



1 Estimation of Metabolic Dynamics of Restored Seagrass Meadows in a Southeast Asia Islet: 2 Insights from Ex Situ Benthic Incubation

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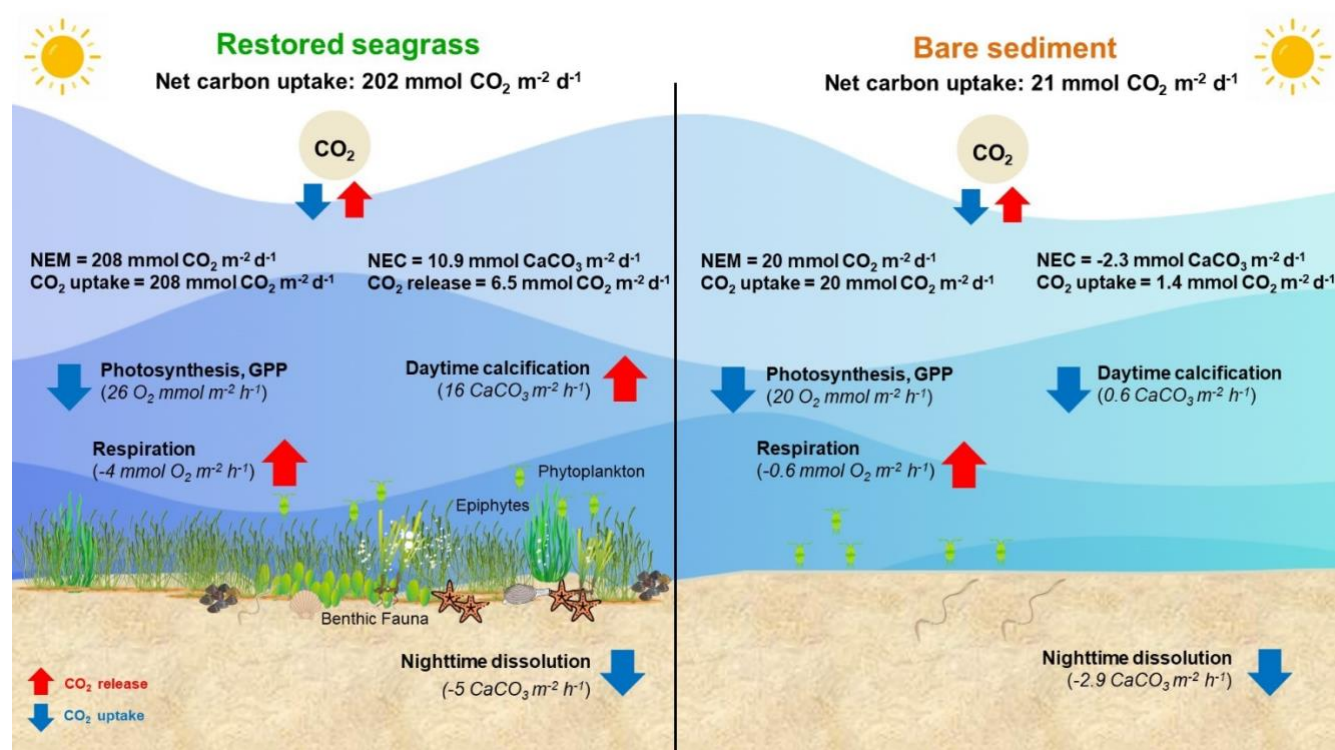
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17 **Abstract.** Seagrass meadows are vital carbon sinks, but their function is threatened by rapid decline,
18 driving restoration efforts to enhance coastal recovery and carbon removal. The capacity of these restored
19 seagrass as carbon sources or sinks depends largely on organic carbon metabolism and carbonate
20 dynamics. In this study, we employed ex situ core incubation to investigate the metabolic rates of
21 replanted seagrasses (SG), including gross primary productivity (GPP), community respiration (R), net
22 ecosystem metabolism (NEM), and net ecosystem calcification (NEC) in SG and surrounding bare
23 sediments (BS). SG exhibited higher GPP ($26.0 \pm 1.0 \text{ mmol O}_2 \text{ m}^{-2} \text{ h}^{-1}$ vs $0.7 \pm 0.1 \text{ mmol O}_2 \text{ m}^{-2} \text{ h}^{-1}$) and
24 NEM ($208.2 \pm 6.3 \text{ mmol O}_2 \text{ m}^{-2} \text{ d}^{-1}$ vs $20.1 \pm 2.8 \text{ mmol O}_2 \text{ m}^{-2} \text{ d}^{-1}$) than BS, indicating their potential as
25 carbon sinks by shifting benthic metabolism towards a more autotrophic state. In contrast, SG showed
26 higher daytime carbonate production and nighttime carbonate dissolution, which could offset each other,
27 resulting in no significant difference in NEC between SG and BS. In summary, our results found that the
28 SG exhibited significantly higher NEM compared to BS, while no significant difference was found for
29 NEC. Consequently, the net effect on the carbon uptake capacity of the restored seagrass is likely



30 increased, primarily due to the higher NEM. Our findings highlight the ecological significance of seagrass
 31 restoration in mitigating climate change through carbon removal. Ex situ core incubation method allows
 32 for the simultaneous measurement of organic and inorganic carbon metabolism. While ex situ core
 33 incubation enhances feasibility, in situ assessments are still necessary to validate the results and ensure a
 34 comprehensive understanding of seagrass ecosystem dynamics.



35

36 **Graphical abstract: Illustration of carbon uptake from organic carbon metabolism (GPP-gross**
 37 **primary productivity, R-respiration, NEM-net ecosystem metabolism) and carbonate dynamics**
 38 **(daytime calcification, nighttime dissolution, and NEC-net ecosystem calcification) in restored**
 39 **seagrass and bare sediment. Net Ecosystem Metabolism (NEM).**



40 **1 Introduction**

41 Seagrass meadows, comprising over 72 species, occupy just 0.1% of the ocean's surface, yet they are
42 highly productive and ecologically significant ecosystems in the marine environments (Fourqurean et al.,
43 2012; Short et al., 2011). These meadows play essential roles in nutrient and carbon cycling and serve as
44 key habitats for many marine species (Duarte et al., 2010; Fourqurean et al., 2012). Due to their relatively
45 complex structure, seagrass meadows capture and retain organic carbon (C_{org}) in the sediment, making
46 them one of the major carbon reservoirs globally (Duarte et al., 2005; Mcleod et al., 2011). Previous
47 estimates suggest that seagrasses account for approximately 15% of the total global carbon sequestered
48 in benthic sediments (Duarte et al., 2013), with burial rates 35 times that of tropical rainforests (Mcleod
49 et al., 2011).

50

51 In spite of their ecological significance, seagrass meadows have experienced a global decline, driven
52 primarily by human-induced activities such as coastal development, eutrophication, and deteriorating
53 water quality (Orth et al., 2006; Waycott et al., 2009). Since 1980, the global coverage of seagrass has
54 decreased by 110 km² annually, with the rate of decline increasing (Waycott et al., 2009). The loss is
55 frequently associated with increased water column turbidity and epiphytic shading, which reduce the light
56 for seagrass photosynthesis, leading to meadow degradation (Campbell et al., 2003; Orth et al., 2006).
57 Degradation also diminishes their capacity to modify local pH and influence the dynamics of dissolved
58 oxygen (DO) and dissolved inorganic carbon (DIC) (Hendricks et al., 2014). Moreover, the continued
59 loss of seagrass ecosystems raises concerns that vast amounts of previously sequestered carbon could be
60 released back in the atmosphere, converting seagrasses from carbon sinks to carbon sources and
61 intensifying global climate change (Macreadie et al., 2013). The ongoing decline could potentially release
62 up to 299 Tg of carbon annually, contributing roughly 10% of CO₂ emissions associated with
63 anthropogenic land-use changes (Fourqurean et al., 2012).

64

65 In response to these challenges, seagrass restoration has emerged as a critical strategy to mitigate
66 environmental degradation, enhance coastal resilience, and address global climate change (Juska and Berg



67 et al., 2022). Protecting and restoring seagrass meadows aligns with international goals like the Paris
68 Agreement, as these ecosystems offer significant potential for long-term carbon storage and climate
69 regulation (Fourqurean et al., 2012). However, despite growing restoration efforts, there remains limited
70 understanding of their success, particularly regarding benthic metabolism and carbon dynamics
71 (Kindeberg et al., 2024). While studies from temperate regions, such as the *Zostera marina* restoration in
72 the Virginia Coast (Rheuban et al., 2014), have provided valuable insights, data from tropical regions —
73 including Southeast Asia, a global hotspot for seagrass diversity — remain scarce (Duarte et al., 2010;
74 Ward et al., 2022; Chou et al., 2023). It represents a critical gap in our knowledge of the impact of
75 restoration efforts on carbon removal and ocean acidification mitigation.

76

77 Although there is increasing consensus on the potential of “Blue Carbon” storage in seagrass meadows
78 as a climate change mitigation strategy, the biogeochemical cycling within these ecosystems is complex.
79 Several processes, including ecosystem calcification, anaerobic metabolism, and bioturbation, can
80 counteract net organic carbon (OC) sequestration (Van Dam et al., 2021). These processes regulate local
81 DIC and total alkalinity (TA) budgets, adding complexity to accurately quantifying carbon sequestration
82 (Kindeberg et al., 2024). Overlooking these processes can result in significant overestimates of local
83 carbon sequestration rates and misinterpretations of the role seagrass meadows play in mitigating climate
84 change, potentially leading to inaccurate assessments of their carbon sink capacity (Johansen et al., 2023;
85 Chen et al., 2024; Fan et al., 2024).

86

87 Several methodologies were developed to quantify benthic metabolism, which is a crucial component of
88 biogeochemical cycling, including photosynthesis-irradiance curve (Kraemer and Alberte, 1993), the
89 open water O₂ mass balance approach (Odum, 1956; Chou et al., 2023), and aquatic eddy covariance
90 (Berg et al., 2022; Juska and Berg, 2022). While these methods provide important data, they might
91 overlook the complexities of bioturbation, remineralization, and carbonate dynamics (Olive et al., 2016;
92 Ward et al., 2022; Juska and Berg, 2022). In this study, we aim to address these knowledge gaps by
93 quantifying organic carbon metabolism (net ecosystem metabolism, NEM) and carbonate dynamics (i.e.,

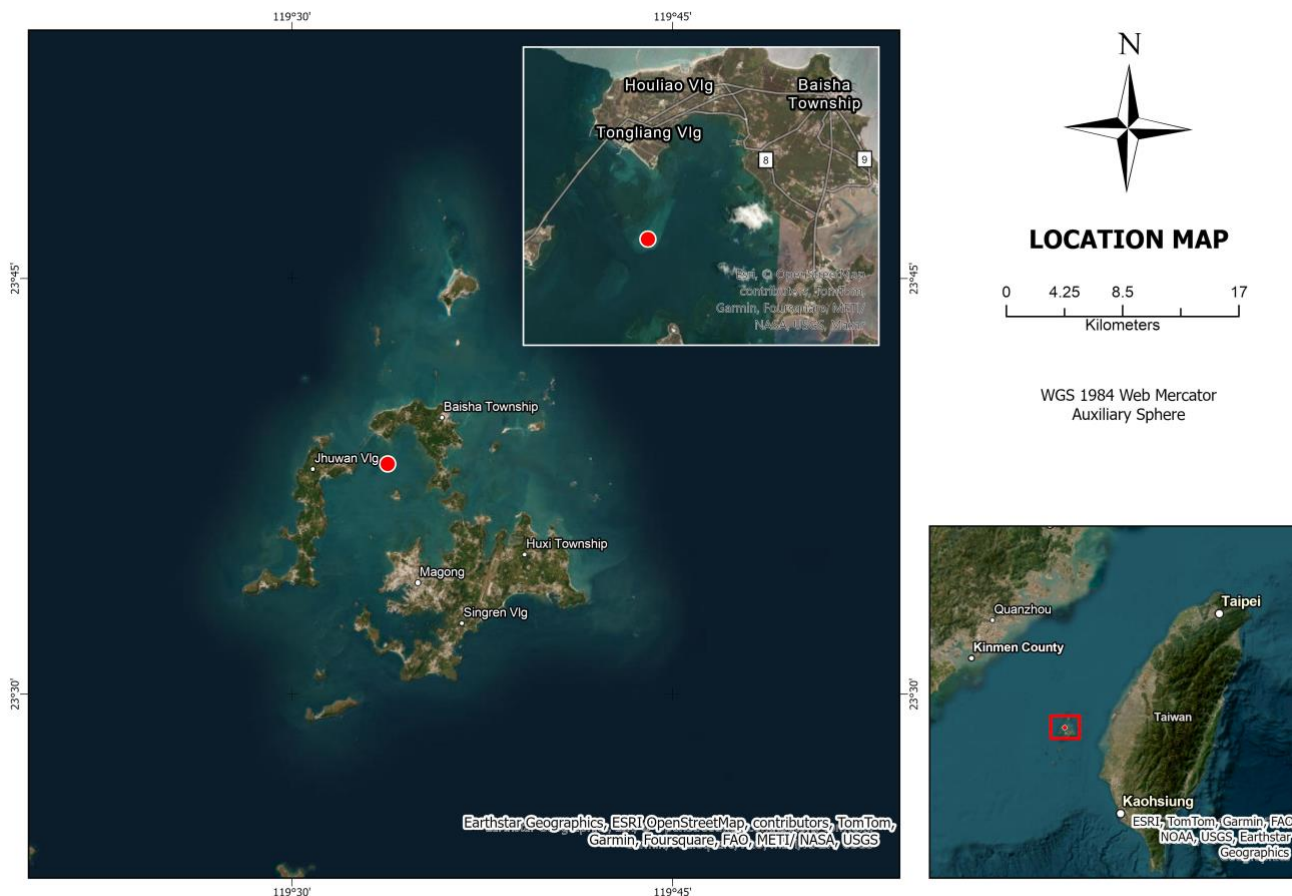


94 net ecosystem calcification, NEC) in restored seagrass meadows (SG) and adjacent bare sediment (BS)
95 habitats on a Southeast Asia islet, using an innovative ex situ benthic incubation.

96 **2 Materials and Methods**

97 **2.1 Study site**

98 The Penghu Islands, located in the southern part of Taiwan Strait (Fig. 1), host a range of seagrass species.
99 Notably, four species have been reported: *Halophila ovalis*, *Halodule pinifolia*, *Halodule uninervis*, and
100 *Zostera japonica* (Yang et al., 2002). The sampling location (23° 38' 18.38" N and 119° 33' 46.48" E) is
101 a restoration meadow dominated by *H. uninervis* and *H. ovalis*. This restoration site encompasses
102 approximately 3 hectares as per Coral Alen Atlas, with seagrass percent cover varying from 20% to 90%.
103 These seagrasses are subtidal, with water depths ranging from 1.7 meters to 4.4 meters. The substrate in
104 this area is composed of carbonate sand. The area supports a diverse community of bivalves (e.g., *Pinna*
105 sp.), gastropods, echinoderms, and various fish species, all of which were observed during the sampling.



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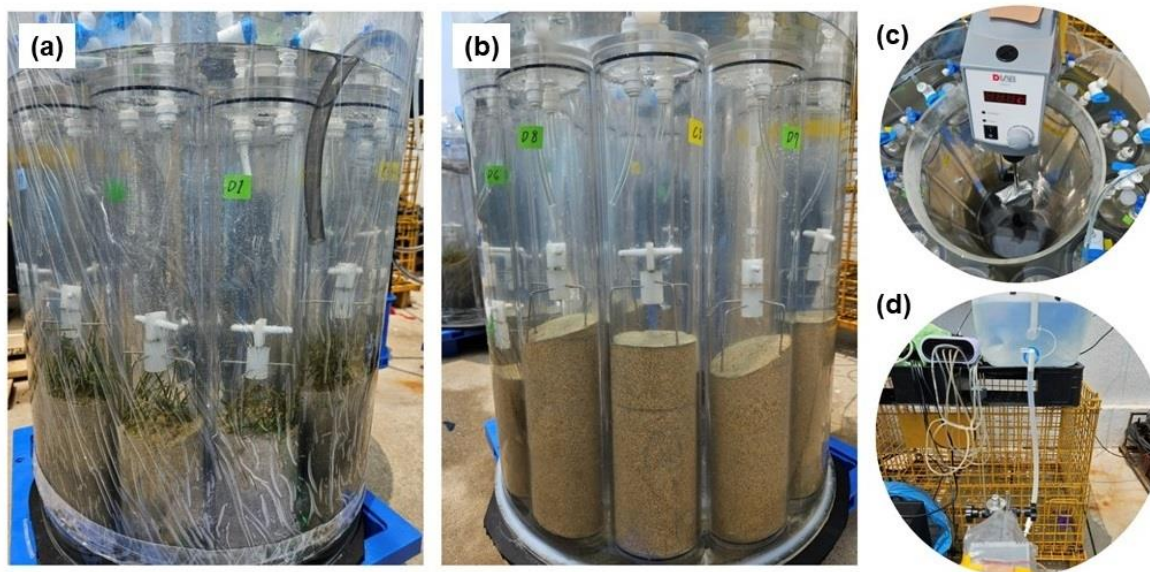
107 **Figure 1: Location map of sampling stations in restored seagrass in Penghu Island, Taiwan (Map**
108 **created in ArcGIS Pro. Source: Earthstar Geographics, ESRI OpenStreetMap, Contributors,**
109 **TomTom, Garmin, Foursquare, FAO, METI/NASA, USGS, NOAA).**

110 2.2 Ex situ core incubation system

111 The ex situ benthic core methodology used in this study was adapted from Chen et al. (2019) (Fig. 2).
112 This approach has been widely employed in various studies to assess nutrient concentrations and benthic
113 metabolism in coastal ecosystems and estuaries (Eyre & Ferguson, 2005; Maher & Eyre, 2011). Typically,
114 the ex situ core incubation involves 150-L treatment tanks containing aerated water. Each tank can
115 accommodate 10 plexiglass cores made of polycarbonate material, 10 cm in diameter and 50 cm in height.
116 The tanks were equipped with magnetic stir bars driven by a centrally located rotating motor fitted with



117 a magnet. The core has a plexiglass lid which contains two ports, one for probe insertion (Eyre &
118 Ferguson, 2005).



119

120 **Figure 2: Ex situ benthic chamber setup for measuring metabolic rates and carbonate dynamics in**
121 **seagrass meadows and bare sediment. The chambers contain seagrass samples (a), while the**
122 **chambers contain bare sediment (b). Insets show close-ups of the central rotating motor with a**
123 **magnet setup for water circulation (c), and the setup for continuous seawater supply (d).**

124 2.3 Sediment core collection and pre-incubation

125 The incubation was conducted on April 12-13, 2024. Twenty intact sediment cores, comprising both
126 seagrass and bare sediment, were collected on-site using the plexiglass tubes. The cores were inserted
127 about 20 cm into the sediment, keeping approximately 1.9 liters of water. Each core was sealed with a
128 gas-tight plexiglass plate at the bottom. The samples were brought back to the incubation site within two
129 hours of collection and allowed to settle for 24 hours. Additionally, 150 liters of water were collected on-
130 site for continuous supply during the experiment.

131



132 At the incubation site, the cores were uncovered and placed in 150-liter tanks filled with aerated seawater.
133 They were kept at in situ temperature, exposed to natural sunlight, and continuously recirculated. The
134 stirring rate was controlled to prevent sediment resuspension (Ferguson et al., 2004). The cores underwent
135 a 24-hour pre-incubation period to promote stable sediment profiles. The seagrass composition within
136 the collected cores for ex situ core incubation was dominated by *H. uninervis* and *H. ovalis*. The shoot
137 count of *H. uninervis* ranged from 20 to 40 shoots per 0.008 m², while *H. ovalis* ranged from 2 to 20
138 shoots per 0.008 m².

139 **2.4 Sample collection and analysis**

140 Following pre-incubation, the cores was tightly closed using a plexiglass lid. Temperature, salinity, and
141 pH were determined using a YSI ProDSS Multiparameter water quality checker, while DO (mg l⁻¹) was
142 measured with a thermo DO probe. Both probes were calibrated with calibration standards. Measurements
143 were taken at midnight (24:00 h) with 2-hour intervals and ended at noon. Photosynthetically active
144 radiation (PAR) levels were measured using SQ-420X Smart Quantum Sensor positioned atop the
145 incubation tank.

146

147 After measurements, three 150 ml seawater samples were collected separately from the SG and BS cores
148 using a syringe for DIC and pH analysis. The water samples were processed with 60 µL HgCl₂ solution
149 to stop any biological activity. DIC analysis was performed using a non-dispersive infrared method with
150 a DIC analyzer (AS-C3, Apollo SciTech Inc.), following the approach of Dickson et al. (2007) and our
151 past studies (Chou et al., 2018; 2021; Fan et al., 2024). For each DIC run, we used certified reference
152 material (Batch no. 206) sourced from A. G. Dickson at Scripps Institution of Oceanography to check for
153 drift and systematic bias. pH values were measured spectrophotometrically in total scale at 25 °C
154 following Clayton and Byrne (1993). Data from DIC and pH, along with actual temperature and salinity,
155 were used to calculate the TA, partial pressure of CO₂ (*p*CO₂), and aragonite saturation state (Ω_{Ar}) using
156 the Excel macro CO2SYS version 2.1 (Pelletier et al. 2011). The dissociation constants for carbonic acid
157 applied in these calculations were obtained from Mehrbach et al. (1973) and subsequently refined by
158 Dickson and Millero (1987).



159 **2.5 Benthic flux rate calculations**

160 Areal rates of R, GPP, NPP, and NEM were calculated based on changes in DO concentrations, following
161 equation 1 (Eyre et al. 2011). Respiration rates were determined from concentration data collected during
162 the initial dark period (midnight to dawn) (eq. 2). NPP was calculated based on light O₂ flux
163 measurements from dawn to noon (eq. 3). Hourly GPP rates were computed as the difference between R
164 and NPP rates (eq. 4). NEM was calculated using equation 5. Positive values indicate autotrophic, while
165 negative values represent heterotrophic.

$$166 \quad F = [(C_{t1} - C_{t0}) \times V/A]/T \quad (\text{eq. 1})$$

167 Where F = flux rate ($\mu\text{mol m}^{-2} \text{h}^{-1}$), C_{t0} and C_{t1} = concentration in the overlying water at the start and end
168 of the time period ($\mu\text{mol l}^{-1}$), respectively, V = volume of overlying water in the core (l), A = surface area
169 in the sediment core (m²), and T = incubation period (h).

$$170 \quad R = \text{dark O}_2 \text{ flux (negative)} \quad (\text{eq. 2})$$

$$171 \quad \text{NPP} = \text{light O}_2 \text{ flux (positive)} \quad (\text{eq. 3})$$

$$172 \quad \text{GPP} = \text{NPP} - R \quad (\text{eq. 4})$$

$$173 \quad \text{NEM} = (\text{GPP} \times 12) - (R \times 24 \text{ h} \times -1) \quad (\text{eq. 5})$$

174

175 NEC rates ($\text{mmol CaCO}_3 \text{ m}^{-2} \text{ h}^{-1}$) were estimated from the change of total alkalinity, assuming these
176 changes are only due to CaCO₃ precipitation and dissolution (eq. 6) (Roth et al., 2019; Van Dam et al.,
177 2019):

$$178 \quad \text{NEC} = -0.5 \frac{\Delta n\text{TA}}{\Delta t} \times hp \quad (\text{eq. 6})$$

179 Here, $\Delta n\text{TA}$ = change in $n\text{TA}$ ($n\text{TA} = \text{TA} \times \text{SSS}_{\text{average}}/\text{SSS}$) over the Δt (time), h = Volume/Area, and p
180 = water density. The -0.5 scalar factor was applied to account for the stoichiometric relationship, where
181 2 moles of TA produce 1 mole of CaCO₃. Day and night incubations were conducted to obtain daily NEC
182 fluxes. NEC is positive with TA consumption, indicating CaCO₃ precipitation, and negative with TA
183 production, indicating CaCO₃ dissolution.



184 **2.5 Statistical analysis**

185 Independent sample T-tests were applied to compare metabolic rates (R, NPP, GPP, NEM, NEC) between
186 SG and BS using SPSS v. 17. Data were subjected to a normality test before performing the analysis.
187 Least-squares linear regression was employed to assess the correlation between changes in DO in the SG
188 and BS. The Mann-Whitney U test was applied for carbonate chemistry analysis due to the non-normal
189 distribution of data.

190 **3 Results**

191 **3.1 Water quality and carbonate chemistry**

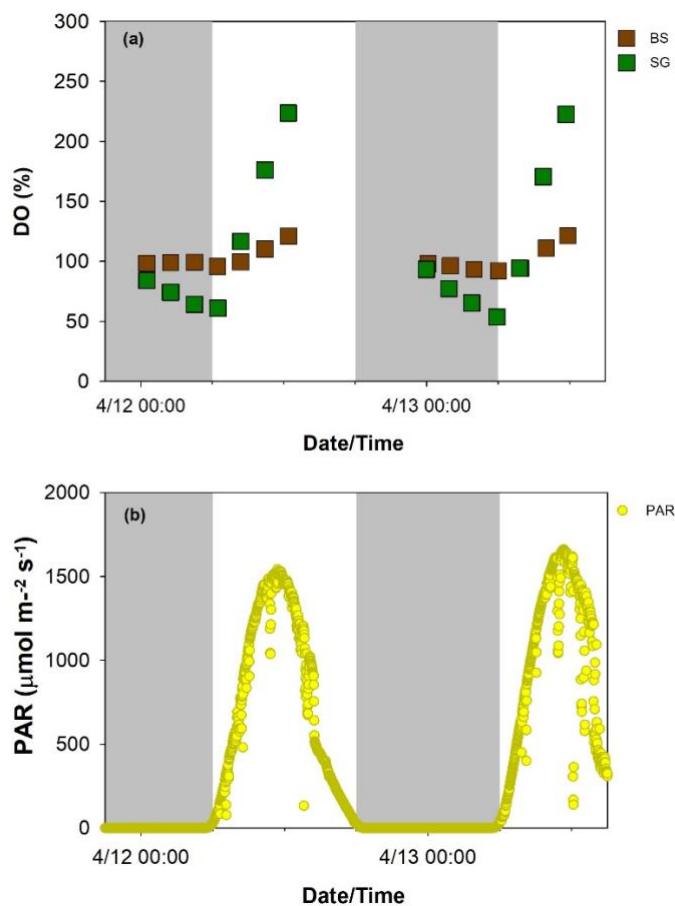
192 Diurnal patterns of water quality and carbonate parameters for SG and BS during the two-day ex situ core
193 incubation are illustrated in Figs. 3 and 4, respectively. The temperature in both treatments ranged from
194 22 to 29 °C, while salinity levels spanned from 35 to 36. These values were similar to in situ measurements
195 obtained from the seagrass beds using a CTD profiler. During the daytime (6:00 AM to 12:30 PM), PAR
196 levels ranged from 26 $\mu\text{mol m}^{-2} \text{s}^{-1}$ to a peak of 1662 $\mu\text{mol m}^{-2} \text{s}^{-1}$, with the highest intensities observed
197 at midday. The average PAR measured 953 $\mu\text{mol m}^{-2} \text{s}^{-1}$ on the first day of incubation, increasing slightly
198 to 1026 $\mu\text{mol m}^{-2} \text{s}^{-1}$ on the second day. DO saturation levels were more variable in SG than BS, with
199 values ranging from 54% to 224% and 92% to 123%, respectively. DO saturation levels in both treatments
200 followed a diel pattern, with lower nighttime and higher daytime values.

201

202 Both $n\text{DIC}$ ($n\text{DIC} = \text{DIC} \times \text{SSS}_{\text{average}}/\text{SSS}$) and pH_T displayed greater diurnal fluctuations at SG compared
203 to the BS. At SG, $n\text{DIC}$ ranged from 1660 to 2118 $\mu\text{mol kg}^{-1}$ (mean \pm SE: $1963 \pm 41 \mu\text{mol kg}^{-1}$), and
204 followed a diel pattern. pH_T ranged from 7.81 to 8.37 at SG (mean \pm SE: 7.99 ± 0.05), following the
205 opposite trend to $n\text{DIC}$, with values decreasing at night and increasing during the day. This daytime
206 increase in pH_T at SG indicated the potential role of seagrass in mitigating ocean acidification effects
207 during daylight hours. At the BS site, these parameters were less variable, with $n\text{DIC}$ values ranging from
208 1948 to 2029 $\mu\text{mol kg}^{-1}$ and pH_T from 7.84 to 7.99, with mean values of $1993 \pm 7 \mu\text{mol kg}^{-1}$ and $7.93 \pm$
209 0.01, respectively. Similarly, the calculated $n\text{TA}$ was also more fluctuating in SG than BS, with mean

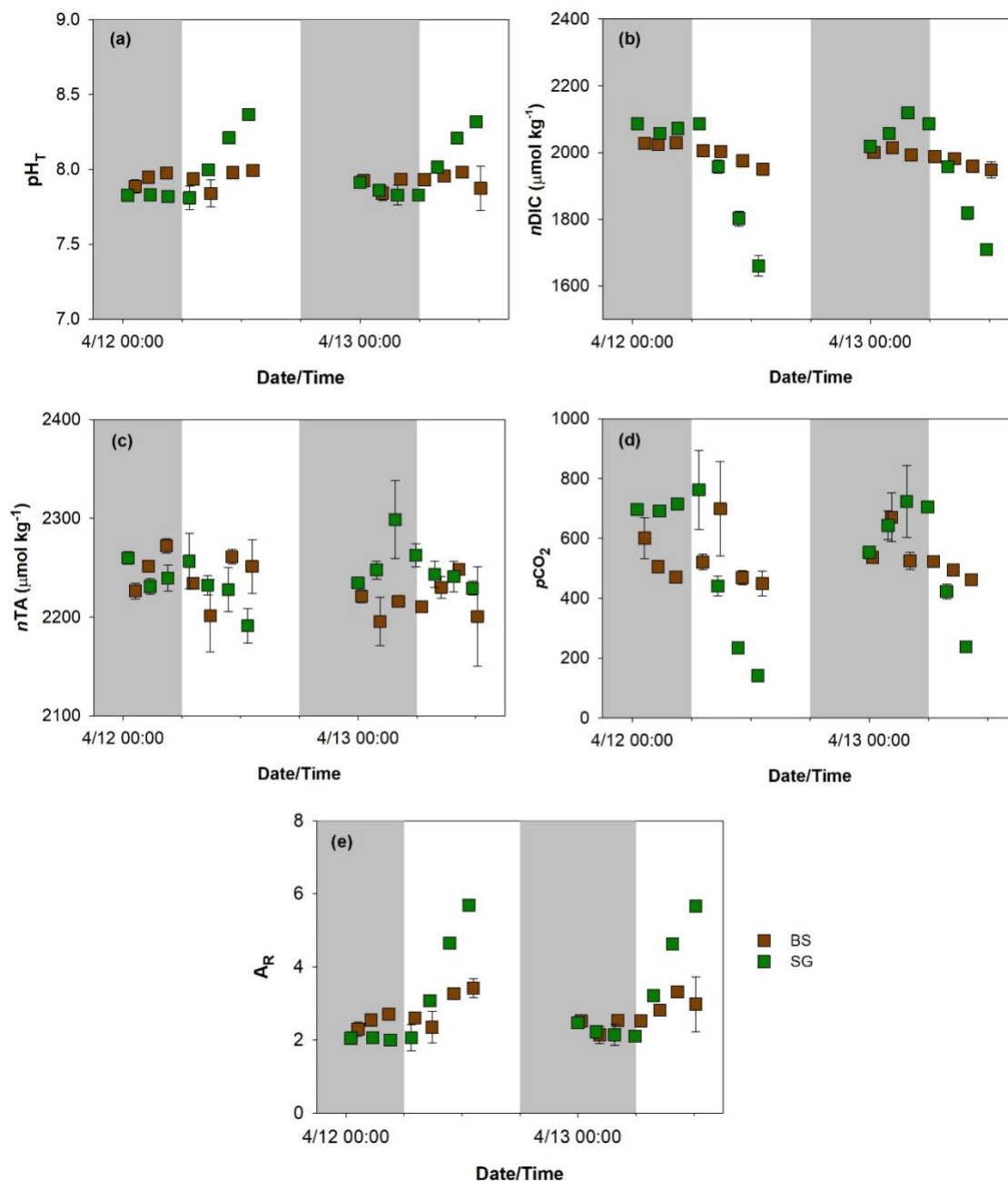


210 values of $2243 \pm 6 \mu\text{mol kg}^{-1}$ and $2230 \pm 6 \mu\text{mol kg}^{-1}$, respectively. The calculated $p\text{CO}_2$ displayed a
211 broader range at SG (142 to $762 \mu\text{atm}$; mean \pm SE: 510 ± 62) compared to BS (450 to $699 \mu\text{atm}$; mean \pm
212 SE: 524 ± 22), suggesting a more dynamic carbon cycling potentially driven by seagrass metabolic
213 activity. The mean Ω_{Ar} was higher in SG (3.14 ± 0.37) compared to BS (2.72 ± 0.11), indicating more
214 favorable conditions for calcification at the seagrass site. Mann–Whitney test on carbonate chemistry
215 revealed no significant distinction between SG and BS ($p\text{H}_T$ $p = 0.713$; $n\text{DIC}$ $p = 0.419$; $n\text{TA}$ $p = 0.679$;
216 Ω_{Ar} $p = 0.511$).



217

218 **Figure 3: Diurnal pattern of dissolved oxygen (DO, a) in replanted seagrass (SG, green square) and**
219 **bare sediment (BS, brown square), and photosynthetically active radiation (PAR, b) during the**
220 **two-day (April 12-13, 2024) incubation.**



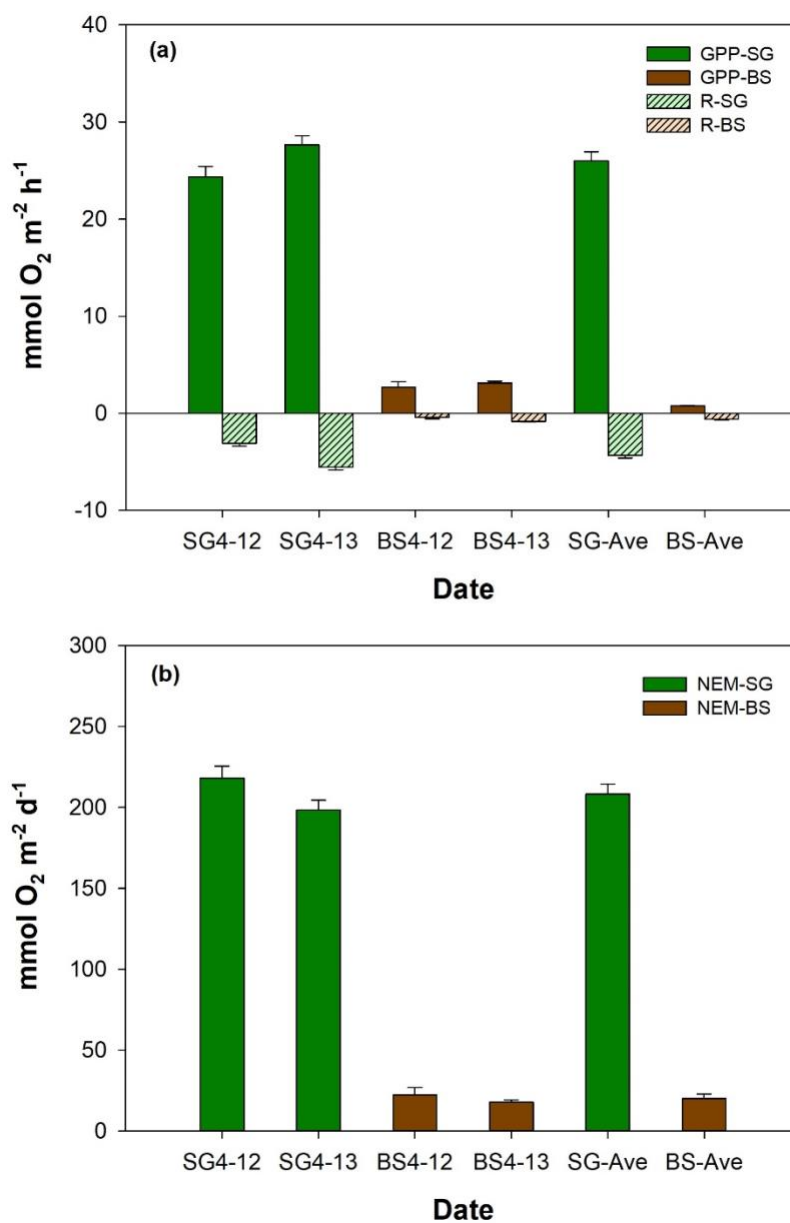
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222 **Figure 4: Total scale pH (pH_T , a), normalized dissolved inorganic carbon ($nDIC$, b), normalized**
 223 **total alkalinity (nTA , c), partial pressure of carbon dioxide (pCO_2 , d), and aragonite saturation state**
 224 **(Ω_{AR} , e) in replanted seagrass (SG, green square) and bare sediment (BS, brown square) during**
 225 **the two-day (April 12-13, 2024) incubation. $n=3$, mean \pm SE.**



226 **3.2 Respiration, gross primary production, and net ecosystem metabolism**

227 Figure 5 illustrates the comparison of metabolic rates between SG and BS. The mean respiration rates in
228 SG ($-4.3 \pm 0.3 \text{ mmol O}_2 \text{ m}^{-2} \text{ h}^{-1}$) were significantly higher than in BS ($-0.6 \pm 0.1 \text{ mmol O}_2 \text{ m}^{-2} \text{ h}^{-1}$), by
229 approximately 8-fold difference ($p < 0.01$). The mean GPP in SG was $26.0 \pm 1.0 \text{ mmol O}_2 \text{ m}^{-2} \text{ h}^{-1}$, which
230 is 35-fold higher than in BS ($0.7 \pm 0.1 \text{ mmol O}_2 \text{ m}^{-2} \text{ h}^{-1}$) ($p < 0.01$). GPP was always higher than R in both
231 systems, with mean GPP/R ratios of 3.4 and 1.9 in SG and BS, respectively. For NEM, both systems
232 displayed positive values, indicating net autotrophy, with SG being 10-fold higher ($208.2 \pm 6.3 \text{ mmol O}_2$
233 $\text{m}^{-2} \text{ d}^{-1}$) compared to BS ($20.1 \pm 2.8 \text{ mmol O}_2 \text{ m}^{-2} \text{ d}^{-1}$) ($p < 0.01$). Both R and GPP in SG and BS increased
234 on the second day of incubation, while NEM showed a slight decrease, but these changes were not
235 statistically significant.



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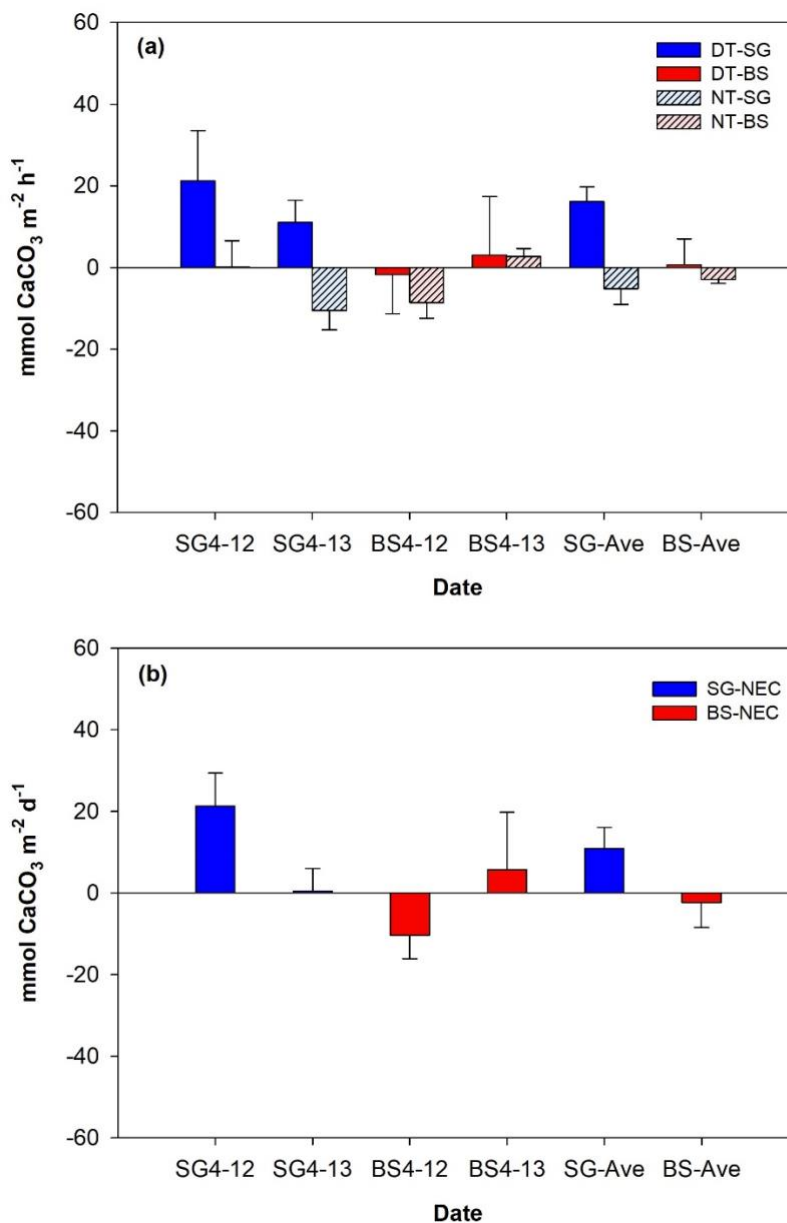
237 **Figure 5: Mean (\pm SE, standard error) values of (a) metabolic rates such as respiration (R), gross**
238 **primary productivity (GPP), and (b) net ecosystem metabolism (NEM,) of restored seagrass (SG,**
239 **green bars) and bare sediment (BS, brown bars) in Penghu during the two-day (April 12-13, 2024)**
240 **incubation (n=9).**

241



242 **3.2 Calcium carbonate precipitation, dissolution, and net ecosystem calcification**

243 The NEC values over a diel cycle for SG and BS demonstrated differences in their overall carbonate
244 dynamics (Fig. 6). Over the two-day incubation period, SG exhibited a net calcifying system with a mean
245 positive daily NEC means ($10.9 \pm 5.1 \text{ mmol CaCO}_3 \text{ m}^{-2} \text{ d}^{-1}$), driven by daytime calcification (16.1 ± 3.7
246 $\text{mmol CaCO}_3 \text{ m}^{-2} \text{ h}^{-1}$) despite nighttime dissolution ($-5.2 \pm 3.9 \text{ mmol CaCO}_3 \text{ m}^{-2} \text{ h}^{-1}$). In contrast, BS
247 supported a net-dissolving system with mean daily NEC ($-2.3 \pm 6.2 \text{ mmol CaCO}_3 \text{ m}^{-2} \text{ d}^{-1}$). Mean daytime
248 calcification and nighttime dissolution were $0.6 \pm 6.4 \text{ mmol CaCO}_3 \text{ m}^{-2} \text{ h}^{-1}$ and $-3.0 \pm 0.9 \text{ mmol CaCO}_3$
249 $\text{m}^{-2} \text{ h}^{-1}$, respectively. Both systems followed a general diurnal pattern, with positive NEC during the day
250 (calcifying) and negative at night (dissolving).



251

252 **Figure 6: Mean (\pm SE, standard error) values of daytime (DT) and nighttime (NT) calcification (a),**
253 **and net ecosystems calcification (NEC, b) of restored seagrass (SG, blue bars) and bare sediment**
254 **(BS, red bars) in Penghu during the two-day (April 12-13, 2024) incubation (n=9).**



255 **4 Discussion**

256 Seagrass meadows are widely recognized as an important blue carbon ecosystem with substantial
257 potential to mitigate anthropogenic CO₂ emissions. Although research on seagrass ecosystems has grown
258 in recent years, significant gaps remain in understanding their carbon dynamics. In particular, the balance
259 of organic and inorganic carbon processes within these systems is not fully understood. Meanwhile, global
260 seagrass coverage continues to decline, which has increased the urgency of restoration efforts (Waycott
261 et al. 2009). Restoring seagrass meadows to enhance carbon sequestration has become increasingly
262 important. Currently, most studies on restored seagrass meadows focus primarily on the burial of
263 particulate organic carbon (Greiner et al. 2013), with far fewer exploring both organic metabolism and
264 carbonate cycling in restored seagrass meadows. Here, we present the first dataset on carbon uptake
265 through metabolic rates and calcification measurements in restored seagrass meadows within tropical
266 regions.

267 **4.1 Restoration of seagrass enhances metabolic rates**

268 The metabolic rates estimated in present study were comparable to those recorded in other seagrass
269 meadows (Table 1). Our GPP in SG was 24% and 37% higher than the tropical and global averages,
270 respectively, but 38% lower than Dongsha Island, Taiwan (Chou et al., 2023). It is also comparable to
271 measurements reported for *H. uninervis* in Tropical Australia (Table 1). Conversely, the R values
272 estimated in this study were roughly half lower than the tropical and global averages (Duarte et al., 2010).
273 Our NEM (214 mmol O₂ m⁻² d⁻¹) is within the range of previous estimates for tropical seagrass meadows
274 (-477.28 to 484.20 mmol O₂ m⁻² d⁻¹) and global estimates (-477.28 to 531.63 mmol O₂ m⁻² d⁻¹). In addition
275 to these global comparisons, our study reveals a clear distinction in metabolic rates (e.g. GPP, R, NEM)
276 between SG and BS. The GPP and R in restored seagrass meadows were 35 and 7 times greater than in
277 BS. The relatively higher metabolic rates in seagrass meadows compared to bare sediments have also
278 been observed in other studies (Table 1). For instance, a two-year-old restored *Halodule wrightii* meadow
279 demonstrated a 13-fold increase in NEM relative to bare sediment (Egea et al., 2023). Similarly,
280 *Posidonia oceanica* exhibited a notable 70-fold increase in metabolic rates compared to bare sediment
281 (Barron et al., 2006). Furthermore, *Zostera marina* exhibits net autotrophy while bare sediments are net



282 heterotrophy (Attard et al., 2019; Chen et al., 2019). Such patterns highlight the fundamental ecological
 283 functions restored seagrass meadows play relative to unvegetated/bare sediments. The increase in GPP
 284 reflects the enhanced carbon fixation capacity of seagrass meadows, while the elevated R indicates active
 285 organic matter decomposition and microbial respiration (Duarte and Krause-Jensen, 2017). According to
 286 Duarte et al. (2010), seagrass meadows generally become autotrophic ($NEM > 0$) when GPP is greater
 287 than $186 \text{ mmol O}_2 \text{ m}^{-2} \text{ d}^{-1}$, shifting to heterotrophy ($NEM < 0$) at lower levels. Based on this threshold,
 288 our mean GPP for restored seagrass exceeded the value for autotrophy, resulting in a positive NEM which
 289 is consistent with their global assessment. The NEM observed in SG was 10 times higher than in BS,
 290 suggesting that SG sequesters significantly more carbon than BS. These findings highlight that seagrass
 291 restoration significantly boosts metabolic rates and enhances carbon cycling. Given the increasing loss of
 292 global seagrass cover, restoration not only boosts ecosystem productivity but also strengthens the ability
 293 of coastal systems to remove carbon, thereby contributing to climate change mitigation efforts.

294

295 **Table 1. Comparison of metabolic rates from global estimates. GPP and R values are expressed in**
 296 **$\text{mmol O}_2 \text{ m}^{-2} \text{ h}^{-1}$ units, while NEM in $\text{mmol O}_2 \text{ m}^{-2} \text{ d}^{-1}$.**

Location	Method	Seagrass Community	GPP	R	NEM	References
Taiwan	Ex situ benthic chambers	Bare sediment <i>H. uninervis</i> , <i>H. ovalis</i>	0.74 ± 0.09 25.99 ± 0.96	0.62 ± 0.09 4.32 ± 0.26	20.10 ± 2.84 208.21 ± 6.33	This study
Taiwan	Open water mass balance	<i>Thalassia</i> , <i>Cymodocea</i>	42.25 ± 14.42	20.71 ± 7.13	8 ± 61	Chou et al., 2023
Mexico	In situ benthic chambers	Bare sediment 2-year <i>H. wrightii</i> 4-year <i>H. wrightii</i> 4-year <i>H. wrightii</i>	2.13 ± 0.58 13.76 ± 3.35 9.24 ± 2.34 9.34 ± 0.35	0.73 ± 0.16 2.61 ± 0.40 1.60 ± 0.19 2.15 ± 0.25	8.1 ± 10.9 102.4 ± 31.5 72.5 ± 27.9 60.7 ± 4.7	Egea et al., 2023
Sweden	Aquatic eddy covariance and benthic chambers	3-year-old restored seagrass (<i>Z. marina</i>) 7-year-old restored seagrass (<i>Z. marina</i>)			-5 to -15 -21	Kindeberg et al., 2024
Finland		Bare sediment	1.60	0.82	-0.14	

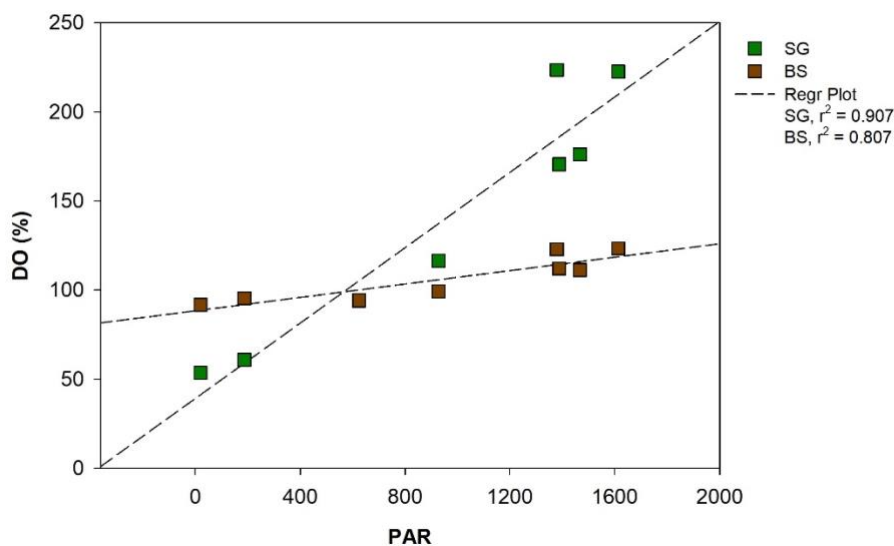


	Aquatic eddy covariance	<i>Z. marina</i>	3.74	1.71	4.17	Attard et al., 2019
Australia	Ex situ benthic	Bare sediment	2.28	1.26	-2.74	Chen et al., 2019
		<i>Zostera</i> sp.	6.94	2.74	7.12	
		<i>Halophila</i> sp.	2.05	1.60	-13.70	
Tropical Australia	Combined methods	<i>H. uninervis</i>	23.42 ± 3.67	9.63 ± 4.04	50 ± 53	Duarte et al., 2010
Tropical	Combined methods	All species	21 ± 0.6	9 ± 0.6	24 ± 8	Duarte et al., 2010
Global	Combined methods	All species	19 ± 0.5	8 ± 0.4	27 ± 6	Duarte et al., 2010
Spain	In situ benthic chambers	Bare sediment	0.43	0.22	0.27	Barron et al., 2006
		<i>P. oceanica</i>	7.72	3.18	16.44	

297 *The daily values of R and GPP reported in the literature were divided by 24 and 12, respectively, to calculate the*
 298 *hourly values.*

299

300 Key drivers of elevated metabolic rates in tropical meadows include greater PAR availability,
 301 aboveground biomass, and higher temperatures (Ganguly et al., 2017; Ward et al., 2022). Many tropical
 302 species grow near their optimal photosynthetic and physiological conditions (Lee et al., 2007; Koch et al.,
 303 2012), efficiently capturing light in shallow, clear waters, which contributes to higher NEP (Ralph et al.,
 304 2007). In our study, DO variation corresponds to light intensity (Figs. 3 and 7), suggesting that the
 305 elevated GPP observed in seagrass meadows could be driven by higher light intensity. This is likely due
 306 to the relatively lower canopy cover of *H. uninervis* and density in SG, which reduces shading within the
 307 seagrass. As a result, more light penetrates to the leaves, increasing their photosynthetic surface area and
 308 contributing to NEM (Ralph et al., 2007).



309

310 **Figure 7: Regression plot between photosynthetically active radiation (PAR, $\mu\text{mol m}^{-2} \text{s}^{-1}$) vs**
311 **dissolved oxygen (DO, %) in restored seagrass (SG, green square) and bare sediment (BS, brown**
312 **square).**

313

314 Several studies indicate that restored seagrass can achieve primary productivity and carbon sequestration
315 levels comparable to natural meadows, although recovery depends on the extent of degradation,
316 restoration success, and site-specific habitat conditions (Oreska et al., 2017; Marbà et al., 2015). For
317 example, long-term research in Florida Bay demonstrated that sediment carbon sequestration rates and
318 plant biomass took nearly a decade to match those of natural meadows (Greiner et al., 2013). The ability
319 of restored meadows to maintain net autotrophy is crucial for their role as carbon sinks (Kennedy et al.,
320 2010). This is particularly relevant for climate change mitigation strategies, where the conservation and
321 rehabilitation of this ecosystem are recognized as natural climate solutions (Griscom et al., 2017).
322 Nonetheless, a recent investigation on restored seagrass exhibits net heterotrophy, as observed by
323 Kindeberg et al. (2024) in both 3-year and 7-year-old meadows in Sweden. A similar pattern also reported
324 in some natural seagrass meadows in Australia (Chen et al., 2019) (Table 1). This discrepancy underscores
325 the variability in seagrass productivity and metabolic processes based on geographical location and



326 environmental conditions, highlighting the need for region-specific assessments to fully understand
327 seagrass ecosystem dynamics. Long-term studies should also consider temporal and annual variations.

328 **4.2 Calcification dynamics in restored seagrass**

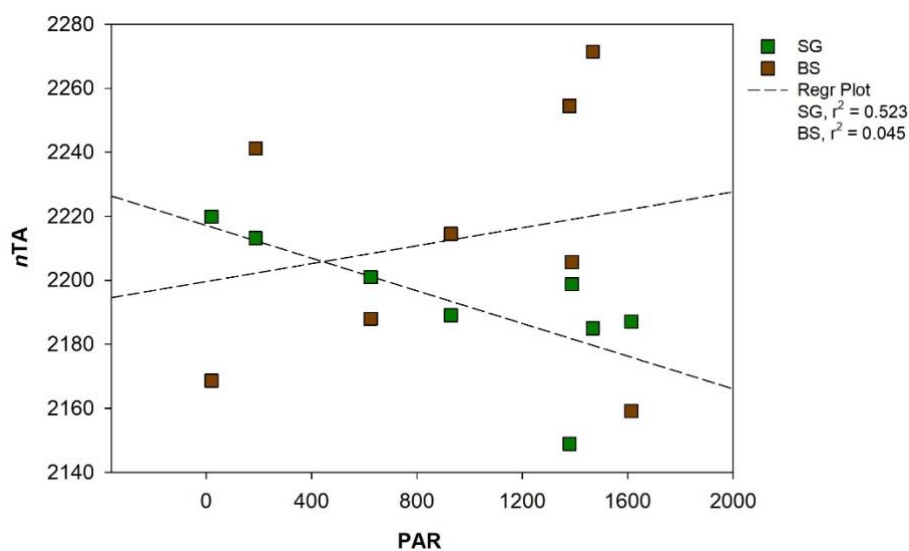
329 Our results show that restored seagrass meadows exhibit significantly higher CaCO_3 cycling — both
330 formation and dissolution — compared to bare sediments. This corroborates with prior studies, which
331 documented enhanced carbonate dynamics in vegetated habitats relative to unvegetated sediments. For
332 instance, *P. oceanica* and *Thalassia testudinum* meadows have been shown to promote both CaCO_3
333 production and dissolution (Burdige and Zimmerman, 2002; Barrón et al., 2006), with tropical seagrass
334 ecosystems displaying similar patterns (Chou et al., 2021; Fan et al., 2024). Further, our data revealed a
335 typical diurnal pattern, with positive values during daytime (net calcifying) and negative values during
336 nighttime (net dissolving). These findings align with previous estimates, such as those in Florida Bay,
337 which reported similar diurnal calcification dynamics (Yates and Halley, 2006).

338

339 The variations of CaCO_3 production and dissolution in surface waters and sediment are related to the
340 carbon cycle through photosynthesis and respiration (Yates and Halley 2006). During daylight hours,
341 photosynthesis raises pH and reduces CO_2 levels in the water, creating favorable conditions for calcium
342 carbonate precipitation—a process referred to as light-enhanced calcification (Schneider et al., 2009).
343 We found a significant positive correlation between PAR and $n\text{TA}$ changes ($r^2=0.52, p<0.05$), suggesting
344 that increased light availability may enhance calcification by photoautotrophs in restored seagrass areas
345 during the day (Fig. 8). Additionally, our data showed a significant negative correlation between $n\text{TA}$
346 flux and NEM ($r^2=0.54, p<0.01$), indicating that higher photosynthetic activity (positive NEM) promotes
347 calcification by consuming TA, while lower NEM or net heterotrophy contributes to TA production,
348 likely through carbonate dissolution or anaerobic decomposition (Fig. 9). Similar relationships between
349 photosynthesis and calcification have been reported in marine calcifiers (Mallon et al., 2022), and the
350 influence of epiphytic organisms in promoting calcification during active photosynthesis has been
351 highlighted in seagrass meadows such as *P. oceanica* (Barrón et al., 2006). At night, carbonate dissolution
352 predominates as aerobic respiration produces CO_2 and carbonic acid in sediment porewater (Eyre et al.,

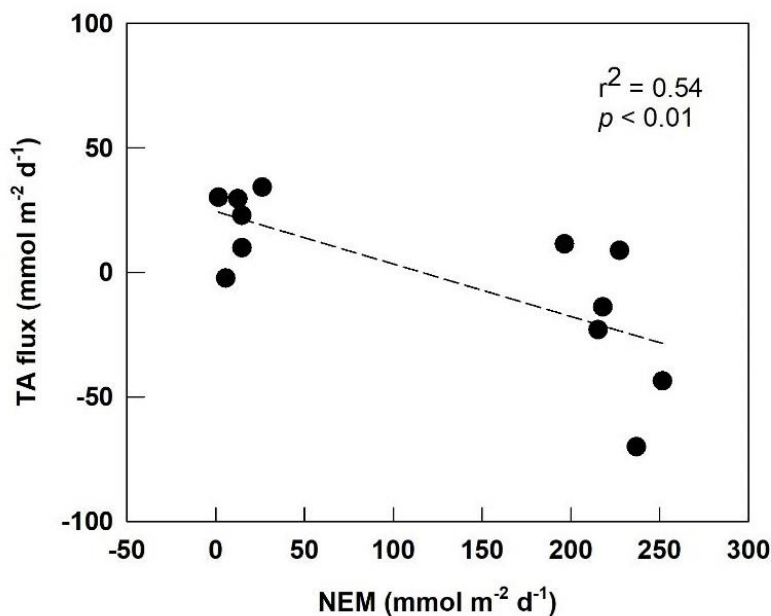


353 2014), lowering carbonate saturation and driving mineral dissolution (Burdige and Zimmerman, 2002;
354 Burdige et al., 2008; Chou et al., 2021; Fan et al., 2024). The degree of dissolution is directly link to the
355 rate of organic matter decomposition, which depends on the quantity of organic matter, its reactivity, and
356 oxygen availability (Anderson et al., 2005; Morse et al., 2006). The presence of shoot density and root
357 biomass in restored seagrass meadows enhances organic matter supply and decomposition in sediment,
358 further driving nighttime dissolution (Holmer et al., 2013).



359

360 **Figure 8: Regression plot between photosynthetically active radiation (PAR, $\mu\text{mol m}^{-2} \text{s}^{-1}$) vs**
361 **normalized total alkalinity (nTA, $\mu\text{mol kg}^{-1}$) in restored seagrass (SG, green square) and bare**
362 **sediment (BS, brown square).**



363

364 **Figure 9: Linear regression showing the relationship between total alkalinity (TA) flux and net**
365 **ecosystem metabolism (NEM) in this study.**

366

367 Over cumulative days, our NEC measurements indicate that restored seagrass meadows support overall
368 net calcification, whereas BS supports net dissolution. Our estimates are similar to those from Australia
369 (Walker et al., 1988) and seven times higher than Mediterranean seagrass net calcification rates (Barrón
370 et al., 2006), which are 295 g CaCO₃ m⁻² yr⁻¹ (8.8 mmol CaCO₃ m⁻² d⁻¹) and 51 g CaCO₃ m⁻² yr⁻¹ (1.40
371 mmol CaCO₃ m⁻² d⁻¹), respectively. In contrast, our findings are lower than those reported in the
372 Caribbean region of Mexico, where ex situ estimates ranged from 14 to 153 mmol CaCO₃ m⁻² d⁻¹
373 (Enriquez and Schubert, 2014). This highlights the enhanced carbonate production potential in tropical
374 seagrass meadows. A positive net calcification system occurs when CaCO₃ precipitation exceeds
375 dissolution within the system (Kleypas et al., 2001; Eyre et al., 2014). Restoration of seagrass meadows
376 provides a substrate for diverse calcifying organisms, including crustose coralline algae, bryozoans,
377 foraminifera, and serpulids, which enhance carbonate production (Beavington-Penney et al., 2005).
378 Epiphytes on seagrass leaves significantly contribute to CaCO₃ production, with tropical seagrass



379 meadows typically supporting higher carbonate loads than temperate ones. Reported production rates
380 span from 180 g CaCO₃ m⁻² yr⁻¹ in Jamaica (Land, 1970) to 2800 g CaCO₃ m⁻² yr⁻¹ in Barbados (Patriquin,
381 1972), underscoring regional variability in seagrass-associated calcification. Moreover, fluctuations in
382 CO₃²⁻ concentrations are crucial in regulating the capacity of calcifying organisms to form CaCO₃. Our
383 data reveals a consistently higher aragonite saturation state (Ω_a) compared to bare sediment (BS). Notably,
384 SG environments exhibit significant peaks in aragonite saturation, with a maximum value of 5.686,
385 whereas the highest Ω_{Ar} in BS is 3.419. Seagrass photosynthesis raises pH and Ω_{Ar} , enhancing the
386 calcification of surrounding calcifying organisms (De Beer and Lakrum, 2001). However, the
387 consumption of TA by calcifiers during the calcification process releases CO₂, potentially counteracting
388 pH increases and partially offsetting the net carbon uptake potential of seagrass ecosystems (Alongi et
389 al., 2008; Mazarrasa et al., 2015; Saderne et al., 2019). This highlights the dual role of seagrass restoration
390 in supporting biodiversity and CO₂ uptake while influencing carbonate and carbon flux dynamics.
391 Although the restored seagrass meadow in our study functions as a net calcifying system, TA fluxes
392 between SG and BS showed no significant difference.

393 **4.3 Net carbon uptake of seagrass restoration**

394 In order to estimate the net carbon uptake potential of seagrass restoration, we applied the photosynthesis-
395 quotient of 1 to calculate CO₂ uptake from organic carbon metabolism (Gattuso et al., 1998; Ward et al.,
396 2022). In terms of carbonate dynamics, each mole of CaCO₃ formed releasing 0.6 mmol of CO₂ into the
397 seawater was used (Frankignoulle et al., 1994). The calculated results show that total carbon uptake from
398 NEM was 208 mmol CO₂ m² d⁻¹ in SG and 20 mmol CO₂ m² d⁻¹ BS. For NEC, the carbon release in SG
399 was 6.52 CO₂ m² d⁻¹, while for BS, an additional CO₂ uptake was -1 mmol CO₂ m² d⁻¹. Consequently, the
400 net carbon uptake is 202 and 21 mmol CO₂ m² d⁻¹ for SG and BS, respectively. Our results demonstrate
401 that the primary productivity of restored seagrass through photosynthesis exceeds the rates of calcification
402 by 31-fold, suggesting that restored seagrass can act as a net carbon sink. However, further assessments
403 are necessary to capture temporal variations, as our current measurements are based on daily observations
404 and one season only.

405



406 **4.4 Limitations of ex situ benthic incubation and future research**

407 We tested the ex situ benthic core incubation approach for restored seagrass meadows, drawing from the
408 existing utilities in some coastal areas and freshwater ecosystems for sulfate and nutrient fluxes (Eyre, et
409 al., 2005, Chen et al., 2019). Overall, the ex situ benthic incubation method provides a significant
410 advantage by measuring both organic and inorganic carbon dynamics simultaneously, addressing a
411 critical gap in previous methods that often overlook carbonate dynamics (Johanssen, 2023). This approach
412 is also useful for assessing seagrass metabolism in subtidal meadows, where collecting data is challenging
413 due to high labor costs and weather conditions. Moreover, some in situ autonomous methods are often
414 expensive and constrained operational periods of only a few weeks due to challenges like sensor error
415 and biofouling (Yates and Halley, 2003; Takeshita et al., 2016). While this approach provides several
416 advantages, one notable limitation is its applicability. Currently, the design is primarily suited for small
417 seagrass, like *H. ovalis*, *H. uninervis*, and *Z. japonica*. It may not be adequate for larger species, like
418 *Enhalus acoroides* and large *Thalassia hemprichii*, due to differences in size and growth characteristics.
419 Moreover, we suggest validating the ex situ results with in situ data to ensure comparability with natural
420 conditions. Future research should integrate ex situ results with in situ data with different geographic and
421 environmental settings to enhance the generalizability of the findings. This will provide a more accurate
422 assessment of seagrass ecosystems' role in global carbon cycling and inform more effective coastal
423 management and conservation practices.

424 **5 Conclusion**

425 This study investigates the organic carbon metabolism and carbonate dynamics of replanted SG compared
426 to BS using the ex situ core incubation method. The results show that SG has higher GPP and NEM, while
427 exhibiting similar NEC, making it a stronger carbon sink than BS. The findings highlight the role of
428 seagrass restoration in enhancing carbon removal and contribute to a growing body of literature that
429 highlights the ecological value of restored seagrass meadows. This study represents the first simultaneous
430 quantitative estimate of the effect of both organic carbon metabolism and carbonate dynamics on carbon
431 sequestration of restored seagrass in Southeast Asia, providing valuable insights into the region's carbon



432 dynamics. We emphasize the need for long-term research on metabolic rates and carbonate dynamics to
433 account for temporal variations and to fully understand the implications of these processes in carbon
434 sequestration. This will also help optimize restoration strategies aimed at maximizing carbon sink
435 potential and mitigating ocean acidification. Furthermore, ex situ benthic incubation proves to be a
436 valuable tool for assessing carbon fluxes in seagrass meadows, particularly those dominated by pioneering
437 species, although further in situ assessments are necessary for comprehensive validation.

438 **Author contribution**

439 Wen-Chen Chou (WCC) and Jian-Jhih Chen (JJC) conceptualized the research and spearheaded the
440 implementation. JJC, Mariche B. Natividad (MBN), and Hsin-Yu Chou facilitated sample collection and
441 analysis. MBN and JJC performed the data analysis, drafted the manuscript, and its revision. WCC and
442 Lan-Feng Fan reviewed and revised the manuscript. All authors were involved in the finalization of the
443 manuscript.

444 **Competing interest**

445 The authors declare that they have no conflict of interest.

446 **Data availability**

447 The datasets in this study will be deposited in DRYAD Data Repository.

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454 **References**

- 455 Alongi, D.M., Trott, L.A., Undu, M.C., and Tirendi, F.: Benthic microbial metabolism in seagrass
456 meadows along a carbonate gradient in Sulawesi, Indonesia. *Aquat. Microb. Ecol.*, 51:141–152.
457 doi:10.3354/ame01191, 2008.
- 458 Apostolaki, E. T., Holmer, M., Marbà, N., and Karakassis, I.: Metabolic Imbalance in Coastal Vegetated
459 (*Posidonia oceanica*) and Unvegetated Benthic Ecosystems. *Ecosystems*, 13(3), 459–471.
460 <https://doi.org/10.1007/s10021-010-9330-9>, 2010.
- 461 Barrón, C., Duarte, C. M., Frankignoulle, M., and Borges, A. V.: Organic carbon metabolism and
462 carbonate dynamics in a Mediterranean seagrass (*Posidonia oceanica*), meadow. *Estuar. Coasts.*, 29
463 (3), 417–426. <https://doi.org/10.1007/bf02784990>, 2006.
- 464 Beavington-penney, S. J., Wright, V. P., & Racey, A.: Sediment production and dispersal on foraminifera-
465 dominated early Tertiary ramps: The Eocene El Garia Formation, Tunisia. *Sedimentology*, 52(3),
466 537–569. <https://doi.org/10.1111/j.1365-3091.2005.00709.x>, 2005.
- 467 Berg, P., Huettel, M., Glud, R. N., Reimers, C. E., and Attard, K. M. Aquatic Eddy Covariance: The
468 method and its contributions to defining oxygen and carbon fluxes in marine environments. *Ann. Rev.*
469 *Mar. Sci.*, 14(1), 431–455. <https://doi.org/10.1146/annurev-marine-042121-012329>, 2022.
- 470 Burdige, D. J. and Zimmerman, R.C.: Impact of sea grass density on carbonate dissolution in Bahamian
471 sediments. *Limnol. Oceanogr.*, 47(6), 1751–1763. doi: 10.4319/lo.2002.47.6.1751, 2002.
- 472 Burdige, D. J., Zimmerman, R. C., and Hu, X.: Rates of carbonate dissolution in permeable sediments
473 estimated from porewater profiles: the role of sea grasses. *Limnol. Oceanogr.*, 53: 549–565.
474 doi:10.2307/40006440, 2008.
- 475 Campbell, S., Miller, C., Steven, A., and Stephens, A.: Photosynthetic responses of two temperate
476 seagrasses across a water quality gradient using chlorophyll fluorescence. *J. Exp. Mar. Biol. Ecol.*,
477 291(1), 57–78. [https://doi.org/10.1016/s0022-0981\(03\)00090-x](https://doi.org/10.1016/s0022-0981(03)00090-x), 2003.
- 478 Chen, J., Wells, N., Erler, D., and Eyre, B.: Importance of habitat diversity to changes in benthic
479 metabolism over land-use gradients: evidence from three subtropical estuaries. *Mar. Ecol. Prog. Ser.*,
480 631, 31–47. <https://doi.org/10.3354/meps13147>, 2019.



- 481 Chou, W., Fan, L., Hung, C., Shih, Y., Huang, W., Lui, H., and Chen, T.: Dynamics of O₂ and pCO₂ in a
482 Southeast Asia seagrass meadow: Metabolic rates and carbon sink capacity. *Front. Mar. Sci.*, 10,
483 <https://doi.org/10.3389/fmars.2023.1076991>, 2023.
- 484 Chou, W., Fan, L., Yang, C., Chen, Y., Hung, C., Huang, W., Shih, Y., Soong, K., Tseng, H., Gong, G.,
485 Chen, H., and Su, C.: A unique DIEL pattern in carbonate chemistry in the seagrass meadows of
486 Dongsha Island: the enhancement of metabolic carbonate dissolution in a semienclosed lagoon. *Front.*
487 *Mar. Sci.*, 8. <https://doi.org/10.3389/fmars.2021.717685>, 2021.
- 488 Chou, W.-C., Chu, H.-C., Chen, Y.-H., Syu, R.-W., Hung, C.-C., and Soong, K.: Short-term variability of
489 carbon chemistry in two contrasting seagrass meadows at Dongsha island: implications for pH
490 buffering and CO₂ sequestration. *Estuar. Coast. Shelf Sci.*, 210, 36–44. doi:
491 10.1016/j.ecss.2018.06.006, 2018.
- 492 Clayton, T. D. and Byrne, R. H.: Spectrophotometric seawater pH measurements: total hydrogen ion
493 concentration scale calibration of m-cresol purple and at-sea results. *Deep-Sea Res. I: Oceanogr. Res.*
494 *Pap.*, 40(10), 2115–2129. [https://doi.org/10.1016/0967-0637\(93\)90048-8](https://doi.org/10.1016/0967-0637(93)90048-8), 1993.
- 495 De Beer, D., and A. W. D. Larkum.: Photosynthesis and calcification in the calcifying algae *Halimeda*
496 *discoidea* studied with microsensors. *Plant Cell Environ.* 24: 1209– 1217. doi:10.1046/j.1365-
497 3040.2001.00772.x, 2001.
- 498 Dickson, A.G. and Millero, F.J.: A Comparison of the Equilibrium Constants for the Dissociation of
499 Carbonic Acid in Seawater Media. *Deep-Sea Res. I: Oceanogr. Res. Pap.*, 34, 1733-1743.
500 [http://dx.doi.org/10.1016/0198-0149\(87\)90021-5](http://dx.doi.org/10.1016/0198-0149(87)90021-5), 1987.
- 501 Dickson, A.G., Sabine, C.L. and Christian, J.R. (Eds.): Guide to best practices for ocean CO₂
502 measurements. PICES Special Publication 3, 191 pp., 2007.
- 503 Duarte, C. M. and Krause-Jensen, D.: Export from Seagrass Meadows Contributes to Marine Carbon
504 Sequestration, *Frontiers in Marine Science*, 4, 13, <https://doi.org/10.3389/fmars.2017.00013>, 2017.
- 505 Duarte, C. M., Sintes, T., and Marbà, N.: Assessing the CO₂ capture potential of seagrass restoration
506 projects. *J. Appl. Ecol.*, 50(6), 1341–1349. <https://doi.org/10.1111/1365-2664.12155>, 2013.



- 507 Duarte, C. M., Marbà, N., Gacia, E., Fourqurean, J. W., Beggins, J., Barrón, C., and Apostolaki, E. T.:
508 Seagrass community metabolism: Assessing the carbon sink capacity of seagrass meadows. *Global*
509 *Biogeochem. Cy.*, 24(4). <https://doi.org/10.1029/2010gb003793>, 2010.
- 510 Duarte, C. M., Middelburg, J. J., and Caraco, N.: Major role of marine vegetation on the oceanic carbon
511 cycle. *Biogeosciences*, 2(1), 1–8. <https://doi.org/10.5194/bg-2-1-2005>, 2005.
- 512 Egea, L., Infantes, E., & Jiménez-Ramos, R. (2023). Loss of POC and DOC on seagrass sediments by
513 hydrodynamics. *Sci. Total Environ.*, 901, 165976. <https://doi.org/10.1016/j.scitotenv.2023.165976>
- 514 Enríquez, S. and Schubert, N.: Direct contribution of the seagrass *Thalassia testudinum* to lime mud
515 production. *Nat. Commun.* **5**, 3835. <https://doi.org/10.1038/ncomms4835>, 2014.
- 516 Eyre, B. D. and Ferguson, A. J. P.: Benthic metabolism and nitrogen cycling in a subtropical east
517 Australian estuary (Brunswick): Temporal variability and controlling factors. *Limnol. Oceanogr.*,
518 50(1), 81–96. <https://doi.org/10.4319/lo.2005.50.1.0081>, 2005.
- 519 Eyre, B. D., A. J. P. Ferguson, A. Webb, D. Maher, and J. M. Oakes.: Denitrification, N-fixation and
520 nitrogen and phosphorus fluxes in different benthic habitats and their contribution to the nitrogen and
521 phosphorus budgets of a shallow oligotrophic sub-tropical coastal system (southern Moreton Bay,
522 Australia). *Biogeochemistry* 102: 111–133. doi:10.1007/s10533-010-9425-6, 2011.
- 523 Eyre, B. D., Andersson, A. J., & Cyronak, T.: Benthic coral reef calcium carbonate dissolution in an
524 acidifying ocean. *Nat. Clim. Change*, 4(11), 969–976. <https://doi.org/10.1038/nclimate2380>, 2014.
- 525 Fan, L.-F., Kang, E.-C., Natividad, M. B., Hung, C.-C., Shih, Y.-Y., Huang, W.-J., & Chou, W.-C.: The
526 role of benthic TA and DIC fluxes on carbon sequestration in seagrass meadows of Dongsha Island.
527 *J. Mar. Sci. Eng.*, 12, 2061. <https://doi.org/10.3390/jmse12112061>, 2024.
- 528 Ferguson, A., Eyre, B., and Gay, J.: Benthic nutrient fluxes in euphotic sediments along shallow sub-
529 tropical estuaries, northern New South Wales, Australia. *Aquat. Microb. Ecol.*, 37, 219–235.
530 <https://doi.org/10.3354/ame037219>, 2004.
- 531 Frankignoulle, M., Canon, C., and Gattuso, J.-P.: Marine calcification as a source of carbon dioxide:
532 Positive feedback of increasing atmospheric CO₂, *Limnol. Oceanogr.*, 39(2), 458–462, 1994.
- 533 Fourqurean, J. W., Duarte, C. M., Kennedy, H., Marbà, N., Holmer, M., Mateo, M. A., Apostolaki, E. T.,
534 Kendrick, G. A., Krause-Jensen, D., McGlathery, K. J., & Serrano, O.: Seagrass ecosystems as a



- 535 globally significant carbon stock. *Nat. Geosci.*, 5(7), 505–509. <https://doi.org/10.1038/ngeo1477>,
536 2012.
- 537 Ganguly, D., Singh, G., Ramachandran, P., Selvam, A.P., Banerjee, K., and Ramachandran, R.: Seagrass
538 metabolism and carbon dynamics in a tropical coastal embayment. *Ambio*. Oct;46(6):667-679. doi:
539 10.1007/s13280-017-0916-8. Epub 2017 Mar 31. PMID: 28364264; PMCID: PMC5595744., 2017
- 540 Gazeau, F., Duarte, C. M., Gattuso, J., Barrón, C., Navarro, N., Ruiz, S., Prairie, Y. T., Calleja, M., Delille,
541 B., Frankignoulle, M., and Borges, A. V.: Whole-system metabolism and CO₂ fluxes in a
542 Mediterranean Bay dominated by seagrass beds (Palma Bay, NW Mediterranean). *Biogeosciences*,
543 2(1), 43–60. <https://doi.org/10.5194/bg-2-43-2005>, 2005
- 544 Greiner, J. T., McGlathery, K. J., Gunnell, J., and McKee, B. A.: Seagrass restoration enhances “Blue
545 carbon” sequestration in coastal waters. *PLoS ONE*, 8(8), e72469.
546 <https://doi.org/10.1371/journal.pone.0072469>, 2013.
- 547 Griscom, B. W., Adams, J., Ellis, P. W., Houghton, R. A., Lomax, G., Miteva, D. A., Schlesinger, W. H.,
548 Shoch, D., Siikamäki, J. V., Smith, P., Woodbury, P., Zganjar, C., Blackman, A., Campari, J., Conant,
549 R. T., Delgado, C., Elias, P., Gopalakrishna, T., Hamsik, M. R., . . . Fargione, J.: Natural climate
550 solutions. *P. Natl. A Sci*, 114(44), 11645–11650. <https://doi.org/10.1073/pnas.1710465114>, 2017.
- 551 Hendriks, I. E., Olsen, Y. S., Ramajo, L., Basso, L., Steckbauer, A., Moore, T. S., Howard, J. and Duarte,
552 C. M. 2014. Photosynthetic activity buffers ocean acidification in seagrass meadows, *Biogeosciences*,
553 11(2), 333.
- 554 Howard, J. L., Creed, J. C., Aguiar, M. V. P., and Fourqurean, J. W.: CO₂ released by carbonate sediment
555 production in some coastal areas may offset the benefits of seagrass “Blue Carbon” storage. *Limnol.*
556 *Oceanogr.*, 63(1), 160–172. <https://doi.org/10.1002/lno.10621>, 2017.
- 557 Johannessen, S. C.: How to quantify blue carbon sequestration rates in seagrass meadow sediment:
558 geochemical method and troubleshooting. *Carbon Footprints*, 2(4).
559 <https://doi.org/10.20517/cf.2023.37>, 2023.
- 560 Juska, I. and Berg, P.: Variation in seagrass meadow respiration measured by aquatic eddy covariance.
561 *Limnol. Oceanogr. Lett.* 7(5), 410–418. <https://doi.org/10.1002/lol2.10276>, 2022.



- 562 Kennedy, H., Beggins, J., Duarte, C. M., Fourqurean, J. W., Holmer, M., Marbà, N., and Middelburg, J.
563 J.: Seagrass sediments as a global carbon sink: Isotopic constraints. *Global Biogeochem. Cy.*, 24(4).
564 <https://doi.org/10.1029/2010gb003848>, 2010.
- 565 Kindeberg, T., Attard, K. M., Hüller, J., Müller, J., Quintana, C. O., and Infantes, E.: Structural complexity
566 and benthic metabolism: resolving the links between carbon cycling and biodiversity in restored
567 seagrass meadows. *Biogeosciences*, 21(7), 1685–1705. <https://doi.org/10.5194/bg-21-1685-2024>,
568 2024.
- 569 Kindeberg, T., Bates, N. R., Courtney, T. A., Cyronak, T., Griffin, A., Mackenzie, F. T., et al.: Porewater
570 carbonate chemistry dynamics in a temperate and a subtropical seagrass system. *Aquat. Geochem.* 26,
571 375–399. doi: 10.1007/s10498-020-09378-8, 2020.
- 572 Kleypas, J. A., Buddemeier, R. W., and Gattuso, J. P.: The future of coral reefs in an age of global
573 change. *Int. J. Earth Sci.*, 90, 426–437, <https://doi.org/10.1007/s005310000125>, 2001.
- 574 Koch, M., Bowes, G., Ross, C., and Zhang, X.: Climate change and ocean acidification effects on
575 seagrasses and marine macroalgae. *Glob. Change Biol.*, 19(1), 103–132.
576 <https://doi.org/10.1111/j.1365-2486.2012.02791.x>, 2012
- 577 Kraemer, G.P. and Alberte, R.S.: Age-related patterns of metabolism and biomass in subterranean tissues
578 of *Zostera marina* L. (eelgrass). *Mar. Ecol. Prog. Ser.*, 95: 193–203, 1993.
- 579 Land, L. S. Carbonate mud; production by epibiont growth on *Thalassia testudinum*. *J. Sediment.*
580 *Res.*, 40, 1361–1363. <https://doi.org/10.1306/74D721B7-2B21-11D7-8648000102C1865D>, 1970.
- 581 Lee, K., Park, S. R., and Kim, Y. K.: Effects of irradiance, temperature, and nutrients on growth dynamics
582 of seagrasses: A review. *J. Exp. Mar. Biol. Ecol.*, 350(1–2), 144–175.
583 <https://doi.org/10.1016/j.jembe.2007.06.016>, 2007.
- 584 Macreadie, P. I., Serrano, O., Maher, D. T., Duarte, C. M., and Beardall, J.: Addressing calcium carbonate
585 cycling in blue carbon accounting. *Limnol. Oceanogr. Lett.*, 2(6), 195–201.
586 <https://doi.org/10.1002/lol2.10052>, 2017.
- 587 Maher, D., & Eyre, B.: Benthic carbon metabolism in southeast Australian estuaries: habitat importance,
588 driving forces, and application of artificial neural network models. *Mar. Ecol. Prog. Ser.*, 439, 97–
589 115. <https://doi.org/10.3354/meps09336>, 2011.



- 590 Mallon, J., Cyronak, T., Hall, E. R., Banaszak, A. T., Exton, D. A., and Bass, A. M.: Light-driven dynamics
591 between calcification and production in functionally diverse coral reef calcifiers. *Limnol. and*
592 *Oceanogr.*, 67(2), 434–449. <https://doi.org/10.1002/lno.12002>, 2002.
- 593 Marbà, N., Arias-Ortiz, A., Masqué, P., Kendrick, G. A., Mazarrasa, I., Bastyan, G. R., Garcia-Orellana,
594 J., and Duarte, C. M.: Impact of seagrass loss and subsequent revegetation on carbon sequestration
595 and stocks. *J. Ecol.*, 103(2), 296–302. <https://doi.org/10.1111/1365-2745.12370>, 2015.
- 596 Mazarrasa, I., Marbà, N., Lovelock, C. E., Serrano, O., Lavery, P. S., Fourqurean, J. W., Kennedy, H.,
597 Mateo, M. Á., Krause-Jensen, D., Steven, A. D. L., and Duarte, C. M.: Seagrass meadows as a globally
598 significant carbonate reservoir. *Biogeosciences*, 12(16), 4993–5003. [https://doi.org/10.5194/bg-12-](https://doi.org/10.5194/bg-12-4993-2015)
599 [4993-2015](https://doi.org/10.5194/bg-12-4993-2015), 2015.
- 600 Mcleod, E., Chmura, G. L., Bouillon, S., Salm, R., Björk, M., Duarte, C. M., Lovelock, C. E., Schlesinger,
601 W. H., and Silliman, B. R.: A blueprint for blue carbon: toward an improved understanding of the role
602 of vegetated coastal habitats in sequestering CO₂. *Front. Ecol. Environ.*, 9(10), 552–560.
603 <https://doi.org/10.1890/110004>, 2011
- 604 Mehrbach, C., Culberson, C. H., Hawley, J. E., and Pytkowicz, R. M.: Measurement of the apparent
605 dissociation constants of carbonic acid in seawater at atmospheric pressure¹. *Limnol. Oceanogr.*,
606 18(6), 897–907. <https://doi.org/10.4319/lo.1973.18.6.0897>, 1973.
- 607 Odum, H. T.: Primary production in flowing waters, *Limnol. Oceanogr.*, 1(2), 102–117, 1956.
- 608 Olivé, I., Silva, J., Costa, M. M., and Santos, R.: Estimating seagrass community metabolism using benthic
609 chambers: The effect of incubation time. *Estuar. Coasts*, 39(1), 138–144.
610 <https://doi.org/10.1007/s12237-015-9973-z>, 2015.
- 611 Oreska, M. P. J., Wilkinson, G. M., McGlathery, K. J., Bost, M., and McKee, B. A.: Non-seagrass carbon
612 contributions to seagrass sediment blue carbon. *Limnol. Oceanogr.*, 63(S1).
613 <https://doi.org/10.1002/lno.10718>, 2017.
- 614 Ortegón-Aznar, I., Chuc-Contreras, A., & Collado-Vides, L. Calcareous green algae standing stock in a
615 tropical sedimentary coast. *J. Appl. Phycol.*, 29, 2685–2693. [https://doi.org/10.1007/s10811-017-](https://doi.org/10.1007/s10811-017-1057-y)
616 [1057-y](https://doi.org/10.1007/s10811-017-1057-y), 2017.



- 617 Orth, R. J., Carruthers, T. J. B., Dennison, W. C., Duarte, C. M., Fourqurean, J. W., Heck, K. L., Hughes,
618 A. R., Kendrick, G. A., Kenworthy, W. J., Olyarnik, S., Short, F. T., Waycott, M., and Williams, S.
619 L.: A global crisis for seagrass ecosystems. OUP Academic. [https://doi.org/10.1641/0006-](https://doi.org/10.1641/0006-3568(2006)56)
620 [3568\(2006\)56](https://doi.org/10.1641/0006-3568(2006)56), 2006.
- 621 Patriquin, D.G: The origin of nitrogen and phosphorus for growth of the marine angiosperm *Thalassia*
622 *testudinum*. Mar. Biol. 15, 35–46. <https://doi.org/10.1007/BF00347435>, 1972.
- 623 Pelletier, G., Lewis, E., and Wallace, D.: *CO2SYS. XLS: A Calculator for the CO₂ System in Seawater for*
624 *Microsoft Excel/VBA. Version 16*. Washington, DC: Washington State Department of Ecology, 2011
- 625 Perry, C., & Beavington-Penney, S.: Epiphytic calcium carbonate production and facies development
626 within sub-tropical seagrass beds, Inhaca Island, Mozambique. *Sedimentary Geol.*, 174(3–4), 161–
627 176. <https://doi.org/10.1016/j.sedgeo.2004.12.003>, 2005.
- 628 Ralph, P., Durako, M., Enríquez, S., Collier, C., and Doblin, M.: Impact of light limitation on seagrasses.
629 *J. Exp. Mar. Biol. Ecol.*, 350(1–2), 176–193. <https://doi.org/10.1016/j.jembe.2007.06.017>, 2007.
- 630 Rheuban, J. E., Berg, P. and McGlathery, K. J.: Ecosystem metabolism along a colonization gradient of
631 eelgrass (*Zostera marina*) measured by eddy correlation, *Limnol. Oceanogr.*, 59(4), 1376–1387, 2014.
- 632 Roth, F., Wild, C., Carvalho, S., Rådecker, N., Voolstra, C. R., Kürten, B., Anlauf, H., El-Khaled, Y. C.,
633 Carolan, R., and Jones, B. H.: An in situ approach for measuring biogeochemical fluxes in structurally
634 complex benthic communities. *Methods Ecol. Evol.*, 10(5), 712–725. [https://doi.org/10.1111/2041-](https://doi.org/10.1111/2041-210x.13151)
635 [210x.13151](https://doi.org/10.1111/2041-210x.13151), 2019.
- 636 Saderne, V., Geraldi, N.R., Macreadie, P.I. *et al.*: Role of carbonate burial in Blue Carbon budgets. *Nat*
637 *Commun.*, 10, 1106, <https://doi.org/10.1038/s41467-019-08842-6>, 2019.
- 638 Schneider, K., Levy, O., Dubinsky, Z., & Erez, J. (2009). In situ diel cycles of photosynthesis and
639 calcification in hermatypic corals. *Limnol. Oceanogr.*, 54(6), 1995–2002.
640 <https://doi.org/10.4319/lo.2009.54.6.1995>
- 641 Takeshita, Y., W. McGillis, E.M. Briggs, A.L. Carter, E.M. Donham, T.R. Martz, N.N. Price, and J.E.
642 Smith.: Assessment of net community production and calcification of a coral reef using a boundary
643 layer approach. *J. Geophys. Res-Oceans*. 121: 5655–5671, 2016.



- 644 Van Dam, B. R., Lopes, C., Osburn, C. L., and Fourqurean, J. W.: Net heterotrophy and carbonate
645 dissolution in two subtropical seagrass meadows. *Biogeosciences*, 16(22), 4411–4428.
646 <https://doi.org/10.5194/bg-16-4411-2019>, 2019.
- 647 Van Dam, B. R., Zeller, M. A., Lopes, C., Smyth, A. R., Böttcher, M. E., Osburn, C. L., Zimmerman, T.,
648 Präfrock, D., Fourqurean, J. W., and Thomas, H.: Calcification-driven CO₂ emissions exceed “Blue
649 Carbon” sequestration in a carbonate seagrass meadow. *Sci. Adv.*, 7(51).
650 <https://doi.org/10.1126/sciadv.abj1372>, 2021.
- 651 Walker, D. & Woelkerling, W. Quantitative study of sediment contribution by epiphytic coralline red algae
652 in seagrass meadows in Shark Bay, Western Australia. *Mar. Ecol. Prog. Ser.* **43**, 71–77 (1988).
- 653 Ward, M., Kindinger, T. L., Hirsh, H. K., Hill, T. M., Jellison, B. M., Lummis, S., Rivest, E. B.,
654 Waldbusser, G. G., Gaylord, B., and Kroeker, K. J.: Reviews and syntheses: Spatial and temporal
655 patterns in seagrass metabolic fluxes, *Biogeosciences*, 19, 689–699, [https://doi.org/10.5194/bg-19-](https://doi.org/10.5194/bg-19-689-2022)
656 [689-2022](https://doi.org/10.5194/bg-19-689-2022), 2022.
- 657 Waycott, M., Duarte, C. M., Carruthers, T. J. B., Orth, R. J., Dennison, W. C., Olyarnik, S., Calladine, A.,
658 Fourqurean, J. W., Heck, K. L., Hughes, A. R., Kendrick, G. A., Kenworthy, W. J., Short, F. T., and
659 Williams, S. L.: Accelerating loss of seagrasses across the globe threatens coastal ecosystems. *P. Natl.*
660 *A. Sci.*, 106(30), 12377–12381. <https://doi.org/10.1073/pnas.0905620106>, 2009.
- 661 Yang, Y.-P., Fong, S.-C., and Liu H.-Yih.: Taxonomy and distribution of seagrasses in Taiwan. *Taiwania*,
662 47(1):54–61, 2002.
- 663 Yates, K.K. and Halley, R.B.: Measuring coral reef community metabolism using new benthic chamber
664 technology. *Coral Reefs*, 22, 247–255. <https://doi.org/10.1007/s00338-003-0314-5>, 2003.
- 665 Yates, K.K. and Halley, R.B.: Diurnal variation in rates of calcification and carbonates sediment
666 dissolution in Florida Bay. *Estuar. Coasts*, 29:24–39. 2006.