

Importance of Aerosol Composition and Aerosol Vertical Profile in Global Spatial Variation ~~in~~ of the PM_{2.5} to AOD Relationship ~~Strongly Influenced~~ by Aerosol Composition

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Abstract Ambient fine particulate matter (PM_{2.5}) is the leading global environmental determinant of mortality. However, large gaps exist in ground-based PM_{2.5} monitoring. Satellite remote sensing of aerosol optical depth (AOD) offers information to fill these gaps worldwide, when augmented with a modeled PM_{2.5} to AOD relationship (η). This study aims to understand the spatial pattern and driving factors of η ~~the relationship by examining~~ η ($= \frac{PM_{2.5}}{AOD}$), from both observations and modeling. A global observational estimate of η for the year 2019 is inferred from 6,118–870 ground-based PM_{2.5} measurement sites and satellite retrieved AOD ~~from the MAIAC algorithm~~. ~~The GEOS-Chem~~ global chemical transport model, ~~GEOS-Chem~~, in its high performance configuration (GCHP), is used to interpret the observed spatial pattern of annual mean η . Measurements and the GCHP simulation consistently identify a global population-weighted mean η of ~~92–100~~ 96–98 $\mu\text{g}/\text{m}^3$, with regional values ranging from ~~60.359.8~~ 60.359.8 $\mu\text{g}/\text{m}^3$ ~~for in~~ North America to more than ~~130–190~~ 130–190 $\mu\text{g}/\text{m}^3$ in Africa. The highest η is found in arid regions where aerosols are less hygroscopic due to mineral dust, followed by regions strongly influenced by surface aerosol sources. Relatively low η is found over regions distant from strong aerosol sources. The spatial correlation of observed η with meteorological fields, aerosol vertical profiles, and aerosol chemical composition reveals that the spatial variation of η is strongly influenced by aerosol composition driven by its effects on aerosol hygroscopicity and aerosol vertical profile. Sensitivity tests with globally uniform parameters ~~reveal that aerosol composition leads to the strongest η spatial~~

29 ~~variability~~ quantify their effects on η spatial variability, with a population-weighted ~~normalized~~
30 mean difference of $12.3 \mu\text{g}/\text{m}^3$; ~~for aerosol composition that higher than that from aerosol vertical~~
31 ~~profile ($8.4 \mu\text{g}/\text{m}^3$)~~, reflecting ~~ings~~ the determinant composition effects on aerosol hygroscopicity and
32 aerosol optical properties; ~~and a population-weighted mean difference .-of $8.4 \mu\text{g}/\text{m}^3$ for aerosol~~
33 ~~vertical profile that reflects spatial variation in the column to surface relationship.~~

34 **1 Introduction**

35 Exposure to ambient fine particulate matter ($\text{PM}_{2.5}$) has been recognized as the predominant
36 environmental risk factor for the global burden of disease, leading to millions of deaths annually
37 (Brauer et al., 2024). Even at low $\text{PM}_{2.5}$ concentrations, long-term exposure can increase
38 circulatory and respiratory related mortality (Christidis et al., 2019; Pinault et al., 2016;
39 Weichenthal et al., 2022). Despite the importance of $\text{PM}_{2.5}$, many of the world's countries do not
40 provide publicly accessible $\text{PM}_{2.5}$ data (Martin et al., 2019). Satellite remote sensing of aerosol
41 optical depth (AOD), an optical measure of aerosol abundance, offers information about the
42 distribution of $\text{PM}_{2.5}$ (Kondragunta et al., 2022). A large community relies upon the spatial
43 distribution of $\text{PM}_{2.5}$ concentrations inferred from ~~satellite AOD and a modeled $\text{PM}_{2.5}$ to AOD~~
44 ~~relationship, satellite remote sensing~~ for health impact assessment and epidemiological analyses
45 of long-term exposure (Brauer et al., 2024; Burnett et al., 2018; Cohen et al., 2017; Hao et al.,
46 2023). Quantitative application of satellite AOD for long-term characterization of the spatial
47 distribution of $\text{PM}_{2.5}$ would benefit from a better understanding of the factors affecting the $\text{PM}_{2.5}$ -
48 AOD relationship.

49 The relationship between satellite AOD and surface $\text{PM}_{2.5}$ can be established through a statistical
50 method, a geophysical method, or their combination. A statistical method uses ground-based
51 monitors for training and is well suited for regions with dense monitors (Di et al., 2016; Hu et al.,
52 2014; Xin et al., 2014). A geophysical approach utilizes a chemical transport model to simulate
53 the relationship (η) between $\text{PM}_{2.5}$ and AOD for application to satellite AOD (van Donkelaar et
54 al., 2006, 2010; He et al., 2021), and thus depends on accurate model representation of η . ~~V~~van
55 Donkelaar et al. (2015, 2016) combined the two methods by applying geographically weighted
56 regression (GWR) on the geophysical $\text{PM}_{2.5}$, which further constrains geophysical $\text{PM}_{2.5}$ using
57 ground measurements and other predictors. However, accuracy of geophysical $\text{PM}_{2.5}$ remains

58 critical over vast areas with sparse monitoring, and knowledge about the factors affecting η spatial
59 variability are needed to guide improvements of modeled η and geophysical $PM_{2.5}$.

60 Previous studies have identified several factors that affect η variability, including aerosol vertical
61 distribution, aerosol hygroscopicity, aerosol optical properties, and ambient meteorological factors
62 such as relative humidity (RH), planetary boundary layer height (PBLH), wind speed, temperature,
63 and fire events (van Donkelaar et al., 2013; Ford and Heald, 2015; Guo et al., 2017; Jin et al., 2019;
64 Li et al., 2015; Wendt et al., 2023). Most studies focused on the temporal variability of η and found
65 association with meteorological variables such as PBLH (Chu et al., 2015; Damascena et al., 2021;
66 Gupta et al., 2006; He et al., 2021; Yang et al., 2019; Zhang et al., 2009). A few studies have
67 examined the regional-scale spatial variation of η with meteorological, land type variables, and
68 aerosol vertical profile in North America (van Donkelaar et al., 2006; Jin et al., 2020; Li et al.,
69 2015) and China (Yang et al., 2019). To our knowledge, none have examined the factors at the
70 global scale affecting the spatial variation of η or the effects of chemical composition.

71 In this work, we examine this knowledge gap about the spatial variation in η at a global scale. We
72 first collect data from more than 6,000 $PM_{2.5}$ monitoring sites provided by ~~nine-ten~~ networks and
73 satellite AOD to obtain an observationally based map of η . We further interpret the global η
74 distribution using the GEOS-Chem model of atmospheric composition with recent improvements
75 in aerosol size representation, $PM_{2.5}$ diel variation, and vertical allocation. By decomposing the
76 simulated η , we identify 2 strong drivers of η spatial variability: aerosol composition and aerosol
77 vertical profile. We conduct sensitivity tests using GEOS-Chem to study how the two factors vary
78 globally and how they contribute to the spatial variation in η .

79 **2 Methods**

80 **2.1 Ground Measured $PM_{2.5}$**

81 We collect ground-based measurements of $PM_{2.5}$ for the year 2019 ~~from which~~ to produce
82 observational constraints on η ($\frac{PM_{2.5}}{AOD}$), the spatially and temporally varying ratio between 24-hour
83 surface $PM_{2.5}$ concentrations and total column AOD at satellite sampling time. At the time of
84 manuscript preparation, the year 2019 offered the greatest density of measurements and the most
85 current emission inventory. We obtain $PM_{2.5}$ measurements from 7 regional networks and 32

86 global networks, as shown in Figure A1. For the United States, we access data from the United
87 States Environmental Protection Agency's Air Quality System ([https://www.epa.gov/outdoor-air-](https://www.epa.gov/outdoor-air-quality-data/download-daily-data)
88 [quality-data/download-daily-data](https://www.epa.gov/outdoor-air-quality-data/download-daily-data)), including both Federal Reference Method and non-Federal
89 Reference Methods PM_{2.5} (e.g. IMPROVE network). PM_{2.5} data for Canada are from the
90 Environment Canada's National Air Pollution Surveillance (NAPS) program. PM_{2.5} data for
91 Europe are from the European Environment Agency Air Quality e-Reporting system
92 (<https://www.eea.europa.eu/data-and-maps/data/aqereporting>). Over mainland China, PM_{2.5}
93 measurements ~~from the National and Provincial Environmental Protection Agencies~~ are
94 downloaded from <http://beijingair.sinaapp.com/>, ~~which provides instantaneous air quality data~~
95 ~~records from the National and Provincial Environmental Protection Agencies~~. Over India, PM_{2.5}
96 data are originally from the Central Pollution Control Board Continuous Ambient Air Quality
97 Monitoring network and the U.S. embassies. Over Australia, observations are downloaded for the
98 Northern Territory (<http://ntepa.webhop.net/NTEPA/>), Queensland
99 (<https://www.data.qld.gov.au/dataset/>), and New South Wales ([https://www.dpie.nsw.gov.au/air-](https://www.dpie.nsw.gov.au/air-quality/air-quality-data-services/data-download-facility)
100 [quality/air-quality-data-services/data-download-facility](https://www.dpie.nsw.gov.au/air-quality/air-quality-data-services/data-download-facility)). We require at least 5 days of
101 measurements for each month for a monitor to be included. Additionally, we obtain PM_{2.5}
102 measurements over other regions provided by the World Health Organization (WHO) Global
103 Ambient Air Quality Database ([https://www.who.int/data/gho/data/themes/air-pollution/who-air-](https://www.who.int/data/gho/data/themes/air-pollution/who-air-quality-database/2022)
104 [quality-database/2022](https://www.who.int/data/gho/data/themes/air-pollution/who-air-quality-database/2022)), OpenAQ (<https://openaq.org/>), and ~~by~~ the Surface PARTiculate mAtter
105 Network ~~Network~~ (SPARTAN, <https://www.spartan-network.org/>), which is co-located with the
106 Aerosol Robotic Network (AERONET). SPARTAN also provides filter based PM_{2.5} chemical
107 composition, which is initially described in Snider et al., (2016). Subsequent developments to the
108 sampling and analysis procedure of SPARTAN include an upgrade to the AirPhoton SS5 sampling
109 station to use a cyclone inlet, an automated weighing system (MTL AH500E) to improve precision
110 and throughput, additional black carbon analysis by Hybrid Integrating Plate/Sphere (White et al.,
111 2016), ~~trace metal~~ elements measured by X-ray Fluorescence (Liu et al., 2024) and a global mineral
112 dust equation (Liu et al., 2022). We require at least 50 days of coincident PM_{2.5} and AERONET
113 AOD measurements for a SPARTAN site to be included in our analysis.

114 We also collected publicly available PM_{2.5} compositional data to assess GCHP simulated
115 composition. Long-term PM_{2.5} compositional data are included from the United States

116 [Environmental Protection Agency’s Air Quality System, the European Environment Agency Air](#)
117 [Quality e-Reporting system, and SPARTAN, with a total of 365 sites covering the U.S. \(306\),](#)
118 [Europe \(37\), and the Global South \(22\).](#)

119 **2.2 Satellite AOD**

120 We obtain AOD at 550 nm from the Multi-Angle Implementation of Atmospheric Correction
121 (MAIAC) algorithm, which offers AOD at a high spatial resolution of 1 km worldwide over both
122 land and coastal regions (Lyapustin et al., 2018). The radiances used in the retrieval are measured
123 by the twin MODerate resolution Imaging Spectroradiometer (MODIS) instruments onboard the
124 Terra and Aqua satellites. Terra follows a descending orbital path, crossing the equator at 10:30
125 local time, while Aqua is on an ascending orbit with 13:30 equatorial crossing local time. Both
126 MODIS instruments offer a wide swath width of 2330 km, enabling nearly global daily coverage
127 of the Earth (Sayer et al., 2014). PM_{2.5} monitoring sites with annual mean satellite AOD less than
128 0.05 (background AOD level over land) are excluded to reduce the influence of retrieval
129 uncertainties on our analysis.

130 **2.3 AERONET AOD**

131 AERONET is a worldwide sun photometer network that provides long-term measurement of AOD.
132 We use the Version 3 Level 2 database, which includes an improved cloud screening algorithm
133 (Giles et al., 2019). We sample AERONET AOD within ± 15 min of the satellite overpass time and
134 interpolate to 550 nm wavelength, based on the local Ångström exponent at 440 and 670 nm. For
135 SPARTAN sites, we sample AERONET data coincidentally with SPARTAN aerosol composition
136 to obtain the ground-based observation of η .

137 **2.4 GEOS-Chem Simulation**

138 We simulate η with the GEOS-Chem chemical transport model (www.geos-chem.org, last access:
139 26 October 2023), driven by offline meteorological data, MERRA-2, from the Goddard Earth
140 Observing System (GEOS) of the NASA Global Modeling and Assimilation Office (Schubert et
141 al., 1993). We use the high-performance configuration of GEOS-Chem (GCHP) (Eastham et al.,
142 2018) version 13.4.0 (DOI: 10.5281/zenodo.7254268), which includes advances in performance
143 and usability (Martin et al., 2022). The simulation is conducted for the year 2019, on a C90 cubed-

144 sphere grid corresponding to a horizontal resolution of about 100 km, with a spin-up time of 1
145 month.

146 The GEOS-Chem aerosol simulation includes the sulfate-nitrate-ammonium (SNA) system
147 (Fountoukis and Nenes, 2007), primary and secondary carbonaceous aerosols (Pai et al., 2020;
148 Park et al., 2003; Wang et al., 2014), sea salt (Jaeglé et al., 2011), and natural (Fairlie et al., 2007;
149 Meng et al., 2021) and anthropogenic (Philip et al., 2017) dust. Emissions are processed with the
150 Harmonized Emissions Component (HEMCO) (Lin et al., 2021). The primary emission data are
151 from the Community Emissions Data System version 2 (CEDS~~EBD-MAPSv2~~; (Hoesly et al., 2018;
152 CEDS, 2024) McDuffie et al., 2020) for the year 2019. Emissions from stacks are distributed
153 vertically (Bieser et al., 2011). Diel variation of anthropogenic emissions is included (Li et al.,
154 2023). Resolution-dependent soil NO_x, sea salt, biogenic VOC, and natural dust emissions are
155 calculated offline at native meteorological resolution to produce consistent emissions across
156 resolutions (Meng et al., 2021; Weng et al., 2020). Biomass burning emissions use the Global Fire
157 Emissions Database, version 4 (GFED4) at daily resolution (van der Werf et al., 2017) for the year
158 2019. We estimate organic matter (OM) from primary organic carbon using an OM/OC
159 parameterization (Canagaratna et al., 2015; Philip et al., 2014b). For secondary aerosol
160 components, the concentration at 2 m above the surface is used to calculate PM_{2.5}, following Li et
161 al. (2023). A 50% reduction of the surface nitrate concentration is applied to account for the long-
162 persisting-standing bias in surface nitrate simulated by GEOS-Chem (Heald et al., 2012; Miao et
163 al., 2020; Travis et al., 2022; Zhai et al., 2021; Zhang et al., 2012); also Figure A22 in this
164 manuscript) and other models- such as CMAQ (Shimadera et al., 2014), WRF-Chem (Sha et al.,
165 2019), and EMEP/MSC-W (Prank et al., 2016) (Zakoura and Pandis, 2018; Shimadera et al., 2014).
166 Despite this bias, GEOS-Chem can sufficiently represent the variability of nitrate for applications
167 to studies at global (McDuffie et al., 2021; Weagle et al., 2018) and regional (Geng et al., 2017;
168 Kim et al., 2015; Philip et al., 2014a; Zhai et al., 2021) scales. Dry and wet deposition follows
169 Amos et al. (2012), with a standard resistance-in-series dry deposition scheme (Wang et al., 1998).
170 Wet deposition includes scavenging processes from convection and large-scale precipitation (Liu
171 et al., 2001).

172 Global RH-dependent aerosol optical properties are based on the Global Aerosol Data Set (GADS)
173 (Kopke et al., 1997), as originally implemented by Martin et al. (2003), with updates for SNA and

174 OM dry size (Zhu et al., 2023), hygroscopicity (Latimer and Martin, 2019), mineral dust size
 175 distribution (Zhang et al., 2013), and absorbing brown carbon (Hammer et al., 2016). These
 176 updates enable GEOS-Chem to capture 74% of the AOD spatial variability versus AERONET
 177 (Zhu et al., 2023). A slight systematic low bias against MAIAC AOD is found, with an intercept
 178 of -0.05 and a population-weighted mean difference (PWMD) of -0.04. Low bias in simulated
 179 AOD is also reported for other models, such as CMAQ (Jin et al., 2019) and WRF-Chem
 180 (Benavente et al., 2023). We artificially increase simulated AOD by 0.04 globally to address this
 181 poorly understood systematic bias that, although minor, is useful for the representation of η
 182 (PWMD reduced from 20.6 $\mu\text{g}/\text{m}^3$ to 1.9 $\mu\text{g}/\text{m}^3$). $\text{PM}_{2.5}$ is calculated as the sum of each component
 183 at 35% RH to align with common measurement protocols. ~~$\text{PM}_{2.5}$ is calculated as the sum of each~~
 184 ~~component at 35% RH to align with common measurement protocols.~~

185 2.5 Population

186 Global population information is obtained from the Gridded Population of the World provided by
 187 the NASA Socioeconomic Data and Applications Center (Center for International Earth Science
 188 Information Network—~~CIESIN~~, 2018).

189 2.6 Sensitivity Tests with Globally Uniform Parameters

190 We conduct sensitivity tests of factors affecting the spatial variation of η , with a focus on aerosol
 191 composition and aerosol vertical profile. To understand the relative importance of these factors,
 192 we impose a constant for each factor and simulate the corresponding η . The difference between
 193 the test scenario and the base scenario reflects the change due to variation of the factor. We use
 194 the global population-weighted mean (PWM) and population-weighted mean difference (PWMD)
 195 to summarize changes with a focus on relevance to population exposure:

$$X_{PWM} = \frac{\sum_j \sum_i P_{i,j} X_{i,j}}{\sum_j \sum_i P_{i,j}}$$

196

$$PWMD = \frac{\sum_j \sum_i P_{i,j} |X_{i,j} - Y_{i,j}|}{\sum_j \sum_i P_{i,j}}$$

197 where i and j are grid box identifiers. X and Y could be any variable of interest. $|X_{i,j} - Y_{i,j}|$ is the
 198 absolute value of their difference. P represents population density in each grid box.

199 The first test imposes globally uniform aerosol chemical composition calculated as the global
200 PWM aerosol component fraction ($F_{s,k,s,PWM}$):

$$F_{s,k,s,PWM} = \frac{\sum_j \sum_i P_{i,j} F_{i,j,s,k,s}}{\sum_j \sum_i P_{i,j}}$$

201 where i , j , and k are grid box identifiers along latitude, longitude, and vertical layer. ~~P represents~~
202 ~~population density in each grid box. F_s P represents population density in each grid box. F_s is the~~
203 fraction of aerosol component S_s in total aerosol mass. This test keeps the total columnar aerosol
204 mass and aerosol vertical profile unchanged.

205 The second test imposes a globally uniform aerosol vertical profile calculated as the PWM column
206 relative vertical profile ($R_{s,k,s,PWM}$):

$$R_{s,k,s,PWM} = \frac{\sum_j \sum_i P_{i,j} R_{i,j,k,s}}{\sum_j \sum_i P_{i,j}}$$

207 where $R_{i,j,k,s}$ is the relative dry mass ratio compared to the surface. The total mass loading and
208 relative chemical composition are unchanged.

209 We analyze global and regional variations of η , as well as that for the driving factors. The definition
210 of each region used in this study is summarized in Figure A3.

211 3 Results and Discussion

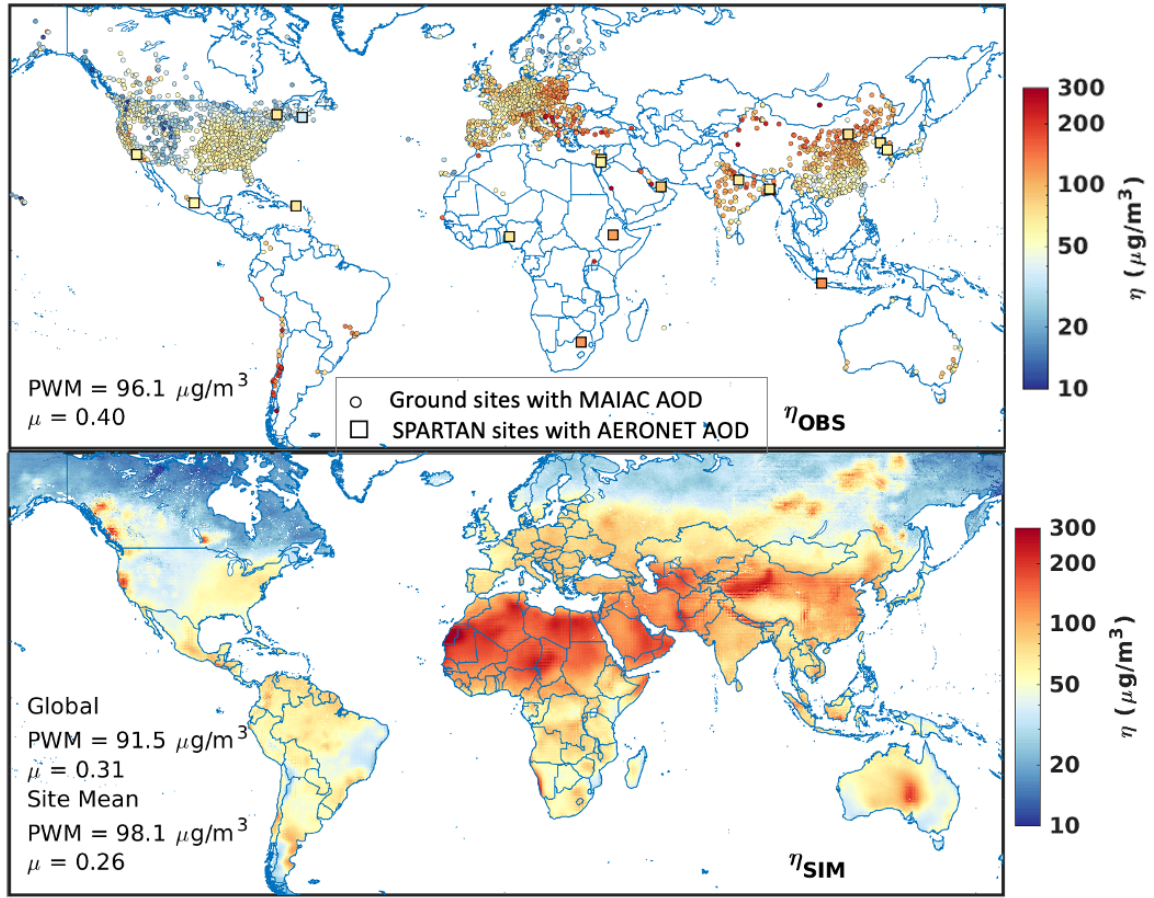
212 3.1 Global Spatial Pattern of η

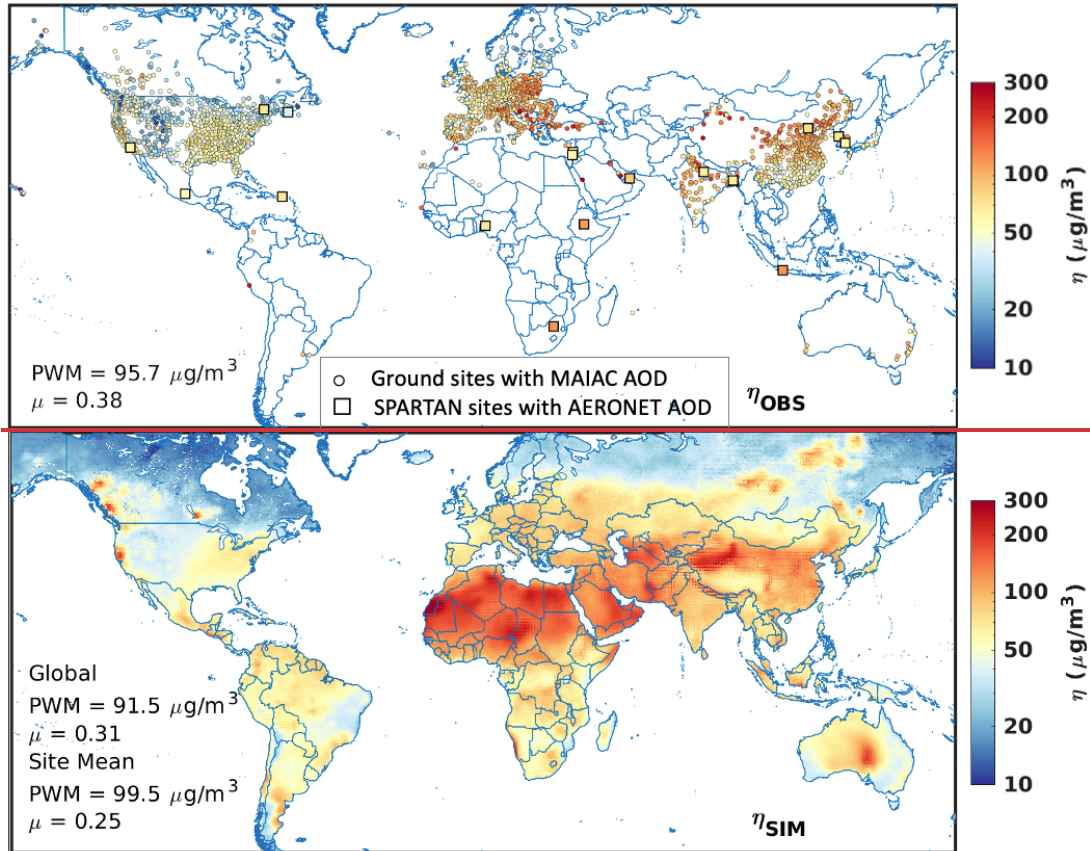
213 The top panel of ~~Figure 1~~ ~~Figure 1~~ shows the observationally based annual mean η , inferred from
214 the ratio of ground-measured $PM_{2.5}$ to MAIAC AOD. Measurements are most dense in North
215 America, Europe, and East Asia. The annual mean η varies substantially, from $7.8 \mu\text{g}/\text{m}^3$ in Hawaii
216 to ~~321-504~~ $\mu\text{g}/\text{m}^3$ in ~~Central Asia~~ ~~Mongolia~~, with a PWM of ~~95.796.1~~ $\mu\text{g}/\text{m}^3$ ~~and standard deviation~~
217 ~~(σ) of $36.6 \mu\text{g}/\text{m}^3$. Higher PWM η of ~~132-196~~ $\mu\text{g}/\text{m}^3$ to ~~154-154~~ $\mu\text{g}/\text{m}^3$ exist over desert regions
218 such as Africa and West Asia, followed by PWM η of $97 \mu\text{g}/\text{m}^3$ to ~~121-119~~ $\mu\text{g}/\text{m}^3$ ~~by over~~
219 strongly influenced by anthropogenic aerosols, such as East Asia ~~and~~ South Asia (Figure A4 and
220 Table A1). Over North America, η is around $60 \mu\text{g}/\text{m}^3$ in the east and in California, which is more
221 than double that in the Rockies, driven by the spatial pattern of surface $PM_{2.5}$ (Figure A4). The
222 PWM η in North America of ~~60.359.8~~ $\mu\text{g}/\text{m}^3$ is about 30% lower than the global PWM. The η~~

223 pattern found here is similar to that reported by Jin et al. (2020) for the U.S. In Europe, η also
224 varies noticeably between the east and the west, driven by the spatial pattern of surface $\text{PM}_{2.5}$, as
225 $\text{PM}_{2.5}$ increases by 60% from west to east while AOD increases by only 8%. The PWM η in Europe
226 is ~~94.0~~2.3 $\mu\text{g}/\text{m}^3$, slightly lower than the global PWM. In Asia, measured η is concentrated in
227 China and India. In China, the η spatial pattern shows a clear distinction between the northern and
228 southern regions, driven by the higher AOD in the south ([Figure A5](#)), where relative humidity is
229 high. A similar η spatial pattern and a negative correlation between η and RH are reported by Yang
230 et al. (2019). In India, η is highest in the northwest, with a PWM η of ~~128~~129 $\mu\text{g}/\text{m}^3$, and decreases
231 to about 80 $\mu\text{g}/\text{m}^3$ toward the east and the south. Both $\text{PM}_{2.5}$ and AOD follow the same spatial
232 pattern, while $\text{PM}_{2.5}$ exhibits a stronger decreasing tendency ([Figure A4](#) and [Figure A5](#)). PWM η
233 in Asia is ~~100~~102 $\mu\text{g}/\text{m}^3$, the highest among the populous regions and ~~4.56~~0% higher than the
234 global PWM. Globally, from west to east, η increases by about ~~627~~0%, despite that both $\text{PM}_{2.5}$ and
235 AOD increased more than threefold ([Figure A6](#)). The coefficient of variation (standard deviation
236 divided by mean) in η is higher in Europe ($\mu = 0.31$) and Asia ($\mu = 0.$ ~~343~~6), than North America
237 ($\mu = 0.$ ~~253~~5, [Figure A6](#)).

238 The bottom panel in [Figure 1](#)~~Figure 1~~ shows the GCHP simulated η , the ratio between simulated
239 24-hour mean surface $\text{PM}_{2.5}$ and simulated total column AOD at satellite overpass time. The
240 simulation generally reproduces the global observations of η with a tendency for high values in
241 arid regions influenced by dust and low values in regions distant from strong surface sources. The
242 ~~simulated~~ global ~~simulated~~-PWM η is ~~24~~% higher than the observations (~~99.598.1~~ $\mu\text{g}/\text{m}^3$ vs.
243 ~~95.796.1~~ $\mu\text{g}/\text{m}^3$), mostly driven by an overestimation in [East](#) Asia (~~1087~~ $\mu\text{g}/\text{m}^3$ vs. ~~400~~96.9 $\mu\text{g}/\text{m}^3$),
244 that reflects an overestimation of PWM $\text{PM}_{2.5}$ (~~47.03.3~~ $\mu\text{g}/\text{m}^3$ vs. ~~43.638.0~~ $\mu\text{g}/\text{m}^3$). The simulation
245 generally reproduces the regional spatial pattern in North America and Asia but underestimates
246 the η variability in Europe as it overestimates η in central Europe and underestimates η in Eastern
247 Europe, due to positive bias in simulated $\text{PM}_{2.5}$ in central Europe and positive bias in simulated
248 AOD in Eastern Europe. Nonetheless, the PWM η in Europe (~~84.183.6~~ $\mu\text{g}/\text{m}^3$) is within ~~119.4~~%
249 of observations. Globally, there is overall consistency between the simulated η and observed η ,
250 with a correlation of ~~0.6459~~, resulting in a high degree of consistency between geophysical $\text{PM}_{2.5}$
251 and measured $\text{PM}_{2.5}$ ($r = 0.$ ~~9089~~9, [Figure A6](#)). [Evaluation of the simulation of \$\text{PM}_{2.5}\$ chemical](#)

252 composition versus ground-based measurements reveals a high degree of consistency (Figure A2)
253 that supports their further assessment of the factors affecting η .





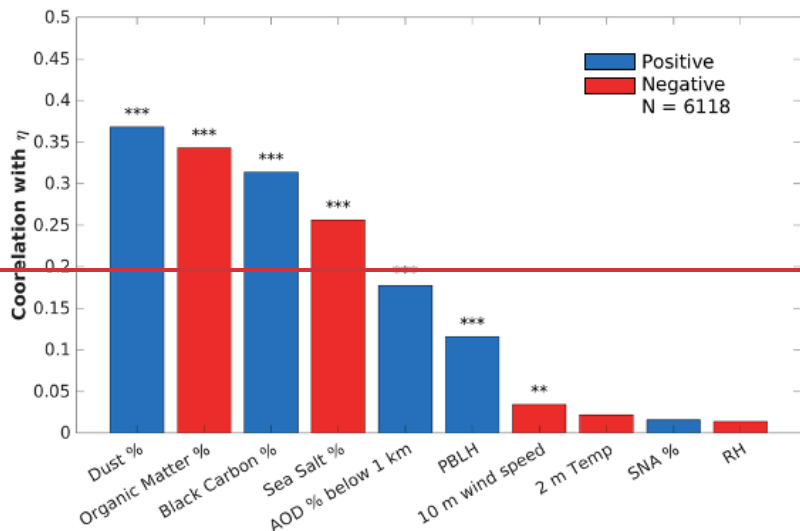
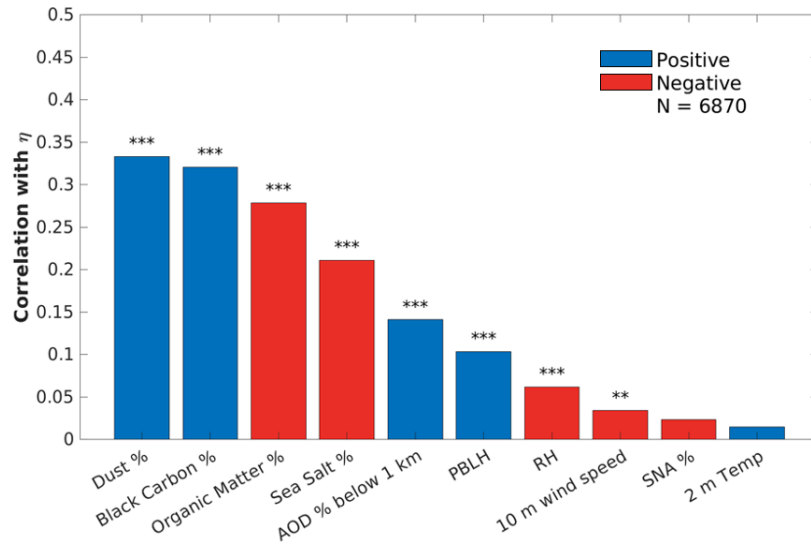
255

256 Figure 1. Observed (top) and simulated (bottom) annual mean η for 2019. Circles represent ground
 257 measurement sites from regional networks or the World Health Organization. Squares represent co-
 258 located ground measured $\text{PM}_{2.5}$ from SPARTAN and AOD from AERONET. PWM = population-
 259 weighted mean, μ = coefficient of variation (standard deviation divided by mean).

260 We explore the dominant driving factors for η spatial variation by calculating the spatial
 261 correlation between each candidate factor and the observation-based η . Candidate factors
 262 examined include meteorological fields (MERRA-2), aerosol vertical profile, and aerosol
 263 composition as collected from the GCHP simulation or SPARTAN. Meteorological fields include
 264 those commonly considered to represent the temporal variation in η , such as PBLH, RH at 700
 265 hPa, wind speed at 10 m, and temperature at 2 m (Chu et al., 2015; Damascena et al., 2021; He et
 266 al., 2021; Yang et al., 2019). The aerosol vertical profile is represented as the AOD fraction below
 267 1 km (AOD % below 1 km). Aerosol composition includes SNA, OM, dust, black carbon, and sea
 268 salt, all represented as the fractional contributions (%) to surface $\text{PM}_{2.5}$. Figure 2 shows the spatial
 269 correlation of annual mean factors versus observation-based η . Aerosol components, particularly

270 those with strong primary sources (dust, OM, and black carbon), exhibit the strongest correlations
271 (>0.327) with observationally based η . Significant positive correlations are found for mineral dust
272 and black carbon, both of which are non- or weakly-hygroscopic. Significant negative correlations
273 are found for organic matter and sea salt, reflecting a weak connection between surface
274 concentrations and AOD aloft. Processes are further discussed in sections 3.2 and 3.4. The aerosol
275 vertical profile exhibits a moderate correlation with η (0.4814), which is notably higher than any
276 meteorological factors ($<0.12 \leq 0.10$). Ground-based data from SPARTAN and AERONET
277 corroborate the correlation between aerosol composition and η (Figure A7). We thus focus further
278 analysis in Sections 3.2-3.4 on the two main drivers in η : aerosol composition and aerosol vertical
279 profile.

280 The ~~indicators~~ drivers of spatial variation in η found here differ from that for temporal variation
281 of η in prior work (e.g. He et al. 2021), reflecting the different processes involved. Meteorological
282 parameters drive short-term variability in the aerosol vertical profile, such as day-to-day variation
283 in mixed layer depth or in advection from a point source. In contrast, the spatial variation in annual
284 mean η reflects the spatial variation in processes affecting the long-term relation of surface $PM_{2.5}$
285 at controlled RH of 35% with AOD at ambient RH. Aerosol composition and the aerosol vertical
286 profile reflect spatial variation in aerosol hygroscopicity, mass extinction efficiency, and sources.
287 The following sections explore how aerosol composition and aerosol vertical profile vary globally
288 and examine how they affect the spatial pattern of η by conducting two sensitivity tests. In each
289 sensitivity test, we replace the spatial variability of a factor with a globally uniform value. The
290 variability of aerosol composition and aerosol vertical profile are discussed in sections 3.2 and 3.3,
291 respectively. The sensitivity test results are discussed in section 3.4.



292

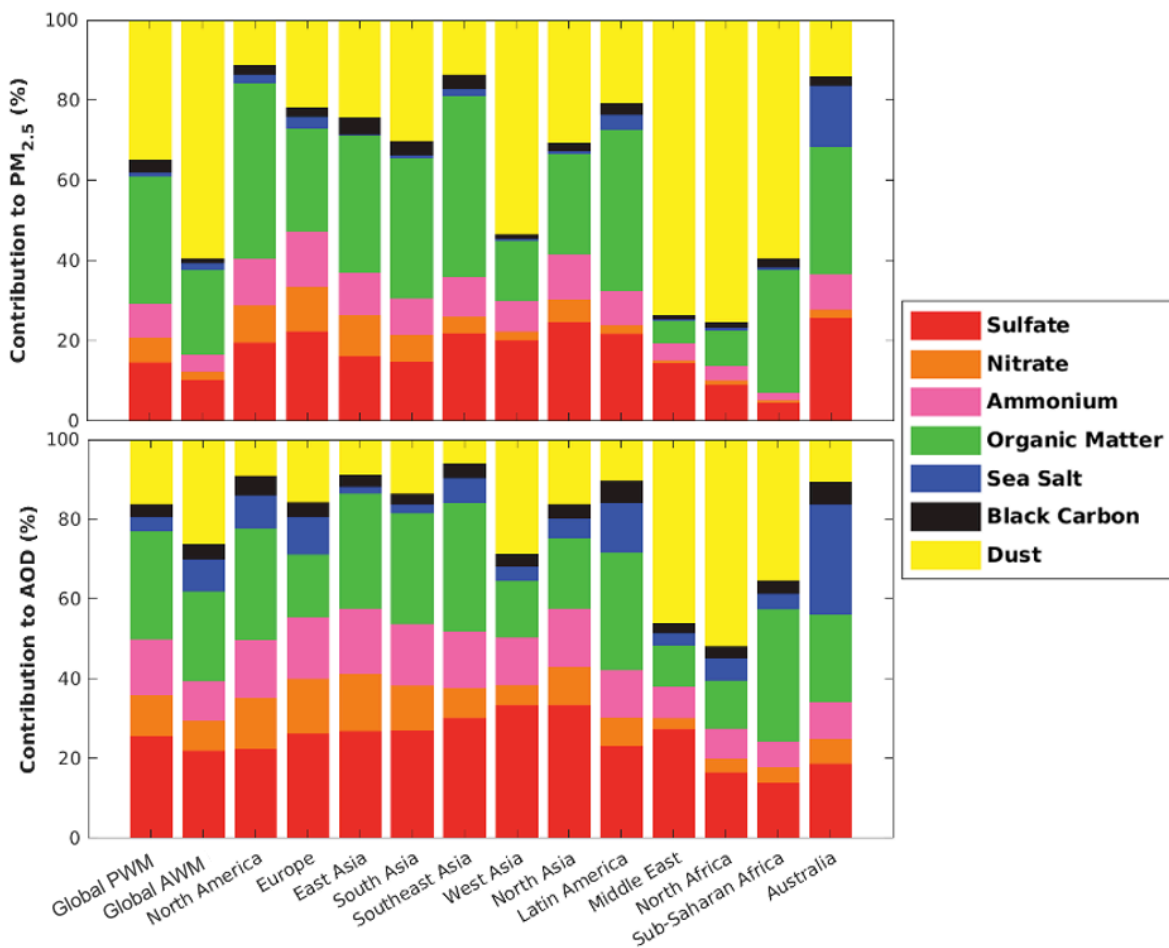
293

294 Figure 2. Spatial correlation between annual mean modeled parameters and observationally-based η . Blue
 295 bars indicate positive correlations. Red bars indicate negative correlations. Stars above each bar indicate
 296 the p-value associated with each correlation. ‘***’ indicates the p-value is lower than 0.001 and ‘**’
 297 indicates lower than 0.01.

298 3.2 Spatial Variability in Aerosol Composition

299 ~~Figure 3~~ ~~Figure 3~~ shows the simulated PWM aerosol composition globally and regionally, as well
 300 as the global area-weighted mean (AWM). The top panel shows the compositional contribution to
 301 PM_{2.5}. Globally, dust is the leading PWM PM_{2.5} component (34.7%), followed by OM (31.9%)
 302 and SNA (29.3%). The bottom panel shows the compositional contribution to AOD. PWM AOD

303 composition is more evenly distributed, with more contribution from SNA (49.9%), followed by
 304 OM (27.2%) and dust (16.1%). Overall, more hygroscopic aerosols such as SNA tend to contribute
 305 a larger fraction of AOD which is at ambient RH, while less hygroscopic aerosols, such as mineral
 306 dust tend to contribute a larger fraction of PM_{2.5} which is at controlled RH of 35%. The AWM
 307 PM_{2.5} and AOD composition exhibit weaker contributions from SNA, primarily reflecting a larger
 308 contribution from dust in remote regions than in more densely populated areas. Over populous
 309 regions such as North America, Europe, and Southeast Asia, there are greater SNA and OM
 310 fractions than the global mean ([Figure 3](#)). Arid regions, such as West Asia, the Middle
 311 East, North Africa, and Sub-Saharan Africa, have large fractions of non-hygroscopic mineral dust
 312 that (1) reduce aerosol mass extinction efficiency, yielding less AOD per unit mass, and (2) are
 313 unaffected by the controlled RH of PM_{2.5}. Both of these factors increase η in dusty regions
 314 compared with regions dominated by hygroscopic SNA aerosols.

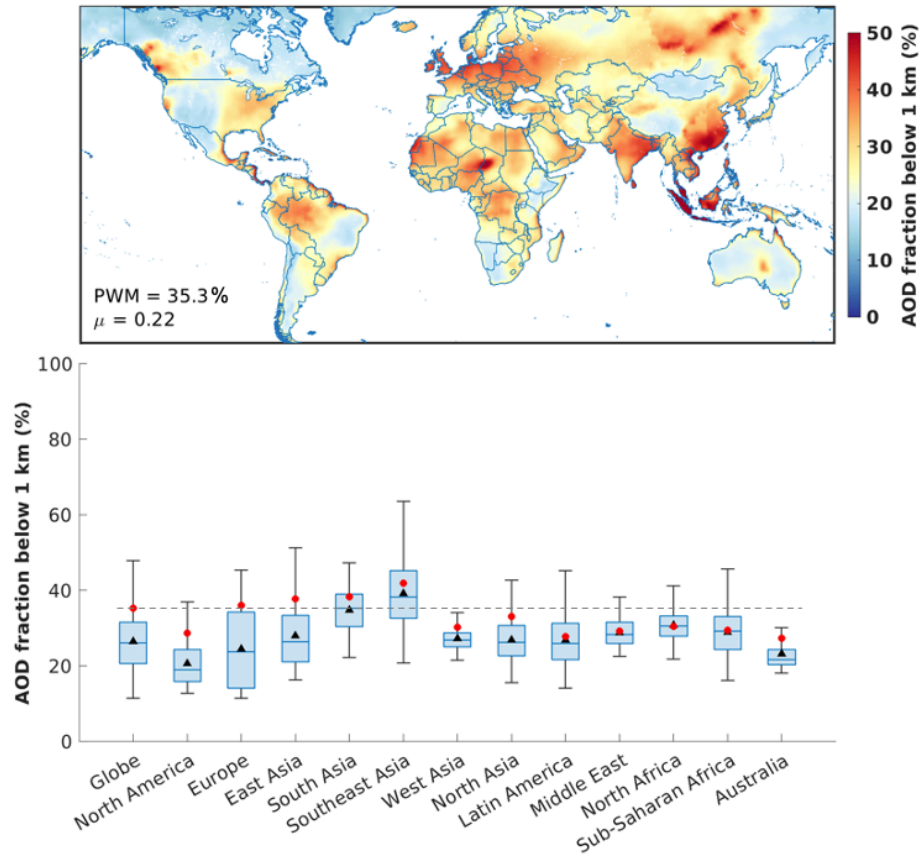


315

316 Figure 3. Global and regional PWM contributions of aerosol composition to surface PM_{2.5} (top) and AOD
317 (bottom). The global area-weighted mean (AWM) over land is also included as the second column.

318 **3.3 Spatial Variability in Aerosol Vertical Profile**

319 Figure 4 shows the AOD fraction below 1 km in the GEOS-Chem simulation. Globally, 35.3% of
320 the PWM AOD is below 1 km. The PWM value is greater than the AWM value since populated
321 areas tend to have more surface emissions of particles and precursors. Over North America, Europe,
322 and East Asia, the PWM surface AOD fractions are much higher than the medians and AWM,
323 indicating high spatial heterogeneity between urban and remote areas. Europe exhibits the highest
324 variation and the largest discrepancy between PWM and AWM, reflecting the largest spatial
325 heterogeneity in aerosol vertical profile, driven by influences from regional pollution, marine
326 aerosols, and transported dust (Zhao et al., 2018). Southeast Asia has the highest surface AOD
327 fraction and a large variation. Local sources, long-range transported dust, and the influence of
328 trade winds all contribute to the unique spatial variation in aerosol vertical profile in this region
329 (Banerjee et al., 2021; Nguyen et al., 2019). Globally, PWM values exhibit less variation than
330 AWM, indicating moderate variation in aerosol profile across populous areas.



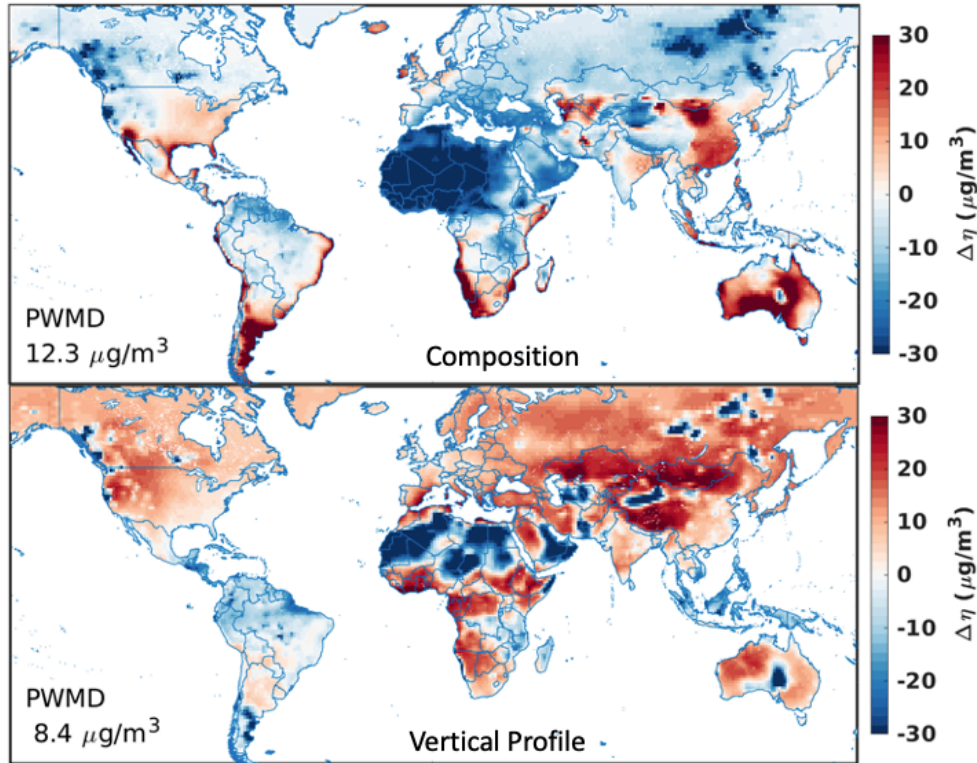
331
 332 Figure 4. (Top) Map of AOD fraction below 1 km. (Bottom) Global and regional statistics for AOD
 333 fraction below 1 km. Black triangles show the area-weighted mean. Red circles show the PWM. The line
 334 inside each box is the sample median. Each box's top and bottom edges are the 75 and 25 quartiles,
 335 respectively. Vertical bars are the maximum and minimum values within 1.5 times the interquartile range.
 336 The dashed line indicates global PWM.

337 3.4 Sensitivity Tests with Globally Uniform Parameters

338 Figure 5 shows the global changes in the spatial variation in η due to variations in aerosol
 339 chemical composition (top) and aerosol vertical profile (bottom), the two main drivers found in
 340 Figure 2. Globally, neglect of spatial variation in aerosol composition induces a $12.3 \mu\text{g}/\text{m}^3$
 341 PWMD in η spatial variation. Both $\text{PM}_{2.5}$ and AOD are strongly affected by aerosol composition,
 342 following a similar spatial pattern (Figure A8). Over mid- and low-latitude areas, the change in
 343 AOD is stronger than in $\text{PM}_{2.5}$, since AOD at ambient RH is more sensitive to hygroscopicity
 344 changes. which gives This yields the opposite pattern in the η . Neglect of spatial variation in
 345 chemical composition reduces η over North Africa and the Middle East, desert regions where

346 aerosols contain more weakly hygroscopic components such as mineral dust, compared to
347 populous areas, which contain more secondary inorganic aerosol ([Figure 3](#)~~Figure 3~~). For smaller
348 deserts in the Southwest U.S., Argentina, and Southwest Africa, the dust fractions of surface
349 aerosols are higher than the global mean (36%, 76%, and 49%, respectively), but the dust fraction
350 for AOD is similar to the global mean (15%, 25%, and 14%, respectively). Therefore, neglect of
351 the spatial variation of chemical composition increases η over these small deserts by increasing
352 the fraction of hygroscopic components in $PM_{2.5}$ and leaving AOD almost unchanged (Figure A8).
353 ~~†Neglect of spatial variation in chemical composition~~ also reduces η over the boreal forests, ~~and~~
354 ~~the Amazon~~, where surface aerosols ~~contain little dust and~~ are more hygroscopic compared to
355 populous areas ~~and show strong changes, while less so for column aerosol (Figure A8). (Figure 3).~~
356 Neglect of spatial variation in chemical composition increases η over the eastern U.S. and eastern
357 China, where $PM_{2.5}$ contains more hygroscopic SNA and less dust than the global mean. It also
358 increases η in coastal regions where aerosol contains more hygroscopic sea salt than the global
359 mean.

360 Neglect of spatial variation in the aerosol vertical profile induces an $8.4 \mu\text{g}/\text{m}^3$ PWMD in η spatial
361 variation ([Figure 5](#)~~Figure 5~~), following the spatial pattern of the change in surface $PM_{2.5}$ (Figure
362 A9). The most apparent feature is an increase in η throughout the remote northern hemisphere,
363 driven by an increased aerosol fraction near the surface where the fraction is normally small
364 (Figure 4). The uniform aerosol vertical profile decreases η over northern Africa and biomass
365 burning regions of the boreal forests, the Amazon, and Indonesia, driven by a decreased aerosol
366 fraction near the surface in regions where that fraction is normally high.



367

368 Figure 5. Changes in η (test -base) for each sensitivity test. In the first test, a global PWM aerosol
 369 composition replaces the actual composition (top). In the second test, a global PWM aerosol profile
 370 replaces the actual profiles (bottom). Numbers inset indicate population-weighted mean difference
 371 (PWMD).

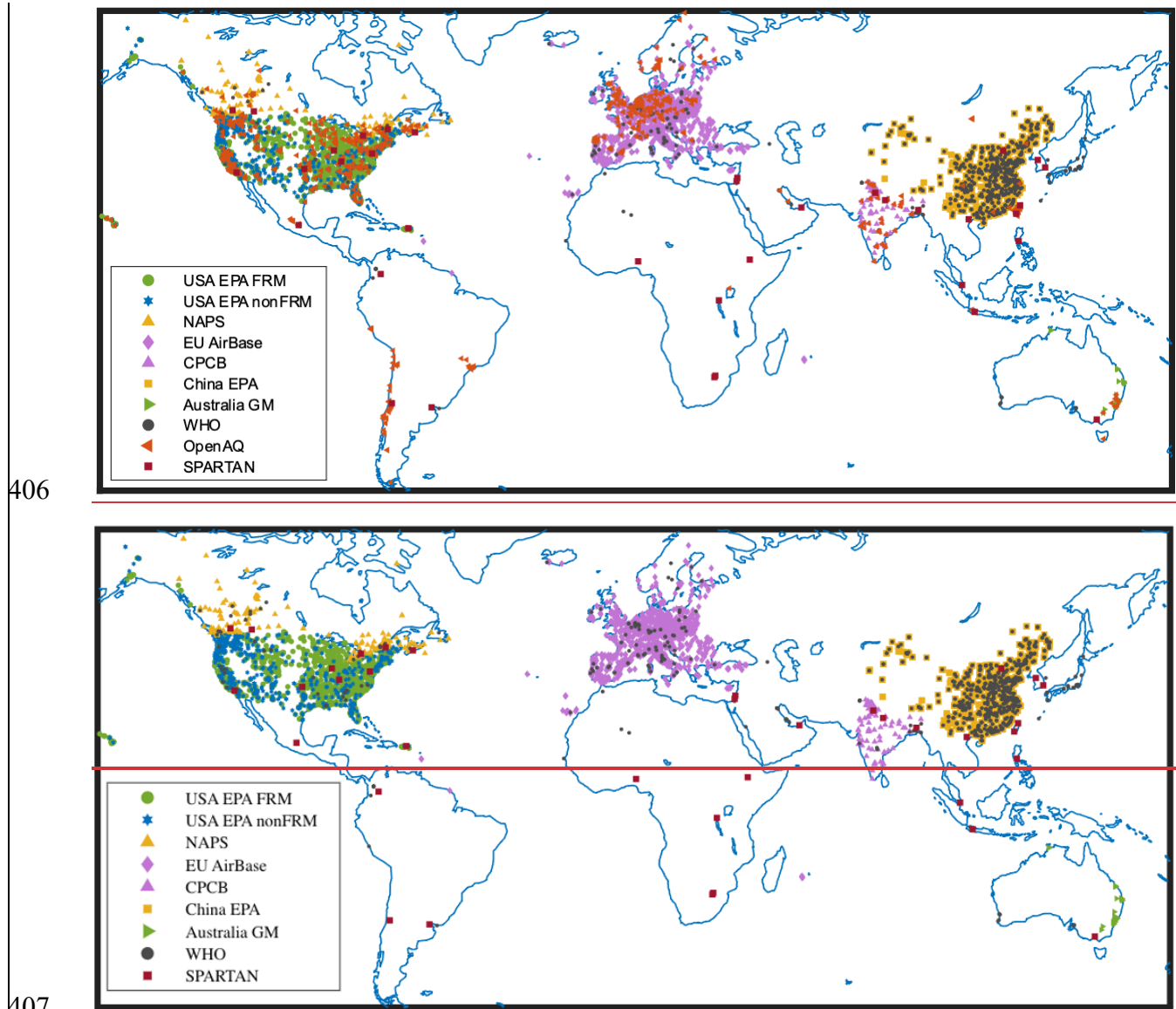
372 Conclusion

373 Understanding the global variation of the $PM_{2.5}$ and AOD relationship (η) offers insight into the
 374 geophysical inference of $PM_{2.5}$ from satellite AOD observations. We collected ground-based $PM_{2.5}$
 375 measurements from [6188-6,870](#) sites and MODIS MAIAC satellite AOD throughout the year 2019
 376 to obtain, for the first time, a global scale observationally based η map. Observed annual mean η
 377 ranges from 7.8 $\mu\text{g}/\text{m}^3$ in Hawaii to [321-504](#) $\mu\text{g}/\text{m}^3$ in [Central AsiaMongolia](#). We observed
 378 enhanced η of [132-196](#) $\mu\text{g}/\text{m}^3$ to 154 $\mu\text{g}/\text{m}^3$ over arid regions such as Africa and West Asia, due to
 379 their low aerosol extinction efficiency. Moderate η of 97 $\mu\text{g}/\text{m}^3$ to [121-119](#) $\mu\text{g}/\text{m}^3$ was found in
 380 industrial areas such as East Asia and South Asia, where anthropogenetic emissions increase the
 381 near-surface $PM_{2.5}$ concentrations. Over remote areas, low η ($< 50 \mu\text{g}/\text{m}^3$) was usually observed.

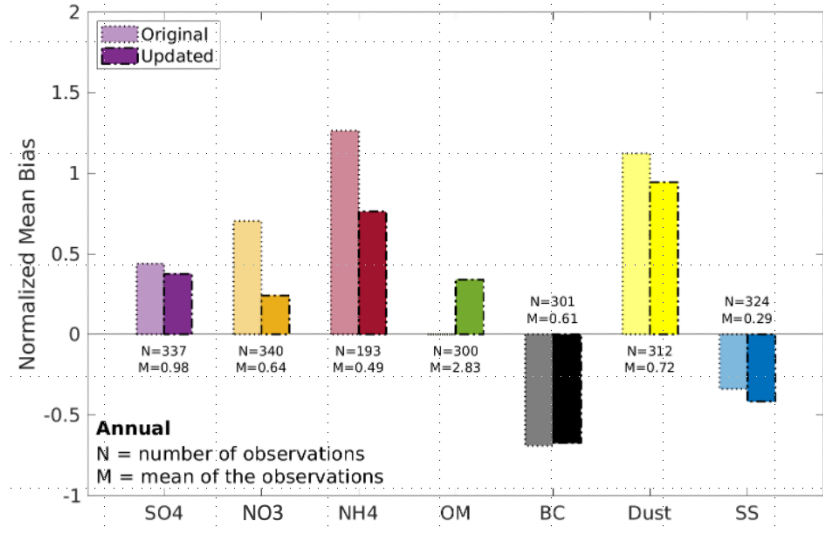
382 We simulated the global annual mean η with the GEOS-Chem chemical transport model in its high
383 performance configuration (GCHP). The simulation generally represented observed η with PWM
384 within 34% (99.598.1 $\mu\text{g}/\text{m}^3$ vs 95.796.1 $\mu\text{g}/\text{m}^3$) and, a correlation of 0.5964 over the 6,118-780
385 measurement sites, ~~and a slope of 0.81~~. We examined the correlation between simulation and
386 measurements to identify the two most impactful drivers for η spatial variation - aerosol
387 composition and aerosol vertical profile, both of which strongly affect the annual mean relation of
388 columnar AOD at ambient RH with surface $\text{PM}_{2.5}$ at controlled RH of 35%. -We subsequently
389 conducted sensitivity tests by eliminating the spatial variation of each of the two drivers and
390 quantified the impact on η spatial variability. Imposing a globally uniform aerosol composition led
391 to ~~more~~ pronounced changes ($\text{PWMD} = 12.3 \mu\text{g}/\text{m}^3$), reflecting how changes in aerosol
392 composition affect both AOD and surface $\text{PM}_{2.5}$, due to the effects of aerosol hygroscopicity on
393 both quantities. Imposing a globally uniform aerosol vertical profile had a moderate effect (PWMD
394 $= 8.4 \mu\text{g}/\text{m}^3$), reflecting changes in the fraction of aerosol near the surface.

395 These findings motivate additional efforts to develop the simulation of aerosol composition and
396 aerosol vertical profile. Promising avenues include: (1) enhancing global long-term measurements
397 of $\text{PM}_{2.5}$ chemical composition to evaluate and improve simulations, (2) exploiting new and
398 emerging information about aerosol type from satellite remote sensing (e.g. PACE, MAIA), (3)
399 advancing simulations at finer spatial resolution to better represent processes affecting aerosol
400 composition and vertical profile, (4) leveraging aircraft, lidar, and collected AOD-to- $\text{PM}_{2.5}$
401 measurements for constraints on the vertical profile, and (5) exploiting nascent capabilities in
402 applying satellite remote sensing (e.g. TROPOMI, TEMPO, GEMS) for top-down constraints on
403 emissions that affect aerosol composition.

404



408 Figure A1. PM_{2.5} measurement sites from publicly available networks.



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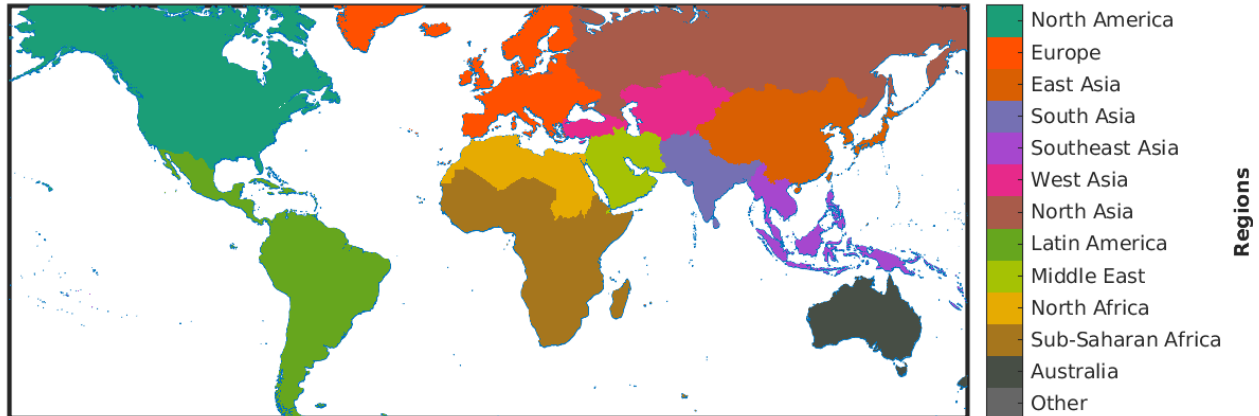
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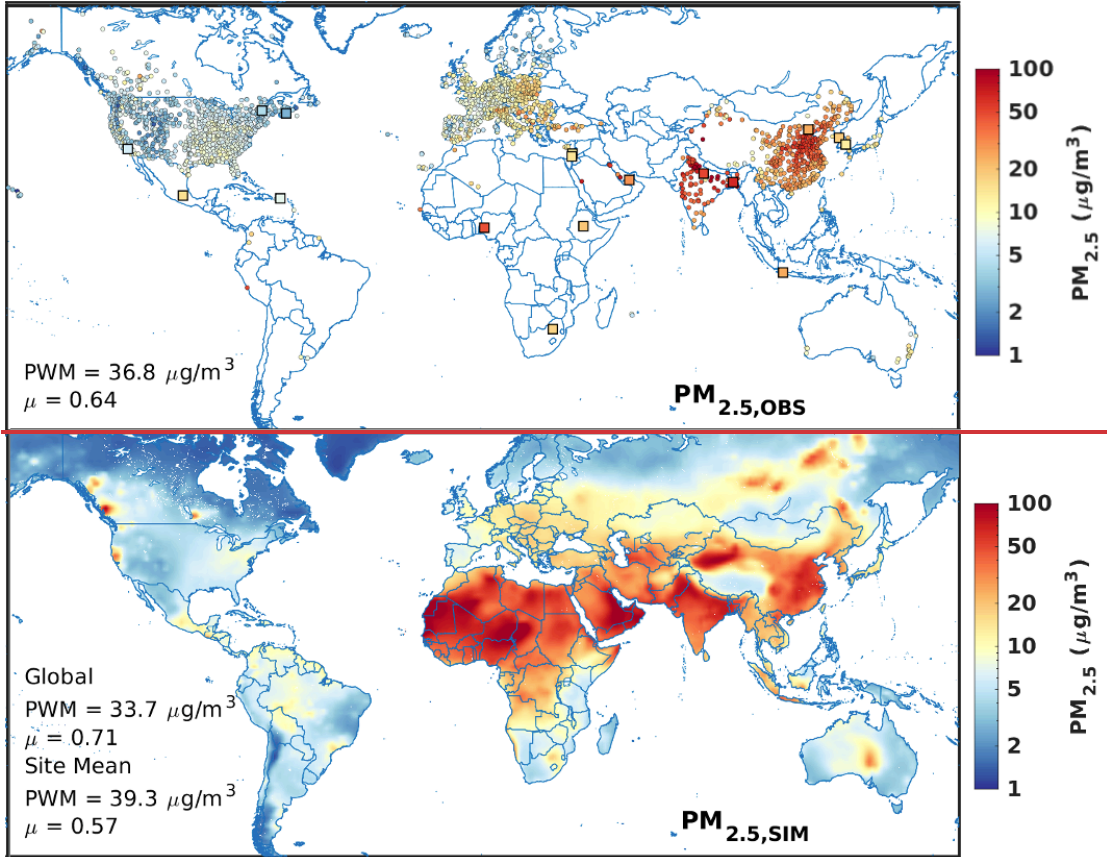
Figure A2. Normalized mean bias (NMB) between simulated $PM_{2.5}$ chemical composition and ground measurements from CSN, IMPROVE, EBAS, and SPARTAN. The original simulation is the out-of-box version of GCHP v13.4.0, the updated simulation includes adjustments such as GFED4.1s emission at daily scale, diel variation and vertical distribution of anthropogenic emissions, and 50% reduction in nitrate concentration.



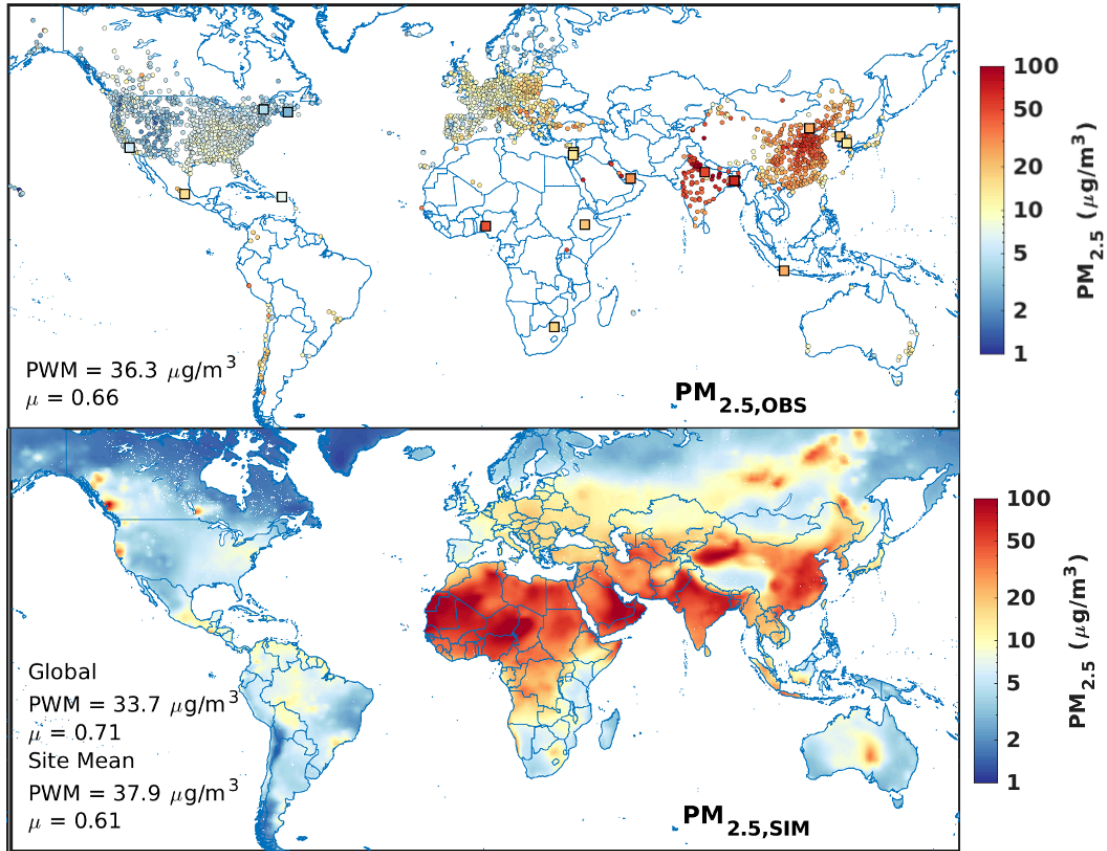
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Figure A3. Region definition.



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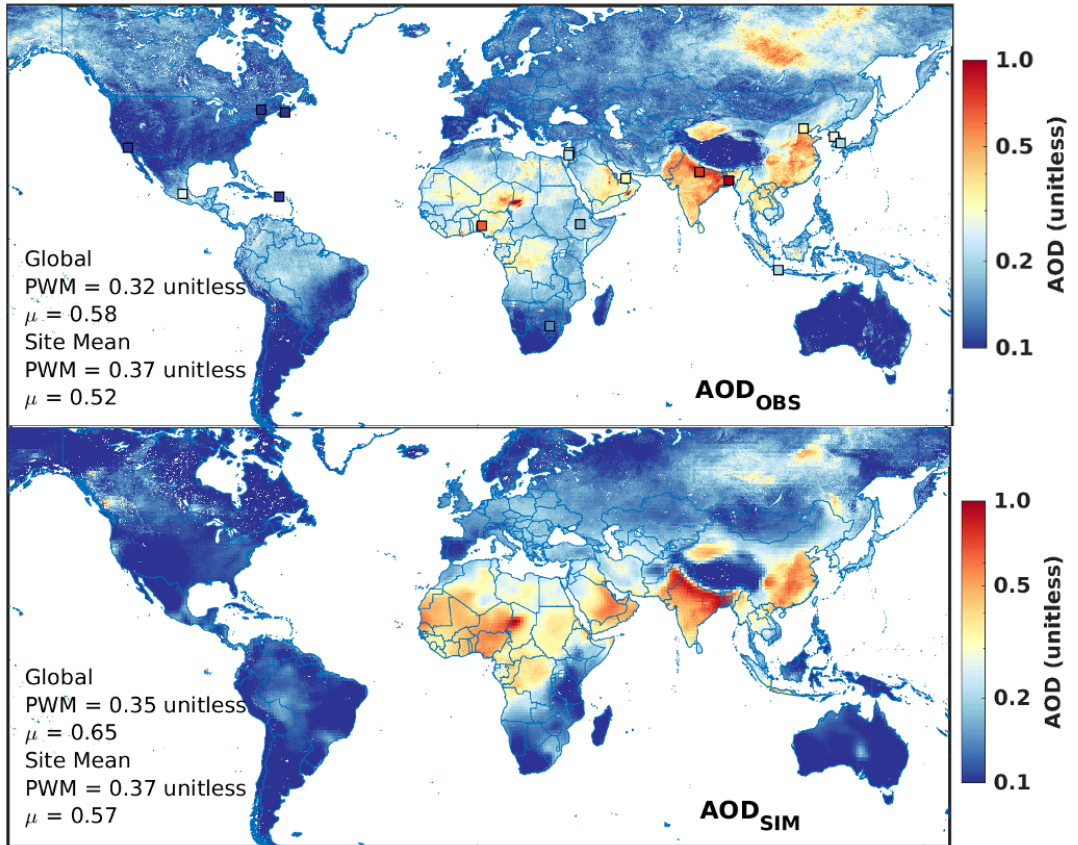
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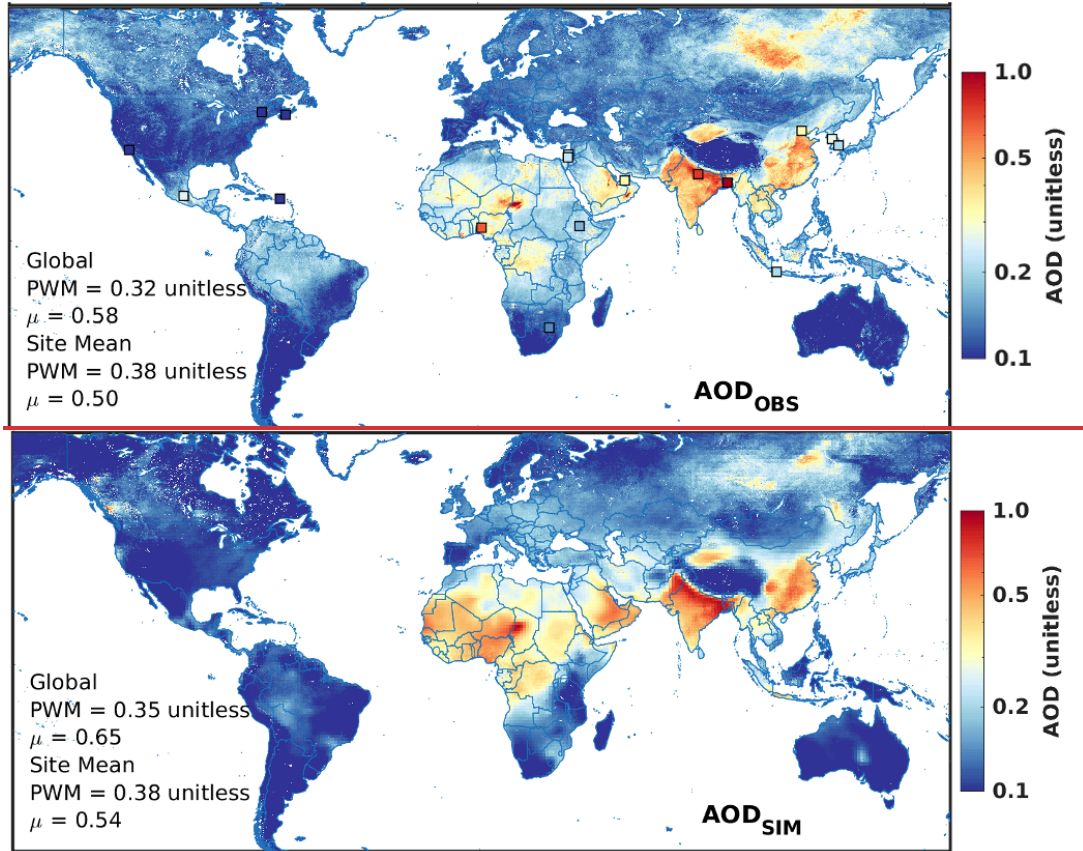
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Figure A4. Observed (top) and simulated (bottom) annual mean $PM_{2.5}$ for 2019. Circles represent measurement sites from regional networks or reported by the WHO. Squares represent measured $PM_{2.5}$ from SPARTAN. PWM = population-weighted mean, μ = coefficient of variation.



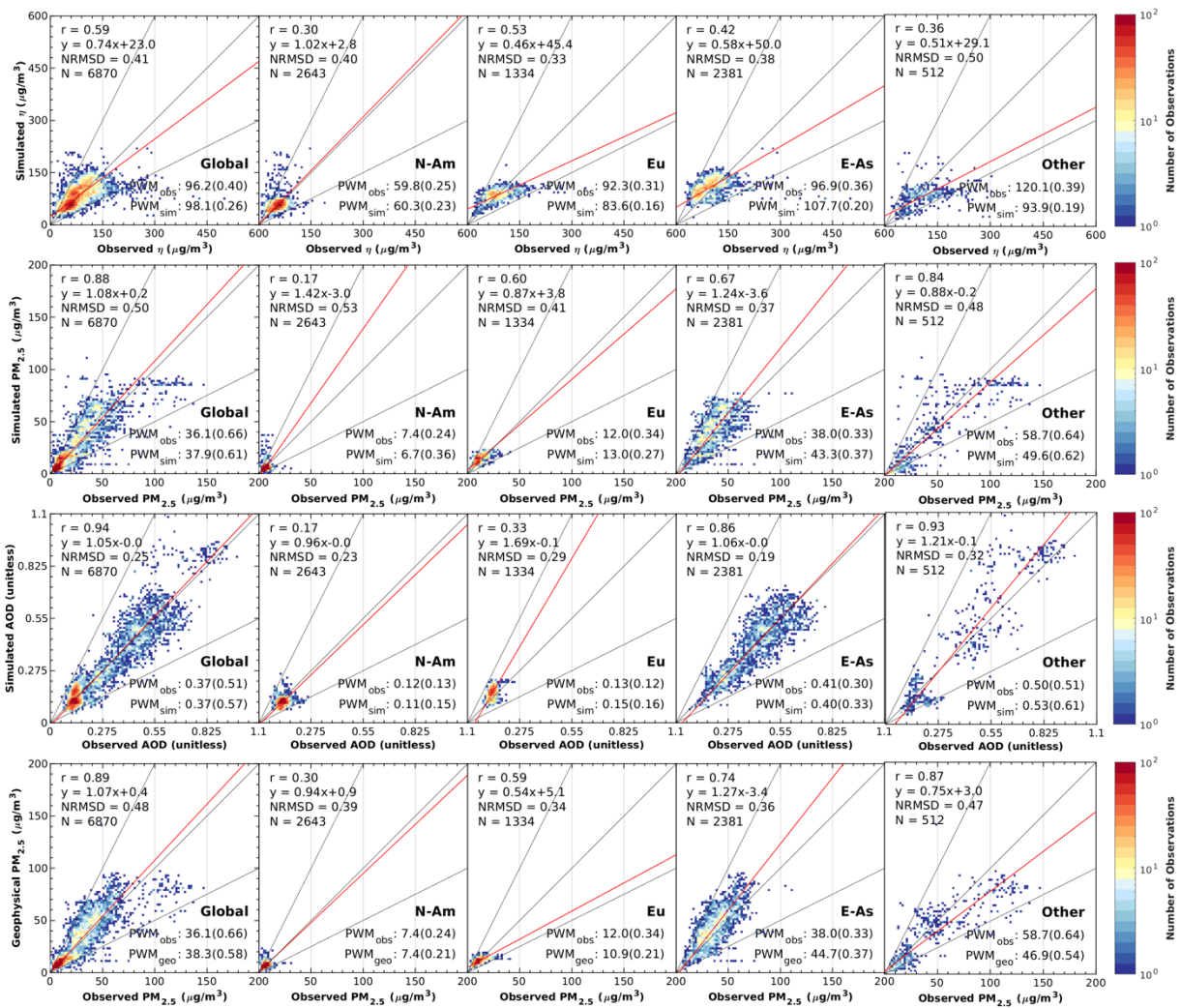
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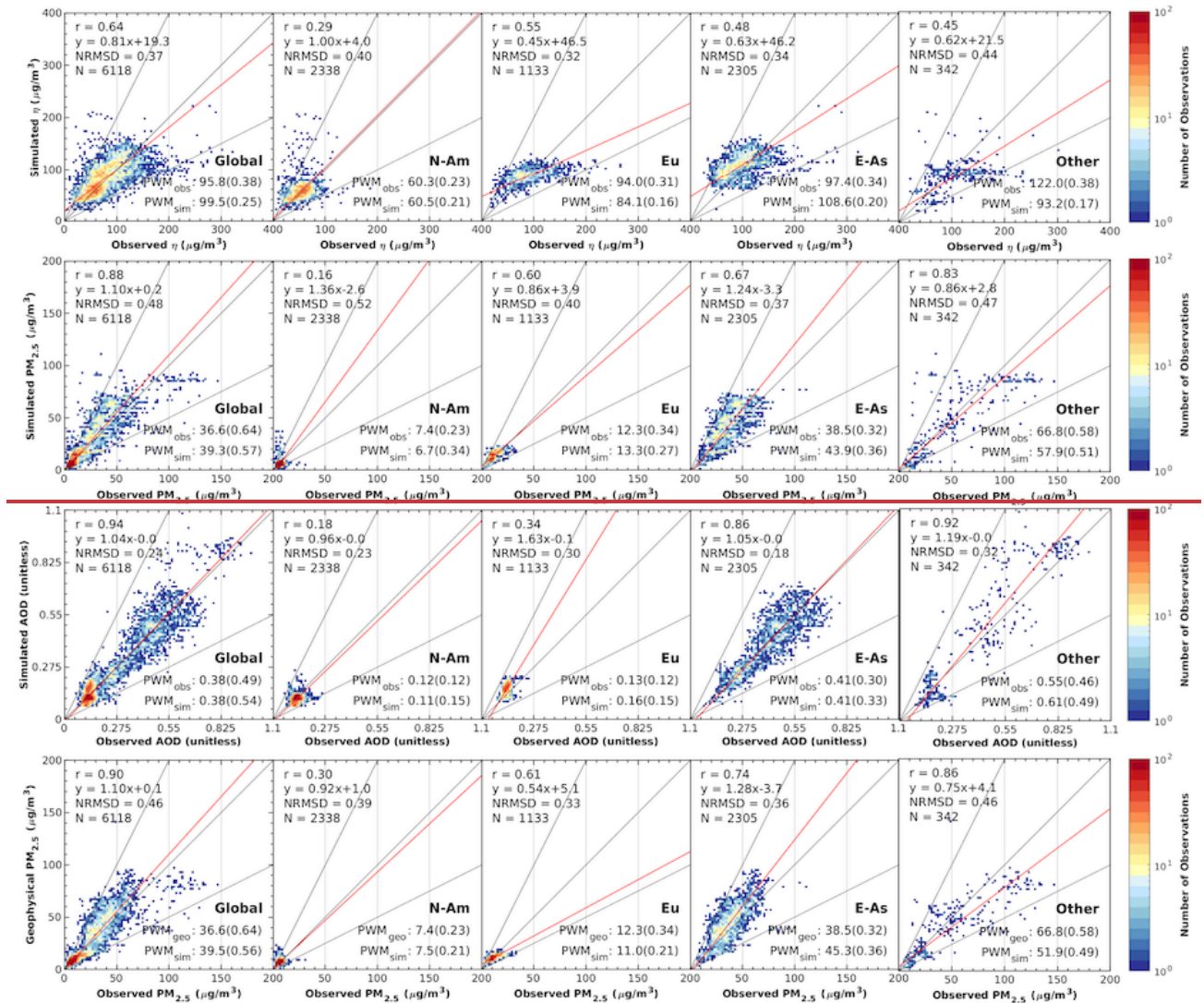


424

425 Figure A5. Satellite retrieved (top) and GCHP simulated (bottom) annual mean AOD for 2019. Squares
 426 represent ground-measured AOD from AERONET. PWM = population-weighted mean, μ = coefficient of
 427 variation.

428





430

431 Figure A6. Scatter plots of simulated and observed η (top row), simulated and ground measured $PM_{2.5}$
 432 (second row), simulated and MAIAC AOD (third row), and geophysical and observed $PM_{2.5}$ (bottom
 433 row). The red line shows the line of best fit using Reduced Major Axis Linear Regression. Insets on the
 434 top left show the coefficient of determination (R^2), line of best fit, normalized root mean square deviation
 435 (NRMSD), and total number of data points (N). The bottom right insets show the population-weighted
 436 mean of observed, simulated, or geophysical estimation of each dataset, coefficients of variation are
 437 bracketed. Detailed regional mean and coefficients of variation for other regions can be found in
 438 A1.

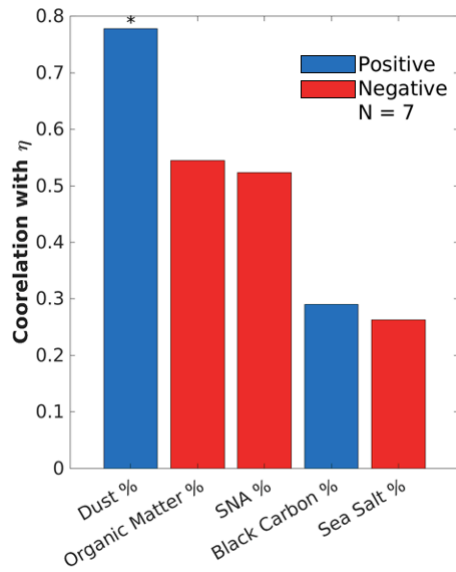
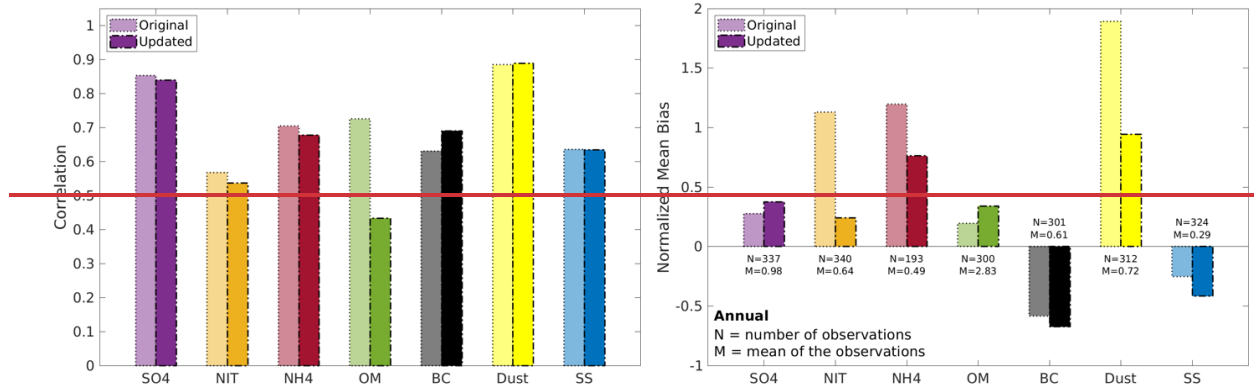
439 Table A1. Regional population-weighted mean η , PM_{2.5}, and AOD from both observation and
 440 simulations. Geophysical PM_{2.5} is also included. Coefficients of variation are bracketed. Regional mean
 441 and coefficients of variation for North America, Europe, and East Asia can be found in Figure A6.

Region	South Asia	Southeast Asia	West Asia	Latin America	Middle East	North Africa	Sub-Saharan Africa	Australia	
Number of sites	16222 0	35	43	2	46142	2932	3	56	
η [$\mu\text{g}/\text{m}^3$]	Observed	121.61 19.5 (0.367)	128111.4 .6 (0.2142)	154.0 (0.23)	72.0 (0.29)	94.111 7.5 (0.516)	135.0 2.3 (0.325)	133.9187 .8 (0.34)	
		Simulated	95.03 4 (0.140)	82.493.8 (0.1809)	93.4 (0.03)	74.1 (0.04)	83.686 6 (0.1821)	126.6 135.8 (0.197)	187.3128 4 (0.5426)
	Observed	81.075 .7 (0.451)	35.740.6 (0.4426)	22.0 (0.21)	12.0 (0.23)	21.720 4 (0.3651)	28.73 2.2 (0.536)	24.0 (0.00)	35.546.3 (0.29)
PM_{2.5} [$\mu\text{g}/\text{m}^3$]	Simulated	70.264 .9 (0.370)	31.838.1 (0.230)	20.8 (0.08)	20.9 (0.06)	10.12 (0.3025)	38.34 7.2 (0.525)	16.7 (0.03)	90.056.6 (0.3187)
		Geo-physical	62.759 .9 (0.310)	22.736.1 (0.4329)	13.9 (0.08)	12.4 (0.08)	20.417 6 (0.397)	27.33 3.0 (0.404)	12.9 (0.03)
	Observed	0.637 (0.295)	0.2738 (0.305)	0.14 (0.08)	0.17 (0.03)	0.204 (0.3221)	0.231 (0.302)	0.12 (0.01)	0.2730 (0.5266)
AOD [unitless]	Simulated	0.6973 (0.3628)	0.4038 (0.128)	0.22 (0.09)	0.28 (0.02)	0.2142 (0.2314)	0.29 33 (0.342)	0.16 (0.01)	0.3751 (0.4726)

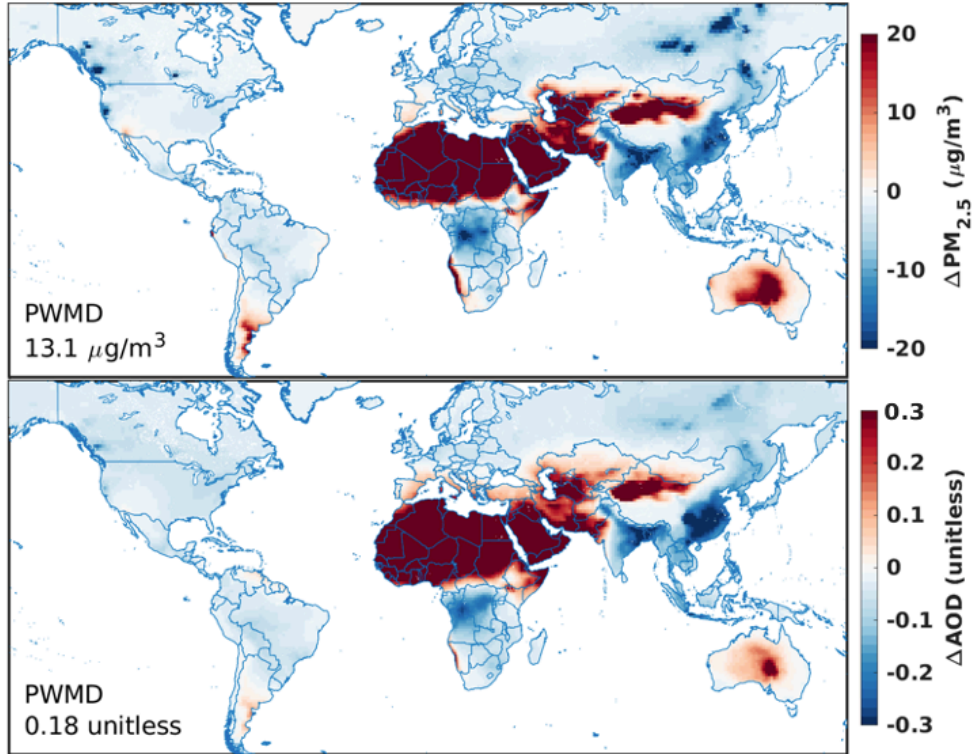
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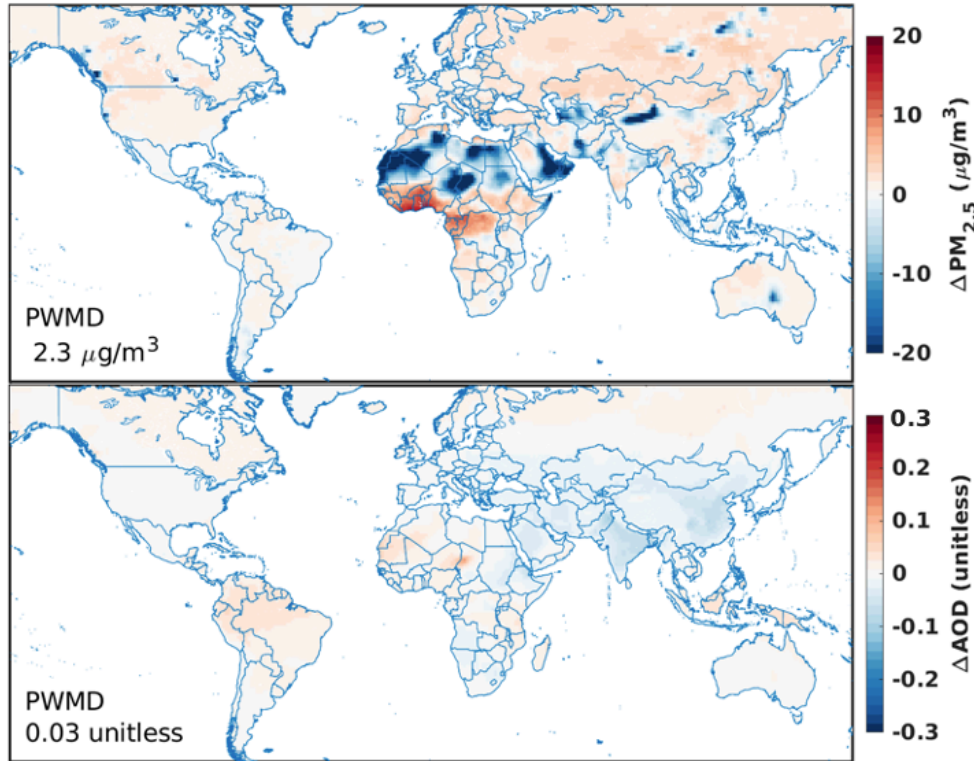


445
 446 Figure A7. Correlation with η of ground-measured aerosol fractional composition from SPARTAN.
 447 Organic matter is inferred through residual (Snider et al., 2016). Blue bars indicate positive correlations.
 448 Red bars indicate negative correlations. Stars above each bar indicate the p-value associated with each
 449 correlation. ‘***’ means the p-value is lower than 0.001, ‘**’ means lower than 0.01, and ‘*’ means
 450 lower than 0.5.



451

452 Figure A8. Changes in PM_{2.5} (top) and AOD (bottom) (test - base) when imposing a global PWM aerosol
453 composition.



454

455 Figure A9. Changes in $PM_{2.5}$ (top) and AOD (bottom) (test - base) when imposing a global PWM aerosol
 456 profile.

457 *Data availability.* GEOS-Chem in its high-performance configuration version 13.4.0 can be
 458 downloaded at <https://zenodo.org/records/6564711>.

459 *Author contributions.* HZ and RVM designed the study. HZ performed the data analysis and model
 460 simulation with contributions from AvD, CL, YL, DZ, JM, MH & IS. AvD contributed to the
 461 compiled the MAIAC AOD dataset and ground-based observation datasets for $PM_{2.5}$. AL
 462 contributed to the original MAIAC AOD dataset. CRO and XL contributed to the SPARTAN data
 463 utilization and analysis. The manuscript was written by HZ and RVM with contributions from all
 464 authors.

465 *Competing interests.* The authors declare no competing financial interest.

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 467 Zhou from Princeton University for providing ground based $PM_{2.5}$ data over India.

468

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