



1 The effect of a short oxygen exposure period on algal biomass degradation and
2 methane release from eutrophic and oligotrophic lake sediments

3

4 Sigrid van Grinsven^{1*}, Natsumi Maeda¹, Clemens Glombitza², Mark A. Lever²,
5 Carsten J. Schubert^{1,2}

6

7 ¹Department of Surface Waters – Research and Management, Swiss Federal
8 Institute of Aquatic Science and Technology (EAWAG), Seestrasse 79,
9 Kastanienbaum, 6047, Switzerland.

10 ²Institute of Biogeochemistry and Pollutant Dynamics, Swiss Federal Institute of
11 Technology, Zurich (ETH Zurich), Universitätstrasse 16, Zurich, 8092, Switzerland.

12

13 *Current address: Geomicrobiology, Department of Geosciences, University of
14 Tübingen, Germany, sigrid.van-grinsven@geo.uni-tuebingen.de

15



16 **Abstract**

17

18 Algal blooms in lakes result in large amounts of labile carbon being transported down
19 the water column towards the sediments, often resulting in temporary water column
20 hypoxia. The algal biomass is deposited at the surface sediments, where it is degraded
21 by the microbial community. Negative effects of algal blooms and biomass depositions
22 are sometimes mitigated by pumping air or oxygen into the bottom waters of lakes.
23 The fate of the algal biomass, in terms of greenhouse gas release, is however often
24 unknown. We investigated methane emissions from sediments originating from both
25 a eutrophic and oligotrophic lake and tested the effect of additional algal C inputs.
26 Additionally, we investigated the effect of a pulse supply of oxygen, a mediating
27 measure that is currently being used in the investigated eutrophic lake. Our results
28 show a difference in the control experiments based on the state of eutrophication, but
29 the methane release from new algal biomass additions was the same, although the
30 process proceeded more rapidly in the eutrophic sediments. A 3-week pulse of oxygen
31 lowered the emitted methane from both types of sediments by 50%, not only reducing
32 the emissions of algal biomass additions, but also reducing methane emissions from
33 the experiments without fresh organic matter inputs. This effect was relatively long-
34 lasting: its effects were visible for several weeks after anoxic conditions were re-
35 established, making it a potentially interesting measure to lower methane emissions
36 over a longer period.

37



38 **Introduction**

39

40 Lakes are known to be significant contributors to global methane emissions, despite
41 their relatively small surface area (Bastviken et al. 2004). Methane emissions are the
42 result of the net outcome of two processes: the production of methane, called
43 methanogenesis, and the consumption of methane, methanotrophy. Methane
44 emissions from aquatic environments originate mostly from sediments. Organic matter
45 is delivered from either internal lacustrine (autochthonous) or external, e.g. riverine
46 and terrestrial (allochthonous) sources. Autochthonous primary production, e.g. by
47 planktonic microalgae in the water column, captures CO₂ from the atmosphere to
48 produce biomass. After cell death, this biomass sinks down the water column, and
49 becomes part of the sediment. The decomposition of this biomass lowers dissolved
50 oxygen concentrations in bottom water and surface sediments and frequently
51 enhances sedimentary methane production (Fiskal et al. 2019; van Grinsven et al.
52 2022). Eutrophication, the increase in (mainly algal) primary production due to
53 increased nutrient concentrations in lakes, has thus been shown to increase methane
54 emissions from lakes (Beaulieu, DelSontro, and Downing 2019).

55

56 Most sublittoral lake sediments are anoxic from a depth of a few mm to cm below the
57 sediment surface. This is due to the mainly diffusive transport of oxygen into sediment
58 and the high rates of aerobic decomposition processes at the sediment surface. In the
59 underlying anoxic sediment, organic matter breakdown is performed by a community
60 of hydrolytic, fermentative, and respiring microorganisms. Generally, methanogenesis
61 is expected to occur only after the depletion of other, more energy-rich anaerobic
62 oxidants, such as nitrate, nitrite, metal-oxides and sulfate (Bastviken et al. 2004),



63 though strong overlaps in the distribution of methanogenesis with other anaerobic
64 respiration reactions have also been observed (Fiskal et al. 2019).

65

66 Generally, the anaerobic breakdown of organic matter follows three steps: first, the
67 large, complex organic compounds (e.g. macromolecules, polymers) are broken down
68 to their building blocks (e.g. monomers, oligomers, fatty acids) by extracellular
69 reactions (e.g. hydrolysis). Subsequently, these building blocks are taken up by
70 microbial cells and fermented to smaller chemical compounds, such as H₂, alcohols
71 (e.g. methanol, ethanol) and volatile fatty acids (VFAs; e.g., acetate, propionate,
72 butyrate, isovalerate, formate, and pyruvate). These smaller compounds can then
73 undergo a secondary fermentation step by syntrophic microorganisms, prior to
74 respiration, or be directly respired to CO₂ or methane using nitrate, nitrite, metal-
75 oxides, sulfate, or CO₂ as electron acceptors. Methanogens, which are respiring
76 organisms that gain energy through the production of methane, are mostly obligate
77 anaerobes, although their tolerance to oxygen is debated and may differ between
78 clades (Kato, Field, and Lettinga 1993; Zinder 1993; Kiener and Leisinger 1983).
79 Methanogenesis in sediments is believed to proceed mainly via three different
80 pathways: CO₂ reduction using H₂ or formate as electron sources, acetoclastic
81 involving the disproportionation of acetate into CO₂ and methane, or methylotrophic
82 methanogenesis, which involves the conversion of methylated compounds, such as
83 methanol, methylamines and methyl sulfides to methane. With few exceptions, these
84 pathways are performed by distinct taxa of methanogens.

85

86 A significant fraction of methane is consumed by methanotrophy, the process of
87 methane consumption. Methanotrophy can occur both aerobically or anaerobically via



88 the reduction of various anaerobic electron acceptors. Methanotrophs are found both
89 within the archaeal and bacterial domain and include strict aerobes, facultative
90 anaerobes, and obligate anaerobes. Methanotrophic activity often peaks at the oxic-
91 anoxic interface in either the sediment or water column, presumably due to the high
92 energy yields of aerobic methanotrophy.

93

94 Past research indicates that high organic matter inputs to lake sediments following
95 algal blooms result in increases in methane concentrations in the sediment (Schulz
96 and Conrad 1994). Both eutrophication and reduced water column mixing due to
97 increased thermal stratification as a result of global warming will likely increase the
98 frequency of algal blooms and contribute to more widespread bottom water anoxia in
99 the future (Hou et al. 2022). While this promotes the deposition and burial of organic
100 carbon in lake sediments, it will also increase methane emissions by increasing
101 methanogenesis rates and lowering rates of aerobic methanotrophy. In addition to
102 future algal blooms, legacy effects, e.g. continued high rates of methane production
103 sustained by the decomposition of older organic carbon from past periods of
104 eutrophication may contribute to these elevated methane emissions. These increases
105 in methane emissions may be a lesser concern for oligotrophic sediments, which are
106 generally lower in organic carbon content and hence methane production rates.

107

108 In order to mediate the effects of current and past eutrophication, artificial aeration is
109 applied to lakes in Switzerland. This aeration reduces the detrimental ecological and
110 socioeconomic consequences of seasonal anoxia. In addition, artificial aeration may
111 lower methane emissions from lakes by promoting aerobic methanotrophy and
112 reducing methane production in deep water columns and surface sediments, though



113 the efficacy of artificial aeration in achieving lower methane emissions is not known.
114 Here, we experimentally test the impact of artificial aeration on methane emissions
115 under different trophic regimes by short, pulse-wise, oxygen supply to sediments from
116 oligotrophic Lake Lucerne and eutrophic Lake Baldegg (both Switzerland). Based on
117 slurry and whole-core incubations, we study the impact of oxygen pulses on methane
118 emissions from sediments with and without an initial spike of algal biomass to mimic
119 the situation shortly after an algal bloom. Samples were extracted for gas
120 concentration and isotope analysis, as well as VFA and microbial community analyses.
121 Our results show that oxygen exposure had effects lasting past the oxic period such
122 as that methane emission rates remained lower for up to 10 weeks. The methanogenic
123 and methanotrophic communities did not seem affected by the oxygen exposure.

124

125

126 **Methods**

127

128 Study sites

129 Two lakes with a different trophic state were sampled for various experiments. A map
130 showing the location of both lakes is shown in Fig. S1.

131 Lake Lucerne is located at the northern alpine front in Central Switzerland (47°N, 8°E,
132 434 m a.s.l). It has a surface area of 116 km², and is fed by four alpine rivers that
133 provide ±80% of the lakes total water supply (Schnellmann et al. 2002). It is
134 oligotrophic, with a maximum P-concentration within the past century of 1.7 μM (Bürgi
135 and Stadelmann 2002). Further details on the trophic history of both lakes can be
136 found in (Fiskal et al. 2019).



137 Lake Baldegg is in an area with intensive cattle and pig farming within Central
138 Switzerland. It has a surface area of 5.22 km². Eutrophication has been ongoing, with
139 a peak in the 1970s, reaching P concentrations of 15.4 μM before remediation
140 measures were put in place. It has had an anoxic hypolimnion for almost 100 years
141 until artificial aeration was started in 1982 (Gächter and Wehrli 1998).

142

143 Field sampling

144 Lake Lucerne sediment cores were taken in November 2020 and February 2021 at a
145 location near the village of Kastanienbaum (47.00085N, 8.33697E). Lake Baldegg
146 sediment cores were taken in June 2021 from the center of the lake (47.193071N,
147 8.265238E). Both lakes were sampled with a multicorer device containing 10 cm
148 diameter, transparent butyrate plastic core liners of 65 cm, which were never filled
149 more than 3/4rd. The average core length was 40 cm. All sediment cores were brought
150 into a climate room of 10°C within 2 hours after core collection and stored until further
151 processing. Bottom water temperatures are between 5 and 9°C, according to (Fiskal
152 et al. 2019).

153

Slurry experiments in bottles

Oxygen regime	Biomass addition*	Lake	Headspace gas analysis	VFA	Microbial community
N ₂ flushed at start	0.1 g	Lucerne	+	+	+
N ₂ flushed after 1 week	0.1 g	Lucerne	+	+	+
N ₂ flushed after 3 weeks	0.1 g	Lucerne	+	+	+
N ₂ flushed at start	-	Lucerne	+	+	+
N ₂ flushed after 1 week	-	Lucerne	+	+	+
N ₂ flushed after 3 weeks	-	Lucerne	+	+	+
N ₂ flushed at start	0.1 g	Baldegg	+	+	+
N ₂ flushed after 3 weeks	0.1 g	Baldegg	+	+	+



N ₂ flushed at start	-	Baldegg	+	+	+
N ₂ flushed after 3 weeks	-	Baldegg	+	+	+

Whole sediment core experiments

Oxygen regime	Biomass addition*	Lake	Headspace gas analysis	VFA	Microbial community
N ₂ flushed at start	0.3 g	Lucerne	+	-	-
No N ₂ flushing, air headspace	0.3 g	Lucerne	+	-	-
N ₂ flushed at start	-	Lucerne	+	-	-
No N ₂ flushing, air headspace	-	Lucerne	+	-	-

* freeze-dried Spirulina algae (slurries) or 1:1 mixture of freeze-dried Spirulina + Chlorella algae (whole cores)

154

155 *Table 1. Overview of experiments. Further details are provided in Table 2 and the*
 156 *Methods section.*

157

158 Experimental setups

159 The experimental setups were designed to mimic oxygen intrusion from overlying oxic
 160 water. Two different experimental setups were used: slurry experiments (Lakes
 161 Baldegg and Lucerne) and whole core experiments (Lake Lucerne only). An overview
 162 of the experiments and analyses is presented in Table 1. Oxygen concentrations were
 163 followed and are shown in Fig. S2 and S3.

164

165 Slurry incubation experiments

166 The oxygen penetration depth of aquatic sediments generally does not exceed 1 cm
 167 (Fiskal et al. 2019; Horppila et al. 2015), with outliers in very organic material poor
 168 (marine) sediments up to 6 cm (Cai and Sayles 1996). It is highly unlikely that
 169 sediments deeper than 5 cm will experience oxygen intrusion. Therefore, we choose



170 a setup which does not expose the sediments deeper than 5 cm to oxygen, as this
171 would inhibit methanogenesis in a way that is highly unlikely to appear in lake
172 sediments.

173 To ensure this, sediment cores were separated in a surface part (0 – 5 cm depth) and
174 a deep part (5 – 15 cm depth). Lake bottom water was used to dilute sediment material
175 1:1, to create slurries. For the experiments that started completely anoxically directly
176 from the start, both the surface sediment and deeper sediment were added at the
177 same time and flushed with N₂. For the experiments that started under oxic conditions
178 (see Table 1), only surface sediments were placed into the incubation vials. The
179 deeper sediments were added upon O₂ removal, which was after either 1 or 3 weeks.
180 An overview is provided in Table 2. The oxygen concentrations inside a subset of
181 incubation bottles was followed to ensure the oxic and anoxic periods were indeed
182 established as aimed for (Fig. S2).

Treatment	Sediment provided at start	Sediment added after 1 week	Sediment added after 3 weeks	Final contents
N ₂ flushed at start	0 – 15 cm	-	-	0 – 15 cm
N ₂ flushed after 1 week	0 – 5 cm	5 – 15 cm	-	0 – 15 cm
N ₂ flushed after 3 weeks	0 – 5 cm	-	5 – 15 cm	0 – 15 cm

183

184 *Table 2. Overview of sediment additions within different experimental treatments of*
185 *the slurry experiment.*

186

187 Slurry experiments were performed in triplicate in 0.5L Schott bottles, filled with a final
188 volume of 258 ml sediment slurry. Each bottle was closed with an adapted stopper,
189 containing a three-way-stopcock that could be opened to allow throughflow of
190 sediment slurry and gas without exchange with the air. Freeze-dried algal biomass
191 (0.1 g of freeze-dried *Chlorella*; DietFoods CH) was added to a subset of the bottles.



192 Incubations proceeded at 10°C in the dark. The oxygen concentration was measured
193 daily to bi-weekly in 8 out of the 12 oxic bottles, using optical oxygen sensor spots
194 (Pyroscience, UK).

195

196 After 1 week or 3 weeks, the oxic bottles were flushed with N₂ for 15 minutes to remove
197 oxygen, after which they received additional anoxic sediments, according to Table 2.
198 The resulting oxygen trends are shown in Fig. S2. Sediments were added via the
199 sampling ports in the stoppers.

200

201 At each gas sampling time point, headspace gas was sampled via the sampling port
202 in the bottle caps. Prior to gas extraction, 10 ml of N₂ or air was pushed into the bottle.
203 Immediately after, 10 ml of headspace gas was sampled and stored in N₂ flushed 70
204 ml serum bottles. At each VFA (volatile fatty acids) and DNA sampling time point, 1.5
205 ml overlying water and 1.5 ml mixed slurry were sampled as described below. For the
206 VFA samples, samples without particulates were required. To achieve this, the bottles
207 were carefully tipped over, so the natural layering of sediment at the bottom and water
208 at the top remained. The water was then sampled via the sampling port. For the DNA
209 sample, the bottle was briefly shaken and then held upside down to take a mixed water
210 + sediment DNA sample. Both VFA and DNA samples were put directly on ice, and
211 after the sampling series was finished, moved to -20°C (VFA) or -80°C (DNA) freezers.

212

213 Whole core incubation experiments

214 In contrast to the slurry experiments, the cores for the whole core incubation
215 experiments were not disturbed or opened. Whole core experiments were only
216 performed with Lake Lucerne sediments for practical reasons, as shown in Table 1.



217 The core retrieval resulted in cores that were filled with on average 40 cm of sediment
218 and 25 cm of overlying water. To allow for headspace gas extraction over the
219 experiment duration, 12 cm of overlying water was carefully removed without
220 disturbing the sediment surface, leaving on average 13 cm of overlying water above
221 the sediment-water interface of each core. Whole core experiments were performed
222 in quadruplicate, an overview of the treatments is provided in Table 1.

223 To set up the treatments, freeze-dried algal biomass was added to selected cores ca.
224 18 hours after sampling (0.3 g per core, 1:1 mixture of freeze-dried *Chlorella* and
225 *Spirulina*; DietFoods, CH), corresponding to the same amount of algal biomass per
226 sediment mass as in the slurry incubations, and following earlier studies (Hiltunen,
227 Nykänen, and Syväranta 2021; Dai et al. 2005). The algal biomass was carefully
228 deposited on the sediment surface using a pipet, without disturbing the sediment-water
229 interface. The headspace and overlying water of the cores for the anoxic incubations
230 were flushed for at least 10 minutes with N₂.

231 Adjusted rubber stoppers with sampling ports were used to seal the cores on the top,
232 the bottoms were sealed off with rubber stoppers plus plastic caps. All cores were then
233 placed at 10°C in the dark. One core of each oxic treatment contained an oxygen
234 sensor spot (Pyroscience, UK), glued to the inner wall of the core liner.

235 Oxygen concentrations were measured at the start of the experiment and at irregular
236 intervals over the course of the experiment (shown in Fig. S3) and showed oxic
237 conditions were indeed retained over the full course of the oxic experiment, as was
238 aimed for.

239

240 Over the course of the experiment, gas samples for CH₄, CO₂ and N₂O analysis were
241 taken via the sampling ports. 10 ml of the gas headspace was extracted and placed



242 into 70 ml N₂-flushed serum bottles with butyl stoppers. After the gas sampling, 10-15
243 ml of N₂ gas (anoxic cores) or N₂ or air (oxic cores) was added via the sampling ports,
244 to equilibrate the internal and external pressure and limit the risk of leakage or
245 contamination. The gas pressure prior to sampling was determined with a pressure
246 meter and noted for each timepoint (not shown). One core was discarded due to water
247 leakage during the experiment.

248

249 Gas concentration and stable isotope analysis

250 Gas samples were analyzed for the concentration of CH₄ and CO₂ by gas
251 chromatography (GC; Agilent 6890N, Agilent Technologies) using a Carboxen 1010
252 column (30 m x 0.53 mm, Supelco), a flame ionization detector and an auto-sampler
253 (Valco Instruments Co. Inc.) for both the slurry and whole core experiments. Isotopic
254 ratios of methane ¹³C/¹²C (presented in the standard δ¹³C-notation relative to the
255 Vienna Pee Dee Belemnite (VPDB) reference) were measured in selected headspace
256 samples by isotope ratio mass spectrometry (IRMS; GV Instruments, Isoprime). To
257 purify, concentrate and combust the CH₄ to CO₂, injected samples were passed
258 through a trace gas unit (T/GAS PRECON, Micromass UK Ltd).

259

260 VFA analysis

261 Volatile fatty acids (VFA) analysis was only performed on samples from the slurry
262 experiment. Samples were filtered through pre-cleaned syringe filters Acrodisc™, 0.2
263 μm PES membrane, Supor™) and analyzed by two dimensional ion chromatography
264 (2D IC) at ETHZ according to the method described in (Glombitza et al. 2014) with
265 some modifications. The instrument used was a Dinonex™ ICS6000 (Thermo Fisher
266 Scientific) equipped with two 2.5-mm columns (AS24 for the first dimension and



267 AC11HC for the second dimension). The Retention time window on the first IC
268 dimension to collect the bulk VFAs for injection onto the second IC column was set to
269 3 min – 6.5 min to account for the low salinity of the freshwater samples compared to
270 the original method, as described in (Schaedler et al. 2018; Vuillemin et al. 2023).
271 Likewise, the VFA standards for quantification (mixed standards of formate, acetate,
272 propionate, butyrate, valerate, isovalerate and pyruvate at 1, 5, 10, 50 and 100 μmol
273 L^{-1}) were prepared in Milli-Q[®] water instead of IAPSO seawater as described in the
274 original method. Quantification was done using the conductivity detector signal of the
275 second IC dimension.

276

277 Microbial community analysis

278 Microbial community analysis was only performed on samples from the slurry
279 experiment. Each DNA sampled was stored at -80°C until processing. DNA was
280 extracted using the Qiagen Powersoil DNeasy kit without adaptations. The DNA
281 concentration of all extracts was measured on a Nanodrop device (Thermo Scientific).
282 When the concentration was below $2 \text{ ng}/\mu\text{l}$, an additional extraction was performed,
283 and samples were pooled. DNA extracts from all experiments were combined in two
284 lanes, including extraction blanks, and send for 16S rRNA NovaSeq PE250
285 sequencing (30K tags per sample) to Novogene UK, using the general 16S rRNA
286 archaeal and bacteria primer pair 515F and 806R, targeting the V4 region (Caporaso
287 et al. 2012). Quality control and species annotation were performed using the standard
288 Novogene pipelines ([https://www.novogene.com/eu-en/services/research-
289 services/metagenome-sequencing/16s-18s-its-amplicon-metagenomic-sequencing/](https://www.novogene.com/eu-en/services/research-services/metagenome-sequencing/16s-18s-its-amplicon-metagenomic-sequencing/)).
290 Raw sequencing data is deposited in the public repository [available upon publication,
291 or on reviewers' request].



294 **Results**

295

296 *Slurry experiments oligotrophic Lake Lucerne*

297 All oligotrophic slurry experiments that received algal biomass emitted significant
298 quantities of methane. The control experiment, without additional carbon source, only
299 emitted methane in the permanently anoxic setup (ca. 9 μmol per week). The control
300 slurries in which the top 5 cm was initially exposed to oxygen did not emit methane,
301 also not after anoxic conditions were established and the deeper anoxic sediments
302 were added (Fig. 1; Fig. 2). All given concentrations and concentration increases are
303 given in μM per liter headspace volume.

304 The addition of algal biomass to the oligotrophic slurries increased the methane
305 emission almost 25-fold under permanently anoxic conditions (to 161 μM per week),
306 but only 17-fold and 14-fold under the 1-week and 3-week oxic start conditions (118
307 μM and 93 μM per week), respectively (Fig. 1, Fig. 2).

308

309 Methane emissions started directly in the first week after the start of the anoxic
310 experiments (Fig. 2). After oxygen removal from the oxic start slurries, thus re-
311 establishing anoxic conditions and adding deeper sediments, methane emission also
312 started immediately. Net methane emission continued until week 13 in the anoxic
313 oligotrophic slurries, and two weeks longer in both types of the oxic slurries, despite
314 the 2 weeks difference between the start of methane emission in these two oxic
315 incubations (Fig. 2). The methane concentration in the oligotrophic slurries plateaued
316 and remained constant between weeks 13 or 15 to week 28, at a concentration of
317 1900 (anoxic), 1500 (1-week oxic start) or 1200 (3-weeks oxic start) μmol per L
318 headspace.

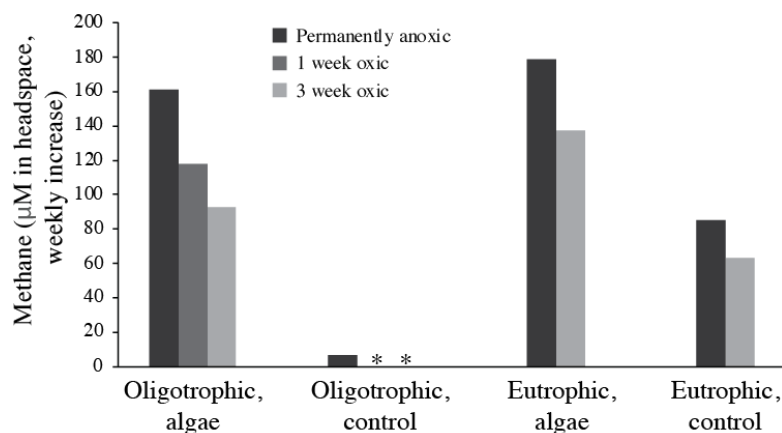


319 *Slurry experiments eutrophic Lake Baldegg*

320 Methane emission was observed in all eutrophic slurry setups, both the control and
321 algal biomass addition setups, after establishment of anoxic conditions and the
322 addition of deeper sediments. Under oxic conditions, and with only the top 5 cm, no
323 methane emission was observed. However, similar to the oligotrophic setup, the onset
324 of methane emission happened directly after establishment of anoxic conditions and
325 the addition of the deeper sediments.

326 The anoxic background methane emission in the eutrophic control slurries exceeded
327 those of the oligotrophic control slurries over 12-fold (85 versus 6.9 μM per week,
328 respectively, Fig. 1; Table S1).

329



330

331

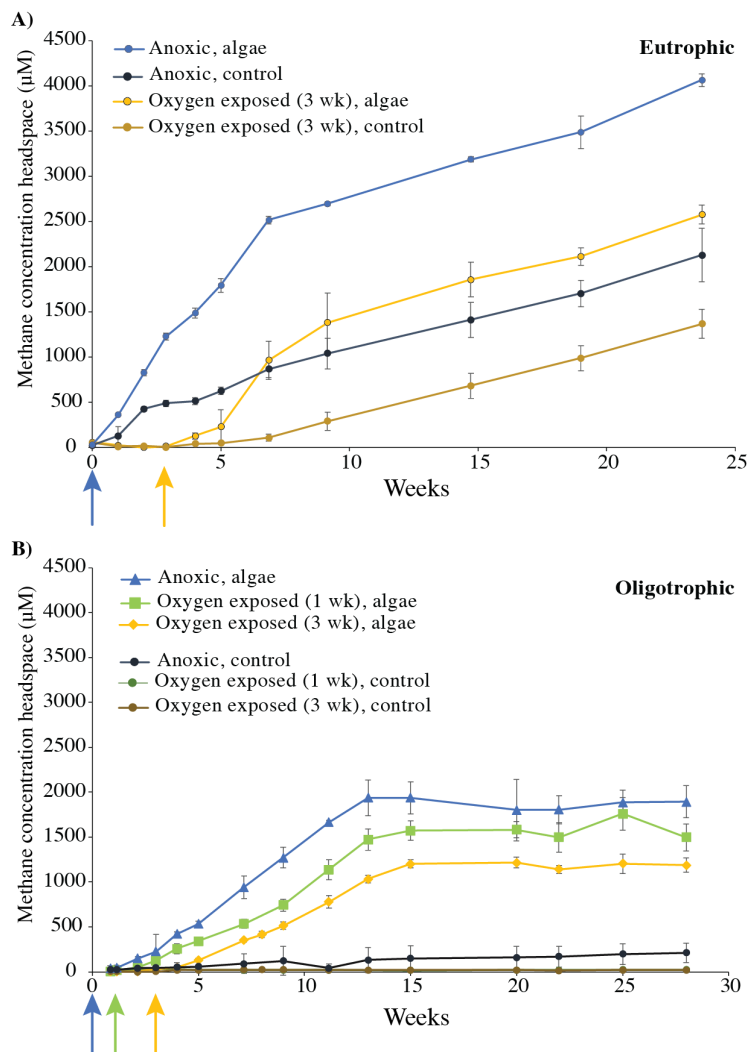
332 Fig. 1. Weekly increase in headspace methane concentration in the oligotrophic and
333 eutrophic slurry experiments, respectively, as derived from the linear phase of the methane
334 concentration plots (Fig. 2). * no increase in concentration detected.

335

336 The methane concentration in the eutrophic slurries did not plateau, although two
337 phases could be identified in the algae-fed slurries: A phase of rapid increase in



338 headspace CH₄ was observed from t₀ until week 7 in the anoxic algae-fed slurries, and
339 from week 3 to week 9 in the 3-week oxic start algae-fed slurries (Fig. 2). After this
340 initial phase, the methane emission rate stabilized at a similar rate as was observed
341 in the control experiments, as can be observed by the parallel lines in the graph of Fig.
342 2.
343



344

345



346 **Fig. 2.** Methane concentration in the headspace of the slurry experiments with A) eutrophic
347 and B) oligotrophic sediments. Arrows indicate the moment of oxygen removal, by flushing
348 with N₂, and the addition of anoxic sediments (see methods).

349

350 Even though the addition of algal biomass did increase the methane emission from
351 the eutrophic slurries, it was partly diminished when oxygen was present at the start
352 of the experiment. The background methane emission under anoxic control conditions
353 was similar to that of the oxic conditions with additional algal biomass, resulting in a
354 total amount of methane emitted of 2600 and 2100 μM for the oxic with algae and
355 anoxic control, respectively, corresponding to an increase of only 20% in emitted
356 methane (Table S1; Fig. 2). The methane emission rate was initially higher in the oxic
357 incubations with algae, but because the phase of high emission was of a short duration
358 (6 weeks, from week 3 to week 9), the total emission did not strongly exceed the total
359 anoxic control emission. The high emission phase in the anoxic algae experiment was
360 also short (t₀ until week 7), but due to the high weekly emission rate of 180 μmol, the
361 total amount of methane produced after 24 weeks was twice as high as the methane
362 emission in the anoxic control experiment, and 1.5 times higher than in the oxic algae
363 experiment (Fig. 2; Table S1).

364

365 Both the oligotrophic and eutrophic incubation experiments received equal amounts
366 of algal biomass. The methane emission rate in the oligotrophic anoxic experiments
367 increased from 6.9 to 161 μM per week due to the algae addition, whereas the
368 eutrophic anoxic rate increased from 85 to 179 μM per week, showing a much larger
369 increase in the oligotrophic experiments. The same holds for the 3-weeks oxic
370 experiments, which increased in weekly rate from 0 (control) to 93 (with algae) in the



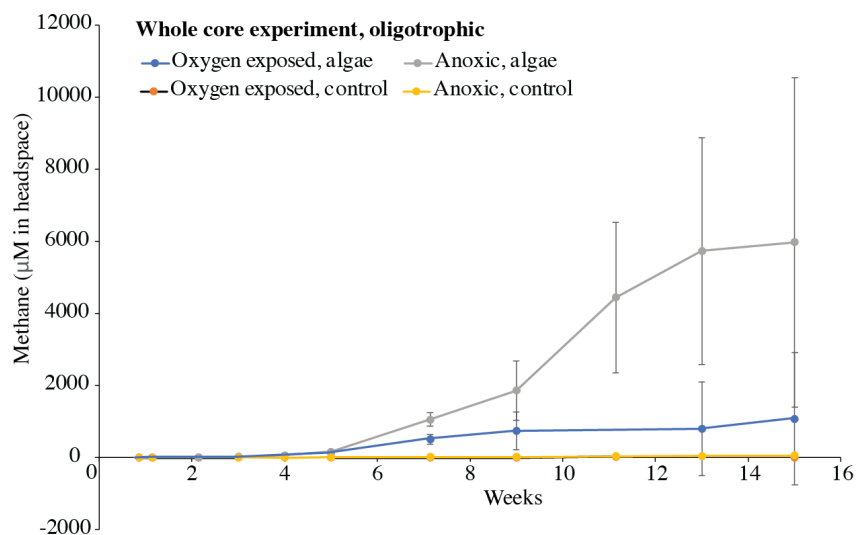
371 oligotrophic experiments and from 63 to 137 in the eutrophic experiments,
372 respectively. Due to the shorter duration of this high-rate methane emission the total
373 methane produced as a result of the algal biomass was, however, similar between the
374 oligotrophic and eutrophic sediments (Fig. 2).

375

376 *Intact core experiments oligotrophic Lake Lucerne*

377 Experiments with whole sediment cores, rather than sediment slurries, showed a
378 similar effect of oxygen exposure and algal biomass additions as the oligotrophic slurry
379 experiments. Although the variation within the experiments with whole cores was much
380 larger than in the more controlled slurry experiments, still a significant effect of oxygen
381 exposure, and of algal biomass addition, was observed (Fig. 3).

382



383

384

385 **Fig. 3.** Methane concentration in the headspace of the whole core experiment under
386 oxygen exposed (air headspace above overlying water) or anoxic (N₂ headspace)
387 conditions, with and without the addition of algal biomass, respectively. A cut-out of



388 the lower values that highlights the methane concentrations at the start of the
389 experiments is available in Fig. S4. The oxygen exposed control line is hidden from
390 view behind the anoxic control line in this graph.

391

392 **Microbial community in slurry experiments**

393

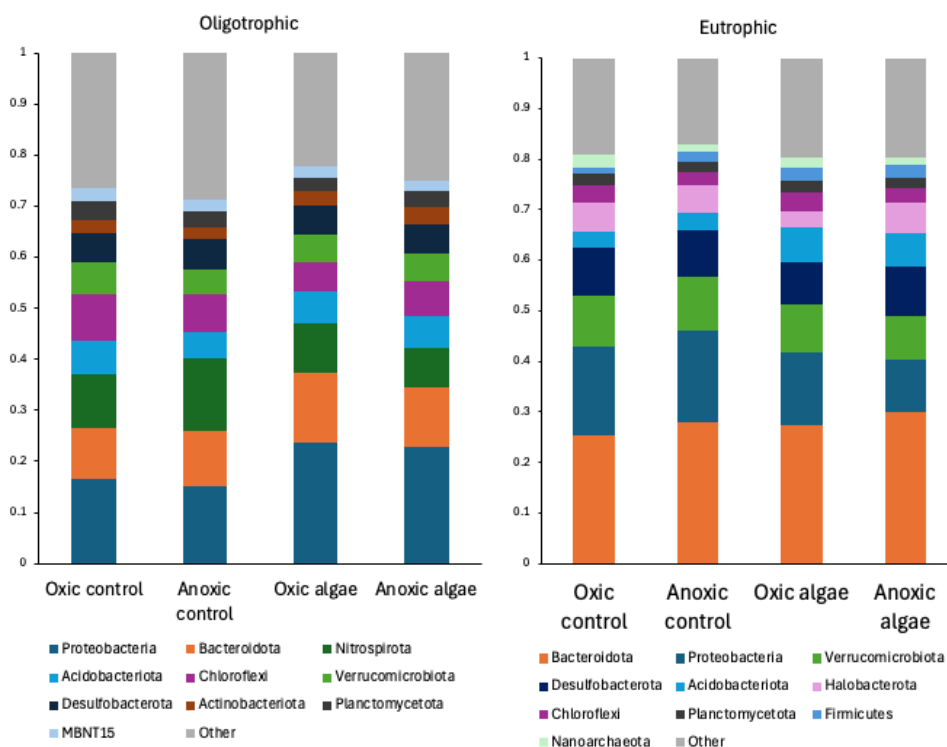
394 *Effect of algal biomass on microbial community*

395 The phyla of the Proteobacteria and the Bacteroidota were abundant in both the
396 eutrophic and oligotrophic sediment incubations. Other phyla that were found among
397 the 10 most abundant phyla in both setups were the Verrucomicrobiota, Chloroflexi,
398 Acidobacteriota, Desulfobacterota and Planctomycetota.

399 The microbial community was similar in the permanently anoxic and initially oxic
400 treatments, in both the oligotrophic and eutrophic sediment incubations. The addition
401 of the algal biomass influenced the microbial community composition. In the
402 oligotrophic incubations, the relative abundance of the Proteobacteria was higher in
403 all incubations with algae, both with and without oxygen exposure. In the eutrophic
404 experiments, the Proteobacteria abundance was actually lower in incubations with
405 algal biomass than without, both oxygen exposed and permanently anoxic. The
406 abundance of the Bacteroidota was higher in incubations with algal biomass than the
407 control, in all treatments in both setups. The Nitrospirota had a high abundance in
408 the oligotrophic sediments (11 – 14% in control setups), that was lowered under
409 conditions with algal biomass (8 – 10%, Fig. 4). In the eutrophic sediments, the
410 relative abundance of Nitrospirota was <1%. The abundance of the Acidobacteriota
411 was the same in oligotrophic sediments with and without algal biomass (5 – 7%). In



412 the eutrophic sediments, the relative abundance in the control incubations was lower
 413 (3%) than in the algal addition incubations (7%).



414

415 **Fig. 4.** Microbial phyla in the oligotrophic (A) and eutrophic (B) sediment incubations,
 416 after 15 weeks and 9 weeks, respectively. The 10 most abundant phyla, as detected
 417 by 16S rRNA sequencing, are shown. The timepoints correspond to the start of the
 418 stationary methane release phase in both setups, as shown in Fig. 2.

419

420 *Methanogenic/methanotrophic clades*

421 Surprisingly, nor the oxygen exposure or the algal biomass additions had a significant
 422 effect on the relative abundance of methanogenic and methanotrophic clades in the
 423 microbial community (Fig. S5). There was a profound difference between the
 424 oligotrophic and eutrophic sediments, but only for the methanogen relative abundance.



425 In the oligotrophic incubations, the relative abundance of operational taxonomic units
426 (OTUs) assigned to methanogenic orders was between 0.5 and 3% of the total 16S
427 rRNA detected sequences (Fig. S5) at all tested timepoints between 0 and 27 weeks.

428 In the eutrophic experiments, the relative abundance of OTUs assigned to
429 methanogenic orders was mostly between 4 and 7% of the total 16S rRNA detected
430 sequences in all selected timepoints between 0 and 15 weeks (Fig. S5).

431

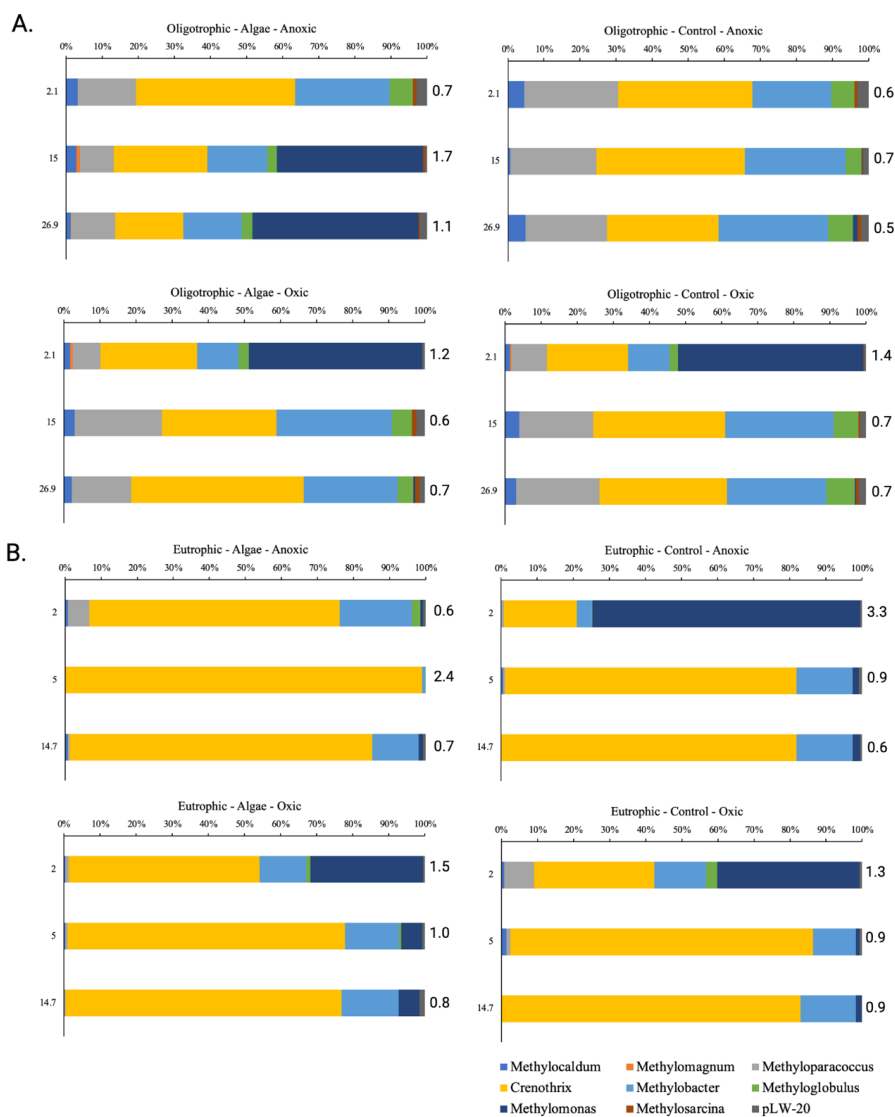
432 The methanogenic community was dominated by OTUs assigned to the order
433 Methanomicrobiales (Fig. S6). The contribution of Methanomicrobiales to the total 16S
434 rRNA sequences assigned to methanogenic orders was, however, higher in the
435 eutrophic (average of 75%) than in the oligotrophic experiments (average of 65%).
436 Methanomassiliicoccales made up a larger fraction in the oligotrophic experiments
437 (average of 29%, versus 12% in eutrophic experiments). Methanosarcinales were
438 relatively more abundant in the eutrophic experiments (Fig. S6). In the oligotrophic
439 experiments, no patterns were observed over time, nor differences between
440 treatments. In the eutrophic experiments, Methanomicrobiales dominated both the
441 oxic and anoxic experiments. However, in the permanently anoxic incubations, more
442 OTUs are assigned to Methanosarcinales, whereas in the temporarily oxic
443 incubations, more OTUs are assigned to the order Methanomassiliicoccales.

444

445 In both the oligotrophic and eutrophic incubations, the relative abundance of
446 methanotrophs belonging to the order Methylococcales was around 1 – 3% for all
447 treatments (Fig. 5). The majority of OTUs within the Methylococcales order were
448 assigned to the genus *Crenothrix* in all eutrophic treatments (20 – 99% of
449 Methylococcales reads), followed by the genus *Methylobacter* (0.9 – 30.6% of



450 *Methylococcales* reads, Fig. 5). However, recent research on the SILVA annotation
451 within the *Methylococcales* order has shown that a differentiation between *Crenothrix*
452 and certain groups of *Methylococcaceae* cannot be supported, and the weight of the
453 differentiation between these two groups is thus limited, currently (van Grinsven et al.
454 2022). The methanotrophic community in the oligotrophic experiments was less
455 dominated by sequences assigned to “*Crenothrix*”, and rather than a higher
456 abundance of *Methylobacter* assigned OTUs (Fig. 5) and a higher abundance of
457 genera that were marginal in the eutrophic incubations, such as *Methyloparacoccus*.
458 At specific time points, the genus *Methylomonas* showed particularly high peaks in its
459 relative abundance (up to 74% of *Methylococcales* reads) in both the eutrophic and
460 oligotrophic incubations (Fig. 5). Overall, the methanotrophic communities looked
461 relatively similar and no trends could be established over time, nor differences
462 between the oxygen or algae treatments (Fig. 5, Fig. S7).



463

464 **Fig. 5.** Relative abundance of OTUs assigned to methanotrophic genera in the
 465 oligotrophic (A) and eutrophic (B) incubation experiments. The bar plots show the
 466 abundance of specific genera relative to the total methanotroph abundance, whereas
 467 the number behind each bar indicates the abundance of methanotrophs relative to



468 the total microbial community (in % of 16S rRNA reads of each sample). The y-axis

469 shows the time in weeks since the start of the experiment.

470

471 **Volatile fatty acids in oligotrophic incubation experiments**

472

473 To compare the release of different volatile fatty acids (VFAs) after the algal biomass

474 additions, VFA concentrations were traced in the oligotrophic slurry experiments

475 during the first 8 weeks, as shown in Fig. 6. VFA concentrations were highest in anoxic

476 incubations, and were significantly lowered by temporary oxygen exposure. The

477 addition of algal biomass led to a strong increase in VFA concentrations (5 - 500 fold

478 increase) compared to control incubations. One of the anoxic control incubations had

479 20-fold higher VFA concentrations than the other two bottles of this treatment, resulting

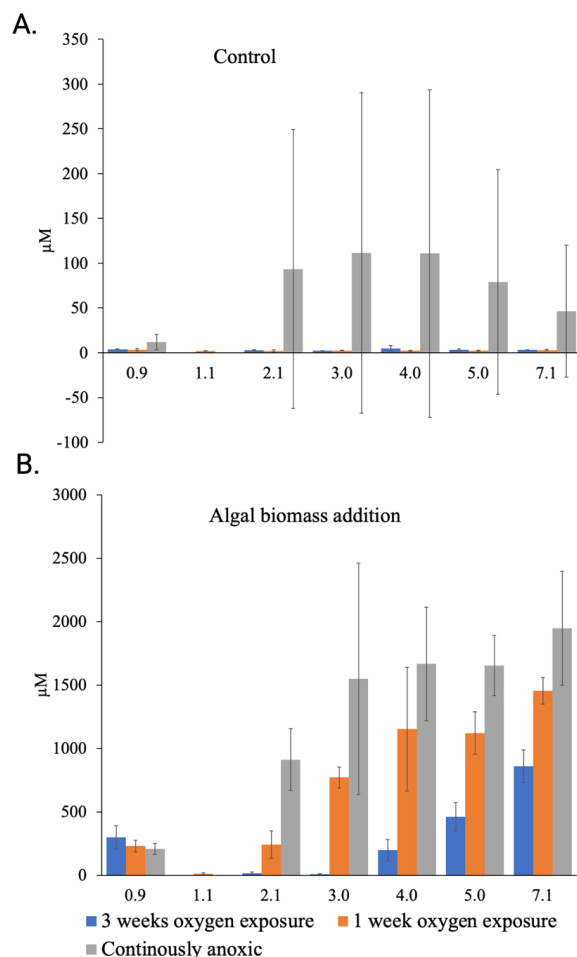
480 in the large error bars (Fig. 6A).

481 Acetate was the dominant VFA, with concentrations 10-100x higher than the other

482 VFAs (Fig. S8), and the key VFA to differ between oxic and anoxic treatments. The

483 concentrations of formate and pyruvate were not significantly affected by the oxygen

484 exposure.



485

486 **Fig. 6.** Total volatile fatty acid concentrations during the first 8 weeks of the oligotrophic
487 control (A) and algal biomass addition (B) experiment. The x-axis indicates weeks
488 since the start of the experiment. Note the different y-axis for A and B. Each bar
489 represents the average of triplicate samples at each timepoint. No samples were taken
490 at 1.1 weeks (8 days) of the anoxic and 3 weeks-exposure experiments.

491

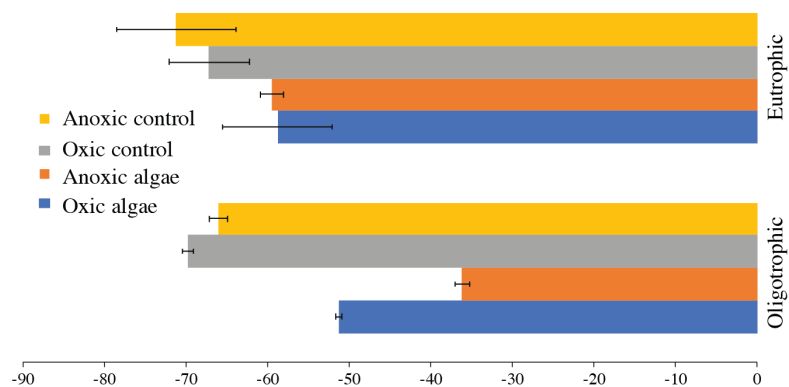
492 **Carbon isotopes of methane in slurry incubations**



493 The stable isotope profile of the headspace methane was determined only in the
494 stable, post-algal biomass degradation phase in the slurry experiments, when
495 methane emission rates no longer increased over time. This was at 17 weeks in the
496 oligotrophic, and at 11 weeks in the eutrophic experiments. The stable isotope ratio is
497 given in the standard $\delta^{13}\text{C}$ -notation, relative to the Vienna Pee Dee Belemnite (VPDB)
498 reference. When comparing the treatments, we see more negative $\delta^{13}\text{CH}_4$ values in
499 the algal biomass experiments in both the eutrophic and oligotrophic slurry
500 experiments, compared to the controls. A significant difference between the oxygen
501 exposed and anoxic treatments was only visible in the oligotrophic sediments.

502

503



504

505 *Fig. 7. $\delta^{13}\text{CH}_4$ values of headspace methane, after 11 weeks (eutrophic sediments)*

506 *or 17 weeks (oligotrophic) of slurry incubations.*

507

508



509 **Discussion**

510

511 *Trophic state and legacy effects on methane emissions and methane-cycling*
512 *communities*

513 The two investigated lakes differ in their current and historical trophic status. Lake
514 Lucerne is currently oligotrophic and has a history of low phosphate inputs. Lake
515 Baldegg is eutrophic and has been receiving high phosphate loading in the past,
516 resulting in water column anoxia and algal blooms in the period of 1910 – 1985(Fiskal
517 et al. 2019). Our results show that the eutrophication state of the lake affects the
518 methane emissions throughout the entire incubation period, both with and without
519 fresh organic matter inputs. The importance of legacy effects on biogeochemical
520 processes and communities in these lakes has been shown in earlier studies as well
521 (Fiskal et al. 2019; Han et al. 2020) and was also shown for other lakes along a trophic
522 gradient (Zhou et al. 2024). The methane emission was 12 times higher in control
523 experiments with eutrophic sediments than oligotrophic sediments, though no fresh
524 material input was delivered over 160 days (Fig. 1, Table S1). Contrastingly to their
525 historic and current carbon inputs, the TOC concentrations in sediments of both lakes
526 are comparable, according to a recent study at the same sampling locations as used
527 in this study (Fiskal et al. 2019).

528

529 Although these legacy effects are clearly visible in the methane emissions of the
530 control setups, the response to the input of new algal material was similar in magnitude
531 in both oligotrophic and eutrophic sediments (Fig. 1; Fig. 2). The addition of easily
532 degradable carbon compounds to environmental samples can spark a priming effect,
533 in which carbon stocks in the original sample are degraded more rapidly upon addition



534 of additional fresh material (Y. Wang et al. 2021; Guenet et al. 2010). Although we
535 cannot separate the contribution of algal carbon and sedimentary carbon to the
536 emitted CH₄, the equal emission responses from both lakes to the addition of fresh
537 organic matter suggest that the OM degradation response was also similar, and likely
538 primarily driven by degradation of the algal biomass rather than by older sedimentary
539 organic matter. A study with a similar setup regarding the addition of algal biomass
540 also showed a direct response in methane production rates during the first 60 days,
541 after which methane production rates stabilized (T. Wang et al. 2023).

542

543 Previous research has shown higher abundances and diversities of methanogens and
544 methanotrophs in eutrophic than in oligotrophic sediments (Yang et al. 2019), although
545 other studies indicate that the effect of sediment depth is of stronger effect (Yang et al.
546 2017) than the trophic state of the lake. The continuous high methane emission in our
547 eutrophic sediments (weekly increase of 85 μM in control experiments), in contrast to
548 the low methane emission in the oligotrophic sediments (weekly increase of 7 μM in
549 control experiments), also suggests that a more abundant and more active
550 methanogenic microbial community may exist in the eutrophic sediments. The
551 methanogenic community was indeed higher in its relative abundance (3 – 7%) in the
552 eutrophic experiments than in the oligotrophic sediments (0.5 – 4%).

553

554 *Effect of algal biomass and oxygen exposure on substrate availability to the microbial*
555 *community*

556 The conversion of algal biomass to gaseous methane emissions requires an initial
557 step of carbon degradation by fermenters, and a second step in which the reaction
558 products are converted into methane. The reaction products can consist of various



559 organic molecules, of which part can be used by methanogens directly, but others
560 need to be degraded further to become a suitable substrate for methane production.
561 Here, we found that the concentration of organic compounds indeed increased
562 strongly after the addition of algal biomass to the slurries, up to 500-fold (Fig. 6). This
563 was similar to a study by Schwarz et al. (Schwarz, Eckert, and Conrad 2008), who
564 found increased acetate and propionate concentrations in lake sediment incubations
565 with algal additions. A study by (Zhou et al. 2024) showed algal deposition on top of
566 the surface sediments led to a distinct increase in TOC in the top 8 cm of the sediment
567 cores, also without active mixing. The same two compounds as found in the Schwarz
568 et al. 2008 study, acetate and propionate, were also the major compounds detected
569 in our experiments. Surprisingly, these compounds were also produced (and build up)
570 under oxic conditions (200 μM acetate and 13 μM propionate under oxic conditions at
571 day 6, Fig. S8). However, the production of both acetate and propionate did not reach
572 the same values as the concentrations reached in the continuously anoxic incubations,
573 and a clear difference in the VFA buildup was also visible between 1 or 3 weeks
574 oxygen exposure treatments. The short exposure to oxic condition did lower the
575 acetate and propionate buildup, but did not diminish it. A recent study by (Kallistova et
576 al. 2023) showed that acetate additions strongly enhanced methane production from
577 surface sediments, showing it had an active function as methane precursor and higher
578 concentrations of acetate are likely directly correlated to higher methane emissions
579 from the sediments into the water column. In our experiments, substrates for acetate-
580 consuming methanogens were present in both the oxygen-exposed and permanently
581 anoxic experiments, but the concentrations were significantly lowered by oxygen
582 exposure at the start of the experiment. This corresponds to the methane production
583 in each of these treatments (Fig. 2). The methanogenic community did not show



584 similar patterns, suggesting that the substrate concentrations rather than the microbial
585 presence determines and predicts the methane emission rates in lake sediments.

586

587 *Microbial community*

588 A recent study by (Yang et al. 2021) followed the succession of a sedimentary
589 microbial community during an algal bloom, and found that both the archaeal and
590 bacterial community transitioned, taking part in biomass degradation steps that
591 changed over the time since the start of the algal bloom. (Schwarz, Eckert, and
592 Conrad 2008) noted that in their experiments with algal biomass additions, the
593 Deltaproteobacteria and Clostridiales increased immediately, and the Bacteroidetes
594 after 6 days. Methanogens, specifically acetate-using methanogens of the type
595 Methanosaetaceae increased in abundance after 6 days (Schwarz, Eckert, and
596 Conrad 2008). In our experiments, we noticed a similar increase in the
597 Proteobacterial relative abundance in the oligotrophic incubations due to algal
598 biomass additions, but a decrease in the eutrophic sediments (Fig. 4). The
599 Bacteroidota also increased in abundance in the oligotrophic incubations only. We
600 did not see an increase in the relative abundance in methanogenic clades in
601 response to the algal biomass addition, despite the much higher emission rates. A
602 similar effect was observed in experiments with algal biomass additions by (T. Wang
603 et al. 2023), who also saw large effects on the methane emission, but no increase in
604 methanogen copy numbers.

605 The effect of the oxygen exposure on the microbial community composition was
606 limited, both on the total prokaryotic community and on the specific methane-cycling
607 community. The methanogenic archaea are predominantly present in deeper
608 sediment layers (> 5 cm depth, as shown for these lakes by (Meier et al. 2024). They



609 will therefore likely not be affected by oxygen supply to lake bottom waters, which
610 was the process that was mimicked here. In our experimental setup, only sediments
611 of 0 – 5 cm depth were oxygen exposed, and 5 – 15 cm depth sediments were
612 added after oxygen removal. Therefore, the methanogenic community was
613 predominantly affected via the availability of substrates, and not due to direct oxygen
614 toxicity. Methanotrophs were presented throughout the sediment in these lakes (van
615 Grinsven et al. 2022). The methanotrophic bacteria found in our experiments are all
616 known as aerobic methanotrophs. However, these methanotrophs have been found in
617 anoxic environments, in these lakes and others, more often (van Grinsven et al.
618 2022). Although the methanotrophic community was diverse, especially in the
619 oligotrophic sediments, and showed changes in structure over time, these did not
620 seem related to oxygen exposure (Fig. 5).

621

622 *Oxygen exposure decreases methane emissions*

623 The effect of oxygen penetration depth on methane emission from lake sediments is
624 well established. However, these studies generally address long term stable oxygen
625 conditions ((Sobek et al. 2009; Huttunen et al. 2006). Here, we look at short oxygen
626 pulses, as a potential mediative measure for lakes with anoxic bottom water. The
627 presence of oxygen for a short, 3-week period at the start of the incubation had major
628 implications for methane emissions over the course of the entire experiments. The
629 total release of methane was significantly lower in the treatments that had experienced
630 an oxic period (Fig. 1; Fig. 2, Table S1). Most likely, part of the algal biomass was
631 converted to CO₂ and/or biomass during the oxic period and was therefore not directly
632 available for methanogenesis anymore. This is supported by the peak in CO₂
633 emissions that was observed during the oxic period of the experiments (Fig. S9; S10).



634 However, due to difficulties in translating headspace CO₂ concentrations to dissolved
635 CO₂, it is not possible to make a carbon mass balance, to see how much is indeed
636 released as CO₂. Part of the produced CO₂ will again be converted prior to release to
637 the headspace, leading to underestimates that cannot be sufficiently quantified. The
638 bubbling with N₂ to remove oxygen, that occurred at different timepoints in the different
639 experiments, removed CO₂ and may therefore has changed the pH in the system. pH
640 was not measured. Given the immediate production of CO₂ after bubbling (Fig. S9,
641 S10), we however assume that a [CO₂] close to natural conditions was rapidly
642 established following N₂ bubbling.

643 As CO₂ has a much lower warming potential per mole than methane (approximately
644 28 times lower on a hundred year basis, (Forster et al. 2021)) the release of CO₂ is
645 strongly preferred over that of methane in light of global warming. Besides CO₂, part
646 of the carbon may have been converted to microbial biomass during the oxic period,
647 and is stored as such in the sediments. (Sobek et al. 2009) published a weak linear
648 relationship between the diffusive methane flux from lake sediments, and the oxygen
649 penetration depth at those locations. A direct comparison with this study is, however,
650 difficult to make, as there are likely other factors involved that affect both the oxygen
651 penetration depth and the methane production, such as carbon content of the
652 sediments.

653

654 Directly after oxygen was removed from the incubation bottles and sediments from 5
655 – 15 cm depth were added, methane started to build up (Fig. 2). Algal biomass was
656 directly available for methane production, or the fresh organic matter enabled the
657 production of methane from previously present organic compounds (priming) or CO₂.
658 Methanotrophy may have been electron acceptor limited under the anoxic conditions,



659 and could not consume all methane produced. Even though the sediments recovered
660 directly after the establishment of anoxic conditions, and emitted methane, oxygen
661 pulse additions did decrease the methane release from the algal inputs.

662 A similar effect of an oxic-anoxic switch was observed by (Frenzel, Thebrath, and
663 Conrad 1990), who observed an abrupt increase in sedimentary methane emissions
664 when the oxygen concentration in the water overlying their core experiments dropped
665 below 18 μM . They assigned the difference between oxic and anoxic methane
666 emissions solely to an increased activity of methanotrophs under oxic bottom water
667 conditions.

668 Stable isotope analysis of the headspace methane in the stable, post-algal biomass
669 degradation phase (17 weeks of oligotrophic, and 11 weeks of eutrophic experiments,
670 Fig. 7), showed more negative $\delta^{13}\text{CH}_4$ values in the algal biomass experiments. The

671 $\delta^{13}\text{C}$ signal of the algal biomass likely decreased the $\delta^{13}\text{CH}_4$ values in the algal
672 addition experiments, with a larger effect in the oligotrophic lake, where the relative
673 contribution of algal biomass was largest, compared to the organic matter already
674 present in the sediments. Another potential factor is the shift in methanogenesis
675 pathway due to the algal biomass availability. (Zhou et al. 2022) showed that
676 cyanobacteria accumulation in lake sediments shifted the availability of organic
677 compounds for methane production and increased the potential for methylotrophic
678 methane production. Methylotrophic methanogenesis results in more depleted $\delta^{13}\text{CH}_4$
679 values compared to hydrogenotrophic methanogenesis (Summons, Franzmann, and
680 Nichols 1998). When comparing the oxic and anoxic experiments, only the oligotrophic
681 experiment showed significant differences: the $\delta^{13}\text{CH}_4$ values were lower (more



682 depleted) in the anoxic than in the oxic incubations, both with and without algal
683 biomass additions. This could also be caused by differences in methanogenesis
684 pathways, as hydrogenotrophic methane production (from CO₂) yields more ¹³C-
685 depleted methane than acetoclastic methanogenesis (Conrad 2005). As no changes
686 in the methanogenic community were observed between the oxic and anoxic
687 oligotrophic treatments, it is unlikely that a change in the community caused the
688 dominant methanogenesis pathway to swap and to cause the differences in the δ
689 ¹³CH₄ values. A further explanation is that differences in rates of methanotrophy
690 caused the observed changes in ¹³C-compositions of methane. Indeed, increased
691 rates of methanotrophy under oxic conditions would be expected to contribute to a
692 less depleted isotopic composition of the remaining methane (Barker and Fritz 1981).

693

694 *Methane emissions and implications*

695 Sedimentation of (algal) biomass is a key factor in the magnitude and seasonal
696 variation in lake methane emission rates (Gruca-Rokosz and Cieśła 2021). Our
697 experiments with intact sediment cores, rather than slurries, showed a significant
698 decrease in methane emissions under oxic bottom water conditions compared to
699 anoxic bottom waters, similar to our slurry experiments. Algal biomass led to a strong
700 increase in methane emissions, which was dampened by oxygen exposure. Oxygen
701 was not actively mixed into the sediments: only the overlying water and headspace
702 were made oxic, oxygen penetration into the sediments was due to natural occurring
703 diffusion. Algal biomass was deposited on top of the sediments, and not mixed in
704 either, to mimic natural algal deposition. Both in the oxic and anoxic algal-addition
705 experiments, methane emission started almost immediately after algal biomass
706 addition (Fig. 3; Fig. S4). The weekly methane release was however lower under oxic



707 conditions and resulted in lower concentrations at the end of the 15-week experiment,
708 despite the methanogenic zone of the sediments (> 5 cm depth) not being in direct
709 contact with either the oxygen or algal biomass.

710

711 Generally, only the sediment surface is affected by the oxygen conditions in the bottom
712 water; deeper sediments are anoxic, due to the low diffusion coefficient through
713 sediments. (Maerki et al. 2009) investigated the oxygen, carbon and nitrogen
714 dynamics of lake sediments, and stated that short term (weeks to months) oxygen
715 exposure is insufficient to change the reactivity spectrum of eutrophic Lake Zug
716 sediments, that the exposure times are too short for that. Our whole core experiment
717 shows, however, that despite the fact that the methanogenic layer is deeper in the
718 sediments than the bottom water affected layer, the conditions in the bottom water are
719 still of key importance for the methane emissions from the sediments following the
720 deposition of algal material, for example after an algae bloom in the surface waters.
721 (Maerki et al. 2009) also state that over 95% of the anaerobic mineralization in Lake
722 Zug sediments was due to methanogenesis, and that methane oxidation was
723 responsible for over half of the oxygen consumption at the sediment surface. If a
724 similar situation is the case in our eutrophic Lake Baldegg, changes in the methane
725 cycling are likely to have substantial effects on the carbon and oxygen cycling in the
726 shallow sediments.

727

728 Our experiments show that the effects of a short (1-3 week) oxygen exposure can
729 last for several months, i.e. decreasing methane emissions without changing the
730 methane-related microbial community (Fig. 5, Fig. S6, S5, S10). We believe these
731 findings should be further explored in environmental settings. In certain Swiss lakes,



732 artificial aeration is already applied to combat bottom water anoxia. If brief pulses of
733 oxygen, like the 1- and 3-week oxygen exposure periods tested here, have the
734 capacity to reduce longer-term methane emissions, we believe this could be
735 promising, especially if applied directly after an algal bloom, as tested here. Given
736 the expectations of ongoing eutrophication in the upcoming decades, plus the global
737 warming of lakes that further draws down oxygen levels, we believe this should be a
738 topic for further research.

739

740 **Data availability statement**

741 Raw reads of the 16S rRNA sequencing data is deposited and made publicly
742 available in the online repository NCBI SRA, under accession number XXX (in
743 progress).

744

745 **Author contribution statement**

746 Conceptualization by SvG, MAL and CJS. Original draft preparation by SvG, review
747 and editing by SvG, NM, CG, MAL and CJS. Investigation and Methodology by SvG,
748 NM and CG.

749

750 **Acknowledgements**

751 The authors thank Patrick Kathriner, Karen Beck, Kathrin Baumann, Cameron
752 Callbeck and Dimitri Meier for help in the field and the lab.

753

754 **Competing interests**

755 At least one of the authors is a member of the editorial board of BG. The authors
756 have no further conflicts of interest to declare.

756



757 **References**

- 758 Barker, James F., and Peter Fritz. 1981. "Carbon Isotope Fractionation during
759 Microbial Methane Oxidation." *Nature* 293 (5830): 289–91.
760 <https://doi.org/10.1038/293289a0>.
- 761 Bastviken, David, Jonathan Cole, Michael Pace, and Lars Tranvik. 2004. "Methane
762 Emissions from Lakes: Dependence of Lake Characteristics, Two Regional
763 Assessments, and a Global Estimate." *Global Biogeochemical Cycles* 18 (4):
764 1–12. <https://doi.org/10.1029/2004GB002238>.
- 765 Beaulieu, Jake J., Tonya DelSontro, and John A. Downing. 2019. "Eutrophication
766 Will Increase Methane Emissions from Lakes and Impoundments during the
767 21st Century." *Nature Communications* 10 (1): 3–7.
768 <https://doi.org/10.1038/s41467-019-09100-5>.
- 769 Bürgi, H. R., and P. Stadelmann. 2002. "Alteration of Phytoplankton Structure in
770 Lake Lucerne Due to Trophic Conditions." *Aquatic Ecosystem Health and
771 Management* 5 (1): 45–59. <https://doi.org/10.1080/14634980260199954>.
- 772 Cai, Wei-Jun, and Frederick L. Sayles. 1996. "Oxygen Penetration Depths and
773 Fluxes in Marine Sediments." *Marine Chemistry* 52 (2): 123–31.
774 [https://doi.org/10.1016/0304-4203\(95\)00081-X](https://doi.org/10.1016/0304-4203(95)00081-X).
- 775 Caporaso, J. Gregory, Christian L. Lauber, William A. Walters, Donna Berg-Lyons,
776 James Huntley, Noah Fierer, Sarah M. Owens, et al. 2012. "Ultra-High-
777 Throughput Microbial Community Analysis on the Illumina HiSeq and MiSeq
778 Platforms." *ISME Journal* 6 (8): 1621–24.
779 <https://doi.org/10.1038/ismej.2012.8>.
- 780 Conrad, Ralf. 2005. "Quantification of Methanogenic Pathways Using Stable Carbon
781 Isotopic Signatures: A Review and a Proposal." *Organic Geochemistry* 36 (5):
782 739–52. <https://doi.org/10.1016/j.orggeochem.2004.09.006>.
- 783 Dai, Jihong, Ming Yi Sun, Randolph A. Culp, and John E. Noakes. 2005. "Changes
784 in Chemical and Isotopic Signatures of Plant Materials during Degradation:
785 Implication for Assessing Various Organic Inputs in Estuarine Systems."
786 *Geophysical Research Letters* 32 (13): 1–4.
787 <https://doi.org/10.1029/2005GL023133>.
- 788 Fiskal, Annika, Longhui Deng, Anja Michel, Philip Eickenbusch, Xingguo Han,
789 Lorenzo Lagostina, Rong Zhu, et al. 2019. "Effects of Eutrophication on
790 Sedimentary Organic Carbon Cycling in Five Temperate Lakes."
791 *Biogeosciences* 16 (19): 3725–46. <https://doi.org/10.5194/bg-16-3725-2019>.
- 792 Forster, P., T. Storelvmo, K. Armour, W. Collins, J. -L. Dufresne, D. Frame, D. Lunt,
793 et al. 2021. "Chapter 7: The Earth's Energy Budget, Climate Feedbacks, and
794 Climate Sensitivity. Climate Change 2021: The Physical Science Basis.
795 Contribution of Working Group I to the Sixth Assessment Report of the
796 Intergovernmental Panel on Climate Change."
797 <https://doi.org/10.25455/wgtn.16869671.v1>.
- 798 Frenzel, Peter, Bernward Thebrath, and Ralf Conrad. 1990. "Oxidation of Methane in
799 the Oxic Surface Layer of a Deep Lake Sediment (Lake Constance)." *FEMS
800 Microbiology Letters* 73 (2): 149–58. [https://doi.org/10.1016/0378-1097\(90\)90661-9](https://doi.org/10.1016/0378-1097(90)90661-9).
- 801
- 802 Gächter, René, and Bernhard Wehrli. 1998. "Ten Years of Artificial Mixing and
803 Oxygenation: No Effect on the Internal Phosphorus Loading of Two Eutrophic
804 Lakes." *Environmental Science and Technology* 32 (23): 3659–65.
805 <https://doi.org/10.1021/es980418l>.



- 806 Glombitza, Clemens, Jeanette Pedersen, Hans Røy, and Bo Barker Jørgensen.
807 2014. "Direct Analysis of Volatile Fatty Acids in Marine Sediment Porewater
808 by Two-Dimensional Ion Chromatography-Mass Spectrometry." *Limnology*
809 *and Oceanography: Methods* 12 (JUL): 455–68.
810 <https://doi.org/10.4319/lom.2014.12.455>.
- 811 Grinsven, Sigrid van, Dimitri V Meier, Anja Michel, Xingguo Han, Carsten J Schubert,
812 and Mark A Lever. 2022. "Redox Zone and Trophic State as Drivers of
813 Methane-Oxidizing Bacterial Abundance and Community Structure in Lake
814 Sediments." *Frontiers in Environmental Science* 10 (March): 1–22.
815 <https://doi.org/10.3389/fenvs.2022.857358>.
- 816 Gruca-Rokosz, Renata, and Maksymilian Cieśla. 2021. "Sediment Methane
817 Production within Eutrophic Reservoirs: The Importance of Sedimenting
818 Organic Matter." *Science of The Total Environment* 799 (December):149219.
819 <https://doi.org/10.1016/j.scitotenv.2021.149219>.
- 820 Guenet, Bertrand, Michael Danger, Luc Abbadie, and Gérard Lacroix. 2010. "Priming
821 Effect: Bridging the Gap between Terrestrial and Aquatic Ecology." *Ecology*
822 91 (10): 2850–61. <https://doi.org/10.1890/09-1968.1>.
- 823 Han, Xingguo, Carsten Johnny Schubert, Annika Fiskal, Nathalie Dubois, and Mark
824 Alexander Lever. 2020. "Eutrophication as a Driver of Microbial Community
825 Structure in Lake Sediments." *Environmental Microbiology* 22 (8): 3446–62.
826 <https://doi.org/10.1111/1462-2920.15115>.
- 827 Hiltunen, Minna, Hannu Nykänen, and Jari Syväranta. 2021. "The Influence of Lipid
828 Content and Taxonomic Affiliation on Methane and Carbon Dioxide
829 Production from Phytoplankton Biomass in Lake Sediment." *Limnology and*
830 *Oceanography*, 1–11. <https://doi.org/10.1002/lno.11732>.
- 831 Horppila, Jukka, Petrina Köngäs, Juha Niemistö, and Susanna Hietanen. 2015.
832 "Oxygen Flux and Penetration Depth in the Sediments of Aerated and Non-
833 Aerated Lake Basins." *International Review of Hydrobiology* 100 (3–4): 106–
834 15. <https://doi.org/10.1002/iroh.201401781>.
- 835 Hou, Xuejiao, Lian Feng, Yanhui Dai, Chuanmin Hu, Luke Gibson, Jing Tang,
836 Zhongping Lee, et al. 2022. "Global Mapping Reveals Increase in Lacustrine
837 Algal Blooms over the Past Decade." *Nature Geoscience* 15 (2): 130–34.
838 <https://doi.org/10.1038/s41561-021-00887-x>.
- 839 Huttunen, Jari T., Tero S. Väisänen, Seppo K. Hellsten, and Pertti J. Martikainen.
840 2006. "Methane Fluxes at the Sediment-Water Interface in Some Boreal
841 Lakes and Reservoirs." *Boreal Environment Research* 11 (1): 27–34.
- 842 Kallistova, A. Yu., A. I. Kosyakova, I. I. Rusanov, V. V. Kadnikov, A. V. Beletsky, D.
843 D. Koval', S. K. Yusupov, I. Zekker, and N. V. Pimenov. 2023. "Methane
844 Production in a Temperate Freshwater Lake during an Intense Cyanobacterial
845 Bloom." *Microbiology* 92 (5): 638–49.
846 <https://doi.org/10.1134/S0026261723601586>.
- 847 Kato, Mario T., Jim A. Field, and Gatzke Lettinga. 1993. "High Tolerance of
848 Methanogens in Granular Sludge to Oxygen." *Biotechnology and*
849 *Bioengineering* 42 (11): 1360–66. <https://doi.org/10.1002/BIT.260421113>.
- 850 Kiener, Andreas, and Thomas Leisinger. 1983. "Oxygen Sensitivity of Methanogenic
851 Bacteria." *Systematic and Applied Microbiology*.
852 [https://doi.org/10.1016/S0723-2020\(83\)80017-4](https://doi.org/10.1016/S0723-2020(83)80017-4).
- 853 Maerki, Martin, Beat Müller, Christian Dinkel, and Bernhard Wehrli. 2009.
854 "Mineralization Pathways in Lake Sediments with Different Oxygen and



- 855 Organic Carbon Supply." *Limnology and Oceanography* 54 (2): 428–38.
856 <https://doi.org/10.4319/lo.2009.54.2.0428>.
- 857 Meier, Dimitri, Sigrid Van Grinsven, Anja Michel, Philip Eickenbusch, Clemens
858 Glombitza, Xingguo Han, Annika Fiskal, Stefano Bernasconi, Carsten J
859 Schubert, and Mark A Lever. 2024. "Hydrogen–Independent CO₂ Reduction
860 Dominates Methanogenesis in Five Temperate Lakes That Differ in Trophic
861 States." *ISME Communications* 4 (1): ycae089.
862 <https://doi.org/10.1093/ismeco/ycae089>.
- 863 Schaedler, Franziska, Cindy Lockwood, Ulf Lueder, Clemens Glombitza, Andreas
864 Kappler, and Caroline Schmidt. 2018. "Microbially Mediated Coupling of Fe
865 and N Cycles by Nitrate- Reducing Fe(II)-Oxidizing Bacteria in Littoral
866 Freshwater Sediments." *Applied and Environmental Microbiology* 84 (2): 1–
867 14.
- 868 Schnellmann, Michael, Flavio S. Anselmetti, Domenico Giardini, Judith A. McKenzie,
869 and Steven N. Ward. 2002. "Prehistoric Earthquake History Revealed by
870 Lacustrine Slump Deposits." *Geology* 30 (12): 1131–34.
871 [https://doi.org/10.1130/0091-7613\(2002\)030<1131:PEHRBL>2.0.CO;2](https://doi.org/10.1130/0091-7613(2002)030<1131:PEHRBL>2.0.CO;2).
- 872 Schulz, Silke, and Ralf Conrad. 1994. "Effect of Algal Deposition on Acetate and
873 Methane Concentrations in the Profundal Sediment of a Deep Lake (Lake
874 Constance)." *FEMS Microbiology Ecology* 6496.
- 875 Schwarz, Julia I.K., Werner Eckert, and Ralf Conrad. 2008. "Response of the
876 Methanogenic Microbial Community of a Profundal Lake Sediment (Lake
877 Kinneret, Israel) to Algal Deposition." *Limnology and Oceanography* 53 (1):
878 113–21. <https://doi.org/10.4319/lo.2008.53.1.0113>.
- 879 Sobek, Sebastian, Edith Durisch-Kaiser, Roland Zurbrugg, Nuttakan Wongfun,
880 Martin Wessels, Natacha Pasche, and Bernhard Wehrli. 2009. "Organic
881 Carbon Burial Efficiency in Lake Sediments Controlled by Oxygen Exposure
882 Time and Sediment Source." *Limnology and Oceanography* 54 (6): 2243–54.
883 <https://doi.org/10.4319/lo.2009.54.6.2243>.
- 884 Summons, Roger E., Peter D. Franzmann, and Peter D. Nichols. 1998. "Carbon
885 Isotopic Fractionation Associated with Methylotrophic Methanogenesis."
886 *Organic Geochemistry* 28 (7): 465–75. [https://doi.org/10.1016/S0146-6380\(98\)00011-4](https://doi.org/10.1016/S0146-6380(98)00011-4).
- 888 Vuillemin, Aurèle, Christoph Mayr, Jan A. Schuessler, André Friese, Kohen W.
889 Bauer, Andreas Lücke, Verena B. Heuer, et al. 2023. "A One-Million-Year
890 Isotope Record from Siderites Formed in Modern Ferruginous Sediments."
891 *Bulletin of the Geological Society of America* 135 (1–2): 504–22.
892 <https://doi.org/10.1130/B36211.1>.
- 893 Wang, Tong, Maidina Zhumabieke, Nan Zhang, Cheng Liu, Jicheng Zhong,
894 Qianjiahua Liao, and Lei Zhang. 2023. "Variable Promotion of Algae and
895 Macrophyte Organic Matter on Methanogenesis in Anaerobic Lake Sediment."
896 *Environmental Research* 237 (November):116922.
897 <https://doi.org/10.1016/j.envres.2023.116922>.
- 898 Wang, Yarui, Muhua Feng, Jianjun Wang, Xinfang Chen, Xiangchao Chen, Xian Du,
899 Fan Xun, and Bryne Tendelo Ngwenya. 2021. "Algal Blooms Modulate
900 Organic Matter Remineralization in Freshwater Sediments: A New Insight on
901 Priming Effect." *Science of The Total Environment* 784 (August):147087.
902 <https://doi.org/10.1016/j.scitotenv.2021.147087>.
- 903 Yang, Yuyin, Jianfei Chen, Xiuli Chen, Qingsong Jiang, Yong Liu, and Shuguang
904 Xie. 2021. "Cyanobacterial Bloom Induces Structural and Functional



- 905 Succession of Microbial Communities in Eutrophic Lake Sediments.”
906 *Environmental Pollution* 284 (September):117157.
907 <https://doi.org/10.1016/j.envpol.2021.117157>.
- 908 Yang, Yuyin, Jianfei Chen, Tianli Tong, Baoqin Li, Tao He, Yong Liu, and Shuguang
909 Xie. 2019. “Eutrophication Influences Methanotrophic Activity, Abundance and
910 Community Structure in Freshwater Lakes.” *Science of The Total
911 Environment* 662 (April):863–72.
912 <https://doi.org/10.1016/j.scitotenv.2019.01.307>.
- 913 Yang, Yuyin, Ningning Li, Wei Wang, Bingxin Li, Shuguang Xie, and Yong Liu. 2017.
914 “Vertical Profiles of Sediment Methanogenic Potential and Communities in
915 Two Plateau Freshwater Lakes.” *Biogeosciences* 14 (2): 341–51.
916 <https://doi.org/10.5194/bg-14-341-2017>.
- 917 Zhou, Chuanqiao, Yu Peng, Miaotong Yu, Yang Deng, Li Chen, Lanqing Zhang,
918 Xiaoguang Xu, Siyuan Zhang, Yan Yan, and Guoxiang Wang. 2022. “Severe
919 Cyanobacteria Accumulation Potentially Induces Methylophilic Methane
920 Producing Pathway in Eutrophic Lakes.” *Environmental Pollution* 292
921 (January):118443. <https://doi.org/10.1016/j.envpol.2021.118443>.
- 922 Zhou, Chuanqiao, Yu Peng, Muchun Zhou, Ruoyu Jia, Huazu Liu, Xiaoguang Xu, Li
923 Chen, Jie Ma, Tsuyoshi Kinouchi, and Guoxiang Wang. 2024. “Cyanobacteria
924 Decay Alters CH₄ and CO₂ Produced Hotspots along Vertical Sediment
925 Profiles in Eutrophic Lakes.” *Water Research* 265 (November):122319.
926 <https://doi.org/10.1016/j.watres.2024.122319>.
- 927 Zinder, Stephen H. 1993. “Physiological Ecology of Methanogens.” In
928 *Methanogenesis: Ecology, Physiology, Biochemistry {&} Genetics*, edited by
929 James G Ferry, 128–206. Boston, MA: Springer US.
930 https://doi.org/10.1007/978-1-4615-2391-8_4.
- 931
- 932
- 933
- 934