

The significant role of snow in shaping alpine treeline responses in modelled boreal forests

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Abstract. Treelines across the Northern Hemisphere are shifting upward and northward in response to global warming, particularly in boreal forests, where climate change progresses more rapidly at high elevations and latitudes. These shifts intensify competition for resources, threaten endemic alpine species, and disrupt established ecological relationships, leading to biodiversity loss. However, significant heterogeneity and regional variation exist in how treelines respond to environmental changes, with many underlying drivers and constraints still poorly understood.

This study aims to enhance understanding of alpine-treeline dynamics at three alpine study sites – located in Canada, Russia, and Alaska – and to improve vegetation model predictions under changing climatic conditions. We evaluated the relative impact of key factors influencing treeline migration velocity and examined the effects of varying snow regimes on treeline migration within the alpine treeline ecotone. To achieve this, we incorporated a novel snow module into the vegetation model LAVESI (*Larix* Vegetation Simulator), enabling the integration of precipitation outside the growing season, snow accumulation, and snowmelt processes. This module allows for explicit modelling of the positive and negative impacts of snow depth on tree growth and treeline migration, while accounting for stochastically occurring extreme events and capturing full weather variability.

Our findings reveal site-specific responses to factors driving treeline shifts and forest expansion, with localised conditions playing a critical role in shaping migration dynamics. The Canadian and the Russian sites demonstrate clear insights into primary migration drivers, while the high variability at the Alaskan site indicates more complex local dynamics and greater predictive uncertainty. The study highlights the significant role of snow in modulating migration potential, as snow accumulation creates favourable conditions for seedling germination and growth while also posing risks of increased mortality from snow loads or avalanches. These results underscore the importance of incorporating snow-related processes into vegetation models to improve the accuracy of predictions for boreal forest dynamics.

Overall, this study provides valuable insights into tree migration processes, highlighting the varied predictability of treeline responses across regions. These findings carry significant implications for refining vegetation models and guiding conservation strategies to sustain alpine tundra resilience in the face of accelerating climate change.

1 Introduction

1.1 Importance of treeline research

Treelines across the Northern Hemisphere are shifting upward and northward in response to global warming (Harsch et al., 2009; Holtmeier and Broll, 2010; Lu et al., 2020). This trend is particularly evident in ~~the~~ alpine treelines ~~within~~ the boreal forests, as climate change is progressing more ~~quickly~~ rapidly at high elevations and latitudes than ~~in other regions~~ elsewhere (IPCC, 2022). Several studies report elevational treeline advancement in the mountainous regions of Eurasia, such as the Altai Mountains (Cazzolla Gatti et al., 2019) and Polar Urals (Devi et al., 2008; 2020), ~~and as well as~~ in North America, including the Kenai Mountains (Dial et al., 2007) and Yukon Territory (Danby and Hik, 2007).

The upward expansion of trees into higher elevations intensifies competition for ~~resources such as~~ light, water, nutrients, and habitat ~~availability~~, threatening endemic alpine species that ~~thrive independent on~~ open tundra or meadow ecosystems. This ~~increased~~ competition can limit the growth and survival of native alpine plants adapted to specific environmental conditions now altered by ~~the presence of trees~~ encroachment (Greenwood and Jump, 2014). ~~Consequently~~ As a result, treeline advancement may ~~lead to the extinction of~~ drive specialised flora and fauna ~~to extinction if they cannot~~ unable to adapt or migrate quickly enough (He et al., 2023), ~~resulting in~~ leading to biodiversity loss and ~~the~~ disruption of ~~established~~ ecological relationships (Hansson et al., 2021). Current studies reveal ~~significant heterogeneity and~~ substantial regional variation ~~in the responses of~~ how both elevational and northern treelines ~~respond to~~ changing environmental ~~conditions~~ changes (Holtmeier and Broll, 2010). Despite ongoing research, the specific drivers and, especially, the constraints ~~affecting~~ shaping how alpine treeline ~~dynamics~~ in boreal forests ~~respond to a changing~~ under climate ~~change~~ are not yet fully understood (He et al., 2023).

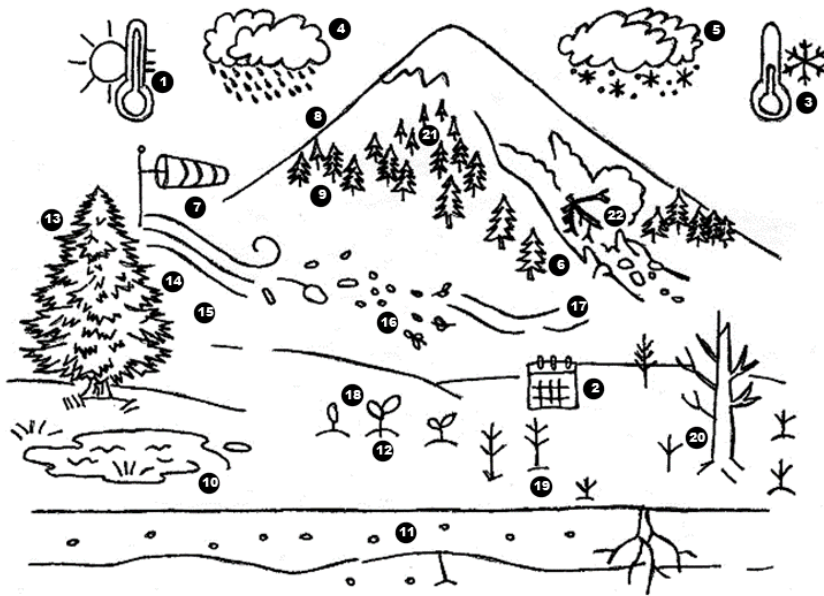
1.2 Time lag in treeline response to climate warming

~~Many factors, both~~ Numerous favourable and unfavourable, influence ~~the tree~~ establishment ~~of trees~~ within the treeline ecotone and, ~~consequently, thus the~~ treeline advance ~~of the treeline~~. First and foremost, climatic factors determine the position of the treeline. During his ~~research~~ expeditions in the late 18th and early 19th centuries, Alexander von Humboldt observed a pronounced correlation between ~~the~~ latitudinal and elevational treelines and corresponding temperatures. Humboldt integrated this revelation into his pioneering concept of isotherms, ~~delineating~~ lines that connecting elevations ~~experiencing with~~ equivalent temperatures. This ~~innovative~~ framework ~~allowed~~ enabled him to systematically connect mountains across the globe based on their ~~respective~~ treelines, ~~thereby~~ revolutionising our understanding of climate ~~ie~~ patterns and their spatial distribution on a global scale. Körner (1998) ~~introduced~~ formulated the "growth limitation hypothesis" (GLH), which explains cool air temperatures during the growing season as the primary factor determining the upper limits of treelines. Contemporary observations ~~of supporting~~ the GLH ~~indicate~~ show that the global treeline ~~follows~~ aligns with a Humboldtian isotherm, ~~where~~ with mean growing season air temperatures (GSAT) ~~are~~ around 6.7°C (Körner and Paulsen, 2004) and 6.4°C (Körner, 2012; Paulsen and Körner, 2014). Treelines warmer than this mean GSAT 6-7°C isotherm (Millar et al., 2020) indicate that ~~factors beyond cold temperatures also significantly influence the treeline~~ tree growth is not solely constrained by cold GSAT and that

other factors may play significant roles in determining the position of the treeline (Maher et al., 2021). Numerous Multiple studies indicate highlight a significant notable time lag between climatic shifts e change progression and the corresponding latitudinal or elevational advance of treeline responses (Holtmeier, 1985; Hofgaard and Wilmann, 2002; Lloyd, 2005; Zheng et al., 2024). The fact that only a little more than half of all global treelines studied around the world have advanced in response to anthropogenic warming (Harsch et al., 2009) further underlines the need to consider integrate factors other than beyond GSAT into vegetation models.

The use of vVegetation models areis crucial for predicting how trees responsesed to changing climateic conditions. Numerous simulation studies have been carried out usingemployed dynamic global vegetation models (DGVMs), such as SEIB-DGVM (Spatially Explicit Individual Based DGVM) (Smith et al., 2001) or LPJ-GUESS (Lund-Potsdam-Jena General Ecosystem Simulator) (Sato et al., 2007). In tThis study, we use the employs LAVESI (Larix Vegetation Simulator) (Kruse et al., 2016), an individual-based, and spatially explicit model LAVESI (Larix Vegetation Simulator) (Kruse et al., 2016), which was originally developed to simulate the latitudinal migration of larch forest-treelines in northern Siberia.

Studies using LAVESI, which already incorporates many of the factorsvariables influencing treeline migration rates, also indicate that vegetation models often overestimate the transition from treeless tundra to forest (Kruse et al., 2016). While temperature is widely recognised as the primary driver of tree growth in this region, dispersal constraints may significantly hinder the northward migration of treelines. These findings underscore the pronounced time lag exhibited byin tree population s in responses to climate change. However, despite the inclusion of many relevant factors in LAVESI, critical essential fine-scale processes that are essential forcritical to accurately estimating simulating vegetation dynamics may still be overlookedunderrepresented.



Abiotic climatic factors

1. Summer temperature
2. Length of the growing season
3. Winter temperature
4. Summer precipitation
5. Winter precipitation
6. Facilitation on tree growth through snow
7. Wind speed

Abiotic non-climatic factors

8. Slope angle
9. Wind exposure
10. Topographic wetness index
11. Active layer depth
12. Availability of suitable seedbeds

Species-dependent factors

13. Age at which a tree becomes mature
14. Factor of seed productivity
15. Permanent seed introduction
16. Seed and pollen dispersal
17. Seed and pollen dispersal over long distances
18. Establishment of seedlings
19. Mortality of seedlings
20. Mortality of trees caused by the effects of ageing

Other biotic factors

21. Intra- and inter-specific competition
22. Natural and anthropogenic disturbances

Figure 1: Overview graphic for the key factors influencing alpine treeline migration.

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1.3 Factors influencing alpine treeline migration

The establishment and upward expansion of trees are governed by a complex interplay of abiotic climatic influences, species-specific traits, abiotic non-climatic constraints, and biotic interactions (Figure 1).

Abiotic climatic factors

90 ~~Treeline migration is strongly linked to rising~~Tree establishment and upward expansion of trees correlate closely with the
~~increase in~~ summer temperatures (Almqvist et al., 1998; Holtmeier and Broll, 2007; Kullman, 2007; Devi et al., 2008; 2020;
Grigoriev et al., 2022). ~~Longer and prolonged~~ growing seasons, ~~where with~~ an increase in the number of warm days lead to
~~an increase in~~enhances photosynthetic performance (Holtmeier and Broll, 2007). In addition to GSAT, winter temperatures
also plays a critical role, ~~as in influencing treeline migration. M~~mild winters ~~can reduce seedling mortality rates among trees,~~
95 ~~thereby promoting the establishment of tree seedlings (Kulman, 2007). Conversely, whereas late and early and late~~ frosts
hinder tree establishment (Kulman, 2007; Holtmeier and Broll, 2007). Furthermore, the frequency of freezing ~~Frequent freeze-~~
~~thaw~~ cycles may pose a greater risk of frost damage to trees than ~~prolonged the duration of exposure to~~ low temperatures
(Gross et al., 1991).
~~It is worth noting that the continued increase in~~Elevated atmospheric CO₂ levels, ~~however, is unlikely to do not~~ significantly
100 ~~enhance tree growth or the advancement of the treeline on a global scale, as trees within the treeline ecotone are not limited~~
~~by carbon availability is not a limiting factor~~ (Körner, 2003; Wiesner et al., 2019).
~~Optimal s~~Soil moisture conditions are ~~is a~~ crucial for facilitating the upward migration of treelines, particularly given the
substantial ~~determinant, as increased summer precipitation enhances seedling emergence and enhanced survival of seedlings~~
~~associated with increased summer precipitation (Devi et al., 2020). Drought conditions present a substantial obstacle to~~
105 ~~seedling establishment, with limitations from, whereas~~ drought stress likely outweighing the potential benefits of sharply
~~rising temperatures imposes a significant constraint~~ (Bailey et al., 2021). Otherwise, both paludification and waterlogging also
impede ~~the advancement of~~ treelines ~~advance~~ (Crawford et al., 2003; Holtmeier and Broll, 2007). ~~The amount of w~~Winter
precipitation also has a major influence on the elevational treeline. ~~Winters with little~~Low snow ~~fall,~~ especially in wind-
exposed sites, ~~may lead to setbacks to~~can expose seedlings and saplings ~~that are not fully protected by snow cover to~~ climatic
110 ~~stress and herbivory, hindering their survival~~ (Gross et al., 1991; Holtmeier, 2003; 2005a; 2005b; 2007; Kullman, 2005;
Holtmeier and Broll, 2007). ~~In contrast, s~~Snow cover ~~during winter can offer~~provides essential protection, ~~promoting to~~
~~seedlings and saplings, shielding them from climatic injuries and herbivores (Holtmeier and Broll, 2007). As a result of~~
~~increased winter precipitation, forest expansion may commence earlier in snow-rich and consequently, sheltered sites~~
(Holtmeier and Broll, 2007; Kirilyanov et al., 2012; Hagedorn et al., 2014). Increased ~~depth of snow cover during winter~~ snow
115 ~~depth can facilitate the~~enhances seedling establishment of additional seedlings, thereby contributing to the upward expansion
~~of trees aiding treeline advance~~ (Germino and Smith, 2000; Smith et al., 2003; Holtmeier, 2005a; Grigoriev et al., 2022).
~~However, persistent~~Late-lying snow ~~cover resulting from increased precipitation can significantly shorten the growing season~~

and delay tree, which may, in turn, inhibit the establishment of new generations of trees (Holtmeier and Broll, 2007). Nevertheless, ~~although the delayed onset may coincide with warmer snowmelt causes the growth period to begin later in the year, when temperatures are already warmer, while additional snowfall and~~ increased moisture availability during the growing season (Walker et al., 1999). ~~In extremely snow-rich winters with increased snow load and~~ Excessive snowfall can also impede tree growth through destructive snow movements like such as avalanches, snow slides, and snow creep, the potential destruction of tree-damaging crowns and branches can impede tree establishment (Holtmeier and Broll, 2007; Autio, 2006; Holtmeier, 2005a). Furthermore, ~~the presence of~~ wet snow exacerbates damage and promotes the spread of snow fungi, ~~potentially leading to an increase in~~ seedling and sapling mortality (Holtmeier, 2005a; 2005b; Holtmeier and Broll, 2007; Barbeito et al., 2013).

Wind can have very different effects on the advance of the treeline. ~~The dispersal of seeds and pollen depends on wind speed and direction, with~~ low to moderate wind speeds promoting the dispersal of seeds and pollen, ~~but and frequent~~ high wind speeds causing physiological and mechanical stress and frost-drying (Holtmeier, 1985; Holtmeier and Broll, 2007; Kruse et al., 2018). ~~Seed accumulation occurs primarily~~ In areas of extreme wind exposure, exacerbated by higher topography, seeds may disperse and accumulate exclusively within in wind-protected shelters where seed entry is possible (Kullman, 2005; Resler et al., 2005; Anschlag, 2006; Holtmeier and Broll, 2007; Beloiu et al., 2022). ~~Reduced growth on the windward sides of trees, combined with frost-drying damage, underlines the damaging effects of wind exposure (Holtmeier, 1985).~~

Abiotic non-climatic factors

Topography, geomorphological processes, and local site conditions, including microclimate and soil characteristics such as temperature, moisture, and mineralisation, can exert a greater influence on treeline advancement than rising temperatures (Holtmeier and Broll, 2005; 2007). In steep-sided high mountain valleys, treeline elevation is largely shaped by regional orography, with factors like avalanches, mass wasting, unstable slope debris, and fragmented or absent soil cover preventing forests from reaching their thermal limit (Holtmeier and Broll, 2005; 2007). Permafrost thawing further constrains treeline advance by destabilising steep slopes, and intensifying erosion in potential forest habitats (Lloyd, 2005; Holtmeier and Broll, 2007). Wildfires can generate favourable seedbeds, but also disrupt treelines through altered soil temperatures and snow regimes (Brehaut and Brown, 2022).

Species-dependent factors

As long as there are no disturbances, the height of the treeline is climatically determined by climatic factors (Körner, 2020); ~~though~~ recruitment-related factors are only likely to influence the local appearance and demography of the treeline (Wiesner et al., 2019; Körner, 2020). Nevertheless, ~~the probability of tree recruitment in the current treeline ecotone is also of high importance for the elevation of the treeline and its advance (Wiesner et al., 2019).~~ The success of development from seed to seedling to sapling is contingent upon various factors including the availability of mature, germinable, and viable seeds, seed dispersal and dispersal method, the presence of suitable seedbeds, and the rates of establishment, growth, and survival of

150 seedlings and saplings (Holtmeier, 1995; Juntunen and Neuvonen, 2006; Holtmeier and Broll, 2007; Holtmeier and Broll, 2019). In cold climates, short vegetation ~~can thrive due to~~ the relatively warm microenvironment near the ground, while ~~advanced tree individuals achieve tree size reach maturity~~ only if they ~~can tolerate withstand~~ the surrounding air temperature (Wilson et al., 1987; Grace, 1989). ~~Relatively warm summers promote enhance the production of viable seeds viability, the seedling establishment, of seedlings and the achievement of growth to~~ a size that ~~enables ensures winter~~ survival ~~in the following winter~~ (Karlsson and Weih, 2001; Holtmeier and Broll, 2007). In arid continental climates, ~~particularly on fast draining substrates, the presence of soil moisture becomes is~~ the ~~most critical dominant limiting~~ factor for the development of tree seedlings and saplings (Holtmeier and Broll, 2005; 2007). ~~Variability in traits variability enables the emergence of fosters~~ diverse variants and ~~thus the~~ adaptability to environmental changes. ~~Through inheritance, populations drives adaptation, to their environments and promoting~~ successful traits, ~~enhancing leading to increased~~ population spread and resilience (Holtmeier and Broll, 2007; Gloy, 2023). ~~The more trees in the treeline ecotone suffer from overage~~ ~~aging trees exhibit reduced reproductive capacity (mortality caused by the effects of ageing), the more they are hindered in their ability to reproduce and at a certain age the trees die, limiting treeline advance~~ (Holtmeier and Broll, 2007). Global and regional studies ~~have often overemphasised~~ broad drivers such as temperature, ~~overlooking other crucial variables~~ (Holtmeier and Broll, 2019).

165 ~~The adverse effects of topography, geomorphological processes, local site conditions encompassing microclimate and soil characteristics including temperature, moisture, and mineralisation, as well as the influences of natural and anthropogenic disturbances such as mass outbreaks of leaf-eating insects or autumnal moth, and grazing activity through reindeer or other pasturing animals, may have a more pronounced impact on treeline advancement than the positive effects of increased temperatures (Holtmeier et al., 2004; Holtmeier and Broll, 2005; 2007; Wiesner et al., 2019; Hansson et al., 2021; Brehaut et al., 2022). In many steep-sided high mountain valleys, the elevation at which treelines are found is primarily determined by the orography of the region. Factors such as high avalanche activity, mass wasting, unstable slope debris, and fragmented or absent soil cover hinder forests from reaching their maximum thermal elevational limit in these regions (Holtmeier and Broll, 2005; 2007). The treeline's advance is additionally governed by the occurrence of permafrost. Specifically, the thawing of permafrost in elevated regions induces instability in steep mountain slopes, consequently escalating erosion processes in potential forest habitats (Lloyd, 2005; Holtmeier and Broll, 2007). Under certain conditions, wildfires can generate seedbed environments favourable for seedling emergence. However, low-severity wildfires generate diverse microsite conditions, reducing the likelihood of seeds landing on suitable seedbeds. Moreover, burned treelines may experience alterations in extreme soil temperatures and snow depths, introducing new obstacles to seedling establishment (Brehaut and Brown, 2022).~~

180 Other biotic factors

During ~~the initial phase of early~~ treeline ~~advancement expansion~~, open tree stands ~~are prone to experiencing higher snowpack accumulation compared to previously more snow than~~ treeless areas (Holtmeier, 2005a). ~~The increasing~~ As tree population density ~~increases, in the treeline ecotone creates a warmer microclimate develops, reducing the risk of mitigating~~ summer frost damage by ~~limiting exposure to shielding seedlings and saplings from~~ intense solar radiation after cold nights ~~and/or~~ strong winds, ~~thus facilitating the survival of seedlings and saplings~~ (Örlander, 1993; Germino and Smith, 2000; Germino et al., 2002; Smith et al., 2003; Johnson et al., 2004; Holtmeier 2005a ; Holtmeier and Broll, 2007; Sigdel et al., 2020). ~~In ecotones with~~

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sparse tree cover during early tree invasion, snow depth and its effects may vary abruptly (Holtmeier 2005a). With increasing tree density, intra- and inter-specific competition, including above- and below-ground resource competition, which may ultimately limit further tree increase. Thus, as tree density increases, the favourable effects tend to change to unfavourable effects on tree development and thus treeline migration (Kruse et al., 2016; Zheng et al., 2024). In addition to competition from trees, another significant challenge for tree migration is competition with non-arboreal vegetation, such as grassland and dwarf shrub vegetation, further restricts seedling access to nutrients, moisture, and light as they impede the access of seeds to suitable seedbeds and of seedlings and saplings to light, nutrients and moisture (Holtmeier, 1995; Löffler et al., 2004; Dullinger et al., 2003; 2004; Holtmeier and Broll, 2007; Liang et al., 2016). In contrast, bare mineral soil surfaces offer favourable seedling establishment by reducing competition beds for wind-mediated tree seeds (Hobbie and Chapin, 1998; Löffler et al., 2004), primarily due to the absence or reduction of competition with dwarf shrubs or grass vegetation (Hobbie and Chapin, 1998). Therefore, moderate grazing may actually provide benefits for facilitate tree establishment growth by reducing competing vegetation due to trampling, especially if exclusion of cattle or a reduction in grazing pressure follows tree seed germination (Holtmeier, 1995; Löffler et al., 2004; Holtmeier and Broll, 2007). Beyond plant competition, treeline dynamics are also shaped by natural and anthropogenic disturbances, including mass outbreaks of leaf-eating insects like the autumnal moth and grazing by reindeer and other pasturing animals, both of which affect seedling survival and forest expansion (Holtmeier et al., 2004; Holtmeier and Broll, 2007; Wiesner et al., 2019; Hansson et al., 2021; Brehaut et al., 2022).

The aim of the present study is to enhance our understanding of treeline dynamics and to improve the predictive capabilities of vegetation models in the context of climate change. Consequently, we this study examines evaluate and contrast the relative impact of various factors on the migration velocity of the alpine treeline migration. Furthermore, Specifically, we integrate extended the impact of varying snow depths on trees in the treeline ecotone through the incorporation of a novel snow module into the LAVESI vegetation model by integrating a novel snow module that. This integration enables us to assess the influence of diverse-varying snow regimes-depths on the rate of treeline migration within the treeline ecotone advancement.

This leads us to investigate two key research questions:

1. What are the most significant factors influencing the migration rates of alpine treelines, and how do these factors compare in terms of their respective impacts?
2. Does the rate of upward migration of alpine treelines within the alpine treeline ecotone vary in response to changes in snow regime?

2 Methods

2.1 Model description

In our study, we used the individual-based and spatially explicit model LAVESI (*Larix* Vegetation Simulator) (Kruse et al., 2016). LAVESI, initially developed for *Larix gmelinii* in northeastern Siberia, was designed with the aim of simulating the life cycle of larch species as completely as possible. Following a pattern-oriented modeling, a simple model was iteratively refined to more accurately reproduce observed patterns (Grimm and Railsback, 2005; Grimm et al., 2005). Within one simulation step, which is one year, the relevant processes are incorporated consecutively as submodules. The previous version of LAVESI already incorporates many of the forementioned factors influencing treeline migration rates. At the beginning of each annual cycle, an 'Environment update' takes place. On the one hand, climate variables and a weather index, namely monthly means of temperature and monthly precipitation sums, are processed to estimate daily values. On the other hand, competition is assessed using an annually updated map of tree population density. LAVESI then calculates an annual cycle of seed production, seed dispersal, establishment and competition-dependent tree growth. Finally, ageing and mortality complete a LAVESI simulation step. As part of the validation process, model outputs are compared with actual forest characteristics. While treelines are fitted, this is not done in a spatially explicit manner. LAVESI has been extended to include dispersal processes related to wind speed and direction (Kruse et al., 2018), landscape topography and further boreal forest species (Kruse et al., 2022), adaptability to environmental changes by implementing trait variation and inheritance (Gloy et al., 2023), and forest fire (Glückler et al., 2024). For North America, disturbance effects caused by insects were also included in the model.

2.2 Description of impact factors

In the initialisation phase of each simulation step, which is one year, the monthly weather data are processed by LAVESI (Kruse et al., 2016). The length of the growing season is calculated from the sum of days with a temperature greater than 0°C, called "net degree days" (NDD0), for each simulation year. The summer temperature is based on the sum of the temperatures of the days above 10°C, the so-called "active air temperature" (AAT10). The average January temperature is used as an indicator for the winter temperature. Summer precipitation is the mean precipitation for the months of May to August; winter precipitation – e.g. temperature-dependent snowfall and accumulation – is part of the new snow module. A wind module, where seed and pollen dispersal are controlled by wind speed and direction, has been implemented in LAVESI-Wind (Kruse et al., 2018). Increased tree mortality due to wind exposure is implemented in the snow module of this new version of LAVESI. In LAVESI, the mortality rate is determined by a combination of abiotic and individual factors; once mortality occurs, the tree is removed from the simulated plot (Kruse et al., 2016). Species-dependent factors of seed production, age of maturity, number of seeds introduced, seedling establishment, and seedling mortality are already included as sub-modules in the initial model (Kruse et al., 2016). Wind-dependent seed dispersal distance, which is positively correlated with the height of the releasing

tree, as well as long-distance dispersal, are part of the wind module in LAVESI-Wind (Kruse et al., 2018). The impact of overageing is considered as a specific mortality rate, which is included as a sub-module in the initial version of the model (Kruse et al., 2016). As part of the implementation of topography and landscape sensing of individuals, the slope and the topographic wetness index (TWI) derived from a Digital Elevation Model (DEM) were first implemented in Kruse et al. (2022). The availability of the seedbed is derived from the litter layer height, the dynamics of which were first added to LAVESI in Kruse et al. (2022). The estimation of the maximum active layer thickness, which is strongly coupled to air temperature, was introduced in the original setup of LAVESI (Zhang et al., 2005, Kruse et al., 2016). Tree density dependent mortality due to competition was also included in the original version of LAVESI (Kruse et al., 2016), whereas tree density dependent facilitation is part of the new snow module. Competition with non-arboreal vegetation such as grassland and dwarf shrub vegetation is only indirectly implemented in the seedbed availability factor.

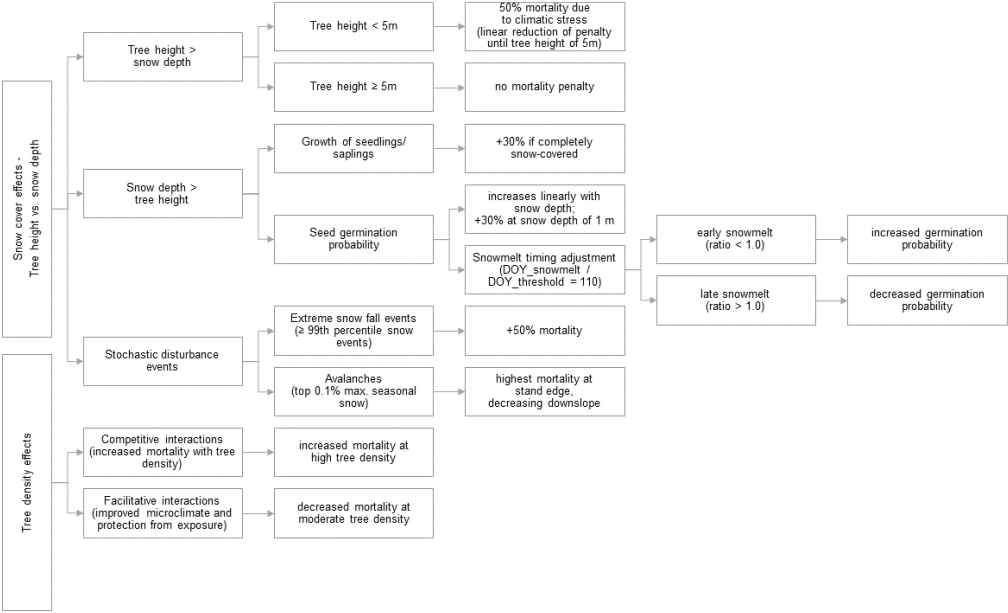


Figure 2: Flowchart illustrating snow-related processes influencing alpine treeline migration, as implemented in the snow module of the vegetation model LAVESI.

2.3 New model development: The snow module

In this study, we examine the driving forces and, in particular, the constraints that influence the responses of alpine treelines. The current version of the LAVESI model successfully simulates the dynamics of the Siberian and North American latitudinal treelines. Since many factors influencing treeline migration are common to both latitudinal and elevational treelines, we adapted LAVESI to account for the specific characteristics of alpine treelines. As part of this adaptation, various effects of snow on treeline migration are incorporated into the model to better capture the unique dynamics of alpine environments.

In recent decades, there have been a number of efforts to incorporate snow dynamics into individual-based gap models. In the SEIB-DGVM (Spatially Explicit Individual-Based DGVM) model, Sato et al. (2007) introduce a method for distinguishing between rainfall and snowfall based on an empirical function of daily mean air temperature. This approach allows for the accumulation of snowfall in a designated snow pool, which subsequently melts in response to soil temperature variations. UVAFME (University of Virginia Forest Model Enhanced) (Wang et al., 2017) has been significantly updated from its original version FAREAST (forest gap model to simulate dynamics and patterns of eastern Eurasian forests) (Xiaodong and Shugart, 2005). However, neither individual-based gap models account for snow accumulation or melt processes. The revised version of UVAFME by Foster et al. (2017) addresses this gap by implementing a simplified snowmelt submodel using the degree-day method proposed by Singh et al. (2000). In ~~this~~^{their} framework, if daily air temperatures drop below 5 °C, precipitation is classified as snowfall and added to the snow cover; conversely, temperatures above 0 °C trigger the melting process for that day.

In previous versions of the individual-based, spatially explicit vegetation model LAVESI (Kruse et al., 2016; 2018; 2022; Gloy et al., 2023; Glückler et al., 2024), precipitation was limited to the growing season months of May to August, neglecting any contributions from precipitation outside this period. Consequently, processes related to snow accumulation and melt were not incorporated into the model. As part of the adaptation of LAVESI for elevational treelines, ~~a range of snow-influenced processes were incorporated into the model, addressing both positive and negative ecological effects across varying snow conditions was enhanced to simulate variable snow scenarios.~~ Snow depth is calculated from precipitation using a snow-to-water ratio, which can vary significantly depending on factors such as the amount of ice in the snow-producing clouds, the structure of snowflakes, and ambient temperature. As a general guideline, we apply an average snow-to-water ratio of 10:1 (National Weather Service, 2024). Following the approach of Sato et al. (2007), ~~which builds on the precipitation partitioning method of Ito and Oikawa (2002),~~ processes of snow accumulation and snowmelt ~~have also been~~^{were} integrated into the LAVESI model. ~~Ito and Oikawa (2002) applied a gradual transition in which 50% of total precipitation is considered snow at 2°C. We initially adopted this approach, but subsequently refined the temperature threshold during model tuning. By comparing simulated and observed snow accumulation patterns, the threshold was iteratively adjusted to improve model accuracy.~~ Unlike models such as SEIB-DGVM and UVAFME, which primarily focus on temperature and moisture dynamics without fully integrating snow effects, the revised version of LAVESI explicitly considers ~~how different~~^{different} snow depths ~~influence variability and its impact on tree growth and treeline migration dynamics.~~ Snow-related processes are governed by the interplay

between snow depth, timing of snowmelt, tree height, and stand structure, all of which influence tree establishment, growth, and mortality (Figure 2).

The effects of varying snow depths on trees within the treeline ecotone have been integrated into the new snow module of LAVESI as follows.

Tree mortality is influenced by the interaction between tree height and snow depth. If the height of the tree is higher than the current snow depth is less than the height of the tree, the mortality rate of the tree increases by 50% due to climatic stress. This penalty decreases linearly with increasing tree height and is eliminated once the tree reaches a height of 5 m.

Conversely, if When seedling or sapling is completely covered by snow, it benefits from insulation, resulting in a 30% increase in growth for that year.

Snow cover also influences germination processes. When seeds are covered by snow, their germination is enhanced, exhibiting a linear probability increases linearly in germination rates with greater snow depth. Specifically, this relationship begins at zero and reaching a maximum of a $\pm 30\%$ increase in germination at a snow depth of 1 m. If seedlings and saplings are covered by snow, their growth is enhanced by 30%. Additionally, the timing of snowmelt influences germination success: The model applies a linear response function that adjusts germination probability germination process is influenced linearly, either positively or negatively, based depending on whether the snow melts occurs before or after the 110th day 110 of the year. The actual day of snowmelt (DOY_snowmelt) is divided by the threshold value (DOY_threshold = 110), and the deviation from 1.0 is added linearly to the germination probability. Early snowmelt (ratio < 1.0) increases the probability, whereas late snowmelt (ratio > 1.0) decreases it. The actual day of snowmelt is divided by 110, and the deviation from 1.0 is added to the germination probability.

Extreme snowfall and avalanche events are incorporated as stochastic disturbances influencing tree mortality and reproduction.

Increased snow load and avalanche occurrences during a given year contribute to higher mortality rates and reduced seed production in the following years. The probability of damage due to extreme snow load damage is calculated based on stochastic daily snowfall events that exceeding the 99th percentile for the respective a given location. When this threshold is surpassed If such an event occurs, affected trees experience a 50% increase in mortality.

Avalanche risk is determined by the from a maximum snow depth during a winter season. The A dynamic threshold is derived from the top 0.1% of total snow depths values across the modelled climate occurring over a historical time series, in the simulation run forms a threshold, which is compared with the current winter season's maximum snow depth to determine the avalanche risk. The risk increases linearly as snow depth rises above the threshold. Mortality is added where a modelled When an avalanche occurs, it induces mortality primary at impacts the leading edge of tree stands, with its destructive force diminishing damage decreasing downslope due to the protective tree growth effect of surrounding vegetation. Trees that are affected by an avalanche will exhibit reduced produce fewer seeds production for up to 10 years, depending on the avalanche's force and the corresponding damage to the tree severity of the impact.

The influence of wind exposure on tree mortality has been included in the snow module. The wind exposure of a tree depends on its elevation and the density of surrounding trees. If trees are highly exposed without surrounding groups of trees to protect them from the wind, their mortality will increase and vice versa.

In the updated version of the vegetation model, we incorporated new processes to account for how tree density modulates both competitive and facilitative interactions under varying snow conditions (Zheng et al., 2024). Tree mortality due to competition – i.e., the negative effect of increasing tree density on tree individual survival – was already included in the previous model versions using competitive strength evaluations (Kruse et al., 2016). Tree mortality depends on tree density due to competition, i.e., as tree density increases, competition for light, water, and nutrients increases, intensifies, resulting in higher tree mortality rates.

However, in this study, we also account for facilitative effects, which may mitigate mortality under certain conditions. Specifically, the positive effect of tree density can improve the microclimate and physical environment – for instance on tree survival, e.g. by providing a protective wind shelter or promoting snow accumulation; – which in turn may reduce exposure-related mortality. This is first included in this study. The facilitation strength is modelled as a function on tree density: it, i.e. as tree density increases, facilitation increases. With increasing facilitation, tree density-dependent mortality is reduced linearly, reaching its maximum reduction at a tree density up to a value of 0.5 (dimensionless units), where it reaches its maximum. Within this range, the facilitative effect linearly reduces tree density-dependent mortality. With further increasing tree density, facilitation strength decreases again until it ceases entirely at a tree density of 1.0. Beyond this a tree density of 0.5, the facilitative effect gradually diminishes and vanishes completely. When the tree density exceeds 1.0, only competitive effects remain active in the model, and facilitation no longer occurs.

This implementation allows the model to represent a realistic balance between competition and facilitation, particularly under snow-affected conditions. It captures how sparse stands may suffer from exposure (e.g. to wind and snow abrasion), while moderately dense stands benefit from protective interactions, and overly dense stands experience intensified competition.

Enhanced winter snow fall also affects soil moisture availability during the growing season, resulting from enhanced winter snowfall. This feedback is incorporated into the general weather processing during the initialisation phase. This adjustment modifies the species-specific weather index, thereby influencing. As a result, the moisture-related climate-growth response of individual tree species is adjusted accordingly.

The LAVESI-Mountain Treelines 1.0 model code, along with the input data used in this study, is available for review on GitHub (https://github.com/StefanKruse/LAVESI/tree/circumboreal_mountain). Upon manuscript acceptance for publication, the code, input data, and simulation output will be permanently archived on Zenodo.

2.4 Study sites

Table 1: Characteristics of the study sites.

	Site CA – Fort McPherson West	Site RU – Lake Ilirney	Site AK – Road to Central
Country	Canada	Russia	Alaska/ USA
Location (Coordinates)	Longitude xmin = -136.001; Longitude xmax = -135.867; Latitude ymin = 67.135; Latitude ymax = 67.187	xLongitude min = 168.128; Longitude xmax = 168.501; Latitude ymin = 67.301; Latitude ymax = 67.437	Longitude xmin = - 145.552; Longitude xmax = - 145.440; Latitude ymin = 65.420; Latitude ymax = 65.467
Mean annual temperature (°C) (1990-2020)	-5.36	-8.61	-0.84
Total annual precipitation (mm) (1990-2020)	221	222	178
Vegetation type	White spruce single trees on steep slope with avalanche damage	Tundra–taiga ecotone, characterised by diverse vegetation: open tundra, shrublands and forest tundra	West to North upslope, a little bit wet, White and Black spruce mixed forest and many <i>Salix</i> and <i>Betula</i> shrubs
Tree species in the region	<i>Larix laricina</i> <i>Picea glauca</i> <i>Picea mariana</i> <i>Betula papyrifera</i> <i>Populus tremuloides</i> <i>Populus balsamifera</i> <i>Pinus contorta</i>	<i>Larix gmelinii</i> <i>Larix sibirica</i> <i>Larix cajanderi</i> <i>Picea obovata</i> <i>Pinus sylvestris</i> <i>Pinus sibirica</i> <i>Betula pendula</i>	<i>Larix laricina</i> <i>Betula pendula</i> subsp. <i>mandshurica</i> <i>Tsuga mertensiana</i> <i>Picea mariana</i> <i>Picea sitchensis</i> <i>Picea glauca</i> <i>Populus tremuloides</i> <i>Populus balsamifera</i>
Total simulation area	1000 x 100 m	1000 x 100 m	1000 x 100 m

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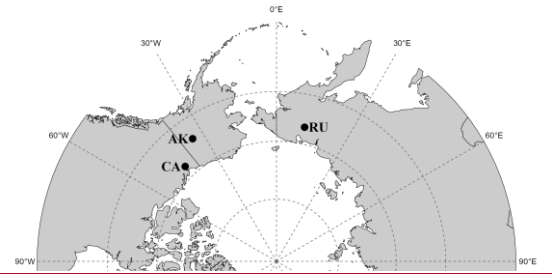
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Table 2: Adjustments to LAVESI simulation settings across study sites.

Plot code / weatherchoice (parameters.txt)	roi (parameters.txt)	seedintroarea (parameters.txt)	simulation area (inc/declaration.h)
5004 - Fort McPherson West	2 – CA	seedintro_min=600; seedintro_max=700; seedintro_minx=0; seedintro_maxx=50;	constexpr unsigned int treerows = 990; constexpr unsigned int treecols = 105;

5005 - Lake Ilirney	1 – RU	seedintro_min=600; seedintro_max=700; seedintro_minx=0; seedintro_maxx=50;	constexpr unsigned int treerows = 1005; constexpr unsigned int treecols = 105;
5009 - Road to Central	3 – AK	seedintro_min=900; seedintro_max=1000; seedintro_minx=25; seedintro_maxx=75;	constexpr unsigned int treerows = 1005; constexpr unsigned int treecols = 105;



365 **Figure 3: Map of the study sites: Fort McPherson West, Canada (CA); Lake Ilirney, Russia (RU); and Road to Central, Alaska, USA (AK).**



370 **Figure 4: Photographs of the study sites: a) Fort McPherson West, Canada (CA); b) Lake Ilirney, Russia (RU); and c) Road to Central, Alaska, USA (AK).**

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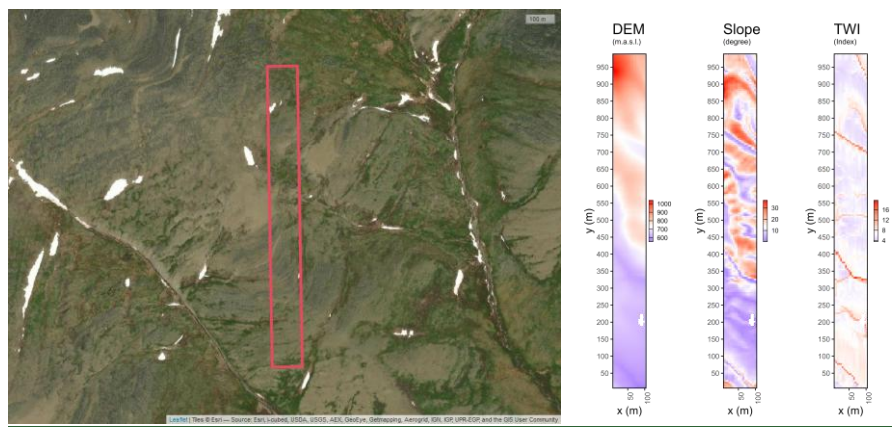


Figure 5A1: Satellite image of the study site Fort McPherson, Canada (CA), including the simulation area (red box). Digital elevation model (DEM), slope, and topographic wetness index (TWI) for the simulation area. Esri, i-cubed, USDA, USGS, AEX, Geoeye, Getmapping, Aerogrid, IGN, IGP, UPR-EGP, and the GIS User Community.

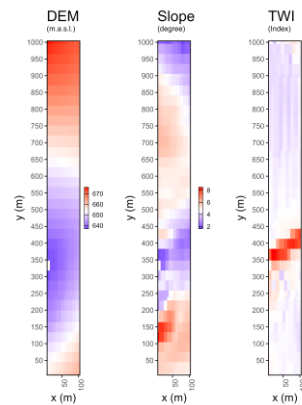
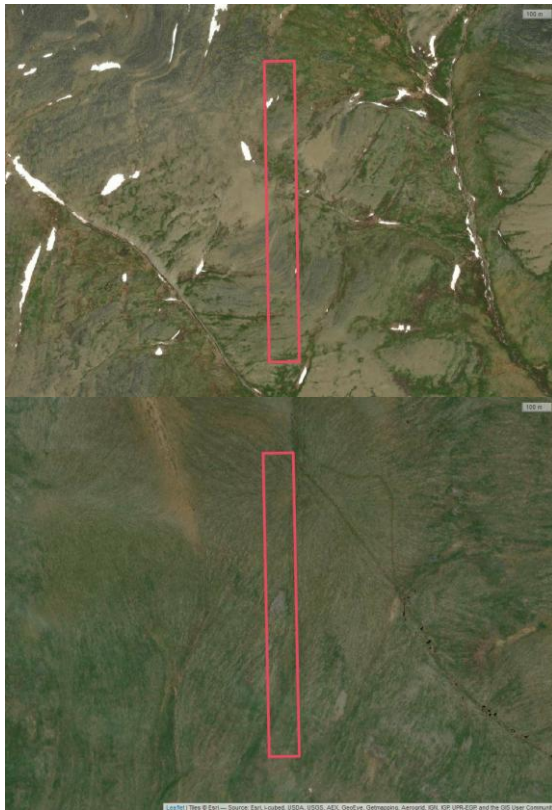


Figure 6B1: Satellite image of the study site Lake Ilirney, Russia (RU), including the simulation area (red box). Digital elevation model (DEM), slope, and topographic wetness index (TWI) for the simulation area. Esri, i-cubed, USDA, USGS, AEX, Geoeve, Getmapping, Aerogrid, IGN, IGP, UPR-EGP, and the GIS User Community.

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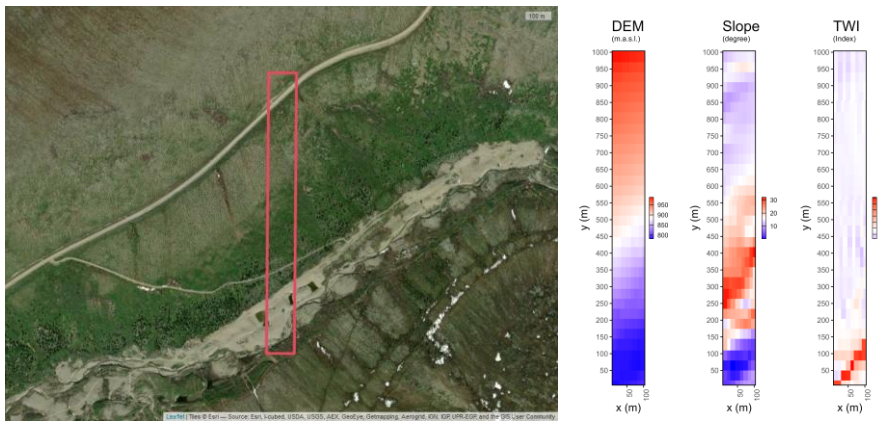


Figure 7C1: Satellite image of the study site Road to Central, Alaska/ USA (AK), including the simulation area (red box), Digital elevation model (DEM), slope, and topographic wetness index (TWI) for the simulation area. Esri, i-cubed, USDA, USGS, AEX, Geoeve, Getmapping, Aerogrid, IGN, IGP, UPR-EGP, and the GIS User Community.

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The model had to be parameterised for the three study sites in Fort McPherson West, Canada (CA), at Lake Ilirney, Russia (RU), and at the Road to Central, Alaska, USA (AK) (see Table 1; Figure 3–4), for which the respective adjustments were made in the input files parameter.txt and declaration.h (see Table 2). Satellite images of the study sites, along with a digital elevation model (DEM), slope, and topographic wetness index (TWI) for the simulation area, are provided in Figures A45–C1 in the Appendix 7.

2.5 Model inputs and simulation scenarios

Monthly precipitation and temperature data for the historical period (1400–1950 CE) were obtained from the Max Planck Institute Earth System Model (Transient MPI-ESM GlacID_P3) (Kapsch et al., 2022). This past climate data was bias-corrected to align with modern data from the CRU-TS dataset (Harris et al., 2020; available for 1900–2020 CE), ensuring consistent means during the overlap period (1900–1950 CE) to avoid discontinuities. The corrected dataset was extended to 2020 CE using CRU-TS. The first 620 years are both to simulate the current conditions of the populations and as the spin up-phase to allow the forests to establish from seeds introduced over the entire area. Climate projections for 2020–2100 CE were derived from the CMIP6 MPI-ESM-2-LR model (Eyring et al., 2016; specific version: r1i1p1f1) and similarly bias-corrected to CRU-TS, using the overlap period (2015–2020 CE) to ensure consistency.

Monthly wind data for 1400–1950 CE was obtained from the transient MPI-ESM Glac1D_P3 model. Since the vegetation model LAVESI required 6-hourly wind data, a downscaling approach was applied. Modern wind data from the ERA5 dataset (Hersbach et al., 2023), covering 1940–2020 CE, was used to create a 6-hourly resolution dataset for the overlapping period and extend the dataset to 2020 CE. ERA5's detailed temporal resolution facilitated realistic downscaling and interpolation. Future wind data from the CMIP6 MPI-ESM-2-LR model (2020–2100 CE), provided at daily resolution, was interpolated to 6-hourly means to meet LAVESI's requirements.

The environmental input data for LAVESI was derived from the Ensemble Digital Terrain Model (EDTM) with a 30 m resolution (Ho et al., 2023). Key information, including elevation, slope, and the topographic wetness index (TWI), was extracted from the DEM of the simulation area.

Simulation runs are performed using a transect representing a slice of the alpine treeline with a size of 10 km x 0.1 km.

To evaluate and compare the influence of various factors on the migration rate of the alpine treeline and to identify key drivers and assess how variations in these factors influence treeline dynamics, a sensitivity analysis was performed. Most of the factors examined in this study have already been implemented in the original LAVESI model (Kruse et al., 2016) or its subsequent versions. These factors are detailed extensively in the corresponding publications. [A comprehensive list of species traits and model parameter values used in this study is provided in the supplementary material.](#) The only newly introduced component in this study was the snow module, which is specifically designed to simulate the impact of snow cover dynamics on the treeline. A description of this snow module is provided earlier in this manuscript. For the sensitivity analysis following Kruse et al. (2018), simulations were run in which the following 17 factors were increased by 10% and decreased by 10% compared to an average reference value: growing season length, summer temperature, winter temperature, summer precipitation, wind speed, seed production, maturation age, seed intro number permanent, dispersal distance, long distance dispersal, seedling establishment, seedling mortality, overageing mortality, slope, twi, seed bed availability, and active layer ([see Figure 1](#)).

In addition, further simulations were run with the forest fire module (Glückler et al., 2024), the insect disturbance module, or the module for trait variation and inheritance (Gloy et al., 2023). The results of each of the simulations were compared with the results of the reference run.

Finally, simulations were conducted using the new snow module, with a 10% increase and decrease applied to the following factors within the module: winter precipitation, wind exposure, and facilitation. The results were then compared to those obtained from the reference snow module run.

We compared the effect of all these settings on treeline migration rate, where the treeline is defined as the northernmost position where forest density does not fall below a threshold of one hundred trees per hectare, where a tree must be >1.3 m tall.

In total, 45 different simulation scenarios were computed (for a structured overview, see Table [A-D1](#) in the Appendix). These include those evaluating model sensitivity to [predicting variables](#)~~impact factors~~, as described before, as well as the simulations with and without the forest fire module, the insect disturbance module, the module for trait variation and inheritance, and the new snow module including the evaluation of the model sensitivity to [predicting variables](#)~~impact factors~~ within the snow module.

Each of the simulation scenarios was repeated 40 times.

2.6 Statistical analyses of simulation

The output files generated by the LAVESI model provide information on stem count and tree density for each 5×5 m grid cell within the simulation area. The statistical analysis was performed using the statistical tool R (R Core Team, 2023).

Stem counts for all tree species were aggregated across spatial positions for every fifth year between 2020 and 2095. The aggregated data are visualised as line graphs using the 'ggplot2' R package (Wickham, 2016) to illustrate changes over time across different scenarios. Additionally, spatial distributions of tree densities are depicted with separate raster plots created for the years 2020 CE, 2055 CE, and 2095 CE. Treelines were determined using threshold-based criteria of one hundred trees per hectare.

The migration rate of the alpine treeline was recorded for each simulation by calculating the change in treeline position between 2020 and 2095 for each scenario. Using the *diff* function in R, we measured incremental changes between consecutive treeline positions, capturing the position change between each measured interval. These differences were then averaged to obtain the mean migration rate, which represents the average advancement of the treeline over time. The significance of the migration rates was assessed using a paired t-test, conducted with the *t.test* function in R.

The relative change in migration rate was calculated for both the increased and decreased scenarios to quantify the sensitivity of the treeline to each factor. All resulting migration rates were computed to sensitivity values that present the percentage of change in migration rate scaled to the 10% change in the factor. The sensitivity values are calculated from the following Eq. (1):

$$S = \frac{\Delta Y \times 100}{\Delta P}, \quad (1)$$

where ΔY represents the change in the migration rate derived from each simulation run and ΔP is the 10% change in the parameter. The significance of the sensitivity values was evaluated using a one-sample t-test, performed with the *t.test* function in R (R Core Team, 2023).

We conducted a redundancy analysis (RDA) using the *rda* function from the 'vegan' R package (Oksanen et al., 2018) to assess the relationships between environmental variables and treeline migration rates across the three study sites (CA, RU, and AK). The response variable consisted of treeline migration rates from 45 different simulation scenarios. In these scenarios, key factors influencing treeline migration were manipulated by switching them on or off, or by increasing or decreasing their magnitude compared to an average reference value. Each scenario was simulated across 40 runs per study site. Environmental predictor variables were derived from weather data for each study site, spanning the period from 1990 to 2020. The climate variables, including mean annual temperature, mean January temperature, mean July temperature, mean summer (JJA) temperature, accumulated annual temperature (AAT), degree days above threshold (DDT), total annual precipitation, and

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480 To quantify and compare the influence of key predicting variables on simulated treeline migration rates, a sensitivity analysis
 was conducted across all three study sites, based on $\pm 10\%$ variations in individual variables. In general, the sensitivity of
 treeline migration rates to $\pm 10\%$ changes in impact factors was assessed across all three study sites. On average, a 10% increase
 in predicting variables impact factors results in a +2.98% change in migration rate (SD = 8.76) at site CA, +1.89% (SD = 7.14)
 at site RU, and +5.22% (SD = 32.45) at site AK. A 10% decrease leads to changes of -2.30% (SD = 8.43) at site CA, -0.97%
 (SD = 6.39) at site RU, and -7.21% (SD = 39.51) at site AK. The minimum and maximum significant sensitivities observed
 485 are +0.43% and +4.50% for site CA and +0.50% and +2.90% for site RU. For site AK, no significant values were detected
 (Table 3).

490 **Table 3: Sensitivity analysis results for treeline migration rates. Mean sensitivity values (in %/%) and standard deviations for 17
 predicting variables, each adjusted by $\pm 10\%$ from an average reference value, resulting in 34 distinct scenarios across three different
 study sites. * Highly significant with $p < 0.05$, and the remaining values non-significantly different from the reference run. Values
 are the mean of 40 simulations.**

-	Fort McPherson West (Canada)		Lake Ilimey (Russia)		Road to Central (Alaska)	
-	10%		10%		10%	
Predicting variable	S-	S+	S-	S+	S-	S+
activelayer	-3.65±9.41*	2.54±10.01	-1.23±6.64	0.98±6.71	-8.29±40.82	6.64±29.44
dispersaldistance	-0.75±8.02	2.99±9.85	-0.37±5.24	0.60±5.48	-6.70±28.06	6.36±39.29
growingseasonlength	-3.12±9.66*	2.78±8.30*	-2.24±7.02	1.59±6.79	-3.55±14.66	-0.40±10.06
longdistancedispersal	-1.87±8.85	1.73±7.88	-0.44±6.22	2.90±6.08*	-9.98±51.50	2.91±15.23
maturationage	-3.84±8.03*	0.30±7.18	-3.10±7.39*	1.91±7.81	-5.03±40.20	7.93±40.25
overageingmortality	-2.69±10.35	3.22±8.86*	0.12±4.32	1.97±7.72	-10.70±65.20	6.40±39.20
seedbedavailability	-2.06±8.58	4.32±9.66*	0.50±6.08	2.51±7.71*	-5.95±35.82	4.60±29.80
seedintronumberpermanent	-0.77±6.45	2.12±8.57	-0.95±6.32	1.65±7.03	0.17±7.64	5.14±27.14
seedlingestablishment	-1.01±7.86	3.39±9.39*	-1.31±8.65	2.83±7.84*	-7.38±46.34	5.68±36.30
seedlingmortality	-2.34±9.50	2.35±7.15*	-2.11±6.13*	1.45±7.22	-9.40±46.04	4.89±25.42
seedproduction	-1.18±7.63	4.50±8.84*	-1.33±8.13	1.52±6.57	-8.59±37.65	1.51±15.73
slope	-1.96±7.90	0.24±8.06	-2.03±7.01	1.25±6.60	-7.70±57.74	8.37±44.68
summerprecipitation	-3.03±8.99*	0.43±6.28	-1.19±5.16	2.25±7.21	-6.08±25.64	3.94±14.03
summertemperature	-2.38±8.85	2.73±9.19	-0.67±5.73	1.24±8.19	-3.54±18.97	4.33±32.80
twiWtwi	-2.74±9.71	4.00±10.29*	-1.22±7.09	1.71±8.66	-4.51±29.31	3.28±16.66
windspeed	-2.52±9.80	2.99±8.99*	-1.31±6.41	2.61±7.53	-10.05±62.16	4.09±22.37
wintertemperature	-2.18±7.38	2.61±7.35*	-0.20±4.67	2.10±7.38	-9.32±43.66	12.38±76.15
Mean (all variables)	-2.30 ± 8.43	2.98 ± 8.76	-0.97 ± 6.39	1.89 ± 7.14	-7.21 ± 39.51	5.22 ± 32.45

Site AK exhibits the highest mean sensitivity values; however, these results are characterised by substantial variability,
 preventing any significant findings. In contrast, site RU generally shows lower sensitivity magnitudes across all scenarios,

495 although some scenarios do reach significance. Site CA displays several scenarios with significant positive or negative sensitivities, indicating a stronger influence on migration dynamics compared to the other sites.

500 The direction of sensitivity across all study sites is consistent for the majority of predicting variables~~impact factors~~, with positive and negative sensitivities aligning similarly between sites. There are four exceptions where the direction of influence differs between study sites, but these deviations are minimal and not statistically significant.

505 Notably, four scenarios demonstrate significant sensitivity at both sites CA and RU ($p < 0.05$). Increased seedbed availability has the most pronounced positive influence on migration, with mean sensitivities of 4.32% at site CA and 2.51% at site RU (seedbedavailability_plus10%). In contrast, reduced maturation age strongly hinders migration potential, as reflected by negative mean sensitivities of -3.84% at site CA and -3.10% at site RU (maturationage_minus10%). Favourable conditions for seedling establishment also plays an important role, promoting migration with mean sensitivities of 3.39% at site CA and 2.83% at site RU (seedlingestablishment_plus10%). Lastly, higher wind speeds facilitate migration through enhanced seed dispersal, as indicated by mean sensitivities of 2.99% at site CA and 2.61% at site RU (windspeed_plus10%).

3.2 Impact of snow-related processes on alpine treeline migration rates

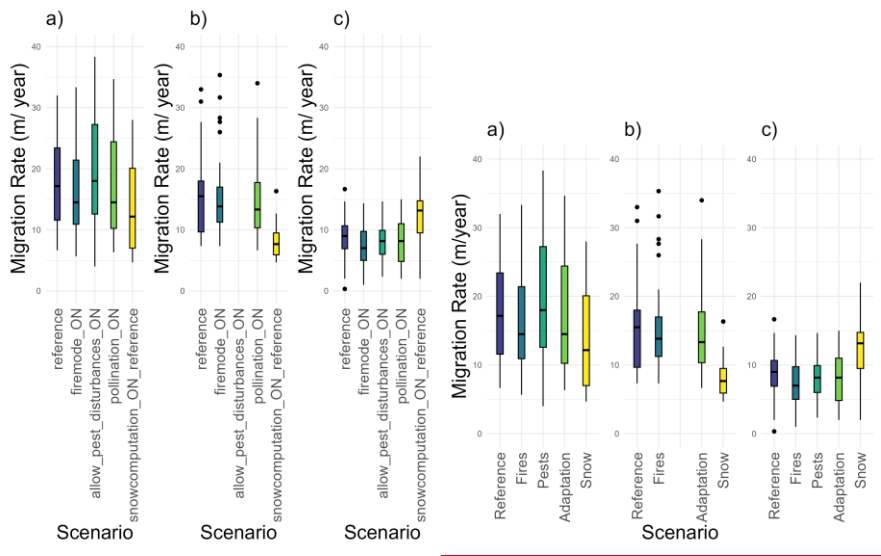


Figure 81: Box plots showing migration rates from 40 simulation runs at three study sites: a) Fort McPherson, Canada (CA), b) Lake Ilirney, Russia (RU), and c) Road to Central, Alaska, USA (AK). Each panel compares the reference scenario with four scenarios in which specific environmental modules were activated. Note that for the Russian study site, the insect disturbance module was not implemented in LAVESI.

The comparison of the results from additional simulations, which were performed with the forest fire module (Glückler et al., 2024), the insect disturbance module, the trait variation and inheritance module (Gloy et al., 2023), and the snow module implemented in this study, were compared with the results of the reference run revealed that (Figure 1). Notably, across all three study sites, only the simulations incorporating the influence of snow-related factors resulted in a significant change in the treeline migration rate (Figure 8).

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Table 4: Sensitivity analysis results for treeline migration rates within snow computation simulations. Mean sensitivity values (in %) and standard deviations for three predicting variables, each adjusted by ±10% from an average reference value, resulting in six distinct snow computation scenarios across three different study sites. * Highly significant with p < 0.05, and the remaining values non-significantly different from the reference run. Values are the mean of 40 simulations.

	Fort McPherson West (Canada)		Lake Ilirney (Russia)		Road to Central (Alaska)	
	10%		10%		10%	
Predicting variable	S-	S+	S-	S+	S-	S+
facilitation	-5.26±12.64*	6.53±14.38*	-2.03±6.79	4.15±8.68*	-4.48±15.69	2.37±13.34
windexposure	-5.19±13.42*	2.62±10.53	-6.07±9.05*	-1.23±2.85*	-1.54±6.60	4.79±14.34*
winterprecipitation	-7.64±14.99*	3.91±11.41*	-2.43±7.73	2.99±9.09*	-1.96±14.64	5.30±12.15*
Mean	-6.49 ± 14.47	4.35 ± 12.44	-3.51 ± 7.86	1.30 ± 6.87	-2.66 ± 12.31	4.15 ± 13.28

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The sensitivity of treeline migration rates to ±10% changes in snow-related predicting variables, relative to the snow module reference run, was analysed across all three study sites. On average, a 10% increase in predicting variables leads to a +4.35% change in migration rate (SD = 12.44) at site CA, +1.30% (SD = 6.87) at site RU, and +4.15% (SD = 13.28) at site AK. Conversely, a 10% decrease in predicting variables results in a -6.49% change (SD = 14.47) at site CA, -3.51% (SD = 7.86) at site RU, and -2.66% (SD = 12.31) at site AK. The minimum and maximum significant sensitivities observed are +3.91% and +6.53% for site CA, +2.99% and +4.15% for site RU, and +4.79% and +5.30% for site AK (Table 4).

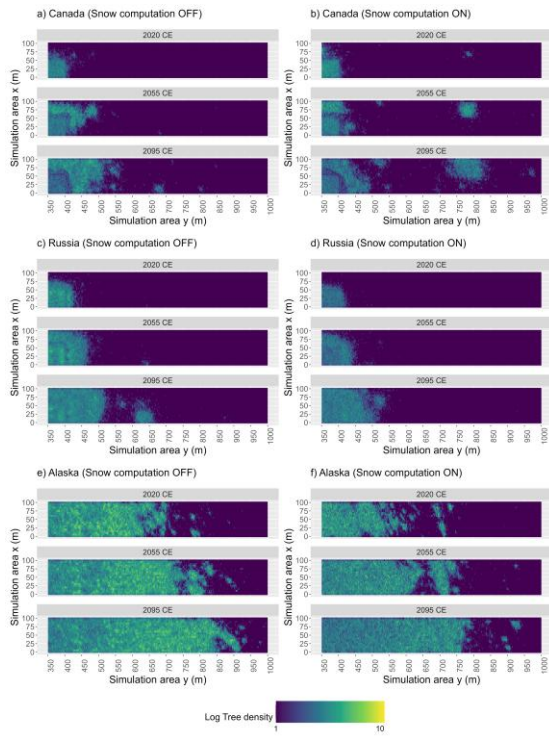
At site CA, several scenarios demonstrate strong positive or negative sensitivities, with nearly all results reaching statistical significance (p < 0.05). In contrast, site RU exhibits more moderate sensitivities, although some scenarios still achieve

statistical significance. Despite relatively high mean sensitivity values at site AK, substantial variability in the results causes fewer statistically significant outcomes

540 The sensitivity analysis for snow computation reveals consistent directional influences across the three study sites, with one notable exception: windexposure_plus10%. While sites CA and AK exhibit positive sensitivity values, site RU displays a negative sensitivity value. The sensitivity values at sites RU and AK are statistically significant ($p < 0.05$).

The scenario winterprecipitation_plus10% has statistically significant impacts across all sites, suggesting that increased snow accumulation during winter enhances migration potential. Both facilitation_plus10% and windexposure_minus10% show significant sensitivity at sites CA and RU ($p < 0.05$). Increased facilitation positively influences migration potential, with the strongest effects observed at site CA, while reduced wind exposure negatively impacts migration at both sites. Although site AK displays similar patterns, the variability in the data prevented statistical significance.

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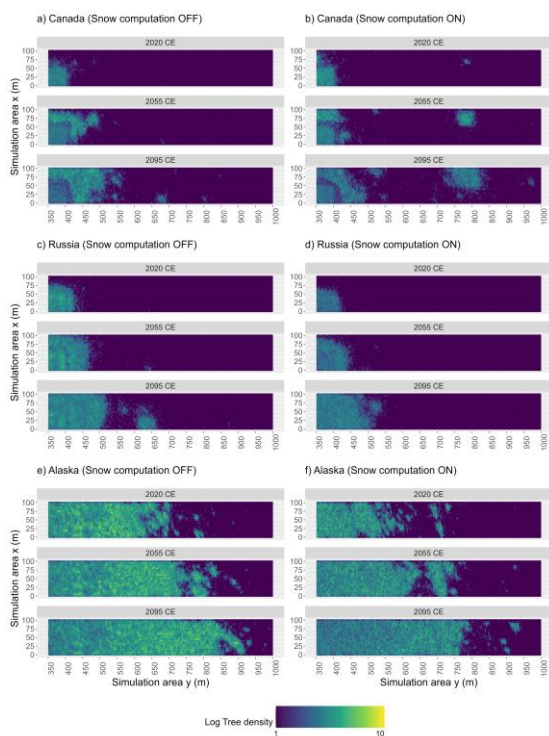


Figure 92: Tree density over time. Tree density in 2020, 2055, and 2095 CE at study site Fort McPherson West, Canada (CA) under a) the reference scenario (run 34) and b) the scenario with activated snow computation (run 2). Tree density in 2020, 2055, and 2095 CE at study site Lake Ilirney, Russia (RU) under c) the reference scenario (run 13) and d) the scenario with activated snow computation (run 32). Tree density in 2020, 2055, and 2095 CE at study site Road to Central, Alaska, USA (AK) under e) the reference scenario (run 15) and f) the scenario with activated snow computation (run 3). (Run with the median migration rate calculated from 40 runs for each simulation scenario). Each plot represents data from the simulation run with the median migration rate derived from 40 simulation runs per scenario. Tree density values were log-transformed to enhance data visualisation.

Across all three study sites, simulations without the snow module (Snow computation OFF) produce different tree density values over time compared to those incorporating snow-related processes (Snow computation ON). For all sites, the inclusion of snow processes results in an increase in tree density by the years 2055 and 2095 (Figure 92).

3.3 Relationships between environmental variables and treeline migration rates

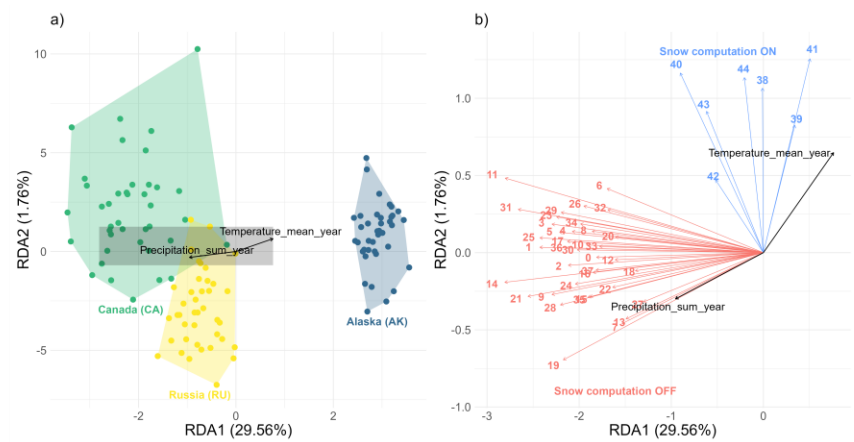


Figure 103: Redundancy Analysis (RDA) biplot showing a) site scores and b) species scores along the first two principal axes. Plot b provides a close-up of the centre region (grey rectangle) of plot a to highlight tightly grouped species. Site scores (migration rates) are represented by points, colour-coded according to their study site. Species scores (simulation scenarios) are depicted as arrows, colour-coded to indicate whether snow computation is activated (blue: Snow computation ON) or not (red: Snow computation OFF). Black arrows in both plots represent the direction and strength of environmental variables influencing migration rates.

The redundancy analysis (RDA) identified two significant constrained axes, collectively explaining 31.32% of the variance in treeline migration rates across simulation scenarios and study sites (Figure 103). The site scores reveal distinct clusters of study sites along the RDA1 axis, reflecting geographical differences. Sites CA and RU group closely together on one side of the axis, while site AK is distinctly separate on the opposite side (Figure 103a). Along the RDA2 axis, the species scores indicate that scenarios incorporating snow-related processes (Snow computation ON) are clearly distinguished from those without snow effects (Snow computation OFF) (Figure 103b).

The stepwise forward selection in RDA identified total annual precipitation and mean annual temperature as the most significant climate predictors of treeline migration rates across the simulation scenarios and study sites. Total annual precipitation is the first variable included in the model (AIC = 903.76, F = 43.80, p = 0.001), followed by mean annual temperature (AIC = 898.55, F = 7.25, p = 0.001). Total annual precipitation shows a positive correlation with treeline migration rates at sites CA and RU, but a negative correlation at site AK. Conversely, mean annual temperature exhibits an opposite effect: positively correlating with migration rates at site AK and negatively correlating at sites CA and RU.

Table 5: Expected and observed direction of influence of **predicting variables**~~impact factors~~ on migration rate.

Predicting variable Impact factor	Expected influence	Observed influence	Verification status
Active layer depth (activelayer)	Positive	Positive	Confirmed
Seed and pollen dispersal (dispersaldistance)	Positive	Positive	Confirmed
Length of the growing season (growingseasonlength)	Positive	Positive	Confirmed
Seed and pollen dispersal over long distances (longdistancedispersal)	Positive	Positive	Confirmed
Age at which a tree becomes mature (maturationage)	Negative	Positive	Not confirmed
Mortality of trees caused by the effects of ageing (overageingmortality)	Negative	Positive	Not confirmed
Availability of suitable seedbeds (seedbedavailability)	Positive	Positive	Confirmed
Permanent seed introduction (seedintronumberpermanent)	Positive	Positive	Confirmed
Establishment of seedlings (seedlingestablishment)	Positive	Positive	Confirmed
Mortality of seedlings (seedlingmortality)	Negative	Positive	Not confirmed
Factor of seed productivity (seedproduction)	Positive	Positive	Confirmed
Slope angle (slope)	Negative	Positive	Not confirmed
Summer precipitation (summerprecipitation)	Variable, magnitude-dependent	Positive	Partly confirmed
Summer temperature (summertemperature)	Positive	Positive	Confirmed
Topographic wetness index (twi)	Positive	Positive	Confirmed
Wind speed (windspeed)	Variable, magnitude-dependent	Positive	Partly confirmed
Winter temperature (wintertemperature)	Positive	Positive	Confirmed
Facilitation on tree growth through snow (Snowcomputation_ON_facilitation)	Positive	Positive	Confirmed
Wind exposure (Snowcomputation_ON_windexposure)	Variable, magnitude-dependent	Variable, magnitude-dependent	Confirmed
Temperature-dependent snow fall and accumulation (Snowcomputation_ON_winterprecipitation)	Variable, magnitude-dependent	Positive	Partly confirmed

Formatiert: Schriftart: 9 Pt.

4.1 Site-specific responses to factors influencing alpine treeline migration

4.1.1 Spatial variability in migration rates

The RDA analysis reveals a clear spatial grouping of locations based on migration rates, with sites CA and RU clustering together, while site AK is distinctly separate. **Notably, different tree species dominate at the respective sites (see Figure B1 in the Appendix). Based on species-specific traits, one would expect *Picea glauca* to exhibit a faster migration rate in Alaska than *Picea mariana* in Canada based on faster and taller growth. However, our findings contradict this expectation, indicating that local climatic conditions likely exert a stronger influence on migration dynamics than intrinsic species traits alone. This**

585

~~highlights the importance~~~~This pattern underscores the influence~~ of unique environmental conditions at each site on migration dynamics. Such geographical and ecological variability in treeline responses aligns with previous findings, reinforcing site-specific factors as the central factor to shaping treeline migration (Harsch et al., 2009; Hansson et al., 2021; 2023).

4.1.2 Sensitivity analysis: site-specific patterns

The sensitivity analysis further uncovered marked differences in responses to $\pm 10\%$ changes in ~~predicting variables~~~~impact factors~~ across the three sites (CA, RU, AK). Site AK exhibited high mean sensitivity values but with substantial variability, reflecting the influence of complex and less predictable interactions between ecological and climatic factors. ~~At this site, no single factor emerged as statistically significant, which may be due to stronger interactions among multiple drivers that obscure the effects of individual parameters. Another possible explanation is a greater influence of stochastic processes, further contributing to the uncertainty observed. Although studies explicitly linking stochastic processes to treeline dynamics are currently lacking, Shoemaker et al. (2020) and Lindo et al. (2023) emphasise their broader ecological relevance, particularly with regard to biodiversity patterns and ecosystem functioning. Consequently,~~~~This variability makes~~ migration dynamics at site AK ~~are~~ inherently more uncertain compared to the more consistent and significant sensitivities observed at sites CA and RU. These findings underscore the need to account for localised environmental conditions when predicting treeline responses to climate change.

4.1.3 Consistency and anomalies in sensitivities

Despite these site-specific differences, the overall consistency in the direction of sensitivities across all sites enhances confidence in the key drivers of treeline migration identified by this study.

While many observed sensitivities align with ecological expectations, some findings are unexpected and warrant further investigation to clarify these interactions (see Table 5). For instance, the positive relationship between maturation age and migration rate is counterintuitive, as earlier maturation typically shortens generation time and reduces exposure to juvenile mortality before the first reproductive event (Stearns, 2000). However, later maturation offers advantages such as larger tree size at maturity and higher fecundity (Stearns, 2000). Moreover, taller, older trees likely enhance migration potential through improved pollen and seed dispersal distances, consistent with the dispersal model used (Kruse et al., 2018). This extended dispersal likely reduces local competition among seedlings, enhancing migration potential, even when maturation age increases.

Another unexpected finding is the association of higher seedling and overageing mortality with increased migration rates. One explanation could be that, while higher mortality intuitively reduces population growth, it may simultaneously reduce competition and increase resource availability for surviving seedlings, thereby enhancing their establishment and growth (Kruse et al., 2016). This highlights the complex balance between population dynamics and resource competition in determining migration outcomes, ~~as also demonstrated by Zheng et al. (2024).~~

620 The positive sensitivity of migration rates to slope angles also stands out, as steeper slopes are typically considered physical barriers to migration. This may be attributed to the creation of diverse microhabitats in complex terrain, which reduce competition and provide unique niches for tree establishment, as observed by Soliveres et al. (2011).

4.1.4 Important factors for treeline migration

625 The significant sensitivity observed at sites CA and RU underscores the importance of certain environmental factors in promoting treeline migration. For example, reducing maturation age negatively impacts migration potential, likely because of the reduced tree height at maturity, which limits pollen and seed dispersal distances as discussed earlier (Stearns, 2000; Kruse et al. 2018).

Seedbed availability emerges as a crucial factor for migration potential, with sensitivity analysis confirming that increases in suitable seedbeds significantly enhance the migration rate, as expected. Suitable seedbeds provide favourable conditions for tree establishment, especially bare mineral soils, with their reduced competition from dwarf shrubs or grasses, which offer ideal conditions for wind-mediated tree seeds (Löffler et al., 2004; Hobbie and Chapin, 1998; Holtmeier and Broll, 2007; [Liang et al., 2016](#)).

Seedling establishment strongly supports treeline advancement, as higher rates of successful seedling establishment drive treeline progression and maintain population stability (Holtmeier and Broll, 2007), though delays in processes such as seed dispersal, establishment, or growth may slow migration rates, as noted by Holtmeier and Broll (2019).

Wind speed is another important factor influencing treeline migration, as moderate wind velocities enhance seed and pollen dispersal, aligning with findings from Holtmeier and Broll (2007) and Kruse et al. (2018).

640 These findings align closely with existing literature, highlighting the complex interplay between factors such as maturation age, seedbed availability, seedling establishment, and wind speed in shaping treeline migration processes. The observed patterns are consistent with broader ecological principles related to tree migration under changing environmental conditions. However, the site-specific variations in sensitivity emphasise the critical need to account for localised conditions in predictive simulation models.

Overall, the simulations analysing various factors affecting alpine treeline migration reveal statistically significant influences on migration rates at sites CA and RU, offering valuable insights into the key factors driving migration potential under current conditions. In contrast, the variability observed at site AK underscores the need for further investigation into local factors influencing migration sensitivity. These results collectively reinforce the importance of integrating both broad-scale ecological drivers and site-specific conditions to enhance our understanding of treeline dynamics.

4.2 Climate drivers of alpine treeline migration across study sites

650 Our findings align with previous research showing that climate-induced treeline shifts are shaped by a complex interplay of warming temperatures and precipitation variability, which jointly influence seedling establishment and growth (Kullman, 2005, 2007; Holtmeier and Broll, 2005, 2007; Körner, 2012). Here, total annual precipitation and mean annual temperature

emerge as the most influential drivers of treeline migration accounting for over 30% of the observed variance in migration rates across the different scenarios. Traditionally, growing season temperature has been regarded as a critical factor driving treeline shifts and forest expansion (Dial et al., 2007; Devi et al., 2008; 2020). However, our findings underscore the importance of considering year-round climate conditions in shaping treeline dynamics, a perspective consistent with studies emphasising the importance of mean annual temperature in influencing treeline migration (Cazzolla Gatti et al., 2019; Danby and Hik, 2007).

In the Polar Urals, winter precipitation has more than doubled over the 20th century (Devi et al., 2020), while in the Kenai Mountains, Alaska, mean annual temperatures have risen by 2°C since the 1950s, with winter temperatures increasing by as much as 4°C (Dial et al., 2007). These trends underscore the necessity of considering winter climate, in addition to growing season conditions, when studying treeline dynamics.

Although precipitation outside the growing season is often simply added to its annual sum, primarily in relation to soil water balance, our results highlight the multifaceted role of snow, which extends beyond its contribution to water availability (see section 4.3 The role of snow in alpine treeline dynamics).

4.2.1 Site-specific responses to climate drivers

The migration responses to the climate drivers temperature and precipitation vary between the sites (see Table 1). Site AK is characterised by relatively higher mean annual temperatures and lower precipitation compared to sites CA and RU.

At site AK, warmer mean annual temperatures positively influence migration. This effect is likely due to an extended growing season and improved soil warmth, which enhance seedling germination and establishment. These findings align with evidence suggesting that warmer conditions promote tree germination, establishment, and survival during the growing season (Almquist et al., 1998; Holtmeier and Broll, 2007; Kullman, 2007; Devi et al., 2008; 2020; Grigoriev et al., 2022). In contrast, at the cooler, wetter sites CA and RU, higher temperatures are associated with reduced migration rates. This negative effect may be attributed to drought stress, that may override temperature benefits at upper treeline boundaries, as warming-induced soil moisture limitations could hinder tree growth and prevent treeline advance in a warmer world (Bailey et al., 2021; Gruber et al., 2022).

The influence of precipitation also varies across the sites. At sites CA and RU, higher precipitation is positively correlated with accelerated migration rates. This is likely due to improved soil moisture, which supports seedling establishment and early growth in these regions (Holtmeier and Broll, 2007). These findings are consistent with studies showing that treeline migration and forest expansion are linked to increased moisture availability, particularly from higher winter precipitation (Hagedorn et al., 2014; Grigoriev et al., 2022). Conversely, at site AK, higher precipitation is associated with slower migration rates, possibly due to negative factors such as waterlogging or paludification implicitly modelled by an optimum precipitation and adverse effects on potential growth in the model (Kruse et al., 2016). These findings highlight the potential for moisture-related limitations on tree expansion under certain site conditions (Crawford et al., 2003; Holtmeier and Broll, 2007).

685 Additionally, the different tree species present across sites likely play a role in shaping these site-specific responses. Species-specific stand structure dynamics influence how various species respond to environmental conditions, further contributing to the observed variation in migration rates (Chhetri et al., 2020). For instance, the deciduous species in Siberia (RU) contrast with the evergreen species in North America (CA and AK), which may affect their climatic sensitivities. Traits such as seed production, growth rate, and the ability to recover from disturbances can modulate species' responses.

4.3 The role of snow in alpine treeline dynamics

690 **4.3.1 Significant impact of snow-related processes on alpine treeline migration**

Our study further demonstrates that snow-related processes, incorporated into the LAVESI snow module, consistently reveal significant changes in migration rates across all three study sites. This finding aligns with the work of Huang et al. (2023), who show that seasonal snow-cover patterns – particularly spring snowline elevation and snow-cover duration – affect treeline elevation in the eastern Himalaya. Similarly, Niittynen and Luoto (2018) demonstrate that incorporating snow persistence data into species distribution models enhances predictive accuracy and refines distribution mapping in high-elevation and high-latitude ecosystems. Additionally, Autio (2006) predicts that warmer temperatures could lead to increased snowfall, potentially resulting in heavier snow loads that may damage or even destroy tree crowns and branches. Collectively, these findings underscore the critical role of snow as an environmental predictor shaping treeline dynamics in alpine and subarctic ecosystems.

700 In contrast, recent studies evaluating environmental modules with the model LAVESI, such as explicit forest fire (Glückler et al., 2024), biotic (insect) disturbances, and trait variation and inheritance (Gloy et al., 2023), have highlighted the significance of various processes in influencing tree migration. However, our simulations do not reveal notable impacts from these factors at alpine treeline ecotones. This discrepancy may be attributed to the relatively short duration of our simulations, which might not have allowed enough time for these processes to fully manifest. For the study sites of our study, findings emphasise the immediate and dominant influence of snow-related processes on treeline dynamics, compared to longer-term processes like fire or trait evolution.

705

4.3.2 The dual role of snow and model enhancements

The separation between scenarios with and without snow effects along the second RDA axis further underscores the substantial impact of snow-related processes on migration rates. Snow acts as both a facilitator and a hindrance, depending on environmental context. Incorporating snow effects is thus critical for accurately assessing treeline migration in snow-dominated regions (e.g. Niittynen and Luoto, 2018; Huang et al., 2023).

710 Compared to earlier implementations of snow effects in vegetation models, our new snow module represents a significant advancement. Instead of merely adding precipitation outside the growing season to the annual total, the module captures the complex and multifaceted role of snow, extending beyond its contribution to water availability. Hansson et al. (2023) note that

715 average precipitation changes often show weak correlations with treeline migration, likely due to model limitations that neglect
variables like moisture stress and extreme precipitation events. Similarly, Kharuk et al. (2022) observe that under harsh
conditions, factors such as wind and soil moisture often outweigh the role of snow. By incorporating previously overlooked
factors such as wind exposure and microtopographical features, the enhanced module provides a more nuanced representation
of treeline dynamics, capturing the complex interactions between snow and other environmental drivers. These processes –
720 modelled on findings from other studies and existing knowledge – are not explicitly validated with data in this study but are
tested to better estimate their influences.

4.3.3 Sensitivity analysis and site-specific variations

The sensitivity analysis conducted with the snow module reveals significant site-specific variations in migration rates in
response to $\pm 10\%$ changes in ~~predicting variables~~~~impact factors~~ across the three study sites (CA, RU, AK). The higher
725 sensitivity values observed at site CA likely reflect more extreme or variable environmental conditions, whereas the relatively
lower sensitivity at site RU suggests a more stable environment where migration processes are less reactive to changes. Site
AK exhibits high sensitivity values, but the lack of statistically significant results in some scenarios suggests complex
interactions and local factors not fully captured by the model.

The contrasting sensitivity observed in the scenario with a 10% increase in wind exposure highlights the complex role of wind
730 in tree migration. At sites CA and AK, positive sensitivity values indicate that wind may enhance migration by aiding seed
and pollen dispersal, which are crucial drivers of migration potential in these regions (Holtmeier and Broll, 2007; Kruse et al.,
2018). Wind-driven dispersal facilitates seed movement over long distances, likely explaining the positive association between
wind exposure and migration at these sites (Holtmeier and Broll, 2007). However, the negative sensitivity observed at site RU
suggests that wind exposure here may exert physiological stress that inhibits treeline migration. High wind velocities, for
735 example, can damage trees through frost-drying and limit growth, a pattern documented in other wind-exposed treeline areas
(Holtmeier, 1985). Moreover, seeds carried by strong winds tend to accumulate in sheltered areas, complicating migration in
exposed environments (Anschlag, 2006). Under such conditions, topographic factors such as wind exposure can outweigh
potential dispersal benefits and prevent the advance of the treeline despite significant climate warming, as observed in other
wind-exposed environments (Kullman, 2005; Beloiu et al., 2022). Additionally, wind erosion of exposed soils can reduce
740 available seedbeds beyond the treeline, further inhibiting migration (Holtmeier et al., 2004; Anschlag, 2006). This duality
underscores the complex role of wind in shaping treeline dynamics across varying environments.

Positive sensitivity to increased winter precipitation across all three study sites aligns with established ecological principles
emphasising the role of moisture availability (Walker et al., 1999) and protective cover during critical early growth stages
(Gross et al., 1991; Holtmeier, 2005a). Snow cover can provide crucial shelter for seedlings and saplings, protecting them from
745 frost damage and herbivory during winter months, thereby supporting seedling survival (Juntunen and Neuvonen, 2006;
Holtmeier and Broll, 2007). When snow melts, the resulting increase in soil moisture enhances seedling establishment during
the growing season, especially in regions where winter precipitation forms a significant portion of the annual water supply

(Germino and Smith, 2000; Smith et al., 2003; Hagedorn et al., 2014). However, excessive snow accumulation can also have negative effects, such as snow creep or slides that may threaten seedling survival (Holtmeier, 2003; 2005a; 2005b; Kullman, 2005). [Moreover, wetter snow conditions may promote snow fungi infections, resulting in increased mortality among seedlings of evergreen conifers \(Holtmeier, 2005a; 2005b; Holtmeier and Broll, 2007; Barbeito et al., 2013\).](#) This dual role of snow as both a protective factor and a potential hazard could explain why increased winter precipitation generally supports migration but may pose risks under extreme conditions (Holtmeier and Broll, 2007). Warmer temperatures, combined with increased snowfall, as observed in some regions, can promote forest expansion but also lead to snow-related damage (Autio, 2006).

These findings underscore how snow-rich environments can facilitate forest advancement by providing shelter and moisture during critical growth stages (Hagedorn et al., 2014; Kirdyanov et al., 2012; Devi et al., 2020; Grigoriev et al., 2022). Enhanced facilitation by surrounding trees, as opposed to competition, has a strong positive effect on migration at sites CA and RU. This likely arises from positive interactions within the growing tree population, where clusters of trees can moderate harsh conditions by creating warmer microclimates, dampening wind, and accumulating snowpack, thereby insulating young trees from extreme cold and herbivory (Germino and Smith, 2000; Germino et al., 2002; Holtmeier and Broll, 2007; [Bader et al. 2021; Zheng et al., 2024](#)). Conversely, reduced wind exposure consistently shows a negative impact on migration at sites CA and RU. This could be due to the important role wind plays in dispersing seeds across the landscape, as less wind exposure may reduce the ability of seeds to travel efficiently, thereby hindering migration (Holtmeier and Broll, 2007). Although reduced wind exposure may alleviate physiological stress on established trees, it can also delay or even prevent treeline advance by reducing seed dispersal in wind-exposed environments where natural shelter is limited (Holtmeier, 2005a; Kullman, 2005).

The sensitivity analysis highlights the complex interactions between snow accumulation, wind exposure, and facilitation processes in driving treeline migration. These processes are further shaped by site-specific factors such as topography and climatic variability, which can either enhance or inhibit migration under changing environmental conditions. These findings highlight the necessity of a comprehensive approach when assessing treeline migration in snow-dominated regions.

Our enhanced model represents a major advancement in simulating alpine treeline processes. By integrating snow effects on seedling survival and soil conditions, alongside factors like wind exposure, facilitation by neighbouring trees, and microtopography, the model provides a more comprehensive and accurate framework for understanding treeline migration in high latitudes and elevations. This holistic approach is crucial for improving predictions and informing conservation strategies in sensitive ecosystems. To refine our understanding, we challenged the model by incorporating various processes derived from general models and previous studies to assess their effects. Although explicit validation was beyond the scope of this study, this approach identifies key areas for future investigation to improve the accuracy of treeline dynamics modelling.

4.4 Constraints on treeline advance

The complexity of treeline migration is highlighted in our results, which show that site-specific climatic factors have the strongest influence on migration rates, though the relative importance of these factors varies by location. This aligns with the

780 widely accepted view that the elevational position of treelines is climate-dependent in the absence of significant disturbance regimes (Körner et al., 2020; Maher et al., 2021). However, our study underscores that treeline dynamics are not driven by climate alone but also by additional environmental variables, such as local site conditions and the impacts of disturbances, both natural and anthropogenic. These non-climatic influences complicate the relationship between climate warming and treeline advance, as certain detrimental factors cannot be offset by even the positive effects of a warmer environment

785 (Holtmeier and Broll, 2005; 2007; 2019; Hansson et al., 2021).

A key finding of our research is the importance of snow-related processes, which have been incorporated into the LAVESI snow module, in affecting tree migration rates within the treeline ecotone. Wiesner et al. (2019) suggest that, as mean surface temperatures rise, the influence of extreme weather events (e.g., early or late frosts, or summer soil dryness) may become more significant than gradual temperature increases. These extreme events, coupled with increasing snow cover due to climate

790 humidification, may act as a constraint on treeline advance in alpine regions.

Moreover, research on forest dynamics in high-elevation ecosystems highlights the need to consider both long-term warming trends and climate variability, as these drive episodic dieback and tree invasions (Bugmann and Pfister, 2000; Milar et al., 2004). While changes in climate averages are important, episodic events such as extreme cold episodes or changes in moisture availability can profoundly impact treeline dynamics. These episodic effects have often been underrepresented in vegetation

795 models.

This points to a potential explanation for the time lag between climate warming and observed treeline migration, as these short-term events and longer ecological processes complicate the response of tree populations to warming. The updated LAVESI version addresses this gap by explicitly accounting for the influence of varying snow depths on tree growth and treeline migration. By incorporating stochastically occurring extreme events, the model captures the complete spectrum of weather

800 variability.

Another significant contributor to the time lag in treeline migration is the slow pace of population processes, such as seed dispersal, establishment, and growth. These slow processes can result in centuries-long delays between individual tree development and broader forest establishment (Lloyd, 2005; Holtmeier and Broll, 2007; 2019). Favourable conditions for seedling establishment may not align with good seed years, further delaying the upward shift of the treeline (Gruber et al.,

805 2022). Despite these delays, over time, slow processes at the advancing forest edge will eventually bring tree populations into alignment with the climatic treeline and its fundamental niche (Körner, 2021). The persistence of relict populations from past warm periods, such as northernmost dwarf spruces in the Canadian tundra, further supports this idea of delayed treeline responses (Nichols, 1976; Holtmeier, 1985). Similarly, the legacy of past environmental conditions strongly shapes the distribution of plant communities in both the treeline ecotone and lower alpine regions (Hofgaard and Wilmann, 2002). The

810 integration of snow processes into the LAVESI model enhances the accuracy of simulations related to seed dispersal, establishment, and growth, ultimately improving predictions of treeline advancement. However, the localised legacy effects of past populations vary significantly at small spatial scales. Addressing these variations requires further region-specific research to refine vegetation models and enhance their predictive capabilities.

815 Dial et al. (2022) propose that part of the delay in treeline advance may be due to conifers being on the brink of a stochastic,
climate-driven invasion of tundra after centuries of stability. However, many treelines worldwide show slower upward shifts
compared to the rate of densification, with tree populations rapidly increasing under favourable conditions (Kruse et al., 2016;
Wang et al., 2016; Shi et al., 2022). The LAVESI model is capable of simulating not only treeline migration but also tree
density within the specified simulation area. Additionally, shifts in reproductive methods – such as a move toward seedling-
based regeneration in species like Engelmann spruce (*Picea engelmannii*) and subalpine fir (*Abies lasiocarpa*) – could facilitate
820 greater dispersal distances and potentially accelerate treeline advancement in the future (Holtmeier and Broll, 2007). An
upward shift of the alpine treeline in cold climate zones may also be limited by the fact that while short vegetation thrives in
the relatively warmer microenvironment near the ground, trees can only reach full size if they can withstand the surrounding
ambient temperatures (Wilson et al., 1987, Grace, 1989). Incorporating such characteristics that are partly species-specific and
partly population-specific would further enhance the accuracy of vegetation models. Moreover, integrating insights into the
825 ecological processes that shape spatial patterns within treeline ecotones – particularly those related to tree cover, clustering,
and stature – could further improve the prediction of treeline dynamics under changing climatic conditions (Bader et al.; 2021).
Our study sites are characterised by conifer-dominated treelines, comprising predominantly evergreen species in North
America and deciduous conifers in Siberia. In ecotones with mixed forest composition, warming-induced acceleration of
successional shifts from broadleaved to coniferous dominance warrants particular attention (Sigdel et al., 2024). Our findings
830 underscore the complexity of treeline dynamics, suggesting that both climate factors, including snow-related processes, and
slow ecological processes, such as seedling establishment, growth, survival, reproduction, and forest development, play crucial
roles in shaping treeline responses to climate change. The gaps highlight the need for further refinement of vegetation models
to better explain the time lag between climate warming and treeline advance.

5 Conclusions

835 In conclusion, this study highlights the distinct site-specific responses to factors influencing treeline shift and forest expansion,
emphasising the critical role of localised conditions in shaping migration dynamics. While the findings reveal commonalities
across sites, they also underscore unique interactions between environmental factors and local contexts. Notably, the responses
observed at the Canadian and the Russian sites provide clearer insights into the primary drivers of migration, whereas the high
variability at the Alaskan site suggests more complex or less predictable local dynamics. This variability highlights the need
840 for further investigation to reduce uncertainties in model predictions.
A key finding of this study is the significant role of snow in modulating migration potential, as snow accumulation creates
favourable moisture conditions and protects seedlings from extreme cold, particularly in snow-dominated environments. These
results emphasise the importance of incorporating snow-related processes into vegetation models to improve predictions of
boreal forest dynamics under changing environmental conditions.

845 In boreal regions, altered snow regimes and increasingly frequent climatic extremes may affect treeline advance, with cascading effects on forest structure, biodiversity, surface albedo, and the long-term carbon sequestration potential of these ecosystems. As warming accelerates, these dynamics are likely to intensify, highlighting the need for refined, region-specific models to anticipate ecological feedbacks.

Overall, this study enhances our understanding of tree migration processes and underscores the varied predictability of migration responses across different sites. These insights carry valuable implications for enhancing model accuracy and guiding conservation strategies aimed at sustaining alpine tundra resilience in the face of rapid environmental change.

Code availability. The source code of the host model is available from GitHub at <https://github.com/StefanKruse/LAVESI/tree/circumboreal>. The updated version presented here is named LAVESI-Mountain Treelines and the first version 1.0 is available for review on GitHub at https://github.com/StefanKruse/LAVESI/tree/circumboreal_mountain. Upon manuscript acceptance for publication, the code, input data, and simulation output will be permanently archived on Zenodo.

Author contributions. Sarah Haupt: Conceptualisation; methodology; software; validation; formal analysis; investigation; resources; data curation; writing - original draft; writing - review & editing; visualisation. Josias Gloy: Software; data curation. Luca Farkas: Software; data curation; writing - review & editing. Katharina Schildt: Software; data curation; writing - review & editing. Lisa Trimborn: Software; data curation; writing - review & editing. Stefan Kruse: Conceptualisation; methodology; software; validation; formal analysis; investigation; resources; data curation; writing - review & editing; supervision; project administration; funding acquisition.

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Appendix

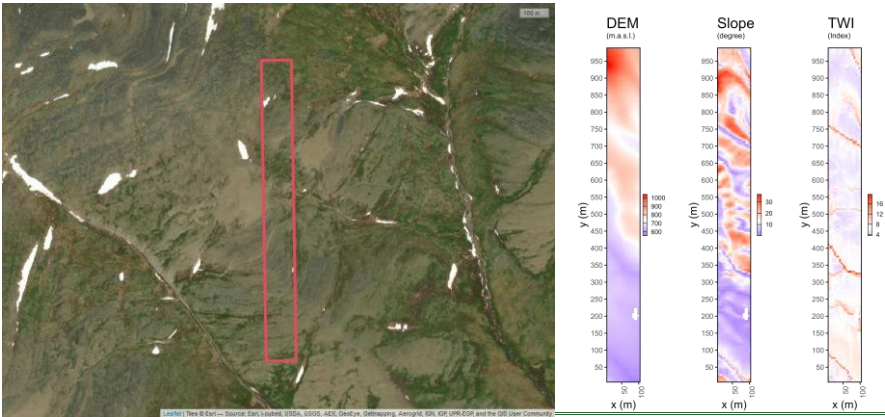


Figure A1: Satellite image of the study site Fort McPherson, Canada (CA), including the simulation area (red box). Digital elevation model (DEM), slope, and topographic wetness index (TWI) for the simulation area. Esri, i cubed, USDA, USGS, AEX, Geoeye, Getmapping, Aerogrid, IGN, IGP, UPR-EGP, and the GIS User Community.

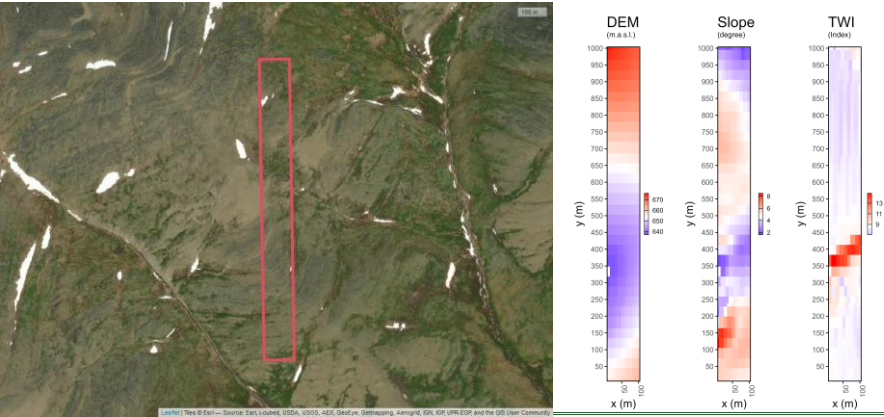


Figure B1: Satellite image of the study site Lake Ilirney, Russia (RU), including the simulation area (red box). Digital elevation model (DEM), slope, and topographic wetness index (TWI) for the simulation area. Esri, i-cubed, USDA, USGS, AEX, Geoeye, Getmapping, Aerogrid, IGN, IGP, UPR-EGP, and the GIS User Community.

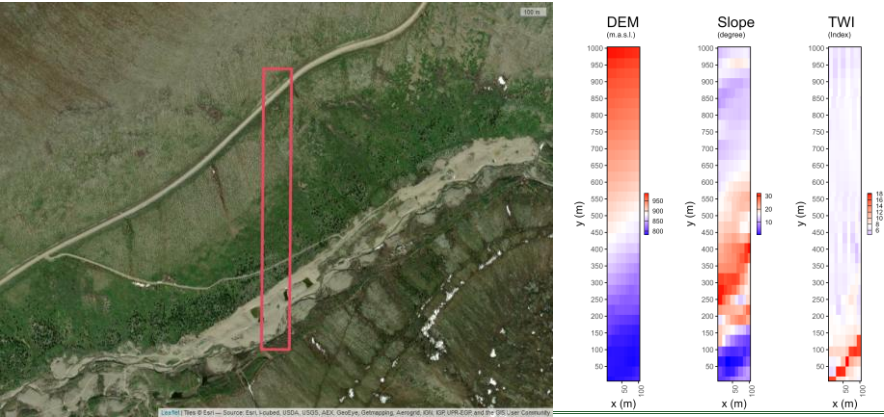


Figure C1: Satellite image of the study site Road to Central, Alaska/ USA (AK), including the simulation area (red box). Digital elevation model (DEM), slope, and topographic wetness index (TWI) for the simulation area. Esri, i-cubed, USDA, USGS, AEX, Geoeye, Getmapping, Aerogrid, IGN, IGP, UPR-EGP, and the GIS User Community.

Table AD1: Overview of simulation scenarios. Summary of the simulation scenarios and their respective multipliers for predicting variablesimpact-factors used in the sensitivity analysis.

Scenario ID	Scenario name	Predicting variableImpact factor	Multiplier
0	reference	growingseasonlength	1.00
		summertemperature	1.00
		wintertemperature	1.00
		summerprecipitation	1.00
		windspeed	1.00
		seedproduction	1.00
		maturationage	1.00
		seedintronumberpermanent	1.00
		dispersaldistance	1.00
		longdistancedispersal	1.00
		seedlingestablishment	1.00
		seedlingmortality	1.00
		overageingmortality	1.00
		slope	1.00
		twi	1.00
		seedbedavailability	1.00
		activelayer	1.00
		firemode	0.00
		allow_pest_disturbances	0.00
		pollination	0.00
		snowcomputation	0.00
1	growingseasonlength_plus10%	growingseasonlength	1.10
2	growingseasonlength_minus10%	growingseasonlength	0.90
3	summertemperature_plus10%	summertemperature	1.10
4	summertemperature_minus10%	summertemperature	0.90
5	wintertemperature_plus10%	wintertemperature	1.10
6	wintertemperature_minus10%	wintertemperature	0.90
7	summerprecipitation_plus10%	summerprecipitation	1.10
8	summerprecipitation_minus10%	summerprecipitation	0.90
9	windspeed_plus10%	windspeed	1.10
10	windspeed_minus10%	windspeed	0.90
11	seedproduction_plus10%	seedproduction	1.10
12	seedproduction_minus10%	seedproduction	0.90

13	<i>maturationage_plus10%</i>	<i>maturationage</i>	1.10
14	<i>maturationage_minus10%</i>	<i>maturationage</i>	0.90
15	<i>seedintronumberpermanent_plus10%</i>	<i>seedintronumberpermanent</i>	1.10
16	<i>seedintronumberpermanent_minus10%</i>	<i>seedintronumberpermanent</i>	0.90
17	<i>dispersaldistance_plus10%</i>	<i>dispersaldistance</i>	1.10
18	<i>dispersaldistance_minus10%</i>	<i>dispersaldistance</i>	0.90
19	<i>longdistancedispersal_plus10%</i>	<i>longdistancedispersal</i>	1.10
20	<i>longdistancedispersal_minus10%</i>	<i>longdistancedispersal</i>	0.90
21	<i>seedlingestablishment_plus10%</i>	<i>seedlingestablishment</i>	1.10
22	<i>seedlingestablishment_minus10%</i>	<i>seedlingestablishment</i>	0.90
23	<i>seedlingmortality_plus10%</i>	<i>seedlingmortality</i>	1.10
24	<i>seedlingmortality_minus10%</i>	<i>seedlingmortality</i>	0.90
25	<i>overageingmortality_plus10%</i>	<i>overageingmortality</i>	1.10
26	<i>overageingmortality_minus10%</i>	<i>overageingmortality</i>	0.90
27	<i>slope_plus10%</i>	<i>slope</i>	1.10
28	<i>slope_minus10%</i>	<i>slope</i>	0.90
29	<i>twi_plus10%</i>	<i>twi</i>	1.10
30	<i>twi_minus10%</i>	<i>twi</i>	0.90
31	<i>seedbedavailability_plus10%</i>	<i>seedbedavailability</i>	1.10
32	<i>seedbedavailability_minus10%</i>	<i>seedbedavailability</i>	0.90
33	<i>activelayer_plus10%</i>	<i>activelayer</i>	1.10
34	<i>activelayer_minus10%</i>	<i>activelayer</i>	0.90
35	<i>firemode_ON</i>	<i>firemode</i>	112.00
36	<i>allow_pest_disturbances_ON</i>	<i>allow_pest_disturbances</i>	1.00
37	<i>pollination_ON</i>	<i>pollination</i>	1.00
38	<i>snowcomputation_ON_reference</i>	<i>snowcomputation</i>	1.00
		<i>winterprecipitation</i>	1.00
		<i>windexposure</i>	1.00
		<i>facilitation</i>	0.00
39	<i>snowcomputation_ON_winterprecipitation_plus10%</i>	<i>snowcomputation</i>	1.00
		<i>winterprecipitation</i>	1.10
40	<i>snowcomputation_ON_winterprecipitation_minus10%</i>	<i>snowcomputation</i>	1.00
		<i>winterprecipitation</i>	0.90
41	<i>snowcomputation_ON_windexposure_plus10%</i>	<i>snowcomputation</i>	1.00

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		windexposure	1.10
42	snowcomputation_ON_windexposure_minus10%	snowcomputation	1.00
		windexposure	0.90
43	snowcomputation_ON_facilitation_plus10%	snowcomputation	1.00
		facilitation	0.10
44	snowcomputation_ON_facilitation_minus10%	snowcomputation	1.00
		facilitation	-0.10

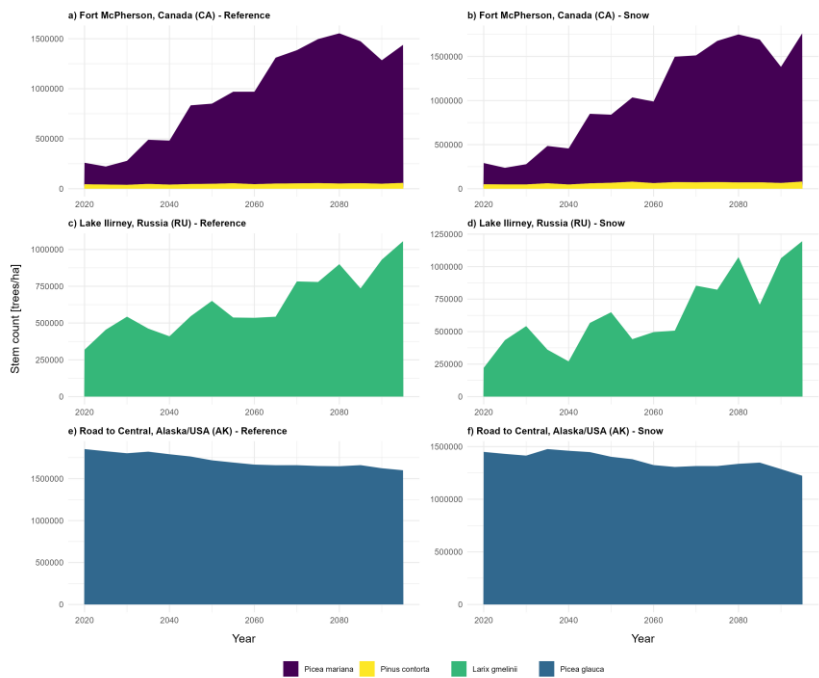


Figure B1: Stem count [trees/ha] over time. Stem count from 2020 to 2095 CE at the study site Fort McPherson West, Canada (CA), under a) the reference scenario (run 34) and b) the scenario with snow computation enabled (run 2). Stem count from 2020 to 2095 CE at the study site Lake Ilirney, Russia (RU), under c) the reference scenario (run 13) and d) the scenario with snow computation enabled (run 3). Stem count from 2020 to 2095 CE at the study site Road to Central, Alaska, USA (AK), under e) the reference scenario (run 15) and f) the scenario with snow computation enabled (run 3). (Runs based on the median migration rate calculated

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from 40 simulations per scenario). Each plot represents data from the simulation run with the median migration rate derived from 40 simulation runs per scenario. Stem count values are shown for the tree species occurring at the respective study site.