

1 Integrated Geophysical Analysis of Rangpur Saddle: Insights on Tectonics and 2 Magnetic Mineral Potential of North-Western Bangladesh

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12 13 Abstract

14
15 The northwestern region of Bangladesh holds untapped potential for magnetic mineral deposits at
16 shallow depths. Unlike much of Bangladesh, characterized by thick sediments of the Bengal Basin,
17 this area is an extension of the Indian Shield, often referred to as the Stable Platform. It is also
18 geologically distinct, hosting structures related to the breakup of Pangea. The geology and
19 tectonics of this region have remained largely understudied. To address this gap, this study
20 integrates gravity, magnetic, seismic, and drilling data to investigate the subsurface structure and
21 evaluate the resource potential of the area. We utilize advanced filtering and modeling techniques,
22 including tilt derivatives and horizontal gradient methods, to understand the tectonic framework
23 and geometry of the subsurface structures. Our spatial analysis, using multiple geophysical
24 datasets, reveals dipolar magnetic anomalies, which we attribute to gabbroic intrusions along
25 extensional faults that define the region's horst and graben structures. To validate our
26 interpretations, we developed an integrated 2-D subsurface model that aligns with the observed
27 geophysical data. However, the study is limited by the availability of high-resolution seismic data
28 and the sparse distribution of drilling locations, which may affect the precision of our subsurface
29 characterization. Our findings provide crucial insights into the tectonic evolution of the stable
30 platform and underscore the economic potential of the Rangpur Saddle, the shallowest part of the
31 stable platform, for mineral exploration. These insights pave the way for further exploration and
32 development initiatives focused on uncovering the mineral wealth of this underexplored region.

33 34 1. Introduction

35
36 The northwestern part of Bangladesh is rich in potential mineral resources. The geological
37 diversity of this area suggests the presence of other valuable minerals that remain largely
38 unexplored (Akhtar, 2005; Hasan et al., 2023; Moon, 2022). In this area, basement rocks are
39 present at relatively shallow depths of 128 meters, hosting a spectrum of economic mineral
40 resources, including coal, limestone, white clay, and hard rock (Khan and Rahman, 1992).
41 Notably, Pirganj in the Rangpur district (**Figure 1**) records the country's highest magnetic
42 anomaly, suggesting the potential for magnetic mineral ore deposits. In the 1990s, drilling
43 activities were conducted in the region after developing a 2D subsurface model by the Geological
44 Survey of Bangladesh (GSB) in collaboration with the United States Geological Survey (USGS).
45 Despite these efforts, no noteworthy ore body was identified during the drilling process (Rahman

46 and Ullah, 2009). Moreover, the tectonic evolution of the Paleo-Proterozoic basement in the
47 northwest region of the Bengal Basin still needs to be studied.

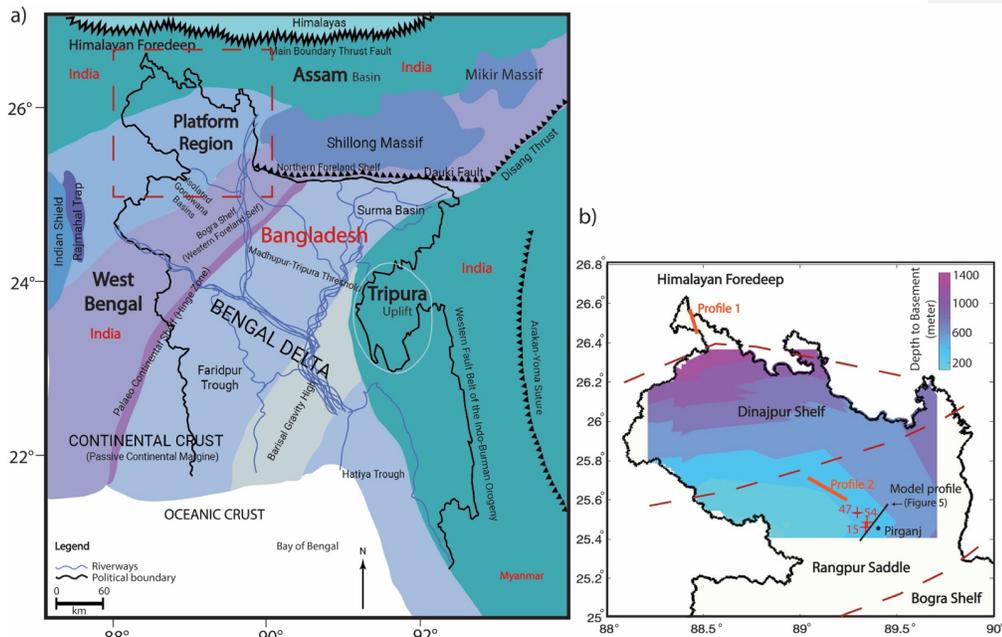
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49 The study area is located in a region commonly referred to as the Rangpur platform or saddle
50 (Masum et al., 2021), an eastern extension of the Indian Shield (Alam et al., 2003). The northern
51 part of the Rangpur Saddle rests on a shallow Precambrian basement, ranging from 130 to 1000
52 meters in depth. This area of Bangladesh is geologically stable, characterized predominantly by
53 horsts and grabens, which were formed during the Cretaceous rifting of the Indian plate from the
54 Antarctica-Australia section of Gondwanaland (Curry, 1991; Curry and Moore, 1974). Although
55 the basement primarily consists of diorite, tonalite, and granodiorite, it is also intersected by
56 pegmatite and mafic/ultramafic dykes (Hossain et al., 2007; Kabir et al., 2001).

57

58 In this study, we explore the possibility of mafic dykes as the source of high magnetic anomalies
59 using multiple geophysical datasets while establishing the regional geological structure of this
60 significantly understudied area. We aim to conduct integrated spatial analysis and develop
61 subsurface models with gravity, magnetic, seismic, and drilling data to understand the regional
62 geological features. Gravity and magnetic data are widely used to understand and characterize
63 geological formations, particularly for identifying thin magnetic layers or faults (Adebiyi et al.,
64 2023; Jaffal et al., 2010). High-resolution potential field data help identify structures related to
65 local-scale mineralization that are covered by shallow alluvium or other unconsolidated sediments
66 (Hendrickson, 2016; McCafferty et al., 2014). Since potential field data can yield multiple
67 solutions (Filina et al., 2019), we also incorporate seismic and drilling data to constrain the
68 subsurface framework and produce more reliable results (Sundararajan, 2012). Our study applies
69 various filters to magnetic and gravity data to enhance and delineate regional crustal structures,
70 with a focus on highlighting structure edges where metallic mineral deposits are likely to be found
71 (Hildenbrand, 2000).

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73
 74 Figure 1: (a) Regional geological map of Bangladesh adopted from Hossain et al. (2019), illustrating various
 75 geological and tectonic zones with distinct color coding. The names of neighboring countries are labeled in red. The
 76 red dashed box indicates the study area in the northwestern part of Bangladesh. (b) The map of the northwestern
 77 Bangladesh region shows basement depth with its tectonic subdivisions. Regional basement depths are collected from
 78 GSB and constructed from various seismic reflection surveys. Red dashed lines mark the approximate boundaries of
 79 the tectonic divisions. Seismic profiles (orange lines) and drilling locations (red plus signs) utilized in this study are
 80 also highlighted. Our specific study area, Pirganj, is also shown on the map, where the highest magnetic anomaly is
 81 observed. The integrated 2-D modeling profile of this study is also shown in this map with a solid black line.
 82

83 2. Geological setting

84
 85 Bangladesh, though geographically compact, possesses a complex and diverse geological
 86 framework shaped largely by its position within the Bengal Basin (Roy and Chatterjee, 2015). This
 87 region is primarily divided into two major tectonic units (Morgan and McIntire, 1959). To the
 88 northwest of the hinge zone (see **Figure 1a** for location), the stable platform region features a
 89 shallow basement composed predominantly of Precambrian-aged rocks (Uddin and Lundberg,
 90 1998). Conversely, the southeastern portion of the Bengal Basin comprises a geosynclinal basin,
 91 distinguished by significant sediment accumulation, with sediment thicknesses exceeding 12
 92 kilometers (Alam, 1989). This tectonic configuration highlights the geological contrasts between
 93 the stable platform and the more dynamic, subsiding basin to the southeast.
 94

95 The Bengal Basin, encompassing Bangladesh and parts of the neighboring Indian states of West
 96 Bengal, Assam, and Tripura, is situated at the northeastern edge of the Indian craton. It is one of

97 South Asia's largest peripheral collisional foreland basins, with a sedimentary sequence spanning
98 from the Early Cretaceous to the Holocene (DeCelles, 2011). The study area, located in the
99 northwest part of the basin and known as the stable platform, consists of three geological
100 components (Hossain et al., 2019): the Dinajpur Shelf, Rangpur Saddle, and the Bogra Shelf
101 (**Figure 1b**). The Himalayan Foredeep region, located just north of the Dinajpur Shelf, contains
102 the deepest basement in the stable platform (**Figure 1b**) and hosts numerous faults associated with
103 extensional tectonics (**Figure 2a**). South of the Foredeep region, the Dinajpur Shelf gently slopes
104 northward toward the Himalayan Foredeep at an angle of 1-3 degrees and is covered by recent
105 sedimentary deposits. The Rangpur Saddle, the southern block of the Dinajpur Shelf, connects the
106 Indian Shield to the Shillong Plateau and contains the shallowest basement in the Bengal Basin
107 (Jain et al., 2020) (**Figure 1b**). The Bogra Shelf, on the southern slope of the Rangpur Saddle, was
108 formed during the Early Cretaceous rifting of the Indian plate from Gondwana (Alam, 1989; Alam
109 et al., 2003). Both the Rangpur Saddle and Bogra Shelf host several horsts and graben, and half-
110 graben basins (**Figure 2b**).
111

112 The entire northwestern Bangladesh is geologically stable with minimal folding impact that can
113 be traced back to Rodinia and Nuna or Columbia supercontinents (Ameen et al., 2007; Zhang et
114 al., 2012). Ameen et al. (2007) suggest that the basement is a separate micro-continental fragment
115 trapped during the northward migration of the Indian Plate, while Hossain et al. (2007) propose
116 that it is a continuation of the central Indian tectonic zone. During the Precambrian, only the stable
117 shelf of the Bengal basin was part of the Indian Plate within Gondwana. By the Middle Jurassic
118 (~170–175 Ma), the Indian Plate began drifting and became isolated by the end of the Paleocene
119 (~55.9 Ma) (Hossain et al., 2019). During the Late Paleozoic–Mid Mesozoic, the stable shelf was
120 developed as an intra-cratonic rift basin with Gondwana sediments in graben structures, followed
121 by Kerguelen igneous activity and widespread Rajmahal Trap volcanism (Hossain et al., 2019;
122 Valdiya, 2016).
123

124 The basement in our study area is the shallowest part of the Stable Shelf, which is uplifted to a
125 depth of 128 m from the surface, overlain by the Plio-Pleistocene Dupi Tila Sandstone and
126 Madhupur Clay, and is mainly composed of crystalline rocks, including granite, granodiorite, and
127 gneiss (Alam et al., 2003; Hossain et al., 2007). There is no outcrop of Precambrian basement in
128 this area, and the commonly observed horst and graben structures control the stratigraphic
129 subdivision. The Precambrian basement in this region lies beneath thick Cenozoic clastic deposits
130 and is primarily felsic in composition, intersected by mafic-ultramafic and occasional felsic dykes
131 (Chowdhury et al., 2022). Fault-bound graben basins within the basement contain Carboniferous
132 rock units from the Permian Period (286 to 245 million years ago) called the Gondwana formation,
133 marking the oldest sedimentary rocks in Bangladesh (Alam et al., 2003; Jain et al., 2020). Above
134 the Permian Gondwana formation is the Jurassic Rajmahal Trap Formation, consisting of volcanic
135 basalt strata (Alam, 1989; Roy and Chatterjee, 2015). The Shibganj Trapwash Formation overlays
136 it, formed through weathering and erosion of the underlying igneous rocks (Khan, 1991).
137

138 The Rangpur Saddle serves as the subsurface extension of the Indian shield, stretching between
139 the Shillong Plateau to the east and the Rajmahal Hills to the west. Geophysical studies have
140 identified two major faults framing the Garo-Rajmahal gap: the Dhubri-Jamuna Fault (western
141 edge of Garo Hills) and the Rajmahal Fault (eastern edge of Rajmahal Hills), which encloses the
142 Rangpur saddle (Hossain et al., 2019). Tectonic activities, particularly extensional tectonics during

143 continental rifting, have significantly disrupted the basement topography, forming numerous
 144 horsts and grabens (Ahamed et al., 2020; Khan and Rahman, 1992). Despite this, the study area
 145 remains predominantly flat with sediment covers (Khan, 1991) from the Pleistocene to the
 146 Holocene period (Reimann and Hiller, 1993). Near the Rangpur Saddle, the eastern part of the
 147 Indian Shield includes three major tectonic domains: Singhbhum Craton, Singhbhum Mobile Belt,
 148 and Chhotanagpur Gneissic Complex (Mukhopadhyay and Matin, 2020). Singhbhum Craton is
 149 characterized by prolonged crustal evolution during the Archean, comprising lithologies such as
 150 granitoids and metamorphic rocks. Singhbhum Mobile Belt underwent accretion and modification
 151 through volcanics, dyke swarms, and various intrusive bodies in the Proterozoic. Chhotanagpur
 152 Gneissic Complex comprises mainly of gneisses, amphibolites, and granulites with mafic dyke
 153 swarms, forming a structurally complex mobile belt. The tectonics of this region can be
 154 characterized by intra-cratonic structural depressions between the uplifted tectonic blocks. Seismic
 155 reflection studies indicate that the Moho in this region is complex and laminated, suggesting a
 156 history of tectonic and magmatic activity (Valdiya, 2016). In the West Bengal basin, the Moho is
 157 relatively shallow and horizontal (at a depth of 36 km), while other nearby regions (e.g., Kutch
 158 basin) have a dipping Moho influenced by tectonic processes (Rangin and Sibuet, 2017).

159
 160 Magnetic minerals, particularly iron ore, are widely found across the Indian Shield. These iron
 161 deposits typically appear as metamorphosed banded iron or silica formations. In eastern India, the
 162 Precambrian iron ore of the Singhbhum-North Orissa region is a horseshoe-shaped synclinorium
 163 that contains the most significant iron deposits near Bangladesh. The first discovery of iron ore in
 164 Bangladesh is located on the Dinajpur slope of the Rangpur platform (Alam et al., 2003). Around
 165 30 km southwest of our study area, Masum et al. (2021) report iron ore-bearing basement rock
 166 about 30 km southwest of our study area, with the iron ores occurring as a thin, metamorphosed
 167 laminated layer.

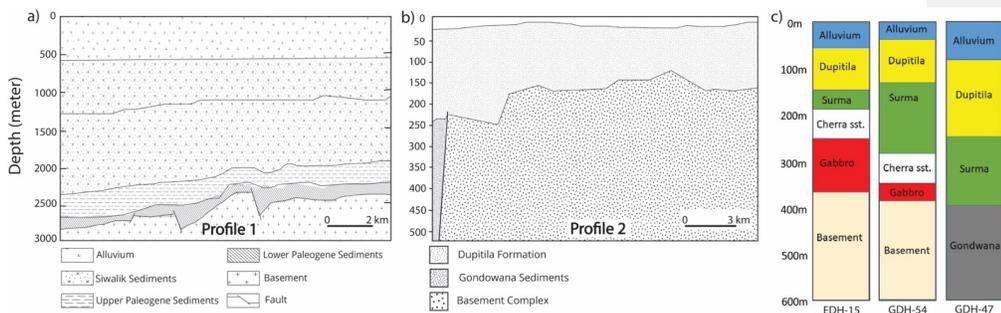
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 169 To understand the geological succession in Pirganj and its surroundings, we present a generalized
 170 lithological depiction using data from EDH-15 and GDH-54 drill holes (Table 1).

171 Table 1: Stratigraphic Succession of Pirganj and its adjoining Areas according to drill holes EDH-15 and GDH-54
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 173

Age	Rock Units	Lithology	Thickness (m)
Recent	Alluvium	Loose sand, medium to coarse-grained	52
Unconformity			
Pliocene	Dupitila	Sandstone (SST), silty SST, pebbly SST, pebbly bed; SST: medium to coarse-grained, pebbles: quartzite, gneiss and schist	33
Unconformity			
Late Oligocene to Early Miocene	Surma	Alteration of SST and shales and their combination, sand and silty shale. SST: fine to medium-grained; shale: soft and sticky	125
Unconformity			

Late Cretaceous to Paleocene	Cherra Sandstone	Sandstone: medium to coarse-grained	50
Unconformity			
	Basement	Gabbro; Diorite, quartz diorite, granodiorite, quartz monazite.	180+ (base not seen)

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Figure 2: Interpretation of seismic and drilling data used for integrated geophysical analysis in this study. (a) and (b) present interpretations of the two seismic profiles analyzed. (c) shows the drilling log data, which correlates with the 2D subsurface models and assists in spatial analysis. Refer to Figure 1b for their locations. Seismic profile 1 is from Himalayan Foredeep region while profile 2 represents Rangpur Saddle. All the drilling logs are situated near the highest observed magnetic anomaly in Pirganj, located in Northwestern Bangladesh.

3. Data and Methodology

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3.1 Geophysical Data

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In this study, we apply an integrative analytical approach utilizing multiple geophysical datasets for spatial analysis and 2-D subsurface modeling. Our primary datasets are potential field data, specifically gravity and magnetic data, provided by the Geological Survey of Bangladesh (GSB). We use Bouguer gravity data and total magnetic intensity data for our analysis.

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The gravity data presented in this paper represent a cumulative dataset acquired through successive land-based surveys initiated in the late 1970s and continuing to the present day. These surveys were primarily conducted during the dry season (November–April) to ensure accessibility, gradually covering approximately 8,000 km² in northwestern Bangladesh. Data acquisition methodology evolved over the decades to incorporate technological advancements. During the earlier campaigns (starting in the 1970s), gravity measurements were taken using an analog Sodin Worden gravimeter (Model WS 410), while surveys conducted after 2007 also deployed the digital Scintrex CG-5 AutoGrav. Cross-verification between the analog and digital instruments was performed regularly in later phases to ensure consistency across the historical dataset. Similarly, positioning methods were modernized over time; while earlier observation points relied on standard benchmarks from the Survey of Bangladesh, later campaigns utilized handheld GPS receivers and the CG-5's integrated GPS for precise coordinate determination.

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211 Throughout the survey history, observation points were generally spaced 1–1.5 km apart.
212 Elevations were determined using digital leveling referenced to benchmarks from the Survey of
213 Bangladesh, maintaining high vertical accuracy crucial for gravity corrections. The Sylhet Gravity
214 Base Station, connected to the IGSN 71 network, served as the primary national reference, with
215 local sub-bases established to correct for instrument drift. Standard geophysical corrections—
216 including those for instrumental drift, tidal and latitude variations, and elevation differences—
217 were applied to the entire dataset. Bouguer corrections were calculated using a crustal density of
218 2.0 kg/m³.

219
220 Between 1979 and 1980, Hunting Geology & Geophysics Ltd. completed a nationwide
221 aeromagnetic survey for the Government of Bangladesh, under the auspices of the Geological
222 Survey of Bangladesh and Petrobangla. Flying a Geometrics G-803 proton magnetometer just 500
223 ft (≈152 m) above ground, the crew collected total-field data along flight lines oriented N 45° W
224 on a nominal 3 km grid (locally tightened to 1 km) and crossed them with tie-lines oriented N 45°
225 E at 5 km spacing. Measurements were recorded every two seconds with a resolution of ±0.05 nT,
226 within a regional field that varied from 44 848 nT to 47 086 nT (inclinations 28° 30'–38° 30',
227 declinations 13° W–37° W). All readings were archived in both digital and analogue form, and the
228 data were uniformly shifted upward by 900 nT so that every value is positive.

229
230 The topography data used in our geophysical modelling (Figure S1) are obtained from the online
231 repository of the Scripps Institution of Oceanography, which are derived from satellite altimetry
232 (Smith and Sandwell, 1997). Additionally, we incorporate interpretations of seismic images
233 obtained from GSB to correlate our findings from gravity and magnetic data. However, raw
234 seismic images are unavailable, as private entities originally collected them. To further refine our
235 2-D integrated modeling, we use drilling data from three boreholes near our study area, also
236 provided by GSB (see Figure 1b for the locations).

237 3.2 Methods

238
239 For geophysical spatial analysis, we use gravity and magnetic anomaly maps. The Bouguer gravity
240 data from GSB are corrected for elevation. For the total magnetic intensity map, we apply a
241 differential reduction to the pole (RTP) to adjust the magnetic grid (Arkani-Hamed, 2007). This
242 correction involves computing magnetic inclination, declination, and total magnetic field values
243 based on the International Geomagnetic Reference Field (Alken et al., 2021) with a magnetic epoch
244 of 1980.

245
246 The next step in our spatial analysis methodology involves removing the regional trend from both
247 the gravity and magnetic data. This process, known as regional-residual separation, is essential for
248 isolating local anomalies by filtering out the broader, long-wavelength trends associated with
249 large-scale geological structures (Ashraf and Filina, 2023b; Kheyrollahi et al., 2021; Núñez-
250 Demarco et al., 2023). Total magnetic intensity anomalies arise from the combined effect of
251 induced and remanent magnetization in rocks. Induced magnetization is caused by the Earth's
252 ambient field acting on magnetic minerals, so it is aligned parallel to the present field, whereas
253 remanent magnetization is a permanent magnetization inherent to the rocks (acquired in the past)
254 that often points in a different direction (reflecting the Earth's field at the time of rock formation).
255 The total magnetization is the vector sum of the induced and remanent contributions. As a result,
256 the direction and relative magnitude of each component strongly influence the observed anomaly.

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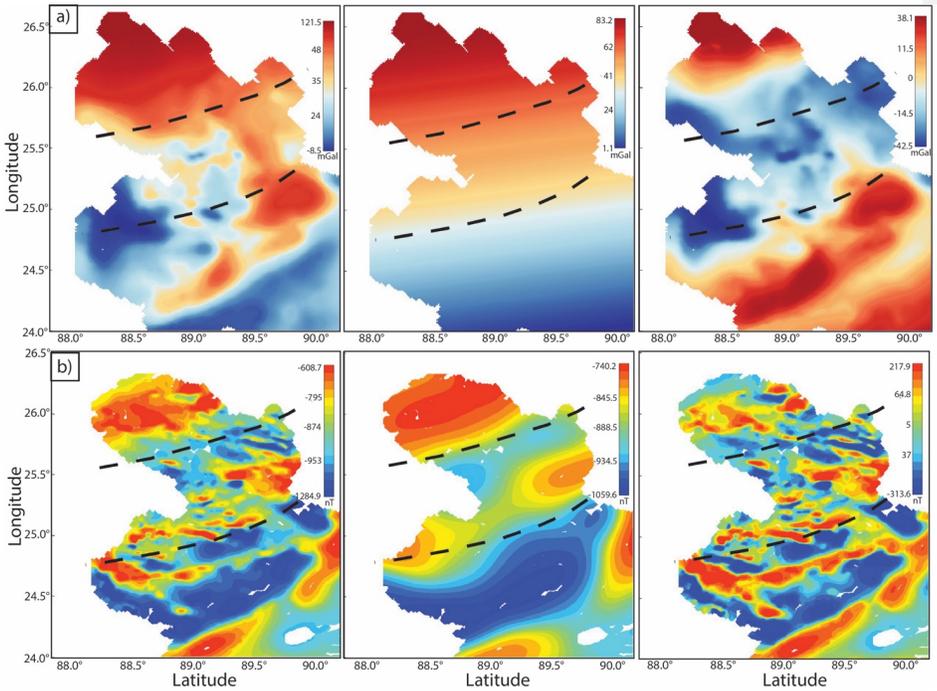
271 If the induced and remanent magnetization vectors are aligned, they reinforce each other to
272 produce a stronger (high-amplitude) anomaly; if they are opposed or significantly misaligned, they
273 partially cancel or reorient the net magnetization, which can weaken the anomaly or even yield
274 one of opposite polarity compared to what an induced-only model would predict. This interplay
275 complicates data interpretation, since assuming all magnetization is induced (parallel to today's
276 field) can lead to errors in locating or characterizing sources.

277
278 By removing these regional trends, we enhance the visibility of smaller-scale or high-frequency
279 anomalies, allowing subtle features and variations in the subsurface to be highlighted more
280 effectively. To remove regional anomalies, we apply an upward continuation of 1000 m to the
281 Bouguer gravity data (**Figure 3a**). This approach simulates measuring the gravity field at a higher
282 elevation—1000 m in our case—effectively smoothing out high-frequency anomalies associated
283 with shallow or near-surface geological features (Balogun et al., 2023). For RTP magnetic
284 anomaly, we use a Gaussian filter to calculate the regional trend (**Figure 3b**). This Gaussian filter
285 acts as a low-pass filter, smoothing out high-frequency components in the dataset. After extracting
286 the regional anomaly from the potential field, we subtract it from the unfiltered total anomaly,
287 yielding the residual anomaly (**Figure 3**).

288
289 We apply several filters to the residual potential field data to enhance specific features and improve
290 interpretability (**Figure 3**). Under the framework of Poisson's theorem in potential-field theory,
291 taking a vertical derivative of gravity and performing a reduction-to-the-pole on magnetic data are
292 mathematically analogous operations and are equivalent Fourier-domain operations: both sharpen
293 source-edge contrasts while suppressing deep, long-wavelength signals, thus yielding mutually
294 consistent structural imagery (Blakely, 1996). The filters we have used involve various forms of
295 derivative operations, which help to highlight changes in the data that correspond to geological
296 boundaries, faults, or other structural details (Ibraheem et al., 2023; Ma et al., 2016; Nasuti et al.,
297 2019; Núñez-Demarco et al., 2023). For the residual RTP magnetic data, we apply and show two
298 filters: the horizontal derivative and the tilt derivative. The horizontal derivative filter, applied in
299 x-direction (i.e., across the longitudes) and y-direction (i.e., across the latitudes), accentuates
300 lateral changes in the magnetic field, helping to reveal abrupt variations. The tilt derivative filter,
301 however, is especially effective as an edge detector. It operates by combining both vertical and
302 horizontal gradients, effectively highlighting the edges of anomalous bodies. The tilt derivative
303 produces values that tend toward zero over magnetically flat regions, positive over rising areas,
304 and negative over falling areas, creating a clear demarcation of the edges of magnetic sources. As
305 a result, this filter enhances the boundaries of anomalies and helps pinpoint the locations and
306 shapes of features with minimal distortion across varying depths (Ashraf and Filina, 2023b; Pham
307 and Oliveira, 2023). We apply the lineament mapping techniques to map major structural
308 boundaries from the filtered magnetic data (Ashraf and Filina, 2023a, 2023b; Ogah and Abubakar,
309 2024; Zhang et al., 2024). Our approach focused on identifying gaps between magnetic stripes,
310 changes in the stripe orientation, and a significant reduction in stripe width. We also calculate the
311 analytical signals of the magnetic anomalies to highlight the areas with high magnetizing
312 amplitude (Nabighian, 1972; Roest and Pilkington, 1993) that may illuminate magnetic mineral
313 deposits (Mohamed et al., 2022). To calculate the analytical signal of the residual magnetic data,
314 we first compute the horizontal and vertical derivatives of the magnetic field in the x, y, and z
315 directions. The analytical signal was then derived by taking the square root of the sum of the
316 squares of these derivatives, yielding a map that represents the amplitude of the magnetic field

317 independent of direction and highlights the edges of magnetic sources. To validate our
318 interpretations, we cross-reference the magnetic lineaments with gravity data. Before validating
319 with gravity data, we filter the residual gravity data by applying the first vertical derivative and tilt
320 derivative to map major tectonic structures and delineate their boundaries.
321

322 We also develop 2-D integrated models of the subsurface to examine the variations in the physical
323 properties of the rocks (density and magnetic susceptibility). We build our models using the GM-
324 SYS module within the Geosoft software suite, employing a 2-D approximation (Geosoft, 2021).
325 The GM-SYS model is extended to +/- 30000 km (i.e., infinity) along the X-axis and 90 km along
326 the Z-axis to eliminate edge effects. Due to the absence of reliable magnetic susceptibility logs for
327 the sedimentary cover and granitic basement, we treat these lithologies as a non-magnetic
328 background (0 SI). This approximation reduces the non-uniqueness of the potential field problem,
329 ensuring that the modeled response is driven primarily by the lateral susceptibility contrast of the
330 target magnetic unit. By integrating gravity and magnetic data within this 2D context, we can
331 delineate major geological boundaries and assess regional structural trends with sufficient
332 accuracy for our research objectives. Our goal is to develop a simple subsurface structure that
333 satisfactorily aligns with gravity, magnetic, and logging data without introducing excessive
334 complexity that might overfit the potential field anomaly. Instead, we aim to replicate the general
335 pattern of the observed anomaly in our 2D modeling, seeking to generate computed anomalies
336 with comparable amplitude, wavelength, and phase to those observed in the potential fields.
337



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340 Figure 3: Potential field data maps of northwestern Bangladesh used in this study to analyze geological and tectonic
341 features. (a) Gravity anomaly maps: from left to right, the Bouguer gravity anomaly, regional Bouguer gravity anomaly
342 with 1000 m upward continuation, and residual Bouguer gravity anomaly. (b) Magnetic anomaly maps: from left to
343 right, the RTP total magnetic anomaly, regional magnetic anomaly after Gaussian filtering, and residual magnetic
344 anomaly. The black dashed lines show the boundary of Rangpur Saddle.

345

346 4. Result

347

348 4.1 Integrated spatial analysis

349

350 In this study, we first establish the regional tectonic structures of northwestern Bangladesh through
351 spatial analysis of multiple geophysical datasets. We utilize gravity, magnetic, seismic image
352 interpretations, and drilling log data to characterize the tectonic setup of this region. Our analysis
353 of residual magnetic and gravity anomalies reveals that the Rangpur Saddle exhibits distinctly
354 different geophysical characteristics compared to its northern counterpart, the Himalayan Foredeep
355 region. These differences are discernible in the long-wavelength or broad-scale geophysical
356 signatures across the two regions.

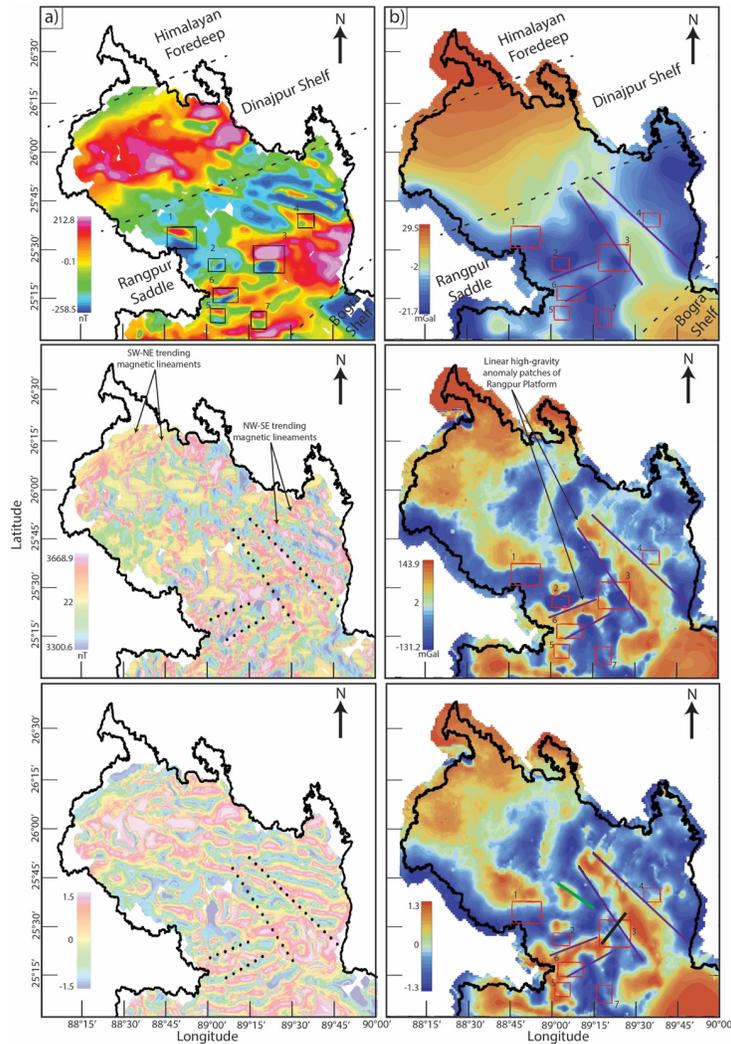
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358 In the Himalayan Foredeep, north of 26°N latitude, we observe a high residual Bouguer gravity
359 anomaly, ranging from approximately 5 to 30 mGal. In contrast, the Rangpur Saddle, located south
360 of this latitude, is marked by a low residual Bouguer gravity anomaly, typically between -22 and
361 0 mGal. Similarly, the Rangpur Saddle shows lower residual RTP total magnetic anomalies,
362 ranging from 0 to 260 nT compared to mostly high magnetic anomalies in north (~ 0 to 215 nT).
363 The spatial patterns of these anomalies also differ across the regions. South of 26°N latitude, in
364 the Rangpur Saddle, the horizontal and tilt derivative magnetic anomalies show a predominance
365 of NW-SE trending magnetic lineaments. North of this latitude, in the Himalayan Foredeep, the
366 magnetic anomaly trend shifts predominantly to a SW-NE orientation. Additionally, in the vertical
367 and tilt derivatives of the gravity data, we observe the most pronounced high gravity linear
368 anomaly patch trends NW-SE south of 26°N latitude, consistent with the magnetic anomaly trend
369 in this area. South of 26°30'N latitude, we see another linear high-gravity patch in the filtered
370 Bouguer gravity data trending NW-SE, indicating a different orientation than the other high-
371 gravity patch of the Rangpur Saddle region. These two high-gravity patches are also traceable in
372 the filtered magnetic anomalies, following the magnetic lineament mapping procedure. Notably,
373 some areas exhibit strong inverse gravity and magnetic field trends, where high gravity coincides
374 with low magnetic anomalies, or vice versa. One might expect a corresponding increase in the
375 gravity signal due to the higher density of the gabbroic intrusions relative to the overlying rocks;
376 however, this expected gravity anomaly is not observed in our data. We believe this discrepancy
377 arises from two interconnected factors. First, gravity and magnetic methods are most sensitive to
378 lateral variations in subsurface properties, and in our study area, the subsurface structures are
379 predominantly flat-lying. At depth, the gabbroic intrusions are laterally adjacent to felsic basement
380 rocks, and while the density contrast between these lithologies is modest (~0.1 g/cm³), their
381 magnetic susceptibilities differ significantly. This contrast produces a strong magnetic response
382 but only a subtle gravity anomaly that may be indistinguishable from background variations. Also,
383 the intrusions occur in fault-bounded graben fill that causes a localized low density (due to thick
384 low-density Gondwana sediments in the graben). Second, the spatial resolution of our gravity
385 survey, with a station spacing of approximately 3 km, is relatively coarse compared to the scale of

386 the gabbroic intrusions. As a result, any high-frequency gravity signals associated with these
387 smaller or more localized bodies are likely undersampled and thus not adequately captured in the
388 final gravity dataset. In brief, the high gravity areas correspond to uplifted blocks of dense
389 basement, whereas the magnetic highs occur on the flanks where intrusions have come up along
390 faults and the graben is filled with lighter sediment, yielding a relative gravity low.

391
392 Within the Rangpur Saddle region, we observe multiple oval dipolar patterns. From the RTP
393 residual magnetic anomaly data, we identify seven distinct patterns. The identification of these
394 patterns follows several consistent criteria. First, the dipolar magnetic anomalies are generally
395 oriented in a north-south direction. Second, each pattern features a magnetic high in the northern
396 section and a corresponding low to the south. Third, within each pattern, the size, area, and
397 amplitude of the magnetic high and low anomalies are comparable, contributing to a symmetrical
398 structure. The filtered magnetic data also reveals possible signatures of remanently magnetized
399 sources, some of which may be reversely magnetized. At the Earth's magnetic field inclination of
400 approximately 45° in northwestern Bangladesh (based on the IGRF 1980 epoch used for RTP
401 correction), reversely magnetized sources such as gabbroic intrusions formed during past
402 geomagnetic reversals produce negative anomalies after RTP where positive anomalies would
403 align with the current field. Mapping these magnetic patterns onto the filtered gravity maps reveals
404 that they consistently lie near the boundary between high and low gravity anomalies. While most
405 of these magnetic patterns (patterns 1, 4, 5, 6, and 7) intersect only one boundary, a few patterns
406 (specifically patterns 2 and 3) touch boundaries on both sides.

407



408
 409 Figure 4: a) Magnetic anomaly maps of northwestern Bangladesh. The top panel displays the residual RTP total
 410 magnetic anomaly, the middle panel shows the first horizontal derivative in the x-direction (across longitudes), and
 411 the bottom panel illustrates the tilt derivative. Black boxes highlight areas where oval dipolar magnetic patterns are
 412 observed. Dotted black lines in the middle and bottom panels indicate boundaries of high gravity regions. b) Bouguer
 413 gravity anomaly maps of northwestern Bangladesh. The top panel presents the residual Bouguer gravity anomaly, the
 414 middle panel shows the first vertical derivative, and the bottom panel displays the tilt derivative. High gravity regions
 415 are outlined by purple solid lines, and red boxes indicate areas with dipolar magnetic patterns. The solid green line in
 416 the bottom panel represents seismic profile 2 from Figure 2. The thick solid black line shows the 2-D integrated
 417 modelling profile of Figure 5.

418 **4.2 Integrated subsurface modeling**

419
420 We develop a 2-D subsurface model utilizing gravity, magnetic, and drilling log data (**Figure 5**).
421 The goal of this model is to approximate the observed anomalies on a local scale while aligning
422 with the available drilling log information (**Figure 2c**). The magnetic anomaly modeled in **Figure**
423 **5** is the residual RTP anomaly, reflecting local high-frequency features after removing regional
424 trends. Importantly, we do not aim to achieve a precise fit between the observed data and calculated
425 anomalies. This decision is driven by the fact that the observed potential field data are influenced
426 by regional structures, and without sufficient seismic information, constructing a reliable regional
427 model is not feasible. Instead, we focus on developing a simplified version of the subsurface model
428 that reasonably fits the observed anomalies and drilling data.

429
430 To ground the 2-D model in reality, we incorporated available drilling log information (**Figure 2**).
431 Based on these logs, the model includes five sedimentary layers above a crystalline basement. The
432 shallowest layer is assigned as alluvium, extending to a depth of 50 meters or less. Beneath the
433 alluvium, we sequentially include the Dupitila, Surma, and Cherra sandstone layers. The variable
434 thickness of these sedimentary layers comes from the drilling log information, except for the
435 Gondwana layer. The thickness of Gondwana is approximated based on the gravity fit. Below
436 these layers, we model a mafic intrusion of gabbro, which occurs within fault structures that have
437 developed within the underlying felsic basement.

438
439 The resulting model indicates that a horst and graben structure is the most plausible configuration
440 when considering the regional geological context. This structural interpretation also provides a
441 reasonable match to the observed geophysical anomalies. Additionally, based on information from
442 drilling log GDH 47, we assign a layer of Gondwana sediments within the graben basin, which is
443 contributing to the low observed gravity in this region.

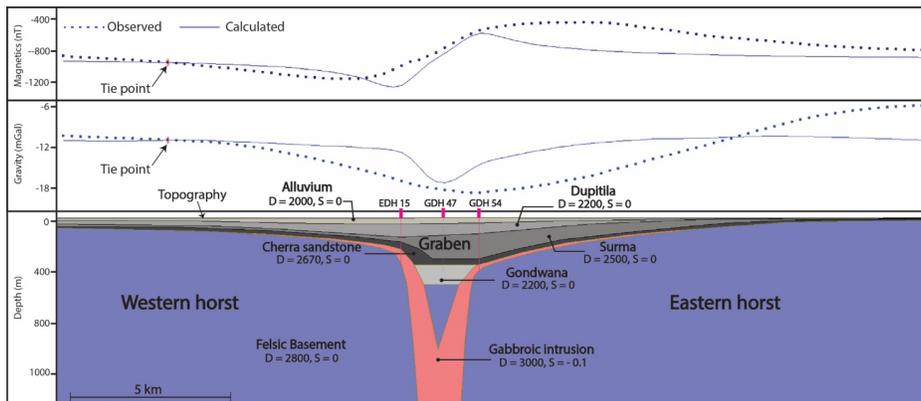
444
445 Densities in the model are derived from drilling data, reflecting the unique lithological
446 characteristics of each unit. To get an initial guess on the density, we developed a gravity model
447 for profile 1 that has seismic-derived subsurface information (**Figure S2**). The alluvium, Dupitila,
448 Surma, and Cherra Sandstone layers are modeled with densities of 2000, 2200, 2500, and 2670
449 kg/m^3 , capturing their progressive compaction and mineral composition. The Gondwana unit,
450 enriched with coal deposits, exhibits a notably lower density of 2200 kg/m^3 , consistent with its
451 organic-rich composition. Beneath these layers, the felsic basement and gabbroic intrusion stand
452 out with densities of 2800 and 3000 kg/m^3 , highlighting their denser crystalline structure and mafic
453 origins.

454
455 Between 400 and 800 meters depth, the gabbroic intrusion begins to follow the fault plane. The
456 thickness of both the Gondwana layer and the underlying felsic basement is determined from the
457 amplitude of the calculated gravity and magnetic anomalies. The thickness of the Gondwana layer
458 influences the minimum value of the calculated gravity anomaly, while the amplitude of the
459 magnetic anomaly helps determine the depth of the gabbroic intrusion. Notably, our calculated
460 anomalies are of high frequency. This occurs because our modeling does not incorporate all
461 regional structures; instead, it focuses only on local structures. As a result, the calculated anomalies
462 reflect primarily local or high-frequency potential field anomalies.

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467 Our 2D subsurface model reveals two horst structures with a graben basin in the middle. We refer
 468 to these as the eastern and western horsts based on their geographic locations. In the model, the
 469 western horst structure is depicted as deeper, consistent with the observed gravity anomaly, which
 470 shows lower gravity in the western horst compared to the eastern horst. Our model does not include
 471 the broader horst structure on the eastern side, which likely explains the discrepancy between the
 472 calculated and observed anomalies over the eastern horst.
 473



474
 475
 476 Figure 5: Integrated 2D geophysical model over the highest magnetic anomaly in northwestern Bangladesh. See Figure
 477 1b and 4c for the location of the model profile. The top panel presents the total magnetic intensity anomaly, and the
 478 middle panel shows the gravity anomaly, with observed (dotted) and calculated (solid) data. Tie points, marked by red
 479 stars, indicate where the calculated anomalies are vertically shifted to align with the observed data. The bottom panel
 480 illustrates the subsurface model, with geological units represented by distinct colors. 'D' denotes density (kg/m^3), and
 481 'S' denotes magnetic susceptibility (SI units). Bold pink vertical dashes mark the locations of drilling logs used to
 482 constrain subsurface units. Thin vertical pink lines extend downward from these surface locations. Solid pink lines
 483 beneath EDH 15 and GDH 54 indicate where the forward model is directly tied to lithological boundaries from well
 484 logs. In contrast, the dashed pink line beneath GDH 47 reflects an interpretative connection, as this well does not lie
 485 directly along the modeling profile.
 486

487 We acknowledge the mismatch between the observed and calculated anomalies in our forward
 488 model. We aim to adopt the simplest model that captures the essential features of the observed
 489 anomaly without introducing unnecessary complexity. While even basic 2-D models can yield
 490 multiple valid solutions, incorporating intricate structures without strong geological constraints
 491 would increase interpretive ambiguity. Additionally, our model is inherently two-dimensional and
 492 thus cannot fully represent the three-dimensional geological variations present in the study area,
 493 such as the broad, shallow eastern horst, whose lateral extent likely contributes to higher observed
 494 gravity values. Finally, our modeling profile spans a smaller area than the full potential field survey
 495 (refer to section 3.1), resulting in calculated anomalies that emphasize localized, high-frequency
 496 features, whereas the observed data reflect broader, lower-frequency trends. Our forward model is
 497 calculated on 250 m spacing, whereas the gravity data were acquired every 1–1.5 km and the
 498 magnetic data every 5 km.
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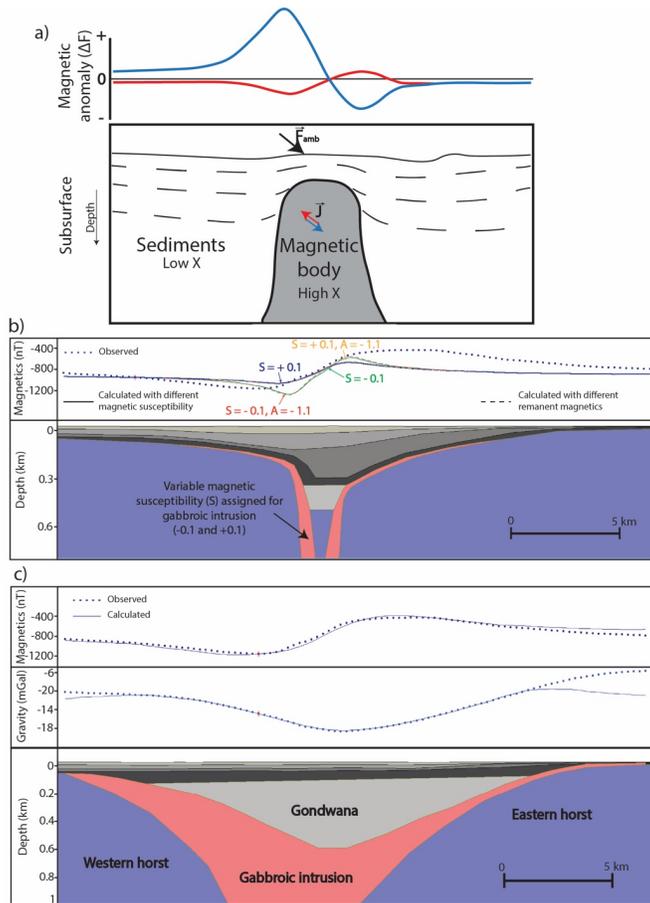
501 Furthermore, we also test alternative models, which further support that the presented version
502 offers the most concise and geologically reasonable solution. First, we test how the polarization
503 direction of the remanent magnetization affects the total magnetic intensity reading of a dome-
504 shaped magnetic body (**Figure 6a**). We incorporate both remanent magnetization (A) and
505 magnetic susceptibility (S) in the forward model, which affects the total magnetic intensity and its
506 polarity (**Figure 6b**). When we assign only a positive susceptibility contrast ($S = +0.1$ SI) without
507 remanence, the calculated anomaly captures the general shape of the observed data more
508 accurately in terms of slope, but the amplitude (highs and lows) is notably lower than the observed
509 values. In contrast, using a reverse-polarity remanent magnetization yields calculated amplitudes
510 that closely match the observed highs and lows; however, the overall shape shows a poorer fit,
511 with significant mismatches in the slopes. Because of the better match for the amplitudes, we prefer
512 the reverse polarity for the magnetic body in our forward model. We also implemented an
513 additional component of remanent magnetization (A) to the magnetic body, fixed at -1.1 A/m, and
514 tested combinations such as $S = -0.1$ SI with $A = -0.1$ A/m and $S = +0.1$ SI with $A = -1.1$ A/m.
515 These combinations yielded a similar response, primarily increasing the amplitude of the anomaly
516 without changing the overall trend. In conclusion, while the additional remanent magnetization
517 component amplifies the anomaly magnitude, it does not significantly affect the trend or alter the
518 interpretation of the model. We have also developed a model that shows a perfect fit between the
519 observed and calculated anomaly (**Figure 6c**). However, to achieve that, we need to completely
520 abandon the constraints from well-logs and regional structural context based on gravity and
521 magnetic data. For these reasons, we adopt a simplified, geologically grounded model that
522 balances interpretive clarity with structural realism.

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Figure 6: a) Conceptual illustration showing polarization direction of the remanent magnetization affecting the total magnetic intensity reading of a dome-shaped magnetic body. The synthetic total-field anomaly (ΔF) profiles are calculated for a concealed high-susceptibility intrusion beneath low-susceptibility sediments. The blue curve represents the response when magnetization is purely induced parallel to the present-day ambient field F_{amb} , producing a simple positive-negative (dipolar) signature. The red curve shows how the anomaly shape and amplitude are modified when an oblique remanent magnetization component J is added to the induced vector. b) Investigation on how incorporating both remanent magnetization (A) and magnetic susceptibility (S) in the forward model affects the total magnetic intensity and its polarity. c) 2-D forward modelling along the same profile as Figure 5, where we have achieved a perfect fit between the calculated and observed anomalies except for the model edges.

542 **5. Discussion**

543

544 **5.1 Basement structure**

545

546 One of the primary goals of this paper is to understand the basement structure and tectonic setup
547 of the Rangpur Saddle. Our findings reveal significant differences in the basement structures of
548 the Rangpur Saddle compared to its northern counterparts, evident in both gravity and magnetic
549 data. A key distinction is the depth to basement between these regions (**Figure 1b**). Previous
550 studies consistently suggest that the basement depth in the Rangpur Saddle is shallower than in the
551 Dinajpur Shelf. However, those studies were based on limited seismic and drilling data. Our
552 results, however, indicate a more complex basement geometry than previously understood.
553 Filtered gravity and seismic data presented here suggest that the Rangpur Saddle region contains
554 both shallow and deep basement features (**Figure 4b**). Considering regional tectonics, we propose
555 that high gravity values correspond to shallow horst structures, while low gravity values indicate
556 deeper graben structures. However, the frequency of horsts and grabens is significantly lower in
557 the Rangpur Saddle compared to the northern Himalayan Foredeep and the Dinajpur Shelf (**Figure**
558 **2a** and **2b**). In the Rangpur region, we identified only two horsts of notable width and length. The
559 boundaries of these horst structures also appear as lineaments in the filtered magnetic data (**Figure**
560 **4a**), suggesting the presence of major structural boundaries.

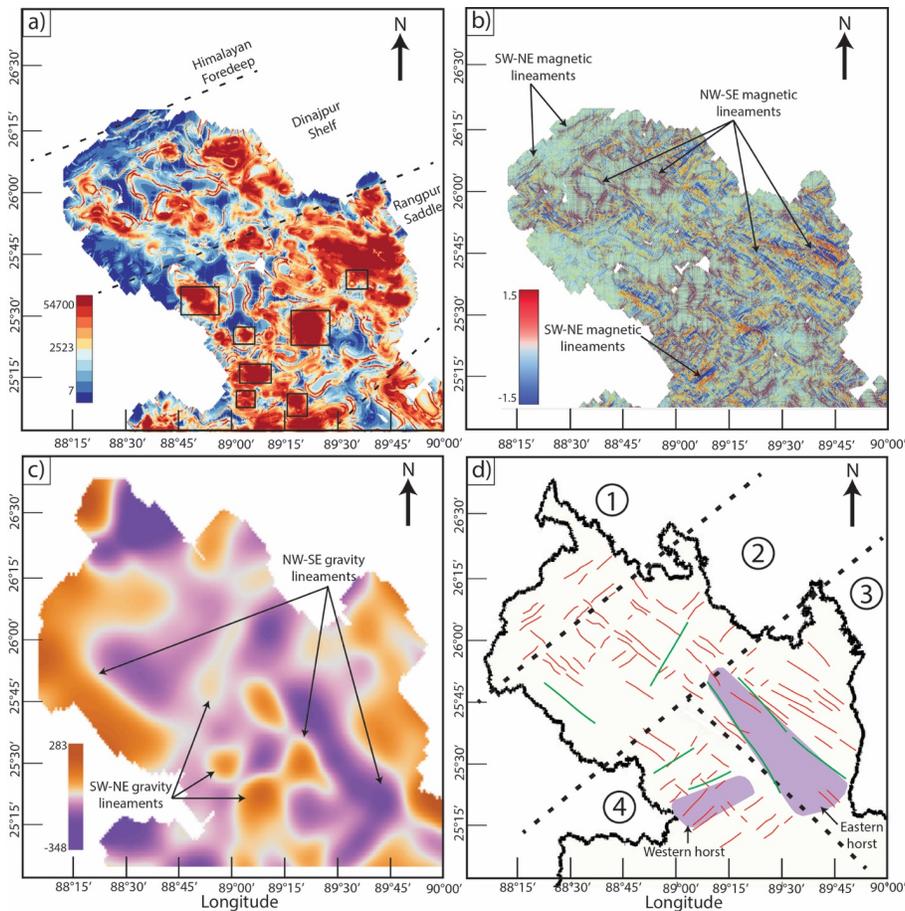
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562 Furthermore, differences in the directions of magnetic lineaments between the Rangpur Saddle
563 region and its northern counterparts—the Dinajpur Shelf and the Himalayan Foredeep—suggest
564 variations in paleo-tectonic stress orientation (**Figure 4a, 7a** and **7b**). Existing literature
565 characterizes northwestern Bangladesh predominantly with extensional tectonics (Gani and Alam,
566 2003). Near the Himalayan Foredeep boundary, SW-NE trending magnetic lineaments indicate a
567 paleo-principal stress direction oriented NW-SE. Moving southward into the Dinajpur Shelf, NW-
568 SE trending magnetic lineaments become more common (**Figure 7b**), pointing to a shift in the
569 principal stress direction associated with extensional tectonics. In the Rangpur Saddle, these NW-
570 SE trending lineaments are even more frequent, especially in the eastern part, indicating a greater
571 intensity of paleo-tectonic stress in this region compared to the Dinajpur Shelf. Based on traced
572 magnetic and gravity lineaments, the stable platform can be divided into four distinct zones
573 (**Figure 7d**). In Zone 1, located in the northern part, lineaments predominantly trend SW-NE, with
574 only a few exceptions. Moving southward into Zone 2, there is a mixture of two differently
575 trending lineaments: SW-NE and NW-SE. In Zone 3, in the easternmost section of the stable
576 platform, lineaments primarily trend NW-SE, aligning with the eastern horst. Finally, in Zone 4,
577 the southernmost region, lineaments generally follow an SW-NE trend, corresponding with the
578 western horst. This subdivision highlights the complex nature of past tectonic processes in this
579 region.

580

581 We also identify seven oval dipolar patterns of magnetic highs and lows. When mapped onto
582 gravity data, these patterns consistently align with the boundaries between high and low gravity
583 patches, indicating the possible boundaries between horsts and grabens. Drilling log data near the
584 most prominent magnetic anomaly reveals the presence of gabbro, which may explain these
585 dipolar magnetic patterns. Furthermore, these regions of dipolar magnetic patterns also spatially
586 correlate with overall high magnetization (**Figure 7a**). We propose that these gabbro formations

587 resulted from intrusions along normal faults, formed by extensional tectonics, which define the
 588 boundaries between horsts and grabens.
 589



590
 591 Figure 7: Filtered potential field maps of the northwestern stable platform region of Bangladesh, highlighting major
 592 tectonic structures and boundaries. (a) Analytical signal map derived from the magnetic anomaly (methodology
 593 detailed in Sect. 3), showing magnetization amplitudes across the region. Black boxes indicate areas of oval dipolar
 594 magnetic patterns, as shown in Figure 4a. Dashed lines represent boundaries between established tectonic
 595 subdivisions. (b) Total horizontal gradient map of the magnetic anomaly, calculated by applying a first horizontal
 596 derivative filter in the X-direction, followed by a first horizontal derivative filter in the Y-direction. (c) Total horizontal
 597 derivative (similar method as in 'b') of the residual gravity anomaly. This map is developed after applying upward
 598 continuation to the residual gravity data to a level where high-resolution artifacts are no longer present in the derivative
 599 map. (d) Map displaying traced magnetic lineaments (in red) and gravity lineaments (in green) from 'c'.
 600 Black dashed lines indicate regional subdivisions based on mapped tectonic lineaments. Purple polygons mark horst
 601 structures in the Rangpur Saddle, interpreted from filtered gravity data (refer to Figure 4b).

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5.2 Economic mineral potential

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Our integrated spatial analysis and 2-D modeling reveal that the Rangpur Saddle contains at least two prominent horst and graben structures. The 2-D model results (**Figure 5**), supported by drilling log data, confirm the presence of gabbroic bodies along normal faults within these structures. This suggests that the most prominent oval-shaped dipolar magnetic signature has strong potential to host iron-bearing rocks such as gabbro. When the other similar patterns are mapped onto the gravity data, they consistently correspond to the boundaries between high and low gravity zones, marking the transitions between horsts and grabens. We interpret these gabbroic intrusions as products of magmatic emplacement along extensional faults (**Figure 8**), consistent with processes associated with the early stages of continental rifting. Furthermore, analysis of lineament orientations (**Figure 7d**) indicates that, except in the western horst region, most of the interpreted magmatic emplacement appears to be controlled by NW–SE oriented lineaments.

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The breakup process involved rifting events that progressively separated the Indian subcontinent from the other Gondwanaland constituents (Veevers, 2004). The breakup of Gondwanaland involved multiple rifting episodes and magmatic activities. While the Central Atlantic Magmatic Province (CAMP) exemplifies rift-related magmatism during the Triassic-Jurassic breakup of Pangea, the Rangpur Saddle's tectonic evolution is more directly tied to the Cretaceous rifting of the Indian plate from Antarctica-Australia (Curry, 1991; Hossain et al., 2019). This rifting facilitated extensional faulting and mafic intrusions, analogous to processes observed in CAMP but occurring later in the Mesozoic. In such tectonic settings, normal faulting and rift-related subsidence create pathways for mafic magmas to ascend (Brune et al., 2023; Pirajno and Santosh, 2014; Ruppel, 1995), leading to the emplacement of gabbroic and other mafic intrusions within the crust (Magee et al., 2019).

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Mafic magma intrusions are commonly observed along extensional fault planes in tectonically active regions worldwide. Troll et al. (2021) shows that in NW Scotland, Long Loch Fault acted as dynamic magma conduit with its movements facilitating the ascent of ultrabasic magmas. These magmas intruded along faults and fractures, leading to the destabilization and collapse of existing cumulate layers, forming extensive breccias. A study on Mesozoic gabbroic intrusions in the High Atlas Mountains highlights their emplacement along extensional faults during the rifting associated with the Central Atlantic Magmatic Province, where more than 50% of gabbro samples exhibit stable magnetization from magnetite which is identified as the primary component (Calvin et al., 2017). Fuller and Waters (1929) have extensively studied horsts and graben structures in southern Oregon where they found emplacement of volcanic and intrusive rocks are closely associated with extensional faulting. These normal faults and the resulting grabens provide pathways for magma to ascend and intrude, which is evident from the prevalence of rhyolitic, dacitic, and andesitic vents as well as basaltic dikes aligned with the faults.

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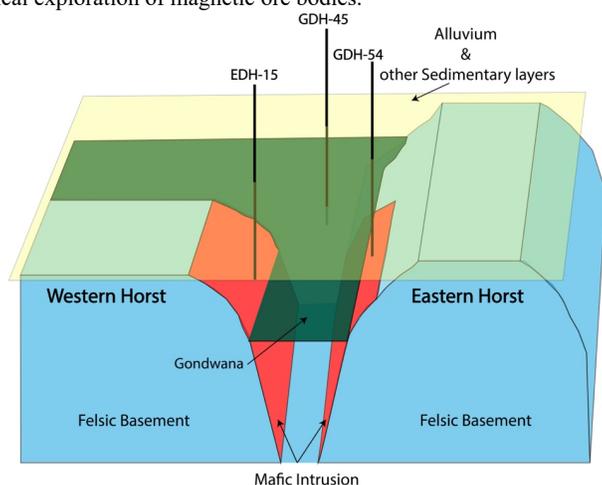
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Mafic intrusions along fault zones are often associated with magnetic mineral deposits, which can be economically significant as ore bodies (Zhou et al., 2005). For instance, magnetite deposits are found within mafic intrusions along the fault planes in Egypt, which exhibit high magnetic anomalies, facilitating their detection and delineation (Kharbish et al., 2022; Mousa et al., 2020).

647 Also, in Central Iran, iron-bearing magnetic mineral deposits are observed to exhibit high magnetic
648 and gravity anomalies (Kheyrollahi et al., 2021).

649
650 Based on current geochemistry data, the Rangpur Saddle holds significant potential for magnetic
651 mineral ore deposits, driven by its mafic-ultramafic sequences and iron-rich dykes (Ameen et al.,
652 2021). High Fe₂O₃ content in hornblendites and associated minerals like magnetite and ilmenite
653 indicate strong magnetization. The tectonic setting, with extensive magmatic intrusions, provides
654 ideal conditions for mineralization. These findings highlight the region as a possible target for
655 future geophysical exploration of magnetic ore bodies.



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657
658 Figure 8: 3D schematic diagram illustrating the tectonic structures in Pirganj, where the highest magnetic anomaly is
659 observed in Bangladesh. The felsic basement, depicted in blue blocks, includes two horsts, as shown in Figure 5d.
660 Possible mafic intrusions along the faults are indicated in red. Sedimentary layers are represented by a single horizontal
661 layer in specific colors, with the graben basin likely filled with Gondwana sediments shown in green, and other
662 sedimentary layers, including alluvium, shown in yellow, which also represents the topography of the region. The
663 approximate locations of the three drilling logs used in this study (refer to Figure 1b for their locations) are also
664 indicated.

665 5.3 Limitations and Future Research

666
667 Compared to other regions globally, limited published research exists for our study area, restricting
668 our ability to draw on established models and findings. Situated at the eastern edge of the Indian
669 Shield, this area transitions into a different tectonic subdivision, creating a tectonic complexity
670 that cannot be fully resolved with the available low-resolution data. The scarcity of high-resolution
671 datasets, such as regional-scale seismic surveys, further limits our capacity to delineate regional
672 structures accurately. Currently, no regional or localized high-resolution datasets from ground-
673 based surveys are widely available, though such data could enhance future studies. Additionally,
674 the available well log data are predominantly shallow, reaching depths of approximately 500
675 meters, which limits insights into deeper geological and tectonic features beyond 2 kilometers.
676
677 Petrographic descriptions of the basement rocks are also rudimentary, as they are primarily based

678 on a limited number of well logs, most of which were drilled before the 1990s, leaving significant
679 gaps in our understanding of the area's deeper subsurface geology.

680
681 Future research in this region will focus on constructing detailed three-dimensional (3D) models
682 of the subsurface to enhance our understanding of the geological framework. This will involve
683 compiling a comprehensive dataset repository that integrates historical seismic reflection data,
684 including those acquired by private entities. Given the shallow nature of the basement, which lies
685 between approximately 128 and 1000 meters depth, high-resolution gravity and magnetic data will
686 be critical for resolving fine-scale structural features. To achieve the necessary spatial resolution,
687 precision ground-based gravity and magnetic surveys are recommended. Specifically, we propose
688 a targeted geophysical survey in areas exhibiting oval-shaped dipolar magnetic anomalies, using
689 station spacing of 500 meters or less. Such high-density measurements will improve our ability to
690 detect and delineate narrow mafic intrusions at shallow depths. In addition, developing a regional
691 tomographic model using either local earthquake data or ambient seismic noise will provide
692 complementary constraints on subsurface structures and further refine interpretations of the
693 geological setting.

694
695 The study area is characterized by complex tectonic structures shaped by a diverse tectonic history.
696 While the current study has focused on the Rangpur Saddle, future research will expand to cover
697 the entire northwestern part of Bangladesh, including the Himalayan Foredeep and the Eocene
698 Hinge region, which represents a Palecontinental shelf. Such expanded coverage will provide
699 deeper insights into the tectonic evolution and geological complexity of the area.

700

701 **Conclusion**

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703 This study provides a geophysical investigation of the northwestern region of Bangladesh,
704 revealing its significant potential for mineral resource exploration. The integration of gravity,
705 magnetic, seismic, and drilling data has allowed us to identify key tectonic features, such as
706 gabbroic intrusions along extensional faults, some which may host of valuable magnetic mineral
707 deposits. Despite the limitations in data resolution and coverage, our findings offer important
708 insights into the region's geological evolution and resource potential. Future work should prioritize
709 acquiring more detailed geophysical data and expanding drilling campaigns to refine the
710 understanding of subsurface structures and evaluate the feasibility of mineral extraction. The
711 Rangpur Saddle, as the shallowest part of the stable platform, emerges as a promising target for
712 future mineral exploration endeavors.

713

714 **Data availability**

715

716 All raw data can be provided by the corresponding authors upon request.

717

718 **Author contribution**

719

720 Mohammad Tawhidur Rahman Tushar was involved in conceptualization, methodology, software,
721 validation, formal analysis, investigation, resources, data curation, writing – original draft, writing
722 – review and editing, visualization. Asif Ashraf contributed to conceptualization, methodology,

723 software, validation, formal analysis, investigation, resources, data curation, writing – original
724 draft, writing – review and editing, and visualization. Md. Mahfuz Alam worked on software
725 development, validation, and formal analysis. Md Nasif Jamil contributed to writing – review and
726 editing and visualization. Saba Karim was involved in writing – review and editing and
727 visualization. Md. Shahjahan contributed to conceptualization, investigation, resources, project
728 administration, and supervision. Md. Anwar Hossain Bhuiyan was involved in writing – review
729 and editing, project administration, supervision, and funding acquisition.

730

731 **Competing interests**

732

733 The authors declare that they have no conflict of interest.

734

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736

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749

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