

1 **Estimation of Metabolic Dynamics of Restored Seagrass Meadows in a Southeast Asia Islet:**
2 **Insights from Ex Situ Benthic Incubation**

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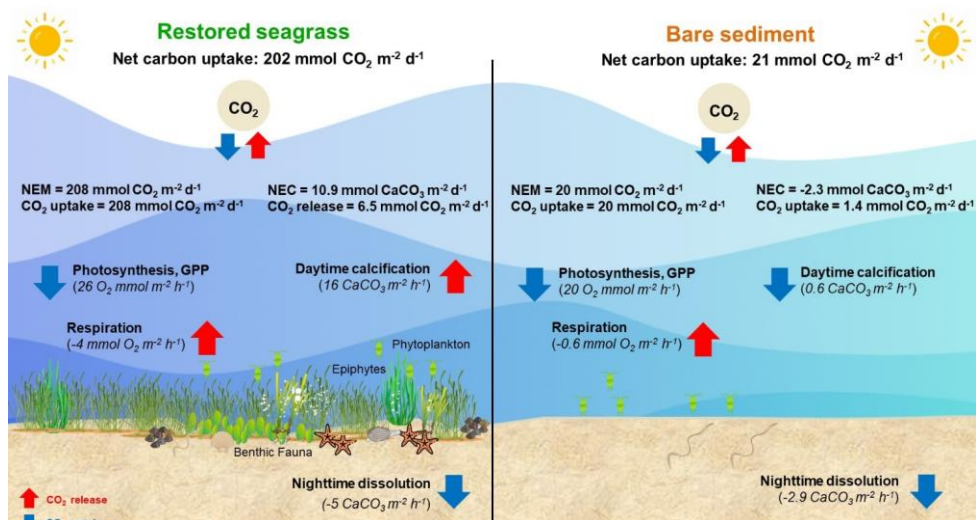
17 **Abstract.** Seagrass meadows are vital carbon sinks, but their function is threatened by rapid decline,
18 driving restoration efforts to enhance coastal recovery and carbon removal. The capacity of these restored
19 seagrass as carbon sources or sinks depends largely on organic carbon metabolism and carbonate
20 dynamics. In this study, we employed ex situ core incubation to investigate the metabolic rates of
21 replanted seagrasses (SG), including gross primary productivity (GPP), community respiration (R), net
22 ecosystem metabolism (NEM), and net ecosystem calcification (NEC) in SG and surrounding bare
23 sediments (BS). SG exhibited higher GPP (26.0 ± 4.3 mmol O₂ m⁻² h⁻¹ vs 0.7 ± 1.3 mmol O₂ m⁻² h⁻¹) and NEM (208.2 ± 6.3 mmol O₂ m⁻² d⁻¹ vs 20.1 ± 2.9 mmol O₂ m⁻² d⁻¹) than BS, indicating their
24 potential as carbon sinks by shifting benthic metabolism towards a more autotrophic state. In contrast,
25 SG exhibited net calcification with positive NEC values (10.9 ± 15.7 mmol CaCO₃ m⁻² d⁻¹), driven by
26 higher daytime carbonate production than nighttime dissolution, while BS showed net dissolution with
27 negative NEC values (-2.3 ± 18.8 mmol CaCO₃ m⁻² d⁻¹).~~In contrast,~~ Despite this, high variability in
28 carbonate fluxes led to no significant difference between SG and BS ($p > 0.05$).~~SG showed higher daytime~~

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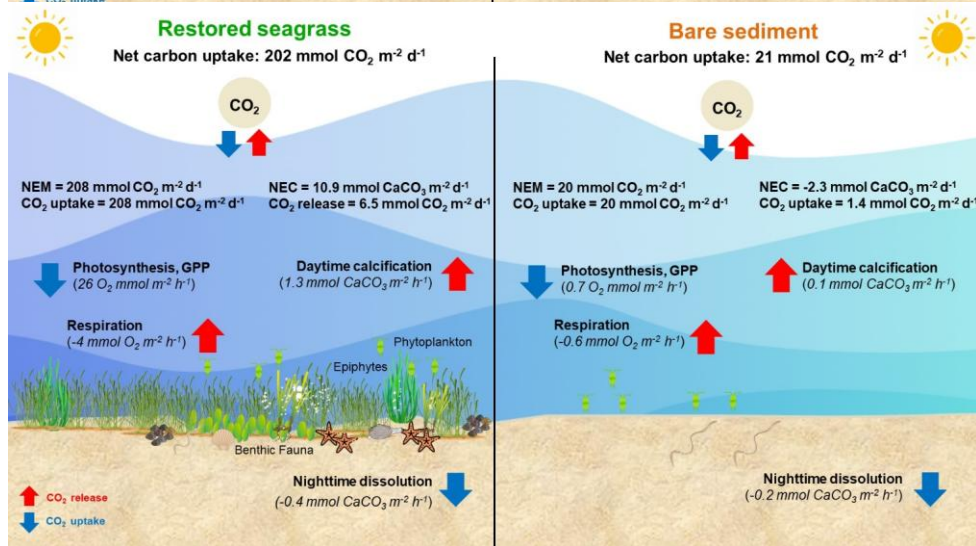
30 ~~carbonate production and nighttime carbonate dissolution, which could offset each other, resulting in no~~
31 ~~significant difference in NEC between SG and BS.~~ In summary, our results found that the SG exhibited
32 significantly higher NEM compared to BS (*p*<0.01), while no significant difference was found for NEC.
33 Consequently, the net effect on the carbon uptake capacity of the restored seagrass is likely increased,
34 primarily due to the higher NEM. Our findings highlight the ecological significance of seagrass
35 restoration in mitigating climate change through carbon removal. ~~The Ex~~ *Ex* situ core incubation method
36 allows for the simultaneous measurement of organic and inorganic carbon metabolism. While ex situ core
37 incubation enhances feasibility, in situ assessments are still necessary to validate the results and ensure a
38 comprehensive understanding of seagrass ecosystem dynamics.

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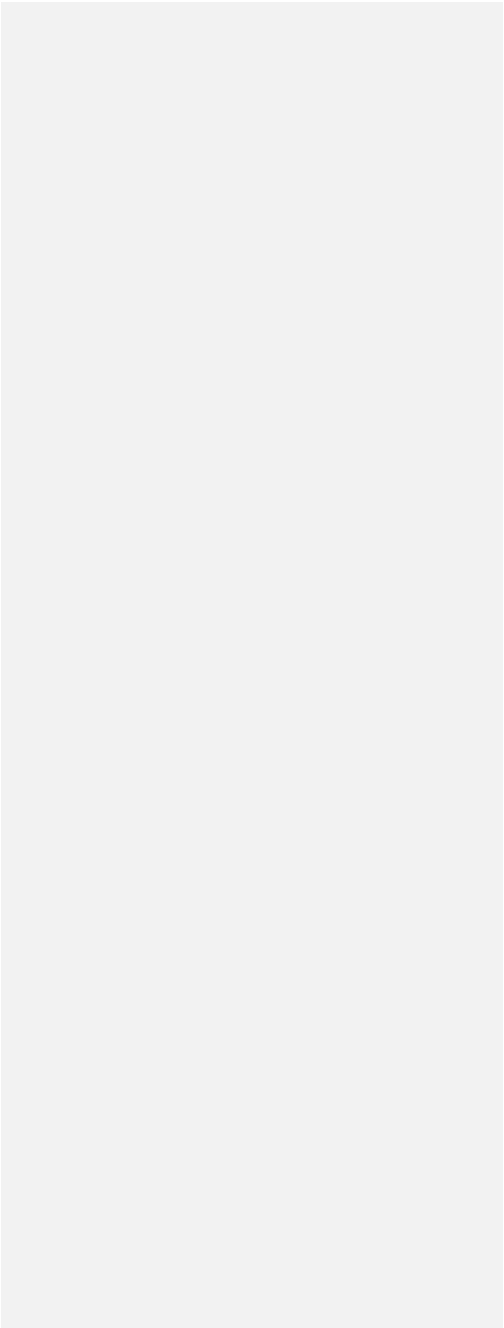


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43 **Graphical abstract: Illustration of carbon uptake from organic carbon metabolism (GPP-gross**
44 **primary productivity, R-respiration, NEM-net ecosystem metabolism) and carbonate dynamics**
45 **(daytime calcification, nighttime dissolution, and NEC-net ecosystem calcification) in restored**
46 **seagrass and bare sediment. Net Ecosystem Metabolism (NEM).**



47 **1 Introduction**

48 Seagrass meadows, comprising over 72 species, occupy just 0.1% of the ocean's surface, yet they are
49 highly productive and ecologically significant ecosystems in the marine environments (Fourqurean et al.,
50 2012; Short et al., 2011). These meadows play essential roles in nutrient and carbon cycling and serve as
51 key habitats for many marine species (Duarte et al., 2010; Fourqurean et al., 2012). Due to their relatively
52 complex structure, seagrass meadows capture and retain organic carbon (C_{org}) in the sediment, making
53 them one of the major carbon reservoirs globally (Duarte et al., 2005; Mcleod et al., 2011). Previous
54 estimates suggest that seagrasses account for approximately 15% of the total global carbon sequestered
55 in benthic sediments (Duarte et al., 2013), with burial rates 35 times that of tropical rainforests (Mcleod
56 et al., 2011).

57
58 In spite of their ecological significance, seagrass meadows have experienced a global decline, driven
59 primarily by human-induced activities such as coastal development, eutrophication, and deteriorating
60 water quality (Orth et al., 2006; Waycott et al., 2009). Since 1980, the global coverage of seagrass has
61 decreased by 110 km² annually, with the rate of decline increasing (Waycott et al., 2009). The loss is
62 frequently associated with increased water column turbidity and epiphytic shading, which reduce the light
63 for seagrass photosynthesis, leading to meadow degradation (Campbell et al., 2003; Orth et al., 2006).
64 Degradation also diminishes their capacity to modify local pH and influence the dynamics of dissolved
65 oxygen (DO) and dissolved inorganic carbon (DIC) (Hendricks et al., 2014). Moreover, the continued
66 loss of seagrass ecosystems raises concerns that vast amounts of previously sequestered carbon could be
67 released back in the atmosphere, converting seagrasses from carbon sinks to carbon sources and
68 intensifying global climate change (Macreadie et al., 2013). The ongoing decline could potentially release
69 up to 299 Tg of carbon annually, contributing roughly 10% of CO₂ emissions associated with
70 anthropogenic land-use changes (Fourqurean et al., 2012).

71
72 In response to these challenges, seagrass restoration has emerged as a critical strategy to mitigate
73 environmental degradation, enhance coastal resilience, and address global climate change (Juska and Berg

et al., 2022). Protecting and restoring seagrass meadows aligns with international goals like the Paris Agreement, as these ecosystems offer significant potential for long-term carbon storage and climate regulation (Fourqurean et al., 2012). However, despite growing restoration efforts, there remains limited understanding of their success, particularly regarding benthic metabolism and carbon dynamics (Kindeberg et al., 2024). While studies from temperate regions, such as the *Zostera marina* restoration in the Virginia Coast (Rheuban et al., 2014), have provided valuable insights, data from tropical regions — including Southeast Asia, a global hotspot for seagrass diversity — remain scarce (Duarte et al., 2010; Ward et al., 2022; Chou et al., 2023). It represents a critical gap in our knowledge of the impact of restoration efforts on carbon removal and ocean acidification mitigation.

Although there is increasing consensus on the potential of “Blue Carbon” storage in seagrass meadows as a climate change mitigation strategy, the biogeochemical cycling within these ecosystems is complex. Several processes, including ecosystem calcification, anaerobic metabolism, and bioturbation, can counteract net organic carbon (OC) sequestration (Van Dam et al., 2021). These processes regulate local DIC and total alkalinity (TA) budgets, adding complexity to accurately quantifying carbon sequestration (Kindeberg et al., 2024). Overlooking these processes can result in significant overestimates of local carbon sequestration rates and misinterpretations of the role seagrass meadows play in mitigating climate change, potentially leading to inaccurate assessments of their carbon sink capacity (Johansen et al., 2023; Chen et al., 2024; Fan et al., 2024).

Several methodologies were developed to quantify benthic metabolism, which is a crucial component of biogeochemical cycling, including photosynthesis-irradiance curve (Kraemer and Alberte, 1993), the open water O₂ mass balance approach (Odum, 1956; Chou et al., 2023), and aquatic eddy covariance (Berg et al., 2022; Juska and Berg, 2022). While these methods provide important data, they might overlook the complexities of bioturbation, remineralization, and carbonate dynamics (Olive et al., 2016; Ward et al., 2022; Juska and Berg, 2022). In this study, we aim to address these knowledge gaps by quantifying organic carbon metabolism (net ecosystem metabolism, NEM) and carbonate dynamics

101 (i.e., net ecosystem calcification, NEC) in restored seagrass meadows (SG) and adjacent bare sediment
102 (BS) habitats on a Southeast Asia islet, using an innovative ex situ benthic incubation.

103 **2 Materials and Methods**

104 **2.1 Study site**

105 The Penghu Islands, located in the southern part of Taiwan Strait (Fig. 1), host a range of seagrass species.
106 Notably, four species have been reported: *Halophila ovalis*, *Halodule pinifolia*, *Halodule uninervis*, and
107 *Zostera japonica* (Yang et al., 2002). The sampling location (23° 38' 18.38" N and 119° 33' 46.48" E) is
108 a restoration meadow dominated by *H. uninervis* and *H. ovalis*. This restoration site encompasses
109 approximately 3 hectares (~~as per Coral Allen~~ [Allen Coral Atlas, 2020](#)), with seagrass percent cover varying
110 from 20% to 90%. These seagrasses are subtidal, with water depths ranging from 1.7 meters to 4.4 meters.
111 The substrate in this area is composed of carbonate sand. The area supports a diverse community of
112 bivalves (e.g., *Pinna* sp.), gastropods, echinoderms, and various fish species, all of which were observed
113 during the sampling.

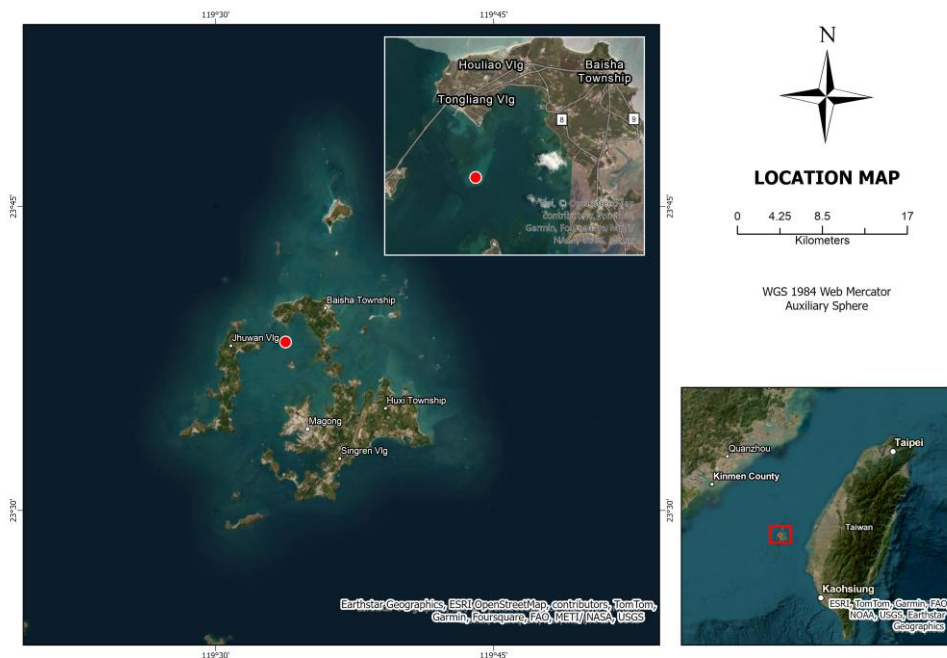
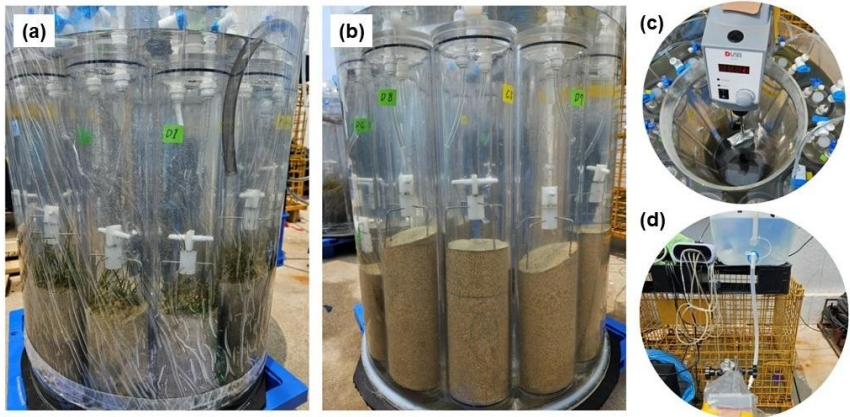


Figure 1: Location map of sampling stations in restored seagrass in Penghu Island, Taiwan (Map created in ArcGIS Pro. Source: Earthstar Geographics, ESRI OpenStreetMap, Contributors, TomTom, Garmin, Foursquare, FAO, METI/NASA, USGS, NOAA).

2.2 Ex situ core incubation system

The ex situ benthic core methodology used in this study was adapted from Chen et al. (2019) (Fig. 2). This approach has been widely employed in various studies to assess nutrient concentrations and benthic metabolism in coastal ecosystems and estuaries (Eyre & Ferguson, 2005; Maher & Eyre, 2011). Typically, the ex situ core incubation involves 150-L treatment tanks containing aerated water. Each tank can accommodate 10 plexiglass cores made of polycarbonate material, 10 cm in diameter and 50 cm in height. The tanks were equipped with magnetic stir bars driven by a centrally located rotating motor fitted with

125 a magnet. The core has a plexiglass lid which contains two ports, one for probe insertion (Eyre &
126 Ferguson, 2005). This method offers a feasible approach for quantifying seagrass metabolism, especially
127 in subtidal systems where in situ measurements are often logistically challenging. While ex situ conditions
128 may differ from natural underwater environments, we carefully designed our setup to closely replicate
129 field conditions, including natural light exposure and ambient temperature, to ensure ecological relevance.



130

131 **Figure 2: Ex situ benthic chamber setup for measuring metabolic rates and carbonate dynamics in**
132 **seagrass meadows and bare sediment. The chambers contain seagrass samples (a), while the**
133 **chambers contain bare sediment (b). Insets show close-ups of the central rotating motor with a**
134 **magnet setup for water circulation (c), and the setup for continuous seawater supply (d).**

135 **2.3 Sediment core collection and pre-incubation**

136 The incubation was conducted on April 12-13, 2024. Twenty intact sediment cores, comprising both
137 seagrass and bare sediment, were collected on-site using the plexiglass tubes. The cores were inserted
138 about 20 cm into the sediment, keeping approximately 1.9 liters of water. Each core was sealed with a
139 gas-tight plexiglass plate at the bottom. The samples were brought back to the incubation site within two
140 hours of collection and allowed to settle for 24 hours. Additionally, 150 liters of water were collected on-
141 site for continuous supply during the experiment.

142

143 At the incubation site, the cores were uncovered and placed in 150-liter tanks filled with aerated seawater.
144 They were kept at in situ temperature, exposed to natural sunlight, and continuously recirculated. The
145 stirring rate was controlled to prevent sediment resuspension (Ferguson et al., 2004). The cores underwent
146 a 24-hour pre-incubation period to promote stable sediment profiles. The seagrass composition within
147 the collected cores for ex situ core incubation was dominated by *H. uninervis* and *H. ovalis*. The shoot
148 count of *H. uninervis* ranged from 20 to 40 shoots per 0.008 m², while *H. ovalis* ranged from 2 to 20
149 shoots per 0.008 m².

150 **2.4 Sample collection and analysis**

151 Following pre-incubation, the cores were tightly closed using a plexiglass lid. Temperature, salinity, and
152 pH were determined using a YSI ProDSS Multiparameter water quality ~~checker~~sonde, while DO (mg l⁻¹)
153 was measured with a thermo DO probe. Both probes were calibrated with calibration standards.
154 Measurements were taken at midnight (24:00 h) with 2-hour intervals and ended at noon.
155 Photosynthetically active radiation (PAR) levels were measured using SQ-420X Smart Quantum Sensor
156 positioned atop the incubation tank.

157

158 After measurements, three 150 ml seawater samples were collected separately from the SG and BS cores
159 using a syringe for DIC and pH analysis. The water samples were processed with 60 µL HgCl₂ solution
160 to stop any biological activity. DIC analysis was performed using a non-dispersive infrared method with
161 a DIC analyzer (AS-C3, Apollo SciTech Inc.), following the approach of Dickson et al. (2007) and our
162 past studies (Chou et al., 2018; 2021; Fan et al., 2024). For each DIC run, we used certified reference
163 material (Batch no. 206) sourced from A. G. Dickson at Scripps Institution of Oceanography to check for
164 drift and systematic bias. pH values were measured spectrophotometrically in total scale at 25 °C
165 following Clayton and Byrne (1993). Data from DIC and pH, along with actual temperature and salinity,
166 were used to calculate the TA, partial pressure of CO₂ (pCO₂), and aragonite saturation state (Ω_{Ar}) using
167 the Excel macro CO2SYS version 2.1 (Pelletier et al. 2011). The dissociation constants for carbonic acid

168 applied in these calculations were obtained from Mehrbach et al. (1973) and subsequently refined by
169 Dickson and Millero (1987).

170 2.5 Benthic flux rate calculations

171 Areal rates of R, GPP, NPP, and NEM were calculated based on changes in DO concentrations, following
172 equation 1 (Eyre et al. 2011). Respiration rates were determined from concentration data collected during
173 the initial dark period (midnight to dawn) (eq. 2). NPP was calculated based on light O₂ flux
174 measurements from dawn to noon (eq. 3). We implemented a 6-hour dark incubation period to ensure
175 oxygen concentrations remained above 80% (Eyre et al., 2002) and a 6-hour light incubation period to
176 prevent oxygen from reaching supersaturated levels (Olivé et al., 2016). Hourly GPP rates were computed
177 as the difference between R and NPP rates (eq. 4). NEM was calculated using equation 5. Positive values
178 indicate autotrophic, while negative values represent heterotrophic.

$$179 \quad F = [(C_{t1} - C_{t0}) \times V/A]/T \quad (\text{eq. 1})$$

180 Where F = flux rate ($\mu\text{mol m}^{-2} \text{h}^{-1}$), C_{t0} and C_{t1} = concentration in the overlying water at the start and end
181 of the time period ($\mu\text{mol l}^{-1}$), respectively, V = volume of overlying water in the core (l), A = surface area
182 in the sediment core (m²), and T = incubation period (h).

$$183 \quad R = \text{dark O}_2 \text{ flux (negative)} \quad (\text{eq. 2})$$

$$184 \quad \text{NPP} = \text{light O}_2 \text{ flux (positive)} \quad (\text{eq. 3})$$

$$185 \quad \text{GPP} = \text{NPP (positive)} - R \text{ (negative)} \quad (\text{eq. 4})$$

$$186 \quad \text{NEM} = (\text{GPP} \times 12) - (R \times 24 \text{ h} \times -1) \quad (\text{eq. 5})$$

187

188 NEC rates ($\text{mmol CaCO}_3 \text{ m}^{-2} \text{h}^{-1}$) were estimated from the change of total alkalinity, assuming these
189 changes are only due to CaCO₃ precipitation and dissolution (eq. 6) (Roth et al., 2019; Van Dam et al.,
190 2019):

$$191 \quad \text{NEC} = -0.5 \frac{\Delta n\text{TA}}{\Delta t} \times hp \quad (\text{eq. 6})$$

192 Here, $\Delta n\text{TA}$ = change in nTA ($n\text{TA} = \text{TA} \times \text{SSS}_{\text{average}}/\text{SSS}$) over the Δt (time), h =
193 Volume/volume/Area/area, and p = water density. The -0.5 scalar factor was applied to account for the

194 stoichiometric relationship, where 2 moles of TA produce 1 mole of CaCO_3 . Day and night incubations
195 (lasting 12 hours) were conducted simultaneously with organic carbon metabolism to obtain daily NEC
196 fluxes. The dark period (midnight to dawn) was used to measure nighttime dissolution, while the light
197 period (dawn to noon) was used for daytime calcification. Alkalinity was measured every 3 hours
198 throughout the incubation period. NEC is positive with TA consumption, indicating CaCO_3 precipitation,
199 and negative with TA production, indicating CaCO_3 dissolution.
200 In this study, both hourly and daily rates were reported. Hourly rates allow us to examine diel variations
201 in metabolic processes, while daily rates provide an integrated view of overall carbon dynamics,
202 facilitating comparison with existing literature.

203 2.5 Statistical analysis

204 Independent sample T-tests were applied to compare metabolic rates (R, NPP, GPP, NEM, NEC) between
205 SG and BS using SPSS v. 17. Data were subjected to a normality test before performing the analysis.
206 Least-squares linear regression was employed to assess the correlation between changes in DO in the SG
207 and BS. The Mann-Whitney U test was applied for carbonate chemistry analysis due to the non-normal
208 distribution of data.

209 3 Results

210 3.1 Water quality and carbonate chemistry

211 Diurnal patterns of water quality and carbonate parameters for SG and BS during the two-day ex situ core
212 incubation are illustrated in Figs. 3 and 4, respectively. The temperature in both treatments ranged from
213 22 to 29 °C, while salinity levels spanned from 35 to 36. These values were similar to in situ measurements
214 obtained from the seagrass beds using a CTD profiler. During the daytime (6:00 AM to 12:30 PM), PAR
215 levels ranged from $26 \mu\text{mol m}^{-2} \text{s}^{-1}$ to a peak of $1662 \mu\text{mol m}^{-2} \text{s}^{-1}$, with the highest intensities observed
216 at midday. The average PAR measured $953 \mu\text{mol m}^{-2} \text{s}^{-1}$ on the first day of incubation, increasing slightly
217 to $1026 \mu\text{mol m}^{-2} \text{s}^{-1}$ on the second day. DO saturation levels were more variable in SG than BS, with

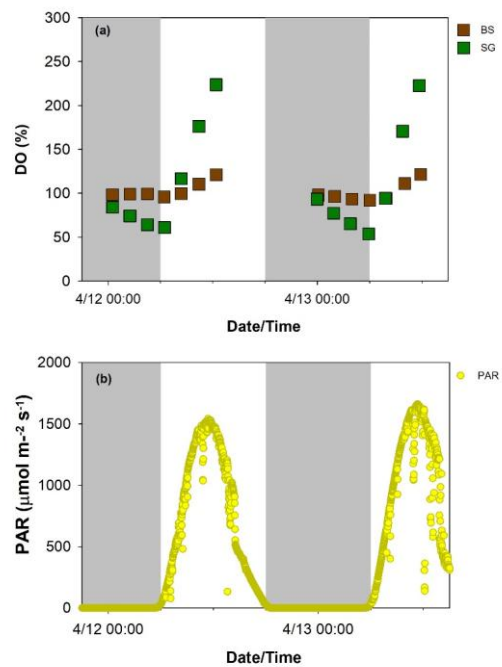
218 values ranging from 54% to 224% and 92% to 123%, respectively. DO saturation levels in both treatments
219 followed a diel pattern, with lower nighttime and higher daytime values.

220

221 Both $nDIC$ ($nDIC = DIC \times SSS_{\text{average}}/SSS$) and pH_T displayed greater diurnal fluctuations at SG compared
222 to the BS. At SG, $nDIC$ ranged from 1660 to 2118 $\mu\text{mol kg}^{-1}$ (mean \pm ~~SESD~~: $1963 \pm 44-153$ $\mu\text{mol kg}^{-1}$),
223 and followed a diel pattern. pH_T ranged from 7.81 to 8.37 at SG (mean \pm ~~SESD~~: $7.99 \pm 0.050.2$), following
224 the opposite trend to $nDIC$, with values decreasing at night and increasing during the day. ~~This daytime~~
225 ~~increase in pH_T at SG indicated the potential role of seagrass in mitigating ocean acidification effects~~
226 ~~during daylight hours.~~ At the BS site, these parameters were less variable, with $nDIC$ values ranging from
227 1948 to 2029 $\mu\text{mol kg}^{-1}$ and pH_T from 7.84 to 7.99, with mean values of 1993 ± 27 $\mu\text{mol kg}^{-1}$ and $7.93 \pm$
228 0.041 , respectively. Similarly, the calculated nTA was also more fluctuating in SG than BS, with mean
229 values of $2243 \pm 6-24$ $\mu\text{mol kg}^{-1}$ and $2230 \pm 6-24$ $\mu\text{mol kg}^{-1}$, respectively. The calculated pCO_2 displayed
230 a broader range at SG (142 to 762 μatm ; mean \pm ~~SED~~: 510 ± 62231) compared to BS (450 to 699 μatm ;
231 mean \pm ~~SED~~: 524 ± 2282), suggesting a more dynamic carbon cycling potentially driven by seagrass
232 metabolic activity. The mean Ω_{Ar} was higher in SG (3.14 ± 0.371) compared to BS ($2.72 \pm 0.140.4$),
233 indicating more favorable conditions for calcification at the seagrass site. Mann–Whitney test on
234 carbonate chemistry revealed no significant distinction between SG and BS (pH_T $p = 0.713$; $nDIC$ $p =$
235 0.419 ; nTA $p = 0.679$; Ω_{Ar} $p = 0.511$).

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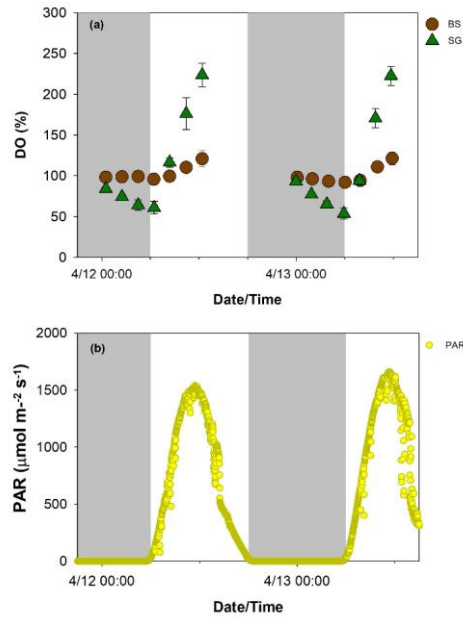
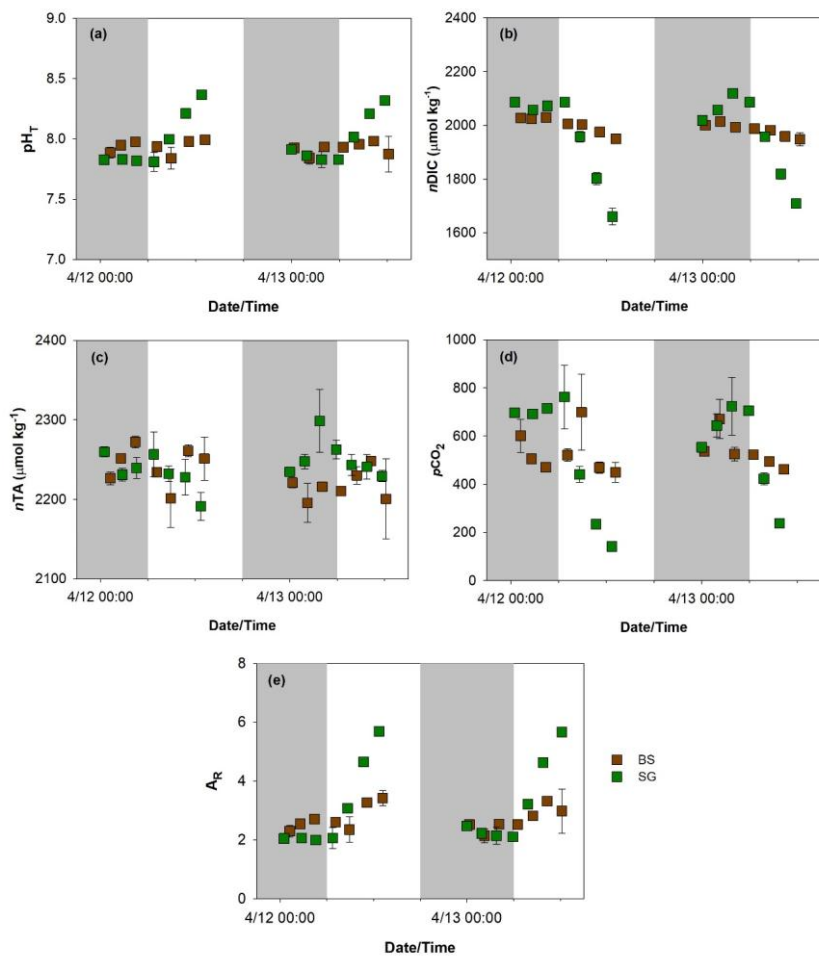


Figure 3: Diurnal pattern of dissolved oxygen (DO, a) in replanted seagrass (SG, green squaretriangle) and bare sediment (BS, brown squarecircle) ($n=9$, $\text{mean} \pm \text{SD}$), and photosynthetically active radiation (PAR, b) during the two-day (April 12-13, 2024) incubation.



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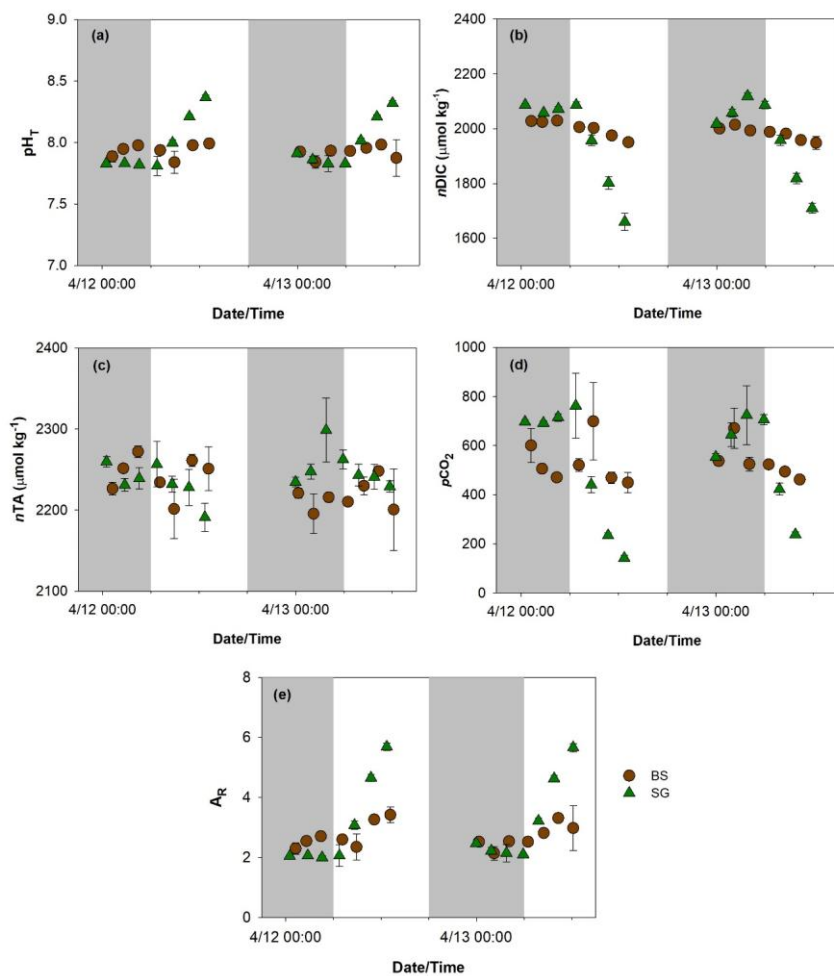
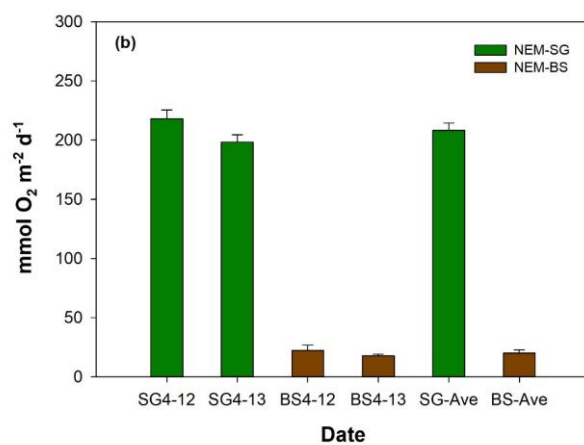
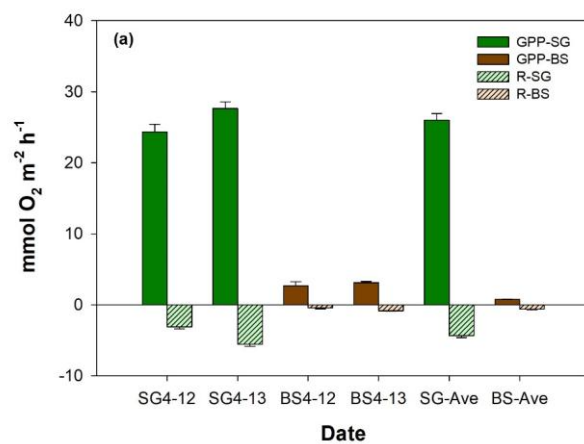


Figure 4: Total scale pH (pH_T , a), normalized dissolved inorganic carbon ($nDIC$, b), normalized total alkalinity (nTA , c), partial pressure of carbon dioxide (pCO_2 , d), and aragonite saturation state

248 (Ω_{AR} , e) in replanted seagrass (SG, green ~~square~~triangle) and bare sediment (BS, brown
249 ~~square~~circle) during the two-day (April 12-13, 2024) incubation. n=3, mean \pm ~~SE~~SD.

250 **3.2 Respiration, gross primary production, and net ecosystem metabolism**

251 Figure 5 illustrates the comparison of metabolic rates (mean \pm SD) between SG and BS. The mean
252 respiration rates in SG ($-4.3 \pm \del{0.3}1.5$ mmol O₂ m⁻² h⁻¹) were significantly higher than in BS ($-0.6 \pm 0.\del{4}4$
253 mmol O₂ m⁻² h⁻¹), by approximately 8-fold difference ($p < 0.01$). The mean GPP in SG was $26.0 \pm \del{1.0}3.4$
254 mmol O₂ m⁻² h⁻¹, which is 35-fold higher than in BS ($0.7 \pm \del{0.1}1.3$ mmol O₂ m⁻² h⁻¹) ($p < 0.01$). GPP was
255 always higher than R in both systems, with mean GPP/R ratios of 3.4 and 1.9 in SG and BS, respectively.
256 For NEM, both systems displayed positive values, indicating net autotrophy, with SG being 10-fold
257 higher ($208.2 \pm \del{6.3}22.2$ mmol O₂ m⁻² d⁻¹) compared to BS ($20.1 \pm \del{2.8}9.9$ mmol O₂ m⁻² d⁻¹) ($p < 0.01$). Both
258 R and GPP in SG and BS increased on the second day of incubation [SG (R: -3.1 vs -5.6 mmol O₂ m⁻²
259 h⁻¹; GPP: 23.3 vs 24.7 mmol O₂ m⁻² h⁻¹); BS (R: -0.4 vs -0.81 mmol O₂ m⁻² h⁻¹; GPP: 2.7 vs 3.1 mmol
260 O₂ m⁻² h⁻¹)], while NEM in SG (218.04 vs 198.4 mmol O₂ m⁻² d⁻¹) and BS (22.3 vs 17.8 mmol O₂ m⁻²
261 d⁻¹) showed a slight decrease. However, these changes were not statistically significant.~~Both R and GPP~~
262 ~~in SG and BS increased on the second day of incubation, while NEM showed a slight decrease, but these~~
263 ~~changes were not statistically significant.~~



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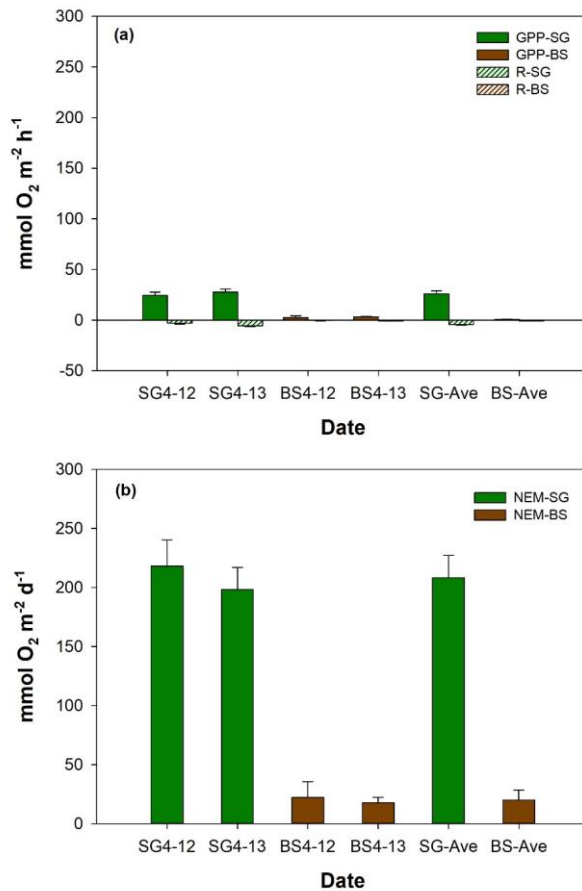
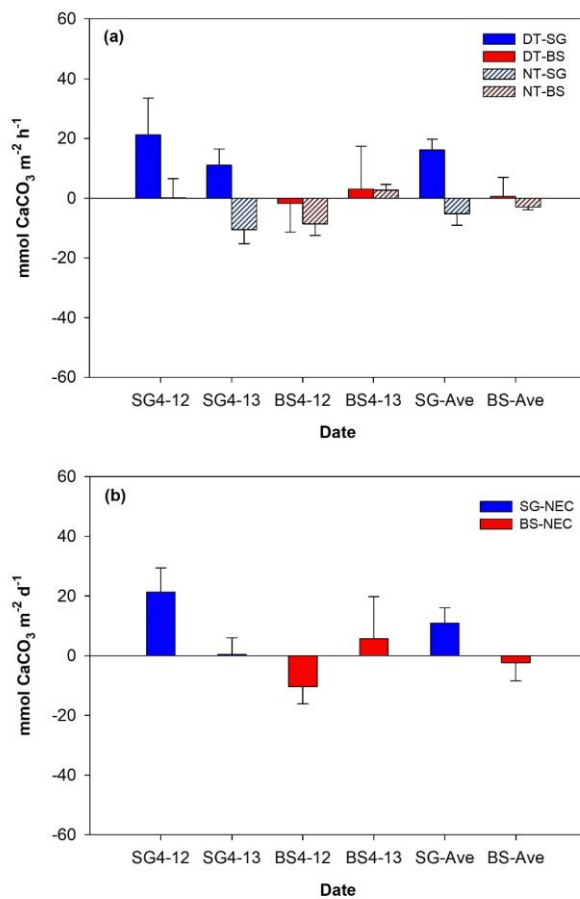


Figure 5: Mean (\pm SED, standard error deviation) values of (a) metabolic rates such as respiration (R), gross primary productivity (GPP), and (b) net ecosystem metabolism (NEM,) of restored seagrass (SG, green bars) and bare sediment (BS, brown bars) in Penghu during the two-day (April 12-13, 2024) incubation (n=9).

272 **3.2 Calcium carbonate precipitation, dissolution, and net ecosystem calcification**

273 The NEC values (mean ± SD) over a diel cycle for SG and BS demonstrated differences in their overall
274 carbonate dynamics (Fig. 6). Over the two-day incubation period, SG exhibited a net calcifying system
275 with a mean positive daily NEC means ($10.9 \pm \textcolor{red}{5.1} \textcolor{red}{15.7}$ mmol CaCO₃ m⁻² d⁻¹), driven by daytime
276 calcification (~~16.1 ± 1.3~~ $\pm \textcolor{red}{3.7} \textcolor{red}{1.3}$ mmol CaCO₃ m⁻² h⁻¹) despite nighttime dissolution (~~-5.2 ± 0.4~~ $\pm \textcolor{red}{3.9} \textcolor{red}{0.9}$ mmol
277 CaCO₃ m⁻² h⁻¹). In contrast, BS supported a net-dissolving system with mean daily NEC ($-2.3 \pm \textcolor{red}{6.2} \textcolor{red}{18.8}$
278 mmol CaCO₃ m⁻² d⁻¹). Mean daytime calcification and nighttime dissolution were $0.6 \pm \textcolor{red}{6.4} \textcolor{red}{1.6}$ mmol
279 CaCO₃ m⁻² h⁻¹ and ~~-3.0 ± 0.2~~ $\pm \textcolor{red}{0.9} \textcolor{red}{0.6}$ mmol CaCO₃ m⁻² h⁻¹, respectively. Both systems followed a general
280 diurnal pattern, with positive NEC during the day (calcifying) and negative at night (dissolving).



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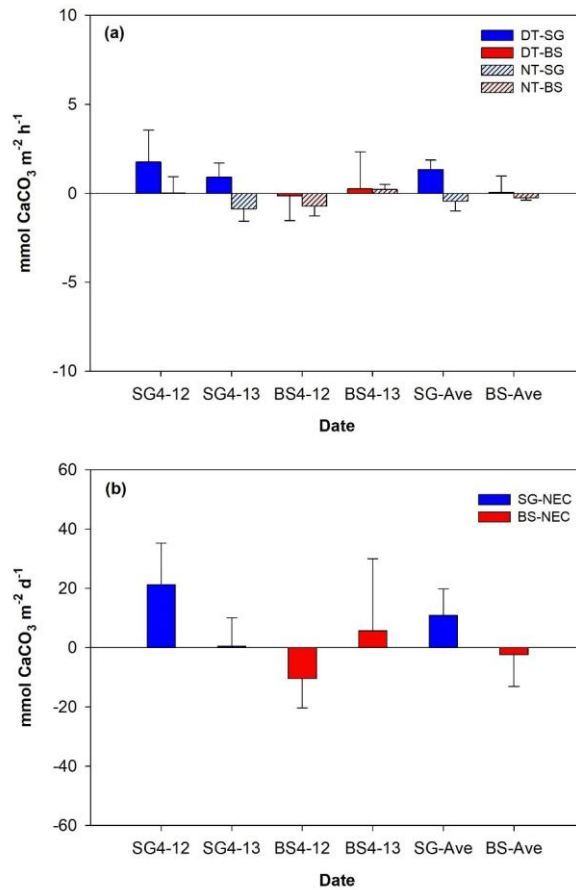


Figure 6: Mean (\pm SED, standard error deviation) values of daytime (DT) and nighttime (NT) calcification (a), and net ecosystems calcification (NEC, b) of restored seagrass (SG, blue bars) and bare sediment (BS, red bars) in Penghu during the two-day (April 12-13, 2024) incubation (n=93).

287 **4 Discussion**

288 Seagrass meadows are widely recognized as an important blue carbon ecosystem with substantial
289 potential to mitigate anthropogenic CO₂ emissions. Although research on seagrass ecosystems has grown
290 in recent years, significant gaps remain in understanding their carbon dynamics. In particular, the balance
291 of organic and inorganic carbon processes within these systems is not fully understood. Meanwhile, global
292 seagrass coverage continues to decline, which has increased the urgency of restoration efforts (Waycott
293 et al. 2009). Restoring seagrass meadows to enhance carbon sequestration has become increasingly
294 important. Currently, most studies on restored seagrass meadows focus primarily on the burial of
295 particulate organic carbon (Greiner et al. 2013), with far fewer exploring both organic metabolism and
296 carbonate cycling in restored seagrass meadows. Here, we present the first dataset on carbon uptake
297 through metabolic rates and calcification measurements in restored seagrass meadows within tropical
298 regions.

299 **4.1 Restoration of seagrass enhances metabolic rates**

300 The metabolic rates estimated in present study were comparable to those recorded in other seagrass
301 meadows (Table 1). Our GPP in SG was 24% and 37% higher than the tropical and global averages,
302 respectively, but 38% lower than Dongsha Island, Taiwan (Chou et al., 2023). It is also comparable to
303 measurements reported for *H. uninervis* in Tropical Australia (Table 1). Conversely, the R values
304 estimated in this study were roughly half lower than the tropical and global averages (Duarte et al., 2010).
305 Our NEM (214 mmol O₂ m⁻² d⁻¹) is within the range of previous estimates for tropical seagrass meadows
306 (-477.28 to 484.20 mmol O₂ m⁻² d⁻¹) and global estimates (-477.28 to 531.63 mmol O₂ m⁻² d⁻¹). In addition
307 to these global comparisons, our study reveals a clear distinction in metabolic rates (e.g. GPP, R, NEM)
308 between SG and BS. The GPP and R in restored seagrass meadows were 35 and 7 times greater than in
309 BS. The relatively higher metabolic rates in seagrass meadows compared to bare sediments have also
310 been observed in other studies (Table 1). For instance, a two-year-old restored *Halodule wrightii* meadow
311 demonstrated a 13-fold increase in NEM relative to bare sediment (Egea et al., 2023). Similarly,
312 *Posidonia oceanica* exhibited a notable 70-fold increase in metabolic rates compared to bare sediment
313 (Barron et al., 2006). Furthermore, *Zostera marina* exhibits net autotrophy while bare sediments are net

heterotrophy (Attard et al., 2019; Chen et al., 2019). Such patterns highlight the fundamental ecological functions restored seagrass meadows play relative to unvegetated/bare sediments. The increase in GPP reflects the enhanced carbon fixation capacity of seagrass meadows, while the elevated R indicates active organic matter decomposition and microbial respiration (Duarte and Krause-Jensen, 2017). According to Duarte et al. (2010), seagrass meadows generally act as autotrophic (NEM > 0) CO₂ sinks when GPP exceeds 186 mmol O₂ m⁻² d⁻¹, and shift to heterotrophy (NEM < 0) when GPP falls below this threshold~~become autotrophic (NEM > 0) when GPP is greater than 186 mmol O₂ m⁻² d⁻¹, shifting to heterotrophy (NEM < 0) at lower levels.~~ Based on this threshold, our mean GPP for restored seagrass exceeded the value for autotrophy, resulting in a positive NEM which is consistent with their global assessment. The NEM observed in SG was 10 times higher than in BS, suggesting that SG sequesters significantly more carbon than BS. These findings highlight that seagrass restoration significantly boosts metabolic rates and enhances carbon cycling. Given the increasing loss of global seagrass cover, restoration not only boosts ecosystem productivity but also strengthens the ability of coastal systems to remove carbon, thereby contributing to climate change mitigation efforts.

Table 1. Comparison of metabolic rates from global estimates. GPP and R values are expressed in mmol O₂ m⁻² h⁻¹ units, while NEM in mmol O₂ m⁻² d⁻¹.

Location	Method	Seagrass Community	GPP	R	NEM	References
Taiwan	Ex situ benthic chambers	Bare sediment	0.74 ± 0.09	0.62 ± 0.09	20.10 ± 2.84	This study
		<i>H. uninervis</i> , <i>H. ovalis</i>	25.99 ± 0.96	4.32 ± 0.26	208.21 ± 6.33	
Taiwan	Open water mass balance	<i>Thalassia</i> , <i>Cymodocea</i>	42.25 ± 14.42	20.71 ± 7.13	8 ± 61	Chou et al., 2023
Mexico	In situ benthic chambers	Bare sediment	2.13 ± 0.58	0.73 ± 0.16	8.1 ± 10.9	Egea et al., 2023
		2-year <i>H. wrightii</i>	13.76 ± 3.35	2.61 ± 0.40	102.4 ± 31.5	
		4-year <i>H. wrightii</i>	9.24 ± 2.34	1.60 ± 0.19	72.5 ± 27.9	
		4-year <i>H. wrightii</i>	9.34 ± 0.35	2.15 ± 0.25	60.7 ± 4.7	
Sweden	Aquatic eddy	3-year-old restored seagrass (<i>Z. marina</i>)			-5 to -15	Kindeberg et al., 2024

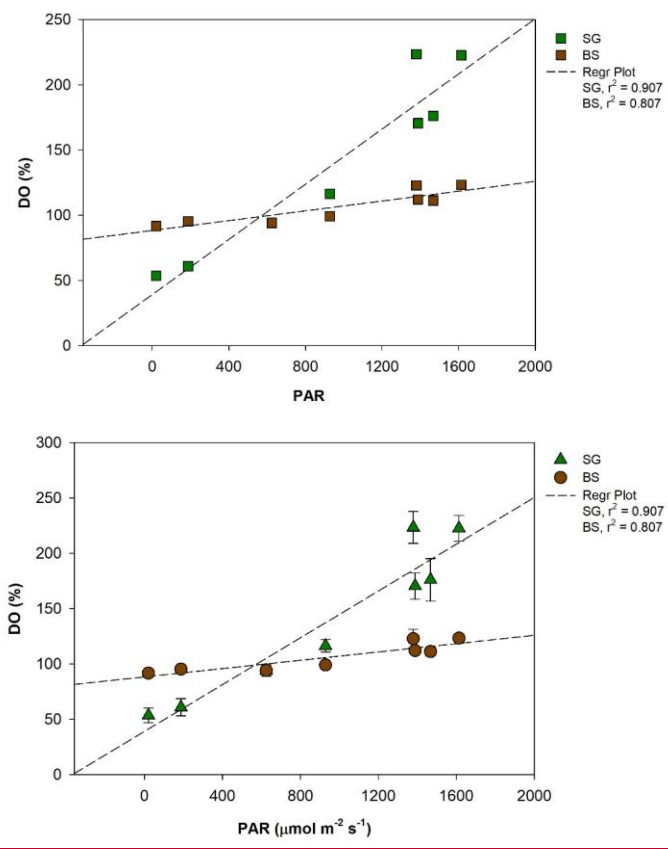
	covariance and benthic chambers	7-year-old restored seagrass (<i>Z. marina</i>)			-21	
Finland	Aquatic eddy covariance	Bare sediment	1.60	0.82	-0.14	Attard et al., 2019
		<i>Z. marina</i>	3.74	1.71	4.17	
Australia	Ex situ benthic	Bare sediment	2.28	1.26	-2.74	Chen et al., 2019
		<i>Zostera</i> sp.	6.94	2.74	7.12	
		<i>Halophila</i> sp.	2.05	1.60	-13.70	
Tropical Australia	Combined methods	<i>H. uninervis</i>	23.42 ± 3.67	9.63 ± 4.04	50 ± 53	Duarte et al., 2010
Tropical	Combined methods	All species	21 ± 0.6	9 ± 0.6	24 ± 8	Duarte et al., 2010
Global	Combined methods	All species	19 ± 0.5	8 ± 0.4	27 ± 6	Duarte et al., 2010
Spain	In situ benthic chambers	Bare sediment	0.43	0.22	0.27	Barron et al., 2006
		<i>P. oceanica</i>	7.72	3.18	16.44	

The daily values of *R* and *GPP* reported in the literature were divided by 24 and 12, respectively, to calculate the hourly values.

Key drivers of elevated metabolic rates in tropical meadows include greater PAR availability, aboveground biomass, and higher temperatures (Ganguly et al., 2017; Ward et al., 2022). Many tropical species grow near their optimal photosynthetic and physiological conditions (Lee et al., 2007; Koch et al., 2012), efficiently capturing light in shallow, clear waters, which contributes to higher NEP (Ralph et al., 2007). In our study, DO variation corresponds to light intensity (Figs. 3 and 7), suggesting that the elevated GPP observed in seagrass meadows could be driven by higher light intensity. This is likely due to the relatively lower canopy cover of *H. uninervis* and density in SG, which reduces shading within the seagrass. As a result, more light penetrates to the leaves, increasing their photosynthetic surface area and contributing to NEM (Ralph et al., 2007). In contrast, lower respiration rates in the SG area were primarily likely due to the sediment characteristics and organic matter quality in this habitat. The seagrass beds are situated in carbonate-rich sediments, which typically contain less organic matter than siliciclastic or muddy sediments (Belshe et al., 2018; Kindeberg et al., 2018). This limits the availability of substrates for microbial decomposition. Moreover, the organic matter derived from seagrass detritus is generally

347 more refractory and less labile, further reducing its accessibility for microbial breakdown and thereby
348 suppressing heterotrophic respiration (Ren et al., 2024). Although seagrasses are capable of transporting
349 oxygen to their belowground tissues via internal aerenchyma (Borum et al., 2006), which can support
350 aerobic respiration, the combined effect of low organic content and poor substrate lability limits microbial
351 activity and oxygen consumption.

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356 **Figure 7: Regression plot between photosynthetically active radiation (PAR, $\mu\text{mol m}^{-2} \text{s}^{-1}$) vs**
357 **dissolved oxygen (DO, %) in restored seagrass (SG, green ~~square~~triangle) and bare sediment (BS,**
358 **brown ~~square~~circle). Error bars represent standard deviation (SD).**

359

360 Several studies indicate that restored seagrass can achieve primary productivity and carbon sequestration
361 levels comparable to natural meadows, although recovery depends on the extent of degradation,
362 restoration success, and site-specific habitat conditions (Oreska et al., 2017; Marbà et al., 2015). For
363 example, long-term research in Florida Bay demonstrated that sediment carbon sequestration rates and
364 plant biomass took nearly a decade to match those of natural meadows (Greiner et al., 2013). The ability
365 of restored meadows to maintain net autotrophy is crucial for their role as carbon sinks (Kennedy et al.,
366 2010). This is particularly relevant for climate change mitigation strategies, where the conservation and
367 rehabilitation of this ecosystem are recognized as natural climate solutions (Griscom et al., 2017).
368 Nonetheless, a recent investigation on restored seagrass exhibits net heterotrophy, as observed by
369 Kindeberg et al. (2024) in both 3-year and 7-year-old meadows in Sweden. A similar pattern also reported
370 in some natural seagrass meadows in Australia (Chen et al., 2019) (Table 1). This discrepancy underscores
371 the variability in seagrass productivity and metabolic processes based on geographical location and
372 environmental conditions, highlighting the need for region-specific assessments to fully understand
373 seagrass ecosystem dynamics. Long-term studies should also consider temporal and annual variations.

374 **4.2 Calcification dynamics in restored seagrass**

375 Our results show that restored seagrass meadows exhibit significantly higher CaCO_3 cycling — both
376 formation and dissolution — compared to bare sediments. This corroborates with prior studies, which
377 documented enhanced carbonate dynamics in vegetated habitats relative to unvegetated sediments. For
378 instance, *P. oceanica* and *Thalassia testudinum* meadows have been shown to promote both CaCO_3
379 production and dissolution (Burdige and Zimmerman, 2002; Barrón et al., 2006), with tropical seagrass
380 ecosystems displaying similar patterns (Chou et al., 2021; Fan et al., 2024). Further, our data revealed a
381 typical diurnal pattern, with positive values during daytime (net calcifying) and negative values during

382 nighttime (net dissolving). These findings align with previous estimates, such as those in Florida Bay,
383 which reported similar diurnal calcification dynamics (Yates and Halley, 2006).
384
385 The variations of CaCO_3 production and dissolution in surface waters and sediment are related to the
386 carbon cycle through photosynthesis and respiration (Yates and Halley 2006). During daylight hours,
387 photosynthesis raises pH and reduces CO_2 levels in the water, creating favorable conditions for calcium
388 carbonate precipitation—a process referred to as light-enhanced calcification (Schneider et al., 2009).
389 We found a significant positive correlation between PAR and $n\text{TA}$ changes ($r^2=0.52$, $p<0.05$), suggesting
390 that increased light availability may enhance calcification by photoautotrophs in restored seagrass areas
391 during the day (Fig. 8). Additionally, our data showed a significant negative correlation between $n\text{TA}$
392 flux and NEM ($r^2=0.54$, $p<0.01$), indicating that higher photosynthetic activity (positive NEM) promotes
393 calcification by consuming TA, while lower NEM or net heterotrophy contributes to TA production,
394 likely through carbonate dissolution or anaerobic decomposition (Fig. 9). Similar relationships between
395 photosynthesis and calcification have been reported in marine calcifiers (Mallon et al., 2022), and the
396 influence of epiphytic organisms in promoting calcification during active photosynthesis has been
397 highlighted in seagrass meadows such as *P. oceanica* (Barrón et al., 2006). At night, carbonate dissolution
398 predominates as aerobic respiration produces CO_2 and carbonic acid in sediment porewater (Eyre et al.,
399 2014), lowering carbonate saturation and driving mineral dissolution (Burdige and Zimmerman, 2002;
400 Burdige et al., 2008; Chou et al., 2021; Fan et al., 2024). The degree of dissolution is directly link to the
401 rate of organic matter decomposition, which depends on the quantity of organic matter, its reactivity, and
402 oxygen availability (Anderson et al., 2005; Morse et al., 2006). ~~The presence of High~~ shoot density and
403 root biomass in restored seagrass meadows enhances organic matter supply and decomposition in
404 sediment, further driving nighttime dissolution (Holmer et al., 2013).

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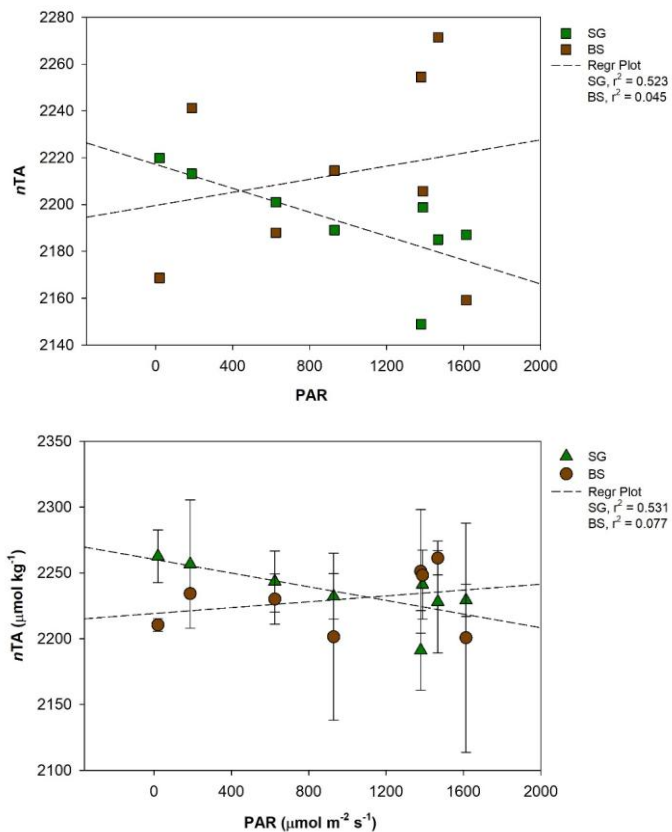


Figure 8: Regression plot between photosynthetically active radiation (PAR, $\mu\text{mol m}^{-2} \text{s}^{-1}$) vs normalized total alkalinity (nTA, $\mu\text{mol kg}^{-1}$) in restored seagrass (SG, green squaretriangle) and bare sediment (BS, brown squarecircle). Error bars represent standard deviation (SD).

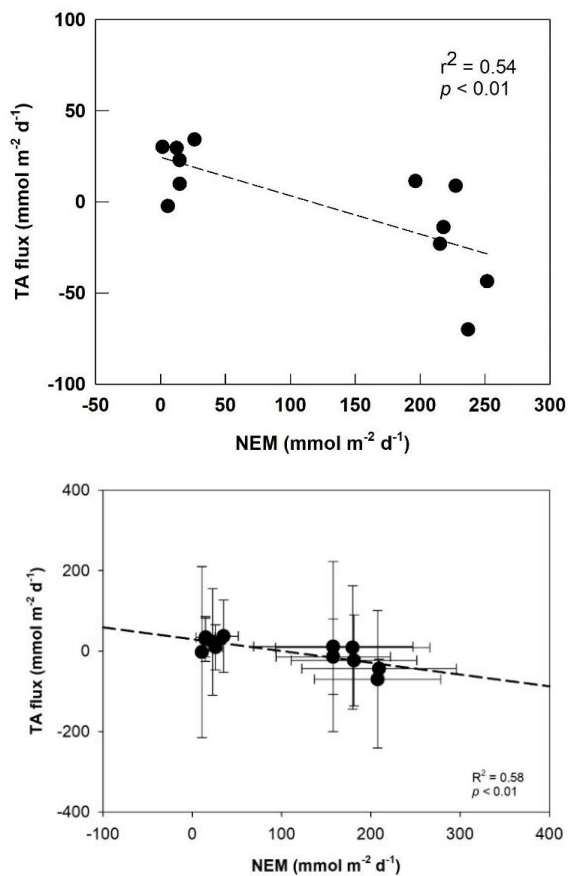


Figure 9: Linear regression showing the relationship between total alkalinity (TA, mmol m⁻² d⁻¹) flux and net ecosystem metabolism (NEM, mmol m⁻² d⁻¹) in this study in restored seagrass meadows and bare sediment. Error bars represent standard deviation (SD).

419

420 Over cumulative days, our NEC measurements indicate that restored seagrass meadows support overall
421 net calcification, whereas BS supports net dissolution. Our estimates are similar to those from Australia
422 (Walker et al., 1988) and seven times higher than Mediterranean seagrass net calcification rates (Barrón
423 et al., 2006), which are 295 g CaCO₃ m⁻² yr⁻¹ (8.8 mmol CaCO₃ m⁻² d⁻¹) and 51 g CaCO₃ m⁻² yr⁻¹ (1.40
424 mmol CaCO₃ m⁻² d⁻¹), respectively. In contrast, our findings are lower than those reported in the
425 Caribbean region of Mexico, where ex situ estimates ranged from 14 to 153 mmol CaCO₃ m⁻² d⁻¹
426 (Enriquez and Schubert, 2014). This highlights the enhanced carbonate production potential in tropical
427 seagrass meadows. A positive net calcification system occurs when CaCO₃ precipitation exceeds
428 dissolution within the system (Kleypas et al., 2001; Eyre et al., 2014). Restoration of seagrass meadows
429 provides a substrate for diverse calcifying organisms, including crustose coralline algae, bryozoans,
430 foraminifera, and serpulids, which enhance carbonate production (Beavington-Penney et al., 2005).
431 Epiphytes on seagrass leaves significantly contribute to CaCO₃ production, with tropical seagrass
432 meadows typically supporting higher carbonate loads than temperate ones. Reported production rates
433 span from 180 g CaCO₃ m⁻² yr⁻¹ in Jamaica (Land, 1970) to 2800 g CaCO₃ m⁻² yr⁻¹ in Barbados (Patriquin,
434 1972), underscoring regional variability in seagrass-associated calcification. Moreover, fluctuations in
435 CO₃²⁻ concentrations are crucial in regulating the capacity of calcifying organisms to form CaCO₃. Our
436 data reveals a ~~consistently higher mean Ω_{Ar} aragonite saturation state (Ω_{Ar}) in SG (3.14 ± 1) compared to~~
437 ~~bare sediment (BS) BS (2.72 ± 0.4). Notably, SG environments exhibit significant peaks in aragonite~~
438 ~~saturation, with a maximum value of 5.686, whereas the highest Ω_{Ar} in BS is 3.419.~~ Seagrass
439 photosynthesis raises pH and Ω_{Ar} , enhancing the calcification of surrounding calcifying organisms (De
440 Beer and Lakrum, 2001). However, the consumption of TA by calcifiers during the calcification process
441 releases CO₂, potentially counteracting pH increases and partially offsetting the net carbon uptake
442 potential of seagrass ecosystems (Alongi et al., 2008; Mazarrasa et al., 2015; Saderne et al., 2019). This
443 highlights the dual role of seagrass restoration in supporting biodiversity and CO₂ uptake while
444 influencing carbonate and carbon flux dynamics. Although the restored seagrass meadow in our study
445 functions as a net calcifying system, TA fluxes between SG and BS showed no significant difference.

4.3 Net carbon uptake of seagrass restoration

In order to estimate the net carbon uptake potential of seagrass restoration, we applied the photosynthesis-quotient (PQ) of 1 to calculate CO₂ uptake from organic carbon metabolism (Gattuso et al., 1998; Ward et al., 2022). In terms of carbonate dynamics, we applied Φ , as described by Humphreys et al. (2018), to calculate the size of CO₂ source or sink for each system. In the SG system, which is net calcifying, Φ indicates a CO₂ source, with 0.61 moles of CO₂ released into the seawater for each mole of CaCO₃ precipitated. In contrast, the BS system, which is net dissolving, Φ represents a CO₂ sink, with 0.65 moles of CO₂ absorbed for each mole of CaCO₃ dissolved. These values are comparable to previous findings, which reported a CO₂ flux-to-CaCO₃ precipitation ratio of 0.63 (Frankignoulle et al., 1994; Smith, 2013; Mazarrasa et al., 2015). In terms of carbonate dynamics, each mole of CaCO₃ formed releasing 0.6 mmol of CO₂ into the seawater was used (Frankignoulle et al., 1994). The calculated results show that total carbon uptake from NEM was 208 mmol CO₂ m² d⁻¹ in SG and 20 mmol CO₂ m² d⁻¹ BS. For NEC, the carbon release in SG was 6.52 CO₂ m² d⁻¹, while for BS, an additional CO₂ uptake was -1.5 mmol CO₂ m² d⁻¹. Consequently, the net carbon uptake is 202 and 224 mmol CO₂ m² d⁻¹ for SG and BS, respectively. Our results demonstrate that the primary productivity of restored seagrass through photosynthesis exceeds the rates of calcification by 31-fold, suggesting that restored seagrass can act as a net carbon sink. However, further assessments are necessary to capture temporal variations, as our current measurements are based on daily observations and one season only.

4.4 Limitations of ex situ benthic incubation and future research

We tested the ex situ benthic core incubation approach for restored seagrass meadows, drawing from the existing utilities in some coastal areas and freshwater ecosystems for sulfate and nutrient fluxes (Eyre, et al., 2005, Chen et al., 2019). Overall, the ex situ benthic incubation method provides a significant advantage by measuring both organic and inorganic carbon dynamics simultaneously, addressing a critical gap in previous methods that often overlook carbonate dynamics (Johanssen, 2023). This approach is also useful for assessing seagrass metabolism in subtidal meadows, where collecting data is challenging due to high labor costs and weather conditions. Moreover, some in situ autonomous methods are often

expensive and constrained operational periods of only a few weeks due to challenges like sensor error and biofouling (Yates and Halley, 2003; Takeshita et al., 2016). While this approach provides several advantages, one notable limitation is its applicability. Currently, the design is primarily suited for small seagrass, like *H. ovalis*, *H. uninervis*, and *Z. japonica*. It may not be adequate for larger species, like *Enhalus acoroides* and large *Thalassia hemprichii*, due to differences in size and growth characteristics. Moreover, we suggest validating the ex situ results with in situ data to ensure comparability with natural conditions, particularly the effects of light attenuation. Our measurements were obtained under ex-situ conditions in a shallow water column, which likely exposed the cores to higher irradiance than would be encountered in situ at different seagrass depths (2–4 m). While previous research has shown that ex situ and in situ incubations can yield comparable metabolic estimates, supporting the validity of our approach (Maher and Eyre, 2011), we acknowledge the need for future in situ incubations to more accurately capture the natural light environment experienced by seagrass leaves. Future research should integrate ex situ results with in situ data with different geographic and environmental settings to enhance the generalizability of the findings. This will provide a more accurate assessment of seagrass ecosystems' role in global carbon cycling and inform more effective coastal management and conservation practices.

5 Conclusion

This study investigates the organic carbon metabolism and carbonate dynamics of replanted SG compared to BS using the ex situ core incubation method. The results show that SG has higher GPP and NEM, while exhibiting similar NEC, making it a stronger carbon sink than BS. The findings highlight the role of seagrass restoration in enhancing carbon removal and contribute to a growing body of literature that highlights the ecological value of restored seagrass meadows. This study represents the first simultaneous quantitative estimate of the effect of both organic carbon metabolism and carbonate dynamics on carbon sequestration of restored seagrass in Southeast Asia, providing valuable insights into the region's carbon dynamics. We emphasize the need for long-term research on metabolic rates and carbonate dynamics to account for temporal variations and to fully understand the implications of these processes in carbon sequestration. This will also help optimize restoration strategies aimed at maximizing carbon sink

499 potential and mitigating ocean acidification. Furthermore, ex situ benthic incubation proves to be a
500 valuable tool for assessing carbon fluxes in seagrass meadows, particularly those dominated by pioneering
501 species, although further in situ assessments are necessary for comprehensive validation.

502 **Author contribution**

503 Wen-Chen Chou (WCC) and Jian-Jhih Chen (JJC) conceptualized the research and spearheaded the
504 implementation. JJC, Mariche B. Natividad (MBN), and Hsin-Yu Chou facilitated sample collection and
505 analysis. MBN and JJC performed the data analysis, drafted the manuscript, and its revision. WCC and
506 Lan-Feng Fan reviewed and revised the manuscript. All authors were involved in the finalization of the
507 manuscript.

508 **Competing interest**

509 The authors declare that they have no conflict of interest.

510 **Data availability**

511 The data supporting the findings of this study are available in the DRYAD repository at
512 <https://doi.org/10.5061/dryad.d7wm37qd0> (Natividad et al., 2025).~~The datasets in this study will be~~
513 ~~deposited in DRYAD Data Repository.~~

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