

# 1 **Reviews and Syntheses: Variable Inundation**

## 2 **Across Earth's Terrestrial Ecosystems**

3 James Stegen<sup>1,2</sup>, Amy J Burgin<sup>3</sup>, Michelle H. Busch<sup>3</sup>, Joshua B. Fisher<sup>4</sup>, Joshua Ladau<sup>5</sup>, Jenna  
4 Abrahamson<sup>6</sup>, Lauren Kinsman-Costello<sup>7</sup>, Li Li<sup>8</sup>, Xingyuan Chen<sup>1</sup>, Thibault Datry<sup>9</sup>, Nate  
5 McDowell<sup>1</sup>, Corianne Tatariw<sup>10</sup>, Anna Braswell<sup>11</sup>, Jillian M. Deines<sup>1</sup>, Julia A. Guimond<sup>12</sup>, Peter  
6 Regier<sup>1</sup>, Kenton Rod<sup>1</sup>, Edward K. P. Bam<sup>13</sup>, Etienne Fluet-Chouinard<sup>1</sup>, Inke Forbrich<sup>14</sup>, Kristin L.  
7 Jaeger<sup>15</sup>, Teri O'Meara<sup>16</sup>, Tim Scheibe<sup>1</sup>, Erin Seybold<sup>3</sup>, Jon N. Sweetman<sup>8</sup>, Jianqiu Zheng<sup>1</sup>,  
8 Daniel C Allen<sup>8</sup>, Elizabeth Herndon<sup>16</sup>, Beth A. Middleton<sup>17</sup>, Scott Painter<sup>16</sup>, Kevin Roche<sup>18</sup>,  
9 Julianne Scamardo<sup>19</sup>, Ross Vander Vorste<sup>20</sup>, Kristin Boye<sup>21</sup>, Ellen Wohl<sup>22</sup>, Margaret Zimmer<sup>23</sup>,  
10 Kelly Hondula<sup>24</sup>, Maggi Laan<sup>1</sup>, Anna Marshall<sup>22</sup>, and Kaizad F. Patel<sup>1</sup>

11  
12 <sup>1</sup>*Pacific Northwest National Laboratory, Richland, WA, USA*

13 <sup>2</sup>*School of the Environment, Washington State University, Pullman, WA, USA*

14 <sup>3</sup>*University of Kansas, Lawrence, KS, USA*

15 <sup>4</sup>*Chapman University, Orange, CA, USA*

16 <sup>5</sup>*University of California San Francisco, San Francisco, CA, USA*

17 <sup>6</sup>*North Carolina State University, Raleigh, NC, USA*

18 <sup>7</sup>*Kent State University, Kent, OH, USA*

19 <sup>8</sup>*Penn State University, State College, PA, USA*

20 <sup>9</sup>*National Institute for Agriculture, Food, and Environment (INRAE), Villeurbanne, France*

21 <sup>10</sup>*Rowan University, Glassboro, NJ, USA*

22 <sup>11</sup>*University of Florida, Gainesville, FL, USA*

23 <sup>12</sup>*Woods Hole Oceanographic Institution, Woods Hole, MA, USA*

24 <sup>13</sup>*International Water Research Institute (IWRI), Mohamed VI Polytechnic University, Benguerir, Morocco*

25 <sup>14</sup>*University of Toledo, Woods Hole, MA, USA*

26 <sup>15</sup>*U.S. Geological Survey, Washington Water Science Center, Tacoma, WA, USA*

27 <sup>16</sup>*Oak Ridge National Laboratory, Oak Ridge, TN, USA*

28 <sup>17</sup>*U.S. Geological Survey, Wetland and Aquatic Research Center, Lafayette, LA, USA*

29 <sup>18</sup>*Boise State University, Boise, ID, USA*

30 <sup>19</sup>*University of Vermont, Burlington, VT, USA*

31 <sup>20</sup>*University of Wisconsin, La Crosse, WI, USA*

32 <sup>21</sup>*SLAC National Acceleratory Laboratory, Menlo Park, CA, USA*

33 <sup>22</sup>*Colorado State University, Fort Collins, CO, USA*

34 <sup>23</sup>*U.S. Geological Survey Upper Midwest Water Science Center, Madison, WI, USA*

35 <sup>24</sup>*Arizona State University, Tempe, AZ, USA*

36  
37 Journal: EGU Biogeosciences

38 Type of Manuscript: Reviews & Synthesis

39  
40 Correspondence: James C. Stegen, E-mail: [james.stegen@pnnl.gov](mailto:james.stegen@pnnl.gov); Phone: (509) 371-6763

41

## 42 **Abstract**

43 The structure, function, and dynamics of Earth's terrestrial ecosystems are profoundly  
44 influenced by the frequency and duration that they are inundated with water. A diverse array of  
45 natural and human-engineered systems experience temporally variable inundation whereby  
46 they fluctuate between inundated and non-inundated states. Variable inundation spans from  
47 extreme events to predictable sub-daily cycles. Variably inundated ecosystems (VIEs) include  
48 hillslopes, non-perennial streams, wetlands, floodplains, temporary ponds, tidal systems, storm-  
49 impacted coastal zones, and human-engineered systems. VIEs are diverse in terms of  
50 inundation regimes, water chemistry and flow velocity, soil and sediment properties, vegetation,  
51 and many other properties. The spatial and temporal scales of variable inundation are vast,  
52 ranging from sub-meter to whole landscapes and from sub-hourly to multi-decadal. The broad  
53 range of system types and scales makes it challenging to predict the hydrology,  
54 biogeochemistry, ecology, and physical evolution of VIEs. Despite all experiencing the loss and  
55 gain of an overlying water column, VIEs are rarely considered together in conceptual,  
56 theoretical, modeling, or measurement frameworks/approaches. Studying VIEs together has the  
57 potential to generate mechanistic understanding that is transferable across a much broader  
58 range of environmental conditions, relative to knowledge generated by studying any one VIE  
59 type. We postulate that enhanced transferability will be important for predicting changes in VIE  
60 function in response to global change. Here we aim to catalyze cross-VIE science that studies  
61 drivers and impacts of variable inundation across Earth's VIEs. To this end, we complement  
62 expert mini-reviews of eight major VIE systems with overviews of VIE-relevant methods and  
63 challenges associated with scale. We conclude with perspectives on how cross-VIE science can  
64 derive transferable understanding via unifying conceptual models in which the impacts of  
65 variable inundation are studied across multi-dimensional environmental space.

## 66 **Introduction**

67 The chemical and biological processes within terrestrial ecosystems hinge on the presence,  
68 residence time, volume, and chemistry of water (Schimel et al. 1991, Lohse et al. 2009, Arias-  
69 Real et al. 2024). A variety of factors influence water retention, infiltration, flow, and surface  
70 expression within an ecosystem, such as land surface relief, topographic slope, subsurface  
71 permeability, evapotranspiration, and human-based modifications of the landscape (Horton  
72 1940, Ribolzi et al. 2011, Appels et al. 2016, McGrane 2016, Orozco-López et al. 2018,  
73 McDowell et al. 2023). Water supply is most commonly 'top down' in the form of precipitation  
74 and overland flow or 'bottom up' due to rising water tables and transient saturation in the  
75 subsurface (Freeze 1974, Smith et al. 2017, Stewart et al. 2019). Inundation, however, may also  
76 occur from lateral inputs, as is common in tidal systems, or from upslope inputs, as in  
77 floodplains. Regardless of where water comes from, its expression at the land-atmosphere  
78 interface occurs when the rate of water supply is greater than the rate of export via infiltration,  
79 evapotranspiration, and runoff (Tromp-van Meerveld and McDonnell 2006).

80 Here, we define inundation as occurring when there is an uninterrupted aqueous barrier that  
81 limits diffusive gas exchange at the land-atmosphere interface (Elberling et al. 2011, Smith et al.  
82 2018). This conceptualization includes diverse hydrological conditions ranging from free  
83 standing water to soil surface saturation. Hence, our broad definition spans from extreme events

84 such as hurricane-driven inundation to shallow intermittent overland runoff across hillslopes.  
85 This definition does not attempt to separate ‘inundation’ from ‘flooding’ based on temporal  
86 frequency/duration, as has been proposed elsewhere (Flick et al. 2012). To avoid confusion  
87 from interchangeable use of these two words (as in USACE 2024), we exclusively use  
88 ‘inundation’ and avoid references to ‘flooding’ in this paper. We define variably inundated  
89 ecosystems (VIEs) as areas of any spatial and temporal scale that experience transitions  
90 between the presence and absence of inundated conditions. Variable inundation is natural in  
91 many systems and can be critical to system function (Shaeri Karimi et al. 2022, Tsoi et al.  
92 2022), while in other systems it represents a disturbance (Sun et al. 2022a, Hopple et al. 2023).  
93 Variably inundated ecosystems cover at least 5-9 million km<sup>2</sup>, or 4-7% of the Earth’s land  
94 surface excluding Greenland and Antarctica. These estimates are according to monthly data  
95 over multiple decades (Zhang et al. 2017, 2021, Davidson et al. 2018). Current areal estimates  
96 of VIEs may, however, be underestimates as many VIEs are not detectable with current remote  
97 sensing techniques.

98 Variable inundation occurs across a wide range of terrestrial ecosystems, but the factors  
99 governing its influences are typically studied independently without cross-ecosystem  
100 comparisons. Some examples of VIEs are hillslopes with overland flow, non-perennial streams,  
101 floodplains and parafluvial zones, variably inundated wetlands, vernal ponds/pools/playas, tidal  
102 systems, coastal systems impacted by storm-driven inundation, and human-engineered  
103 systems intended to shift inundation dynamics (e.g., flood-irrigated agriculture, stormwater  
104 infrastructure, and constructed wetlands) (**Fig. 1**). A given system may not fit clearly into a  
105 single VIE category and/or may transition across categories through time. For example, when  
106 flow ceases and isolated pools form in a non-perennial stream network, the stream begins to  
107 behave more like a wetland or vernal pond as opposed to a flowing stream (Day et al. 2019).  
108 Further, while VIEs may be classified as wetlands under the broadest definition from the  
109 Ramsar Convention (Secretariat 2016), there is significant variation in how wetlands are defined  
110 (Finlayson and Van Der Valk 1995) and we do not attempt to rectify or clarify variation in those  
111 definitions. Here, when using the term ‘wetland’ we simply align with the perspective that  
112 wetlands are similar to marshes, swamps, and bogs.

113 Inundation dynamics are changing due to increased variability and magnitudes of  
114 precipitation and evapotranspiration, accelerated sea level rise, and human modifications to the  
115 Earth’s land surface, including an increase in extreme events (Konapala et al. 2020, Li et al.  
116 2022a). For example, extreme events such as coastal inundation are increasingly frequent  
117 (Vitousek et al. 2017). However, inundation patterns are changing in different ways across  
118 different VIEs (Zipper et al. 2021, Londe et al. 2022). For example, in river systems seasonal  
119 drying is becoming more common in multiple biomes (Sweet et al. 2014, Zipper et al. 2021).  
120 While some rivers are shifting from non-perennial to perennial (Döll and Schmied 2012, Datry et  
121 al. 2018a) and others have fewer no-flow days than they did historically (Zipper et al. 2021).  
122 Divergence in the direction of change, with some systems inundating less and others inundating  
123 more, is likely linked to diverse drivers of change associated with changing climates and/or  
124 direct human impacts (Datry et al. 2023). Therefore, researchers and decision makers cannot  
125 rely exclusively on historical trends to predict future impacts (e.g., on species diversity) of  
126 changing inundation dynamics (Culley et al. 2016, Quinn et al. 2018, Rameshwaran et al. 2021,  
127 Li et al. 2022b).



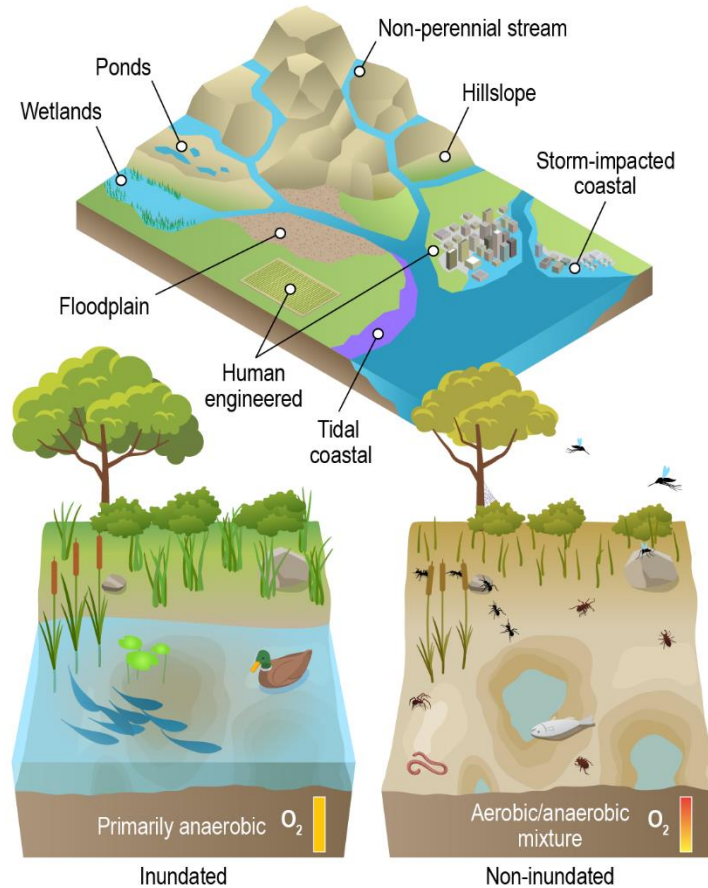
128  
 129  
 130  
 131  
 132  
 133  
 134  
 135  
 136  
 137  
 138  
 139  
 140  
 141

**Figure 1. Variably inundated ecosystems (VIEs) span numerous ecosystem types and are globally distributed across the Earth's land surface.** There are few places across Earth's land surfaces that do not experience variable inundation, which is defined here as the loss/gain of an aqueous barrier between the atmosphere and porous media (e.g., soil) that inhibits gas phase transport. Due to global changes in the dynamics of variable inundation, there is a need to integrate knowledge into models that are predictive across VIEs. This will require intentionally studying VIEs together to understand how the details of any given VIE modulate the impacts of variable inundation. Credit: Nathan Johnson. There are several photos from different sources and permissions granted as follows: (a) Sullivan et al 2019; (b) Jon Sweetman, co-author; (c) Shutterstock; (d) @WeirdBristol [Twitter] 2018; (global image, e, f, g) Shutterstock; (h) Mikac et al 2018.

142 Mechanistic knowledge that is transferable (per Schuwirth et al. 2019) across inundation  
143 regimes (i.e., from extreme events to predictable cycling) and across VIEs is required to develop  
144 hydrologic, biogeochemical, and ecological models that are predictive across contemporary and  
145 future conditions. We envision the impacts of variable inundation as dependent on the location  
146 of any given VIE within multi-dimensional environmental space. This space can be defined with  
147 a variety of environmental variables such as inundation return interval and duration, topographic  
148 slope, geology, vegetation composition, precipitation, salinity, and temperature. Similar to multi-  
149 dimensional niche space (Hutchinson 1978), many other variables could be used, but  
150 regardless, environmental change will alter the position of VIEs within continuous, multi-  
151 dimensional environmental space. Predicting future impacts of variable inundation requires  
152 mechanistic understanding of how the location of a VIE in this multi-dimensional space  
153 influences those potential impacts. We propose that our best chance to achieve such  
154 understanding is to generate knowledge of variable inundation impacts that is transferable  
155 across VIEs.

156 Here we aim to catalyze cross-VIE science for the pursuit of transferable knowledge and  
157 ultimately models that are predictive across and aid in conserving contemporary and future  
158 VIEs. We briefly summarize high-level divergences in drivers of variable inundation,  
159 commonalities in the impacts of variable inundation, and then present expert mini-reviews of  
160 eight major VIE systems. Variable inundation occurs across vast ranges in spatial and temporal  
161 scales, which presents challenges to cross-VIE science. As such, we overview these challenges  
162 and offer suggested solutions along with a summary of methods that are most relevant to VIE  
163 science. We conclude with perspectives on how cross-VIE science can use conceptual models  
164 based on environmental continuums to derive transferable understanding useful for protecting  
165 these systems and their biodiversity.

166



167  
 168  
 169  
 170  
 171  
 172  
 173  
 174  
 175  
 176  
 177

**Figure 2. Conceptual overview of where different types of VIEs are often found within watersheds and some common shifts in system states across inundated and non-inundated conditions.** VIEs are found from headwaters to coastal environments (Top) and the impacts of variable inundation have some consistencies across these diverse landscapes (Bottom). Organismal ecology, physiology, and demographics are altered by variable inundation, leading to shifts in community composition. Biogeochemical processes also shift, such as greater gas-phase transport of oxygen into soil/sediment when surface water is lost. . A key goal for cross-VIE science is to mechanistically understand variation in the impacts of variable inundation across multi-dimensional environmental space. Credit: Nathan Johnson.

178 **Divergent Drivers, Common Responses, and VIE Mini-Reviews**

179 The drivers of variable inundation differ markedly across VIEs and are linked to factors such as  
 180 long-term drought, heavy precipitation, evapotranspiration, changing groundwater storage,  
 181 soil/sediment properties, extreme climatic events, and dam (Glaser et al. 2021, Shanafield et al.  
 182 2021, Arnold et al. 2023, Bourke et al. 2023, Swenson et al. 2024). This leads to significant  
 183 variation across VIEs in inundation regimes, which includes inundation timing, return interval,  
 184 duration, spatial extent, depth, and flow rate (Celi and Hamilton 2020, Dee and Tank 2020, Van  
 185 Appledorn et al. 2021). For example, sediments within the active channel of tidal rivers can  
 186 experience sub-daily losses and gains of surface water (Tagestad et al. 2021), while other

187 coastal zones may experience extreme inundation events on a 100 year return interval (Slater  
188 et al. 2021, Clementson et al. 2021). Other systems, such as non-perennial streams and vernal  
189 ponds, also experience a broad range of inundation regimes, ranging from sporadic and  
190 extreme inundation following rain events to more regular seasonal cycles (Allen et al. 2020,  
191 Barczok et al. 2023). Variation in the spatial scale of inundation is also large, with floodplains  
192 and storm-impacted coastal zones experiencing inundation over tens of kilometers, whereas  
193 non-perennial streams and ponds can experience changes across a few meters (Hamilton et al.  
194 2002, Vousdoukas et al. 2016, Allen et al. 2020). As discussed below within the series of VIE  
195 mini-reviews, the temporal and spatial scales of inundation also vary substantially within each  
196 type of VIE. Variation within a given type of VIE is large enough that we suggest it cannot be  
197 used to clearly differentiate VIEs into named categories. As discussed in the “Toward cross-VIE  
198 transferable understanding” section, this is one motivation for pursuing VIE conceptual models  
199 and investigations that span broad continuums of environmental conditions.

200 Variable inundation impacts physical [e.g., sediment transport (Peruccacci et al. 2017, Siev  
201 et al. 2019)], chemical [e.g, water quality (Whitworth et al. 2013)], and biological/ecological [e.g.,  
202 invertebrate communities (Plum 2005)] attributes of both natural and anthropogenically modified  
203 ecosystems, in addition to human society (Dube et al. 2021) (**Fig. 2**). Due to intense periods of  
204 inundation and drought, these systems are often referred to as hotspots or ecosystem control  
205 points (Bernhardt et al. 2017, Arias-Real et al. 2024), with disproportionately high reaction rates  
206 or areas of high diversity (Davidson et al. 2012, Palta et al. 2014). In a qualitative sense, some  
207 of these impacts are common across VIEs even if the quantitative details vary.

208 During inundated periods, biogeochemical processes in VIEs often shift from a dominance  
209 of aerobic respiration during drier periods to a diverse suite of anaerobic processes, such as  
210 methanogenesis (Datry et al. 2018b, Hondula et al. 2021b). Changes in the frequency of  
211 inundation events change the dynamics of dry-wet, hot-cold, and aerobic-anaerobic transitions  
212 (Valett et al. 2005). Such dynamics can challenge existing theories. For example, while rates of  
213 soil respiration are expected to peak under aerobic conditions, periodic anaerobic conditions  
214 can lead to unexpectedly high rates of soil carbon loss (Huang et al. 2021) and the anaerobic  
215 process of methanogenesis can be fastest in well-oxygenated dry soils (Angle et al. 2017). More  
216 broadly, variable inundation can alter fluxes of greenhouse gasses to the atmosphere such as  
217 the common observation of soil rewetting leading to significant carbon loss arising from sudden  
218 intensification of soil respiration (Schimel 2018, Shumilova et al. 2019). Variation in inundation  
219 also has large impacts on the global CH<sub>4</sub> budget (Zhang et al. 2017, Peng et al. 2022) and  
220 rewetting of dry sediment in intermittent streams can contribute considerably to the total CO<sub>2</sub>  
221 emissions from streams (von Schiller et al. 2019). More generally, top down and bottom up  
222 hydrologic inundation events broadly influence biogeochemical cycles (Smith et al. 2017) and  
223 can result in hysteretic responses to wetting and drying (Patel et al. 2022).

224 Across VIEs, inundation impacts the structure, composition, and function of vegetation  
225 communities. Growth and survival can either increase or decrease with inundation depending  
226 on local aridity and the impacts on soil hypoxia. Hypoxia kills roots, leading to reduced water  
227 uptake, reduced photosynthesis, mortality (Pedersen et al. 2021, McDowell et al. 2022, Cubley  
228 et al. 2023), and shifts in vegetation composition. More broadly, inundation dynamics impact  
229 organismal ecology (Datry et al. 2023) across all VIEs, such as herbivores responding to  
230 inundation-induced shifts in vegetation (De Sassi et al. 2012). Inundation can also alter

231 arthropod communities leading to reductions in diversity, abundance, and biomass (Plum  
232 2005). Changes at the base of food webs can have further, cascading effects (Chen and Wise  
233 1999).

234 To pursue cross-VIE science requires knowledge of the diverse array of ecosystems that  
235 can be considered VIEs. Researchers that design and carry out cross-VIE studies may be  
236 considered generalists in terms of the breadth of systems they study, even if their science  
237 questions are highly specialized. To facilitate such researchers in the pursuit of cross-VIE  
238 science, we go beyond the high-level summaries of divergences and commonalities (above)  
239 and provide expert mini-reviews of eight primary VIE types. The following subsections present  
240 these mini-reviews which summarize system characteristics, drivers, and impacts of variable  
241 inundation with an emphasis on biogeochemistry and organismal ecology, and opportunities to  
242 better understand spatiotemporal patterns and impacts of variable inundation. Each mini-review  
243 is accompanied by a graphic that either provides a conceptual overview or imagery-based  
244 examples, with the goal of collectively touching on key drivers, dynamics, impacts, and tangible  
245 system examples. The collection is not meant to be a comprehensive classification of all  
246 possible VIE types. It does cover a broad range of VIEs and is meant to serve as an overview of  
247 individual VIEs to provide context for later sections of this manuscript. The sequence of mini-  
248 reviews roughly follows the flow of water moving from hillslopes to coastal environments (**Fig. 2**)  
249 and includes variably inundated components of: (i) hillslopes, (ii) non-perennial streams, (iii)  
250 riverine floodplains and parafluvial zones, (iv) wetlands, (v) temporary ponds, (vi) storm-  
251 impacted coastal zones, and (vii) tidal systems. The final mini-review (viii) is focused on  
252 ecosystems that have been engineered to modify inundation regimes, which occur throughout  
253 the continuum from hillslopes to coasts.

254 We separate VIEs into categories as a heuristic simplification that allows for an appreciation  
255 of variation and commonalities in drivers, impacts, and opportunities. We anticipate that the  
256 disciplinary foci of individual researchers will align most closely with a subset of the summarized  
257 VIE types. One goal of this manuscript is to facilitate researchers thinking about how their  
258 science applies across multiple VIEs. We emphasize that in many (and maybe all) cases there  
259 is not a clear distinction among the types of VIEs we discuss below (e.g., non-perennial streams  
260 can be inundated due to storm surge, resulting in floodplains or parafluvial zones). Ultimately,  
261 we encourage a continuum perspective that does not rely on discrete system names or hard  
262 boundaries, and instead views VIEs across multi-dimensional environmental space based on  
263 inundation regimes and physical settings such as topographic slope.

264 This continuum perspective is more fully developed as a conceptual model in the final  
265 section of the paper, titled "Towards Cross-VIE Transferable Understanding." However, we  
266 briefly summarize here that it is based on two continuous environmental axes: inundation return  
267 interval and topographic slope. These variables can be used to define a two-dimensional  
268 environmental space that contains all VIE systems. With this model, impacts of variable  
269 inundation can be studied across environment space instead of within discrete named types of  
270 VIEs. When going through the following mini-reviews, we encourage the reader to conceptualize  
271 each VIE type in context of return interval and slope (e.g., hillslopes may have a long return  
272 interval and steep slopes relative to tidal systems, while coastal systems inundated by storms  
273 may have similar slopes as tidal systems but much longer return intervals). When VIEs are

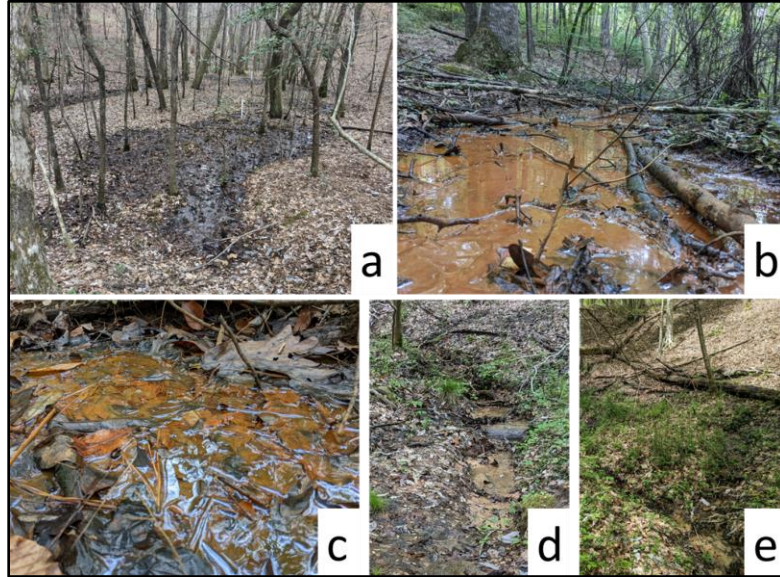


274 viewed through a unified lens of environmental continuums, larger interdisciplinary questions  
275 may be answered.

### 276 **Hillslopes with Surface Runoff**

277 Hillslopes provide water to lower-lying areas, often concentrating the water in gullies and  
278 depressions (**Fig. 3**). Hillslopes produce relatively transient VIE features and may often be seen  
279 as extensions of other VIEs, such as hillslope seeps co-located with a wetland or the  
280 unchannelized swales that contribute to a non-perennial network. In cold regions, snow, ice and  
281 permafrost can create an impermeable layer resulting in near-surface soil being inundated for  
282 days to weeks during spring thaw (Coles et al. 2017, Patel et al. 2020). In dry regions, intense  
283 precipitation that exceeds the local infiltration capacity can result in water ponding on the  
284 surface of hillslopes or overland flow generation down hillslopes, which can be exacerbated by  
285 initial hydrophobicity of dry soil (Kirkby et al. 2002). Exceeding the infiltration capacity is more  
286 likely on hillslopes with low-permeability, such as clay-rich soil or when near-surface soils are  
287 frozen. This can be exacerbated by restrictive soil horizons located at shallow depths across  
288 hillslopes that generate seasonal perched water tables and lead to inundation (McDaniel et al.  
289 2008). Overland flow can be spatially heterogeneous due to variations in soil characteristics as  
290 well as flow accumulation, leading to infiltration or exfiltration along the hillslope (Betson and  
291 Marius 1969).

292 In forested hillslopes, soil infiltration often exceeds rainfall intensity (McDonnell 2009, Burt  
293 and Swank 2010) and lateral flow towards topographic depressions can lead to saturation and  
294 ponding (Anderson and Burt 1978) (**Fig. 3a**). Microtopography within hillslopes (**Fig. 3b**) can  
295 also lead to temporary ponding, e.g., from rain in tropical environments and from spring  
296 snowmelt in colder environments (Clark et al. 2014). Toe slopes can generate wedges of  
297 saturation that grow upslope (Weyman 1973, Choularton and Perry 1986), although subsurface  
298 saturation and ponding can also occur on upper slopes where the soil is thinner [e.g., (Tromp-  
299 van Meerveld and McDonnell 2006)]. Finally, spatial variation in topographic characteristics  
300 (e.g., aspect, slope, curvature) can result in differences in soil moisture, incoming energy, and  
301 vegetation, affecting evapotranspiration and inundation patterns (McVicar et al. 2007).  
302



303  
304  
305  
306  
307  
308  
309

**Figure 3. Examples of variable inundation along hillslopes.** a) looking downslope at an inundated slope; b) ponding with no flow due to microtopography; c) sheet wash with directional flow across the surface of a hillslope; d) rill formation with turbid water from erosion; e) vegetation community change on slope due to differences in soil moisture. All photos taken by Corianne Tatariw at Tanglewood Forest, Alabama.

310  
311  
312  
313  
314  
315  
316  
317  
318  
319  
320  
321  
322  
323  
324

Surface runoff and inundation on hillslopes can result in the export of soil nutrients, salinization of soil from groundwater seeps, erosion, and landslides. There is a balance between the effects of variable inundation on hillslope vegetation and erosion. In water-limited systems, inundation can increase plant productivity and diversity, as well as increased rooting strength of soils (Zhao et al. 2022) (**Fig. 3e**). However, increased inundation can also lead to increased chemical weathering and lower shear strength in hillslope soils during storms, leading to higher erosion and landslide potential. Along with erosion, landslides and soil compaction are inherent to many hillslopes, which also can create areas ripe for inundation (Bogaard and Greco 2016). At shoulder and midslope positions, increased overland flow due to saturation- or infiltration-excess increases sediment detachment, which is then deposited in foot and toe slopes (Huang et al. 2002). The transport of particles also leads to the transport of nutrients that are sorbed to the particles, such as phosphorus. Erosion can be concentrated in rills and gullies or can spread out across a slope as ‘sheet wash’ that impacts large areas of hillslopes (**Fig. 3c,d**). Impacts of erosion are dependent on interactions between the persistence of inundation and soil properties (Thomas et al. 2020).

325  
326  
327  
328  
329  
330  
331  
332

The aqueous chemistry of water that is transported over hillslope surfaces reflects the chemistries of contributing water sources such as precipitation, shallow soil water, and exfiltrating groundwater. Shallow soils in hillslopes have abundant organic materials and nutrients (Herndon et al. 2015), whereas organic matter decreases with depth, solutes derived from the parent rock material increase with depth (Brantley et al. 2017). These stratifications collectively regulate source water chemistry in hillslopes. Dry to wet transitions shift flow paths from groundwater to soil water dominance in streams, therefore shaping stream chemistry (Zhi and Li 2020, Stewart et al. 2022). Dry to wet transitions also shift water content and pore space

333 oxygen concentrations (Jarecke et al. 2016, Smyth et al. 2019), often triggering the release of a  
334 cascade of solutes produced under anaerobic conditions (Schlesinger and Bernhardt 2020).  
335 These entangled, complex interactions among hydrological and biogeochemical processes  
336 often challenge the differentiation of individual processes and mechanistic understanding on  
337 how variable inundation regulates flow paths, reactions, stream chemistry, and solute and gas  
338 export fluxes (Li et al. 2021).

339 Investigations of variably inundated hillslopes present significant and challenging research  
340 opportunities due to their inherently dynamic nature. One key challenge is quantifying the  
341 occurrence and spatial extent of hillslope VIEs across the globe. Remote sensing could be used  
342 to identify and quantify these areas, spatially and temporally, based on sky-visible vegetation  
343 (e.g., plant morphologies, leaf nutrient contents) (Hwang et al. 2012, Tai et al. 2020) and  
344 topographic signatures (e.g., erosional patterns) (Trochim et al. 2016) caused by variable  
345 inundation. To fully understand the ecological and biogeochemical impacts of variable  
346 inundation on hillslopes, research needs to focus on shallow subsurface physical properties,  
347 hydrology, and their linkage to biogeochemical processes. This can be pursued via  
348 environmental geophysics to map and characterize the influence of subsurface restrictive layers  
349 (Fan et al. 2019 p. 201). Understanding the subsurface soil architecture is key to predicting  
350 variable inundation from bottom-up and top-down water sources, along with the follow-on  
351 impacts to ecology and biogeochemistry.

## 352 **Non-Perennial Streams**

353 Non-perennial streams, defined as rivers and streams that cease to flow at some point in either  
354 space or time (Busch et al. 2020), are ubiquitous and comprise 50-60% of the global river length  
355 (Messenger et al. 2021). These systems occur across all continents and biomes (Messenger et al.  
356 2021). Streamflow in non-perennial streams ranges from nearly perennial (year-round) flow, to  
357 seasonal flow, responding to drivers like snowmelt, to daily or sub-daily flow events responding  
358 to rainfall events or evapotranspiration (Price et al. 2021). At the reach scale, non-perennial  
359 streams shift between three main states - flowing, ponded/pooled, or no-surface water present  
360 (**Fig. 4**). As reaches become hydrologically connected (or disconnected), the spatial  
361 footprint/extent of the connected stream network can grow or shrink over sub-daily to seasonal  
362 to interannual timescales (Xiao et al. 2019). Spatial and temporal shifts among the three  
363 hydrologic states strongly influence the network's capacity to process, transport, and export  
364 material to downstream systems (Allen et al. 2020).

365 The high variability in the spatial and temporal scales of streamflow intermittency is  
366 indicative of the complex set of interacting drivers that induce stream drying. At the global and  
367 regional scales, the degree of aridity is a primary control on the abundance of non-perennial  
368 streams (Hammond et al. 2021, Zipper et al. 2021). At smaller scales, catchment properties  
369 exert strong control over both the capacity of water delivery to the channel and the subsequent  
370 balance between the channel and near subsurface capacity to transport water (Hammond et al.  
371 2021, Zipper et al. 2021, Price et al. 2021). Non-perennial flow can occur anywhere in the stream  
372 network, from headwaters to higher order rivers. While some networks display longitudinal  
373 transitions from non-perennial to perennial flow (or vice versa), other networks exhibit more  
374 complex patterns in surface water flow and connectivity, which may be driven by topography,  
375 geology, vegetation, or groundwater abstraction/use (Costigan et al. 2015, 2016).

376 The variable inundation dynamics in non-perennial streams have cascading implications for  
377 biogeochemical cycling, water quality, ecosystem function, and community ecology. Under non-  
378 flowing conditions, riverbeds are characterized by dry conditions or discontinuous and stagnant  
379 water pools, often with high temperatures, low dissolved oxygen levels, and long residence  
380 times, functioning more like soils (Arce et al. 2019), as described also in the hillslope section.  
381 Pooled, non-flowing conditions can lead to steep redox gradients in the shallow subsurface that  
382 drive nutrient processing (Datry and Larned 2008, Gómez-Gener et al. 2021, DelVecchia et al.  
383 2022). During dry/non-flowing states, terrestrial organic matter accumulates in the channel and  
384 is subjected to varying degrees of breakdown (Datry et al. 2018c, Del Campo et al. 2021).  
385 Rewetting of accumulated substrates can stimulate microbial activity, nutrient attenuation  
386 (Saltarelli et al. 2022), and generate pulses of greenhouse gasses such as CO<sub>2</sub> and N<sub>2</sub>O (Datry  
387 et al. 2018a, Song et al. 2018). During re-wetting and resumption of flow, non-perennial streams  
388 can contain large amounts of terrestrial and aquatic organisms that can be flushed downstream  
389 (Corti and Datry 2012, Rosado et al. 2015), with high sediment, dissolved organic carbon, and  
390 solutes (Laronne and Reid 1993, Hladyz et al. 2011, Herndon et al. 2018, Wen et al. 2020,  
391 Fortesa et al. 2021, Blaurock et al. 2021).

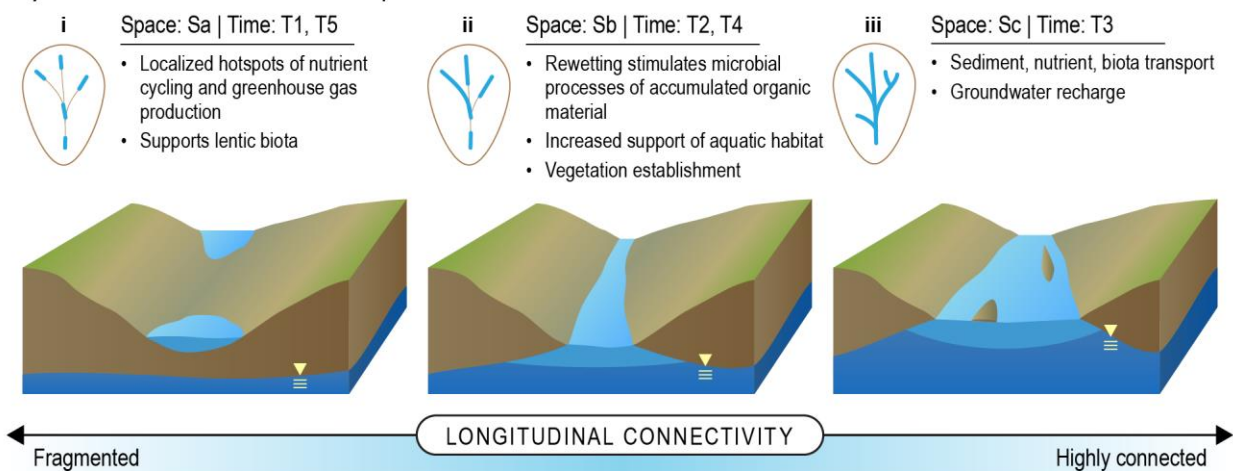
392 Biological responses to rewetting depend on the distribution of habitats and biota at the  
393 watershed scale and the duration of the preceding dry phase. In highly dynamic river systems,  
394 such as braided rivers, drying and wetting cycles can be spatially patchy and short-lived but  
395 frequent, and thus ecological recovery following wetting can be very rapid due to the very active  
396 hyporheic zones characterizing these systems (Arscott et al. 2002, Vorste et al. 2016). In other  
397 systems recovery can be slow, depending on the proximity of refuges, such as springs, isolated  
398 pools, and perennial reaches (Sarremejane et al. 2021, Fournier et al. 2023). Systems with  
399 frequent and severe drying events are more likely to be colonized by aerial or other overland  
400 dispersers than by aquatic dispersers (Bonada et al. 2007, Bogan et al. 2017a, Sarremejane et  
401 al. 2021). Life-history events of some species coincide with predictable rewetting events, such  
402 as post-snowmelt fish spawning (Hooley-Underwood et al. 2019) and amphibian and insect life  
403 histories (Bogan et al. 2017a). Rewetting also partly determines the germination success and  
404 establishment of riparian vegetation (Merritt and Wohl 2002).

405 Compared to their perennial counterparts, non-perennial streams have received less  
406 research and monitoring attention and tend to be undervalued relative to ecological/functional  
407 performance of perennial streams (Palmer and Hondula 2014). As such, many of the pressing  
408 research needs in non-perennial streams are limited by data availability (Van Meerveld et al.  
409 2020, Zimmer et al. 2022). Non-perennial streams are systematically under-represented in  
410 global gaging networks (Messenger et al. 2021, Krabbenhoft et al. 2022), leading to major gaps  
411 in our understanding of the timing, magnitude, and duration of flow in diverse non-perennial  
412 streams. In addition, our ability to predict the onset or cessation of flowing periods is limited by a  
413 lack of gaging. Infrequent grab sampling for water chemistry tends to undersample non-  
414 perennial streams specifically, leading to an even greater paucity of biogeochemical data from  
415 these systems, particularly during rapid re-wetting events. Spatially explicit data on streamflow  
416 intermittency and subsurface conditions at fine spatial scales (10s of meters) remain limited to a  
417 few intensively studied catchments [e.g., (Zimmer and McGlynn 2017)]. While some global scale  
418 datasets on streamflow intermittency have been developed (Messenger et al. 2021), the  
419 resolution of these products necessarily omit smaller, headwater reaches, hindering our ability

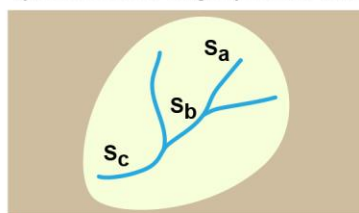
420 to quantify hydrologic and biogeochemical processes in non-perennial streams broadly  
 421 (Benstead and Leigh 2012).

422 Major challenges and opportunities include accurate mapping of non-perennial streams and  
 423 accurate predictions of flow timing at annual, seasonal, and shorter time scales across scales.  
 424 Headwaters, which are small, numerous, and often non-perennial (Kampf et al. 2021), are  
 425 difficult to map and understand hydrologically, leading to knowledge gaps in the hydrological  
 426 integrity of ecosystems at regional scales (Benstead and Leigh 2012, Dugdale et al. 2022).  
 427 While challenges remain, the use of drones and thermal infrared remote sensing could connect  
 428 field observations with modeling to better understand the hydrology of these valuable systems  
 429 (Dugdale et al. 2022). In addition to mapping issues, limited time series data makes predictions  
 430 of flow in terms of duration, frequency, and spatial extent challenging. How the timing and  
 431 frequency of flow will change under climate change remains an open question. It is expected  
 432 that an increased frequency and duration of droughts will shift streams toward more non-  
 433 perennial flow states (Döll and Schmied 2012). In contrast, flow permanence may increase in  
 434 select areas where streams are fed by melting glaciers or snowpack, or where anthropogenic  
 435 intervention occurs (Datry et al. 2023). The changing frequency of extreme flow events and  
 436 rapid no-flow/high-flow oscillations also have the potential to further alter streamflow,  
 437 biogeochemical processes, and organismal ecology in non-perennial streams, necessitating  
 438 further integrated hydro-biogeochemical studies in these dynamic systems.  
 439

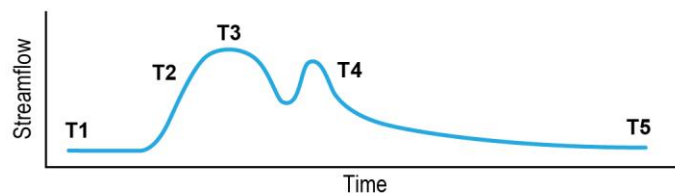
**a) Inundation mechanisms in non-perennial streams**



**b) Spatial variation in inundation dynamics at a single point in time**



**c) Temporal variation in inundation dynamics at the watershed outlet**



440  
 441 **Figure 4. Conceptual model of variable inundation in non-perennial streams.** a) Water  
 442 connections between groundwater, near surface, and surface regions at locations within a given

443 *network result in varying degrees of longitudinal connectivity with associated biogeochemical*  
444 *processes. b) At a single snapshot in time, water connections result in spatial variation in*  
445 *surface water inundation. c) Under time varying flow states, extent of surface inundation will*  
446 *also vary at a given location. Inundation mechanisms depicted in a) represent a losing system*  
447 *that is transitioning to a flowing state. We acknowledge that in some systems, a low flow*  
448 *fragmented state also occurs in gaining streams with locally connected groundwater. Spatial*  
449 *variation is signified by Sa - Sc and temporal variation is signified by T1 - T5. Credit: Nathan*  
450 *Johnson.*

## 451 **Floodplains and Parafluvial Zones**

452 Rivers, both perennial and non-perennial, create two types of VIEs, floodplains and parafluvial  
453 zones (**Fig. 5**). Floodplains are alluvial landforms generated by river erosion and deposition and  
454 hydrologically connected to the contemporary active channel but outside the active river channel  
455 (Nanson and Croke 1992). Parafluvial zones are areas in the active channel without surface  
456 water at low flow, i.e., at higher-elevation areas within an active channel that contains perennial  
457 flow (Goldman et al. 2017). Nearly all rivers have parafluvial zones and adjacent floodplains,  
458 although these may be longitudinally discontinuous (e.g., absent where the river flows through a  
459 narrow bedrock gorge or descends into the subsurface). Consequently, the global distribution of  
460 these environments is extensive, as few terrestrial surfaces do not include a river network.

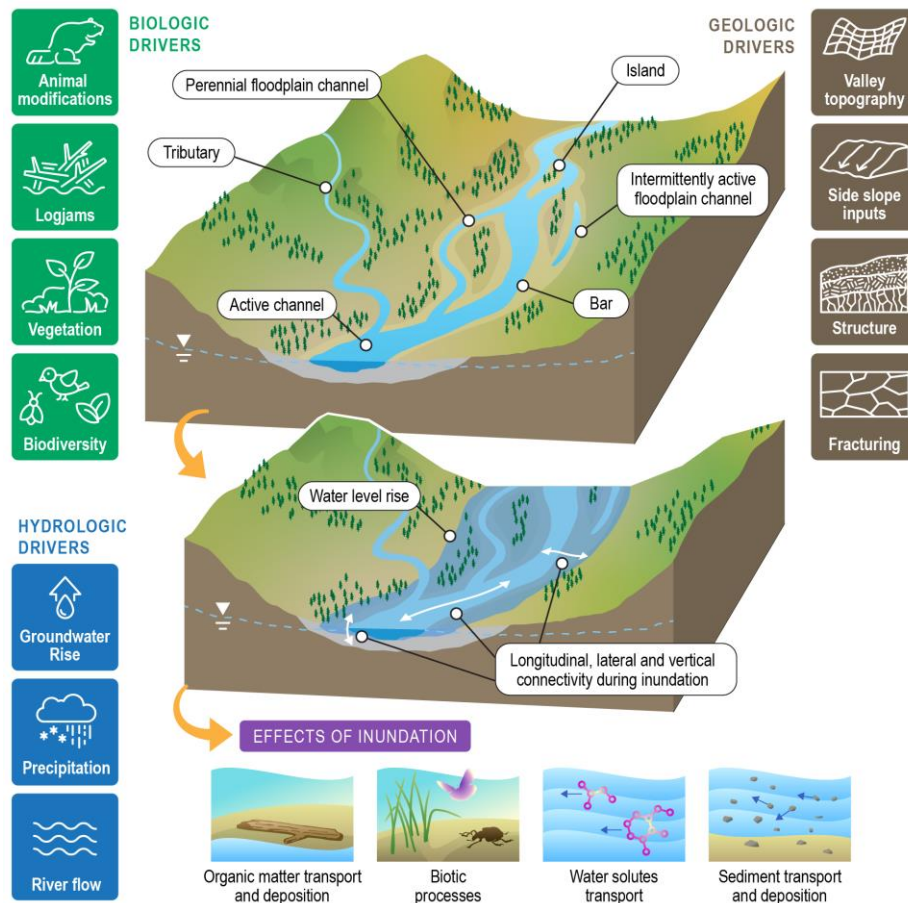
461 Spatial scales of inundation in floodplains and parafluvial zones are variable between rivers  
462 and through time along a river. Fundamentally, spatial scales are governed by the interaction  
463 between the magnitude of flow and available space as defined by topography (Nardi et al.,  
464 2006). Floodplains of the world's largest rivers such as the Amazon, Congo, or Mississippi can  
465 extend laterally for kilometers on both sides of the active channel (Arnesen et al. 2013). In  
466 contrast, the floodplain of a headwater channel in high-relief terrain may be only 1-2 m wide on  
467 each side of the channel (Adams and Spotila 2005).

468 Temporal scales of inundation (e.g., frequency, periodicity, intensity) vary substantially  
469 across climates, topographic regions, and river network position. A snowmelt-dominated or  
470 monsoon-fed river will have regular annual inundation that lasts for weeks, whereas a small  
471 stream dominated by convective rainfall or tropical depressions may have irregular floods that  
472 only last for hours. Although precipitation-driven over bank flow from the main and tributary  
473 channels is the primary driver of inundation on floodplains and parafluvial zones, inundation also  
474 results from direct precipitation, rising water tables, and overland flow from adjacent uplands  
475 (Mertes 2011). Thus, inundation of floodplains may be directly related to their proximity to  
476 variably inundated hillslopes and streams.

477 The nature of floodplain/parafluvial inundation affects the dynamics of surface and  
478 subsurface water, solutes, particulate organic matter, sediment, and biota (Junk et al. 1989).  
479 Dynamics include volume and duration of storage; rate of movement; direction of movement  
480 between surface, hyporheic, and groundwater; and biogeochemical alterations that in turn  
481 impact river water quality, greenhouse gas emissions, plant function, and organismal ecology.  
482 The duration, frequency, and areal extent of floodplain/parafluvial inundation control ecosystem  
483 function, and the types and abundances of organismal communities, including both aquatic and  
484 terrestrial species (Ward et al. 1999). Species distribution, movement, and biological  
485 interactions, such as predator-prey, are intricately tied to these inundation patterns (Robinson et

486 al. 2002, Stanford et al. 2005). Fish species, for example, can migrate from dry season refugia  
 487 into floodplains during inundation, influencing food web structure and ecosystem productivity  
 488 (Crook et al. 2020).

489 Among the primary challenges to answering questions regarding the variation in  
 490 floodplain/parafluvial inundation are limited monitoring data and a lack of numerical models that  
 491 integrate knowledge across disciplines and processes. Measurements and models of hydrology  
 492 commonly treat floodplains as flat, impermeable surfaces, which ignores surface-subsurface  
 493 water exchanges that influence hydrology and ecosystem function (Wohl 2021). Models also  
 494 often ignore the micro-heterogeneities that influence spatially and temporally variable patterns  
 495 of inundation, biogeochemical cycling, and ecology in both floodplains and parafluvial zones.  
 496 The degree of physical detail represented in models often involves tradeoffs in spatiotemporal  
 497 extent; a one-dimensional model might ignore microtopography that influences important  
 498 inundation details, whereas a more representative two-dimensional or three-dimensional model  
 499 becomes computationally intensive for larger spatial extents. This problem gives rise to the  
 500 challenges and opportunities for (i) designing measurement campaigns across disciplines that  
 501 can create integrative data for diverse floodplains and parafluvial zones to adequately represent  
 502 the physical complexity of variable inundation processes at broad scales, and (ii) developing  
 503 floodplain/parafluvial functional groups [e.g., (Fryirs and Brierley 2022)] that can facilitate  
 504 understanding of scaling and transferability of data.  
 505  
 506



507

508 **Figure 5. Conceptual model of variable inundation in floodplain and parafluvial systems.**  
509 *Across floodplains and parafluvial zones a suite of biological, hydrologic, and geologic factors*  
510 *drive inundation regimes in terms of spatiotemporal duration, timing, depth, flow rate, etc. These*  
511 *systems include diverse subsystems as summarized in the top panel. Rising water levels, due*  
512 *to one or more drivers, can inundate these subsystems as shown in the middle panel, resulting*  
513 *in a variety of biogeochemical, ecological, and physical effects (bottom sub-panels). Credit:*  
514 *Nathan Johnson.*

## 515 **Variably Inundated Wetlands**

516 While not all wetlands are variably inundated, variable inundation is a common feature of many  
517 wetland ecosystems (Arias-Real et al. 2024). Here we focus primarily on wetlands that are  
518 similar to swamps, marshes, and bogs (**Fig. 6**). Wetlands cover about 10% of the global land  
519 area, and nearly half of global wetland area (46%) is temporarily inundated (Davidson et al.  
520 2018). Generally, wetland inundation regimes are shaped by the wetland's connectivity to  
521 surface and subsurface hydrologic sources and landscape position (Åhlén et al. 2022). The  
522 landscape position of wetlands is a first order indicator of the water source and chemistry,  
523 ranging from headwater depressional locally-fed wetlands, to flow-through and fringing wetlands  
524 to groundwater-fed low-lying wetlands (Fan and Miguez-Macho 2011, Tiner 2013). Wetland  
525 typologies applied in several national inventories generally rely on a combination of three  
526 criteria: soil type, hydrophytic vegetation and hydrology (Cowardin and Golet 1995).  
527 Alternatively, hydrogeomorphic classification systems propose to exclusively draw on physical  
528 drivers, such as geomorphology, hydrology and substrate to allow for cross-site comparisons of  
529 biota and serve functional assessments (Brinson 1993, Semeniuk and Semeniuk 1995, 2011,  
530 Davis et al. 2013).

531 While changes to inundation extent and depth can occur at time scales ranging from days to  
532 decades, the most conspicuous inundation patterns occur on event (e.g., following rain events),  
533 seasonal (e.g. snow melt or wet/dry seasons), and interannual time scales. Primary drivers of  
534 inundation in unmanaged wetlands come from subsurface groundwater discharge and surface  
535 flows including rainfall or snowmelt runoff that occur when antecedent soil moisture conditions  
536 are high, preventing quick infiltration of water (Rasmussen et al. 2016). Many wetlands are  
537 actively managed, such as to provide bird habitat, so that inundation can vary based on  
538 management decisions [see below and (Fredrickson and Taylor 1982)].

539 The spatial scales of variable inundation are shaped both by wetland size and  
540 geomorphology. Wetlands can be shallow over large spatial scales, and thus the size of variably  
541 inundated wetland area can range from microtopographic (i.e., hummock/hollow, ~m<sup>2</sup> scales) to  
542 larger ecosystem scales. Large wetland areas, especially in the tropics, experience strong  
543 seasonal inundation cycles which depend on changes in water balance and local topography  
544 (Zhang et al. 2021). While the largest variably inundated wetlands are connected to floodplains,  
545 like the 130,000 km<sup>2</sup> Pantanal located in Brazil and extending into Bolivia and Paraguay (Ivory  
546 et al. 2019), non-floodplain wetlands surrounded by upland (also known as geographically  
547 isolated wetlands) as large as ~6 ha may also experience whole-system drying and rewetting  
548 (Lane and D'Amico 2016).

549 Embedded within wetland ecosystems, microtopographic structures can create within-  
550 system mosaics of inundation regimes. Microtopography in peaty wetlands is particularly



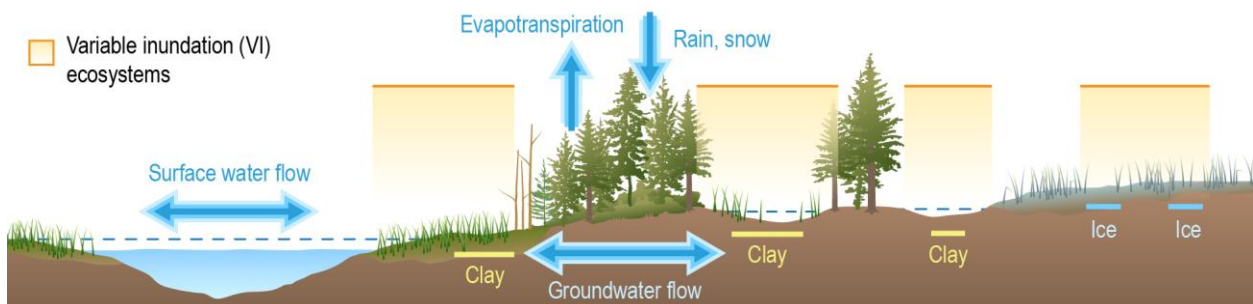
551 pronounced, ranging from several tens of meters [e.g., ridges and sloughs (Larsen et al. 2011)]  
552 to meters [e.g. hummock-hollows (Shi et al. 2015)], These spatial patterns result from dynamic  
553 feedbacks between ecological processes (e.g. peat accumulation) and hydrology that reinforce  
554 these patterns (Belyea and Baird 2006, Eppinga et al. 2008, Larsen et al. 2011).

555 Wetlands are widely acknowledged to be biogeochemical hot spots and ecosystem control  
556 points (McClain et al. 2003, Bernhardt et al. 2017) because of the confluence in space and time  
557 of allochthonous substrates into reactive environments (e.g., nitrate produced under oxic  
558 conditions entering anaerobic environments where denitrification can occur). In addition,  
559 variable inundation is associated with nutrient influx into wetlands that replenishes nutrient pools  
560 and can drive productivity and organic matter decomposition (Venterink et al. 2002). The depth  
561 and duration of inundation shapes the wetland vegetation community by controlling germination  
562 success, modifying oxygen availability and changing concentrations of toxins and nutrients, by  
563 desiccating aquatic plants or inundating terrestrial plants, and by changing the light availability  
564 (Casanova and Brock 2000). Wetland vegetation is structurally adapted to low oxygen  
565 environments, for example, some vegetation has developed air channels in leaves, stems, and  
566 roots to transport oxygen belowground (Tiner 2017). Alternatively, wetland vegetation can also  
567 respond to shifts in oxygen levels physiologically on shorter time scales (Colmer 2003).

568 Variable inundation provides an environmental filter for biota adapted to live either under dry  
569 or inundated conditions, resulting in distinct communities including wetland obligate and  
570 facultative species (Gleason and Rooney 2018). The temporal duration of inundation (i.e.,  
571 hydroperiod) indirectly controls the bird community composition through absence and presence  
572 of wetland vegetation and availability of aquatic macroinvertebrate prey (Daniel and Rooney  
573 2021). Amphibian communities are particularly impacted by hydroperiod: It needs to be long  
574 enough for eggs to hatch and tadpoles to reach metamorphosis, but should not allow the  
575 establishment of many predator species (Resetarits 1996).

576 Predicting how complex inundation patterns in wetlands will change under changing climate  
577 is a major research challenge. Predictions span the range from a decrease in inundation in  
578 some regions (Londe et al. 2022) to an increase in others (Watts et al. 2014), with uncertain  
579 consequences for wetland persistence overall. To improve regional or global predictions,  
580 accurate maps of wetland extent on different scales that can be incorporated into mechanistic  
581 models will be necessary (Melton et al. 2013). This is particularly challenging for non-permanent  
582 wetlands, which are hard to reliably map and are generally understudied (Gallant 2015, Calhoun  
583 et al. 2017), but which are, by definition, VIEs. As climate change alters wetland inundation  
584 regimes, the net impacts to carbon storage and greenhouse gas fluxes are of particular concern  
585 (Moomaw et al. 2018), because together they will determine the net climatic impact of changes  
586 in wetland area and dynamics (Neubauer and Megonigal 2015).

587



588

589 **Figure 6. Conceptual model of variable inundation in wetland systems.** *Different wetland*  
590 *types are influenced and shaped by variable inundation. Absence and presence of surface*  
591 *water is driven by (e.g., seasonally) changing water supply and the hydrologic function of the*  
592 *wetland in the landscape. Sediment characteristics (e.g., clay or ice) and topographic positions*  
593 *of wetlands in the landscape influence water loss to infiltration or gain from groundwater. Credit:*  
594 *Nathan Johnson.*

## 595 **Freshwater Ponds**

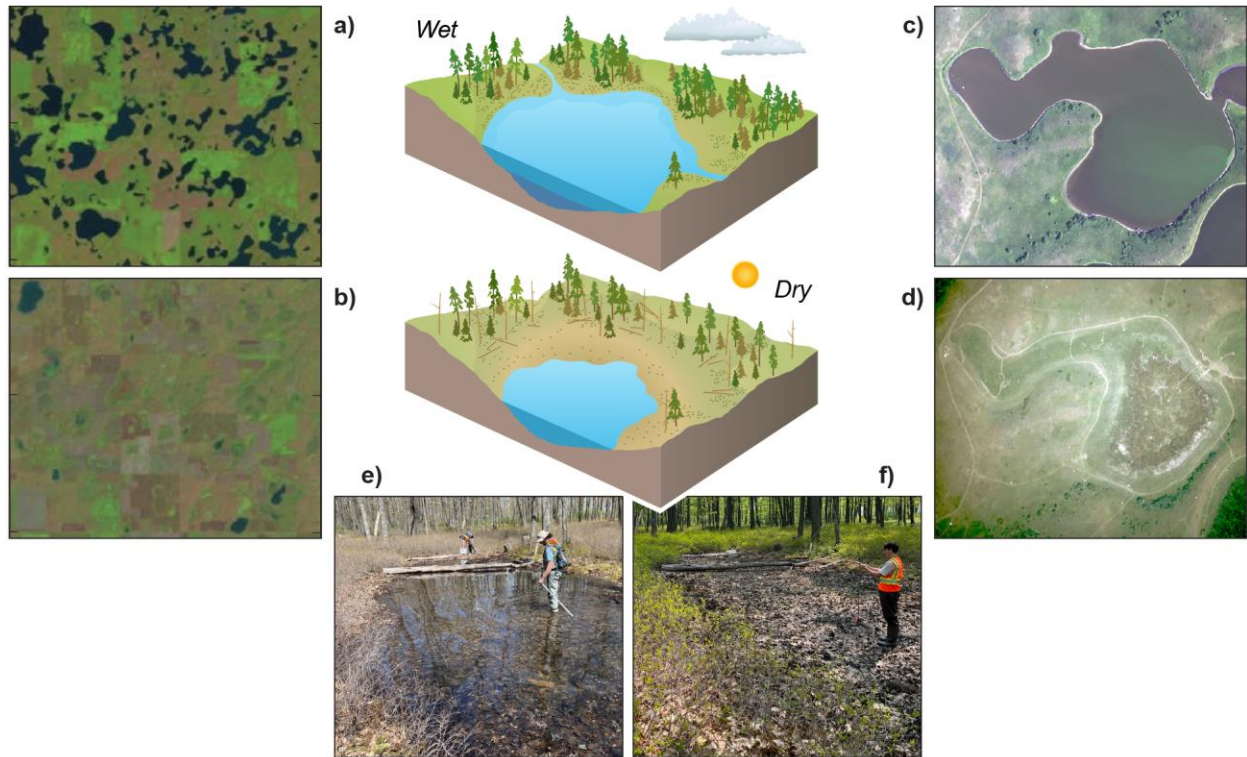
596 Freshwater ponds are among the most abundant and common freshwater ecosystems  
597 worldwide, with estimates between 500 million and 3.2 billion ponds globally (Davidson et al.  
598 2018, Hill et al. 2021). Ponds are generally small (less than 5 ha) and shallow (less than 5 m),  
599 and consequently, are highly sensitive to changes in water levels that can result in highly  
600 variable inundation regimes (Gendreau et al. 2021, Richardson et al. 2022a). Pond ecosystems  
601 are extremely diverse, and include arctic thermokarst ponds, prairie potholes, vernal pools,  
602 playas, rock pools and agricultural dugouts. The numbers of ponds globally are likely  
603 underestimated, as their size and ephemeral/temporary nature has meant they are often  
604 excluded from physical inventories and they are below the resolution of many remote sensing  
605 techniques (Hayashi et al. 2016, Calhoun et al. 2017, Hill et al. 2021).

606 As in many other VIEs, inundation of freshwater ponds can be highly variable, and the  
607 timing, duration and frequency of inundation can vary considerably (Williams 2006). Many  
608 temporary or ephemeral ponds can become intermittently or seasonally inundated (**Fig. 7**). For  
609 some ponds, particularly vernal pools, seasonal inundation is relatively predictable, as these  
610 systems become inundated following snowmelt or spring runoff, and are subsequently drawn  
611 down with increasing summer evapotranspiration (Zedler 2003, Brooks 2004). Variation in the  
612 hydroperiod can alter the composition of biotic communities (Brooks 2004, Gleason and Rooney  
613 2018), as well as impact biogeochemical and hydrological processes (Bam et al. 2020, Hondula  
614 et al. 2021b). In more temperate regions, the timing of inundation is often driven by heavy  
615 rainfall, and periods of inundation can be highly variable, with inundation durations lasting from  
616 days to months, and sometimes occurring intermittently as ephemeral systems dry and rewet  
617 multiple times in a year (Ripley and Simovich 2009, Kneitel 2014, Florencio et al. 2020). For  
618 nearly permanent ponds, the pattern of wet and dry periods are more predictable, but the  
619 initiation and length of the hydroperiod can vary spatially as water levels fluctuate, inundating  
620 and exposing shallower areas (Brendonck et al. 2017). Freshwater ponds often demonstrate  
621 both high inter- and intra-annual variability, and diurnal, annual and multidecadal periods of  
622 inundation can occur due to changes in evapotranspiration, drought, drainage, and / or  
623 hydrologic function of the pond on the landscape (Brooks 2004, Gendreau et al. 2021).  
624 Modifications to ponds by humans (e.g. irrigation ponds, urban stormwater ponds; see section  
625 on human-engineered systems) or other organisms, such as beavers, can also impact  
626 hydroperiod and inundation regimes (Renwick et al. 2006, Brazier et al. 2021).

627 Like many of the other ecosystems that experience variable inundation, freshwater ponds  
628 are also considered biodiversity and biogeochemical hotspots, providing many critical  
629 ecosystem services (Capps et al. 2014, Marton et al. 2015). Despite their relatively small size,  
630 ponds can have considerable variability in both community composition and in biogeochemical  
631 processes, in part due to differences in inundation regimes, where pond margins are more likely

632 to be more frequently desiccated for longer periods than central regions (Reverey et al. 2018).  
633 Models that explicitly incorporate remotely sensed variable inundation predict that ephemeral  
634 systems with shorter hydroperiods retain nitrogen at greater rates than larger systems with less  
635 variable inundation and longer hydroperiods, particularly in semi-arid regions like the Prairie  
636 Potholes of the North American northern Great Plains and playas in the south-central United  
637 States (Cheng et al. 2023). In addition, research suggests reproduction is largely impacted by  
638 inundation. Salamanders, for example, tend to lay more eggs during years with greater rainfall  
639 while hatching success decreases with desiccation (Della Rocca et al. 2005). Variable  
640 inundation across ponds can result in ecosystem heterogeneity at the landscape scale,  
641 increasing local abiotic and biotic variation (Jeffries 2008), but the number and distribution of  
642 inundated ponds can also impact regional biodiversity through processes like dispersal  
643 (Brendonck et al. 2017).

644 Climate change will likely alter the inundation regimes in freshwater ponds in terms of timing,  
645 frequency, duration, and extent. Decreases in precipitation and increases in extreme drought  
646 can result in shortened hydroperiods, and increasing temperatures can alter water temperatures  
647 and evaporation rates (Matthews 2010). The persistence of freshwater ponds may, therefore, be  
648 reduced with climate change (Londe et al. 2022). Understanding how future changes in  
649 inundation regimes impact freshwater ponds will be critical. Similar to wetland ecosystems,  
650 improved remote sensing methods, including incorporating multispectral imagery and radar  
651 along with finer spatial resolution mapping approaches may improve the mapping, counting and  
652 inclusion of small ponds in freshwater inventories (Bie et al. 2020, Rosentreter et al. 2021,  
653 Hofmeister et al. 2022). As inundation regimes become more variable, increasing conservation  
654 and protection efforts for maintaining ephemeral and temporary ponds will become more  
655 essential.  
656



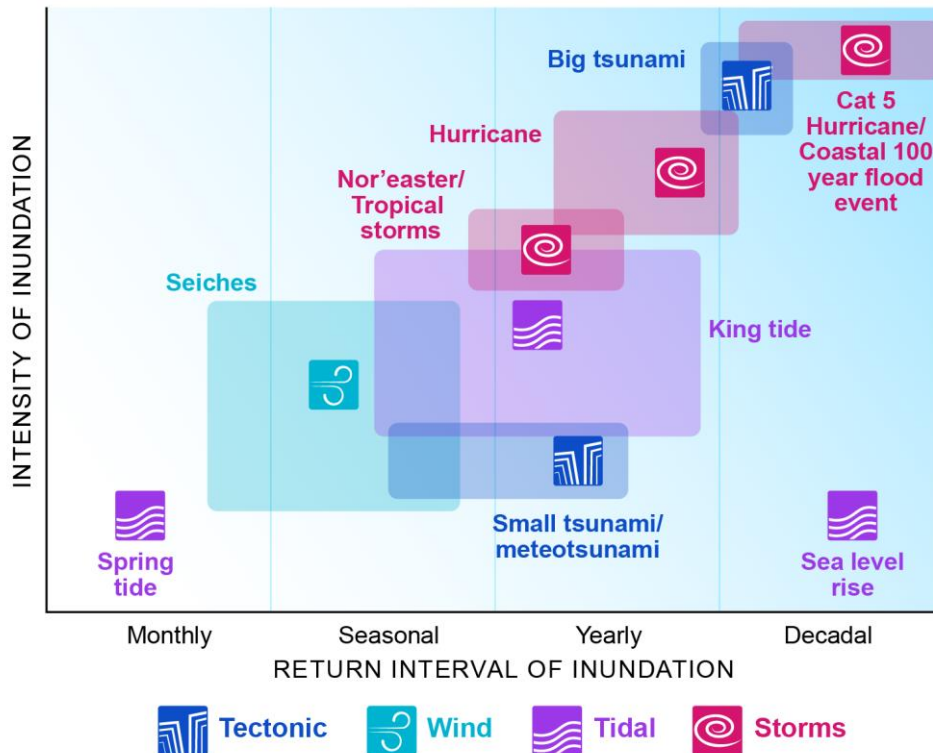
657  
658  
659  
660  
661  
662  
663  
664  
665  
666

**Figure 7. Examples of variable inundation across scales in pond systems.** Satellite imagery of the Prairie Pothole Region, North Dakota, USA illustrating decadal variable inundation at a landscape scale a) September 2, 1992; b) May 23, 2013 [modified from (Scientific Investigations Report 2015)] and at the pond scale; Aerial Imagery of Pond P1, Cottonwood Lake Study Area, North Dakota c) September, 2002 d) September, 1992 (Images from (U.S. Geological Survey 2017)). Seasonal changes in a vernal pond in Moshannon State Forest, Pennsylvania, USA inundated (May 11, 2023) non-inundated (May 23, 2023) (J.N. Sweetman). Conceptual drawings by Nathan Johnson.

## 667 Storm-Impacted Coastal Zones

668 The coastal zone includes ecosystems and communities (cities/towns) that are adjacent and  
669 hydrologically connected to a large water body (e.g., ocean, Great Lakes). These systems  
670 influence, are impacted by, and are dependent on coastal zone hydrologic processes, such as  
671 inundation, that occur at the interface between terrestrial and aquatic domains. Unlike tidal  
672 environments, inundation that affects the coastal zone is driven by temporary, often stochastic  
673 events including storms, seiches, and king tides. The impact and areal extent of coastal  
674 inundation varies across events, depending on topography, infrastructure, and event size (**Fig.**  
675 **8**). The frequency of these events ranges from multiple times a season to decadal (**Fig. 8**).  
676 Tropical storms and cyclones develop in tropical regions during seasonal periods of warm water  
677 each year. Due to their high energy and movement, they influence more temperate regions as  
678 well (Colbert and Soden 2012). In temperate or cold regions, storms develop in the winter time  
679 due to large temperature differences between land and ocean (Liberato et al. 2013). Natural  
680 systems will display some form of resilience and recovery to storm impacts (Lugo 2008, Wang

681 et al. 2016), but human settlements and infrastructure are vulnerable to both intense winds and  
 682 inundation (Lane et al. 2013, Hinkel et al. 2014, Braswell et al. 2022). Land use development  
 683 also alters the natural resilience of coastal environments through the proliferation of gray  
 684 infrastructure such as jetties and seawalls (Gittman et al. 2015). Systems in low-lying regions  
 685 are particularly vulnerable to inundation as opposed to rocky shores with steep slopes. While  
 686 regional or global data sets based on elevation data exist, the extent at any given time of storm  
 687 surges, king tides, and other high water episodes depend locally/regionally on where the event  
 688 hits, infrastructure, and topography of the area.  
 689



690  
 691 **Figure 8. Coastal VIEs experience inundation events with different frequencies and**  
 692 **intensities.** Some events occur rarely, but are very high intensity events (category 5 hurricanes;  
 693 large tsunamis), increasing the area of inundation and affecting areas that seldom experience  
 694 inundation. The impacted systems are often less adapted to inundation, increasing the extent of  
 695 destruction or reorganization of the system. Other events occur more regularly and/or are lower  
 696 in intensity (spring tide, seiches), leading to less extensive inundation and impacting coastal  
 697 systems that are more adapted to inundation. Credit: Nathan Johnson.  
 698

699 Inundation in the coastal zone impacts sediment transport, solute and nutrient mobilization,  
 700 vegetation distribution, biological diversity, and biogeochemical processes. Erosion and  
 701 sediment deposition alter ecosystem geomorphology (e.g., dune shape, marsh accretion)  
 702 (Houser and Hamilton 2009, Dissanayake et al. 2015) and ecosystem nutrient pools [e.g.,  
 703 (O'Mara et al. 2019, Castañeda-Moya et al. 2020)]. In coastal zones adjacent to marine and  
 704 estuarine waters, saltwater intrusion changes surface (Schaffer-Smith et al. 2020) and  
 705 groundwater (Cantelon et al. 2022) quality and mobilizes nutrients through porewater ionic

706 exchange processes (Herbert et al. 2018). Coastal zone inundation as a natural process alters  
707 dune systems, which generates a mosaic of habitats that increase biodiversity (Smith et al.  
708 2021) and alter distributions of vegetation and animals. For example, the frequency of overwash  
709 events affects plant composition and diversity on sand dunes (Stallins and Parker 2003) and  
710 regular inundation is thought to provide necessary habitats for some insects and birds (Smith et  
711 al. 2021). Increased salinity and associated geochemical changes alter microbial community  
712 diversity and population heterogeneity (Nelson et al. 2015), shifting to more specialized  
713 communities as an adaptation to anaerobic conditions, redox fluctuation, and salt stress.  
714 Previous studies found high variability in relationships between salinity and ecosystem carbon  
715 dioxide fluxes (Morrissey and Franklin 2015, van Dijk et al. 2015, Dang et al. 2019, Hopple et al.  
716 2022).

717 Human communities within the coastal zone are impacted by inundation events as well.  
718 Inundation of coastal agricultural lands from storm surge and sea level rise reduces agricultural  
719 productivity (Lei et al. 2016). In particular, risk is high to coastal zone communities in developing  
720 nations, where inundation events can lead to food insecurity, loss of livelihood, and increased  
721 transmission of waterborne diseases. As climate change alters the magnitude and frequency of  
722 inundation in the coastal zone, it will be necessary to integrate both natural and human  
723 adaptations, such as enabling salt marsh transgression (marsh migration upland) to mitigate  
724 storm surge impacts on crop yield (Guimond and Michael 2021).

725 While we understand many of the linkages between the ecology, biogeochemistry,  
726 hydrology, and geomorphology that regulate ecosystem structure and function in coastal  
727 systems (Fagherazzi et al. 2012, Hinshaw et al. 2017, Braswell and Heffernan 2019, Cantelon  
728 et al. 2022), we know little of how to predict the future effects of the interacting stressors  
729 associated with climate change (O'Meara et al. 2017, Ward et al. 2020, Arrigo et al. 2020). Our  
730 ability to predict is reliant on our understanding of shifting inundation regimes in the context of  
731 elevated CO<sub>2</sub>, nutrient pollution, and coastal development which can generate antagonistic,  
732 synergistic, or additive effects. These knowledge gaps stem from the dynamic and  
733 unpredictable nature of events that drive coastal inundation. Observational data to inform  
734 mechanistic models is limited and governed by where and when events happen (not necessarily  
735 within monitored sites), funding periods, and accessible coastlines. This difficulty is exacerbated  
736 by the fact that 40% of the world's population lives within 100 km of the coast (Maul and Duedall  
737 2019), which heightens social impacts of variable inundation while also adding logistical  
738 difficulty to coastal monitoring. When events do overlap with instrumented sites, the extreme  
739 nature of inundation events threaten the physical integrity of instrumentation. Lastly, high-  
740 latitude coastlines are also susceptible to coastal inundation, yet few models incorporate  
741 physical, biogeochemical, and ecological implications of inundation on permafrost bound  
742 coastlines and environments (Ekici et al. 2019, Bevacqua et al. 2020). Opportunities of critical  
743 knowledge advancement exist in 1) monitoring events through *in-situ* or remotely sensed  
744 monitoring data, 2) model development that integrates more robust process-based  
745 understanding, and 3) expansion into urban and permafrost-bound coastlines.

## 746 **Tidally Driven Coastal Zones**

747 Tidally-influenced coastal zones exist at the intersection of terrestrial and marine environments  
748 and encompass diverse intertidal ecosystems such as marshes, mangroves, ghost forests, and

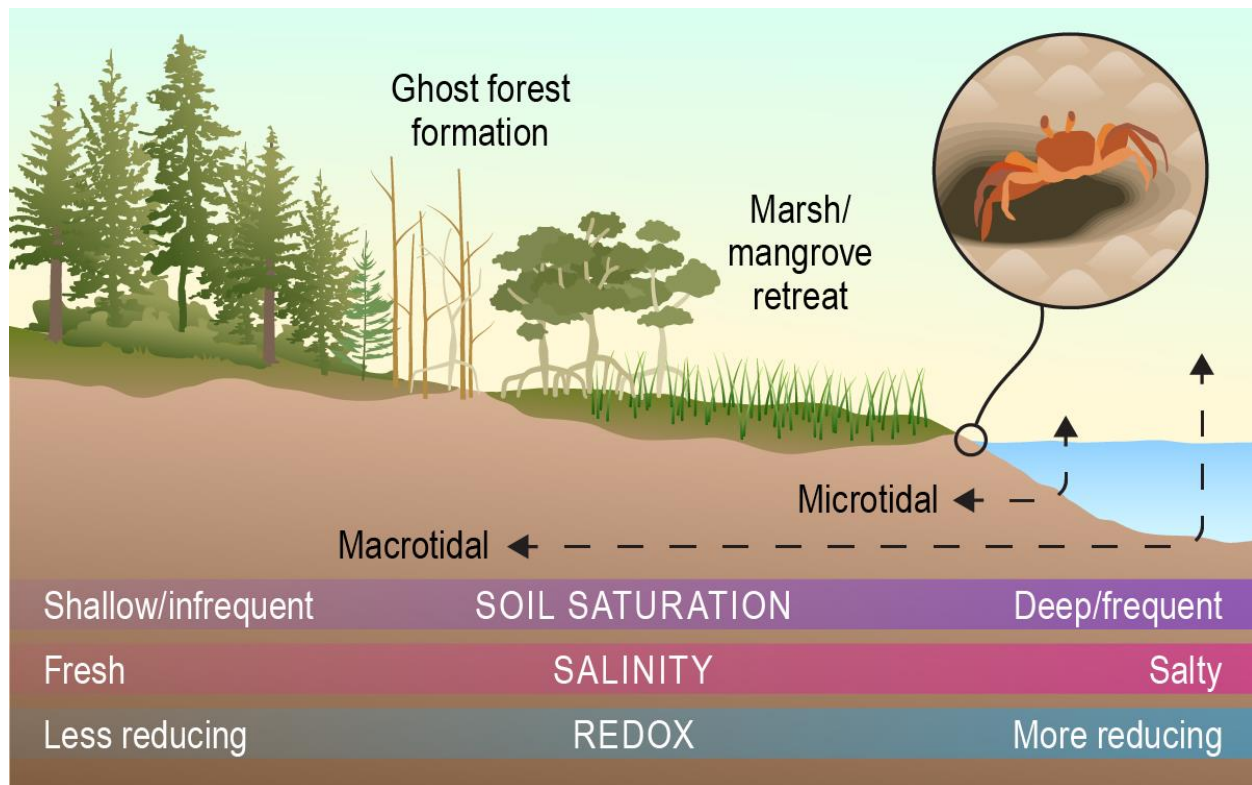
749 beaches (**Fig. 9**). Globally, tidal wetlands exist on 6 of 7 continents, and are spread across  
750 tropical, temperate, and polar latitudes (Murray et al. 2022a). Tidal flats are predominantly found  
751 along low sloping coastlines with approximately 70% of global tidal flat area existing in Asia,  
752 North America, and South America (Murray et al. 2022b), while beaches encompass 31% of  
753 ice-free shorelines (Luijendijk et al. 2018).

754 Tidally-driven coastal zones are inundated semi-diurnally (i.e., twice a day) or diurnally (i.e.,  
755 once a day). Unlike VIE systems summarized above, where inundation events may be difficult  
756 to predict, inundation in tidally-driven coastal zones varies primarily based on predictable  
757 drivers. For example, high tide and low tide water levels dictate the spatial extent and duration  
758 of inundation. In addition, intra-annual tidal dynamics are largely controlled by lunar cycles  
759 which drive approximately monthly highest (spring) and lowest (neap) tides, as well as annual  
760 high (king) and low tides. Inter-annual tidal dynamics are linked to sea level rise, which is  
761 shifting the zone of variable inundation inland (Ensign and Noe 2018, Tagestad et al. 2021). We  
762 note that while the timing of king tides is predictable (perigean spring tide), their impacts can be  
763 difficult to predict, as mentioned in the storm-impacted coastal zones section. In addition,  
764 topography (e.g., slope) and other natural physical factors, including wind speed and direction,  
765 waves, and even localized high and low pressure events mediate the lateral extent of surface  
766 water inundation in tidal ecosystems. Human modifications further alter both vertical and  
767 longitudinal extent of tidal inundation via control structures which may exclude tides (gates,  
768 weirs, etc.) and channels that transport tidal waters well inland of the natural intertidal zone.

769 The extent of tidal influence, which spans microtidal (< 2 meter tidal range) to macrotidal (>  
770 10 meter tidal range in some locations), controls water quality, terrestrial-aquatic interactions  
771 and resulting biogeochemical and ecological responses [e.g., (Tweedley 2016)]. Estuaries,  
772 where tides mix saltwater and freshwater, are dynamic biogeochemical mixing zones  
773 characterized by sharp chemical gradients that regulate biological activity [e.g., (Crump et al.  
774 2017)]. Shifts in tidal zones associated with sea-level rise are predicted to alter the extent of key  
775 intertidal habitats, with potential disruptions to coastal food webs (Rullens et al. 2022). Changes  
776 in duration and extent of inundation associated with tides control soil saturation and salinity,  
777 which influence redox dynamics, and hydrologically driven exchange of carbon, nutrients, and  
778 pollutants (Pezeshki and DeLaune 2012, Bogard et al. 2020, Regier et al. 2021). Biological  
779 activity, including crab burrows that alter hydrologic flow paths (Crotty et al. 2020), also  
780 influence tidal exchanges across the coastal terrestrial-aquatic interface (Crotty et al. 2020).  
781 Increased saltwater exposure due to shifting tidal ranges can alter the stability of coastal soils  
782 [e.g., (Chambers et al. 2019)], which represent a globally important carbon sink (McLeod et al.  
783 2011). In addition, tidal regimes structure vegetation gradients, where salt-sensitive  
784 communities including low-lying forests and freshwater marsh species are replaced by salt-  
785 tolerant communities including mangroves and saltmarsh species (Kirwan and Gedan 2019,  
786 Lovelock and Reef 2020). This shift in tidal range leads to the creation of ghost forests (Kirwan  
787 and Gedan 2019), which can impact coastal biogeochemical cycles [e.g., (Cawley et al. 2014)].  
788 Similarly, sea level rise may lead to mangrove or marsh retreat as inundation patterns change  
789 (Xie et al. 2020).

790 Due to the frequency of inundation, tidally inundated ecosystems are hydrologically,  
791 biogeochemically, and geomorphologically dynamic, creating challenges for scientists and land  
792 managers seeking accurate estimations of land surface area, elevation, and carbon storage.

793 These challenges are exacerbated by sea level rise, which exerts heterogeneous and non-linear  
 794 influences on tidal ranges (Du et al. 2018). Methodological approaches to assess tidal  
 795 ecosystem area and elevation that are based on satellite imagery will be critical for present and  
 796 future management and decision making. Similarly, complex feedbacks exist among hydrology,  
 797 biogeochemistry, ecology, and geomorphology (Xin et al. 2022); these dynamics may need to  
 798 be considered in future ecosystem projections. Thus, a deeper understanding of feedbacks and  
 799 their variability in space and time in response to tidal activity is needed (Ward et al. 2020).  
 800 Lastly, with sea-level rise, tidal constituents may change, with nonlinear impacts on tidal range  
 801 and inundation extent (Pickering et al. 2017). Tidally inundated VIEs represent the interface  
 802 between marine and terrestrial ecosystems, and to predict their future will require understanding  
 803 bi-directional connections among physical, chemical, and biological system components.  
 804



805  
 806 **Figure 9. Conceptual model of variable inundation in tidal systems.** Tidally driven coastal  
 807 zones span sediments exposed at low tide to marshes and coastal forests inundated at high  
 808 tide. This lateral gradient of tidal exposure across micro to macro-tidal systems (dotted black  
 809 lines) alters physical (e.g., particle deposition), biological (e.g., species composition), and  
 810 chemical (e.g., nutrient transformations) factors. Organisms can impact conditions along the  
 811 gradient, such as flow path alteration by crab burrowing. Credit: Nathan Johnson.

## 812 Human-Engineered Systems

813 Human-engineered systems are environments where inundation magnitude, frequency, timing,  
 814 and duration are either actively managed or have been dramatically altered by structural  
 815 modifications to the landscape (Fig. 10). Human-engineered VIEs rival natural systems in area



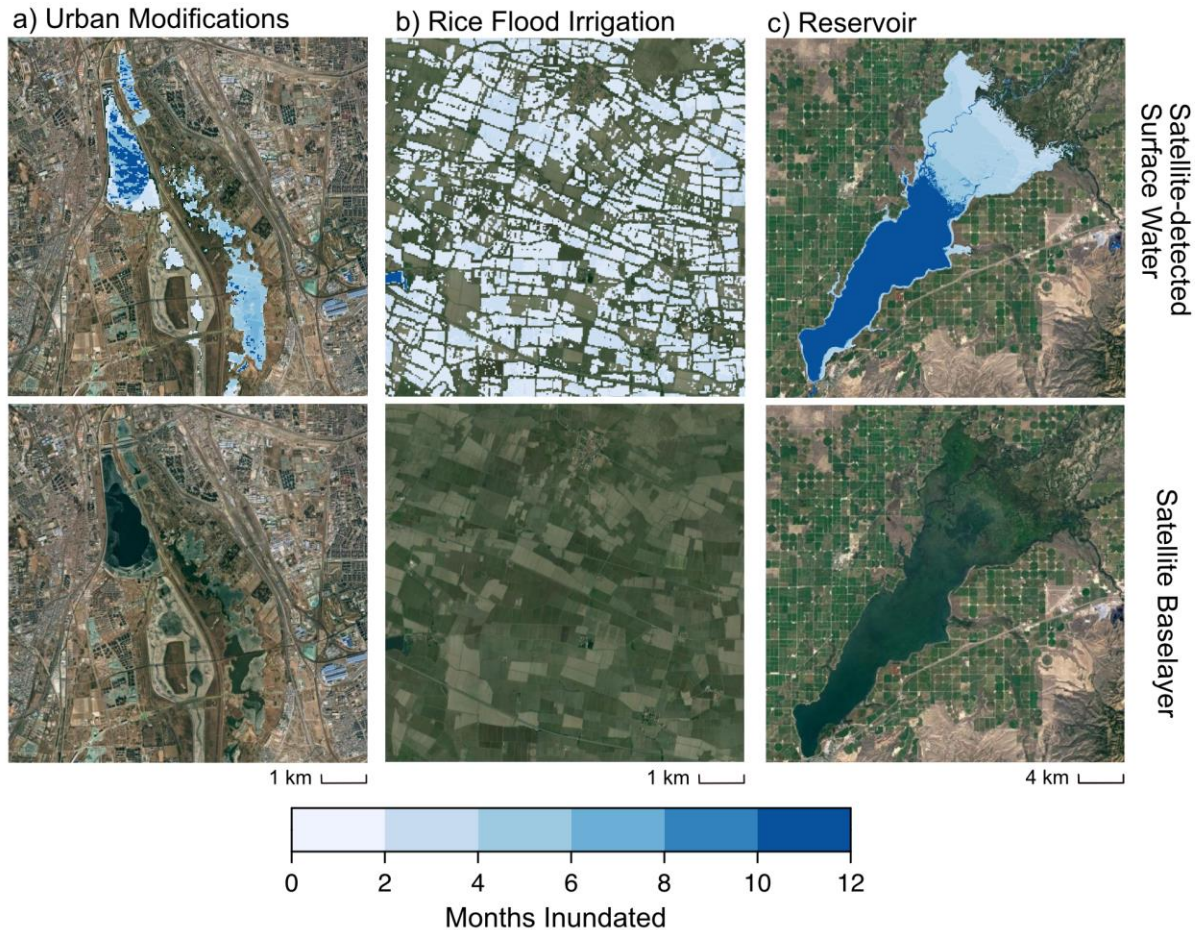
816 and extent (Clifford and Heffernan 2018), yet the significance of engineered VIEs in influencing  
817 landscape processes is relatively unexplored compared to natural systems (Koschorreck et al.  
818 2020) and they are historically excluded from water and nutrient budgets (Abbott et al. 2019).  
819 The primary drivers of human-engineered VIE formation explored here are land use change and  
820 restoration (including those for nature-based solutions), though hydrologic modifications impact  
821 inundation regimes of the natural VIEs explored earlier in the manuscript. Examples of land-use  
822 driven human-engineered VIEs include, but are not limited to: croplands irrigated to the point of  
823 inundation (e.g., rice paddies, cranberry bogs), canals for irrigation, drainage and stormwater  
824 (e.g., roadside ditches, retention ponds), and unintentional VIE formation following landscape  
825 modification (e.g., “accidental” urban wetlands (Palta et al. 2017) and ponds in agricultural fields  
826 (Saadat et al. 2020). Whereas the purpose of land-use driven engineered VIEs is to redistribute  
827 water for human purposes, the goal of VIEs engineered for restoration is to either replace or  
828 enhance ecosystems lost or damaged as a result of human activity. VIE restoration efforts vary  
829 in scope and form, spanning local (e.g., residential living shorelines, individual stream reaches,  
830 agricultural ditch wetlands) to ecosystem (e.g., adding sediment to degrading marshes), to  
831 regional (e.g., dam removal) scales (VanZomeren et al. 2018, Baptist et al. 2021).

832 While the full extent of human-engineered VIEs is difficult to quantify, key examples highlight  
833 their significance in the landscape. Agriculture covers nearly 40% of the earth’s land surface  
834 (Siebert et al. 2010), and nearly a quarter of that is variably inundated by flood irrigation (Wu et  
835 al. 2023). In urban systems, the extent of stormwater control networks rival those of natural  
836 systems. For example, the total linear length of residential canals in North America nearly  
837 equals that of the Mississippi River (Waltham and Connolly 2011). While restoration efforts are  
838 not as widely distributed as land-use change, restoration still contributes to extensive VIE  
839 creation. For example, restoration accounts for 14% of areal gain of tidal wetlands globally  
840 (Murray et al. 2022b). Inundation regimes in human-engineered VIEs can be driven by natural  
841 hydrologic processes, such as connectivity with the water table or tidal inputs. This is  
842 particularly important in VIEs built for restoration, as establishing natural inundation regimes  
843 enhances landscape connectivity and mediates ecosystem functions (Reis et al. 2017, Jones et  
844 al. 2018). However, unlike the previously discussed natural systems, the drivers and duration of  
845 inundation in human-engineered VIEs may be decoupled from natural hydrologic processes.  
846 Controlling drainage, such as for stormwater management, land reclamation, or effluent  
847 releases, is a key motivation for VIE construction and system design, resulting in inundation  
848 periods largely driven by precipitation that persist at event to seasonal scales depending on  
849 local hydrology and climate. Inundation duration may also occur on longer timescales, such as  
850 seasonally in paddy systems (De Vries et al. 2010). Finally, direct human interventions, such as  
851 floodgates, weirs, and dams, may affect water residence time at timescales that are  
852 asynchronous from natural drivers, such as seasonality or tides.

853 Human-engineered VIEs fundamentally alter the landscape, changing the spatial and  
854 temporal patterns of ecosystem processes. Agricultural inundation, such as flood irrigation or  
855 ponding, alters redox conditions, greenhouse gas emissions, groundwater recharge,  
856 evapotranspiration fluxes, plant growth, and pollutant export to natural water bodies (Hale et al.  
857 2015, Pan et al. 2017, Pool et al. 2021, Buszka and Reeves 2021). For example, a recent study  
858 showed that variably inundated depressions in agricultural fields can account for ~30% of  
859 nitrous oxide emissions across cultivated areas despite comprising ~1% of the land surface

860 (Elberling et al. 2023). The creation of drainage canals increases waterborne carbon fluxes from  
861 VIEs by producing a newly decomposed stock of labile soil carbon to be leached as well as by  
862 increasing the hydrological runoff rate through the soil and receiving canals and ditches (Stanley  
863 et al. 2012). Human-engineered VIEs can also provide ecosystem services that supplement or  
864 replace those of natural VIEs in the landscape (Clifford and Heffernan 2018). For example, they  
865 can enhance habitat (Connolly 2003, Herzon and Helenius 2008), nitrogen removal (Bettez and  
866 Groffman 2012, Reisinger et al. 2016), and recreation (Beckingham et al. 2019). Further, the  
867 services these systems provide can be improved through targeted management [e.g.,  
868 vegetation composition; (Castaldelli et al. 2015)] or restoration practices [i.e., two-stage ditches;  
869 (Speir et al. 2020)].

870 Including human-engineered systems in our conceptualization of VIEs emphasizes the  
871 growing significance of these systems as human landscape modifications continue to alter and  
872 eliminate natural VIEs. Recent efforts have synthesized the role and impacts of human-  
873 engineered VIEs at large scales (Peacock et al. 2021, Li et al. 2022b) but, as with many natural  
874 systems, the majority of studies on human-engineered VIEs are based in North America and  
875 Europe (González et al. 2015, Zhang et al. 2018, Bertolini and da Mosto 2021). Thus, our  
876 knowledge may not reflect the social, political, and economic challenges of developing areas  
877 where the highest rates of VIE modification are occurring (Wantzen et al. 2019). The knowledge  
878 gaps surrounding human-engineered VIEs will become increasingly important to address as  
879 global change continues to alter the spatial and temporal patterns of inundation. Given that  
880 human-engineered VIEs can enhance or disrupt hydrologic connectivity, they potentially  
881 magnify the effects of human driven changes such as sea level rise and impacts of  
882 contamination from anthropogenic “chemical cocktails” (Kaushal et al. 2022). We lack a  
883 baseline standard for how human-engineered VIEs function in the landscape, even as global  
884 change continues to shift existing baselines [e.g., (Palmer et al. 2014)]. A baseline  
885 understanding would also enable the restoration and repurposing of engineered VIEs as nature-  
886 based solutions (Clifford and Heffernan 2023)(Clifford et al., 2023). Addressing these  
887 knowledge gaps will require the incorporation of human-engineered VIEs into large-scale  
888 synthesis and modeling efforts, particularly those that address hydrologic and biogeochemical  
889 fluxes. Conclusive definitions and inventories of human-engineered VIEs is essential for  
890 estimating their ecological and biogeochemical roles at the global scale. Finally, human-  
891 engineered VIEs need to be conceptualized within an ecological, rather than managerial,  
892 context for comparison with natural systems and to be integrated into a more continuum-based  
893 approach for VIE science. Human-engineered VIEs rival the range of natural VIEs in structure,  
894 inundation regime, and global distribution. Understanding their role in the Earth system is,  
895 therefore, critical for understanding both the impacts of and potential solutions to global change.  
896



897  
898  
899  
900  
901  
902  
903  
904  
905

**Figure 10. Examples of human-engineered Variably Inundated Ecosystems.** a) Yongding River in Beijing, China; b) Paddy rice fields in northern Italy; c) American Falls Reservoir on the Snake River in Idaho, United States. These three examples emphasize significant variation in the degree of variable inundation across human-engineered VIEs, with some regions being perennially inundated. Top row: Satellite-derived map data on months inundated is derived from the “seasonality” product in the Global Surface Water Mapping Layers v1.4 (Pekel et al. 2016). Credit: Jillian Deines.

906 **Inundation Processes are Relevant at the Scale of the Beholder**

907 VIEs span broad spatiotemporal scales of variable inundation, from small wetlands and vernal  
908 ponds to the floodplains of the world’s largest rivers. While examples in the mini-reviews focus  
909 on eight different ecosystems, variably inundated ecosystems are even broader such as  
910 mosses and pore spaces that are periodically covered by droplets of water and vast endorheic  
911 lakes and rivers. Inundation volumes and surface areas of VIEs vary by at least sixteen orders  
912 of magnitude, from under  $10^{-3}$  L to over  $10^{13}$  L (Bonython and Mason 1953), and  $10^{-6}$  m<sup>2</sup> to over  
913  $10^{10}$  m<sup>2</sup> (Hess et al. 2015), respectively. The duration of inundation varies by up to eight orders  
914 of magnitude, spanning a few seconds, in the case of droplets, to decades, in the case of  
915 endorheic lakes, and centuries in the case of sea level rise. Non-inundated periods likewise

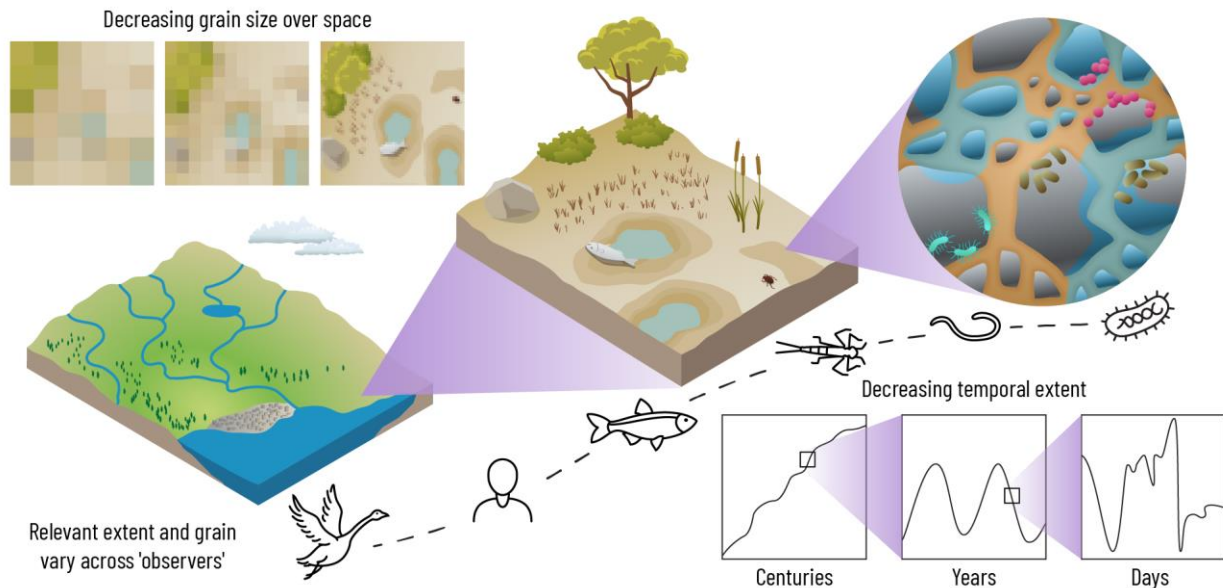
916 span seconds to centuries and longer. This variability in spatial and temporal extent has  
917 profound consequences for the ecology and biogeochemistry of VIEs. This section highlights  
918 the importance of considering scale and explores hypotheses regarding how scale drives  
919 variability in drivers, processes, and impacts across VIEs and how we study them.

920 Spatial and temporal scales of VIEs can be categorized along two axes – extent and  
921 granularity. Extent comprises the total size of the spatial domain or time duration of a defined  
922 system, while granularity pertains to the spatial or temporal intervals of system transitions  
923 (Ladau and Eløe-Fadrosch 2019). For example, the dynamics of water droplets across North  
924 America would represent a large extent with fine granularity, relative to the inundation dynamics  
925 of a several square meter desert playa (smaller extent but coarser grain). The impacts of  
926 variable inundation are dependent ‘on the scale of the beholder’ relative to the extent and grain  
927 of variable inundation, where a ‘beholder’ may be a molecule, organism, population, community,  
928 land manager, or otherwise (**Fig. 11**). The expressed metabolism of an individual microbe will  
929 be influenced by inundation down to the spatial scale of water films and on hourly or shorter  
930 time scales. An individual microbe may not, however, be influenced by whether variable  
931 inundation occurs only within a square meter or across many square kilometers because it does  
932 not perceive these larger scales. In contrast, macroinvertebrate behavior is influenced by  
933 variable inundation down to scales of meters and days, and is likely further influenced by larger  
934 and longer scales of stream network connectivity (Bogan et al. 2017b, Sarremejane et al. 2017).

935 VIEs can be viewed as habitat patches of different sizes that vary in how long they persist in  
936 a given state and that have dynamic connectivity among patches. Terrestrial and aquatic biota  
937 respond on ecological and evolutionary time scales to the expansion and contraction cycles of  
938 inundation (Bornette et al. 1998, Ward et al. 2002). Biotic diversity is influenced by productivity,  
939 connectivity, disturbance severity and disturbance frequency, all of which operate at hierarchical  
940 scales (Ward et al. 1999). Biogeographical and ecological theories posit that patch size (e.g.,  
941 species area scaling) and disturbance regimes (e.g., intermediate disturbance hypothesis) are  
942 strong determinants of community composition (Adler et al. 2005, Svensson et al. 2012),  
943 suggesting that VIE community composition may vary predictably with these factors. The  
944 duration, predictability, and frequency of inundation likely have consistent community-level  
945 consequences that vary predictably with VIE extent and grain. Different extents and grains of  
946 inundation have the potential to change habitat connectivity in addition to directly selecting for  
947 different groups of organisms. Isolated marshes may, for example, become merged during a  
948 flood, thereby enhancing dispersal of aquatic organisms. The scale of variable inundation has  
949 numerous influences over ecological processes and dynamics that need to be understood.

950 From a biogeochemical perspective, variable inundation generates spatial and temporal  
951 variation in rates and patterns of biogeochemical processes. This variability is important for  
952 scaling biogeochemical rates because of process nonlinearity and Jensen’s inequality (Ruel and  
953 Ayres 1999). That is, a rate based on average conditions differs systematically from the average  
954 rate across variable conditions. This is important because the scales of processes (e.g.,  
955 microbial activity occurring within pore channels) are typically not aligned with the scales of  
956 measurements and models (e.g., core-scale or above). The lack of clear understanding for how  
957 variable inundation influences variation in biogeochemical processes and how these  
958 relationships change with extent and grain of inundation can, therefore, lead to unreliable  
959 predictions for the scaling of biogeochemical processes.

960 Understanding the biogeochemical influences of variable inundation across a broad range of  
 961 scales is important for informing a diverse suite of needs across models, decision makers, and  
 962 other interested parties. Our ability to inform these needs depends on our ability to rigorously  
 963 understand and predict influences of variable inundation across scales. This is a challenge as  
 964 variable inundation likely has direct, but unknown, influences over the scaling of biogeochemical  
 965 function. For example, cumulative metabolism in streams is predicted to increase faster than  
 966 their upstream drainage area for perennial stream networks (Wollheim et al. 2022). The  
 967 influence of variable inundation on biogeochemical processes cannot yet be accounted for in  
 968 such scaling theory. More generally, perturbations like variable inundation can drive systems  
 969 away from steady-state assumptions from which scaling relationships are derived (McCarthy et  
 970 al. 2019), therefore, we expect significant changes in scaling behavior across inundation  
 971 regimes. A research frontier is to quantify the direction, magnitude, and duration of changes in  
 972 scaling patterns in response to variable inundation and to modified variable inundation regimes  
 973 wrought by climate, land-use, and other environmental changes.  
 974



975 **Figure 11. Variable inundation can be observed at different spatiotemporal granularities**  
 976 **and extents.** (Upper left) Granularity is based on the resolution of observations in space or  
 977 time. (Lower right) Extent is based on the cumulative breadth of observations in space or time.  
 978 (Middle panels) Granularity and extent of observations are often correlated, such as barely  
 979 resolving individual trees when extent spans a watershed and resolving individual microbes  
 980 when extent spans a few soil particles. A given beholder observes variable inundation at a given  
 981 scale and will, in turn, make changes to behavior, physiology, and/or aspects of life history. For  
 982 example, migratory waterfowl select habitats based on inundation state as they move across  
 983 watersheds, humans plan cities based on regional patterns, fish move across stream reaches  
 984 based on continuity of inundation, macroinvertebrates lay eggs on individual rocks based on  
 985 inundation state, nematodes experience variable inundation as they move through porous  
 986 media, and soil microbes separated by microns likely experience vastly different inundation  
 987 dynamics linked to water films on soil particles.  
 988

## 989 **Summary of Primary Methods used to Study VIEs**

990 The multi-scale nature of VIE systems has led to experimental and observational studies that  
991 span from point-scale lab-based characterization, to reach- or watershed-scale monitoring  
992 networks, and to regional- and global-scale remote sensing. Point-scale measurements at the  
993 smallest scales help reveal processes that underlie larger scale dynamics. For example, point  
994 measures of water presence, water absence, and low flow detection within a watershed are  
995 increasingly available with the development of small, inexpensive, and easily deployable  
996 sensors, meters, and time-lapse cameras [e.g., (Soupir et al. 2009, Chapin et al. 2014, Costigan  
997 et al. 2017, Zimmer et al. 2020)] (**Fig. 12**). While these measurements are easy to take and can  
998 provide a long temporal dataset for little effort, they are not always detailed and require regular  
999 calibrations.

1000 A broad range of methods can be used to link the hydrologic dynamics to ecological and  
1001 biogeochemical responses. Standardized field surveys and biomolecular methods (e.g., isotopic  
1002 ratios, including compound specific analyses) are commonly used to study organismal,  
1003 population, and community ecology across multiple taxa [e.g., (Ode et al. 2016, Gates et al.  
1004 2020)] and can be standardized for both inundated and non-inundated states. There is  
1005 increasing use of crowdsourcing for biogeochemical characterization to consistently obtain  
1006 samples across diverse systems (von Schiller et al. 2019, Garayburu-Caruso et al. 2020).  
1007 Sample collection can be followed by a variety of laboratory measurements of properties (e.g.,  
1008 carbon content, redox potential and redox-active elements, microbial genetic potential, sediment  
1009 grain size) and processes, such as CO<sub>2</sub> production and methanogenesis related to variable  
1010 inundation. Point-scale measurements often operate at instantaneous to daily scales.  
1011 Conversely, larger scale measurements integrate across finer-scale processes to quantify  
1012 ecosystem dynamics and properties, but without necessarily revealing what governs those  
1013 processes. Spatially distributed monitoring networks using *in situ* sensors (e.g., the United  
1014 States Geological Survey, USGS, gage network) can connect event-scale responses across  
1015 hydrologically linked locations as well as reveal long-term trends [e.g., (Zipper et al. 2021)].  
1016 Long-term field manipulations are another complementary *in situ* technique that can reveal  
1017 mechanisms underlying system responses to changes in inundation state. There are numerous  
1018 configurations of such experiments that directly or indirectly impact inundation dynamics, such  
1019 as intentional inundation (Hopple et al. 2023), water exclusion (Kundel et al. 2018) and heating  
1020 (Hanson et al. 2017). Despite the plethora of data produced by such large scale projects, these  
1021 are expensive and require deep buy-in of researchers and landowners.

1022 Remote sensing can complement *in situ* measurements to facilitate more spatially  
1023 continuous characterization of surface water dynamics and their impacts. There are different  
1024 types of remote sensing techniques, from drones to satellites and optical to microwave sensors,  
1025 that can capture different aspects of VIEs. For example, soil surface saturation may be captured  
1026 by a passive microwave radiometer as well as C and L-band radar backscatter, which can also  
1027 penetrate through thin canopies, clouds, and through the top few centimeters of the soil  
1028 (Schumann and Moller 2015). Recent satellite missions such as the Surface Water and Ocean  
1029 Topography (SWOT) mission provide increased capabilities for monitoring changes in surface  
1030 water over time with radar data (Biancamaria et al. 2016), while NASA's forthcoming NISAR  
1031 mission will allow for detection of inundation even under tree canopy. Thermal infrared  
1032 measurements can indirectly reveal saturation at very high spatiotemporal resolutions, as well

1033 as evapotranspiration associated with water table depth, soil moisture, and rooting depth (Fisher  
1034 et al. 2020, Lalli et al. 2022). Long time series from moderate resolution (~30 m) optical  
1035 satellites can document multi-decadal open water trends and seasonal regimes across the  
1036 globe (Pekel et al. 2016), while some combinations of indices have shown success in detecting  
1037 mixed vegetation and inundation cover (Jones 2019). Commercial satellite constellations  
1038 provide daily global imagery at < 4 m resolution, enabling monitoring of more dynamic water  
1039 bodies [e.g., Arctic lakes, (Cooley et al. 2017) and forested wetlands (Hondula et al. 2021a)].  
1040 Deep groundwater and changes in the total water column storage are detectable through  
1041 measurements of gravitational anomalies at very high precision but low spatial resolution  
1042 (Bloom et al. 2010, 2017, Richey et al. 2015, Pascolini-Campbell et al. 2021). Fine-scale  
1043 inundation dynamics, which have been historically hard to measure, can be captured using  
1044 unmanned aerial vehicles (UAVs), which are often useful during or immediately after a  
1045 significant inundation event (Perks et al. 2016), to capture small-scale spatial dynamics that are  
1046 difficult to detect with satellite or airborne methods (Manfreda et al. 2018, Dugdale et al. 2022),  
1047 or to derive detailed data for input into hydrologic models and surface water calculations  
1048 (Acharya et al. 2021).

1049



1050

1051 **Figure 12. Monitoring inundation regimes is increasingly possible via in situ sensors.**  
1052 *Stream Temperature, Intermittency, and Conductivity Sensors (STICs) (Chapin et al. 2014), one*  
1053 *of the types of increasingly available sensors to measure water presence/absence in an*  
1054 *inexpensive and easily deployable manner. These sensors can be used across all types of VIEs.*  
1055 *Credit: Amy Burgin.*

1056

1057 To advance predictive understanding requires integration of data with models. Process-  
1058 based models can be used to simulate hydrological and biogeochemical processes under dry  
1059 and wet conditions (Fatichi et al. 2016, Li et al. 2017). These models are often built upon mass  
1060 conservation principles, with ordinary or partial differential equations that describe coupled  
1061 hydrological, ecological, and biogeochemical processes. They rely on existing knowledge on  
1062 processes, including, for example, theories or empirical relationships between discharge and  
1063 water storage (Wittenberg 1999), biogeochemical reaction rate dependence on temperature and  
1064 water content (Davidson et al. 1998, Mahecha et al. 2010) and redox reactions (Borch et al.  
1065 2010). Among process-based models, there are spatially distributed models that couple surface  
1066 and subsurface flow dynamics explicitly (Kollet and Maxwell 2006, Coon et al. 2020). This class

1067 of models has recently been extended to include reactive transport (Wu et al. 2021), which may  
1068 be considered as a set of tools to understand the biogeochemical effects of variable inundation  
1069 (Molins et al. 2022). However, spatial resolution and data requirements for the integrated  
1070 surface and subsurface models are high, which places practical limits on the spatial scales that  
1071 can be addressed. Semi- or fully-distributed models with coarse spatial resolution are able to  
1072 work at larger scales, but require theories or empirical relationships to represent processes and  
1073 impacts at subgrid-scales. Data-driven machine learning methods present new opportunities to  
1074 blend models with various levels of mechanistic representations into hybrid models (Reichstein  
1075 et al. 2019). Increases in the volume of observational data sets combined with advances in high  
1076 performance computing have triggered a shift towards machine learning applications for  
1077 capturing inundation dynamics. More recently, integration of physics-based models with  
1078 machine learning have improved the interpretability of machine learning methods and increased  
1079 their ability to model complex ecosystem processes (Sun et al. 2022b). These hybrid  
1080 approaches have the potential to optimize the characterization and prediction of inundation  
1081 dynamics by incorporating the strengths of multiple models to achieve predictions with  
1082 minimized uncertainty and greater accuracy than either model alone.

1083 Coordinated integration (Patel et al. 2023) between model development and data generation  
1084 is key to deepening our understanding of VIEs and increasing our ability to predict their future  
1085 ecosystem function and ecological integrity. More specifically, we promote iterating between  
1086 model-guided data generation and observation-informed model development. This iterative  
1087 cycle between models and ‘experiments’ (i.e., real-world data generation) has previously been  
1088 termed ‘ModEx’ (Atchley et al. 2015) and is similar to approaches used in ‘ecological  
1089 forecasting’ (Dietze et al. 2017, 2018). It also aligns generally with the scientific method based  
1090 on continuous iteration between conjectures (hypotheses / models) and refutation (falsification  
1091 of hypothesis using observations and data) to drive scientific discovery and knowledge growth  
1092 (Popper 2014). The ModEx approach often starts by using experimental or field data to  
1093 parameterize and calibrate models and/or generate new data based on known model input  
1094 needs. This can be expanded whereby models generate hypotheses via *in silico* experiments,  
1095 and field or lab studies can be designed to test those hypotheses. Models can also be used to  
1096 optimize the design of real-world experiments by indicating when, where, and what to measure  
1097 to provide the strongest hypothesis evaluation.

1098 In the context of VIEs, we expect ModEx to touch scales ranging from molecular  
1099 microbiology to landscape ecology to regional ecosystem function to Earth system elemental  
1100 cycles. As a landscape-scale example of ModEx, physical models could first be used to predict  
1101 variable inundation across a watershed. Spatial and/or temporal uncertainty in those predictions  
1102 could then be used to optimize collection of commercial remote sensing data. Those data  
1103 would, in turn, be used to evaluate model predictions, leading to updated guidance from the  
1104 model on where/when to collect additional remote sensing data. Further cycles could be  
1105 pursued and model uncertainties could also guide collection of *in situ* data on variable  
1106 inundation, organismal ecology, and/or biogeochemical processes. Many other examples  
1107 across a variety of scales can be envisioned, and key to enabling this approach is the further  
1108 development of models and measurement techniques that can capture system states in both  
1109 inundated and non-inundated conditions. Techniques/models designed for specific kinds of  
1110 ecosystems (e.g., perennial rivers) may be difficult to adapt. This emphasizes a need to do



1111 ModEx using models and measurements intentionally designed to span inundated and non-  
1112 inundated system states.

1113 Across the continuum of ModEx, it is important to consider the scales at which models and  
1114 measurements operate, as discussed above. The issues around scale could, in part, be  
1115 addressed by Integrated Coordinated Open Networked (ICON) science principles (Goldman et  
1116 al. 2022). ICON is based on intentional design of research efforts to be Integrated across  
1117 disciplines and scales, Coordinated across research efforts via consistent methods, Open  
1118 throughout the research lifecycle, and Networked across stakeholders to understand collective  
1119 needs. We propose using ICON principles for *in situ* data generation and remote sensing, jointly  
1120 guided by model-generated predictions (i.e., ModEx). Embedding ICON throughout the research  
1121 life cycle can help to ensure that new data are at the right scale and can be used to link  
1122 disciplines (e.g., hydrology, biogeochemistry, and community ecology). This can also ensure  
1123 that data are interoperable across VIEs, are available to everyone and connected to deep  
1124 metadata, and are useful to a broad range of stakeholders with interests spanning different  
1125 types and locations of VIEs. The use of ICON in cross-VIE science could bridge existing data  
1126 across multiple spatial and temporal scales, and potentially bridge gaps among VIEs.

## 1127 **Towards Cross-VIE Transferable Understanding**

1128 We propose that a key goal for VIE science is the development and open sharing of knowledge,  
1129 models, algorithms, and data that transcend individual system types. Knowledge that crosses  
1130 VIE systems will inherently span scales and levels of certainty from predictable, sub-daily  
1131 inundation regimes to rare extreme events; integrating perspectives of these dynamic systems  
1132 can aid in understanding and anticipating tipping points of physical, chemical, and biological  
1133 components across VIEs. Development of such knowledge should be done via ModEx  
1134 approaches coupled with ICON principles, which can generate models that can be used across  
1135 VIEs. Similar to the perspectives of Arias-Real et al. (2024), we suggest this can be facilitated  
1136 through the development of conceptual models based on continuous environmental axes that  
1137 modulate system responses to re-inundation (e.g., greenhouse gas production and changes in  
1138 biological diversity).

1139 Such continuum-based conceptual models necessitate going beyond discrete VIE  
1140 categories by treating key physical characteristics as continuous variables that influence all VIE  
1141 systems. One realization of such a conceptual model is summarized in Figure 13. Related  
1142 approaches that are based on a suite of temporally variable ecological and geomorphological  
1143 characteristics have proven useful for wetlands (Euliss et al. 2004, Lisenby et al. 2019). These  
1144 wetlands frameworks have improved the understanding of human impacts on wetlands and led  
1145 to more effective management (Wierzbicki et al. 2020, Mandishona and Knight 2022). These  
1146 successes emphasize the potential effectiveness of continuum-based conceptual models for  
1147 cross-VIE science.

1148 The impacts of variable inundation depend on multiple characteristics of inundation regimes  
1149 (e.g., return interval and duration) and factors that influence those regimes (e.g., subsurface  
1150 permeability, topography, climate, and vegetation) (Banach et al. 2009, De Jager et al. 2012).  
1151 Furthermore, there are dynamic attributes such as water residence time and hydrologic  
1152 connectivity that influence process rates (Covino 2017). We hypothesize that despite this

1153 complexity, cross-VIE science can make progress towards transferable understanding through  
1154 the evaluation of conceptual models that focus on impacts of variable inundation across  
1155 relatively simple physical variables that can be easily measured. Two such variables are  
1156 inundation return interval and topographic slope (**Fig. 13**). As suggested above, we encourage  
1157 studies that examine responses to variable inundation (e.g., biogeochemical rates and  
1158 ecological community composition) across VIEs that collectively span a broad range of return  
1159 intervals and slopes.

1160 While many environmental variables could be used in this conceptual model (**Fig. 13**), here  
1161 we propose using inundation return interval and topographic slope, as both are well known to  
1162 impact ecological communities. For example, inundation return interval has been shown to alter  
1163 plant composition (Arim et al. 2023) and biogeochemical function such as CH<sub>4</sub> fluxes (Batson et  
1164 al. 2015). We view it as an integrated proxy for variables with direct impacts (e.g., desiccation)  
1165 that are linked to the temporal scale of non-inundated conditions. The other axis of our  
1166 conceptual model is topographic slope (**Fig. 13**), which we also view as an integrated proxy, but  
1167 for variables linked to how much time water spends in a system (Anderson and Burt 1978,  
1168 McGuire et al. 2005). Slope and the variables it represents (e.g., water residence time and  
1169 velocity) are also well known to influence ecological communities (e.g., by altering fish  
1170 composition, as in (Bain et al. 1988)) and biogeochemistry (e.g., by altering nitrate reductions as  
1171 in (Gomez et al. 2012)).

1172 At a high-level, return interval and slope are two key dimensions of temporal scale: how long  
1173 it takes water to return and how long a parcel of water spends in the system. Similarly, these  
1174 variables encompass differences across spatial scales, capturing differences in timing of  
1175 inundation and how water flows through and is connected to different components of VIEs (e.g.,  
1176 differences in drying across branches of a river network). While these two components should  
1177 jointly influence nearly every physical, chemical, and biological aspect of VIEs through time and  
1178 across space, we do not imply that these two variables will capture all relevant processes. Other  
1179 variables such as sediment/soil mineralogy and climate also have strong influences over  
1180 biogeochemistry and community ecology of VIEs (e.g., Pumo et al. 2016). We may learn that  
1181 additional axes are needed and these may be linked to other conceptual models, such as  
1182 whether inundation emerges through infiltration-excess (Hortonian flow generation) or through  
1183 saturation-excess (Dunnian flow generation) (Freeze 1974). Nonetheless, we propose that  
1184 significant progress can be made towards cross-VIE understanding of the controls over  
1185 biogeochemistry and ecology by further developing and testing the high-level conceptual model  
1186 proposed here linked to inundation return interval and topographic slope. In doing so, we  
1187 encourage careful attention towards the spatial and temporal scales of modeling and data  
1188 generation efforts linked to return interval and slope.

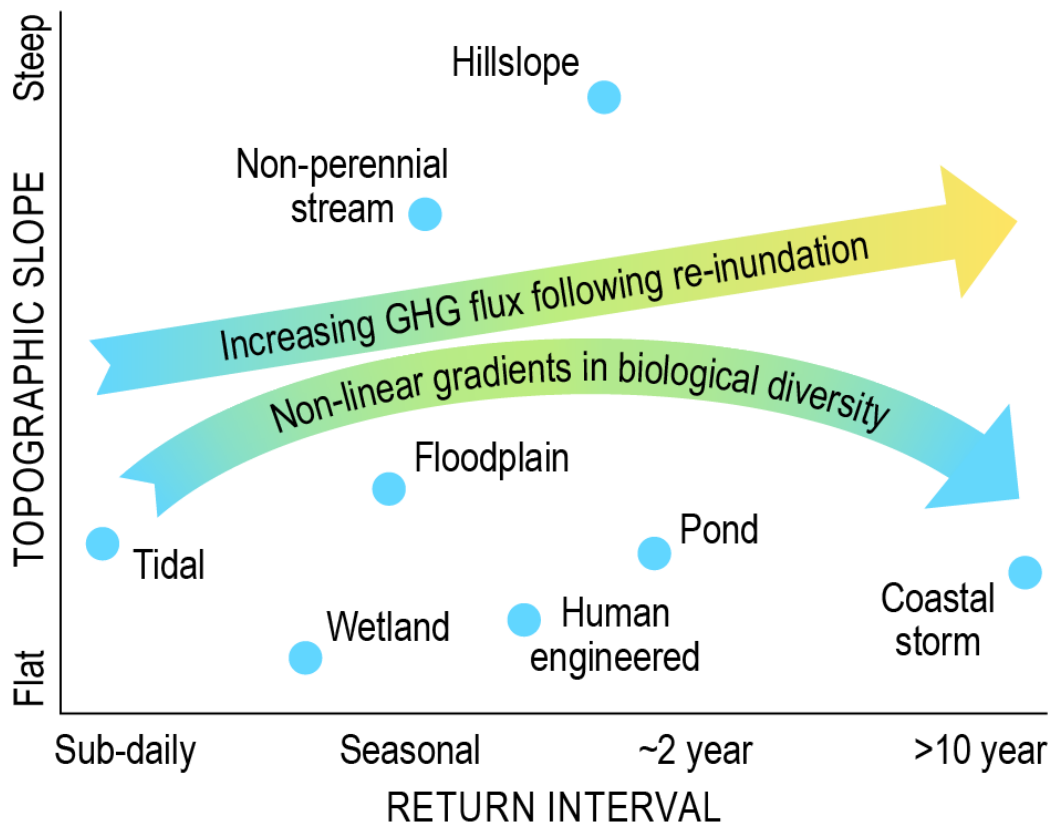
1189 Our conceptual model can be used to frame and study questions representing science  
1190 challenges that span all VIEs, such as how greenhouse gas fluxes and biological diversity  
1191 respond to variable inundation (**Fig. 13**). Similarly, metabolism research has suggested using a  
1192 continuum of flow predictability and light availability to better unify river metabolism research  
1193 (Bernhardt et al. 2022). In this approach there is no need to bin VIEs into discrete categories  
1194 (Euliss et al. 2004), many of which have varying definitions and levels of overlap. A given  
1195 system may also not fit clearly into a single VIE category and/or may transition across  
1196 categories through time and across space. Rather, we can observe and study continuous

1197 response surfaces across multiple physical axes and identify patterns within this quantitative  
1198 space.

1199 In addition to generating transferable understanding, bringing all VIEs together via studies  
1200 focused on unifying conceptual models could help raise awareness of VIE diversity, importance,  
1201 vulnerabilities, and how they may change in the future. This may, in turn, help address the fact  
1202 that VIEs are often overlooked in terms of conservation and monitoring efforts (Calhoun et al.  
1203 2017, Hill et al. 2018, Krabbenhoft et al. 2022, Zimmer et al. 2022). Studying diverse VIEs  
1204 across broad ranges of key environmental axes can also be used to learn where, along  
1205 environmental continuums, functional thresholds exist that could help with categorizations  
1206 important for policy and management (Richardson et al. 2022b).

1207 Cross-VIE understanding of the drivers, patterns, and processes linking inundation to  
1208 system responses can greatly improve with increased collaboration and communication across  
1209 scientific fields and systems. Our experience is that communities working in VIEs are scattered  
1210 across different societies and funding programs. Studying VIEs together via unifying conceptual  
1211 models tied to environmental continuums can bring these science communities together. To this  
1212 end, we encourage training and collaborations focused on consistent data generation methods  
1213 that may be adopted across the VIE community and in pursuit of conceptual unification. In  
1214 addition, disciplinary conferences could also recognize VIE commonalities with special sessions  
1215 to bring people together from across the VIE continuum to discuss research needs.

1216 Cross-VIE knowledge and models are needed to address human impacts to environments  
1217 across the globe. Humans both directly (i.e., dams, weirs, surface water and groundwater  
1218 abstraction, channelization, draining, invasive species introduction and spread, etc.) and  
1219 indirectly (i.e., climate change) alter VIEs (Maris et al. 2016, Pumo et al. 2016, Kiss et al. 2019).  
1220 As climate change and other anthropogenic impacts increasingly alter these already dynamic  
1221 systems, it is imperative that knowledge and models transcend VIEs. Future environmental  
1222 change can alter the position of a given VIE within environmental space, including what is  
1223 depicted in our conceptual model (**Fig. 13**) (e.g., increasing frequency of storm surges changing  
1224 the inundation return interval). The ability to predict impacts of such environmental change can  
1225 be facilitated by mechanistic knowledge that is transferable across the environmental space  
1226 occupied by VIEs. We hypothesize that unifying VIEs across environmental continuums can  
1227 help achieve this mechanistic, transferable knowledge.  
1228



1229  
 1230 **Figure 13. We encourage unifying conceptual models of VIEs based on hypotheses**  
 1231 **linked to continuous environmental axes, across which these systems can be studied**  
 1232 **without regard for what system names may be attached to a given studied place and**  
 1233 **time.** In our proposed conceptual model, two key are topographic slope and inundation return  
 1234 interval. Points represent approximate locations of where each VIE type may lie. Each VIE type  
 1235 spans a range of slopes and inundation return intervals, but we do not define these ranges as  
 1236 the conceptual model is based on how study systems fall across the environmental space  
 1237 represented here, rather than within specific nomenclature. Two priority research directions are  
 1238 greenhouse gas (GHG) fluxes and biological diversity, and the arrows represent possible  
 1239 hypotheses that could be evaluated with cross-VIE studies. We propose that knowledge and  
 1240 models that are transferable across VIEs can be achieved through evaluation of such  
 1241 hypotheses across broad ranges in slope and return interval. Credit: Nathan Johnson.

1242  
 1243

1244 **Acknowledgements**

1245 Any use of trade, firm, or product names is for descriptive purposes only and does not imply  
 1246 endorsement by the U.S. Government. The research described in this paper was supported by  
 1247 the Earth & Biological Sciences Program Development Office at Pacific Northwest National  
 1248 Laboratory, a multiprogram national laboratory operated by Battelle for the U.S. Department of  
 1249 Energy. We thank Jon Chorover, Sarah Godsey, Jesus Gomez-Velez, Wei Huang, Roser  
 1250 Matamala, Hyun Song and Kristen Underwood for contributions to the conceptual directions of

1251 this manuscript. This manuscript was an outgrowth of the VIE Workshop and we greatly thank  
1252 the participants for their contributions.

1253

1254 **Competing Interests:** The authors declare that they have no conflict of interest.

1255

## 1256 **References**

1257

1258 Abbott, B. W., K. Bishop, J. P. Zarnetske, C. Minaudo, F. S. Chapin, S. Krause, D. M. Hannah,

1259 L. Conner, D. Ellison, S. E. Godsey, S. Plont, J. Marçais, T. Kolbe, A. Huebner, R. J.

1260 Frei, T. Hampton, S. Gu, M. Buhman, S. Sara Sayedi, O. Ursache, M. Chapin, K. D.

1261 Henderson, and G. Pinay. 2019. Human domination of the global water cycle absent

1262 from depictions and perceptions. *Nature Geoscience* 12:533–540.

1263 Acharya, B. S., M. Bhandari, F. Bandini, A. Pizarro, M. Perks, D. R. Joshi, S. Wang, T.

1264 Dogwiler, R. L. Ray, G. Kharel, and S. Sharma. 2021. Unmanned aerial vehicles in

1265 hydrology and water management: Applications, challenges, and perspectives. *Water*

1266 *Resources Research* 57:e2021WR029925.

1267 Adams, R. K., and J. A. Spotila. 2005. The form and function of headwater streams based on

1268 field and modeling investigations in the southern Appalachian Mountains. *Earth Surface*

1269 *Processes and Landforms* 30:1521–1546.

1270 Adler, P. B., E. P. White, W. K. Lauenroth, D. M. Kaufman, A. Rassweiler, and J. A. Rusak.

1271 2005. Evidence for a General Species-Time-Area Relationship. *Ecology* 86:2032–2039.

1272 Åhlén, I., J. Thorslund, P. Hambäck, G. Destouni, and J. Jarsjö. 2022. Wetland position in the

1273 landscape: Impact on water storage and flood buffering. *Ecohydrology* 15.

1274 Allen, D. C., T. Datry, K. S. Boersma, M. T. Bogan, A. J. Boulton, D. Bruno, M. H. Busch, K. H.

1275 Costigan, W. K. Dodds, K. M. Fritz, S. E. Godsey, J. B. Jones, T. Kaletova, S. K. Kampf,

1276 M. C. Mims, T. M. Neeson, J. D. Olden, A. V. Pastor, N. L. Poff, B. L. Ruddell, A. Ruhi,

1277 G. Singer, P. Vezza, A. S. Ward, and M. Zimmer. 2020. River ecosystem conceptual

1278 models and non-perennial rivers: A critical review. *WIREs. Water* 7:e1473.

1279 Anderson, M. G., and T. P. Burt. 1978. The role of topography in controlling throughflow

1280 generation. *Earth Surface Processes* 3:331–344.

1281 Angle, J. C., T. H. Morin, L. M. Solden, A. B. Narrowe, G. J. Smith, M. A. Borton, C. Rey-  
1282 Sanchez, R. A. Daly, G. Mirfenderesgi, D. W. Hoyt, W. J. Riley, C. S. Miller, G. Bohrer,  
1283 and K. C. Wrighton. 2017. Methanogenesis in oxygenated soils is a substantial fraction  
1284 of wetland methane emissions. *Nature Communications* 8:1567.

1285 Appels, W. M., P. W. Bogaart, and S. E. A. T. M. van der Zee. 2016. Surface runoff in flat  
1286 terrain: How field topography and runoff generating processes control hydrological  
1287 connectivity. *Journal of Hydrology* 534:493–504.

1288 Arce, M. I., C. Mendoza-Lera, M. Almagro, N. Catalán, A. M. Romaní, E. Martí, R. Gómez, S.  
1289 Bernal, A. Foulquier, M. Mutz, R. Marcé, A. Zoppini, G. Gionchetta, G. Weigelhofer, R.  
1290 Del Campo, C. T. Robinson, A. Gilmer, M. Rulik, B. Obrador, O. Shumilova, S.  
1291 Zlatanović, S. Arnon, P. Baldrian, G. Singer, T. Datry, N. Skoulikidis, B. Tietjen, and D.  
1292 Von Schiller. 2019. A conceptual framework for understanding the biogeochemistry of  
1293 dry riverbeds through the lens of soil science. *Earth-Science Reviews* 188:441–453.

1294 Arias-Real, R., M. Delgado-Baquerizo, S. Sabater, C. Gutiérrez-Cánovas, E. Valencia, G.  
1295 Aragón, Y. Cantón, T. Datry, P. Giordani, N. G. Medina, A. de los Ríos, A. M. Romaní,  
1296 B. Weber, and P. Hurtado. 2024. Unfolding the dynamics of ecosystems undergoing  
1297 alternating wet-dry transitional states. *Ecology Letters* 27:e14488.

1298 Arim, M., V. Pinelli, L. Rodríguez-Tricot, E. Ortiz, M. Illarze, C. Fagúndez-Pachón, and A. I.  
1299 Borthagaray. 2023. Chance and necessity in the assembly of plant communities:  
1300 Stochasticity increases with size, isolation and diversity of temporary ponds. *Journal of*  
1301 *Ecology* 111:1641–1655.

1302 Arnesen, A. S., T. S. F. Silva, L. L. Hess, E. M. L. M. Novo, C. M. Rudorff, B. D. Chapman, and  
1303 K. C. McDonald. 2013. Monitoring flood extent in the lower Amazon River floodplain  
1304 using ALOS/PALSAR ScanSAR images. *Remote Sensing of Environment* 130:51–61.

1305 Arnold, W., J. Z. Salazar, A. Carlino, M. Giuliani, and A. Castelletti. 2023. Operations eclipse

1306 sequencing in multipurpose dam planning. *Earth's Future* 11:e2022EF003186.

1307 Arrigo, K. R., G. L. Van Dijken, M. A. Cameron, J. Van Der Grient, L. M. Wedding, L. Hazen, J.  
1308 Leape, G. Leonard, A. Merkl, F. Micheli, M. M. Mills, S. Monismith, N. T. Ouellette, A.  
1309 Zivian, M. Levi, and R. M. Bailey. 2020. Synergistic interactions among growing  
1310 stressors increase risk to an Arctic ecosystem. *Nature Communications* 11:6255.

1311 Arscott, D. B., K. Tockner, D. Van Der Nat, and J. V. Ward. 2002. Aquatic habitat dynamics  
1312 along a braided Alpine river ecosystem (Tagliamento River, Northeast Italy). *Ecosystems*  
1313 5:0802–0814.

1314 Atchley, A. L., S. L. Painter, D. R. Harp, E. T. Coon, C. J. Wilson, A. K. Liljedahl, and V. E.  
1315 Romanovsky. 2015. Using field observations to inform thermal hydrology models of  
1316 permafrost dynamics with ATS (v0.83). *Geoscientific Model Development* 8:2701–2722.

1317 Bain, M. B., J. T. Finn, and H. E. Booke. 1988. Streamflow regulation and fish community  
1318 structure. *Ecology* 69:382–392.

1319 Bam, E. K. P., A. M. Ireson, G. Kamp, and J. M. Hendry. 2020. Ephemeral ponds: Are they the  
1320 dominant source of depression-focused groundwater recharge? *Water Resources*  
1321 *Research* 56.

1322 Banach, A. M., K. Banach, E. J. W. Visser, Z. Stępniewska, A. J. M. Smits, J. G. M. Roelofs,  
1323 and L. P. M. Lamers. 2009. Effects of summer flooding on floodplain biogeochemistry in  
1324 Poland; implications for increased flooding frequency. *Biogeochemistry* 92:247–262.

1325 Baptist, M. J., P. Dankers, J. Cleveringa, L. Sittoni, P. W. J. M. Willemsen, M. E. B. van  
1326 Puijenbroek, B. M. L. de Vries, J. R. F. W. Leuven, L. Coumou, H. Kramer, and K.  
1327 Elschot. 2021. Salt marsh construction as a nature-based solution in an estuarine social-  
1328 ecological system. *Nature-Based Solutions* 1:100005.

1329 Barczok, M., C. Smith, N. Di Domenico, L. Kinsman-Costello, and E. Herndon. 2023. Variability  
1330 in soil redox response to seasonal flooding in a vernal pond. *Frontiers in Environmental*  
1331 *Science* 11.

- 1332 Batson, J., G. B. Noe, C. R. Hupp, K. W. Krauss, N. B. Rybicki, and E. R. Schenk. 2015. Soil  
1333 greenhouse gas emissions and carbon budgeting in a short-hydroperiod floodplain  
1334 wetland. *Journal of Geophysical Research: Biogeosciences* 120:77–95.
- 1335 Beckingham, B., T. Callahan, and V. Vulava. 2019. Stormwater Ponds in the Southeastern U.S.  
1336 Coastal Plain: Hydrogeology, Contaminant Fate, and the Need for a Social-Ecological  
1337 Framework. *Frontiers in Environmental Science* 7:1–14.
- 1338 Belyea, L. R., and A. J. Baird. 2006. Beyond "The limits to peat bog growth: cross-scale  
1339 feedback in peatland development. *Ecological Monographs* 76:299–322.
- 1340 Benstead, J. P., and D. S. Leigh. 2012. An expanded role for river networks. *Nature Geoscience*  
1341 5:678–679.
- 1342 Bernhardt, E. S., J. R. Blaszczak, C. D. Ficken, M. L. Fork, K. E. Kaiser, and E. C. Seybold.  
1343 2017. Control points in ecosystems: Moving beyond the hot spot hot moment concept.  
1344 *Ecosystems* 20:665–682.
- 1345 Bernhardt, E. S., P. Savoy, M. J. Vlah, A. P. Appling, L. E. Koenig, R. O. Hall, M. Arroita, J. R.  
1346 Blaszczak, A. M. Carter, M. Cohen, J. W. Harvey, J. B. Heffernan, A. M. Helton, J. D.  
1347 Hosen, L. Kirk, W. H. McDowell, E. H. Stanley, C. B. Yackulic, and N. B. Grimm. 2022.  
1348 Light and flow regimes regulate the metabolism of rivers. *Proceedings of the National*  
1349 *Academy of Sciences* 119:e2121976119.
- 1350 Bertolini, C., and J. da Mosto. 2021. Restoring for the climate: a review of coastal wetland  
1351 restoration research in the last 30 years. *Restoration Ecology* 29:e13438.
- 1352 Betson, R. P., and J. B. Marius. 1969. Source areas of storm runoff. *Water Resources Research*  
1353 5:574–582.
- 1354 Bettez, N. D., and P. M. Groffman. 2012. Denitrification potential in stormwater control  
1355 structures and natural riparian zones in an urban landscape. *Environmental Science and*  
1356 *Technology* 46:10909–10917.
- 1357 Bevacqua, E., M. I. Vousdoukas, G. Zappa, K. Hodges, T. G. Shepherd, D. Maraun, L.



1358 Mentaschi, and L. Feyen. 2020. More meteorological events that drive compound  
1359 coastal flooding are projected under climate change. *Communications Earth &*  
1360 *Environment* 1:47.

1361 Biancamaria, S., D. P. Lettenmaier, and T. M. Pavelsky. 2016. The SWOT Mission and its  
1362 capabilities for land hydrology. *Surveys in Geophysics* 37:307–337.

1363 Bie, W., T. Fei, X. Liu, H. Liu, and G. Wu. 2020. Small water bodies mapped from Sentinel-2  
1364 MSI (MultiSpectral Imager) imagery with higher accuracy. *International Journal of*  
1365 *Remote Sensing* 41:7912–7930.

1366 Blaurock, K., P. Garthen, B. S. Gilfedder, J. H. Fleckenstein, S. Peiffer, and L. Hopp. 2021.  
1367 Elucidating sources and pathways of dissolved organic carbon in a small, forested  
1368 catchment: A qualitative assessment of stream, soil and shallow groundwater. *other,*  
1369 *pico.*

1370 Bloom, A. A., K. W. Bowman, M. Lee, A. J. Turner, R. Schroeder, J. R. Worden, R. Weidner, K.  
1371 C. McDonald, and D. J. Jacob. 2017. A global wetland methane emissions and  
1372 uncertainty dataset for atmospheric chemical transport models (WetCHARTs version  
1373 1.0). *Geoscientific Model Development* 10:2141–2156.

1374 Bloom, A. A., P. I. Palmer, A. Fraser, D. S. Reay, and C. Frankenberg. 2010. Large-scale  
1375 controls of methanogenesis inferred from methane and gravity spaceborne data.  
1376 *Science* 327:322–325.

1377 Bogaard, T. A., and R. Greco. 2016. Landslide hydrology: from hydrology to pore pressure.  
1378 *WIREs Water* 3:439–459.

1379 Bogan, M. T., E. T. Chester, T. Datry, A. L. Murphy, B. J. Robson, A. Ruhi, R. Stubbington, and  
1380 J. E. Whitney. 2017a. Resistance, Resilience, and Community Recovery in Intermittent  
1381 Rivers and Ephemeral Streams. Pages 349–376 *Intermittent Rivers and Ephemeral*  
1382 *Streams*. Elsevier.

1383 Bogan, M. T., E. T. Chester, T. Datry, A. L. Murphy, B. J. Robson, A. Ruhi, R. Stubbington, and

1384 J. E. Whitney. 2017b. Resistance, Resilience, and Community Recovery in Intermittent  
1385 Rivers and Ephemeral Streams. Pages 349–376 Intermittent Rivers and Ephemeral  
1386 Streams. Elsevier.

1387 Bogard, M. J., B. A. Bergamaschi, D. E. Butman, F. Anderson, S. H. Knox, and L. Windham-  
1388 Myers. 2020. Hydrologic export is a major component of coastal wetland carbon  
1389 budgets. *Global Biogeochemical Cycles* 34.

1390 Bonada, N., M. Rieradevall, and N. Prat. 2007. Macroinvertebrate community structure and  
1391 biological traits related to flow permanence in a Mediterranean river network.  
1392 *Hydrobiologia* 589:91–106.

1393 Bonython, C. W., and B. Mason. 1953. The Filling and Drying of Lake Eyre. *The Geographical*  
1394 *Journal* 119:321–330.

1395 Borch, T., R. Kretzchmar, A. Kappler, P. Van Cappellen, M. Ginder-Vogel, and K. Campbell.  
1396 2010. Biogeochemical redox processes and their impact on contaminant dynamics.  
1397 *Environmental Science & Technology* 44:15–23.

1398 Bornette, G., C. Amoros, H. Piegay, J. Tachet, and T. Hein. 1998. Ecological complexity of  
1399 wetlands within a river landscape. *Biological Conservation* 85:35–45.

1400 Bourke, S. A., M. Shanafield, P. Hedley, S. Chapman, and S. Dogramaci. 2023. A hydrological  
1401 framework for persistent pools along non-perennial rivers. *Hydrology and Earth System*  
1402 *Sciences* 27:809–836.

1403 Brantley, S. L., M. I. Lebedeva, V. N. Balashov, K. Singha, P. L. Sullivan, and G. Stinchcomb.  
1404 2017. Toward a conceptual model relating chemical reaction fronts to water flow paths in  
1405 hills. *Geomorphology* 277:100–117.

1406 Braswell, A. E., and J. B. Heffernan. 2019. Coastal wetland distributions: Delineating domains of  
1407 macroscale drivers and local feedbacks. *Ecosystems* 22:1256–1270.

1408 Braswell, A. E., S. Leyk, D. S. Connor, and J. H. Uhl. 2022. Creeping disaster along the U.S.  
1409 coastline: Understanding exposure to sea level rise and hurricanes through historical

1410 development. PLOS ONE 17:e0269741.

1411 Brazier, R. E., A. Puttock, H. A. Graham, R. E. Auster, K. H. Davies, and C. M. L. Brown. 2021.

1412 Beaver: Nature's ecosystem engineers. WIREs Water 8.

1413 Brendonck, L., T. Pinceel, and R. Ortells. 2017. Dormancy and dispersal as mediators of

1414 zooplankton population and community dynamics along a hydrological disturbance

1415 gradient in inland temporary pools. Hydrobiologia 796:201–222.

1416 Brinson, M. 1993. A Hydrogeomorphic classification for wetlands. Technical Report, U.S. Army

1417 Corps of Engineers, Washington, DC.

1418 Brooks, R. T. 2004. Weather-related effects on woodland vernal pool hydrology and

1419 hydroperiod. Wetlands 24:104–114.

1420 Burt, T. P., and W. T. Swank. 2010. Hursh CR and Brater EF (1941) Separating storm-

1421 hydrographs from small drainage-areas into surface- and subsurface-flow. Transactions,

1422 American Geophysical Union 22: 863-871. Progress in Physical Geography: Earth and

1423 Environment 34:719–726.

1424 Busch, M. H., K. H. Costigan, K. M. Fritz, T. Datry, C. A. Krabbenhoft, J. C. Hammond, M.

1425 Zimmer, J. D. Olden, R. M. Burrows, W. K. Dodds, K. S. Boersma, M. Shanafield, S. K.

1426 Kampf, M. C. Mims, M. T. Bogan, A. S. Ward, M. Perez Rocha, S. Godsey, G. H. Allen,

1427 J. R. Blaszcak, C. N. Jones, and D. C. Allen. 2020. What's in a name? Patterns, trends,

1428 and suggestions for defining non-perennial rivers and streams. Water 12:1980.

1429 Buszka, T. T., and D. M. Reeves. 2021. Pathways and timescales associated with nitrogen

1430 transport from septic systems in coastal aquifers intersected by canals. Hydrogeology

1431 Journal 29:1953–1964.

1432 Calhoun, A. J. K., D. M. Mushet, K. P. Bell, D. Boix, J. A. Fitzsimons, and F. Isselin-Nondedeu.

1433 2017. Temporary wetlands: challenges and solutions to conserving a 'disappearing'

1434 ecosystem. Biological Conservation 211:3–11.

1435 Cantelon, J. A., J. A. Guimond, C. E. Robinson, H. A. Michael, and B. L. Kurylyk. 2022. Vertical

1436 saltwater intrusion in coastal aquifers driven by episodic flooding: A review. *Water*  
1437 *Resources Research* 58:e2022WR032614.

1438 Capps, K. A., R. Rancatti, N. Tomczyk, T. B. Parr, A. J. K. Calhoun, and M. Hunter. 2014.  
1439 Biogeochemical hotspots in forested landscapes: The role of vernal pools in  
1440 denitrification and organic matter processing. *Ecosystems* 17:1455–1468.

1441 Casanova, M. T., and M. A. Brock. 2000. How do depth, duration and frequency of flooding  
1442 influence the establishment of wetland plant communities? *Plant Ecology* 147:237–250.

1443 Castaldelli, G., E. Soana, E. Racchetti, F. Vincenzi, E. A. Fano, and M. Bartoli. 2015. Vegetated  
1444 canals mitigate nitrogen surplus in agricultural watersheds. *Agriculture, Ecosystems &*  
1445 *Environment* 212:253–262.

1446 Castañeda-Moya, E., V. H. Rivera-Monroy, R. M. Chambers, X. Zhao, L. Lamb-Wotton, A.  
1447 Gorsky, E. E. Gaiser, T. G. Troxler, J. S. Kominoski, and M. Hiatt. 2020. Hurricanes  
1448 fertilize mangrove forests in the Gulf of Mexico (Florida Everglades, USA). *Proceedings*  
1449 *of the National Academy of Sciences* 117:4831–4841.

1450 Cawley, K. M., Y. Yamashita, N. Maie, and R. Jaffé. 2014. Using optical properties to quantify  
1451 fringe mangrove inputs to the dissolved organic matter (DOM) pool in a subtropical  
1452 estuary. *Estuaries and Coasts* 37:399–410.

1453 Celi, J. E., and S. K. Hamilton. 2020. Measuring Floodplain Inundation Using Diel Amplitude of  
1454 Temperature. *Sensors* 20:6189.

1455 Chambers, L. G., H. E. Steinmuller, and J. L. Breithaupt. 2019. Toward a mechanistic  
1456 understanding of “peat collapse” and its potential contribution to coastal wetland loss.  
1457 *Ecology* 100.

1458 Chapin, T. P., A. S. Todd, and M. P. Zeigler. 2014. Robust, low-cost data loggers for stream  
1459 temperature, flow intermittency, and relative conductivity monitoring. *Water Resources*  
1460 *Research* 50:6542–6548.

1461 Chen, B., and D. H. Wise. 1999. Bottom-up limitation of predaceous arthropods in a detritus-

1462 based terrestrial food web. *Ecology* 80:761–772.

1463 Cheng, F. Y., J. Park, M. Kumar, and N. B. Basu. 2023. Disconnectivity matters: the outsized  
1464 role of small ephemeral wetlands in landscape-scale nutrient retention. *Environmental*  
1465 *Research Letters* 18:024018.

1466 Choularton, T. W., and S. J. Perry. 1986. A model of the orographic enhancement of snowfall by  
1467 the seeder-feeder mechanism. *Quarterly Journal of the Royal Meteorological Society*  
1468 112:335–345.

1469 Clark, K. E., M. A. Torres, A. J. West, R. G. Hilton, M. New, A. B. Horwath, J. B. Fisher, J. M.  
1470 Rapp, A. Robles Caceres, and Y. Malhi. 2014. The hydrological regime of a forested  
1471 tropical Andean catchment. *Hydrology and Earth System Sciences* 18:5377–5397.

1472 Clementson, L. A., A. J. Richardson, W. A. Rochester, K. Oubelkheir, B. Liu, E. J. D'Sa, L. F. M.  
1473 Gusmão, P. Ajani, T. Schroeder, P. W. Ford, M. A. Burford, E. Saeck, and A. D. L.  
1474 Steven. 2021. Effect of a once in 100-year flood on a subtropical coastal phytoplankton  
1475 community. *Frontiers in Marine Science* 8.

1476 Clifford, C. C., and J. B. Heffernan. 2023. North Carolina coastal plain ditch types support  
1477 distinct hydrophytic communities. *Wetlands* 43:56.

1478 Clifford, C., and J. Heffernan. 2018. Artificial aquatic ecosystems. *Water* 10:1096.

1479 Colbert, A. J., and B. J. Soden. 2012. Climatological variations in North Atlantic tropical cyclone  
1480 tracks. *Journal of Climate* 25:657–673.

1481 Coles, A. E., B. G. McConkey, and J. J. McDonnell. 2017. Climate change impacts on hillslope  
1482 runoff on the northern Great Plains, 1962–2013. *Journal of Hydrology* 550:538–548.

1483 Colmer, T. D. 2003. Long-distance transport of gases in plants: A perspective on internal  
1484 aeration and radial oxygen loss from roots: Gas transport in plants. *Plant, Cell &*  
1485 *Environment* 26:17–36.

1486 Connolly, R. M. 2003. Differences in trophodynamics of commercially important fish between  
1487 artificial waterways and natural coastal wetlands. *Estuarine, Coastal and Shelf Science*

1488 58:929–936.

1489 Cooley, S., L. Smith, L. Stepan, and J. Mascaro. 2017. Tracking dynamic northern surface water  
1490 changes with high-frequency planet CubeSat imagery. *Remote Sensing* 9:1306.

1491 Coon, E. T., J. D. Moulton, E. Kikinzon, M. Berndt, G. Manzini, R. Garimella, K. Lipnikov, and S.  
1492 L. Painter. 2020. Coupling surface flow and subsurface flow in complex soil structures  
1493 using mimetic finite differences. *Advances in Water Resources* 144:103701.

1494 Corti, R., and T. Datry. 2012. Invertebrates and sestonic matter in an advancing wetted front  
1495 travelling down a dry river bed (Albarine, France). *Freshwater Science* 31:1187–1201.

1496 Costigan, K. H., M. D. Daniels, and W. K. Dodds. 2015. Fundamental spatial and temporal  
1497 disconnections in the hydrology of an intermittent prairie headwater network. *Journal of*  
1498 *Hydrology* 522:305–316.

1499 Costigan, K. H., K. L. Jaeger, C. W. Goss, K. M. Fritz, and P. C. Goebel. 2016. Understanding  
1500 controls on flow permanence in intermittent rivers to aid ecological research: integrating  
1501 meteorology, geology and land cover: Integrating Science to Understand Flow  
1502 Intermittence. *Ecohydrology* 9:1141–1153.

1503 Costigan, K. H., M. J. Kennard, C. Leigh, E. Sauquet, T. Datry, and A. J. Boulton. 2017. Flow  
1504 regimes in intermittent rivers and ephemeral streams. Pages 51–78 *Intermittent Rivers*  
1505 *and Ephemeral Streams*. Elsevier.

1506 Covino, T. 2017. Hydrologic connectivity as a framework for understanding biogeochemical flux  
1507 through watersheds and along fluvial networks. *Geomorphology* 277:133–144.

1508 Cowardin, L. M., and F. C. Golet. 1995. US Fish and Wildlife Service 1979 wetland  
1509 classification: A review. *Vegetatio* 118:139–152.

1510 Crook, D. A., D. J. Buckle, J. R. Morrongiello, Q. A. Allsop, W. Baldwin, T. M. Saunders, and M.  
1511 M. Douglas. 2020. Tracking the resource pulse: Movement responses of fish to dynamic  
1512 floodplain habitat in a tropical river. *Journal of Animal Ecology* 89:795–807.

1513 Crotty, S. M., C. Ortals, T. M. Pettengill, L. Shi, M. Olabarrieta, M. A. Joyce, A. H. Altieri, E.

1514 Morrison, T. S. Bianchi, C. Craft, M. D. Bertness, and C. Angelini. 2020. Sea-level rise  
1515 and the emergence of a keystone grazer alter the geomorphic evolution and ecology of  
1516 southeast US salt marshes. *Proceedings of the National Academy of Sciences*  
1517 117:17891–17902.

1518 Crump, B. C., L. M. Fine, C. S. Fortunato, L. Herfort, J. A. Needoba, S. Murdock, and F. G.  
1519 Prah. 2017. Quantity and quality of particulate organic matter controls bacterial  
1520 production in the Columbia River estuary. *Limnology and Oceanography* 62:2713–2731.

1521 Cubley, E. S., D. J. Cooper, and D. M. Merritt. 2023. Are riparian vegetation flow response  
1522 guilds transferable between rivers? *Freshwater Biology* 68:406–424.

1523 Culley, S., S. Noble, A. Yates, M. Timbs, S. Westra, H. R. Maier, M. Giuliani, and A. Castelletti.  
1524 2016. A bottom-up approach to identifying the maximum operational adaptive capacity of  
1525 water resource systems to a changing climate. *Water Resources Research* 52:6751–  
1526 6768.

1527 Dang, C., E. M. Morrissey, S. C. Neubauer, and R. B. Franklin. 2019. Novel microbial  
1528 community composition and carbon biogeochemistry emerge over time following  
1529 saltwater intrusion in wetlands. *Global Change Biology* 25:549–561.

1530 Daniel, J., and R. C. Rooney. 2021. Wetland hydroperiod predicts community structure, but not  
1531 the magnitude of cross-community congruence. *Scientific Reports* 11:429.

1532 Datry, T., A. J. Boulton, N. Bonada, K. Fritz, C. Leigh, E. Sauquet, K. Tockner, B. Hugueny, and  
1533 C. N. Dahm. 2018a. Flow intermittence and ecosystem services in rivers of the  
1534 Anthropocene. *Journal of Applied Ecology* 55:353–364.

1535 Datry, T., A. Foulquier, R. Corti, D. von Schiller, K. Tockner, C. Mendoza-Lera, J. C. Clément,  
1536 M. O. Gessner, M. Moleón, R. Stubbington, B. Gücker, R. Albariño, D. C. Allen, F.  
1537 Altermatt, M. I. Arce, S. Arnon, D. Banas, A. Banegas-Medina, E. Beller, M. L.  
1538 Blanchette, J. F. Blanco-Libreros, J. J. Blessing, I. G. Boëchat, K. S. Boersma, M. T.  
1539 Bogan, N. Bonada, N. R. Bond, K. C. Brintrup Barría, A. Bruder, R. M. Burrows, T.

1540 Cancellario, C. Canhoto, S. M. Carlson, S. Cauvy-Fraunié, N. Cid, M. Danger, B. de  
1541 Freitas Terra, A. M. De Girolamo, E. de La Barra, R. del Campo, V. D. Diaz-Villanueva,  
1542 F. Dyer, A. Elozegi, E. Faye, C. Febria, B. Four, S. Gafny, S. D. Ghate, R. Gómez, L.  
1543 Gómez-Gener, M. a. S. Graça, S. Guareschi, F. Hoppeler, J. L. Hwan, J. I. Jones, S.  
1544 Kubheka, A. Laini, S. D. Langhans, C. Leigh, C. J. Little, S. Lorenz, J. C. Marshall, E.  
1545 Martín, A. R. McIntosh, E. I. Meyer, M. Miliša, M. C. Mlambo, M. Morais, N. Moya, P. M.  
1546 Negus, D. K. Niyogi, A. Papatheodoulou, I. Pardo, P. Pařil, S. U. Pauls, V. Peřić, M.  
1547 Polářek, C. T. Robinson, P. Rodríguez-Lozano, R. J. Rolls, M. M. Sánchez-Montoya, A.  
1548 Savić, O. Shumilova, K. R. Sridhar, A. L. Steward, R. Storey, A. Taleb, A. Uzan, R.  
1549 Vander Vorste, N. J. Waltham, C. Woelfle-Erskine, D. Zak, C. Zarfl, and A. Zoppini.  
1550 2018b. A global analysis of terrestrial plant litter dynamics in non-perennial waterways.  
1551 Nature Geoscience 11:497–503.

1552 Datry, T., A. Foulquier, R. Corti, D. Von Schiller, K. Tockner, C. Mendoza-Lera, J. C. Clément,  
1553 M. O. Gessner, M. Moleón, R. Stubbington, B. Gücker, R. Albariño, D. C. Allen, F.  
1554 Altermatt, M. I. Arce, S. Arnon, D. Banas, A. Banegas-Medina, E. Beller, M. L.  
1555 Blanchette, J. F. Blanco-Libreros, J. J. Blessing, I. G. Boëchat, K. S. Boersma, M. T.  
1556 Bogan, N. Bonada, N. R. Bond, K. C. Brintrup Barría, A. Bruder, R. M. Burrows, T.  
1557 Cancellario, C. Canhoto, S. M. Carlson, S. Cauvy-Fraunié, N. Cid, M. Danger, B. De  
1558 Freitas Terra, A. M. De Girolamo, E. De La Barra, R. Del Campo, V. D. Diaz-Villanueva,  
1559 F. Dyer, A. Elozegi, E. Faye, C. Febria, B. Four, S. Gafny, S. D. Ghate, R. Gómez, L.  
1560 Gómez-Gener, M. A. S. Graça, S. Guareschi, F. Hoppeler, J. L. Hwan, J. I. Jones, S.  
1561 Kubheka, A. Laini, S. D. Langhans, C. Leigh, C. J. Little, S. Lorenz, J. C. Marshall, E.  
1562 Martín, A. R. McIntosh, E. I. Meyer, M. Miliša, M. C. Mlambo, M. Morais, N. Moya, P. M.  
1563 Negus, D. K. Niyogi, A. Papatheodoulou, I. Pardo, P. Pařil, S. U. Pauls, V. Peřić, M.  
1564 Polářek, C. T. Robinson, P. Rodríguez-Lozano, R. J. Rolls, M. M. Sánchez-Montoya, A.  
1565 Savić, O. Shumilova, K. R. Sridhar, A. L. Steward, R. Storey, A. Taleb, A. Uzan, R.



1566 Vander Vorste, N. J. Waltham, C. Woelfle-Erskine, D. Zak, C. Zarfl, and A. Zoppini.  
1567 2018c. A global analysis of terrestrial plant litter dynamics in non-perennial waterways.  
1568 Nature Geoscience 11:497–503.

1569 Datry, T., and S. T. Larned. 2008. River flow controls ecological processes and invertebrate  
1570 assemblages in subsurface flowpaths of an ephemeral river reach. Canadian Journal of  
1571 Fisheries and Aquatic Sciences 65:1532–1544.

1572 Datry, T., A. Truchy, J. D. Olden, M. H. Busch, R. Stubbington, W. K. Dodds, S. Zipper, S. Yu,  
1573 M. L. Messenger, J. D. Tonkin, K. E. Kaiser, J. C. Hammond, E. K. Moody, R. M.  
1574 Burrows, R. Sarremejane, A. G. DelVecchia, M. L. Fork, C. J. Little, R. H. Walker, A. W.  
1575 Walters, and D. Allen. 2023. Causes, responses, and implications of anthropogenic  
1576 versus natural flow intermittence in river networks. BioScience 73:9–22.

1577 Davidson, Eric A., E. Belk, and R. D. Boone. 1998. Soil water content and temperature as  
1578 independent or confounded factors controlling soil respiration in a temperate mixed  
1579 hardwood forest. Global Change Biology 4:217–227.

1580 Davidson, N. C., E. Fluet-Chouinard, and C. M. Finlayson. 2018. Global extent and distribution  
1581 of wetlands: trends and issues. Marine and Freshwater Research 69:620.

1582 Davidson, T. A., A. W. Mackay, P. Wolski, R. Mazebedi, M. Murray-Hudson, and M. Todd. 2012.  
1583 Seasonal and spatial hydrological variability drives aquatic biodiversity in a flood-pulsed,  
1584 sub-tropical wetland. Freshwater Biology 57:1253–1265.

1585 Davis, C. A., D. Dvoretz, and J. R. Bidwell. 2013. Hydrogeomorphic classification and functional  
1586 assessment. Pages 29–68 in J. T. Anderson and C. A. Davis, editors. Wetland  
1587 Techniques: Volume 3: Applications and Management. Springer Netherlands, Dordrecht.

1588 Day, J. A., H. L. Malan, E. Malijani, and A. P. Abegunde. 2019. Review: Water quality in non-  
1589 perennial rivers (with erratum). Water SA 45.

1590 De Jager, N. R., M. Thomsen, and Y. Yin. 2012. Threshold effects of flood duration on the  
1591 vegetation and soils of the Upper Mississippi River floodplain, USA. Forest Ecology and

1592 Management 270:135–146.

1593 De Sassi, C., O. T. Lewis, and J. M. Tylianakis. 2012. Plant-mediated and nonadditive effects of  
 1594 two global change drivers on an insect herbivore community. *Ecology* 93:1892–1901.

1595 De Vries, M. E., J. Rodenburg, B. V. Bado, A. Sow, P. A. Leffelaar, and K. E. Giller. 2010. Rice  
 1596 production with less irrigation water is possible in a Sahelian environment. *Field Crops  
 1597 Research* 116:154–164.

1598 Dee, M. M., and J. L. Tank. 2020. Inundation time mediates denitrification end products and  
 1599 carbon limitation in constructed floodplains of an agricultural stream. *Biogeochemistry*  
 1600 149:141–158.

1601 Del Campo, R., R. Corti, and G. Singer. 2021. Flow intermittence alters carbon processing in  
 1602 rivers through chemical diversification of leaf litter. *Limnology and Oceanography Letters*  
 1603 6:232–242.

1604 Della Rocca, F., L. Vignoli, and M. A. Bologna. 2005. The reproductive biology of *Salamandrina*  
 1605 *terdigitata* (Caudata, Salamandridae). *The Herpetological Journal* 15:273–278.

1606 DelVecchia, A. G., M. Shanafield, M. A. Zimmer, M. H. Busch, C. A. Krabbenhoft, R.  
 1607 Stubbington, K. E. Kaiser, R. M. Burrows, J. Hosen, T. Datry, S. K. Kampf, S. C. Zipper,  
 1608 K. Fritz, K. Costigan, and D. C. Allen. 2022. Reconceptualizing the hyporheic zone for  
 1609 nonperennial rivers and streams. *Freshwater Science* 41:167–182.

1610 Dietze, M. C., C. Averill, J. Foster, and K. Wheeler. 2017. *Ecological Forecasting*. Princeton  
 1611 University Press.

1612 Dietze, M. C., A. Fox, L. M. Beck-Johnson, J. L. Betancourt, M. B. Hooten, C. S. Jarnevich, T.  
 1613 H. Keitt, M. A. Kenney, C. M. Laney, L. G. Larsen, H. W. Loescher, C. K. Lunch, B. C.  
 1614 Pijanowski, J. T. Randerson, E. K. Read, A. T. Tredennick, R. Vargas, K. C. Weathers,  
 1615 and E. P. White. 2018. Iterative near-term ecological forecasting: Needs, opportunities,  
 1616 and challenges. *Proceedings of the National Academy of Sciences* 115:1424–1432.

1617 van Dijk, G., A. J. P. Smolders, R. Loeb, A. Bout, J. G. M. Roelofs, and L. P. M. Lamers. 2015.

1618 Salinization of coastal freshwater wetlands; effects of constant versus fluctuating salinity  
1619 on sediment biogeochemistry. *Biogeochemistry* 126:71–84.

1620 Dissanayake, P., J. Brown, P. Wisse, and H. Karunaratna. 2015. Effects of storm clustering on  
1621 beach/dune evolution. *Marine Geology* 370:63–75.

1622 Döll, P., and H. M. Schmied. 2012. How is the impact of climate change on river flow regimes  
1623 related to the impact on mean annual runoff? A global-scale analysis. *Environmental*  
1624 *Research Letters* 7:014037.

1625 Du, J., J. Shen, Y. J. Zhang, F. Ye, Z. Liu, Z. Wang, Y. P. Wang, X. Yu, M. Sisson, and H. V.  
1626 Wang. 2018. Tidal response to sea-level rise in different types of estuaries: The  
1627 importance of length, bathymetry, and geometry. *Geophysical Research Letters* 45:227–  
1628 235.

1629 Dube, K., G. Nhamo, and D. Chikodzi. 2021. Flooding trends and their impacts on coastal  
1630 communities of Western Cape Province, South Africa. *Geojournal*:1–16.

1631 Dugdale, S. J., J. Klaus, and D. M. Hannah. 2022. Looking to the Skies: Realising the  
1632 Combined Potential of Drones and Thermal Infrared Imagery to Advance Hydrological  
1633 Process Understanding in Headwaters. *Water Resources Research* 58.

1634 Ekici, A., H. Lee, D. M. Lawrence, S. C. Swenson, and C. Prigent. 2019. Ground subsidence  
1635 effects on simulating dynamic high-latitude surface inundation under permafrost thaw  
1636 using CLM5. *Geoscientific Model Development* 12:5291–5300.

1637 Elberling, B., L. Askaer, C. J. Jørgensen, H. P. Joensen, M. Kühl, R. N. Glud, and F. R.  
1638 Lauritsen. 2011. Linking Soil O<sub>2</sub>, CO<sub>2</sub>, and CH<sub>4</sub> Concentrations in a Wetland Soil:  
1639 Implications for CO<sub>2</sub> and CH<sub>4</sub> Fluxes. *Environmental Science & Technology* 45:3393–  
1640 3399.

1641 Elberling, B. B., G. M. Kovács, H. F. E. Hansen, R. Fensholt, P. Ambus, X. Tong, D. Gominski,  
1642 C. W. Mueller, D. M. N. Poultney, and S. Oehmcke. 2023. High nitrous oxide emissions  
1643 from temporary flooded depressions within croplands. *Communications Earth &*

1644 Environment 4:463.

1645 Ensign, S. H., and G. B. Noe. 2018. Tidal extension and sea-level rise: recommendations for a  
 1646 research agenda. *Frontiers in Ecology and the Environment* 16:37–43.

1647 Eppinga, M. B., M. Rietkerk, W. Borren, E. D. Lapshina, W. Bleuten, and M. J. Wassen. 2008.  
 1648 Regular surface patterning of peatlands: Confronting theory with field data. *Ecosystems*  
 1649 11:520–536.

1650 Euliss, N. H., J. W. LaBaugh, L. H. Fredrickson, D. M. Mushet, M. K. Laubhan, G. A. Swanson,  
 1651 T. C. Winter, D. O. Rosenberry, and R. D. Nelson. 2004. The wetland continuum: A  
 1652 conceptual framework for interpreting biological studies. *Wetlands* 24:448–458.

1653 Fagherazzi, S., M. L. Kirwan, S. M. Mudd, G. R. Guntenspergen, S. Temmerman, A. D’Alpaos,  
 1654 J. Van De Koppel, J. M. Rybczyk, E. Reyes, C. Craft, and J. Clough. 2012. Numerical  
 1655 models of salt marsh evolution: Ecological, geomorphic, and climatic factors. *Reviews of*  
 1656 *Geophysics* 50:RG1002.

1657 Fan, Y., M. Clark, D. M. Lawrence, S. Swenson, L. E. Band, S. L. Brantley, P. D. Brooks, W. E.  
 1658 Dietrich, A. Flores, G. Grant, J. W. Kirchner, D. S. Mackay, J. J. McDonnell, P. C. D.  
 1659 Milly, P. L. Sullivan, C. Tague, H. Ajami, N. Chaney, A. Hartmann, P. Hazenberg, J.  
 1660 McNamara, J. Pelletier, J. Perket, E. Rouholahnejad-Freund, T. Wagener, X. Zeng, E.  
 1661 Beighley, J. Buzan, M. Huang, B. Livneh, B. P. Mohanty, B. Nijssen, M. Safeeq, C.  
 1662 Shen, W. Verseveld, J. Volk, and D. Yamazaki. 2019. Hillslope hydrology in global  
 1663 change research and earth system modeling. *Water Resources Research* 55:1737–  
 1664 1772.

1665 Fan, Y., and G. Miguez-Macho. 2011. A simple hydrologic framework for simulating wetlands in  
 1666 climate and earth system models. *Climate Dynamics* 37:253–278.

1667 Fatichi, S., E. R. Vivoni, F. L. Ogden, V. Y. Ivanov, B. Mirus, D. Gochis, C. W. Downer, M.  
 1668 Camporese, J. H. Davison, B. Ebel, N. Jones, J. Kim, G. Mascaro, R. Niswonger, P.  
 1669 Restrepo, R. Rigon, C. Shen, M. Sulis, and D. Tarboton. 2016. An overview of current

1670 applications, challenges, and future trends in distributed process-based models in  
1671 hydrology. *Journal of Hydrology* 537:45–60.

1672 Finlayson, C. M., and A. G. Van Der Valk, editors. 1995. *Classification and Inventory of the*  
1673 *World's Wetlands*. Springer Netherlands, Dordrecht.

1674 Fisher, J. B., B. Lee, A. J. Purdy, G. H. Halverson, M. B. Dohlen, K. Cawse-Nicholson, A. Wang,  
1675 R. G. Anderson, B. Aragon, M. A. Arain, D. D. Baldocchi, J. M. Baker, H. Barral, C. J.  
1676 Bernacchi, C. Bernhofer, S. C. Biraud, G. Bohrer, N. Brunsell, B. Cappelaere, S. Castro-  
1677 Contreras, J. Chun, B. J. Conrad, E. Cremonese, J. Demarty, A. R. Desai, A. De Ligne,  
1678 L. Foltýnová, M. L. Goulden, T. J. Griffis, T. Grünwald, M. S. Johnson, M. Kang, D.  
1679 Kelbe, N. Kowalska, J. Lim, I. Maïnassara, M. F. McCabe, J. E. C. Missik, B. P.  
1680 Mohanty, C. E. Moore, L. Morillas, R. Morrison, J. W. Munger, G. Posse, A. D.  
1681 Richardson, E. S. Russell, Y. Ryu, A. Sanchez-Azofeifa, M. Schmidt, E. Schwartz, I.  
1682 Sharp, L. Šigut, Y. Tang, G. Hulley, M. Anderson, C. Hain, A. French, E. Wood, and S.  
1683 Hook. 2020. ECOSTRESS: NASA's next generation mission to measure  
1684 evapotranspiration from the international space station. *Water Resources Research* 56.

1685 Flick, R. E., D. B. Chadwick, J. Briscoe, and K. C. Harper. 2012. "Flooding" versus "inundation."  
1686 *Eos, Transactions American Geophysical Union* 93:365–366.

1687 Florencio, M., R. Fernández-Zamudio, M. Lozano, and C. Díaz-Paniagua. 2020. Interannual  
1688 variation in filling season affects zooplankton diversity in Mediterranean temporary  
1689 ponds. *Hydrobiologia* 847:1195–1205.

1690 Fortesa, J., G. F. Ricci, J. García-Comendador, F. Gentile, J. Estrany, E. Sauquet, T. Datry, and  
1691 A. M. De Girolamo. 2021. Analysing hydrological and sediment transport regime in two  
1692 Mediterranean intermittent rivers. *CATENA* 196:104865.

1693 Fournier, R. J., G. De Mendoza, R. Sarremejane, and A. Ruhi. 2023. Isolation controls  
1694 reestablishment mechanisms and post-drying community structure in an intermittent  
1695 stream. *Ecology* 104:e3911.

1696 Fredrickson, L., and T. S. Taylor. 1982. Management of seasonally flooded impoundments for  
1697 wildlife. Resource Publication 148, U.S. Fish and Wildlife Service.

1698 Freeze, R. A. 1974. Streamflow generation. *Reviews of Geophysics* 12:627–647.

1699 Fryirs, K., and G. Brierley. 2022. Assemblages of geomorphic units: A building block approach  
1700 to analysis and interpretation of river character, behaviour, condition and recovery. *Earth*  
1701 *Surface Processes and Landforms* 47:92–108.

1702 Gallant, A. 2015. The challenges of remote monitoring of wetlands. *Remote Sensing* 7:10938–  
1703 10950.

1704 Garayburu-Caruso, V. A., R. E. Danczak, J. C. Stegen, L. Renteria, M. Mccall, A. E. Goldman,  
1705 R. K. Chu, J. Toyoda, C. T. Resch, J. M. Torgeson, J. Wells, S. Fansler, S. Kumar, and  
1706 E. B. Graham. 2020. Using community science to reveal the global chemogeography of  
1707 river metabolomes. *Metabolites* 10:518.

1708 Gates, J. B., P. M. Chittaro, and K. B. Veggerby. 2020. Standard operating procedures for  
1709 measuring bulk stable isotope values of nitrogen and carbon in marine biota by isotope  
1710 ratio mass spectrometry (IRMS).

1711 Gendreau, K. L., V. Buxton, C. E. Moore, and M. Mims. 2021. Temperature loggers capture  
1712 intraregional variation of inundation timing for intermittent ponds. preprint, *Hydrology*.

1713 Gittman, R. K., F. J. Fodrie, A. M. Popowich, D. A. Keller, J. F. Bruno, C. A. Currin, C. H.  
1714 Peterson, and M. F. Piehler. 2015. Engineering away our natural defenses: an analysis  
1715 of shoreline hardening in the US. *Frontiers in Ecology and the Environment* 13:301–307.

1716 Glaser, B., L. Hopp, D. Partington, P. Brunner, R. Therrien, and J. Klaus. 2021. Sources of  
1717 surface water in space and time: Identification of delivery processes and geographical  
1718 sources with hydraulic mixing-cell modeling. *Water Resources Research*  
1719 57:e2021WR030332.

1720 Gleason, J. E., and R. C. Rooney. 2018. Pond permanence is a key determinant of aquatic  
1721 macroinvertebrate community structure in wetlands. *Freshwater Biology* 63:264–277.

1722 Goldman, A. E., S. R. Emani, L. C. Pérez-Angel, J. A. Rodríguez-Ramos, and J. C. Stegen.  
1723 2022. Integrated, coordinated, open, and networked (ICON) science to advance the  
1724 geosciences: Introduction and synthesis of a special collection of commentary articles.  
1725 Earth and Space Science 9:e2021EA002099.

1726 Goldman, A. E., E. B. Graham, A. R. Crump, D. W. Kennedy, E. B. Romero, C. G. Anderson, K.  
1727 L. Dana, C. T. Resch, J. K. Fredrickson, and J. C. Stegen. 2017. Biogeochemical cycling  
1728 at the aquatic–terrestrial interface is linked to parafluvial hyporheic zone inundation  
1729 history. Biogeosciences 14:4229–4241.

1730 Gomez, J. D., J. L. Wilson, and M. B. Cardenas. 2012. Residence time distributions in sinuosity-  
1731 driven hyporheic zones and their biogeochemical effects. Water Resources Research  
1732 48.

1733 Gómez-Gener, L., A. R. Siebers, M. I. Arce, S. Arnon, S. Bernal, R. Bolpagni, T. Datry, G.  
1734 Gionchetta, H.-P. Grossart, C. Mendoza-Lera, V. Pohl, U. Risse-Buhl, O. Shumilova, O.  
1735 Tzoraki, D. Von Schiller, A. Weigand, G. Weigelhofer, D. Zak, and A. Zoppini. 2021.  
1736 Towards an improved understanding of biogeochemical processes across surface-  
1737 groundwater interactions in intermittent rivers and ephemeral streams. Earth-Science  
1738 Reviews 220:103724.

1739 González, E., A. A. Sher, E. Tabacchi, A. Masip, and M. Poulin. 2015. Restoration of riparian  
1740 vegetation: A global review of implementation and evaluation approaches in the  
1741 international, peer-reviewed literature. Journal of Environmental Management 158:85–  
1742 94.

1743 Guimond, J. A., and H. A. Michael. 2021. Effects of marsh migration on flooding, saltwater  
1744 intrusion, and crop yield in coastal agricultural land subject to storm surge inundation.  
1745 Water Resources Research 57.

1746 Hale, R. L., L. Turnbull, S. R. Earl, D. L. Childers, and N. B. Grimm. 2015. Stormwater  
1747 infrastructure controls runoff and dissolved material export from arid urban watersheds.

1748 Ecosystems 18:62–75.

1749 Hamilton, S. K., S. J. Sippel, and J. M. Melack. 2002. Comparison of inundation patterns among  
1750 major South American floodplains. *Journal of Geophysical Research: Atmospheres*  
1751 107:LBA 5-1-LBA 5-14.

1752 Hammond, J. C., M. Zimmer, M. Shanafield, K. Kaiser, S. E. Godsey, M. C. Mims, S. C. Zipper,  
1753 R. M. Burrows, S. K. Kampf, W. Dodds, C. N. Jones, C. A. Krabbenhoft, K. S. Boersma,  
1754 T. Datry, J. D. Olden, G. H. Allen, A. N. Price, K. Costigan, R. Hale, A. S. Ward, and D.  
1755 C. Allen. 2021. Spatial patterns and drivers of nonperennial flow regimes in the  
1756 contiguous United States. *Geophysical Research Letters* 48:e2020GL090794.

1757 Hanson, P. J., J. S. Riggs, W. R. Nettles, J. R. Phillips, M. B. Krassovski, L. A. Hook, L. Gu, A.  
1758 D. Richardson, D. M. Aubrecht, D. M. Ricciuto, J. M. Warren, and C. Barbier. 2017.  
1759 Attaining whole-ecosystem warming using air and deep-soil heating methods with an  
1760 elevated CO<sub>2</sub> atmosphere. *Biogeosciences* 14:861–883.

1761 Hayashi, M., G. Van Der Kamp, and D. O. Rosenberry. 2016. Hydrology of prairie wetlands:  
1762 Understanding the integrated surface-water and groundwater processes. *Wetlands*  
1763 36:237–254.

1764 Herbert, E. R., J. Schubauer-Berigan, and C. B. Craft. 2018. Differential effects of chronic and  
1765 acute simulated seawater intrusion on tidal freshwater marsh carbon cycling.  
1766 *Biogeochemistry* 138:137–154.

1767 Herndon, E. M., A. L. Dere, P. L. Sullivan, D. Norris, B. Reynolds, and S. L. Brantley. 2015.  
1768 Landscape heterogeneity drives contrasting concentration–discharge relationships in  
1769 shale headwater catchments. *Hydrology and Earth System Sciences* 19:3333–3347.

1770 Herndon, E. M., G. Steinhoefel, A. L. D. Dere, and P. L. Sullivan. 2018. Perennial flow through  
1771 convergent hillslopes explains chemodynamic solute behavior in a shale headwater  
1772 catchment. *Chemical Geology* 493:413–425.

1773 Herzon, I., and J. Helenius. 2008. Agricultural drainage ditches, their biological importance and



1774 functioning. *Biological Conservation* 141:1171–1183.

1775 Hess, L. L., J. M. Melack, A. G. Affonso, C. Barbosa, M. Gastil-Buhl, and E. M. L. M. Novo.  
1776 2015. Wetlands of the Lowland Amazon Basin: Extent, vegetative cover, and dual-  
1777 season inundated area as mapped with JERS-1 synthetic aperture radar. *Wetlands*  
1778 35:745–756.

1779 Hill, M. J., H. M. Greaves, C. D. Sayer, C. Hassall, M. Milin, V. S. Milner, L. Marazzi, R. Hall, L.  
1780 R. Harper, I. Thornhill, R. Walton, J. Biggs, N. Ewald, A. Law, N. Willby, J. C. White, R.  
1781 A. Briers, K. L. Mathers, M. J. Jeffries, and P. J. Wood. 2021. Pond ecology and  
1782 conservation: research priorities and knowledge gaps. *Ecosphere* 12:e03853.

1783 Hill, M. J., C. Hassall, B. Oertli, L. Fahrig, B. J. Robson, J. Biggs, M. J. Samways, N. Usio, N.  
1784 Takamura, J. Krishnaswamy, and P. J. Wood. 2018. New policy directions for global  
1785 pond conservation. *Conservation Letters* 11:e12447.

1786 Hinkel, J., D. Lincke, A. T. Vafeidis, M. Perrette, R. J. Nicholls, R. S. J. Tol, B. Marzeion, X.  
1787 Fettweis, C. Ionescu, and A. Levermann. 2014. Coastal flood damage and adaptation  
1788 costs under 21st century sea-level rise. *Proceedings of the National Academy of*  
1789 *Sciences* 111:3292–3297.

1790 Hinshaw, S. E., C. Tatariw, N. Flournoy, A. Kleinhuizen, C. Taylor, P. A. Sobecky, and B.  
1791 Mortazavi. 2017. Vegetation loss decreases salt marsh denitrification capacity:  
1792 Implications for marsh erosion. *Environmental Science & Technology* 51:8245–8253.

1793 Hladyz, S., S. C. Watkins, K. L. Whitworth, and D. S. Baldwin. 2011. Flows and hypoxic  
1794 blackwater events in managed ephemeral river channels. *Journal of Hydrology* 401:117–  
1795 125.

1796 Hofmeister, K. L., S. L. Eggert, B. J. Palik, D. Morley, E. Creighton, M. Rye, and R. K. Kolka.  
1797 2022. The identification, mapping, and management of seasonal ponds in forests of the  
1798 Great Lakes Region. *Wetlands* 42:9.

1799 Hondula, K. L., B. DeVries, C. N. Jones, and M. A. Palmer. 2021a. Effects of using high

1800 resolution satellite-based inundation time series to estimate methane fluxes from  
1801 forested wetlands. *Geophysical Research Letters* 48:e2021GL092556.

1802 Hondula, K. L., C. N. Jones, and M. A. Palmer. 2021b. Effects of seasonal inundation on  
1803 methane fluxes from forested freshwater wetlands. *Environmental Research Letters*  
1804 16:084016.

1805 Hooley-Underwood, Z. E., S. B. Stevens, N. R. Salinas, and K. G. Thompson. 2019. An  
1806 intermittent stream supports extensive spawning of large-river native fishes.  
1807 *Transactions of the American Fisheries Society* 148:426–441.

1808 Hopple, A. M., K. O. Doro, V. L. Bailey, B. Bond-Lamberty, N. McDowell, K. A. Morris, A. Myers-  
1809 Pigg, S. C. Pennington, P. Regier, R. Rich, A. Sengupta, R. Smith, J. Stegen, N. D.  
1810 Ward, S. C. Woodard, and J. P. Megonigal. 2023. Attaining freshwater and estuarine-  
1811 water soil saturation in an ecosystem-scale coastal flooding experiment. *Environmental*  
1812 *Monitoring and Assessment* 195:425.

1813 Hopple, A. M., S. C. Pennington, J. P. Megonigal, V. Bailey, and B. Bond-Lamberty. 2022.  
1814 Disturbance legacies regulate coastal forest soil stability to changing salinity and  
1815 inundation: A soil transplant experiment. *Soil Biology and Biochemistry* 169:108675.

1816 Horton, R. E. 1940. An approach toward a physical interpretation of infiltration capacity. Page 24  
1817 *Soil science Society of America proceedings*. Madison.

1818 Houser, C., and S. Hamilton. 2009. Sensitivity of post-hurricane beach and dune recovery to  
1819 event frequency. *Earth Surface Processes and Landforms* 34:613–628.

1820 Huang, C., C. Gascuel-Oudou, and S. Cros-Cayot. 2002. Hillslope topographic and hydrologic  
1821 effects on overland flow and erosion. *CATENA* 46:177–188.

1822 Huang, W., K. Wang, C. Ye, W. C. Hockaday, G. Wang, and S. J. Hall. 2021. High carbon  
1823 losses from oxygen-limited soils challenge biogeochemical theory and model  
1824 assumptions. *Global Change Biology* 27:6166–6180.

1825 Hutchinson, G. E. 1978. *Introduction to population ecology*. Yale University Press, New Haven,

1826 CT, USA.

1827 Hwang, T., L. E. Band, J. M. Vose, and C. Tague. 2012. Ecosystem processes at the watershed  
1828 scale: Hydrologic vegetation gradient as an indicator for lateral hydrologic connectivity of  
1829 headwater catchments. *Water Resources Research* 48.

1830 Ivory, S. J., M. M. McGlue, S. Spera, A. Silva, and I. Bergier. 2019. Vegetation, rainfall, and  
1831 pulsing hydrology in the Pantanal, the world's largest tropical wetland. *Environmental*  
1832 *Research Letters* 14:124017.

1833 Jarecke, K. M., T. D. Loecke, and A. J. Burgin. 2016. Coupled soil oxygen and greenhouse gas  
1834 dynamics under variable hydrology. *Soil Biology and Biochemistry* 95:164–172.

1835 Jeffries, M. 2008. The spatial and temporal heterogeneity of macrophyte communities in thirty  
1836 small, temporary ponds over a period of ten years. *Ecography* 31:765–775.

1837 Jones, C. N., G. R. Evenson, D. L. McLaughlin, M. K. Vanderhoof, M. W. Lang, G. W. McCarty,  
1838 H. E. Golden, C. R. Lane, and L. C. Alexander. 2018. Estimating restorable wetland  
1839 water storage at landscape scales. *Hydrological Processes* 32:305–313.

1840 Jones, J. 2019. Improved automated detection of subpixel-scale inundation: Revised dynamic  
1841 surface water extent (DSWE) partial surface water tests. *Remote Sensing* 11:374.

1842 Junk, W., P. Bayley, and R. Sparks. 1989. The flood pulse concept in river-floodplain systems.  
1843 *Page Can. Spec. Public Fish. Aquat. Sci.*

1844 Kampf, S. K., K. A. Dwire, M. P. Fairchild, J. Dunham, C. D. Snyder, K. L. Jaeger, C. H. Luce, J.  
1845 C. Hammond, C. Wilson, M. A. Zimmer, and M. Sidell. 2021. Managing nonperennial  
1846 headwater streams in temperate forests of the United States. *Forest Ecology and*  
1847 *Management* 497:119523.

1848 Kaushal, S. S., J. E. Reimer, P. M. Mayer, R. R. Shatkay, C. M. Maas, W. D. Nguyen, W. L.  
1849 Boger, A. M. Yaculak, T. R. Doody, M. J. Pennino, N. W. Bailey, J. G. Galella, A.  
1850 Weingrad, D. C. Collison, K. L. Wood, S. Haq, T. A. Newcomer-Johnson, S. Duan, and  
1851 K. T. Belt. 2022. Freshwater salinization syndrome alters retention and release

1852 of chemical cocktails along flowpaths: From stormwater management to  
1853 urban streams. *Freshwater Science* 41:420–441.

1854 Kirkby, M., L. Bracken, and S. Reaney. 2002. The influence of land use, soils and topography  
1855 on the delivery of hillslope runoff to channels in SE Spain. *Earth Surface Processes and*  
1856 *Landforms* 27:1459–1473.

1857 Kirwan, M. L., and K. B. Gedan. 2019. Sea-level driven land conversion and the formation of  
1858 ghost forests. *Nature Climate Change* 9:450–457.

1859 Kiss, T., J. Nagy, I. Fehérváry, and C. Vaszkó. 2019. (Mis) management of floodplain  
1860 vegetation: The effect of invasive species on vegetation roughness and flood levels.  
1861 *Science of The Total Environment* 686:931–945.

1862 Kneitel, J. M. 2014. Inundation timing, more than duration, affects the community structure of  
1863 California vernal pool mesocosms. *Hydrobiologia* 732:71–83.

1864 Kollet, S. J., and R. M. Maxwell. 2006. Integrated surface–groundwater flow modeling: A free-  
1865 surface overland flow boundary condition in a parallel groundwater flow model.  
1866 *Advances in Water Resources* 29:945–958.

1867 Konapala, G., A. K. Mishra, Y. Wada, and M. E. Mann. 2020. Climate change will affect global  
1868 water availability through compounding changes in seasonal precipitation and  
1869 evaporation. *Nature Communications* 11:3044.

1870 Koschorreck, M., A. S. Downing, J. Hejzlar, R. Marcé, A. Laas, W. G. Arndt, P. S. Keller, A. J. P.  
1871 Smolders, G. van Dijk, and S. Kosten. 2020. Hidden treasures: Human-made aquatic  
1872 ecosystems harbour unexplored opportunities. *Ambio* 49:531–540.

1873 Krabbenhoft, C. A., G. H. Allen, P. Lin, S. E. Godsey, D. C. Allen, R. M. Burrows, A. G.  
1874 DeIvecchia, K. M. Fritz, M. Shanafield, A. J. Burgin, M. A. Zimmer, T. Datry, W. K.  
1875 Dodds, C. N. Jones, M. C. Mims, C. Franklin, J. C. Hammond, S. Zipper, A. S. Ward, K.  
1876 H. Costigan, H. E. Beck, and J. D. Olden. 2022. Assessing placement bias of the global  
1877 river gauge network. *Nature Sustainability* 5:586–592.

1878 Kundel, D., S. Meyer, H. Birkhofer, A. Fließbach, P. Mäder, S. Scheu, M. van Kleunen, and K.  
1879 Birkhofer. 2018. Design and manual to construct rainout-shelters for climate change  
1880 experiments in agroecosystems. *Frontiers in Environmental Science* 6.

1881 Ladau, J., and E. A. Elie-Fadrosh. 2019. Spatial, temporal, and phylogenetic scales of microbial  
1882 ecology. *Trends in Microbiology* 27:662–669.

1883 Lalli, K., S. Soenen, J. B. Fisher, J. McGlinchy, T. Kleynhans, R. Eon, and L. M. Moreau. 2022.  
1884 VanZyl-1: demonstrating SmallSat measurement capabilities for land surface  
1885 temperature and evapotranspiration. Page 8 in C. D. Norton and S. R. Babu, editors.  
1886 *CubeSats and SmallSats for Remote Sensing VI*. SPIE, San Diego, United States.

1887 Lane, C. R., and E. D’Amico. 2016. Identification of putative geographically isolated wetlands of  
1888 the conterminous United States. *JAWRA Journal of the American Water Resources*  
1889 *Association* 52:705–722.

1890 Lane, K., K. Charles-Guzman, K. Wheeler, Z. Abid, N. Graber, and T. Matte. 2013. Health  
1891 effects of coastal storms and flooding in urban areas: A review and vulnerability  
1892 assessment. *Journal of Environmental and Public Health* 2013:1–13.

1893 Laronne, J. B., and L. Reid. 1993. Very high rates of bedload sediment transport by ephemeral  
1894 desert rivers. *Nature* 366:148–150.

1895 Larsen, L., N. Aumen, C. Bernhardt, V. Engel, T. Givnish, S. Hagerthey, J. Harvey, L. Leonard,  
1896 P. McCormick, C. Mcvoy, G. Noe, M. Nungesser, K. Rutchey, F. Sklar, T. Troxler, J.  
1897 Volin, and D. Willard. 2011. Recent and historic drivers of landscape change in the  
1898 Everglades Ridge, slough, and tree island mosaic. *Critical Reviews in Environmental*  
1899 *Science and Technology* 41:344–381.

1900 Lei, Y., C. Liu, L. Zhang, and S. Luo. 2016. How smallholder farmers adapt to agricultural  
1901 drought in a changing climate: A case study in southern China. *Land Use Policy* 55:300–  
1902 308.

1903 Li, L., K. Maher, A. Navarre-Sitchler, J. Druhan, C. Meile, C. Lawrence, J. Moore, J. Perdrial, P.

1904 Sullivan, A. Thompson, L. Jin, E. W. Bolton, S. L. Brantley, W. E. Dietrich, K. U. Mayer,  
1905 C. I. Steefel, A. Valocchi, J. Zachara, B. Kocar, J. McIntosh, B. M. Tutolo, M. Kumar, E.  
1906 Sonnenthal, C. Bao, and J. Beisman. 2017. Expanding the role of reactive transport  
1907 models in critical zone processes. *Earth-Science Reviews* 165:280–301.

1908 Li, L., P. L. Sullivan, P. Benettin, O. A. Cirpka, K. Bishop, S. L. Brantley, J. L. A. Knapp, I.  
1909 Meerveld, A. Rinaldo, J. Seibert, H. Wen, and J. W. Kirchner. 2021. Toward catchment  
1910 hydro-biogeochemical theories. *WIREs Water* 8.

1911 Li, S., G. Wang, C. Zhu, J. Lu, W. Ullah, D. F. T. Hagan, G. Kattel, and J. Peng. 2022a.  
1912 Attribution of global evapotranspiration trends based on the Budyko framework.  
1913 *Hydrology and Earth System Sciences* 26:3691–3707.

1914 Li, Z., S. Gao, M. Chen, J. J. Gourley, and Y. Hong. 2022b. Spatiotemporal characteristics of  
1915 US floods: Current status and forecast under a future warmer climate. *Earth's Future*  
1916 10:e2022EF002700.

1917 Liberato, M. L. R., J. G. Pinto, R. M. Trigo, P. Ludwig, P. Ordóñez, D. Yuen, and I. F. Trigo.  
1918 2013. Explosive development of winter storm Xynthia over the subtropical North Atlantic  
1919 Ocean. *Natural Hazards and Earth System Sciences* 13:2239–2251.

1920 Lisenby, P. E., S. Tooth, and T. J. Ralph. 2019. Product vs. process? The role of  
1921 geomorphology in wetland characterization. *Science of The Total Environment* 663:980–  
1922 991.

1923 Lohse, K. A., P. D. Brooks, J. C. McIntosh, T. Meixner, and T. E. Huxman. 2009. Interactions  
1924 between biogeochemistry and hydrologic systems. *Annual Review of Environment and*  
1925 *Resources* 34:65–96.

1926 Londe, D. W., D. Dvoretz, C. A. Davis, S. R. Loss, and E. P. Robertson. 2022. Inundation of  
1927 depressional wetlands declines under a changing climate. *Climatic Change* 172:27.

1928 Lovelock, C. E., and R. Reef. 2020. Variable impacts of climate change on blue carbon. *One*  
1929 *Earth* 3:195–211.

- 1930 Lugo, A. E. 2008. Visible and invisible effects of hurricanes on forest ecosystems: an  
1931 international review. *Austral Ecology* 33:368–398.
- 1932 Luijendijk, A., G. Hagenars, R. Ranasinghe, F. Baart, G. Donchyts, and S. Aarninkhof. 2018.  
1933 The state of the world's beaches. *Scientific Reports* 8:6641.
- 1934 Mahecha, M. D., M. Reichstein, N. Carvalhais, G. Lasslop, H. Lange, S. I. Seneviratne, R.  
1935 Vargas, C. Ammann, M. A. Arain, A. Cescatti, I. A. Janssens, M. Migliavacca, L.  
1936 Montagnani, and A. D. Richardson. 2010. Global convergence in the temperature  
1937 sensitivity of respiration at ecosystem level. *Science* 329:838–840.
- 1938 Mandishona, E., and J. Knight. 2022. Inland wetlands in Africa: A review of their typologies and  
1939 ecosystem services. *Progress in Physical Geography: Earth and Environment* 46:547–  
1940 565.
- 1941 Manfreda, S., M. F. McCabe, P. E. Miller, R. Lucas, V. Pajuelo Madrigal, G. Mallinis, E. Ben  
1942 Dor, D. Helman, L. Estes, G. Ciraolo, J. Müllerová, F. Tauro, M. I. De Lima, J. L. M. P.  
1943 De Lima, A. Maltese, F. Frances, K. Caylor, M. Kohv, M. Perks, G. Ruiz-Pérez, Z. Su, G.  
1944 Vico, and B. Toth. 2018. On the use of unmanned aerial systems for environmental  
1945 monitoring. *Remote Sensing* 10:641.
- 1946 Maris, S. C., M. R. Teira-Esmatges, and M. M. Català. 2016. Influence of irrigation frequency on  
1947 greenhouse gases emission from a paddy soil. *Paddy and Water Environment* 14:199–  
1948 210.
- 1949 Marton, J. M., I. F. Creed, D. B. Lewis, C. R. Lane, N. B. Basu, M. J. Cohen, and C. B. Craft.  
1950 2015. Geographically isolated wetlands are important biogeochemical reactors on the  
1951 landscape. *BioScience* 65:408–418.
- 1952 Matthews, J. 2010. Anthropogenic climate change impacts on ponds: a thermal mass  
1953 perspective. *BioRisk* 5:193–209.
- 1954 Maul, G. A., and I. W. Duedall. 2019. Demography of coastal populations. Pages 692–700 *in* C.  
1955 W. Finkl and C. Makowski, editors. *Encyclopedia of Coastal Science*. Springer

- 1956 International Publishing, Cham.
- 1957 McCarthy, J. K., J. M. Dwyer, and K. Mokany. 2019. A regional-scale assessment of using  
1958 metabolic scaling theory to predict ecosystem properties. *Proceedings of the Royal  
1959 Society B: Biological Sciences* 286:20192221.
- 1960 McClain, M. E., E. W. Boyer, C. L. Dent, S. E. Gergel, N. B. Grimm, P. M. Groffman, S. C. Hart,  
1961 J. W. Harvey, C. A. Johnston, E. Mayorga, W. H. McDowell, and G. Pinay. 2003.  
1962 Biogeochemical hot spots and hot moments at the interface of terrestrial and aquatic  
1963 ecosystems. *Ecosystems* 6:301–312.
- 1964 McDaniell, P. A., M. P. Regan, E. Brooks, J. Boll, S. Barndt, A. Falen, S. K. Young, and J. E.  
1965 Hammel. 2008. Linking fragipans, perched water tables, and catchment-scale  
1966 hydrological processes. *CATENA* 73:166–173.
- 1967 McDonnell, J. J. 2009. Hewlett, J.D. and Hibbert, A.R. 1967: Factors affecting the response of  
1968 small watersheds to precipitation in humid areas. In Sopper, W.E. and Lull, H.W.,  
1969 editors, *Forest hydrology*, New York: Pergamon Press, 275—90. *Progress in Physical  
1970 Geography: Earth and Environment* 33:288–293.
- 1971 McDowell, N. G., K. Anderson-Teixeira, J. A. Biederman, D. D. Breshears, Y. Fang, L.  
1972 Fernández-de-Uña, E. B. Graham, D. S. Mackay, J. J. McDonnell, G. W. Moore, M. F.  
1973 Nehemy, C. S. Stevens Rumann, J. Stegen, N. Tague, M. G. Turner, and X. Chen.  
1974 2023. Ecohydrological decoupling under changing disturbances and climate. *One Earth*  
1975 6:251–266.
- 1976 McDowell, N. G., M. Ball, B. Bond-Lamberty, M. L. Kirwan, K. W. Krauss, J. P. Megonigal, M.  
1977 Mencuccini, N. D. Ward, M. N. Weintraub, and V. Bailey. 2022. Processes and  
1978 mechanisms of coastal woody-plant mortality. *Global Change Biology* 28:5881–5900.
- 1979 McGrane, S. J. 2016. Impacts of urbanisation on hydrological and water quality dynamics, and  
1980 urban water management: a review. *Hydrological Sciences Journal* 61:2295–2311.
- 1981 McGuire, K. J., J. J. McDonnell, M. Weiler, C. Kendall, B. L. McGlynn, J. M. Welker, and J.



- 1982 Seibert. 2005. The role of topography on catchment-scale water residence time. *Water*
- 1983 *Resources Research* 41.
- 1984 Mcleod, E., G. L. Chmura, S. Bouillon, R. Salm, M. Björk, C. M. Duarte, C. E. Lovelock, W. H.
- 1985 Schlesinger, and B. R. Silliman. 2011. A blueprint for blue carbon: toward an improved
- 1986 understanding of the role of vegetated coastal habitats in sequestering CO<sub>2</sub>. *Frontiers in*
- 1987 *Ecology and the Environment* 9:552–560.
- 1988 McVicar, T. R., T. G. Van Niel, L. Li, M. F. Hutchinson, X. Mu, and Z. Liu. 2007. Spatially
- 1989 distributing monthly reference evapotranspiration and pan evaporation considering
- 1990 topographic influences. *Journal of Hydrology* 338:196–220.
- 1991 Melton, J. R., R. Wania, E. L. Hodson, B. Poulter, B. Ringeval, R. Spahni, T. Bohn, C. A. Avis,
- 1992 D. J. Beerling, G. Chen, A. V. Eliseev, S. N. Denisov, P. O. Hopcroft, D. P. Lettenmaier,
- 1993 W. J. Riley, J. S. Singarayer, Z. M. Subin, H. Tian, S. Zürcher, V. Brovkin, P. M. Van
- 1994 Bodegom, T. Kleinen, Z. C. Yu, and J. O. Kaplan. 2013. Present state of global wetland
- 1995 extent and wetland methane modelling: conclusions from a model inter-comparison
- 1996 project (WETCHIMP). *Biogeosciences* 10:753–788.
- 1997 Merritt, D. M., and E. E. Wohl. 2002. Processes governing hydrochory along rivers: Hydraulics,
- 1998 hydrology and dispersal phenology. *Ecological Applications* 12:1071–1087.
- 1999 Mertes, L. A. K. 2011. Inland flood hazards: Human, riparian, and aquatic communities. Pages
- 2000 145–166 *Inundation hydrology*. Cambridge University Press, Cambridge, UK.
- 2001 Messenger, M. L., B. Lehner, C. Cockburn, N. Lamouroux, H. Pella, T. Snelder, K. Tockner, T.
- 2002 Trautmann, C. Watt, and T. Datry. 2021. Global prevalence of non-perennial rivers and
- 2003 streams. *Nature* 594:391–397.
- 2004 Molins, S., D. Svyatsky, Z. Xu, E. T. Coon, and J. D. Moulton. 2022. A multicomponent reactive
- 2005 transport model for integrated surface-subsurface hydrology problems. *Water Resources*
- 2006 *Research* 58:e2022WR032074.
- 2007 Moomaw, W. R., G. L. Chmura, G. T. Davies, C. M. Finlayson, B. A. Middleton, S. M. Natali, J.

2008 E. Perry, N. Roulet, and A. E. Sutton-Grier. 2018. Wetlands In a changing climate:  
2009 Science, policy and management. *Wetlands* 38:183–205.

2010 Morrissey, E. M., and R. B. Franklin. 2015. Evolutionary history influences the salinity  
2011 preference of bacterial taxa in wetland soils. *Frontiers in Microbiology* 6.

2012 Murray, N. J., P. Bunting, R. F. Canto, L. Hilarides, E. V. Kennedy, R. M. Lucas, M. B. Lyons, A.  
2013 Navarro, C. M. Roelfsema, A. Rosenqvist, M. D. Spalding, M. Toor, and T. A.  
2014 Worthington. 2022a. coastTrain: A global reference library for coastal ecosystems.  
2015 *Remote Sensing* 14:5766.

2016 Murray, N. J., T. A. Worthington, P. Bunting, S. Duce, V. Hagger, C. E. Lovelock, R. Lucas, M. I.  
2017 Saunders, M. Sheaves, M. Spalding, N. J. Waltham, and M. B. Lyons. 2022b. High-  
2018 resolution mapping of losses and gains of Earth’s tidal wetlands. *Science* 376:744–749.

2019 Nanson, G., and J. Croke. 1992. A genetic classification of floodplains. *Faculty of Science,*  
2020 *Medicine and Health - Papers: part A*:459–486.

2021 Nelson, T. M., C. Streten, K. S. Gibb, and A. A. Chariton. 2015. Saltwater intrusion history  
2022 shapes the response of bacterial communities upon rehydration. *Science of The Total*  
2023 *Environment* 502:143–148.

2024 Neubauer, S. C., and J. P. Megonigal. 2015. Moving beyond global warming potentials to  
2025 quantify the climatic role of ecosystems. *Ecosystems* 18:1000–1013.

2026 Ode, P. R., A. E. Fetscher, and L. B. Busse. 2016. Standard operating procedures for the  
2027 collection of field data for bioassessments of California wadeable streams: Benthic  
2028 macroinvertebrates, algae, and physical habitat. Page 80. California State Water  
2029 Resources Control Board Surface Water Ambient Monitoring Program (SWAMP)  
2030 Bioassessment SOP 004.

2031 O’Mara, K., J. M. Olley, B. Fry, and M. Burford. 2019. Catchment soils supply ammonium to the  
2032 coastal zone - Flood impacts on nutrient flux in estuaries. *Science of The Total*  
2033 *Environment* 654:583–592.

2034 O'Meara, T. A., J. R. Hillman, and S. F. Thrush. 2017. Rising tides, cumulative impacts and  
2035 cascading changes to estuarine ecosystem functions. *Scientific Reports* 7:10218.

2036 Orozco-López, E., R. Muñoz-Carpena, B. Gao, and G. A. Fox. 2018. Riparian Vadose Zone  
2037 preferential flow: Review of concepts, limitations, and perspectives. *Vadose Zone*  
2038 *Journal* 17:180031.

2039 Palmer, M. A., and K. L. Hondula. 2014. Restoration as mitigation: Analysis of stream mitigation  
2040 for coal mining impacts in Southern Appalachia. *Environmental Science & Technology*  
2041 48:10552–10560.

2042 Palmer, M. A., K. L. Hondula, and B. J. Koch. 2014. Ecological restoration of streams and rivers:  
2043 Shifting strategies and shifting goals. *Annual Review of Ecology, Evolution, and*  
2044 *Systematics* 45:247–269.

2045 Palta, M. M., J. G. Ehrenfeld, and P. M. Groffman. 2014. “Hotspots” and “Hot Moments” of  
2046 denitrification in urban Brownfield Wetlands. *Ecosystems* 17:1121–1137.

2047 Palta, M. M., N. B. Grimm, and P. M. Groffman. 2017. “Accidental” urban wetlands: Ecosystem  
2048 functions in unexpected places. *Frontiers in Ecology and the Environment* 15:248–256.

2049 Pan, J., Y. Liu, X. Zhong, R. M. Lampayan, G. R. Singleton, N. Huang, K. Liang, B. Peng, and  
2050 K. Tian. 2017. Grain yield, water productivity and nitrogen use efficiency of rice under  
2051 different water management and fertilizer-N inputs in South China. *Agricultural Water*  
2052 *Management* 184:191–200.

2053 Pascolini-Campbell, M., J. B. Fisher, and J. T. Reager. 2021. GRACE-FO and ECOSTRESS  
2054 synergies constrain fine-scale impacts on the water balance. *Geophysical Research*  
2055 *Letters* 48.

2056 Patel, K. F., K. A. Rod, J. Zheng, P. J. Regier, F. Machado-Silva, B. Bond-Lamberty, X. Chen,  
2057 D. Day, K. O. Doro, M. Kaufman, M. Kovach, N. McDowell, S. A. McKeever, P. J.  
2058 Megonigal, C. G. Norris, T. O'Meara, R. Rich, P. Thornton, K. M. Kemner, N. D. Ward,  
2059 M. N. Weintraub, and V. L. Bailey. 2023. Time to anoxia: Observations and predictions

2060 of oxygen drawdown following coastal flood events.

2061 Patel, K. F., C. Tatariw, J. D. MacRae, T. Ohno, S. J. Nelson, and I. J. Fernandez. 2020.

2062 Snowmelt periods as hot moments for soil N dynamics: a case study in Maine, USA.

2063 Environmental Monitoring and Assessment 192:777.

2064 Patel, N., S. Gahlaud, A. Saxena, B. Thakur, N. Bharti, A. Dabhi, R. Bhushan, and R. Agnihotri.

2065 2022. Revised chronology and stable isotopic (carbon and nitrogen) characterization of

2066 Lahuradewa lake sediment (Ganga-plain, India): Insights into biogeochemistry leading to

2067 peat formation in the lake. Journal of the Palaeontological Society of India Volume

2068 67(1):113–125.

2069 Peacock, M., J. Audet, D. Bastviken, M. N. Futter, V. Gauci, A. Grinham, J. A. Harrison, M. S.

2070 Kent, S. Kosten, C. E. Lovelock, A. J. Veraart, and C. D. Evans. 2021. Global

2071 importance of methane emissions from drainage ditches and canals. Environmental

2072 Research Letters 16.

2073 Pedersen, O., M. Sauter, T. D. Colmer, and M. Nakazono. 2021. Regulation of root adaptive

2074 anatomical and morphological traits during low soil oxygen. New Phytologist 229:42–49.

2075 Pekel, J.-F., A. Cottam, N. Gorelick, and A. S. Belward. 2016. High-resolution mapping of global

2076 surface water and its long-term changes. Nature 540:418–422.

2077 Peng, S., X. Lin, R. L. Thompson, Y. Xi, G. Liu, D. Hauglustaine, X. Lan, B. Poulter, M.

2078 Ramonet, M. Saunois, Y. Yin, Z. Zhang, B. Zheng, and P. Ciais. 2022. Wetland emission

2079 and atmospheric sink changes explain methane growth in 2020. Nature 612:477–482.

2080 Perks, M. T., A. J. Russell, and A. R. G. Large. 2016. Technical Note: Advances in flash flood

2081 monitoring using unmanned aerial vehicles (UAVs). Hydrology and Earth System

2082 Sciences 20:4005–4015.

2083 Peruccacci, S., M. T. Brunetti, S. L. Gariano, M. Melillo, M. Rossi, and F. Guzzetti. 2017.

2084 Rainfall thresholds for possible landslide occurrence in Italy. Geomorphology 290:39–57.

2085 Pezeshki, S. R., and R. D. DeLaune. 2012. Soil oxidation-reduction in wetlands and Its impact

2086 on plant functioning. *Biology* 1:196–221.

2087 Pickering, M. D., K. J. Horsburgh, J. R. Blundell, J. J.-M. Hirschi, R. J. Nicholls, M. Verlaan, and  
2088 N. C. Wells. 2017. The impact of future sea-level rise on the global tides. *Continental*  
2089 *Shelf Research* 142:50–68.

2090 Plum, N. 2005. Terrestrial invertebrates in flooded grassland: A literature review. *Wetlands*  
2091 25:721–737.

2092 Pool, S., F. Francés, A. Garcia-Prats, M. Pulido-Velazquez, C. Sanchis-Ibor, M. Schirmer, H.  
2093 Yang, and J. Jiménez-Martínez. 2021. From flood to drip irrigation under climate change:  
2094 Impacts on evapotranspiration and groundwater recharge in the mediterranean region of  
2095 Valencia (Spain). *Earth's Future* 9:e2020EF001859.

2096 Popper, K. R. 2014. *Conjectures and refutations: the growth of scientific knowledge*. Repr.  
2097 Routledge, London.

2098 Price, A. N., C. N. Jones, J. C. Hammond, M. A. Zimmer, and S. C. Zipper. 2021. The drying  
2099 regimes of non-perennial rivers and streams. *Geophysical Research Letters*  
2100 48:e2021GL093298.

2101 Pumo, D., D. Caracciolo, F. Viola, and L. V. Noto. 2016. Climate change effects on the  
2102 hydrological regime of small non-perennial river basins. *Science of The Total*  
2103 *Environment* 542:76–92.

2104 Quinn, J. D., P. M. Reed, M. Giuliani, A. Castelletti, J. W. Oyler, and R. E. Nicholas. 2018.  
2105 Exploring how changing monsoonal dynamics and human pressures challenge  
2106 multireservoir management for flood protection, hydropower production, and agricultural  
2107 water supply. *Water Resources Research* 54:4638–4662.

2108 Rameshwaran, P., V. A. Bell, H. N. Davies, and A. L. Kay. 2021. How might climate change  
2109 affect river flows across West Africa? *Climatic Change* 169:21.

2110 Rasmussen, T. C., J. B. Deemy, and S. L. Long. 2016. Wetland Hydrology. Pages 1–16 *in* C. M.  
2111 Finlayson, M. Everard, K. Irvine, R. J. McInnes, B. A. Middleton, A. A. Van Dam, and N.

2112 C. Davidson, editors. *The Wetland Book*. Springer Netherlands, Dordrecht.

2113 Regier, P., N. D. Ward, J. Indivero, C. Wiese Moore, M. Norwood, and A. Myers-Pigg. 2021.

2114 Biogeochemical control points of connectivity between a tidal creek and its floodplain.

2115 *Limnology and Oceanography Letters* 6:134–142.

2116 Reichstein, M., G. Camps-Valls, B. Stevens, M. Jung, J. Denzler, N. Carvalhais, and Prabhat.

2117 2019. Deep learning and process understanding for data-driven Earth system science.

2118 *Nature* 566:195–204.

2119 Reis, V., V. Hermoso, S. K. Hamilton, D. Ward, E. Fluet-Chouinard, B. Lehner, and S. Linke.

2120 2017. A global assessment of inland wetland conservation status. *BioScience* 67:523–

2121 533.

2122 Reisinger, A. J., P. M. Groffman, and E. J. Rosi-Marshall. 2016. Nitrogen cycling process rates

2123 across urban ecosystems. *FEMS Microbiology Ecology* 92:fiw198.

2124 Renwick, W., R. Sleezer, R. Buddemeier, and S. Smith. 2006. Small artificial ponds in the

2125 United States: Impacts on sedimentation and carbon budget. Pages 738–744

2126 *Proceedings of the Eighth Federal Interagency Sedimentation Conference*.

2127 Resetarits, W. J. 1996. Oviposition site choice and life history evolution. *American Zoologist*

2128 36:205–215.

2129 Reverey, F., L. Ganzert, G. Lischeid, A. Ulrich, K. Premke, and H.-P. Grossart. 2018. Dry-wet

2130 cycles of kettle hole sediments leave a microbial and biogeochemical legacy. *Science of*

2131 *The Total Environment* 627:985–996.

2132 Ribolzi, O., J. Patin, L. M. Bresson, K. O. Latsachack, E. Mouche, O. Sengtaheuanghoung, N.

2133 Silvera, J. P. Thiébaux, and C. Valentin. 2011. Impact of slope gradient on soil surface

2134 features and infiltration on steep slopes in northern Laos. *Geomorphology* 127:53–63.

2135 Richardson, D. C., M. A. Holgerson, M. J. Farragher, K. K. Hoffman, K. B. S. King, M. B.

2136 Alfonso, M. R. Andersen, K. S. Cheruveil, K. A. Coleman, M. J. Farruggia, R. L.

2137 Fernandez, K. L. Hondula, G. A. López Moreira Mazacotte, K. Paul, B. L. Peierls, J. S.

2138 Rabaey, S. Sadro, M. L. Sánchez, R. L. Smyth, and J. N. Sweetman. 2022a. A  
2139 functional definition to distinguish ponds from lakes and wetlands. *Scientific Reports*  
2140 12:10472.

2141 Richardson, D. C., M. A. Holgerson, M. J. Farragher, K. K. Hoffman, K. B. S. King, M. B.  
2142 Alfonso, M. R. Andersen, K. S. Cheruveil, K. A. Coleman, M. J. Farruggia, R. L.  
2143 Fernandez, K. L. Hondula, G. A. López Moreira Mazacotte, K. Paul, B. L. Peierls, J. S.  
2144 Rabaey, S. Sadro, M. L. Sánchez, R. L. Smyth, and J. N. Sweetman. 2022b. A  
2145 functional definition to distinguish ponds from lakes and wetlands. *Scientific Reports*  
2146 12:10472.

2147 Richey, A. S., B. F. Thomas, M. Lo, J. T. Reager, J. S. Famiglietti, K. Voss, S. Swenson, and M.  
2148 Rodell. 2015. Quantifying renewable groundwater stress with GRACE. *Water Resources*  
2149 *Research* 51:5217–5238.

2150 Ripley, B. J., and M. A. Simovich. 2009. Species richness on islands in time: Variation in  
2151 ephemeral pond crustacean communities in relation to habitat duration and size.  
2152 *Hydrobiologia* 617:181–196.

2153 Robinson, C. T., K. Tockner, and J. V. Ward. 2002. The fauna of dynamic riverine landscapes:  
2154 Fauna of riverine landscapes. *Freshwater Biology* 47:661–677.

2155 Rosado, J., M. Morais, and K. Tockner. 2015. Mass dispersal of terrestrial organisms during first  
2156 flush events in a temporary stream: Mass dispersal of terrestrial organisms. *River*  
2157 *Research and Applications* 31:912–917.

2158 Rosentreter, J. A., A. V. Borges, B. R. Deemer, M. A. Holgerson, S. Liu, C. Song, J. Melack, P.  
2159 A. Raymond, C. M. Duarte, G. H. Allen, D. Olefeldt, B. Poulter, T. I. Battin, and B. D.  
2160 Eyre. 2021. Half of global methane emissions come from highly variable aquatic  
2161 ecosystem sources. *Nature Geoscience* 14:225–230.

2162 Ruel, J. J., and M. P. Ayres. 1999. Jensen's inequality predicts effects of environmental  
2163 variation. *Trends in Ecology & Evolution* 14:361–366.

2164 Rullens, V., S. Mangan, F. Stephenson, D. E. Clark, R. H. Bulmer, A. Berthelsen, J. Crawshaw,  
2165 R. V. Gladstone-Gallagher, S. Thomas, J. I. Ellis, and C. A. Pilditch. 2022.  
2166 Understanding the consequences of sea level rise: the ecological implications of losing  
2167 intertidal habitat. *New Zealand Journal of Marine and Freshwater Research* 56:353–370.

2168 Saadat, S., J. Frankenberger, L. Bowling, and S. Ale. 2020. Evaluation of surface ponding and  
2169 runoff generation in a seasonally frozen drained agricultural field. *Journal of Hydrology*  
2170 588:124985.

2171 Saltarelli, W. A., D. G. F. Cunha, A. Freixa, N. Perujo, J. C. López-Doval, V. Acuña, and S.  
2172 Sabater. 2022. Nutrient stream attenuation is altered by the duration and frequency of  
2173 flow intermittency. *Ecohydrology* 15:e2351.

2174 Sarremejane, R., H. Mykrä, N. Bonada, J. Aroviita, and T. Muotka. 2017. Habitat connectivity  
2175 and dispersal ability drive the assembly mechanisms of macroinvertebrate communities  
2176 in river networks. *Freshwater Biology* 62.

2177 Sarremejane, R., R. Stubbington, J. England, C. E. M. Sefton, M. Eastman, S. Parry, and A.  
2178 Ruhi. 2021. Drought effects on invertebrate metapopulation dynamics and quasi-  
2179 extinction risk in an intermittent river network. *Global Change Biology* 27:4024–4039.

2180 Schaffer-Smith, D., S. W. Myint, R. L. Muenich, D. Tong, and J. E. DeMeester. 2020. Repeated  
2181 hurricanes reveal risks and opportunities for social-ecological resilience to flooding and  
2182 water quality problems. *Environmental Science & Technology* 54:7194–7204.

2183 von Schiller, D., T. Datry, R. Corti, A. Foulquier, K. Tockner, R. Marcé, G. García-Baquero, I.  
2184 Odriozola, B. Obrador, A. Elosegi, C. Mendoza-Lera, M. O. Gessner, R. Stubbington, R.  
2185 Albariño, D. C. Allen, F. Altermatt, M. I. Arce, S. Arnon, D. Banas, A. Banegas-Medina,  
2186 E. Beller, M. L. Blanchette, J. F. Blanco-Libreros, J. Blessing, I. G. Boëchat, K. S.  
2187 Boersma, M. T. Bogan, N. Bonada, N. R. Bond, K. Brintrup, A. Bruder, R. M. Burrows, T.  
2188 Cancellario, S. M. Carlson, S. Cauvy-Fraunié, N. Cid, M. Danger, B. de Freitas Terra, A.  
2189 Dehedin, A. M. De Girolamo, R. del Campo, V. Díaz-Villanueva, C. P. Duerdoth, F. Dyer,



2190 E. Faye, C. Febria, R. Figueroa, B. Four, S. Gafny, R. Gómez, L. Gómez-Gener, M. a. S.  
2191 Graça, S. Guareschi, B. Gücker, F. Hoppeler, J. L. Hwan, S. Kubheka, A. Laini, S. D.  
2192 Langhans, C. Leigh, C. J. Little, S. Lorenz, J. Marshall, E. J. Martín, A. McIntosh, E. I.  
2193 Meyer, M. Miliša, M. C. Mlambo, M. Moleón, M. Morais, P. Negus, D. Niyogi, A.  
2194 Papatheodoulou, I. Pardo, P. Pařil, V. Peřić, C. Piscart, M. Polářek, P. Rodríguez-  
2195 Lozano, R. J. Rolls, M. M. Sánchez-Montoya, A. Savić, O. Shumilova, A. Steward, A.  
2196 Taleb, A. Uzan, R. Vander Vorste, N. Waltham, C. Woelfle-Erskine, D. Zak, C. Zarfl, and  
2197 A. Zoppini. 2019. Sediment respiration pulses in intermittent rivers and ephemeral  
2198 streams. *Global Biogeochemical Cycles* 33:1251–1263.

2199 Schimel, D. S., T. G. F. Kittel, and W. J. Parton. 1991. Terrestrial biogeochemical cycles: global  
2200 interactions with the atmosphere and hydrology. *Tellus A* 43:188–203.

2201 Schimel, J. P. 2018. Life in dry soils: Effects of drought on soil microbial communities and  
2202 processes. *Annual Review of Ecology, Evolution, and Systematics* 49:409–432.

2203 Schlesinger, W. H., and E. S. Bernhardt. 2020. The atmosphere. Pages 51–97  
2204 *Biogeochemistry*. Elsevier.

2205 Schumann, G. J.-P., and D. K. Moller. 2015. Microwave remote sensing of flood inundation.  
2206 *Physics and Chemistry of the Earth, Parts A/B/C* 83–84:84–95.

2207 Schuwirth, N., F. Borgwardt, S. Domisch, M. Friedrichs, M. Kattwinkel, D. Kneis, M.  
2208 Kuemmerlen, S. D. Langhans, J. Martínez-López, and P. Vermeiren. 2019. How to make  
2209 ecological models useful for environmental management. *Ecological Modelling*  
2210 411:108784.

2211 Scientific Investigations Report. 2015. . Scientific Investigations Report.

2212 Secretariat, R. 2016. An Introduction to the Convention on Wetlands (previously The Ramsar  
2213 Convention Manual). 7th edition.

2214 Semeniuk, C. A., and V. Semeniuk. 1995. A geomorphic approach to global classification for  
2215 inland wetlands. Pages 103–124 *Advances in Vegetation Science*.

2216 Semeniuk, C., and V. Semeniuk. 2011. A comprehensive classification of inland wetlands of  
2217 Western Australia using the geomorphic-hydrologic approach. *Journal of the Royal*  
2218 *Society of Western Australia* 94:449–464.

2219 Shaeri Karimi, S., N. Saintilan, L. Wen, and J. Cox. 2022. Spatio-temporal effects of inundation  
2220 and climate on vegetation greenness dynamics in dryland floodplains. *Ecohydrology*  
2221 15:e2378.

2222 Shanafield, M., S. A. Bourke, M. A. Zimmer, and K. H. Costigan. 2021. An overview of the  
2223 hydrology of non-perennial rivers and streams. *WIREs Water* 8:e1504.

2224 Shi, X., P. E. Thornton, D. M. Ricciuto, P. J. Hanson, J. Mao, S. D. Sebestyen, N. A. Griffiths,  
2225 and G. Bisht. 2015. Representing northern peatland microtopography and hydrology  
2226 within the community land model. *Biogeosciences* 12:6463–6477.

2227 Shumilova, O., D. Zak, T. Datry, D. von Schiller, R. Corti, A. Foulquier, B. Obrador, K. Tockner,  
2228 D. C. Allan, F. Altermatt, M. I. Arce, S. Arnon, D. Banas, A. Banegas-Medina, E. Beller,  
2229 M. L. Blanchette, J. F. Blanco-Libreros, J. Blessing, I. G. Boëchat, K. Boersma, M. T.  
2230 Bogan, N. Bonada, N. R. Bond, K. Brintrup, A. Bruder, R. Burrows, T. Cancellario, S. M.  
2231 Carlson, S. Cauvy-Fraunié, N. Cid, M. Danger, B. de Freitas Terra, A. M. D. Girolamo,  
2232 R. del Campo, F. Dyer, A. Elosegi, E. Faye, C. Febria, R. Figueroa, B. Four, M. O.  
2233 Gessner, P. Gnohossou, R. G. Cerezo, L. Gomez-Gener, M. A. S. Graça, S. Guareschi,  
2234 B. Gücker, J. L. Hwan, S. Kubheka, S. D. Langhans, C. Leigh, C. J. Little, S. Lorenz, J.  
2235 Marshall, A. McIntosh, C. Mendoza-Lera, E. I. Meyer, M. Miliša, M. C. Mlambo, M.  
2236 Moleón, P. Negus, D. Niyogi, A. Papatheodoulou, I. Pardo, P. Paril, V. Pešić, P.  
2237 Rodriguez-Lozano, R. J. Rolls, M. M. Sanchez-Montoya, A. Savić, A. Steward, R.  
2238 Stubbington, A. Taleb, R. V. Vorste, N. Waltham, A. Zoppini, and C. Zarfl. 2019.  
2239 Simulating rewetting events in intermittent rivers and ephemeral streams: A global  
2240 analysis of leached nutrients and organic matter. *Global Change Biology* 25:1591–1611.

2241 Siebert, S., F. T. Portmann, and P. Döll. 2010. Global patterns of cropland use intensity.

2242 Remote Sensing 2:1625–1643.

2243 Siev, S., E. C. Paringit, C. Yoshimura, and S. Hul. 2019. Modelling inundation patterns and  
2244 sediment dynamics in the extensive floodplain along the Tonle Sap River. *River*  
2245 *Research and Applications* 35:1387–1401.

2246 Slater, L., G. Villarini, S. Archfield, D. Faulkner, R. Lamb, A. Khouakhi, and J. Yin. 2021. Global  
2247 Changes in 20-Year, 50-Year, and 100-Year River Floods. *Geophysical Research*  
2248 *Letters* 48:e2020GL091824.

2249 Smith, A. P., B. Bond-Lamberty, B. W. Benscoter, M. M. Tfaily, C. R. Hinkle, C. Liu, and V. L.  
2250 Bailey. 2017. Shifts in pore connectivity from precipitation versus groundwater rewetting  
2251 increases soil carbon loss after drought. *Nature Communications* 8:1335.

2252 Smith, J. A. M., K. J. Rossner, and D. P. Duran. 2021. New opportunities for conservation of a  
2253 rare tiger beetle on developed barrier island beaches. *Journal of Insect Conservation*  
2254 25:733–745.

2255 Smith, K. A., T. Ball, F. Conen, K. E. Dobbie, J. Massheder, and A. Rey. 2018. Exchange of  
2256 greenhouse gases between soil and atmosphere: interactions of soil physical factors and  
2257 biological processes. *European Journal of Soil Science* 69:10–20.

2258 Smyth, A. R., T. D. Loecke, T. E. Franz, and A. J. Burgin. 2019. Using high-frequency soil  
2259 oxygen sensors to predict greenhouse gas emissions from wetlands. *Soil Biology and*  
2260 *Biochemistry* 128:182–192.

2261 Song, X., X. Chen, J. Stegen, G. Hammond, H. Song, H. Dai, E. Graham, and J. M. Zachara.  
2262 2018. Drought Conditions Maximize the Impact of High-Frequency Flow Variations on  
2263 Thermal Regimes and Biogeochemical Function in the Hyporheic Zone. *Water*  
2264 *Resources Research* 54:7361–7382.

2265 Soupir, M. L., S. Mostaghimi, and C. E. Mitchem, Jr. 2009. A comparative study of stream-  
2266 gaging techniques for low-flow measurements in two Virginia tributaries. *JAWRA Journal*  
2267 *of the American Water Resources Association* 45:110–122.

2268 Speir, S. L., J. L. Tank, and U. H. Mahl. 2020. Quantifying denitrification following floodplain  
2269 restoration via the two-stage ditch in an agricultural watershed. *Ecological Engineering*  
2270 155:105945.

2271 Stallins, J. A., and A. J. Parker. 2003. The influence of complex systems interactions on barrier  
2272 island dune vegetation pattern and process. *Annals of the Association of American*  
2273 *Geographers* 93:13–29.

2274 Stanford, J. A., M. S. Lorang, and F. R. Hauer. 2005. The shifting habitat mosaic of river  
2275 ecosystems. *SIL Proceedings, 1922-2010* 29:123–136.

2276 Stanley, E. H., S. M. Powers, N. R. Lottig, I. Buffam, and J. T. Crawford. 2012. Contemporary  
2277 changes in dissolved organic carbon (DOC) in human-dominated rivers: is there a role  
2278 for DOC management? *Freshwater Biology* 57:26–42.

2279 Stewart, B., J. B. Shanley, J. W. Kirchner, D. Norris, T. Adler, C. Bristol, A. A. Harpold, J. N.  
2280 Perdrial, D. M. Rizzo, G. Sterle, K. L. Underwood, H. Wen, and L. Li. 2022. Streams as  
2281 mirrors: Reading subsurface water chemistry from stream chemistry. *Water Resources*  
2282 *Research* 58.

2283 Stewart, R. D., A. S. Bhaskar, A. J. Parolari, D. L. Herrmann, J. Jian, L. A. Schifman, and W. D.  
2284 Shuster. 2019. An analytical approach to ascertain saturation-excess versus infiltration-  
2285 excess overland flow in urban and reference landscapes. *Hydrological Processes*  
2286 33:3349–3363.

2287 Sun, B., M. Jiang, G. Han, L. Zhang, J. Zhou, C. Bian, Y. Du, L. Yan, and J. Xia. 2022a.  
2288 Experimental warming reduces ecosystem resistance and resilience to severe flooding  
2289 in a wetland. *Science Advances* 8:eabl9526.

2290 Sun, Z., L. Sandoval, R. Crystal-Ornelas, S. M. Mousavi, J. Wang, C. Lin, N. Cristea, D. Tong,  
2291 W. H. Carande, X. Ma, Y. Rao, J. A. Bednar, A. Tan, J. Wang, S. Purushotham, T. E.  
2292 Gill, J. Chastang, D. Howard, B. Holt, C. Gangodagamage, P. Zhao, P. Rivas, Z.  
2293 Chester, J. Orduz, and A. John. 2022b. A review of Earth artificial intelligence.

2294 Computers & Geosciences 159:105034.

2295 Svensson, J. R., M. Lindegarth, P. R. Jonsson, and H. Pavia. 2012. Disturbance–diversity  
2296 models: what do they really predict and how are they tested? Proceedings of the Royal  
2297 Society B: Biological Sciences 279:2163–2170.

2298 Sweet, W., J. Park, J. Marra, C. Zervas, and S. Gill. 2014. Sea level rise and nuisance flood  
2299 frequency changes around the United States.

2300 Swenson, L. J., S. Zipper, D. M. Peterson, C. N. Jones, A. J. Burgin, E. Seybold, M. F. Kirk, and  
2301 C. Hatley. 2024. Changes in Water Age During Dry-Down of a Non-Perennial Stream.  
2302 Water Resources Research 60:e2023WR034623.

2303 Tagestad, J., N. D. Ward, D. Butman, and J. Stegen. 2021. Small streams dominate US tidal  
2304 reaches and will be disproportionately impacted by sea-level rise. Science of The Total  
2305 Environment 753:141944.

2306 Tai, X., W. R. L. Anderegg, P. D. Blanken, S. P. Burns, L. Christensen, and P. D. Brooks. 2020.  
2307 Hillslope hydrology influences the spatial and temporal patterns of remotely sensed  
2308 ecosystem productivity. Water Resources Research 56:e2020WR027630.

2309 Thomas, M. A., B. B. Mirus, and J. B. Smith. 2020. Hillslopes in humid-tropical climates aren't  
2310 always wet: Implications for hydrologic response and landslide initiation in Puerto Rico.  
2311 Hydrological Processes 34:4307–4318.

2312 Tiner, R. W. 2013. Tidal wetlands primer: An introduction to their ecology, natural history, status,  
2313 and conservation. University of Massachusetts Press, Amherst.

2314 Tiner, R. W. 2017. Wetland indicators: A guide to wetland identification, delineation,  
2315 classification, and mapping. Second edition. Taylor & Francis, Boca Raton.

2316 Trochim, E. D., A. Prakash, D. L. Kane, and V. E. Romanovsky. 2016. Remote sensing of water  
2317 tracks. Earth and Space Science 3:106–122.

2318 Tromp-van Meerveld, H. J., and J. J. McDonnell. 2006. Threshold relations in subsurface  
2319 stormflow: 2. The fill and spill hypothesis: Threshold flow relations. Water Resources

2320 Research 42.

2321 Tsoi, W. (Iris), I. Gowns, M. Southwell, S. Mika, S. Lewis, D. Ryder, and P. Frazier. 2022.

2322 Effects of inundation on water quality and invertebrates in semiarid floodplain wetlands.

2323 Inland Waters 12:397–406.

2324 Tweedley, J. 2016. The contrasting ecology of temperate macrotidal and microtidal estuaries.

2325 U.S. Geological Survey. 2017. Cottonwood Lake Study Area - Aerial Imagery: U.S. Geological

2326 Survey data release, <https://doi.org/10.5066/F7DZ06GR>.

2327 USACE. 2024. US Army Corps of Engineers, Definitions.

2328 <https://www.nap.usace.army.mil/Missions/Regulatory/Definitions/>.

2329 Valett, H. M., M. A. Baker, J. A. Morrice, C. S. Crawford, M. C. Molles Jr., C. N. Dahm, D. L.

2330 Moyer, J. R. Thibault, and L. M. Ellis. 2005. Biogeochemical and metabolic responses to

2331 the flood pulse in a semiarid floodplain. *Ecology* 86:220–234.

2332 Van Appledorn, M., N. R. De Jager, and J. J. Rohweder. 2021. Quantifying and mapping

2333 inundation regimes within a large river-floodplain ecosystem for ecological and

2334 management applications. *River Research and Applications* 37:241–255.

2335 Van Meerveld, H. J. I., E. Sauquet, F. Gallart, C. Sefton, J. Seibert, and K. Bishop. 2020. Aqua

2336 temporaria incognita. *Hydrological Processes* 34:5704–5711.

2337 VanZomeren, C. M., J. F. Berkowitz, C. D. Piercy, and J. R. White. 2018. Restoring a degraded

2338 marsh using thin layer sediment placement: Short term effects on soil physical and

2339 biogeochemical properties. *Ecological Engineering* 120:61–67.

2340 Venterink, H. O., N. M. Pieterse, J. D. M. Belgers, M. J. Wassen, and P. C. De Ruiter. 2002. N,

2341 P, and K budgets along nutrient availability and productivity gradients in wetlands.

2342 *Ecological Applications* 12:1010–1026.

2343 Vitousek, S., P. L. Barnard, C. H. Fletcher, N. Frazer, L. Erikson, and C. D. Storlazzi. 2017.

2344 Doubling of coastal flooding frequency within decades due to sea-level rise. *Scientific*

2345 *Reports* 7:1399.

2346 Vorste, R. V., R. Corti, A. Sagouis, and T. Datry. 2016. Invertebrate communities in gravel-bed,  
2347 braided rivers are highly resilient to flow intermittence. *Freshwater Science* 35:164–177.

2348 Vousdoukas, M. I., E. Voukouvalas, L. Mentaschi, F. Dottori, A. Giardino, D. Bouziotas, A.  
2349 Bianchi, P. Salamon, and L. Feyen. 2016. Developments in large-scale coastal flood  
2350 hazard mapping. *Natural Hazards and Earth System Sciences* 16:1841–1853.

2351 Waltham, N. J., and R. M. Connolly. 2011. Global extent and distribution of artificial, residential  
2352 waterways in estuaries. *Estuarine, Coastal and Shelf Science* 94:192–197.

2353 Wang, X., W. Wang, and C. Tong. 2016. A review on impact of typhoons and hurricanes on  
2354 coastal wetland ecosystems. *Acta Ecologica Sinica* 36:23–29.

2355 Wantzen, K., C. Alves, S. Badiane, R. Bala, M. Blettler, M. Callisto, Y. Cao, M. Kolb, G. Kondolf,  
2356 M. Leite, D. Macedo, O. Mahdi, M. Neves, M. Peralta, V. Rotgé, G. Rueda-Delgado, A.  
2357 Scharager, A. Serra-Llobet, J.-L. Yengué, and A. Zingraff-Hamed. 2019. Urban stream  
2358 and wetland restoration in the Global South—A DPSIR analysis. *Sustainability* 11:4975.

2359 Ward, J. V., K. Tockner, D. B. Arscott, and C. Claret. 2002. Riverine landscape diversity.  
2360 *Freshwater Biology* 47:517–539.

2361 Ward, J. V., K. Tockner, and F. Schiemer. 1999. Biodiversity of floodplain river ecosystems:  
2362 ecotones and connectivity<sup>1</sup>. *Regulated Rivers: Research & Management* 15:125–139.

2363 Ward, N. D., J. P. Megonigal, B. Bond-Lamberty, V. L. Bailey, D. Butman, E. A. Canuel, H.  
2364 Diefenderfer, N. K. Ganju, M. A. Goñi, E. B. Graham, C. S. Hopkinson, T. Khangaonkar,  
2365 J. A. Langley, N. G. McDowell, A. N. Myers-Pigg, R. B. Neumann, C. L. Osburn, R. M.  
2366 Price, J. Rowland, A. Sengupta, M. Simard, P. E. Thornton, M. Tzortziou, R. Vargas, P.  
2367 B. Weisenhorn, and L. Windham-Myers. 2020. Representing the function and sensitivity  
2368 of coastal interfaces in Earth system models. *Nature Communications* 11:2458.

2369 Watts, J. D., J. S. Kimball, A. Bartsch, and K. C. McDonald. 2014. Surface water inundation in  
2370 the boreal-Arctic: potential impacts on regional methane emissions. *Environmental*  
2371 *Research Letters* 9:075001.

2372 Wen, H., J. Perdrial, B. W. Abbott, S. Bernal, R. Dupas, S. E. Godsey, A. Harpold, D. Rizzo, K.  
2373 Underwood, T. Adler, G. Sterle, and L. Li. 2020. Temperature controls production but  
2374 hydrology regulates export of dissolved organic carbon at the catchment scale.  
2375 Hydrology and Earth System Sciences 24:945–966.

2376 Weyman, D. R. 1973. Measurements of the downslope flow of water in a soil. Journal of  
2377 Hydrology 20:267–288.

2378 Whitworth, K. L., J. L. Kerr, L. M. Mosley, J. Conallin, L. Hardwick, and D. S. Baldwin. 2013.  
2379 Options for managing hypoxic blackwater in river systems: case studies and framework.  
2380 Environmental Management 52:837–850.

2381 Wierzbicki, G., P. Ostrowski, and T. Falkowski. 2020. Applying floodplain geomorphology to  
2382 flood management (The Lower Vistula River upstream from Plock, Poland). Open  
2383 Geosciences 12:1003–1016.

2384 Williams, D. D. 2006. The biology of temporary waters. Oxford University Press, Oxford ; New  
2385 York.

2386 Wittenberg, H. 1999. Baseflow recession and recharge as nonlinear storage processes.  
2387 Hydrological Processes 13:715–726.

2388 Wohl, E. 2021. An integrative conceptualization of floodplain storage. Reviews of Geophysics  
2389 59.

2390 Wollheim, W. M., T. K. Harms, A. L. Robison, L. E. Koenig, A. M. Helton, C. Song, W. B.  
2391 Bowden, and J. C. Finlay. 2022. Superlinear scaling of riverine biogeochemical function  
2392 with watershed size. Nature Communications 13:1230.

2393 Wu, B., F. Tian, M. Nabil, J. Bofana, Y. Lu, A. Elnashar, A. N. Beyene, M. Zhang, H. Zeng, and  
2394 W. Zhu. 2023. Mapping global maximum irrigation extent at 30m resolution using the  
2395 irrigation performances under drought stress. Global Environmental Change 79:102652.

2396 Wu, R., X. Chen, G. Hammond, G. Bisht, X. Song, M. Huang, G.-Y. Niu, and T. Ferre. 2021.  
2397 Coupling surface flow with high-performance subsurface reactive flow and transport



2398 code PFLOTRAN. *Environmental Modelling & Software* 137:104959.

2399 Xiao, D., Y. Shi, S. L. Brantley, B. Forsythe, R. DiBiase, K. Davis, and L. Li. 2019. Streamflow  
2400 generation from catchments of contrasting lithologies: The role of soil properties,  
2401 topography, and catchment size. *Water Resources Research* 55:9234–9257.

2402 Xie, D., C. Schwarz, M. Z. M. Brückner, M. G. Kleinhans, D. H. Urrego, Z. Zhou, and B. Van  
2403 Maanen. 2020. Mangrove diversity loss under sea-level rise triggered by bio-  
2404 morphodynamic feedbacks and anthropogenic pressures. *Environmental Research*  
2405 *Letters* 15:114033.

2406 Xin, P., A. Wilson, C. Shen, Z. Ge, K. B. Moffett, I. R. Santos, X. Chen, X. Xu, Y. Y. Y. Yau, W.  
2407 Moore, L. Li, and D. A. Barry. 2022. Surface water and groundwater interactions in salt  
2408 marshes and their impact on plant ecology and coastal biogeochemistry. *Reviews of*  
2409 *Geophysics* 60:e2021RG000740.

2410 Zedler, P. H. 2003. Vernal pools and the concept of “isolated wetlands.” *Wetlands* 23:597–607.

2411 Zhang, Y. S., W. R. Cioffi, R. Cope, P. Daleo, E. Heywood, C. Hoyt, C. S. Smith, and B. R.  
2412 Silliman. 2018. A Global Synthesis Reveals Gaps in Coastal Habitat Restoration  
2413 Research. *Sustainability* 10:1040.

2414 Zhang, Z., E. Fluet-Chouinard, K. Jensen, K. McDonald, G. Hugelius, T. Gumbricht, M. Carroll,  
2415 C. Prigent, A. Bartsch, and B. Poulter. 2021. Development of the global dataset of  
2416 Wetland Area and Dynamics for Methane Modeling (WAD2M). *Earth System Science*  
2417 *Data* 13:2001–2023.

2418 Zhang, Z., N. E. Zimmermann, A. Stenke, X. Li, E. L. Hodson, G. Zhu, C. Huang, and B.  
2419 Poulter. 2017. Emerging role of wetland methane emissions in driving 21st century  
2420 climate change. *Proceedings of the National Academy of Sciences* 114:9647–9652.

2421 Zhao, Y., X. Wang, S. Jiang, J. Xiao, J. Li, X. Zhou, H. Liu, Z. Hao, and K. Wang. 2022. Soil  
2422 development mediates precipitation control on plant productivity and diversity in alpine  
2423 grasslands. *Geoderma* 412:115721.

2424 Zhi, W., and L. Li. 2020. The shallow and deep hypothesis: Subsurface vertical chemical  
2425 contrasts shape nitrate export patterns from different land uses. *Environmental Science*  
2426 & *Technology* 54:11915–11928.

2427 Zimmer, M. A., A. J. Burgin, K. Kaiser, and J. Hosen. 2022. The unknown biogeochemical  
2428 impacts of drying rivers and streams. *Nature Communications* 13:7213.

2429 Zimmer, M. A., K. E. Kaiser, J. R. Blaszczak, S. C. Zipper, J. C. Hammond, K. M. Fritz, K. H.  
2430 Costigan, J. Hosen, S. E. Godsey, G. H. Allen, S. Kampf, R. M. Burrows, C. A.  
2431 Krabbenhoft, W. Dodds, R. Hale, J. D. Olden, M. Shanafield, A. G. DelVecchia, A. S.  
2432 Ward, M. C. Mims, T. Datry, M. T. Bogan, K. S. Boersma, M. H. Busch, C. N. Jones, A.  
2433 J. Burgin, and D. C. Allen. 2020. Zero or not? Causes and consequences of zero-flow  
2434 stream gage readings. *WIREs Water* 7.

2435 Zimmer, M. A., and B. L. McGlynn. 2017. Ephemeral and intermittent runoff generation  
2436 processes in a low relief, highly weathered catchment. *Water Resources Research*  
2437 53:7055–7077.

2438 Zipper, S. C., J. C. Hammond, M. Shanafield, M. Zimmer, T. Datry, C. N. Jones, K. E. Kaiser, S.  
2439 E. Godsey, R. M. Burrows, J. R. Blaszczak, M. H. Busch, A. N. Price, K. S. Boersma, A.  
2440 S. Ward, K. Costigan, G. H. Allen, C. A. Krabbenhoft, W. K. Dodds, M. C. Mims, J. D.  
2441 Olden, S. K. Kampf, A. J. Burgin, and D. C. Allen. 2021. Pervasive changes in stream  
2442 intermittency across the United States. *Environmental Research Letters* 16:084033.