sulfur as anomalies in the PHC period relative to a baseline period of 1850~1900. As
156 shown in Fig.3a, the sulfate burden anomaly is negatively correlated with sulfate
157 deposition anomaly but the correlation is not significant, partly because of the presence
158 of a group of 5 models which prevent a robust linear fitting of the remaining models by
159 having both slow sulfate deposition rates and low sulfate concentrations. There is no
160 clear statistical relationship between sulfate burden and SO₂ deposition (Fig. 3b).
161 However, the correlation between sulfate burden and total sulfur sink (deposition of
162 sulfate and SO₂) increases to -0.69, significant at the 5% level using a Student's t-test
163 (Fig.3c). In particular, most of the 5 models with low anomalous sulfate deposition rates
164 and low sulfate burden have higher than average SO₂ deposition, which influences their
165 total sulfur sink enough for a correlation to emerge. We can hypothesize that in these 5
166 models' oxidation of precursors to sulfate proceeds slower than in the other models.
167 This is reflected by their larger SO₂ deposition rates, and leads to less sulfate in the
168 atmosphere. That is, both the sulfate deposition and the SO₂ deposition (via its
169 relationship with oxidation rates) are responsible for the sulfate burden anomalies,
170 although the relative ratio of both deposition processes is different among the models.
171 Further examination indicates that the anomalous SO₂ deposition rate among the
172 models is highly negatively correlated with the anomalous sulfate deposition rate with
173 correlation coefficient of -0.80 (Fig.3d). The total sulfur sink is examined and discussed
174 hereafter.

175 Considering the importance of anomalous total sulfur sink to sulfate burden,
176 Fig.4a examines their relationship during the PHC period in each model. Generally, the
177 anomalous sulfate burden and total sink are positively correlated and co-vary almost
178 linearly in all the models. We further examine the ratio between sulfate burden and total
179 sink instead of their anomalies in Fig.4b, i.e., Effective Sulfur Retention Timescale
180 (ESRT), defined by Eq. (1) in Appendix A1. The ESRT is also approximately constant
181 during the PHC. Comparisons indicate that the ratios between anomalous sulfate burden
182 and anomalous total sink are highly correlated with ESRT (Fig.4c). However, the ESRT



183 is independent of preindustrial mean state and will be practically more useful. The
184 ESRT is generally longer in models with interactive chemistry (color dot in Fig.4c) than
185 without (color circle), which is consistent with the results demonstrated by two UK
186 models with and without interactive chemistry (Mulcahy et al., 2020). In the two UK
187 models sharing the same aerosol scheme, sulfate lifetime is 5.57 days in UKESM1-0-
188 LL (which has interactive chemistry) but shorter in HadGEM3-GC3.1 (which uses
189 prescribed concentrations for oxidants).

190 2.3 The recommended ESRT value



191

192 **Figure 4.** (a) Scatter plots of yearly total sulfur sink (x-axis) versus sulfate burden anomaly (y-axis)



193 in PHC period in relative to 1850~1900 mean. Number in legend shows the mean and standard
194 deviation of ratio between sulfate burden anomaly and total sulfur sink anomaly in PHC period,
195 units: days. (b) is the same as (a), but for sulfate burden and total sulfur sink. Number in legend
196 shows the ratio defined as ESRT, units: days. (c) the ratio calculated in (a) versus ESRT in (b). (d)
197 The SATa ($^{\circ}\text{C}$, x-axis) versus ESRT (days, y-axis) in PHC period for each model. The black solid
198 line is the linear fitting. SAT anomaly in HadCRUT5 and its 0.175°C boundaries are shown by the
199 red solid line and parallel gray dashed lines, respectively. The constraint range for ESRT based on
200 HadCRUT5 is 0.96 ± 0.63 days. The red and blue asterisks are the results in the two post-CMIP6
201 models BCC-ESM1-1 and UKESM1-1-LL, respectively. (e) Evolutions of SAT anomalies relative
202 to 1850~1900 mean for HadCRUT5, BCC-ESM models, and UKESM models. The numbers in
203 legend are the corresponding correlation coefficients with HadCRUT5.

204

205 The ESRT varies from 1.21 days in MIROC6 to 3.03 days in EC-Earth3-AerChem
206 (Fig. 4b). The standard deviation of ESRT for each model ranges from 0.03 to 0.12
207 days, about 3.0% of the ESRT. That is, although the sulfate burden increased
208 significantly in the PHC period, the ESRT hardly changed. This is an important sign
209 that ESRT is a robust index for evaluating the sulfur cycle in model development. Our
210 finding that a single value of ESRT is capable of characterizing the anomalous cooling
211 for each model, makes it a convenient target for model tuning. Focusing on a single,
212 representative parameter can make tuning more efficient and help to reduce the
213 computation cost, especially when the model resolutions become relatively high.
214 Moreover, because ESRT has a clear physical interpretation as a globally defined
215 efficiency factor for sulfur removal processes, tuning based on ESRT can give
216 confidence that SAT biases are reduced for a ‘right’ (i.e., physically sensible) reason.

217 Tuning based on ESRT requires an empirical best-estimate ESRT value to aim for.
218 Therefore, a further question is how to estimate the reasonable values for ESRT. Here
219 we try to constrain the ESRT using the SATa in observations. Fig. 4d shows the ESRT
220 and SATa in PHC in each model. The SATa is highly correlated with ESRT with a
221 correlation coefficient of -0.85. The SAT anomaly in PHC is 0.34°C , shown by the
222 vertical red dash line (HadCRUT5). Considering the internal variability in the climate
223 system and the uncertainty in observation, the observed uncertainty is suggested to be
224 0.175°C (the vertical gray dash line parallel with the red dash line). The observed



uncertainty is estimated as the standard deviation of observed annual mean globally averaged SAT in HadCRUT5 from 1850 to 2014 after removing the least squares linear trend. After the constraint, the ESRT is suggested to be 0.33 to 1.59 days. As shown in the figure, most of the models have longer ESRT than expected. The ESRT in CESM2, MIROC6 and MIROC-ES2L, ranging from 1.21 to 1.55 days, is at the higher end of the suggested ESRT values, associated with their relatively low SATa. That is, the SO₂ oxidation or deposition terms in CMIP6 models may need to be modified.

Here we use this metric to modify the sulfur cycle in BCC-ESM1, more specifically we quadruple the SO₂ dry deposition over land and multiply the SO₂ dry deposition over the ocean by 1.5. This effect is similar to that in UKESM1-0-LL by modifying SO₂ dry deposition parameterization (Hardacre et al., 2021; Mulcahy et al., 2023). The impact of changes to the SO₂ dry deposition parameterization in UKESM1-0-LL is an increase of SO₂ dry deposition by a factor of 2 to 4. As shown by the red asterisk in Fig.4d, the ESRT reduced from 2.83 days to 0.92 days in updated BCC-ESM1 (BCC-ESM1-1), falling right within the recommended estimations. The new ESRT is 36% of its previous values. Accordingly, the SATa in PHC falls within the reasonable range, rising from -0.11°C to 0.39°C. We also examine the ESRT in UKESM1-1-LL with modified SO₂ dry deposition parameterization. The ESRT is shortened from 2.54 days to 0.80 days, also falling within the recommended ESRT range. Accordingly, the SATa in PHC period increases by about 0.25°C. As shown by the global mean SATa in BCC-ESM and UKESM models (Fig.4e), the BCC-ESM1-1 and UKESM1-1-LL on average tracked the instrumental record quite well with higher correlation coefficients with observation. That is, in both models the improvements to the sulfur sink processes has reduced the ESRT and improved the representation of historical surface temperature observations.

250

251 3. Conclusions and discussion

Aerosol cooling effect is considered to be the second most important anthropogenic forcing over the 20th Century. Our study, based on the 11 CMIP6 models with aerosol



254 schemes, demonstrates that the anomalous cooling bias in the PHC period is closely
255 related to the sulfate burden changes in the atmosphere. Sulfate burden in the models,
256 and hence the strength of the anomalous cooling, is determined by sulfur deposition.
257 We introduce a metric, called the ESRT index, which incorporates the effects of sulfur
258 removal processes on sulfate concentration. The index is highly correlated with cooling,
259 and can be used to constrain sulfur removal processes in models, on a global scale. In
260 particular, we show that most models overestimate ESRT, compared to residence time-
261 scale that would be consistent with near-zero temperature biases during the PHC period.
262 A constraint on ESRT, derived using observed SATa, is used to inform a choice of
263 tunable parameters for deposition in BCC-ESM1 and UKESM1-0-LL. Modifying
264 sulfur deposition properties leads to an improved ESRT in BCC-ESM1-1 and
265 UKESM1-1-LL, as well as SATa reproductions.

266 Given that CMIP6 models overestimate the cooling effects of sulfate during the
267 PHC period, when emissions were rising, it is reasonable to assume that they will
268 underestimate the rate of warming during periods when sulfur emissions are falling.
269 This has potential implications for the use of CMIP6 in scenarios that incorporate clean-
270 air measures to inform the Paris Agreement goals of limiting warming to below 2 or
271 1.5°C, i.e., SSP1-2.6 and SSP1-1.9 in CMIP6 (O'Neill et al., 2016). To improve the
272 reliability of projections for such scenarios, sulfur cycle processes in models should be
273 improved. The ESRT metric introduced in this paper provides a physically meaningful
274 measure of the activity of the sulfur cycle, at a global scale, which we have shown can
275 be used to improve modeled sulfur processes.

276

277 **Appendix A: Methods and Data**

278 **A1 The Effective Sulfur Retention Timescale (ESRT) index**

279 Atmospheric sulfate concentrations are determined by the emission and oxidation
280 of sulfate precursors, as well as deposition processes. Together these processes make
281 up the atmospheric part of the Earth's sulfur cycle. Anthropogenic SO₂ emissions are



the major source of sulfate aerosol over land in polluted regions. Given that the same anthropogenic SO₂ emissions are used in all the CMIP6 models, most of the differences in simulated atmospheric sulfate concentrations occur due to oxidation of SO₂ and sulfate deposition processes. Since much of the loss of SO₂ occurs locally to its emission source by oxidation and deposition, faster SO₂ deposition is associated with weaker SO₂ oxidation. Hence, for sulfate itself, the faster the sulfate deposition rate, the less the sulfate will reside in the atmosphere. That is, both the SO₂ deposition and sulfate deposition are important for the sulfate concentrations in the atmosphere, directly or indirectly.

Here we define the Effective Sulfur Retention Timescale (ESRT) index. This is calculated as the ratio between sulfate burden and sulfur deposition. The sulfur deposition combines the deposition rate of sulfate and its major precursor sulfur-dioxide gas (SO₂):

$$\text{ESRT} = \text{loadSO}_4 / (\text{Dso}_4 + \text{Dso}_2), \quad (1),$$

where loadSO₄ is total sulfate loading (i.e., the sulfate burden) in the atmosphere, Dso₄ and Dso₂ denote the total deposition (wet deposition and dry deposition) of sulfate and SO₂, respectively.

A2 The transient Climate Response (TCR) index

The transient Climate Response (TCR) index is calculated as the mean SAT anomaly of a 1pctCO₂ simulation in a 20-year period centered on year-number 70, by which a doubling CO₂ concentration has occurred. It is an important metric representing CO₂-related historical warming and has been widely used for model evaluations and comparisons (Bevacqua et al., 2025; O'Neill et al., 2016).

A3 CMIP6 models datasets.

Table A1. Information of the eleven CMIP6 models with aerosol schemes.



Model	Country	Interactive Chemistry	Members	Reference
BCC-ESM1	China	Yes	3	Wu et al., 2020; Zhang et al., 2021b
CESM2	US	No	11	Danabasoglu et al. (2020)
CESM2-FV2	US	No	3	Danabasoglu et al. (2020)
EC-Earth3-AerChem	European consortium	Yes	2	Döscher et al. (2021)
GFDL-ESM4	US	Yes	3	Dunne et al. (2020)
MIROC6	Japan	No	50	Tatebe et al. (2019)
MIROC-ES2L	Japan	No	30	Hajima et al. (2020)
MPI-ESM-1-2-HAM	Germany	Yes	3	Mauritsen et al. (2019)
MRI-ESM2-0	Japan	Yes	10	Yukimoto et al. (2019)
NorESM2-LM	Norway	Yes	3	Seland et al. (2020)
UKESM1-0-LL	UK	Yes	19	Sellar et al. (2019)

309

310 Eleven CMIP6 climate models with interactive aerosol schemes are utilized in this
311 study, including seven models with interactive chemistry and four without (Table A1).
312 The outputs from two CMIP6 experiments are used: (1) the historical experiment of
313 climate change over the period 1850~2014, forced by time-varying external forcings
314 that are based on observations of natural processes (e.g., solar activity, volcanic
315 eruptions) and human-induced changes (e.g., greenhouse gas, aerosol emissions, land-
316 use changes). All the available realizations for each model were used to minimize the
317 uncertainty from internal variability in the climate system; (2) the 1pctCO2 simulations,
318 in which CO₂ is gradually increased at a rate of 1% per year. The 1pctCO2 experiment
319 is designed for studying model responses to CO₂ and is somewhat more realistic than
320 rapidly increasing CO₂ such as in the abrupt-4×CO₂ experiment.

321 Model outputs used in this study include surface air temperature (SAT) and the



322 following five sulfur-cycle variables: total atmospheric burden of sulfate (loadSO₄),
323 sulfate wet deposition and sulfate dry deposition, SO₂ wet deposition and SO₂ dry
324 deposition. Uncertainty of the five sulfur cycle related variables is relatively smaller
325 than SAT uncertainty among different members. For example, the standard deviation
326 of loadSO₄ in PHC among 11 CESM2 members is 4% of the interannual variability,
327 whereas the value for SAT is about 21%. Similar results are also evident in 19
328 UKESM1 members, 3% for sulfate burden and 32% for SAT. Therefore, we use the
329 first realization of historical simulations and neglect the inter-member differences for
330 sulfur cycle related variables.

331

332 **A4 HadCRUT**

333 The monthly mean SAT from the Met Office Hadley Centre/Climatic Research
334 Unit global surface temperature (HadCRUT) data version 5 from 1850 to 2014 is used
335 for model evaluations (Morice et al., 2021). Considering the lack of long-term reliable
336 observations in polar regions, we focus on SAT changes between 60°S to 65°N and the
337 ‘global’ mean is calculated as the area-weighted mean in this latitudinal belt.

338

339 **Code availability**

340 All data processing codes are available if a request is sent to the corresponding authors.

341

342 **Data availability**

343 The HadCRUT5 dataset is accessible through Met Office Hadley Centre observations
344 database (<https://www.metoffice.gov.uk/hadobs/hadcrut5/>). All the model data can be
345 freely downloaded from the Earth System Grid Federation (ESGF) nodes
346 (<https://aims2.llnl.gov/search/cmip6/>).

347

348



349 **Author contributions**

350 The main ideas were formulated by J.Z. and K.F. J.Z. wrote the original draft. The
351 results were supervised by K.F. and S.T.T. All the authors discussed the results and
352 contributed to the final manuscript.

353

354 **Competing interests**

355 The authors declare no competing financial and/or non-financial interests.

356

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366



367 **References**

- 368 Aas, W., Mortier, A., Bowersox, V., Cherian, R., Faluvegi, G., Fagerli, H., Hand, J.,
369 Klimont, Z., Galy-Lacaux, C., Lehmann, C. M. B., Myhre, C. L., Myhre, G., Olivié,
370 D., Sato, K., Quaas, J., Rao, P. S. P., Schulz, M., Shindell, D., Skeie, R. B., Stein,
371 A., Takemura, T., Tsyro, S., Vet, R., and Xu, X.: Global and regional trends of
372 atmospheric sulfur, *Scientific Reports*, 9, 953, 10.1038/s41598-018-37304-0,
373 2019.
- 374 Bevacqua, E., Schleussner, C., and Zscheischler, J.: A year above 1.5 °C signals that
375 Earth is most probably within the 20-year period that will reach the Paris
376 Agreement limit, *Nature Climate Change*, 1-4, 10.1038/s41558-025-02246-9,
377 2025.
- 378 Chylek, P., Folland, C., Klett, J. D., and Dubey, M. K.: CMIP5 Climate Models
379 Overestimate Cooling by Volcanic Aerosols, *Geophysical Research Letters*, 47,
380 e2020GL087047, <https://doi.org/10.1029/2020GL087047>, 2020.
- 381 Danabasoglu, G., Lamarque, J. F., Bacmeister, J., Bailey, D. A., DuVivier, A. K.,
382 Edwards, J., Emmons, L. K., Fasullo, J., Garcia, R., Gettelman, A., Hannay, C.,
383 Holland, M. M., Large, W. G., Lauritzen, P. H., Lawrence, D. M., Lenaerts, J. T.
384 M., Lindsay, K., Lipscomb, W. H., Mills, M. J., Neale, R., Oleson, K. W., Otto-
385 Bliesner, B., Phillips, A. S., Sacks, W., Tilmes, S., van Kampenhout, L.,
386 Vertenstein, M., Bertini, A., Dennis, J., Deser, C., Fischer, C., Fox-Kemper, B.,
387 Kay, J. E., Kinnison, D., Kushner, P. J., Larson, V. E., Long, M. C., Mickelson,
388 S., Moore, J. K., Nienhouse, E., Polvani, L., Rasch, P. J., and Strand, W. G.: The
389 Community Earth System Model Version 2 (CESM2), *J. Adv. Model. Earth Syst.*,
390 12, 35, 10.1029/2019ms001916, 2020.
- 391 Dittus, A. J., Hawkins, E., Wilcox, L. J., Sutton, R. T., Smith, C. J., Andrews, M. B.,
392 and Forster, P. M.: Sensitivity of Historical Climate Simulations to Uncertain
393 Aerosol Forcing, *Geophysical Research Letters*, 47, e2019GL085806,
394 10.1029/2019gl085806, 2020.
- 395 Döscher, R., Acosta, M., Alessandri, A., Anthoni, P., Arneth, A., Arsouze, T.,



- 396 Bergmann, T., Bernadello, R., Bousetta, S., Caron, L. P., Carver, G., Castrillo, M.,
397 Catalano, F., Cvijanovic, I., Davini, P., Dekker, E., Doblas-Reyes, F. J., Docquier,
398 D., Echevarria, P., Fladrich, U., Fuentes-Franco, R., Gröger, M., v. Hardenberg,
399 J., Hieronymus, J., Karami, M. P., Keskinen, J. P., Koenigk, T., Makkonen, R.,
400 Massonnet, F., Ménégos, M., Miller, P. A., Moreno-Chamarro, E., Nieradzic, L.,
401 van Noije, T., Nolan, P., O'Donnell, D., Ollinaho, P., van den Oord, G., Ortega,
402 P., Prims, O. T., Ramos, A., Reerink, T., Rousset, C., Ruprich-Robert, Y., Le Sager,
403 P., Schmith, T., Schrödner, R., Serva, F., Sicardi, V., Sloth Madsen, M., Smith, B.,
404 Tian, T., Tourigny, E., Uotila, P., Vancoppenolle, M., Wang, S., Wärlind, D.,
405 Willén, U., Wyser, K., Yang, S., Yepes-Arbós, X., and Zhang, Q.: The EC-Earth3
406 Earth System Model for the Climate Model Intercomparison Project 6, *Geosci.*
407 *Model Dev. Discuss.*, 2021, 1-90, 10.5194/gmd-2020-446, 2021.
- 408 Dunne, J. P., Horowitz, L. W., Adcroft, A. J., Ginoux, P., Held, I. M., John, J. G.,
409 Krasting, J. P., Malyshev, S., Naik, V., Paulot, F., Shevliakova, E., Stock, C. A.,
410 Zadeh, N., Balaji, V., Blanton, C., Dunne, K. A., Dupuis, C., Durachta, J., Dussin,
411 R., Gauthier, P. P. G., Griffies, S. M., Guo, H., Hallberg, R. W., Harrison, M., He,
412 J., Hurlin, W., McHugh, C., Menzel, R., Milly, P. C. D., Nikonov, S., Paynter, D.
413 J., Ploshay, J., Radhakrishnan, A., Rand, K., Reichl, B. G., Robinson, T.,
414 Schwarzkopf, D. M., Sentman, L. T., Underwood, S., Vahlenkamp, H., Winton,
415 M., Wittenberg, A. T., Wyman, B., Zeng, Y., and Zhao, M.: The GFDL Earth
416 System Model Version 4.1 (GFDL-ESM 4.1): Overall Coupled Model Description
417 and Simulation Characteristics, *J. Adv. Model. Earth Syst.*, 12,
418 10.1029/2019ms002015, 2020.
- 419 EB, U.: Protocol to Abate Acidification, Eutrophication and Ground-level Ozone, 1999.
- 420 Flynn, C. M. and Mauritsen, T.: On the climate sensitivity and historical warming
421 evolution in recent coupled model ensembles, *Atmos. Chem. Phys.*, 20, 7829-7842,
422 10.5194/acp-20-7829-2020, 2020.
- 423 Hajima, T., Watanabe, M., Yamamoto, A., Tatebe, H., Noguchi, M. A., Abe, M.,
424 Ohgaito, R., Ito, A., Yamazaki, D., Okajima, H., Ito, A., Takata, K., Ogochi, K.,
425 Watanabe, S., and Kawamiya, M.: Development of the MIROC-ES2L Earth



- 426 system model and the evaluation of biogeochemical processes and feedbacks,
427 Geoscientific Model Development, 13, 2197-2244, 10.5194/gmd-13-2197-2020,
428 2020.
- 429 Hand, J. L., Schichtel, B. A., Malm, W. C., and Pitchford, M. L.: Particulate sulfate ion
430 concentration and SO₂ emission trends in the United States from the
431 early 1990s through 2010, Atmos. Chem. Phys., 12, 10353-10365, 10.5194/acp-
432 12-10353-2012, 2012.
- 433 Hansen, J., Kharecha, P., Sato, M., Tselioudis, G., Kelly, J., Bauer, S., Ruedy, R., Jeong,
434 E., Jin, Q., Rignot, E., Velicogna, I., Schoeberl, M., Schuckmann, K., Amponsem,
435 J., Cao, J., Keskinen, A., Li, J., and Pokela, A.: Global Warming Has Accelerated:
436 Are the United Nations and the Public Well-Informed?, Environment: Science and
437 Policy for Sustainable Development, 67, 6-44, 10.1080/00139157.2025.2434494,
438 2025.
- 439 Hardacre, C., Mulcahy, J. P., Pope, R. J., Jones, C. G., Rumbold, S. T., Li, C., Johnson,
440 C., and Turnock, S. T.: Evaluation of SO₂,
441 SO₄²⁻ and an updated SO₂ dry
442 deposition parameterization in the United Kingdom Earth System Model,
443 Atmospheric Chemistry and Physics, 21, 18465-18497, 10.5194/acp-21-18465-
444 2021, 2021.
- 445 Hienola, A., Partanen, A.-I., Pietikäinen, J.-P., O'Donnell, D., Korhonen, H., Matthews,
446 H. D., and Laaksonen, A.: The impact of aerosol emissions on the 1.5 °C pathways,
447 Environmental Research Letters, 13, 044011, 10.1088/1748-9326/aab1b2, 2018.
- 448 Hoesly, R. M., Smith, S. J., Feng, L., Klimont, Z., Janssens-Maenhout, G., Pitkanen,
449 T., Seibert, J. J., Vu, L., Andres, R. J., Bolt, R. M., Bond, T. C., Dawidowski, L.,
450 Kholod, N., Kurokawa, J. I., Li, M., Liu, L., Lu, Z., Moura, M. C. P., O'Rourke,
451 P. R., and Zhang, Q.: Historical (1750–2014) anthropogenic emissions of reactive
452 gases and aerosols from the Community Emissions Data System (CEDS), Geosci.
453 Model Dev., 11, 369-408, 10.5194/gmd-11-369-2018, 2018.
- 454 IPCC. Climate Change 2021 – The Physical Science Basis: Working Group I
455 Contribution to the Sixth Assessment Report of the Intergovernmental Panel on



- 456 Climate Change, 10.1017/9781009157896, 2023.
- 457 Likens, G. E., Butler, T. J., and Buso, D. C.: Long- and short-term changes in sulfate
458 deposition: Effects of the 1990 Clean Air Act Amendments, *Biogeochemistry*, 52,
459 1-11, 10.1023/a:1026563400336, 2001.
- 460 Mauritsen, T., Bader, J., Becker, T., Behrens, J., Bittner, M., Brokopf, R., Brovkin, V.,
461 Claussen, M., Crueger, T., Esch, M., Fast, I., Fiedler, S., Flaeschner, D., Gayler,
462 V., Giorgetta, M., Goll, D. S., Haak, H., Hagemann, S., Hedemann, C.,
463 Hohenegger, C., Ilyina, T., Jahns, T., Jimenez-de-la-Cuesta, D., Jungclaus, J.,
464 Kleinen, T., Kloster, S., Kracher, D., Kinne, S., Kleberg, D., Lasslop, G.,
465 Kornbluch, L., Marotzke, J., Matei, D., Meraner, K., Mikolajewicz, U., Modali,
466 K., Moebis, B., Muellner, W. A., Nabel, J. E. M. S., Nam, C. C. W., Notz, D.,
467 Nyawira, S.-S., Paulsen, H., Peters, K., Pincus, R., Pohlmann, H., Pongratz, J.,
468 Popp, M., Raddatz, T. J., Rast, S., Redler, R., Reick, C. H., Rohrschneider, T.,
469 Schemann, V., Schmidt, H., Schnur, R., Schulzweida, U., Six, K. D., Stein, L.,
470 Stemmler, I., Stevens, B., von Storch, J.-S., Tian, F., Voigt, A., Vrese, P., Wieners,
471 K.-H., Wilkenskjaeld, S., Winkler, A., and Roeckner, E.: Developments in the MPI-
472 M Earth System Model version 1.2 (MPI-ESM1.2) and Its Response to Increasing
473 CO₂, *J. Adv. Model. Earth Syst.*, 11, 998-1038, 10.1029/2018ms001400, 2019.
- 474 Morice, C. P., Kennedy, J. J., Rayner, N. A., Winn, J. P., Hogan, E., Killick, R. E.,
475 Dunn, R. J. H., Osborn, T. J., Jones, P. D., and Simpson, I. R.: An Updated
476 Assessment of Near-Surface Temperature Change From 1850: The HadCRUT5
477 Data Set, *Journal of Geophysical Research-Atmospheres*, 126,
478 10.1029/2019jd032361, 2021.
- 479 Mulcahy, J. P., Jones, C. G., Rumbold, S. T., Kuhlbrodt, T., Dittus, A. J., Blockley, E.
480 W., Yool, A., Walton, J., Hardacre, C., Andrews, T., Bodas-Salcedo, A., Stringer,
481 M., de Mora, L., Harris, P., Hill, R., Kelley, D., Robertson, E., and Tang, Y.:
482 UKESM1.1: development and evaluation of an updated configuration of the UK
483 Earth System Model, *Geosci. Model Dev.*, 16, 1569-1600, 10.5194/gmd-16-1569-
484 2023, 2023.
- 485 Mulcahy, J. P., Johnson, C., Jones, C. G., Povey, A. C., Scott, C. E., Sellar, A., Turnock,



486 S. T., Woodhouse, M. T., Abraham, N. L., Andrews, M. B., Bellouin, N., Browse,
487 J., Carslaw, K. S., Dalvi, M., Folberth, G. A., Glover, M., Grosvenor, D. P.,
488 Hardacre, C., Hill, R., Johnson, B., Jones, A., Kipling, Z., Mann, G., Mollard, J.,
489 O'Connor, F. M., Palmieri, J., Reddington, C., Rumbold, S. T., Richardson, M.,
490 Schutgens, N. A. J., Stier, P., Stringer, M., Tang, Y., Walton, J., Woodward, S.,
491 and Yool, A.: Description and evaluation of aerosol in UKESM1 and HadGEM3-
492 GC3.1 CMIP6 historical simulations, *Geoscientific Model Development*, 13,
493 6383-6423, 10.5194/gmd-13-6383-2020, 2020.

494 O'Neill, B. C., Tebaldi, C., van Vuuren, D. P., Eyring, V., Friedlingstein, P., Hurtt, G.,
495 Knutti, R., Kriegler, E., Lamarque, J. F., Lowe, J., Meehl, G. A., Moss, R., Riahi,
496 K., and Sanderson, B. M.: The Scenario Model Intercomparison Project
497 (ScenarioMIP) for CMIP6, *Geosci. Model Dev.*, 9, 3461-3482, 10.5194/gmd-9-
498 3461-2016, 2016.

499 Ohara, T., Akimoto, H., Kurokawa, J., Horii, N., Yamaji, K., Yan, X., and Hayasaka,
500 T.: An Asian emission inventory of anthropogenic emission sources for the period
501 1980-2020, *Atmospheric Chemistry and Physics*, 7, 4419-4444, 10.5194/acp-7-
502 4419-2007, 2007.

503 Samset, B. H., Sand, M., Smith, C. J., Bauer, S. E., Forster, P. M., Fuglestad, J. S.,
504 Osprey, S., and Schleussner, C.-F.: Climate Impacts From a Removal of
505 Anthropogenic Aerosol Emissions, *Geophysical Research Letters*, 45, 1020-1029,
506 <https://doi.org/10.1002/2017GL076079>, 2018.

507 Seland, Ø., Bentsen, M., Olivie, D., Toniazzo, T., Gjermundsen, A., Graff, L. S.,
508 Debernard, J. B., Gupta, A. K., He, Y. C., Kirkevåg, A., Schwinger, J., Tjiputra,
509 J., Aas, K. S., Bethke, I., Fan, Y., Griesfeller, J., Grini, A., Guo, C., Ilicak, M.,
510 Karset, I. H. H., Landgren, O., Liakka, J., Moseid, K. O., Nummelin, A.,
511 Spensberger, C., Tang, H., Zhang, Z., Heinze, C., Iversen, T., and Schulz, M.:
512 Overview of the Norwegian Earth System Model (NorESM2) and key climate
513 response of CMIP6 DECK, historical, and scenario simulations, *Geosci. Model*
514 *Dev.*, 13, 6165-6200, 10.5194/gmd-13-6165-2020, 2020.

515 Sellar, A. A., Jones, C. G., Mulcahy, J. P., Tang, Y., Yool, A., Wiltshire, A., O'Connor,



- 516 F. M., Stringer, M., Hill, R., Palmieri, J., Woodward, S., de Mora, L., Kuhlbrodt,
517 T., Rumbold, S. T., Kelley, D. I., Ellis, R., Johnson, C. E., Walton, J., Abraham,
518 N. L., Andrews, M. B., Andrews, T., Archibald, A. T., Berthou, S., Burke, E.,
519 Blockley, E., Carslaw, K., Dalvi, M., Edwards, J., Folberth, G. A., Gedney, N.,
520 Griffiths, P. T., Harper, A. B., Hendry, M. A., Hewitt, A. J., Johnson, B., Jones,
521 A., Jones, C. D., Keeble, J., Liddicoat, S., Morgenstern, O., Parker, R. J., Predoi,
522 V., Robertson, E., Siahann, A., Smith, R. S., Swaminathan, R., Woodhouse, M. T.,
523 Zeng, G., and Zerroukat, M.: UKESM1: Description and Evaluation of the UK
524 Earth System Model, *J. Adv. Model. Earth Syst.*, 11, 4513-4558,
525 10.1029/2019ms001739, 2019.
- 526 Smith, C. J. and Forster, P. M.: Suppressed Late-20th Century Warming in CMIP6
527 Models Explained by Forcing and Feedbacks, *Geophysical Research Letters*, 48,
528 10.1029/2021gl094948, 2021.
- 529 Steffen, W., Crutzen, P. J., and McNeill, J. R.: The Anthropocene: Are humans now
530 overwhelming the great forces of nature, *Ambio*, 36, 614-621, 10.1579/0044-
531 7447(2007)36[614:taahno]2.0.co;2, 2007.
- 532 Tatebe, H., Ogura, T., Nitta, T., Komuro, Y., Ogochi, K., Takemura, T., Sudo, K.,
533 Sekiguchi, M., Abe, M., Saito, F., Chikira, M., Watanabe, S., Mori, M., Hirota, N.,
534 Kawatani, Y., Mochizuki, T., Yoshimura, K., Takata, K., O'Ishi, R., Yamazaki, D.,
535 Suzuki, T., Kurogi, M., Kataoka, T., Watanabe, M., and Kimoto, M.: Description
536 and basic evaluation of simulated mean state, internal variability, and climate
537 sensitivity in MIROC6, *Geoscientific Model Development*, 12, 2727-2765,
538 10.5194/gmd-12-2727-2019, 2019.
- 539 Vestreng, V., Myhre, G., Fagerli, H., Reis, S., and Tarrasón, L.: Twenty-five years of
540 continuous sulphur dioxide emission reduction in Europe, *Atmos. Chem. Phys.*, 7,
541 3663-3681, 10.5194/acp-7-3663-2007, 2007.
- 542 Wang, Z., Lin, L., Xu, Y., Che, H., Zhang, X., Zhang, H., Dong, W., Wang, C., Gui,
543 K., and Xie, B.: Incorrect Asian aerosols affecting the attribution and projection
544 of regional climate change in CMIP6 models, *Npj Climate and Atmospheric*
545 *Science*, 4, 10.1038/s41612-020-00159-2, 2021.



- 546 Wilcox, L. J., Highwood, E. J., and Dunstone, N. J.: The influence of anthropogenic
547 aerosol on multi-decadal variations of historical global climate, *Environmental*
548 *Research Letters*, 8, 10.1088/1748-9326/8/2/024033, 2013.
- 549 Wu, T., Zhang, F., Zhang, J., Jie, W., Zhang, Y., Wu, F., Li, L., Yan, J., Liu, X., Lu,
550 X., Tan, H., Zhang, L., Wang, J., and Hu, A.: Beijing Climate Center Earth System
551 Model version 1 (BCC-ESM1): model description and evaluation of aerosol
552 simulations, *Geosci. Model Dev.*, 13, 977-1005, 10.5194/gmd-13-977-2020, 2020.
- 553 Yukimoto, S., Kawai, H., Koshiro, T., Oshima, N., Yoshida, K., Urakawa, S., Tsujino,
554 H., Deushi, M., Tanaka, T., Hosaka, M., Yabu, S., Yoshimura, H., Shindo, E.,
555 Mizuta, R., Obata, A., Adachi, Y., and Ishii, M.: The Meteorological Research
556 Institute Earth System Model Version 2.0, MRI-ESM2.0: Description and Basic
557 Evaluation of the Physical Component, *Journal of the Meteorological Society of*
558 *Japan*, 97, 931-965, 10.2151/jmsj.2019-051, 2019.
- 559 Zhang, J., Furtado, K., Turnock, S. T., Mulcahy, J. P., Wilcox, L. J., Booth, B. B.,
560 Sexton, D., Wu, T., Zhang, F., and Liu, Q.: The role of anthropogenic aerosols in
561 the anomalous cooling from 1960 to 1990 in the CMIP6 Earth System Models,
562 *Atmos. Chem. Phys. Discuss.*, 2021, 1-39, 10.5194/acp-2021-570, 2021a.
- 563 Zhang, J., Wu, T., Zhang, F., Furtado, K., Xin, X., Shi, X., Li, J., Chu, M., Zhang, L.,
564 Liu, Q., Yan, J., Wei, M., and Ma, Q.: BCC-ESM1 Model Datasets for the CMIP6
565 Aerosol Chemistry Model Intercomparison Project (AerChemMIP), *Advances in*
566 *Atmospheric Sciences*, 38, 317-328, 10.1007/s00376-020-0151-2, 2021b.