



1 **Contributions of primary anthropogenic sources and rapid secondary**
2 **transformations to organic aerosol pollution in Nanchang, Central China**

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10 **Abstract**

11 Owing to the complex composition of organic aerosols (OAs), it is challenging to elucidate their sources and
12 dynamics, particularly in urban environments in China, where natural and anthropogenic influences converge.
13 We attempted to clarify the relative contributions of primary emissions and secondary formations to urban
14 OAs and confirm the sources and influencing factors of OA pollution. To achieve this, we conducted a
15 comprehensive analysis of major polar organic compounds in fine particulate matter (PM_{2.5}) samples collected
16 over a year in Nanchang, Central China. The results indicated that the concentrations of fatty acids, fatty
17 alcohols, and saccharides were relatively high, whereas lignin and resin products, sterols, glycerol, hydroxy
18 acids, and aromatic acids were present at low concentrations. An analysis of molecular characteristics and
19 concentration ratios revealed that they originate from anthropogenic and natural sources. Using the tracer-
20 based method, we observed that the primary organic carbon (POC) and primary organic aerosols (POA)
21 contributed 53% of OC and 21% of PM_{2.5} mass, respectively, compared with a mere 8% and 4% from
22 secondary organic carbon (SOC) and secondary organic aerosols (SOA). Anthropogenic sources were the most
23 dominant determinant, contributing approximately 89% of POC and POA and 60% of SOC and SOA.
24 Seasonal variations indicated that biogenic emissions exerted a stronger influence during spring and summer,
25 whereas anthropogenic emissions were more pronounced in autumn and winter. Short-term winter pollution



26 episodes were characterized by rapid secondary transformation, promoted by elevated primary emissions and
27 favorable oxidation conditions, including increased light intensity and nitrogen oxides.

28

29 **1 Introduction**

30 Fine particulate matter (PM), specifically those with an aerodynamic diameter less than 2.5 μm ($\text{PM}_{2.5}$), affects
31 the environment and climate, posing severe threats to human health (Huebert et al., 2003; Kanakidou, 2005;
32 Riipinen et al., 2012; Shiraiwa et al., 2017; Yazdani et al., 2021). This is evident in China, which has become a
33 focal point for $\text{PM}_{2.5}$ pollution, characterized by recurrent haze episodes that have intensified since 2011 (Cao
34 et al., 2012; Zhang et al., 2012; Huang et al., 2014). To tackle this pressing challenge, the Chinese government
35 implemented the "Air Pollution Prevention and Control Action Plan" in 2013, revisions to the "Air Pollution
36 Prevention and Control Law" in 2014, and the "Three-year Action Plan for Winning the Blue Sky Defense
37 War" in 2018. Such initiatives have significantly reduced $\text{PM}_{2.5}$ levels, with concentrations below 35 $\mu\text{g m}^{-3}$ as
38 of 2020. However, the challenge persists, particularly during winter, as exemplified by the winter of 2023–
39 2024, when $\text{PM}_{2.5}$ concentrations in many cities exceeded the national air quality standards (Figures 1a and S1).
40 This underscores an urgent need for effective air quality management strategies (An et al., 2019; Cao and Cui,
41 2021; Chen et al., 2021; Wu et al., 2021; Zhang et al., 2024). To address urban $\text{PM}_{2.5}$ pollution, a
42 comprehensive investigation of pollutant components, an identification of pollution sources, and an evaluation
43 of influencing factors are required.

44 Organic components make up an important part of $\text{PM}_{2.5}$, accounting for 30%–50% of its mass, with
45 primary and secondary sources (Huang et al., 2014; Zhang et al., 2016; Haque et al., 2022). Primary organic
46 carbon (POC) or primary organic aerosol (POA) is directly emitted from various sources, including biomass
47 burning, coal combustion, vehicle exhaust, cooking, plant debris, and fungal spores (Guo et al., 2012; Yazdani
48 et al., 2021). Secondary organic carbon (SOC) or secondary organic aerosol (SOA) forms in the atmosphere
49 through the photooxidation of biogenic and anthropogenic volatile organic compounds (BVOCs and AVOCs)
50 (Lewandowski et al., 2008; Stone et al., 2009). Common biogenic precursors include hemiterpenes,
51 monoterpenes, and sesquiterpenes, whereas typical anthropogenic precursors include toluene and polycyclic
52 aromatic hydrocarbons (PAHs) (Claeys et al., 2004; Al-Naiema and Stone, 2017; Jaoui et al., 2019). Owing to



53 the complex composition of OAs, their diverse emission sources, and the impact of meteorological conditions
54 and photochemical oxidation processes, the identification of OA sources is complicated (Ding et al., 2013;
55 Zhang et al., 2024). For years, research on the composition and source apportionment of OAs has attracted
56 considerable attention. However, a unified conclusion regarding OA sources, particularly in China's intricate
57 urban environments influenced by natural and anthropogenic factors, remains elusive (Wang et al., 2006; Fu et
58 al., 2008; Guo et al., 2012; Ding et al., 2014; Xu et al., 2022). Various source apportionment methods have
59 been employed. Such methods include the elemental carbon and water-soluble OC methods (EC-based and
60 WSOC-based method; Xu et al., 2021), compound tracer method (tracer-based method; Ding et al., 2014; Ren
61 et al., 2021; Haque et al., 2023), isotope signature method (Tang et al., 2022; Xu et al., 2022 and 2023; Zhang
62 et al., 2024), and source apportionment models, such as the chemical mass balance (CMB) and positive matrix
63 factorization models (Xu et al., 2021; Zhang et al., 2023). However, these methods yielded different source
64 contributions. The relative significance of primary emissions versus secondary formation in urban OAs
65 continues to be controversial. Further investigation is required to clarify the sources of OA pollution and the
66 influencing factors.

67 Thus, this paper presents a comprehensive assessment of OAs in Nanchang, a city in Central China
68 characterized by moderate $PM_{2.5}$ pollution levels that reflect broader urban atmospheric conditions across
69 China. Employing the EC- and tracer-based methods, as well as the CMB model, we quantitatively evaluated
70 the contributions of POC and SOC to OC, as well as the contributions of the corresponding POA and SOA to
71 the mass of $PM_{2.5}$. The results showed that urban OAs were predominantly influenced by anthropogenic
72 sources, with primary contributions exceeding secondary contributions. Based on a continuous year-long
73 observation (November 1, 2020 to October 31, 2021, 365 daily samples), we discovered that the anthropogenic
74 contributions significantly increased in autumn and winter. However, the biogenic contributions increased in
75 spring and summer. Short-term winter pollution episodes were promoted by rapid secondary transformation,
76 primarily due to high primary emissions and favorable oxidation conditions, including increased light intensity
77 and nitrogen oxides (NO_x). This study integrates multiple source apportionment methods and accounts for the
78 seasonal characteristics of OA pollution, providing a robust framework for understanding urban air quality
79 dynamics. The findings have implications for governmental strategies for mitigating air pollution and



80 preventing haze episodes in Nanchang and its environs. The insights obtained serve as a future reference for
81 investigating the impacts of organic compounds in PM_{2.5} on visibility, public health, and climate change.

82 **2 Materials and methods**

83 **2.1 Sampling sites and sample collection**

84 The study area was Nanchang, Central China, with sampling at the East China University of Technology
85 (ECUT) located in the northwest region of the city (28.72°N, 115.83°E; Figure 1b). As detailed in our previous
86 investigations (Guo et al., 2024a and 2024b), the sampling site was in a mixed-use area characterized by
87 educational, commercial, transportation, and residential activities, devoid of significant local pollution sources.
88 The sampler was strategically placed on the rooftop of a six-story building, approximately 20 m in height,
89 ensuring an unobstructed sampling environment.

90 Sampling was conducted from November 1, 2020, to October 31, 2021, with daily PM_{2.5} sample collection.
91 An 8-inch × 10-inch quartz fiber filter (Pall Tissuquartz, USA) was used in a high-volume air sampler for
92 sample collection. The meteorological parameters and gaseous pollutant data for the sampling period were
93 sourced from publicly available online monitoring platforms (<https://weatherandclimate.info> and
94 <http://www.aqistudy.cn/>). Text S1 and Figure S2 provide a comprehensive overview of the prevalent
95 meteorological conditions and air quality during the sampling period.

96 **2.2 Chemical analyses**

97 The OC and EC concentrations in the PM_{2.5} samples were quantified using the Desert Research Institute Model
98 2001 Carbon Analyzer, following the thermal/optical reflectance protocol established by the Interagency
99 Monitoring of Protected Visual Environments (IMPROVE). A 1.0-cm² filter sample was placed in a quartz
100 boat in the analyzer and subjected to incremental heating at predetermined temperatures. Repeated analyses
101 were performed, demonstrating an analytical uncertainty of ±10%.

102 We employed methodologies outlined by Wang et al. (2005, 2006) and Fu et al. (2008, 2010) for the
103 extraction and derivatization of organic compounds. The filter samples underwent three consecutive ultrasonic
104 extractions using a dichloromethane–methanol mixture (2:1 v/v), followed by the concentration and drying of
105 the resulting extracts. Prior to instrumental analyses, N,O-bis-(trimethylsilyl) trifluoroacetamide and pyridine
106 were introduced to derivatize the polar compounds in the extract, with C13 n-alkanes added as internal



107 standards for quantitative analyses.

108 The identification and quantification of organic compounds were achieved via gas chromatography-mass
109 spectrometry (GC-MS), using a Thermo Scientific TRACE GC coupled to a Thermo Scientific ISQ QD single
110 quadrupole mass spectrometer. GC separation was performed using a DB-5MS fused silica capillary column.
111 Text S2 and Table S1 present the parameters for the GC-MS analysis, including temperature elevation
112 procedures, qualitative and quantitative methods for the compounds, and quality assurance and control
113 protocols. Polar and nonpolar organic compounds in the extracts were simultaneously analyzed; however, this
114 study focused on the results regarding the polar compounds; the findings related to nonpolar compounds are
115 provided in our earlier publications (Guo et al., 2024a and 2024b).

116 **2.3 Source apportionment methods**

117 To assess the contributions of various primary and secondary sources to OC and PM_{2.5}, we employed the
118 tracer-based approach, a well-established method for source apportionment. This study focused on two types
119 of organic tracers that serve as indicators for POC and SOC or POA and SOA. By integrating the reported
120 ratios of specific tracers to OC and aerosol mass concentrations from various emission sources with the actual
121 tracer concentrations measured in the present samples, we calculated the contributions of these sources to OC
122 and PM_{2.5}. The calculation equation was as follows (Kleindienst et al., 2007 and 2012):

$$123 \quad [\text{POC}] = \frac{\sum_i [\text{tr}_i]}{f_{\text{POC}}}, \quad (1)$$

124

$$125 \quad [\text{POA}] = \frac{\sum_i [\text{tr}_i]}{f_{\text{POA}}}, \quad (2)$$

126

$$127 \quad [\text{SOC}] = \frac{\sum_i [\text{tr}_i]}{f_{\text{SOC}}}, \quad (3)$$

128

$$129 \quad [\text{SOA}] = \frac{\sum_i [\text{tr}_i]}{f_{\text{SOA}}}, \quad (4)$$

130 where $\sum_i [\text{tr}_i]$ is the total concentration of the selected tracers in the sample, denoting representative
131 compounds from specific emission categories; f_{POC} and f_{POA} are the mass fractions of the tracers in OC and



132 $PM_{2.5}$ from primary emissions, respectively. Similarly, f_{SOC} and f_{SOA} are the mass fractions of the tracers in OC
133 and $PM_{2.5}$ from secondary emissions, respectively. The calculated [POC], [POA], [SOC], and [SOA] are the
134 contributions of different primary and secondary sources to OC and $PM_{2.5}$. Here, we identified mass fractions
135 for representative tracers from several significant primary and secondary emission sources. The primary
136 sources included biomass burning, coal combustion, vehicle exhaust, cooking activities, plant debris, and
137 fungal spores. The secondary sources were categorized into biogenic (isoprene, monoterpenes, and
138 sesquiterpenes) and anthropogenic (toluene and naphthalene) sources. Table S2 presents information on the
139 representative tracers, their corresponding f_{OC} and f_{OA} , and relevant references.

140 We employed the CMB model (version 8.2) provided by the US Environmental Protection Agency to verify
141 the results calculated using the tracer-based method. This model assumes a CMB between the emission sources
142 and environmental receptors. Thus, the mass of pollutants is not lost during transport from the source to the
143 receptor, and the measured chemical component concentration at the receptor is the linear superposition of the
144 contribution of each source class to the concentration (Stone et al., 2009). Furthermore, we employed the EC-
145 based method to estimate the total POC and SOC concentrations, enabling a comparison with the POC and
146 SOC concentrations obtained via the tracer-based approach. This method uses EC as a tracer for POC and
147 approximates the observed minimum ratio of OC to EC $(OC/EC)_{min}$ as an equivalent to the primary emission
148 OC/EC value $(OC/EC)_{pri}$ (Turpin and Huntzicker, 1995; Castro et al., 1999). The equation used for estimating
149 SOC is

$$150 \text{SOC}_{\text{EC-based}} = \text{OC} - \text{EC} \times (OC/EC)_{\text{min}}. \quad (5)$$

151 Here, the minimum OC/EC values observed in spring, summer, autumn, and winter were 2.01, 2.20, 2.15, and
152 2.31, respectively.

153 **3 Results and discussion**

154 **3.1 Carbonaceous components**

155 The OC and EC concentrations were in the ranges of 1.01–13.66 $\mu\text{g m}^{-3}$ (mean: $5.88 \pm 2.57 \mu\text{g m}^{-3}$) and 0.27–
156 4.55 $\mu\text{g m}^{-3}$ (mean: $1.75 \pm 0.88 \mu\text{g m}^{-3}$), respectively. Seasonal variations in the OC and EC concentrations
157 closely mirrored those observed for $PM_{2.5}$, with elevated levels in autumn and winter, and diminished levels in
158 spring and summer (Figure 2). The OC/EC ratio is a widely employed metric for characterizing emissions from



159 fossil fuels and biomass combustion. OC/EC values below 1.1 typically indicate vehicle exhaust emissions;
160 values within the range of 2–3 suggest coal combustion emissions, and values above 7 indicate biomass
161 burning (Saarikoski et al., 2008). Here, the OC/EC range was 2.01–16.95, with an average of 3.86 ± 2.05 .
162 These OC/EC ratios fell within the ranges associated with coal and biomass burning emissions. This suggests
163 that the carbonaceous material was predominantly derived from combustion. Furthermore, owing to its
164 stability and resistance to chemical transformations in the atmosphere, EC is frequently used as a tracer for
165 primary emissions. The OC/EC ratio is a useful tool for assessing the relative contributions of primary and
166 secondary sources (Castro et al., 1999; Turpin and Huntzicker, 1995). Generally, an OC/EC value above 2
167 indicates a predominant contribution of secondary sources to OC, whereas values below 2 suggest greater
168 contributions from primary sources (Kunwar and Kawamura, 2014). Here, the majority of the OC/EC values
169 exceeded 2, implying that the OC may have been significantly influenced by secondary sources. The OC/EC
170 values were relatively higher in summer (4.33 ± 2.07) and winter (4.18 ± 2.57) compared with those observed
171 in spring (3.36 ± 1.98) and autumn (3.59 ± 1.22). This indicates an increased contribution to OC from
172 secondary sources during the summer and winter.

173 **3.2 Major polar components**

174 **3.2.1 Fatty acids**

175 A range of homologous straight-chain C10:0–C32:0 fatty acids (FAs) and unsaturated C16:1 and C18:1 FAs
176 were detected in the PM_{2.5} samples (Figure 3a). Their distribution exhibited a strong even carbon number
177 predominance, as indicated by a Carbon Preference Index (CPI) of 7.59 ± 7.42 (Figure 4a), with peaks at
178 C16:0 and C18:0. The total FA concentration was in the range of 3.24–657.86 ng m⁻³, with an average
179 concentration of 196.50 ± 110.92 ng m⁻³. Similar molecular distribution patterns and concentrations have been
180 documented in urban aerosol studies across China (59 – $2,090$ ng m⁻³; Wang et al., 2006; Fu et al., 2008; Haque
181 et al., 2019; Fan et al., 2020). This average concentration generally exceeds those observed for coastal and
182 marine aerosol samples (0.1 – 160 ng m⁻³; Kawamura et al., 2003; Wang et al., 2007; Fu et al., 2011). High
183 molecular weight FAs (HFAs, \geq C20:0) are typically derived from the waxes of terrestrial higher plants or
184 biomass burning, whereas low molecular weight FAs (LFAs, $<$ C20:0) originate from more diverse sources,
185 including vascular plants, microorganisms, marine phytoplankton, and kitchen emissions (Fu et al., 2010).



186 Consequently, the ratio of LFAs to HFAs (LFAs/HFAs) is an indicator of the relative contributions from
187 various sources. Here, the LFA/HFA range was 0.34–10.67, with an average of 2.45 ± 1.53 (Figure 4b),
188 suggesting that the FAs may have originated from a mixture of natural and anthropogenic sources. The specific
189 FAs C16:0 and C18:0 are associated with different sources. C16:0 is primarily derived from biomass burning
190 or plant emissions, whereas C18:0 predominantly originates from vehicle exhaust, cooking emissions, or road
191 dust (Wang et al., 2007; Fu et al., 2010). Thus, the ratio of C18:0 to C16:0 (C18:0/C16:0) is frequently used to
192 assess the FA source. A C18:0/C16:0 ratio below 0.25 suggests a primary contribution from biomass burning or
193 plant emissions, whereas ratios between 0.25 and 0.5 indicate a predominant influence from vehicle exhaust.
194 C18:0/C16:0 ratios within the range of 0.5–1 suggest a significant contribution from cooking emissions or road
195 dust (Rogge et al., 2006). Here, the C18:0/C16:0 varied from 0.05 to 9.53, with an average of 0.59 ± 0.88
196 (Figure 4c), indicating that the FAs in Nanchang originated from a mixture of sources.

197 The average concentrations of unsaturated FAs, specifically C16:1 and C18:1, were 6.94 ± 4.35 and $3.79 \pm$
198 2.86 ng m^{-3} , respectively. Such unsaturated FAs are primarily derived from biomass burning, plant leaf
199 emissions, cooking activities, and release from marine organisms (Kawamura and Gagosian, 1987; Rogge et
200 al., 1993; Nolte et al., 1999). Upon emission into the atmosphere, unsaturated FAs undergo rapid oxidation by
201 ozone (O_3), hydrogen peroxide (H_2O_2), or hydroxyl (OH) radicals. Consequently, the ratio of unsaturated FAs
202 to saturated FAs, represented as $(\text{C16:1} + \text{C18:1})/(\text{C16:0} + \text{C18:0})$, is an indicator of the reactivity and degree
203 of aging of unsaturated FAs (Rudich et al., 2007; Kawamura and Gagosian, 1987). Here, this ratio was $0.12 \pm$
204 0.06 (Figure 4d), indicating that the unsaturated FAs underwent significant photochemical degradation and that
205 secondary organic compounds may be common in $\text{PM}_{2.5}$.

206 The LFAs/HFAs and C18:0/C16:0 were relatively low during autumn and winter compared with spring and
207 summer (Figures 4b and 4c). This indicated that in autumn and winter, FAs were significantly influenced by
208 terrestrial plant or biomass burning, whereas in spring and summer, they were more influenced by vehicular
209 emissions, marine phytoplankton, and cooking emissions. Conversely, the $(\text{C16:1} + \text{C18:1})/(\text{C16:0} + \text{C18:0})$
210 values were relatively higher in autumn and winter than in spring and summer (Figure 4d). This suggested that
211 unsaturated FAs experienced a greater degree of photochemical degradation during the warm seasons.

212 3.2.2 Fatty alcohols



213 A series of straight-chain n-alkanols (C14–C32) were detected (Figure 3b), exhibiting a predominance of even
214 carbon numbers, as indicated by a CPI of 10.28 ± 17.05 (Figure 4e). Low molecular weight alcohols (LMW_{alc} ,
215 $\leq \text{C}20$) exhibited peaks at C16 and C18, and high MW alcohols (HMW_{alc} , $> \text{C}20$), exhibited peaks at C26,
216 C28, and C30 (Figure 3b). The total concentration of n-alkanols was in the range of 5.80–572.01 ng m^{-3}
217 (average = $113.99 \pm 92.50 \text{ ng m}^{-3}$). This concentration fell within the range of reported n-alkanol
218 concentrations in urban aerosols in China ($3.1\text{--}1,301.0 \text{ ng m}^{-3}$; Wang et al., 2006; Fu et al., 2008; Haque et al.,
219 2019; Fan et al., 2020) and generally exceeded that of coastal and marine aerosol samples ($0.1\text{--}19.7 \text{ ng m}^{-3}$;
220 Kawamura et al., 2003; Fu et al., 2011). HMW_{alc} mainly originates from higher plant leaf waxes, loess
221 deposits, or biomass burning, and LMW_{alc} primarily originates from soil and marine microorganisms
222 (Simoneit, 2002; Kawamura et al., 2003). Here, the LMW_{alc} concentration ($75.12 \pm 61.28 \text{ ng m}^{-3}$) relatively
223 exceeded that of HMW_{alc} ($39.09 \pm 34.07 \text{ ng m}^{-3}$), yielding an average LMW/HMW of 2.50 ± 1.73 (Figure 4f),
224 indicating a mixture of n-alkanols from soil, marine organisms, and biomass burning. The LMW/HMW values
225 were relatively lower in autumn and winter and higher in spring and summer (Figure 4f). This suggests an
226 increased contribution of soil and marine organisms to n-alkanols in spring and summer. Contributions from
227 plant or biomass burning increased in autumn and winter.

228 3.2.3 Saccharides

229 We investigated three saccharide classes: anhydrosugars, primary sugars, and sugar alcohols. Among the
230 anhydrosugars, levoglucosan had the highest concentration ($102.48 \pm 78.63 \text{ ng m}^{-3}$). We observed markedly
231 lower concentrations of mannosan ($4.33 \pm 4.04 \text{ ng m}^{-3}$) and galactosan ($2.41 \pm 2.67 \text{ ng m}^{-3}$) (Figure 3c).
232 Levoglucosan, a prominent biomarker for biomass burning, has been extensively investigated in urban aerosol
233 samples, with concentrations in the range of 22–2,706 ng m^{-3} (Wang and Kawamura, 2005; Wang et al., 2006;
234 Fu et al., 2010; Haque et al., 2019; Fan et al., 2020), indicating the significant influence of biomass burning on
235 the urban atmosphere. The widespread detection of levoglucosan across diverse environments, including
236 suburban ($10\text{--}482 \text{ ng m}^{-3}$; Yttri et al., 2007; Fu et al., 2008), marine ($0.2\text{--}30 \text{ ng m}^{-3}$; Simoneit et al., 2004a; Fu
237 et al., 2011), and polar ($0\text{--}3 \text{ ng m}^{-3}$; Stohl et al., 2007; Fu et al., 2009) regions, suggests its potential for long-
238 range atmospheric transport. Seasonal variations of anhydrosugars exhibited consistent patterns, with elevated
239 concentrations in autumn and winter and reduced levels in spring and summer (Figure 3c). This indicated



240 increased biomass burning during the cold months. The sources were characterized using levoglucosan-to-
241 mannosan (L/M) and mannosan-to-galactosan (M/G) ratios. Research suggests distinct L/M ranges for
242 different biomass-burning sources: softwood (3–10), hardwood (13–35), and agricultural crop burning (40–56)
243 (Oros and Simoneit, 2001; Sheesley et al., 2003; Engling et al., 2014). Coal combustion can produce
244 significant amounts of levoglucosan (Sheesley et al., 2003; Fabbri et al., 2009), typically with higher L/M
245 values (12–189) because of the lower cellulose content compared with that of plant materials (Rybicki et al.,
246 2020a and 2020b). Conversely, M/G values are generally higher for softwood and hardwood combustion at
247 3.6–7.0 and 1.2–2.0, respectively. However, M/G values are relatively lower for agricultural crop burning (0.3–
248 0.6) (Oros and Simoneit, 2001; Sheesley et al., 2003). Here, most samples exhibited relatively high L/M ratios
249 (4.63–466.61, average = 40.22 ± 49.82), and relatively low M/G ratios (0.30–9.41, average = 2.34 ± 1.36). The
250 L/M and M/G ratios were similar to those of previous observations in the urban area of Nanchang (L/M = 7.9–
251 359.1; average = 59.9; Zhu et al., 2022) and fell within the ranges for crop residue burning and coal
252 combustion. This suggests that the dehydrosugars in the PM_{2.5} of Nanchang were likely primarily derived from
253 crop residue burning and coal combustion. In the autumn and winter, the contributions of crop residue burning
254 and coal combustion were more prominent, as evidenced by the higher L/M ratios and lower M/G ratios
255 observed compared with those in the spring and summer (Figures 4g and 4h).

256 Primary sugars and sugar alcohols serve as tracers for primary biological aerosol particles, originating from
257 biogenic emissions, including the release of microorganisms, plants, and flowers, as well as the resuspension
258 of surface soils and unpaved road dust containing biological materials (Graham et al., 2003; Simoneit et al.,
259 2004b; Yttri et al., 2007). Furthermore, biomass burning is a significant source of primary sugars and sugar
260 alcohols (Fu et al., 2012). We identified nine primary sugars (ribose, fructose, galactose, glucose, sucrose,
261 lactulose, maltose, turanose, and trehalose; Figure 3c). Glucose ($14.90 \pm 6.62 \text{ ng m}^{-3}$) and sucrose ($14.68 \pm$
262 6.48 ng m^{-3}) exhibited the highest concentrations, followed by fructose ($10.82 \pm 4.54 \text{ ng m}^{-3}$), trehalose (7.92
263 $\pm 4.40 \text{ ng m}^{-3}$), and maltose ($5.39 \pm 3.43 \text{ ng m}^{-3}$). The concentrations of galactose ($1.88 \pm 1.24 \text{ ng m}^{-3}$), ribose
264 ($2.87 \pm 2.18 \text{ ng m}^{-3}$), turanose ($0.87 \pm 0.61 \text{ ng m}^{-3}$), and lactulose ($0.83 \pm 0.54 \text{ ng m}^{-3}$) were comparatively
265 low. We detected four sugar alcohols (mannitol, arabitol, pinitol, and inositol). Mannitol ($9.88 \pm 4.57 \text{ ng m}^{-3}$)
266 exhibited the highest concentration, followed by pinitol ($6.05 \pm 4.23 \text{ ng m}^{-3}$) and arabitol ($4.96 \pm 3.39 \text{ ng m}^{-3}$).



267 The concentration of inositol ($2.05 \pm 1.66 \text{ ng m}^{-3}$) was low.

268 Levoglucosan serves as a specific biomarker for biomass burning, enabling the assessment of biomass-
269 burning contributions to sugar compounds through the ratio of levoglucosan to OC (Lev/OC) and the
270 proportion of levoglucosan in the total sugar compounds (Lev%). Typically, a Lev/OC ratio above 0.048 and a
271 Lev% above 68% indicate a dominance of biomass burning. Conversely, values below these thresholds suggest
272 that, in addition to biomass burning, other sources, such as biogenic emissions, significantly contribute to the
273 presence of sugar compounds (Yan et al., 2019 and references therein). Here, the Lev% and Lev/OC were
274 $53.57\% \pm 17.93\%$ and 0.016 ± 0.009 (Figures 4i and 4j), respectively, indicating that the sugar compounds
275 were influenced by biomass burning and biogenic emissions, particularly the primary sugars and sugar
276 alcohols. Primary sugars, such as glucose and trehalose, as well as the mannitol in sugar alcohol, positively
277 correlated with levoglucosan (Figure S3a). This confirmed that biomass burning may be a potential source of
278 primary sugars and sugar alcohols. However, the correlation was not statistically significant, suggesting that
279 biological sources remain an important contributor to these compounds in Nanchang. The Lev/OC and Lev%
280 values were higher in autumn and winter than in spring and summer (Figures 4i and 4j), further indicating that
281 biomass-burning contributions increased during the cold months.

282 **3.3 Minor polar components**

283 **3.3.1 Lignin, resin products, and sterols**

284 Here, we detected four types of lignin and resin-derived compounds: 4-hydroxybenzoic acid, dehydroabietic
285 acid, vanillic acid, and syringic acid (Figure 5a). These are generally considered to originate from natural plant
286 release and burning. The burning of coniferous trees, such as pine, releases high concentrations of
287 dehydroabietic acid compared with other lignin and resin-derived compounds (Simoneit, 2002). Here, the
288 concentrations of 4-hydroxybenzoic acid ($5.06 \pm 3.57 \text{ ng m}^{-3}$) and dehydroabietic acid ($2.98 \pm 1.91 \text{ ng m}^{-3}$)
289 were high, whereas those of vanillic acid ($1.50 \pm 1.08 \text{ ng m}^{-3}$) and syringic acid ($1.53 \pm 1.11 \text{ ng m}^{-3}$) were
290 relatively low. A significant positive correlation was observed between the total concentration of lignin and
291 resin-derived compounds and levoglucosan ($r = 0.80$, $p < 0.01$, Figure S3a), suggesting that biomass burning
292 was a potential source of lignin and resin-derived compounds. The concentration of dehydroabietic acid was
293 significantly lower than that of 4-hydroxybenzoic acid, similar to the pattern observed for Nanjing aerosols



294 (Haque et al., 2019). This implies that coniferous tree combustion may not be the primary source of lignin and
295 resin acids in this region.

296 We identified several sterols (cholesterol, stigmasterol, and β -sitosterol, average concentrations = $4.63 \pm$
297 3.17 , 4.89 ± 3.68 , and 10.11 ± 8.26 ng m⁻³, respectively; Figure 5a). These sterols originate from distinct
298 sources. Cholesterol, an animal-derived sterol, is primarily associated with meat cooking in the atmosphere
299 (Rogge et al., 1991; Nolte et al., 1999; He et al., 2004). Contrarily, stigmasterol and β -sitosterol are plant-
300 derived sterols, typically originating from plant leaves, cooking, or biomass burning (Nolte et al., 2001; He et
301 al., 2004; Zhao et al., 2007). Here, stigmasterol and β -sitosterol exhibited a significant positive correlation with
302 levoglucosan (Figure S3a). This suggests that biomass burning may be a potential source of plant-derived
303 sterols. The concentration of stigmasterol and β -sitosterol was higher in autumn and winter than in spring and
304 summer, indicating an increased contribution of biomass burning to sterols in the autumn and winter.

305 3.3.2 Glycerol and hydroxy acids

306 Glycerol and three hydroxy acids, glycolic acid, malic acid, and citric acid, were detected in the PM_{2.5} samples
307 (Figure 5b). The concentration of glycerol was relatively high (9.99 ± 4.13 ng m⁻³), whereas the three hydroxy
308 acids exhibited comparable and relatively lower concentrations: glycolic acid (7.16 ± 1.90 ng m⁻³), malic acid
309 (6.60 ± 1.85 ng m⁻³), and citric acid (5.92 ± 1.99 ng m⁻³). These concentrations aligned with reported ranges
310 for urban aerosols (2 – 146 and 1 – 180 ng m⁻³ for glycerol and hydroxy acids, respectively; Wang et al., 2006;
311 Fu et al., 2010; Haque et al., 2019; Fan et al., 2020). The glycerol in the aerosol is primarily derived from
312 fungal metabolism in suspended soil (Simoneit et al., 2004b), whereas glycolic, malic, and citric acids mainly
313 originate from the secondary photooxidation of organic compounds in the atmosphere (Kawamura and
314 Ikushima, 1993; Kawamura and Sakaguchi, 1999; Claeys et al., 2004). No significant correlation was observed
315 between glycerol and the hydroxy acids; however, a notable correlation was observed among the hydroxy acids
316 ($r = 0.49$ – 0.66 , $p < 0.01$, Figure S3b). This confirmed the divergent sources of glycerol and hydroxy acids and
317 similar origins for the hydroxy acids. In summer, the concentrations of the hydroxy acids exceeded those in
318 other seasons (Figure 5b). A positive correlation was observed between the polyacids and temperature ($r =$
319 0.35 – 0.51 , $p < 0.01$), suggesting enhanced secondary photochemical oxidation processes in the warm months.

320 3.3.3 Aromatic acids



321 We detected four aromatic acids in the PM_{2.5} samples: benzoic acid, phthalic acid, isophthalic acid, and
322 terephthalic acid at concentrations of 0.76 ± 0.34 , 6.40 ± 4.05 , 1.10 ± 0.74 , and 10.06 ± 3.07 ng m⁻³,
323 respectively (Figure 5b). Aromatic acids significantly contribute to the formation of atmospheric particles.
324 Benzoic acid is recognized as a major pollutant in vehicle exhaust and the secondary photochemical oxidation
325 product of aromatic hydrocarbons from vehicle emissions (Kawamura and Kaplan, 1987; Rogge et al., 1993;
326 Kawamura et al., 2000). Phthalic acid typically originates from the secondary transformation of PAHs,
327 including naphthalene and other PAHs. Terephthalic acid is primarily produced through the hydrolysis of
328 terephthalate during the combustion of urban plastics (Fu et al., 2010; Haque et al., 2019). Here, phthalic and
329 terephthalic acids jointly accounted for approximately 90% of the total aromatic acids detected. A significant
330 positive correlation was observed between phthalic acid and PAHs ($r = 83$, $p < 0.01$), as well as between
331 terephthalic acid and the total phthalates ($r = 0.72$, $p < 0.01$). The PAH and phthalate data have been published
332 by Guo et al. (2024a and 2004b). This suggests that the secondary conversion of PAHs and the hydrolysis of
333 phthalates (specifically terephthalate) during plastic burning significantly contribute to aromatic acid levels.
334 The phthalic acid concentrations were higher in autumn and winter, whereas terephthalic acid levels peaked in
335 spring and summer (Figure 5b). This indicates that the secondary transformation of PAHs was more
336 pronounced during the cold months, whereas high temperatures in spring and summer promoted the
337 volatilization and transformation of phthalates from plastic sources.

338 **3.4 Secondary organic aerosol tracers in PM_{2.5}**

339 **3.4.1 Biogenic secondary organic aerosol tracers**

340 Biogenic and anthropogenic SOAs play critical roles in influencing the atmospheric radiation balance and
341 regional air quality (Ding et al., 2014). Biogenic SOA tracers include oxidation products from isoprene,
342 monoterpenes, sesquiterpenes, and other oxygenated hydrocarbons. Isoprene is the dominant component of
343 BVOC emissions, accounting for approximately 45% of the total emissions (annual release estimated at 600
344 Tg C) (Piccot et al., 1992; Guenther et al., 1995; Sharkey et al., 2008). Here, we detected six isoprene
345 oxidation products (2-methylglyceric acid (2-MGA), three C5-alkene triols, and two 2-methyltetrols (MTLs) at
346 concentrations of 1.83 ± 1.11 , 2.84 ± 1.88 , and 4.14 ± 3.27 ng m⁻³, respectively; Figure 6). The concentrations
347 of the C5-alkene triols and MTLs exceeded that of 2-MGA. Generally, C5-alkene triols primarily originate



348 from the secondary transformation of VOCs released during biomass burning and higher plant waxes (Fu et al.,
349 2010 and 2014). Throughout the sampling period, a strong linear correlation was observed between the MTLs
350 and C5-alkene triols ($R^2 = 0.82$, $P < 0.01$), indicating their similar sources. Nevertheless, certain differences
351 were observed in the formation processes of C5-alkene triols and MTLs. Both compounds are formed through
352 the acid-catalyzed ring-opening reaction of isoprene epoxydiols (Surratt et al., 2006 and 2010); however, the
353 formation of C5-alkene triols is enhanced in acidic aerosol environments (Yee et al., 2020). Here, the ratios of
354 C5-alkene triols to MTLs (C5/MTLs) were elevated in autumn and winter compared with those observed in
355 spring and summer (Figure 7a), suggesting that increased aerosol acidity in the cold months may facilitate the
356 production of C5-alkene triols. C5-alkene triols and MTLs are believed to form via isoprene oxidation under
357 low- NO_x conditions (Surratt et al., 2010; Lin et al., 2013). Contrarily, high concentrations of NO_x favor the
358 further oxidation of isoprene to yield 2-MGA (Lin et al., 2013; Nguyen et al., 2015). Here, the 2-MGA/MTL
359 ratios in autumn and winter exceeded those observed in spring and summer (Figure 7b), indicating that
360 elevated NO_x concentrations in the autumn and winter may enhance the formation of 2-MGA. In MTLs, 2-
361 methylthreitol and 2-methylerythritol (MTL1 and MTL2) exhibited a significant linear correlation ($R^2 = 0.68$,
362 $P < 0.01$), indicating that they may originate from similar sources and/or share similar formation pathways.
363 The ratio of MTL2 to MTLs (MTL2/MTLs) can reflect the transformation pathways of MTLs. The ratio of
364 MTL2/MTLs = 0.35, 0.61, 0.76, and 0.90 corresponded to isoprene secondary transformation from biogenic
365 sources, OH-rich conditions (low NO_x), NO_x -rich conditions, and liquid-phase oxidation by H_2O_2 , respectively
366 (Kleindienst et al., 2009; Nozière et al., 2011). The results showed that the MTL2/MTLs values in spring (0.64
367 ± 0.12) and summer (0.59 ± 0.11) were relatively low and closer to the values for biogenic sources and OH-
368 rich secondary transformation. However, the values in autumn and (0.66 ± 0.10) winter (0.69 ± 0.06) were
369 relatively high, approaching those for NO_x -rich secondary transformation and liquid-phase oxidation by H_2O_2 .
370 This indicates that in spring and summer, MTLs were more likely derived from biogenic sources and OH-rich
371 secondary transformation, whereas in autumn and winter, a shift was observed toward NO_x -rich secondary
372 transformations and liquid-phase oxidation by H_2O_2 .

373 Monoterpenes represent a crucial component of BVOC emissions, accounting for approximately 11% of the
374 annual emissions, with an annual emission of 110 Tg C (Guenther et al., 1995). Here, we identified four



375 oxidation products of monoterpene (pinonic acid (PNA); pinic acid (PA); 3-methyl-1,2,3-butanetricarboxylic
376 acid (MBTCA); and 3-hydroxyglutaric acid (3-HGA)) at concentrations of 1.69 ± 0.99 , 1.47 ± 0.86 , $0.79 \pm$
377 0.61 , and $0.77 \pm 0.52 \text{ ng m}^{-3}$, respectively (Figure 6). The concentrations of PNA and PA exceeded those of
378 MBTCA and 3-HGA. PNA and PA are produced via the oxidation of α/β -pinene via reactions with O_3 and OH
379 radicals, and the α/β -pinene detected in the aerosol samples is mainly derived from biomass burning and higher
380 plant release (Hallquist et al., 2009; Eddingsaas et al., 2012; Zhang et al., 2015; Iyer et al., 2021). The
381 predominance of PNA over PA is attributable to its relatively higher vapor pressure, consistent with previous
382 findings (Fu et al., 2008 and 2010). PNA and PA are recognized as first-generation oxidation products of α/β -
383 pinene, whereas MBTCA represents a second-generation oxidation product formed via the further
384 photooxidation of PNA and PA with OH radicals. The relative concentrations of such first- and second-
385 generation oxidation products (M/P) reflect the oxidation degree and aging status of monoterpene compounds
386 (Ding et al., 2014). Here, the M/P values were higher in spring (0.25 ± 0.11) and summer (0.32 ± 0.10) than in
387 autumn (0.16 ± 0.09) and winter (0.24 ± 0.10). This suggests that higher temperatures and/or stronger radiation
388 in spring and summer promoted the oxidation of monoterpene compounds (Figure 7d). The formation of 3-
389 HGA is considered to occur via a ring-opening mechanism, probably linked to heterogeneous reactions of
390 monoterpenes with irradiation in NO_x -rich environments (Jaoui et al., 2005; Claeys et al., 2007). The ratio of
391 MBTCA to 3-HGA (MBTCA/3-HGA) can be used to distinguish monoterpenes, as the secondary
392 transformation of α -pinene yields MBTCA at significantly higher rates relative to 3-HGA compared with β -
393 pinene (Jaoui et al., 2005). Here, the annual MBTCA/3-HGA value was 1.18 ± 0.65 , close to those observed in
394 urban environments in the United States (0.81 ± 0.32) and China (0.68 ± 0.65) (Lewandowski et al., 2013; Ding
395 et al., 2014). The MBTCA/3-HGA value was higher in spring (1.31 ± 0.66) and summer (1.45 ± 0.70) than in
396 autumn (1.09 ± 0.72) and winter (0.88 ± 0.32). This suggests that the contribution of α -pinene to monoterpene
397 was higher in spring and summer than in autumn and winter (Figure 7e). The ratio of the total isoprene to
398 monoterpene oxidation products (Iso/Pine) was lower in autumn and winter than in spring and summer (Figure
399 7f), indicating that the higher temperatures in spring and summer favored isoprene release.

400 BVOC emissions include a class of sesquiterpene compounds, among which β -caryophyllene is the most
401 abundant and frequently reported (Ding et al., 2014; Fan et al., 2020). β -caryophyllinic acid, a product of the



402 ozonolysis or photooxidation of β -caryophyllene, predominantly originates from biomass burning and natural
403 plant emissions, including those from pine and birch trees (Helmig et al., 2006; Duhl et al., 2008). Here, β -
404 caryophyllinic acid was detected at a concentration of $2.74 \pm 1.92 \text{ ng m}^{-3}$. The β -caryophyllinic acid
405 concentrations in autumn ($3.16 \pm 1.77 \text{ ng m}^{-3}$) and winter ($4.63 \pm 1.61 \text{ ng m}^{-3}$) exceeded those observed in
406 spring ($1.49 \pm 0.68 \text{ ng m}^{-3}$) and summer ($1.14 \pm 0.47 \text{ ng m}^{-3}$) (Figure 6). The higher concentrations of
407 caryophyllinic acid in autumn and winter may be associated with increased biomass burning and subsequent
408 secondary transformation of β -caryophyllene. During the sampling period, a significant positive linear
409 correlation was observed between β -caryophyllinic acid and levoglucosan, supporting this inference ($R^2 =$
410 0.66 , $P < 0.01$).

411 **3.4.2 Anthropogenic secondary organic aerosol tracers**

412 Anthropogenic SOA tracers include hydroxy and aromatic acids, which primarily originate from the secondary
413 transformation of AVOCs. Although global AVOC emissions are relatively modest at 110 Tg C yr^{-1} (Piccot et
414 al., 1992) compared with BVOC emissions, which reach $1,150 \text{ Tg C yr}^{-1}$ (Guenther et al., 1995), the
415 contribution of anthropogenic sources to SOA is frequently more pronounced in urban environments because
416 of the impact of human activities (von Schneidmesser et al., 2010; Ding et al., 2012; Li et al., 2019). AVOC
417 emissions in urban settings can enhance the oxidation of BVOCs, promoting SOA formation (Carlton et al.,
418 2010; Hoyle et al., 2011). Here, we identified two primary anthropogenic SOA tracers, 2,3-dihydroxy-4-
419 oxopentanoic acid (DHOPA) and phthalic acid at concentrations of 1.73 ± 1.10 and $6.38 \pm 4.05 \text{ ng m}^{-3}$,
420 respectively (Figures 5 and 6). DHOPA and phthalic acid are recognized as important markers for
421 anthropogenic SOAs, produced via the oxidation of toluene and PAHs, such as naphthalene (Kawamura and
422 Ikushima, 1993; Kleindienst et al., 2007 and 2012). Their concentrations were significantly elevated during
423 autumn and winter compared with spring and summer (Figures 5 and 6). This indicated a marked increase in
424 the contribution of anthropogenic sources to SOA during the cold months.

425 **3.5 Source apportionment of organic carbon and aerosol**

426 Employing a tracer-based methodology, we quantified the contributions of POC and SOC to the total OC
427 (Figure 8). The results showed that the POC and SOC concentrations were 3.22 ± 1.81 and $0.50 \pm 0.32 \text{ } \mu\text{g m}^{-3}$
428 (Figure 8a and 8b), accounting for $53.30\% \pm 12.53\%$ and $8.33\% \pm 3.29\%$ of the measured OC, respectively



429 (Figure 8f). These results were within the reported ranges of POC (5% to 76%) and SOC (3% to 56%)
430 proportions reported in other studies (Stone et al., 2009; Guo et al., 2012; Fan et al., 2020; Xu et al., 2021;
431 Haque et al., 2023). Anthropogenic sources, including biomass burning, coal combustion, motor vehicle
432 emissions, and cooking, contributed the majority of POC ($2.89 \pm 1.72 \mu\text{g m}^{-3}$, accounting for $88.98\% \pm$
433 6.24%). Natural sources, such as plant debris and fungal spores, contributed relatively little ($0.33 \pm 0.17 \mu\text{g}$
434 m^{-3} , accounting for $11.01\% \pm 6.24\%$). For SOC, anthropogenic contributions ($0.33 \pm 0.25 \mu\text{g m}^{-3}$, accounting
435 for $60.67\% \pm 21.29\%$) exceeded biogenic contributions ($0.17 \pm 0.09 \mu\text{g m}^{-3}$, accounting for $39.05\% \pm$
436 21.15%). The concentrations and relative abundances of anthropogenic POC and SOC were higher in autumn
437 and winter, whereas biogenic POC and SOC exhibited greater concentrations and relative abundances during
438 spring and summer (Figures 8a, 8b, 8d, and 8e). These results underscore the significant influence of biogenic
439 emissions during the warm months, whereas anthropogenic emissions exert a pronounced effect during the
440 cold months. This situation is primarily attributed to the elevated temperatures, and increased solar radiation
441 during the warm months facilitates VOC release from vegetation, thereby promoting biogenic POC emissions
442 and SOC formation. Contrarily, the rise in anthropogenic POC emissions from biomass burning and coal
443 combustion during the cold months fosters the development of anthropogenic SOC in the atmosphere.
444 Similarly, using a tracer-based methodology, we quantified the contributions of POA and SOA to $\text{PM}_{2.5}$ (Figure
445 S4). The findings indicated that POA and SOA contributed 6.13 ± 3.58 and $1.02 \pm 0.63 \mu\text{g m}^{-3}$, accounting for
446 $21.02\% \pm 6.50\%$ and $3.47\% \pm 1.41\%$ of the observed $\text{PM}_{2.5}$ mass in Nanchang, respectively. The
447 anthropogenic POA ($5.50 \pm 3.41 \mu\text{g m}^{-3}$, accounting for $88.74\% \pm 6.46\%$) and SOA ($0.65 \pm 0.48 \mu\text{g m}^{-3}$,
448 accounting for $58.79\% \pm 21.30\%$) exceeded the natural POA ($0.63 \pm 0.33 \mu\text{g m}^{-3}$, accounting for $11.26\% \pm$
449 6.46%) and SOA ($0.37 \pm 0.20 \mu\text{g m}^{-3}$, accounting for $40.93\% \pm 21.19\%$), respectively. The POA and SOA
450 exhibited patterns analogous to those of POC and SOC, with greater contributions recorded in autumn and
451 winter compared with spring and summer. The tracer-based method inherently bears a degree of uncertainty,
452 primarily due to the variability in the mass fractions of representative tracers in OC and OA (f_{oc} and f_{oa}) in
453 different observational studies. Nevertheless, this method has been widely employed to estimate the
454 contributions of various primary and secondary sources to OC and aerosol, with relatively reasonable results.
455 Additional observational studies to determine the representative tracers in emission sources and their mass



456 fractions in OC and OA are crucial to reducing this uncertainty (Oros and Simoneit, 2000; He et al., 2004;
457 Zhao et al., 2007; Zhang et al., 2008; Kleindienst et al., 2012; Andreae, 2019).

458 We employed the CMB model to calculate the contributions of different sources of POC and SOC to OC.
459 The results derived from the CMB approach were largely consistent with those obtained using the tracer-based
460 method (Figure S5). The CMB model operates on principles analogous to the tracer-based method, relying on
461 the mass fractions of characteristic tracers in OC and OAs from emission sources to ascertain the contributions
462 of different sources (Stone et al., 2009). Thus, theoretically, the results generated by the CMB model should be
463 consistent with those obtained from the tracer-based method. The present results confirmed that both methods
464 exhibited similar reliabilities. We employed the EC-based method to estimate the total POC and SOC
465 concentrations. Dissimilar to the tracer-based method, which quantifies partial POC and SOC concentrations
466 from specific sources, the EC-based method simply partitions the OC into POC and SOC, which calculates the
467 total POC and SOC concentrations. The results demonstrated that the overall trends of the total POC and SOC
468 calculated by the EC-based method were consistent with those obtained via the tracer-based method (Figure
469 S6).

470 **3.6 Characteristics of organic aerosols during winter pollution**

471 From an annual timescale perspective, anthropogenic POC and SOC are synchronous with the measured OC
472 and observed PM_{2.5} mass (Figure S6). Correlation analysis revealed that anthropogenic POC and SOC
473 exhibited a significant positive correlation ($r = 0.72\text{--}0.80$, $P < 0.01$) with the measured OC (Figure 9a).
474 Redundancy analysis revealed that anthropogenic POC and SOC significantly contributed (40%–65%, $P <$
475 0.01) to the variations in OC concentration (Figure 9a). This suggests that anthropogenic OC and SOC are
476 important factors influencing changes in the OC and PM_{2.5} mass throughout the sampling period. Variations in
477 POC, SOC, OC, and PM_{2.5} mass were associated with specific meteorological conditions and gaseous pollutant
478 concentrations. For instance, high POC, SOC, OC, and PM_{2.5} concentrations corresponded to elevated
479 atmospheric pressure, reduced precipitation, and increased NO₂ concentrations (Figure S2). This suggests that
480 meteorological conditions and gaseous pollutant concentrations could also influence POC, SOC, OC, and
481 PM_{2.5} mass.

482 However, winter presents a different scenario. During the entire winter pollution period and separate PM



483 pollution episodes, POC did not maintain good consistency with the changes in OC and PM_{2.5} mass, whereas
484 SOC remained synchronized with their variations (Figure S6). Correlation and redundancy analyses further
485 revealed that source-specific SOC, derived using the tracer methods, and the total SOC, calculated using the
486 EC-based method, exhibited significant positive correlations with the measured OC ($r > 0.6$, $P < 0.05$) and
487 significantly contributed ($>40\%$, $P < 0.05$) to their variations (Figures 9b–9k). These findings indicate that on
488 shorter timescales, particularly during brief PM pollution episodes lasting several days, SOC was a critical
489 factor influencing the OC and PM_{2.5} mass. The increased SOC concentration may be associated with elevated
490 temperatures and NO_x concentrations during winter pollution episodes. This inference was supported by a
491 significant linear positive correlation between temperature, NO₂ concentrations, and SOC concentrations
492 observed during several winter pollution episodes (Figure 10). An increase in short-term solar radiation
493 intensity and oxidant levels have been shown to accelerate SOC formation (Fry et al., 2009; Ng et al., 2017; Li
494 et al., 2018; Ren et al., 2019). Nevertheless, the contribution of POC should not be underestimated, as its levels
495 remained relatively high throughout the winter. This elevated POC concentration can also promote SOC
496 transformation (Weber et al., 2007; Carlton et al., 2010; Hoyle et al., 2011; Srivastava et al., 2022).

497 **4 Conclusions**

498 We investigated the composition and concentration of major polar organic compounds in PM_{2.5} samples
499 collected over a year in Nanchang, Central China. The results revealed relatively high concentrations of FAs,
500 fatty alcohols, and saccharides, whereas lignin, resin products, sterols, glycerol, hydroxy acids, and aromatic
501 acids were detected at low levels.

502 The study findings indicate that the organic components in the PM_{2.5} of Nanchang are predominantly
503 derived from anthropogenic and natural sources. Anthropogenic sources were the primary contributors to OC
504 and OA, and primary sources contributed more than secondary sources. Throughout the sampling period, we
505 observed that the anthropogenic contributions significantly increased during autumn and winter, whereas
506 biogenic contributions increased in spring and summer.

507 This study highlights the critical role of anthropogenic POC and SOC in influencing atmospheric PM_{2.5}
508 pollution over an annual sampling period. During short-term pollution episodes in winter (lasting several
509 days), the rapid secondary transformation emerged as the primary driver of OC increment and PM_{2.5} pollution.



510 Elevated primary emissions and favorable oxidation conditions, such as increased light intensity and NO_x
511 levels, were identified as key factors facilitating the rapid secondary transformation of OC during the winter
512 pollution episodes.

513 The study findings underscore the necessity for targeted management strategies that consider primary and
514 secondary anthropogenic emission sources across different seasons and pollution periods. Although the tracer-
515 based methodology provided preliminary insights into the OC and OA from diverse sources, the study
516 encountered inherent limitations in characteristic compound identification. The source apportionment analysis
517 potentially underestimated contributions from unidentified primary and secondary sources because of the
518 restricted range of tracers used. To address this, future research should prioritize comprehensive observational
519 studies aimed at identifying representative molecular tracers across a broader spectrum of emissions and
520 quantifying the precise mass fractions of such tracers in OC and OAs.

521

522 **Data availability.** The dataset for this paper is available upon request from the corresponding author
523 (xiaohuayun@sjtu.edu.cn).

524

525 **Competing interests.** The authors declare that they have no conflict of interest.

526

527 **Author contributions.** Huayun Xiao conceptualized and designed the research. Sampling was conducted by
528 Zicong Li, while laboratory analyses were carried out by Wei Guo, Zicong Li, and Renguo Zhu. Data
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531

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538 **References**

- 539 Al-Naiema, I. M., and Stone, E. A.: Evaluation of anthropogenic secondary organic aerosol tracers from
540 aromatic hydrocarbons, *Atmos. Chem. Phys.*, 17, 2053–2065, <https://doi.org/10.5194/acp-17-2053-2017>,
541 2017.
- 542 An, Z. S., Huang, R. J., Zhang, R. Y., Tie, X. X., Li, G. H., Cao, J. J., Zhou, W. J., Shi, Z. G., Han, Y. M., Gu,
543 Z. L., and Ji, Y. M.: Severe haze in northern China: A synergy of anthropogenic emissions and atmospheric
544 processes, *Proc. Natl. Acad. Sci. U.S.A.*, 116, 8657–8666, <https://doi.org/10.1073/pnas.1900125116>, 2019.
- 545 Andreae, M. O.: Emission of trace gases and aerosols from biomass burning – an updated assessment, *Atmos.*
546 *Chem. Phys.*, 19, 8523–8546, <https://doi.org/10.5194/acp-19-8523-2019>, 2019.
- 547 Cao, J. J., and Cui, L.: Current Status, Characteristics and Causes of Particulate Air Pollution in the Fenwei
548 Plain, China: A Review, *J. Geophys. Res.-Atmos.*, 126, e2020JD034472,
549 <https://doi.org/10.1029/2020JD034472>, 2021.
- 550 Cao, J. J., Xu, H. M., Xu, Q., Chen, B. H., and Kan, H. D.: Fine particulate matter constituents and
551 cardiopulmonary mortality in a heavily polluted Chinese city, *Environ. Health Perspect.*, 120, 373–378,
552 <https://doi.org/10.1289/ehp.1103671>, 2012.
- 553 Carlton, A. G., Pinder, R. W., Bhave, P. V., and Pouliot, G. A.: To What Extent Can Biogenic SOA be
554 Controlled?, *Environ. Sci. Technol.*, 44, 3376–3380, <https://doi.org/10.1021/es903506b>, 2010.
- 555 Castro, L. M., Pio, C. A., Harrison, R. M., and Smith, D. J. T.: Carbonaceous aerosol in urban and rural
556 European atmospheres: estimation of secondary organic carbon concentrations, *Atmos. Environ.*, 33, 2771–
557 2781, [https://doi.org/10.1016/S1352-2310\(98\)00331-8](https://doi.org/10.1016/S1352-2310(98)00331-8), 1999.
- 558 Chen, C. R., Zhang, H. X., Yan, W. J., Wu, N. N., Zhang, Q., and He, K. B.: Aerosol water content
559 enhancement leads to changes in the major formation mechanisms of nitrate and secondary organic aerosols
560 in winter over the North China Plain, *Environ. Pollut.*, 287, 117625,
561 <https://doi.org/10.1016/j.envpol.2021.117625>, 2021.
- 562 Cheng, Y. B., Ma, Y. Q., and Hu, D.: Tracer-based source apportioning of atmospheric organic carbon and the
563 influence of anthropogenic emissions on secondary organic aerosol formation in Hong Kong, *Atmos. Chem.*
564 *Phys.*, 21, 10589–10608, <https://doi.org/10.5194/acp-21-10589-2021>, 2021.



- 565 Claeys, M., Graham, B., Vas, G., Wang, W., Vermeylen, R., Pashynska, V., Cafmeyer, J., Guyon, P., Andreae,
566 M. O., Artaxo, P., and Maenhaut, W.: Formation of Secondary Organic Aerosols Through Photooxidation
567 of Isoprene, *Science*, 303, 1173–1176, <https://doi.org/doi:10.1126/science.1092805>, 2004.
- 568 Claeys, M., Szmigielski, R., Kourtchev, I., Van der Veken, P., Vermeylen, R., Maenhaut, W., Jaoui, M.,
569 Kleindienst, T. E., Lewandowski, M., Offenberg, J. H., and Edney, E. O.: Hydroxycarboxylic Acids:
570 Markers for Secondary Organic Aerosol from the Photooxidation of α -Pinene, *Environ. Sci. Technol.*, 41,
571 1628–1634, <https://doi.org/10.1021/es0620181>, 2007.
- 572 Ding, X., Wang, X. M., Gao, B., Fu, X. X., He, Q. F., Zhao, X. Y., Yu, J. Z., and Zheng, M.: Tracer-based
573 estimation of secondary organic carbon in the Pearl River Delta, south China, *J. Geophys. Res-Atmos.*, 117,
574 <https://doi.org/10.1029/2011JD016596>, 2012.
- 575 Ding, X., Wang, X. M., Xie, Z. Q., Zhang, Z., and Sun, L. G.: Impacts of Siberian Biomass Burning on
576 Organic Aerosols over the North Pacific Ocean and the Arctic: Primary and Secondary Organic Tracers,
577 *Environ. Sci. Technol.*, 47, 3149–3157, <https://doi.org/10.1021/es3037093>, 2013.
- 578 Ding, X., He, Q. F., Shen, R. Q., Yu, Q. Q., and Wang, X. M.: Spatial distributions of secondary organic
579 aerosols from isoprene, monoterpenes, β -caryophyllene, and aromatics over China during summer, *J.*
580 *Geophys. Res-Atmos.*, 119, 11877–811891, <https://doi.org/10.1002/2014JD021748>, 2014.
- 581 Duhl, T. R., Helmig, D., and Guenther, A.: Sesquiterpene emissions from vegetation: a review, *Biogeosciences*,
582 5, 761–777, <https://doi.org/10.5194/bg-5-761-2008>, 2008.
- 583 Eddingsaas, N. C., Loza, C. L., Yee, L. D., Chan, M., Schilling, K. A., Chhabra, P. S., Seinfeld, J. H., and
584 Wennberg, P. O.: α -pinene photooxidation under controlled chemical conditions-Part 2: SOA yield and
585 composition in low- and high- NO_x environments, *Atmos. Chem. Phys.*, 12, 7413–7427,
586 <https://doi.org/10.5194/acp-12-7413-2012>, 2012.
- 587 Engling, G., He, J., Betha, R., and Balasubramanian, R.: Assessing the regional impact of Indonesian biomass
588 burning emissions based on organic molecular tracers and chemical mass balance modeling, *Atmos. Chem.*
589 *Phys.*, 14, 8043–8054, <https://doi.org/10.5194/acp-14-8043-2014>, 2014.
- 590 Fabbri, D., Torri, C., Simoneit, B. R. T., Marynowski, L., Rushdi, A. I., and Fabiańska, M. J.: Levoglucosan
591 and other cellulose and lignin markers in emissions from burning of Miocene lignites, *Atmos. Environ.*, 43,
592 2286–2295, <https://doi.org/10.1016/j.atmosenv.2009.01.030>, 2009.
- 593 Fan, Y. B., Liu, C. Q., Li, L. J., Ren, L. J., Ren, H., Zhang, Z. M., Li, Q. K., Wang, S., Hu, W., Deng, J. J., Wu,
594 L. B., Zhong, S. J., Zhao, Y., Pavuluri, C. M., Li, X. D., Pan, X. L., Sun, Y. L., Wang, Z. F., Kawamura, K.,
595 Shi, Z. B., and Fu, P. Q.: Large contributions of biogenic and anthropogenic sources to fine organic



- 596 aerosols in Tianjin, North China, *Atmos. Chem. Phys.*, 20, 117–137, [https://doi.org/10.5194/acp-20-117-](https://doi.org/10.5194/acp-20-117-597)
597 2020, 2020.
- 598 Fry, J. L., Kiendler-Scharr, A., Rollins, A. W., Wooldridge, P. J., Brown, S. S., Fuchs, H., Dubé W., Mensah,
599 A., dal Maso, M., Tillmann, R., Dorn, H. P., Brauers, T., and Cohen, R. C.: Organic nitrate and secondary
600 organic aerosol yield from NO₃ oxidation of β-pinene evaluated using a gas-phase kinetics/aerosol
601 partitioning model, *Atmos. Chem. Phys.*, 9, 1431–1449, <https://doi.org/10.5194/acp-9-1431-2009>, 2009.
- 602 Fu, P. Q., Kawamura, K., Okuzawa, K., Aggarwal, S. G., Wang, G., Kanaya, Y., and Wang, Z.: Organic
603 molecular compositions and temporal variations of summertime mountain aerosols over Mt. Tai, North
604 China Plain, *J. Geophys. Res-Atmos.*, 113, <https://doi.org/10.1029/2008JD009900>, 2008.
- 605 Fu, P. Q., Kawamura, K., and Barrie, L. A.: Photochemical and other sources of organic compounds in the
606 Canadian high arctic aerosol pollution during winter-spring, *Environ. Sci. Technol.*, 43, 286–292,
607 <https://doi.org/10.1021/es803046q>, 2009.
- 608 Fu, P. Q., Kawamura, K., Pavuluri, C. M., Swaminathan, T., and Chen, J.: Molecular characterization of urban
609 organic aerosol in tropical India: contributions of primary emissions and secondary photooxidation, *Atmos.*
610 *Chem. Phys.*, 10, 2663–2689, <https://doi.org/10.2663-2689>, [10.5194/acp-10-2663-2010](https://doi.org/10.5194/acp-10-2663-2010), 2010.
- 611 Fu, P. Q., Kawamura, K., and Miura, K.: Molecular characterization of marine organic aerosols collected
612 during a round-the-world cruise, *J. Geophys. Res-Atmos.*, 116, <https://doi.org/10.1029/2011JD015604>,
613 2011.
- 614 Fu, P. Q., Kawamura, K., Kobayashi, M., and Simoneit, B. R. T.: Seasonal variations of sugars in atmospheric
615 particulate matter from Gosan, Jeju Island: Significant contributions of airborne pollen and Asian dust in
616 spring, *Atmos. Environ.*, 55, 234–239, <https://doi.org/10.1016/j.atmosenv.2012.02.061>, 2012.
- 617 Fu, P. Q., Kawamura, K., Chen, J., and Miyazaki, Y.: Secondary Production of Organic Aerosols from
618 Biogenic VOCs over Mt. Fuji, Japan, *Environ. Sci. Technol.*, 48, 8491–8497,
619 <https://doi.org/10.1021/es500794d>, 2014.
- 620 Graham, B., Guyon, P., Taylor, P. E., Artaxo, P., Maenhaut, W., Glovsky, M. M., Flagan, R. C., and Andreae,
621 M. O.: Organic compounds present in the natural Amazonian aerosol: Characterization by gas
622 chromatography–mass spectrometry, *J. Geophys. Res-Atmos.*, 108, <https://doi.org/10.1029/2003JD003990>,
623 2003.
- 624 Guenther, A., Hewitt, C. N., Erickson, D., Fall, R., Geron, C., Graedel, T., Harley, P., Klinger, L., Lerdau, M.,
625 Mckay, W. A., Pierce, T., Scholes, B., Steinbrecher, R., Tallamraju, R., Taylor, J., and Zimmerman, P.: A
626 global model of natural volatile organic compound emissions, *Res-Atmos.*, 100, 8873–8892,
627 <https://doi.org/10.1029/94JD02950>, 1995.



- 628 Guo, S., Hu, M., Guo, Q., Zhang, X., Zheng, M., Zheng, J., Chang, C. C., Schauer, J. J., and Zhang, R.:
629 Primary Sources and Secondary Formation of Organic Aerosols in Beijing, China, *Environ. Sci. Technol.*,
630 46, 9846–9853, 10.1021/es2042564, 2012.
- 631 Guo, W., Li, Z. C., Zhang, Z. Y., Zhu, R. G., Xiao, H. W., and Xiao, H. Y.: Sources and influences of
632 atmospheric nonpolar organic compounds in Nanchang, central China: Full-year monitoring with a focus
633 on winter pollution episodes, *Sci. Total Environ.*, 912, 169216,
634 <https://doi.org/10.1016/j.scitotenv.2023.169216>, 2024a.
- 635 Guo, W., Zhang, Z. Y., Zhu, R. G., Li, Z. C., Liu, C., Xiao, H. W., and Xiao, H. Y.: Pollution characteristics,
636 sources, and health risks of phthalate esters in ambient air: A daily continuous monitoring study in the
637 central Chinese city of Nanchang, *Chemosphere*, 353, 141564,
638 <https://doi.org/10.1016/j.chemosphere.2024.141564>, 2024b.
- 639 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue,
640 N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M.,
641 Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod, A.,
642 Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The formation, properties and
643 impact of secondary organic aerosol: current and emerging issues, *Atmos. Chem. Phys.*, 9, 5155–5236,
644 10.5194/acp-9-5155-2009, 2009.
- 645 Haque, M. M., Kawamura, K., Deshmukh, D. K., Fang, C., Song, W., Mengying, B., and Zhang, Y. L.:
646 Characterization of organic aerosols from a Chinese megacity during winter: predominance of fossil fuel
647 combustion, *Atmos. Chem. Phys.*, 19, 5147–5164, 10.5194/acp-19-5147-2019, 2019.
- 648 Haque, M. M., Zhang, Y., Bikkina, S., Lee, M., and Kawamura, K.: Regional heterogeneities in the emission
649 of airborne primary sugar compounds and biogenic secondary organic aerosols in the East Asian outflow:
650 evidence for coal combustion as a source of levoglucosan, *Atmos. Chem. Phys.*, 22, 1373–1393,
651 10.5194/acp-22-1373-2022, 2022.
- 652 Haque, M. M., Verma, S. K., Deshmukh, D. K., Kunwar, B., and Kawamura, K.: Seasonal characteristics of
653 biogenic secondary organic aerosol tracers in a deciduous broadleaf forest in northern Japan, *Chemosphere*,
654 311, 136785, <https://doi.org/10.1016/j.chemosphere.2022.136785>, 2023.
- 655 He, L. Y., Hu, M., Huang, X. F., Yu, B. D., Zhang, Y. H., and Liu, D. Q.: Measurement of emissions of fine
656 particulate organic matter from Chinese cooking, *Atmos. Environ.*, 38, 6557–6564,
657 <https://doi.org/10.1016/j.atmosenv.2004.08.034>, 2004.



- 658 Helmig, D., Ortega, J., Guenther, A., Herrick, J. D., and Geron, C.: Sesquiterpene emissions from loblolly pine
659 and their potential contribution to biogenic aerosol formation in the Southeastern US, *Atmos. Environ.*, 40,
660 4150–4157, <https://doi.org/10.1016/j.atmosenv.2006.02.035>, 2006.
- 661 Hoyle, C. R., Boy, M., Donahue, N. M., Fry, J. L., Glasius, M., Guenther, A., Hallar, A. G., Huff Hartz, K.,
662 Petters, M. D., Petäjä T., Rosenoern, T., and Sullivan, A. P.: A review of the anthropogenic influence on
663 biogenic secondary organic aerosol, *Atmos. Chem. Phys.*, 11, 321–343, [https://doi.org/10.5194/acp-11-321-](https://doi.org/10.5194/acp-11-321-2011)
664 2011, 2011.
- 665 Huang, R. J., Zhang, Y., Bozzetti, C., Ho, K. F., Cao, J. J., Han, Y., Daellenbach, K. R., Slowik, J. G., Platt, S.
666 M., Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Bruns, E. A., Crippa, M., Ciarelli, G., Piazzalunga, A.,
667 Schwikowski, M., Abbaszade, G., Schnelle-Kreis, J., Zimmermann, R., An, Z., Szidat, S., Baltensperger,
668 U., El Haddad, I., and Prevot, A. S.: High secondary aerosol contribution to particulate pollution during
669 haze events in China, *Nature*, 514, 218–222, <https://doi.org/10.1038/nature13774>, 2014.
- 670 Huebert, B., Bates, T., Russell, P., Shi, G., Kim, Y., Kawamura, K., Carmichael, G., and Nakajima, T.: An
671 overview of ACE-Asia: Strategies for quantifying the relationships between Asian aerosols and their
672 climatic impacts, *J. Geophys. Res-Atmos.*, <https://doi.org/10.1029/2003JD003550>, 2003.
- 673 Iinuma, Y., Brüggemann, E., Gnauk, T., Müller, K., Andreae, M. O., Helas, G., Parmar, R., and Herrmann, H.:
674 Source characterization of biomass burning particles: The combustion of selected European conifers,
675 African hardwood, savanna grass, and German and Indonesian peat, *J. Geophys. Res-Atmos.*, 112,
676 <https://doi.org/10.1029/2006JD007120>, 2007.
- 677 Iyer, S., Rissanen, M. P., Valiev, R., Barua, S., Krechmer, J. E., Thornton, J., Ehn, M., and Kurtán, T.:
678 Molecular mechanism for rapid autoxidation in α -pinene ozonolysis, *Nat. Commun.*, 12, 878,
679 [10.1038/s41467-021-21172-w](https://doi.org/10.1038/s41467-021-21172-w), 2021.
- 680 Jaoui, M., Kleindienst, T. E., Lewandowski, M., Offenberg, J. H., and Edney, E. O.: Identification and
681 Quantification of Aerosol Polar Oxygenated Compounds Bearing Carboxylic or Hydroxyl Groups. 2.
682 Organic Tracer Compounds from Monoterpenes, *Environ. Sci. Technol.*, 39, 5661–5673,
683 [10.1021/es048111b](https://doi.org/10.1021/es048111b), 2005.
- 684 Jaoui, M., Szmigielski, R., Nestorowicz, K., Kolodziejczyk, A., Sarang, K., Rudzinski, K. J., Konopka, A.,
685 Bulska, E., Lewandowski, M., and Kleindienst, T. E.: Organic Hydroxy Acids as Highly Oxygenated
686 Molecular (HOM) Tracers for Aged Isoprene Aerosol, *Environ. Sci. Technol.*, 53, 14516–14527,
687 [10.1021/acs.est.9b05075](https://doi.org/10.1021/acs.est.9b05075), 2019.
- 688 Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C., Van Dingenen, R.,
689 Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E., Putaud, J. P., Balkanski, Y., Fuzzi, S., Horth, J.,



- 690 Moortgat, G. K., Winterhalter, R., Myhre, C. E. L., Tsigaridis, K., Vignati, E., Stephanou, E. G., and
691 Wilson, J.: Organic aerosol and global climate modelling: a review, *Atmos. Chem. Phys.*, 5, 1053–1123,
692 <https://doi.org/10.5194/acp-5-1053-2005>, 2005.
- 693 Kawamura, K., and Gagosian, R. B.: Implications of ω -oxocarboxylic acids in the remote marine atmosphere
694 for photo-oxidation of unsaturated fatty acids, *Nature*, 325, 330–332, <https://doi.org/10.1038/325330a0>,
695 1987.
- 696 Kawamura, K., and Kaplan, I. R.: Motor exhaust emissions as a primary source for dicarboxylic acids in Los
697 Angeles ambient air, *Environ. Sci. Technol.*, 21, 105–110, <https://doi.org/10.1021/es00155a014>, 1987.
- 698 Kawamura, K., and Ikushima, K.: Seasonal changes in the distribution of dicarboxylic acids in the urban
699 atmosphere, *Environ. Sci. Technol.*, 27, 2227–2235, <https://doi.org/10.1021/ES00047A033>, 1993.
- 700 Kawamura, K., and Sakaguchi, F.: Molecular distributions of water soluble dicarboxylic acids in marine
701 aerosols over the Pacific Ocean including tropics, *J. Geophys. Res-Atmos.*, 104, 3501–3509,
702 <https://doi.org/10.1029/1998JD100041>, 1999.
- 703 Kawamura, K., Steinberg, S., and Kaplan, I. R.: Homologous series of C1–C10 monocarboxylic acids and C1–
704 C6 carbonyls in Los Angeles air and motor vehicle exhausts, *Atmos. Environ.*, 34, 4175–4191,
705 [https://doi.org/10.1016/S1352-2310\(00\)00212-0](https://doi.org/10.1016/S1352-2310(00)00212-0), 2000.
- 706 Kawamura, K., Ishimura, Y., and Yamazaki, K.: Four years' observations of terrestrial lipid class compounds
707 in marine aerosols from the western North Pacific, *Global Biogeochem. Cy.*, 17, 1003,
708 <https://doi.org/10.1029/2001GB001810>, 2003.
- 709 Kleindienst, T. E., Jaoui, M., Lewandowski, M., Offenberg, J. H., Lewis, C. W., Bhave, P. V., and Edney, E.
710 O.: Estimates of the contributions of biogenic and anthropogenic hydrocarbons to secondary organic
711 aerosol at a southeastern US location, *Atmos. Environ.*, 41, 8288–8300,
712 <https://doi.org/10.1016/j.atmosenv.2007.06.045>, 2007.
- 713 Kleindienst, T. E., Lewandowski, M., Offenberg, J. H., Jaoui, M., and Edney, E. O.: The formation of
714 secondary organic aerosol from the isoprene ^+OH reaction in the absence of NO_x , *Atmos. Chem. Phys.*, 9,
715 6541–6558, [10.5194/acp-9-6541-2009](https://doi.org/10.5194/acp-9-6541-2009), 2009.
- 716 Kleindienst, T.E., Jaoui, M., Lewandowski, M., Offenberg, J.H., and Docherty, K.S.: The formation of SOA
717 and chemical tracer compounds from the photooxidation of naphthalene and its methyl analogs in the
718 presence and absence of nitrogen oxides, *Atmos. Chem. Phys.*, 12, 8711–8726 [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-12-8711-2012)
719 [12-8711-2012](https://doi.org/10.5194/acp-12-8711-2012), 2012.



- 720 Kunwar, B., and Kawamura, K.: One-year observations of carbonaceous and nitrogenous components and
721 major ions in the aerosols from subtropical Okinawa Island, an outflow region of Asian dusts, *Atmos.*
722 *Chem. Phys.*, 14, 1819–1836, <https://doi.org/10.5194/acp-14-1819-2014>, 2014.
- 723 Lewandowski, M., Jaoui, M., Offenberg, J. H., Kleindienst, T. E., Edney, E. O., Sheesley, R. J., and Schauer, J.
724 J.: Primary and secondary contributions to ambient PM in the midwestern United States, *Environ. Sci.*
725 *Technol.*, 42, 3303–3309, <https://doi.org/doi:10.1021/es0720412>, 2008.
- 726 Lewandowski, M., Piletic, I. R., Kleindienst, T. E., Offenberg, J. H., Beaver, M. R., Jaoui, M., Docherty, K. S.,
727 and Edney, E. O.: Secondary organic aerosol characterisation at field sites across the United States during
728 the spring–summer period, *Int. J. Environ. An. Ch.*, 93, 1084–1103,
729 <https://doi.org/10.1080/03067319.2013.803545>, 2013.
- 730 Li, J. J., Wang, G. H., Wu, C., Cao, C., Ren, Y. Q., Wang, J. Y., Li, J., Cao, J. J., Zeng, L. M., and Zhu, T.:
731 Characterization of isoprene-derived secondary organic aerosols at a rural site in North China Plain with
732 implications for anthropogenic pollution effects, *Sci. Rep-Uk.*, 8, 535, [https://doi.org/10.1038/s41598-017-](https://doi.org/10.1038/s41598-017-18983-7)
733 18983-7, 2018.
- 734 Li, J. J., Wang, G. H., Zhang, Q., Li, J., Wu, C., Jiang, W. Q., Zhu, T., and Zeng, L. M.: Molecular
735 characteristics and diurnal variations of organic aerosols at a rural site in the North China Plain with
736 implications for the influence of regional biomass burning, *Atmos. Chem. Phys.*, 19, 10481–
737 <https://doi.org/10496>, [10.5194/acp-19-10481-2019](https://doi.org/10.5194/acp-19-10481-2019), 2019.
- 738 Lin, Y. H., Zhang, Z., Docherty, K. S., Zhang, H., Budisulistiorini, S. H., Rubitschun, C. L., Shaw, S. L.,
739 Knipping, E. M., Edgerton, E. S., Kleindienst, T. E., Gold, A., and Surratt, J. D.: Isoprene Epoxydiols as
740 Precursors to Secondary Organic Aerosol Formation: Acid-Catalyzed Reactive Uptake Studies with
741 Authentic Compounds, *Environ. Sci. Technol.*, 46, 250–258, <https://doi.org/10.1021/es202554c>, 2012.
- 742 Lin, Y. H., Knipping, E. M., Edgerton, E. S., Shaw, S. L., and Surratt, J. D.: Investigating the influences of
743 SO₂ and NH₃ levels on isoprene-derived secondary organic aerosol formation using conditional sampling
744 approaches, *Atmos. Chem. Phys.*, 13, 8457–8470, <https://doi.org/10.5194/acp-13-8457-2013>, 2013.
- 745 Ng, N. L., Brown, S. S., Archibald, A. T., Atlas, E., Cohen, R. C., Crowley, J. N., Day, D. A., Donahue, N. M.,
746 Fry, J. L., Fuchs, H., Griffin, R. J., Guzman, M. I., Herrmann, H., Hodzic, A., Iinuma, Y., Jimenez, J. L.,
747 Kiendler-Scharr, A., Lee, B. H., Luecken, D. J., Mao, J., McLaren, R., Mutzel, A., Osthoff, H. D., Ouyang,
748 B., Picquet-Varrault, B., Platt, U., Pye, H. O. T., Rudich, Y., Schwantes, R. H., Shiraiwa, M., Stutz, J.,
749 Thornton, J. A., Tilgner, A., Williams, B. J., and Zaveri, R. A.: Nitrate radicals and biogenic volatile
750 organic compounds: oxidation, mechanisms, and organic aerosol, *Atmos. Chem. Phys.*, 17, 2103–2162,
751 <https://doi.org/10.5194/acp-17-2103-2017>, 2017.



- 752 Nguyen, T. B., Bates, K. H., Crounse, J. D., Schwantes, R. H., Zhang, X., Kjaergaard, H. G., Surratt, J. D., Lin,
753 P., Laskin, A., Seinfeld, J. H., and Wennberg, P. O.: Mechanism of the hydroxyl radical oxidation of
754 methacryloyl peroxyxynitrate (MPAN) and its pathway toward secondary organic aerosol formation in the
755 atmosphere, *Chem. Phys.*, 17, 17914–17926, <https://doi.org/10.1039/c5cp02001h>, 2015.
- 756 Nolte, C. G., Schauer, J. J., Cass, G. R., and Simoneit, B. R. T.: Highly Polar Organic Compounds Present in
757 Meat Smoke, *Environ. Sci. Technol.*, 33, 3313–3316, <https://doi.org/10.1021/es990122v>, 1999.
- 758 Nolte, C. G., Schauer, J. J., Cass, G. R., and Simoneit, B. R. T.: Highly Polar Organic Compounds Present in
759 Wood Smoke and in the Ambient Atmosphere, *Environ. Sci. Technol.*, 35, 1912–1919,
760 <https://doi.org/10.1021/es001420r>, 2001.
- 761 Nozière, B., González, N. J. D., Borg-Karlson, A. K., Pei, Y. X., Redeby, J. P., Krejci, R., Dommen, J., Prévôt,
762 A. S. H., and Anthonsen, T.: Atmospheric chemistry in stereo: A new look at secondary organic aerosols
763 from isoprene, *Geophys. Res. Lett.*, 38, <https://doi.org/10.1029/2011GL047323>, 2011.
- 764 Oros, D. R., and Simoneit, B. R. T.: Identification and emission rates of molecular tracers in coal smoke
765 particulate matter, *Fuel*, 79, 515–536, [https://doi.org/10.1016/S0016-2361\(99\)00153-2](https://doi.org/10.1016/S0016-2361(99)00153-2), 2000.
- 766 Oros, D. R., and Simoneit, B. R. T.: Identification and emission factors of molecular tracers in organic aerosols
767 from biomass burning Part 2. Deciduous trees, *Appl. Geochem.*, 16, 1545–1565,
768 [https://doi.org/10.1016/S0883-2927\(01\)00022-1](https://doi.org/10.1016/S0883-2927(01)00022-1), 2001.
- 769 Piccot, S. D., Watson, J. J., and Jones, J. W.: A global inventory of volatile organic compound emissions from
770 anthropogenic sources, *J. Geophys. Res.*, 97, 9897–9912, <https://doi.org/10.1029/92JD00682>, 1992.
- 771 Puxbaum, H., and Tenze-Kunit, M.: Size distribution and seasonal variation of atmospheric cellulose, *Atmos.*
772 *Environ.*, 37, 3693–3699, [https://doi.org/10.1016/S1352-2310\(03\)00451-5](https://doi.org/10.1016/S1352-2310(03)00451-5), 2003.
- 773 Ren, H., Hu, W., Wei, L. F., Yue, S. Y., Zhao, J., Li, L. J., Wu, L. B., Zhao, W. Y., Ren, L. J., Kang, M. J., Xie,
774 Q. Y., Su, S. H., Pan, X. L., Wang, Z. F., Sun, Y. L., Kawamura, K., and Fu, P. Q.: Measurement report:
775 Vertical distribution of biogenic and anthropogenic secondary organic aerosols in the urban boundary layer
776 over Beijing during late summer, *Atmos. Chem. Phys.*, 21, 12949–12963, [https://doi.org/10.5194/acp-21-](https://doi.org/10.5194/acp-21-12949-2021)
777 [12949-2021](https://doi.org/10.5194/acp-21-12949-2021), 2021.
- 778 Ren, Y. Q., Wang, G. H., Tao, J., Zhang, Z. S., Wu, C., Wang, J. Y., Li, J. J., Wei, J., Li, H., and Meng, F.:
779 Seasonal characteristics of biogenic secondary organic aerosols at Mt. Wuyi in Southeastern China:
780 Influence of anthropogenic pollutants, *Environ. Pollut.*, 252, 493–500,
781 <https://doi.org/10.1016/j.envpol.2019.05.077>, 2019.



- 782 Riipinen, I., Yli-Juuti, T., Pierce, J. R., Petaja, T., Worsnop, D. R., Kulmala, M., and Donahue, N. M.: The
783 contribution of organics to atmospheric nanoparticle growth, *Nat. Geosci.*, 5, 453–458,
784 <https://doi.org/10.1038/ngeo1499>, 2012.
- 785 Rogge, W. F., Hildemann, L. M., Mazurek, M. A., Cass, G. R., and Simoneit, B. R. T.: Sources of fine organic
786 aerosol. 1. Charbroilers and meat cooking operations, *Environ. Sci. Technol.*, 25, 1112–1125,
787 <https://doi.org/10.1021/es00018a015>, 1991.
- 788 Rogge, W. F., Hildemann, L. M., Mazurek, M. A., Cass, G. R., Simoneit, B. R. T. J. E. S., and Technology:
789 Sources of fine organic aerosol. 2. Nuncatalyst and catalyst-equipped automobiles and heavy-duty diesel
790 trucks, *Environ. Sci. Technol.*, 27, 636–651, <https://doi.org/10.1021/es00041a007>, 1993.
- 791 Rogge, W. F., Medeiros, P. M., and Simoneit, B. R. T.: Organic marker compounds for surface soil and
792 fugitive dust from open lot dairies and cattle feedlots, *Atmos. Environ.*, 40, 27–49,
793 <https://doi.org/10.1016/j.atmosenv.2005.07.076>, 2006.
- 794 Rudich, Y., Donahue, N. M., and Mentel, T. F.: Aging of organic aerosol: bridging the gap between laboratory
795 and field studies, *Annu. Rev. Phys. Chem.*, 58, 321–352,
796 <https://doi.org/10.1146/annurev.physchem.58.032806.104432>, 2007.
- 797 Rybicki, M., Marynowski, L., Bechtel, A., and Simoneit, B. R. T.: Variations in $\delta^{13}\text{C}$ values of levoglucosan
798 from low-temperature burning of lignite and biomass, *Sci. Total Environ.*, 733, 138991,
799 <https://doi.org/10.1016/j.scitotenv.2020.138991>, 2020a.
- 800 Rybicki, M., Marynowski, L., and Simoneit, B. R. T.: Composition of organic compounds from low-
801 temperature burning of lignite and their application as tracers in ambient air, *Chemosphere*, 249, 126087,
802 <https://doi.org/10.1016/j.chemosphere.2020.126087>, 2020b.
- 803 Saarikoski, S., Timonen, H., Saarnio, K., Aurela, M., Järvi, L., Keronen, P., Kerminen, V. M., and Hillamo, R.:
804 Sources of organic carbon in fine particulate matter in northern European urban air, *Atmos. Chem. Phys.*, 8,
805 6281–6295, <https://doi.org/10.5194/acp-8-6281-2008>, 2008.
- 806 Sharkey, T. D., Wiberley, A. E., and Donohue, A. R.: Isoprene emission from plants: why and how, *Annals of*
807 *botany*, 101, 5–18, <https://doi.org/10.1093/aob/mcm240>, 2008.
- 808 Sheesley, R. J., Schauer, J. J., Chowdhury, Z., Cass, G. R., and Simoneit, B. R. T.: Characterization of organic
809 aerosols emitted from the combustion of biomass indigenous to South Asia, *J. Geophys. Res-Atmos.*, 108,
810 <https://doi.org/10.1029/2002JD002981>, 2003.
- 811 Shiraiwa, M., Ueda, K., Pozzer, A., Lammel, G., Kampf, C. J., Fushimi, A., Enami, S., Arangio, A. M.,
812 Fröhlich-Nowoisky, J., Fujitani, Y., Furuyama, A., Lakey, P. S. J., Lelieveld, J., Lucas, K., Morino, Y.,
813 Pöschl, U., Takahama, S., Takami, A., Tong, H., Weber, B., Yoshino, A., and Sato, K.: Aerosol Health



- 814 Effects from Molecular to Global Scales, *Environ. Sci. Technol.*, 51, 13545–13567,
815 10.1021/acs.est.7b04417, 2017.
- 816 Simoneit, B. R. T.: Biomass burning — a review of organic tracers for smoke from incomplete combustion,
817 *Applied Geochemistry*, 17, 129–162, [https://doi.org/10.1016/S0883-2927\(01\)00061-0](https://doi.org/10.1016/S0883-2927(01)00061-0), 2002.
- 818 Simoneit, B. R. T., Kobayashi, M., Mochida, M., Kawamura, K., and Huebert, B. J.: Aerosol particles
819 collected on aircraft flights over the northwestern Pacific region during the ACE-Asia campaign:
820 Composition and major sources of the organic compounds, *J. Geophys. Res-Atmos.*, 109,
821 <https://doi.org/10.1029/2004JD004565>, 2004a.
- 822 Simoneit, B. R., Elias, V. O., Kobayashi, M., Kawamura, K., Rushdi, A. I., Medeiros, P. M., Rogge, W. F.,
823 and Didyk, B. M.: Sugars-dominant water-soluble organic compounds in soils and characterization as
824 tracers in atmospheric particulate matter, *Environ. Sci. Technol.*, 38, 5939–5949,
825 <https://doi.org/10.1021/es0403099>, 2004b.
- 826 Srivastava, D., Vu, T. V., Tong, S., Shi, Z., and Harrison, R. M.: Formation of secondary organic aerosols
827 from anthropogenic precursors in laboratory studies, *Npj. Clim. Atmos. Sci.*, 5, 22,
828 <https://doi.org/10.1038/s41612-022-00238-6>, 2022.
- 829 Stohl, A., Berg, T., Burkhart, J. F., Fjærraa, A. M., Forster, C., Herber, A., Hov, Ø., Lunder, C., McMillan, W.
830 W., Oltmans, S., Shiobara, M., Simpson, D., Solberg, S., Stebel, K., Ström, J., Tørseth, K., Treffeisen, R.,
831 Virkkunen, K., and Yttri, K. E.: Arctic smoke – record high air pollution levels in the European
832 Arctic due to agricultural fires in Eastern Europe in spring 2006, *Atmos. Chem. Phys.*, 7, 511–534,
833 <https://doi.org/10.5194/acp-7-511-2007>, 2007.
- 834 Stone, E. A., Zhou, J., Snyder, D. C., Rutter, A. P., Mieritz, M., and Schauer, J. J.: A comparison of
835 summertime secondary organic aerosol source contributions at contrasting urban locations, *Environ. Sci.*
836 *Technol.*, 43, 3448–3454, <https://doi.org/10.1021/es8025209>, 2009.
- 837 Surratt, J. D., Murphy, S. M., Kroll, J. H., Ng, N. L., Hildebrandt, L., Sorooshian, A., Szmigielski, R.,
838 Vermeylen, R., Maenhaut, W., Claeys, M., Flagan, R. C., and Seinfeld, J. H.: Chemical Composition of
839 Secondary Organic Aerosol Formed from the Photooxidation of Isoprene, *J. Phys. Chem. A*, 110, 9665–
840 9690, 10.1021/jp061734m, 2006.
- 841 Surratt, J. D., Chan, A. W. H., Eddingsaas, N. C., Chan, M., Loza, C. L., Kwan, A. J., Hersey, S. P., Flagan, R.
842 C., Wennberg, P. O., and Seinfeld, J. H.: Reactive intermediates revealed in secondary organic aerosol
843 formation from isoprene, *Proc. Natl. Acad. Sci. U.S.A.*, 107, 6640–6645,
844 <https://doi.org/doi:10.1073/pnas.0911114107>, 2010.



- 845 Tang, T. G., Cheng, Z. N., Xu, B. Q., Zhang, B. L., Li, J., Zhang, W., Wang, K. L., and Zhang, G.: Source
846 Diversity of Intermediate Volatility n-Alkanes Revealed by Compound-Specific $\delta^{13}\text{C}$ - δD Isotopes, *Environ.*
847 *Sci. Technol.*, 56, 14262–14271, 10.1021/acs.est.2c02156, 2022.
- 848 Turpin, B. J., and Huntzicker, J. J.: Identification of secondary organic aerosol episodes and quantitation of
849 primary and secondary organic aerosol concentrations during SCAQS, *Atmos. Environ.*, 29, 3527–3544,
850 [https://doi.org/10.1016/1352-2310\(94\)00276-Q](https://doi.org/10.1016/1352-2310(94)00276-Q), 1995.
- 851 von Schneidmesser, E., Zhou, J., Stone, E. A., Schauer, J. J., Shpund, J., Brenner, S., Qasrawi, R., Abdeen, Z.,
852 and Sarnat, J. A.: Spatial Variability of Carbonaceous Aerosol Concentrations in East and West Jerusalem,
853 *Environ. Sci. Technol.*, 44, 1911–1917, <https://doi.org/10.1021/es9014025>, 2010.
- 854 Wang, G. H., and Kawamura, K.: Molecular Characteristics of Urban Organic Aerosols from Nanjing: A Case
855 Study of A Mega-City in China, *Environ. Sci. Technol.*, 39, 7430–7438, <https://doi.org/10.1021/es051055>,
856 2005.
- 857 Wang, G. H., Kawamura, K., Lee, S., Ho, K., and Cao, J. J.: Molecular, Seasonal, and Spatial Distributions of
858 Organic Aerosols from Fourteen Chinese Cities, *Environ. Sci. Technol.*, 40, 4619–4625,
859 <https://doi.org/10.1021/es060291x>, 2006.
- 860 Wang, G. H., Kawamura, K., Hatakeyama, S., Takami, A., Li, H., and Wang, W.: Aircraft Measurement of
861 Organic Aerosols over China, *Environ. Sci. Technol.*, 41, 3115–3120, <https://doi.org/10.1021/es062601h>,
862 2007.
- 863 Weber, R. J., Sullivan, A. P., Peltier, R. E., Russell, A. G., Yan, B., Zheng, M., Gouw, J. d., Warneke, C.,
864 Brock, C. A., Holloway, J. S., Atlas, E. L., and Edgerton, E. S.: A study of secondary organic aerosol
865 formation in the anthropogenic-influenced southeastern United States, *J. Geophys. Res-Atmos.*, 112,
866 <https://doi.org/10.1029/2007JD008408>, 2007.
- 867 Wu, J. R., Bei, N. F., Wang, Y., Li, X., Liu, S. X., Liu, L., Wang, R. N., Yu, J. Y., Le, T. H., Zuo, M., Shen, Z.
868 X., Cao, J. J., Tie, X. X., and Li, G. H.: Insights into particulate matter pollution in the North China Plain
869 during wintertime: local contribution or regional transport?, *Atmos. Chem. Phys.*, 21, 2229–2249,
870 <https://doi.org/10.5194/acp-21-2229-2021>, 2021.
- 871 Xu, B. Q., Tang, J., Tang, T. G., Zhao, S. Z., Zhong, G. C., Zhu, S. Y., Li, J., and Zhang, G.: Fates of
872 secondary organic aerosols in the atmosphere identified from compound-specific dual-carbon isotope
873 analysis of oxalic acid, *Atmos. Chem. Phys.*, 23, 1565–1578, <https://doi.org/10.5194/acp-23-1565-2023>,
874 2023.
- 875 Xu, B. Q., Zhang, G., Gustafsson, Ö., Kawamura, K., Li, J., Andersson, A., Bikkina, S., Kunwar, B., Pokhrel,
876 A., Zhong, G. C., Zhao, S. Z., Li, J., Huang, C., Cheng, Z. N., Zhu, S. Y., Peng, P. A., and Sheng, G. Y.:



- 877 Large contribution of fossil-derived components to aqueous secondary organic aerosols in China, *Nat.*
878 *Commun.*, 13, 5115, <https://doi.org/10.1038/s41467-022-32863-3>, 2022.
- 879 Xu, J. S., Liu, D., Wu, X. F., Vu, T. V., Zhang, Y. L., Fu, P. Q., Sun, Y. L., Xu, W. Q., Zheng, B., Harrison, R.
880 M., and Shi, Z. B.: Source apportionment of fine organic carbon at an urban site of Beijing using a
881 chemical mass balance model, *Atmos. Chem. Phys.*, 21, 7321–7341, [https://doi.org/10.5194/acp-21-7321-](https://doi.org/10.5194/acp-21-7321-2021)
882 2021, 2021.
- 883 Yan, C. C., Sullivan, A. P., Cheng, Y., Zheng, M., Zhang, Y. H., Zhu, T., and Collett, J. L.: Characterization of
884 saccharides and associated usage in determining biogenic and biomass burning aerosols in atmospheric fine
885 particulate matter in the North China Plain, *Sci. Total Environ.*, 650, 2939–2950,
886 <https://doi.org/10.1016/j.scitotenv.2018.09.325>, 2019.
- 887 Yazdani, A., Dudani, N., Takahama, S., Bertrand, A., Prévôt, A. S. H., El Haddad, I., and Dillner, A. M.:
888 Characterization of primary and aged wood burning and coal combustion organic aerosols in an
889 environmental chamber and its implications for atmospheric aerosols, *Atmos. Chem. Phys.*, 21, 10273–
890 10293, [https://doi.org/10.5194/acp-21-10273-](https://doi.org/10.5194/acp-21-10273-2021)2021, 2021.
- 891 Yee, L. D., Isaacman-VanWertz, G., Wernis, R. A., Kreisberg, N. M., Glasius, M., Riva, M., Surratt, J. D., de
892 S á S. S., Martin, S. T., Alexander, M. L., Palm, B. B., Hu, W., Campuzano-Jost, P., Day, D. A., Jimenez, J.
893 L., Liu, Y., Misztal, P. K., Artaxo, P., Viegas, J., Manzi, A., de Souza, R. A. F., Edgerton, E. S., Baumann,
894 K., and Goldstein, A. H.: Natural and Anthropogenically Influenced Isoprene Oxidation in Southeastern
895 United States and Central Amazon, *Environ. Sci. Technol.*, 54, 5980–5991,
896 <https://doi.org/10.1021/acs.est.0c00805>, 2020.
- 897 Yttri, K. E., Dye, C., and Kiss, G.: Ambient aerosol concentrations of sugars and sugar-alcohols at four
898 different sites in Norway, *Atmos. Chem. Phys.*, 7, 4267–4279, [https://doi.org/10.5194/acp-7-4267-](https://doi.org/10.5194/acp-7-4267-2007)2007,
899 2007.
- 900 Zhang, Q., He, K. B., and Huo, H.: Cleaning China's air, *Nature*, 484, 161–162,
901 <https://doi.org/10.1038/484161a>, 2012.
- 902 Zhang, Q., Hu, W., Ren, H., Yang, J. B., Deng, J. J., Wang, D. W., Sun, Y. L., Wang, Z. F., Kawamura, K.,
903 and Fu, P. Q.: Diurnal variations in primary and secondary organic aerosols in an eastern China coastal city:
904 The impact of land-sea breezes, *Environ. Pollut.*, 319, 121016,
905 <https://doi.org/10.1016/j.envpol.2023.121016>, 2023.
- 906 Zhang, X., McVay, R. C., Huang, D. D., Dalleska, N. F., Aumont, B., Flagan, R. C., and Seinfeld, J. H.:
907 Formation and evolution of molecular products in α -pinene secondary organic aerosol, *Proc. Natl. Acad.*
908 *Sci. U. S. A.*, 112, 14168–14173, <https://doi.org/10.1073/pnas.1517742112>, 2015.



- 909 Zhang, Y. X., Cao, F., Song, W. H., Jia, X. F., Xie, T., Wu, C.L., Yan, P., Yu, M. Y., Rauber, M., Salazar, G.,
910 Szidat, S., and Zhang, Y. L.: Fossil and Nonfossil Sources of Winter Organic Aerosols in the Regional
911 Background Atmosphere of China, *Environ. Sci. Technol.*, 58, 1244–1254,
912 <https://doi.org/10.1021/acs.est.3c08491>, 2024.
- 913 Zhang, Y. X., Schauer, J. J., Zhang, Y. H., Zeng, L. M., Wei, Y. J., Liu, Y., and Shao, M.: Characteristics of
914 Particulate Carbon Emissions from Real-World Chinese Coal Combustion, *Environ. Sci. Technol.*, 42,
915 5068–5073, <https://doi.org/10.1021/es7022576>, 2008.
- 916 Zhang, Y. L., Kawamura, K., Agrios, K., Lee, M., Salazar, G., and Szidat, S.: Fossil and Nonfossil Sources of
917 Organic and Elemental Carbon Aerosols in the Outflow from Northeast China, *Environ. Sci. Technol.*, 50,
918 6284–6292, <https://doi.org/10.1021/acs.est.6b00351>, 2016.
- 919 Zhao, Y. L., Hu, M., Slanina, S., and Zhang, Y. H.: Chemical Compositions of Fine Particulate Organic Matter
920 Emitted from Chinese Cooking, *Environ. Sci. Technol.*, 41, 99–105, <https://doi.org/10.1021/es0614518>,
921 2007.
- 922 Zhu, C. M., Kawamura, K., and Fu, P. Q.: Seasonal variations of biogenic secondary organic aerosol tracers in
923 Cape Hedo, Okinawa, *Atmos. Environ.*, 130, 113–119, <https://doi.org/10.1016/j.atmosenv.2015.08.069>,
924 2016.
- 925 Zhu, R. G., Xiao, H. Y., Cheng, L., Zhu, H., Xiao, H., and Gong, Y. Y.: Measurement report: Characterization
926 of sugars and amino acids in atmospheric fine particulates and their relationship to local primary sources,
927 *Atmos. Chem. Phys.*, 22, 14019–14036, [10.5194/acp-22-14019-2022](https://doi.org/10.5194/acp-22-14019-2022), 2022.

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937 **Figure Captions**

938 Figure 1. (a) Average $PM_{2.5}$ concentrations across various cities in China during winter 2023–2024 (December
939 2023 to February 2024). (b) Location of sampling sites in Nanchang.

940 Figure 2. One-year time series of organic carbon (OC), elemental carbon (EC), OC/EC ratios, $PM_{2.5}$
941 concentrations, and concentrations of polar organic compounds.

942 Figure 3. Molecular characteristics and seasonal variations of the major polar compounds in $PM_{2.5}$.

943 Figure 4. The concentration ratios and its seasonal variations of the major polar compounds. Each box plot
944 illustrates the mean (centerline), interquartile range (box encompassing the 25th to 75th percentiles), and
945 standard deviation (whiskers).

946 Figure 5. Molecular characteristics and seasonal variations of the minor polar compounds in $PM_{2.5}$.

947 Figure 6. Molecular characteristics and seasonal variations of the SOA tracers in $PM_{2.5}$.

948 Figure 7. The concentration ratios and its seasonal variations of the SOA tracers. Box plots represent the mean
949 (centerline), interquartile range (box encompassing the 25th to 75th percentiles), and standard deviation
950 (whiskers).

951 Figure 8. Concentrations and relative abundances of primary organic carbon (POC) from biomass burning
952 (POC_{bb}), coal combustion (POC_{cc}), vehicle exhaust (POC_{ve}), cooking (POC_c), plant debris (POC_{pd}), and fungal
953 spores (POC_{fs}), alongside secondary organic carbon (SOC) from isoprene (SOC_i), α/β -pinene (SOC_p), β -
954 caryophyllene (SOC_c), toluene (SOC_t), and naphthalene (SOC_n). POC and SOC concentrations were estimated
955 by the tracer-based method, and other OC (OOC) was obtained by subtracting estimated OC from measured
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957 Figure 9. Pearson correlation between POC, SOC, and the measured OC during the annual sampling period
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960 redundancy analyses that are less than 0.05, while gray circles indicate P-values greater than 0.05. Definitions:
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963 SOC_{EC-based} refer to SOC from isoprene, α / β -pinene, β -caryophyllene, toluene, naphthalene, and EC-based
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965 Figure 10. Linear correlation between temperature, NO₂, and SOC during specific winter pollution episode.

966 SOC concentrations were estimated by the tracer-based method.

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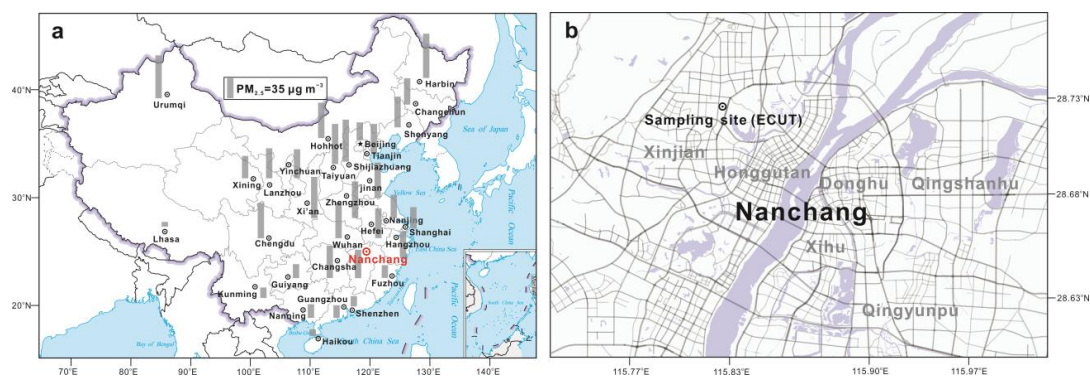
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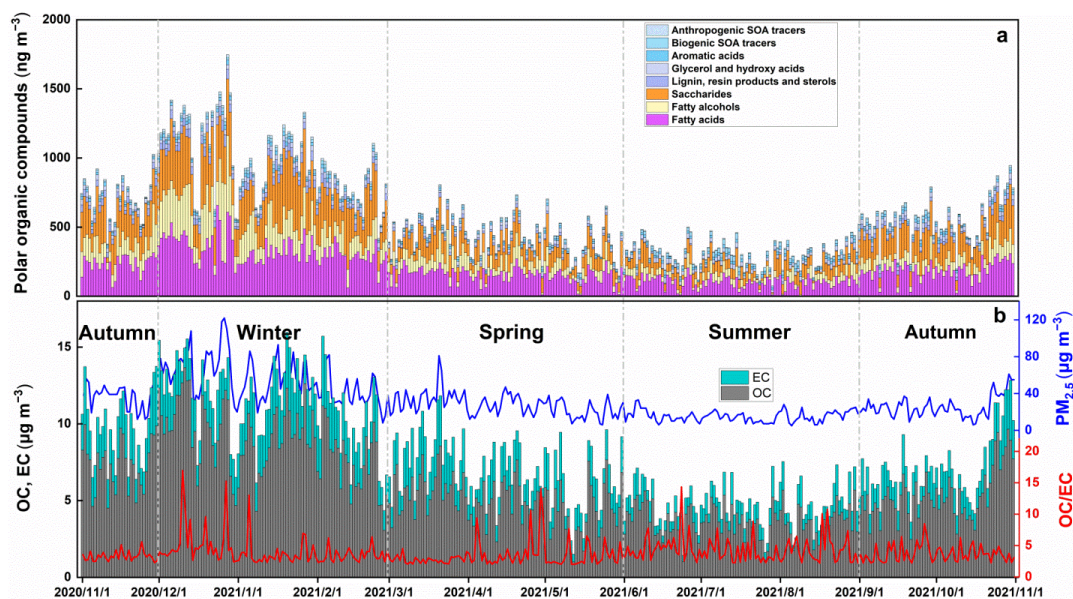
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1019 Figure 3. Molecular characteristics and seasonal variations of the major polar compounds in PM_{2.5}.

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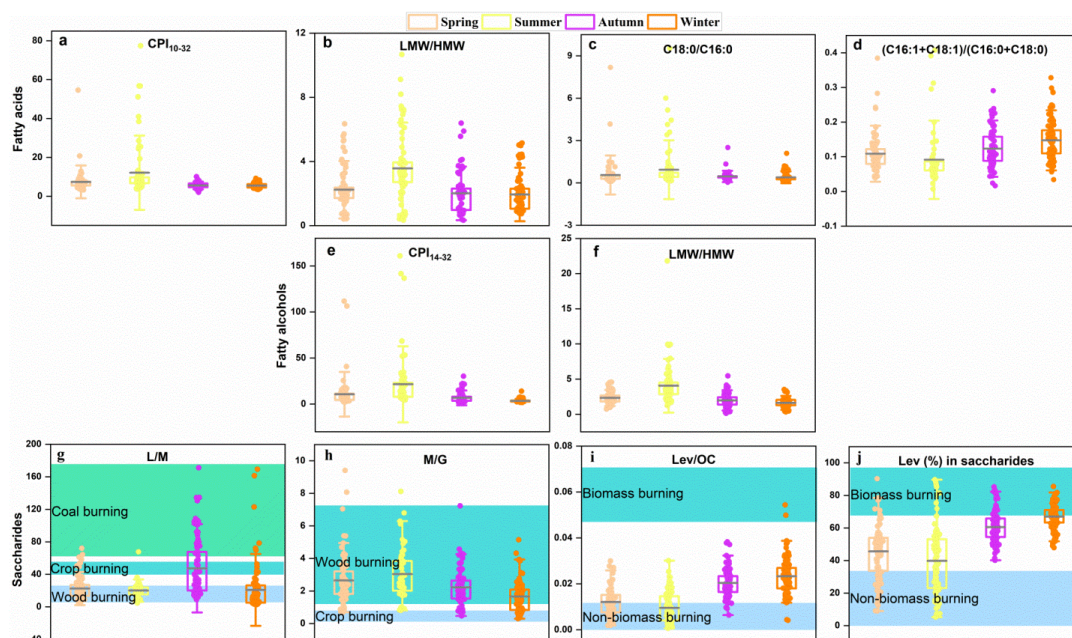
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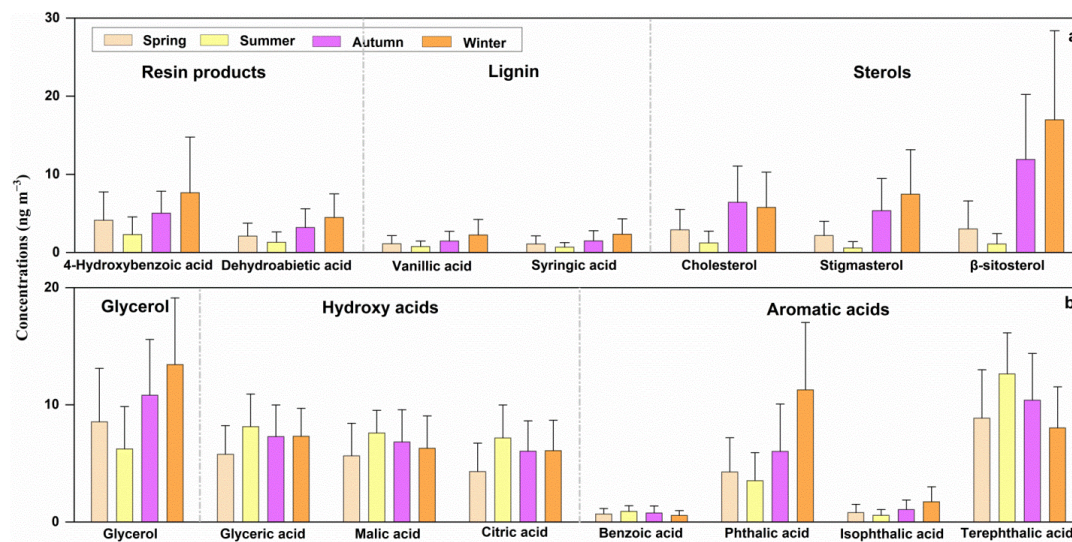
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1032 illustrates the mean (centerline), interquartile range (box encompassing the 25th to 75th percentiles), and
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Figure 5. Molecular characteristics and seasonal variations of the minor polar compounds in PM_{2.5}.

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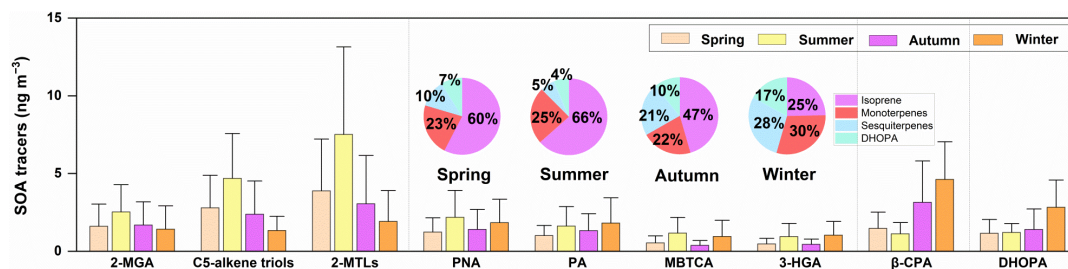
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1076 Figure 6. Molecular characteristics and seasonal variations of the SOA tracers in PM_{2.5}.

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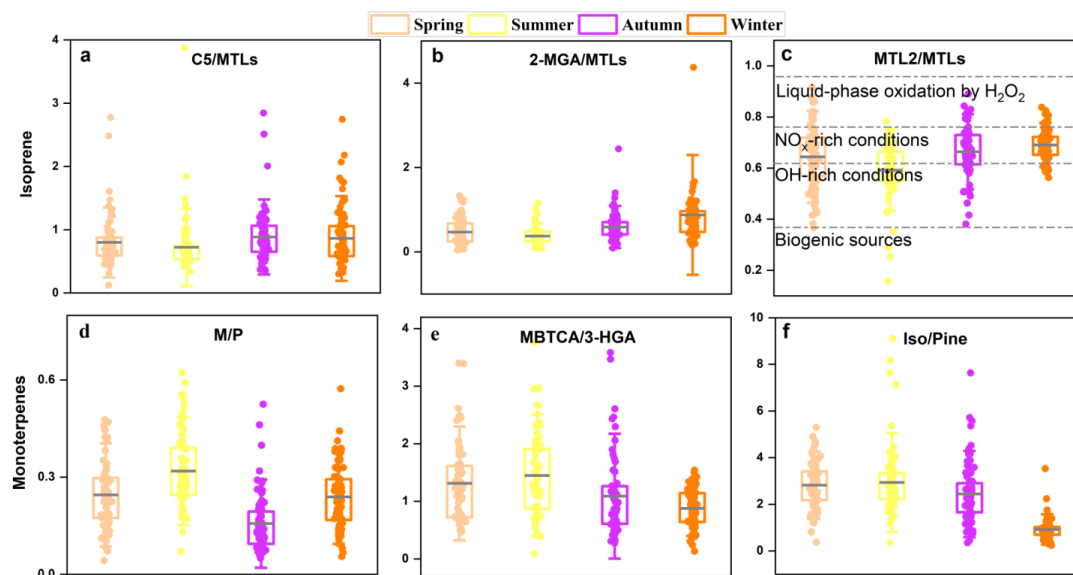
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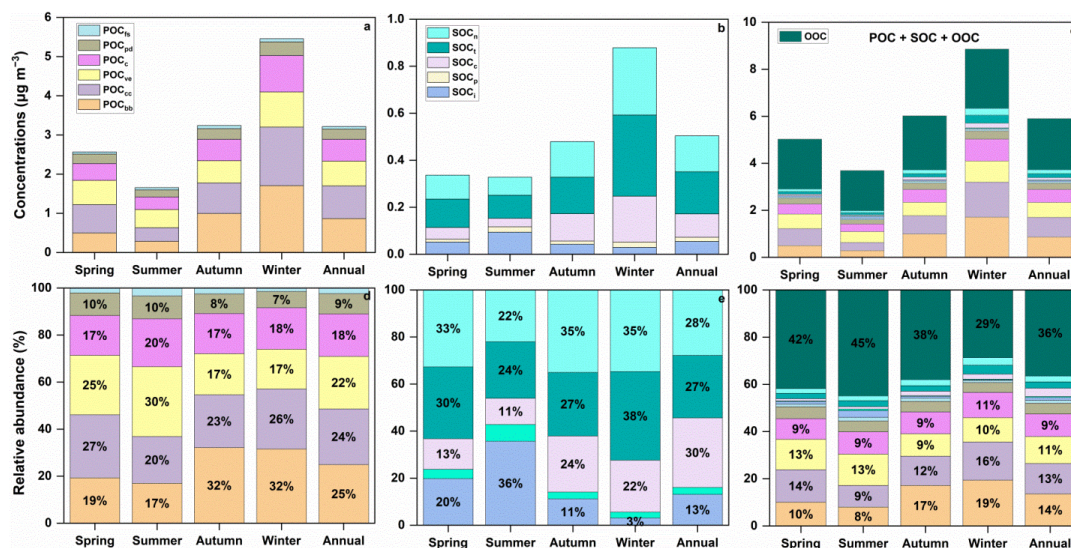
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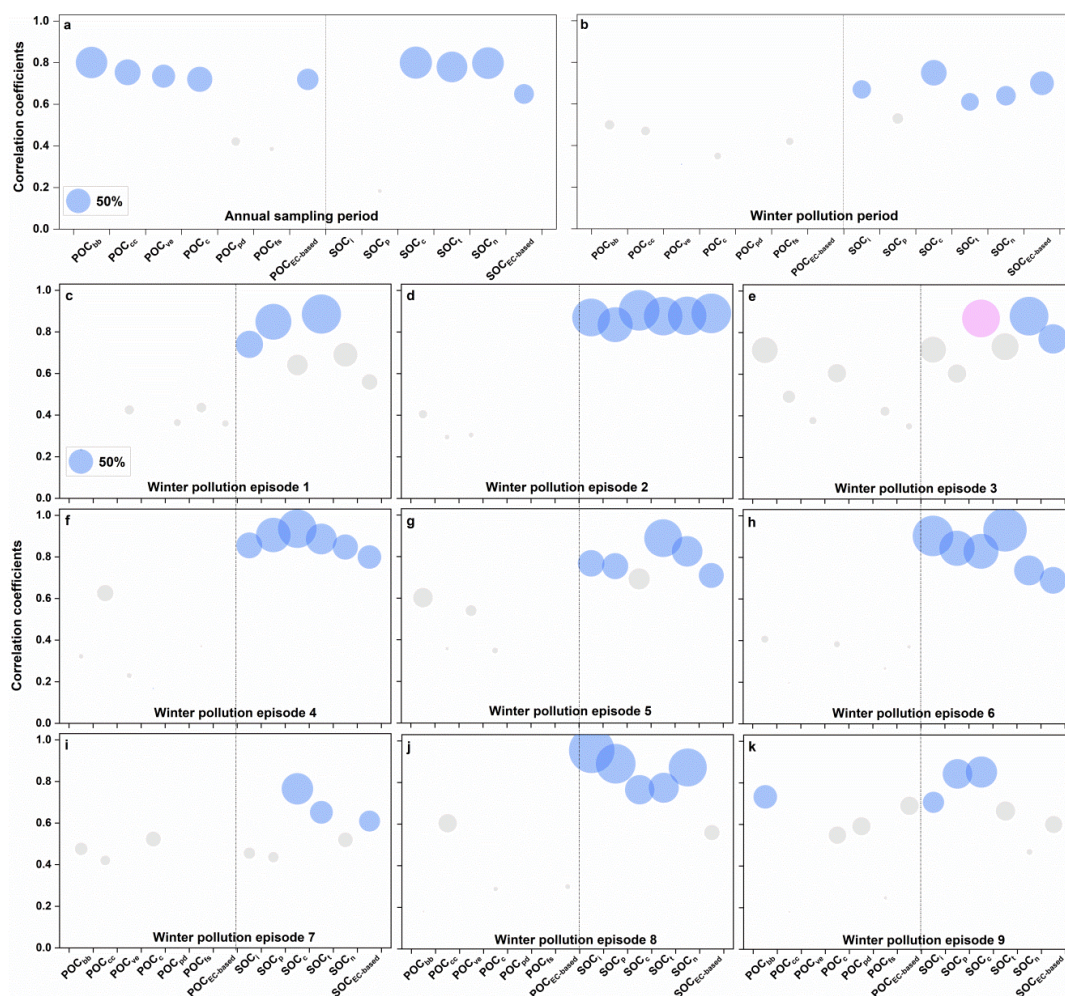
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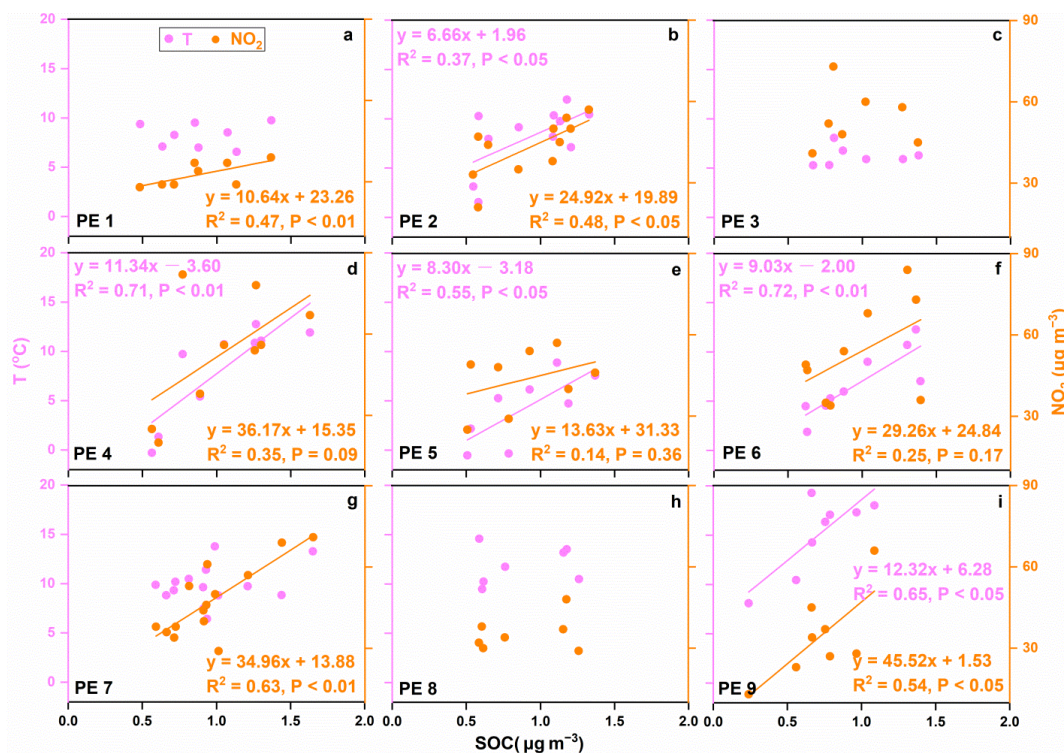
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