

1 **Drivers of soil C quality and stability: Insights from a topsoil**
2 **dataset at landscape scale in Ontario, Canada**

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9 **Abstract**

10 Although soil C is a critical component of soil health, studies robustly exploring the agronomic
11 and pedoclimatic effects on soil C are limited, especially at the landscape scale. Therefore, a
12 dataset of 1490 topsoil samples from agricultural fields across Ontario was used to evaluate the
13 impacts of agronomic and pedoclimatic factors on eight soil C indicators including chemistry and
14 thermal stability of soil C using the programmed pyrolysis approach. Soil C quality and stability
15 were largely controlled by the inherent soil characteristics such as soil texture. Significant
16 interactive effects of cropping system and tillage intensity on soil C indicators were observed;
17 however, the number of significant effects varied among the three soil textural classes. All soil C
18 indicators were significantly different among the cropping systems for the coarse textured soils,
19 but the cropping system differences decreased under medium and fine textured soils. From the
20 pyrolysis analysis, the hydrogen index (HI) and oxygen index (OI) also confirmed that the soil C
21 chemistry was influenced by the cropping system. For instance, orchard systems had stable pools
22 of soil C whereas vegetable systems were associated with less advanced degree of soil C
23 decomposition. Remaining soil management variables (cover crop use, tillage intensity, and
24 organic amendments) had a weaker influence than cropping systems and soil textural classes on

25 soil C indicators. Principal component analysis revealed a close association of soil C indicators
26 with the mean annual precipitation (MAP) and cropping system; suggesting that the quantity and
27 quality of soil C inputs associated with different cropping systems and increase in precipitation
28 had a large influence on soil C. Our results confirm the significant effects of agronomic and
29 pedoclimatic variables on chemistry, thermal stability, and composition of soil C pools, which
30 have long-term implications on soil C storage, mitigating global climate change, and improving
31 soil health.

32 **1. Introduction**

33 Soil C dynamics is estimated based on the amount of C in various soil C pools and the
34 transformation rate of C within these pools due to microbial processes (Cotrufo et al., 2013; Parton
35 et al., 1988). Several indicators, therefore, have been proposed to assess soil C dynamics. Soil
36 organic C (SOC) content or concentration is one of the most measured parameters of soil health
37 (Bunemann et al., 2018) and plays a significant role in many soil functions like nutrient and water
38 cycling and greenhouse gas emissions (Lal, 2016). Soil organic C comprises C compounds with a
39 wide range of stabilities, and so it takes a long time (5 to 10 years) to detect changes after
40 implementation of new land use management practices (Poeplau and Don, 2015). Total organic C
41 includes a labile C pool, which is easily metabolized by soil microbes and has a rapid turnover
42 time. Measurements of the labile soil C pool, by using the permanganate oxidizable C (POXC;
43 Moebius-Clune et al., 2016), potential C mineralization and soil respiration (Haney et al., 2008)
44 tests, are useful assessments of the quantity of C that is metabolized by the soil microbes and may
45 offer early detection of SOC changes from land use changes. Likewise, autoclaved citrate
46 extractable protein (ACE) may be used as a predictor of labile soil N and C (Agnihotri et al., 2022).
47 Assessment of these soil C indicators, hence, provide direct or indirect information on the soil C

48 storage and functions across diverse soil textures, management practices, and climatic conditions
49 (Liptzin et al., 2022). Most of the research on soil C indicators and its response to management
50 practices is conducted in field experiments and at a small-plot scale (Chahal et al., 2021; Culman
51 et al., 2013; Mesgar et al., 2024). In this study, we apply these techniques to samples obtained at
52 the landscape scale from operational agricultural fields. This is important to comprehensively
53 characterize the relationship between soil C indicators and the farm management strategies that
54 drive soil health.

55 Thermal analysis methods, such as programmed pyrolysis, is a novel technique in soil science
56 and is used to assess the molecular composition and the thermal stability of SOC (Gillespie et al.,
57 2014). Programmed pyrolysis subjects soil samples to a temperature ramp under an inert
58 atmosphere and measures the organic and inorganic C released as a function of increasing
59 temperature (Lafargue et al., 1998; Sebag et al., 2016). The thermal stability is related to the
60 biodegradation potential of SOC (Peltre et al., 2013; Sebag et al., 2016; Soucémariadin et al.,
61 2018), which is inferred from the hydrogen index (HI), oxygen index (OI), and T50 (temperature
62 at which 50% of the pyrolyzable C has been released). The HI primarily represents fresh,
63 hydrogenated organic matter, and is related to the labile pool of soil C whereas the OI represents
64 organic matter that has been oxidized through microbial metabolism, and is a more resistant and
65 stable pool of soil C (Carrie et al., 2012; Mesgar et al., 2024). The T50 is inversely related to the
66 decomposition potential, in that increasing T50 indicates lower decomposition potential and thus
67 more biologically stable organic matter (Gillespie et al., 2014; Gregorich et al., 2015). While the
68 HI, OI, and T50 are not direct indicators of soil health, the assessment of SOC stability and quality
69 using these indicators (pyrolysis method) provides valuable knowledge on how to build soil C and
70 to develop effective strategies to reduce the C loss under different land use management practices.

71 Soil C quality and stability is controlled by numerous variables such as soil texture, tillage,
72 crops grown, cover crops, use of organic amendments, changes in temperature, precipitation, and
73 the interactions among these factors (McDaniel and Grandy, 2016). Furthermore, these controlling
74 variables influence the composition of SOC and can potentially alter the stable and labile pools of
75 C (McDaniel and Grandy, 2016; Soon et al., 2007). Adding manure, for instance, to an intensively
76 managed long-term sorghum-wheat cropping system increased the labile fraction of soil C (Datta
77 et al., 2018). Soil C mineralization and respiration was increased by using no-tillage, cover crops,
78 and a diverse crop rotation (Balota et al., 2004; Chahal and Van Eerd, 2020; Viaud et al., 2011).
79 Adopting reduced tillage along with cover crops or perennial crops increased SOC content and
80 soil microbial biomass C (Sun et al., 2023). Likewise, POXC increased with the reduction in tillage
81 intensity and cover cropping (Liptzin et al., 2022). These studies from long-term experiments
82 confirm that different management practices impact the SOC composition which in turn, affects
83 the labile and stable pools of soil C. Yet, it remains uncertain which management or environmental
84 variables exert the strongest influence on soil C stability under the complexity of agricultural
85 fields.

86 Here, we used a large dataset of mineral topsoil samples collected from agricultural fields
87 across Ontario through the Ontario Topsoil Sampling Project (OTSP). Previously, the OTSP
88 dataset was used to assess the soil health scoring functions (Chahal et al., 2023) and SOC:clay
89 ratio as an indicator of soil functionality (Chahal et al., 2024). The goal of the present study was
90 to evaluate the impact of agricultural management and environmental variables on soil C indicators
91 (SOC, 96-h C mineralization potential (C_{\min} -96h), POXC, Solvita CO₂-burst, and ACE) and
92 indicators of thermal stability of soil C using programmed pyrolysis (HI, OI, and T50). We also
93 assessed the associations among these soil C indicators at the landscape scale to comprehensively

94 assess the major drivers of soil C storage and stability. The study results will contribute to making
95 improved recommendations regarding selection of soil C indicators and help growers and
96 researchers to adjust management practices to increase soil C storage.

97 **2. Materials and methods**

98 2.1 Soil sample collection

99 Topsoil samples (n=1511) for this study were collected as a part of the OTSP from 2019 to
100 2022. The soil samples for this project were collected from multiple locations throughout southern
101 Ontario. The OTSP was a collaborative project between the Ontario Ministry of Agriculture, Food,
102 and Agribusiness (formerly known as Ontario Ministry of Agriculture, Food, and Rural Affairs)
103 and the School of Environmental Sciences at the University of Guelph, to assess the soil physical,
104 chemical, and biological characteristics in agricultural soils in Ontario (Chahal et al., 2023; Chahal
105 et al., 2024). Soil samples were collected from the Ap horizon (agricultural tilled layer). To ensure
106 all samples were restricted to this horizon depth, the sampling depth was terminated at 30 cm if
107 the Ap horizon exceeded this depth. The median thickness of the sampled Ap horizons across all
108 the sites was 25 cm. Details about soil sample collection and the selection of the locations is
109 explained in detail in Chahal et al. (2023). Briefly, for each site, three soil samples were collected,
110 georeferenced, and a comprehensive land management survey with the grower was conducted to
111 document information on the crop rotation, type of crops grown, tillage, use of cover crops, and
112 application of organic amendments. The agricultural management factors identified in our study
113 were consistent with the commonly adopted practices by the growers and were representative of
114 the geographical area. Mean annual temperature (MAT) and mean annual precipitation (MAP) data
115 were collected using the WorldClim version 2.1 from 1970 to 2000
116 (<http://www.worldclim.org/data/worldclim21.html>). The MAP and MAT were grouped into two

117 (intermediate zone with precipitation between 800 and 1000 mm and wet zone with precipitation
118 greater than 1000 mm) and three classes (0 to 5°C, 5.01 to 10°C, and 10.01 to 15°C), respectively.
119 All the soil samples were classified into three soil textural classes (coarse with sand % ranging
120 between 52 to 94%), medium (between 2 to 78% sand) , and fine (between 1 to 45% sand);
121 Moebius-Clune et al., 2016), five cropping system (annual grain, forage, vegetable, orchard, and
122 perennial), five tillage intensity (conventional tillage, moderate disturbance, light disturbance, no
123 disturbance, and no-tillage), two cover crop (yes or no), and two organic amendment (yes or no)
124 classes. For the tillage intensity classes, conventional tillage represented the moldboard plow
125 tillage with full soil disturbance, moderate tillage represented more than 2 passes with disk or
126 chisel plow, light disturbance represented 1 or 2 passes with disk or cultivator, no disturbance
127 referred to the little or no soil movement as observed in pasture and perennial cropping systems,
128 and no-tillage represented minimal or no disturbance to the soil such as slot-tillage during planting.

129 We focused on the mineral soils for this study; thus, 21 soil samples with topsoil SOC
130 concentration more than 8.7% were removed from the dataset (Chahal et al., 2024). Therefore,
131 total number of samples used for the soil C analysis were 1490. It is important to note that the
132 number of observations for each soil C indicator (given in section 2.2) varied by the indicator and
133 hence the management practices. For instance, programmed pyrolysis was restricted to 151 soil
134 samples selected via conditioned Latin hypercube approach (Minasny and McBratney, 2006) as
135 representative of full dataset (Chahal et al., 2023).

136 2.2. Laboratory analyses

137 After quality assurance and quality control (QA/QC) analysis, the database consisted of eight
138 soil C indicators: SOC (n=1490), C_{min}-96h (n=1017), POXC (n=1413), Solvita CO₂-burst (n=768),
139 ACE protein (n=151), HI (n=151), OI (n=151) and T50 (n=151). As described above, 151 soil

140 samples used to measure ACE protein and pyrolysis parameters were representative of the full
141 dataset and refer to the same set of soil samples. Soil organic C concentration was calculated as
142 the difference between total C and inorganic C. Total C was estimated using the dry combustion
143 method (samples were combusted at 1300 °C) on LECO 828 Series CN analyzer (Skjemstad et al.,
144 2008). Inorganic C was determined by subjecting a ground subsample of soil to combustion in a
145 muffle furnace at 470°C to remove organic C (Krom and Berner, 1983). Soil C mineralization
146 (C_{\min} -96h) was quantified using the KOH trap method and was determined by measuring the
147 concentration of CO₂ evolved (mg CO₂-C 20 g⁻¹ soil) when 7.5 mL water was added to 20 g air
148 dried soil placed in an air-tight jar with 9 mL of 0.5 M KOH at room temperature (Schindelbeck
149 et al., 2016). Permanganate oxidizable C (µg g⁻¹) was measured using spectrophotometer where
150 the colorimetric change due to the reduction of manganese in a potassium permanganate solution
151 on air dried soil was quantified (Moebius-Clune et al., 2016). Solvita CO₂-burst was quantified as
152 per the updated protocol by the Woods End[®] Laboratories Inc., Mt. Vernon, ME. A 30-cc scoop
153 (around 25 to 35 g) of oven-dried (40°C) soil was placed in a vial and was wetted using 9 to 10
154 mL distilled water. The vials were transferred to 475 mL glass jars and Solvita CO₂-burst paddles
155 were inserted into each jar. The jars were sealed and left undisturbed for 24 h at room temperature
156 and evolved CO₂ concentration (mg kg⁻¹) was determined using a digital colorimeter reader
157 (Brinton, 2019). Autoclaved citrate extractable protein (mg kg⁻¹) was quantified on air-dried soils
158 by autoclaving, centrifuging, and treating a sodium citrate soil extract with bicinchoninic acid
159 (Schindelbeck et al., 2016). Soil particle size analysis was done using the pipette method where
160 soil organic matter was removed by treating the soil with hydrogen peroxide (Sheldrick and Wang,
161 1993). Consistent with soil textural grouping recommended by Moebius-Clune et al. (2016), the
162 dataset was divided into three soil textural classes (coarse, medium, and fine).

163 The programmed pyrolysis analysis was conducted using a HAWK pyrolyzer (Wildcat
164 Technologies, Humble, Texas, USA) at the Canadian Geological Survey in Calgary Alberta
165 according to Gillespie et al. (2014) and Gregorich et al. (2015). The standard manufacturer
166 recommended settings (with the helium flow rate at 100 mL/min) were followed to conduct the
167 programmed pyrolysis. About 70 mg air-dried ground soil sample is subjected to a constant
168 temperature of 300°C for 3 min under helium gas, to quantify free hydrocarbons using a flame
169 ionization detector (mg HC g⁻¹) and referred to as S1. Next, the soil sample was heated at a rate of
170 25°C per minute until 650°C where hydrocarbons were released (i.e., cracking SOC) and
171 quantified as S2. The concentration of CO₂ (mg CO₂ g⁻¹) released during S1 and S2 was determined
172 using an infrared detector and represented as S3. Hydrogen index of a soil sample represents the
173 ratio of all hydrocarbons measured as S1+S2 divided by SOC, whereas OI refers to the ratio of
174 CO₂ released (S3) divided by SOC (Lafargue et al., 1998). The stability of soil C is represented by
175 the T50 which is the temperature at which 50% of the SOC was pyrolyzed (Gillespie et al., 2014).
176 It is important to note that T50 is measured under S1 and S2 only.

177 2.3 Data analysis

178 All soil data were analyzed using the SAS (SAS Institute version 9.4, Cary, NC, USA). A
179 variance component analysis was conducted to test the relative importance of the agronomic and
180 pedoclimatic variables for each soil C indicator. Prior to conducting the variance component
181 analysis, the assumptions of normality were assessed. Given the relatively large sample sizes for
182 most of the variables studied, variance component estimates were considered robust to minor
183 deviations from normality. For each soil textural class, the main as well the interactive effect of
184 the agronomic management practices (cropping system, tillage intensity, cover crops, and organic
185 amendments) on soil C indicators were tested using PROC GLIMMIX in SAS. The relationship

186 among the soil C indicators was tested using the Pearson correlation analysis (PROC CORR). To
187 linearize the relationships among indicators, all the indicators were log-transformed for correlation
188 analysis. Consistent with Liptzin et al. (2022) and Mesgar et al. (2024), a principal component
189 analysis using PRINCOMP procedure was also conducted to explore how the soil C indicators
190 interacted with the site characteristics (i.e., agronomic management practices (cropping system,
191 tillage, cover crop, organic amendment) and pedoclimatic conditions (sand, silt, clay, mean annual
192 temperature and precipitation)). In addition to the site characteristics and pedoclimatic conditions,
193 the variables included in the PCA were SOC concentration, POXC, $C_{\min-24h}$, Solvita CO₂-burst,
194 ACE, HI, OI, and T50. The first two PCs were selected based on scree plots and the eigenvalues
195 of the soil C indicators were used to create the biplots to better understand the interdependence
196 among the soil C indicators and how they interacted with the site characteristics and pedoclimatic
197 conditions. The statistical significance of all the tests was assessed at $P < 0.05$.

198 **3. Results and discussion**

199 3.1 Agronomic and pedoclimatic effects on soil C indicators

200 The results of variance partitioning revealed that soil textural classes and cropping system
201 explained a larger percentage of variance for all soil C indicators compared to the other variables
202 (tillage intensity, cover crop, organic amendments, MAP and MAT; Table 1). For SOC, the amount
203 of variance explained by soil texture was higher than for cropping system (60.5% and 26.4%,
204 respectively), and for POXC, it was 53.5% and 21.7%, respectively (Table 1). Our results of strong
205 dependence of soil texture on SOC confirm the texture mediated soil C stabilization processes. For
206 instance, soil C retention is higher in fine textured than coarse textured soils mainly due to mineral-
207 organic associations and through the formation of microaggregates which physically protect the
208 soil C from microbial decomposition. This finding also aligned with Figure 1 where fine textured

209 soils exhibited greatest T50 values (i.e. indicating greater thermal stability); thus, further
210 highlighting the role of clay induced protection of SOC. Interestingly, the percentage of variance
211 explained by texture was comparable or less than for cropping system in C_{\min} -96h (38.8% and
212 36.9%), Solvita CO₂-burst (37.5% and 38.7%), and ACE (4.87% and 50.7%) (Table 1). The type
213 of main crops grown (i.e., the cropping system) was found to be a key driver of labile pools of soil
214 C (such as C_{\min} and ACE) in a study by Amsili et al. (2021). Soil texture also explained a large
215 amount of variance (75.2%) in the thermal-based parameters of soil C (Table 1). Measures of SOC
216 stability (T50) and quality (HI and OI) were largely controlled by soil texture for T50 and cropping
217 system for HI and OI. Unlike cropping system and soil textural classes, tillage intensity was not
218 found to be an important predictor of any of the tested soil C indicators (Table 1). While tillage
219 intensity was not found to be an important variable impacting soil C indicators when the bulk
220 dataset was used, its effects were detected for some soil C indicators when the data were
221 categorized based on soil textural classes. In particular, tillage intensity significantly influenced
222 POXC in fine-textured soils (Table 2) suggesting that effects of tillage are more pronounced on
223 labile soil C pools in soils with high clay content than the stable fractions of soil C. The least
224 amount of variance in soil C indicators was explained by MAT, use of cover crops, and organic
225 amendments (Table 1), suggesting a minor influence of these factors on soil C variability in our
226 study. One possible explanation for this result could be that cover crop and organic amendment
227 effects varies with soil type, climatic conditions, and baseline fertility which might have potentially
228 masked their overall impact on soil C indicators in our multi-site study. Therefore, consideration
229 of soil textural class and cropping system is needed when interpreting the soil C indicators (Nunes
230 et al., 2021).

231 It is important to note that the model R^2 values in our study were relatively low ranging
232 between 0.16 to 0.37 (Table 1). This is not unexpected for agronomic studies on SOC that integrate
233 variable land use management practices and complex biochemical processes. The unexplained
234 variance likely reflects the additional variables which were not captured in our model such as the
235 crop residue inputs, soil microbial community composition, and/or the historical land use intensity.
236 Although these variables are known to influence the SOC dynamics, the data on these factors were
237 not consistently available for all the sites studied. Incorporating these variables in the future
238 research might improve the model R^2 values; however, the current results still clearly highlight the
239 importance of soil texture as the major predictor of SOC in our study.

240 Table 2 shows the variance analysis on soil C indicators broken out by texture class and
241 parameter, and Table 3 shows the mean values and groupings. In all soil textural classes, significant
242 differences in SOC, $C_{\min-96h}$, and POXC concentration due to cropping systems were observed
243 (Table 2). Perennial and forages when grown with annual crops exhibited greater or comparable
244 SOC concentrations relative to the other cropping systems in all soil textural classes (Table 3).
245 Likewise, forages and perennial systems had greater or comparable concentrations of $C_{\min-96h}$ and
246 POXC than the remaining systems in all soil textural classes (Table 3). Less soil disturbance due
247 to tillage, greater diversity of crop species, continuous presence of living roots in perennial systems
248 perhaps contributed to the greatest concentration of SOC, $C_{\min-96h}$, and POXC observed (Amsili
249 et al., 2021; Congreves et al., 2015; Nunes et al., 2020). Compared to annual grain, perennial and
250 forage systems provide a more temporally consistent (i.e., no fallow period) source of substrate
251 quality to microbial communities mainly due to high C:N ratio, lower lignin content, and high
252 concentration of mineralizable C (Mesgar et al., 2024). For most of the soil C indicators, annual
253 grain, orchard, and vegetable cropping systems had the lowest values (Table 3). Vegetable and

254 annual grain systems are intensively managed with high intensity of tillage and have lower soil C
255 inputs (Norris and Congreves, 2018; Nunes et al., 2020), which negatively impact soil C.
256 Therefore, in all soil textural classes, diversification of cropping systems and addition of organic
257 amendments are critical components of building soil C. While our findings suggest that
258 diversification of cropping systems results in greater soil C, we did not measure the quantity and
259 quality of crop residue inputs, which limits our ability to confirm the exact mechanisms of soil C
260 accumulation in our study. Nevertheless, previous studies by King and Blesh (2018), McDaniel
261 and Grandy (2016) have reported that diversifying crop rotations with cover crops, perennials, or
262 forages tend to increase the quantity and biochemical diversity of soil C inputs than the
263 conventional monocultures (e.g. simple annual grain systems). Similarly, studies by Adhikari and
264 Hartemink (2017) and Presley et al. (2004) demonstrated that adoption of conservational
265 agricultural practices such as reduced tillage, diversified cropping systems and addition of organic
266 amendments contributed to a build up of SOC even on coarse textured soils.

267 While T50 was not statistically different due to cropping system and tillage practices in all soil
268 textures (Table 2), annual grain had the highest whereas forage had the lowest T50 when averaged
269 across all soil textures (Figure 1a). These results suggest that annual grain perhaps contribute to
270 more resistant organic matter additions to soil whereas forage systems might add relatively easily
271 decomposable residue. Usually, forage cropping systems (specifically legumes) have a higher
272 residue quality (high biomass N and low C:N) and result in larger proportion of labile fractions of
273 soil C. Mesgar et al. (2024) found similar results where crop rotations with forages (such as alfalfa
274 or red clover) contributed to labile components of soil C. Furthermore, fine textured soils had the
275 highest whereas coarse and medium textured soils had the lowest T50 (Table 3 and Figure 1b),
276 confirming that the soil texture and clay content influence the thermal stability of soil organic

277 matter. It is likely related to the greater organo-mineral associations in clay rich fine textured soils
278 than the coarse textured soils, which contributed to the protection of soil organic matter from
279 microbial decomposition and increase its thermal stability. Previous studies by Simkovic et al.
280 (2025) and Stoner et al. (2023) have also confirmed a positive relationship between clay content
281 and stabilization of soil organic matter.

282 The other thermal-based parameters of soil C characterization were the HI and OI, which
283 represented the maturity level of soil organic matter. Typically, a high HI represents a thermally
284 labile pool of organic matter which is enriched with hydrogen and the freshly added carbohydrates
285 and lignin (Mesgar et al., 2024). Conversely, OI represents a more resistant pool of organic matter
286 following the oxidation processes occurring during the soil organic matter decomposition (Carrie
287 et al., 2012; Mesgar et al., 2024; Saenger et al., 2013). Simultaneous reduction in both the HI and
288 OI indicates aromatization. The HI was not different among the agronomic management practices
289 in coarse and fine textured soils, but significant differences were observed in medium textured
290 soil. Among the cropping systems in medium textured soils, perennial (200 mg HC g⁻¹ OC) had
291 greatest while annual grain (132 mg HC g⁻¹ OC) had the least HI (Table 3). These results confirm
292 that the diversified cropping systems with forages and perennial crops had a higher quantity and
293 quality of H-rich labile components of soil organic matter and represented a more labile state of
294 organic matter decomposition than the intensively managed systems with less C inputs (Ding et
295 al., 2006; Mesgar et al., 2024). Furthermore, differences in OI due to cropping system were
296 detected in coarse textured soil only (Table 2). Orchard (203 mg CO₂ g⁻¹ OC) had greatest whereas
297 vegetable (147 mg CO₂ g⁻¹ OC) had the least OI in coarse textured soils; suggesting that the orchard
298 systems represent a more advanced state of soil organic matter decomposition than the other
299 cropping system categories.

300 A significant interaction between cropping system and tillage intensity was detected for some
301 of soil C indicators in all soil textural classes (Table 2), but trends varied among indicators and
302 texture. For the thermal based soil C indicators, significant interaction between cropping system
303 and tillage intensity was observed for the HI in medium-textured soil and for OI in coarse textured
304 soil (Table 2, Table S1 to S4). Given that the annual grain and forage cropping systems showed
305 the strongest contrast for these indicators with sufficient number of observations, we selected only
306 these two cropping systems to evaluate the effects of tillage intensity (Figure 2, Table S1). In
307 coarse-textured soils, $C_{\min-96h}$, POXC and Solvita CO_2 -burst concentrations were significantly
308 impacted by the intensity of tillage adopted in annual grain and forage systems, while the
309 remaining indicators were comparable across all the tillage and cropping system combinations
310 (Table S1). In medium-textured soils, all soil C indicators except Solvita CO_2 -burst had a
311 significant interaction between cropping system and tillage intensity (Table S1). In fine textured
312 soils, all but ACE were significantly different among the cropping system and tillage treatment
313 combinations (Table S1). Therefore, medium and fine textured soils had a greater number of
314 interactions than coarse textured soils. One possible mechanism might be that fine textured soils
315 have high clay content which stabilizes soil C via mineral adsorption and formation of
316 microaggregates. Fine textured soils also promote and support diverse soil microbial communities
317 which play a critical role in supporting the complex microbial mediated soil C transformation
318 processes (Six et al., 2002). In contrast, coarse textured soils have lower surface area, lower water
319 holding capacity, and less soil microbial activity, which perhaps contributed to lesser number of
320 detectable interactions between the management practices on soil C indicators. Overall, the
321 significant interaction of cropping system with tillage demonstrates that the changes in soil C pools
322 brought on by various tillage intensity treatments is dependent on the type of the crop species

323 grown and the soil texture. Similar interactive effects of tillage and cropping system on soil health
324 were reported by Angon et al. (2023).

325 A pseudo-Van Krevelen diagram was created by plotting HI against OI (Figure 3) to visually
326 characterize the composition of soil organic matter across various cropping systems (Carrie et al.,
327 2012; Mesgar et al., 2024). We found that most of the orchard systems have oxygenated products
328 and represented the more stable pool of soil organic matter whereas for vegetable systems, the
329 organic matter composition mainly consisted of hydrogenated products and was associated with a
330 less advanced stage of organic matter decomposition (Figure 3). While the visual representation
331 of HI vs OI between both systems (i.e. vegetable and orchards) may appear similar due to
332 variability and sample size, the underlying data distribution supports our interpretation (Figure 3).
333 Interestingly, the OI values for vegetable systems were not significantly different than the annual
334 grain systems in coarse textured soils (Table 3), suggesting that the slow decomposition of organic
335 matter in vegetable systems is dependent on the soil texture and agronomic management practices
336 followed. For the remaining systems, points found closer to the origin, such as in the annual grain
337 site data, suggests that organic matter structures are undergoing aromatization processes compared
338 to sites with forages or with perennial crops.

339 It is important to note that the frequency count of observations within the various tillage
340 intensity classes among the cropping system categories was not equal nor balanced (Figure 4) but
341 are reflective of typical management practices employed within the various cropping systems. For
342 instance, and as expected, the number of observations collected from the 'no disturbance' category
343 was greatest in perennial systems (n=142, Figure 4). Vegetable (n=5) and orchard (n=4) cropping
344 systems had the least number of observations for no-tillage (Figure 4). Orchards had the least
345 number of total observations (n=33, Figure 4), which is consistent with Ontario agriculture census

346 data where orchards represent 7.03% of farmland (Fruit and Vegetable Survey, Statistics Canada
347 2023). Furthermore, the frequency count of samples collected from coarse textured soils (n=308)
348 was lower than medium (n=642) and fine textured soils (n=540, Figure 5) which is largely
349 attributed to glacier deposits that shaped the region and the topography where more sand on top
350 and less fine particles as they are more prone to loss due to erosion. While clearly reflective of
351 Ontario soils and agriculture, the discrepancy in the count of observations suggests the need to be
352 cautious in directly attributing the results to a system.

353 3.2 Relationships of soil C indicators with agronomic and pedoclimatic variables

354 Principal component analysis was conducted where first and second PCs explained 32% and
355 17% of the variance in the data, respectively. Based on the PCA, soil C indicators and the measures
356 of SOC quality (HI and OI) were closely associated with the cropping system, MAP, and organic
357 amendments, and were negatively associated with MAT (Figure 6), suggesting a higher value for
358 soil C indicators under cooler temperatures. Although precipitation and organic amendments did
359 not explain a large amount of variance in our dataset (Table 1), the PCA demonstrated that the soil
360 C indicators increased with an increase in precipitation (Figure 6). It is important to note that PCA
361 and variance component analysis differ in both the statistical structure and objectives, which
362 perhaps led to differences in the results between both approaches. For instance, variance analysis
363 evaluates the independent effect of each predictor variable on soil C indicators, whereas PCA
364 simultaneously assesses the covariance among the multiple soil C indicators. Climate has been a
365 key determinant of soil C (Jenny, 1941). The interactions among the soil microbes, crop residues,
366 and plant root exudates mainly control the influence of climatic variables on soil C (Schmidt et al.,
367 2011). Our results of high values of soil C indicators with an increase in MAP and decrease in
368 MAT were consistent with studies conducted in North America (Burke et al., 1989; Liptzin et al.,

369 2022) and globally (Jobbagy and Jackson, 2000). While relationships of soil C indicators with
370 temperature and precipitation were consistent with expectations, it was surprising to see an effect
371 given the relative minimal differences within the province. In Ontario agricultural zones, MAT
372 varies by only 1⁰C, and MAP by approx. 100 mm. Given current climate change predictions, these
373 results have powerful implications for C sequestration and soil functioning under future climate
374 conditions.

375 Furthermore, silt and clay content were clustered together in PCA on one side of the second axis
376 whereas sand content was positioned on the opposite side of the second axis (Figure 6). This result
377 was consistent with the well-established associations between soil C dynamics and soil texture
378 where soils rich in clay content have higher C retention capacity than coarse textured soil (von
379 Lutzow et al., 2006). Accordingly, the positive loading displayed by the clay and silt rich soils on
380 the second axis corresponds to greater values of soil C indicators observed in our study. Although
381 important, the relationship between soil texture and soil C indicators (particularly POXC and
382 respiration) has not been explored enough in the literature (Nunes et al., 2020; Sinsabaugh et al.,
383 2008).

384 Soil C indicators and HI were negatively associated with tillage intensity on the first PC axis
385 (Figure 6). Increase in tillage intensity reduces soil C (mainly the topsoil C) by increasing the
386 mineralization of soil organic matter, disrupting the soil structure, and decreasing soil microbial
387 populations and communities (Nunes et al., 2020). Therefore, adopting reduced or minimum
388 tillage practices might contribute to building soil C and help to mitigate the negative impacts of
389 climate change. Numerically greater SOC concentration observed with reducing tillage intensity
390 in our study suggests the potential of the sustainable land use management practices on

391 sequestering C, reducing CO₂ emissions, and mitigating the global warming effect (Melland et al.,
392 2017) and greenhouse gas emissions (Mangalassery et al., 2014).

393 Except POXC in fine-textured soils, we did not find a significant effect of tillage intensity on
394 soil C indicators (Table 2 and S5). It is important to note that in our study, the participants were
395 asked to choose one of the tillage intensity categories, which might have caused a variability in
396 the recorded data (i.e. descriptive terms and actual disturbance on the farm) and might not be a
397 true reflection of tillage intensity over the long-term. Moreover, the producer interpretation of the
398 tillage intensity categories could have added uncertainty. Additionally, cover crops were negatively
399 associated with soil C indicators (Figure 6) but explained <5% of the variance in the dataset (Table
400 1) confirming a very small response of soil C indicators to cover crops (Table S6). It is not entirely
401 clear but might be attributed to a smaller number of observations with cover crops in our study
402 (n=55). The cover crop effects on soil C indicators are largely dependent on the cover crop
403 management factors such as type of cover crop species grown, frequency and duration of cover
404 cropping, planting date, and termination time of cover crops (Blanco-Canqui et al., 2015; Peng et
405 al., 2024). Due to the unavailability of the cover crop management factors in our study, the
406 interpretation of cover crop effect is challenging but clearly demonstrates that other management
407 factors (cropping system and tillage system) have a greater effect on soil C indicators.

408 3.3. Relationships among the soil C indicators

409 To better understand associations among soil C indicators, correlation analysis was conducted
410 on the soil C indicators and the indicators characterizing the chemical composition of soil organic
411 matter (Table 4). Interestingly, despite having a range of soil textural classes and cropping system
412 categories, strong positive significant relationships were observed among the soil C indicators
413 (Table 4). Among the indicators, SOC and POXC had the strongest positive relationship ($r=0.81$),

414 which was consistent with Liptzin et al. (2022) and Culman et al. (2012). Consistent with Amsili
415 et al. (2021) and Nunes et al. (2020), our results confirm that an increase in SOC positively impacts
416 the soil microbial activity (as demonstrated by Solvita CO₂-burst and C_{min}-96h) and the quality of
417 soil organic matter (as demonstrated by POXC and ACE). Significant moderate negative
418 associations ($r=-0.24$ to -0.35) were observed between HI and OI, HI and T50 (Table 4). Consistent
419 with Mesgar et al. (2024), our results suggest that the indicators representing the labile components
420 of soil organic matter such as HI were negatively associated with OI (an indicator of resistant pool
421 of soil organic matter) and T50. We also observed that the indicators defining the stable and labile
422 pools of soil C via the thermal analysis had a positive relationship with the soil C indicators such
423 as SOC and C mineralization (Table 4). The correlation analysis also confirmed that the easily
424 decomposable component of soil C (e.g., HI) was closely related to the labile indicators of soil C
425 such as soil respiration and ACE (Table 4). The positive association of ACE with HI confirms that
426 H-rich aliphatic C and protein like N compounds are concomitantly present in fresh organic matter
427 and are co-metabolized by the soil microbes during the early stages of organic matter
428 decomposition. Collectively, these results confirm the efficacy and applicability of the
429 programmed pyrolysis method as a valuable tool to study the biochemical composition and
430 decomposition potential of soil C.

431 **4. Conclusions**

432 This study focused on understanding the key drivers of soil C quality and stability in agricultural
433 production systems at a landscape scale. To the best of our knowledge, this is the first study to
434 evaluate the agronomic and pedoclimatic effects on the measures of soil C quality and C stability
435 (particularly with the pyrolysis approach) in North America at the landscape scale. Our results
436 revealed that soil textural classes and cropping system had a strong influence on both quality and

437 stability of soil C indicators evaluated in this study. The cropping system differences on soil C are
438 mainly related to the quantity and quality of residue C inputs, which in turn are primarily dependent
439 on the cropping system, cover crops, tillage intensity, and organic amendments. Among the
440 cropping systems, we found a greater concentration of soil C in forage and perennial cropping
441 systems than the annual grain and vegetable systems. Likewise, forage cropping systems had a
442 greater preservation of labile components while orchards had a more stable pool of soil C. Our
443 results, therefore, confirmed that agricultural management-induced factors play a crucial role in
444 understanding the chemical composition of soil organic matter. The indicators representing the
445 labile pools of soil C (such as POXC and $C_{\min-96h}$) were positively correlated with the parameters
446 indicating readily decomposable fractions of soil C (i.e., HI).

447 Furthermore, all indicators had a positive association with precipitation and a negative
448 relationship with temperature suggesting an increase in the indicator values at cool and wet
449 conditions despite low differences in values. Increase in the tillage intensity also negatively
450 impacted the soil C indicators. Overall, the management-induced differences in soil C indicators
451 in our study imply the benefits of adopting sustainable agricultural practices on building soil C,
452 reducing CO₂ emissions, and mitigating the negative effects of global climate change. While the
453 findings of our study suggest a significant impact of agronomic and pedoclimatic variables on the
454 soil C, it is important to note that the results of this study pertain only to the topsoil. It is possible
455 that the same trends or effects on soil C quality and stability might not be observed at the deeper
456 soil depths; hence, suggesting a need for future research. Additionally, tillage intensity categories
457 in our study were based on subjective producer descriptions and lacked standardized metrics such
458 as tillage depth, number of passes which might have introduced a potential bias in the interpretation
459 of study results. The small sample sizes of cover crops (n=55) and orchards (n=33) limit the

460 generalizability of results for these management practices. Future studies should employ more
461 standardized tillage measurements and ensure a more balanced sample sizes across the agronomic
462 management categories to improve the robustness of study results.

463 **Code and data availability**

464 Data will be made available upon request but is not available in an online repository to protect
465 the privacy of the participants in this project.

466 **Supplement**

467 The supplementary tables and figures related to this article are attached along with the
468 manuscript text.

469 **Author contributions**

470 IC, writing, data analysis, data interpretation and presentation, and editing the manuscript; DD,
471 data organization and curation, reviewing and editing the manuscript; AW, funding acquisition,
472 supervising, reviewing, and editing the manuscript; LVE, funding acquisition, supervising,
473 reviewing, and editing the manuscript.

474 **Competing interests**

475 The authors declare that they have no conflict of interest.

476 **Acknowledgments**

477 Authors are grateful to the financial support provided by Ontario Ministry of Agriculture, Food
478 and Agribusiness (OMAFRA) and the Ontario Agri-Food Innovation Alliance, a collaboration
479 between the Government of Ontario and the University of Guelph. We would like to thank the
480 undergraduate student Jane Bellefleur who conducted the lab analysis for ACE.

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617 **Table 1** Partitioning of variance of soil C indicators^z in the Ontario Topsoil Sampling Project database
 618 from 2019 to 2022 into soil and crop management, soil texture, and climatic variables.

		SOC	C _{min} -96h	POXC	Solvita CO ₂ -burst	ACE	T50	HI	OI
Variables	df								
Cropping system	4	26.4	36.9	21.8	38.8	50.7	10.0	59.8	43.0
Tillage intensity	4	4.82	18.0	6.30	11.8	22.3	10.4	11.0	10.7
Use of cover crop	1	0.23	2.61	1.07	5.07	13.3	4.21	1.02	2.14
Use of organic amendment	1	0.55	0.67	10.8	6.22	5.68	0.01	3.94	5.54
Soil textural classes	2	60.5	38.8	53.5	37.5	4.87	75.2	24.1	35.6
Mean annual precipitation	1	7.35	2.63	6.52	0.11	0.98	0.43	0.02	1.70
Mean annual temperature	1	0.15	0.39	0.01	0.59	1.82	0.84	0.01	1.22
Model R ²		0.16	0.28	0.13	0.22	0.23	0.37	0.19	0.21

^z Number of observations used were: SOC= 1490, C_{min}-96h=1017, Solvita CO₂-burst=768, and POXC=1413, whereas the number of observations for ACE and programmed pyrolysis parameters (HI, OI, and T50) were 151.

SOC=soil organic carbon; C_{min}-96h= 96-hr carbon mineralization; POXC=permanganate oxidizable carbon; Solvita CO₂-burst= 24 h soil respiration test;

ACE=autoclaved citrate extractable protein index; T50=temperature at which 50% of SOC was pyrolyzed; HI=hydrogen index; and OI=oxygen index.

For each soil C indicator, bold font indicates the top two variables explaining the variance.

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625 **Table 2** Within each soil textural class, variance analysis (*P* values) of soil and crop management
 626 parameters on soil C indicators^z analyzed in the Ontario Topsoil Sampling Project from 2019 to 2022.

		SOC	C _{min} -96h	POXC	Solvita CO ₂ - burst	ACE	T50	HI	OI
		mg g ⁻¹	mg CO ₂ - C 20 g ⁻¹ soil	µg g ⁻¹	mg kg ⁻¹	mg kg ⁻¹	°C	mg HC g ⁻¹ OC	mg CO ₂ g ⁻¹ OC
		<i>P</i> values							
Parameter	df	Coarse- textured							
Cropping system ^y	4 (3)	0.0403	0.0003	0.0044	0.0077	0.0298	0.8275	0.5131	0.0109
Tillage intensity	4	0.4035	0.6495	0.9057	0.4571	0.0505	0.7480	0.7405	0.6670
Cover crop	1	0.5625	0.2332	0.4637	0.3627	0.0055	0.0129	0.0887	0.4202
Use of organic amendment	1	0.9626	0.6656	0.1358	0.9150	0.6184	0.1869	0.3257	0.4489
Crop x Tillage	16 (12)	0.1535	0.0059	0.0071	0.0144	0.4207	0.4651	0.9249	0.0066
		Medium-textured							
Cropping system	4	0.0020	<0.0001	0.0003	0.7583	0.3830	0.4477	0.0005	0.3245
Tillage intensity	4	0.7602	0.2161	0.1347	0.0674	0.6528	0.5661	0.0595	0.2743
Cover crop	1	0.9172	0.0091	0.8947	0.2020	0.8529	0.6813	0.1646	0.0123
Use of organic amendment	1	0.7148	0.0735	0.4686	0.2045	0.6543	0.5044	0.0555	0.9541
Crop x Tillage	16	0.0025	<0.0001	0.0025	0.0610	0.0298	0.5861	0.0105	0.5871
		Fine-textured							
Cropping system ^y	4 (3)	0.0018	<0.0001	0.0485	<0.0001	0.1199	0.8067	0.2488	0.9702
Tillage intensity	4	0.7322	0.7080	0.0210	0.1948	0.4784	0.9968	0.5926	0.5507
Cover crop	1	0.1991	0.4240	0.0962	0.1037	0.1585	0.3945	0.8834	0.9502
Use of organic amendment	1	0.8311	0.2033	0.1543	0.5495	0.0905	0.0739	0.3373	0.3781
Crop x Tillage	16 (12)	0.0006	<0.0001	0.0007	<0.0001	0.2743	0.5131	0.6867	0.3253

627 Bold font indicates statistically significant treatment differences at *P*<0.05.

628 ^z Number of observations used were: SOC= 1490, C_{min}-96h=1017, Solvita CO₂-burst=768, and POXC=1413, whereas the number of observations for ACE and
 629 programmed pyrolysis parameters (HI, OI, and T50) were 151.

630 ^y For the coarse and fine-textured soil, number of cropping system categories analyzed for HI, OI, and T50 were 4. For instance, no soil samples were collected from
 631 the perennial cropping system in coarse-textured soil and from vegetable system in fine-textured soils. Hence, the degree of freedom for cropping system and
 632 interaction between cropping system and tillage treatments was adjusted accordingly in the statistical model.

633 SOC=soil organic C; C_{min}-96h= 96-hr carbon mineralization; POXC=permanganate oxidizable carbon; Solvita CO₂-burst= 24 h soil respiration test; ACE=autoclaved
 634 citrate extractable protein index; T50=temperature at which 50% of SOC was pyrolyzed; HI=hydrogen index; and OI=oxygen index.

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Table 3 Within each soil textural class, mean (SE) values of the soil C indicators^z by the cropping system category sampled in the Ontario Topsoil Sampling Project from 2019 to 2022.

Cropping system	n	SOC mg g ⁻¹	C _{min} -96h mg CO ₂ -C 20 g ⁻¹ soil	POXC µg g ⁻¹	Solvita CO ₂ -burst mg kg ⁻¹	ACE mg kg ⁻¹	T50 °C	HI mg HC g ⁻¹ OC	OI mg CO ₂ g ⁻¹ OC
Coarse-textured									
Annual grain	183	16.9ab (0.70)	14.2bc (0.96)	473ab (16.0)	53.1b (3.35)	5.49ab (0.78)	416 (1.58)	152 (5.94)	168bc (4.26)
Forage	53	19.4a (1.20)	20.4a (1.49)	547a (25.4)	72.0a (6.39)	8.45a (1.47)	412 (4.19)	163 (15.7)	207a (11.2)
Vegetable	32	15.4ab (1.50)	10.9c (1.61)	441ab (31.6)	89.4a (12.2)	6.60ab (1.75)	415 (5.13)	162 (19.2)	147c (13.8)
Orchard	9	11.7b (2.70)	14.4abc (5.70)	342b (46.0)	51.4ab (9.23)	4.71b (1.75)	419 (5.13)	132 (19.2)	203ab (13.8)
Perennial	31	19.4ab (1.70)	18.5ab (1.78)	510ab (34.5)	53.6ab (16.1)	--	--	--	--
All	308 (29) ^z	17.2	15.6	478	61.4	6.80	415	154	172
Medium-textured									
Annual grain	335	20.8b (0.40)	20.1b (0.57)	591a (9.74)	68.3 (2.76)	5.55ab (0.43)	421 (1.65)	132b (5.98)	196 (5.20)
Forage	129	20.2b (0.60)	22.6ab (0.92)	584a (15.2)	74.4 (3.66)	6.72ab (0.50)	417 (2.24)	159ab (8.12)	195 (7.07)
Vegetable	54	20.0b (1.00)	19.7b (1.27)	575a (22.3)	65.4 (7.74)	4.46b (1.02)	415 (5.72)	134ab (20.7)	164 (18.0)
Orchard	20	16.6b (1.50)	17.7b (2.58)	439b (34.4)	81.4 (7.94)	7.63ab (1.32)	420 (5.72)	184ab (20.7)	180 (18.0)
Perennial	104	24.0a (0.80)	25.2a (1.03)	560a (18.7)	81.1 (6.88)	10.1a (1.38)	412 (3.61)	200a (13.1)	188 (11.4)
All	642 (46)	21.1	21.6	581	72	6.30	419	149	193
Fine-textured									
Annual grain	315	23.1b (0.50)	21.2c (0.68)	574c (10.7)	76.2c (2.20)	5.62 (0.34)	428 (1.11)	132 (4.23)	186 (3.14)
Forage	131	25.5ab (0.80)	25.3ab (1.05)	623ab (15.8)	86.9b (3.36)	5.93 (0.64)	427 (2.46)	140 (9.38)	191 (6.97)
Vegetable	15	22.5b (2.50)	21.6bc (2.38)	606b (47.9)	--	--	--	--	--
Orchard	4	25.6ab (4.20)	35.0a (3.99)	694a (80.8)	107ab (13.6)	9.02 (2.05)	426 (8.16)	161 (31.3)	195 (23.1)
Perennial	72	28.9a (1.30)	30.5a (1.56)	613ab (26.7)	108a (5.81)	7.91 (0.82)	427 (2.58)	161 (9.84)	202 (7.31)
All ^y	540 (76)	24.5	23.4	603	83.9	6.27	427	138	189

^zThe number in the parenthesis represents the number of observations for ACE and the programmed parameters (T50, HI, and OI) in each soil textural class.

^yThere were 3 observations within the fine-textured soils for which cropping system details were missing.

^{a-c}Within each soil textural class and for each parameter, treatment means followed by a different letter indicate statistical significance at $P < 0.05$

SOC=soil organic carbon; C_{min}-96h= 96-hr carbon mineralization; POXC=permanganate oxidizable carbon; Solvita CO₂-burst= 24 h soil respiration test; ACE=autoclaved citrate extractable protein index; HI=hydrogen index; and OI=oxygen index.; --=not applicable.

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646 Table 4 Pearson correlation coefficients (r) among soil C indicators sampled in the Ontario Topsoil Sampling
647 Project from 2019 to 2022.

Indicators [^]	SOC	C _{min} -96h	POXC	Solvita CO ₂ -burst	ACE	HI	OI	T50
	mg g ⁻¹	mg CO ₂ -C 20 g ⁻¹ soil	μg g ⁻¹	mg CO ₂ -C kg ⁻¹	mg kg ⁻¹	mg HC g ⁻¹ OC	mg CO ₂ g ⁻¹ OC	°C
SOC	1.00							
C _{min} -24h	0.67***	1.00						
POXC	0.81***	0.63***	1.00					
Solvita CO ₂ -burst	0.36***	0.46***	0.34***	1.00				
ACE	0.72***	0.58***	0.66***	0.17**	1.00			
HI	0.28**	0.42***	0.24**	NS	0.59***	1.00		
OI	0.21**	0.27**	NS	NS	NS	-0.30**	1.00	
T50	NS	NS	NS	NS	NS	-0.24**	NS	1.00

648 SOC=soil organic carbon; C_{min}-96h= 96-hr carbon mineralization; POXC=permanganate oxidizable carbon; Solvita CO₂-burst= 24 h soil respiration test;

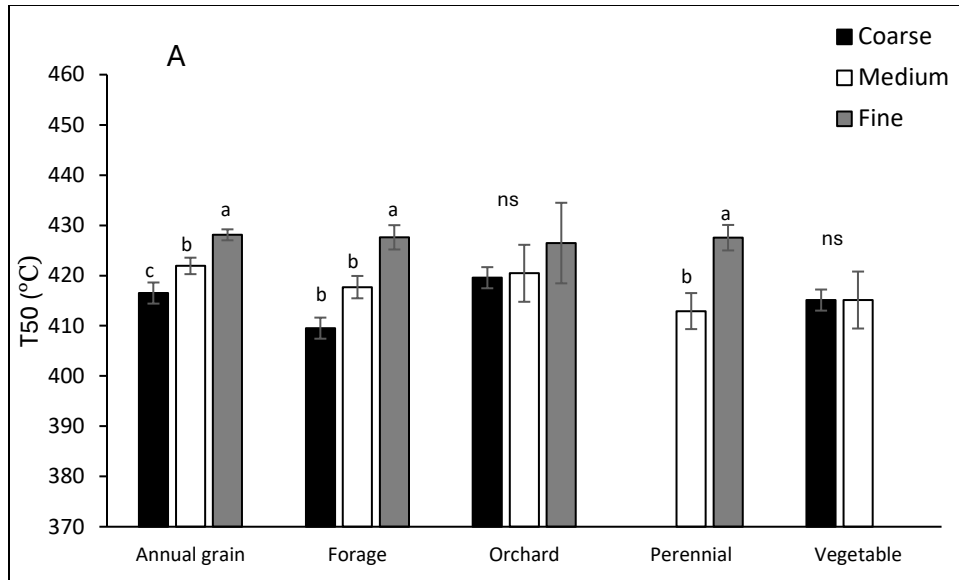
649 ACE=autoclaved citrate extractable protein index; HI=Hydrogen Index; OI=Oxygen Index; and NS=Non-significant.

650 n=1490 for soil organic C; n=1017 for C_{min}-96; n=1413 for POXC; n=768 for Solvita CO₂-burst; n=151 for ACE, HI, OI, T50.

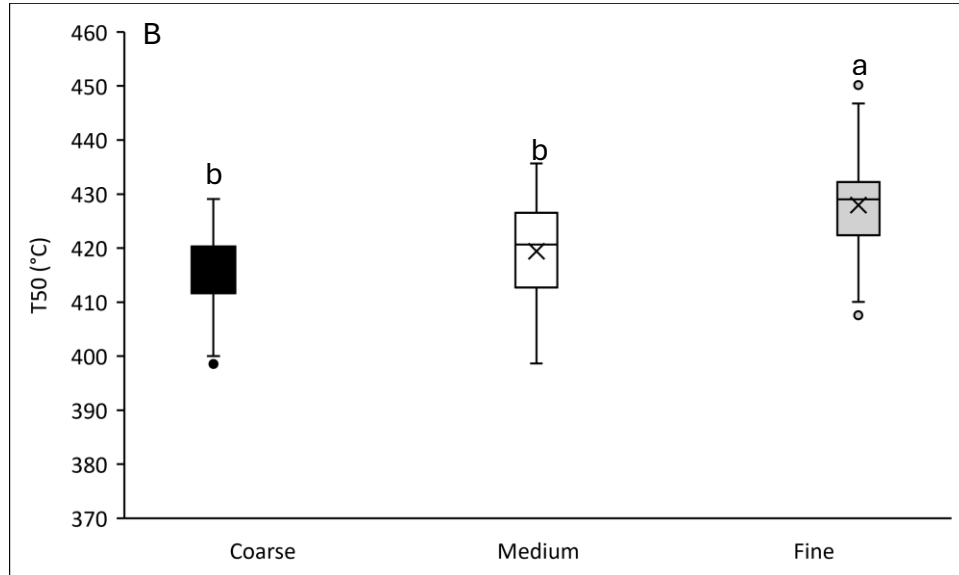
651 **,*** indicates statistical significance at P<0.05 and P<0.0001, respectively.

652 [^]All indicators were log-transformed prior to analysis.

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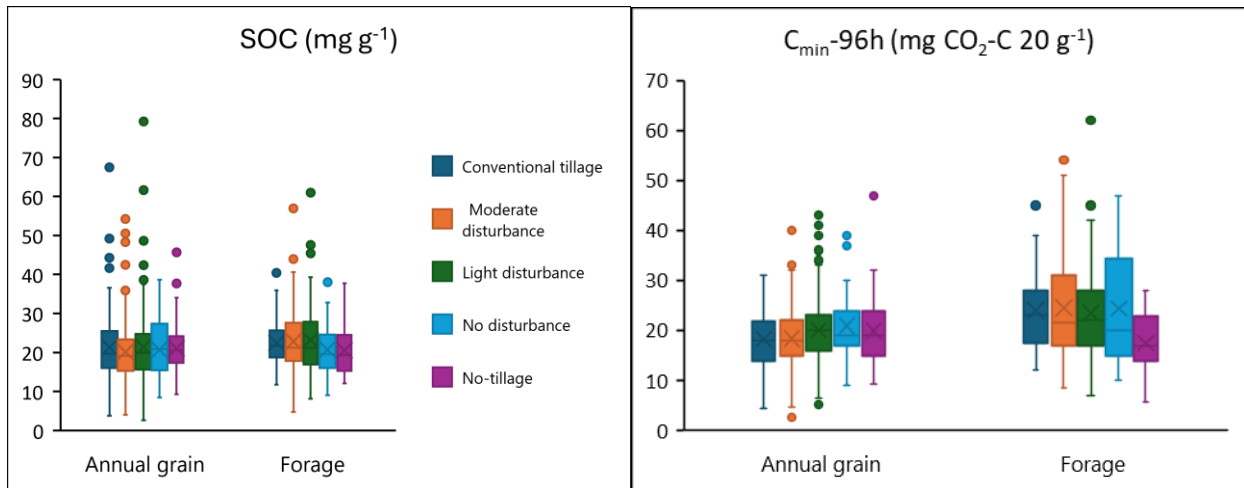
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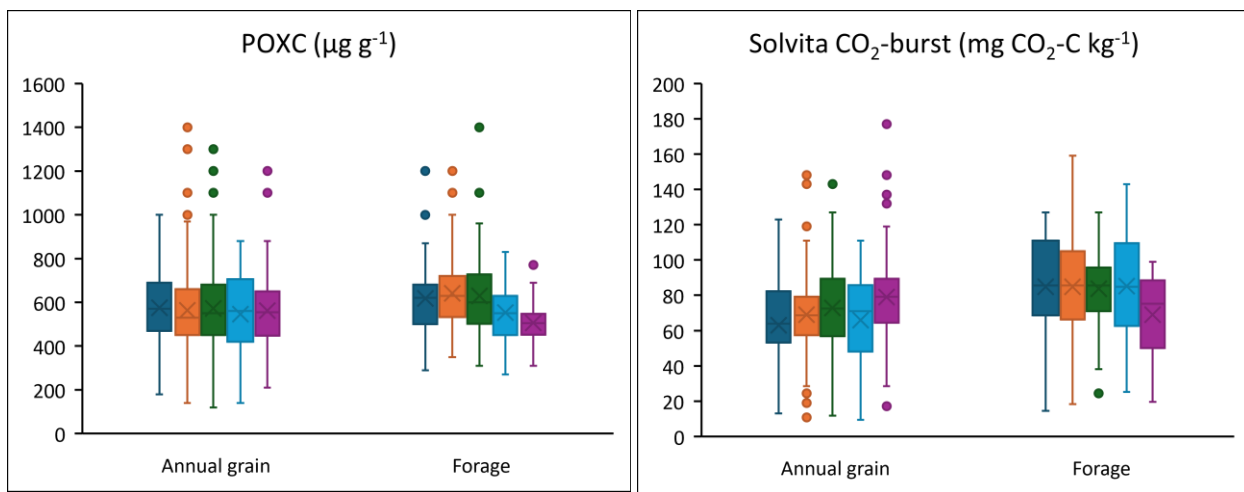
655 **Figure 1** Plots of T50 demonstrating differences due to soil textural class within each cropping system category (A)
656 and soil textural classes (B) for the soils in the Ontario Topsoil Sampling Project in 2019 (n=151). Different letters
657 indicate statistically significant differences at $P < 0.05$. ns represents non-significant statistical differences among soil
658 textural classes within the cropping system category.

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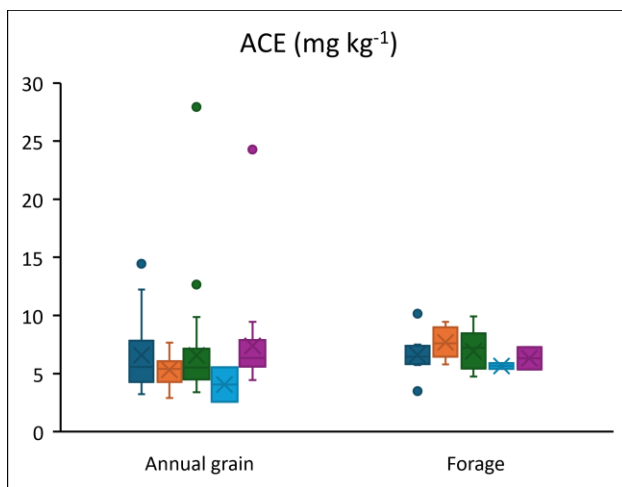
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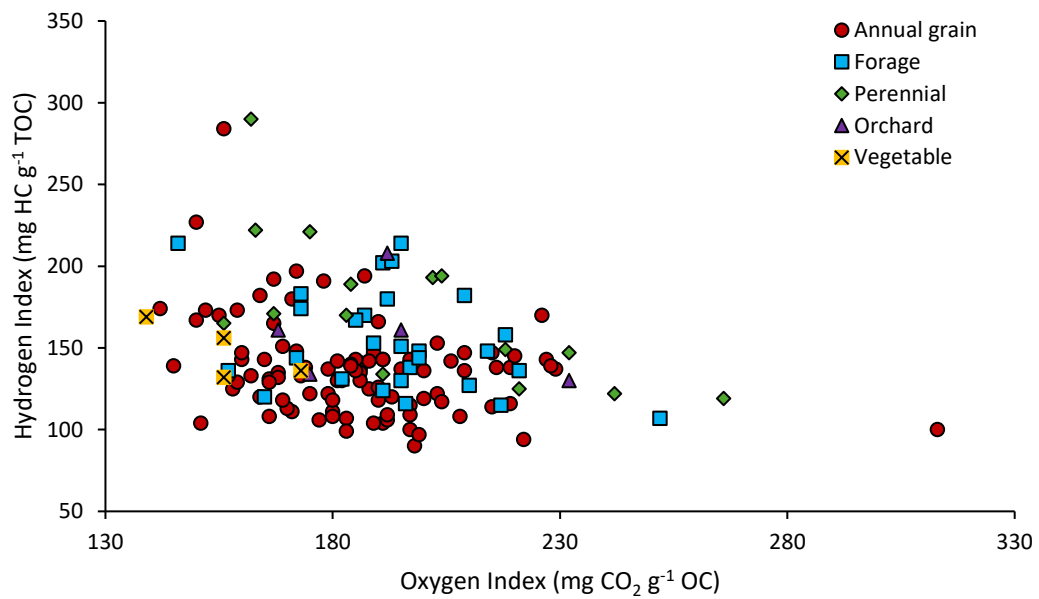
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663 **Figure 2** Box plots demonstrating the interactive effects of cropping system category and tillage intensity on soil C
664 indicators sampled in the Ontario Topsoil Sampling Project from 2019 to 2022. SOC=soil organic carbon; $C_{\min-96h}$ =
665 96-hr carbon mineralization; POXC=permanganate oxidizable carbon; Solvita CO_2 -burst= 24 h soil respiration test;
666 ACE=autoclaved citrate extractable protein index. Due to insufficient number of observations, data for other
667 cropping systems not shown.

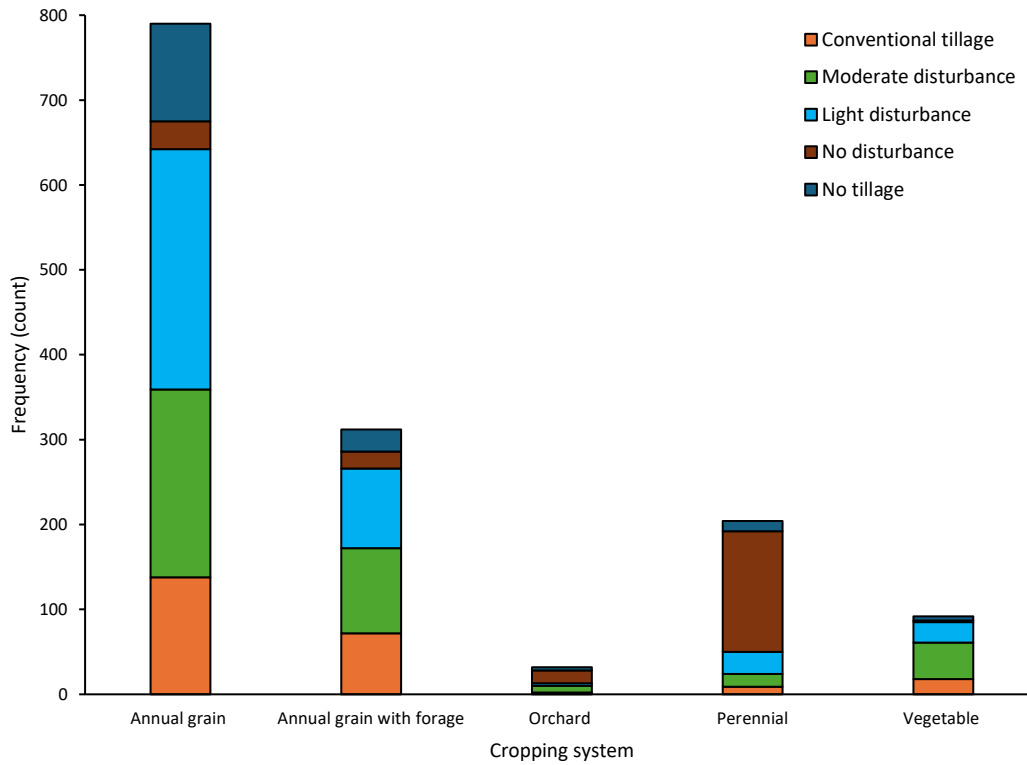


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669 **Figure 3** Pseudo van Krevelan diagram from programmed pyrolysis data for soils indicating cropping system
 670 category sampled in the Ontario Topsoil Sampling Project in 2019 (n=151).

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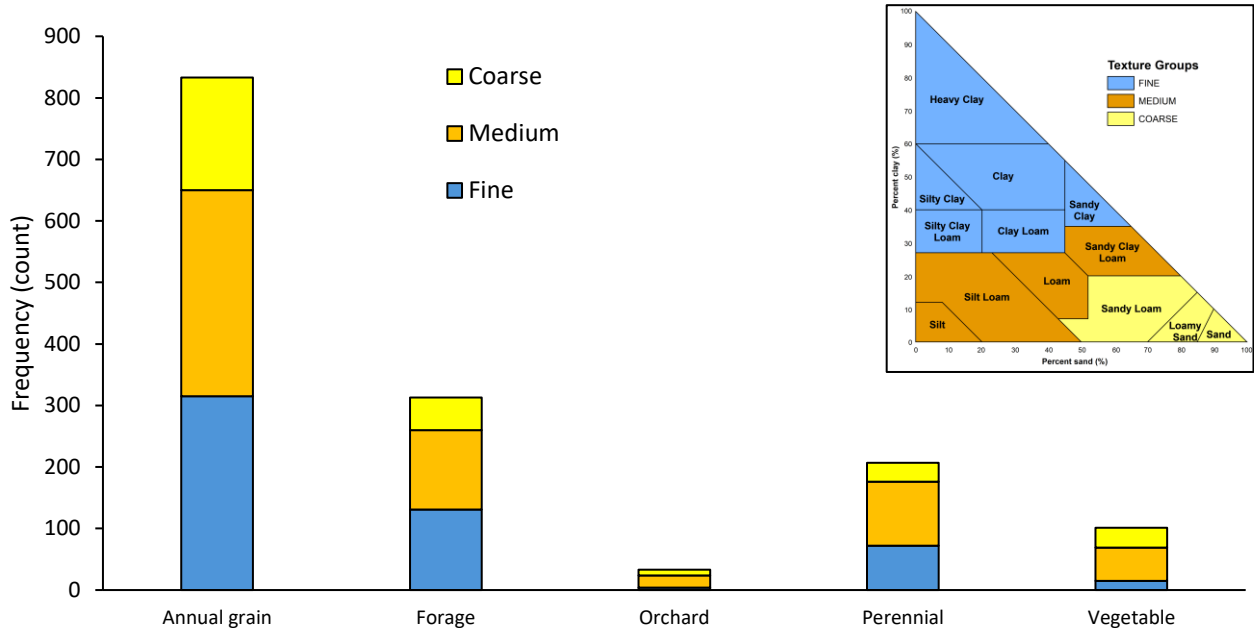


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674 **Figure 4** Frequency distribution of soils partitioned by tillage intensity within each cropping system category
675 sampled in the Ontario Topsoil Sampling Project from 2019 to 2022. Conventional tillage represents the plow tillage
676 in our study. No disturbance represented little to no soil movement and was associated mainly with pastures and
677 perennial forages.

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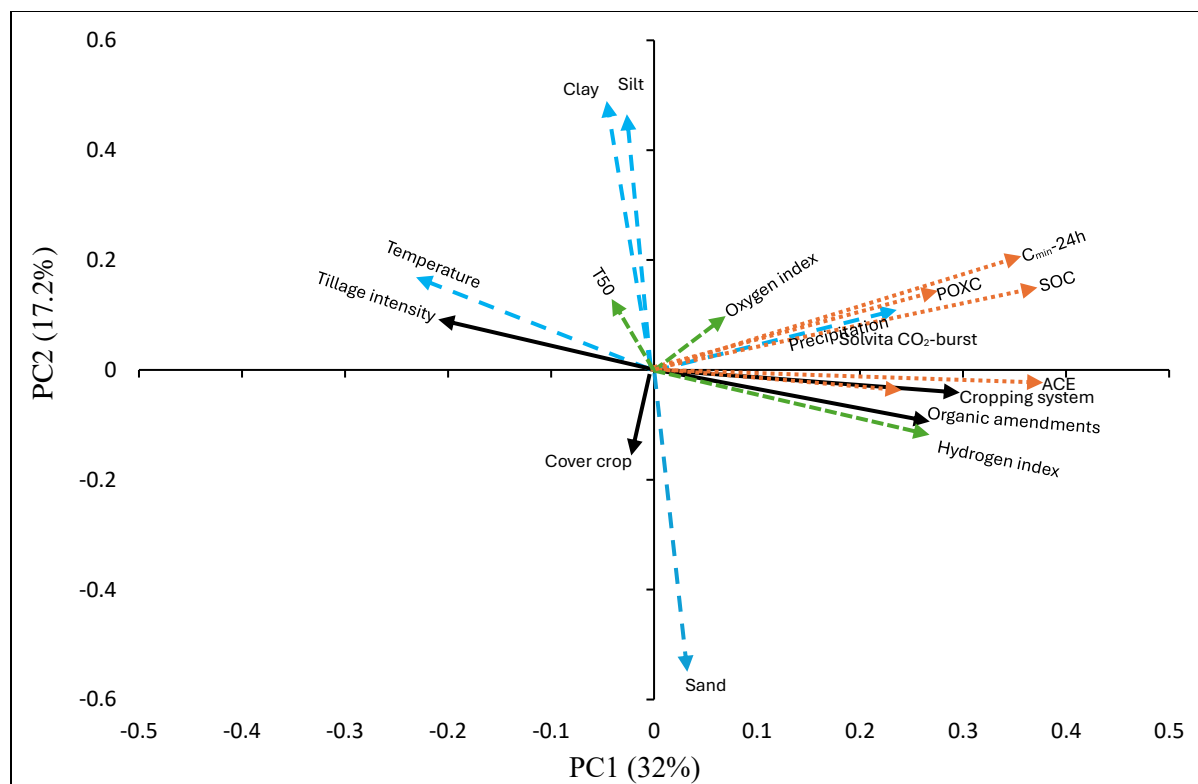
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682 **Figure 5** Frequency distribution of soils partitioned by soil textural classes within each cropping system category
683 sampled in the Ontario Topsoil Sampling Project from 2019 to 2022.

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687 **Figure 6** Principal component analysis (PCA) demonstrating relationships between site characteristics (blue dash-line
 688 vectors), management practices (black solid-line vectors), soil C indicators (orange dash-line vectors), and programmed
 689 pyrolysis parameters (green dash-line vectors) sampled in the Ontario Topsoil Sampling Project from 2019 (n=151).
 690 ACE=autoclaved citrate extractable protein index; C_{min}-96h= 96-h carbon mineralization; POXC=permanganate
 691 oxidizable carbon; SOC=soil organic carbon; and Solvita CO₂-burst= 24 h soil respiration test.
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