



30 to +1.19 °C and a Root Mean Square Error in the range of 0.8 to 5.9 °C among the stations.
31 The model showed enhanced performance for the 10-m wind speed and downward shortwave
32 radiation flux, reflecting advancements over previous studies. Therefore, the WRF-Chem
33 model effectively simulates key meteorological parameters and pollutants over the UAE,
34 demonstrating significant regional-scale prediction skills. Areas for further model refinement
35 are also identified and discussed. Integrating model predictions with satellite and ground-based
36 data is emphasized for advancing air quality monitoring and enhancing predictive accuracy of
37 atmospheric pollutants in this region.

38 **Keywords:** Air quality modelling, gaseous pollutants, TROPOMI satellite retrievals, WRF-
39 Chem, UAE.

40 **Key points:**

- 41 ● First high-resolution WRF-Chem air quality modelling study over the United Arab
42 Emirates (UAE)
- 43 ● WRF-Chem's ability to simulate meteorological parameters and pollutant levels over
44 the UAE is assessed during summer and winter in 2018.
- 45 ● The model showed a strong correlation with TROPOMI satellite data, achieving
46 correlation coefficients of 0.71-0.95 in summer and 0.86-0.94 in winter for different
47 gaseous pollutants.
- 48 ● Lower model skill in simulating total CO columns, in contrast to the more accurate
49 modelling of tropospheric NO₂ and total O₃ columns as compared to TROPOMI data.
- 50 ● Meteorological analysis revealed a tendency to underestimate surface temperature by
51 0.5 °C in summer and overestimate it by 1.3 °C in winter.
- 52 ● Surface wind speed is overestimated by 0.1-0.9 m/s in both seasons across various
53 regimes.

54



55 **1. Introduction**

56 The United Arab Emirates (UAE), a federation of seven emirates, has undergone rapid
57 urbanization and industrialization over the last five decades, which has had a profound impact
58 on its air quality (Ramadan, 2015). The major factors affecting air quality in the UAE include
59 emissions from industrial activities, vehicular traffic, construction projects, and occasionally,
60 natural phenomena such as dust storms, which are quite prevalent in the region due to its desert
61 climate (Environment Agency – Abu Dhabi, 2018; Francis et al., 2020; 2022b; Karagulian et
62 al., 2019). The rapid economic growth of the UAE, especially in cities like Dubai and Abu
63 Dhabi, has led to a surge in energy demand and desalinated water, largely met through the
64 burning of fossil fuels (Shahbaz et al., 2014). This has resulted in increased emissions of
65 pollutants like oxides of nitrogen (NO_x), sulfur dioxide (SO₂), particulate matter (PM), and
66 volatile organic compounds (VOCs). Moreover, the heavy traffic in urban areas contributes to
67 the elevated levels of ground-level ozone and particulate pollution (Abuelgasim & Farahat,
68 2020; Li et al., 2010). Understanding the dynamics of air quality in the UAE involves
69 considering both the environmental challenges posed by rapid development and the steps being
70 taken to mitigate these impacts. The pursuit of balancing economic growth with environmental
71 sustainability is central to this discourse. This area of study is not only vital for ensuring the
72 health and well-being of the population but also plays a crucial role in the UAE's vision for a
73 sustainable future.

74

75 The swift urban expansion in the UAE could intensify air pollution sources. With surface
76 observations sparse in this region, satellite remote sensing becomes a crucial method for air
77 quality monitoring (Chudnovsky et al., 2014; Fonseca et al., 2023; Francis et al., 2023). What
78 is more, satellite measurements themselves fall short in clarifying the different atmospheric
79 processes responsible for peak pollution levels. Consequently, integrating chemistry transport
80 models with satellite-derived and ground-based observations can significantly improve our
81 understanding of pollutant emissions, distribution, transport, and transformation in the targeted
82 regions (Eltahan et al., 2018; Li et al., 2018; Yarragunta et al., 2020; Yin et al., 2021). Air
83 quality (AQ) modelling is dedicated to unravelling the complicated aspects of atmospheric
84 chemistry and transport across both global and regional levels, as explored in numerous studies
85 conducted around the world (Emmons et al., 2010; Kumar et al., 2011, 2018; Tie et al., 2001;
86 Yarragunta et al., 2019, 2020, 2021). Despite facing limitations due to the often low spatial and
87 temporal resolution of observational data, AQ models effectively generate detailed air quality



88 information for remote regions. They predict the formation and removal of air pollutants and
89 facilitate a thorough examination of the transport and photo-chemical transformation of trace
90 gases following their emission into the atmosphere (Archer-Nicholls et al., 2015; Georgiou et
91 al., 2018; Nhu et al., 2021; Sicard et al., 2021). They are also employed globally for operational
92 air quality forecasting (Jena et al., 2021; Koo et al., 2012; Kumar et al., 2012, 2021; Srinivas
93 et al., 2016; Zhang et al., 2012). Air quality models are categorized into two types: 'fully
94 coupled' models, which integrate interactions between chemistry and meteorology, and 'offline'
95 models, where chemistry and meteorology simulations are conducted independently (Gao &
96 Zhou, 2024). Some of state of the art AQ models include the Weather Research and
97 Forecasting (WRF) model coupled with chemistry (WRF-Chem; Grell et al., 2005; Skamarock
98 et al., 2008), WRF-Chem-MADRID (Model of Aerosol Dynamics, Reaction, Ionization and
99 Dissolution, Zhang et al., 2010), CESM2 (Community Earth System Model version 2, Emmons
100 et al., 2020), CHIMERE (Menut et al., 2021), LOTOS-EUROS(v2.0) (Long Term Ozone
101 Simulation European Operational Smog, Manders et al., 2017) and COSMO/MESSy
102 (Consortium for Small Scale Modelling/ Modular Earth Submodel System, Kerkweg & Jöckel,
103 2012). However, before using these AQ models for future applications, it is crucial to conduct
104 thorough evaluations to assess the quality of their simulations. The AQ model chosen for the
105 current study is the WRF-Chem with its foundational meteorological component, WRF.

106

107 The majority of studies conducted in the UAE and similar arid regions have primarily
108 focused on evaluation of meteorological parameter including temperature, humidity, wind, and
109 solar radiation (Parajuli et al., 2019; Nelli et al., 2020; Fonseca et al., 2020, 2021) with a few
110 others investigating the particulate matter dynamics, especially mineral dust. For instance,
111 Ukhov et al., (2021) noted inaccuracies in the WRF-Chem model related to the GOCART
112 aerosol module, affecting $PM_{2.5}$ and PM_{10} diagnostics. Karagulian et al., (2019) highlighted the
113 effectiveness of integrating WRF-chem model simulations with satellite and ground
114 observations to understand and predict the impact of severe dust storms on air quality.
115 Karumuri et al., (2022) reported significant air quality changes due to COVID-19 lockdown
116 measures, with reduced trace gas concentrations but increased particulate matter from dust
117 activities, the latter stressed by Francis et al. (2022a) who attributed it to changes in the
118 atmospheric circulation. Moreover, Parajuli et al., (2022; 2023) utilized high-resolution WRF-
119 Chem simulations and advanced aerosol schemes to analyse the dust and rainfall dynamics,
120 providing insights into the direct and indirect effects of dust on rainfall which aids in better
121 regional water resource planning through accurate rainfall predictions. However, no



122 assessment to date for the gaseous pollutants model performance over the region despite the
123 complex dynamics between anthropogenic and natural factors in air quality management and
124 the necessity of tailored model configurations for accurate environmental assessments in arid
125 regions.

126

127 This study represents the first evaluation of the WRF-Chem model in the region, specifically
128 examining concentrations of gaseous pollutants along with crucial meteorological parameters
129 relevant to air quality studies. The primary objective of this study is twofold:

130

- 131 ● Evaluate the WRF-Chem's ability to replicate meteorological conditions. This involves
132 comparing the model's simulation of temperature, wind speed, downward short-wave
133 radiation and boundary layer height against ground-based observations and data from
134 the ERA5 (Hersbach et al., 2020) reanalysis.
- 135 ● Assess the model's performance in simulating concentrations of gaseous pollutants,
136 specifically NO₂, O₃, and CO. The skill of the WRF-Chem in simulating these
137 pollutants is evaluated by comparing its simulations against data from the
138 TROPospheric Monitoring Instrument (TROPOMI; Veekfind et al., 2012) on the
139 Sentinel-5 Precursor satellite.

140

141 The structure of the paper is as follows: Section 2 describes the configuration of the WRF-
142 Chem considered in this work. Section 3 elaborates on the methodology and datasets used in
143 this study. Section 4 provides a comprehensive assessment of the WRF-Chem's simulated data
144 with observational datasets, reanalysis and satellite-derived products. Section 5 concludes by
145 outlining the main findings.

146 **2. WRF-Chem configuration**

147 The central objective of this study is to apply a regional chemistry/dynamical model to
148 simulate the atmospheric conditions and transport of pollutants in the UAE, whose forecasts
149 will be evaluated against in-situ, space-based measurements and a state-of-the-art reanalysis
150 dataset. To this end, the WRF-Chem version 4.3.1 is employed. WRF-Chem is a mesoscale
151 regional chemistry transport model, developed by the National Oceanic and Atmospheric
152 Administration (NOAA) Earth System Research Laboratory (ESRL), and has been contributed
153 to by the global science community. In WRF-Chem, air quality components and meteorological

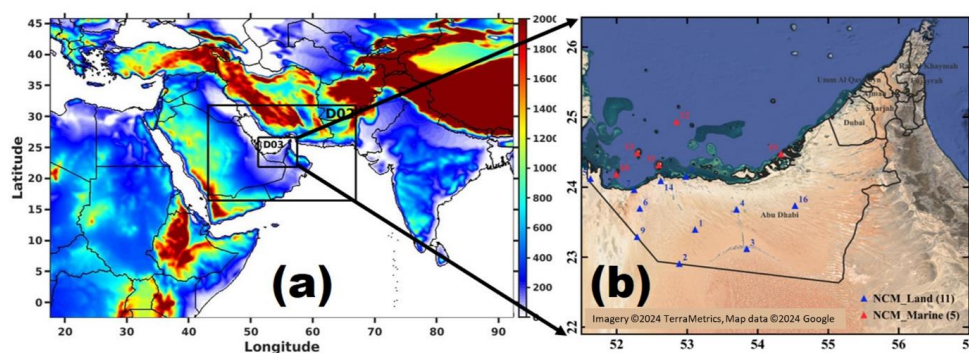


154 components are predicted simultaneously using the same grid coordinates, transport, timestep,
155 and sub-grid scale physics. A detailed description of the model is found in Grell et al., (2005)
156 and Skamarock et al., (2008). The physics schemes employed in the simulations are the Rapid
157 Radiative Transfer Model for Global Circulations Models (RRTMG) for radiation
158 parametrization of both short and long wave radiation (Iacono et al., 2008), the cloud
159 microphysics is represented by the Morrison 2-moment (Morrison et al., 2009), and the Kain-
160 Fritsch scheme is used for convective parameterisation (Kain, J.S, 2004). The Unified Noah
161 model is used to represent the land surface model (Tewari et al., 2004) with an improved
162 representation of soil texture and land use/land cover (LULC) over the UAE (Temimi et al.,
163 2020). The boundary layer dynamics are represented by the Yonsei University (YSU) scheme
164 (Hong, 2010). Other chosen physics schemes are listed in Table 1. Simulated mesoscale
165 meteorology is kept in line with analysed meteorology through spectral nudging to the National
166 Centre for Environmental Prediction (NCEP) Global Forecast System (GFS) analyses used to
167 drive the model, in an attempt to limit errors in the mesoscale transport. During the simulations,
168 horizontal and vertical wind, potential temperature and water vapour mixing ratio are nudged
169 to GFS analyses in all model layers above the planetary boundary layer on a time-scale of 6
170 hours. Meteorological conditions were initialised by NCEP GFS 6-hourly analyses at 0.25°
171 resolution.

172 This study utilised the Model for Ozone and Related Chemical Tracers, version 4
173 (MOZART-4) chemical mechanism for calculating gas-phase chemistry which includes 81
174 chemical species with 159 gas-phase reactions and 38 photolysis processes (Emmons et al.,
175 2010). Aerosol chemistry is represented by the Goddard Chemistry Aerosol Radiation and
176 Transport (GOCART; Chin et al., 2002), along with the Tropospheric, Ultraviolet and Visible
177 (TUV) full photolysis scheme (Madronich, 1987; Tie, 2003), which deploys climatological O₃
178 and O₂ columns. Dry deposition was calculated using Wesely (1989). Anthropogenic emissions
179 were taken from the Emission Database for Global Atmospheric Research version 5
180 (EDGARv5) at 0.1 × 0.1° horizontal resolution (Crippa et al., 2020). Emissions include SO₂,
181 NO_x, CO, NMVOC, NH₃, black carbon (BC) and organic carbon (OC). Biogenic emissions
182 were calculated online by the Model of Emissions of Gases and Aerosol from Nature
183 (MEGAN; Guenther et al., 2012). Model simulation uses CAM-chem model results as
184 chemical boundary conditions (BCs) for the outer domain D01 and initial conditions (ICs) for
185 all domains (Emmons et al., 2020). In this present work, we run the WRF-Chem model using
186 the aforementioned physical and chemical processes on the three nested domains with



187 horizontal resolutions 27-, 9- and 3-km corresponding to 283×205, 271×193 and 256×178 grid
 188 points and 45 vertical layers. The outermost domain covers the vast majority of the Middle
 189 East and surrounding region while the innermost domain covers the entire UAE (Fig. 1(a)).
 190 The analysis in this research article exclusively utilizes results from the inner domain (D03).
 191 The spatial distribution of ground-based observations from NCM are depicted in Fig. 1(b).



192

193 **Figure 1: Model Configuration:** (a) The WRF domain configuration consists of three telescoping
 194 nests, with the outermost boundaries denoting the parent grid (D01). D02 and D03 are the nested
 195 domains. Right panel (b) is a zoom of the innermost domain (D03) showing the spatial distribution of
 196 the 16 meteorological stations (land stations are denoted by blue triangles, and marine stations are
 197 represented by red triangles). The shading in (a) represents the orography (m). Further details about the
 198 stations are given in Tables 2.

199

200 **Table 1: WRF-chem model setup**

201

Model set-up	Option
Model version	4.3.1
Domain	3 domains
Horizontal resolution	D01:27km, D02:9km and D03:3km
Simulation period	Monthly runs from June 2018 and December 2018
Model spin-up period	2 days in each month
Vertical resolution	45 eta levels up to 50 hPa.
Domain size	D01: 283×205 grids, D02: 271×193 grids and D03: 256×178 grids
Meteorological boundary	NCEP FNL reanalysis (0.25°, 6-hourly)
Chemical boundary	CAM-Chem (Emmons, Fasullo, et al., 2020)
Physical Process	Parameterization Scheme
Microphysics	Morrison double moment (Morrison et al., 2009)
Cumulus parameterization	Kain-Fritsch (Kain, J.S, 2004)
Shortwave radiation	Rapid Radiative Transfer Model for GCMs (RRTMG) (Iacono et al. 2008)



Longwave radiation	Rapid Radiative Transfer Model for GCMs (RRTMG) (Iacono et al. 2008)
Land surface	Unified Noah land surface model (Tewari et al., 2004)
Planetary boundary layer	Yonsei University scheme (Hong, 2010)
Chemistry option	Scheme used
Gas phase chemistry	MOZART-4 (Emmons et al., 2010).
Aerosol chemistry	GOCART (Chin et al., 2002)
Photolysis	Madronich F-TUV (Madronich, 1987; Tie, 2003)
Biogenic emissions	MEGAN (Guenther et al. 2012)
Dry deposition	Wesely (Wesely 1989)

202

203 3. Data Sets and methodology

204 3.1 Meteorology observations

205 In this study, meteorological data from 16 automatic weather stations (AWS) operated by
 206 the National Center of Meteorology (NCM), UAE were utilized to assess the WRF-Chem
 207 simulations for air temperature at 2 meters above ground (T2m), wind speed at 10 meters
 208 (WS10m), and downward shortwave radiation flux at the surface (SR) during June and
 209 December of 2018. The spatial distribution of the stations across the UAE is illustrated in Fig.
 210 1(b) (refer to Table 2 for details). These locations were categorically divided into two regions—
 211 land stations (station with ID number: 1-9,14 and 16) and marine stations (station with ID
 212 number: 10-13 and 15)—following the criteria outlined in Branch et al., (2021). Subsequent
 213 analyses are based on these two primary categories, with the land region comprising 11 stations
 214 (marked with green triangles) and the marine region comprising 5 stations (marked with yellow
 215 triangles) in Fig. 1(b). Additional information on the specifics, quality control measures, and
 216 other research studies based on NCM data can be found in the referenced literature (Branch et
 217 al., 2021; Fonseca et al., 2020, 2021, 2022; Temimi et al., 2020a).

218

219 **Table 2 List of Automatic Weather Stations (AWS) utilized for evaluating the WRF-Chem model.**

ID	Name	Lat.	Lon.	Altitude (m)	Region
1	Owtaid	23.40	53.11	160	Land
2	Mukhariz	22.91	52.89	130	Land
3	Mezaria	23.12	53.84	110	Land
4	Madinat Zayed	23.68	53.70	110	Land
5	Al Ghweifat	24.12	51.63	47	Land



6	Bu Hamrah	23.51	54.53	136	Land
7	Barakah	23.96	52.25	5	Land
8	Al Qlaa	24.16	52.98	150	Land
9	Al Jazeera	23.29	52.29	70	Land
10	Yasat	24.19	52.00	115	Marine
11	Sri Bani Yas	24.32	52.60	101	Marine
12	Qarnen	24.94	52.85	26	Marine
13	Dalma	24.49	52.29	10	Marine
14	Al Ruwais	24.09	52.62	33	Land
15	Abu Dhabi	24.48	54.33	3	Marine
16	Al Tawiyen	25.56	56.07	186	Land

220

221 3.2 ERA-5 Reanalysis data

222 The fifth-generation European Centre for Medium-Range Weather Forecasts (ECMWF)
223 reanalysis, known as ERA-5 (Hersbach et al., 2020), represents a significant advancement over
224 its predecessor, the ERA-Interim reanalysis, introduced by Dee et al., (2011). ERA-5
225 incorporates a sophisticated four-dimensional variational (4D-Var) data assimilation method,
226 utilizing the 41r2 cycle of the Integrated Forecast System (IFS). This system is enhanced by
227 the integration of both a soil model and an ocean wave model, offering a comprehensive
228 approach to climate data analysis. For the purposes of this research, we accessed ERA-5 data
229 through the Copernicus Climate Change Service Climate Data Store (CDS). The dataset
230 provides atmospheric observations across 137 hybrid vertical levels, with data available on the
231 CDS interpolated onto 37 distinct pressure levels. These levels span from 1000 hPa, close to
232 the Earth's surface, up to 1 hPa, reaching altitudes of approximately 80 km. Further details on
233 the ERA-5 dataset are available in Dee et al., (2011) and Hersbach et al., (2020). Our study
234 specifically utilized hourly data for a selection of meteorological parameters: air temperature
235 at 2 meters above the ground (T2m), wind speed at 10 meters (WS10m), downward shortwave
236 radiation flux at the surface (SR), and planetary boundary layer height (PBL), for the months
237 of June and December 2018.

238 3.3 Satellite-borne observations

239 Launched by the European Space Agency (ESA) on October 13, 2017, the TROPOspheric
240 Monitoring Instrument (TROPOMI) is aboard the Sentinel-5 Precursor (S5P) satellite,
241 operating in a near-polar sun-synchronous orbit. Positioned at an altitude of 817 km, the S5P
242 satellite crosses the equator at a local solar time of 13:30, boasting a wide swath of



243 approximately 2600 km, and providing daily global coverage. TROPOMI features four distinct
244 spectrometers that measure the ultraviolet (UV) and UV-visible (UV-VIS) range (270 to 500
245 nm), near-infrared (NIR) range (675 to 775 nm), and short-wave infrared (SWIR) range (2305
246 to 2385 nm) spectral bands (Veefkind et al., 2012). Notably, the last two spectral bands, NIR
247 and SWIR, are newly introduced in TROPOMI compared to its predecessor OMI (Ozone
248 Monitoring Instrument). TROPOMI's data products encompass daily observations of trace
249 gases, including CO, O₃, NO₂, CH₄, HCHO, aerosols, and cloud properties. The present study
250 utilized daily NO₂, CO, and ozone column density level 2 products from TROPOMI,
251 downloaded from the GES DISC website (<https://disc.gsfc.nasa.gov/>) for the period of June 1-
252 30 and December 1-31, 2018. The specific data sets employed for the present study includes
253 S5P_OFFL_L2__O₃ for O₃, S5P_OFFL_L2__CO for CO, and S5P_OFFL_L2__NO₂ for NO₂,
254 covering the study region bounded by longitudes [51°,58°] and latitudes [21°, 27°]. Further
255 details regarding each product, retrieval algorithm, and validation results are summarized in
256 the subsequent section.

257

258 TROPOMI retrieval of NO₂ columns are derived using UV-VIS spectrometer backscattered
259 solar radiation measurements in the wavelength range of 405-465 nm and provides total and
260 tropospheric NO₂ vertical column density with a near-nadir resolution of 7x3.5 km. The total
261 NO₂ slant column density (SCD) is retrieved from the measured solar irradiance spectra using
262 the Differential Optical Absorption Spectroscopy (DOAS) method. Tropospheric and
263 stratospheric slant column densities are separated from SCD by a data assimilation system
264 based on the chemistry transport model V5 (TM5-MP). Afterwards, they are converted to
265 vertical column densities (VCDs) with the help of look-up table of altitude-dependent air-mass
266 factors (AMFs) and information on the vertical distribution of NO₂ from TM5-MP a priori
267 profile with a horizontal resolution of 1° x 1° and a time step of 30 min (Boersma et al., 2018;
268 Van Geffen et al., 2022). The TROPOMI NO₂ product has been extensively evaluated using
269 ground-based and aircraft observations and found to have a high correlation and low bias of
270 less than 30% with respect to in-situ measurements (Griffin et al., 2019; Ialongo et al., 2020).
271 We used the reprocessed (RPRO) TROPOMI NO₂ data files with processor version of 1.2.2,
272 for the study period. Additionally, two more NO₂ products are available such as offline (OFFL)
273 and near-real time (NRTI). NRTI data files are generated using TM5-MP forecast data rather
274 than analysis data as with REPO and OFFL files (Van Geffen et al., 2022). The differences
275 between the OFFL/REPO and NRTI NO₂ products are generally very small (references therein
276 Ialongo et al., 2020).



277

278 The Shortwave Infrared Carbon Monoxide Retrieval (SICOR) algorithm is used to retrieve
279 CO column densities from TROPOMI in the spectral range of 2305 to 2385 nm (Landgraf et
280 al., 2016). The SICOR algorithm accounts for a profile-scaling approach that scales retrieved
281 CO total column to the a priori reference profile. The a priori reference profiles are taken from
282 the global chemistry transport model simulations of TM5-MP, and they vary based on the
283 location, month and year (Krol et al., 2005). The detailed outline of all settings and other
284 auxiliary data sets used for CO retrievals are outlined in the Landgraf et al., (2016). This study
285 limits the analysis to CO pixels corresponding to clear-sky conditions and mid-level clouds by
286 filtering the data using the quality flag variable (qa_value). The scenes corresponding to
287 $qa_value > 0.5$ are used in this current analysis as suggested in the ATBD (algorithm theoretical
288 baseline document; Landgraf et al., 2016). In this present work, TROPOMI CO measurements
289 for the period from 1-30 June and 1-31 December, 2018 have been analysed. Moreover, we use
290 either the reprocessed (RPRO) or offline (OFFL) data files from most recent processor versions
291 depending on availability for a given day of observations. Wizenberg et al., (2021) compared
292 global TROPOMI retrieved CO total columns with corresponding ACE-FTS (Atmospheric
293 Chemistry Experiment- Fourier transform spectrometer) columns for the period from
294 November 2017 to May 2020 and found a small relative bias of -0.83% with a correlation
295 coefficient of 0.93 between two data sets. Similar results were also found between TROPOMI
296 CO with corresponding CO fields from the ECMWF assimilation system: Borsdorff et al.
297 (2018) reported a small mean difference between the two data sets of 3.2% with a correlation
298 coefficient of 0.97.

299

300 TROPOMI also provides total ozone column (TOC) and ozone profile data at 15 pressure
301 levels. It measures radiances and irradiances in the ultraviolet wavelength of 270-330 nm and
302 provides the ozone profile information. The Optimal Estimation (OE) algorithm is used to
303 retrieve the ozone profile data. Before this stage, various pre-processing steps are applied to
304 the measured spectra before the estimation of the ozone profile. The main process of the
305 algorithm is the OE method, which combines the information from the measured spectra with
306 the a-priori information. The a-priori information is based on climatology as described in the
307 Labow et al., (2015). The description of the various pre-processing steps performed to retrieve
308 ozone profiles is presented in the Algorithm Theoretical Basis Document (Veefkind, et al.,
309 2021). The validation of TROPOMI retrieved ozone profile data against the ground-based



310 measurements reported a median bias of 0.3% for OFFL/REPO products while 0.8 % for NRTI
 311 ozone products (Lambert et al., 2023).

312 **3.4 Satellite data processing**

313 In order to quantitatively compare the WRF-chem simulations with satellite measurements,
 314 the model outputs must be processed using the appropriate method as described in the literature
 315 (Kumar et al., 2012). Direct comparison between satellite retrievals and model outputs is not
 316 recommended as satellite measurements depend on column averaging kernels (AK) and a-
 317 priori profiles. The AK vector, representing the vertical sensitivity of the retrieved column to
 318 the partial column at different vertical levels, should be employed to convolve the model
 319 simulations.

320 The column density from the WRF-Chem model is re-gridded to match the TROPOMI
 321 instrument's grids and is vertically interpolated to the TROPOMI pressure levels before it is
 322 multiplied by the AK. This treatment of the WRF-Chem-simulated profile with the column
 323 averaging kernels allows for a comparison that is independent of the chemical transport model
 324 (CTM) a-priori assumptions and the vertical sensitivity of the retrieval process; therefore, it
 325 can be directly compared with the TROPOMI-derived tropospheric column of NO₂. The
 326 TROPOMI-NO₂ products also provide a column averaging kernel matrix. In the case of
 327 TROPOMI-NO₂, the application of the column AK averaging kernel accounts for the vertical
 328 distribution and sensitivity of the measurements, as classically done by Borsdorff et al., (2014)
 329 as:

330

$$331 \quad X_{ret} = X_{a\ prior} + AK \times (X_{true} - X_{a\ prior}) + e_x \text{-----(1)}$$

332 where, X_{true} is model simulation profile of trace gas; X_{ret} is the retrieved profile or smoothed
 333 model profile; e_x represents the error on the retrieved trace gas profile; $X_{a\ prior}$ is the a-priori
 334 information provided in the TROPOMI data set. For TROPOMI-NO₂ data, the contribution of
 335 the a priori profile and error on the retrieved profile can be eliminated, as explained in Borsdorff
 336 et al., (2014). The eq. (1) simplifies to

$$337 \quad X_{ret} = AK \times (X_{true}) \text{----- (2)}$$

338

339 For validation of ozone and CO total column, we have used the TROPOMI ozone and CO
 340 profile level 2 data product S5P_OFFL_L2__O₃__ and S5P_OFFL_L2__CO__ that provides
 341 the ozone and CO concentrations at 15 and 50 pressure levels, respectively. This data product



342 also includes the a priori information and column averaging kernel for each pressure level. In
343 order to compare our model profile with this dataset, the model output is horizontally and
344 vertically interpolated to TROPOMI grids and vertical levels. The final model profile was
345 calculated by the Eq. (3)

$$346 \quad X_{ret} = X_{a\ prior} + AK \times (X_{true} - X_{a\ prior}) \text{-----}(3)$$

347 Since the highest vertical level in WRF-Chem-simulated trace gas concentration is 50 hPa, the
348 remaining vertical layers of ozone and CO were made equal to the a priori concentration of
349 respective trace gases as described by ATBD (Landgraf et al., 2016).

350 **3.5 Evaluation methodology**

351 Meteorological parameters from the WRF-Chem model were extracted for the grid points
352 closest to the surface observation sites of NCM. Meteorological parameters were categorized
353 and averaged for land and marine regions for the regional analysis. Consequently, further
354 analyses based on these categories are presented in subsequent sections of the article. To enable
355 comparison of atmospheric column data from the TOPOMI satellite retrievals with WRF-
356 Chem outputs, the data must undergo smoothing through an appropriate method described in
357 Section 3.4, as direct comparison between satellite retrievals and simulations is not feasible
358 due to discrepancies highlighted in previous literature. Additionally, owing to the spatial
359 resolution differences between WRF-Chem and ERA5 datasets, it is necessary to remap the
360 model data to the ERA5 grids for accurate comparison. A wide range of statistical parameters
361 is available for evaluating model simulations. In this study, we employed statistical skill scores
362 including the Pearson correlation coefficient (r), the Mean Bias (MB), the Root Mean Square
363 Error (RMSE), and the Mean Absolute Error (MAE), which have been extensively discussed
364 and applied in similar contexts (Fonseca et al., 2021; Ivatt & Evans, 2020; Temimi et al.,
365 2020b).

366 The following equations (eq. 4 to eq. 7) are used to calculate these statistical matrixes in the
367 present study,

368

$$369 \quad r = \frac{\sum_{i=1}^N [(O_i - \bar{O}_i)(M_i - \bar{M}_i)]}{\sqrt{\sum_{i=1}^N (O_i - \bar{O}_i)^2 \sum_{i=1}^N (M_i - \bar{M}_i)^2}} \text{-----}(4)$$

$$370 \quad RMSE = \left(\frac{1}{N} \sum_{i=1}^N (M_i - O_i)^2 \right)^{\frac{1}{2}} \text{-----}(5)$$

$$371 \quad MB = \frac{1}{N} \sum_{i=1}^N (M_i - O_i) \text{-----}(6)$$



372
$$\text{MAE} = \frac{1}{N} \sum_{i=1}^N |M_i - O_i| \text{-----}(7)$$

373 where O_i denotes the i -th observation, M_i represents the corresponding WRF-chem simulated
 374 value, and N is the number of model and observation pairs. \overline{M}_i and \overline{O}_i are the model and
 375 observational means (i.e. average of 1-30, June and 1-31 December), respectively. The
 376 correlation coefficient (r) is an indication of the phase agreement between the modelled and
 377 observed time-series. The RMSE measures the average error in the model, and the MAE
 378 determines the mean error between the model and observations regardless of whether it is an
 379 under or overestimate. The MB is a measure of the systematic error and gives information
 380 whether the model is over or underpredicting the corresponding observed values.

381 **4. Results and Discussion**

382 **4.1 Model performance for meteorological variables**

383 The general ability of the WRF-Chem model to reproduce realistic spatio-temporal
 384 patterns of the most relevant physical and chemical variables is assessed by comparing the
 385 simulated output with the observational data for June and December for the year 2018,
 386 reflecting the contrasting summer and winter conditions over the UAE. Determining the
 387 accuracy of WRF-Chem simulations by validating meteorological conditions in the study area
 388 is crucial before utilizing the model's output for air quality applications. In this regard, we have
 389 conducted a comparison of the model's T2m, WS10m, and SR outputs with measurements from
 390 observational data sets. Additionally, we have compared the boundary layer height from the
 391 model with the ERA5 reanalysis product. These parameters were chosen due to their
 392 significance in influencing most air pollutants (Ritter et al., 2013). Notably, the ERA5
 393 reanalysis data boasts a high spatial resolution of approximately 28 km, making it superior to
 394 other reanalysis datasets in this aspect. Our comparison involved analysing the hourly results
 395 from both ERA5 and ground-based datasets against WRF-Chem for two distinct months in
 396 2018. Detailed results of this comparison are presented below.

397 **4.1.1 Evaluation against surface-based observations**

398 The WRF-Chem model effectively represented the observed variations in T2m, WS10m,
 399 and SR across all 16 meteorological stations during June and December 2018. The WRF-Chem
 400 model generally underestimated T2m values by less than 0.5 °C in June and overestimated
 401 them in December by less than 1.3 °C across the majority of locations. Correlation coefficients



402 for the observed T2m with model simulations were between 0.66 to 0.99 in June, slightly
403 increasing to a range of 0.70 to 0.99 in December. The MB for T2m varied from -0.04 to +1.19
404 °C in June and -2.17 to +0.50 °C in December, with the RMSE spanning from 0.8 to 5.9 °C in
405 June and 0.9 to 4.1 °C in December. Conversely, the outcomes for WS10m and SR
406 demonstrated variability across different stations. The model performance demonstrates
407 significant enhancements over previous research conducted in this region. For instance,
408 Fonseca et al., (2020) observed a warm bias of 1-3 °C in WRF simulations across the UAE for
409 both winter and summer seasons. This observation aligns with similar findings reported by
410 Schwitalla et al., (2020) and Wehbe et al., (2017). The enhanced performance of the model
411 may be attributed to the present model configuration which differs from that used in previous
412 studies.

413 We concentrate on evaluating the model's performance at a regional scale, as delineated by
414 land (encompassing 11 sites) and marine stations (comprising 5 sites), detailed in section 3.1.1.
415 Table 3 presents an extensive evaluation of the statistical verification scores for essential
416 meteorological variables at these categories within the UAE. In the month of June, the model
417 slightly underestimated the T2m values in both land and marine settings, with a
418 underestimation of 0.37 °C and 0.48 °C, respectively, despite an overprediction of SR. This
419 arises because of colder temperatures in particular in the evening and night-time hours, a bias
420 highlighted by other studies such as Temimi et al. (2020b) and Branch et al. (2021). This has
421 been attributed to deficiencies in the model's physics and/or dynamics, in particular in the land
422 surface model and surface properties, a cold bias in the forcing dataset, and an incorrect
423 representation of the concentration of aerosols and greenhouse gases. Despite this, the model
424 achieves notable correlation coefficients (r) of 0.91 for land regions and 0.83 for marine
425 regions. The lower correlation observed in marine regions possibly arises from the more muted
426 diurnal cycle (Fig. 2) and the model's inability to properly represent the complex land-sea mask
427 even at 3 km spatial resolution. Similar results were reported in Abida et al., (2022), where the
428 WRF model demonstrates improved accuracy in inland areas compared to offshore or coastal
429 regions. The RMSE (MAE) values stand at 3.57 °C (2.68 °C) for land and 1.67 °C (1.47 °C)
430 for marine regions, respectively. In December, the T2m predictions by the model show an
431 overestimation, marked by 0.76 °C in land and 1.30 °C in marine regions. The model maintains
432 strong correlations, with $r = 0.92$ for land and $r = 0.90$ for marine regions, underscoring its
433 consistent performance. The RMSE (MAE) values recorded are 2.87 °C (1.66 °C) for land and
434 2.57 °C (1.37 °C) for marine regions, illustrating the model's accuracy in capturing temperature



435 fluctuations over these regions. For WS10m, the model effectively aligns with observed values,
 436 showing good agreement in both land and marine settings. In June, it slightly overestimated
 437 the wind speed in the marine region by 0.51 m/s, a trend that is also reflected in the RMSE
 438 metrics, which are marginally higher for marine areas compared to land (0.08 m/s). In
 439 December, it notably overestimated wind speeds in marine regions by 0.92 m/s, while the
 440 overestimation was slightly less in land areas, at 0.38 m/s. Despite this, the correlations remain
 441 robust in both seasons, highlighting the model reliability in capturing wind speed variations
 442 across different environments. The model representation of SR demonstrates a similar pattern
 443 of accuracy and overestimation. In June, the model tends to overestimate SR across both
 444 regions, which has been reported in Fonseca et al. (2020) and Temimi et al. (2020b), yet it
 445 achieves a more accurate depiction in December. A possible explanation is a reduced aerosol
 446 loading in the model, with the summer featuring higher atmospheric aerosol amounts than the
 447 winter season (Nelli et al., 2021), with WRF also exhibiting a tendency to underpredict the
 448 observed cloud cover in the region. Although the correlations for SR are slightly lower,
 449 especially in the marine regions, they still indicate a reasonable level of model performance.
 450 Overall, the model tends to overestimate WS10m and SR across both seasons, while it
 451 underestimates the T2m in winter and overestimates it in summer. Such variable performance
 452 of the model has been noted in findings from prior research (for example, Schwitalla et al.,
 453 2020; Wehba et al., 2017; Fonseca et al., 2020; Abida et al., 2022). Furthermore, a more
 454 detailed analysis of the biases identified in T2m and WS10m, including an examination of the
 455 diurnal variation of these parameters, is presented in the following sections.

456 **Table 3: Statistical verification scores for evaluation against weather station data:** skill scores for
 457 air temperature at 2m (T2m), wind speed at 10m (WS10m) and downward shortwave radiation flux
 458 (SR) for 16 meteorological stations (categorised into land and marine regions) over the United Arab
 459 Emirates (UAE).

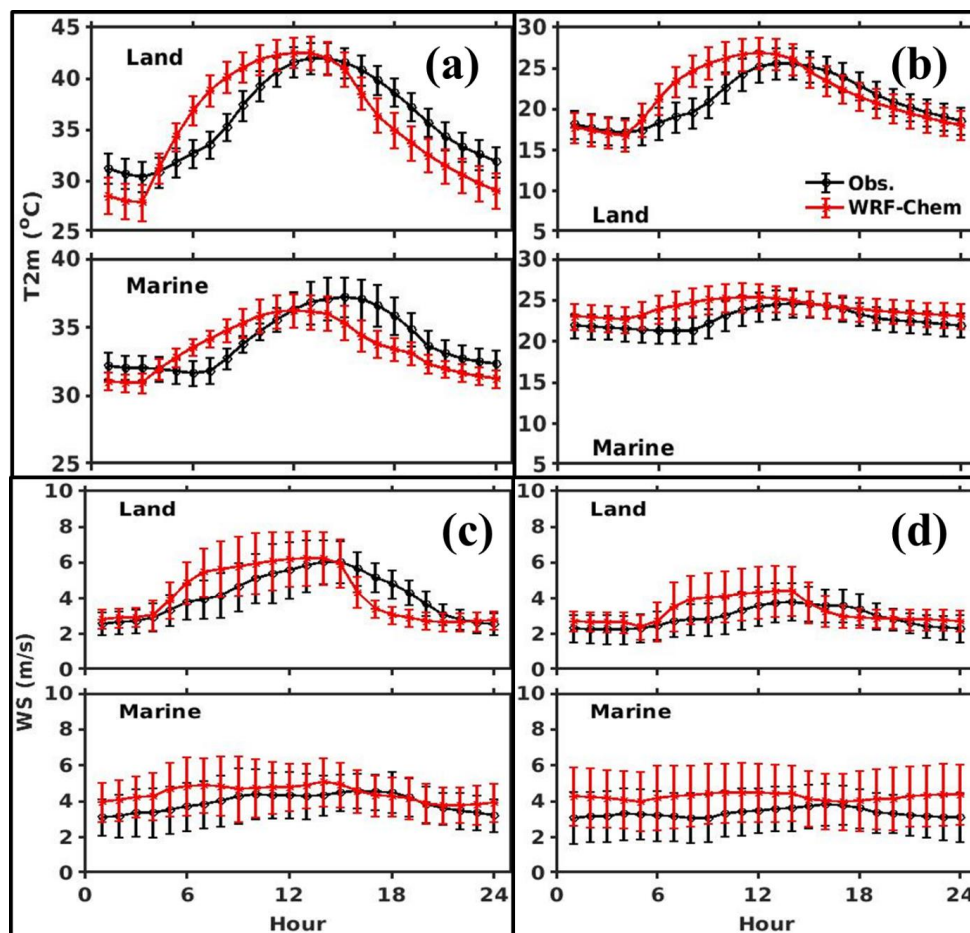
Parameter	Month	Region	MOD	OBS	MB	MAE	R	RMSE
T2m (°C)	June	Land	35.70	36.07	-0.37	2.68	0.91	3.57
		Marine	33.54	34.03	-0.48	1.47	0.83	1.67
	Dec	Land	21.84	21.08	0.76	1.66	0.92	2.87
		Marine	24.02	22.72	1.30	1.37	0.90	2.57
WS10m (m/s)	June	Land	4.24	4.16	0.08	0.90	0.88	1.35
		Marine	4.44	3.92	0.51	1.01	0.78	1.09
	Dec	Land	3.29	2.91	0.38	0.63	0.88	0.95



		Marine	4.26	3.35	0.92	1.12	0.89	1.54
SR (W/m ²)	June	Land	352.0	279.7	72.4	197.1	0.87	327.1
		Marine	349.3	264.9	84.4	273.4	0.68	358.7
	Dec	Land	192.7	177.2	15.5	124.3	0.85	231.2
		Marine	183.8	171.7	12.1	188.8	0.59	240.7

460

461 Figure 2, (a) and (b), presents a comparative analysis of the average diurnal variation in T2m
462 from WRF-Chem simulations and observations at both land and marine sites investigated in
463 this study, for the summer and winter seasons of 2018, respectively. The observed and
464 modelled T2m data exhibit a close alignment over land and marine locations, although some
465 discrepancies are evident. During the daytime, there is a tendency for the model to exhibit a
466 warm bias, while at night and evening, a cold bias is more apparent. Such discrepancies in
467 temperature have been reported before (Abida et al., 2022; Branch et al., 2021; Fonseca et al.,
468 2021; Schwitalla et al., 2020; Temimi et al., 2020a). Overall, the WRF-Chem model displays
469 a consistent cold bias of less than 0.5 °C for both environments during the summer months. In
470 contrast, during winter, the model shows a warm bias ranging from 0.8 to 1.3 °C. This is in
471 contrast to findings by Branch et al. (2021), which indicated an increase in the nocturnal cold
472 bias from winter to summer. Conversely, our study identifies a cold bias in the summer and a
473 warm bias in the winter, persisting throughout the entire day over marine locations. The
474 decrease in cold bias observed during summer in WRF-Chem simulations is a result of
475 enhanced representations of updated surface and soil parameters over the study region.



476

477 **Figure 2: Air temperature and wind speed diurnal cycle:** Diurnal cycles of spatial mean values of
 478 WRF-chem simulated (red) and observed (blue) air temperature at 2m (T2m; °C) in (a) (summer) and
 479 (b) (winter) for the regional categories of land and marine sites (c)-(d) are as (a)-(b) but for the wind
 480 speed at 10 m (WS10m; m/s). The averaged spatial standard deviation is represented by an error bar at
 481 each hour.

482 Figure 2, (c) and (d), showcase a comparative analysis of the mean diurnal variation in WS10m
 483 from model simulations and observations at both land and marine sites examined in this study,
 484 during the summer and winter of 2018, respectively. In both seasons, higher wind speeds are
 485 observed over marine sites, while lower wind speeds are found over land sites, reflecting sea
 486 and land circulations, respectively. It is indicated that wind speeds are higher during the
 487 daytime and lower during the night and evening hours. This pattern is especially pronounced
 488 over land sites compared to marine sites during both seasons. WRF-Chem tends to overestimate



489 WS10m during both day and night, across all regions and seasons. Nonetheless, the model
490 shows the smallest discrepancies over land, with biases being the least significant at 0.1 m/s
491 during summer and 0.4 m/s in winter. In contrast, the biases over marine areas are more
492 pronounced, at 0.5 m/s in the summer and 0.9 m/s in the winter. WRF-Chem tends to
493 overestimate WS10m more significantly during winter, with less overestimation observed
494 during summer. This discrepancy is linked to alterations in wind direction driven by land and
495 sea breeze circulations. Consequently, numerous studies have previously emphasized the
496 model's tendency for wind speed overprediction (Abida et al., 2022; Branch et al., 2021;
497 Fonseca et al., 2021; Schwitalla et al., 2020; Temimi et al., 2020a).

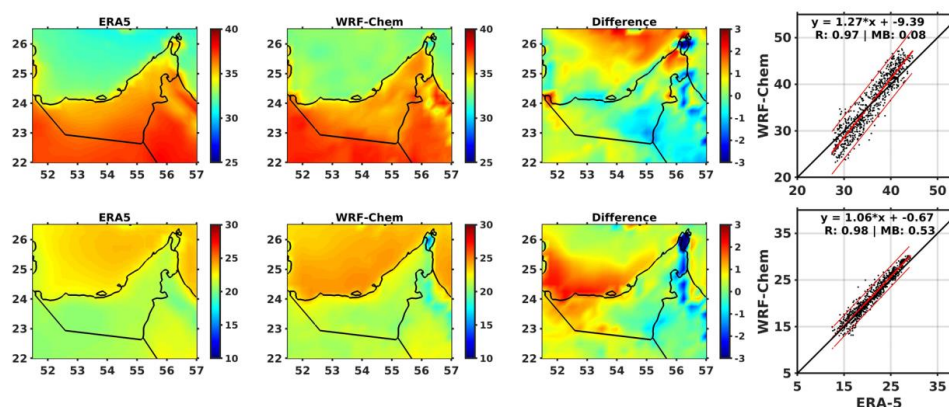
498 **4.1.2 Evaluation against ERA5 reanalysis data**

499 In Fig. 3, a spatial comparison is presented between the averaged ERA5 T2m and the
500 corresponding WRF-chem simulation output across the simulation domain during June and
501 December of 2018. The model adeptly captures regional temperature variations, displaying
502 underestimation in the southern regions and overestimation in the north-western region of the
503 UAE. This observation suggests a comprehensive portrayal of temperature dynamics by the
504 model, with specific tendencies in certain geographical areas. This observation is also
505 supported by NCM data, for instance, at Mezaria (ID No: 3), which represents a southern land
506 site, and at Abu Dhabi (ID No: 15), representing a northern marine site within the emirate of
507 Abu Dhabi. The southern land site found an underestimation of 1 °C, while the northern marine
508 site exhibited an overprediction of T2m by WRF-chem. WRF-Chem overestimates the area-
509 averaged temperature (T2m) over the UAE compared to ERA5 in both seasons. In contrast,
510 NCM observations indicate an underestimation during the summer and an overestimation
511 during the winter across the majority of sites. Kishta et al., (2023) reported that, minor
512 discrepancies in temperature measurements between observational data and ERA5 reanalysis,
513 identifying a strong correlation coefficient of 0.89 over Abu Dhabi. The spatial average of
514 WRF-Chem and ERA5 values are 35.8 °C and 35.7 °C, respectively, with a small
515 underestimation of 0.08 °C over the UAE. The model displays a high correlation (r) of 0.97 and
516 a RMSE of 2.3 °C, MAE of 2.2 °C in June. For December, the model showed a similar pattern,
517 with a underestimation of 0.53 °C which is slightly higher as compared to June, r of 0.98, MAE
518 of 1.0 °C and RMSE of 1.1 °C (Table 4).

519 Moreover, the analysis of the absolute differences between the two datasets highlighted the
520 most pronounced discrepancies over the Arabian Gulf region, observable in both the summer
521 and winter months. However, these discrepancies are notably more emphasised during the



522 warmer months. WRF-simulated Sea Surface Temperatures (SSTs) are compared with both
 523 ERA5 and Group for High Resolution Sea Surface Temperature (GHRSSST) data over the
 524 Arabian Gulf region (not shown). The comparison involved area-averaged daily values,
 525 considering that the diurnal amplitude of SST is 0.5 °K over this region as reported by Nesterov
 526 et al., (2021). The model showed an overestimation of 1.4 °K compared to both ERA5 and
 527 GHRSSST during the summer. Similarly, in winter, its overestimated SSTs by 1.5 °K compared
 528 to ERA5 and by 1.3 °K compared to GHRSSST. Furthermore, the model exhibited a
 529 significantly higher correlation in winter, achieving a correlation coefficient of 0.9 with both
 530 datasets. However, during the summer, it displayed variable correlations, with $r=0.38$ for ERA5
 531 and $r=0.20$ for GHRSSST. This observation suggests potential inaccuracies in the model
 532 simulation of temperature and wind speed in this region, which could be due to the sea surface
 533 temperature data utilized for model forcing. The temperature gradient plays a pivotal role in
 534 driving the land-sea breeze circulation. Higher temperatures observed over the Gulf could
 535 potentially weaken this circulation pattern, resulting in reduced transportation of cleaner
 536 marine air towards inland areas. Consequently, this reduction in the influx of marine air could
 537 obstruct the effective dispersion of pollutants across terrestrial regions, negatively impacting
 538 air quality and the spatial distribution of pollutants.



539

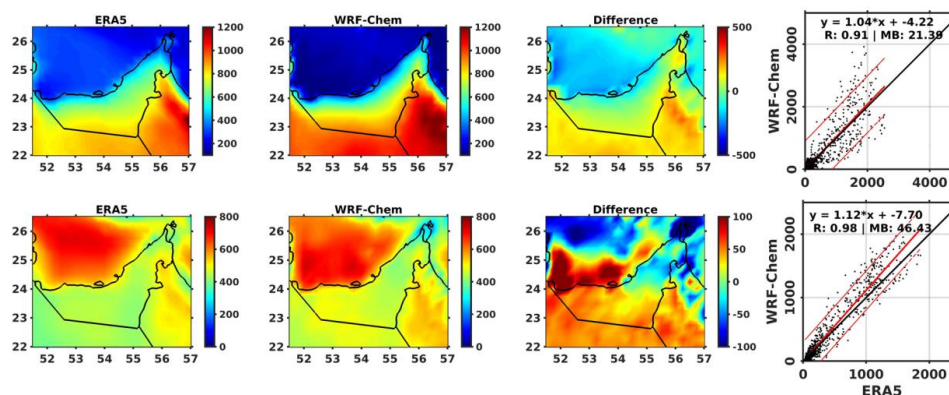
540 **Figure 3: ERA-5 and WRF-Chem Air Temperature:** Average 2-m air temperature (°C) obtained
 541 from ERA5 reanalysis (first panel), simulated by WRF-Chem (second panel), and the corresponding
 542 absolute differences (third panel) and scatter plots between the two datasets (fourth panel) during June
 543 (top) and December (bottom) 2018.

544 It is widely recognized that the Planetary boundary layer (PBL) plays a crucial role in the
 545 pollution transport process over the region. It constitutes the lowest part of the troposphere and
 546 is directly influenced by the Earth's surface. The PBL reaches higher elevations during summer,



547 with its altitudes being lower in winter. There are noticeable differences in the PBL between
548 land areas (approximately 2400–2500 m) and marine regions (about 1200–1500 m) (Basha et
549 al., 2019). Basha et al. (2019) also discovered that ERA-Interim reanalysis data tend to
550 underestimate PBL when compared with data obtained from Global Positioning System Radio
551 Occultation (GPSRO) in most regions and in all the seasons. Chen et al., (2022) emphasized
552 the critical role of the boundary layer in influencing air quality and facilitating the
553 transboundary transport of pollutants. They noted that a higher boundary layer enhances the
554 potential for pollutant transport to the Tibetan Plateau. Wang et al., (2022) highlighted the
555 critical role of meteorological conditions in severe PM_{2.5} pollution episodes. They noted that
556 rapid cold air movement can quickly disperse pollutants, in contrast to the slow accumulation
557 of pollutants under weak high-pressure systems. This slow build-up is characterized by low
558 wind speeds, and low atmospheric boundary layer heights, which lead to prolonged heavy
559 pollution periods.

560 In this study, we aim to compare the PBL as simulated by WRF-Chem with the ERA5
561 reanalysis, providing further specifics of model accuracy and performance. Fig. 4 shows a
562 comparison of the mean ERA5 PBL with corresponding WRF-chem simulated values over the
563 UAE for the months of June and December 2018. The absolute difference and scatter plot for
564 these data sets are also shown. The spatial distribution of PBL across the UAE, as from ERA5
565 data, exhibits a consistent spatial pattern that aligns with the PBL simulated by WRF-Chem.
566 There is a notable trend of increased PBL during the summer months and decreased PBL in the
567 winter. This pattern generally corresponds with the seasonal temperature variations, where
568 warmer summer temperatures contribute to an elevation in PBL, and cooler winter
569 temperatures result in a reduction of PBL (Basha et al., 2019). In terms of PBL (averaged
570 spatially for the UAE), the model exhibits good performance in capturing the regional
571 variations. In June, the modelled PBL is at 669.8 m compared to 646.7 m in ERA5, with a
572 correlation coefficient of 0.91 and a RMSE of 450.1 m. In December, the modelled PBL is
573 490.5 m compared to the ERA5 of 444.2 m, with a high correlation coefficient of 0.98 and an
574 RMSE of 152.8 m (Table 4).



575

576 **Figure 4: ERA-5 and WRF-Chem Boundary Layer Height:** Same as Fig. 3, but for planetary
 577 boundary layer height (PBL).

578

579 In addition to T2m and PBL, Table 4 also summarizes the spatially averaged statistical
 580 verification scores for WS10m and SR over UAE. Regarding WS10m, it is accurately
 581 simulated by the model with small differences in MB (June: 0.08 m/s, Dec: 0.01 m/s), which
 582 are slightly larger compared to observations from land-based sites in Abu Dhabi and good
 583 correlations (June: 0.79, Dec: 0.80). The RMSE values are 1.7 m/s for June and 1.1 m/s for
 584 December. For SR, the model performs well, capturing the variability in radiation flux. In June,
 585 the modelled SR is 643.6 W/m² compared to the ERA5 of 576.5 W/m², with a high correlation
 586 of 0.99 and an RMSE of 75.3 W/m². Similarly, in December, the modelled SR is 460.8 W/m²
 587 compared to the ERA5 of 438.1 W/m², with a correlation of 0.97 and an RMSE of 76.1 W/m².
 588 Overall, these results indicate a very good performance of the WRF-chem model in simulating
 589 meteorological parameters over the UAE during the specified months. This rigorous evaluation
 590 of meteorological parameters showed that WRF-Chem's simulated values closely align with
 591 both ground-based and reanalysis datasets. Since WRF-Chem simulates meteorology and
 592 chemistry simultaneously, accurate meteorological simulations are crucial for the precise
 593 computation of chemistry within the model domain.

594

595 **Table 4: Statistical verification scores for evaluation against ERA-5 data:** skill scores calculated
 596 for model simulations for air temperature at 2m (T2m), wind speed at 10m (WS10m), downward
 597 shortwave radiation flux (SR) and planetary boundary layer (PBL) during June and December of 2018
 598 over the United Arab Emirates.

599



Parameter	Month	MOD	ERA5	MB	MAE	R	RMSE
T2m (°C)	June	35.82	35.73	0.08	2.17	0.97	2.28
	Dec	21.61	21.08	0.53	0.99	0.98	1.12
WS10m (m/s)		4.34	4.26	0.08	1.26	0.79	1.7
		3.05	3.07	0.01	0.87	0.8	1.1
SR (W/m ²)		643.6	576.5	67.1	85	0.99	75.3
		460.8	438.1	22.8	69.5	0.97	76.1
PBL (m)		669.8	646.7	21.4	271.6	0.91	450.1
		490.5	444.2	46.4	113.8	0.98	152.8

600

601 **4.2 Model performance with respect to gaseous pollutants**

602 The study incorporates comparative assessments with satellite data from the TROPOMI
 603 instrument. This includes evaluations of the tropospheric column of NO₂ (denoted as
 604 TROPOMI-NO₂), total column CO (TROPOMI-CO), and total column ozone (TROPOMI-O₃)
 605 for the corresponding periods within the UAE. Detailed outcomes of these comprehensive
 606 assessments are discussed in the following subsections. The WRF-Chem model exhibited
 607 commendable proficiency in replicating the satellite-derived measurements of these pollutants
 608 throughout the UAE over the summer and winter seasons of 2018. The satellite overpass takes
 609 place daily at 13:30 local time; therefore, model simulations corresponding to this time are
 610 utilized here for comparison over the study area. After smoothing the model concentrations
 611 using the a priori and averaging kernel matrix as detailed in Section 3.4, the results were
 612 compared with the corresponding TROPOMI products.

613

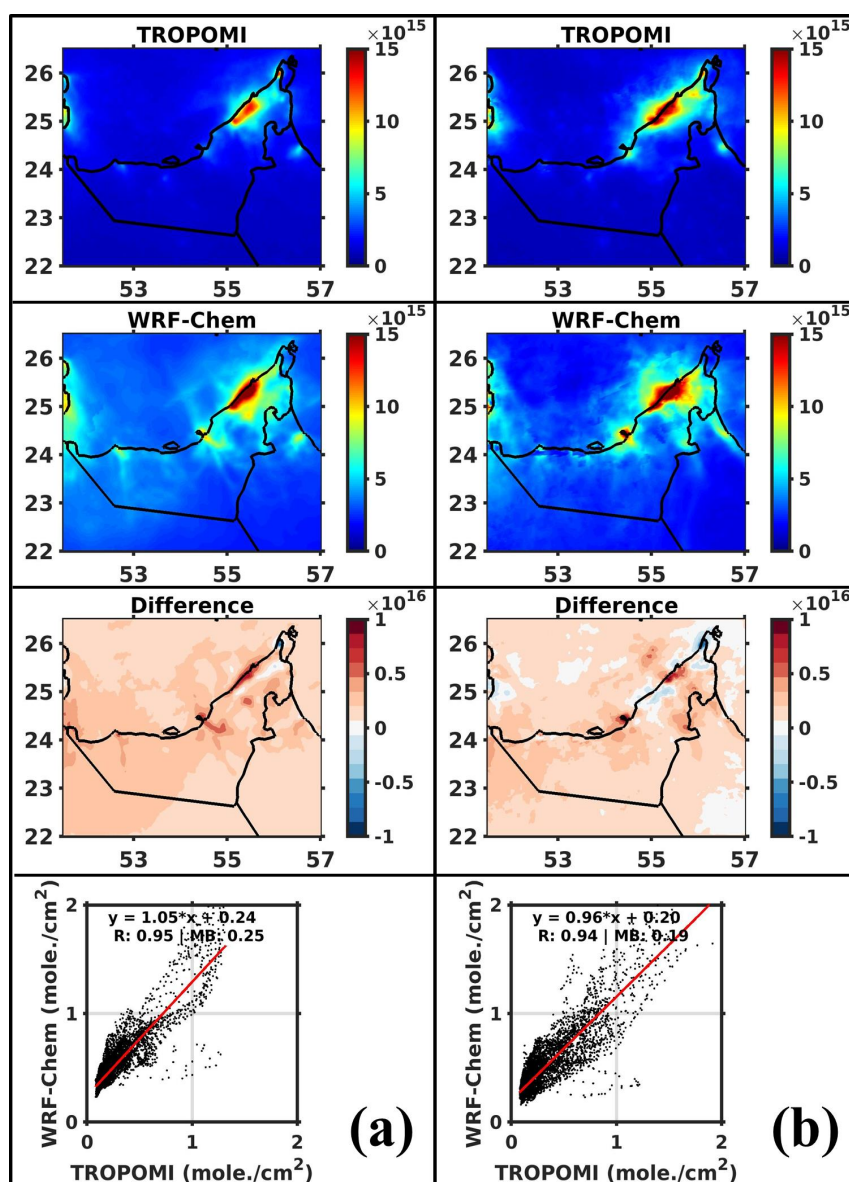
614 In the troposphere, oxides of nitrogen (NO_x=NO+NO₂) are crucial for the mechanisms of
 615 ozone production and depletion in the presence of sunlight. Due to their shorter lifespan, their
 616 concentrations are primarily linked to emission sources. As a result, NO_x is more susceptible
 617 to inaccuracies in emission estimates compared to other criteria pollutants. The Environment
 618 Agency – Abu Dhabi (2018) reported that oil and gas, road transport, and electricity generation
 619 are the primary sectors contributing to NO_x total emissions, accounting for 42%, 34%, and
 620 13% respectively, for the base year of 2015 in the Emirate of Abu Dhabi. In Fig. 5, the average
 621 spatial distributions of both model-simulated and TROPOMI-retrieved tropospheric column
 622 NO₂ are presented. Additionally, the spatial discrepancies between simulated and retrieved



623 columns are illustrated by absolute differences (see third row) and scatter plots between the
624 two datasets are depicted (see fourth row) for June (left) and December (right) 2018 across the
625 study region. The satellite retrievals indicated elevated levels of NO₂ column, exceeding
626 12×10^{15} molecules/cm², in densely populated and industrial areas and the adjacent regions of
627 Dubai and Abu Dhabi in both summer and winter. Conversely, lower NO₂ values, less than
628 5×10^{15} molecules/cm², were observed over the less urbanized areas. The higher columns are
629 associated with significant economic development driven by a high demand in power
630 generation and water desalination projects, which primarily depends on the combustion of
631 fossil fuels in big cities like Dubai and Abu Dhabi (Abuelgasim & Farahat, 2020; Li et al.,
632 2010). The model effectively reproduced the spatial distributions of NO₂ during summer and
633 winter of 2018 as depicted in Fig. 5. Although, the model overestimation is close to zero in
634 rural areas, it can be as high as 10^{16} molecules/cm² in areas of high pollution, specifically over
635 Dubai and Abu Dhabi. Conversely, it underestimates up to 10^{16} molecules/cm² in the Ras Al
636 Khaimah emirate; the sixth-largest city by population and home to the global ceramic
637 manufacturing company, RAK Ceramics. This observation is not unexpected, as urban and
638 industrial areas frequently report elevated pollutant emissions stemming from urban activities,
639 which are significantly high and present challenges that models often struggle to accurately
640 capture these changes. This discrepancy also suggests that anthropogenic and industrial
641 emissions might be improperly represented in the EDGAR emission inventory. Challenges
642 range from the incomplete characterization of emissions in source regions to the impact of
643 model resolution on capturing sub-grid emission sources. Additionally, Hoshyaripour et al.,
644 (2016) found that the PBL is shallower and more stable at night when simulated with the YSU
645 boundary layer scheme, resulting in a higher accumulation of NO_x in the surface layers. Such
646 insights were constrained in the present model evaluation, which is primarily focusing on
647 temporal variability of gaseous pollutants on a daily basis, and did not encompass diurnal
648 variations. Incorporating these diurnal variations in future model simulations over this region
649 may enhance the assessment's accuracy. Additionally, the existing model configuration does
650 not include the formation of secondary aerosols in its simulations, indicating a potential area
651 for improvement in future versions. Additionally, the absence of vertical distribution of
652 anthropogenic emissions in the model simulations also plays a pivotal role in these model
653 discrepancies. The satellite retrieved TROPOMI-NO₂ averaged for the UAE is 0.21×10^{16}
654 molecules/cm² in summer and 0.24×10^{16} molecules/cm² in winter. The corresponding model
655 simulated column is 0.46×10^{16} and 0.43×10^{16} molecules/cm² respectively. The model
656 demonstrated a strong correlation with satellite NO₂ column measurements, achieving



657 correlation coefficients of 0.95 for summer and 0.94 for winter (refer to Table 5). It showed a
 658 slight tendency to overestimate NO_2 levels more in summer, with a discrepancy of 0.24×10^{15} ,
 659 compared to 0.19×10^{15} molecules/ cm^2 in winter. Moreover, the evaluation shows RMSE
 660 values of 0.1×10^{15} to 0.12×10^{15} molecules/ cm^2 and MAE values of 0.20 to 0.25×10^{15}
 661 molecules/ cm^2 during the seasons.
 662



663



664 **Figure 5: Evaluation of WRF-chem against satellite-derived NO₂:** average spatial pattern of
665 tropospheric column NO₂ (mole. /cm²) obtained from TROPOMI satellite (1st row), simulated by WRF-
666 Chem (2nd row), corresponding absolute difference (model minus TROPOMI) (3rd row) and scatter plots
667 between two daily data sets (4th row) during (a) June and (b) December in 2018.

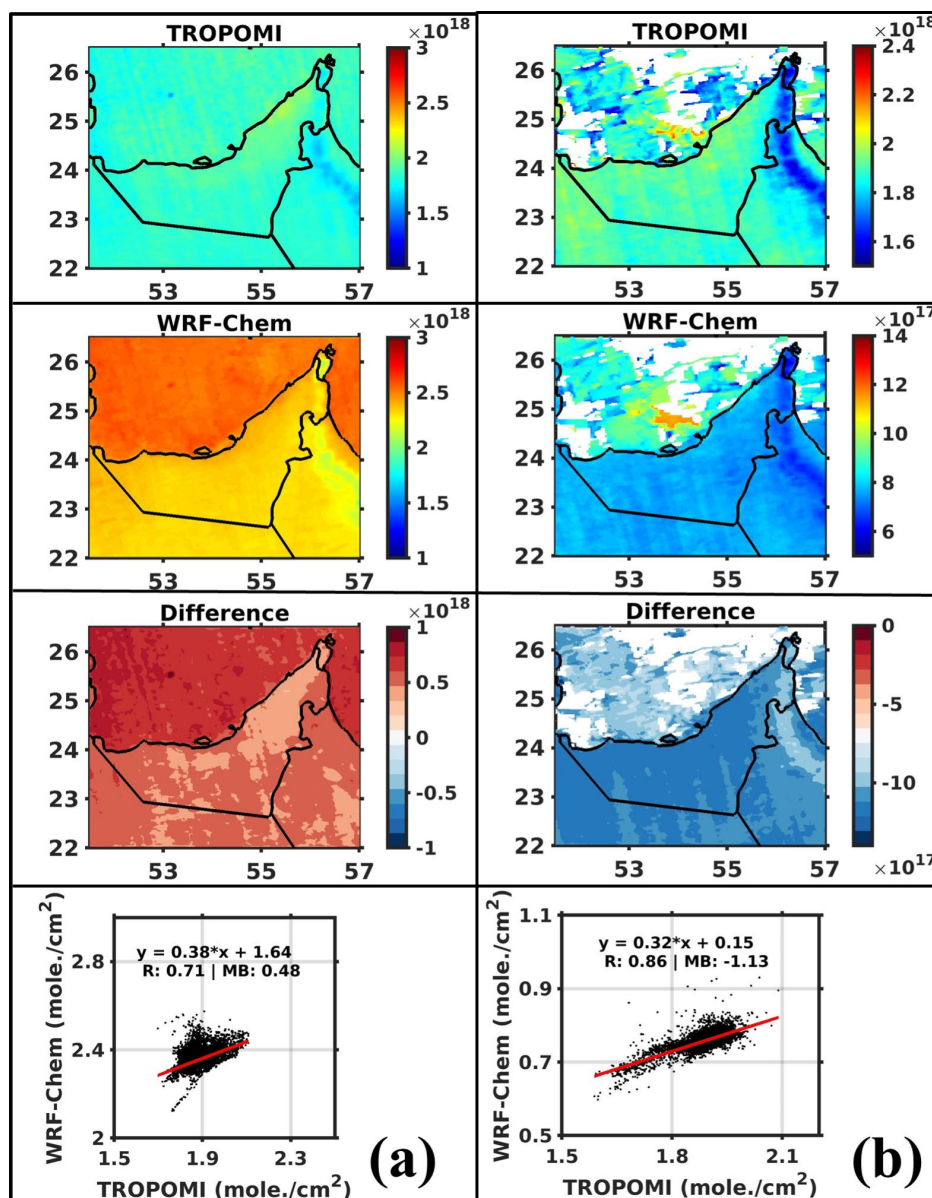
668

669 In Fig. 6, the average spatial distributions of both model-simulated and TROPOMI-retrieved
670 total CO column are presented. Also, the absolute difference of WRF-Chem simulations with
671 TROPOMI-CO is depicted along with scatter plots between them during summer and winter
672 of 2018 over UAE. The statistical metrics comparing these datasets are provided in Table 5.
673 The TROPOMI-retrieved CO columns display values of 1.87 and 1.89 x 10¹⁸ molecules/cm²
674 for summer and winter, respectively. In contrast, the simulated columns show values of 2.35
675 for summer and 0.76 x 10¹⁸ molecules/cm² for winter. Thus, comparing WRF-Chem and
676 TROPOMI-CO data reveals more pronounced discrepancies, with a minor overestimation of
677 0.48 x 10¹⁸ molecules/cm² in summer and a significant underestimation of 1.13 x 10¹⁸
678 molecules/cm² in winter. Shami et al., (2022) discovered that the EDGAR emissions inventory
679 underestimates CO emissions when compared to Lebanon's national emission inventory,
680 identifying the road transport sector as the primary source of CO emissions. Consequently,
681 EDGAR's estimates for CO emissions are lower than those provided by Waked et al., (2012)
682 for the same region. The Environment Agency – Abu Dhabi (2018) reported that the road
683 transport sector is the primary source of CO emissions in Abu Dhabi, accounting for 74% of
684 the total CO emissions. Additionally, the industrial sector contributes 21% to the total CO
685 emissions. Kumar et al. (2022) observed an underestimation of CO by the WRF-Chem model,
686 attributing it to an inaccurate representation of anthropogenic emissions on the vertical scale.
687 This could result in a more rapid deposition of CO molecules at the surface, thereby leading to
688 the observed underestimation.

689 The model output correlates reasonably well with TROPOMI-CO with r of 0.71 and 0.86 while
690 RMSE of 0.04 to 0.02 x 10¹⁸ molecules/cm² in summer and winter respectively (Table 5). In
691 both seasons, the lower correlation coefficients for TROPOMI-CO as compared to TROPOMI-
692 NO₂ suggest a less robust linear relationship between the TROPOMI and WRF-chem CO
693 levels. This variation in performance might be attributed to the complexities inherent in
694 modelling and observing CO distributions, which can be influenced by local emission sources,
695 atmospheric chemistry, and transport processes. These findings are consistent with research
696 conducted in India, where Dekker et al. (2019) reported a correlation of 0.81 between
697 TROPOMI and WRF-Chem CO levels during a high pollution episode during November 2017.



698 Similarly, in East Asia, Zhang et al. (2016a) documented correlations between WRF-Chem
 699 simulated and MOPITT retrieved CO columns, with r of 0.59 and RMSE of 4.6×10^{17}
 700 molecules/cm² for summer, and 0.69 with RMSE of 5.2×10^{17} molecules/cm² for winter,
 701 respectively.



702



703 **Figure 6: Evaluation of WRF-Chem against satellite-derived CO:** Same as Fig. 5 but for total
704 column of CO

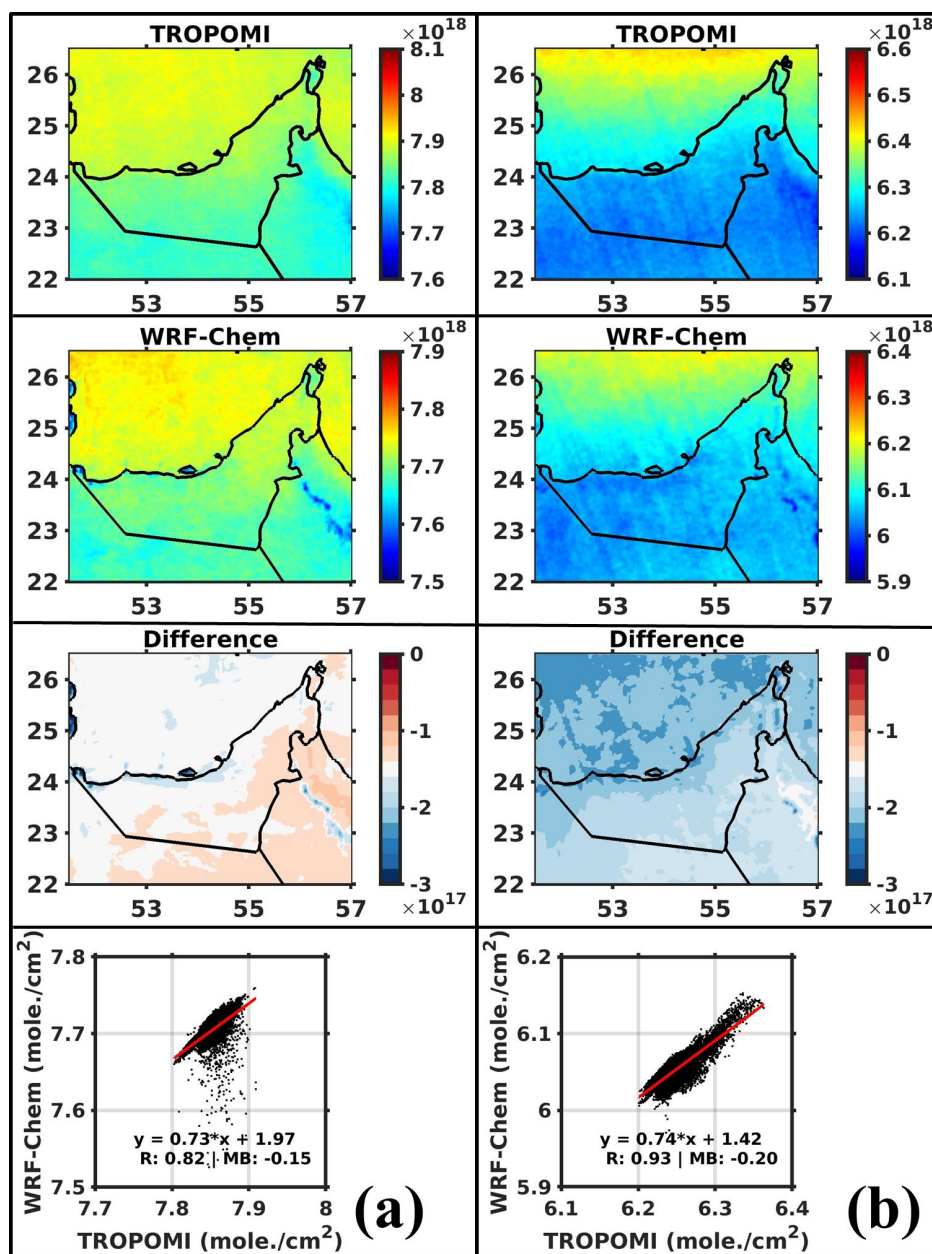
705 We also conducted a comparison of WRF-Chem simulated ozone levels with TROPOMI-
706 retrieved total columns (TROPOMI-O₃), as illustrated in Fig. 7. This figure also presents both
707 the absolute differences (3rd row) and scatter plots (4th row) between the two datasets for both
708 seasons. The statistical comparisons between these datasets are detailed in Table 5. The
709 TROPOMI-O₃ columns show higher values in summer, at 7.85×10^{18} molecules/cm², and
710 lower values in winter, at 6.25×10^{18} molecules/cm². The WRF-Chem simulations closely
711 match these variations, with values of 7.70×10^{18} molecules/cm² for summer and 6.06×10^{18}
712 molecules/cm² for winter, respectively. Therefore, model output is strongly correlated to
713 TROPOMI-O₃ columns with correlation of $r=0.82$ and 0.93 while RMSE (MAE) of $0.01(0.15$
714 and 0.20×10^{18} molecules/cm²) during summer and winter respectively. Many studies
715 commonly report higher ozone concentrations in the summer and lower concentrations in the
716 winter, a phenomenon primarily attributed to increased photochemical activity during the
717 summer months (Reddy et al., 2012; Coates et al., 2016; Badia & Jorba 2015; Abdallah et al.
718 2018; Baldasano et al. 2011). The WRF-Chem model systematically underestimates ozone
719 levels, with 0.15 and 0.20×10^{18} molecules/cm² both seasons respectively. Hu et al., (2021)
720 highlighted that meteorological factors have a considerable effect on ozone production, noting
721 from studies in China that temperature, relative humidity, and sunshine duration significantly
722 influence ozone concentrations in descending order of importance. They also noted that strong
723 solar radiation and elevated temperatures could enhance photochemical reactions, thereby
724 increasing ozone formation. Zhang et al., (2020) pointed out that low wind speeds and high
725 atmospheric pressure can impede the dispersion and dilution of pollutants, which in turn can
726 lead to higher ozone accumulation. Lu et al., (2019) observed that high humidity conditions,
727 with increased water vapor, could cause more significant chemical depletion of O₃, as water
728 vapor interacts with excited ozone molecules to produce OH radicals. Hence, the
729 meteorological conditions are conducive to ozone formation in the model but are insufficient
730 to fully account for the model's significance underprediction of O₃. Sillman, (1999)
731 demonstrated the ozone formation potential by its precursors being highly nonlinear rather than
732 linear. Ozone formation can be either NO_x-sensitive, meaning O₃ formation increases with an
733 increase in NO_x concentration, or VOC-sensitive, where O₃ formation increases with an
734 increase in VOC concentration. However, Geng et al., (2007) observed that high NO_x
735 concentrations in urban environments result in reduced OH radical levels, consequently



736 decreasing ozone production, as loss of OH is evidenced by the chemical reaction $\text{NO}_2 + \text{OH}$
737 $\rightarrow \text{HNO}_3$. This observation is consistent with model simulations showing increased NO_2 levels
738 but markedly lower ozone concentrations at an urban area in the UAE, illustrating the
739 significant impact of NO_x on urban ozone formation. However, drawing such conclusions
740 requires careful analysis of model simulations, suggesting that future work, particularly in the
741 refinement of WRF-Chem evaluations, is essential. Future simulations should not only
742 critically assess these findings but also aim to improve model fidelity by enhancing the
743 representation of chemical processes and emissions. Adopting this approach will lead to more
744 precise forecasts and a more profound grip of atmospheric chemistry, thereby enhancing air
745 quality projections and fostering a more detailed understanding of pollution patterns over this
746 region.

747 The disparities between WRF-Chem and TROPOMI data highlight the intrinsic challenges
748 in air quality monitoring and prediction. WRF-Chem's limitations may stem from its
749 dependency on emissions inventories, meteorological data, and the representation of
750 atmospheric chemistry. TROPOMI, while offering high-resolution satellite observations, is
751 subject to constraints related to retrieval algorithms and the influence of atmospheric conditions
752 on measurement accuracy. Liu et al., (2022) identified that uncertainties in column
753 observations stem from the challenges in differentiating between stratospheric and tropospheric
754 contributions, as well as uncertainties in the tropospheric air mass factor and its spectral fitting.
755 The integration of model predictions with satellite observations, alongside ground-based
756 measurements, is crucial for enhancing our understanding of air quality dynamics and
757 improving predictive capabilities. This synergistic approach can help mitigate biases, enhance
758 accuracy, and provide a more comprehensive view of atmospheric pollutants' distribution over
759 this region.

760



761

762 **Figure 7: Evaluation of WRF-Chem against satellite-derived O_3 :** Same as Fig. 5 but for total column

763 of ozone.

764

765



766 **Table 5: Statistical verification scores for evaluation against TROPOMI measurements:** skill
 767 scores between TROPOMI columns (mole. /cm²), tropospheric column NO₂ (TROPOMI-NO₂), total
 768 column carbon monoxide (TROPOMI-CO) and total column ozone (TROPOMI-O₃) with
 769 corresponding WRF-chem simulated columns during June and December of 2018 over UAE. Means
 770 and MB, MAE and RMSE are given in units of (x10¹⁵ mole. / cm² for TROPOMI-NO₂.
 771

Parameter	Month	MOD	SAT	MB	MAE	R	RMSE
NO ₂ (x10 ¹⁶)	June	0.46	0.21	0.25	0.25	0.95	0.10
	Dec	0.43	0.24	0.19	0.20	0.94	0.12
O ₃ (x10 ¹⁸)		7.70	7.85	-0.15	0.15	0.82	0.01
		6.06	6.25	-0.2	0.20	0.93	0.01
CO (x10 ¹⁸)		2.35	1.87	0.48	0.48	0.71	0.04
		0.76	1.89	-1.13	1.13	0.86	0.02

772

773 5. Conclusions

774 This study rigorously evaluates the performance of the Weather Research and Forecasting
 775 model coupled with chemistry (WRF-Chem). The model ability to simulate meteorological
 776 parameters and gaseous pollutants over the United Arab Emirates (UAE) is assessed during
 777 June and December 2018 to reflect contrasting summer and winter conditions. The model
 778 performance is assessed through comparison with ground-based observations and ERA-5
 779 reanalysis data for meteorological parameters, and TROPOMI satellite observations for
 780 gaseous pollutants.

781

782 We evaluated WRF-Chem model's accuracy in simulating meteorological parameters, in
 783 particular 2-meter temperature (T2m), 10-meter wind speed (WS10m), and solar radiation
 784 (SR), across 16 locations in the UAE. The model generally underestimates T2m in summer by
 785 less than 0.5 °C and overestimates it in winter by less than 1.3 °C, with correlation coefficients
 786 ranging from 0.7 to 0.9 among the stations. WRF-chem performance for WS10m and SR has
 787 shown high scores, indicating enhanced accuracy across the locations. Regionally, it slightly
 788 underpredicts T2m in summer (by 0.37 °C for land and 0.48 °C for marine) mainly due to
 789 colder nights, and overestimates in winter (by 0.76 °C for land and 1.30 °C for marine), both
 790 with strong correlations above 0.83. Higher SR values in summer and winter, suggest reduced



791 cloud cover and aerosol loading in WRF-Chem. For WS10m, the model's bias is within ± 1 m/s,
792 and correlation coefficients range between 0.78 and 0.89, indicating good agreement for both
793 land and marine areas.

794 The comparison of ERA5 reanalysis data with WRF-Chem simulations revealed regional
795 variations in T2m, specifically underestimation in the UAE's south and overestimation in the
796 north-west. The most significant differences were observed over the Arabian Gulf region,
797 especially during warmer months. These temperature discrepancies are crucial for the land-sea
798 breeze circulation, with higher Gulf temperatures potentially weakening this pattern. This could
799 lead to diminished transport of cleaner marine air inland, thereby hindering pollutant dispersion
800 over land and adversely affecting air quality and pollutant distribution. Statistical metrics for
801 summer shows an overestimation of 0.08 °C and a correlation coefficient (r) of 0.97, while
802 winter's follows a similar pattern with an overestimation of 0.53 °C and r of 0.98 over land
803 mass region of UAE. The fact that WRF-Chem performs well against in-situ data and ERA5
804 reanalysis with respect to air temperature is an indication the reanalysis dataset performs well
805 in this region. The mean PBL from ERA5 is largely consistent with that from the WRF-Chem
806 outputs, with both data sets displaying a clear seasonal variation—increased PBL during
807 summer and decreased in winter, correlating with temperature changes. June's modelled PBL
808 has a correlation of 0.91, and December's correlation of 0.98 with ERA5.

809 Regarding gaseous pollutants, both WRF-Chem and satellite data show higher TROPOMI-
810 NO₂ columns greater than 12×10^{15} molecules/cm² in urban and industrial regions such as
811 Dubai, Abu Dhabi and Ras Al Khaimah emirate, and reflecting emissions from economic
812 activities like power generation, water desalination and industries. Lower concentrations of
813 less than 5×10^{15} molecules/cm² are noted in less urbanized areas. The WRF-Chem model
814 closely reproduces TROPOMI-NO₂ spatial patterns. However, it overestimates NO₂ in the Abu
815 Dhabi region and underestimates it in north-eastern UAE. High correlation coefficients (0.95
816 in summer and 0.94 in winter) confirm the model's effectiveness in capturing NO₂'s day-to-
817 day variability. The model shows minimal MB and high r values, indicating small discrepancies
818 in NO₂ estimations. Moreover, the WRF-Chem underestimates TROPOMI-O₃ columns, as
819 indicated by negative MB values, yet maintains high correlation coefficients (0.82 in summer
820 and 0.93 in winter), suggesting accurate ozone concentration simulations. TROPOMI-CO
821 column simulations, however, exhibit significant discrepancies and lower correlation
822 coefficients (0.71 in summer and 0.86 in winter), highlighting challenges in accurately



823 modelling CO levels. This analysis stresses the WRF-Chem model's strengths in simulating
824 NO₂ and O₃ columns with high fidelity to TROPOMI observations but also points out its
825 limitations in estimating CO columns accurately.

826 The WRF-Chem model exhibits satisfactory capability in simulating key meteorological
827 parameters and gaseous pollutants over the UAE, showcasing significant improvements in
828 regional-scale dynamics. This is evidenced by strong correlation coefficients, variable MB,
829 RMSE and MAE values, and a clear enhancement over previous research outcomes. This
830 comprehensive assessment validates the model's effectiveness and identifies potential areas for
831 improvement in simulating gaseous pollutant concentrations across the UAE. The
832 discrepancies between model simulations and various observational data sets may arise from
833 improper emission inventories, particularly anthropogenic emissions, model parameterizations,
834 and meteorological inputs. Integrating model predictions with satellite observations and
835 ground-based measurements is crucial for advancing air quality monitoring and enhancing the
836 predictive accuracy of atmospheric pollutant distributions in the UAE. This collective approach
837 aids in addressing biases and improving the overall understanding of regional air quality
838 dynamics.

839

840 **Code and Data Availability**

841 The authors would like to thank the United Arab Emirates' National Center of Meteorology for
842 providing meteorological observations at 16 weather stations for the months of June and
843 December 2018 under an agreement with clauses for non-disclosure of data. Access to these
844 data is restricted and readers should request them through contacting research@ncms.ae. The
845 remaining products considered in this study are freely available online: (i) ERA-5 reanalysis
846 data is extracted from the Copernicus Climate Change Service Climate Data Store (Hersbach
847 et al. 2023a,b); (ii) Nitrogen Dioxide (NO₂), Ozone (O₃) and Carbon Monoxide (CO) column
848 concentrations estimated from the measurements collected by the Tropospheric Monitoring
849 Instrument (TROPOMI) onboard the Sentinel 5-P satellite are extracted from the National
850 Aeronautics and Space Administration's (NASA's) website; (iii) National Centers for
851 Environmental Prediction (NCEP) Final (FNL) Operational Global Analysis meteorological
852 data used to drive the WRF-chem simulations is downloaded from the National Center for
853 Atmospheric Research (NCAR) Research Data Archive website (NCEP/NWS/NOAA/USDC,
854 2000), with the chemistry data used to force WRF-Chem, the output of the Community



855 Atmosphere Model with Chemistry (CAM-chem) model, extracted from NCAR’s website
856 (Bucholz et al., 2019); (iv) the WRF-Chem model used, version 4.3.1, is freely available from
857 the developers’ website (WRF, 2023), with the pre-processor tools available at NCAR’s
858 website (NCAR, 2023). All figures displayed in this manuscript were generated with the Matrix
859 Laboratory (MATLAB) software version 2023 (Mathworks, 2023).

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873 **Conflict of interest**

874 The authors declare they do not have any conflict of interest.

875 **Author contribution**

876 Conceptualisation and methodology: D.F. and Y.Y.; Data curation and visualization: Y.Y.; formal
877 analysis and interpretation: Y.Y., R.F., N.N., and D.F.; project administration and supervision: D.F.;
878 writing—original draft: Y.Y.; review and editing: all authors.

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