



Integrated Geophysical Analysis of Rangpur Saddle: Insights on Tectonics and Magnetic Mineral Potential of North-Western Bangladesh

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13 Abstract

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15 The northwestern region of Bangladesh holds untapped potential for magnetic mineral deposits at shallow depths. Unlike much of Bangladesh, characterized by thick sediments of Bengal Basin, 16 this area is an extension of the Indian Shield, often referred to as the Stable Platform. It is also 17 geologically distinct, hosting structures related to the breakup of Pangea. The geology and 18 tectonics of this region have remained largely understudied. To address this gap, this study 19 integrates gravity, magnetic, seismic, and drilling data to investigate the subsurface structure and 20 evaluate the resource potential of the area. We utilize advanced filtering and modeling techniques, 21 including tilt derivatives and horizontal gradient methods, to understand the tectonic framework 22 23 and geometry of the subsurface structures. Our spatial analysis, using multiple geophysical datasets, reveals distinct magnetic anomalies characterized by alternating magnetic highs and lows, 24 which we attribute to gabbroic intrusions along extensional faults that define the region's horst and 25 26 graben structures. To validate our interpretations, we developed an integrated 2-D subsurface model that aligns with the observed geophysical data. However, the study is limited by the 27 availability of high-resolution seismic data and the sparse distribution of drilling locations, which 28 may affect the precision of our subsurface characterization. Our findings provide crucial insights 29 into the tectonic evolution of the stable platform and underscore the economic potential of the 30 Rangpur Saddle, the shallowest part of the stable platform, for mineral exploration. These insights 31 pave the way for further exploration and development initiatives focused on uncovering the 32 mineral wealth of this underexplored region. 33

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35 1. Introduction

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The northwestern part of Bangladesh is rich in potential mineral resources. The geological 37 diversity of this area suggests the presence of other valuable minerals that remain largely 38 39 unexplored (Akhtar, 2005; Hasan et al., 2023; Moon, 2022). In this area, basement rocks are 40 discernible at relatively shallow depths of 128 meters, hosting a spectrum of economic mineral resources, including coal, limestone, white clay, and hard rock (Khan & Rahman, 1992). Notably, 41 Pirganj in the Rangpur district (Figure 1) records the country's highest magnetic anomaly 42 suggesting the potential for magnetic mineral ore deposits. In the 1990s, drilling activities were 43 conducted in the region after developing a 2D subsurface model by the Geological Survey of 44 Bangladesh (GSB) in collaboration with the United States Geological Survey (USGS). Despite 45 46 these efforts, no noteworthy ore body was identified during the drilling process (Rahman & Ullah,





47 2009). Moreover, the tectonic evolution of the Paleo-Proterozoic basement in the northwest region
48 of the Bengal Basin still needs to be studied.

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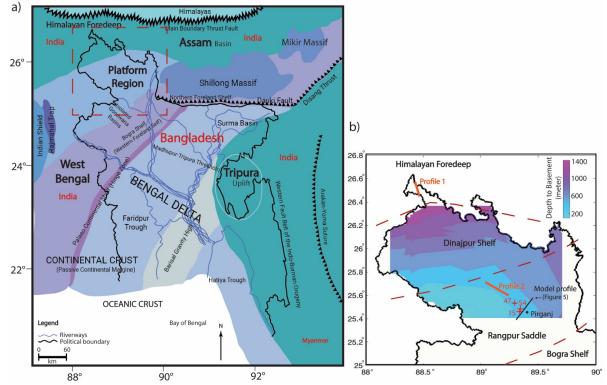
The study area is located in a region commonly referred to as the Rangpur platform or saddle 50 (Masum et al., 2021), an eastern extension of the Indian Shield (Alam et al., 2003). The northern 51 part of the Rangpur Saddle rests on a shallow Precambrian basement, ranging from 130 to 1000 52 meters in depth. This area of Bangladesh is geologically stable, characterized predominantly by 53 54 horsts and grabens, which were formed during the Cretaceous rifting of the Indian plate from the Antarctica-Australia section of Gondwanaland (Curray, 1991; Curray & Moore, 1974). Although 55 the basement primarily consists of diorite, tonalite, and granodiorite, it is also intersected by 56 pegmatite and mafic/ultramafic dykes (Hossain et al., 2007; Kabir et al., 2001). 57

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59 In this study, we explore the possibility of mafic dykes as the source of high magnetic anomalies using multiple geophysical datasets while establishing the regional geological structure of this 60 significantly understudied area. We aim to conduct integrated spatial analysis and develop 61 subsurface models with gravity, magnetic, seismic, and drilling data to understand the regional 62 geological features. Gravity and magnetic data are widely used to understand and characterize 63 64 geological formations, particularly for identifying thin magnetic layers or faults (Adebiyi et al., 2023; Jaffal et al., 2010). High-resolution potential field data help identify structures related to 65 local-scale mineralization that are covered by shallow alluvium or other unconsolidated sediments 66 (Hendrickson, 2016; McCafferty et al., 2014). Since potential field data can yield multiple 67 solutions (Filina et al., 2019). We also incorporate seismic and drilling data to constrain the 68 subsurface framework and produce more reliable results (Sundararajan, 2012). Our study applies 69 70 various filters to magnetic and gravity data to enhance and delineate regional crustal structures, with a focus on highlighting structure edges where metallic mineral deposits are likely to be found 71 (Hildenbrand, 2000). 72







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Figure 1: (a) Regional geological map of Bangladesh adopted from Hossain et al. (2019), illustrating various geological and tectonic zones with distinct color coding. The names of neighboring countries are labeled in red. The red dashed box indicates the study area in the northwestern part of Bangladesh. (b) The map of the northwestern Bangladesh region shows basement depth with its tectonic subdivisions. Regional basement depths are collected from GSB and constructed from various seismic reflection surveys. Red dashed lines mark the approximate boundaries of the tectonic divisions. Seismic profiles (orange lines) and drilling locations (red plus signs) utilized in this study are also highlighted. Our specific study area, Pirganj, is also shown on the map, where the highest magnetic anomaly is observed. The integrated 2-D modeling profile of this study is also shown in this map with a solid black line.

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2. Geological setting

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Bangladesh, though geographically compact, possesses a complex and diverse geological 86 87 framework shaped largely by its position within the Bengal Basin (Roy and Chatterjee, 2015). This region is primarily divided into two major tectonic units (Morgan and McINTIRE, 1959). To the 88 northwest of the hinge zone (see Figure 1a for location), the stable platform region features a 89 shallow basement composed predominantly of Precambrian-aged rocks (Uddin and Lundberg, 90 1998). Conversely, the southeastern portion of the Bengal Basin comprises a geosynclinal basin, 91 distinguished by significant sediment accumulation, with sediment thicknesses exceeding 12 92 kilometers (Alam, 1989). This tectonic configuration highlights the geological contrasts between 93 the stable platform and the more dynamic, subsiding basin to the southeast. 94

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96 The Bengal Basin, encompassing Bangladesh and parts of the neighboring Indian states of West 97 Bengal, Assam, and Tripura, is situated at the northeastern edge of the Indian craton. It is one of





98 South Asia's largest peripheral collisional foreland basins, with a sedimentary sequence spanning from the Early Cretaceous to the Holocene (DeCelles, 2011). The study area, located in the 99 northwest part of the basin and known as the stable platform, consists of three geological 100 components (Hossain et al., 2019): the Dinajpur Shelf, Rangpur Saddle, and the Bogra Shelf 101 (Figure 1b). The Himalayan Foredeep region, located just north of the Dinajpur Shelf, contains 102 the deepest basement in the stable platform (Figure 1b) and hosts numerous faults associated with 103 extensional tectonics (Figure 2a). South of the Foredeep region, the Dinajpur Shelf gently slopes 104 northward toward the Himalayan Foredeep at an angle of 1-3 degrees and is covered by recent 105 sedimentary deposits. The Rangpur Saddle, the southern block of the Dinajpur Shelf, connects the 106 Indian Shield to the Shillong Plateau and contains the shallowest basement in the Bengal Basin 107 (Jain et al., 2020) (Figure 1b). The Bogra Shelf, on the southern slope of the Rangpur Saddle, is 108 formed during the Early Cretaceous rifting of the Indian plate from Gondwana (Alam, 1989; Alam 109 et al., 2003). Both the Rangpur Saddle and Bogra Shelf host several horsts and graben, and half-110 graben basins (Figure 2b). 111

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The entire north-eastern Bangladesh is geologically stable with minimal folding impact that can 113 be traced back from Rodinia and Luna or Colombia supercontinents (Ameen et al., 2007; Zhang 114 et al., 2012). Ameen et al. (2007) suggest that the basement is a separate micro-continental 115 fragment trapped during the northward migration of the Indian Plate, while Hossain et al. (2007) 116 propose that it is a continuation of the central Indian tectonic zone. During the Precambrian, only 117 the stable shelf of the Bengal basin was part of the Indian Plate within Gondwana. By the Middle 118 Jurassic ($\sim 170-175$ Ma), the Indian Plate began drifting, and becoming isolated by the end of the 119 Paleocene (~55.9 Ma) (Hossain et al., 2019). During the Late Paleozoic-Mid Mesozoic, the stable 120 121 shelf was developed as an intra-cratonic rift basin with Gondwana sediments in graben structures, followed by Kerguelen igneous activity and widespread Rajmahal Trap volcanism (Hossain et al., 122 2019; Valdiya, 2016). 123

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The basement in our study area is the shallowest part of the Stable Shelf, which is uplifted to a 125 depth of 128 m from the surface, overlain by the Plio-Pleistocene Dupi Tila Sandstone and 126 Madhupur Clay (Rahman 1987), and is mainly composed of crystalline rocks, including granite, 127 128 granodiorite, and gneiss (Alam et al., 2003; Hossain et al., 2007). There is no outcrop of Precambrian basement in this area, and the commonly observed horst and graben structures control 129 the stratigraphic subdivision. The Precambrian basement in this region lies beneath thick Cenozoic 130 clastic deposits and is primarily felsic in composition, intersected by mafic-ultramafic and 131 occasional felsic dykes (Chowdhury et al., 2022). Fault-bound graben basins within the basement 132 contain Carboniferous rock units from the Permian Period (286 to 245 million years ago) called 133 134 Gondwana formation, marking the oldest sedimentary rocks in Bangladesh (Alam et al., 2003; Jain 135 et al., 2020). Above the Permian Gondwana formation is the Jurassic Rajmahal Trap Formation, consisting of volcanic basalt strata (Alam, 1989; Roy and Chatterjee, 2015). The Shibganj 136 Trapwash Formation overlays it, formed through weathering and erosion of the underlying igneous 137 rocks (Khan, 1991). 138

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140 The Rangpur Saddle serves as the subsurface extension of the Indian shield, stretching between 141 the Shillong Plateau to the east and the Rajmahal Hills to the west. Geophysical studies have 142 identified two major faults framing the Garo-Rajmahal gap: the Dhubri-Jamuna Fault (western 143 edge of Garo Hills) and the Rajmahal Fault (eastern edge of Rajmahal Hills) which encloses the





Rangpur saddle (Hossain et al., 2019). Tectonic activities, particularly extensional tectonics during 144 continental rifting, have significantly disrupted the basement topography, forming numerous 145 horsts and grabens (Ahamed et al., 2020; Khan & Rahman, 1992). Despite this, the study area 146 remains predominantly flat with sediment covers (Khan, 1991) from the Pleistocene to Holocene 147 period (Reimann and Hiller, 1993). Near the Rangpur Saddle, the eastern part of the Indian Shield 148 includes three major tectonic domains: Singhbhum Craton, Singhbhum Mobile Belt, and 149 Chhotanagpur Gneissic Complex (Mukhopadhyay and Matin, 2020). Singhbhum Craton is 150 characterized by prolonged crustal evolution during the Archean, comprising lithologies such as 151 granitoids and metamorphic rocks. Singhbhum Mobile Belt underwent accretion and modification 152 through volcanics, dyke swarms, and various intrusive bodies in the Proterozoic. Chhotanagpur 153 Gneissic Complex comprised mainly of gneisses, amphibolites, and granulites with mafic dyke 154 swarms, forming a structurally complex mobile belt. The tectonics of this region can be 155 characterized by intra-cratonic structural depressions between the uplifted tectonic blocks. Seismic 156 reflection studies indicate that the Moho in this region is complex and laminated, suggesting a 157 history of tectonic and magmatic activity (Valdiya, 2016). In the West Bengal basin, the Moho is 158 relatively shallow and horizontal (at a depth of 36 km), while other nearby regions (e.g., Kutch 159 basin) have a dipping Moho influenced by tectonic processes (Rangin and Sibuet, 2017). 160

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Magnetic minerals, particularly iron ore, are widely found across the Indian Shield. These iron 162 deposits typically appear as metamorphosed banded iron or silica formations. In eastern India, the 163 Precambrian iron ore of the Singhbhum-North Orissa region is a horseshoe-shaped synclinorium 164 165 that contains the most significant iron deposits near Bangladesh. The first discovery of iron ore in Bangladesh is located on the Dinajpur slope of the Rangpur platform (Alam et al., 2003). Around 166 30 km southwest of our study area, Masum et al. (2021) report iron ore-bearing basement rock 167 about 30 km southwest of our study area, with the iron ores occurring as a thin, metamorphosed 168 laminated layer. 169

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To understand the geological succession in Pirganj and its surroundings, we present a generalized lithological depiction using data from EDH-15 and GDH-54 drill holes (Table 1).

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Table 1: Stratigraphic Succession of Pirganj & its adjoining Areas according to drill holes EDH-15 and GDH-54

Age	Rock Units	Lithology	Thickness (m)
Recent	Alluvium	Loose sand, medium to coarse-grained	52
Unconformity			
Pliocene	Dupitila	Sandstone(SST), silty SST, pebbly SST, pebbly bed; SST: medium to coarse-grained, pebbles: quartzite, gneiss & schist	33
Unconformity			
Late Oligocene to Early Miocene	Surma	Alteration of SST and shales and their combination, sand and silty shale. SST: fine to medium-grained; shale: soft and sticky	125
Unconformity			





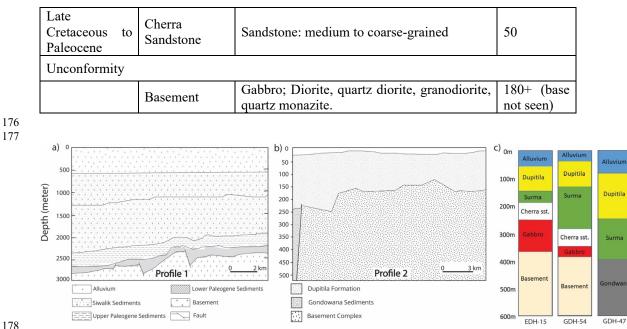


Figure 2: Seismic and drilling data used for integrated geophysical analysis in this study. (a) and (b) present interpretations of the two seismic profiles analyzed. (c) shows the drilling log data, which correlates with the 2D subsurface models and assists in spatial analysis. Refer to Figure 1b for their locations. Seismic profile 1 is from Himalayan Foredeep region while profile 2 represents Rangpur Saddle. All the drilling logs are situated near the highest observed magnetic anomaly in Pirganj, located in Northwestern Bangladesh.

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185 **3. Data and Methodology**

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In this study, we apply an integrative analytical approach utilizing multiple geophysical datasets 187 for spatial analysis and 2-D subsurface modeling. Our primary datasets are potential field data, 188 specifically gravity and magnetic data, provided by the Geological Survey of Bangladesh (GSB). 189 We use Bouguer gravity data and total magnetic intensity data for our analysis. Topography data 190 191 used in our geophysical modeling are obtained from the online repository of the Scripps Institution 192 of Oceanography, which are derived from satellite altimetry (Smith and Sandwell, 1997). Additionally, we incorporate interpretations of seismic images obtained from GSB to correlate our 193 findings from gravity and magnetic data. However, raw seismic images are unavailable, as private 194 entities originally collected them. To further refine our 2-D integrated modeling, we use drilling 195 data from three boreholes near our study area, also provided by GSB (see Figure 1b for the 196 197 locations).

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For geophysical spatial analysis, we use gravity and magnetic anomaly maps. The Bouguer gravity data from GSB are corrected for elevation. For the total magnetic intensity map, we apply a differential reduction to the pole (RTP) to adjust the magnetic grid (Arkani-Hamed, 2007). This correction involves computing magnetic inclination, declination, and total magnetic field values based on the International Geomagnetic Reference Field (Alken et al., 2021) with a magnetic epoch of 1980.





206 The next step in our spatial analysis methodology involves removing the regional trend from both the gravity and magnetic data. This process, known as regional-residual separation, is essential for 207 isolating local anomalies by filtering out the broader, long-wavelength trends associated with 208 large-scale geological structures (Ashraf and Filina, 2023b; Kheyrollahi et al., 2021; Núñez-209 Demarco et al., 2023). By removing these regional trends, we enhance the visibility of smaller-210 scale or high-frequency anomalies, allowing subtle features and variations in the subsurface to be 211 highlighted more effectively. To remove regional anomalies, we apply an upward continuation of 212 1000 m to the Bouguer gravity data (Figure 3a). This approach simulates measuring the gravity 213 field at a higher elevation—1000 m in our case—effectively smoothing out high-frequency 214 anomalies associated with shallow or near-surface geological features (Balogun et al., 2023). For 215 RTP magnetic anomaly, we use a Gaussian filter to calculate the regional trend (Figure 3b). This 216 Gaussian filter acts as a low-pass filter, smoothing out high-frequency components in the dataset. 217 After extracting the regional anomaly from the potential field, we subtract it from the unfiltered 218 total anomaly, yielding the residual anomaly (Figure 3). 219

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We apply several filters to the residual potential field data to enhance specific features and improve 221 interpretability (Figure 3). These filters involve various forms of derivative operations, which help 222 223 to highlight changes in the data that correspond to geological boundaries, faults, or other structural details (Ibraheem et al., 2023; Ma et al., 2016; Nasuti et al., 2019; Núñez-Demarco et al., 2023). 224 For the residual RTP magnetic data, we apply and show two filters: the horizontal derivative and 225 the tilt derivative. The horizontal derivative filter, applied in x-direction (i.e., across the longitudes) 226 and v-direction (i.e., across the latitudes), accentuates lateral changes in the magnetic field, helping 227 to reveal abrupt variations. The tilt derivative filter, however, is especially effective as an edge 228 229 detector. It operates by combining both vertical and horizontal gradients, effectively highlighting the edges of anomalous bodies. The tilt derivative produces values that tend toward zero over flat 230 regions, positive over rising areas, and negative over falling areas, creating a clear demarcation of 231 the edges of magnetic sources. As a result, this filter enhances the boundaries of anomalies and 232 helps pinpoint the locations and shapes of features with minimal distortion across varying depths 233 (Ashraf and Filina, 2023b; Pham and Oliveira, 2023). We apply the lineament mapping techniques 234 to map major structural boundaries from the filtered magnetic data (Ashraf & Filina, 2023a, 2023b; 235 236 Ogah & Abubakar, 2024; Zhang et al., 2024). Our approach focused on identifying gaps between magnetic stripes, changes in the stripe orientation, and a significant reduction in stripe width. We 237 also calculate the analytical signals of the magnetic anomalies to highlight the areas with high 238 magnetizing amplitude (Nabighian, 1972; Roest and Pilkington, 1993) that may illuminate 239 magnetic mineral deposits (Mohamed et al., 2022). To calculate the analytical signal of the residual 240 magnetic data, we first compute the horizontal and vertical derivatives of the magnetic field in the 241 242 x, y, and z directions. The analytical signal was then derived by taking the square root of the sum of the squares of these derivatives, yielding a map that represents the amplitude of the magnetic 243 field independent of direction and highlights the edges of magnetic sources. To validate our 244 interpretations, we cross-reference the magnetic lineaments with gravity data. Before validating 245 with gravity data, we filter the residual gravity data by applying the first vertical derivative and tilt 246 derivative to map major tectonic structures and delineate their boundaries. 247 248

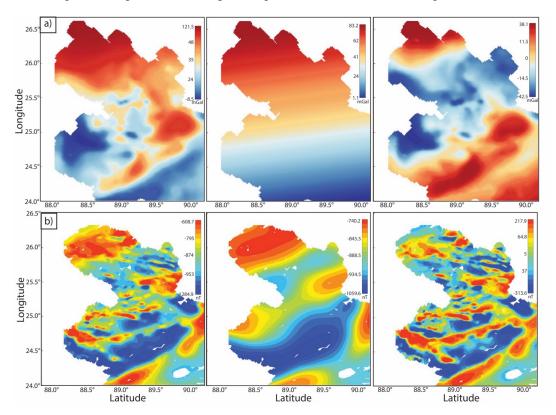
We also develop 2-D integrated models of the subsurface to examine the variations in the physical properties of the rocks (density and magnetic susceptibility). We build our models using the GM-SYS module within the Geosoft software suite, employing a 2-D approximation (Geosoft, 2021).





The GM-SYS model is extended to +/- 30000 km (i.e., infinity) along the X-axis and 90 km along 252 the Z-axis to eliminate edge effects. By integrating gravity and magnetic data within this 2D 253 context, we can delineate major geological boundaries and assess regional structural trends with 254 sufficient accuracy for our research objectives. Our goal is to develop a simple subsurface structure 255 that satisfactorily aligns with gravity, magnetic, and logging data without introducing excessive 256 complexity that might overfit the potential field anomaly. Instead, we aim to replicate the general 257 pattern of the observed anomaly in our 2D modeling, seeking to generate computed anomalies 258 with comparable amplitude, wavelength, and phase to those observed in the potential fields. 259

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Figure 3: Potential field data maps of northwestern Bangladesh used in this study to analyze geological and tectonic features. (a) Gravity anomaly maps: from left to right, the Bouguer gravity anomaly, regional Bouguer gravity anomaly with 1000 m upward continuation, and residual Bouguer gravity anomaly. (b) Magnetic anomaly maps: from left to right, the RTP total magnetic anomaly, regional magnetic anomaly after Gaussian filtering, and residual magnetic anomaly.

269 **4. Result**

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271 4.1 Integrated spatial analysis

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In this study, we first establish the regional tectonic structures of northwestern Bangladesh through spatial analysis of multiple geophysical datasets. We utilize gravity, magnetic, seismic image





interpretations, and drilling log data to characterize the tectonic setup of this region. Our analysis
 of residual magnetic and gravity anomalies reveals that the Rangpur Saddle exhibits distinct
 geophysical characteristics compared to its northern counterpart, the Himalayan Foredeep region.

These differences are discernible in the long-wavelength or broad-scale geophysical signatures

- 279 across the two regions.
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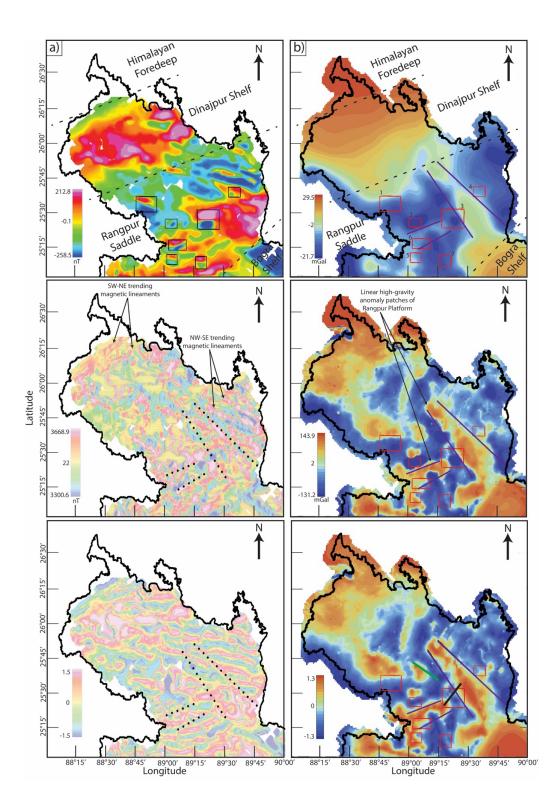
In the Himalayan Foredeep, north of 26°N latitude, we observe a high residual Bouguer gravity 281 anomaly, ranging from approximately 5 to 30 mGal. In contrast, the Rangpur Saddle, located south 282 of this latitude, is marked by a low residual Bouguer gravity anomaly, typically between -22 and 283 0 mGal. Similarly, the Rangpur Saddle shows lower residual RTP total magnetic anomalies, 284 ranging from 0 to 260 nT compared to mostly high magnetic anomalies in north (~ 0 to 215 nT). 285 The spatial patterns of these anomalies also differ across the regions. South of 26°N latitude, in 286 the Rangpur Saddle, the horizontal and tilt derivative magnetic anomalies show a predominance 287 of NW-SE trending magnetic lineaments. North of this latitude, in the Himalayan Foredeep, the 288 magnetic anomaly trend shifts predominantly to a SW-NE orientation. Additionally, In the vertical 289 and tilt derivatives of the gravity data, we observe the most pronounced high-gravity linear 290 anomaly patch trends NW-SE south of 26°N latitude, consistent with the magnetic anomaly trend 291 in this area. South of 26°30'N latitude, we see another linear high-gravity patch in the filtered 292 293 Bouguer gravity data trending NW-SE, indicating a different orientation than the other highgravity patch of the Rangpur Saddle region. These two high-gravity patches are also traceable in 294 the filtered magnetic anomalies, following the magnetic lineament mapping procedure. 295

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Within the Rangpur Saddle region, we observe multiple semi-circular patterns characterized by 297 298 alternating magnetic highs and lows. From the RTP residual magnetic anomaly data, we identify seven distinct patterns. The identification of these patterns follows several consistent criteria. First, 299 the alternating high and low anomalies are generally oriented in a north-south direction. Second, 300 each pattern features a magnetic high in the northern section and a corresponding low to the south. 301 Third, within each pattern, the size, area, and amplitude of the magnetic high and low anomalies 302 are comparable, contributing to a symmetrical structure. Mapping these magnetic patterns of 303 alternating highs and lows onto the filtered gravity maps reveals that they consistently lie near the 304 305 boundary between high and low gravity anomalies. While most of these magnetic patterns (patterns 1, 4, 5, 6, and 7) intersect only one boundary, a few patterns (specifically patterns 2 and 306 3) touch boundaries on both sides. 307











310 Figure 4: a) Magnetic anomaly maps of northwestern Bangladesh. The top panel displays the residual RTP total 311 magnetic anomaly, the middle panel shows the first horizontal derivative in the x-direction (across longitudes), and 312 the bottom panel illustrates the tilt derivative. Black boxes highlight areas where semi-circular magnetic patterns of 313 alternating highs and lows are observed. Dotted black lines in the middle and bottom panels indicate boundaries of 314 high gravity regions. b) Bouguer gravity anomaly maps of northwestern Bangladesh. The top panel presents the 315 residual Bouguer gravity anomaly, the middle panel shows the first vertical derivative, and the bottom panel displays 316 the tilt derivative. High gravity regions are outlined by purple solid lines, and red boxes indicate areas with magnetic patterns of alternating highs and lows. The solid green line in the bottom panel represents seismic profile 2 from Figure 317 318 2. The thick solid black line shows the 2-D integrated modelling profile of Figure 5.

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320 4.2 Integrated subsurface modeling

322 We develop a 2-D subsurface model utilizing gravity, magnetic, and drilling log data (Figure 5). The goal of this model is to approximate the observed anomalies on a local scale while aligning 323 with the available drilling log information (Figure 2c). Importantly, we do not aim to achieve a 324 precise fit between the observed data and calculated anomalies. This decision is driven by the fact 325 326 that the observed potential field data are influenced by regional structures, and without sufficient seismic information, constructing a reliable regional model is not feasible. Instead, we focus on 327 developing a simplified version of the subsurface model that reasonably fits the observed 328 anomalies and drilling data. 329

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To ground the 2-D model in reality, we incorporated available drilling log information (Figure 2). 331 332 Based on these logs, the model includes five sedimentary layers above a crystalline basement. The 333 shallowest layer is assigned as alluvium, extending to a depth of 50 meters or less. Beneath the alluvium, we sequentially include the Dupitila, Surma, and Cherra sandstone layers. The variable 334 thickness of these sedimentary layers comes from the drilling log information, except for the 335 Gondwana layer. The thickness of Gondwana is approximated based on the gravity fit. Below 336 these layers, we model a mafic intrusion of gabbro, which occurs within fault structures that have 337 338 developed within the underlying felsic basement.

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The resulting model indicates that a horst and graben structure is the most plausible configuration when considering the regional geological context. This structural interpretation also provides a reasonable match to the observed geophysical anomalies. Additionally, based on information from drilling log GDH 47, we assign a layer of Gondwana sediments within the graben basin which is contributing to the low observed gravity in this region.

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Densities in the model are derived from drilling data, reflecting the unique lithological characteristics of each unit. The alluvium, Dupitila, Surma, and Cherra Sandstone layers are modeled with densities of 2000, 2200, 2500, and 2670 g/cm³, capturing their progressive compaction and mineral composition. The Gondwana unit, enriched with coal deposits, exhibits a notably lower density of 2200 g/cm³, consistent with its organic-rich composition. Beneath these layers, the felsic basement and gabbroic intrusion stand out with densities of 2800 and 3000 g/cm³, highlighting their denser crystalline structure and mafic origins.

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Between 400 and 800 meters depth, the gabbroic intrusion begins to follow the fault plane. The thickness of both the Gondwana layer and the underlying felsic basement is determined from the amplitude of the calculated gravity and magnetic anomalies. The thickness of the Gondwana layer influences the minimum value of the calculated gravity anomaly, while the amplitude of the



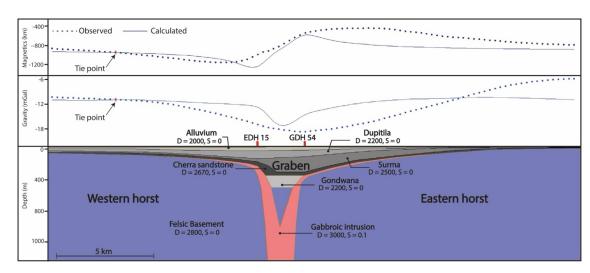


magnetic anomaly helps determine the depth of the gabbroic intrusion. Notably, our calculated anomalies are of high frequency. This occurs because our modeling does not incorporate all regional structures; instead, it focuses only on local structures. As a result, the calculated anomalies reflect primarily local or high-frequency potential field anomalies.

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Our 2D subsurface model reveals two horst structures with a graben basin in the middle. We refer to these as the eastern and western horsts based on their geographic locations. In the model, the western horst structure is depicted as deeper, consistent with the observed gravity anomaly, which shows lower gravity in the western horst compared to the eastern horst. Our model does not include the larger horst structure on the eastern side, which likely explains the discrepancy between the

- 368 calculated and observed anomalies over the eastern horst.
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Figure 5: Integrated 2D geophysical model over the highest magnetic anomaly in northwestern Bangladesh. See Figure b and 4c for the location of the model profile. The top panel presents the magnetic anomaly, and the middle panel shows the gravity anomaly, with observed (dotted) and calculated (solid) data. Tie points, marked by red stars, indicate where the calculated anomalies are vertically shifted to align with the observed data. The bottom panel illustrates the subsurface model, with geological units represented by distinct colors. 'D' denotes density (g/cm³), and 'S' denotes magnetic susceptibility (SI units). The red vertical dashes show the location of the drilling logs used to constrain the subsurface units.

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380 **5. Discussion**

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382 **5.1 Basement structure**

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One of the primary goals of this paper is to understand the basement structure and tectonic setup of the Rangpur Saddle. Our findings reveal significant differences in the basement structures of the Rangpur Saddle compared to its northern counterparts, evident in both gravity and magnetic data. A key distinction is the depth to basement between these regions (**Figure 1b**). Previous studies consistently suggest that the basement depth in the Rangpur Saddle is shallower than in the Dinajpur Shelf. However, those studies were based on limited seismic and drilling data. Our results, however, indicate a more complex basement geometry than previously understood.





391 Filtered gravity and seismic data presented here suggest that the Rangpur Saddle region contains both shallow and deep basement features (Figure 4b). Considering regional tectonics, we propose 392 that high gravity values correspond to shallow horst structures, while low gravity values indicate 393 deeper graben structures. However, the frequency of horsts and grabens is significantly lower in 394 the Rangpur Saddle compared to the northern Himalayan Foredeep and the Dinajpur Shelf (Figure 395 2a & 2b). In the Rangpur region, we identified only two horsts of notable width and length. The 396 boundaries of these horst structures also appear as lineaments in the filtered magnetic data (Figure 397 4a), suggesting the presence of major structural boundaries. 398

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Furthermore, differences in the directions of magnetic lineaments between the Rangpur Saddle 400 region and its northern counterparts-the Dinajpur Shelf and the Himalayan Foredeep-suggest 401 variations in paleo-tectonic stress orientation (Figure 4a, 6a & 6b). Existing literature 402 403 characterizes northwestern Bangladesh predominantly with extensional tectonics (Royhan Gani and Mustafa Alam, 2003). Near the Himalayan Foredeep boundary, SW-NE trending magnetic 404 lineaments indicate a paleo-principal stress direction oriented NW-SE. Moving southward into the 405 Dinajpur Shelf, NW-SE trending magnetic lineaments become more common (Figure 6b), 406 pointing to a shift in the principal stress direction associated with extensional tectonics. In the 407 408 Rangpur Saddle, these NW-SE trending lineaments are even more frequent, especially in the eastern part, indicating a greater intensity of paleo-tectonic stress in this region compared to the 409 Dinajpur Shelf. Based on traced magnetic and gravity lineaments, the stable platform can be 410 divided into four distinct zones (Figure 6d). In Zone 1, located in the northern part, lineaments 411 predominantly trend SW-NE, with only a few exceptions. Moving southward into Zone 2, there is 412 a mixture of two differently trending lineaments: SW-NE and NW-SE. In Zone 3, in the 413 414 easternmost section of the stable platform, lineaments primarily trend NW-SE, aligning with the eastern horst. Finally, in Zone 4, the southernmost region, lineaments generally follow an SW-NE 415 trend, corresponding with the western horst. This subdivision highlights the complex nature of past 416 417 tectonic processes in this region.

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We also identify seven semi-circular patterns of alternating magnetic highs and lows. When 419 mapped onto gravity data, these patterns consistently align with the boundaries between high and 420 421 low gravity patches, indicating the boundaries between horsts and grabens. Drilling log data near the most prominent magnetic anomaly reveals the presence of gabbro, which may explain these 422 alternating magnetic highs and lows. Furthermore, these regions of alternating magnetic patterns 423 also spatially correlate with overall high magnetization (Figure 6a). We propose that these gabbro 424 formations resulted from intrusions along normal faults, formed by extensional tectonics, which 425 define the boundaries between horsts and grabens. 426





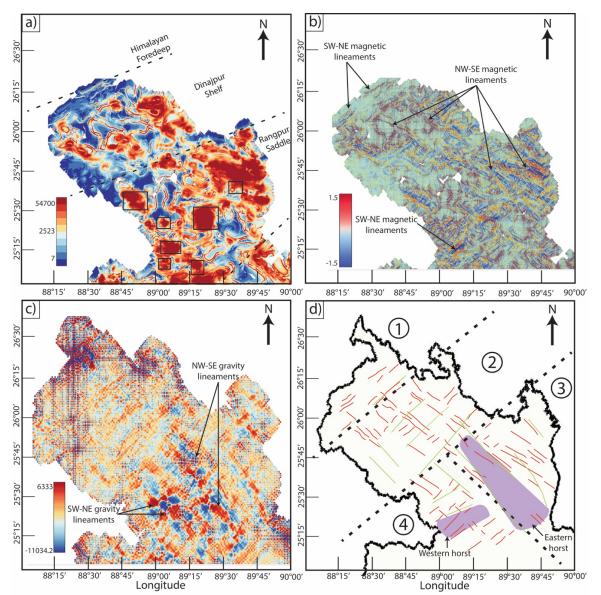




Figure 6: Filtered potential field maps of the northwestern stable platform region of Bangladesh, highlighting major 430 tectonic structures and boundaries. (a) Analytical signal map derived from the magnetic anomaly (methodology 431 detailed in Sect. 3), showing magnetization amplitudes across the region. Black boxes indicate areas of semi-circular magnetic patterns with alternating highs and lows, as shown in Figure 4a. Dashed lines represent boundaries between 432 433 established tectonic subdivisions. (b) Total horizontal gradient map of the magnetic anomaly, calculated by applying 434 a first horizontal derivative filter in the X-direction, followed by a first horizontal derivative filter in the Y-direction. 435 (c) Total horizontal derivative of the residual gravity anomaly. (d) Map displaying traced magnetic lineaments (in red) from 'b' and gravity lineaments (in green) from 'c'. Black dashed lines indicate regional subdivisions based on mapped 436 437 tectonic lineaments. Purple polygons mark horst structures in the Rangpur Saddle, interpreted from filtered gravity 438 data (refer to Figure 4b).





440 **5.2 Economic mineral potential**

441

Our integrated spatial analysis and 2-D modeling indicate that the Rangpur Saddle hosts at least 442 two significant horst and graben structures. The 2-D modeling results, corroborated by drilling log 443 data, confirm the presence of gabbroic bodies along the normal faults within these structures 444 (Figure 7). We propose that these gabbroic intrusions result from magmatic emplacement along 445 extensional faults, consistent with processes typically observed during the early stages of 446 continental breakup. The breakup process involved rifting events that progressively separated the 447 Indian subcontinent from the other Gondwanaland constituents (Veevers, 2004). Key rifting 448 episodes and magmatic activities, such as the formation of the Central Atlantic Magmatic Province 449 and related tectonic movements, facilitated this separation (Thompson et al., 2019; Veevers, 2004). 450 451 In such tectonic settings, normal faulting and rift-related subsidence create pathways for mafic magmas to ascend (Brune et al., 2023; Pirajno and Santosh, 2014; Ruppel, 1995), leading to the 452 emplacement of gabbroic and other mafic intrusions within the crust (Magee et al., 2019). 453

454

Mafic magma intrusions are commonly observed along extensional fault planes in tectonically 455 active regions worldwide. Troll et al. (2021) shows that in NW Scotland, Long Loch Fault acted 456 as dynamic magma conduit with its movements facilitating the ascent of ultrabasic magmas. These 457 magmas intruded along faults and fractures, leading to the destabilization and collapse of existing 458 cumulate layers, forming extensive breccias. A study on Mesozoic gabbroic intrusions in the High 459 Atlas Mountains highlights their emplacement along extensional faults during the rifting 460 associated with the Central Atlantic Magmatic Province, where more than 50% of gabbro samples 461 exhibit stable magnetization from magnetite which is identified as the primary component (Calvín 462 463 et al., 2017). Fuller & Waters (1929) have extensively studied horsts and graben structures in southern Oregon where they found emplacement of volcanic and intrusive rocks are closely 464 associated with extensional faulting. These normal faults and the resulting grabens provide 465 pathways for magma to ascend and intrude, which is evident from the prevalence of rhyolitic, 466 dacitic, and andesitic vents as well as basaltic dikes aligned with the faults. 467

468

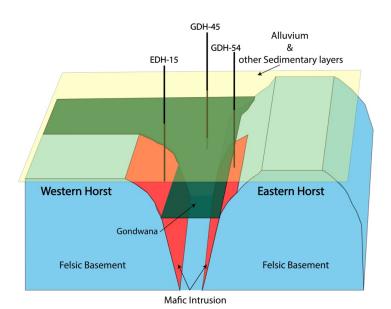
Mafic intrusions along fault zones are often associated with magnetic mineral deposits, which can
be economically significant as ore bodies (Zhou et al., 2005). For instance, magnetite deposits are
found within mafic intrusions along the fault planes in Egypt, which exhibit high magnetic
anomalies, facilitating their detection and delineation (Kharbish et al., 2022; Mousa et al., 2020).
Also, in Central Iran, iron-bearing magnetic mineral deposits are observed to exhibit high magnetic
and gravity anomalies (Kheyrollahi et al., 2021).

475

Based on current geochemistry data, the Rangpur Saddle holds significant potential for magnetic mineral ore deposits, driven by its mafic-ultramatic sequences and iron-rich dykes (Ameen et al., 2021). High Fe₂O₃ content in hornblendites and associated minerals like magnetite and ilmenite indicate strong magnetization. The tectonic setting, with extensive magmatic intrusions, provides ideal conditions for mineralization. These findings highlight the region as a promising target for future exploration of magnetic ore bodies.







482 483

Figure 7: 3D schematic diagram illustrating the tectonic structures in Pirganj, where the highest magnetic anomaly is observed in Bangladesh. The felsic basement, depicted in blue blocks, includes two horsts, as shown in Figure 5d. Possible mafic intrusions along the faults are indicated in red. Sedimentary layers are represented by a single horizontal layer in specific colors, with the graben basin likely filled with Gondwana sediments shown in green, and other sedimentary layers, including alluvium, shown in yellow. The approximate locations of the three drilling logs used in this study (refer to Figure 1b for their locations) are also indicated.

490

491 5.3 Limitations and Future Research

492

Compared to other regions globally, limited published research exists for our study area, restricting 493 our ability to draw on established models and findings. Situated at the eastern edge of the Indian 494 Shield, this area transitions into a different tectonic subdivision, creating a tectonic complexity 495 496 that cannot be fully resolved with the available low-resolution data. The scarcity of high-resolution 497 datasets, such as regional-scale seismic surveys, further limits our capacity to delineate regional structures accurately. Additionally, the available well log data are predominantly shallow, reaching 498 499 depths of approximately 500 meters, which limits insights into deeper geological and tectonic features beyond 2 kilometers. Petrographic descriptions of the basement rocks are also 500 rudimentary, as they are primarily based on a limited number of well logs, most of which were 501 502 drilled before the 1990s, leaving significant gaps in our understanding of the area's deeper 503 subsurface geology.

504

Future research in this region will focus on developing 3D models of the subsurface, providing a 505 more detailed understanding of its geological structure. This effort will include creating a 506 comprehensive dataset repository that incorporates historical seismic reflection data from private 507 entities. Given the shallow depth of the region, high-resolution gravity and magnetic data will also 508 509 be essential for a thorough analysis. This may require precision measurements through ground surveys to accurately collect gravity and magnetic data. The study area is characterized by complex 510 tectonic structures shaped by a diverse tectonic history. While the current study has focused on the 511 Rangpur Saddle, future research will expand to cover the entire northwestern part of Bangladesh, 512





including the Himalayan Foredeep and the Eocene Hinge region, which represents a
Paleocontinental shelf. Such expanded coverage will provide deeper insights into the tectonic
evolution and geological complexity of the area.

516

517 Conclusion

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519 This study provides a comprehensive geophysical investigation of the northwestern region of Bangladesh, revealing its significant potential for mineral resource exploration. The integration of 520 gravity, magnetic, seismic, and drilling data has allowed us to identify key tectonic features, such 521 as gabbroic intrusions along extensional faults, which suggest the presence of valuable magnetic 522 mineral deposits. Despite the limitations in data resolution and coverage, our findings offer 523 important insights into the region's geological evolution and resource potential. Future work 524 should prioritize acquiring more detailed geophysical data and expanding drilling campaigns to 525 refine the understanding of subsurface structures and evaluate the feasibility of mineral extraction. 526 527 The Rangpur Saddle, as the shallowest part of the stable platform, emerges as a promising target for future mineral exploration endeavors. 528

529

530 Data availability

- 531
- All raw data can be provided by the corresponding authors upon request.
- 532 533

534 Author contribution

535

Mohammad Tawhidur Rahman Tushar was involved in conceptualization, methodology, software, 536 validation, formal analysis, investigation, resources, data curation, writing - original draft, writing 537 538 - review and editing, visualization. Asif Ashraf contributed to conceptualization, methodology, software, validation, formal analysis, investigation, resources, data curation, writing - original 539 draft, writing - review and editing, and visualization. Md. Mahfuz Alam worked on software 540 development, validation, and formal analysis. Md Nasif Jamil contributed to writing - review and 541 editing and visualization. Saba Karim was involved in writing - review and editing and 542 visualization. Md. Shahjahan contributed to conceptualization, investigation, resources, project 543 544 administration, and supervision. Md. Anwar Hossain Bhuiyan was involved in writing - review and editing, project administration, supervision, and funding acquisition. 545

546

547 **Competing interests**

548

549 The authors declare that they have no conflict of interest.

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- 552

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560

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