



1 **Representation of a two-way coupled irrigation system in the Common Land Model**

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10

11 **Abstract**

12 Human land–water management, especially irrigation water withdrawal and use, significantly
13 impacts the global and regional water cycle, energy budget, and near-surface climate. While land
14 surface models are widely used to explore and predict the impacts of irrigation, the irrigation
15 system representation in these models is still in its early stages. This study enhances the
16 Common Land Model (CoLM) by introducing a two-way coupled irrigation module. This
17 module includes an irrigation water demand scheme based on soil moisture deficit, an irrigation
18 application scheme considering four major irrigation methods, and an irrigation water
19 withdrawal scheme that incorporates multiple water source constraints by integrating CoLM with
20 a river routing model and a reservoir operation scheme. Crucially, it explicitly accounts for the
21 feedback between irrigation water demand and supply, which is constrained by available surface
22 water (i.e., runoff, streamflow, reservoir storage) and groundwater. Simulations conducted from
23 2001 to 2016 at a 0.25° spatial resolution across the contiguous United States reveal that the
24 model effectively reproduces irrigation withdrawals, their spatial distribution, and water source
25 proportions, aligning well with reported state-level statistics. Comprehensive validation
26 demonstrates that the new module significantly improves model accuracy in simulating regional
27 energy dynamics (sensible heat, latent heat, and surface temperature), hydrology (river flow),
28 and agricultural outputs (yields for maize, soybean, and wheat). Application analyses highlight
29 the potential of the enhanced CoLM as a valuable tool for predicting irrigation-driven climate
30 impacts and assessing water use and scarcity. This research offers a pathway for a more holistic
31 representation of fluxes in irrigated areas and human-water interactions within land surface
32 models. It is valuable for exploring the interconnected evolution of climate, water resources,
33 agricultural production, and irrigation activities, while supporting sustainable water management
34 decisions in a changing climate.



35 1. Introduction

36 Freshwater resources are indispensable for human society. Since 1900, the global population has
37 increased more than fourfold, leading to a nearly sixfold rise in water withdrawals, from
38 approximately 500 km³ per year in 1900 to about 3000 km³ per year in 2000, with agriculture
39 being the dominant water user (Pokhrel et al., 2016). Around 70% of global freshwater has been
40 withdrawn for irrigation (Campbell et al., 2017), accounting for 90% of consumptive water use
41 (Siebert and Döll, 2010), with irrigated areas providing approximately 40% of global food
42 production on just 2.5% of global land (Abdullah, 2006). Accompanied by significant
43 socioeconomic benefits, these intense human land-water management practices have profoundly
44 altered Earth's surface and impacted terrestrial water and energy cycles (Ketchum et al., 2023;
45 Nocco et al., 2019; Rappin et al., 2022; Thiery et al., 2017; de Vrese et al., 2016). The demand
46 for irrigation water is anticipated to rise with the growing global population and food demand,
47 while climate-warming-induced droughts are likely to exacerbate this need (McDermid et al.,
48 2023; Mehta et al., 2024; Yang et al., 2023). Therefore, understanding and quantifying the
49 impacts of irrigation water management in human-Earth system interactions are crucial for
50 developing strategies to sustainably manage these resources amidst changing climatic and
51 demographic conditions.

52 Irrigation practices transfer water from various sources, such as rivers, lakes, reservoirs, and
53 aquifers, into agricultural systems, directly affecting the magnitude and timing of runoff and
54 river flow (Ketchum et al., 2023). The rising irrigation demand has spurred increased
55 construction of reservoirs and diversions, resulting in both local and downstream impacts. In
56 some regions, water extraction for irrigation has reduced the availability of both surface and
57 groundwater (Döll et al., 2014). Besides modifying water fluxes, irrigation also influences
58 regional climates both locally and remotely. Locally, it alters surface albedo, evapotranspiration,
59 and surface soil moisture, impacting regional radiation and energy balances and affecting
60 temperature, humidity, and precipitation through land-atmosphere feedback (Chen and Dirmeyer,
61 2019; Kang and Eltahir, 2018; Li et al., 2022; McDermid et al., 2017; Nocco et al., 2019).
62 Remotely, it affects climate through complex interactions between altered temperature and
63 moisture gradients and larger-scale processes such as atmospheric circulation and wave activity
64 (Douglas et al., 2009; Phillips et al., 2022; de Vrese et al., 2016).

65 Earth system models (ESMs) are powerful tools for examining the interactions and feedback
66 among the intricately intertwined processes of the Earth system, both in the past and future. Land
67 surface models (LSMs) are a crucial component of ESMs. Due to the complex dynamics of
68 natural hydrological processes and anthropogenic activities, describing human-water interactions
69 has been recognized as a significant challenge in Land surface modeling (Nazemi and Wheeler,
70 2015). In recent years, targeted efforts have aimed to address this deficiency, yet water use
71 remains largely underrepresented or in a nascent stage within LSMs (Blyth et al., 2021; Taranu et
72 al., 2024). Meanwhile, global hydrological models (GHMs), originally designed for water



73 resource assessment, have undergone continuous improvements over the last three decades to
74 explicitly represent human water use (Hanasaki et al., 2018; Liang et al., 1994; Müller Schmied
75 et al., 2021; Sood and Smakhtin, 2015; Sutanudjaja et al., 2018; Tang et al., 2007). These models
76 enable the determination of the spatial distribution and temporal evolution of water resources and
77 water stress for both humans and other biota under the pressures of global change (Döll et al.,
78 2018; Schewe et al., 2014; Schlosser et al., 2014). These advancements have offered valuable
79 insights for incorporating human water use into LSMs.

80 Parameterizing irrigation water use and modeling its impacts in GHMs and LSMs has been
81 approached using different assumptions and simplifications in three key aspects: irrigation
82 demands, irrigation methods, and irrigation water supplies/withdrawals. The first aspect is
83 estimating irrigation water demands. Models estimate these demands using either a root-zone
84 soil moisture deficit approach or a crop-specific potential evapotranspiration approach. The root-
85 zone soil moisture deficit approach estimates irrigation demand as the water needed to keep root-
86 zone soil moisture (usually within the top meter of soil) above a certain threshold during the
87 growing season (normally a certain percentage of field capacity or soil saturation) (Ozdogan et
88 al., 2010). The crop-specific potential evapotranspiration approach estimates irrigation needs
89 based on the difference between crop-specific potential evapotranspiration and simulated
90 unirrigated evapotranspiration, or between potential and effective precipitation under well-
91 watered conditions where crops transpire at their maximum rate (Müller Schmied et al., 2021).
92 Notably, LSMs generally do not use potential evapotranspiration to estimate irrigation demand.

93 The second aspect concerns the representation of irrigation methods. Many models simplify
94 irrigation application by directly modifying soil moisture or treating it as additional rainfall
95 across all irrigated land, overlooking the diversity of irrigation techniques employed in various
96 parts of the world or by different farmers (Li et al., 2024; Lu et al., 2015; de Vrese et al., 2018).
97 Recently, some models have started integrating specific irrigation techniques for certain crops or
98 regions. For instance, LPJmL includes sprinkler, drip, and surface irrigation methods, and CLM
99 incorporates drip, sprinkler, flood, and paddy irrigation methods (Jägermeyr et al., 2015; Yao et
100 al., 2022). Different irrigation techniques affect farmland hydrological processes and irrigation
101 efficiency in distinct ways. For example, drip and surface irrigation methods avoid interception
102 losses observed with sprinkler (Nair et al., 2013).

103 Third is the representation of irrigation water supplies/withdrawals, which is particularly critical
104 as it involves the interaction between multiple processes or modules, such as hydrological and
105 agricultural systems. However, explicit representation of these interactions remains largely
106 absent in LSMs, despite the extensive modeling experience provided by GHMs. Such modeling
107 first requires identifying the sources of irrigation water, typically categorized into surface water
108 and groundwater. Surface water sources are normally constrained by available runoff,
109 streamflow, and storage such as lakes and reservoirs. Accessing these sources, such as rivers and
110 reservoirs, necessitates coupling with river routing and reservoir modules, which are well-



111 represented in many GHMs (Biemans et al., 2011; Hanasaki et al., 2018). Groundwater is
112 typically divided into renewable sources (baseflow or dynamic groundwater levels) and
113 nonrenewable sources (fossil groundwater). Some models assume an inexhaustible supply of
114 nonrenewable groundwater to meet irrigation demands, neglecting irrigation shortages caused by
115 water scarcity (Zhou et al., 2020). Additionally, some GHMs incorporate alternative sources,
116 such as inter-regional water transfers and seawater desalination (Hanasaki et al., 2018;
117 Sutanudjaja et al., 2018). A second critical aspect of irrigation water supply modeling is
118 determining the allocation of irrigation water among different sources, including the
119 prioritization of water usage. Various models adopt different assumptions for this allocation. For
120 example, H08 prioritizes surface water (Hanasaki et al., 2018), while WBMplus prioritizes
121 reservoirs and groundwater (Wisser et al., 2010). PCR-GLOBWB uses an empirical approach
122 that allocates groundwater use based on comparisons between baseflow conditions and long-term
123 historical climatology, capturing feedback between water supply and demand (Sutanudjaja et al.,
124 2018). Another common approach is to assume a predefined allocation ratio based on water
125 withdrawal infrastructure (e.g., Siebert et al., 2010), using this ratio to divide total irrigation
126 abstractions between groundwater and surface water (Arboleda-Obando et al., 2024; Leng et al.,
127 2015). Despite these advances, the representation of water extraction and the coupling of
128 irrigation and hydrological systems in LSMs is still in its early stages. Most irrigation-enabled
129 models still assume an unlimited water supply, failing to account for constraints imposed by
130 water availability (Druel et al., 2022; Yao et al., 2022; Zhou et al., 2020).

131 The Common Land Model (CoLM; Dai et al., 2003), derived from the Community Land Model
132 (CLM), is a widely used land surface model that integrates ecological, hydrological, and
133 biophysical processes. In recent years, it has further incorporated various physical processes such
134 as lakes, wetlands, and dynamic vegetation, enhancing the representation of energy and water
135 exchanges among soil, vegetation, snow, and atmosphere. CoLM has been successfully
136 implemented in global atmospheric models, such as GRAPES, CWRP, and CAS-ESM2.0 (Shen
137 et al., 2021; Yuan and Liang, 2011; Zhang et al., 2020a). Despite significant advancements in
138 parameterizing natural land surface processes, the representation of human activities in CoLM
139 remains at an early stage. Recently, CoLM has further integrated a crop module, providing a
140 foundation for considering irrigation and its interactions with natural water systems.

141 To enhance the representation of human–water interactions in land surface models, we introduce
142 a new irrigation module for CoLM. This module provides a comprehensive framework for
143 simulating the entire irrigation water system, including water demand, withdrawal, and
144 utilization. It incorporates an irrigation water demand scheme based on soil moisture deficits, an
145 irrigation application scheme accounting for four major irrigation methods, and an irrigation
146 water withdrawal scheme that incorporates multiple water source constraints by integrating
147 CoLM with a river routing model and a reservoir operation scheme. A key focus of this module
148 is the bidirectional coupling between irrigation water demand and supply, alongside a detailed
149 representation of water withdrawals from different sources. Section 2 provides a detailed



150 description of the module and its implementation, including an overview of CoLM, the datasets
151 used for simulation and validation, and the experimental design. Section 3 validates the module's
152 performance in simulating irrigation water withdrawals using reported data and compares its
153 results to other hydrological models. It also assesses improvements in model accuracy for
154 regional energy dynamics (sensible heat, latent heat, and surface temperature), hydrology (river
155 flow), and agricultural outputs (maize, soybean, and wheat yields). Section 4 demonstrates two
156 key applications of the module: analyzing irrigation impacts on the energy budget and evaluating
157 irrigation water security. Finally, we discuss the module's current limitations and propose
158 potential future improvements.

159 **2. Materials and Methods**

160 **2.1 Description of CoLM and its crop module**

161 The Common Land Model (CoLM) is one of the most advanced land surface models widely used
162 to simulate the Water–Energy–Carbon Nexus. The original version of CoLM (Dai et al., 2003)
163 combines the three land surface models: the Land Surface Model (LSM; Bonan, 1996), the
164 Biosphere-Atmosphere Transfer Scheme (BATS; Dickinson et al., 1993), and the 1994 version of
165 the Chinese Academy of Sciences Institute of Atmospheric Physics LSM (IAP94; Dai and Zeng,
166 1997). CoLM2014 integrates the Catchment-Based Macro-Scale Floodplain model (CaMa-
167 Flood; Yamazaki et al., 2011), enabling river routing calculations within the model. Specifically,
168 runoff generated by CoLM is transferred to CaMa-Flood for routing through the river network.
169 CaMa-Flood represents the river network as a series of irregular unit catchments, defined
170 through sub-grid topographic parameters. River discharge and other flow characteristics are
171 computed using the local inertial equations along the river network, allowing for detailed flow
172 dynamics across catchments.

173 The CoLM2024 version incorporates substantial updates over CoLM2014, particularly by
174 introducing representations of biogeochemical cycles and human activity processes (e.g., crop
175 growth and reservoir management). The new crop module introduces a phenological
176 development scheme based on accumulated temperature, a biomass allocation scheme among
177 different plant organs, and fertilization schemes (Drewniak et al., 2013). Crops are categorized
178 into four organ pools: leaves, stems, fine roots, and grains. The growth stages are divided into
179 three phases: sowing to emergence, emergence to grain filling, and grain filling to maturity, with
180 carbon allocation ratios to roots, stems, leaves, and grains varying across these phases. Upon
181 maturation, crops are harvested, with part of the carbon from the grains contributing to the yield,
182 while a small portion (3g) is reserved as seeds for the next growing season. For carbon
183 assimilation, the module employs Farquhar's photosynthesis scheme (Collatz et al., 1992;
184 Farquhar et al., 1980) and Ball-Berry's stomatal model (Ball et al., 1987; Collatz et al., 1991),
185 treating maize as a C4 crop and other crops as C3. Additionally, the module accounts for the
186 effects of heat stress, water stress, nitrogen stress, and ozone stress on yield (Li et al., 2024;



187 Lombardozzi et al., 2020). The module has been calibrated for various crops, including maize,
188 soybean, spring and winter wheat, rice, cotton, and sugarcane, enabling accurate simulation of
189 crop yields.

190 **2.2 Two-way coupled irrigation water use module**

191 **2.2.1 Irrigation demand**

192 The irrigation demand is calculated using the soil moisture deficit method (Leng et al., 2017;
193 Ozdogan et al., 2010; Yao et al., 2022). During the crop growth stage, irrigation is triggered at 6
194 a.m. local time if the soil moisture in the root zone ($z_{\text{irrig}}=1\text{m}$) falls below the threshold value
195 (ω_{thresh}). The total irrigation water demand (D_{irrig} , mm) is then calculated using Equation (1):

$$196 \quad D_{\text{irrig}} = \begin{cases} \omega_{\text{irrig}} - \omega_{\text{avail}} & \omega_{\text{avail}} \leq \omega_{\text{thresh}} \\ 0 & \omega_{\text{avail}} > \omega_{\text{thresh}} \end{cases} \quad (1)$$

197 where ω_{avail} is the total soil water amount in the root zone (mm); ω_{irrig} is the irrigation target
198 threshold (mm), calculated using Equation (2):

$$199 \quad \omega_{\text{irrig}} = f_{\text{irrig}}(\omega_{\text{target}} - \omega_{\text{wilt}}) + \omega_{\text{wilt}} \quad (2)$$

200 where ω_{wilt} is the wilting point soil water amount in the root zone (mm), calculated as the sum of
201 soil water at the wilting point for each soil layer ($\sum_{j=1}^{N_{\text{irr}}} \theta_{\text{wilt}} \Delta z_j$); ω_{target} is the target soil water
202 amount in the root zone (mm), calculated as the sum of target soil water for each soil layer
203 ($\sum_{j=1}^{N_{\text{irr}}} \theta_{\text{target}} \Delta z_j$). N_{irr} is the number of soil layers in the root zone and Δz_j is the thickness of each
204 soil layer (m). The target (θ_{target}) and wilting point (θ_{wilt}) soil moisture (m^3/m^3) for each layer are
205 calculated based on the corresponding soil water potential (Φ_{target} and Φ_{wilt}). f_{irrig} is a weighting
206 coefficient ranging from 0 to 1, controlling the extent to which soil water amount approaches the
207 target level ω_{target} during irrigation (default value = 1). In some cases, it can represent the
208 efficiency of the irrigation system, accounting for water losses due to evaporation, seepage, or
209 other factors.

210 The irrigation trigger threshold (ω_{thresh}) in Equation (1) is calculated as:

$$211 \quad \omega_{\text{thresh}} = f_{\text{thresh}}(\omega_{\text{trigger}} - \omega_{\text{wilt}}) + \omega_{\text{wilt}} \quad (3)$$

212 where ω_{trigger} is the trigger water amount in the root zone (mm); f_{thresh} is also a weighting
213 coefficient ranging from 0 to 1 that controls the proximity of soil water amount to the trigger
214 level ω_{trigger} (default value = 1).



215 The values of ω_{trigger} and ω_{target} are set according to the irrigation application method. For drip
216 and sprinkler irrigation, both ω_{trigger} and ω_{target} are set to the soil field capacity water amount. For
217 flood irrigation, ω_{trigger} is set to the soil field capacity water amount and ω_{target} to the saturation
218 water amount. For paddy irrigation, both ω_{trigger} and ω_{target} are set to the saturation water amount.

219 2.2.2 Irrigation application

220 The model incorporates four different irrigation application methods: drip irrigation, sprinkler
221 irrigation, flood irrigation, and paddy irrigation, each with unique triggering conditions, water
222 demand requirements, and application processes. Drip irrigation is triggered when soil moisture
223 in the root zone falls below field capacity, with the irrigation goal being to restore soil moisture
224 to field capacity. This method applies water directly to the surface soil, allowing it to percolate
225 into deeper soil layers. Sprinkler irrigation shares the same triggering condition and demand
226 requirement as drip irrigation but applies water above the canopy. In this method, water can be
227 intercepted and evaporated before reaching the soil surface, resulting in relatively lower
228 irrigation efficiency. This method is the most commonly used in the United States. Flood
229 irrigation is triggered when soil moisture falls below field capacity, to raise soil moisture to the
230 point of saturation. Paddy irrigation is applied whenever soil moisture drops below saturation,
231 aiming to restore soil moisture to saturation without causing runoff loss. Paddy fields are
232 typically maintained with a specific water level on the surface (10 cm) during the growing
233 season. A global irrigation method map (Yao et al., 2022; Figure S3) is used to determine the
234 irrigation method for each grid. In addition, irrigation is implemented daily at 6 a.m., if
235 necessary, with water supply evenly distributed across each time step throughout the next 4
236 hours.

237 2.2.3 Irrigation water supply/withdrawal

238 The model incorporates two distinct irrigation water supply/withdrawal schemes. The first
239 scheme, Unlimited Supply (irrig-unlim), assumes that irrigation demand is fully met without
240 accounting for specific water sources, a common approach in most land surface models (Yao et
241 al., 2022). The second scheme, Limited Supply (irrig-lim), divides total irrigation demand
242 between surface water and groundwater sources, labeled as surface water demand (D_{surf}) and
243 groundwater demand (D_{grnd}), respectively. Both demands are constrained by the available water
244 within each respective system. This distribution is based on the spatial extent of groundwater
245 irrigation equipment, as provided by Siebert et al. (2010), and is formulated as follows:

$$246 \quad D_{\text{surf}} = D_{\text{irrig}} \times (1 - f_{\text{grnd}}) \quad (4)$$

$$247 \quad D_{\text{grnd}} = D_{\text{irrig}} \times f_{\text{grnd}} \quad (5)$$



248 where D_{surf} and D_{grnd} represent the demand from surface water and groundwater systems, and
249 f_{grnd} denotes the area fraction covered by groundwater equipment. In this scheme, surface water
250 demand (D_{surf}) is sourced sequentially from local grid cell runoff, local river streamflow, and
251 upstream reservoirs, while groundwater demand (D_{grnd}) is drawn from groundwater aquifers.

252 2.2.3.1 Surface water supply

253 In our two-way coupled irrigation system (Figure 1), the daily surface water supply for irrigation
254 is constrained by surface water availability, which is simulated by CoLM (runoff) and CaMa-
255 Flood (local streamflow and upstream reservoirs). We first examine whether the runoff from the
256 local grid cell (S_{ro}) can meet the daily surface water demand (D_{surf}) for that cell. If runoff is
257 insufficient, additional water is sourced from local streamflow and upstream reservoirs. River
258 streamflow availability (S_{riv}) is determined by CaMa-Flood. For each irrigated grid cell, the river
259 grid with the highest flow within a 250 km radius is selected as the source. To prevent excessive
260 water extraction, a withdrawal limit is imposed, ensuring that the remaining flow in each river
261 grid cell does not drop below 20% of its average daily volume. Before conducting irrigation
262 simulations, natural river flow simulations are performed to establish essential parameters for
263 both river and reservoir water withdrawal schemes.

264 Reservoir water availability (S_{res}) is also determined by CaMa-Flood, which now includes a
265 reservoir module. This module consists of the following components: (i) a reservoir dataset that
266 provides reservoir location information matched with the river network, along with reservoir
267 parameters (e.g., characteristic storage capacity); (ii) a reservoir operation scheme designed for
268 flood control; and (iii) a routing scheme that integrates reservoir operations into river flow
269 simulations. For more details, refer to Hanazaki et al. (2022). In this study, we further propose a
270 new scheme for sourcing irrigation water from reservoirs (Figure 2), which involves the
271 following steps:

272 (i) Identifying the irrigation area served by each reservoir. It is challenging to accurately define
273 the true irrigation extent/area for each reservoir, especially across large spatial domains.
274 Therefore, a simplified approach is adopted here: larger reservoirs are assumed to cover a
275 proportionately larger irrigation area, restricted to downstream regions only (since upstream
276 water transfer is economically infeasible). Based on the relationship between reservoir size and
277 the corresponding irrigation area provided in Table S1, we calculate the irrigation area for each
278 reservoir according to its storage capacity by linear interpolation. Downstream irrigation grids
279 are selected sequentially, from nearest to farthest, until the cumulative grid area closely matches
280 the calculated irrigation area. If multiple reservoirs serve the same irrigation grid, a sharing
281 proportion (f_{share} , ranging from 0 to 1) is assigned to the irrigation grid based on the degree of
282 shared usage.

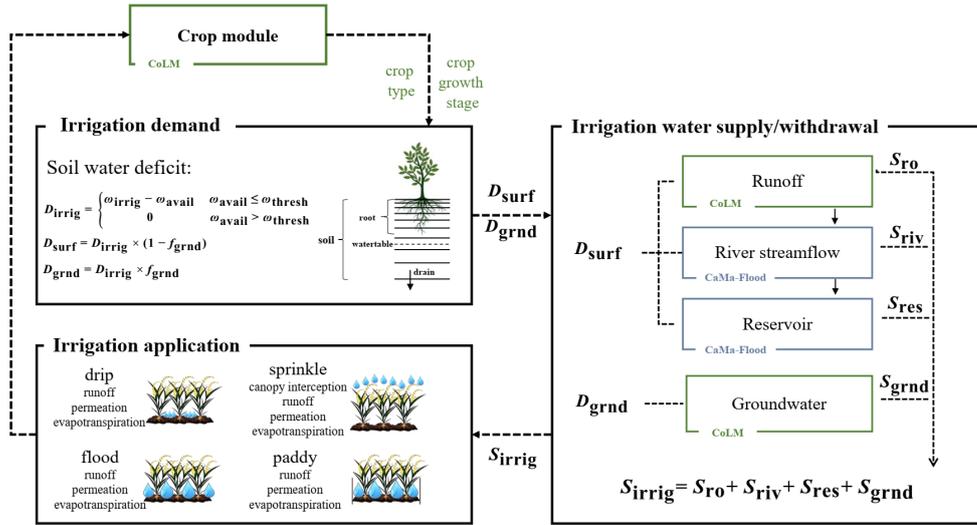


283 (ii) Calculating the irrigation demand for each reservoir by aggregating the demands of
284 associated irrigation grids. This is expressed as: $D_{\text{res-total}} = \sum_{i=1}^N (D_{\text{irrig-unmet}}^i \times f_{\text{share}}^i)$, where
285 $D_{\text{irrig-unmet}}^i$ and f_{share}^i represent the irrigation demand (i.e., the portion of D_{surf} not met by local
286 runoff and river streamflow) and sharing proportion of grid i , respectively. N denotes the number
287 of irrigation grids served by the reservoir.

288 (iii) Executing reservoir withdrawals for irrigation based on demands. Water is then withdrawn
289 ($S_{\text{res-total}}$) from the reservoir's effective storage (V_{eff}) — the portion between the current water
290 level and dead water level—according to the required demand. This is expressed as
291 $S_{\text{res-total}} = \min(D_{\text{res-total}}, V_{\text{eff}})$. After updating the reservoir storage, the reservoir operation and
292 subsequent river routing are calculated following the approach outlined in Hanazaki et al. (2022).

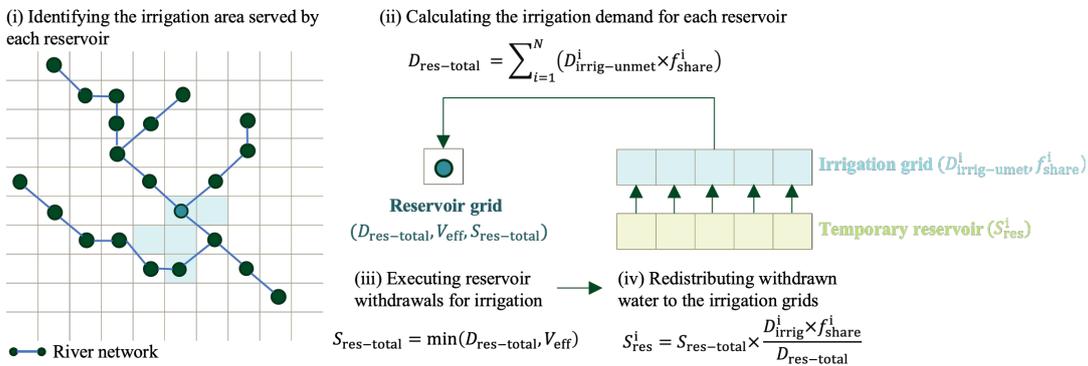
293 (iv) Redistributing withdrawn water to the irrigation grids. Based on each irrigation grid's
294 contribution to the total reservoir irrigation demand, the total withdrawal volume is
295 proportionally allocated across the associated grids (S_{res}^i). This is expressed as
296 $S_{\text{res}}^i = S_{\text{res-total}} \times \frac{D_{\text{irrig-unmet}}^i \times f_{\text{share}}^i}{D_{\text{res-total}}}$. Notably, this water is not applied directly to irrigation but is stored in
297 a temporary reservoir (i.e., a temporary variable) for each irrigation grid in CoLM. This approach
298 addresses the response delay in water supply from the river routing model to the land model's
299 irrigation demands, as the time step for CoLM is 60 minutes, while CaMa-Flood operates with a
300 6-hour time step and exchanges information with CoLM every 6 hours. Moreover, if the
301 reservoir cannot fully meet the irrigation demand within the initial time step, any unmet demand
302 is carried forward to the next time step. This process continues over a 24-hour cycle, after which
303 new water demands for the next day are received.

304 Thus, the computational sequence proceeds as follows: Step (i) is completed before model
305 execution, with its results serving as an essential input for the irrigation module. During model
306 operation, CoLM calculates the irrigation demand at 6 a.m. local time. The unmet demand (after
307 subtracting the water supplied by local runoff and streamflow) is then sent to CaMa-Flood, as
308 described in Step (ii). CaMa-Flood supplies water from reservoir to meet this demand, as
309 described in Step (iii), and returns the supplied water to CoLM according to Step (iv), over the
310 next 24 hours. During this process, the water supplied by reservoir is stored in the temporary
311 reservoir (variable) for each irrigation grid within CoLM. The following day, when irrigation
312 begins again at 6 a.m., water is withdrawn directly from the temporary reservoir if the demand
313 cannot be met by local runoff and streamflow.



314
 315

Figure 1. Diagram of the two-way coupled irrigation water system in the Common Land Model.



316
 317

Figure 2. Diagram of the reservoir water supply scheme.

318 2.2.3.1 Groundwater supply

319 Groundwater supply is constrained by the availability of water within the aquifer. In CoLM, the
 320 groundwater table interacts with soil layers through vertical water exchange, allowing recharge
 321 or withdrawal of water from the aquifer (Li et al., 2017). The evolution of the groundwater table
 322 is determined by the balance of soil water recharge and subsurface outflow, with the specific
 323 yield dynamically linking the water table position to changes in soil moisture and aquifer
 324 storage. When irrigation is required, water is directly extracted from the top of the simulated
 325 aquifer, and the water table depth is updated accordingly. This process continues until either the
 326 irrigation demand is fully met, or the water table falls below a predefined threshold, set as 1
 327 meter below the initial depth at the beginning of the year (Jasechko et al., 2024; Russo and Lall,
 328 2017). Groundwater supply is immediately available upon demand, with no temporal lag



329 between the request and its availability for irrigation. Changes in the water table depth can then
330 affect subsurface drainage and recharge from the bottom soil layer to the aquifer.

331 **2.3 Materials**

332 **2.3.1 Input datasets**

333 In this study, CoLM was implemented across the contiguous United States at a 0.25° spatial
334 resolution for the period 2001–2016. Meteorological input data were derived from the WATCH
335 Forcing Data methodology applied to ERA-Interim data (WFDEI) (Weedon et al., 2014), which
336 has also been utilized in the Inter-Sectoral Impact Model Intercomparison Project Phase 2a
337 (ISIMIP2a; Gosling et al., 2019). Soil property data were sourced from the Global Soil Dataset
338 for Earth System Modeling (GSDE), originally provided at a spatial resolution of 30 arc-seconds
339 (Dai et al., 2019; Shangguan et al., 2014). Land cover data were derived from the MODIS
340 dataset (MCD12Q1; Friedl and Sulla-Menashe, 2022), providing detailed global land
341 classification information at a spatial resolution of 500 meters.

342 The simulation of irrigation processes also required detailed data on crop areas, planting dates,
343 irrigation areas and irrigation methods. Crop planting areas were derived from the 30-meter
344 resolution CropScape and Cropland Data Layer (CDL) datasets (2008–2020) and aggregated to a
345 spatial resolution of 5 arcminutes for analysis (USDA, 2019). These datasets, produced by the
346 U.S. Department of Agriculture, provide annual, crop-specific land cover information using
347 satellite imagery and ground reference data. For each pixel, we calculated the proportion of
348 cropland relative to the pixel's area (PCT_CROP) and the proportions of maize, wheat, and
349 soybean relative to the cropland area (PCT_CFT). Pixels with a cropland percentage
350 (PCT_CROP) exceeding zero were classified as crop pixels. The Plant Functional Type (PFT)
351 approach employed in CoLM allowed different crops and vegetation types coexist within the
352 same grid cell according to their percentages (PCT_CFT). To define planting and harvesting
353 dates, we utilized an observation-based crop calendar dataset from the Global Gridded Crop
354 Model Intercomparison (GGCMI), which provided information for 20 major crops under both
355 rainfed and irrigated conditions at each grid cell for 1980–2010 (Jägermeyr et al., 2021).

356 The irrigation map was derived using the 5' resolution data from the FAO Global Map of
357 Irrigation Areas - Version 5 (Siebert et al., 2013). Since the CropScape data does not distinguish
358 between rainfed and irrigated crops, we combined it with the irrigation map to determine the
359 proportions of rainfed and irrigated crops. Irrigation water withdrawals were classified into
360 surface water and groundwater sources following FAO data on regions equipped for groundwater
361 extraction, which informed the allocation of irrigation demand across sources (Siebert et al.,
362 2010). The irrigation application method data was obtained from Yao's global irrigation map,
363 which details irrigation methods (drip, sprinkler, or flood) for 32 crop types, each assigned a
364 single method (Yao et al., 2022). Jägermeyr et al. (2015) originally used a decision tree to refine



365 AQUASTAT's data, classifying irrigation methods for 14 Crop Functional Types (CFTs) based
366 on crop area, soil characteristics, and socio-economic conditions. Yao et al. (2022) then matched
367 these CFTs to 32 crop types in CLM5 and incorporated an additional irrigation method, *paddy*,
368 specifically for rice-growing regions, creating a more detailed global irrigation dataset.

369 For river routing simulations in CaMa-Flood, the baseline topography was derived from the
370 Multi-Error-Removed Improved-Terrain Hydrography dataset (MERIT Hydro; Yamazaki et al.,
371 2019). Fundamental information on dams/reservoirs in the river network, including dam name,
372 coordinates, total storage capacity, and drainage area, was obtained from the GRanD database
373 (Lehner et al., 2011). GRanD version 1.3 contains data on 7,320 dams globally, along with their
374 associated reservoirs. The locations of the dams in the 0.25° river map were determined
375 following the method outlined by Hanazaki et al. (2022), which enabled the identification of
376 1464 reservoirs across the contiguous United States (Figure S2). In addition to GRanD, the
377 Global Reservoir Surface Area Data Set (GRSAD; Zhao and Gao, 2018) and the Global
378 Reservoir Geometry Database (ReGeom; Yigzaw et al., 2018) were used to estimate reservoir
379 parameters, such as storage capacity at emergency, flood control, and critical levels (Hanazaki et
380 al., 2022). GRSAD provides a monthly time series of surface areas for 6,817 GRanD reservoirs
381 from 1984 to 2015, based on global surface water occurrence data (Pekel et al., 2016). ReGeom
382 contains storage-area-depth information for 6,824 reservoirs in GRanD, with geometry estimates
383 derived from assumed surface and cross-sectional shapes, as well as data on reservoir extent,
384 total storage, and surface area.

385 **2.3.2 Validation datasets**

386 To evaluate the scheme developed in this study, we focused on validating irrigation water
387 withdrawal volumes, land fluxes (including energy fluxes and river flows) and crop yields in
388 irrigated areas. We used hydrological survey data from the U.S. Geological Survey (USGS,
389 2023), which provided detailed statistics on total irrigation water withdrawals, categorized by
390 surface and groundwater sources, every five years since 2000. Within the timeframe of this
391 study, data were available for the years 2005, 2010, and 2015. Building on this, Ruess et al.
392 (2024) employed a global hydrological model (PCR-GLOBWB) to estimate annual, crop-
393 specific irrigation water withdrawals from 2008 to 2020. Additionally, we compared the
394 irrigation water withdrawal volumes simulated by our model with those generated by six other
395 hydrological models—VIC, PCR-GLOBWB, MATSIRO, LPJmL, H08, and DBH—that
396 participated in ISIMIP2a (Gosling et al., 2019). Although more hydrological models were
397 included in ISIMIP2a, our comparison was limited to these six because they provided irrigation
398 water withdrawal outputs. The simulations were driven by the WFDEI climate dataset, with a
399 spatial resolution of 0.5° and covering the period from 1971 to 2010.

400 For land surface flux validation, we used monthly latent and sensible heat fluxes provided by
401 FLUXCOM at a resolution of 0.5° (Jung et al., 2019). FLUXCOM leveraged FLUXNET site



402 observations and extended these globally through machine learning algorithms, resulting in a
403 global dataset for latent heat, sensible heat, and carbon fluxes. For temperature validation, we
404 used land surface temperature data from 2001 to 2016 at a spatial resolution of 0.1° from the
405 ERA5-Land reanalysis dataset (Muñoz-Sabater et al., 2021).

406 For streamflow validation, we utilized monthly streamflow data from the Global Runoff Data
407 Centre (GRDC, 2023) for the period 2001–2016. To ensure robust validation, we excluded
408 catchments with fewer than five years of data during the study period and focused on catchments
409 significantly influenced by irrigation while minimizing the impacts of other anthropogenic
410 activities. These selection criteria ultimately resulted in 77 catchments being included in the
411 analysis (Figure S7).

412 For crop yield validation, we relied on annual yield reports for irrigated and rainfed crops from
413 the USDA NASS at the county level, which is regarded as a reliable source of yield statistics
414 (USDA/NASS, 2023). The data for irrigated crops primarily covered the Central Plains of the
415 United States, with limited coverage in the eastern and western regions. We aggregated our grid-
416 based yield simulation results to the county level and performed validation only for regions and
417 years with available USDA data.

418 **2.4 Experimental Design**

419 This study conducted three simulation experiments to evaluate the effectiveness of the newly
420 developed module by comparing their performance: (i) Non-Irrigation Experiment (abbreviated
421 as *noirrig*): This scenario assumes all crops in the region are rainfed, with no irrigation applied. It
422 serves as a baseline to represent natural surface water and energy balance conditions. (ii)
423 Unlimited Irrigation Experiment (abbreviated as *noirrig-unlim*): This scenario distinguishes
424 between irrigated and rainfed areas based on crop maps. In irrigated areas, crop water demands
425 are fully satisfied throughout the growing season, without considering the limitations of water
426 resources. (iii) Limited Irrigation Experiment (abbreviated as *irrig-lim*): In this scenario,
427 irrigation water is supplied proportionally from surface water and groundwater based on
428 availability, as illustrated in Figure 1. Here, irrigation is constrained by the availability of surface
429 and groundwater, which may result in unmet crop water demands.

430 The non-irrigation experiment was first simulated for the period 2001–2010 to stabilize
431 vegetation carbon and nitrogen pools, soil moisture, and the groundwater table. This stabilized
432 state served as the initial condition for all three experiments. The main simulation period spanned
433 2001–2016, covering the contiguous United States at a spatial resolution of $0.25^\circ \times 0.25^\circ$. In the
434 subsequent analysis, key evaluation metrics—bias, root-mean-square error (RMSE), Pearson
435 correlation coefficient (r), and Kling-Gupta efficiency (KGE)—were employed to assess the
436 performance of the simulations.



437 3. Results

438 3.1 Evaluation of simulated irrigation water withdrawal

439 3.1.1 Comparison with observations

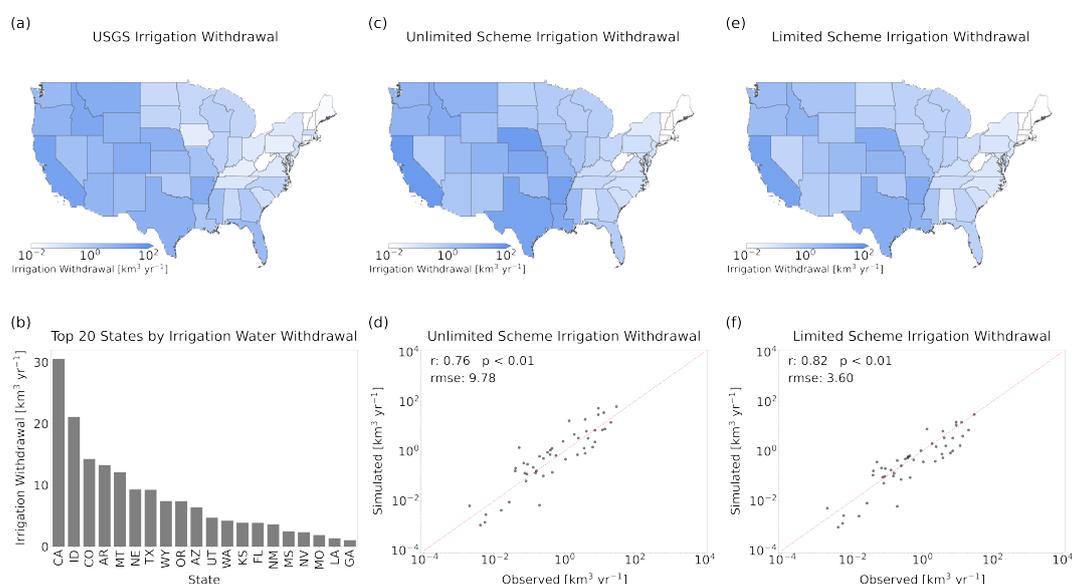
440 Based on annual irrigation withdrawal data from the USGS, states in the western and central
441 United States withdraw significantly more water for irrigation than those in the eastern regions
442 (Figure 3a). This is primarily due to the relatively low precipitation in the western and central
443 regions, where the majority of irrigated areas are located, while crops in the eastern U.S. are
444 predominantly rainfed. The top five states with the highest annual irrigation withdrawals—
445 California (CA), Idaho (ID), Colorado (CO), Arkansas (AR), and Montana (MT)—are all
446 situated in the Midwest and West (Figure 3b). Nationally, the total annual irrigation withdrawal
447 averages approximately $166.23 \text{ km}^3 \text{ yr}^{-1}$, based on data from 2005, 2010 and 2015. In
448 comparison, the irrig-unlim and irrig-lim schemes simulate national total withdrawals of 290.94
449 $\text{km}^3 \text{ yr}^{-1}$ and $120.81 \text{ km}^3 \text{ yr}^{-1}$, respectively. As illustrated in Figure 3c-f, the simulations capture
450 the spatial patterns of water withdrawals across different states effectively, with the irrig-lim
451 scheme yielding better performance. The root-mean-square-error (RMSE) and correlation
452 coefficient (r) for the irrig-lim scheme are $3.60 \text{ km}^3 \text{ yr}^{-1}$ and 0.82 , respectively, slightly
453 outperforming the corresponding values for the irrig-unlim scheme ($9.78 \text{ km}^3 \text{ yr}^{-1}$ and 0.76).

454 Irrigation water withdrawals draw from both surface water and groundwater sources. According
455 to USGS reports, most irrigation withdrawals in the central U.S. come from groundwater (Dieter
456 et al., 2018). In states such as Missouri (MO), Kansas (KS), Iowa (IA), Illinois (IL), Rhode
457 Island (RI), and Mississippi (MS), the share of groundwater withdrawals exceeds 90% (Figure
458 4c). In contrast, states with high surface water withdrawals are primarily in the eastern and
459 western U.S., with states like Wyoming (WY), Connecticut (CT), Kentucky (KY), and Montana
460 (MT) reporting surface water withdrawal proportions greater than 90%. These spatial variations
461 in water source usage are primarily attributed to the central U.S.'s abundant groundwater
462 resources and widespread groundwater extraction infrastructure.

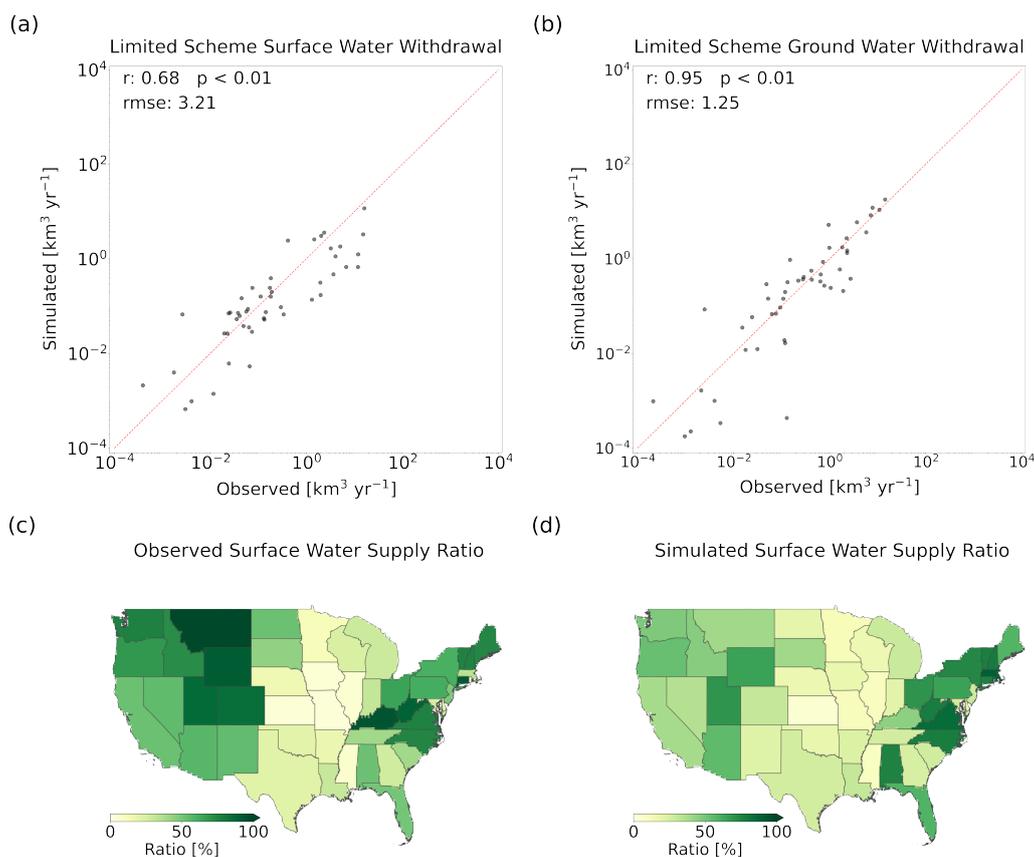
463 In our simulations, the irrig-lim scheme effectively accounts for irrigation water withdrawals
464 from different sources, constrained by their availability. Encouragingly, the scheme generally
465 reproduces observed annual surface water and groundwater withdrawals across states (Figure 4a-
466 b), achieving correlation coefficients of 0.68 and 0.95 , respectively. Furthermore, the simulated
467 proportions of water sources closely align with observed data (Figure 4c-d), with a correlation
468 coefficient of 0.64 ($p < 0.01$). However, the model tends to underestimate the surface water
469 withdrawal proportions in the northwestern regions of the U.S. (particularly in Montana and
470 Colorado; Figure 4d), while slightly overestimating them in some central and eastern states. This
471 discrepancy may stem from limitations in the data used to allocate water demand. Specifically,
472 the model relies on pre-determined groundwater extraction infrastructure proportions, which may



473 not accurately reflect actual extraction practices, particularly as the dataset was published in
 474 2005 and may not account for subsequent changes in groundwater infrastructure in certain states.
 475 Alternatively, the discrepancy could arise from model biases in simulating surface and
 476 groundwater availability. For example, in the northwestern region, surface runoff is heavily
 477 influenced by snowmelt and glacial meltwater (Li et al., 2017), and biases in simulating these
 478 processes could lead to an underestimation of surface water availability.



479 **Figure 3.** Comparison of reported and simulated irrigation water withdrawal in the United States.
 480 (a) Annual irrigation water withdrawal reported by the USGS for individual states. (b) Annual
 481 withdrawal amounts for the top 20 states by irrigation water withdrawal. (c) Annual irrigation
 482 water withdrawal simulated by CoLM using the unrestricted water supply (irrig-unlim) scheme
 483 for individual states. (d) Comparison of reported and simulated irrigation water withdrawal
 484 (using the irrig-unlim scheme) for individual states, with Pearson correlation coefficient (r) and
 485 root mean square error (RMSE) displayed, along with statistical significance (two-tailed
 486 Student's t -test). (e) Annual irrigation water withdrawal simulated by CoLM using the restricted
 487 water supply (irrig-lim) scheme for individual states. (f) Comparison of reported and simulated
 488 irrigation water withdrawal (using the irrig-lim scheme) for individual states.



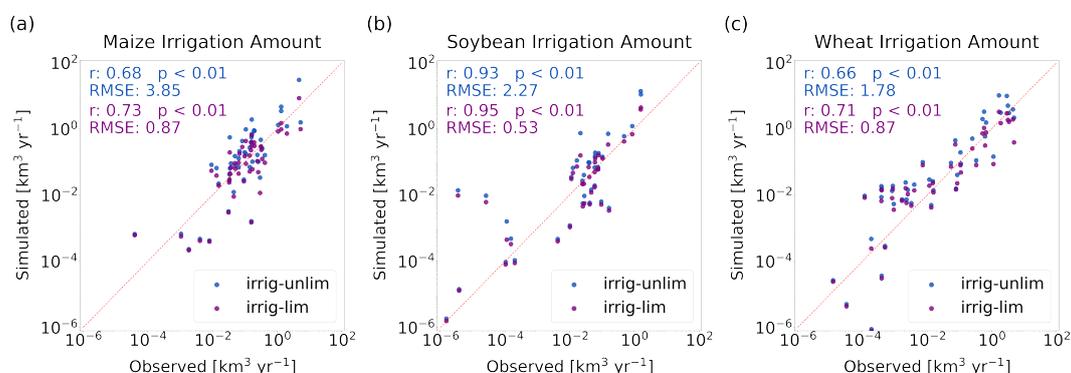
490

491 **Figure 4.** Comparison of reported and simulated irrigation water withdrawal in the United States
 492 by water source. (a) Comparison of reported and simulated surface water withdrawal volumes for
 493 individual states. (b) Same as (a), but for groundwater withdrawal volumes. (c) Proportion of
 494 surface water in irrigation withdrawal, based on USGS reports for individual states. (d)
 495 Proportion of surface water in irrigation withdrawal, simulated by CoLM using the irrig-lim
 496 scheme for individual states.

497 Ruess et al. (2024), using data from the USGS and model outputs from PCR-GLOBWB 2,
 498 generated an irrigation water withdrawal dataset that included withdrawal volumes for major
 499 crops in the U.S. According to this dataset (Figure 5), wheat is the largest consumer of irrigation
 500 water, with an average annual withdrawal of approximately $27.29 \text{ km}^3 \text{ yr}^{-1}$, followed by maize at
 501 about $20.91 \text{ km}^3 \text{ yr}^{-1}$. In contrast, soybean requires considerably less irrigation (i.e., $5.89 \text{ km}^3 \text{ yr}^{-1}$),
 502 partly due to its greater drought tolerance and smaller planted area compared to the other two
 503 crops. Under the irrig-unlim (irrig-lim) simulation scheme, the annual irrigation withdrawals for
 504 maize, wheat, and soybean are $53.98 \text{ km}^3 \text{ yr}^{-1}$, $47.53 \text{ km}^3 \text{ yr}^{-1}$, and $29.99 \text{ km}^3 \text{ yr}^{-1}$ ($19.19 \text{ km}^3 \text{ yr}^{-1}$,



505 17.95 km³ yr⁻¹, and 11.05 km³ yr⁻¹), respectively. Once again, the irrig-lim scheme provides a
 506 closer alignment with observation-based data, as indicated by a lower RMSE (Figure 5). These
 507 results suggest that our irrigation module generally performs well in simulating total national
 508 annual water withdrawals, the spatial distribution of withdrawals (Figure 3), the proportion of
 509 water source types (Figure 4), and the irrigation volumes for different crops (Figure 5).



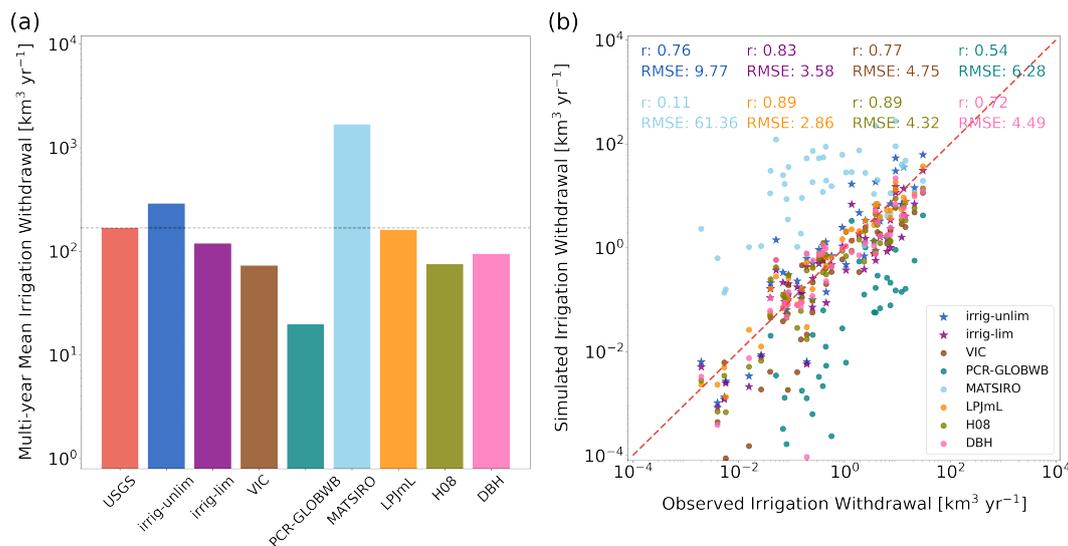
510
 511 **Figure 5.** Comparison of reported and simulated irrigation water withdrawal in the United States
 512 by crop type. (a) Comparison of reported and simulated irrigation water withdrawal for maize,
 513 using both the unrestricted (irrig-unlim, blue dots) and restricted (irrig-lim, purple dots) supply
 514 schemes for individual states. (b-c) Same as (a) but for soybean and wheat.

515 3.1.2 Comparison with other models

516 We further compare the irrigation water withdrawal simulations from this study with outcomes
 517 from six global hydrological models (VIC, PCR-GLOBWB, MATSIRO, LPJmL, H08, and
 518 DBH) that participated in ISIMIP2a. Notably, all simulations used the same climate forcing
 519 (WFDEI), ensuring consistency in the comparison. Our results, particularly from the irrig-lim
 520 scheme, closely align with observed total national annual irrigation withdrawals. By contrast,
 521 five of the six models, excluding LPJmL, exhibit larger absolute deviations from observed value
 522 (Figure 6a). Regarding spatial distribution, most models perform well (Figure 6b), with LPJmL
 523 (orange dots) achieving the highest correlation coefficient (0.89) and the lowest RMSE (2.86 km³
 524 yr⁻¹). The irrig-lim scheme in this study (purple dots) performs comparably to LPJmL,
 525 demonstrating competitive accuracy. In terms of temporal dynamics, comparisons across models
 526 are limited due to the scarcity of observed data. However, the general seasonal patterns are
 527 consistent across models (Figure S5), with the highest irrigation withdrawals occurring in June
 528 and July, and the lowest in January and December. Most models exhibit similar seasonal
 529 fluctuations, with irrigation volumes during peak months approximately ten times greater than
 530 during off-peak months. Overall, these results suggest that our model performs similarly to, or
 531 even better than, existing models in simulating irrigation water withdrawals in the U.S.



532



533

534 **Figure 6.** Comparison of irrigation water withdrawal simulated by CoLM and six global
 535 hydrological models participating in ISIMIP2a. (a) Annual total irrigation water withdrawal
 536 amounts in the United States as reported by the USGS, compared with simulations from CoLM
 537 (using both the irrig-unlim and irrig-lim schemes) and the six global hydrological models. (b)
 538 Comparison of reported and simulated irrigation water withdrawal for individual states, with
 539 Pearson correlation coefficient (r) and root mean square error (RMSE) for each simulation
 540 displayed.

541 3.2 Evaluation of simulated land energy and water fluxes

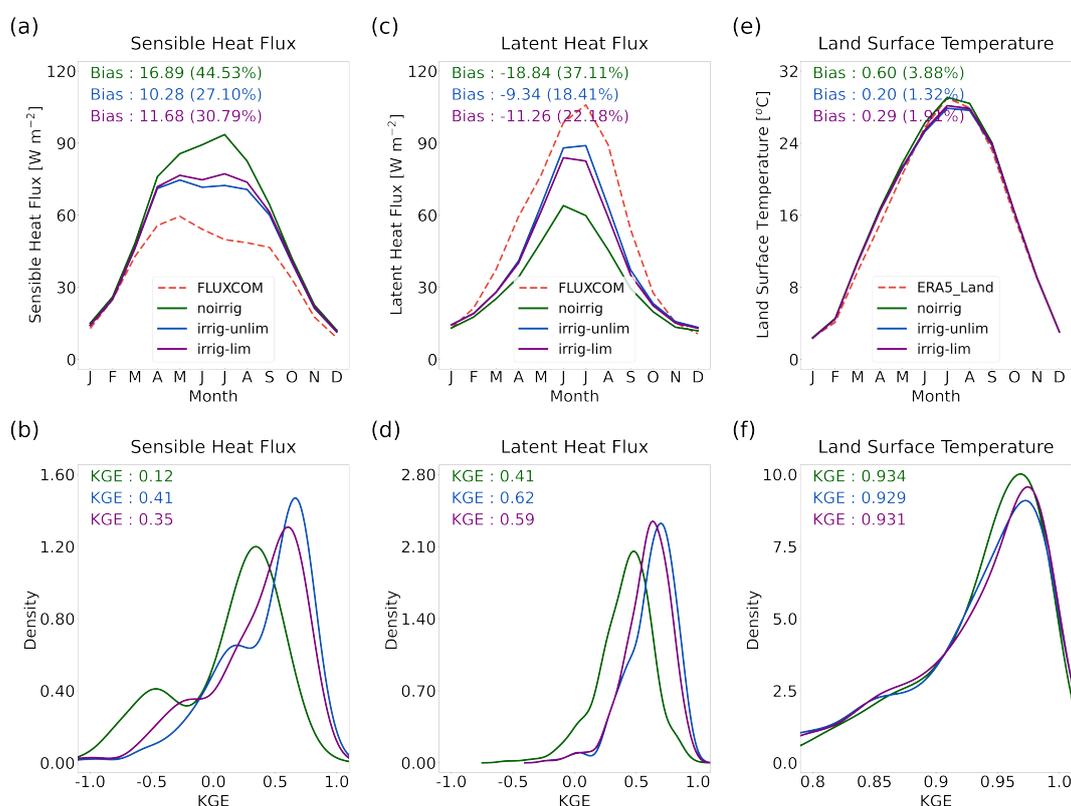
542 3.2.1 Evaluation of simulated energy fluxes

543 We evaluate CoLM's performance in simulating surface energy fluxes over irrigated areas in the
 544 U.S. using different schemes, with FLUXCOM monthly sensible heat (SH) and latent heat (LH)
 545 fluxes as observational references. Figure 7 compares multi-year monthly averages of observed
 546 and simulated SH and LH fluxes across irrigated grid points. Without irrigation (the noirrig
 547 scheme), the model significantly overestimates SH (Figure 7a) with an average bias of 16.89 W
 548 m^{-2} (44.53%) and underestimates LH (Figure 7c) with an average bias of 18.84 W m^{-2} (37.11%).
 549 In contrast, biases over non-irrigated grids are considerably lower, at 3.04% and 17.38% for SH
 550 and LH , respectively (Figure S6). This indicates that CoLM performs satisfactorily in simulating
 551 energy processes over natural vegetation and rainfed areas, but less so over irrigated regions.

552 Upon introducing the irrigation module, the simulation errors in surface energy fluxes over
 553 irrigated areas are significantly reduced. Under the irrig-unlim and irrig-lim schemes, average SH
 554 biases decrease to 27.10% and 30.79%, respectively, while LH biases decrease to 18.41% and



555 22.18%. These improvements are evident across most irrigated grid points, as illustrated by the
 556 KDE curves of KGE, which indicate an increase in grid points with higher KGE values (Figure
 557 7). A KS test confirms that the differences between the irrigation (i.e., the irrig-unlim and irrig-
 558 lim schemes) and noirrig simulations are statistically significant. Although the irrig-unlim
 559 scheme performs slightly better than irrig-lim for *SH* and *LH*, this difference is not significant.



560
 561 **Figure 7.** Evaluation of simulated energy fluxes and land surface temperature in the irrigation
 562 region. (a) Monthly sensible heat flux averaged from 2001 to 2016, based on FLUXCOM dataset
 563 and simulated by CoLM using the noirrig, irrig-unlim, and irrig-lim schemes in irrigation regions
 564 of the United States, with the bias between simulations and observations (i.e., FLUXCOM)
 565 indicated in the panel. (b) Kernel density estimate (KDE) curves for the Kling-Gupta efficiency
 566 (KGE) between observed and simulated monthly sensible heat flux for each irrigation grid, with
 567 mean KGE value indicated in the panel. (c-d) Same as (a-b) but for latent heat flux. (e-f) Same as
 568 (a-b) but for land surface temperature, using data from ERA5-Land reanalysis dataset.

569 Additionally, the FLUXCOM data (red dashed line) show that the highest monthly *SH* and *LH*
 570 occur in May and July, respectively. However, the noirrig simulation (green solid line) fails to
 571 capture this seasonal peak, showing instead that *SH* peaks in July and *LH* in June. This



572 discrepancy is not present in non-irrigated areas (Figure S6), suggesting that irrigation in
573 agricultural regions (and the subsequent crop growth it supports) substantially affects the
574 seasonal pattern of regional energy balance. When the irrigation module is incorporated into the
575 model, these seasonal patterns are more accurately reproduced, with the timing of the simulated
576 peak months aligning more closely with FLUXCOM data (blue and purple solid lines).

577 The incorporation of the irrigation module improves the simulation of energy partitioning in
578 irrigated areas, enabling the model to better capture surface temperature dynamics (Figure 7e).
579 Under the noirrig scheme, the average bias of monthly surface temperature is 0.6°C (3.88%).
580 This bias decreases to 0.20°C (1.32%) with the irrig-unlim scheme and 0.29°C (1.91%) with the
581 irrig-lim scheme. However, even with irrigation included, the simulated total evapotranspiration
582 remains systematically underestimated (Figure 7c). This underestimation is also evident in non-
583 crop areas (Figure S6c), suggesting that it may not be due to limitations in the irrigation module
584 itself but rather to certain deficiencies in CoLM's evapotranspiration simulation approach.

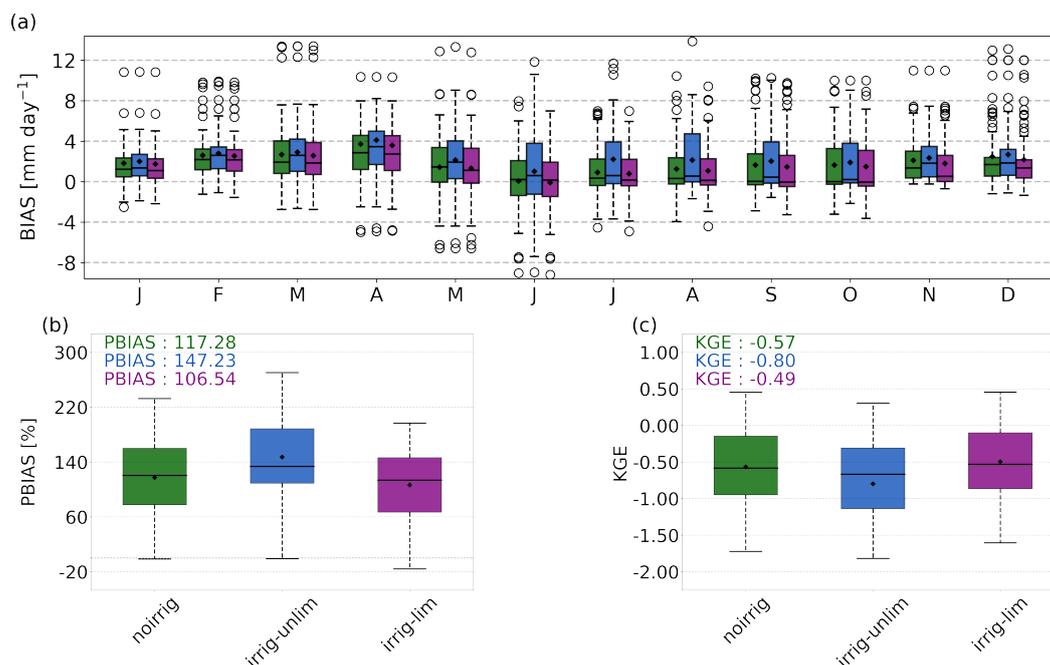
585 **3.2.2 Evaluation of simulated river flow**

586 Irrigation processes can significantly alter natural hydrological dynamics and river flow patterns
587 both temporally and spatially. To evaluate the effectiveness of the irrigation module in capturing
588 these impacts, we compare model outputs with observed catchment streamflow data. We select
589 catchments that are substantially influenced by irrigation while minimizing the effects of other
590 anthropogenic activities. Figure S7 illustrates the locations of the selected 77 catchments. Figure
591 8 shows that CoLM's performance in simulating runoff—and consequently streamflow—
592 remains limited, with relatively low average KGE values across all three schemes. This
593 limitation is likely due to the use of a simplified runoff parameterization scheme in CoLM (Li et
594 al., 2011). However, it is encouraging to note that the irrig-lim scheme notably improves monthly
595 streamflow simulations compared to the noirrig scheme, increasing the average KGE from -0.57
596 to -0.49 and reducing the average percentage bias (PBIAS) from 117.28% to 106.54%. The
597 enhancement can be largely attributed to the incorporation of irrigation effects, which account
598 for reduced streamflow due to increased water use for evapotranspiration. This adjustment
599 effectively mitigates the overestimation of streamflow observed in the noirrig scheme.

600 Furthermore, our analysis reveals that the irrig-unlim scheme significantly reduces the accuracy
601 of streamflow simulations compared to the noirrig scheme, leading to a pronounced
602 overestimation of river discharge. The average relative bias increases substantially from 117.28%
603 to 147.23% (Figure 8b). This issue arises because the irrig-unlim scheme meets any irrigation
604 demand by introducing additional water directly into the system without considering its source.
605 Such an approach is common among crop and land surface models that incorporate irrigation
606 (Malek et al., 2017; Yao et al., 2022; Zhang et al., 2020b). However, our findings indicate that
607 introducing extra water for irrigation without accounting for its specific sources and limitations
608 may lead to an imbalance in the water budget from a comprehensive perspective of the entire



609 water system, undermining the model’s ability to accurately represent the dynamics of the
 610 hydrological system.



611

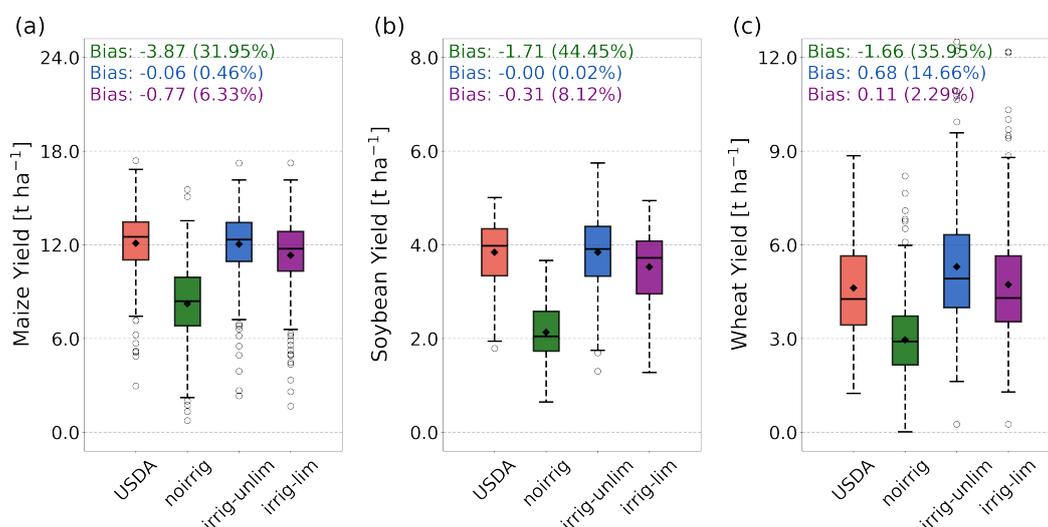
612 **Figure 8.** Evaluation of simulated streamflow in 77 irrigation-affected catchments. (a) Multi-
 613 year average monthly streamflow bias simulated using the noirrig, irrig-unlim, and irrig-lim
 614 schemes in the evaluation catchments. The boxes represent the interquartile range, black lines
 615 indicate median values, black dots show mean values, and dashed black whiskers extend to 1.5
 616 times the interquartile range; points outside the boxes represent outliers. (b) Percentage bias
 617 (PBIAS) between observed monthly streamflow and simulations from CoLM under the noirrig,
 618 irrig-unlim, and irrig-lim schemes, with the average PBIAS value indicated in the panel. (c)
 619 Same as (b) but for the Kling-Gupta efficiency (KGE) between simulated and observed
 620 streamflow.

621 3.3 Evaluation of simulated crop yield

622 Irrigation reflects a direct human influence on crop yields by providing supplemental water. Crop
 623 models primarily focus on this aspect, but they often neglect how irrigation affects other
 624 processes. Conversely, most hydrological models concentrate on the impact of irrigation
 625 withdrawals on the water cycle, with some also addressing energy fluxes, yet pay less attention
 626 to crop yield. From this perspective, land surface models offer distinct advantages; they provide
 627 a more detailed representation of hydrological and surface energy processes compared to crop



628 models, while also presenting more physics-based representations of crop growth than traditional
 629 hydrological models. Therefore, this study further evaluates whether incorporating the developed
 630 irrigation module can enhance crop yield the simulations.



631
 632 **Figure 9.** Evaluation of crop yield simulated using different schemes in the United States. (a)
 633 Maize yield in irrigated maize-growing regions of the United States, as reported by the USDA
 634 (orange boxes), compared with simulations by CoLM using the noirrig (green boxes), irrig-unlim
 635 (blue boxes), and irrig-lim (purple boxes) schemes. Since reported yields are at the county scale,
 636 grid-based simulation results were aggregated to corresponding counties. (b-c) Same as (a) but
 637 for soybean and wheat yields.

638 Using county-scale crop yield data for irrigated and rainfed regions provided by the USDA, we
 639 assess simulated yields under both irrigated and non-irrigated scenarios. The dataset may not
 640 comprehensively cover all irrigated areas in the U.S. or all years during the study period, so
 641 comparisons are limited to regions and years with reported data. In rainfed regions, the model
 642 broadly reproduces average annual yields for the maize, soybean, and wheat (Figure S8).
 643 However, in irrigated regions, the model without irrigation significantly underestimates crop
 644 yields, with average underestimations of 31.95%, 44.45%, and 35.95% for maize, soybean, and
 645 wheat, respectively (Figure 9). Under both the irrig-unlim and irrig-lim schemes, despite slight
 646 differences in performance across crops, the model effectively simulates yield increases under
 647 irrigation, aligning well with observations. Differences between the two irrigation schemes are
 648 minimal: the irrig-unlim scheme performs slightly better for maize and soybean in terms of
 649 average biases, while the irrig-lim scheme shows better performance for wheat.

650 Furthermore, based on limited annual yield data, we observe that considering irrigation generally
 651 improves the model's ability to capture inter-annual yield fluctuations (Figure S9). The KGE of



652 annual yields under the noirrig scheme are -1.342, -1.451, and -1.308 for maize, soybean, and
653 wheat, respectively, while with the irrig-umlim and irrig-lim schemes, the KGE values increase
654 to 0.101, -1.132, and 0.197, and -0.158, -1.449, and -0.144, respectively.

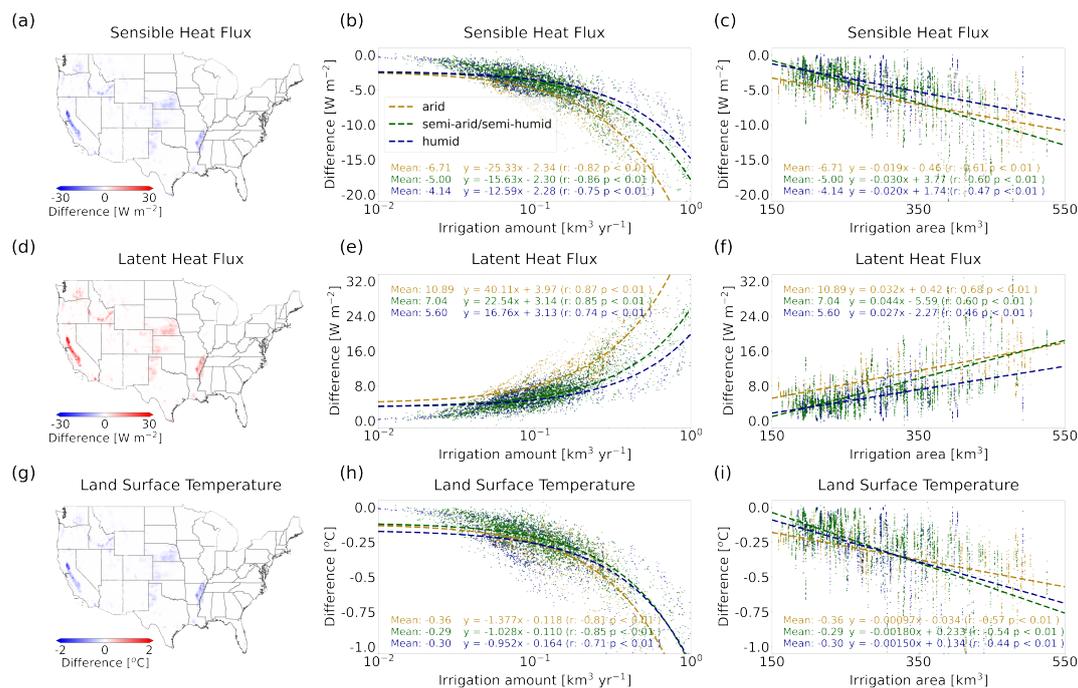
655 4. Discussions

656 4.1 Applications of the developed module

657 4.1.1 Impacts of irrigation on energy budget

658 Numerous studies have highlighted the impacts of irrigation on global and regional energy
659 budgets and near-surface climates. In this study, we similarly examine the effects of irrigation on
660 the energy budget over irrigated areas in the U.S. by comparing results from the irrig-lim and
661 noirrig scheme. Consistent with prior research, we find that irrigation increases latent heat (LH)
662 by 7.53 W m^{-2} (23.25 %) and decreases sensible heat (SH) by 5.18 W m^{-2} (9.48 %) averaged
663 from 2001 to 2016, resulting in an approximately 0.30°C reduction in land surface temperature
664 (Figure 10). Since land-atmosphere coupling is not included, the primary mechanisms driving
665 these impacts are increased soil evaporation due to enhanced soil moisture and greater vegetation
666 transpiration driven by improved crop growth following irrigation (Figure S10 a-b). Annually,
667 these mechanisms contribute roughly equally to the increase in total evapotranspiration in
668 irrigated regions, with pronounced seasonal differences: during the peak growing seasons
669 (summer and autumn), the contribution was dominated by vegetation transpiration, while in other
670 seasons, particularly winter, the increase in soil evaporation plays a larger role in affecting
671 regional energy distribution and temperature (Figure S10c).

672 This study further explores the spatial characteristics of these impacts, analyzing the correlations
673 between irrigated area, irrigation water withdrawal, and changes in LH , SH , and land surface
674 temperature (ΔLH , ΔSH , ΔT_s) across different climate zones. Notably, irrigation has the most
675 substantial impact in arid regions, especially on LH , where ΔLH is more than double that of
676 semi-arid and humid regions, with a larger reduction in temperature by 0.36°C . Interestingly,
677 while previous studies have emphasized irrigated area as the primary determinant of irrigation-
678 induced climate effects (Al-Yaari et al., 2022; Chen and Dirmeyer, 2019), our results indicate
679 that irrigation water withdrawal has a stronger influence on the regional energy budget and
680 temperature. Across all climate zones, ΔSH , ΔLH , and ΔT_s are significantly correlated ($p < 0.01$)
681 with irrigation water withdrawal, with correlation coefficients of -0.81, 0.79, and -0.82,
682 respectively (Figure 10 (b, e and h)), which are higher than the correlations with irrigated area (-
683 0.59, 0.61, and -0.52; Figure 10 (c, f and i)). This emphasizes the critical role of water
684 availability in modulating the climate effects of irrigation.



685

686 **Figure 10.** Impact of irrigation on local energy flux and surface temperature in the United States.
 687 (a) Impact of irrigation on sensible heat flux, quantified by the difference (ΔSH) between the
 688 noirrig and irrig-lim simulation results. (b) Relationship between irrigation amount and ΔSH ,
 689 with grid colors indicating the climate zones (i.e., arid, semi-arid/semi-humid, humid). For each
 690 climate zone, the mean ΔSH , the regression line of irrigation amount versus ΔSH , and the
 691 regression equation are displayed. (c) Same as (b), but for the relationship between irrigation
 692 area and ΔSH . (d-f) Same as (a-c), but for the impact on latent heat flux (ΔLH). (g-i) Same as (a-
 693 c), but for the impact on land surface temperature (ΔT_s).

694 It is important to note that this study employs offline land simulations and does not account for
 695 land-atmosphere interactions, which may introduce biases in the estimated climate impacts.
 696 Future studies should include coupled land-atmosphere simulations to provide a more
 697 comprehensive assessment (Cook et al., 2015; Puma and Cook, 2010; Sacks et al., 2009).
 698 Another aspect worth considering is that some farmers irrigate not only to address water deficits
 699 but also to mitigate heat stress during high-temperature periods (Verma et al., 2020). This
 700 practice can notably affect local temperatures. For instance, surface water temperatures generally
 701 track air temperatures, whereas groundwater temperatures remain relatively stable throughout the
 702 year—typically warmer than air in winter and cooler in summer. This temperature difference,
 703 especially in regions relying on groundwater irrigation, may have non-negligible effects on local
 704 climate that should be incorporated into future modeling efforts.

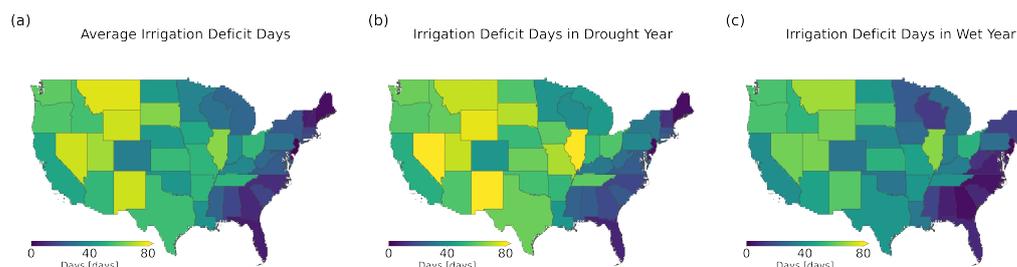


705 **4.1.2 Assessments of irrigation water security**

706 This study compares the irrigation schemes with and without water availability constraints,
707 highlighting the necessity and importance of incorporating water limitations into simulations.
708 Our results demonstrate that including these constraints improves simulation accuracy,
709 particularly in the modeling of water systems. Specifically, irrigation water withdrawal simulated
710 under the irrig-lim scheme aligns more closely with observational data (Figure 3 and Figure 6).
711 Validation against river flow observations further supports the improved performance of the
712 irrig-lim scheme. Importantly, this scheme avoids the risk of potential water imbalances in the
713 modeled hydrological system—an issue commonly associated with non-constrained schemes
714 (Figure 8).

715 Additionally, incorporating water availability constraints more accurately reflects the reality of
716 water resource utilization. By accounting for the interconnections between subsystems within the
717 irrigation water demand-supply system, this approach enables simulation and prediction of
718 irrigation water security issues. Here, we visualize the average number of days when water
719 supply was insufficient to fully meet irrigation demand that simulated by the irrig-lim scheme
720 (Figure 11). Spatially, in humid regions, where irrigation demand is low and water resources are
721 abundant, fewer days of unmet irrigation needs occur. Conversely, in arid regions, where
722 irrigation demand is high and water resources are often limited, the number of unmet irrigation
723 days increases significantly. Figure 11a illustrates that states with a higher number of unmet
724 irrigation days are also those with relatively scarce water resources (e.g., Montana and Nevada).
725 From a temporal perspective, drought years lead to increased irrigation requirements due to
726 reduced precipitation or higher evaporative demand. Although additional water withdrawals can
727 partially address this increased demand, drought conditions often simultaneously result in
728 deficits in both surface and groundwater resources within the water system. As a result, most
729 states experience a substantial increase in unmet irrigation days during drought years (an average
730 of 43 days). In contrast, during wetter years, the number of unmet days decreases significantly
731 (an average of 31 days).

732 Reported disaster data show that even with irrigation, significant crop losses can occur during
733 drought years, aligning with broader water security challenges (Mieno et al., 2024). Our
734 approach effectively captures this phenomenon by describing the connectivity between
735 subsystems in the water demand-supply system and highlighting the impact of water limitations
736 on irrigation. In contrast, ignoring these constraints risks underestimating potential food security
737 issues in a future characterized by more frequent and/or severe droughts. This represents a
738 critical limitation of crop and land surface models that adopt irrigation schemes without
739 considering water availability constraints.



740

741 **Figure 11.** Days per year with unmet irrigation demand (i.e., irrigation deficit days) in the United
742 States simulated by the irrig-lim scheme. (a) Multi-year average irrigation deficit days from 2001
743 to 2015 for individual states. (b) Irrigation deficit days in drought year for individual states. (c)
744 Irrigation deficit days in wet year for individual states. Drought year (wet year) is defined as the
745 year with the lowest (highest) annual precipitation during 2001–2016.

746 **4.2 Limitations and a way forward**

747 While the developed module represents a significant advancement in modeling irrigation water
748 system within land surface models by providing a comprehensive representation of the irrigation
749 processes—including water demand, water withdrawal, and water utilization, several limitations
750 and assumptions should be acknowledged.

751 Irrigation water demand in this study is estimated using a soil moisture deficit method. However,
752 the parameterization of certain key variables (e.g., target and threshold soil moisture levels) is
753 overly simplified and does not account for variations among crop types. These parameters are
754 adjustable, and their calibration could further enhance the model's accuracy in reproducing
755 irrigation water use. Additionally, in some cases, farmers irrigate not only to address soil
756 moisture deficits but also to reduce crop heat stress during high-temperature periods—a factor
757 that should be incorporated into future modeling efforts. Furthermore, this study did not account
758 for water losses during conveyance and application. Irrigation losses, as noted by Jägermeyr et
759 al. (2015), include conveyance losses and on-field application losses. By ignoring conveyance
760 losses, the model assumes that water withdrawn equals water applied, likely leading to an
761 underestimation of total irrigation water use. Field application losses depend on irrigation
762 methods (Leng et al., 2017), and while this study considered four irrigation systems with
763 differentiated efficiencies, the reliance on simplified rules and a coarse irrigation map fails to
764 reflect the diversity of irrigation methods and distributions. For example, actual sprinkler
765 systems distribute water in specific spray patterns rather than uniformly. However, the model
766 assumes uniform water distribution across each Crop Functional Type (CFT). Future models
767 could benefit from parameterizations that capture spatial heterogeneity in irrigation distribution
768 (Jägermeyr et al., 2015; Merriam et al., 1999). Moreover, irrigation water demand also depends
769 on agricultural practices, such as crop types, cropping calendars, and planting intensities. While
770 the model determines crop phenology based on meteorological data, real cropping calendars are



771 influenced by farmers' decisions (Sacks et al., 2010). Incorporating satellite-derived phenology
772 data could better represent these human factors. Addressing these agricultural practices is crucial
773 for improving the accuracy and applicability of irrigation models.

774 In simulations of irrigation water withdrawal, this study provides a detailed representation of
775 reservoir water withdrawal but acknowledges several sources of uncertainty: First, the dataset
776 includes fewer dams than exist, as it focuses primarily on large dams and may lack data due to
777 protection policies. This omission likely contributes to the underestimation of surface water
778 extraction in some states. Second, all dams are assumed to supply irrigation water, although
779 some reservoirs may not serve this purpose. The irrigation areas served by each dam are
780 unknown, and a generalized estimation method is employed in this study, introducing large
781 uncertainties that remain difficult to validate. Third, dam operations are simplified, while in
782 reality, they often involve complex considerations, such as multi-objective operations and
783 coordinated management of multiple reservoirs. Advanced reservoir optimization strategies,
784 which require predictive simulations and prior knowledge of future inflows and demands, are not
785 incorporated into the model, presenting a significant challenge for considering the impacts of
786 complex human decision-making in land surface models.

787 Determining the division of irrigation water withdrawals between surface and groundwater
788 sources, as well as the withdrawal sequence, is also critical. This study allocates irrigation
789 demand based on pre-defined proportions and simultaneously withdraws water from both
790 sources. Surface water demand is met sequentially through local runoff, river discharge, and
791 upstream reservoir storages. This method, employed in models such as ORCHIDEE v2.2
792 (Arboleda-Obando et al., 2024) and E3SM (Zhou et al., 2020), provides satisfactory simulations
793 of water source allocation for irrigation (Figure 4 vs. Figure S11). However, its reliability
794 depends on the accuracy of input data and may underestimate withdrawals if any water source is
795 inadequately represented. Alternatively, some models (e.g., MATSIRO and CLM5; Pokhrel et al.,
796 2012; Yao et al., 2022) do not pre-allocate demand but set a fixed order of water withdrawals—
797 typically prioritizing surface water before groundwater. This method tends to satisfy more
798 irrigation demand and provides better estimates in regions with unreported groundwater
799 extraction. We propose that a hybrid approach, defining surface and groundwater proportions
800 dynamically, warrants consideration in future study. For instance, during wet seasons, surface
801 water extraction proportions could increase to reduce groundwater reliance and associated
802 pumping costs. Conversely, during dry seasons, surface water may be more constrained,
803 necessitating greater reliance on groundwater for irrigation. However, such an approach still
804 needs to address challenges, including unreported groundwater use, data scarcity, and the
805 physical, technical, and socio-economic constraints on groundwater use across regions.

806 Additionally, this study does not account for restrictions beyond water availability, such as local
807 regulations governing water allocation, including water rights and inter-basin water transfers.
808 Alternative water sources, such as desalinated seawater and treated wastewater, also warrant
809 consideration (Vliet et al., 2021). Recent assessments indicate that these unconventional water



810 sources are experiencing exponential growth (Jones et al., 2019). Although their contributions
811 remain low globally, they play a significant role in water-scarce regions. Incorporating these
812 factors into models could further improve simulations of irrigation water security.

813 Finally, the developed module's results and applicability are strongly influenced by the CoLM
814 framework itself. A critical aspect requiring careful consideration is the evaluation and
815 calibration of hydrological variables, such as soil moisture, runoff, river discharge, and
816 groundwater levels, which are essential for water resource modeling. Currently, the CoLM
817 employs the simplified top model (SIMTOP) developed by Niu et al. (2005) for runoff
818 simulations. The excessive simplification of this approach, coupled with the lack of calibration,
819 limits the model's accuracy in runoff simulations. Inadequate representation of snow and glacial
820 melt processes introduces regional biases, particularly in northern and midwestern U.S. states
821 where these factors are pivotal. For instance, surface water extraction is underestimated in some
822 states within these regions, likely because the model fails to accurately capture snowmelt and
823 glacial melt contributions to streamflow, leading to erroneous estimates of surface water
824 availability. Similarly, simulated evapotranspiration is systematically underestimated, even in
825 areas without crops or irrigation, likely due to more complex underlying causes. These biases,
826 when aggregated at the watershed level, result in significant discrepancies in river discharge,
827 thereby constraining the model's applicability for water resource management and its ability to
828 predict irrigation water security. Addressing these issues requires urgent improvements in the
829 representation of related processes, along with further calibration and parameter tuning.

830 **5. Conclusions**

831 The growing challenges posed by increasing global food demand and water scarcity underscore
832 the need for advanced modeling tools capable of accurately capturing human-water interactions.
833 This study contributes to addressing this need by implementing a two-way coupled irrigation
834 water system within the Common Land Model. The developed module provides a
835 comprehensive representation of the entire irrigation water use process, including water demand,
836 withdrawal, and utilization. It introduces a refined multi-source water withdrawal framework and
837 achieves bidirectional coupling between water demand and withdrawal during simulation.

838 The robustness of the new irrigation module is validated through simulations across the
839 contiguous United States, focusing on regional-scale water, energy, and crop yield dynamics. The
840 module effectively simulates total national annual irrigation withdrawals, their spatial
841 distribution, the proportions of different water sources, and irrigation volumes for various crops.
842 Compared to other hydrological models in ISIMIP2a, our model performs similarly or better in
843 simulating U.S. irrigation withdrawals. Incorporating the new irrigation module also
844 significantly improves the accuracy of simulated surface energy fluxes, both in magnitude and
845 seasonal patterns, resulting in more accurate surface temperature predictions. For streamflow, the
846 irrigation scheme accounting for water availability constraints enhances the model's
847 representation of hydrological system dynamics, whereas the unrestricted irrigation scheme



848 introduces potential water budget imbalances. Additionally, the new module markedly improves
849 the model's ability to simulate annual yields and interannual fluctuations of major crops,
850 including maize, soybean, and wheat.

851 We further apply the developed module in two novel analyses. First, the scheme effectively
852 characterizes the climatic impacts of irrigation, revealing a stronger positive correlation between
853 irrigation water volume, rather than irrigated area, and the intensity of irrigation-induced climatic
854 effects. This highlights the critical role of water availability in modulating irrigation-driven
855 climate impacts. Although more accurate simulation of these effects requires land-atmosphere
856 coupled modeling, the enhanced CoLM is clearly ready for such tasks. Second, the module
857 captures irrigation-related water security issues, particularly during drought years, where water
858 shortages across the resource system lead to irrigation water deficits and associated food security
859 challenges. These results demonstrate the promise of CoLM as a valuable tool for future water
860 use and scarcity assessments, paralleling the functionality of global hydrological models and
861 contributing to initiatives such as the Inter-Sectoral Impact Model Intercomparison Project.

862 **Data Availability Statement**

863 The meteorological variables from the WFDEI can be freely accessed from
864 [ftp://rfddata:forceDATA@ftp.iiasa.ac.at](ftp://rfddata.forceDATA@ftp.iiasa.ac.at). The land cover type datasets (MCD12Q1) can be freely
865 accessed from <https://lpdaac.usgs.gov/products/mcd12q1v061/>. The soil characteristics datasets
866 (GSDE) can be freely accessed from <http://globalchange.bnu.edu.cn/research/data/>. The
867 CropScape and Cropland Data Layer (CDL) datasets can be freely accessed from
868 <https://nassgeodata.gmu.edu/CropScape/>. The crop calendar datasets can be freely accessed from
869 <https://zenodo.org/records/5062513/>. The irrigation map and irrigation equipment percentage can
870 be freely accessed from <https://www.fao.org/aquastat/en/geospatial-information/global-maps-irrigated-areas/latest-version/>. The GRanD database can be freely accessed from
871 <https://www.globaldamwatch.org/grand/>. The GRSAD database can be freely accessed from
872 <https://dataverse.tdl.org/dataset.xhtml?persistentId=doi:10.18738/T8/DF80WG/>. The ReGeom
873 database can be freely accessed from <https://zenodo.org/records/1322884/>. The USGS's
874 hydrological survey data can be freely accessed from <https://water.usgs.gov/watuse/data/>. The
875 USDA NASS's agricultural survey data can be freely accessed from
876 <https://quickstats.nass.usda.gov/>. The crop-specific irrigation water withdrawals data can be
877 freely accessed from https://doi.org/10.13012/B2IDB-2656127_V1/. The ISIMIP2a datasets can
878 be freely accessed from <https://data.isimip.org/search/>. The FluxCom datasets can be freely
879 accessed via <ftp.bgc-jena.mpg.de>. The ERA5-Land skin temperature data can be freely accessed
880 from <https://cds.climate.copernicus.eu/datasets/reanalysis-era5-land-monthly-means?tab=download/>. The streamflow data (GRDC) can be freely accessed from
881 <https://www.bafg.de/GRDC/EN/Home/>. CoLM codes are available for download from GitHub
882 (<https://github.com/CoLM-SYSU/CoLM202X/>).



885 **Author contributions**

886 SZ and YD conceptualized and designed the study. SZ and HL collected the data, developed the
887 model, and conducted the analyses. FL and XL provided assistance with model development. HL
888 prepared the figures. SZ drafted the manuscript. All authors contributed to result interpretation
889 and reviewed the final manuscript.

890 **Competing interests**

891 The authors declare no competing interests.

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