



1 Strong wind occurrence in Poland from the 13th to 16th centuries based on documentary evidence

2

3 Rajmund Przybylak^{1,4}, Andrzej Arażny^{1,4}, Janusz Filipiak², Piotr Oliński^{3,4}, Przemysław Wyszyński^{1,4},
4 Artur Szwaba⁵

5

6 ¹Nicolaus Copernicus University in Toruń, Faculty of Earth Sciences and Spatial Management, Poland

7 (rp11@umk.pl, andy@umk.pl, przemyslaw.wyszynski@umk.pl)

8 ²University of Gdansk, Department of Physical Oceanography and Climate Research, Gdańsk, Poland

9 (janusz.filipiak@ug.edu.pl)

10 ³Nicolaus Copernicus University in Toruń, Faculty of History, Poland (olinski@umk.pl)

11 ⁴Centre for Climate Change Research, Nicolaus Copernicus University, Toruń, Poland (cccr@umk.pl)

12 ⁵State Water Management, Polish Waters, Toruń, Poland (artur.szwaba@wody.gov.pl)

13

14 *Correspondence to:* Przemysław Wyszyński (przemyslaw.wyszynski@umk.com)

15 **Abstract.** A comprehensive database of strong winds based on documentary evidence was created for
16 Poland until AD 1600. Three types of documentary sources were used: handwritten and unpublished,
17 published, and “secondary” literature. The database contains detailed information about the
18 occurrence of strong winds (the location/region, time, duration and indexation for intensity, extent
19 and character of damage), as well as the exact textual content of the original weather note, the name
20 of the source, and an evaluation of the source’s quality. Five categories of strong winds were delimited:
21 1 – fresh and strong breeze (Beaufort scale 5–7), 2 – gale (8–9), 3 – storm (10–12), 4 – squall (i.e., gusty
22 wind during a thunderstorm), and 5 – tornadoes. The intensity, extent, and character of damage were
23 estimated based on the proposition given by Brázdil et al. (2004), which we slightly modified to include
24 the Baltic Sea and its influence on coastal parts. In the database, 137 thus-defined strong winds were
25 identified. A reliable estimate of some characteristics of the occurrence of strong winds in Poland
26 seems possible from the mid-15th century onwards. The highest number of strong winds occurred in
27 the second half of the 16th century, with a maximum in the 1570s. For each season, the greatest
28 number of strong winds was found for the Baltic Coast and Pomerania region, and then for Silesia and
29 Lesser Poland. Storms and gales were most common during the cold half-year (mainly in March,
30 November, and December).

31

32 **Keywords:** historical climatology, documentary evidence, Poland, strong winds, database.

33

34

35 1. Introduction

36 Strong winds are among the most significant natural disasters, causing great damage in the entire
37 world and loss of human and animal lives. In Poland, for example, according to Lorenc (2012), they are
38 second the most dangerous natural phenomenon after floods. Results presented recently by the
39 Statista Research Department (Apr 29 2024, [https://www.statista.com/statistics/1269886/most-](https://www.statista.com/statistics/1269886/most-common-natural-disasters-in-europe)
40 [common-natural-disasters-in-europe](https://www.statista.com/statistics/1269886/most-common-natural-disasters-in-europe)) confirm that the same situation is observed in Europe. From
41 2001 to 2020, floods were most frequent (41%), followed by strong winds (27%). According to
42 MunichRe (2011, 2020) estimates, ~60% of all insured losses during 2000–18 were due to extreme
43 meteorological events, primarily extreme winds. Cusack (2023, see Figs 2 and 7) estimated the annual
44 windstorm losses in Europe (12 countries) from 1950 to 2022. The results showed that yearly losses



45 usually (~80%) oscillated between 1 and 5 billion euros. Although no long-term trend is seen in the
46 study period, the greatest losses were observed in the 1980s.

47 Good and reliable knowledge about extreme winds is essential for many economic sectors,
48 e.g., the design and construction of large and high buildings or the wind power sector (Outten and
49 Sokolowski 2021). However, our knowledge about the different characteristics and impacts of extreme
50 winds is still based mainly on results gathered for the instrumental period (more or less for the last
51 100–150 years), and most often only for the last few decades. Therefore, such knowledge is still
52 insufficient and needs improvements, which can be done by using a more extended series of data
53 coming from the early- and pre-instrumental period. Many such data are still undiscovered or
54 discovered but not digitised and exist only on paper stored in numerous archives worldwide (Hawkins
55 et al. 2003). That is why data rescue activity is vital (for details, see e.g. Brönnimann et al. 2019;
56 Lundstad et al. 2023). It can help improve our understanding of historical climate variations, including
57 strong winds. For example, Hawkins et al. (2019) demonstrated that the severe windstorm that
58 occurred in February 1903 in England and Wales (reconstructed by them using documentary evidence)
59 was characterised in some places by stronger winds than were observed in the modern period (1950–
60 2015). They thus suggest that an estimate of risk from severe windstorms based on contemporary data
61 may need to be revised. A longer perspective on changes in intensity and impacts of extreme winds
62 can also be beneficial for their future simulations and also for the more reliable assessment of the risks
63 connected with them for societies. However, significant spatial-temporal changes of strong wind
64 occurrences and also their rarity and often local character significantly hinder the proper identification
65 of the mechanisms responsible for their changes and risks.

66 According to Donat et al. (2011), in future climate simulations (investigated using multi-model
67 simulations from global [GCM] and regional [RCM] climate models), enhanced extreme wind speeds
68 were found over northern parts of Central and Western Europe in most simulations and in the
69 ensemble mean (up to 5%). Consequently, they forecast that the potential losses will be higher in these
70 regions, particularly in Central Europe. In turn, in Southern Europe, according to them, an expected
71 decrease in extreme wind speeds will result in a reduction in loss potential. More recent work (Outten
72 and Sokolowski 2021) partly confirms the above findings. Using a 15-member ensemble of high-
73 resolution Euro-CORDEX simulations (~12 km), they found increases in the return period, i.e. more
74 frequent extreme episodes projected for Northern, Central and Southern Europe throughout the 21st
75 century. At the same time, they underlined, however, that the assessments of future extreme wind
76 changes remain fraught with uncertainty.

77 As seen from all the presented scenarios, strong winds in Poland will be more common. Thus,
78 associated economic and societal consequences may also be more significant than at present.
79 Therefore, the investigation of all the characteristics of strong winds and their impacts should be



80 intensified in Poland. Although we have quite a large number of works describing strong winds in
81 Poland in the contemporary period using instrumental measurements (e.g. Stopa-Boryczka 1989;
82 Paszyński and Niedźwiedz 1991; Krawczyk 1994; Adamczyk 1996; Lorenc 1996, 2012; Araźny et al.
83 2007; Tarnowska 2011; Ustrnul et al. 2014; Chojnacka-Oźga and Oźga 2018; Wibig 2021 and references
84 therein) they all cover some periods since 1950s and are written mainly in Polish and therefore are
85 unknown to the international scientific community. Only in a few works (e.g., Bartnicki 1930; Gumiński
86 1952; Piasecki 1952) is analysis of winds (including strong winds) available for an instrumental period
87 before 1950, starting from the late 19th century. On the other hand, there is a complete lack of such
88 works for historical periods, i.e. before 1800, using documentary evidence and existing visual regular
89 observations of winds. For example, the latter kind of observation with a quantitative estimation of
90 the force of the wind using a seven-degree scale (0–6) exists for Wrocław (Orig. Breslau, source:
91 newspaper *Oekonomische Nachrichten der Patriotischen Gesellschaft in Schlesien*) and a five-degree
92 scale (0–4) for Żagań (Orig. Saganenses, source: *Ephemerides Societatis Meteorologicae Palatinae*,
93 1783–1795, see also Przybylak et al. 2014 and Pappert et al. 2021) for the periods 1773–81 and 1781–
94 92, respectively. In addition, similar detailed observations of wind reporting quantitative force of wind
95 also exist for Gdańsk, where the first regular meteorological observations started in 1655 and were
96 carried out by Búthner, a professor of mathematics, who used to note daily observations of various
97 weather phenomena including the occurrence of strong winds. Unfortunately, his manuscript was lost,
98 probably irretrievably. It is possible to analyse only data for some selected years within his
99 observations covering the years 1655–1701. The 18th century (the Enlightenment) brought to the
100 coast a boom in interest in observing weather conditions and their impact on the economy. Its climax
101 was the beginning of regular instrumental meteorological measurements in Gdańsk by Hanov on the
102 1st of January 1739. Wind observations were made using a nine-degree scale (0–8) in 1739–72 (see
103 Table 5.1 in Przybylak 2010).

104 Knowledge about strong wind occurrences in Europe in the pre-instrumental period is also very
105 limited, although significantly better than in other parts of the world. For example, for the period under
106 study (13th–16th centuries), most of the works containing the most detailed climatic analyses
107 regarding strong winds are available mainly for the Czech Lands (e.g., Brázdil and Dobrovolný 2000,
108 2001; Dobrovolný and Brázdil 2003; Brázdil et al. 2004 and references therein). The last item is
109 particularly valuable for its very detailed analysis of different aspects related to strong winds for the
110 entire last millennium based on documentary evidence. However, even in this publication, some
111 limited information about strong winds for the pre-1500 period is available, directly caused by the
112 small number of existing historical sources. For central Europe, information about strong winds is also
113 contained in monographs analysing different kinds of extremes (Pfister 1999; Glaser 2001, 2013). For
114 the Low Countries (the coastal areas of the southwest Netherlands and Flanders), a valuable paper is



115 available presenting storminess changes in the period 1390–1725 (de Kraker 2013 and references
116 therein). It is essential to add that the author also graded storm events using an eight-degree scale.
117 Similar work, as mentioned for the Low Countries, also exists for the North Sea, the British Isles, and
118 Northwest Europe (Lamb 1991 and references therein). For a smaller area in this European region
119 (Thames estuary), there are also works published by Galloway and Potts (2007) and Galloway (2009).
120 Finally, we should also mention the work of Orme (2015), which analysed late-Holocene storminess in
121 Europe using various proxies.

122 This short review shows that our knowledge about the occurrence of strong winds in Europe,
123 although better than in other parts of the world, is very limited. That is why there is an urgent need to
124 improve and widen this knowledge. The main aim of the present paper is to partly fill this gap by
125 presenting an analysis of strong wind occurrences in Poland for the period from the late 13th century
126 (the first record of strong wind in Poland found in historical sources) to the end of the 16th century.

127

128 2. Area, data and methods

129 2.1. Area

130 The analysis of strong winds in the studied period is conducted for the area of Poland within
131 contemporary boundaries. Poland is a Central European country stretching from the Baltic Sea in the
132 north and the Sudetes and Carpathian Mountains in the south (Fig. 1). To more precisely estimate the
133 spatial changes in strong winds occurrence, the analysis was also made for six historical-geographical
134 regions: Baltic Coast and Pomerania, Masuria and Podlasie, Greater Poland, Masovia, Silesia, and
135 Lesser Poland (Fig. 1). Data from contemporary period representing all the mentioned regions were
136 collected for 12 meteorological stations – two for each region.



137

138 Fig. 1. Geographical location of Poland, main historical-geographical regions and contemporary meteorological
139 stations (red dots) (after Ghazi et al. 2024, modified)



140 At present (1966–2018), the average annual wind speed calculated based on 41 stations was
141 3.6 ms^{-1} (Wibig 2021). Excluding mountainous areas, for which we have no historical data, the largest
142 average wind speeds in Poland are noted in the coastal part of the Baltic Sea ($4\text{--}5 \text{ ms}^{-1}$) and a little
143 smaller in the central part, e.g. in Warszawa (Eng. Warsaw) 4.0 ms^{-1} . The weakest winds in Poland are
144 noted in the foothills of the Sudeten and Carpathian Mountains (less than 3 ms^{-1}). The spatial
145 distribution of extreme winds is slightly different compared to the average ones. The greatest extreme
146 winds occur in the SW part of Poland and then in the coastal part of the Baltic Sea (Wibig 2021).

147 2.2. Sources and data

148 The following three types of documentary sources were used: handwritten and unpublished,
149 published, and “secondary” literature (e.g., articles, monographs) to search for weather notes
150 describing the occurrence of strong winds in Poland. The number of used historical sources in the study
151 period correlates strongly with their availability, which is the greatest in the less-distant centuries. For
152 example, for the 16th century, we used 85 sources, while for the 13th and 14th centuries, we used 1
153 and 27 sources, respectively. For every event of strong wind occurrence in Poland, a detailed reference
154 to the source(s) is given; see <https://doi.org/10.18150/W6PMBQ>.

155 A quality assessment of each used source was conducted by historians using the method called
156 “source criticism” in the historical sciences. This method allows for examining the authenticity of
157 specific human activities (historical source) and reading their meaning (historical event) in the light of
158 the causes and conditions of their creation in the historical process. Three quality categories were
159 distinguished: 1 – weak, 2 – moderate, and 3 – high, to choose an appropriate source and weather
160 note(s) describing strong winds. The following rules were used to stratify sources according to their
161 quality; 1 – weak, if the information was derived from secondary literature rather than the original
162 source; 2 – moderate, if the information was written centuries after the strong wind occurrence; and
163 3 – high, if the information was written in a source in the same period that the strong wind event
164 occurred and provides precise information. For the analysis, we used mainly the last category of
165 sources, which provided the best valuable data.

166

167 2.2.1. Database: historical period

168 For the first time, a comprehensive database of strong winds related to the period before the 19th
169 century is prepared for Poland by a team of climatologists and historians based on all documentary
170 evidence. The database at the present stage is finished until the end of the 16th century (see
171 <https://doi.org/10.18150/W6PMBQ>). It contains detailed information about the occurrence of strong
172 winds (the location/region, time, duration, and indexation for intensity, extent, and character of



173 damage), as well as the exact textual content of the original weather note, the name of the source,
174 and an evaluation of the source's quality. The information is not complete for every case of strong
175 wind. Sometimes, only general information is available about the strong winds in Poland. In such a
176 case, there is no information about the place or region; therefore, we introduced an additional region
177 category called "Poland". Information available for the time of the strong wind occurrences is also
178 varied. For some, we only have information about the year; for others, we only have information about
179 both season and year. However, in most cases, we have information about year, month and day(s)
180 (and often the start and end of the phenomenon). For this reason, it is essential to remember that
181 presented statistics, e.g. frequency of occurrence of strong winds in months, seasons and years, are
182 based on different numbers of cases. In Table 1 we showed some selected examples of entries to the
183 database presenting varying degrees of detail.

184

185

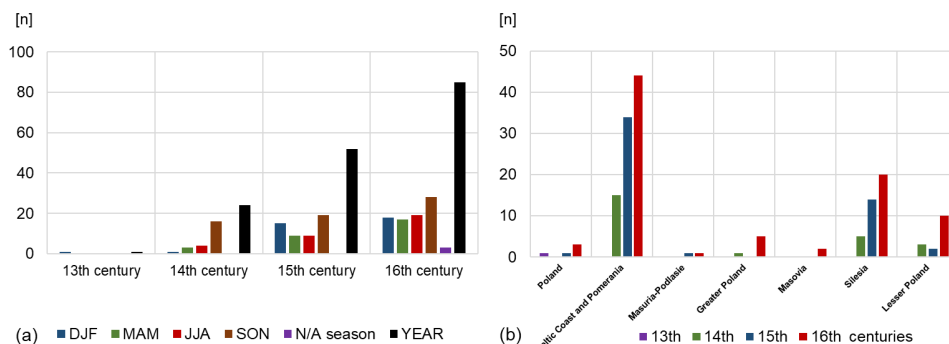


Table 1. Examples of entries in the strong wind database. Explanation of abbreviations (A, B, C, D) below the table. Explanations of numbers and abbreviations in columns A, B, C, and D are provided in the database.

| Region | Place | Date of occurrence | Description: original | Description: translation | Source | A | B | C | D |
|--|----------------------------|--------------------|--|--|---|---|---|---|----------------------------|
| A complete record, including details of the exact date and place of occurrence of the phenomenon and associated damage | Baltic Coast and Pomerania | 11–13 Jan 1558 | 1558 Am Dienstag nach Maria Lichtmess (8. Februar) richtete der Wind in Stralsund grossen Schaden an den Brücken der Stadt an, die alle bis auf eine entzwei brachen. Auch Schiffe und Botte wurden zertrümmert. Es war Nordwestwind, und zwar in der Nacht, sonst wären auch wohl Menschen ertrunken. Auch auf dem Lande wurde den Häusern und Obstbäumen viel Schaden zugefügt. [...] Auch in Rügenwalde richteten zwei grosse Stürme, die vom 11. bis 13. Januar und am 8. Februar wütheten, viele Verheerungen an. Das Wasser soll ellenhoch an der Stadtmauer gestanden haben, in die Speicher bei der Wipper gedungen sein und auf de Münde 18 Wohngebäude zerstört haben, deren Einwohner sich nur dadurch retteten, dass sie auf Bäume kletterten. | 1558 On the Tuesday after Candlemas (February 8th), the wind in Stralsund caused great damage to the city's bridges, all except one broke in two. Ships and boats were also wrecked. It was a north-westerly wind, and at night, otherwise people would probably have drowned. Much damage was also done to houses and fruit trees in the countryside. [...] Two big storms that raged from January 11th to 13th and February 8th also caused a lot of devastation in Rügenwalde. The water is said to have risen several metres above the city walls, infiltrated the granaries near the Wipper and destroyed 18 residential buildings on the estuary, whose residents only saved themselves by climbing trees. | Besch R., Strenge Winter in alter Zeit. Witterungsgeschichtliches aus Pommern, Unser Pommerland, Jg. 8 (1923), H. 1, p. 13. | 1 | 3 | 2 | DB DS DU DV DF |
| Inaccurate record – only fragmentary information on date of occurrence and associated damage | Sllesia | autumn 1578 | Ein heftiger Wind beschaedigte im Spaetherbst den obern Theil des Rathhauses, das schon durch fruehre Stuerme und Gewitter gelitten. | In late autumn, a strong wind damaged the upper part of the town hall, which had already suffered from early storms and thunderstorms. | Minsberg F., Geschichtliche Darstellung 1 der merkwuerdigsten Ereignisse in der Fuerstenthums Stadt Neisse, Neisse | 1 | 1 | 1 | DS |
| General record - very fragmented information on the phenomenon | Baltic Coast and Pomerania | 1531 | Anno 1531 — War ein grosser Sturm Wind, dass der Wind den grossen Wetter Hahn vom Kirchen Thurm abwarf und haben einige dieses vor eine Vorbedeutung gehalten, dass die Catholischen sollten untergehen, so auch kurzze Zeit darauf richtig erfolget, nemlich anno 1534, da das Pabsthum im ganzen Lande abgethan worden. | Anno 1531 — There was a great storm wind that threw the great clock from the church tower, and some thought that this was an omen that the Catholics would perish, which happened a short time later, namely in 1534, when the papacy was abolished in the entire country. | Wendland J.D., Eine Sammlung unterschiedlicher die Historia der Stadt Cöslin betreffende Sachen | 3 | 3 | 2 | DB |



5 The entire database contains 137 records documenting the occurrence of strong winds in
 6 Poland. The first weather note reporting the presence of strong wind was found for the year 1283. The
 7 number of weather notes (162) is not identical to the number of strong wind cases because sometimes
 8 we have more than one weather note describing the same case of strong wind. The statistics of
 9 weather notes documenting the occurrence of strong winds are presented in Fig. 2. The greatest
 10 number of weather notes we found for the 16th (85 cases) and 15th (52) centuries and less for the 13th
 11 century (1). Most of them we found for autumn (63) and winter (35). The majority of weather notes
 12 exist in the documentary evidence for two regions, Baltic Coast and Pomerania (93) and Silesia (39),
 13 and least for the Masuria-Podlasie and Masovia regions, with 1 and 2, respectively (Fig. 2).



14 (a) ■ DJF ■ MAM ■ JJA ■ SON ■ N/A season ■ YEAR (b) ■ 13th ■ 14th ■ 15th ■ 16th centuries
 15 Fig. 2. Number (n) of weather notes for Poland for seasons (a) and regions (b), 1281–1600

16 There is strong coherence between the number of available sources and the number of
 17 weather notes (usually the number of weather notes is greater than the number of sources) on the
 18 one hand and the number of occurrences of strong winds on the other. This pattern is similar to what
 19 we observed for flood records (Ghazi et al. 2023a, b, 2024, 2025) and drought records (Przybylak et al.
 20 2020).
 21

22 2.2.2. Database: contemporary period

23 Sub-daily data (every three hours) of wind speed for the period 1993–2022, for which most
 24 homogeneous and complete series of wind values for Poland are available, were gathered for 12
 25 meteorological stations (see Table 2, Fig. 1). As mentioned earlier, they also represent all six
 26 distinguished historical-geographical regions in the country. The wind data were downloaded from the
 27 website of the Institute of Meteorology and Water Management - National Research Institute (IMGW-
 28 PIB) (<https://danepubliczne.imgw.pl/>). The two data types were collected as average wind speed every
 29 three hours and the highest gust of wind for 3-hour intervals.
 30

31



32

Table 2. Geographical location of meteorological stations used in the work

| No. | Name of meteorological station | H (m a.s.l.) | φ (N) | λ (E) |
|-----|--------------------------------|--------------|---------------|---------------|
| 1. | Świnoujście | 6 | 53°55' | 14°14' |
| 2. | Chojnice | 164 | 53°43' | 17°33' |
| 3. | Olsztyn | 133 | 53°46' | 20°25' |
| 4. | Suwałki | 184 | 54°08' | 22°57' |
| 5. | Poznań | 83 | 52°25' | 16°51' |
| 6. | Kalisz | 138 | 51°47' | 18°05' |
| 7. | Warszawa | 106 | 52°10' | 20°58' |
| 8. | Siedlce | 152 | 52°11' | 22°15' |
| 9. | Wrocław | 120 | 51°06' | 16°53' |
| 10. | Opole | 165 | 50°38' | 17°58' |
| 11. | Kraków | 237 | 50°05' | 19°48' |
| 12. | Rzeszów | 200 | 50°06' | 22°03' |

33

34

2.3. Methods

35

36

2.3.1. Historical period

37

In our analysis, we distinguished four categories of strong winds, which are consistent with the proposition used by Brázdil et al. (2004, their types T1, T3–T5 in Table 6.1) for Czech Lands. In addition, we added one more category (tornadoes), treated separately:

38

39

40

a) Fresh and strong breeze (force according to Beaufort scale [BS] 5–7),

41

b) Gale (BS 8–9),

42

c) Storm (BS 10 and more),

43

d) Squall (i.e., gusty wind during a thunderstorm),

44

e) Tornado.

45

Using the final version of the database, each case of strong wind was analysed in detail. In the first stage, an indexation of its intensity was done by the author of a particular record contributed to the database. In the second stage, the proposed categories of wind intensities by individual contributors were discussed and finally accepted by the whole team of authors of this paper. To investigate the damage caused by strong winds, three categories of extent of damage were utilised (consistent with types E0–E2 in Brázdil et al. 2004, Table 6.2, modified):

46

47

48

49

50

51

a) E0 – no information about damage,

52

b) E1 – small damage, damage of lesser extent,

53

c) E2 – large damage, areally extensive damage.

54

The last category was slightly modified by us to include damages on the sea (destruction or sinking of ships) and losses caused by storm floods – reflecting the coastal location of Poland. Finally, if weather

55



56 notes allowed, we also estimated the damaged character, again using the proposition of Brázdil et al.
57 (2004, Table 3).

58 Similarly as in the case of the extent of damage, we added two more categories to the list of
59 types of the character of damage presented for the Czech Lands. The proposed new categories
60 precisely describe the influence of the Baltic Sea on coastal parts. The “N/A” designation was used
61 when the weather notes did not contain information about the character of damage. Thus, the
62 following nine categories were distinguished to characterise damages and losses:

63

64 DO – casualties (lost lives),

65 DL – wind damage in forests,

66 DP – minor damage to buildings,

67 DB – considerable damage or destruction of buildings,

68 DS – uprooted fruit trees, damage to hop gardens and vineyards,

69 DU – damage to field crops, gardens and orchard harvests,

70 DV – considerable damage/destruction to vessels (including sunk),
71 newly added

72 DJ – other damage (e.g., upturned carriages, vessels, injured persons, minor damage to
73 property),

74 DF – considerable damage/destruction by storm flood or inland flood, newly added,

75 N/A – information not available.

76

77 2.3.2. Contemporary period

78 Based on gathered sub-daily wind data, the following statistics were calculated:

79 1. Average daily, monthly, seasonal, and annual wind speed,

80 2. Highest gust of wind for every day, month, season, and year,

81 3. Frequency of gust winds $> 17.0 \text{ ms}^{-1}$ calculated for every month and year for each station,

82 4. Frequency of gust winds in the following speed intervals:

83 a) $8.0\text{--}17.1 \text{ ms}^{-1}$ (BS 5-7) – quite strong, strong, and very strong wind,

84 b) $17.2\text{--}24.4 \text{ ms}^{-1}$ (BS 8-9) – gale,

85 c) $>24.4 \text{ ms}^{-1}$ (BS 10-12) – very strong and violent storms, and hurricanes.

86 These three categories of strong winds listed in point 4 above, but particularly the categories
87 described in points b and c, were used for comparison with the historical data. The analysis, however,
88 is limited mainly to estimating differences and similarities in the study area's annual course and spatial
89 distribution. We need to add that a reliable comparison of absolute values of the frequency of strong



90 wind occurrences in both periods is impossible due to the undetectability of probably a significant
91 number of events in the study's historical period, particularly those classified to the first category.

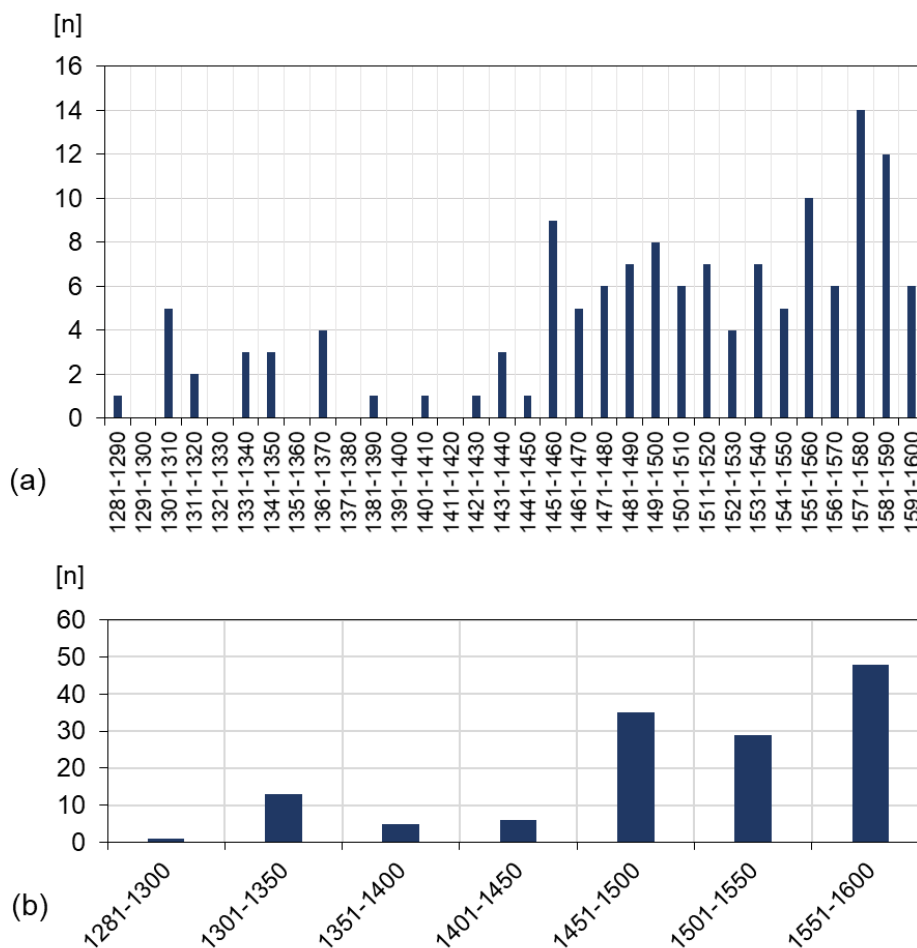
92 For many years the IMGW-PIB did not maintain a separate database on squalls, land- and
93 waterspouts and tornadoes, and the damage caused by the occurrence of such extreme events. In fact,
94 this was a typical measure for many European national meteorological services – the duty to monitor
95 damage caused by the aforementioned phenomena belonged to other state services. However, in
96 2015, the WMO decided to standardize weather, water, and climate hazard information to allow more
97 sophisticated analyses of data on the occurrence of such phenomena and the losses and damage
98 associated with them. Then, many meteorological services, including IMGW-PIB, started to catalogue
99 extreme weather, climate, and water phenomena and develop a dedicated database, but the material
100 collected so far is too sparse for us to make a comprehensive analysis.

101

102 3. Results

103 3.1. Historical

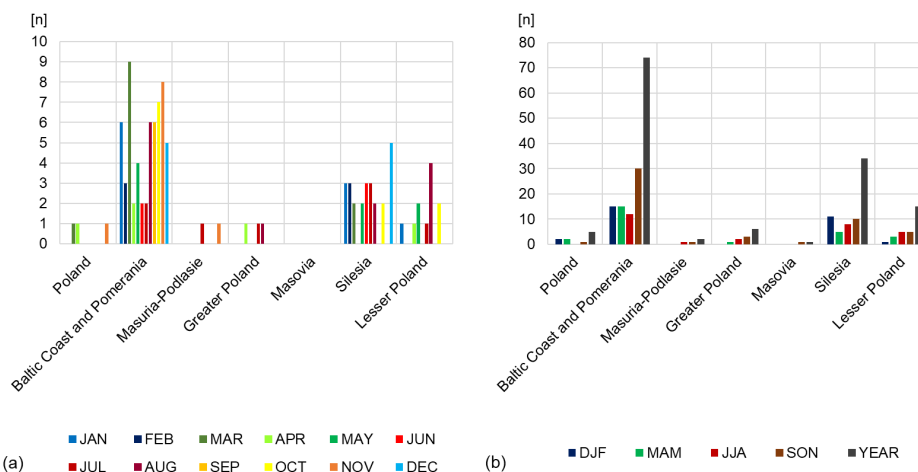
104 According to Przybylak et al. (2023), in Poland, the Medieval Warm Period (MWP, recently also called
105 the “Medieval Climate Anomaly”, MCA) started in the late 12th century and finished between the mid-
106 14th and mid-15th centuries. Then, until the mid-16th century, the Transitional Period (TP) was
107 distinguished (Niedźwiedź et al. 2015), followed by the Little Ice Age (LIA). Thus, our study period
108 covers a large part of the MWP, the entire TP and the early decades of the LIA. As results from analysis
109 conducted by Przybylak et al. (2023) and also from the database accompanied by this paper, the
110 available number of weather notes allow for a reliable estimate of the occurrence of strong winds,
111 mainly from the 1450s onward. Two maxima of greatest storminess can be distinguished in this time:
112 1451–1520 and 1551–1600 (Fig. 3). The latter, however, reveals a greater number of occurrences of
113 strong winds, in particular in two decades: 1571–80 (14 cases) and 1581–90 (12). In the pre-1450
114 period, we should underline the clear maximum of strong winds noted in the first half of the 14th
115 century, followed by a less stormy period until the 1450s.



116

117 Fig. 3. Decadal (a) and 50-year (b) number (n) of occurrence of all categories of strong winds in Poland, 1281–
118 1600

119 Out of the six regions analysed, strong winds were noted most often in the Baltic Coast and
120 Pomerania region (74 cases), and then in the Silesia region (34) (Fig. 4a, b). Quite often, they were
121 registered also in the Lesser Poland region (15). In other regions their occurrence was sporadic, ranging
122 from 1 to 6 cases (Fig. 4b). The annual cycle of strong winds occurrence can be estimated based on
123 regions for which enough information exists. Strong winds were noted most often in autumn and in
124 winter for Silesia and the Baltic Coast and Pomerania regions, and in summer and fall for the Lesser
125 Poland region (Fig. 4b). August (13 cases) and March (12) were most abundant in strong wind
126 occurrences in Poland, but the stormiest period of the year was from October to March, with at least
127 ten cases in each month except February (Fig. 4a).

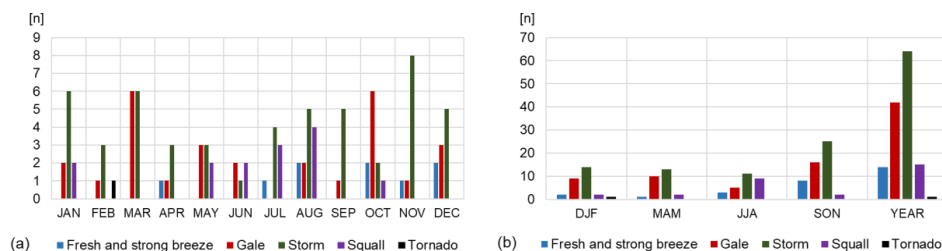


128

129 Fig. 4. Monthly (a) and seasonal (b) number (n) of all categories of strong winds according to regions, 1281–
 130 1600. Explanation: Please note that the total number of strong winds in seasons/years presented in Fig. b is greater than calculated based
 131 only on monthly statistics because, in some weather notes, there is information only about the season or even the year of the strong wind
 132 occurrence (see also text in 2.2.1. and Table 1).

133

134 In the entire study period (1281–1600) (Fig. 5), and also in two subperiods (1281–1500 and
 135 1501–1600) (Fig. S1), the most frequent were storms and gales and the least frequent were fresh and
 136 strong breezes and, in particular, tornadoes (only one case). The storms and gales were most common
 137 in spring (mainly in March), autumn (particularly in November), and winter (mainly in December) (Fig.
 138 5). It is worth noting a big change between the studied two sub-periods in August. In the period 1281–
 139 1500, in this month only storms were registered, whereas in the 16th century, all categories of strong
 140 winds were recorded (except tornadoes) (Fig. S1).



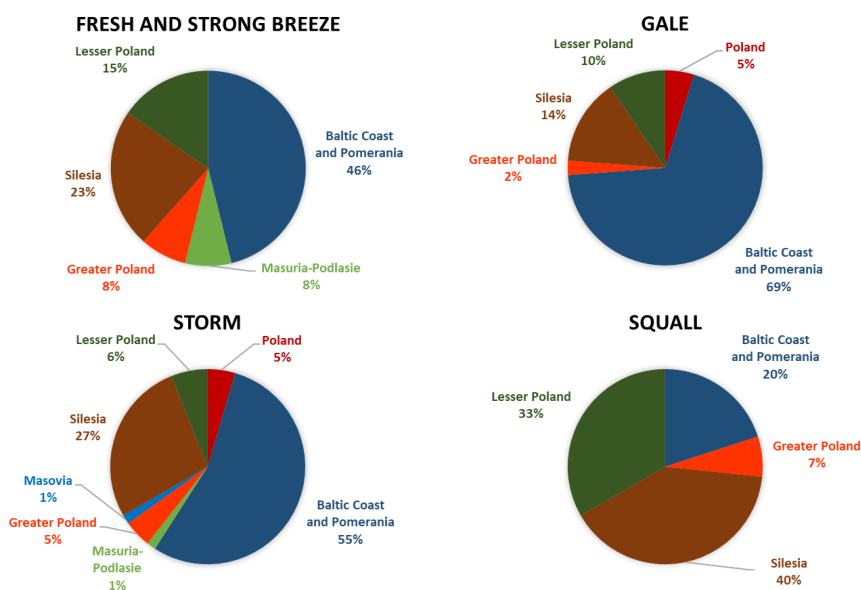
141

142 Fig. 5. Monthly (a) and seasonal (b) number (n) of occurrence of different types of strong winds in Poland in the
 143 period 1281–1600

144 The spatial distribution of occurrence of strong winds (except tornadoes) is presented in Fig. 6
 145 for the entire study period and in Fig. S2 for two sub-periods. The most significant differences in spatial
 146 distribution between the two sub-periods are seen for two categories of strong winds, i.e. fresh and



147 strong breezes and (particularly) for squalls (Fig. S2). In the latter case, the squalls were noted in the
 148 historical sources in the period 1281–1500 only for three regions: Silesia, the Baltic Coast and
 149 Pomerania, and Lesser Poland. In the 16th century, they were also found in Greater Poland. On
 150 average, for the entire study period, all categories of strong winds were most frequent in the Baltic
 151 Coast and Pomerania region (about 46–69%), except the squall category, which was noted most often
 152 in Silesia (40%). The region with the second greatest amount of strong winds (except squalls) was
 153 Silesia (14–27%) (Fig. 6).



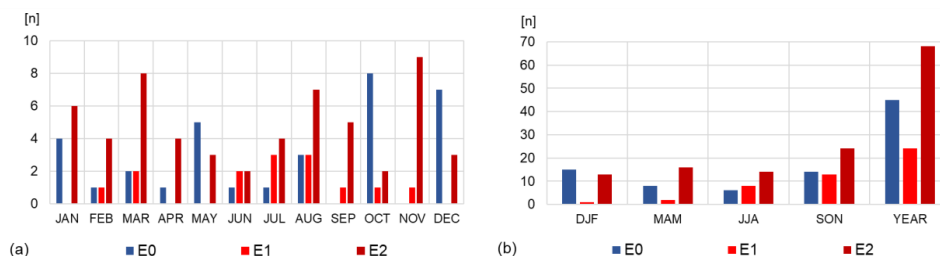
154

155 Fig. 6. Relative frequencies (%) of occurrence of different types of strong winds (tornadoes excluded) in the
 156 studied regions of Poland, 1281–1600

157 The first category of strong winds (category 1) we proposed does not cause damage. The wind
 158 is chaotic in its flow, causing the moving of large tree branches and whole trees; it makes the use of
 159 umbrellas and walking against the wind difficult and causes snowstorms and blizzards during snowfall.
 160 According to Lorenc (2012), a gust wind speed above 17 ms^{-1} in climate conditions in Poland creates a
 161 threat to the population, economy, and environment. It means that all gales and storm winds
 162 (categories 2 and 3) classified by us based on documentary evidence have the potential to cause
 163 various kinds of damage, which we listed in the method section. However, when chroniclers mention
 164 these winds in historical materials, they rarely mention the devastating effects of their impact on the
 165 environment or humans. The lack of information about the destruction probably relates to the
 166 immediate area of their residence and does not rule out possible destruction in other locations. In the
 167 case of squall winds, due to their locally limited character and when there is no information about



168 damage, we cannot say the same, as in the case of gales and storms that are the effect of
 169 macrocirculation. The available weather notes describing the occurrence of strong winds allow us to
 170 classify damage according to its scale, small or large (see Fig. 7 and Fig. S3).

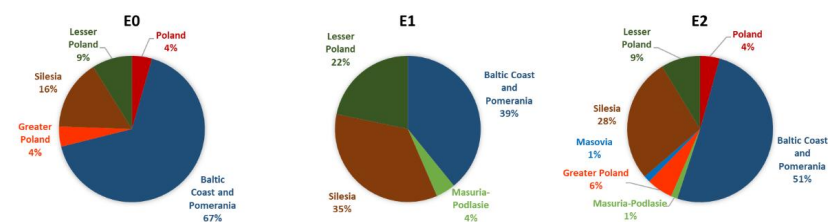


171

172 Fig. 7. Monthly (a) and seasonal (b) number (n) of occurrence of strong winds in Poland, according to the extent
 173 of damage, 1281–1600

174 It is interesting to note that, for the study period, category E2 of damage is most frequent,
 175 then category E0, i.e. no information about damage. Category E2 prevails in all seasons (excluding
 176 winter) and in most of the months, except December, October, June and May (Fig. 7). It is also worth
 177 noting that the frequency of category E0 in relation to E2 decreases from medieval times to the 16th
 178 century (see Fig. S3). This means that the 16th-century chroniclers were inclined to describe weather
 179 events, in this case, strong winds, and their effects more precisely than their predecessors.

180 In the entire dataset, the largest share of strong winds causing damage (categories E1 and E2)
 181 and in which there is no information on damage (category E0) was found, in line with expectations, for
 182 the Baltic Coast and Pomerania (39–67%) and Silesia (16–35%), for which the most abundant datasets
 183 exist (Fig. 8). A similar situation is noted for both sub-periods with one only exception being the
 184 category E1 in the period 1281–1500 (Fig. S4). This category was a little more frequently reported for
 185 Lesser Poland (29%) and Silesia (29%) than the Baltic Coast and Pomerania (28%) region.

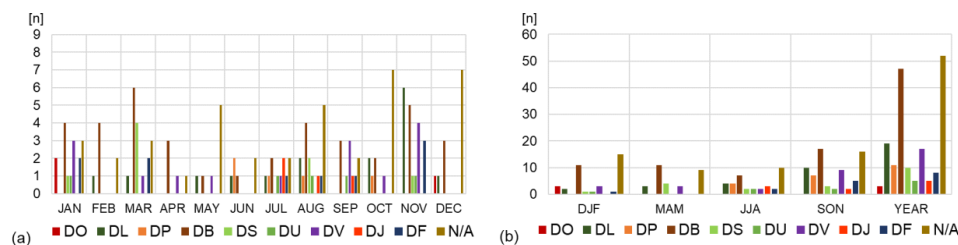


186

187

188 Fig. 8. Relative frequencies (%) of occurrence of strong winds in studied regions of Poland for different types of
 189 damage, 1281–1600. For explanations of abbreviations, see Methods section.

190 The character of damage caused by the strong winds is shown in Fig. 9 for the entire study period and
 191 in Fig. S5 for the two analysed sub-periods. The two figures are roughly similar.

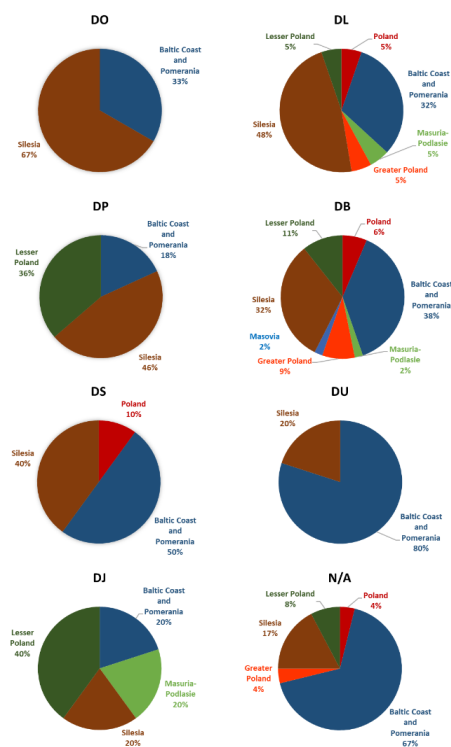


192

193 Fig. 9. Monthly (a), seasonal and annual (b) number (n) of strong winds in Poland for which
 194 information about the character of damage exists (DO, DL, DP, DB, DS., DU, DV, DJ and DF) or does not exist
 195 (N/A), 1281–1600. For explanations of abbreviations, see the Methods section.

196 On average, 28.7% of weather notes did not contain information about the damage caused by the
 197 strong winds (Fig. 9). Fewer such cases occurred in the 16th century (24.5%) than in medieval times
 198 (34.2%) (Fig. S5). Out of all distinguished damage categories, DB was the most frequent (26.0%), with
 199 a maximum in the 16th century (28.4%). The following most frequent categories of damage noted in
 200 the study period were DL (10.5%) and DV (9.4%), while the least frequent was DO (1.7%) (Fig. 9). About
 201 60% of all damages mentioned in the weather notes were found for the cold half-year, but particularly
 202 for autumn (39.2%). The smallest number of weather notes about damages was noted for spring
 203 (16.6%). In the annual course, damages caused by strong winds in Poland occurred most often in
 204 November (20 cases, 14.0%), March and August (17, 11.9%), with the smallest frequency in April (5,
 205 3.5%) (Fig. 9a).

206 The spatial distribution of distinguished categories of damage made by strong winds in Poland
 207 in analysed regions is shown in Fig. 10 for 1281–1600 and in Fig. S6 for two sub-periods: 1281–1500
 208 and 1501–1600. The analysis reveals that the most frequently noted categories of damage in Poland
 209 (DB, DL) were also present in all analysed regions, with a maximum in Silesia and Baltic Coast and
 210 Pomerania regions. For these two regions, the number of information items about other categories of
 211 damages caused by strong winds was also the highest, except for the DJ category, which was noted
 212 most often for the Lesser Poland region (Fig. 10). The greatest number of weather notes describing



213

214 Fig. 10. The relative frequencies (%) of damage categories estimated for particular regions in Poland in the period
 215 1281–1600. Explanations: Two categories (DV and DF) are not shown because they can occur only in the Baltic Coast and Pomerania
 216 region. For explanations of abbreviations, see the Methods section.

217 strong winds not informing about damages was found for the Baltic Coast and Pomerania region (67%)
 218 and then for the Silesia region (17%). The spatial distribution of damages caused by strong winds
 219 presented for the entire study period was more similar to that noted in the 16th century than that in
 220 the medieval period (cf. Figs 10 and S6).

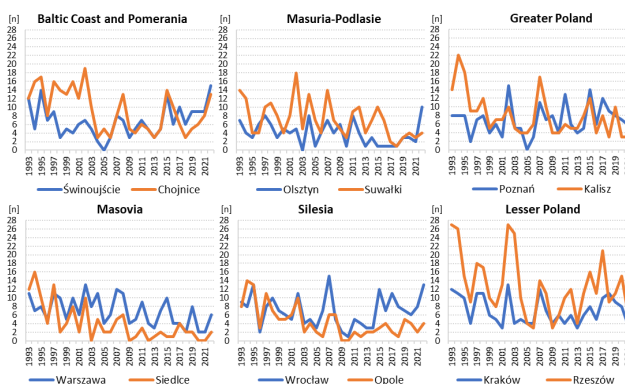
221 3.2. Contemporary period

222 A short analysis of the present occurrence of strong winds in Poland can be helpful for some
 223 comparison purposes with historical results, mainly spatial distribution and run of the frequency of
 224 strong winds in the annual cycle. Also, helpful knowledge can be about the values of the greatest speed
 225 of strong winds observed presently in Poland.

226 The annual number of gust winds $>17 \text{ ms}^{-1}$ in Poland in 1993–2022, which can potentially cause
 227 damage, usually does not exceed 20 cases (Fig. 11), but 30-year average values in all stations are <10 ,
 228 except Rzeszów. The probability of occurrence of that kind of strong wind was the smallest in Masovia
 229 and Silesia (fewer than 16 cases). There are significant fluctuations in the annual number of gust winds

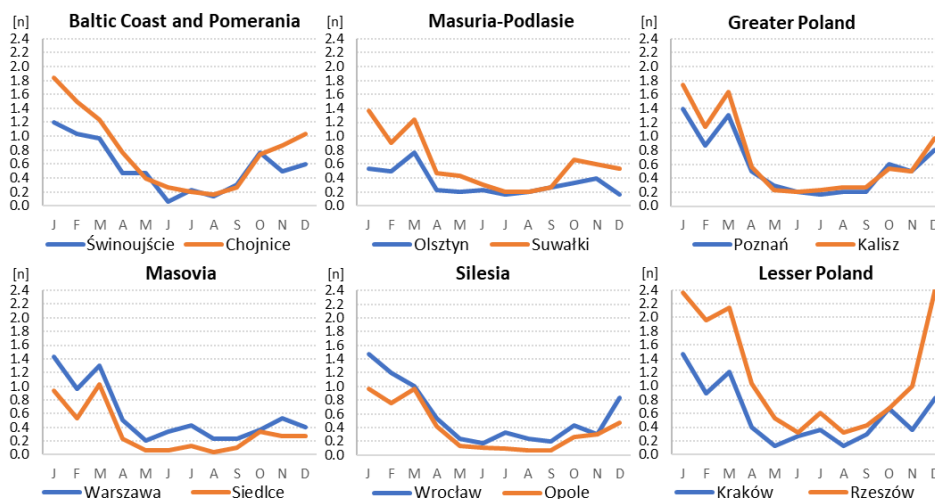


230 from year to year, sometimes exceeding 10 cases. Also, the number of thus-defined strong winds has
 231 decreased since about 2010, particularly in the Masovia and Silesia regions.
 232



233
 234 Fig. 11. Year-to-year course of the annual number (n) of gust winds > 17 ms⁻¹ in Poland, 1993–2022

235 In the annual cycle, the average 30-year (1993–2022) monthly number of strong winds >17
 236 ms⁻¹ oscillated from below 0.4 in the warm half-year to more than 0.8–1.0 in Jan–Mar (Fig. 12). The
 237 first half of the cold season (Oct–Dec) has significantly fewer gust winds, rarely exceeding 0.8 cases in
 238 one month, except December in some stations.

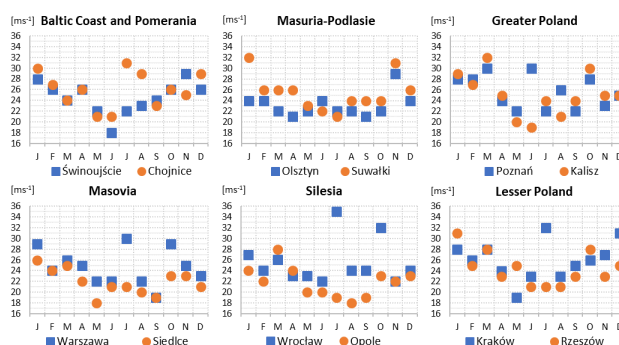


239
 240 Fig. 12. Annual course based on the average monthly number (n) of gust winds >17 ms⁻¹ in Poland, 1993–
 241 2022

242 In the studied period, the highest gust wind speed in the majority of analysed stations
 243 exceeded 30 ms⁻¹. It occurred most often in winter or summer months (Fig. 13). On the other hand,
 244 the lowest, oscillating between 16 and 22 ms⁻¹, were measured only in the period from May to



245 September. The highest strong wind speed (35 ms^{-1}) in Poland was measured in Wrocław (SW Poland)
 246 on July 23, 2017, while the lowest (18 ms^{-1}) was recorded in a few places (Świnoujście – June 18, 2012
 247 and June 28, 2014; Siedlce – May 4, 1996 and May 3, 1997; Opole – August 28, 1994 and August 19,
 248 2022). Gust winds $>17 \text{ ms}^{-1}$ stratified into gales (8–9 BS) and storms (10 BS and more) do not show any
 249 important changes in the annual courses compared to annual courses based on all cases of strong
 250 winds $>17 \text{ ms}^{-1}$ (see Fig. 12 and Fig. S7). The main reason for this is the rare occurrence of storms in
 251 Poland. In the study period, storms were not observed in most of the years, and the maximum of them
 252 in one year reached three cases (not shown).



253

254 Fig. 13. Maximum monthly wind gust speeds (ms^{-1}) at selected stations in Poland in 1993–2022

255

256

257 4. Discussion

258 Five categories of strong winds were distinguished in our database
 259 (<https://doi.org/10.18150/W6PMBQ>, see also Methods) and analysed in the present paper. The first
 260 three categories of strong winds (fresh and strong breezes, gales, and storms) in central Europe are
 261 connected with vigorous moving cyclones born near Iceland, and in particular with a passing cold front.
 262 Cyclonic activity is the greatest in the cold half-year (October to March), and therefore, these
 263 categories of strong winds dominate in this part of the year. On the other hand, the last two categories
 264 (squalls and tornadoes) are typical for the warm half-year (April to September) and are related to
 265 thunderstorm clouds (Cumulonimbus) developing within the thermal turbulence in the one air mass
 266 (sometimes called “isolated” or “local” thunderstorms) or within the zone of passing cold front
 267 (multiple-cell storms, squall lines, or a supercell). As a result, the duration of these kinds of strong
 268 winds is short, and their spatial coverage is local, in particular in the case of the first category.

269 Recognising the types of strong winds based on the available descriptions of this weather
 270 element in historical sources is not always easy and unambiguous, due to the scarcity of information.
 271 Another difficulty in analysing changes in the frequency of occurrence of the phenomenon is the



272 increasing number of sources as we move closer towards the present day. According to Brázdil et al.
273 (2004), this is the main limiting factor for the climatological analysis investigating changes in the
274 occurrence of strong wind in historical times. Moreover, in older periods, attention was focused more
275 on describing only extremely strong winds (see Fig. 5) and mostly those that caused serious material
276 or human damage. On the other hand, at present, the occurrence of strong winds of the 1st category
277 (fresh and strong breezes) in all meteorological stations in Poland is about ten times more frequent
278 than the sum of categories 2 and 3 (see Fig. S5). For this reason, time analysis of occurrences of most
279 extreme wind categories seems most reliable, not only for the study of historical periods but also when
280 we compare their frequencies in historical and contemporary periods. It is also not possible, based on
281 the documentary evidence, to reconstruct strong winds for the individual places in Poland as it is
282 normally done in the case of systematic instrumental measurements. Therefore, all series of strong
283 wind frequencies are presented here for the entire area of Poland.

284 The lower gust wind speed threshold, which was assumed for central Europe as potentially
285 dangerous for the destruction of buildings, forests, gardens, etc., is 17ms^{-1} (Brázdil and Dobrovolny
286 2001; Lorenc 2012). Therefore, it is possible to approximately compare the frequency of strong winds
287 above this threshold obtained from contemporary instrumental observations with the summed
288 frequency of historical winds assigned to categories 2–5. In the case of category 4 (squall), only those
289 cases for which damage was noted (8 cases) were taken. The results of the comparative analyses
290 presented below must, however, be limited only to similarities and differences in the annual cycle and
291 spatial distribution in the study area. A reliable comparison of absolute values of the frequency of
292 strong winds is impossible due to the undetectability of probably a significant number of events in the
293 study's historical period. This will only be possible for some isolated periods for which daily weather
294 records are available. For Poland, such series exist for selected periods in the 17th and 18th centuries
295 (for details see Introduction) and will be the subject of our analysis in a separate article.

296 In Poland, the most frequent winds are from the western sector (from SW to NW), in particular
297 from the NW direction, while the least frequent winds are from the sector from N to ESE (see Fig. 15.4
298 in Wibig 2021). According to investigations of Lorenc (2012), about 80% of strong winds $\geq 17\text{ms}^{-1}$ in
299 Poland in the period 1991–2005 were associated with the occurrence of NWc circulation type (i.e.,
300 wind blowing from NW direction within cyclonic pattern) according to Lityński's (1969) classification.
301 Also, the hurricane winds ($>33\text{ms}^{-1}$) in Poland in the period 1971–2005 were mostly associated with
302 NWc type and then with Wc type. As a result, both the greatest measured winds and the greatest
303 frequency of them are characteristic for north-western, western and south-western Poland, in
304 particular in the cold half-year (Lorenc 2012; Wibig 2021, see also Figs 10–12). Such spatial distribution
305 of strong winds was also found for the study's historical period when the greatest frequencies were
306 noted in the Baltic Coast and Pomerania and Silesia regions (see Fig. 4). In the Greater Poland region



307 (central western Poland) this was not found due to a very small number of historical sources available
308 (see Fig. 2). It is worth also noting some similarity in annual cycles of strong wind occurrences in the
309 historical and modern periods. In both periods, the greatest frequencies occurred in the cold half-year,
310 and the lowest in summer. However, in the historical time, a greater frequency was observed in
311 autumn than in winter, i.e. opposite to the modern time (cf. Figs 4 and 11). This can be related to
312 greater continentality of climate in medieval times than at present (see Sadowski 1991; Przybylak
313 2016; Przybylak et al. 2023). The winters were clearly colder than today (see Przybylak et al. 2005,
314 2023) and were connected with negative values of the NAO (Przybylak et al. 2003). Such a circulation
315 pattern in the Atlantic-European sector allowed more frequent advection of cold air masses to Poland
316 from eastern and northern sectors within anticyclones coming from eastern Asia (Siberia) and the
317 Arctic. On the other hand, positive NAO conditions often bring high storminess in Europe (Ommel 2015
318 and references therein).

319 From the mid-15th century, a decadal number of sources and, as a consequence, the number of
320 occurrences of strong winds seems comparable with those found in the 16th century (see Fig. 3).
321 Therefore, for this period, it is possible to make a comparison with analogical investigations made for
322 other parts of Europe. Unfortunately, the comparison is limited to only a few existing works presenting
323 results, in particular for the 15th century (see Introduction). For Czech Lands, Brázdil et al. (2004) found
324 only 24 cases of strong winds for this century, i.e. 40% fewer than we found for Poland (40). But the
325 main difference is that, in the Czech Lands, the strong winds were connected with thunderstorms
326 (convective storms), while in Poland, this category was noted only five times. A better and more
327 reliable comparison is possible with a number of storm events occurring in the coastal area of Belgium
328 and the south-west Netherlands provided by DeKraker (2013) using town accounts informing about
329 repairing the damage of dikes, piers, groynes, and quay walls after storms. In the second half of the
330 15th century, in both areas (Poland and western Europe), the greatest numbers of strong winds were
331 noted in the first and the last decades (cf. Graph 1 in DeKraker 2013 and Fig. 3 in this study). Also, a
332 good agreement is seen in the 16th century. Both in Poland and coastal areas of Belgium and the
333 Netherlands, strong winds were more common in the second half of the century than in the first.
334 Results presented for the Czech Republic (Brázdil et al. 2004) and Germany (Glaser 2013) also confirm
335 this finding. The greatest difference is noted for the second decade of the 16th century, which was
336 exceptionally rich in storm events in Belgium and the Netherlands in comparison to neighbouring
337 decades. In Poland, the maximum in this decade is also seen, but it is not as great as in the case of
338 Belgium and the Netherlands. On the other hand, in the Czech Republic, this maximum is not seen. In
339 Germany, there was a steady increase in the number of strong winds until about 1575 and then a
340 decrease by the end of the century. Such a tendency in the occurrence of strong winds in the last three



341 decades of the 16th century was also noted in Poland (see Fig. 3). In both countries, the maximum of
342 strong winds in the 16th century occurred in the 1570s.

343 Comparison of potential periods with strong winds (usually reconstructions of high/low storminess
344 periods) with results obtained using other different natural proxies (e.g., sand dune development,
345 windblown sand in peat bogs or marshes, etc.) allows only for a very rough comparison. For example,
346 the reconstruction of the start of the great storminess period in the mid-16th century (also seen in the
347 documentary evidence from Poland, Czech Republic, Germany, and Belgium and the Netherlands) was
348 found for Scandinavia (De Jong et al. 2006; Clemmensen et al. 2008), the northwest Mediterranean
349 (Sabatier et al. 2012) and Portugal (Costas et al. 2012).

350

351 5. Conclusions and final remarks

352 There are quite a few mentions of strong winds in historical periods, but as can be seen from the
353 presented literature review, this important element has been studied in a long-term perspective to
354 only a small extent. One of the most important reasons for this is certainly the great dynamics over
355 time and spatial variability of the occurrence of strong winds in Europe, including Poland. This last
356 feature is especially important in the case of convective storms, which are local in nature. The
357 mentioned features of the occurrence of strong winds, together with the significantly changing
358 number of available historical sources, especially before 1500, significantly hamper the proper
359 recognition of this important element of the climate based on documentary evidence. As a result,
360 reliable examination of changes in the occurrence of strong winds compared to the modern period is
361 extremely difficult and limited but still possible.

362 Taking into account the following reservations, we present below the most important research
363 results:

364 - A reliable estimate of some characteristics of the occurrence of strong winds in Poland seems
365 possible since the mid-15th century onwards,

366 - The highest number of strong winds occurred in the second half of the 15th century and particularly
367 in the second half of the 16th century. The decade with the most significant number of strong winds
368 was 1571–80 (14) (see Fig. 3),

369 - For each season, the greatest number of strong winds was found for the Baltic Coast and Pomerania
370 region, and then for Silesia and Lesser Poland (Fig. 4),

371 - Strong winds were noted most often in autumn and winter in Silesia and in the Baltic Coast and
372 Pomerania regions (two regions for which there is enough information to estimate the annual cycle),



373 and in summer in the Lesser Poland region (Fig. 4b). August (13 cases) and March (12) were most
374 abundant in strong wind occurrences in Poland, but the stormiest period of the year was from October
375 to March, with at least ten cases in each month except February (Fig. 4a),

376 - In the entire study period (1281–1600), and also in two subperiods (1281–1500 and 1501–1600), the
377 most frequent were storms and gales and the least frequent were fresh and strong breezes (Fig. 5, Fig
378 S1), which were most common in autumn and in winter,

379 - All categories of strong winds in the study period occurred with the greatest frequency in the Baltic
380 Coast and Pomerania region, except squalls, which were most frequently noted in the Silesia region
381 (Fig. 6). In the 16th century, however, squalls were most common in Lesser Poland (Fig. S2),

382 - Damage of category E2 (areally extended damage) was more than three times more frequent than
383 category E1 (damage of lesser extent). About one third of the notes mentioned strong wind occurrence
384 and did not contain information about the damage. The greatest and most frequent damage was noted
385 in autumn (37.2%) and the smallest in spring (19.0%) (Fig. 7, Fig. S3),

386 - Out of all distinguished categories of damage, DB (26.0%) and DL (10.5%) were most frequent, while
387 the least frequent was DO (1.7%) (Fig. 9). About 60% of all damages mentioned in the weather notes
388 were found for the cold half-year, but particularly for autumn (39.2%). The analysis reveals that the
389 most frequently noted categories of damage in Poland (DB, DL) were also noted in all analysed regions,
390 with a maximum in Silesia and Baltic Coast and Pomerania regions.

391 - Spatial distribution of strong winds estimated for the studied historical period based on documentary
392 evidence, i.e. the greatest frequencies of strong wind occurrence in the Baltic Coast and Pomerania
393 and Silesia regions (see Fig. 4) as well as the annual cycle (Fig. 5) are similar as today (Lorenc 2012;
394 Wibig 2021; see also Figs 10–12). These facts seem to confirm, to some extent, the reliability of the
395 results obtained for the historical period,

396 - The time changes in occurrences of strong winds in Poland in the period since the mid-15th century
397 were found to correspond well with other European countries (e.g., Czech Republic, Germany, and the
398 Netherlands and Belgium) using documentary evidence. Other proxy data used to reconstruct
399 storminess changes in some parts of Europe in historical times (e.g., in Scandinavia, Portugal, and the
400 south-west Mediterranean) also shows a good correspondence with the presented results for central
401 and western Europe,

402 - The “fresh and strong breeze” category of strong winds in historical times was rarely noted by
403 chroniclers, because winds of this category did not cause any material or human damage. On the other
404 hand, this category of strong winds delimited using measurement data is, at present, about ten times



405 more frequent than gales and storms taken together. Therefore, this category of strong winds
406 recorded sporadically during the historical period cannot be reliably compared to current conditions.
407 This will probably only be possible for isolated periods for which daily weather observations are
408 available.

409 The creation of databases on extreme phenomena and events in historical periods, as in this
410 case of strong winds, is extremely important in order to be able to recognise the natural range of their
411 temporal changes and to learn about spatial variability and their changes over time. The high
412 variability, both in time and space, of the extreme phenomenon under study requires that significant
413 temporal and spatial coverage be obtained for the correct and reliable identification of its features in
414 Europe from a long-term perspective. The review of the state of knowledge on this subject made in
415 this article clearly proves that it is very limited and, therefore, urgently requires intensified work,
416 especially using documentary evidence, which is undoubtedly the most accurate source of information
417 for the last millennium.

418 The same method of processing data on strong winds available in the documentary evidence
419 of individual European and other non-European countries would significantly facilitate the recognition
420 of this extreme phenomenon over a larger area and the comparison of results. For these reasons, in
421 this article, we applied the proposal given in the work of Brázdil et al. (2004), which we only slightly
422 modified to reflect the Baltic Sea's influence on the coastal part of Poland.

423

424 **Competing interests.** The authors declare that they have no known competing financial interests or
425 personal relationships that could have appeared to influence the work reported in this paper.

426 **Acknowledgements.** The research work of RP, JF, PO and PW was supported by grant funded by the
427 National Science Centre, Poland No 2020/37/B/ST10/00710). The work of AA was supported by funds
428 from IDUB Research Group *Weather and Climate: Reconstructions and Future Scenarios*. We would
429 like to thank Dr Babak Ghazi for preparing Fig. 1.

430 **Author contributions.** **Rajmund Przybylak:** Conceptualisation, Methodology, Investigation, Data
431 collection and selection, Database construction, Formal analysis, Literature review, Interpretation of
432 results, Writing – original draft, Writing – review & editing, Funding acquisition, Project
433 administration. **Andrzej Arażny:** Conceptualisation, Methodology, Investigation, Data collection and
434 selection, Formal analysis, Validation, Visualisation, Interpretation of results. **Janusz Filipiak:**
435 Conceptualisation, Methodology, Investigation, Data collection and selection, Database construction,
436 Validation, Interpretation of results, Writing – original draft. **Piotr Oliński:** Conceptualisation,
437 Investigation, Data collection and selection, Database construction, Visualisation, Validation.
438 **Przemysław Wyszynski:** Software, Visualisation, Validation, Formal analysis, Data collection and
439 selection, Database construction. **Artur Szwaba:** Database construction.

440 **Financial support.** The work was supported by the National Science Centre, Poland, project No.
441 2020/37/B/ST10/00710.



442 **Data availability.** Datasets for this research were derived from the following public domain resources:

443 1. Repository for Open Data (RepOD), Nicolaus Copernicus University Centre for Climate

444 Change Research collection, as cited in Przybylak et al. (2025),

445 <https://doi.org/10.18150/W6PMBQ>

446 2. The Institute of Meteorology and Water Management (IMGW-PIB) website:

447 <https://danepubliczne.imgw.pl/>

448

449 **References:**

450 Adamczyk A. B., 1996, Charakterystyka wiatrów silnych i bardzo silnych w Polsce, Zeszyty Instytutu Geografii i
451 Przestrzennego Zagospodarowania PAN, 37, 5-42

452 Araźny A., Przybylak R., Vízi Z., Kejna M., Maszewski R., Uscka-Kowalkowska J., 2007, Mean and extreme wind velocities in
453 Central Europe 1951-2005 (on the basis of data from NCEP/NCAR reanalysis project), *Geographia Polonica*, Vol. 80, No. 2,
454 69-78.

455 Bartnicki L., 1930, Prądy powietrzne dolne w Polsce. *Prace Geofiz.* 3(3), 2–98.

456 Brázdil R., Dobrovolný P., 2000, Chronology of strong wind events in the Czech Lands during the 16th–19th centuries.
457 *Instytut Geograficzny UJ, Prace Geograficzne* 107, 65–70.

458 Brázdil R., Dobrovolný P., 2001, History of strong winds in the Czech Lands: causes, fluctuations, impacts, *Geographia*
459 *Polonica* 74, 11–27.

460 Brázdil R., Dobrovolný P., Štekl J., Kotyza O., Valášek H., Jež J., 2004, History of weather and climate in the Czech Lands VI:
461 Strong winds, Masaryk University, Brno.

462 Brázdil R., Kotyza O., 1995, History of Weather and Climate in the Czech Lands I (Period 1000–1500), *Zürcher Geographische*
463 *Schriften* 62, Zürich, 260 pp.

464
465 *Ephemerides Societatis Meteorologicae Palatinae. 1783–1795. vol. II-XIII, Mannheim*

466
467 Chojnacka-Oźga L., Oźga W., 2018, Silne wiatry jako przyczyna zjawisk kłęskowych w lasach, *Studia i Materiały CEPL w*
468 *Rogowie*, 20 (54), 13-23.

469
470 Clemmensen, L. B., Murray, A., Heinemeier, J. and De Jong, R., 2009, The evolution of Holocene coastal dune fields, Jutland,
471 Denmark: A record of climate change over the past 5000 years, *Geomorphology*, 105: 303-313.

472
473 Costas, S., Jerez, S., Trigo, R. M., Goble, R. and Rebêlo, L., 2012, Sand invasion along the Portuguese coast forced by westerly
474 shifts during cold climate events, *Quaternary Science Reviews*, 42, 15-28.

475 Cusack S., 2023, A long record of European windstorm losses and its comparison to standard climate indices, *Nat. Hazards*
476 *Earth Syst. Sci.*, 23, 2841–2856, <https://doi.org/10.5194/nhess-23-2841-2023>.

477
478 De Jong, R., Björck, S., Björkman, L. and Clemmensen, L. B., 2006, Storminess variation during the last 6500 years as
479 reconstructed from an ombrotrophic peat bog in Halland, southwest Sweden. *Journal of Quaternary Science*, 21, 905-919.

480
481 De Kraker A., 2013, Storminess in the Low Countries, 1390–1725, *Environment and History* 19(2), 149-171,
482 <https://www.liverpooluniversitypress.co.uk/doi/10.3197/096734013X13642082568570>.

483
484 Dobrovolný P., Brázdil R., 2003, Documentary evidence on strong winds related to convective storms in the Czech Republic
485 since AD 1500, *Atmospheric Research* 67–68, 95–116.

486
487 Donat, M.G., Leckebusch, G.C., Wild, S., Ulbrich, U., 2011. Future changes in European winter storm losses and extreme
488 wind speeds inferred from GCM and RCM multimodel simulations. *Nat. Hazards Earth Syst. Sci.* 11, 1351–1370.

489 <http://dx.doi.org/10.5194/nhess-11-1351-2011>.

490



- 491 Galloway, J.A., 2009, Storm Flooding, Coastal Defence and Land Use around the Thames Estuary and Tidal River c. 1250–
492 1450', *Journal of Medieval History* 35, 171–188.
493
494 Galloway, J.A. and Potts, J., 2007, Marine Flooding in the Thames Estuary and Tidal River c.1250–1450: Impact and
495 Response, *Area* 39, 370–379.
496
497 Ghazi B., Przybylak R., Oliński P., Bogdańska K., Pospieszńska A., 2023a, The frequency, intensity, and origin of floods in
498 Poland in the 11th–15th centuries based on documentary evidence, *Journal of Hydrology*, vol. 623, 129778,
499 <https://doi.org/10.1016/j.jhydrol.2023.129778>
500
501 Ghazi B., Przybylak R., Oliński P., Chorążyczewski W., Pospieszńska A., 2023b, An assessment of flood occurrences in
502 Poland in the 16th century, *Journal of Hydrology: Regional Studies* 50, 101597, <https://doi.org/10.1016/j.ejrh.2023.101597>.
503
504 Ghazi B., Przybylak R., Oliński P., Pospieszńska A., 2025, Flood occurrences and characteristics in Poland (Central Europe) in
505 the last millennium, *Global and Planetary Change*, 246, 104706, <https://doi.org/10.1016/j.gloplacha.2025.104706>
506
507 Ghazi B., Przybylak R., Oliński P., Targowski M., Filipiak J., Pospieszńska A., 2024, A comprehensive study of floods in
508 Poland in the 17th–18th centuries, *Journal of Hydrology: Regional Studies* 53, 101796,
509 <https://doi.org/10.1016/j.ejrh.2024.101796>.
510
511 Glaser R., 2001, *Klimageschichte Mitteleuropas. 1000 Jahre Wetter, Klima, Katastrophen*, Primus Verlag, Darmstadt, 227 pp.
512
513 Glaser R., 2013, *Klimageschichte Mitteleuropas. 1200 Jahre Wetter, Klima, Katastrophen*, 3. Auflage, WBG, Darmstadt, 264
514 pp.
515
516 Gumiński R., 1952, Rozkład kierunków i prędkości wiatru na niektórych stacjach meteorologicznych Polski, *Wiad. St. Hydr. i*
517 *Met.* 3(2a), 45–64.
518
519 Hawkins E., Brohan P., Burgess S.N., Burt S., Compo G. P., Gray S.L., Haigh I.D., Hersbach H., Kuyjjer K., Martinez-Alvarado O.,
520 Mc Coll C., Schurer A.P., Slivinski L., Williams J., 2023, Rescuing historical weather observations improves quantification of
521 severe windstorm risks, *Nat. Hazards Earth Syst. Sci.*, 23, 1465–1482, <https://doi.org/10.5194/nhess-23-1465-2023>.
522
523 Krawczyk B., 1994, Średnia liczba dni z wiatrem o prędkości $v > 8$ ms. In: *Atlas zasobów, walorów i zagrożeń środowiska*
524 *geograficznego Polski*, IGiPZ PAN, Warszawa.
525
526 Lamb F., 1991, *Historic Storms of the North Sea, British Isles and Northwest Europe*, Great Britain, Cambridge University
527 Press.
528
529 Lityński J., 1969, Liczbowa klasyfikacja typów cyrkulacji I typów pogody dla Polski, *Prace PIHM*, 97, 3-14.
530
531 Lorenc H., 1996 – Struktura i zasoby energetyczne wiatru w Polsce. *Materiały Badawcze*, s. *Meteorologia* 25. IMGW,
532 Warszawa.
533
534 Lorenc H., 2012, *Maksymalne prędkości wiatru w Polsce*, Monografie, Instytut Meteorologii i Gospodarki Wodnej-
535 Państwowy Instytut Badawczy, Warszawa.
536
537 Lundstad E., Brugnara Y., Pappert D. Kopp J. Samakinwa E., Hürzeler A., Andersson A., Chimani B., Cornes R., Demarée G.,
538 Filipiak J., Gates L., Ives G.L., Jones J.M., Jourdain S., Kiss A., Nicholson S.E., Przybylak R., Jones P., Rousseau D., Tinz B.,
539 Rodrigo F. S., Grab S., Domínguez-Castro F., Slonosky V., Cooper J., Brunet M., Brönnimann S., 2023, The global historical
540 climate database HCLIM, *Scientific Data* 10:44, <https://doi.org/10.1038/s41597-022-01919-w> 1.
541
542 MunichRe, 2011, *Topics GEO, Natural Catastrophes 2011 Analyses Assessments Positions*. Munich Reinsurance Company
Publications, Munich, pp. 1-49.
543
544 MinicheRe, 2020, *Natural Catastrophe statistics online – the NatCatSERVICE analysis tool*.
<https://www.munichre.com/en/solutions/for-industry-clients/natcatservice.html>.
545
546 Orme L. C., 2014, *Reconstructions of Late Holocene storminess in Europe and the role of the North Atlantic Oscillation*, PhD
work, University of Exeter, 306 pp.



- 547 Outten S., Sobolowski S., 2021, Extreme wind projections over Europe from the Euro-CORDEX regional climate models,
548 Weather and Climate Extremes, 33, 100363, <https://doi.org/10.1016/j.wace.2021.100363>.
- 549 Pappert D., Brugnara Y., Jourdain S., Pospieszńska A., Przybylak R., Rohr Ch., and Brönnimann S., 2021, Unlocking weather
550 observations from the Societas Meteorologica Palatina (1781–1792), Clim. Past, 17, 2361–2379,
551 <https://doi.org/10.5194/cp-17-2361-2021>.
- 552 Paszyński J., Niedźwiedź T., 1991, Klimat. In: Starkel L. (ed.) Geografia Polski środowisko przyrodnicze. PWN, Warszawa.
- 553 Piasecki D., 1952, Wiatry o maksymalnych prędkościach na obszarze Polski w latach 1928–1938, Wiad. St. Hydr. i Met. 3(2a),
554 65–101
- 555 Pfister, C., 1999. Winternachhersage. 500 Jahre Klimavariationen und Naturkatastrophen (1496– 1995). Paul Haupt, Bern.
- 556 Przybylak R., 2010, The Climate of Poland in Recent Centuries: A Synthesis of Current Knowledge: Instrumental
557 observations. In: Przybylak R., Majorowicz J., Brázdil R., Kejna M. (eds), The Polish Climate in the European Context: An
558 Historical Overview, Springer, Berlin Heidelberg New York, 129-166.
- 559 Przybylak R., 2016, Poland's Climate in the Last Millennium, in: Oxford Research Encyclopedia, Climate Science, Oxford
560 University Press, USA, <https://doi.org/10.1093/acrefore/9780190228620.013.2>.
- 561 Przybylak R., Majorowicz J., Wójcik G., Zielski A., Chorążyczewski W., Marciniak K. Nowosad W., Oliński P. Syta K., 2005,
562 Temperature changes in Poland from the 16th to the 20th centuries, Int. J. of Climatology, 25, 773-791.
- 563 Przybylak R., Oliński, P., Koprowski, M., Filipiak, J., Pospieszńska, A., Chorążyczewski, W., Puchałka, R., and Dąbrowski, H.
564 P.: Droughts in the area of Poland in recent centuries in the light of multi-proxy data, Clim. Past, 16, 627–661,
565 <https://doi.org/10.5194/cp-16-627-2020>, 2020.
- 566 Przybylak R., Oliński, P., Koprowski, M., Szychowska-Krąpiec, E., Krąpiec, M., Pospieszńska, A., and Puchałka, R.: The
567 climate in Poland (central Europe) in the first half of the last millennium, revisited, Clim. Past, 19, 2389–2408,
568 <https://doi.org/10.5194/cp-19-2389-2023>, 2023.
- 569 Przybylak R., Oliński P., Wyszyński P., Szwabą, A., Filipiak J., Araźny A. 2025, "Database of strong wind occurrence in
570 Poland from the 13th to 16th centuries based on documentary evidence", <https://doi.org/10.18150/W6PMBQ>, RepOD, V1
- 571 Przybylak R., Pospieszńska A., Wyszyński P., Nowakowski M., 2014, Air temperature changes in Żagan (Poland) in the
572 period from 1781 to 1792. Int. J. Climatol. 34: 2408–2426, <https://doi.org/10.1002/joc.3847>.
- 573 Przybylak R., Wójcik G., Marciniak K., 2003, Wpływ Oscylacji Północnoatlantyckiej oraz Arktycznej na warunki termiczne
574 chłodnej pory roku w Polsce w XVI-XX wiekach, Przegl. Geof., 1-2, 61-74.
- 575 Sadowski, M., 1991, Variability of extreme climatic events in Central Europe since the 13th century, Z. Meteorol., 41, 350–
576 356, 1991.
- 577 Stopa-Boryczka M. (ed.), 1989, Atlas współzależności parametrów meteorologicznych i geograficznych w Polsce, 5. Wyd.
578 UW, Warszawa.
- 579 Tarnowska K., 2011, Wiatry silne na polskim wybrzeżu Morza Bałtyckiego, Prace i Studia Geograficzne, 47, 197-214.
- 580 Ustrnul Z., Wypych A., Henek E., Czekierda D., Walawender J., Kubacka D., Pyrc R., Czernecki B., 2014, Atlas zagrożeń
581 meteorologicznych Polski (Meteorological hazard atlas of Poland), Instytut Meteorologii i Gospodarki Wodnej - Państwowy
582 Instytut Badawczy, Kraków
- 583 Wibig J., 2021, Change of Wind. In: Falarz M. (Ed.), Climate Change in Poland: Past, Present, Future, Springer Nature
584 Switzerland, 391-420.
- 585