Assessment framework to predict sensitivity of marine calcifiers to ocean alkalinity
 enhancement - identification of biological thresholds and importance of precautionary
 principle

4 Nina Bednaršek¹*, Hanna van de Mortel², Greg Pelletier³, Marisol García-Reyes⁴, Richard A.

5 Feely⁵, Andrew G. Dickson⁶

- 6 ¹*Cooperative Institute for Marine Ecosystem and Resources Studies, Hatfield Marine Science
- 7 Center, Oregon State University, 2030 SE Marine Science Drive Newport, OR 97365, USA
- 8 ²HvdMortel Consulting, Utrecht, NL
- ³Washington Department of Ecology, Olympia, 300 Desmond Dr SE, WA 98503, USA (Emeritus)
- ⁴Farallon Institute, 101 St. Suite Q, Petaluma, CA 94952, United States
- 11 ⁵NOAA Pacific Marine Environmental Laboratory, Seattle, WA, 98115 USA
- ⁶University of California at San Diego, Scripps Institution of Oceanography, 9500 Gilman Drive,
- 13 La Jolla, CA 92093, USA (Emeritus)
- 14 *Correspondence to*: nina.bednarsek@oregonstate.edu

15 Abstract

Ocean alkalinity enhancement (OAE), one of the marine carbon dioxide removal strategies, is 16 17 gaining recognition in its ability to mitigate climate change and ocean acidification (OA). OAE is 18 based on adding alkalinity to open-ocean and coastal marine systems through a variety of different 19 approaches, which raises carbonate chemistry parameters (such as pH, total alkalinity, aragonite 20 saturation state), and enhances the uptake of carbon dioxide (CO_2) from the atmosphere. There are 21 large uncertainties in both short- and long-term outcomes related to potential environmental 22 impacts, which would ultimately have an influence on the social license and success of OAE as a 23 climate strategy. This paper represents a synthesis effort, leveraging on the OA studies and 24 published data, observed patterns and generalizable responses. Our assessment framework was 25 developed to predict the sensitivity of marine calcifiers to OAE by using data originating from OA 26 studies. The synthesis was done using raw experimental OA data based on 68 collected studies, 27 covering 84 unique species and capturing the responses of eleven biological groups (calcifying 28 algae, corals, dinoflagellates, mollusks, gastropods, pteropods, coccolithophores, annelids, 29 crustacean, echinoderms, and foraminifera), using regression analyses to predict biological 30 responses to NaOH or Na₂CO₃ addition and their respective thresholds. Predicted responses were 31 categorized into six different categories (linear positive and negative, threshold positive and 32 negative, parabolic and neutral) to delineate responses per species. The results show that 34.4% of 33 responses are predicted to be positive (N=33), 26.0% negative (N=25), and 39.2% (N=38) neutral 34 upon alkalinity addition. For the negatively impacted species, biological thresholds, which were 35 based on 50% reduction of calcification rate, were in the range of 50 to 500 µmol/kg NaOH 36 addition. Thus, we emphasize the importance of including much lower additions of alkalinity in 37 experimental trials to realistically evaluate *in situ* biological responses. The primary goal of the 38 research was to provide an assessment of biological rates and thresholds predicted under 39 NaOH/Na₂CO₃ addition that can serve as a tool for delineating OAE risks, guiding and prioritizing 40 future OAE biological research and regional OAE monitoring efforts and communicate the risks 41 with stakeholders. This is important given the fact that at least some of the current OAE approaches 42 do not always assure safe biological space. With 60% of responses being non-neutral, a 43 precautionary approach for OAE implementation is warranted, identifying the conditions where 44 potential negative ecological outcomes could happen, which is key for scaling up and avoiding 45 ecological risks.

46 **1. Introduction**

47 Anthropogenic carbon dioxide (CO₂) emissions have increased at an unprecedented rate and have 48 contributed to global climate change and negative ecological and biogeochemical impacts in the 49 oceans (Feely et al., 2004; Gattuso et al., 2018), to the extent of crossing six different planetary 50 boundaries (Richardson et al., 2023). Oceans play a crucial role in attenuating the increase in 51 atmospheric CO_2 through the absorption of the excess atmospheric CO_2 of roughly a quarter of anthropogenic carbon dioxide (CO₂) emissions, drawing down around 2–3 Pg C yr⁻¹ in recent 52 53 decades (Friedlingstein et al., 2022). However, without substantial CO₂ emissions abatement and 54 CO₂ removal strategies, profound repercussions on climate, extreme weather events, and 55 socioeconomic implications will follow. Ocean-based CO₂ removal and sequestration strategies 56 (broadly referred to as marine CDR) are among the proposed CDR approaches that remove CO_2 57 and store it for geologically relevant times (National Academies of Sciences, Engineering, and Medicine, 2021). These mCDR approaches only complement CO₂ emission reductions and 58 59 contribute to the portfolio of climate response strategies needed to meet the global goal of limiting 60 warming to well below 2°C as established by the Paris Agreement. Various mCDR approaches 61 have unique benefits and costs but differ in their value depending on their state of implementation, 62 and whether they act globally and/or locally (Oschlies et al., 2023).

63 Ocean alkalinity enhancement (OAE) has the potential to mitigate climate change through 64 increasing ocean uptake of CO₂, while simultaneously reversing ocean acidification (OA) and 65 improving marine habitats. Despite mostly being in the concept stage, OAE is viewed with a high 66 level of confidence as to its effectiveness: medium on environmental risk, but low on the 67 underlying knowledge base (Eisaman et al., 2023; Gattuso et al., 2021; National Academies of 68 Sciences, Engineering, and Medicine, 2021). One of the major concerns about OAE is large 69 uncertainties in both short- and long-term OAE outcomes related to potential environmental 70 impacts of OAE (Kheshgi, 1995; Bach et al., 2019), especially if OAE were to induce novel 71 conditions in the marine systems that are outside the range of the natural variability, exposing 72 organisms to conditions not experienced in their evolutionary history. The outcome of OAE as a successful climate strategy depends on a thorough and advanced understanding of the impacts of 73 74 OAE implementation while avoiding or minimizing negative biological effects.

75 1.1 Leveraging ocean acidification research on marine calcifiers

76 Increased CO₂ uptake, which initially is absorbed by the ocean as dissolved CO₂, causes a decline 77 in pH, shoaling of the saturation state horizon (Ω_{ar}) and reduced carbonate ion amount content in 78 a process termed ocean acidification (Feely et al., 2004), causing negative consequences to marine 79 biota, especially marine calcifiers, the structure and function of the vulnerable marine ecosystem, 80 and alteration of the carbon cycle. On the other hand, chemical changes induced by OAE are 81 inherently linked to reversing the OA process: increasing pH, shifting carbonate chemistry 82 speciation towards lower aqueous carbon dioxide (pCO_2) and higher carbonate ion (CO_3^{2-}) content, 83 as well as higher aragonite saturation state (Ω_{ar}). Such changes could either be within the ranges 84 of the variability of the natural systems to which species are acclimatized, or outside them, creating 85 novel conditions for which species might not have developed suitable acclimation strategies. As 86 such, the biological outcomes are, due to their complexity, highly unpredictable.

87 Scientific progress over the past 30+ years of OA research has brought substantial insights into the biological effects, with the most fundamental outcome being that calcifying organisms would be 88 primarily affected (Riebesell and Gattuso, 2015), with the calcification process being one of the 89 90 most susceptible pathways, underpinned by species differences in calcification mechanisms (Ries 91 et al., 2009; 2011; Bach et al., 2013; 2015; Leung et al., 2022). However, OA focused heavily on 92 investigating biological effects on the higher acidity range of the carbonate chemistry conditions 93 predicted under future scenarios and most of the studies focused on manipulating the level of pCO₂ 94 rather than alkalinity. This resulted in poor understanding of the biological effects at the higher pH end of the carbon chemistry range (Renforth and Henderson, 2017). Some biological inferences 95 96 can be made based on the understanding of the physiological mechanisms underlying the 97 calcification mechanisms (Bach et al., 2019), but such insights are not adequate to provide 98 sufficient understanding. Despite the lack of biological data at the upper ranges of pH and Ω_{ar} , this 99 study builds on the premise that previous OA studies could be leveraged for assessment of 100 biological responses under OAE. Comparative experimental work, meta-analyses, and the 101 threshold work (Kroeker et al., 2013; Leung et al., 2022; Bednaršek et al., 2019; 2021b,c) have 102 indicated that even very diverse responses can be grouped into categorical responses.

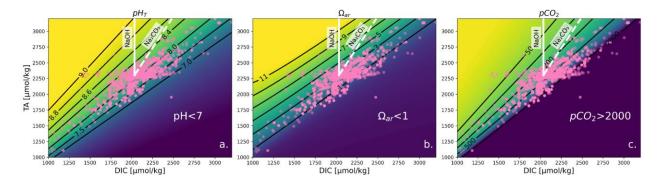
103 Calcification is a primary pathway through which organismal sensitivity to OA is expressed. It is

104 directly involved in growth and (abnormal) development across most marine calcifiers, and it 105 indirectly influences susceptibility to predation. As such, calcification can serve as an early 106 warning indicator of stress, while also playing a crucial role in the ecological success of numerous 107 marine calcifiers. Studies have shown that the thresholds for calcification occur at similar pH and 108 saturation state (Ω) values as those affecting energy metabolism processes (Lutier et al., 2022; Bednaršek et al., 2019; 2021b,c). Furthermore, calcification is directly linked to carbon export, 109 110 which has significant biogeochemical implications that may influence the efficiency of OAE. This 111 study aims to systematically assess the calcification responses of various species under predicted 112 conditions following carbonate-based OAE compound addition.

113 **1.2.** Complex carbonate chemistry changes induced by various OAE compounds

114 Various OAE compounds added to the water change carbonate chemistry in a multifaceted way 115 and require complex calculations of a multi-parameter problem. As the values of TA and DIC change, a variety of other parameters, such as pH, CO_3^{2-} , Ω_{ar} , and pCO₂ exhibit approximately 116 linear relationships, with slopes that vary along these lines (see Fig. 1). This means that if TA and 117 118 DIC vary in proportion to one another, then the values of these displayed parameters hardly change 119 at a particular salinity, temperature, and pressure. With TA, DIC and the hydrographic conditions 120 (salinity, temperature and pressure), one can constrain the carbonate system. Our method requires 121 us to have one variable constraining the entire carbonate system. TA and DIC have the benefit that 122 they can both be directly measured with high precision and accuracy or calculated from other 123 carbonate parameters. They are also both directly linked to OAE, as we are enhancing the TA 124 which then allows DIC to increase over time due to the gradual uptake of atmospheric CO₂.

To demonstrate the changes of the carbonate system in the experimental system, Figure 1 shows the changes in carbonate parameters with the addition of two OAE compounds, i.e. NaOH (solid line) and Na₂CO₃ (dashed line) to seawater. When NaOH is added, only TA increases and when Na₂CO₃ is added, TA and DIC increase at a 2:1 ratio. This results in corresponding changes in pH (Fig. 1a), Ω_{ar} (Fig. 1b) and pCO₂ (Fig. 1c) and shows how much of a change is required to bring the system back to equilibrium with respect to the atmosphere.



133 **Figure 1:** The effect of changes in TA and DIC on the properties of seawater (S = 34.68, $T = 16^{\circ}C$, $[SiO_2] = 50 \ \mu mol/kg, [PO_4^{3-}] = 0.5 \ \mu mol/kg, TA = 2303 \ \mu mol/kg, DIC = 2034 \ \mu mol/kg), adapted$ 134 from Schulz et al. (2023). Pink dots represent experimental TA and DIC data used in our synthesis. 135 136 Subfigures show pH_T , Ω_{ar} and pCO_2 (in μatm). Calculations were carried out with the Python version of CO2SYS (Humphreys et al., 2022) using the stoichiometric dissociation constants for 137 138 carbonic acid from Sulpis et al. (2020), for sulfuric acid by Dickson et al. (1990) and for total 139 boron from Uppström (1974). The solid white line indicates the effect of adding NaOH and the dashed white line indicates the effect of adding Na₂CO₃. This grouping of lines can be translated 140 141 so that its initial position moves elsewhere to visualize different initial conditions. Note that at TA 142 $< 1000 \,\mu$ mol/kg and DIC $< 500 \,\mu$ mol/kg the isolines are no longer straight when considering Ω_{ar} , however, such conditions are rare in the ocean and not widely applicable. The same contour plot 143 144 utilizing GLODAP data plotted instead of experimental data is shown in Supplemental Figure 1.

145 **1.3 Testable conceptual framework based on the existing OA studies**

132

146 Based on Ries et al. (2009), calcification responses can be categorized into six categories (Fig. 2): 147 linear positive or negative response; threshold positive or negative response (exponential fit); 148 parabolic response; and neutral (no significant) response. We hypothesize that these categories of 149 responses based on ocean acidification data and delineated by Ries et al. (2009, 2011), could also 150 be applicable to OAE dosing. For this meta-analysis, we have undertaken three steps: first, synthesize carbonate chemistry data at regional and global scales to obtain TA, DIC and Ω_{ar} 151 152 correlations; second, conduct a literature review and collect available data from OA literature 153 related to the calcification rate responses across the species of eleven groups of marine calcifiers; 154 and third, run regression analyses and determine the category of calcification rate response to 155 TA:DIC, further extending it with addition of NaOH and Na₂CO₃.

The most accurate way of predicting the responses to OAE addition is done based on the mechanistic understanding of calcification response to specific carbonate chemistry parameter(s). The hypothesis was that if mechanistic relationships with identified carbonate chemistry driver(s) are available for species, calcification rate under various feasible OAE scenarios can be predicted with greater accuracy and lower uncertainty. We further focused on investigating if the empirical results were consistent with mechanistic calcification predictions for a few selected species for which the mechanisms were known.

163

Here, we demonstrate the TA:DIC relationship with calcification rates and show the application for the TA:DIC thresholds beyond which the responses become negative. Ultimately, we synthesize which calcifying species or groups are predicted to benefit or lose due to OAE, what constitutes a species-specific safe operating space related to OAE, and we delineate what experiments are most urgently needed to fill in critical knowledge gaps before massive OAE field implementation can be considered.

170 **2. Methodology**

171 2.1 Literature review of data on marine calcification impact by OA

172 To assess the impact of OAE on a range of marine calcifiers, we used existing studies on marine 173 species calcification response to OA that had aligned raw biological (calcification rate) data along 174 with corresponding carbonate chemistry. We searched within Scopus, Web of Science, and 175 PubMed and used datasets that were archived in NCEI, OA-ICC and Pangaea. Through personal 176 correspondence, we have additionally contacted lead authors of the studies whose data are not or 177 are insufficiently archived. Searches for biological datasets relating to calcification rate and 178 corresponding carbonate chemistry were conducted through November 2023, encompassing 68 179 existing studies. The aim was to cover a wide range of calcifying organisms across various 180 functional groups and 84 species. For several functional groups data was easy to find (algae, 181 coccolithophores, corals, foraminifera, mollusks and dinoflagellates), so no new studies were 182 added after 10 to 15 studies were found. Seven studies were found for pteropods, five for 183 gastropods, four for echinoderms, three for crustaceans and one for annelids. When reviewing the 184 literature, we included data from the OA experimental studies related to the physical-chemical 185 parameters (temperature, salinity, TA, DIC) and biological data related to calcification rate.

186 **2.2** Use of TA:DIC instead of Ω_{ar} or pH

Understanding the change in carbonate chemistry upon alkalinity addition is essential for the biological experimentalists who are conducting biological assessments to report on the effects of OAE. However, complex changes in the carbonate chemistry induced by alkalinity addition are not intuitive or straightforward; in fact, they are multi-parameter problems that require complex carbonate chemistry calculations. Using the TA:DIC ratio is a more practical way of looking at the impacts of the OAE treatment instead of using a single carbonate parameter because of the high degree of correlation between TA:DIC and other carbonate system parameters (see Fig. 1).

With TA, DIC and the hydrographic conditions (salinity, temperature and pressure), one can fully constrain the carbonate system. Our method allows *one* variable constraining the entire carbonate system. TA and DIC have the benefit that they can both be directly measured or calculated from other carbonate and physical parameters. They are also both directly linked to OAE, as we are enhancing the TA which then allows DIC to increase over time due to the gradual uptake of atmospheric CO₂ (Fig. 1 shows the changes in the carbonate chemistry system upon NaOH and Na₂CO₃ addition).

201 Our focus was on streamlining the process of expressing experimental results and subsequently 202 reporting responses, with the goal of reducing the multi-parameter complexity into a single-203 parameter simplification. This step reduces multiple degrees of freedom into just two, i.e. TA and 204 DIC, with the ratio allowing us to consider this as a 1-parameter problem. As such, TA:DIC is a 205 simplistic and convenient way of describing the system, where we only need to understand the 206 change in TA and DIC ratio, which is feasible for every OAE compound added to the experimental system. In addition, TA:DIC is also the best approximation for the CO_3^{2-} concentration. The 207 insights from multiple biological experimental studies show that the CO_3^{2-} concentration is the 208 209 representative driver of the calcification process for multiple calcifying groups, although not all, 210 compared to Ω_{ar} , which represents an empirical approximation based on a number of physical and 211 chemical parameters. Furthermore, by using TA:DIC we do not have to choose a particular 212 parameter to describe the changes in calcification. It could also work for the species in which other parameters drive the calcification, e.g. bicarbonate in autotrophic species, Ω_{ar} in bivalves and H⁺ 213

flux in foraminifera. In that way, we standardize all the parameters that would otherwise influence the carbonate system and come up with a more uniform way to express the experimental conditions, which would then be useful for easier comparisons among the conducted experiments. For the ease of comparing TA:DIC with pH and Ω_{ar} , we refer the reader to Supplemental Table 1 and Supplemental Fig. 2.

219 **2.3 Experimental biological and biogeochemical data**

Based on the collected data, the range of pH and Ω_{ar} , experimental conditions used and their TA:DIC relationship was determined (Supplemental Fig. 2 and Supplemental Table 1). Most studies covered pH conditions from 7.5 to 8.5 and Ω_{ar} from <1.0 to values up to 5.0, with a few studies increasing pH up to 9 and exceeding Ω_{ar} of 10. This indicates the potential of leveraging such experimental studies as a baseline for predictive regression models of biological responses to a range of Ω_{ar} conditions, as expected under OAE studies.

226 Once the biological data was compiled, units were standardized where possible. The main issue when compiling data was the lack of standardization of the calcification rates. A variety of 227 228 calcification rate units were used across different studies. Where possible, the units were converted to mmol of CaCO₃ g weight⁻¹ hr⁻¹. However, the data required to do so was not always readily 229 available. Other units used for calcification rate were mmol of CaCO₃ m⁻² h⁻¹ and mmol of CaCO₃ 230 m⁻³ hr⁻¹, and there was also data used as an indication of calcification rate with units mmol #⁻¹ h⁻¹, 231 mmol h⁻¹, mmol cm⁻², % h⁻¹, where '#' indicates one individual. Growth rates and PIC production 232 233 rates were used as indicators of calcification rate for single-cell organisms. For some species, direct 234 calcification rates were not reported in the literature, instead only relevant parameters related to 235 calcification (shell length, density, thickness) over time were available from the experimental 236 studies. The decision was made to also collect these additional datasets because the statistical 237 analyses of this study focus on the trend in the absolute numbers and would not change by being 238 transformed into the rates. Data were analyzed on a species level, wherever rate units were the 239 same. Hereafter, this is referred to as the species rate group. Where there were multiple studies 240 available for the calcification rate of one species using the same rate units, the data were combined 241 (e.g. Emiliania huxleyi).

242 2.4 Sorting species-specific responses into categories per calcification response

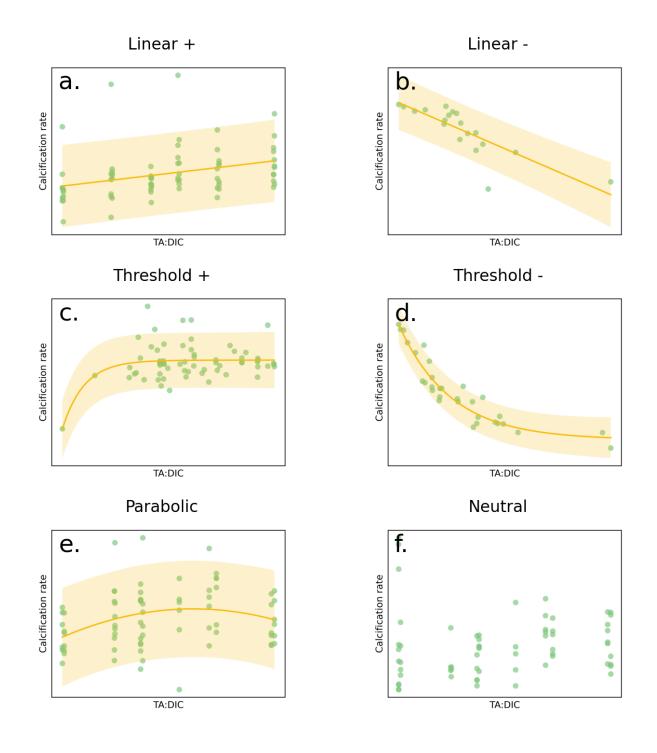
Responses were split into 6 categories: linear positive and linear negative, parabolic, threshold
positive and negative, and neutral. The response was determined with a best-fit regression model,
using the ordinary least squares method in Statsmodels for Python (see Seabold et al., 2010). See
Fig. 2 for examples of these responses of calcification rate to increasing TA:DIC ratio.

247 The final response for each species was determined by the regression with the lowest p-value. This 248 method is in contrast with the Ries et al. (2009) study where they chose the regression analysis 249 that yielded the lowest square root of the mean squared error (RMSE) for a given species, and that 250 was statistically significant ($p \le 0.05$). When applying their method to our data, parabolic and 251 exponential regressions were always favored over linear regressions. When examining these 252 regressions, we found that choosing the best fit based on the lowest p-value yielded better fits, as 253 this method prevents overfitting due to noise in the data. Where a linear regression had the best fit, 254 we assigned a linear response, which could be either positive or negative based on the slope. The 255 species with a significant exponential fit were categorized as threshold positive (+) or threshold 256 negative (-), which was distinguished from the parabolic response with the fitted parabolic curve.

The best fit regression was assigned to each species and plotted, but only if the p-value was considered significant, i.e. lower than 0.05. These regressions were plotted along with a 90% prediction interval, which accounts for the variability of the experimental data. The species with a p-value > 0.05 were categorized as having no correlation (neutral response).

261 When multiple datasets obtained from different studies for the same species and rate units could 262 not be combined, we took each response into consideration and reported the p-value and RMSE 263 for each of the studies. Even when different studies reported varying calcification rates for the 264 same species, we refrained from selecting a single overall species response; rather, we analyzed 265 each species individually. The TA:DIC threshold was computed to indicate the point at which the 266 current calcification rate (i.e. the calcification rate at the baseline) is reduced by a half for linear 267 negative, threshold negative and parabolic responders. The thresholds and the amount of NaOH 268 and Na₂CO₃ required (starting at 10 μ mol/kg and then in steps of 50 μ mol/kg) to reach this 269 threshold were determined. For parabolic responders, the inflection points that tell us when the 270 rate is predicted to change slope are also included in Supplemental Table 2. Once the species' 271 responses were determined, an attempt was made to group them based on functional groups.

However, since species within the same functional group had varying responses, grouping them
together meant these responses were no longer visible due to a wide spread of data. Therefore,
most of the analysis remained on the species level (Table 1).



276 Figure 2: Examples of the categories of responses between carbonate chemistry parameters

277 (TA:DIC) and calcification rate: a) linear positive (calcification increase with increased TA:DIC); 278 b) linear negative (calcification decrease with increased TA:DIC); c) exponential for the threshold 279 positive response (calcification increase, plateauing at higher TA:DIC); d) exponential for the 280 threshold negative response (calcification decline, plateauing at lower TA:DIC), e parabolic 281 (calcification increase followed by a decrease at higher TA:DIC) and f) neutral (non-significant) 282 response. Responses were only considered significant when p < 0.05, otherwise they were 283 categorized as neutral. Yellow shading represents the 90% prediction interval. Note that TA:DIC 284 on the x-axis corresponds to pH_T and Ω_{ar} , as these variables have an approximately linear 285 relationship at a particular salinity, temperature and pressure (see Fig. 1).

286 2.5 Conceptual framework to evaluate increases in TA:DIC

287 The regression models applied to each species could be used to predict calcification rates at higher 288 TA:DIC ratio. We conceptually added alkalinity from the current calcification rate baseline. This 289 baseline was computed for each species using CO2SYS with $pCO_2 = 425$ ppm and $pH_T = 8.1$, for 290 the average temperature and salinity for each species rate group, based on their respective OA 291 dataset(s) (see Supplemental Table 3). All CO2SYS calculations in this study were carried out 292 with the Python version of CO2SYS (Humphreys et al., 2022) using the stoichiometric dissociation 293 constants for carbonic acid from Sulpis et al. (2020), for sulfuric acid by Dickson et al. (1990) and 294 for total boron from Uppström (1974). From this baseline, TA was added in the form of both NaOH 295 and Na₂CO₃ to approximate changes in the carbonate chemistry settings, with NaOH changing 296 TA:DIC in the 1:1 ratio, and Na₂CO₃ inducing a 2:1 TA:DIC change. For example, 10 µmol/kg of 297 NaOH addition will increase TA by 10 µmol/kg and not affect DIC. For Na₂CO₃, 10 µmol/kg 298 addition will increase TA by 10 µmol/kg and increase DIC by 5 µmol/kg. Figure 1 demonstrates 299 the usefulness of this approach. For both NaOH and Na2CO3, 10 µmol/kg was conceptually added 300 using the principles of mass balance approach for the carbonate system via CO2SYS. This was 301 repeated for increments of 50 µmol/kg. We show this incremental addition in the plots up to a total 302 of 500 µmol/kg when generating the plots. When computing the thresholds, we added up to 1400 303 μ mol/kg NaOH. The new TA:DIC ratios were estimated by adding the direct effect of Δ TA and 304 Δ DIC due to chemical additions of NaOH (assume Δ DIC = 0) or Na2CO3 (assume Δ DIC = 305 $0.5^*\Delta TA$). A maximum of 500 µmol/kg was chosen to have more realistic additions of TA that 306 resemble those appropriate within the OAE field trials (e.g. Wang et al., 2023). With the new

TA:DIC ratios after TA addition, the species' regression models based on the fitted OA response
data were used to compute respective calcification rates (note that added points with NaOH or
Na2CO3 were not calculated as part of the regression). These data points were all plotted along
with the experimental data, regression model and prediction intervals as shown in Fig. 3.

We also determine the amount of NaOH and Na₂CO₃ needed to reach pH_T 9 for each study. This was computed for each species rate group using CO2SYS starting from pCO₂ = 425 ppm and pH_T = 8.1, using the average temperature and salinity per species rate group, and by adding NaOH or Na₂CO₃ in increments of 50 μ mol/kg until pH_T 9 was reached. Note that this method does not incorporate gas exchange with the atmosphere, any biological processes, organic matter effects,

316 nitrification/denitrification, complexation, speciation or sediment-water interactions.

317 **2.6** Evaluation of the biological responses based on alkalinity addition

The individual species with significant correlations were grouped visually based on their best-fit regression models and are classified into positive, negative, and neutral as the following:

320 1) *Positive responders*: species with predicted *linear positive* and *threshold positive* calcification
321 rate response with increased TA addition.

322 2) Negative responders: species with predicted *linear negative*, *parabolic* and *threshold negative*

323 *response* in calcification rate upon (a certain amount of) TA addition. For the parabolic responders,

a concentration of NaOH was determined that indicates the threshold in TA:DIC beyond whichthe response becomes negative (see inflection points in Supplemental Table 2).

326 3) Neutral responders: species with no significant correlation (p < 0.05) in calcification rate upon

327 TA addition.

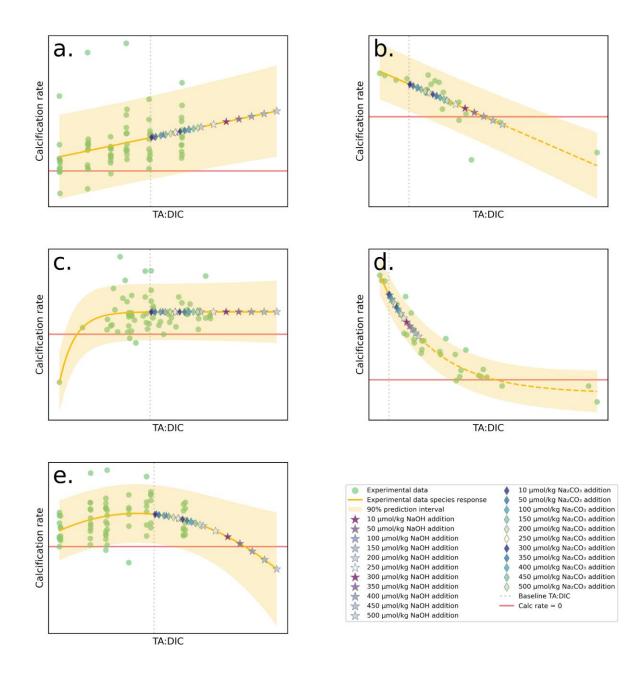


Figure 3: Conceptual diagrams for five types of responses; a) linear positive; b) linear negative;
c) threshold positive; d) threshold negative and e) parabolic response, plotted with experimental
data from OA studies (green dots), predicted values at various additions of alkalinity (stars and
diamonds), the regression line and prediction error margins fitted for a given species. The red
horizontal line indicates zero net dissolution (calcification rate is equal to 0; dissolution rate =
calcification rate). The grey vertical line indicates the baseline from which alkalinity is added.
NaOH and Na₂CO₃ addition is shown up to 500 µmol/kg.

336 2.7 Determining threshold values indicative of negative biological response to OAE

337 The metrics to evaluate the sensitivity of calcification rate of the negative responders in this study 338 were based on the amount of NaOH or Na₂CO₃ addition required to reduce the current calcification 339 rate by a half. The greater the TA:DIC ratio value was required to trigger half calcification rate 340 reduction, the less sensitive species was to NaOH addition. We refer to this TA:DIC ratio as the 341 biological threshold, which we also report along with corresponding pH and Ω_{ar} and the associated 342 uncertainty. TA:DIC thresholds were converted to their respective pH and Ω_{ar} , which are affected 343 by temperature and salinity. To calculate threshold pH and Ω_{ar} we used the average temperature 344 and salinity per species rate group, as done for calculating the baseline.

345 2.8 Extraction of the carbonate chemistry data from the GLODAP dataset

346 We extracted total alkalinity, dissolved inorganic carbon, Ω_{ar} , and pH_T from the Global Ocean 347 Data Analysis Project GLODAPv2.2023 dataset (https://glodap.info). We used the regression 348 application in MATLAB with a second-order polynomial equation to predict Ω_{ar} from the TA:DIC. 349 The regression analysis was performed using data from various depth intervals (0–10m, 0–30m, 350 0-50m, 0-100m, 0-200m) regionally and globally. The regional analysis divided the global oceans 351 into the following groupings: Arctic (north of 65°N), Southern (south of 40°S), North Pacific 352 (north of 40°N), Central Pacific (40°S to 40°N), North Atlantic (North of 40°N), Central Atlantic 353 (40° S to 40° N), and Indian Ocean (north of 40° S).

354 **3. Results**

355 3.1 Data collection for the calcification rate responses of different biological groups

356 We examined 68 datasets, which covered 84 different species that were divided into 11 different 357 groups (Fig. 4). These functional groups were corals (20% of datasets), calcifying algae (18%), 358 mollusks (14%), foraminifera (10%), dinoflagellates (10%), coccolithophores (4%), gastropods 359 (8%), crustaceans (5%), echinoderms (4%), pteropods (5%), and annelids (1%). In the mollusks 360 group, we have separated out the gastropod and pteropod because of a higher number of studies 361 that explicitly cover these two groups. The group of gastropods refers to all gastropods that are not 362 pteropods. If all three groups were combined (mollusks, gastropods, pteropods), this group would 363 be the largest.

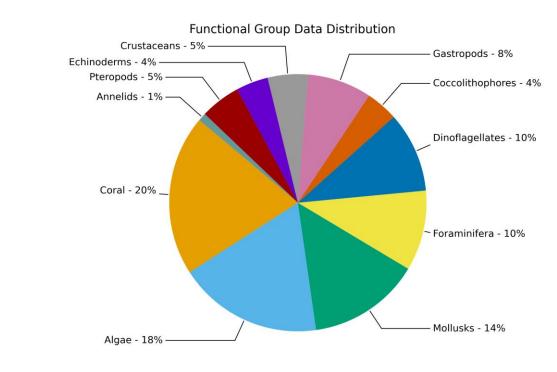


Figure 4: Percent of studies for multiple groups (N=11) with available data for the calcification
rate responses as part of data compilation of 68 studies covering 84 species).

368 **3.2 Species-specific responses to NaOH/Na₂CO₃ addition**

369 Calcification rate responses of species from different groups were correlated to TA:DIC and 370 summarized to obtain calcification rate response. The calcification rate responses encompassed 371 linear (positive and negative), threshold (positive and negative), parabolic, and neutral responses, 372 with the slope and the intercept of the response determining the type and the magnitude of the 373 response. We present fitted responses of calcification rate per TA:DIC ratio for each examined 374 species (Table 1; Supplemental Fig. 4). When possible, we fit a regression to multiple datasets of 375 the same species that used the same calcification units. We also present the response with the 376 additions of NaOH and Na₂CO₃ for each species per examined study and corresponding rate unit 377 and their biological TA:DIC thresholds (Table 2; Supplemental Table 4).

Table 1: The summary of all the OA studies from which the chemical and biological data was
collected, including the name of the species and group and the accompanying calcification rate
unit. The response for each species rate group was determined by the regression with the lowest

p-value, where the *p*-value was smaller than 0.05. We also include the *p*-value, goodness of fit (R^2)

382 and Root Mean Square Error (RMSE). Non-significant responses are categorized as having a

383 'neutral' response. The type of response (linear positive or negative, threshold positive or

negative, parabolic, and neutral) is indicated, as well as if this response is positive, negative or

neutral.

						Pos/Neg/			
Studies	n	Group	Species	Rate unit	Response	Neut	p-value	R ²	RMSE
Vasquez-Elizondo et al. (2016)	4	Algae	Amphiroa tribulus	mmol/m²/hr	neutral	Neutral			
Sinutok et al. (2011)	16	Algae	Halimeda cylindracea	mmol/hr	neutral	Neutral			
Comeau et al. (2013)	71	Algae	Halimeda macroloba	mmol/g/hr	parabolic	Negative	0.0127	0.1200	0.0028
Meyer et al. (2015)	24	Algae	Halimeda macroloba	mmol/m²/hr	neutral	Neutral			
Sinutok et al. (2011)	16	Algae	Halimeda macroloba	mmol/hr	parabolic	Negative	0.0108	0.5000	0.0001
Comeau et al. (2013)	62	Algae	Halimeda minima	mmol/g/hr	neutral	Neutral			
Meyer et al. (2015)	24	Algae	Halimeda opuntia	mmol/m²/hr	linear +	Positive	0.0080	0.2800	0.0222
Comeau et al. (2013)	72	Algae	Hydrolithon reinboldii	mmol/g/hr	linear +	Positive	0.0053	0.1100	0.0026
Cornwall et al. (2018)	23	Algae	Hydrolithon reinboldii	mmol/m²/hr	neutral	Neutral			
Comeau et al. (2013)	72	Algae	Lithophyllum flavescens	mmol/g/hr					
Johnson et al. (2021)	420	Algae	Lithophyllum sp.	mmol/g/hr	linear +	Positive	0.0000	0.1000	0.1136
Vasquez-Elizondo et al. (2016)	4	Algae	Lithothamnion sp.	mmol/m²/hr	neutral	Neutral			
Monserrat et al. (2022)	62	Algae	Neogoniolithon brassica-florida	mmol/m²/hr	neutral	Neutral			
Ries et al. (2009)	42	Algae	Neogoniolithon sp.	mmol/g/hr	parabolic	Negative	0.0000	0.4100	0.0003
Vasquez-Elizondo et al. (2016),									
Comeau et al. (2018)			Neogoniolithon sp.	mmol/m²/hr	neutral	Neutral			
Briggs-Carpenter et al. (2019)	425	Algae	Porolithon onkodes	mmol/m²/hr	linear +	Positive	0.0010	0.0300	0.8093
Comeau et al. (2018, 2019)	64	Algae	Sporolithon durum	mmol/m²/hr	parabolic	Negative	0.0012	0.2000	0.1704
Ries et al. (2009)	41	Annelid	Hydroides crucigera	mmol/g/hr	neutral	Neutral			
Fiorini et al. (2011),									
Langer et al. (2006, 2011)	14	Cocco.	Calcidiscus leptoporus	mmol/#/hr	neutral	Neutral			
*	233	Cocco.	Emiliania huxleyi	mmol/#/hr	parabolic	Negative	0.0000	0.1600	0.0000
Casareto et al. (2009)	14	Cocco.	Pleurochrysis carterae	mmol/m³/hr	neutral	Neutral			
White et al. (2018)	118	Cocco.	Pleurochrysis carterae	mmol/#	neutral	Neutral			
Meyer et al. (2016)	24	Coral	Acropora millepora	mmol/m²/hr	neutral	Neutral			
Camp et al. (2017),									
Comeau et al. (2013)	74	Coral	Acropora pulchra	mmol/m²/hr	parabolic	Negative	0.0000	0.2900	1.3257
Agostini et al. (2021)	18	Coral	Acropora solitaryensis	mmol/m²/hr	neutral	Neutral			
Comeau et al. (2018),									
Comeau et al. (2019)	eau et al. (2019) 81 Coral Acropora		Acropora yongei	mmol/m²/hr	linear +	Positive	0.0000	0.2900	1.9447
Bove et al. (2020)	27	Coral	Duncanopsammia axifuga	mmol/m²/hr	linear +	Positive	0.0016	0.3300	5.0785
Cornwall et al. (2018)	44	Coral	Goniopora sp.	mmol/m²/hr	neutral	Neutral			
Maier et al. (2009)	237	Coral	Lophelia pertusa	mmol/g/hr	linear +	Positive	0.0030	0.0400	0.0002
Bove et al. (2020)	65	Coral	Montastraea cavernosa	mmol/m²/hr	linear +	Positive	0.0154	0.0900	0.5047
Ries et al. (2009)	54	Coral	Oculina arbuscula	mmol/g/hr	parabolic	Negative	0.0000	0.8600	0.0001
Comeau et al. (2013)	72	Coral	Pavona cactus	mmol/m²/hr	parabolic	Negative	0.0002	0.2200	0.9093
Comeau et al. (2019)	49	Coral	Plesiastrea versipora	mmol/m²/hr	linear +	Positive	0.0069	0.1500	0.6003
Brown et al. (2022)	4	Coral	Pocillopora damicornis	mmol/g/hr	neutral	Neutral			
Comeau et al. (2013, 2018),									
Putnam-Gates et al. (2015)	117	Coral	Pocillopora damicornis	mmol/m²/hr	neutral	Neutral			

						Pos/Neg/				
Studies	n	Group	Species	Rate unit	Response	Neut	p-value	R ²	RMSE	
Evensen-Edmunds et al. (2016)	60	Coral	Pocillopora verrucosa	mmol/m²/hr	linear +	Positive	0.0132	0.1000	0.8297	
Agostini et al. (2021)18CoralPorites heronensis				mmol/m²/hr	neutral	Neutral				
Comeau et al. (2013)	72	Coral	Porites rus	mmol/m²/hr	linear +	Positive	0.0020	0.1300	2.0281	
Okazaki et al. (2013)	75	Coral	Siderastrea radians	mmol/m²/hr	linear +	Positive	0.0004	0.1600	2.7886	
Okazaki et al. (2013)	64	Coral	Solenastrea hyades	mmol/m²/hr	threshold +	Positive	0.0004	0.2300	2.0385	
Krueger et al. (2017)	36	Coral	Stylophora pistillata	mmol/m²/hr	neutral	Neutral				
Pansch et al. (2014)	36	Crust.	Amphibalanus improvisus	mmol/g/hr	linear +	Positive	0.0000	0.4300	0.0004	
Ries et al. (2009)	36	Crust.	Callinectes sapidus	mmol/g/hr	linear -	Negative	0.0000	0.4000	0.0082	
Ries et al. (2009)	18	Crust.	Homarus americanus	mmol/g/hr	linear -	Negative	0.0014	0.4800	0.0079	
Ries et al. (2009)	12	Crust.	Penaeus plebejus	mmol/g/hr	linear -	Negative	0.0124	0.4800	0.0006	
Findlay et al. (2010)	6	Crust.	Semibalanus balanoides	mmol/g/hr	neutral	Neutral				
Tatters et al. (2013)	45	Dino.	Alexandrium sp.	1/hr	neutral	Neutral				
Hansen et al. (2007)	19	Dino.	Ceratium lineatum	#/hr	linear -	Negative	0.0000	0.6700	0.0043	
Tatters et al. (2013)	45	Dino.	Gonyaulax sp.	1/hr	neutral	Neutral				
Hansen et al. (2007)		Dino.	Heterocapsa triquetra	#/hr	threshold -	Negative	0.0000	0.9100	0.0027	
Wang et al. (2019)		Dino.	Karenia mikimotoi	1/hr	neutral	Neutral				
Tatters et al. (2013)		Dino.	Lingulodinium polyedrum	1/hr	neutral	Neutral				
Tatters et al. (2013)		Dino.	Prorocentrum micans	1/hr	neutral	Neutral				
			Prorocentrum minimum	#/hr	threshold -	Negative	0.0000	0.8800	0.0019	
			#/hr	linear -	-	0.0000	0.0600	0.0019		
Brading et al. (2011) 175 Dino. Symbiodinium sp.			mmol/hr		Negative	0.0010	0.0600	0.0000		
	n de Waal et al. (2013) 12 Dino. Thoracosphaera heimii		-		parabolic	Negative				
Ries et al. (2009)		Echino.	Arbacia punctulata	mmol/g/hr	parabolic	Negative	0.0000	0.8900	0.0003	
Courtney et al. (2013)		Echino.	Echinometra viridis	%/hr	linear +	Positive	0.0244	0.9500	2.3854	
Courtney et al. (2015)		Echino.	Echinometra viridis	%	linear +	Positive	0.0009	0.3500	13.0388	
Ries et al. (2009)		Echino.	Eucidaris tribuloides	mmol/g/hr	threshold +	Positive	0.0000	0.8400	0.0004	
Keul et al. (2013)		Foram.	Ammonia sp.	mmol/#/hr	linear -	Negative	0.0277	0.0200	0.0000	
Prazeres et al. (2015)		Foram.	Amphistegina lessonii	%/hr	parabolic	Negative	0.0008	0.3900	0.0010	
Kisakurek et al. (2011)		Foram.	Globigerinella siphonifera	mmol/hr	neutral	Neutral				
Kisakurek et al. (2011)		Foram.	Globigerinoides ruber	mmol/#/hr	neutral	Neutral				
Reymond et al. (2013)	179	Foram.	Marginopora rossi	%/hr	linear +	Positive	0.0000	0.1900	0.0090	
Uthicke-Fabricius et al. (2012)	47	Foram.	Marginopora vertebralis	mmol/g/hr	threshold +	Positive	0.0000	0.4000	0.0004	
Sinutok et al. (2011)	16	Foram.	Marginopora vertebralis	mmol/hr	neutral	Neutral				
Prazeres et al. (2015)	32	Foram.	Marginopora vertebralis	%/hr	linear -	Negative	0.0006	0.3300	0.0005	
Manno et al. (2012)	192	Foram.	Neogloboquadrina pachyderma	mmol/#/hr	linear +	Positive	0.0000	0.7100	0.0000	
Oron et al. (2020)	96	Foram.	Operculina ammonoides	mmol/g/hr	linear -	Negative	0.0031	0.0900	0.0017	
Manriquez et al. (2016)	74	Gastropod	Concholepas concholepas	mmol/g/hr	linear +	Positive	0.0000	0.2400	0.0009	
Noisette et al. (2016),										
Ries et al. (2009)	173	Gastropod	Crepidula fornicata	mmol/g/hr	parabolic	Negative	0.0000	0.2100	0.0028	
Garilli et al. (2015) 68 Gastropod Cyclope neritea		mmol/g/hr	linear -	Negative	0.0020	0.1400	0.0037			
Ries et al. (2009)	42	Gastropod	Littorina littorea	mmol/g/hr	linear +	Positive	0.0001	0.3400	0.0002	
				µm (shell						
Bibby et al. (2007)	4	Gastropod	Littorina littorea	thickness)	neutral	Neutral				
Garilli et al. (2015)	315	Gastropod	Nassarius corniculus	mmol/g/hr	parabolic	Negative	0.0000	0.2500	0.0064	
Ries et al. (2009)	21	21 Gastropod Strombus alatus		mmol/g/hr	linear +	Positive	0.0000	0.6400	0.0001	
Ries et al. (2009)		Gastropod	Urosalpinx cinerea	mmol/g/hr	linear +	Positive	0.0000	0.5700	0.0001	
Ries et al. (2009)		Mollusks	Argopecten irradians	mmol/g/hr	linear +	Positive	0.0097	0.3500	0.0002	
Ramajo et al. (2016)		Mollusks	Argopecten purpuratus	mmol/g/hr	neutral	Neutral				
Zhang et al. (2011)		Mollusks	Azumapecten farreri	mmol/g/hr	linear +	Positive	0.0106	0.9200	0.0001	
Ong et al. (2017)		Mollusks	Cerastoderma edule	mmol/g/hr	neutral	Neutral				

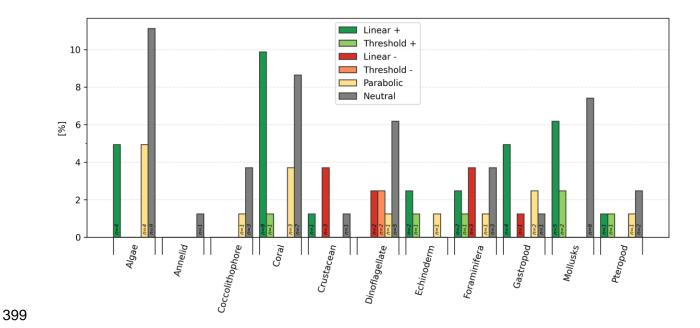
						Pos/Neg/			
Studies	n	Group	Species	Rate unit	Response	Neut	p-value	R ²	RMSE
Sordo et al. (2021)	27	Mollusks	Chamelea gallina	mmol/g/hr	neutral	Neutral			
Gazeau et al. (2007)	nzeau et al. (2007) 20 Mollusks Crassostrea gigas		mmol/g/hr	linear +	Positive	0.0001	0.6100	0.0000	
Ries et al. (2009),									
Waldbusser et al. (2011) 28 Mollusks Crassostrea virginica		mmol/g/hr	mmol/g/hr threshold +		0.0000	0.5600	0.0003		
Ries et al. (2009)	25	Mollusks	Mercenaria mercenaria	mmol/g/hr	threshold +	Positive	0.0000	0.8300	0.0000
Ries et al. (2009) 14 Mollusks Mya arenaria		mmol/g/hr	linear +	Positive	0.0001	0.7300	0.0003		
Ninokawa et al. (2020) 13 Mollusks		Mytilus californianus	s mmol/m²/hr neu		Neutral				
Ries et al. (2009),									
Gazeau et al. (2007)		Mollusks	Mytilus edulis	mmol/g/hr	linear +	Positive	0.0119	0.0700	0.0002
Gazeau et al. (2014)	eau et al. (2014) 11 Mollusks Mytilus galloprovincialis		Mytilus galloprovincialis	mmol/g/hr	neutral	Neutral			
Cameron et al. (2019)	30	Mollusks	Pecten maximus	mmol/g/hr	neutral	Neutral			
				mm (shell					
Comeau et al. (2010b)	5	Pteropod	Cavolinia inflexa	length)	neutral	Neutral			
Comeau et al. (2009, 2010a)	12	Pteropod	Limacina helicina	mmol/g/hr	linear +	Positive	0.0000	0.8500	0.0001
				mm (shell					
Lischka et al. (2011, 2012)	119	Pteropod	Limacina helicina	length)	threshold +	Positive	0.0003	0.1300	0.1303
Bednarsek (2021a),				μm (shell					
Mekkes et al. (2021)	117	Pteropod	Limacina helicina	thickness)	parabolic	Negative	0.0000	0.1800	0.0038
				mm (shell					
Lischka et al. (2012)	28	Pteropod	Limacina retroversa	length)	neutral	Neutral			

386

*Barcelos-Ramos et al. (2010), Fiorini et al. (2011), Iglesias-Rodriguez et al. (2008), Richier et al. (2011), Sciandra et al. (2003), Stoll et al. (2012), 387 Gafar et al. (2018), Bach et al. (2011), Sett et al. (2014).

388 Within each of the 11 functional groups, several categories of calcification response occur within 389 each functional group, with the most varied being the group of dinoflagellates and foraminifera, 390 both showing 4 or 5 different categories of calcification responses (Fig. 5). Of the six types of 391 responses of calcification rate vs. TA:DIC, 28% were linear positive (N=27), 9% linear negative 392 (N=9), 6% threshold positive (N=6), 2% threshold negative (N=2), 15% parabolic (N=14) and 393 40% neutral (N=38).

394 Such responses could be further summed up into positive (linear and threshold positive), negative 395 (linear and threshold negative, parabolic) and neutral responses (Fig. 6) when generalized for 396 calcification rate against TA:DIC ratio. A summary of responses includes 34.4% positive (N=33), 397 26.0% negative (N=25), while 39.6% show a neutral response (N=38).



400 Figure 5: Categories of calcification rate responses and percentage (%) response across eleven
401 groups (calcifying algae, annelids, coccolithophores, corals, crustaceans, dinoflagellate,
402 echinoderms, foraminifera, gastropods, mollusks, pteropods). The number on the bar indicates
403 the number of studies of species included.



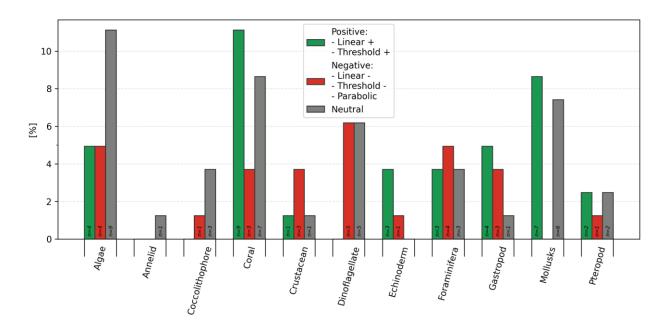


Figure 6: Summary of percentage (%) responses in calcification rates as positive (linear and

threshold positive), negative (linear and threshold negative, parabolic) and neutral across

408 eleven groups (calcifying algae, annelids, coccolithophores, corals, crustaceans, dinoflagellate,

409 echinoderms, foraminifera, gastropods, mollusks, pteropods). The number on the bar indicates
410 the number of studies with species included.

411

412 3.3 Evaluation of the responses to NaOH/Na₂CO₃ addition

413 Upon added TA, the calcification rate in positive responders will increase, either in a linear or 414 threshold positive response, where calcification plateaus, with the concentration being dependent 415 on the species-specific rate of response (Fig. 2; Supplemental Fig. 4). The negative responders 416 (linear or threshold negative and parabolic) will be negatively impacted as follows: first, for the 417 linear negative responders, addition of the Na₂CO₃ will linearly decrease calcification rate, but 418 there is no associated threshold to it; second, for the threshold negative responders, calcification 419 rate will decline in an exponential way until reaching a TA:DIC value where the response plateaus; 420 and third, for the parabolic responders, the calcification rate will initially increase until reaching a 421 certain TA:DIC threshold upon which calcification starts declining. The TA:DIC thresholds for 422 negative responders are species-specific (Table 2; Supplemental Table 4).

423 3.4 Threshold values indicative of negative biological response to OAE

The TA:DIC biological thresholds in Table 2 are determined by the amount of NaOH addition required to reduce calcification rate by a half (see Supplemental Table 4 for Na₂CO₃ thresholds). These thresholds demonstrate the range of carbonate chemistry conditions over which the negative biological effects of OAE deployment might occur and are shown alongside the corresponding pH_T and Ω_{ar} . Uncertainties are higher for the experimental studies where the experimental temperature and salinity ranges were high (see Supplemental Table 5), seeing as we use the average for each species rate group to compute the baseline and thresholds.

For the negative responders, TA:DIC thresholds range from 1.13 to 1.74. The majority of species
have reached their thresholds by an addition of 500 µmol/kg NaOH, though for 3 species a NaOH
addition of more than 500 µmol/kg is required to cross the thresholds in the TA:DIC range of 1.39
to 1.74. *Crepidula fornicata* (gastropod), *Neogoniolithon sp.* (algae), *Homarus americanus*(crustacean) and *Oculina arbuscula* (coral) reach their thresholds by 100 µmol/kg addition of
NaOH, indicating they are more sensitive to alkalinity addition. Foraminifera, dinoflagellates and

437 coccolithophores generally require higher concentrations of NaOH to reach their thresholds, with
438 the linear negative responder *Ammonia sp.* of the foraminifera group requiring 1400 µmol/kg to
439 reduce calcification rate in half.

For some negative responders (*Arbacia punctulata, Nassarius corniculus, Penaeus plebejus, Callinectes sapidus, Cyclope neritea,* and *Symbiodinium sp.*), the baseline from which NaOH addition occurs was outside of the range of the experimental data and very close to a calcification rate of 0. These were omitted from Table 2 since our defined threshold does not give an accurate representation of their sensitivity to alkalinity addition. *Limacina helicina* was also omitted since the indicator of calcification (shell thickness) was not an actual rate.

446 **Table 2:** Studies with negative responders (linear and threshold negative, parabolic) with 447 demonstrated TA:DIC thresholds, indicating the amount of NaOH needed to halve the current 448 calcification rate (i.e. at the baseline). The value for TA:DIC threshold is used to determine the 449 pH_T and Ω_{ar} (at average temperature and average salinity per species). See Supplemental Table 4 450 for Na₂CO₃ thresholds.

Studies	Group	Species	Temp (°C)	Salini ty	Rate unit	Thres hold	TA additio n	рН _Т at threshold	ΔpH _T from baselin e	$\Omega_{ m ar}$ at threshold	Exposure time
Noisette et											
al. (2016),											
Ries et al.		Crepidula									6 months
(2009)	Gastropod	fornicata	15.31	34.33	mmol/g/hr	1.13	50	8.17	0.07	3.77	60 days
Ries et al.		Neogoniolit	25.00	21.70				0.16	0.04	4.05	50 1
(2009)	Algae	hon sp.	25.00	31.70	mmol/g/hr	1.17	50	8.16	0.06	4.87	60 days
Ries et al.	a .	Homarus	25.02	21.00	1/ 0	1.10	100	0.00	0.10	5.40	CO 1
(2009)	Crustacean	americanus	25.02	31.96	mmol/g/hr	1.19	100	8.22	0.12	5.49	60 days
Ries et al. (2009)	Coral	Oculina	25.01	21 (1		1 10	100	8.22	0.12	5.40	(0 1
()	Foraminifer	arbuscula Amphistegin	25.01	51.01	mmol/g/hr	1.19	100	8.22	0.12	5.40	60 days
Prazeres et al. (2015)	a	Ampnistegin a lessonii	24.18	33.46	0/ /han	1.21	150	8.27	0.17	6 10	30 days
Hansen et	a Dinoflagella		24.18	55.40	%0/111	1.21	150	0.27	0.17	0.10	14 d acclimation; 7 days; 14 days exposure;
al. (2007)	te	lineatum	15.00	30.00	#/br	1.18	200	8.38	0.28	5 1 5	22 days stationary growth phase
Sinutok et		Halimeda	15.00	30.00	π/111	1.10	200	0.50	0.28	5.15	22 days stationary growth phase
al. (2011)	Algae	macroloba	27.23	36.27	mmol/g/hr	1.26	200	8.30	0.20	7 38	2 weeks acclimation, 2 weeks incubation
Comeau et	riigue	Sporolithon	21.25	50.27	innioi/g/m	1.20	200	0.50	0.20	7.50	2 weeks acconnation, 2 weeks incubation
al. (2019)	Algae	durum	20.60	35 87	mmol/m²/hr	1.22	200	8.32	0.22	631	27 weeks
Van de	8							0.01			
Waal et al.	Dinoflagella	Thoracosph									21 days acclimation, 8 days experiment =
(2013)	te	aera heimii	15.00	34.00	mmol/hr	1.23	300	8.46	0.36		total of >10 generations
Oron et al.	Foraminifer	Operculina									
(2020)	a	ammonoides	25.00	37.00	mmol/g/hr	1.33	400	8.46	0.36	9.44	65 - 120 hours
Prazeres et	Foraminifer	Marginopor									
al. (2015)	a	a vertebralis	24.18	33.46	%/hr	1.33	450	8.53	0.43	9.78	30 days
Camp et al.											
(2017),											
Comeau et		Acropora									N7A (natural conditions)
al. (2013)	Coral	pulchra	27.30	36.27	mmol/m²/hr	1.38	500	8.52	0.42	11.05	2 weeks acclimation; 2 weeks incubation
											14 d acclimation; 7 days acclimation to
Hansen et	0	Heterocapsa	15.00	20.00		1.00	500	0.55	0.74		experimental conditions; 14 days exposure;
al. (2007)	te	triquetra	15.00	30.00	#/hr	1.30	500	8.66	0.56	8.81	22 days stationary growth phase
Comeau et	Genel	Pavona	27.22	26.20		1.20	500	0.50	0.42	11.02	
al. (2013)	Coral	cactus	27.23	30.28	mmol/m²/hr	1.38	500	8.52	0.42	11.03	2 weeks acclimation; 2 weeks incubation
Hansen et	Dinoflagella	Prorocentru									14 d acclimation; 7 days acclimation to experimental conditions; 14 days exposure ;
al. (2007)	te	Prorocentru m minimum	15.00	30.00	#/br	1.39	700	8.81	0.71	11.25	22 days stationary growth phase
ai. (2007)	Coccolithop	m minimum Emiliania	15.00	50.00	π/1II	1.39	700	0.01	0.71	11.55	22 days stationary growin phase
*	hore	huxleyi	17.30	35.12	mmol/#/hr	1.46	850	8.83	0.73	13.65	**
Keul et al.	Foraminifer	Ammonia	17.50	55.12	111101/#/111	1.40	0.50	0.03	0.75	15.05	
(2013)	a	sp.	26.00	32 75	mmol/#/hr	1.74	1400	9.11	1.01	22.27	59-96 days of culturing

451 *Barcelos-Ramos et al. (2010), Fiorini et al. (2011), Iglesias-Rodriguez et al. (2008), Richier et al. (2011), Sciandra et al. (2003), Stoll et al. (2012),
 452 Gafar et al. (2018), Bach et al. (2011), Sett et al. (2014).

453 **26hrs, Acclimation for 7 generations, experiment/sampling for 2-3 generations, n/a, 8 days, 16 days, Acclimation for 12 generations, Pre 454 acclimation for 8-12 generations, 9 generations, Acclimated for at ~7 generations (5-15 days)

455 3.5 Regulatory pH_T 9 threshold

We also compute how much NaOH and Na₂CO₃ needs to be added before reaching a pH_T threshold
of 9, as per the US Environmental Protection Agency's rule for waste water not exceeding a pH_T
of 9 when entering the coastal ocean (NPDES manual, 2010). This amount averages at 1200
µmol/kg of NaOH and 4700 µmol/kg of Na₂CO₃ for most of the examined species. For some
species (*Amphibalanus improvisus, Neogloboquadrina pachyderma, Limacina helicina, Limacina retroversa, Lophelia pertusa,* and *Semibalanus balanoides*), their threshold was reached below
1000 µmol/kg NaOH and 3000 µmol/kg Na2CO3, with *Amphibalanus improvisus* reaching a

463 threshold at 750 μmol/kg NaOH and 2250 μmol/kg Na₂CO₃.

464 **3.6** Global and regional carbonate chemistry data coverage based on GLODAP datasets

The compilation of chemical observational data (pH, Ω_{ar} , TA, DIC) was done for the GLODAP data across the regional ocean and global scales to determine the range of Ω_{ar} , TA and DIC (as represented by the TA:DIC ratio) and TA:DIC vs Ω_{ar} correlation down to the depths averaged over 200 m. This allowed us to apply the thresholds even for the regions for which we do not have sufficient or reliable data or experimental coverage, making the inferences about the OAE impact even in those regions.

471 Here, we focused on showing the results ranging over the 0-50m because this covers most of the 472 biological habitat for examined species and it is where the OAE enhancement would induce the greatest changes. Over the 0–50 m depth, Ω_{ar} ranges from 0.2 to 5 and TA:DIC ranges from 0.1 to 473 474 1.25 and both parameters are correlated across all the regions, as demonstrated by the fitted secondorder polynomial regressions, with R² of 0.96 or higher, and all the correlations being significant 475 476 (Fig. 7), with regional specific relationships not impacting the fit. All the correlation parameters 477 are presented in Supplemental Table 4. Similar fits were found at different depths. The conditions 478 in the higher latitude regions are located at the lower range of Ω_{ar} vs TA:DIC, while the conditions 479 in the low latitudes and temperate regions are at the upper range, with the highest values present in the central Atlantic and Pacific region. Such strong correlation as observed for Ω_{ar} vs TA:DIC 480 does not exist with pH, regardless of the depth interval examined. While the correlations are still 481 482 significant, they are broadly distributed and represented over a shorter TA:DIC range, with 483 significantly lower goodness of fit (Supplemental Fig. 4), with the correlations being highly 484 regionally dependent due to pH and temperature co-linearity. Because of this, all further biological 485 analyses are only done using the Ω_{ar} vs TA:DIC ratio.

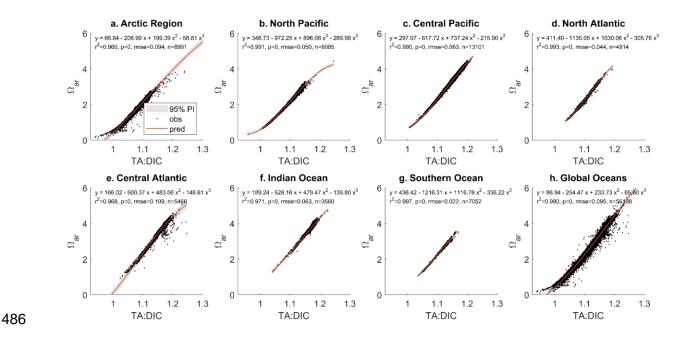
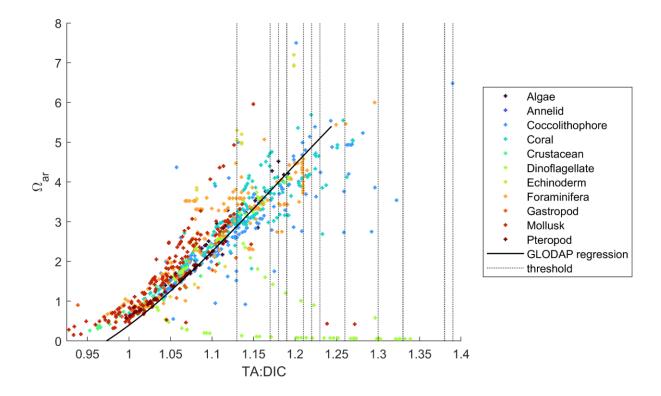


Figure 7: The range of observed Ω_{ar} , TA and DIC values (as represented by the TA:DIC ratio) values and the relationship with the best fitted curve between Ω_{ar} vs TA:DIC across regional (ag) and global (h) scales based on the observational GLODAP dataset averaged over the 0-50 m depth range.

491 **3.7 TA:DIC vs** Ω_{ar} for experimental data and GLODAP

We compared the ranges of TA:DIC and Ω_{ar} of biological experimental data with field biogeochemical data (GLODAP) to examine if similar range of conditions and TA:DIC correlations are applicable over a broader, global dataset. For this, we plotted Ω_{ar} vs TA:DIC along with the GLODAP regression line for Ω_{ar} vs TA:DIC (Fig. 8). For each TA and DIC datapoint, the corresponding salinity and temperature specific values for that data point were used to compute Ω_{ar} . We show the similarity in the conditions, which gives the validity of our experimentally derived thresholds to be extrapolated within the global GLODAP dataset.

Figure 8 also shows that various biological groups are clustered around specific TA:DIC ratios, for example, mollusks, coral and coccolithophores are represented on the lower, mid, and higher TA:DIC spectra, respectively, while dinoflagellates are randomly scattered off the TA:DIC line. This indicates that there is a general lack of data distribution in the upper ranges of TA:DIC ratio, especially for the groups that are lying at the lower and mid end of the TA:DIC ratio spectra. 504 Plotting biological data from the OA datasets against the regional and global TA:DIC gradient 505 derived from GLODAP (Fig. 7), we also observed that experimental data ranges were not always 506 consistent with natural conditions, for example, having a lower Ω_{ar} at a higher TA:DIC ratio.



507

508 Figure 8: Ω_{ar} values from experimental biological studies for eleven investigated functional

509 groups (see legend) plotted against TA:DIC, with the latter being computed using experimental

- 510 *TA and DIC. The black line represents the regression line of TA:DIC and* Ω_{ar} *data from the*
- 511 GLODAP dataset (covering 0-50m depth). See Supplemental Fig. 5 for GLODAP Ω_{ar} vs TA:DIC,
- 512 from which the black regression line shown here is derived. The vertical dotted lines represent

513 *the thresholds shown in Table 2.*

514 **4. Discussion**

515 OAE is a quickly developing strategy that is in the field-testing phase despite extremely limited 516 understanding of the sequestration potential, biological implications and environmental concerns. 517 Hence, gaining insights of potential risks for the biological species and communities is essential 518 and timely. In retrospect, it took decades for the OA research community to get a more accurate 519 and comprehensive understanding leading to predictions of biological responses to OA (Riebesell 520 and Gattuso, 2015). Without a very clear conceptual strategy for the OAE testing, the research 521 community might also need years to decades before OAE-related implications are 522 comprehensively understood. Consequently, there is an essential need to develop an assessment 523 framework of predictive responses and testing strategies that will assist in OAE scaling and risk 524 avoidance. This paper aims at developing such an assessment, where calcification responses 525 against TA:DIC are categorized per species. We propose to use the TA:DIC ratio in the biological 526 studies reporting OAE results, as we believe it simplifies the system and makes it easier to use and 527 translate the carbonate chemistry in the experimental setting. Such a TA:DIC ratio allows to 528 ultimately standardize the biogeochemical and biological data and is useful for easier comparisons 529 among the conducted experiments.

530 4.1. Identified strengths and limitations of the synthesis approach based on OA studies

531 Prior to conducting this study, several drawbacks were identified that could potentially limit such 532 a synthesis work: first, an insufficient amount of data at the upper range of carbonate chemistry 533 conditions (high pH, high Ω_{ar}); second, experimental data under conditions with no relevance to 534 natural settings (Fig. 8); and third, an insufficient number of validation studies under high TA 535 conditions to validate the results of this synthesis. To overcome the first two limitations, the 536 decision was made to combine multiple OA datasets for a single species with the aim to achieve a 537 greater range in carbonate chemistry conditions, including higher pH, Ω_{ar} experimental values, 538 which should reduce the uncertainty of the predictions. However, combining raw data on species 539 calcification rate proved to be more challenging because even across the same species the reporting 540 of the calcification rates was highly variable. The use of different measuring approaches of calcification rates while conducting OA studies generated data with divergent units that do not 541 542 allow for the intercomparison of data and results. As different studies for a single species could 543 not be combined, we chose to increase the number of studies and thus, the number of examined 544 species. Based on the response categories from the OA studies (Ries et al., 2009), our hypothesis 545 was that OAE will elucidate the same categories of responses, i.e. positive, negative and neutral. 546 Within each of the groups examined, multiple categories of predicted calcification response were 547 found. In this way, we demonstrated that it was possible to develop a useful framework for 548 assessing and predicting species-specific OAE responses that can delineate different responders, 549 identify species with greater OAE sensitivity and determine the thresholds where such negative

550 responses could happen.

551 4.2 Synthesizing biological response under OAE identifies positive and negative responders

The responses were summarized across three emerging groups of responses: positive, negative, and neutral (Fig. 6). We observe species-specific variability at the species level, which is related to various calcification mechanisms across the observed groups. The greatest variability upon NaOH addition within each group in calcification rate was evident in corals, dinoflagellates, foraminifera, gastropods and pteropods, where four to five different categories of responses were found.

Positive responders (34%) show an increased calcification rate upon alkalinity addition, observed within all functional groups besides annelids, coccolithophores and dinoflagellates. Corals mostly have positive and neutral responses, suggesting that coral species would not be negatively impacted during OAE field trials. This mostly positive response is validated by increased coral calcification, shown for two coral species of *Acropora* and *Siderastre* in experiments conducted by Palmer et al. (2022).

564 The metrics to evaluate the sensitivity of calcification rate for the negative responders (negative 565 linear and threshold) to alkalinity addition was based on the amount of alkalinity addition required 566 to halve the current calcification rate (Fig. 3; Tables 1, 2). The most negative responses were found 567 in dinoflagellates (6% of all species), algae and foraminifera (both 5% of all species). However, 568 these numbers are affected by the difference in data coverage per functional group. When 569 comparing the ratio of negative to positive and neutral responses, crustaceans and dinoflagellates 570 are expected to be most negatively affected. As such, these groups are one of the priorities for the 571 future OAE experimental work to determine at which TA:DIC negative response happens. 572 Dinoflagellates demonstrate negative response in 5 cases, 5 neutral responses and 0 positive (see 573 Table 1; Supplemental Fig. 4). The reason for negative response to OAE in this group is related to 574 the fact that their growth gets limited at higher pH, with further carbon limitation playing a role at 575 very high pH levels and low DIC concentration (Hansen et al., 2002; 2007). On the other hand, 576 crustaceans only demonstrated positive response in one study (Pansch et al., 2014), while 577 remaining results predict either negative or neutral response. While crustaceans are effective in 578 retaining homeostasis at lower pH, they might be less so at higher pH, which was shown in the

579 OA experiments by Ries et al. (2009) for three crustacean species (*Callinectes sapidus*, *Homarus*) 580 americanus, Penaeus plebejus), confirmed in the OAE study by Cripps et al. (2013) in Carcinus 581 maenas. While studies are still lacking, physiological acid-base regulation at higher pH is 582 associated with higher costs (Cripps et al., 2013). Crustaceans show a disrupted acid-base balance, 583 evident through the increase in hemolymph pH, K⁺, Na⁺ ions and osmolality, coupled with a decrease in extracellular pCO_2 and HCO_3^- , indicative of respiratory alkalosis (Truchot, 584 585 1984;1986). This is often associated with hyperventilation, the aim of which is to flush out the 586 hemolymph CO_2 to increase the affinity of oxygen uptake. However, while this might be a 587 temporary physiological relief it also implies energetic costs, potentially also for calcification.

For the neutral responders or groups with no significant correlation between calcification rates and TA:DIC, it is somewhat uncertain to predict if such responses will be retained under OAE. While parabolic responders show a physiologically understandable parabolic type of dose-response, positioning the TA:DIC values where the thresholds occur is also highly species-specific and potentially uncertain, meaning that it might depend on other environmental factors.

593 With respect to the coccolithophores, we note that this was the only group where data compilation 594 on calcification rate across the group was possible because the OA studies were conducted in a 595 more uniform way, using similar approaches, and reporting the result in the same units. When data 596 for E. huxleyi across the comparable studies was compiled (Barcelos-Ramos et al., 2010; Fiorini 597 et al., 2011; Iglesias-Rodrigues et al., 2008; Sciandra et al., 2003; Stoll et al., 2012; Richier et al., 598 2011), a significant parabolic response was obtained (Table 1), although the goodness of fit was 599 fairly low ($R^2=0.16$). Despite lower R^2 , we decided to use the compiled dataset because of the 600 increased statistical power. The parabolic response obtained aligns with Langer et al. (2006) and 601 also with the parabolic type responses found in the synthesis studies by Paul and Bach (2020) and 602 Bach et al. (2015). The threshold indicates the mechanisms of coccolithophore growth that are 603 driven by CO_2 , which is shown to decline with alkalinity addition. The threshold based on all 604 studies for *E. huxleyi* combined was positioned at a TA:DIC of 1.46 ($\Omega_{ar} = 13.65$, see Table 2), 605 which would be triggered at 850 µmol/kg of added NaOH and at a pCO₂ of 60 µatm. 606 Comparatively with the phytoplanktonic diatoms, such growth limitation is predicted at a pCO₂ 607 amount at 100 µatm (Riebesell et al., 1993). It is important to note that when these studies were 608 analyzed individually, a mixture of different responses was observed. We emphasize the variability

609 within the coccolithophore responses, which are species-specific and inherently related to the 610 strain adaptation to their innate regional settings and dependent on a variety of other factors (Bach 611 et al., 2015; Gafar and Schultz, 2018), including the longevity of the species, the experimental 612 settings used in the study (e.g. nutrient-replete vs nutrient deficient conditions) and the presence 613 or absence of (un)suitable light conditions. Interestingly, for all the coccolithophore species other 614 than E. huxleyi, responses were neutral. For validation purposes, the results of our study could not 615 be compared, either because the calcification rates were not studied or the calcification units were 616 not comparable (e.g. Diner et al., 2015).

617 **4.3 Parameters impacting derivation of thresholds and their application**

618 We developed a set of species-specific thresholds in this study, with demonstrated application 619 across the global Ω_{ar} vs TA:DIC conditions (Table 2; Fig. 8). The range of alkalinity additions to 620 result in a threshold of 50% decline in calcification rate varied significantly between the species 621 and the type of response. The TA:DIC thresholds upon TA application ranged between 50 to 1400 622 µmol/kg of NaOH addition and 2250 to 6500 µmol/kg of Na₂CO₃ addition, and the pH_T 9 623 thresholds averaged at 1200 µmol/kg of NaOH and 4700 µmol/kg of Na₂CO₃for all species. 624 However, there are many parameters that impact threshold derivation and application, which we 625 discuss in greater detail.

First, we note that differences in experimental conditions for different species make it difficult to directly compare different species' thresholds among each other. Instead, they are intended to delineate sensitivity to alkalinity addition of individual species at given experimental conditions. In the case that the lab experimental conditions mimic species' natural habitat, this thresholdrelated sensitivity can be extrapolated to their natural habitats.

Second, we emphasize that the threshold application should not only consider the magnitude of NaOH added, but also the duration or exposure time of the experimental study. As such, when applying the thresholds to respective model outputs or observation data, both duration and exposure time should be considered. For all the derived thresholds, we have added duration exposure information to Table 2. Additional parameters that need to be included when applying these thresholds are related to local temperature and salinity. The extracted threshold values are calculated with the temperature and salinity from the experimental conditions, which means that this threshold should not be applied to very different conditions without adjusting for salinity andtemperature.

640 Third, we assumed global surface ocean conditions to be standardized at a pCO_2 of 425 ppm and 641 a pH_T of 8.1 as a control point for OAE compound additions. However, we note that in different 642 habitats, pH_T 8.1 may not represent the baseline from where OAE should be considered adding, 643 because the average pH might be different. This means that the amount of TA required to reach a 644 certain threshold could vary and is dependent on the baseline carbonate chemistry at the site of deployment and its variability. This is especially relevant in habitats with a lower baseline pH, 645 646 where more TA would need to be added for the threshold to be reached, meaning less negative 647 biological implications.

648

In addition, physical parameters of importance are related to the dilution effect, mixing, retention capacity, as well as the rate of the equilibration effects of the air-sea CO_2 uptake (Ferderer et al., 2022; He and Tyka, 2023; Schulz et al., 2023; Wang et al., 2023), because they determine relevant exposure duration and the variability of carbonate chemistry parameters across spatial and vertical depths. Therefore, to obtain the most accurate and regionally applicable threshold for the species of interest, it is recommended that the baseline for OAE additions be determined based on local conditions.

656 Lastly, if similar conditions as induced by the OAE field trial are present in the habitats that species 657 inhabit, it is more likely that the species might be pre-adapted to such conditions. However, if 658 species have not been exposed to such conditions, OAE might induce rapid change in conditions 659 and species exposure, which could be more challenging for the species. As such, it is worth 660 considering that OAE deployments could be, at least for the most sensitive species, carried out not 661 as a single high dosage deployment, but rather as a more continuous, lower dosage application. 662 This would eliminate the swings and maxima in conditions, while also allowing more time for 663 species acclimation or migration during the initial injection of the OAE deployment. Ultimately, 664 it is the combination of all these factors that creates baseline exposure conditions that are relevant in the context of biological outcomes (Wang et al., 2023). 665

666 4.4 Direction of laboratory OAE experiments should change to incorporate field conditions

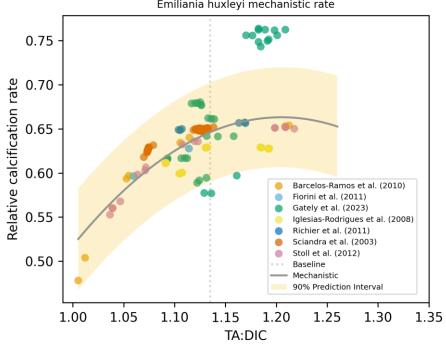
667 The lab OAE experiments that are being conducted right now are done under different conditions 668 than in the field. The former are conducted with the aim of gaining a wide-ranging empirical 669 response, which implies high treatment levels of OAE additions. However, biogeochemical model 670 outputs show that OAE-related concentrations at the injection site are high for a short-time, while 671 the realistic field dosing upon rapid dilution due to mixing is low. Wang et al. (2023) reported that 672 the nearfield maxima in the respective investigation area of the Bering Sea is to increase TA by 673 about 10 µmol/kg in the nearfield and by about 1 µmol/kg of NaOH in the farfield region. As such, 674 we should be more concerned about the threshold of exceedance occurring at the low NaOH 675 dosing, rather than at high NaOH additions, because these are more realistic and point to the most 676 sensitive species. As a result, we explicitly emphasize the importance of including much lower 677 additions of TA in the experimental treatment levels to better support biological understanding and 678 OAE application in the field. In addition, prior to the lab experiments it would be important to identify what type of response is predicted in the experimental species. This is especially pertinent 679 680 for the groups for which OA experimental data is limited and skewed towards the lowest TA:DIC 681 ratio (Fig. 8; Supplemental Fig. 4).

682 What is needed urgently for the safe biological field trial experiments is a set of protocols that are 683 species-, habitat- and local conditions- specific, which would allow for comprehensive and 684 comparative risk analyses and threshold determination. As part of this, we also need to develop 685 regionally specific indicators for biological monitoring. Ideally, such biological and environmental 686 risk monitoring and assessment would be accompanied by the application of the physical mixing 687 models with site-specific biogeochemical processes (Ho et al., 2023; Fennel et al., 2023) that can 688 predict the maximum expected TA increase in the nearfield and farfield regions of the study site, 689 representing a more realistic exposure and better informing further experimental work.

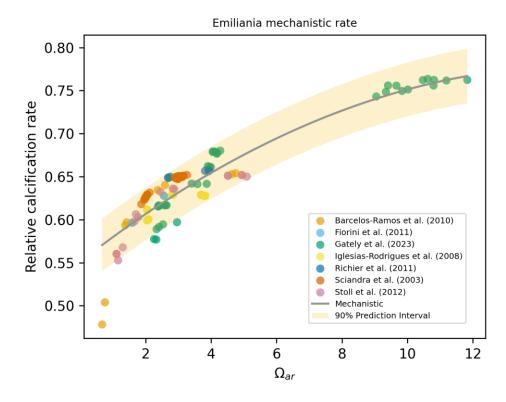
690 4.5 Validating OAE responses based on the mechanistically-derived calcification

This study establishes the predictions of responses that relied upon empirical experimental studies. A good alternative to validating the predicted responses is to use species-specific mechanistic responses, a more accurate representation of responses compared to empirical studies. Here, we conducted a subset synthesis study for the two species of coccolithophores, using the results from this study and compared it to the literature-derived mechanistic responses where the responses are described with a different set of carbonate chemistry parameters. We wanted to determine to what 697 extent mechanistic relationships can contribute to improved, i.e. more accurate and certain, OAE 698 predictions.

699 For *Emiliania huxleyi*, we used experimental TA and DIC data to calculate the [HCO₃⁻], [H⁺] and 700 $[CO_2]$ concentrations to be able to use the mechanistic rate equation from Bach et al. (2015). We 701 calculated and plotted the rate derived via mechanistic approach and applied linear, polynomial 702 (second-order) and exponential regressions and chose the best fit based on the lowest p-value, 703 using the same method as for our experimental calcification rate data regressions. Like the mechanistic rate regression based on three carbonate chemistry parameters was a parabolic fit 704 705 (Bach et al., 2015), we also obtained the same fit using the experimental calcification rate data (see 706 Fig. 9). However, when using the same approach for another coccolithophore species Calcidiscus 707 *leptoporus* (Bach et al., 2015), our best fit did not align with the proposed mechanistic response; 708 instead, a non-significant relationship was obtained using experimental data (Supplemental Fig. 709 5). Such comparisons reveal species-specific relationships are likely dependent on a lot of 710 parameters, with one equation alone not being operable among different species from different 711 experiments or over varied regional settings.



Emiliania huxleyi mechanistic rate



713

Figure 9: *Mechanistic rate equation and parameters* (a = 9.56e-1, b = 7.04e-4 mol/kg, c = 2.1e6kg/mol, d = 8.27e6 kg/mol) taken from Bach et al. (2015) and fitted using experimental data for

E. huxleyi (used data from the studies indicated in legend). Shading represents the 90%

717 *prediction interval.*

718

For most of the species, such mechanistic relationships are not available. The 719 substrate-to-720 inhibitor ratio (SIR) (i.e. the bicarbonate ion to hydrogen ion concentration ratio) has often been 721 used to describe a calcification relationship based on a single parameter relationship. To 722 if species rate group responses based on experimental data using TA:DIC vs calcification see 723 rate could reproduce SIR relationships (TA:DIC vs SIR), we computed and plotted the 724 SIR ratio. This included calculating the bicarbonate and hydrogen ion concentrations in CO2SYS 725 using the experimental TA and DIC, for the mollusk, coral and coccolithophore groups, and 726 applying a best-fit regression model. We categorized these responses (using the categories 727 shown in Figure 2) and compared these SIR regressions to the respective best-fit regressions based 728 on the calcification rate responses from the experiments (shown in Table 1).

730 We found large differences between our calcification rate responses (based on TA:DIC vs 731 calcification) and the SIR-proposed mechanisms (Supplemental Fig. 6). For most of the 732 coccolithophore groups, the experimental rate regressions cannot be explained using SIR 733 mechanisms (i.e. the responses are different). Only in the case of *Calcidiscus leptoporus*, the 734 experimental and mechanistic responses are the same (neutral). For mollusks, a third of the 735 mechanistic rate regressions based on the SIR agreed with the experimental calcification rate 736 regressions. The other two-thirds did not agree, especially for the studies with experimental 737 conditions of $\Omega_{ar} > 1$. For corals, the majority of the coral species (N=14) were classified as having 738 a linear positive mechanistic relationship when using SIR relationships. When comparing this to 739 our experimental rate regressions, we only found agreements between the experimental and 740 mechanistic regressions in 6 out of 18 species. It seems that SIR is a less common principle of 741 calcification and cannot be applied across a variety of species. It is likely that SIR might 742 insufficiently explain the multitude of biological processes involved in the calcification (e.g. how 743 carbon is provisioned or the ability to regulate calcifying fluid pH). Based on these results, the 744 general consensus is that the SIR ratio might actually tend to oversimplify species' calcification 745 rate responses. Ninokawa et al. (2024) and Li et al. (2023) emphasized that using only one 746 parameter to describe the calcification process is insufficient and strongly recommended using at 747 least two parameters for more accurate calcification predictions. Our findings agree with 748 Ninokawa et al. (2024), for example, we observe that using SIR relationships to successfully 749 describe calcification was limited to only a few species and that there are no generalizable patterns 750 that could be applicable across multiple groups.

751

752 Mechanistic models can offer better insights into calcification responses for some species, 753 especially when multiple environmental factors are accounted for, but they are not generally 754 applicable across taxa. Species-specific responses are influenced by unique biological and 755 physiological factors, which can lead to significant deviations between mechanistic and empirical 756 predictions. Therefore, mechanistic approaches will only provide valuable frameworks for species 757 with well-understood calcification processes. However, for many species covered in this study the 758 calcification process is not well-understood. By comparing mechanistic studies with experimental 759 data, we hoped to validate the predictive results of our experimental studies. This clearly delineates 760 a major gap in the mechanistic understanding of calcification so far, the lack of which significantly

761 limits our ability of ecological and biogeochemical predictions to OAE. As such, more research is 762 urgently needed on broader mechanistic understanding of calcification across different species, 763 and additionally, one parameter calcification processes should be replaced with more accurate and 764 comprehensive methods using two or three parameters.

4.6 Unknowns about ecological and biogeochemical implications call for the precautionaryapproach

767 The value of calcification as the response proxy is indicative of organismal fitness, which directly 768 relates to OAE effects as harmful or beneficial for the species. From an ecological perspective, a 769 total of 26.0% negative responders demonstrates a potential for negative implications. In addition, 770 we note that this study did not include diatoms in the analyses, which are predicted to be negatively 771 impacted by carbonate-based OAE (Ferderer et al., 2022), leading to possible community-based 772 ecological shifts (Bach et al., 2019). The possibility of the ecological shifts should not be neglected 773 given the variety of the positive responders, understudied effects of OAE in non-calcifiers and 774 their relationship with the calcifiers through the grazing impact, and lastly, unknown and highly 775 unpredictable indirect effects. In addition, the inferences on the neutral responders should also 776 remain cautious.

777 From a biogeochemical perspective, it is reasonable to infer that OAE will introduce changes in 778 calcification rate across species, potentially resulting in changing the carbon export or carbonate 779 counter pump. Species-specific responses in major carbonate producers, i.e. coccolithophores, 780 foraminifera and pteropods show both negative and positive response, which could have strong 781 effects on biogeochemical fluxes (Riebesell et al., 2017; Bach et al., 2019). Increased calcification 782 could result in thicker and denser shells, contributing to faster sinking and increased carbonate 783 fluxes, while decreased calcification has the opposite effect. This could potentially induce changes 784 on the subsurface total alkalinity at intermediate and deeper depths in the water column, and 785 dissolution at or near the seafloor (Gehlen et al., 2011) or result in a potential feedback of increased 786 CO₂ flux to the atmosphere (Gattuso et al., 2021). The full scope of ecological and biogeochemical 787 shifts remains a high priority topic for future investigations and until these huge uncertainties are 788 resolved, we should exercise a precautionary principle in considering the next steps of OAE field 789 implementations.

790 **4.7 Potential confounding effects**

791 This study only considered the changes in carbonate chemistry due to the addition of NaOH and 792 Na_2CO_3 . However, other OAE feedstocks contain compounds that could induce biological toxicity 793 due to the presence of trace metals (Ni, Cu, Ca, Si; Bach et al., 2019), as well as potential negative 794 environmental impacts due to secondary precipitation (Hartmann et al., 2022; Moras et al., 2022). 795 This study also did not focus on the sensitivity across different life stages, even though stage-796 specific sensitivities to OAE are expected based on previous OA results. Furthermore, we did 797 include data from experimental lab and field studies that involve multiple stressors in their 798 experimental designs. As such, an additional impact of warming, dissolved oxygen, and light 799 intensity on the OAE-induced responses was not determined, although they could elicit different 800 biological pathways than OAE alone or have additional confounding effects.

The synthesis of the experimental studies always includes implicit biases that are based on the published experimental studies, the range and species used, regional coverage and heterogeneity. Important consideration is the adaptation of the species used in the experimental studies because their calcification optimum might be pre-determined based on their local habitat conditions. Given that the baseline for the OAE-compound addition was chosen at the global current surface pH value, some of the thresholds might actually be lower than expected.

4.8 Applications within the existing governmental regulations and the guiding principle

Our results, especially related to the use of biological thresholds or NaOH dosing, could have 808 809 wider applications, most notably with policy-management governmental regulations. For example, 810 we calculated the amount of alkalinity addition required to reach the pH_T threshold of 9, the 811 maximum pH allowed by the US Environmental Protection Agency's for waste water entering the 812 coastal ocean (see NPDES manual, 2010). To reach this threshold, 1200 µmol/kg of NaOH and 813 4700 µmol/kg of Na₂CO₃ was required on average for all species, with the lowest threshold reached 814 at 750 µmol/kg NaOH and 2250 µmol/kg of Na₂CO₃ addition for Amphibalanus improvisus. This 815 is a very high concentration, and the thresholds for most of the negative responders with identified 816 thresholds (Table 2) will be exceeded far below the regulatory standards of pH_T 9 (Table 2), 817 especially if the exposure occurred over a duration period that matters for calcification and for the 818 organism's physiological status. This case demonstrates discrepancy of the current chemical pH regulation and associated biological effects, where safe biological limits are not considered, and
biological harm might not be prevented. Despite the fact that achieving such a high pH through
NaOH/Na₂CO₃ implementation is unlikely to occur in the field, such regulations currently do not
assure safety space for marine biota and they need to be urgently addressed.

823 5. Conclusions and next steps

824 Sufficient certainty in predicting biological responses reduces the risks and supports safe operating 825 space for OAE implementation and scaling up. Overall, given that almost 60% of examined species 826 showed non-neutral response (either positive or negative), this calls for careful implementation of 827 OAE until the safe operational temporal and spatial scales are identified and OA mitigation 828 measures are established. The goal of this study is to serve as a baseline for prioritizing 829 experimental and field OAE research and assess environmental risks. Such prioritization identifies 830 those species for which experimental work needs to be conducted first. This would involve species 831 with the greatest OAE-related sensitivity (negative responders), species with the greatest 832 uncertainty in response, as well as the species with very strong predicted positive response that 833 could potentially introduce a shift on the community level. In addition, it would also recognize the 834 species for which the existing knowledge is sufficient and there is less immediate need for the 835 OAE experiments. We hope that all presented tools provide guidance for the practicing and 836 regulatory community considering OAE field application within the safe limits.

It is important to emphasize that this study is the first comprehensive synthesis of the effects of OAE. Ongoing updates and additional data would enhance its value, particularly when complemented by further experimental research. Similar datasets on OA exist for various biological parameters, including genetics, physiology, and survival data, as well as for noncalcifying organisms. This availability allows for the exploration of ecological implications and contributes to developing an ecosystem-based predictive risk assessment for OAE.

843 Data availability

No new data were generated during this study; all data were collected from previously published
studies. The compiled data is currently available on request. The Python code used for computing
baselines per species, conceptually adding alkalinity in the form of NaOH and Na2CO3, predicting

847 calcification rate response, visualizing data and computing thresholds is available in the GitHub

- 848 repository at https://github.com/hannavdmortel/OAE_calc_response (last access: 1 November
- 849 2024) and is archived on Zenodo at https://doi.org/10.5281/zenodo.14024442 (van de Mortel,
- 850 2024). PyCO2SYS v1.8.0 (Humphreys et al., 2022) was used to solve for the carbonate system,
- with software available at <u>https://doi.org/10.5281/zenodo.3744275</u> (Humphreys et al., 2023).

852 Author contributions

NB designed and conceptualized the research and wrote the first draft of the paper. HvdM collected and curated data, conducted formal analyses and provided visualization. GP provided the analyses using GLODAP data, and also provided visualizations and formal analyses. MGR has provided formal statistical analyses and visuals. RAF and AD have provided insights, suggestions, and generated discussion about specific parts of the paper. All have contributed to the writing of this paper.

859

860 Competing interests

861 The contact author has declared that none of the authors has any competing interests.

862 Financial support

863 This study was funded by the NOAA NOPP project (mCRD 48914-2023 NOAA to AD, NB, and 864 RAF), with the title: mCDR 2023: Assessing chemical and biological implications of alkalinity 865 enhancement using carbonate salts obtained from captured CO₂ to mitigate negative effects of 866 ocean acidification and enable mCDR). This project also fully supported HvdM who worked on 867 the project as an external consultant. This work was supported by NOAA funding from the 868 Inflation Reduction Act and the Ocean Acidification Program (ROR ID: 100018228). NOAA's 869 Ocean Acidification Program supports this project on behalf of the National Oceanographic 870 Partnership Program (Award #NA23OAR0170516). HvdM has also been supported through the 871 Slovenian research Agency (ARRS J1-2468, N1-0359). This is PMEL contribution number 5621.

872 **References**

873 Agostini, S., Harvey, B. P., Milazzo, M., Wada, S., Kon, K., Floc'h, N., Komatsu, K., Kuroyama,

874 M., Hall-Spencer, J. M: Seawater carbonate chemistry and kelp densities and coral coverages at

three study locations and photosynthesis and calcification of corals measured in the laboratory,

876 2021. PANGAEA, https://doi.org/10.1594/PANGAEA.944056, Supplement to: Agostini, S. et al.:

- Simplification, not "tropicalization", of temperate marine ecosystems under ocean warming and
 acidification. *Global Change Biology*, 27(19), 4771-4784, <u>https://doi.org/10.1111/gcb.15749</u>,
 2021.
- Bach, L. T., Riebesell, U., and Schulz, K. G. (2011): Seawater carbonate chemistry, growth rate
 and PIC and POC production during experiments with *Emiliania huxleyi* (B92/11), 2011 [dataset].
 PANGAEA, https://doi.org/10.1594/PANGAEA.771288, Supplement to: Bach, L. T. et al. (2011):
 Distinguishing between the effects of ocean acidification and ocean carbonation in the
 coccolithophore *Emiliania huxleyi*. Limnology and Oceanography, 56(6), 2040-2050,
 https://doi.org/10.4319/lo.2011.56.6.2040, 2011.
- 886 Bach, L. T., and Mackinder, L. C. M.: Experiment: Dissecting the impact of CO2 and pH on the 887 mechanisms of photosynthesis and calcification in the coccolithophore Emiliania huxleyi. 888 PANGAEA, https://doi.org/10.1594/PANGAEA.830627, Supplement to: Bach, L. T., Mackinder, 889 L. C. M., Schulz, K. G., Wheeler, G., Schroeder, D. C., Brownlee, C., Riebesell, U.: Dissecting 890 the impact of CO2 and pH on the mechanisms of photosynthesis and calcification in the 891 coccolithophore Emiliania huxleyi, New Phytol., 199(1), 121-134. 892 https://doi.org/10.1111/nph.12225, 2013.
- Bach, L.T., Riebesell, U., Gutowska, M.A., Federwisch, L. and Schulz, K.G.: A unifying concept
 of coccolithophore sensitivity to changing carbonate chemistry embedded in an ecological
 framework, Prog. Oceanogr., 135, 125-138, 2015.
- Bach, L.T., Gill, S.J., Rickaby, R.E., Gore, S. and Renforth, P.:, 2019. CO2 removal with enhanced
 weathering and ocean alkalinity enhancement: potential risks and co-benefits for marine pelagic
 ecosystems,. Frontiers in Climate, 1, p.7, 2019.
- Barcelos e Ramos, J., Müller, M. N., and Riebesell, U.: Seawater carbonate chemistry and
 processes during experiments with phytoplankton *Emiliania huxleyi* (strain Bergen 2005), 2010.
 PANGAEA, https://doi.org/10.1594/PANGAEA.736022, Supplement to: Barcelos e Ramos, J. et
 al. : Short-term response of the coccolithophore *Emiliania huxleyi* to an abrupt change in seawater
- 903 carbon dioxide concentrations. Biogeosciences, 7(1), 177-186, https://doi.org/10.5194/bg-7-177904 2010, 2010.

- Bednaršek, N., Feely, R.A., Howes, E.L., Hunt, B.P., Kessouri, F., León, P., Lischka, S., Maas,
 A.E., McLaughlin, K., Nezlin, N.P. and Sutula, M.: Systematic review and meta-analysis toward
 synthesis of thresholds of ocean acidification impacts on calcifying pteropods and interactions with
 warming, Frontiers in Marine Science, 6, 227, 2019.
- 909 Bednaršek, N., Naish, K. A., Feely, R. A., Hauri, C., Kimoto, K., Hermann, A. J., Michel, C.,
- 910 Niemi, A., and Pilcher, D.: Integrated Assessment of Ocean Acidification Risks to Pteropods in
- 911 the Northern High Latitudes: Regional Comparison of Exposure, Sensitivity and Adaptive
 912 Capacity, Frontiers in Marine Science, 8, 671497, https://doi.org/10.3389/FMARS.2021.671497,
- 913 2021a.
- 914 Bednaršek, N., Ambrose, R., Calosi, P., Childers, R.K., Feely, R.A., Litvin, S.Y., Long, W.C.,
- 915 Spicer, J.I., Štrus, J., Taylor, J. and Kessouri, F.: Synthesis of thresholds of ocean acidification
- 916 impacts on decapods, Frontiers in Marine Science, 8, 651102, 2021b.
- Bednaršek, N., Calosi, P., Feely, R.A., Ambrose, R., Byrne, M., Chan, K.Y.K., Dupont, S., PadillaGamiño, J.L., Spicer, J.I., Kessouri, F. and Roethler, M.: Synthesis of thresholds of ocean
 acidification impacts on echinoderms, Frontiers in Marine Science, 8, 602601, 2021c.
- 920 Bibby, R., Cleall-Harding, P., Rundle, S., Widdicombe, S., & Spicer, J. I.: Seawater carbonate 921 chemistry during experiments with Littorina littorea, 2007 [dataset], 2007. PANGAEA, 922 https://doi.org/10.1594/PANGAEA.716837, Supplement to: Bibby, R., Cleall-Harding, P., 923 Rundle, S., Widdicombe, S., & Spicer, J. I.: Ocean acidification disrupts induced defences in the 924 intertidal gastropod Littorina littorea. Biology Letters, 3(6), 699-701. 925 https://doi.org/10.1098/rsbl.2007.0457, 2007.
- 926 Bove, C. B., Whitehead, R. F., and Szmant, A. M.: Seawater carbonate chemistry and
- 927 gastrovascular cavity pH, calcification of *Montastraea cavernosa* and *Duncanopsammia axifuga*.
- 928 PANGAEA, https://doi.org/10.1594/PANGAEA.927310, 2020.
- 929 Brading, P., Warner, M. E., Davey, P., Smith, D. J., Achterberg, E. P., and Suggett, D. J.: Seawater
- 930 carbonate chemistry and growth rate during experiments with phylotypes of Symbiodinium
- 931 (Dinophyceae), 2011. PANGAEA, https://doi.org/10.1594/PANGAEA.771293, Supplement to:
- 932 Brading, P. et al.: Differential effects of ocean acidification on growth and photosynthesis among

933 phylotypes of Symbiodinium (Dinophyceae). Limnol. Oceanogr., 56(3), 927-938,
934 https://doi.org/10.4319/lo.2011.56.3.0927, 2011.

Briggs, A. A., and Carpenter, R. C.: Seawater carbonate chemistry and photosynthesis and
photochemical efficiency of *Porolithon onkodes*. PANGAEA,
https://doi.org/10.1594/PANGAEA.920025, 2019.

Brown, K. T., Mello-Athayde, M. A., Sampayo, E. M., Chai, A., Dove, S., and Barott, K. L.:
Seawater carbonate chemistry and endosymbiont density, photosynthesis and net calcification
rates of reef-building coral *Pocillopora damicornis*. PANGAEA,
https://doi.org/10.1594/PANGAEA.953058, 2022.

Cameron, L. P., Reymond, C. E., Müller-Lundin, F., Westfield, I. T., Grabowski, J. H., Westphal,
H., and Ries, J. B.: Seawater carbonate chemistry and physiology and extrapallial fluid pH,
calcification rate, and condition factor of the king scallop *Pecten maximus*. PANGAEA,
https://doi.org/10.1594/PANGAEA.919939, 2019.

946 Camp, E. F., Nitschke, M. R., Rodolfo-Metalpa, R., Houlbrèque, F., Gardner, S. G., Smith, D. J., 947 Zampighi, M., and Suggett, D. J.: Seawater carbonate chemistry and calcification rate, net 948 photosynthesis and respiration rate of reef-building corals. PANGAEA, 949 https://doi.org/10.1594/PANGAEA.880242, Supplement to: Camp, E.F. et al.: Reef-building 950 thrive within hot-acidified and deoxygenated waters, corals Sci. Rep., 7(1), 951 https://doi.org/10.1038/s41598-017-02383-y, 2017.

952 Casareto, B. E., Niraula, M. P., Fujimura, H., and Suzuki, Y.: Seawater carbonate chemistry,
953 primary production, biomass and calcification of plankton and bacteria, 2009. PANGAEA,
954 https://doi.org/10.1594/PANGAEA.756687, 2009.

955 Comeau, S., Gorsky, G., Jeffree, R., Teyssié, Jean-Louis, and Gattuso, Jean-Pierre: Seawater 956 carbonate chemistry, shell linear extension and calcification during calcein staining and ⁴⁵Ca 957 experiments with pteropod Limacina helicina, 2009. PANGAEA. 958 https://doi.org/10.1594/PANGAEA.726856, Supplement to: Comeau, S. et al. : Impact of ocean 959 acidification on a key Arctic pelagic mollusc (Limacina helicina), Biogeosciences, 6(9), 1877-960 1882, https://doi.org/10.5194/bg-6-1877-2009, 2009.

961 Comeau, S., Jeffree, R., Teyssié, Jean-Louis, and Gattuso, Jean-Pierre: Seawater carbonate
962 chemistry and biological processes during experiments with *Limacina helicina*, 2009. PANGAEA,
963 https://doi.org/10.1594/PANGAEA.744720, In: EPOCA Arctic experiment 2009 team (2009):
964 EPOCA 2009 Svalbard benthic experiment. PANGAEA,
965 https://doi.org/10.1594/PANGAEA.745083, 2010a.

966 Comeau, S., Gorsky, G., Alliouane, S., and Gattuso, Jean-Pierre: Seawater carbonate chemistry 967 and shell length of Mediterranean pteropod Cavolinia inflexa larvae during experiments. 968 Laboratoire d'Océanographie de Villefranche, PANGAEA, 969 https://doi.org/10.1594/PANGAEA.733905, Supplement to: Comeau, S. et al. : Larvae of the 970 pteropod Cavolinia inflexa exposed to aragonite undersaturation are viable but shell-less, Marine 971 Biol., 157(10), 2341-2345, https://doi.org/10.1007/s00227-010-1493-6, 2010b.

972 Comeau, S., Edmunds, P. J., Spindel, N. B., and Carpenter, R. C.: The responses of eight coral
973 reef calcifiers to increasing partial pressure of CO2 do not exhibit a tipping point. PANGAEA,
974 https://doi.org/10.1594/PANGAEA.833687, Supplement to: Comeau, S. et al.: The responses of
975 eight coral reef calcifiers to increasing partial pressure of CO2 do not exhibit a tipping point.
976 Limnol. Oceanogr., 58(1), 388-398, https://doi.org/10.4319/lo.2013.58.1.0388, 2013.

977 Comeau, S., Cornwall, C. E., De Carlo, E. H., Krieger, E., and McCulloch, M. T.: Seawater
978 carbonate chemistry and calcification physiology data in coral reef taxa. PANGAEA,
979 https://doi.org/10.1594/PANGAEA.892655, Supplement to: Comeau, S. et al.: Similar controls on
980 calcification under ocean acidification across unrelated coral reef taxa, Glob. Change Biol.,
981 https://doi.org/10.1111/gcb.14379, 2018.

Comeau, S., Cornwall, C. E., Pupier, C. A., DeCarlo, Thomas M., Alessi, C., Trehern, R., and
McCulloch, M. T.: Seawater carbonate chemistry and calcification rate, calcifying fluid pH,
calcifying fluid DIC, photosynthetic rates, metabolic alteration of pH in the DBL of corals and
coralline algae. PANGAEA, https://doi.org/10.1594/PANGAEA.914328, 2019.

986 Cornwall, C. E., Comeau, S., DeCarlo, T. M., Moore, B., D'Alexis, Q., and McCulloch, M. T.:

987 Seawater carbonate chemistry and resistance of corals and coralline algae to ocean acidification.

988 PANGAEA, https://doi.org/10.1594/PANGAEA.914886, 2018.

989 Courtney, T., Westfield, I. T., & Ries, J. B.: Seawater carbonate chemistry and calcification in the 990 tropical urchin Echinometra viridis in a laboratory experiment [dataset], 2013. PANGAEA, 991 https://doi.org/10.1594/PANGAEA.824707, Supplement to: Courtney, T., Westfield, I. T., & Ries, 992 J. B.: CO₂-induced ocean acidification impairs calcification in the tropical urchin *Echinometra* 993 viridis. Journal of Experimental Marine Biology and Ecology, 440, 169-175. 994 https://doi.org/10.1016/j.jembe.2012.11.013, 2013.

- Courtney, T., and Ries, J. B.: Impact of atmospheric pCO2, seawater temperature, and calcification
 rate on the delta 180 and delta 13C composition of echinoid calcite (Echinometra viridis).
 PANGAEA, https://doi.org/10.1594/PANGAEA.862558, Supplement to: Courtney, T., and Ries,
 J.B. : Impact of atmospheric pCO2, seawater temperature, and calcification rate on the delta 180
 and delta 13C composition of echinoid calcite (*Echinometra viridis*), Chem. Geol., 411, 228-239,
 https://doi.org/10.1016/j.chemgeo.2015.06.030, 2015.
- 1001 Cripps, G., Widdicombe, S., Spicer, J.I. and Findlay, H.S. Biological impacts of enhanced
 1002 alkalinity in *Carcinus maenas*. Marine pollution bulletin, 71(1-2), pp.190-198, 2013
- Dickson, A. G.:Standard potential of the reaction: , and and the standard acidity constant of the
 ion HSO4- in synthetic sea water from 273.15 to 318.15 K. J. Chem. Thermodyn., 22(2), 113127, https://doi.org/10.1016/0021-9614(90)90074-z, 1990.
- Diner, R. E., Benner, I., Passow, U., Iglesias-Rodriguez, M. D., & Robertson, D. L.:
 Negative effects of ocean acidification on calcification vary within the coccolithophore
 genus *Calcidiscus. Marine Biology*, 162, 1287–1305. https://doi.org/10.1007/s00227-0152669-x, 2015.
- 1010 Eisaman, M. D., Geilert, S., Renforth, P., Bastianini, L., Campbell, J., Dale, A. W., Foteinis, S.,
- 1011 Grasse, P., Hawrot, O., Löscher, C. R., Rau, G. H., and Rønning, J.: Assessing the technical aspects
- 1012 of ocean-alkalinity-enhancement approaches. State Of The Planet, 2-oae2023, 1-29,
- 1013 https://doi.org/10.5194/sp-2-oae2023-3-2023, 2023.
- Evensen, N. R., and Edmunds, P. J.: Interactive effects of ocean acidification and neighboring
 corals on the growth of Pocillopora vertucosa. PANGAEA,
 https://doi.org/10.1594/PANGAEA.867268, Supplement to: Evensen, N.R., and Edmunds, P.J.:

- 1017 Interactive effects of ocean acidification and neighboring corals on the growth of *Pocillopora*1018 *verrucosa*, Marine Biol., 163(7), https://doi.org/10.1007/s00227-016-2921-z, 2016.
- 1019 Feely, R. A., Sabine, C. L., Lee, K., Berelson, W., Kleypas, J., Fabry, V. J., and Millero, F. J. :
- 1020 Impact of anthropogenic CO2 on the CaCO3 system in the oceans. Science, 305(5682), 362–366,
- 1021 https://doi.org/10.1126/SCIENCE.1097329, 2004.
- 1022 Fennel, K., Long, M.C., Algar, C., Carter, B., Keller, D., Laurent, A., Mattern, J.P., Musgrave, R.,
- 1023 Oschlies, A., Ostiguy, J. and Palter, J.: Modeling considerations for research on Ocean Alkalinity
- 1024 Enhancement (OAE). State of the Planet Discussions, 2023, 1-47, 2023.
- Ferderer, A., Chase, Z., Kennedy, F., Schulz, K.G. and Bach, L.T.: Assessing the influence of
 ocean alkalinity enhancement on a coastal phytoplankton community. Biogeosciences, 19(23),
 5375-5399, 2022.
- Findlay, H. S., Kendall, M. A., Spicer, J. I., and Widdicombe, S. : Seawater carbonate chemistry
 and biological processes during experiments with barnacle Semibalanus balanoides, 2010.
 PANGAEA, https://doi.org/10.1594/PANGAEA.737438, Supplement to: Findlay, H.S. et al.:
 Relative influences of ocean acidification and temperature on intertidal barnacle post-larvae at the
 northern edge of their geographic distribution, Estuar. Coast. Shelf S., 88(4), 675-682,
 https://doi.org/10.1016/j.ecss.2009.11.036, 2010.
- 1034 Fiorini, S., Middelburg, J. J., and Gattuso, Jean-Pierre: Seawater carbonate chemistry, nutrients, 1035 particulate carbon and growth rate of Emiliania huxleyi (AC472), Calcidiscus leptoporus (AC370) 1036 and Syracosphaera pulchra (AC418) during experiments, 2011. PANGAEA, 1037 https://doi.org/10.1594/PANGAEA.773860, Supplement to: Fiorini, S. et al.: Testing the effects 1038 of elevated pCO2 on coccolithophores (Prymnesiophyceae): comparison between haploid and 1039 diploid life stages. J. Phycol., 47(6), 1281–1291, https://doi.org/10.1111/j.1529-1040 8817.2011.01080.x, 2011.
- Friedlingstein, P., O'sullivan, M., Jones, M.W., Andrew, R.M., Gregor, L., Hauck, J., Le Quéré,
 C., Luijkx, I.T., Olsen, A., Peters, G.P. and Peters, W.: Global carbon budget 2022. Earth System
 Science Data Discussions, 2022, 1-15, 2022.

- Gafar, N.A. and Schulz, K.G.: A three-dimensional niche comparison of *Emiliania huxleyi* and
 Gephyrocapsa oceanica: reconciling observations with projections. Biogeosciences, 15(11),
 3541-3560, 2018.
- 1047 Garilli, V., Rodolfo-Metalpa, R., Scuderi, D., Brusca, L., Parrinello, D., Rastrick, S. P. S.,
- 1048 Foggo, A., Twitchett, R. J., Hall-Spencer, J. M., & Milazzo, M.: Physiological advantages
- 1049 of dwarfing in surviving extinctions in high-CO₂ oceans [dataset], 2015. PANGAEA.
- 1050 https://doi.org/10.1594/PANGAEA.847397, Supplement to: Garilli, V., Rodolfo-Metalpa,
- 1051 R., Scuderi, D., Brusca, L., Parrinello, D., Rastrick, S. P. S., et al.: Physiological
- advantages of dwarfing in surviving extinctions in high-CO₂ oceans. *Nature Climate*
- 1053 *Change*. https://doi.org/10.1038/NCLIMATE2616, 2015.
- 1054 Gately, J. A., Kim, S. M., Jin, B., Brzezinski, M. A., and Iglesias-Rodriguez, M. D.:
- 1055 Coccolithophores and diatoms resilient to ocean alkalinity enhancement: A glimpse of hope?
 1056 Science Advances, 9(24), eadg6066. https://doi.org/10.1126/SCIADV.ADG6066, 2023.
- 1057 Gattuso, J.P., Magnan, A.K., Bopp, L., Cheung, W.W., Duarte, C.M., Hinkel, J., Mcleod, E.,
- 1058 Micheli, F., Oschlies, A., Williamson, P. and Billé, R.: Ocean solutions to address climate
- 1059 change and its effects on marine ecosystems, Frontiers in Marine Science, 5, 410554, 2018.
- 1060
- Gattuso, J.P., Williamson, P., Duarte, C.M. and Magnan, A.K.: The potential for ocean-based
 climate action: negative emissions technologies and beyond. Frontiers in Climate, 2, 575716,
 2021.
- Gazeau, F., Quiblier, C., Jansen, J. M., Gattuso, Jean-Pierre, Middelburg, J. J., and Heip, C. H.R.:
 Seawater carbonate chemistry and calcification during incubation experiments with *Mytilus edulis*and *Grassostrea gigas*, 2006. PANGAEA, https://doi.org/10.1594/PANGAEA.718130,
 Supplement to: Gazeau, F. et al.: Impact of elevated CO2 on shellfish calcification. Geophys. Res.
 Lett., 34, https://doi.org/10.1029/2006GL028554, 2007.
- Gehlen, M., Gruber, N., Gangstø, R., Bopp, L. and Oschlies, A.: Biogeochemical consequences of
 ocean acidification and feedback to the earth system. Ocean acidification, 1, 230-248, 2011.
- 1071 González, M.F. and Ilyina, T: 2016. Impacts of artificial ocean alkalinization on the carbon cycle

- and climate in Earth system simulations. Geophys.ical Re.earch Letters, 43(12), pp.6493-6502,2016.
- 1074 González, M.F., Ilyina, T., Sonntag, S. and Schmidt, H.: Enhanced Rates of Regional Warming
- and Ocean Acidification After Termination of Large-Scale Ocean Alkalinization, Geophys. Res.
 Lett., 45(14), 7120-7129, 2018.
- Hansen, P.J., 2002. Effect of high pH on the growth and survival of marine phytoplankton:
 implications for species succession. Aquatic Mmicrobial Eecology, 28(3), pp.279-288, 2002.
- Hansen, P. J., Lundholm, N., and Rost, Björn.: Seawater carbonate chemistry and growth rate
 during experiments with dinoflagellates, 2007. PANGAEA,
 https://doi.org/10.1594/PANGAEA.718182, In supplement to: Hansen, P.J. et al.: Growth
 limitation in marine red-tide dinoflagellates: effects of pH versus inorganic carbon availability.
 Mar. Ecol. Prog. Ser., 334, 63-71, https://doi.org/10.3354/meps334063, 2007.
- Hartmann, J., Suitner, N., Lim, C., Schneider, J., Marín-Samper, L., Arístegui, J., Renforth, P.,
 Taucher, J. and Riebesell, U.: Stability of alkalinity in ocean alkalinity enhancement (OAE)
 approaches–consequences for durability of CO 2 storage, Biogeosciences Discussions, 2022, 129, 2022.
- He, J. and Tyka, M.D.: Limits and CO 2 equilibration of near-coast alkalinity enhancement,
 Biogeosciences, 20(1), 27-43, 2023.
- Ho, D.T., Bopp, L., Palter, J.B., Long, M.C., Boyd, P.W., Neukermans, G. and Bach, L.T.:
 Monitoring, reporting, and verification for ocean alkalinity enhancement, State of the Planet, 2, 112, 2023.
- Humphreys, M. P., Lewis, E. R., Sharp, J. D., and Pierrot, D.: PyCO2SYS v1.8: marine carbonate
 system calculations in Python, Geosci. Model Dev., 15(1), 15–43, https://doi.org/10.5194/gmd15-15-2022, 2022.
- 1096 Iglesias-Rodriguez, M.D., Halloran, P. R., Rickaby, R.E.M., Hall, I. R., Colmenero-Hidalgo, E.,
- 1097 Gittins, J. R., Green, D. R. H., Tyrrell, T., Gibbs S.J., von Dassow, P., Rehm, E., Armbrust, E. V.,
- 1098 and Boessenkool, K. P.: Seawater carbonate chemistry and processes during experiments with

Emiliania huxleyi, 2008. PANGAEA, https://doi.org/10.1594/PANGAEA.718841, Supplement
to: Iglesias-Rodriguez, D. et al.: Phytoplankton calcification in a high-CO2 world. Science,
320(5874), 336-340, https://doi.org/10.1126/science.1154122, 2008.

Johnson, M. D., Bravo, L., Lucey, N. M., and Altieri, A.H. : Seawater carbonate chemistry and
calcification rate of crustose coralline algae. PANGAEA,
https://doi.org/10.1594/PANGAEA.939809, 2021.

Keul, N., Langer, G., de Nooijer, L. J., and Bijma, J.: Seawater carbonate chemistry and benthic
foraminifera *Ammonia sp.* mass, size, and growth rate during experiments, 2013. PANGAEA,
https://doi.org/10.1594/PANGAEA.821209, Supplement to: Keul, N. et al.: Effect of ocean
acidification on the benthic foraminifera *Ammonia* sp. is caused by a decrease in carbonate ion
concentration. Biogeosciences, 10(10), 6185-6198, https://doi.org/10.5194/bg-10-6185-2013, 2013.

1110 Kheshgi, H.S.: Sequestering atmospheric carbon dioxide by increasing ocean alkalinity, Energy,
1111 20(9), 915-922, 1995.

Kisakürek, B., Eisenhauer, A., Böhm, F., Hathorne, E. C., and Erez, J.: Seawater carbonate
chemistry and biological processes of foraminifera, Globigerinoides ruber and Globigerinella
siphonifera during experiments, 2011. PANGAEA, https://doi.org/10.1594/PANGAEA.763297,

1115 Supplement to: Kisakürek, B. et al.: Controls on calcium isotope fractionation in cultured

1116 planktonic foraminifera, *Globigerinoides ruber and Globigerinella siphonifera*. Geochim.

1117 Cosmochim. Ac., 75(2), 427-443, https://doi.org/10.1016/j.gca.2010.10.015, 2011.

Kroeker, K.J., Kordas, R.L., Crim, R., Hendriks, I.E., Ramajo, L., Singh, G.S., Duarte, C.M. and
Gattuso, J.P. Impacts of ocean acidification on marine organisms: quantifying sensitivities and
interaction with warming. Global change biology, 19(6), pp.1884-1896, 2013.

1121 Krueger, T., Horwitz, N., Bodin, J., Giovani, Maria-Evangelia, Escrig, S., Meibom, A., and Fine, 1122 M.: Seawater carbonate chemistry and photosynthesis, respiration and calcification of common 1123 reef-building coral in the Northern Red Sea. PANGAEA, 1124 https://doi.org/10.1594/PANGAEA.880318, Supplement to: Krueger, T. et al.: Common reef-1125 building coral in the Northern Red Sea resistant to elevated temperature and acidification. Roy. 1126 Soc. Open Sci., 4(5), 170038, https://doi.org/10.1098/rsos.170038, 2017.

1127 Langer, G., Geisen, M., Baumann, Karl-Heinz, Kläs, J., Riebesell, U., Thoms, S., and Young, J.: 1128 Seawater carbonate chemistry, growth rate and processes during experiments with Coccolithus 1129 pelagicus and Calcidiscus 2006. PANGAEA, leptoporus, 1130 https://doi.org/10.1594/PANGAEA.721107, Supplement to: Langer, G. et al.: Species-specific 1131 responses of calcifying algae to changing seawater carbonate chemistry. Geochem. Geophys. 1132 Geosys., 7, Q09006, https://doi.org/10.1029/2005GC001227, 2006.

1133 Langer, G., and Bode, M.: Seawater carbonate chemistry, growth rate and morphology of 1134 Calcidiscus leptoporus (RCC1135) during experiments, 2011. PANGAEA, 1135 https://doi.org/10.1594/PANGAEA.763286, Supplement to: Langer, G., and Bode, M.: CO2 1136 mediation of adverse effects of seawater acidification in Calcidiscus leptoporus. Geochem. 1137 Geophys. Geosys., 12(5), Q05001, https://doi.org/10.1029/2010GC003393, 2011.

Leung, J.Y., Zhang, S., and Connell, S.D.: Is ocean acidification really a threat to marine calcifiers?
A systematic review and meta-analysis of 980+ studies spanning two decades. Small, 18(35),
2107407, 2022.

Lischka, S., Büdenbender, J., Boxhammer, T., and Riebesell, U.: Seawater carbonate chemistry
and biological processes of Limacina helicina during experiments, 2011. PANGAEA,
https://doi.org/10.1594/PANGAEA.761910, Supplement to: Lischka, S. et al.: Impact of ocean
acidification and elevated temperatures on early juveniles of the polar shelled pteropod Limacina
helicina: mortality, shell degradation, and shell growth. Biogeosciences, 8(4), 919-932,
https://doi.org/10.5194/bg-8-919-2011, 2011.

Lischka, S., and Riebesell, U.: Synergistic effects of ocean acidification and warming on
overwintering pteropods in the Arctic. PANGAEA, https://doi.org/10.1594/PANGAEA.832422,
Supplement to: Lischka, S., and Riebesell, U.: Synergistic effects of ocean acidification and
warming on overwintering pteropods in the Arctic. Glob. Change Biol., 18(12), 3517-3528,
https://doi.org/10.1111/gcb.12020, 2012.

Lutier, M., Di Poi, C., Gazeau, F., Appolis, A., Luyer, J. L., and Pernet, F.: Revisiting tolerance to
ocean acidification: Insights from a new framework combining physiological and molecular
tipping points of Pacific oyster, Glob. Change Biol., 28(10), 3333–3348.

Maier, C., Hegeman, J., Weinbauer, M. G., and Gattuso, Jean-Pierre: Seawater carbonate
chemistry and calcification of Lophelia pertusa during experiments, 2009. PANGAEA,
https://doi.org/10.1594/PANGAEA.767577, Supplement to: Maier, C. et al.: Calcification of the
cold-water coral Lophelia pertusa under ambient and reduced pH. Biogeosciences, 6(8), 16711680, https://doi.org/10.5194/bg-6-1671-2009, 2009.

Manno, C., Morata, N., and Bellerby, R.G.J.: Seawater carbonate chemistry, survival rate, shell
size, calcification rate of the planktonic foraminifer Neogloboquadrina pachyderma (sinistral) in a
laboratory experiment. PANGAEA, https://doi.org/10.1594/PANGAEA.830908, Supplement to:
Manno, C. et al. : Effect of ocean acidification and temperature increase on the planktonic
foraminifer Neogloboquadrina pachyderma (sinistral). Polar Biol., 35(9), 1311-1319,
https://doi.org/10.1007/s00300-012-1174-7, 2012.

1167 Manríquez, P. H., Jara, M. E., Seguel, M.E., Torres, R., Alarcon, E., Lee, M. R., and Dam, H. G.: 1168 Ocean acidification and increased temperature have both positive and negative effects on early 1169 ontogenetic traits of a shore keystone predator PANGAEA, rocky species. 1170 https://doi.org/10.1594/PANGAEA.869291, 2016.

1171 Mekkes, L., Renema, W., Alin, S. R., Feely, R. A., Huisman, J., Roessingh, P., &

Peijnenburg, K. T. C. A.: Seawater carbonate chemistry and shell thickness, shell dissolution
of *Limacina helicina* pteropods [dataset]. PANGAEA.
https://doi.org/10.1594/PANGAEA.930065, 2021.

Meyer, F. W., Vogel, N., Teichberg, M., Uthicke, S., Wild, C., and Diaz-Pulido, G.: The
physiological response of two green calcifying algae from the great barrier reef towards high
dissolved inorganic and organic carbon (DIC and DOC) availability. PANGAEA,
https://doi.org/10.1594/PANGAEA.868094, 2015.

1179 Meyer, F. W., Vogel, N., Diele, K., Kunzmann, A., Uthicke, S., & Wild, C.: Effects of high 1180 dissolved inorganic and organic carbon availability on the physiology of the hard coral 1181 Acropora millepora from the Great Barrier Reef [dataset]. PANGAEA. 1182 https://doi.org/10.1594/PANGAEA.869416, 2016

50

- 1183 Mongin, M., Baird, M.E., Lenton, A., Neill, C. and Akl, J.: Reversing ocean acidification along
- the Great Barrier Reef using alkalinity injection, Environ. Res. Lett., 16(6), 064068, 2021.
- 1185 Monserrat, M., Comeau, S., Verdura, J., Alliouane, S., Spennato, G., Priouzeau, F., Romero, G.,
- 1186 and Mangialajo, L.: Seawater carbonate chemistry and the recruitment of macroalgal marine
- 1187 forests. PANGAEA, https://doi.org/10.1594/PANGAEA.955425, 2022.
- 1188 Moras, C.A., Bach, L.T., Cyronak, T., Joannes-Boyau, R. and Schulz, K.G.: Ocean alkalinity
- enhancement–avoiding runaway CaCO 3 precipitation during quick and hydrated lime dissolution.
- 1190 Biogeosciences, 19(15), 3537-3557, 2022.
- 1191 National Academies of Sciences, Engineering, and Medicine: A research strategy for ocean-based1192 carbon dioxide removal and sequestration, 2021.
- Ninokawa, A., Takeshita, Y., Jellison, B. M., Jurgens, L. J., and Gaylord, B.: Seawater carbonate
 chemistry and mussel respiration and calcification rates. PANGAEA,
 https://doi.org/10.1594/PANGAEA.915978, 2020.
- Ninokawa, A. T., Saley, A. M., Shalchi, R., and Gaylord, B.: Multiple carbonate system parameters
 independently govern shell formation in a marine mussel. Communications Earth & Environment,
 5(1), https://doi.org/10.1038/s43247-024-01440-5, 2024.
- Noisette, F., Bordeyne, F., Davoult, D., and Martin, S.: Assessing the physiological responses of
 the gastropod Crepidula fornicata to predicted ocean acidification and warming. PANGAEA,
 https://doi.org/10.1594/PANGAEA.860508, Supplement to: Noisette, F. et al.: Assessing the
 physiological responses of the gastropod Crepidula fornicata to predicted ocean acidification and
 warming. Limnol. Oceanogr., 61(2), 430-444, https://doi.org/10.1002/lno.10225, 2016.
- Okazaki, R., Swart, P. K., and Langdon, C.: Stress-tolerant corals of Florida Bay are vulnerable to
 ocean acidification. PANGAEA, https://doi.org/10.1594/PANGAEA.833005, Supplement to:
 Okazaki, R. et al.: Stress-tolerant corals of Florida Bay are vulnerable to ocean acidification. Coral
 Reefs, 32(3), 671-683, https://doi.org/10.1007/s00338-013-1015-3, 2013.
- 1208 Ong, E. Z., Briffa, M., Moens, T., and Van Colen, C.: Seawater carbonate chemistry and 1209 respiration, clearance and calcification rates of the common cockle Cerastoderma edule.

1210 PANGAEA, https://doi.org/10.1594/PANGAEA.949749, 2017.

Oron, S., Evans, D., Abramovich, S., Almogi-Labin, A., and Erez, J.: Seawater carbonate
chemistry and calcification, respiration, and photosynthesis of the widespread diatom-bearing LBF
Operculina ammonoides. PANGAEA, https://doi.org/10.1594/PANGAEA.929866, 2020.

1214 Oschlies, A., Bach, L., Rickaby, R., Satterfield, T., Webb, R.M. and Gattuso, J.P.: Climate targets,

1215 carbon dioxide removal and the potential role of Ocean Alkalinity Enhancement. State of the

1216 Planet Discussions, 2023, 1-11, 2023.

Palmer, R. M.: Alkalinity enhancement, thermal stress and their impacts on the physiology of three
Caribbean coral species: Acropora Cervicornis, Pseudodiploria strigosa and Siderastrea siderea. In
University of Miami. https://scholarship.miami.edu/esploro/outputs/graduate/Alkalinity-

1220 Enhancement-Thermal-Stress-and-their/991031772314902976, 2022.

1221 Pansch, C., Schaub, I., Havenhand, J. N., and Wahl, M.: Habitat traits and food availability 1222 determine the response of marine invertebrates to ocean acidification. PANGAEA, 1223 https://doi.org/10.1594/PANGAEA.831428, Supplement to: Pansch, C., Schaub, M., Havenhand, 1224 J. N., and Wahl, M.: Habitat traits and food availability determine the response of marine 1225 Glob. 765-777, invertebrates to ocean acidification. Change Biol. 20(3),1226 https://doi.org/10.1111/gcb.12478, 2014.

Paul, A.J. and Bach, L.T.: Universal response pattern of phytoplankton growth rates to increasing
CO2. New Phytologist, 228(6), 1710-1716, 2020.

Prazeres, M., Uthicke, S., and Pandolfi, J. M.: Ocean acidification induces biochemical and
morphological changes in the calcification process of large benthic foraminifera. PANGAEA,
https://doi.org/10.1594/PANGAEA.848419, 2015.

Putnam, H. M., and Gates, R. D.: Preconditioning in the reef-building coral Pocillopora damicornis
and the potential for trans-generational acclimatization in coral larvae under future climate change
conditions. PANGAEA, https://doi.org/10.1594/PANGAEA.859356, Supplement to: Putnam,
H.M., and Gates, R.D.: Preconditioning in the reef-building coral Pocillopora damicornis and the
potential for trans-generational acclimatization in coral larvae under future climate change

1237 conditions. J. Exp. Biol., 218(15), 2365-2372, https://doi.org/10.1242/jeb.123018, 2015.

1238 Ramajo, L., Marbà, N., Prado, L., Peron, S., Lardies, M. A., Rodriguez-Navarro, A., Vargas, C. 1239 A., Lagos, N. A., and Duarte, C. M.: Biomineralization changes with food supply confer juvenile 1240 scallops (Argopecten purpuratus) resistance to ocean acidification. PANGAEA, 1241 https://doi.org/10.1594/PANGAEA.860506, Supplement to: Ramajo, L. et al.: Biomineralization 1242 changes with food supply confer juvenile scallops (Argopecten purpuratus) resistance to ocean 1243 acidification. Glob. Change Biol., 22(6), 2025-2037, https://doi.org/10.1111/gcb.13179, 2016.

1244 Renforth, P. and Henderson, G.: Assessing ocean alkalinity for carbon sequestration, Rev.1245 Geophys., 55(3), 636-674, 2017.

Reymond, C. E., Lloyd, A., Kline, D. I., Dove, S., and Pandolfi, J. M.: Decline in growth of
foraminifer Marginopora rossi under eutrophication and ocean acidification scenarios.
PANGAEA, https://doi.org/10.1594/PANGAEA.833683, Supplement to: Reymond, C.E., et al.:
Decline in growth of foraminifer Marginopora rossi under eutrophication and ocean acidification
scenarios. Glob. Change Biol., 19(1), 291-302, https://doi.org/10.1111/gcb.12035, 2013.

- Richardson, K., Steffen, W., Lucht, W., Bendtsen, J., Cornell, S.E., Donges, J.F., Drüke, M.,
 Fetzer, I., Bala, G., von Bloh, W., and Feulner, G.: Earth beyond six of nine planetary boundaries.
 Science Advances, 9(37), p.eadh2458, 2023.
- 1254 Richier, S., Fiorini, S., Kerros, Marie-Emmanuelle, von Dassow, P., and Gattuso, Jean-Pierre: 1255 Seawater carbonate chemistry, particulate inorganic and organic carbon and growth rate of 1256 Emiliana huxleyi (RCC1216) duirng experiments, 2011. PANGAEA, 1257 https://doi.org/10.1594/PANGAEA.770439, In supplement to: Richier, S. et al.: Response of the 1258 calcifying coccolithophore Emiliania huxleyi to low pH/high pCO2: from physiology to molecular 1259 level. Marine Biol., 158(3), 551-560, https://doi.org/10.1007/s00227-010-1580-8, 2011.
- Riebesell, U., Wolf-Gladrow, D.A. and Smetacek, V.: Carbon dioxide limitation of marine
 phytoplankton growth rates. Nature, 361(6409), 249-251, 1993.
- Riebesell, U. and Gattuso, J.P.:, 2015. Lessons learned from ocean acidification research, NatureClimate Change, 5(1), pp.12-14, 2015.

Riebesell, U., Bach, L.T., Bellerby, R.G., Monsalve, J.R.B., Boxhammer, T., Czerny, J., Larsen,
A., Ludwig, A. and Schulz, K.G.: Competitive fitness of a predominant pelagic calcifier impaired
by ocean acidification, Nat. Geosci., 10(1), 19-23, 2017.

1267 Ries, J. B., Cohen, A.L., and McCorkle, D.C.: Seawater carbonate chemistry and biological 1268 experiments with calcifiing organisms, 2009. processes during PANGAEA, 1269 https://doi.org/10.1594/PANGAEA.733947, Supplement to: Ries, J.B., et al.: Marine calcifiers 1270 exhibit mixed responses to CO2-induced ocean acidification. Geology, 37(12), 1131-1134, 1271 https://doi.org/10.1130/G30210A.1, 2009.

Ries, J.B.:, 2011. A physicochemical framework for interpreting the biological calcification
response to CO2-induced ocean acidification. Geochimica et cosmochimica acta, 75(14), pp.40534064, 2011.

Sett, S., Bach, L. T., Schulz, K. G., Koch-Klavsen, S., Lebrato, M., and Riebesell, U.: Temperature
modulates coccolithophorid sensitivity of growth, photosynthesis and calcification to increasing
seawater pCO2 [dataset]. PANGAEA, https://doi.org/10.1594/PANGAEA.835214, Supplement
to: Sett, S et al. (2014): Temperature Modulates Coccolithophorid Sensitivity of Growth,
Photosynthesis and Calcification to Increasing Seawater pCO2. PLoS ONE, 9(2), e88308,
https://doi.org/10.1371/journal.pone.0088308, 2014.

Schulz, K.G., Bach, L.T. and Dickson, A.G.: Seawater carbonate system considerations for ocean
alkalinity enhancement research. State of the Planet Discussions, 2023, 1-24, 2023.

Sciandra, A., Harlay, J., Lefèvre, D., Lemee, R., Rimmelin, P., Denis, M., and Gattuso, JeanPierre: Seawater carbonate chemistry and processes during experiments with Emiliania huxleyi
(TW1), 2003. PANGAEA, https://doi.org/10.1594/PANGAEA.727841, Supplement to: Sciandra,
A. et al.: Response of coccolithophorid Emiliania huxleyi to elevated partial pressure of CO2 under
nitrogen limitation. Mar. Ecol. Prog. Ser., 261, 111-122, https://doi.org/10.3354/meps261111,
2003.

Seabold, S., & Perktold, J.: Statsmodels: Econometric and statistical modeling with Python. In *9th Python in Science Conference* (pp. 57–61). Austin, TX, June 28–July 3, 2010.
https://doi.org/10.25080/Majora-92bf1922-011, 2010.

- Sinutok, S., Hill, R., Doblin, M. A., Wuhrer, R., and Ralph, P.J. : Seawater carbonate chemistry,
 calcification rate, oxygen production, maximum quantum yield, symbiont density, chlorophyll
 concentration and crystal width of Halimeda macroloba, Halimeda cylindracea and Marginopora
 vertebralis during experiments, 2011. PANGAEA, https://doi.org/10.1594/PANGAEA.774792,
 Supplement to: Sinutok, S. et al. : Warmer more acidic conditions cause decreased productivity
 and calcification in subtropical coral reef sediment-dwelling calcifiers. Limnol. Oceanogr., 56(4),
 1200-1212, https://doi.org/10.4319/lo.2011.56.4.1200, 2011.
- Sordo, L., Duarte, C., Joaquim, S., Gaspar, M. B., Matias, D.: Seawater carbonate chemistry and
 growth and survival of juveniles of the striped venus clam Chamelea gallina. PANGAEA,
 https://doi.org/10.1594/PANGAEA.937477, 2021.
- Stoll, H. M., Cruzado, A., Shimizu, N., and Kanamaru, K. : Seawater carbonate chemistry and
 B/Ca, calcification rate of Emiliania huxleyi and Coccolithus braarudii. PANGAEA,
 https://doi.org/10.1594/PANGAEA.949913, 2012.
- Sulpis, O., Lauvset, S. K., and Hagens, M.: Current estimates of K₁^{*} and K₂^{*} appear inconsistent
 with measured CO₂ system parameters in cold oceanic regions. Ocean Sci. 16(4), 847–862,
 https://doi.org/10.5194/os-2020-19, 2020.
- Tatters, A. O., Schnetzer, A., Fu, F., Lie, A.Y.A., Caron, D. A., and Hutchins, D. A.: Short- versus
 long-term responses to changing CO2 in a coastal dinoflagellate bloom. PANGAEA,
 https://doi.org/10.1594/PANGAEA.823381, Supplement to: Tatters, A.O., et al.: Short- versus
 long-term responses to changing CO2 in a coastal dinoflagellate bloom: implications for
 interspecific competitive interactions and community structure. Evolution, 67(7), 1879-1891,
 https://doi.org/10.1111/evo.12029, 2013.
- Truchot, J.-.: Water carbonate alkalinity as a determinant of hemolymph acid-base balance in the
 shore crab, Carcinus maenas: a study at two different ambientP CO 2 andP O 2 levels, J. Comp.
 Physiol.. B, 154(6), 601–606. https://doi.org/10.1007/bf00684414, 1984.
- Truchot, J.: Changes in the Hemolymph Acid-Base State of the Shore Crab, Carcinus maenas,
 Exposed to Simulated Tidepool Conditions, Biol. Bull., 170(3), 506–518,
 https://doi.org/10.2307/1541858, 1986.

Uppström, L. R.: The boron/chlorinity ratio of deep-sea water from the Pacific Ocean, Deep-sea
Research And Oceanographic Abstracts, 21(2), 161–162, https://doi.org/10.1016/00117471(74)90074-6, 1974.

1323 Uthicke, S., and Fabricius, K.E.: Seawater carbonate chemistry, productivity and calcification of 1324 vertebralis experiment. Marginopora in а laboratory PANGAEA. 1325 https://doi.org/10.1594/PANGAEA.831207, Supplement to: Uthicke, S., Fabricius, K.E.: 1326 Productivity gains do not compensate for reduced calcification under near-future ocean 1327 acidification in the photosynthetic benthic foraminifer species Marginopora vertebralis. Glob. 1328 Change Biol., 18. Zenodo. https://doi.org/10.5281/zenodo.12806137, 2024.

Van de Waal, D. B., John, U., Ziveri, P., Reichart, Gert-Jan, Hoins, M., Sluijs, A., and Rost, B.:
Seawater carbonate chemistry and growth, calcification of Thoracosphaera heimii in a laboratory
experiment. PANGAEA, https://doi.org/10.1594/PANGAEA.824705, Supplement to: Van de
Waal, D.B., et al. : Ocean Acidification Reduces Growth and Calcification in a Marine
Dinoflagellateal.pone.0065987, 2013.

1334 Vásquez-Elizondo, R. M., and Enríquez, S.: Coralline algal physiology is more adversely affected 1335 elevated reduced PANGAEA, by temperature than pH. 1336 https://doi.org/10.1594/PANGAEA.860802, Supplement to: Vásquez-Elizondo, R.M., and 1337 Enríquez, S.: Coralline algal physiology is more adversely affected by elevated temperature than reduced pH. Scientific Reports, 6, 19030, https://doi.org/10.1038/srep19030, 2016. 1338

Waldbusser, G. G., Voigt, E. P., Bergschneider, H., Green, M. A., and Newell, R.I. E.: Seawater
carbonate chemistry and calcification rate of eastern oyster Crassostrea virginica, 2011.
PANGAEA, https://doi.org/10.1594/PANGAEA.758181, Supplement to: Waldbusser, G.G., et al.
Biocalcification in the Eastern Oyster (Crassostrea virginica) in Relation to Long-term Trends in
Chesapeake Bay pH. Estuar. Coast., 34(2), 221-231, https://doi.org/10.1007/s12237-010-9307-0,
2011.

Wang, X., Feng, X., Zhuang, Y., Lu, J., Wang, Y., Gonçalves, R.J., Li, X., Lou, Y., and Guan, W.:
Seawater carbonate chemistry and physiology and toxicity of dinoflagellate Karenia mikimotoi.
PANGAEA, https://doi.org/10.1594/PANGAEA.923683, 2019.

- Wang, H., Pilcher, D.J., Kearney, K.A., Cross, J.N., Shugart, O.M., Eisaman, M.D. and Carter,
 B.R. Simulated impact of ocean alkalinity enhancement on atmospheric CO2 removal in the
 Bering Sea. Earth's Future, 11(1), p.e2022EF002816, 2023.
- 1351 White, M. M., Drapeau, D.T., Lubelczyk, L. C., Abel, V. C., Bowler, B. C., and Balch, W. M.:
- 1352 Seawater carbonate chemistry and calcification of an estuarine coccolithophore. PANGAEA,
- 1353 https://doi.org/10.1594/PANGAEA.923623, 2018.
- Zhang, M., Fang, J., Zhang, J., Li, B., Ren, S., Mao, Y., and Gao, Y.: Seawater carbonate chemistry
 and calcification and respiration of Chlamys farreri. PANGAEA,
 https://doi.org/10.1594/PANGAEA.949604, 2011.