

The transport history of African biomass burning aerosols arriving in the remote Southeast Atlantic ~~marine boundary layer~~ and their impacts on cloud properties

Huihui Wu¹, Fanny Peers^{2,a}, Jonathan W Taylor^{1,b}, Chenjie Yu^{1,c}, Steven J Abel³, Paul A. Barrett³, Jamie Trembath⁴, Keith Bower¹, Jim M. Haywood^{2,3} and Hugh Coe¹

¹Department of Earth and Environmental Sciences, University of Manchester, Manchester, UK

²College of Mathematics, Engineering and Physical Science, University of Exeter, UK

³Met Office, Fitzroy Road, Exeter, EX1 3PB, UK

⁴Facility for Airborne Atmospheric Measurements (FAAM), FAAM, Cranfield, UK

^aNow at: Laboratoire d'Optique Atmosphérique, Université de Lille, Villeneuve-d'Ascq, France

^bNow at: Research IT, University of Manchester, Manchester, UK

^cNow at: Université Paris Cité and Univ Paris Est Créteil, CNRS, LISA, F-75013 Paris, France

Correspondence to: Hugh Coe (hugh.coe@manchester.ac.uk)

Abstract. African biomass burning (BB) aerosols transported over the southeast Atlantic (SEA) strongly influence cloud properties but remain a major source of uncertainty in regional climate assessment. This study characterizes vertical profiles of thermodynamic conditions, aerosol properties, and cloud microphysics around Ascension Island during an aircraft campaign (August–September 2017). Backward-dispersion simulations evidence that observed pollution originated from long-range transported African BB plumes. In BB-polluted marine boundary layers (MBL), aerosol number concentrations (N_a) were substantially elevated relative to the clean MBL, driving increased cloud droplet number concentrations (N_d) and reduced cloud effective radii (R_e). Cloud-layer mean N_d correlated strongly with aerosols below the cloud (sub- N_a) but weakly with free-tropospheric (FT) aerosols. Enhanced sub- N_a was due to BB aerosols entrained from the FT into the MBL along long-range transport and/or locally. Droplet activation fractions were similar in clean and moderately BB-polluted (sub- $N_a < 700 \text{ cm}^{-3}$) clouds, while a weaker N_d - N_a correlation was observed in more polluted clouds. Region-specific N_d - N_a parameterizations are necessary for representing BB aerosol-cloud interactions over the remote SEA. A robust inverse N_d - R_e relationship was observed, regardless of BB influence. By coupling backward simulations with satellite retrievals, this study indicates that FT-to-MBL entrainment of African BB aerosols over the SEA occurs several days before arrival at Ascension Island, predominantly west of 0° E for examined cases. These findings provide unique observational constraints for representing aerosol-cloud interactions and vertical transport of African BB aerosols in climate models, offering improved assessments of African BB impacts over the SEA.

1 Introduction

Every year from July to October, seasonal wildfires across central and southern Africa account for about one-third of global carbon emissions from biomass burning (BB) (Roberts et al., 2009). The emitted smoke is frequently transported westward over the southeast Atlantic (SEA), where it often resides in the free troposphere (FT) and may entrain into the marine boundary layer (MBL) during its subsiding transport (Painemal et al., 2014; Adebiyi and Zuidema, 2016; Das et al., 2017). These transported BB aerosols exert complex radiative effects by absorbing and reflecting solar radiations, and by interacting with one of the world's largest semi-permanent stratocumulus cloud decks over the SEA. The SEA region is climatically important, due to the significant net cooling effect of these stratocumulus clouds through their strong reflection of solar radiation but small positive longwave radiative effect (Wood, 2012)

Most previous studies have focused on assessing the direct effect (solar absorption and scattering by BB aerosols) and the semi-direct effect (cloud adjustments in response to the direct radiative effect) of transported African BB aerosols. It is reported that these transported BB aerosols can induce a considerable warming direct radiative effect over the SEA during the fire season (monthly mean = $\sim 15\text{--}35 \text{ W m}^{-2}$) (De Graaf et al., 2014; Zuidema et al., 2018). The semi-direct effect can either reduce or enhance cloud coverage depending on the relative vertical locations of BB aerosols and cloud layers (Boucher et al., 2013). BB aerosols above clouds can strengthen the cloud-top temperature inversion and stabilize the low atmosphere, which weakens entrainment and preserves cloud cover over the SEA (Deaconu et al., 2019; Herbert et al., 2020). Conversely, in the smoky MBL, BB aerosol heating induced by absorbing solar radiation tends to reduce the humidity below the cloud, thereby decreasing the cloud cover (Zhang and Zuidema, 2019). Modeling studies suggest an overall negative semi-direct effect of BB aerosols over the SEA during the fire season (Wilcox et al., 2012; Gordon et al., 2018).

In contrast, the effect of BB aerosols on clouds by acting as cloud condensation nuclei (CCN) (known as the indirect effect) has received less attention compared with their direct and semi-direct effects over the SEA. Contrary to earlier assumptions that BB aerosols are well separated from the cloud deck, recent observations have demonstrated that BB aerosols can be in direct contact with the underlying clouds over the SEA (Painemal et al., 2014; Gupta et al., 2021). Recent modeling studies show that BB aerosols can increase cloud droplet number concentrations (N_d) and reduce droplet sizes, which may enhance cloud lifetime and coverage by reducing drizzle formation (Gordon et al., 2018; Lu et al., 2018). The overall indirect effect of BB aerosols is suggested to enhance the cloud brightness over the SEA and cause a strong cooling effect. Notably, some studies indicate that the indirect effect may dominate the overall aerosol radiative forcing of BB aerosols over the SEA (Lu et al., 2018). Therefore, accurate quantification of cloud responses to BB aerosols is vital for assessments of BB aerosol-cloud-radiation interactions over the SEA.

However, quantifying the cloud response to BB aerosols over the SEA remains uncertain due to observational limitations. Satellite-based observations have been widely used to examine the vertical distance between the smoke layer base and the cloud top, which plays a role in understanding the impact of BB aerosol layers on cloud properties (Costantino and Bréon, 2010, 2013; Painemal et al., 2014). When cloud and aerosol layers nearly overlap, empirical relationships between cloud

65 properties (i.e. N_d or effective radius, R_e) and CCN (or aerosol number concentration) were estimated to quantify the cloud response to BB aerosols. In contrast, clouds and BB aerosols are considered uncorrelated when smoke and clouds are vertically well separated. However, satellites tend to overestimate the base of the smoke layer and imply an unrealistic separation between BB aerosols and underlying clouds. This is because absorbing BB aerosols attenuate the lidar beam, thereby the satellite is not able to detect the lower portions of BB aerosol layers (Rajapakshe et al., 2017). Satellites are also unable to resolve the exact
70 amounts of BB aerosols activated in the clouds, which directly impacts the magnitude of the indirect effect. **More accurate in-situ measurements of aerosol-cloud relationship are needed to quantify the response of clouds to BB aerosols in this region.**

Another uncertainty in assessing BB aerosol-cloud interactions over the SEA arises from model biases in simulating the vertical distribution of transported BB aerosols. **Many simulations using Earth System models show that smoke layers descend too rapidly as they advect away from the African coast, thus the smoke layer heights were underestimated over the SEA (i.e. Das et al., 2017; Gordon et al., 2018; Shinozuka et al., 2020).** The underestimated smoke layer heights would lead to an
75 underestimation of above-cloud aerosols and an overestimation of MBL aerosols on the eastern side of the SEA (Shinozuka et al., 2020). The misrepresentation of aerosol vertical distribution would cause a significant diversity in simulated climate forcing over the SEA, as the radiative effects of aerosols are dependent on their vertical distribution and locations with respect to clouds (Samset et al., 2013). This unrealistically rapid descent of BB aerosol layers may be attributed to an overestimated
80 subsidence in the model-simulated large-scale vertical velocities. Thus, the entrainment process of elevated FT smoke into the MBL is poorly represented in models, despite its importance in determining the MBL BB aerosols in transport regions over the SEA (Johnson et al., 2023). Improved representation of vertical transport and entrainment, potentially aided by higher-resolution meteorological fields, is crucial for simulating BB aerosol-cloud interactions (Das et al., 2017).

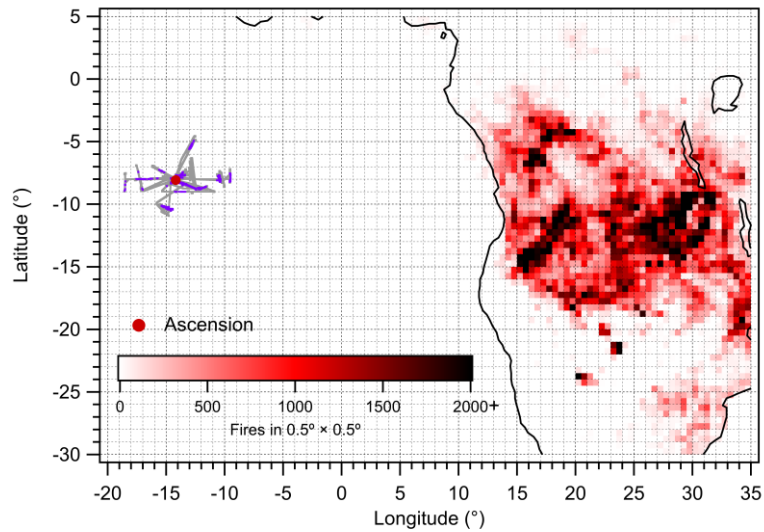
**To address the aforementioned issues, aircraft in-situ measurements are essential to provide unique constraints on the vertical distribution of transported African BB aerosols over the SEA for climate models. Aircraft observations with continuous vertical sampling are also the most reliable source for accurately characterizing the correlations between aerosols and clouds. This study presents airborne observations of BB aerosols and clouds collected during the CLARIFY campaign (Cloud-Aerosol-Radiation Interactions and Forcing; August–September 2017), which was based around Ascension Island (7.96° S, 14.35° W) (Haywood et al., 2021). The CLARIFY campaign addresses a key observational gap over the remote SEA, contributing to a
90 more comprehensive understanding of BB aerosol vertical distributions and interactions with clouds across the SEA. When integrated with complementary campaigns such as ORACLES (ObseRvations of Aerosols above CLouds and their intEractionS), which primarily focused on regions westward of the African continent and eastward of 0° E (Redemann et al., 2021), the CLARIFY campaign provides an integrated wide-scale assessment of BB aerosol transport and its impact on cloud microphysics over the SEA. Observations from these collaborative projects have indicated that the physicochemical properties of BB aerosols continuously evolve during weeklong transport, influenced by aging, cloud processing, and MBL environments,
95 implying different CCN activity after long-range transport (Wu et al., 2020; Dobracki et al., 2023). Additionally, CLARIFY observations have shown that there is often a complex vertical structure of BB aerosol relative to cloud layers, as well as vertical variability of aerosol chemical composition and size distributions (Wu et al., 2020; Haywood et al., 2021). These**

complexities underscore the significance of investigating the mechanisms by which long-range transported BB aerosols modulate cloud microphysics over the remote SEA.

In this study, we first present the pollution conditions over Ascension Island and trace the origins of air masses using backward-dispersion simulations (Sect. 3.1). We then characterize the vertical profiles of thermodynamic variables, aerosol properties, and cloud microphysics over Ascension Island (Sect. 3.2). While the region is influenced by the transition from stratocumulus to cumulus clouds, associated with increasing sea surface temperatures (Gordon et al., 2018), this study focuses on assessing the effects of BB aerosols on stratocumulus cloud microphysics (Sect. 3.3). More details of cloud types (stratocumulus or cumulus clouds) collected during the CLARIFY are provided in Sect. 3.2. Finally, we integrate air parcel analysis with satellite observations (Spinning Enhanced Visible and Infrared Imager, SEVIRI), to identify the efficient entrainment regions where FT air parcels from Africa are likely to enter the MBL over the SEA and to demonstrate their impacts on cloud properties along transport (Sect.3.4).

2 Methodology

The CLARIFY campaign was conducted using the UK FAAM (Facility for Airborne Atmospheric Measurements) BAe-146 Atmospheric Research Aircraft and characterized the aerosol-cloud system centered around Ascension Island (7.96° S, 14.35° W) in the remote SEA. A total of 28 scientific flights (designated flight labels from C028 to C055) took place between August 16 and September 7, 2017. The aircraft was equipped with a range of instruments to derive aerosol and cloud properties, as well as meteorological variables (e.g. ambient temperature (T, K), potential temperature (θ , K), and total water mixing ratio (q_t , g kg⁻¹, the total mass of vapor, liquid, and ice forms of water per unit mass of dry air)). Figure S1 shows the typical flight patterns on example flights. Each flight consisted of a series of straight and level runs at varying altitudes and vertical profiles to characterize aerosols. On cloudy days, saw-tooth and stepped profiles were made to characterize cloud microphysics from cloud base to top. Tracks of flights used in this study are shown in Fig. 1. Transit flights (C040-C041, predominately at high altitudes) and flights with mainly limited cloud samplings are not included in this study. Specific events (pocket of open cell, C052-C054) have been characterized in Abel et al. (2020), which are also excluded from this study. The main aerosol and cloud measurements used in this study are briefly described below. A more detailed overview of the project is provided by Haywood et al. (2021).



125 **Figure 1. Tracks of the CLARIFY flights used in this study, cloud sampling periods during tracks are highlighted in purple colour. Fire density maps are counted over the African continent, showing 2° 2° bins of the number of MODIS-detected fire during the campaign period.**

2.1 Aerosol and cloud measurements

Aerosol size distributions were measured at 1-Hz resolution via two wing-mounted Passive Cavity Aerosol Spectrometer Probes, which resolved number concentrations in 30 diameter bins between 0.1 and 3 μm . The two PCASPs were size-calibrated using di-ethyl-hexylsebacate (DEHS) and polystyrene latex spheres (PSL) with known size and refractive index (Rosenberg et al., 2012). A refractive index of $1.54 - 0.027i$ was assumed for ambient aerosol to determine the bin sizes by using Mie scattering theory. The refractive index was obtained by the method reported by Peers et al. (2019), where the aerosol model is tuned using the refractive index to best represent the PCASP measurements. Aerosol number concentration in the size range of 0.1 – 3 μm (N_a) was obtained by integrating the PCASP distribution. The lognormal fitted count median diameter (CMD) was also calculated from the PCASP aerosol size distribution. A model 3786-LP water-filled condensation particle counter (CPC, TSI Inc.) on board can detect aerosol particles in a size range of 3 nm to 1 μm at 1-Hz resolution (Hering et al., 2005), and it can provide total submicron aerosol number concentration ($CN_{3, > 3 \text{ nm}}$) at an accuracy of $\pm 12\%$. The CPC used in this study is water-based and may undercount hydrophobic particles, particularly soot-rich BB aerosols (Keller et al., 2013). However, most of the sampled BB plumes have undergone week-long transport. Aging processes may have reduced the hydrophobicity of BB organics due to their oxidation and internal mixing with secondary species such as sulfates (Pósfai et al., 2003). Therefore, undercounting is expected to be minimal for the conditions of this study. The CCN number concentrations were measured at 1-Hz resolution by a Droplet Measurement Technologies (DMT) dual-column CCN counter (Roberts and Nenes, 2005). The CCN data was analyzed when measured at a supersaturation (SS) of $\sim 0.2\%$. The relative uncertainty in the supersaturation measurements is $\pm 10\%$, and the uncertainty associated with the flow calibration and counting efficiency of

the particle counter is typically ~6 % (Trembath, 2013). These aerosol instruments covered the typical size range of BB aerosols, primarily within the submicron range.

Cloud droplet number size distributions were measured at 1-Hz resolution using a Cloud Droplet Probe (CDP). The CDP is a forward-scattering optical particle counter which can detect particles in 30 diameter bins between 2 and 52 μm . The operation and calibration of the CDP follow procedures described in Lance et al. (2010). A 10-point glass bead calibration spanning the instrument's detection range was performed before each flight day. The nominal bead size was corrected for the differences between the refractive indices of glass and water, and the water-corrected size was used to calibrate the instrument's sizing response (Barrett et al., 2022). From the CDP's cloud droplet spectrum, three key cloud microphysical parameters were derived: N_d , R_e , and liquid water content (LWC). The N_d was calculated by integrating the droplet size distribution, as described by Eq. (1).

$$N_d = \int n(r) dr \approx \sum_1^m n(r_i) \quad (1)$$

where $n(r_i)$ is the droplet number concentration in a particular size bin, and r_i is the middle radius for each of the size bins.

The R_e was calculated based on the Eq. (2).

$$R_e = \frac{\int r^3 n(r) dr}{\int r^2 n(r) dr} \approx \frac{\sum_1^m r_i^3 n(r_i)}{\sum_1^m r_i^2 n(r_i)} \quad (2)$$

The LWC was calculated based on the Eq. (3), assuming spherical droplets and a liquid water density (ρ_{water}) of 1 g cm^{-3} .

$$\text{LWC} = \frac{4\pi}{3} \rho_{\text{water}} \int r^3 n(r) dr \approx \frac{4\pi}{3} \rho_{\text{water}} \sum_1^m r_i^3 n(r_i) \quad (3)$$

An LWC value over 0.01 g m^{-3} for 1 Hz measurements was used to define the low threshold for the presence of cloud. To eliminate the inclusion of optically thin clouds, a threshold of $N_d > 5 \text{ cm}^{-3}$ and bulk $\text{LWC} > 0.02 \text{ g m}^{-3}$ was used to perform statistical cloud sample analysis. The duration of efficient cloud sampling in each flight using these criteria is listed in Table S1, with corresponding cloud sampling periods highlighted in Fig. 1. Of relevance to this work are the saw-tooths and stepped profiles to characterize cloud regions. A vertical profile that sampled a continuous cloud layer was used to determine the cloud base height as the lowest altitude with $N_d > 5 \text{ cm}^{-3}$ and bulk $\text{LWC} > 0.02 \text{ g m}^{-3}$, and the cloud-top height was identified as the highest altitude satisfying these criteria. The liquid water path (LWP) was also integrated from each profile sampled.

Aerosol concentrations measured aboard the FAAM BAe-146 were corrected for ram air compression at inlet speed using the volumetric correction as described in previous studies (Trembath et al., 2013). Airborne aerosol and cloud measurements reported here were corrected to the standard temperature and pressure condition (273.15 K and 1013.25 hPa). It is noted that the PCASP, CPC and CCN measurements inside clouds are unreliable and thus in-cloud ($\text{LWC} > 0.01 \text{ g m}^{-3}$) data was removed.

We also obtained the above-cloud aerosol optical thickness (AOT), cloud optical thickness (COT), R_e and cloud-top height (CTH) across the SEA region ($20^\circ \text{ W} - 15^\circ \text{ W}$; $30^\circ \text{ S} - 0^\circ \text{ N}$), which was also retrieved from the SEVIRI sensor aboard the geostationary Meteosat-10 satellite, following Peers et al., (2019, 2021). The N_d was calculated assuming an adiabatic-like vertical stratification (Painemal et al., 2012):

$$N_d = 1.4067 \times 10^{-6} \left[\text{cm}^{-\frac{1}{2}} \right] \text{COT}^{\frac{1}{2}} R_e^{-\frac{5}{2}} \quad (4)$$

It is noted that only data from liquid clouds in the MBL (cloud top height less than 3000 m) are included in this analysis. SEVIRI aerosol and cloud retrievals are available for the period of August 16–31, 2017. Observations were collected daily between 06:12 and 17:57 UTC, with a time resolution of 15 min.

2.2 NAME description

To investigate the sources and transport pathways of airmasses arriving at Ascension Island, we performed backward-dispersion simulations using the UK Met Office’s Numerical Atmospheric Modelling Environment (NAME) (Jones et al., 2007). NAME was chosen as an appropriate model for this study because it uses high-resolution meteorological fields (~ 10 km $\times$ 10 km). It can predict transport over distances ranging from a few kilometers to the whole globe, and it has been used successfully in similar studies (i.e. Panagi et al., 2020).

Backward-dispersion simulations were performed at 3-hour (h) intervals throughout the campaign. In each simulation, a certain amount of hypothetical tracer particles (hereafter referred to as “air parcels”) ($\sim 2.7 \times 10^{-7}$ g m $^{-3}$) were released from a $2^\circ \times 2^\circ$ grid box centered around the Ascension Island observation site (14.35°W, 7.96°S). To distinguish source origins of MBL and FT airmasses over Ascension Island, tracer particles were released within the MBL (341 ± 300 m) and FT (2.5–4 km) altitude ranges separately. The transport of released tracer particles was tracked backward over the past 7-days in the NAME, driven by the three-dimensional gridded (3D) meteorological fields from the UK Met Office’s Unified Model (Brown et al., 2012). The meteorological fields have a high resolution of 0.14° longitude by 0.1° latitude and 59 vertical levels up to ~ 29 km. The model output integrated horizontal footprints, representing original air parcels passing through horizontal grids over the past 7 days before arriving at the release area (see results in Sect. 3.1). These footprints indicate horizontal contributions of historical airmasses. The footprints are expressed in units of mass concentration (g m $^{-3}$), based on the predetermined quantity of released air parcels.

We also employed three case studies (see Sect. 3.1 for details of cases), to further investigate vertical distributions and the transport of original air parcels for MBL airmasses over Ascension Island. For these cases, the model output instantaneous 3D air parcel footprints every 3h during the 7-days backward dispersion simulations. The 3D meteorological parameters, e.g. θ , relative humidity, and water mixing ratios, were also output for estimating the capping inversion height (z_i). Over the SEA, there is typically a strong thermodynamic capping inversion, that inhibits turbulent mixing between the BL and overlying FT air (Abel et al., 2020). However, the BL depth in NAME is determined to be the height of the surface mixed layer to which BL turbulence extends (Lock et al., 2020). When getting to Ascension Island, the z_i is typically higher than the BL depth employed in NAME, as the upper MBL tends to be decoupled from the surface mixed layer (Abel et al., 2020). Here, we estimated the z_i over the SEA, using the outputs of 3D meteorological parameters. The z_i over the SEA was quantified as the height at which the vertical gradient of liquid water potential temperature (θ_l) is the largest, and there is also a steep decrease in humidity (Jones et al., 2011). The θ_l was estimated following Eq. (5).

$$\theta_l = \theta - \frac{L}{c_p} q_l \quad (5)$$

210 Where θ is the potential temperature, L is the latent heat of vaporization for water ($2.5 \times 10^6 \text{ J kg}^{-1}$), c_p is the specific heat of dry air at constant pressure ($1005 \text{ J kg}^{-1} \text{ K}^{-1}$), q_l is the liquid water mixing ratio (g kg^{-1}).

For the three case studies, we analyzed the instantaneous 3D air parcel footprints in two approaches (see results in Sect. 3.4). First, vertical layers of instantaneous 3D footprints were integrated across the horizontal domain to derive column-integrated horizontal footprints. Instead of using the BL depth in NAME, the estimated z_i was employed to divide the column-integrated horizontal footprints into the FT and BL separately, distinguishing original air parcels as either from the FT or BL. This enabled an assessment of the relative contributions of original air parcels from FT and BL sources. Second, horizontal grids were integrated within each vertical layer for instantaneous 3D footprints, to derive vertical distributions of original air parcels as a function of time (at 3h intervals) along the backward simulations. This provides insights into the vertical transport and entrainment pathways of African BB plumes.

220 **3 Results and discussions**

3.1 BB pollution conditions and source analysis

Figure 2a shows complex vertical distributions of aerosol number concentrations (N_a , $0.1 - 3 \mu\text{m}$) from PCASP measurements for the CLARIFY flights used in this study, alongside the estimated z_i . The z_i estimates were derived from airborne measurements of θ and q_l , using Eq. (5). Following established criteria used in previous CLARIFY studies, the presence of BB-pollution was defined using thresholds of carbon monoxide above 83 ppb and black carbon above $0.1 \mu\text{g m}^{-3}$ (Wu et al., 2020; Haywood et al., 2021). Table S1 summarizes the BB pollution conditions in the FT and MBL respectively, for the 17 flights used in this study. During the CLARIFY campaign, BB pollution events were observed at different times to be solely in the MBL, or solely in the FT, or in both layers. Concurrent surface observations on Ascension Island (Zuidema et al., 2018) presented the same trend of MBL BB pollution as observed during the CLARIFY campaign. Based on the BB pollution conditions, the CLARIFY campaign can be divided into three periods: Period 1 from August 16 to 20, BB aerosols were observed to exist predominantly in the MBL, and the FT was mainly clean; Period 2 from August 21 to 25, the MBL became considerably clean, while the BB pollution existed predominantly in the FT; Period 3 from August 26 to 31, the MBL was BB-impacted again, and the BB aerosols were observed in both the MBL and FT.

NAME backward-dispersion fields are used to represent the horizontal footprint of air parcels transported over the past 7 days before arriving at the sampling area over Ascension Island. Figure S2 provides three examples of backward-dispersion fields (cases 1 to 3), with release times indicated by red markers in Fig. 2a. These cases illustrate distinct and representative transport pathways of original air parcels for three periods. Case 1 released tracers near the end of Period 1 (12:00 UTC, August 18). Cases 2 and 3 released tracers at the beginning of Periods 2 (12:00 UTC, August 21) and 3 (12:00 UTC, August 26) respectively, coinciding with shifts in MBL pollution conditions. FT dispersion simulations (Figs. S2a–c) indicate that most FT air parcels over Ascension Island originated from westerly flow across the SEA and African continent. Polluted FT cases (Figs. S2b, c) show a substantially greater influence from African continental airmasses compared to the clean FT case (Fig.

S2a). MBL dispersion simulations (Figs. S2d–f) suggest that MBL air parcels over Ascension Island mainly arose from clean oceanic flow that transported from the southeast to northwest over the SEA. Polluted MBL cases (Figs. S2d, f) also show contributions of westerly flow originated from African continent, while negligible continental influence in the clean MBL case (Fig. S2e). We calculated fractional contributions of air mass from African continent ($20^{\circ}\text{S} - 5^{\circ}\text{N}$, $9^{\circ}\text{W} - 35^{\circ}\text{E}$) to Ascension Island area throughout the campaign period, differentiating between FT (Fig. 2b) and MBL (Fig. 2c) simulations. These contributions closely tracked the influence of BB plumes from African wildfire regions (Fig. 1, fire maps) throughout the campaign period. The temporal evolution of continental contributions aligns well with the observed FT and MBL pollutions over Ascension Island. NAME simulations evidence that the FT and MBL pollutions observed during the campaign were attributable to long-range transport of African BB plumes.

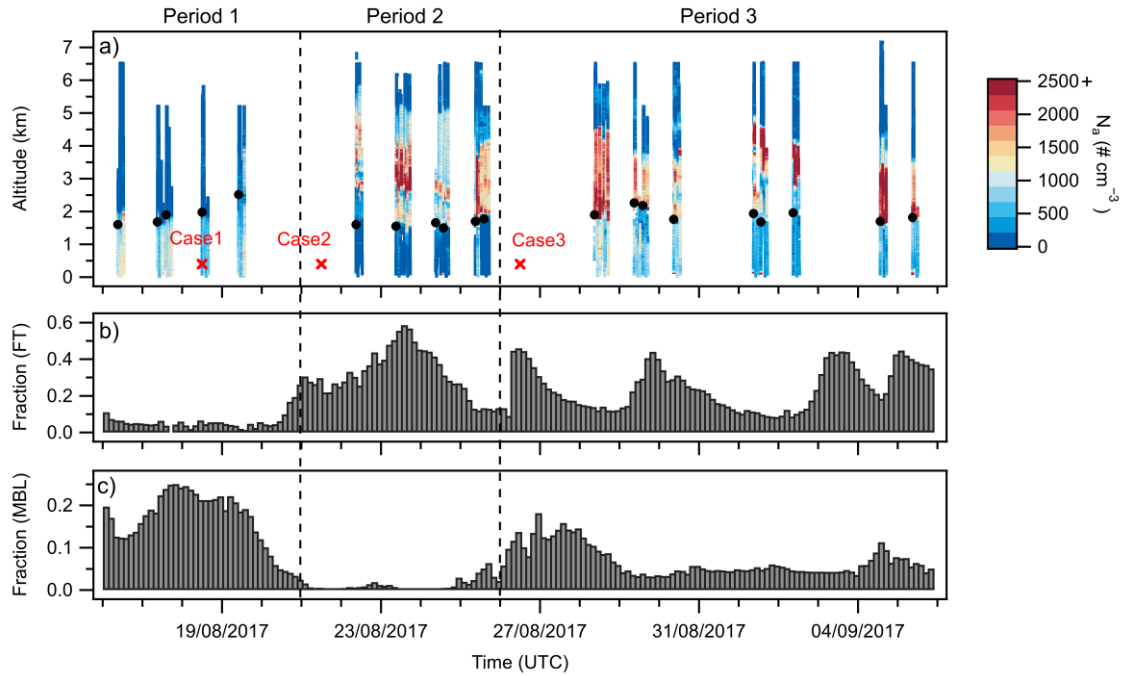


Figure 2. a) Vertical distributions of aerosol number concentrations (N_a , $0.1 - 3 \mu\text{m}$) from PCASP measurements for the CLARIFY flights (C028 to C051), alongside the estimated z_i . The red markers represent the release times for three NAME example cases. **b, c)** Fractional contributions of air mass from African continent ($20^{\circ}\text{S} - 5^{\circ}\text{N}$, $9^{\circ}\text{W} - 35^{\circ}\text{E}$) to Ascension Island area throughout the campaign period, differentiating between **b)** FT and **c)** MBL simulations.

3.2 Vertical profiles of thermodynamics, aerosol and cloud

Here, we characterize the vertical profiles of thermodynamics, aerosol and cloud properties during the campaign. Figure 3 shows the vertical profiles of θ_i (Fig. 3a) and q_i (Fig. 3b) for the flights used in this study, which are colored by three periods. It is suggested that the z_i is an important parameter for understanding the BL structure. The height scale was normalized by the z_i , to exclude the variation of z_i and compare the MBL structure on different days. During CLARIFY, we observed similar vertical structures of thermodynamic parameters between flights. The MBL generally showed a well-mixed layer with near-

constant θ_1 and q_1 between the sea surface and the lifting condensation level (LCL, the height at which the relative humidity of an air parcel reaches 100% when lifted adiabatically), which is at an altitude of ~ 300 to 800 m ($z/z_i = \sim 0.2 - 0.5$). Above the LCL, there was another layer that extended to the base of the inversion. During 6 flight cases (marked as type-d in Table S1), profiles above the LCL exhibited near-constant θ_1 and q_1 , indicating a fairly well-mixed upper layer and the presence of a decoupled stratocumulus deck. By contrast, in the remaining cases, the upper layer above the LCL was conditionally unstable, suggesting the presence of a stratocumulus-over-cumulus MBL (marked as type-e in Table S1) or a cumulus-capped MBL (marked as type-f in Table S1) structure. The BL decoupling parameters α_q and α_θ , were also calculated from the respective profiles following Eq. (6) and (7) respectively, which represent the relative differences of the q_1 and θ_1 between the surface and the upper part of the BL respectively (Zheng et al., 2011).

$$\alpha_q = \frac{q_t(z_i^-) - q_t(0)}{q_t(z_i^+) - q_t(0)} \quad (6)$$

$$\alpha_\theta = \frac{\theta_1(z_i^-) - \theta_1(0)}{\theta_1(z_i^+) - \theta_1(0)} \quad (7)$$

Where $\theta_1(z_i^+)$ and $q_1(z_i^+)$ are θ_1 and q_1 at the level 50m above the inversion height, $\theta_1(z_i^-)$ and $q_1(z_i^-)$ are at the level 50m below the inversion height, $\theta_1(0)$ and $q_1(0)$ are at the lowest measured level near sea surface. It is suggested that a value of α_q over ~ 0.3 indicates a decoupled BL. The average values of α_q and α_θ in this study were calculated to be 0.35 ± 0.11 and 0.32 ± 0.09 respectively, and the upper part of the MBL was clearly drier than the lower surface layer, suggesting a generally decoupled MBL over Ascension Island area during CLARIFY. The strong inversion layer occurred at the top of the MBL ($0.9 < z/z_i < 1.0$), with the average θ_1 increase of approximately 4–9 K across BL top and a sharp decrease in humidity (q_1). Above the inversion layer, the air was much drier compared to the MBL and is regarded as being in the FT. Figure 3b shows that FT humidity was generally higher on days with the presence of FT BB plumes (Periods 2 and 3, red and black profiles) compared to clean FT cases (Period 1, blue profiles). This pattern is consistent with previous studies, suggesting a positive correlation between BB plume strength and atmospheric water vapor content in the FT over the SEA (Pistone et al., 2021). The covariation between plume strength and humidity is attributed to the mixing between the moist, smoky continental boundary layer and the dry, clean FT. Consequently, humid BB plumes above the boundary layer can be advected to the SEA, enhancing FT humidity in the presence of transported smoke plumes (Pistone et al., 2021).

In this study, vertical profiles of aerosol properties (CMD, CCN/CN₃ and CCN/N_a ratios) were summarized under clean (red line and shades in Figs. 3c-e) and BB-impacted (black line and shades in Figs. 3c-e) conditions separately. MBL analysis included both the clean and BB-impacted conditions. However, due to background aerosols in the clean FT during Period 1 being mainly below instrumental detection limits, the FT analysis (CMD, CCN/CN₃, and CCN/N_a ratios) was restricted to BB-impacted conditions. Figure. 3c shows that the CMD of aerosols in the BB-impacted MBL was 10 – 15 % lower than in the FT BB plumes, which is probably attributable to some processes occurring in the MBL such as mixing with marine emitted small particles (i.e. marine sulfate) and aerosol removal by drizzle (Wu et al., 2020). Submicron aerosol chemical compositions under BB-impacted condition during CLARIFY were previously characterized by Wu et al. (2020), as show in Fig. S3. BB-impacted MBL had higher average sulfate mass fractions ($\sim 26\%$) than the FT BB plumes ($\sim 11\%$) around Ascension Island,

295 due to the mixing with marine sulfate. The corresponding profiles of CCN/CN₃ (Fig. 3d) and CCN/N_a (Fig. 3e) ratios under the BB-impacted condition were fairly constant with altitude, suggesting that vertical variations of aerosol size and chemical composition did not apparently affect the BB aerosol ability to act as CCN. In this study, the sampled African BB aerosols have undergone week-long transport and are defined as highly aged African BB aerosols (Wu et al., 2020). Compared to freshly emitted African BB aerosols, these highly aged BB aerosols exhibited larger particle sizes, enhanced fractions of inorganics and the loss of organics after atmospheric aging such as chemical oxidation and cloud processing (Wu et al., 2020; Dobracki et al., 2023). Due to atmospheric aging, the CCN activation fractions of these highly aged African BB aerosols are generally higher than fresher BB aerosols sampled over the African continent (CCN/N_a = 0.68, at SS = 0.3%) (Ross et al., 2003), but are comparable to transatlantic African BB aerosols observed over Amazon area during the dry season (the average ratio of CCN to aerosol concentrations in a range of 20 nm to 1 μm is 0.83 ± 0.06 at SS = 0.5%) (Holanda et al., 2020).

305 In the clean MBL, Wu et al. (2020) presented an average aerosol size distribution, with a CMD of ~30 nm in the Aitken mode and a CMD of ~160 nm in the accumulation mode (red marker in Fig. 3c). The aerosols in the clean MBL were smaller than in the BB-impacted MBL during CLARIFY. The average ratio of N_a/CN₃ in the clean MBL was calculated to be 0.40 ± 0.26 , also suggesting that the majority of submicron aerosols detected by the CPC (3 nm to 1 μm) in the clean MBL were Aitken mode particles (< 0.1 μm). The submicron aerosols from marine emissions in the clean MBL were previously reported to be dominated by sulfate (~ 60%) and organic (~ 24%) (Wu et al., 2020). Although the clean MBL aerosols contained enhanced fractions of sulfate with high hygroscopicity, the average CCN/N_a (0.83 ± 0.20) was close to those in the BB-impacted MBL (0.82 ± 0.17), suggesting that the CCN activated fractions of accumulation-mode aerosols were similar between the clean and BB-impacted MBL conditions. However, average CCN/CN₃ in the clean MBL (0.38 ± 0.18) was much smaller than in the BB-impacted MBL (0.76 ± 0.10), as the dominant Aitken mode particles (< 0.1 μm) in the clean MBL are not sufficiently large to act as CCN (Dusek et al., 2006). Under both clean and BB-impacted conditions, the CCN/CN₃ ratios remained consistently below unity, whereas CCN/N_a occasionally exceeded unity, implying that a subset of Aitken mode particles (< 0.1 μm) contributed to CCN.

315 Figs. 3f-h show the vertical profiles of cloud properties (N_d, LWC and R_e), using a height scale normalized by z_i. In this study, we focus on sawtooth profiles that contain a continuous cloud layer from the cloud base to the cloud top. In Flight C032 with cumulus clouds, stepped profiles and SLRs were carried out. The average vertical profiles of cloud properties from flight C032 are also provided in Figs. 3f-h (grey dashed lines). It should be noted that there were broken cumulus clouds sampled below the continuous stratocumulus cloud layer on some days, which are not included in this study. In total, there are 31 and 12 sufficient profiles sampled in the BB-impacted and clean MBL respectively. The profiles of N_d (Fig. 3f) remain fairly constant, with few cases in the BB-impacted MBL (i.e. profiles in flights C029 and C038) presenting an increase near cloud top. The profiles of LWC (Fig. 3g) generally increased with height from the cloud base and then decreased near the top of the MBL due to droplet evaporation. The profiles of in-cloud R_e (N_d > 5 cm⁻³ and bulk LWC > 0.02 g m⁻³) generally presented enhanced values from the cloud base to the top (Fig. 3h). SEVIRI-retrieved N_d and R_e were averaged along the cloud sampling tracks of the CLARIFY flights. For comparison, cloud-top-layer (uppermost 50 m) N_d and R_e were summarized from CDP

measurements. As shown in Fig. 4, SEVIRI retrievals and CDP measurements show overall good agreement around Ascension Island. The agreement for R_e is weaker in low- N_d cases, since the CDP detects cloud droplets below 50 μm and larger droplets beyond its detection range may lead to discrepancies (Abel et al., 2020). In-situ and SEVIRI observations consistently indicate higher N_d and smaller R_e values in the BB-impacted MBL compared with the clean MBL.

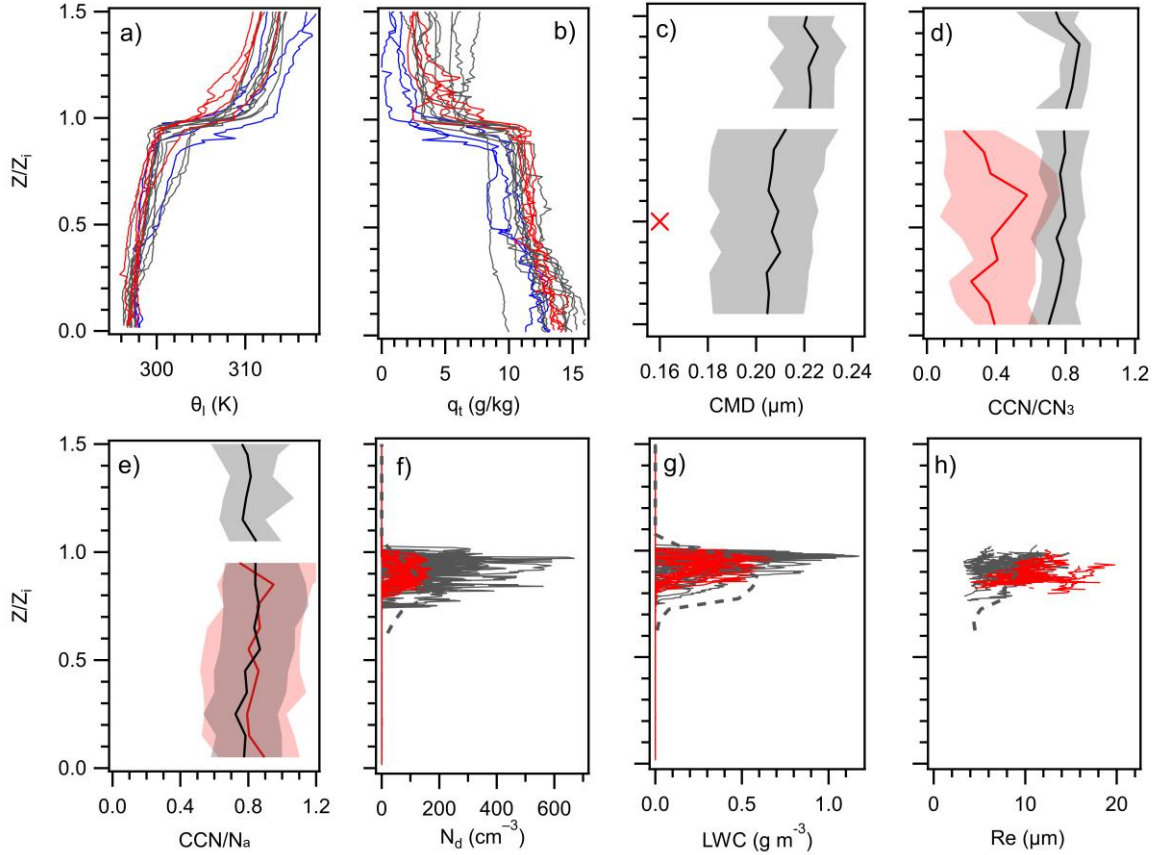


Figure 3. a, b) Average vertical profiles of a) liquid water potential temperature (θ_l , K), b) total water mixing ratio (q_t) for each flight used in this study. Blue, red and black lines represent measurements from Periods 1, 2 and 3 respectively. c-e) Summarized profiles of c) aerosol count median diameter derived from the PCASP (CMD, μm) and d,e) the ratio of CCN ($\sim 0.2\%$) to condensation nuclei ($D_a > 3$ nm from the CPC) (CCN/CN₃) and accumulation aerosol concentration ($D_a > 0.1$ μm from the PCASP) (CCN/N_a) under polluted (black) and clean (red) conditions. Solid lines and shades represent median values and range from 10 to 90 percentiles. The red cross marker in Fig. 3c represents the average CMD in accumulation mode from measurements within the clean MBL. f-h) Vertical profiles of 1-hz, f) cloud droplet number concentration (N_d , $\# \text{ cm}^{-3}$), g) liquid water content (LWC, g m^{-3}) and h) cloud effective radius (R_e , μm) in sampled continuous cloud layers. Red lines represent cloud measurements in the clean MBL, and grey lines represent cloud measurements in the BB-impacted MBL. It is noted that average vertical profiles of cloud properties from flight C032 are also provided in Figs. 3f-h (grey dashed lines). The y-axis uses a height scale normalized by inversion height (z_i).

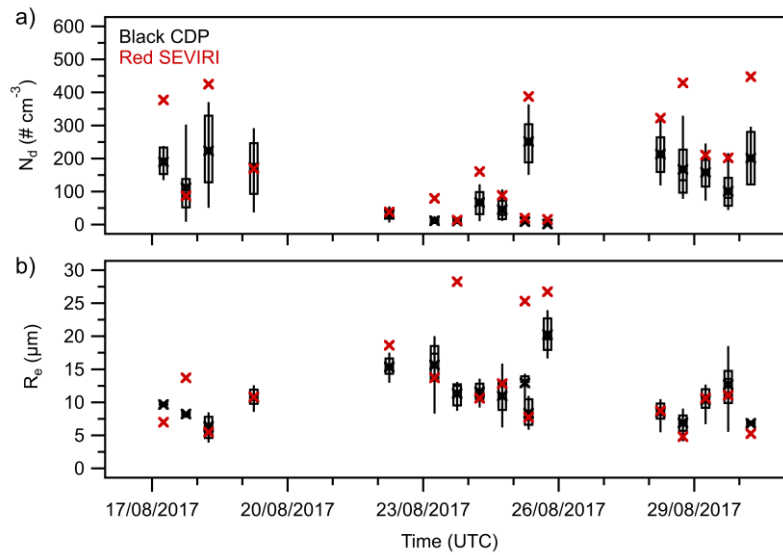


Figure 4. Average SEVIRI-retrieved a) N_d and b) R_e along the cloud sampling tracks of the CLARIFY flights, and comparison the corresponding cloud-top-layer N_d and R_e from CDP measurements. Red markers show average values from SEVIRI. Black markers, boxes and whiskers represent average, 10, 25, 50, 75 and 90 percentiles from CDP measurements.

3.3 Aerosol-cloud interaction around Ascension

In this section, we examine the relationship between aerosols and cloud microphysics over Ascension Island in the remote SEA. Here, we use N_a as the metric instead of CCN concentrations to establish the aerosol impact on cloud microphysics, since CCN concentrations depend on cloud supersaturations. In this study, we focus on aerosols in the size above $0.1 \mu\text{m}$ that behaved as the dominant CCN (as discussed in Sect. 3.2) and impacted cloud microphysics.

3.3.1 Aerosol- N_d relationship

Previous studies suggest two main mechanisms for the activation of aerosols into cloud droplets that typically occur: 1) via updrafts carrying aerosols to the cloud base (cloud base activation) or 2) via entrainment through turbulent mixing at the cloud top (cloud-top activation) (e.g. Hoffmann et al., 2015; Slawinska et al., 2012). Here, we analyzed the cloud-layer mean N_d , N_a below the cloud (sub- N_a) and above the cloud (above- N_a), to examine the relative importance for both activation mechanisms. In each profile with a continuous cloud layer, we averaged the N_a within 200m above the cloud top to obtain the above- N_a , and within 200m below the cloud base to obtain the sub- N_a . Figures 5a and 5b show the relationship between cloud-layer mean N_d and N_a concentrations for all the analyzed profiles, in terms of sub- N_a and above- N_a respectively. The N_d exhibited a strongly positive correlation with sub- N_a , with a Pearson correlation coefficient (r) of 0.89 ($p < 0.01$, statistically significant) for all analyzed profiles. The positive correlation between N_d and sub- N_a was stronger ($r = 0.93$, $p < 0.01$) in the clean or moderately BB-impacted clouds (sub- $N_a < 700 \text{ cm}^{-3}$), while a weaker correlation was observed in more polluted clouds (sub- $N_a > 700 \text{ cm}^{-3}$). In contrast, the overall correlation between N_d and above- N_a was weak but statistically significant

365 when considering all analyzed profiles ($r = 0.41$, $p < 0.01$). For polluted FT cases (circle markers in Fig. 5b), the N_d exhibited a fairly positive correlation with above- N_a ($r = 0.74$, $p < 0.01$) when above- N_a was $< 2000 \text{ cm}^{-3}$, while the correlation became weaker when above- N_a was $> 2000 \text{ cm}^{-3}$. These results suggest that the influence of above- N_a on cloud properties was weaker than sub- N_a at the place/time of observation, indicating that cloud base activation of N_a played a considerably greater role as compared to cloud-top activation of N_a through entrainment. Particularly, the cloud base activation of N_a could play a dominant role, on observation days characterized by a clean FT but BB impacted MBL (triangle markers in Fig. 5).

370 Previous studies similarly reported a higher positive correlation between N_d and below-cloud aerosols than above-cloud aerosols, e.g. over the SEA BB-impacted region near African western offshore (Diamond et al., 2018) and over the Pacific BB-impacted region (Mardi et al., 2019). The disconnectedness between N_d and above- N_a is partly due to the gap between FT aerosol layers and the cloud layer (Mardi et al., 2019). The bottom of the BB layer in the FT was defined as the lowest altitude of the plume where observed N_a exceeded 500 cm^{-3} (Gupta et al., 2021). The distance from the top of cloud layers in the MBL to the bottom of BB layers in the FT is referred to as Cloud Top to Aerosol Base (CTtoAB). Figure 6a shows the relationship between sub- N_a and the CTtoAB. An overall negative correlation was observed between sub- N_a and CTtoAB. Separated profiles (CTtoAB $> 100 \text{ m}$) all occurred within the clean MBL, with a low range of sub- N_a ($56 - 125 \text{ cm}^{-3}$) (Fig. 6a). The large gap observed on these days separated the elevated BB layers from the cloud layer and hence weakened the effects of above- N_a on cloud microphysical properties. Figure 6b shows the relationship between sub- N_a and above- N_a for all analyzed profiles. When contact profiles (CTtoAB $< 100 \text{ m}$) occurred within the clean MBL, they presented higher sub- N_a ($172 - 315 \text{ cm}^{-3}$) than separated profiles ($56 - 125 \text{ cm}^{-3}$) within the clean MBL. Example vertical structures of N_a , LWC, and θ_i in separate and contact profiles within the clean MBL are shown in Fig. S4a and S4b respectively. In the separate profile (Fig. S4a), N_a remained low and fairly constant throughout the entire MBL column, whereas in the contact profile (Fig. S4b), N_a was slightly elevated below the cloud compared to the surface well-mixed layer. These results suggest that the proximity of FT BB plumes to the cloud top allows them to impact the sub-cloud N_a budget at the place of observation. Similar results have also been reported in other SEA studies near the western offshore of Africa (Gupta et al., 2021, 2022). The interconnectedness between sub-cloud and above-cloud regions for these contact profiles is attributed to the entrainment process of BB aerosols from the cloud top to below the cloud at the place of observation, as suggested in previous studies (Diamond et al., 2018). Most contact profiles (34 out of 42) occurred within the BB-impacted MBL, which had a broad range of sub- N_a ($212 - 1183 \text{ cm}^{-3}$). For these contact profiles within the BB-impacted MBL, the increase in sub- N_a is not only related to the entrainment process at the place of observation, but also previously entrained African BB aerosols from above-cloud to below-cloud regions over the SEA during long-range transport (Haywood et al., 2021). It should be noted that the entrainment process is time-dependent and relies on factors such as the entrainment rate and the time duration since the aerosol and cloud layer have come into contact (Diamond et al., 2018), which would prevent an instantaneous correlation and consequently lead to the weak correlation between N_d and above- N_a at the place/time of observation. It is also important to highlight observation days that exhibited a clean FT but a BB-impacted MBL (triangle markers in Figs. 5b and 6b). The increase in amounts of MBL BB aerosols on these days would be predominantly due to previously entrained African BB aerosols along the long-range transport. This

suggests that previously entrained BB aerosols during long-range transport are more likely to impact the sub-cloud N_a budget than the entrainment of local FT layers at the place of observation, further resulting in a weak correlation between N_d and above- N_a .

During CLARIFY, the BB-impacted MBL had substantially enhanced sub- N_a ($212 - 1183 \text{ cm}^{-3}$, black markers in Fig. 5a) compared to the clean MBL ($56 - 315 \text{ cm}^{-3}$, red markers in Fig. 5a). The main contributor to the sub- N_a in the clean MBL was local marine emissions. The increase in the sub- N_a is hypothesized to be the entrained BB aerosols from above-cloud to below-cloud regions during long-range transport or/and at the place of observation. More details on the long-range transport and entrainment processes of African BB aerosols are discussed in Sect. 3.4. The enhanced sub- N_a in the BB-impacted MBL led to a substantially enhanced range of N_d values as compared to the clean MBL (Fig. 5a), showing the importance of transported BB aerosols that affect cloud formation in this region. The relationship between N_d and sub- N_a follows a similar pattern in clean or moderately BB-impacted clouds (sub- $N_a < 700 \text{ cm}^{-3}$). The weaker correlation between N_d and sub- N_a under more polluted conditions (sub- $N_a > 700 \text{ cm}^{-3}$) may be partly due to the variability in the MBL updraft velocity. ORACLES Observations near African west coast suggests that in a more polluted MBL, defined as the velocity-limited regime, cloud droplet formation is more sensitive to fluctuations in updraft velocity than to fluctuations in aerosol concentration (Kacarab et al., 2020). When a high concentration of aerosols or CCN is present, there is a strong competition between them for water vapor. The increase in updraft velocity may augment supersaturation levels, thereby playing a more important role in driving droplet formation than the increase in aerosol concentration under polluted conditions (Kacarab et al., 2020). Additionally, differences in the strength of cloud-top entrainment processes may contribute to observed bias in the response of N_d to sub- N_a . Adiabatic LWC (aLWC) was calculated for each cloud profile, and the ratio of LWC to aLWC (LWC/aLWC) was averaged over the cloud layer to assess the extent of cloud-top entrainment mixing or/and the presence of drizzle (Gupta et al., 2021). When contact profiles had average LWC/aLWC values close to 1, this suggests near-adiabatic profiles with negligible mixing between cloudy and non-cloudy air. In contrast, other contact profiles had lower average LWC/aLWC values ($0.34 - 0.83$), indicating greater entrainment mixing of aerosols from above-cloud into the cloud layer at the place of observation. Notably, some of the near-adiabatic profiles (blue dashed box highlighted in Fig. 5a) displayed lower droplet activation fractions at similar sub- N_a levels, compared to other profiles. Although cloud-top entrainment of warmer, drier FT air could cause cloud droplet evaporation, our observations suggest that contact profiles with greater entrainment generally promoted additional droplet activation and enhanced N_d , compared to near-adiabatic profiles.

For applications in remote sensing retrievals and climate models, parameterizations of N_d based on N_a have been suggested using various schemes, e.g. linear regression (Hegg et al., 2012) or exponential power law (Twohy et al., 2013) depending on in-situ measurements. We quantified the N_d response to sub- N_a by a frequently used power law relationship ($N_d \sim \alpha N_a^\beta$). The exponent β can be thought of as the fractional increase in N_d in response to the fractional increase in N_a ($> 0.1 \mu\text{m}$) (McComiskey and Feingold, 2008). For all BB-impacted MBL profiles, analysis using bootstrap resampling yielded median values (95% confidence interval) of $\alpha = 3.18$ ($0.56 - 13.44$) and $\beta = 0.71$ ($0.42 - 0.92$). For all clean MBL profiles, the values were $\alpha = 2.17$ ($0.22 - 7.32$) and $\beta = 0.74$ ($0.42 - 1.03$). It should be noted that the data from clean MBL profiles exhibit

small-scale variation, which may limit the robustness of the derived power law relationship. Since droplet activation displayed similar behavior in clean and moderately BB-polluted clouds ($\text{sub-}N_a < 700 \text{ cm}^{-3}$) (Fig. 5a), a strong linear correlation ($r^2 = 0.87$) was observed between $\log(N_d)$ and $\log(\text{sub-}N_a)$. Data in this aerosol-limited droplet activation regime ($\text{sub-}N_a < 700 \text{ cm}^{-3}$), yielded $\alpha = 0.64$ ($0.20 - 1.28$) and $\beta = 0.93$ ($0.79 - 1.04$), which we recommend as representative parameters for application in future studies. Our observations suggest a small difference in the response of N_d to N_a between BB-impacted and clean MBL profiles. This is likely due to their comparable CCN activation abilities of accumulation-mode aerosols under two MBL conditions (as discussed in Sect 3.2). In contrast, a previous study reported higher droplet activation fractions for the cleaner MBL compared to the BB-impacted MBL over the Pacific Ocean (Mardi et al., 2019). The discrepancy likely stems from differences in their droplet activation behaviors of transported BB aerosols between the studies. The β value for BB-impacted MBL cases in this study (0.71 ($0.42 - 0.92$)) is in a higher range than that reported for BB-impacted areas off the California coast of North America (0.26 ($0.15 - 0.42$)) in Mardi et al. (2019). A key factor contributing to the different droplet activation behaviors of transported BB aerosols, may be the variability in aerosol chemical composition and CCN activity, which depends on source combustion conditions and aging process (Wu et al., 2020; Farley et al., 2025). Submicron BB aerosols from western U.S. wildfires have been reported to be dominated by organic ($\sim 90\%$) with minimal inorganic content ($< 2\%$) from near-source to regional scales (0.5 hours to several days) (Farley et al., 2025). In contrast, highly aged African BB aerosols were reported to contain $\sim 35\%$ inorganic mass (Wu et al., 2020). This implies that highly aged African BB aerosols are more hygroscopic, as inorganics typically have higher hygroscopicity than organics on a global scale (Pöhlker et al., 2023). Additionally, transported BB aerosols from western US wildfire presented similar accumulation-mode aerosol size distributions to this study (Laing et al., 2016). Consequently, the CCN activation ability of transported African BB aerosols in this study is broadly higher than that reported for aged BB aerosols from Western/Northern American wildfires ($\text{CCN}/\text{CN} = 0.11 - 0.62$, at $SS = 0.2 - 0.5\%$) (Pratt et al., 2011; Zheng et al., 2020), leading to the observed differences in their N_a - N_d relationship between these studies. Overall, these results highlight regional variability in the sensitivity of droplet formation to aerosol concentration in different BB-impacted environments, emphasizing the need for region-specific parameterizations of N_d based on N_a when simulating regional radiative forcing.

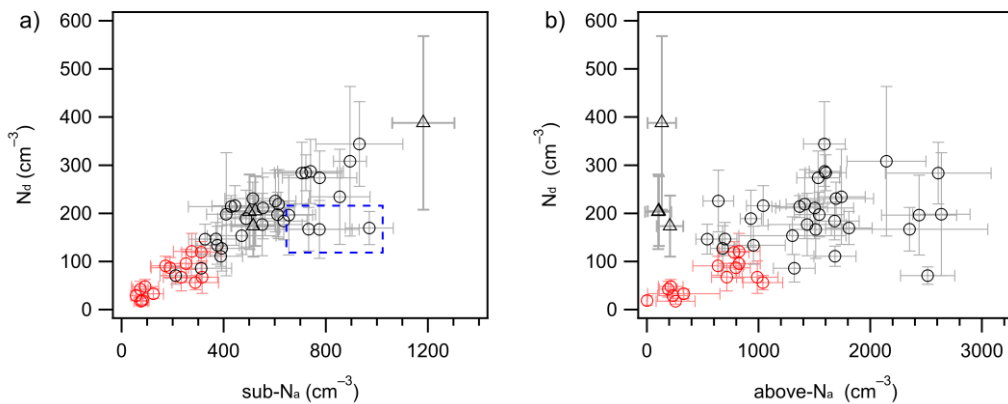


Figure 5. Relationships between cloud-layer mean N_d and N_a for all profiles, in terms of a) sub- N_a and b) above- N_a . The markers and error bars represent the average values and standard deviation for each profile. Black triangle and circle markers are from the BB-impacted MBL in Periods 1 and 3 respectively, and red circle markers are from the clean MBL in Period 2. The blue dashed box represents the contact profiles that outline the central relationship between N_d and sub- N_a and had average LWC/aLWC values close to 1.

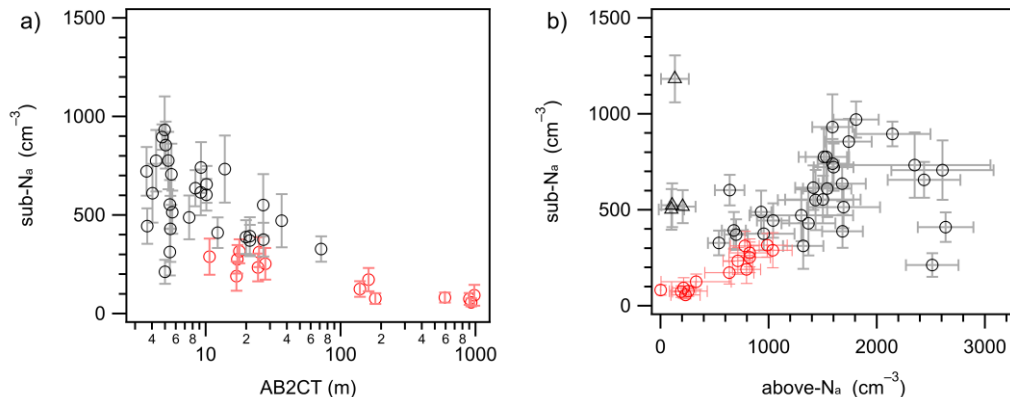


Figure 6. Relationships a) between sub- N_a and the distance from the top of cloud layers in the MBL to the bottom of BB layers in the FT (Cloud Top to Aerosol Base, CTtoAB); b) between sub- N_a and above- N_a . The markers and error bars represent the average values and standard deviation for each profile. Black triangle and circle markers are from the BB-impacted MBL in Periods 1 and 3 respectively, and red circle markers are from the clean MBL in Period 2.

3.3.2 R_e relationship with N_a and N_d

The R_e is one of the key variables that determine the radiative properties of liquid water clouds, and the parameterization of R_e in climate models is critical for assessing global radiative forcing. **Figure 7a** shows the relationship between cloud-layer mean R_e and sub- N_a for all the analyzed profiles, and **Fig. S5a** is for cloud-top-layer R_e and sub- N_a . LWP values integrated from the analyzed cloud profiles were in a larger range for the BB-impacted MBL (11.1 – 161.3 g m⁻²) than the clean MBL (25.1 – 63.0 g m⁻²). Although the LWP is indicated to be generally higher in the BB-impacted MBL than the clean MBL, the intrusion of BB aerosols led to a significantly smaller range of R_e as compared to the clean MBL, consistent with the Twomey effect (Twomey, 1977). The relationship between R_e and sub- N_a (**Fig. 7a**) shows a weaker correlation than that between N_d and sub- N_a (**Fig. 5a**), likely due to the influence of additional atmospheric factors such as the MBL thermodynamic structure, cloud depth, cloud-top entrainment process, etc (Wood, 2012; Herbert et al., 2020). **In Figs. 7a and S5a**, there are some profiles displaying smaller fractional decrease in R_e in response to the fractional increase in N_a ($> 0.1 \mu\text{m}$) compared to the other profiles. These profiles were mainly observed on days with larger temperature differences between the surface and upper layers of the MBL relative to the other profiles, potentially supporting stronger updrafts and thereby the development of thicker cloud layers and larger cloud droplets (Wood, 2012). A generally positive relationship between the cloud-top-layer R_e and cloud-layer depth is shown in **Fig. S5b**. Cloud-top entrainment processes may also explain the variability in the response of R_e to sub- N_a , since the mixing of warm and dry FT air into the clouds would reduce LWC and limit condensational growth, thereby constraining the droplet size (Wood, 2012). **Fig. S5c** shows the relationship between cloud-layer mean R_e and average

485 LWC/aLWC ratio for all the analyzed profiles. Cloud layers with negligible entrainment mixing (average LWC/aLWC values, ~ 1) present generally larger R_e values compared to those with greater entrainment mixing (lower average LWC/aLWC, < 0.83). Herbert et al. (2020) conducted large-eddy simulations and reported diurnal cycles of cloud thickness and LWP over the SEA, with both reaching their maximum at dawn and minimum shortly after midday. Thus, variations in the observation times of the day, which ranged from 9:00 to 16:00 (UTC time) during the campaign, may also contribute to the observed variability

490 in the response of R_e to sub- N_a .

Various studies have generated a power law parameterization between R_e and specific cloud water content (the ratio of LWC/ N_d) ($R_e \sim a(LWC/N_d)^b$), for use in climate models. Figure 7b shows the relationship between cloud-layer mean R_e versus the average ratio of LWC/ N_d for all analyzed profiles. BB-impacted and clean MBL profiles yielded similar exponent (b) values of (0.34 ± 0.01) and (0.33 ± 0.01) respectively, which are close to the exponent (~ 0.33) validated in previous BB aerosol-cloud studies such as in Amazon (e.g. Reid et al., 1999). Previous studies have reported a theoretical exponent of $b = 1/3$ for adiabatic clouds, where droplet growth is dominated by condensation without entrainment (Burnet and Brenguier, 2007). In this study, the estimated ratios of LWC/aLWC ($\sim 0.3 - 1$, Fig. S5c) suggest the occurrence of entrainment mixing, nevertheless, the empirical b values near or slightly exceed $\sim 1/3$. This indicates that entrainment processes were likely dominated by homogeneous mixing, which could proportionally reduce LWC and N_d and thus preserve $b \sim 1/3$, particularly in clean cases (Burnet and Brenguier, 2007). In BB-impacted cases, entrainment of aerosol-rich air may also supply additional CCN, increasing N_d disproportionately to LWC reduction. This entrainment of additional CCN could yield cases with $b > 1/3$ in this study. A key issue in the use of ($R_e \sim a(LWC/N_d)^b$) is the speciation of the pre-factor “a” (Liu and Daum et al., 2000), which is also similar for the BB-impacted (77.06 ± 7.53) and clean MBL profiles (69.54 ± 4.95). When combined, all the data in this study from clean and BB-impacted MBL profiles yielded $a = 68.54 \pm 3.05$ and $b = 0.33 \pm 0.01$. Field observations from this and previous work have tested similar power law schemes, suggesting that the parametrizations of $R_e \sim a(LWC/N_d)$ may exhibit consistent skill for varying degrees of BB pollution influence.

500

505

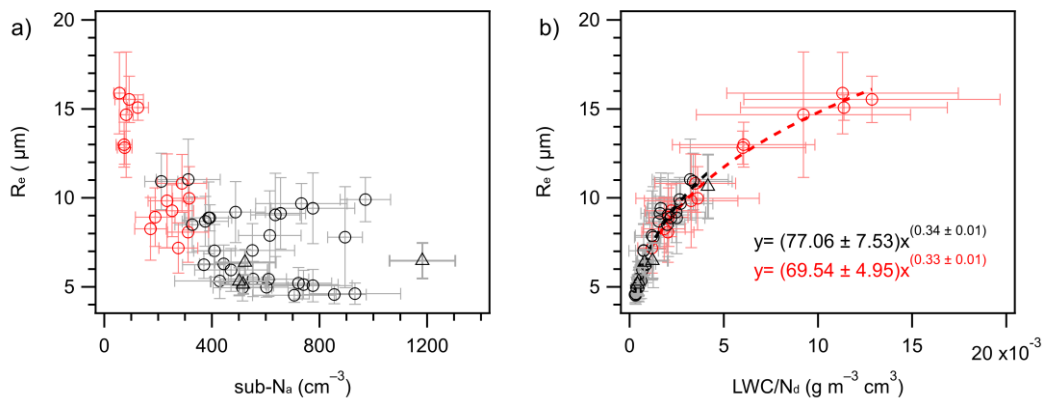


Figure 7. a) Relationship between cloud-layer mean R_e and sub- N_a for all analyzed profiles. b) Relationship between cloud-layer mean R_e and the average ratio of LWC/ N_d for all analyzed profiles. The markers and error bars represent the average values and standard deviation for each profile. Black triangle and circle markers are from the BB-impacted MBL in Periods 1 and 3 respectively, and red circle markers are from the clean MBL in Period 2.

510

3.4 Vertical transport history of observed MBL pollutants

3.4.1 Entrainment area

In previous sections, NAME dispersion simulations evidence that both FT and MBL pollutions observed over Ascension Island were attributable to the long-range transport of African BB aerosols (Sect. 3.1). Observational results around Ascension Island also highlight strong impacts of these transported BB aerosols on cloud properties over the remote SEA (Sects. 3.2 and 3.3). In this section, we further employed instantaneous 3D air parcel footprints from NAME dispersion simulations to examine vertical transport history of air parcels from Africa to the remote SEA, which is under-constrained in current models. Here, we focus on the three cases studies (Cases 1 to 3 illustrated in Sect. 3.1), representative for the three periods observed during the campaign. For the three NAME cases, we employed estimated z_i to divide the column-integrated horizontal footprints (Figs. 8a-c) into the FT (Figs. 8d-f) and BL (Figs. 8g-i) footprints separately, distinguishing original air parcels as either from the FT or BL. Figure 8 indicates that airmasses arriving at Ascension Island area could originate from both FT and BL during long-range transport. Original BL air parcels (Figs. 8g-i) were consistently transported from the southeast to northwest over the SEA and arose from a fairly clean oceanic region. It is the difference of original FT air parcels that contributed to BB pollution conditions. BB-impacted MBL cases (Cases 1 and 3, Figs. 8d, f) received original FT air parcels mostly from the African continent (20°S – 0°N), which could bring BB plumes from seasonal wildfires, while the clean MBL case (Cases 2, Figs. 8e, f) mainly received clean oceanic air.

For BB-impacted MBL cases (Cases 1 and 3), the instantaneous horizontal fields (midday-step footprint in Figs. S6 and S7) show that FT air parcels from the African continent moved westward toward the SEA region. To arrive in the MBL over Ascension Island, there should be African air parcels entraining from the FT into the MBL over the SEA. We co-located NAME transport pathways with SEVIRI retrievals, to identify the efficient entrainment regions where FT air parcels from Africa are likely to enter the MBL over the SEA. First, we analyzed the vertical distributions of air parcels that were dispersed into the FT region (covering the horizontal area of 20°S – 0°N, 15°W – 12°E) along the backward-simulation time, for two BB-impacted MBL cases (Fig. 9). The corresponding exchange rates (black dashed lines in Figs. 9a and 9b) and the cumulative exchange amounts (black dashed lines in Figs. 8Sa and 8Sb) of FT air parcels along the backward-simulation time were also calculated, to indicate the mixing conditions of African air parcels from the FT into the MBL over the SEA. These results suggest that FT air parcels from African sources spanning a wide range of altitudes were generally descending over the SEA before reaching the release area, and the MBL was deepening further offshore, which has been also identified in previous work (e.g. Das et al., 2017). FT air parcels can be entrained into the MBL when meeting the BL top. FT air parcels are suggested to be mostly mixed into the MBL in the last ~2 days for Case 1 (Figs. 9a and 8Sa) and in the last ~3 days for Case 3 (Figs. 9b and 8Sb) before arriving at Ascension Island. The mixing rates of FT air parcels into the MBL are suggested to vary between two BB-impacted MBL cases. There is a marked maximum in the exchange rate of FT air parcels for Case 1, whereas Case 3 shows a steadier and continuous mixing. For a given z_i and FT particle concentration, the rate of increase in MBL particle concentration is suggested to be dependent on the entrainment rate (ω_e) (Wood and Bretherton, 2004; Diamond et al., 2018).

The discrepancy between two cases suggests that entrainment rates in Case 1 were generally higher than Case 3, accounting for the stronger mixing in Case 1. Overall, NAME results of BB-impacted MBL cases indicate that efficient mixing of FT air parcels into the MBL would occur multiple days over the SEA during their westward movement from the African continent to the Ascension Island area. Combining the instantaneous fields (Figs. S6 and S7), the main area where the efficient mixing occurred is suggested to be located west of 0°E for examined cases.

To confirm FT BB pollution conditions along the simulated transport pathways, we performed a statistical analysis of the SEVIRI-retrieved above-cloud AOT (hereafter referred to as AOT) co-located with NAME horizontal footprints, at 3-hourly backward step. Examples of the co-location relationship between AOT and NAME instantaneous horizontal footprints are shown in Fig. S9. The parts of the scene that are co-located with the contemporaneous horizontal footprints from NAME simulations are highlighted. The air-density-weighted AOT (AOT_{weighted}) was calculated for the co-located areas following Eq. (8).

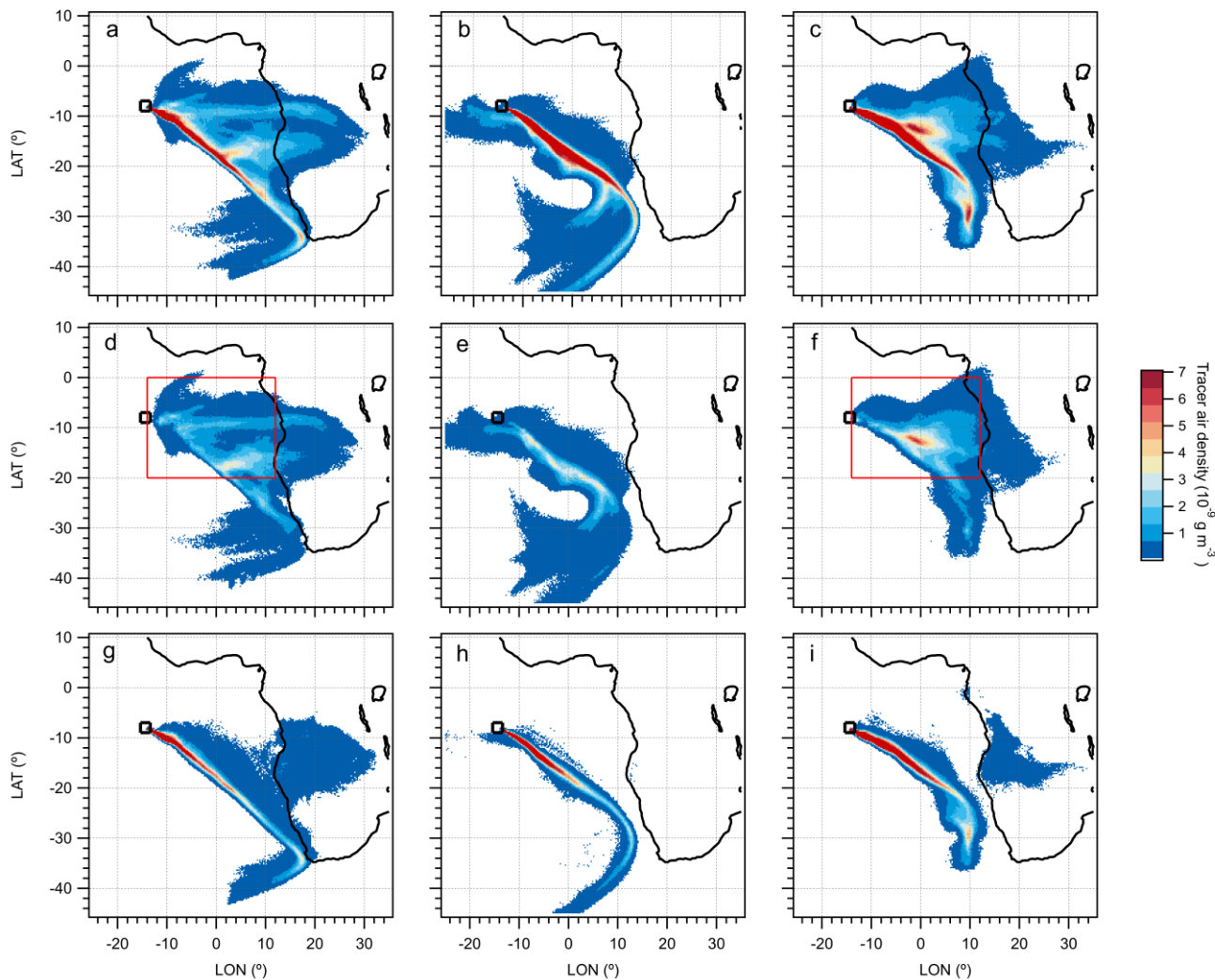
$$AOT_{\text{weighted}} = \sum_{i=1}^N \left(\frac{Mass_i}{\sum_{i=1}^N Mass_i} \times AOT_i \right) \quad (8)$$

Where N is the total number of horizontal air parcel grids for co-located areas, $Mass_i$ is air parcel concentration (g m^{-3}) in each grid, $\frac{Mass_i}{\sum_{i=1}^N Mass_i}$ represents fractional air parcel concentration in each grid relative to the total air parcel density, AOT_i is the SEVIRI-retrieved above-cloud AOT co-located with each air parcel grid. It is noted that the air-density-weighted N_d , R_e and CTH (Sect. 3.4.2) were also calculated following Eq. (8), with AOT replaced by the corresponding SEVIRI-retrieved parameters. Figs. 9c and 9d show the AOT_{weighted} along the FT and BL air parcel transport pathways respectively, for the three case studies. High AOT_{weighted} values show the co-located and co-temporaneous abundance of African BB aerosols within the FT column above the cloud layer along the simulated transport pathway. For BB-impacted MBL cases (Cases 1 and 3), the AOT co-located with FT air parcels (black lines in Fig. 9c) remained mostly high during the periods of efficient mixing indicated by the NAME simulations. This indicates that FT BB aerosols existed over the efficient mixing area and so could be entrained from the FT into the MBL if the bottom of FT BB layer is near the cloud top. Along BL transport from southeast to northwest over the SEA, the AOT co-located with BL air parcels (black lines in Fig. 9d) was enhanced when approaching the region of efficient mixing of FT air parcels (west of 0°E). This further indicates that once FT BB aerosols could entrain into the MBL over efficient mixing periods, they could be subsequently advected by the MBL south-easterly winds toward Ascension Island area. Since the above-cloud AOT is a column integrated abundance and the aerosol layer may be vertically separated from the cloud top, this analysis does not unequivocally demonstrate entrainment of BB aerosols at these locations. Nevertheless, it is a necessary condition that elevated FT aerosol abundance is co-located with regions of efficient entrainment for the MBL to receive inputs of BB aerosols. Meteorological analysis of the SEA in August 2017 supports the connection between FT BB layers and the cloud top (Ryoo et al., 2022). The low-level easterly jet in early and late August allowed the transport of African BB plumes in the low FT and thus the possibility of BB aerosols entrainment into the MBL in efficient mixing area. In comparison, the AOT in the clean-MBL case (Case 2) was continuously low along both FT and BL transport (red lines in Fig. 9c and 9d), demonstrating a negligible contribution of BB pollution to the MBL over Ascension Island. This

580 is due to a strong mid-level easterly jet in mid-August (Ryoo et al., 2022), leading to a disconnection between FT BB layers and the BL top, thereby suppressing the entrainment of African BB aerosols into the MBL.

Based on NAME simulations and AOT along the simulated transport, the occurrence of BB-impacted MBL over Ascension Island is suggested to be related to two factors. Firstly, African BB plumes aloft need to be transported to the efficient mixing areas (west of 0°E) identified in this study. Furthermore, the transported BB plumes can be entrained into the MBL in efficient mixing areas over the SEA, which depends on the altitude of the African easterly jet and the distances between BB layers and the BL top. The significance of the efficient mixing areas identified in this study lies in the potential for the entrained BB aerosols in these regions to be transported by MBL south-easterly winds to the vicinity of Ascension Island.

585



590 **Figure 8. The 7-days backward-dispersion of three cases from NAME simulations. Panels a-c) are column-integrated horizontal footprints for a) Case 1, b) Case 2 and c) Case 3. Panels d-f) are dispersion results attributed to FT for d) Case 1, e) Case 2 and f) Case 3. Panels g-i) are dispersion results attributed to BL for g) Case 1, h) Case 2 and i) Case 3. All plots are shown in the same colour scale. The black boxes represent the release locations of Ascension Island. The red boxes in Fig. e and g represent the**

horizontal grids (covering the area of 20°S – 0°N, 15°W – 12°E) used for integrating within each vertical layer to derive the vertical distribution of dispersed air parcels in Fig. 9.

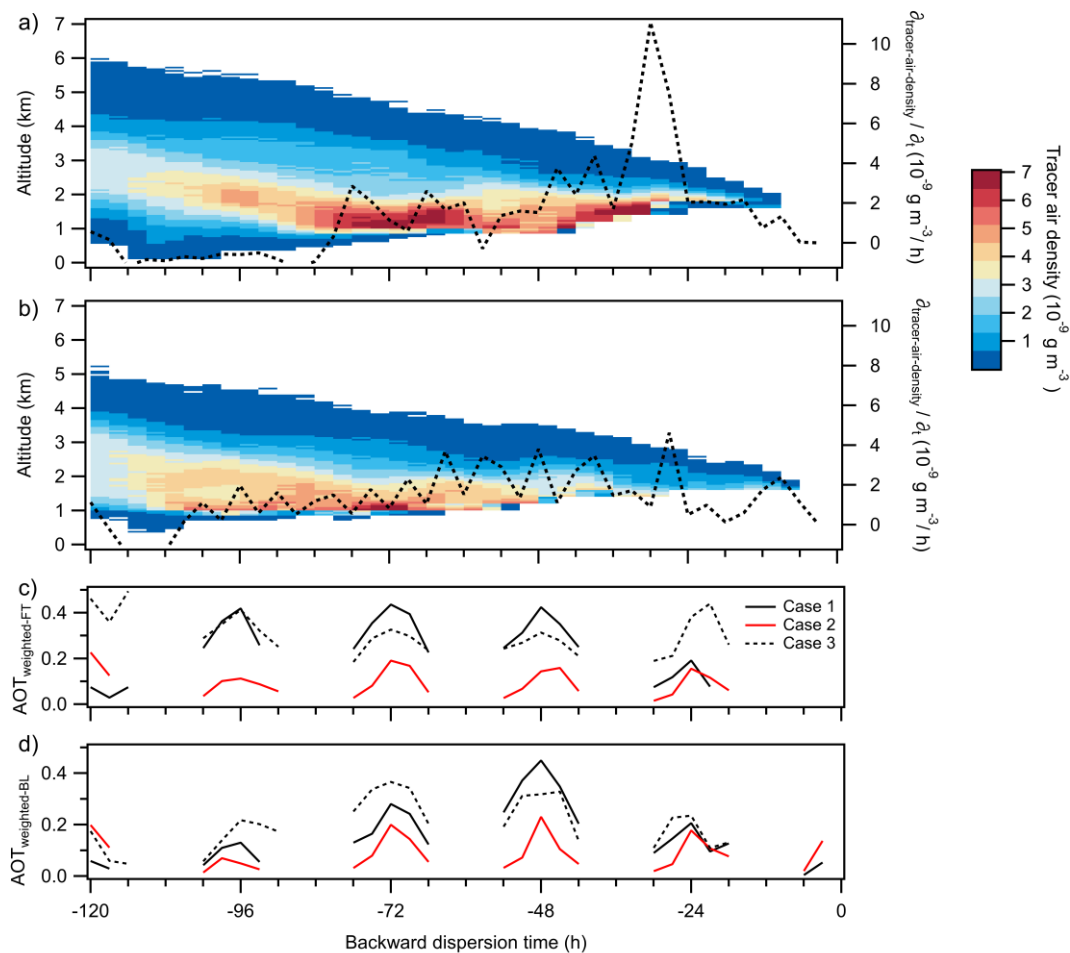


Figure 9. a-b) Time series of the vertical distribution of air parcels dispersed into the FT region (covering the horizontal area of 20°S – 0°N, 15°W – 12°E, see red box in Fig. 8), in terms of backward dispersion time from 3 to 120 h. The right axis (dashed black lines) are the corresponding 3-hourly exchange rates of air parcels dispersed between the FT and MBL, calculated by dividing the increase in dispersed FT air parcels by the dispersion time along backward simulations. a) is for Case 1 and b) is for Case 3. c-d) The air-density-weighted AOT co-located with c) FT and d) BL footprints along the backward-simulation time, for three cases.

3.4.2 Cloud fields along transport

Here, we extend the analysis of SEVIRI-retrieved cloud properties (N_d and R_e) along the simulated BL transport from southeast to northwest over the SEA, before arrival at Ascension Island. Comparisons between BB-impacted and clean MBL cases demonstrate the impacts of entrained BB aerosols on a broad cloud field. Similar to the AOT, examples of the co-location relationships between N_d or R_e and NAME instantaneous horizontal footprints at a given time step are shown in Fig. S9. The air-density-weighted N_d and R_e were calculated for co-located areas along the BL transport, at 3-hourly backward step. Figures.

10a and 10b show the evolution of air-density-weighted N_d and R_e along the BL transport, for three cases. In the clean MBL case (Case 2), the N_d was consistently low throughout the passage of BL air parcels over the SEA, while the R_e generally increased. The evolution of clouds over the SEA is related to meteorological factors such as the evolution of CTH (Painemal et al., 2014). The air-density-weighted CTH was calculated along the BL transport, at 3-hourly backward step (Fig. 10c). In Case 2 (red line in Fig. 10c), the CTH shows an increasing trend along the southeast to northwest transport path over the SEA. Concurrent MBL deepening and enhanced CTH could promote condensational growth, yielding larger droplets at the cloud top (Painemal et al., 2014). This agrees with the observed overall increase in R_e along the BL transport. The deeper MBL also tends to be decoupled, and the decoupled MBL is likely associated with an increased occurrence of drizzle due to the presence of larger droplets (Jones et al., 2011). The occurrence of drizzle and associated deposition may result in reduced N_d and R_e , corresponding to the observed slightly reduction in the vicinity of Ascension Island. For BB-impacted MBL cases (Cases 1 and 3), the N_d (black lines in Fig. 10a) was largely enhanced when FT air parcels from the African continent were suggested to efficiently entrain into the MBL during ~2–3 days prior to arrival at Ascension Island. Although the deepened MBL (black lines in Fig. 10c) with accompanying processes of condensational growth should promote increasing R_e (as in the clean MBL case), the R_e (black lines in Fig. 10b) was similar along the BL transport. This suggests that the enhanced N_d from entrained BB aerosols promoted a reduction in cloud droplet size, which offsets the growth effects from the deepened MBL. Comparisons between BB-impacted and clean MBL cases show that BB aerosol entrainment would lead to significantly enhanced N_d but reduced R_e , modifying cloud fields in broad efficient mixing areas that persist through the SEA region. This modification of cloud properties, which is co-located with the efficient entrainment areas, further supports NAME analysis of the mixing history of African air parcels. Additionally, the modification of the cloud fields by high BB aerosol loadings, imply the important indirect radiative effects of transported African BB aerosols over the SEA. Future studies that integrate aerosol–cloud interaction parameterizations developed in this study will be conducted to improve the assessment of aerosol indirect radiative effects in this region.

By coupling NAME simulations with SEVIRI retrievals (aerosol and cloud fields) along the simulated transport path, we have been able to identify the efficient entrainment areas where African FT air parcels mix into the MBL before reaching Ascension Island. Efficient FT-to-MBL entrainment of African BB aerosols may occur multiple days prior to arrival at Ascension Island and mainly in a region of west of ~0°E over the SEA for the examined cases. Previous model studies have highlighted uncertainty in accurately simulating the vertical distribution of transported African BB aerosols over the SEA and their subsequent climate effects (Das et al., 2017; Shinozuka et al., 2020). These studies emphasized that the transported BB aerosols are predicted to descend to unrealistically low altitudes as they move westward off the African coast. One potential explanation for this sudden subsidence over the SEA is the inadequate representation of a "self-lofting" process in models, where absorption-induced heating from BB aerosols can lift the associated plume layer to higher altitudes (Johnson et al., 2023). The insufficient simulation of this self-lofting process is likely due to the large underestimation of the single scattering albedo (SSA) and subsequent absorption of transported BB aerosols in Earth system models (Mallet et al., 2021). Another potential explanation for the misrepresentation of vertical transport and hence vertical distribution of BB aerosols is related to

the resolution of models, as small-scale processes such as turbulence and self-lofting are not explicitly resolved in coarse-resolution models (Das et al., 2017). To improve the simulation of vertical distributions of African BB aerosols over the remote SEA, it is essential to reduce model biases in aerosol properties (e.g. SSA) and vertical processes. Improving the representation of these processes is complex as reducing bias in one property may lead to greater biases in another. The identified efficient entrainment areas in this study could provide further constraints for testing and refining predictions related to BB aerosol transport and vertical distributions over the SEA. Furthermore, the indicated transport and mixing history from NAME simulations are supported by the evolution of aerosol/cloud fields from SEVIRI retrievals, suggesting that high-resolution meteorological fields from the Unified Model could provide a reasonable simulation of the plume descent, which could be employed in future regional model simulations.

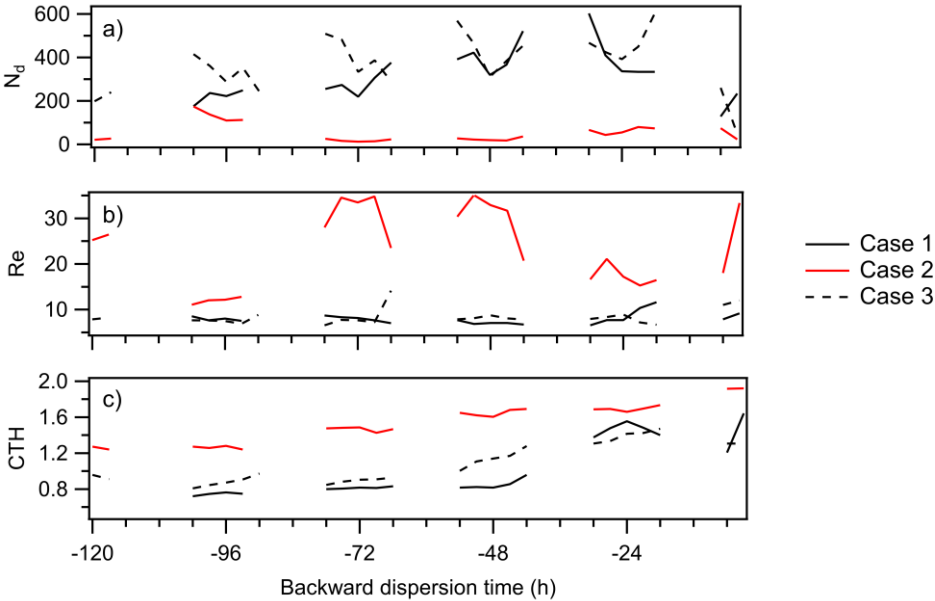


Figure 10. The air-density-weighted a) N_d , b) Re , c) CTH from SEVIRI, which are co-located with BL air parcels along the backward-simulation time, for three cases.

4 Conclusions and implications

In this study, we have characterized the vertical profiles of aerosol and cloud properties around Ascension Island in the remote SEA during the CLARIFY-2017 aircraft campaign and investigated the response of clouds to observed aerosols. We conducted backward-dispersion simulations using the UK Met Office's NAME, which evidences that observed pollutions around Ascension Island originated from long-range transported African BB plumes.

During the campaign, the BB-impacted MBL had substantially enhanced N_a compared to the clean MBL around Ascension Island. This resulted in increased N_d and LWC values but reduced R_e within BB-impacted clouds compared to clean clouds. A stronger correlation was observed between the cloud layer-mean N_d and sub-cloud aerosol properties (sub- N_a), while the N_d was only weakly associated with smoke immediately above the cloud. This suggests that the cloud base activation of sub-cloud N_a was more important in determining N_d values than the cloud-top activation of N_a at the place/time of observation. The observed increase in the sub- N_a is a result of entrained BB aerosols from above-cloud to below-cloud regions along long-range transport or/and the place of observation. We explored possible simplifications to establish relationships between N_d and sub- N_a or R_e from in-situ measurements. Similar droplet activation patterns ($N_d \sim \text{sub-}N_a$) were observed in the clean and moderately BB-impacted (sub- $N_a < 700 \text{ cm}^{-3}$) clouds, likely due to the comparable CCN ability of MBL aerosols under the two conditions. However, a weaker N_d - N_a correlation was noted in more BB-polluted clouds (sub- $N_a > 700 \text{ cm}^{-3}$), possibly influenced by additional atmospheric factors such as the cloud-top entrainment and updraft velocity variations. Comparisons between this study and other BB regions indicate that the droplet formation sensitivity to aerosol concentration is varying across different BB-impacted areas, likely due to the variations in CCN ability of BB aerosols. Thus, region-specific parameterizations of N_d based on N_a are necessary for accurately simulating regional radiative forcing. The relationship between N_d and R_e remained consistent regardless of the levels of BB influence. This study extends previous aerosol-cloud interaction field studies to a wider region over the SEA, which provides important aerosol and cloud parameterizations in climate models for benefiting the assessment of the impacts of transported BB aerosols on clouds and their subsequent climate effects over the remote SEA. Furthermore, the disconnectedness between cloud properties and the above-cloud aerosol layer indicates that satellite-based observations of aerosol-cloud interaction in this region may be subject to large uncertainty, as they are unable to resolve the aerosol properties within and below the cloud layer. This underscores the importance of airborne in-situ measurements in this region to provide direct observations of aerosol-cloud interaction.

NAME simulations allow a discrimination of air parcels in different layers, and this has been used to derive the FT-to-MBL entrainment of African air parcels over the SEA. NAME analysis coupled with SEVIRI-retrievals (aerosol and cloud fields) along the simulated transport suggests that efficient entrainment of African BB smoke from the FT into the MBL occurred multiple days before arrival at Ascension Island, which is mainly in the region to the west of 0°E for examined cases. One significance of the efficient mixing areas identified in this study lies in the potential for the entrained BB aerosols to be further transported by MBL south-easterly winds to the vicinity of Ascension Island. The analysis of SEVIRI-retrieved cloud fields along the simulated transport path also indicates significant modifications in cloud droplet concentrations and sizes

within the broad efficient mixing areas of African BB smoke. The efficient entrainment areas identified in this study can provide an important constraint for refining the transport processes and vertical distributions of African BB aerosols over the SEA. More case studies should be conducted to validate this methodology and examine entrainment patterns across different months and years, ultimately to characterize a climatological area for entrainment over the SEA.

690

Data availability. Airborne measurements are available from the Centre for Environmental Data Analysis <https://catalogue.ceda.ac.uk/uuid/38ab7089781a4560b067dd6c20af3769>.

695

Author contributions. H.C. and J.H. designed the research; J.W.T., S.J.A., P.B., J.T. and K.B. performed field experiments; H.W., J.W.T. and J.T. prepared in-situ datasets; F.P. provided the SEVIRI dataset; H.W. and C.Y. performed NAME analysis; H.W. analyzed combined datasets and led the manuscript writing.

Competing interests. The authors declare no competing interests.

700

Acknowledgements. This work was mainly supported by the UK Natural Environment Research Council (NERC) Large Grant NE/L013584/1 via project CLARIFY (Cloud-Aerosol-Radiation Interactions and Forcing for Year 2017). The first author has also received funding from Horizon Europe programme under Grant Agreement No. 101137680 via project CERTAINTY (Cloud-aERosol inTeractions & their impActs IN The earth sYstem). The staff of Airtask, Avalon Engineering and FAAM are thanked for their thoroughly professional work, before, during and after the deployment. The NAME group in the UK Met

705

Office are thanked for their instructions on backward simulations.

References

- Abel, S. J., Barrett, P. A., Zuidema, P., Zhang, J., Christensen, M., Peers, F., Taylor, J. W., Crawford, I., Bower, K. N., and Flynn, M.: Open cells exhibit weaker entrainment of free-tropospheric biomass burning aerosol into the south-east Atlantic boundary layer, *Atmos. Chem. Phys.*, 20, 4059–4084, <https://doi.org/10.5194/acp-20-4059-2020>, 2020.
- Adebisi, A. and Zuidema, P.: The Role of the Southern African Easterly Jet in Modifying the Southeast Atlantic Aerosol and Cloud Environments, *Q. J. Roy. Meteor. Soc.*, 142, 697, 1574–1589, <https://doi.org/10.1002/qj.2765>, 2016.
- Barrett, P. A., Abel, S. J., Coe, H., Crawford, I., Dobracki, A., Haywood, J., Howell, S., Jones, A., Langridge, J., McFarquhar, G. M., Nott, G. J., Price, H., Redemann, J., Shinozuka, Y., Szpek, K., Taylor, J. W., Wood, R., Wu, H., Zuidema, P., Bauguitte, S., Bennett, R., Bower, K., Chen, H., Cochrane, S., Cotterell, M., Davies, N., Delene, D., Flynn, C., Freedman, A., Freitag, S., Gupta, S., Noone, D., Onasch, T. B., Podolske, J., Poellot, M. R., Schmidt, S., Springston, S., Sedlacek III, A. J., Trembath, J., Vance, A., Zawadowicz, M. A., and Zhang, J.: Intercomparison of airborne and surface-based measurements during the CLARIFY, ORACLES and LASIC field experiments, *Atmos. Meas. Tech.*, 15, 6329–6371, <https://doi.org/10.5194/amt-15-6329-2022>, 2022.
- Boucher, O., Randall, D., Artaxo, P., Bretherton, C., Feingold, G., Forster, P., Kerminen, V.-M., Kondo, Y., Liao, H., Lohmann, U., Rasch, P., Satheesh, S. K., Sherwood, S., Stevens, B., and Zhang, X. Y.: Clouds and aerosols, in: *Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change*, edited by: Stocker, T. F., Qin, D., Plattner, G.-K., Tignor, M., Allen, S. K., Doschung, J., Nauels, A., Xia, Y., Bex, V., and Midgley, P. M.: Cambridge University Press, Cambridge, United Kingdom and New York USA, 571–657, <https://doi.org/10.1017/CBO9781107415324.016>, 2013.
- Brown, A., Milton, S., Cullen, M., Golding, B., Mitchell, J., and Shelly, A.: Unified Modeling and Prediction of Weather and Climate: A 25-Year Journey, *B. Am. Meteorol. Soc.*, 93, 1865–1877, <https://doi.org/10.1175/bams-d-12-00018.1>, 2012.
- Burnet, F. and Brenguier, J.-L.: *Observational Study of the Entrainment-Mixing Process in Warm Convective Clouds*, *J. Atmos. Sci.*, 64, 1995–2011, <https://doi.org/10.1175/JAS3928.1>, 2007.
- Costantino, L. and Bréon, F.-M.: Analysis of aerosol-cloud interaction from multi-sensor satellite observations, *Geophys. Res. Lett.*, 37, L11801, <https://doi.org/10.1029/2009gl041828>, 2010.
- Costantino, L. and Bréon, F.-M.: Aerosol indirect effect on warm clouds over South-East Atlantic, from co-located MODIS and CALIPSO observations, *Atmos. Chem. Phys.*, 13, 69–88, <https://doi.org/10.5194/acp-13-69-2013>, 2013.
- Das, S., Harshvardhan, H., Bian, H., Chin, M., Curci, G., Protonotariou, A. P., Mielonen, T., Zhang, K., Wang, H., and Liu, X.: Biomass burning aerosol transport and vertical distribution over the South African-Atlantic region, *J. Geophys. Res.-Atmos.*, 122, 6391–6415, <https://doi.org/10.1002/2016jd026421>, 2017.
- Deaconu, L. T., Ferlay, N., Waquet, F., Peers, F., Thieuleux, F., and Goloub, P.: Satellite inference of water vapour and above-cloud aerosol combined effect on radiative budget and cloud-top processes in the southeastern Atlantic Ocean, *Atmos. Chem. Phys.*, 19, 11613–11634, <https://doi.org/10.5194/acp-19-11613-2019>, 2019.

740 De Graaf, M., Bellouin, N., Tilstra, L. G., Haywood, J., and Stammes, P.: Aerosol direct radiative effect of smoke over clouds over the southeast Atlantic Ocean from 2006 to 2009, *Geophys. Res. Lett.*, 41, 7723–7730, <https://doi.org/10.1002/2014GL061103>, 2014.

Diamond, M. S., Dobracki, A., Freitag, S., Small Griswold, J. D., Heikkila, A., Howell, S. G., Kacarab, M. E., Podolske, J. R., Saide, P. E., and Wood, R.: Time-dependent entrainment of smoke presents an observational challenge for assessing aerosol–

745 cloud interactions over the southeast Atlantic Ocean, *Atmos. Chem. Phys.*, 18, 14623–14636, <https://doi.org/10.5194/acp-18-14623-2018>, 2018.

Dobracki, A., Zuidema, P., Howell, S. G., Saide, P., Freitag, S., Aiken, A. C., Burton, S. P., Sedlacek III, A. J., Redemann, J., and Wood, R.: An attribution of the low single-scattering albedo of biomass burning aerosol over the southeastern Atlantic, *Atmos. Chem. Phys.*, 23, 4775–4799, <https://doi.org/10.5194/acp-23-4775-2023>, 2023.

750 Dusek, U., Frank, G. P., Hildebrandt, L., Curtius, J., Schneider, J., Walter, S., Chand, D., Drewnick, F., Hings, S., Jung, D., Borrmann, S., and Andreae, M. O.: Size matters more than chemistry for cloud-nucleating ability of aerosol particles, *Science*, 312, 1375–1378, <https://doi.org/10.1126/science.1125261>, 2006.

Farley R., Zhou S., Collier S., Jiang W., Onasch T.B., Shilling J.E., Kleinman L., Sedlacek III A.J., and Zhang Q.: Chemical Evolution of Biomass Burning Aerosols across Wildfire Plumes in the Western US: From Near-Source to Regional Scales.

755 *ACS Es&t Air.*, 2(4), 677–691, <https://doi.org/10.1021/acsestair.5c00002>, 2025.

Gordon, H., Field, P. R., Abel, S. J., Dalvi, M., Grosvenor, D. P., Hill, A. A., Johnson, B. T., Miltenberger, A. K., Yoshioka, M., and Carslaw, K. S.: Large simulated radiative effects of smoke in the south-east Atlantic, *Atmos. Chem. Phys.*, 18, 15261–15289, <https://doi.org/10.5194/acp-18-15261-2018>, 2018.

Gupta, S., McFarquhar, G. M., O'Brien, J. R., Delene, D. J., Poellot, M. R., Dobracki, A., Podolske, J. R., Redemann, J.,

760 LeBlanc, S. E., Segal-Rozenhaimer, M., and Pistone, K.: Impact of the variability in vertical separation between biomass burning aerosols and marine stratocumulus on cloud microphysical properties over the Southeast Atlantic, *Atmos. Chem. Phys.*, 21, 4615–4635, <https://doi.org/10.5194/acp-21-4615-2021>, 2021.

Gupta, S., McFarquhar, G. M., O'Brien, J. R., Poellot, M. R., Delene, D. J., Miller, R. M., and Small Griswold, J. D.: Factors affecting precipitation formation and precipitation susceptibility of marine stratocumulus with variable above- and below-

765 cloud aerosol concentrations over the Southeast Atlantic, *Atmos. Chem. Phys.*, 22, 2769–2793, <https://doi.org/10.5194/acp-22-2769-2022>, 2022.

Haywood, J. M., Abel, S. J., Barrett, P. A., Bellouin, N., Blyth, A., Bower, K. N., Brooks, M., Carslaw, K., Che, H., Coe, H., Cotterell, M. I., Crawford, I., Cui, Z., Davies, N., Dingley, B., Field, P., Formenti, P., Gordon, H., de Graaf, M., Herbert, R., Johnson, B., Jones, A. C., Langridge, J. M., Malavelle, F., Partridge, D. G., Peers, F., Redemann, J., Stier, P., Szpek, K., Taylor, J. W., Watson-Parris, D., Wood, R., Wu, H., and Zuidema, P.: The CLOUD–Aerosol–Radiation Interaction and Forcing: Year 2017 (CLARIFY-2017) measurement campaign, *Atmos. Chem. Phys.*, 21, 1049–1084, <https://doi.org/10.5194/acp-21-1049-2021>, 2021.

- Hegg, D. A., Covert, D. S., Jonsson, H. H., and Woods, R. K.: A simple relationship between cloud drop number concentration and precursor aerosol concentration for the regions of Earth's large marine stratocumulus decks, *Atmos. Chem. Phys.*, 12, 1229–1238, <https://doi.org/10.5194/acp-12-1229-2012>, 2012.
- Herbert, R. J., Bellouin, N., Highwood, E. J., and Hill, A. A.: Diurnal cycle of the semi-direct effect from a persistent absorbing aerosol layer over marine stratocumulus in large-eddy simulations, *Atmos. Chem. Phys.*, 20, 1317–1340, <https://doi.org/10.5194/acp-20-1317-2020>, 2020.
- Hering, S. V., Stolzenburg, M. R., Quant, F. R., Oberreit, D. R., and Keady, P. B.: A Laminar-Flow, Water-Based Condensation Particle Counter (WCPC), *Aerosol Sci. Tech.*, 39, 659–672, <https://doi.org/10.1080/02786820500182123>, 2005.
- Hoffmann, F., Raasch, S., and Noh, Y.: Entrainment of aerosols and their activation in a shallow cumulus cloud studied with a coupled LCM–LES approach, *Atmos. Res.*, 156, 43–57, <https://doi.org/10.1016/j.atmosres.2014.12.008>, 2015.
- Holanda, B. A., Pöhlker, M. L., Walter, D., Saturno, J., Sörgel, M., Ditas, J., Ditas, F., Schulz, C., Franco, M. A., Wang, Q., Donth, T., Artaxo, P., Barbosa, H. M. J., Borrmann, S., Braga, R., Brito, J., Cheng, Y., Dollner, M., Kaiser, J. W., Klimach, T., Knote, C., Krüger, O. O., Fütterer, D., Lavrič, J. V., Ma, N., Machado, L. A. T., Ming, J., Morais, F. G., Paulsen, H., Sauer, D., Schlager, H., Schneider, J., Su, H., Weinzierl, B., Walser, A., Wendisch, M., Ziereis, H., Zöger, M., Pöschl, U., Andreae, M. O., and Pöhlker, C.: Influx of African biomass burning aerosol during the Amazonian dry season through layered transatlantic transport of black carbon-rich smoke, *Atmos. Chem. Phys.*, 20, 4757–4785, <https://doi.org/10.5194/acp-20-4757-2020>, 2020.
- Johnson, B. T. and Haywood, J. M.: Assessing the Impact of Self-Lofting on Increasing the Altitude of Black Carbon in a Global Climate Model, *J. Geophys. Res.-Atmos.*, 128, e2022JD038039, <https://doi.org/10.1029/2022JD038039>, 2023.
- Jones, A., Thomson, D., Hort, M., and Devenish, B.: The UK Met Office's next-generation atmospheric dispersion model, NAME III, *Air Pollut. Model. Appl. XVII*, 17, 580–589, https://doi.org/10.1007/978-0-387-68854-1_62, 2007.
- Jones, C. R., Bretherton, C. S., and Leon, D.: Coupled vs. decoupled boundary layers in VOCALS-REx, *Atmos. Chem. Phys.*, 11, 7143–7153, <https://doi.org/10.5194/acp-11-7143-2011>, 2011.
- Kacarab, M., Thornhill, K. L., Dobracki, A., Howell, S. G., O'Brien, J. R., Freitag, S., Poellot, M. R., Wood, R., Zuidema, P., Redemann, J., and Nenes, A.: Biomass burning aerosol as a modulator of the droplet number in the southeast Atlantic region, *Atmos. Chem. Phys.*, 20, 3029–3040, <https://doi.org/10.5194/acp-20-3029-2020>, 2020.
- Keller, A., Tritscher, T., and Burtscher, H.: Performance of water-based CPC 3788 for particles from a propane-flame soot-generator operated with rich fuel/air mixtures, *J. Aerosol Sci.*, 60, 67–72, <https://doi.org/10.1016/j.jaerosci.2013.02.005>, 2013.
- Laing, J. R., Jaffe, D. A., and Hee, J. R.: Physical and optical properties of aged biomass burning aerosol from wildfires in Siberia and the Western USA at the Mt. Bachelor Observatory, *Atmos. Chem. Phys.*, 16, 15185–15197, <https://doi.org/10.5194/acp-16-15185-2016>, 2016.
- Lance, S., Brock, C. A., Rogers, D., and Gordon, J. A.: Water droplet calibration of the Cloud Droplet Probe (CDP) and in-flight performance in liquid, ice and mixed-phase clouds during ARCPAC, *Atmos. Meas. Tech.*, 3, 1683–1706, <https://doi.org/10.5194/amt-3-1683-2010>, 2010.

- Liu, Y. and Daum, P. H.: Spectral dispersion of cloud droplet size distributions and the parameterization of cloud droplet effective radius, *Geophys. Res. Lett.*, 27, 1903–1906, <https://doi.org/10.1029/1999GL011011>, 2000.
- Lock, A., Edwards, J. and Boutle, I.: Unified Model Documentation Paper 024, The Parametrization of Boundary Layer Processes, Version 11.7, Met Office, available at: <https://code.metoffice.gov.uk/doc/um> (last access: 20 June 2024), Last updated 2020.
- Lu, Zh., Liu, X., Zhang, Z., Zhao, C., Meyer, K., Rajapakshe, C., Wu, C., Yang, Z., and Penner, J. E.: Biomass smoke from southern Africa can significantly enhance the brightness of stratocumulus over the southeastern Atlantic Ocean, *P. Natl. Acad. Sci. USA*, 115, 2924–2929, <https://doi.org/10.1073/pnas.1713703115>, 2018.
- Mallet, M., Nabat, P., Johnson, B., Michou, M., Haywood, J. M., Chen, C., and Dubovik, O.: Climate models generally underrepresent the warming by Central Africa biomass-burning aerosols over the Southeast Atlantic, *Sci. Adv.*, 7, eabg9998, 685 <https://doi.org/10.1126/sciadv.abg9998>, 2021.
- Mardi, A. H., Dadashazar, H., MacDonald, A. B., Crosbie, E., Coggon, M. M., Aghdam, M. A., Woods, R. K., Jonsson, H. H., Flagan, R. C., Seinfeld, J. H., and Sorooshian, A.: Effects of Biomass Burning on Stratocumulus Droplet Characteristics, Drizzle Rate, and Composition, *J. Geophys. Res.-Atmos.*, 124, 12301–12318, <https://doi.org/10.1029/2019jd031159>, 2019.
- McComiskey, A. and Feingold, G.: Quantifying error in the radiative forcing of the first aerosol indirect effect, *Geophys. Res. Lett.*, 35, L02810, <https://doi.org/10.1029/2007GL032667>, 2008.
- Painemal, D., Minnis, P., Ayers, J. K., and O'Neill, L.: GOES-10 microphysical retrievals in marine warm clouds: multi-instrument validation and daytime cycle over the southeast Pacific, *J. Geophys. Res.-Atmos.*, 117, D19212, <https://doi.org/10.1029/2012jd017822>, 2012.
- Painemal, D., Kato, S., and Minnis, P.: Boundary layer regulation in the southeast Atlantic cloud microphysics during the biomass burning season as seen by the A-train satellite constellation, *J. Geophys. Res.-Atmos.*, 119, 11288–11302, <https://doi.org/10.1002/2014JD022182>, 2014.
- Panagi, M., Fleming, Z. L., Monks, P. S., Ashfold, M. J., Wild, O., Hollaway, M., Zhang, Q., Squires, F. A., and Vande Hey, J. D.: Investigating the regional contributions to air pollution in Beijing: a dispersion modelling study using CO as a tracer, *Atmos. Chem. Phys.*, 20, 2825–2838, <https://doi.org/10.5194/acp-20-2825-2020>, 2020.
- Peers, F., Francis, P., Fox, C., Abel, S. J., Szpek, K., Cotterell, M. I., Davies, N. W., Langridge, J. M., Meyer, K. G., Platnick, S. E., and Haywood, J. M.: Observation of absorbing aerosols above clouds over the south-east Atlantic Ocean from the geostationary satellite SEVIRI – Part 1: Method description and sensitivity, *Atmos. Chem. Phys.*, 19, 9595–9611, <https://doi.org/10.5194/acp19-9595-2019>, 2019.
- Peers, F., Francis, P., Abel, S. J., Barrett, P. A., Bower, K. N., Cotterell, M. I., Crawford, I., Davies, N. W., Fox, C., Fox, S., Langridge, J. M., Meyer, K. G., Platnick, S. E., Szpek, K., and Haywood, J. M.: Observation of absorbing aerosols above clouds over the south-east Atlantic Ocean from the geostationary satellite SEVIRI – Part 2: Comparison with MODIS and aircraft measurements from the CLARIFY-2017 field campaign, *Atmos. Chem. Phys.*, 21, 3235–3254, <https://doi.org/10.5194/acp-21-3235-2021>, 2021.

- Pistone, K., Zuidema, P., Wood, R., Diamond, M., da Silva, A. M., Ferrada, G., Saide, P. E., Ueyama, R., Ryoo, J.-M., Pfister, L., Podolske, J., Noone, D., Bennett, R., Stith, E., Carmichael, G., Redemann, J., Flynn, C., LeBlanc, S., Segal-Rozenhaimer, M., and Shinozuka, Y.: Exploring the elevated water vapor signal associated with the free tropospheric biomass burning plume over the southeast Atlantic Ocean, *Atmos. Chem. Phys.*, 21, 9643–9668, <https://doi.org/10.5194/acp-21-9643-2021>, 2021.
- 845 Pöhlker, M. L., Pöhlker, C., Quaas, J., Mülmenstädt, J., Pozzer, A., Andreae, M. O., Artaxo, P., Block, K., Coe, H., Ervens, B., Gallimore, P., Gaston, C. J., Gunthe, S. S., Henning, S., Herrmann, H., Krüger, O. O., McFiggans, G., Poulain, L., Raj, S. S., Reyes-Villegas, E., Royer, H. M., Walter, D., Wang, Y., and Pöschl, U.: Global organic and inorganic aerosol hygroscopicity and its effect on radiative forcing, *Nat. Commun.*, 14, 6139, <https://doi.org/10.1038/s41467-023-41695-8>, 2023.
- 850 Pósfai, M., Simonics, R., Li, J., Hobbs, P. V., and Buseck, P. R.: Individual aerosol particles from biomass burning in southern Africa: 1. Compositions and size distributions of carbonaceous particles, *J. Geophys. Res.*, 108, 8483, <https://doi.org/10.1029/2002jd002291>, 2003.
- Pratt, K. A., Murphy, S. M., Subramanian, R., DeMott, P. J., Kok, G. L., Campos, T., Rogers, D. C., Prenni, A. J., Heymsfield, A. J., Seinfeld, J. H., and Prather, K. A.: Flight-based chemical characterization of biomass burning aerosols within two prescribed burn smoke plumes, *Atmos. Chem. Phys.*, 11, 12549–12565, <https://doi.org/10.5194/acp-11-12549-2011>, 2011.
- 855 Rajapakshe, C., Zhang, Z., Yorks, J. E., Yu, H., Tan, Q., Meyer, K., Platnick, S., and Winker, D. M.: Seasonally transported aerosol layers over southeast Atlantic are closer to underlying clouds than previously reported, *Geophys. Res. Lett.*, 44, 5818–5825, <https://doi.org/10.1002/2017gl073559>, 2017.
- Redemann, J., Wood, R., Zuidema, P., Doherty, S. J., Luna, B., LeBlanc, S. E., Diamond, M. S., Shinozuka, Y., Chang, I. Y., Ueyama, R., Pfister, L., Ryoo, J.-M., Dobracki, A. N., da Silva, A. M., Longo, K. M., Kacenelenbogen, M. S., Flynn, C. J.,
- 860 Pistone, K., Knox, N. M., Piketh, S. J., Haywood, J. M., Formenti, P., Mallet, M., Stier, P., Ackerman, A. S., Bauer, S. E., Fridlind, A. M., Carmichael, G. R., Saide, P. E., Ferrada, G. A., Howell, S. G., Freitag, S., Cairns, B., Holben, B. N., Knobelspiesse, K. D., Tanelli, S., L'Ecuyer, T. S., Dzambo, A. M., Sy, O. O., McFarquhar, G. M., Poellot, M. R., Gupta, S., O'Brien, J. R., Nenes, A., Kacarab, M., Wong, J. P. S., Small-Griswold, J. D., Thornhill, K. L., Noone, D., Podolske, J. R., Schmidt, K. S., Pilewskie, P., Chen, H., Cochrane, S. P., Sedlacek, A. J., Lang, T. J., Stith, E., Segal-Rozenhaimer, M., Ferrare,
- 865 R. A., Burton, S. P., Hostetler, C. A., Diner, D. J., Seidel, F. C., Platnick, S. E., Myers, J. S., Meyer, K. G., Spangenberg, D. A., Maring, H., and Gao, L.: An overview of the ORACLES (ObseRvations of Aerosols above CLouds and their intEractionS) project: aerosol–cloud–radiation interactions in the southeast Atlantic basin, *Atmos. Chem. Phys.*, 21, 1507–1563, <https://doi.org/10.5194/acp-21-1507-2021>, 2021.
- Reid, J. S., Hobbs, P. V., Rangno, A. L., and Hegg, D. A.: Relationships between cloud droplet effective radius, liquid water content and droplet concentration for warm clouds in Brazil embedded in biomass smoke, *J. Geophys. Res.*, 104, 6145–6153, <https://doi.org/10.1029/1998JD200119>, 1999.
- Roberts, G. C. and Nenes, A.: A Continuous-Flow Streamwise Thermal-Gradient CCN Chamber for Atmospheric Measurements, *Aerosol Sci. Tech.*, 39, 206–221, <https://doi.org/10.1080/027868290913988>, 2005.

Roberts, G., Wooster, M. J., and Lagoudakis, E.: Annual and diurnal african biomass burning temporal dynamics, Biogeosciences, 6, 849–866, <https://doi.org/10.5194/bg-6-849-2009>, 2009.

Rosenberg, P. D., Dean, A. R., Williams, P. I., Dorsey, J. R., Minikin, A., Pickering, M. A., and Petzold, A.: Particle sizing calibration with refractive index correction for light scattering optical particle counters and impacts upon PCASP and CDP data collected during the Fennec campaign, Atmos. Meas. Tech., 5, 1147–1163, <https://doi.org/10.5194/amt-5-1147-2012>, 2012.

Ross, K. E., Piketh, S. J., Bruintjes, R. T., Burger, R. P., Swap, R. J., and Annegarn, H. J.: Spatial and seasonal variations in CCN distributions and the aerosol-CCN relationship over southern Africa, J. Geophys. Res., 108, 8481, <https://doi.org/10.1029/2002JD002384>, 2003.

Ryoo, J.-M., Pfister, L., Ueyama, R., Zuidema, P., Wood, R., Chang, I., and Redemann, J.: A meteorological overview of the ORACLES (ObseRvations of Aerosols above CLouds and their intEractionS) campaign over the southeastern Atlantic during 2016–2018: Part 2 – Daily and synoptic characteristics, Atmos. Chem. Phys., 22, 14209–14241, <https://doi.org/10.5194/acp-22-14209-2022>, 2022.

Samset, B. H., Myhre, G., Schulz, M., Balkanski, Y., Bauer, S., Bernsten, T. K., Bian, H., Bellouin, N., Diehl, T., Easter, R. C., Ghan, S. J., Iversen, T., Kinne, S., Kirkevåg, A., Lamarque, J.-F., Lin, G., Liu, X., Penner, J. E., Seland, Ø., Skeie, R. B., Stier, P., Takemura, T., Tsigaridis, K., and Zhang, K.: Black carbon vertical profiles strongly affect its radiative forcing uncertainty, Atmos. Chem. Phys., 13, 2423–2434, <https://doi.org/10.5194/acp13-2423-2013>, 2013.

Shinozuka, Y., Saide, P. E., Ferrada, G. A., Burton, S. P., Ferrare, R., Doherty, S. J., Gordon, H., Longo, K., Mallet, M., Feng, Y., Wang, Q., Cheng, Y., Dobracki, A., Freitag, S., Howell, S. G., LeBlanc, S., Flynn, C., Segal-Rosenhaimer, M., Pistone, K., Podolske, J. R., Stith, E. J., Bennett, J. R., Carmichael, G. R., da Silva, A., Govindaraju, R., Leung, R., Zhang, Y., Pfister, L., Ryoo, J.-M., Redemann, J., Wood, R., and Zuidema, P.: Modeling the smoky troposphere of the southeast Atlantic: a comparison to ORACLES airborne observations from September of 2016, Atmos. Chem. Phys., 20, 11491–11526, <https://doi.org/10.5194/acp-20-11491-2020>, 2020.

Slawinska, J., Grabowski, W. W., Pawlowska, H., and Morrison, H.: Droplet activation and mixing in large-eddy simulation of a shallow cumulus field, J. Atmos. Sci., 69, 444–462, <https://doi.org/10.1175/JAS-D-11-054.1>, 2012.

Trembath, J.: Airborne CCN measurements, University of Manchester, PhD Thesis, 2013.

Twohy, C. H., Anderson, J. R., Toohey, D. W., Andrejczuk, M., Adams, A., Lytle, M., George, R. C., Wood, R., Saide, P., Spak, S., Zuidema, P., and Leon, D.: Impacts of aerosol particles on the microphysical and radiative properties of stratocumulus clouds over the southeast Pacific Ocean, Atmos. Chem. Phys., 13, 2541–2562, <https://doi.org/10.5194/acp-13-2541-2013>, 2013.

Twomey, S.: The Influence of Pollution on the Shortwave Albedo of Clouds, J. Atmos. Sci., 34, 1149–1152, [https://doi.org/10.1175/1520-0469\(1977\)034<1149:TIOPOT>2.0.CO;2](https://doi.org/10.1175/1520-0469(1977)034<1149:TIOPOT>2.0.CO;2), 1977.

Wilcox, E. M.: Direct and semi-direct radiative forcing of smoke aerosols over clouds, Atmos. Chem. Phys., 12, 139–149, <https://doi.org/10.5194/acp-12-139-2012>, 2012.

Wood, R. and Bretherton, C. S.: Boundary Layer Depth, Entrainment, and Decoupling in the Cloud-Capped Subtropical and Tropical Marine Boundary Layer, *J. Climate*, 17, 3576–3588, [https://doi.org/10.1175/1520-0442\(2004\)017<3576:BLDEAD>2.0.CO;2](https://doi.org/10.1175/1520-0442(2004)017<3576:BLDEAD>2.0.CO;2), 2004.

Wood, R.: Stratocumulus clouds, *Mon. Weather Rev.*, 140, 2373–2423, <https://doi.org/10.1175/MWR-D-11-00121.1>, 2012.

Wu, H., Taylor, J. W., Szpek, K., Langridge, J. M., Williams, P. I., Flynn, M., Allan, J. D., Abel, S. J., Pitt, J., Cotterell, M. I., Fox, C., Davies, N. W., Haywood, J., and Coe, H.: Vertical variability of the properties of highly aged biomass burning aerosol transported over the southeast Atlantic during CLARIFY-2017, *Atmos. Chem. Phys.*, 20, 12697–12719, <https://doi.org/10.5194/acp-20-12697-2020>, 2020.

Zhang, J. and Zuidema, P.: The diurnal cycle of the smoky marine boundary layer observed during August in the remote southeast Atlantic, *Atmos. Chem. Phys.*, 19, 14493–14516, <https://doi.org/10.5194/acp-19-14493-2019>, 2019.

Zheng, G., Sedlacek, A. J., Aiken, A. C., Feng, Y., Watson, T. B., Raveh-Rubin, S., Uin, J., Lewis, E. R., and Wang, J.: Long-range transported North American wildfire aerosols observed in marine boundary layer of eastern North Atlantic, *Environ. Int.*, 139, 105680, <https://doi.org/10.1016/j.envint.2020.105680>, 2020.

Zheng, X., Albrecht, B., Jonsson, H. H., Khelif, D., Feingold, G., Minnis, P., Ayers, K., Chuang, P., Donaher, S., Rossiter, D., Ghate, V., Ruiz-Plancarte, J., and Sun-Mack, S.: Observations of the boundary layer, cloud, and aerosol variability in the southeast Pacific near-coastal marine stratocumulus during VOCALS-REx, *Atmos. Chem. Phys.*, 11, 9943–9959, <https://doi.org/10.5194/acp-11-9943-2011>, 2011.

Zuidema, P., Sedlacek III, A. J., Flynn, C., Springston, S., Delgadillo, R., Zhang, J., Aiken, A. C., Koontz, A., and Muradyan, P.: The Ascension Island Boundary Layer in the Remote Southeast Atlantic is Often Smoky, *Geophys. Res. Lett.*, 45, 4456–4465, <https://doi.org/10.1002/2017GL076926>, 2018.

Table S1. CLARIFY aircraft flights used in this study.

Flight	Date	Take-off time	Duration (hours)	Pollution layer (MBL/FT)	In-cloud data (min)	MBL structure type
C029	17/08/2017	08:56	3:23	MBL	5:38	e
C031	18/08/2017	11:59	3:43	MBL	4:00	d
C032	19/08/2017	10:01	3:43	MBL	33:24	f
C033	22/08/2017	08:54	3:45	FT	7:44	e
C036	24/08/2017	09:03	3:02	FT	23:40	e
C037	24/08/2017	13:46	3:07	FT	1:58	e
C038	25/08/2017	09:00	3:49	MBL&FT	14: 24	e
C039	25/08/2017	14:17	3:06	FT	7: 43	d
C042	28/08/2017	08:55	3:28	MBL&FT	14: 18	e
C044	29/08/2017	08:54	3:50	MBL&FT	36: 34	e
C045	29/08/2017	14:10	3:06	MBL&FT	33:12	e
C046	30/08/2017	08:45	4:06	MBL&FT	11:46	d
C047	01/09/2017	08:56	2:50	MBL&FT	2:21	d
C048	01/09/2017	13:26	3:57	MBL&FT	26:02	e
C049	02/09/2017	08:56	3:43	MBL&FT	11:11	d
C050	04/09/2017	13:28	3:46	MBL&FT	29:47	e
C051	05/09/2017	08:58	3:14	MBL&FT	38:41	d

Note: 1) “MBL” represents marine boundary layer, and “FT” represents free troposphere. 2) “d” represents a decoupled stratocumulus MBL, 3) “e” represents stratocumulus-over-cumulus MBL. 4) “f” represents a cumulus-capped MBL.

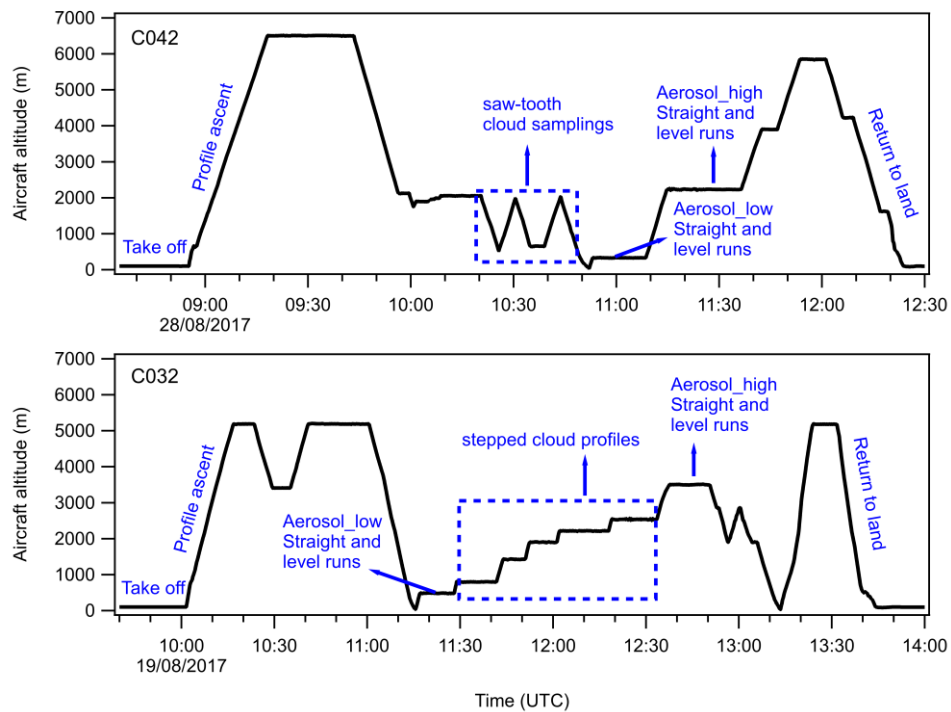


Figure S1. Example flight patterns from flight C042 (top) and C032 (bottom). The example flights provide an illustration of straight and level runs for aerosol characterization and saw-tooth and stepped profiles for cloud samplings.

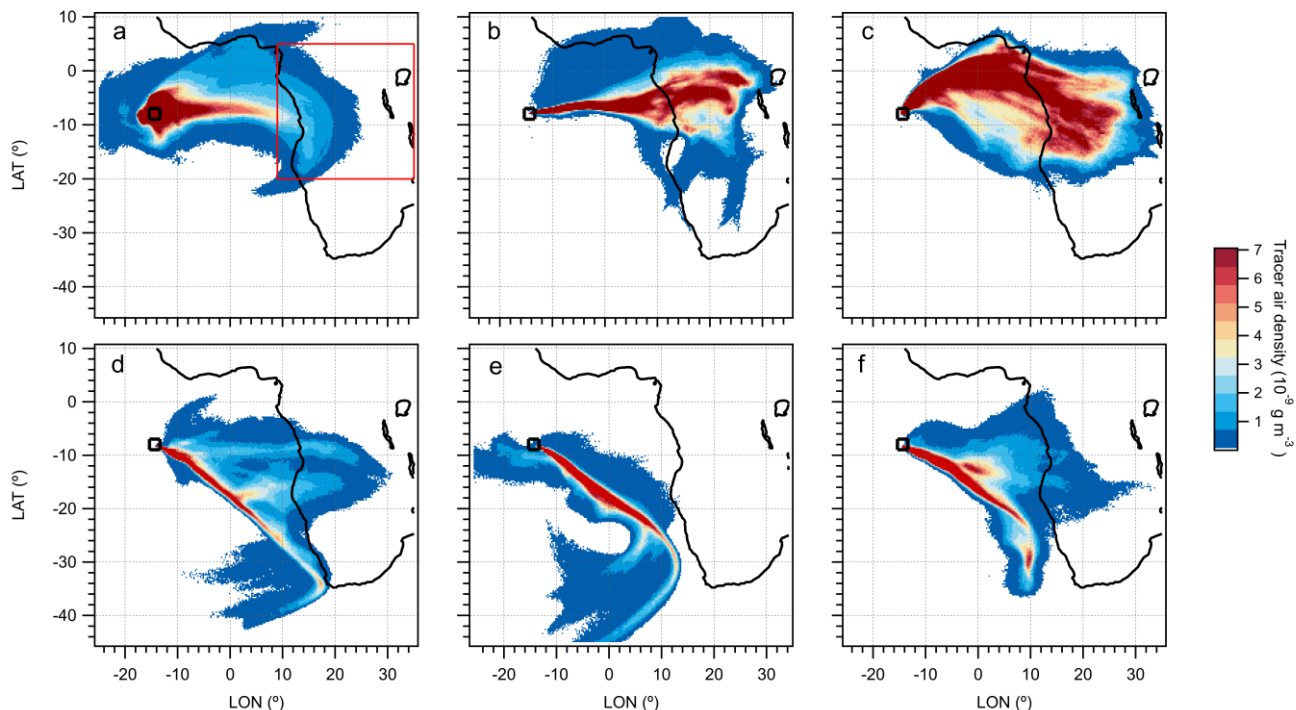


Figure S2. Examples of backward-dispersion fields from NAME simulations, illustrating representative horizontal footprints of original air parcels transported over the past 7 days before reaching Ascension Island area. a-c) Top and d-f) bottom panels are results from FT and MBL simulations respectively, for three example cases. Left two panels (a, d) are Case 1 released on 18 August 2017 at 12:00 UTC, middle two panels (b, e) are Case 2 released on 21 August 2017 at 12:00 UTC, and right two panels (c, f) are Case 3 released on 26 August 2017 at 12:00 UTC. The black boxes represent the release area around Ascension Island. The red box in a) represent the horizontal grids (20°S – 5°N, 9°W – 35°E) used for calculating fractional contributions of airmass from African continent to Ascension Island.

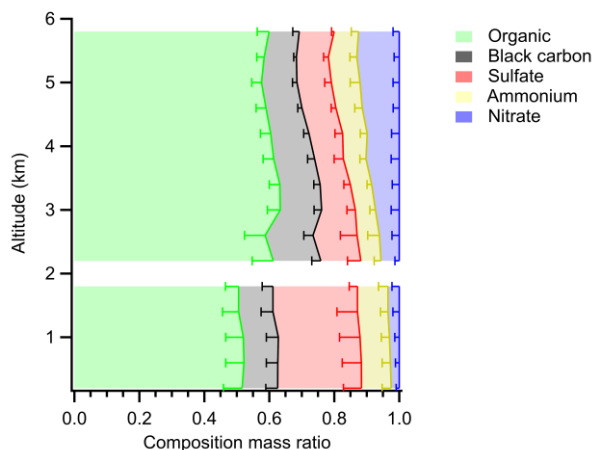


Figure S3. Average vertical profiles of out-of-cloud submicron aerosol chemical composition ratios for the flights used in this study, which is summarized from Wu et al. (2020). The width of colour bars represents the average mass ratios of different species in each vertical bin. The error bars represent standard deviation.

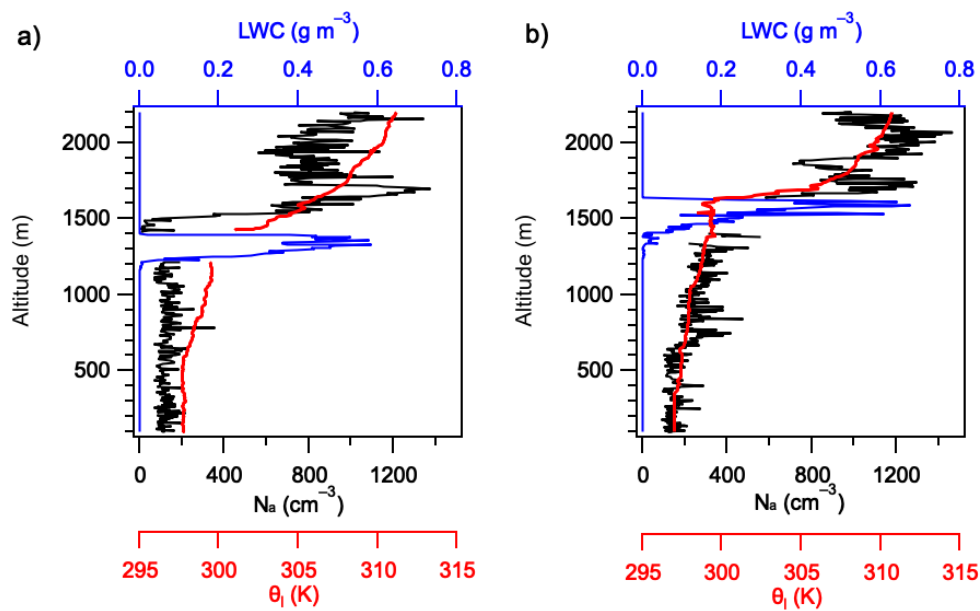


Figure S4. Vertical structures of N_a , LWC, and θ_l in example (a) separate and (b) contact profiles within the clean MBL.

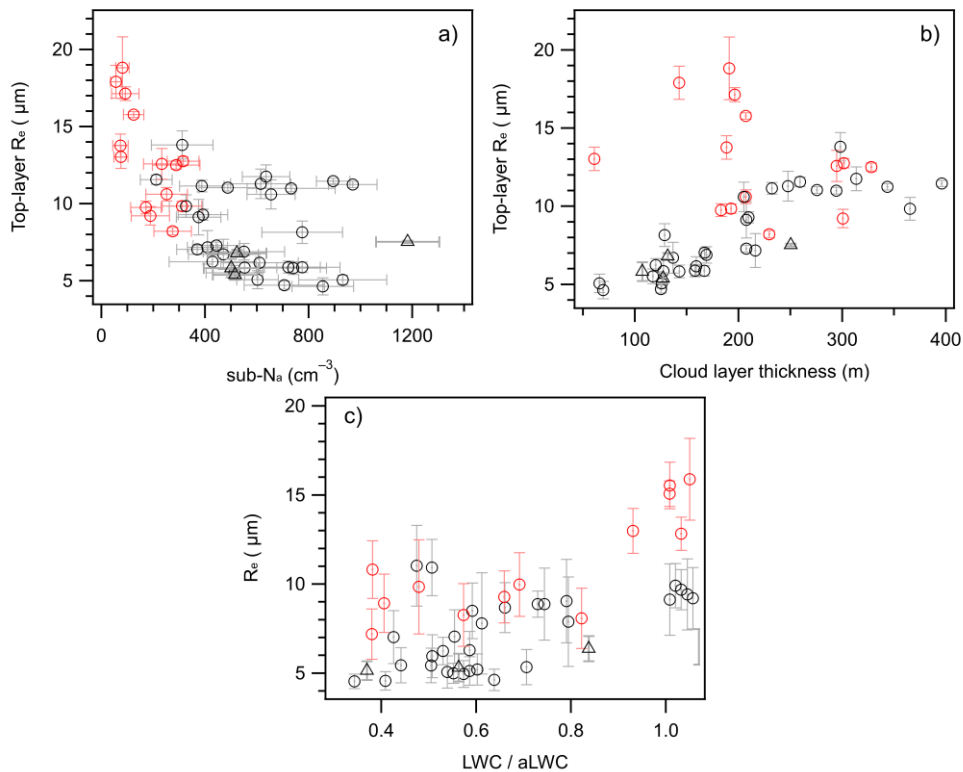


Figure S5. a) Relationship between cloud-top-layer R_e and sub- N_a for all analyzed profiles. b) Relationship between cloud-top-layer R_e and cloud layer depth for all analyzed profiles. c) Relationship between cloud-layer mean R_e and the average ratio of LWC/aLWC for all analyzed profiles. The markers and error bars represent the average values and standard deviation for each profile. Black triangle and circle markers are from the BB-impacted MBL in Periods 1 and 3 respectively, and red circle markers are from the clean MBL in Period 2..

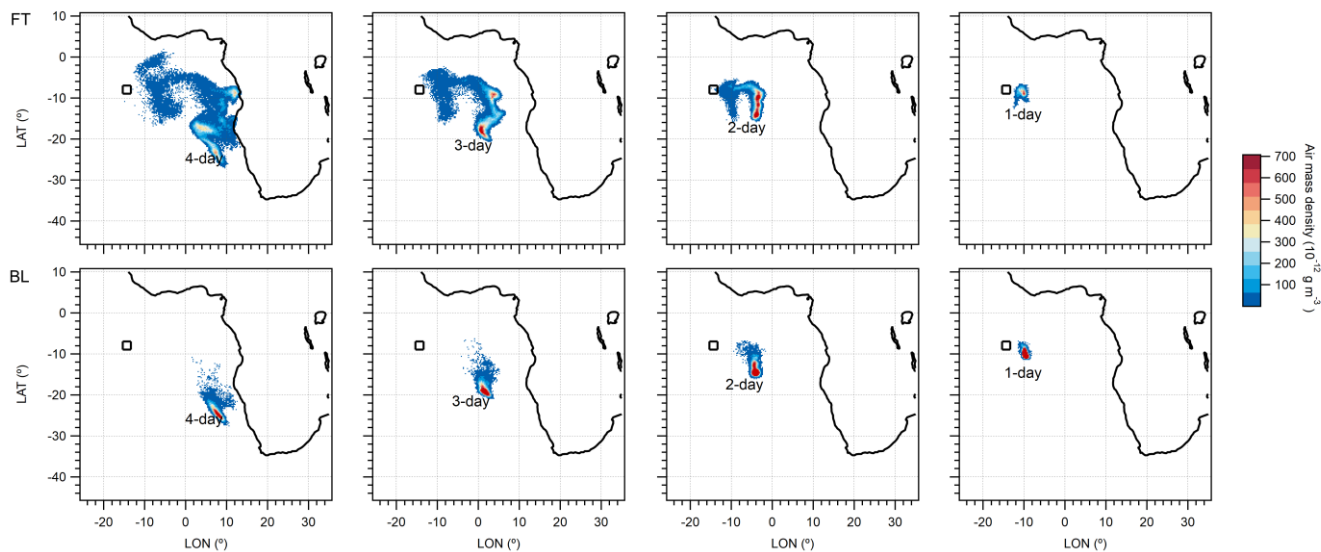


Figure S6. The instantaneous horizontal air parcel fields at each mid-day during 4 days prior to the NAME start time (air mass arrival at Ascension Island) for Case 1 (released from 18 August 2017 at 12:00 UTC). Top panels represent air parcel dispersion results attributed to the FT; bottom panel represent air parcel dispersion results attributed to the BL. From the left to right, the panels indicate the movement of air parcels before arrival at the release location around the Ascension Island area (black box). All plots are shown in same colour scale.

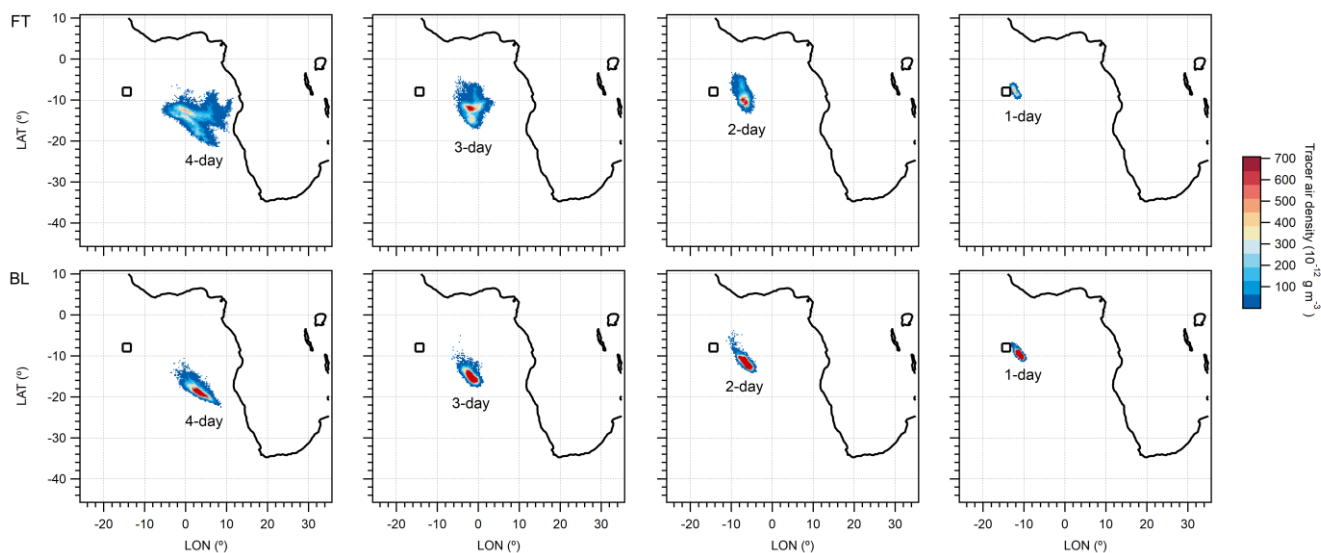


Figure S7. The instantaneous horizontal air parcel fields at each mid-day during 4 days prior to the NAME start time (arrival time at Ascension Island) for Case 3 (released from 26 August 2017 at 12:00 UTC). Top panels represent air parcel dispersion results attributed to the FT; bottom panel represent air parcel dispersion results attributed to the BL. From the left to right, the panels indicate the movement of air parcels before arrival at the release location around the Ascension Island area (black box). All plots are shown in same colour scale.

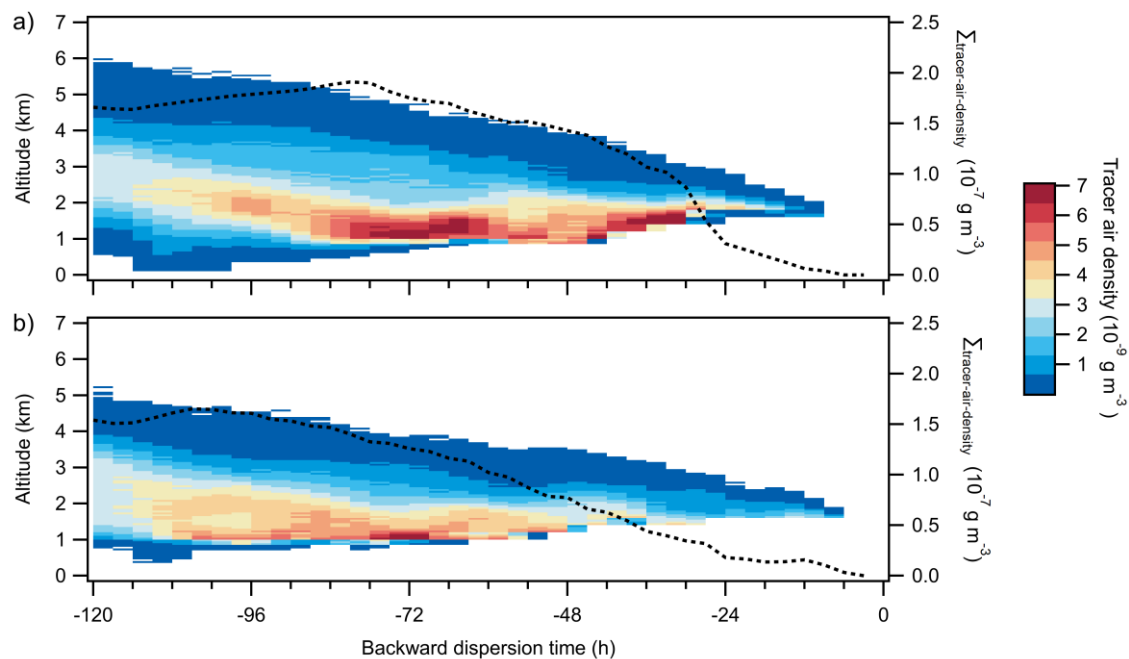


Figure S8. a-b) The 3-hourly time series of vertical distribution of air parcels dispersed into the FT region identified in the horizontal area of $20^{\circ}\text{S} - 0^{\circ}\text{N}$, $15^{\circ}\text{W} - 12^{\circ}\text{E}$ (see red box in Fig. 6), in terms of backward dispersion time from 3 to 120 h. The right axis (dashed black lines) are the corresponding cumulative exchange amounts of air parcels dispersed between the FT and MBL along backward simulations. a) is for Case 1 and b) is for Case 3.

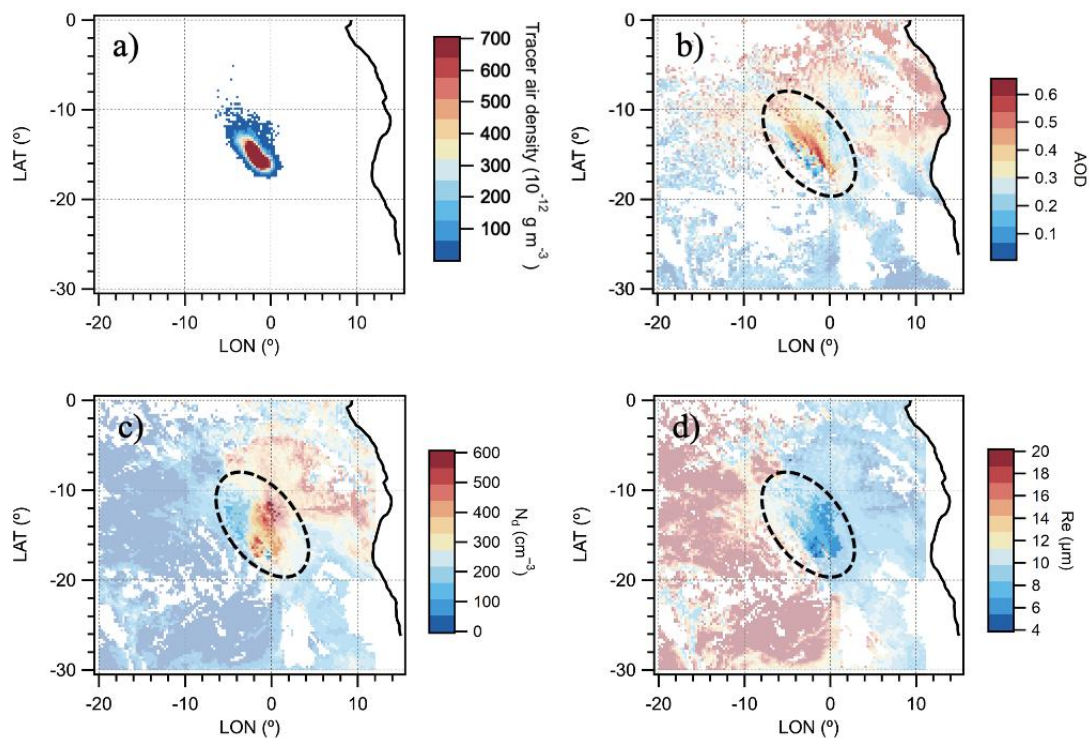


Figure S9. Example co-location relationship between NAME instantaneous horizontal footprints and SEVIRI-retrieved aerosol and cloud fields. a) is an example BL instantaneous horizontal footprint from NAME simulations. b-d) are the example SEVIRI-retrieved AOD, N_d and R_e at the time of the example BL horizontal footprint. The parts of the scene (b-d) that are co-located with the contemporaneous horizontal footprints are highlighted. The air-density-weighted average SEVIRI-retrievals were calculated for these co-located areas at 3-hourly backward step.