

1 **Assessment framework to predict sensitivity of marine calcifiers to ocean alkalinity**
2 **enhancement - identification of biological thresholds and importance of precautionary**
3 **principle**

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15 **Abstract**

16 Ocean alkalinity enhancement (OAE), one of the marine carbon dioxide removal strategies, is
17 gaining recognition in its ability to mitigate climate change and ocean acidification (OA). OAE is
18 based on adding alkalinity to open-ocean and coastal marine systems through a variety of different
19 approaches, which raises carbonate chemistry parameters (such as pH, total alkalinity, aragonite
20 saturation state), and enhances the uptake of carbon dioxide (CO₂) from the atmosphere. There are
21 large uncertainties in both short- and long-term outcomes related to potential environmental
22 impacts, which would ultimately have an influence on the social license and success of OAE as a
23 climate strategy. This paper represents a synthesis effort, leveraging on the OA studies and
24 published data, observed patterns and generalizable responses. Our assessment framework was
25 developed to predict the sensitivity of marine calcifiers to OAE by using data originating from OA
26 studies. The synthesis was done using raw experimental OA data based on 68 collected studies,
27 covering 84 unique species and capturing the responses of eleven biological groups (calcifying
28 algae, corals, dinoflagellates, mollusks, gastropods, pteropods, coccolithophores, annelids,
29 crustacean, echinoderms, and foraminifera), using regression analyses to predict biological
30 responses to NaOH or Na₂CO₃ addition and their respective thresholds. Predicted responses were
31 categorized into six different categories (linear positive and negative, threshold positive and
32 negative, parabolic and neutral) to delineate responses per species. The results show that 34.4% of
33 responses are predicted to be positive (N=33), 26.0% negative (N=25), and 39.2% (N=38) neutral
34 upon alkalinity addition. For the negatively impacted species, biological thresholds, which were
35 based on 50% reduction of calcification rate, were in the range of 50 to 500 μmol/kg NaOH
36 addition. Thus, we emphasize the importance of including much lower additions of alkalinity in
37 experimental trials to realistically evaluate *in situ* biological responses. The primary goal of the
38 research was to provide an assessment of biological rates and thresholds predicted under
39 NaOH/Na₂CO₃ addition that can serve as a tool for delineating OAE risks, guiding and prioritizing
40 future OAE biological research and regional OAE monitoring efforts and communicate the risks
41 with stakeholders. This is important given the fact that at least some of the current OAE approaches
42 do not always assure safe biological space. With 60% of responses being non-neutral, a
43 precautionary approach for OAE implementation is warranted, identifying the conditions where
44 potential negative ecological outcomes could happen, which is key for scaling up and avoiding
45 ecological risks.

46 **1. Introduction**

47 Anthropogenic carbon dioxide (CO₂) emissions have increased at an unprecedented rate and have
48 contributed to global climate change and negative ecological and biogeochemical impacts in the
49 oceans (Feely et al., 2004; Gattuso et al., 2018), to the extent of crossing six different planetary
50 boundaries (Richardson et al., 2023). Oceans play a crucial role in attenuating the increase in
51 atmospheric CO₂ through the absorption of the excess atmospheric CO₂ of roughly a quarter of
52 anthropogenic carbon dioxide (CO₂) emissions, drawing down around 2–3 Pg C yr⁻¹ in recent
53 decades (Friedlingstein et al., 2022). However, without substantial CO₂ emissions abatement and
54 CO₂ removal strategies, profound repercussions on climate, extreme weather events, and
55 socioeconomic implications will follow. Ocean-based CO₂ removal and sequestration strategies
56 (broadly referred to as marine CDR) are among the proposed CDR approaches that remove CO₂
57 and store it for geologically relevant times (National Academies of Sciences, Engineering, and
58 Medicine, 2021). These mCDR approaches only complement CO₂ emission reductions and
59 contribute to the portfolio of climate response strategies needed to meet the global goal of limiting
60 warming to well below 2°C as established by the Paris Agreement. Various mCDR approaches
61 have unique benefits and costs but differ in their value depending on their state of implementation,
62 and whether they act globally and/or locally (Oschlies et al., 2023).

63 Ocean alkalinity enhancement (OAE) has the potential to mitigate climate change through
64 increasing ocean uptake of CO₂, while simultaneously reversing ocean acidification (OA) and
65 improving marine habitats. Despite mostly being in the concept stage, OAE is viewed with a high
66 level of confidence as to its effectiveness: medium on environmental risk, but low on the
67 underlying knowledge base (Eisaman et al., 2023; Gattuso et al., 2021; National Academies of
68 Sciences, Engineering, and Medicine, 2021). One of the major concerns about OAE is large
69 uncertainties in both short- and long-term OAE outcomes related to potential environmental
70 impacts of OAE (Kheshgi, 1995; Bach et al., 2019), especially if OAE were to induce novel
71 conditions in the marine systems that are outside the range of the natural variability, exposing
72 organisms to conditions not experienced in their evolutionary history. The outcome of OAE as a
73 successful climate strategy depends on a thorough and advanced understanding of the impacts of
74 OAE implementation while avoiding or minimizing negative biological effects.

75 **1.1 Leveraging ocean acidification research on marine calcifiers**

76 Increased CO₂ uptake, which initially is absorbed by the ocean as dissolved CO₂, causes a decline
77 in pH, shoaling of the saturation state horizon (Ω_{ar}) and reduced carbonate ion amount content in
78 a process termed ocean acidification (Feely et al., 2004), causing negative consequences to marine
79 biota, especially marine calcifiers, the structure and function of the vulnerable marine ecosystem,
80 and alteration of the carbon cycle. On the other hand, chemical changes induced by OAE are
81 inherently linked to reversing the OA process: increasing pH, shifting carbonate chemistry
82 speciation towards lower aqueous carbon dioxide (pCO₂) and higher carbonate ion (CO₃²⁻) content,
83 as well as higher aragonite saturation state (Ω_{ar}). Such changes could either be within the ranges
84 of the variability of the natural systems to which species are acclimatized, or outside them, creating
85 novel conditions for which species might not have developed suitable acclimation strategies. As
86 such, the biological outcomes are, due to their complexity, highly unpredictable.

87 Scientific progress over the past 30+ years of OA research has brought substantial insights into the
88 biological effects, with the most fundamental outcome being that calcifying organisms would be
89 primarily affected (Riebesell and Gattuso, 2015), with the calcification process being one of the
90 most susceptible pathways, underpinned by species differences in calcification mechanisms (Ries
91 et al., 2009; 2011; Bach et al., 2013; 2015; Leung et al., 2022). However, OA focused heavily on
92 investigating biological effects on the higher acidity range of the carbonate chemistry conditions
93 predicted under future scenarios and most of the studies focused on manipulating the level of pCO₂
94 rather than alkalinity. This resulted in poor understanding of the biological effects at the higher pH
95 end of the carbon chemistry range (Renforth and Henderson, 2017). Some biological inferences
96 can be made based on the understanding of the physiological mechanisms underlying the
97 calcification mechanisms (Bach et al., 2019), but such insights are not adequate to provide
98 sufficient understanding. Despite the lack of biological data at the upper ranges of pH and Ω_{ar} , this
99 study builds on the premise that previous OA studies could be leveraged for assessment of
100 biological responses under OAE. Comparative experimental work, meta-analyses, and the
101 threshold work (Kroeker et al., 2013; Leung et al., 2022; Bednaršek et al., 2019; 2021b,c) have
102 indicated that even very diverse responses can be grouped into categorical responses.

103 Calcification is a primary pathway through which organismal sensitivity to OA is expressed. It is

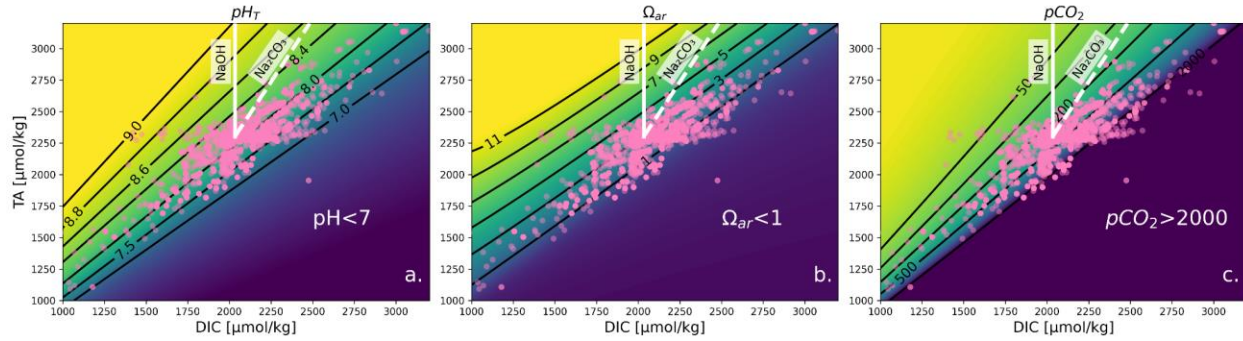
104 directly involved in growth and (abnormal) development across most marine calcifiers, and it
105 indirectly influences susceptibility to predation. As such, calcification can serve as an early
106 warning indicator of stress, while also playing a crucial role in the ecological success of numerous
107 marine calcifiers. Studies have shown that the thresholds for calcification occur at similar pH and
108 saturation state (Ω) values as those affecting energy metabolism processes (Lutier et al., 2022;
109 Bednaršek et al., 2019; 2021b,c). Furthermore, calcification is directly linked to carbon export,
110 which has significant biogeochemical implications that may influence the efficiency of OAE. This
111 study aims to systematically assess the calcification responses of various species under predicted
112 conditions following carbonate-based OAE compound addition.

113 **1.2. Complex carbonate chemistry changes induced by various OAE compounds**

114 Various OAE compounds added to the water change carbonate chemistry in a multifaceted way
115 and require complex calculations of a multi-parameter problem. As the values of TA and DIC
116 change, a variety of other parameters, such as pH, CO_3^{2-} , Ω_{ar} , and pCO_2 exhibit approximately
117 linear relationships, with slopes that vary along these lines (see Fig. 1). This means that if TA and
118 DIC vary in proportion to one another, then the values of these displayed parameters hardly change
119 at a particular salinity, temperature, and pressure. With TA, DIC and the hydrographic conditions
120 (salinity, temperature and pressure), one can constrain the carbonate system. Our method requires
121 us to have *one* variable constraining the entire carbonate system. TA and DIC have the benefit that
122 they can both be directly measured with high precision and accuracy or calculated from other
123 carbonate parameters. They are also both directly linked to OAE, as we are enhancing the TA
124 which then allows DIC to increase over time due to the gradual uptake of atmospheric CO_2 .

125 To demonstrate the changes of the carbonate system in the experimental system, Figure 1 shows
126 the changes in carbonate parameters with the addition of two OAE compounds, i.e. NaOH (solid
127 line) and Na_2CO_3 (dashed line) to seawater. When NaOH is added, only TA increases and when
128 Na_2CO_3 is added, TA and DIC increase at a 2:1 ratio. This results in corresponding changes in pH
129 (Fig. 1a), Ω_{ar} (Fig. 1b) and pCO_2 (Fig. 1c) and shows how much of a change is required to bring
130 the system back to equilibrium with respect to the atmosphere.

131



132

133 **Figure 1:** The effect of changes in TA and DIC on the properties of seawater ($S= 34.68$, $T=16^{\circ}\text{C}$,
 134 $[\text{SiO}_2] = 50 \mu\text{mol/kg}$, $[\text{PO}_4^{3-}] = 0.5 \mu\text{mol/kg}$, $\text{TA} = 2303 \mu\text{mol/kg}$, $\text{DIC} = 2034 \mu\text{mol/kg}$), adapted
 135 from Schulz et al. (2023). Pink dots represent experimental TA and DIC data used in our synthesis.
 136 Subfigures show pH_T , Ω_{ar} and pCO_2 (in μatm). Calculations were carried out with the Python
 137 version of CO2SYS (Humphreys et al., 2022) using the stoichiometric dissociation constants for
 138 carbonic acid from Sulpis et al. (2020), for sulfuric acid by Dickson et al. (1990) and for total
 139 boron from Uppström (1974). The solid white line indicates the effect of adding NaOH and the
 140 dashed white line indicates the effect of adding Na_2CO_3 . This grouping of lines can be translated
 141 so that its initial position moves elsewhere to visualize different initial conditions. Note that at TA
 142 $< 1000 \mu\text{mol/kg}$ and $\text{DIC} < 500 \mu\text{mol/kg}$ the isolines are no longer straight when considering Ω_{ar} ,
 143 however, such conditions are rare in the ocean and not widely applicable. The same contour plot
 144 utilizing GLODAP data plotted instead of experimental data is shown in Supplemental Figure 1.

145 1.3 Testable conceptual framework based on the existing OA studies

146 Based on Ries et al. (2009), calcification responses can be categorized into six categories (Fig. 2):
 147 linear positive or negative response; threshold positive or negative response (exponential fit);
 148 parabolic response; and neutral (no significant) response. We hypothesize that these categories of
 149 responses based on ocean acidification data and delineated by Ries et al. (2009, 2011), could also
 150 be applicable to OAE dosing. For this meta-analysis, we have undertaken three steps: first,
 151 synthesize carbonate chemistry data at regional and global scales to obtain TA, DIC and Ω_{ar}
 152 correlations; second, conduct a literature review and collect available data from OA literature
 153 related to the calcification rate responses across the species of eleven groups of marine calcifiers;
 154 and third, run regression analyses and determine the category of calcification rate response to
 155 TA:DIC, further extending it with addition of NaOH and Na_2CO_3 .

156 The most accurate way of predicting the responses to OAE addition is done based on the
157 mechanistic understanding of calcification response to specific carbonate chemistry parameter(s).
158 The hypothesis was that if mechanistic relationships with identified carbonate chemistry driver(s)
159 are available for species, calcification rate under various feasible OAE scenarios can be predicted
160 with greater accuracy and lower uncertainty. We further focused on investigating if the empirical
161 results were consistent with mechanistic calcification predictions for a few selected species for
162 which the mechanisms were known.

163 Here, we demonstrate the TA:DIC relationship with calcification rates and show the application
164 for the TA:DIC thresholds beyond which the responses become negative. Ultimately, we
165 synthesize which calcifying species or groups are predicted to benefit or lose due to OAE, what
166 constitutes a species-specific safe operating space related to OAE, and we delineate what
167 experiments are most urgently needed to fill in critical knowledge gaps before massive OAE field
168 implementation can be considered.

169 **2. Methodology**

170 **2.1 Literature review of data on marine calcification impact by OA**

171 To assess the impact of OAE on a range of marine calcifiers, we used existing studies on marine
172 species calcification response to OA that had aligned raw biological (calcification rate) data along
173 with corresponding carbonate chemistry. We searched within Scopus, Web of Science, and
174 PubMed and used datasets that were archived in NCEI, OA-ICC and Pangea. Through personal
175 correspondence, we have additionally contacted lead authors of the studies whose data are not or
176 are insufficiently archived. Searches for biological datasets relating to calcification rate and
177 corresponding carbonate chemistry were conducted through November 2023, encompassing 68
178 existing studies. The aim was to cover a wide range of calcifying organisms across various
179 functional groups and 84 species. For several functional groups data was easy to find (algae,
180 coccolithophores, corals, foraminifera, mollusks and dinoflagellates), so no new studies were
181 added after 10 to 15 studies were found. Seven studies were found for pteropods, five for
182 gastropods, four for echinoderms, three for crustaceans and one for annelids. When reviewing the
183 literature, we included data from the OA experimental studies related to the physical-chemical
184 parameters (temperature, salinity, TA, DIC) and biological data related to calcification rate.

185 2.2 Use of TA:DIC instead of Ω_{ar} or pH

186 Understanding the change in carbonate chemistry upon alkalinity addition is essential for the
187 biological experimentalists who are conducting biological assessments to report on the effects of
188 OAE. However, complex changes in the carbonate chemistry induced by alkalinity addition are
189 not intuitive or straightforward; in fact, they are multi-parameter problems that require complex
190 carbonate chemistry calculations. Using the TA:DIC ratio is a more practical way of looking at the
191 impacts of the OAE treatment instead of using a single carbonate parameter because of the high
192 degree of correlation between TA:DIC and other carbonate system parameters (see Fig. 1).

193 With TA, DIC and the hydrographic conditions (salinity, temperature and pressure), one can fully
194 constrain the carbonate system. Our method allows *one* variable constraining the entire carbonate
195 system. TA and DIC have the benefit that they can both be directly measured or calculated from
196 other carbonate and physical parameters. They are also both directly linked to OAE, as we are
197 enhancing the TA which then allows DIC to increase over time due to the gradual uptake of
198 atmospheric CO₂ (Fig. 1 shows the changes in the carbonate chemistry system upon NaOH and
199 Na₂CO₃ addition).

200 Our focus was on streamlining the process of expressing experimental results and subsequently
201 reporting responses, with the goal of reducing the multi-parameter complexity into a single-
202 parameter simplification. This step reduces multiple degrees of freedom into just two, i.e. TA and
203 DIC, with the ratio allowing us to consider this as a 1-parameter problem. As such, TA:DIC is a
204 simplistic and convenient way of describing the system, where we only need to understand the
205 change in TA and DIC ratio, which is feasible for every OAE compound added to the experimental
206 system. In addition, TA:DIC is also the best approximation for the CO₃²⁻ concentration. The
207 insights from multiple biological experimental studies show that the CO₃²⁻ concentration is the
208 representative driver of the calcification process for multiple calcifying groups, although not all,
209 compared to Ω_{ar} , which represents an empirical approximation based on a number of physical and
210 chemical parameters. Furthermore, by using TA:DIC we do not have to choose a particular
211 parameter to describe the changes in calcification. It could also work for the species in which other
212 parameters drive the calcification, e.g. bicarbonate in autotrophic species, Ω_{ar} in bivalves and H⁺
213 flux in foraminifera. In that way, we standardize all the parameters that would otherwise influence
214 the carbonate system and come up with a more uniform way to express the experimental

215 conditions, which would then be useful for easier comparisons among the conducted experiments.
216 For the ease of comparing TA:DIC with pH and Ω_{ar} , we refer the reader to Supplemental Table 1
217 and Supplemental Fig. 2.

218 **2.3 Experimental biological and biogeochemical data**

219 Based on the collected data, the range of pH and Ω_{ar} , experimental conditions used and their
220 TA:DIC relationship was determined (Supplemental Fig. 2 and Supplemental Table 1). Most
221 studies covered pH conditions from 7.5 to 8.5 and Ω_{ar} from <1.0 to values up to 5.0, with a few
222 studies increasing pH up to 9 and exceeding Ω_{ar} of 10. This indicates the potential of leveraging
223 such experimental studies as a baseline for predictive regression models of biological responses to
224 a range of Ω_{ar} conditions, as expected under OAE studies.

225 Once the biological data was compiled, units were standardized where possible. The main issue
226 when compiling data was the lack of standardization of the calcification rates. A variety of
227 calcification rate units were used across different studies. Where possible, the units were converted
228 to mmol of CaCO_3 g weight⁻¹ hr⁻¹. However, the data required to do so was not always readily
229 available. Other units used for calcification rate were mmol of CaCO_3 m⁻² h⁻¹ and mmol of CaCO_3
230 m⁻³ hr⁻¹, and there was also data used as an indication of calcification rate with units mmol #⁻¹ h⁻¹,
231 mmol h⁻¹, mmol cm⁻², % h⁻¹, where ‘#’ indicates one individual. Growth rates and PIC production
232 rates were used as indicators of calcification rate for single-cell organisms. For some species, direct
233 calcification rates were not reported in the literature, instead only relevant parameters related to
234 calcification (shell length, density, thickness) over time were available from the experimental
235 studies. The decision was made to also collect these additional datasets because the statistical
236 analyses of this study focus on the trend in the absolute numbers and would not change by being
237 transformed into the rates. Data were analyzed on a species level, wherever rate units were the
238 same. Hereafter, this is referred to as the species rate group. Where there were multiple studies
239 available for the calcification rate of one species using the same rate units, the data were combined
240 (e.g. *Emiliana huxleyi*).

241 **2.4 Sorting species-specific responses into categories per calcification response**

242 Responses were split into 6 categories: linear positive and linear negative, parabolic, threshold

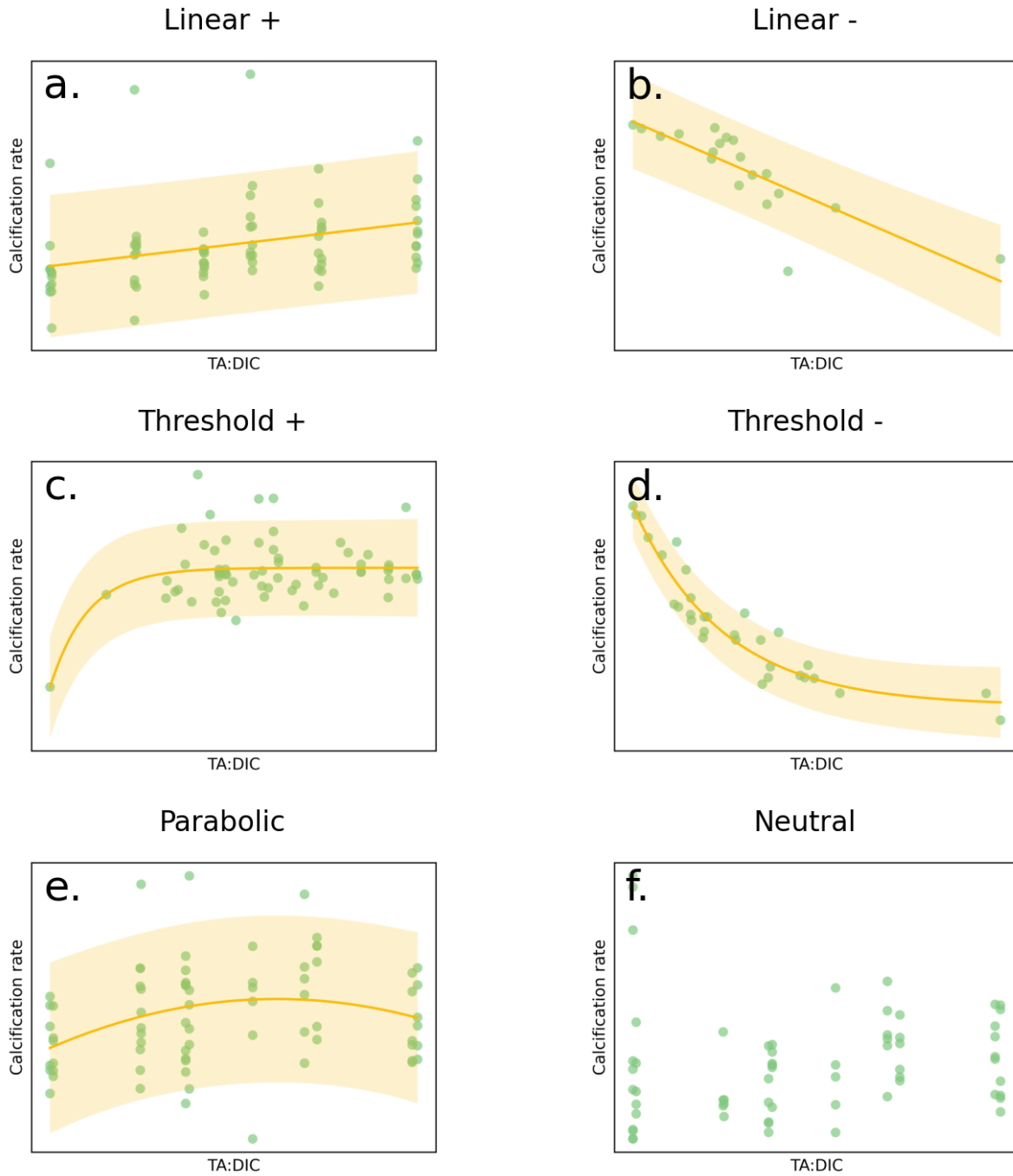
243 positive and negative, and neutral. The response was determined with a best-fit regression model,
244 using the ordinary least squares method in Statsmodels for Python (see Seabold et al., 2010). See
245 Fig. 2 for examples of these responses of calcification rate to increasing TA:DIC ratio.

246 The final response for each species was determined by the regression with the lowest p-value. This
247 method is in contrast with the Ries et al. (2009) study where they chose the regression analysis
248 that yielded the lowest square root of the mean squared error (RMSE) for a given species, and that
249 was statistically significant ($p \leq 0.05$). When applying their method to our data, parabolic and
250 exponential regressions were always favored over linear regressions. When examining these
251 regressions, we found that choosing the best fit based on the lowest p-value yielded better fits, as
252 this method prevents overfitting due to noise in the data. Where a linear regression had the best fit,
253 we assigned a linear response, which could be either positive or negative based on the slope. The
254 species with a significant exponential fit were categorized as threshold positive (+) or threshold
255 negative (-), which was distinguished from the parabolic response with the fitted parabolic curve.

256 The best fit regression was assigned to each species and plotted, but only if the p-value was
257 considered significant, i.e. lower than 0.05. These regressions were plotted along with a 90%
258 prediction interval, which accounts for the variability of the experimental data. The species with a
259 p-value > 0.05 were categorized as having no correlation (neutral response).

260 When multiple datasets were obtained from different studies for the same species and rate units
261 could not be combined, we took each response into consideration and reported the p-value and
262 RMSE for each of the studies. Even when different studies reported varying calcification rates for
263 the same species, we refrained from selecting a single overall species response; rather, we analyzed
264 each species individually. The TA:DIC threshold was computed to indicate the point at which the
265 current calcification rate (i.e. the calcification rate at the baseline) is reduced by a half for linear
266 negative, threshold negative and parabolic responders. The thresholds and the amount of NaOH
267 and Na₂CO₃ required (starting at 10 $\mu\text{mol/kg}$ and then in steps of 50 $\mu\text{mol/kg}$) to reach this
268 threshold were determined. For parabolic responders, the inflection points that tell us when the
269 rate is predicted to change slope are also included in Supplemental Table 2. Once the species'
270 responses were determined, an attempt was made to group them based on functional groups.
271 However, since species within the same functional group had varying responses, grouping them

272 together meant these responses were no longer visible due to a wide spread of data. Therefore,
273 most of the analysis remained on the species level (Table 1).



274

275 **Figure 2:** Examples of the categories of responses between carbonate chemistry parameters
276 (TA:DIC) and calcification rate: a) linear positive (calcification increase with increased TA:DIC);

277 *b) linear negative (calcification decrease with increased TA:DIC); c) exponential for the threshold*
278 *positive response (calcification increase, plateauing at higher TA:DIC); d) exponential for the*
279 *threshold negative response (calcification decline, plateauing at lower TA:DIC), e parabolic*
280 *(calcification increase followed by a decrease at higher TA:DIC) and f) neutral (non-significant)*
281 *response. Responses were only considered significant when $p < 0.05$, otherwise they were*
282 *categorized as neutral. Yellow shading represents the 90% prediction interval. Note that TA:DIC*
283 *on the x-axis corresponds to pH_T and Ω_{ar} , as these variables have an approximately linear*
284 *relationship at a particular salinity, temperature and pressure (see Fig. 1).*

285 **2.5 Conceptual framework to evaluate increases in TA:DIC**

286 The regression models applied to each species could be used to predict calcification rates at higher
287 TA:DIC ratio. We conceptually added alkalinity from the current calcification rate baseline. This
288 baseline was computed for each species using CO2SYS with $pCO_2 = 425$ ppm and $pH_T = 8.1$, for
289 the average temperature and salinity for each species rate group, based on their respective OA
290 dataset(s) (see Supplemental Table 3). All CO2SYS calculations in this study were carried out
291 with the Python version of CO2SYS (Humphreys et al., 2022) using the stoichiometric dissociation
292 constants for carbonic acid from Sulpis et al. (2020), for sulfuric acid by Dickson et al. (1990) and
293 for total boron from Uppström (1974). From this baseline, TA was added in the form of both NaOH
294 and Na_2CO_3 to approximate changes in the the carbonate chemistry settings, with NaOH changing
295 TA:DIC in the 1:1 ratio, and Na_2CO_3 inducing a 2:1 TA:DIC change. For example, 10 $\mu\text{mol/kg}$ of
296 NaOH addition will increase TA by 10 $\mu\text{mol/kg}$ and not affect DIC. For Na_2CO_3 , 10 $\mu\text{mol/kg}$
297 addition will increase TA by 10 $\mu\text{mol/kg}$ and increase DIC by 5 $\mu\text{mol/kg}$. Figure 1 demonstrates
298 the usefulness of this approach. For both NaOH and Na_2CO_3 , 10 $\mu\text{mol/kg}$ was conceptually added
299 using the principles of mass balance approach for the carbonate system via CO2SYS. This was
300 repeated for increments of 50 $\mu\text{mol/kg}$. We show this incremental addition in the plots up to a total
301 of 500 $\mu\text{mol/kg}$ when generating the plots. When computing the thresholds, we added up to 1400
302 $\mu\text{mol/kg}$ NaOH. The new TA:DIC ratios were estimated by adding the direct effect of ΔTA and
303 ΔDIC due to chemical additions of NaOH (assume $\Delta\text{DIC} = 0$) or Na_2CO_3 (assume $\Delta\text{DIC} =$
304 $0.5 * \Delta\text{TA}$). A maximum of 500 $\mu\text{mol/kg}$ was chosen to have more realistic additions of TA that
305 resemble those appropriate within the OAE field trials (e.g. Wang et al., 2023). With the new
306 TA:DIC ratios after TA addition, the species' regression models based on the fitted OA response

307 data were used to compute respective calcification rates (note that added points with NaOH or
308 Na₂CO₃ were not calculated as part of the regression). These data points were all plotted along
309 with the experimental data, regression model and prediction intervals as shown in Fig. 3.

310 We also determine the amount of NaOH needed to reach pH_T 9 for each study. This was computed
311 for each species rate group using CO₂SYS starting from pCO₂ = 425 ppm and pH_T = 8.1, using
312 the average temperature and salinity, and by adding NaOH in increments of 50 μmol/kg until pH_T
313 9 was reached.

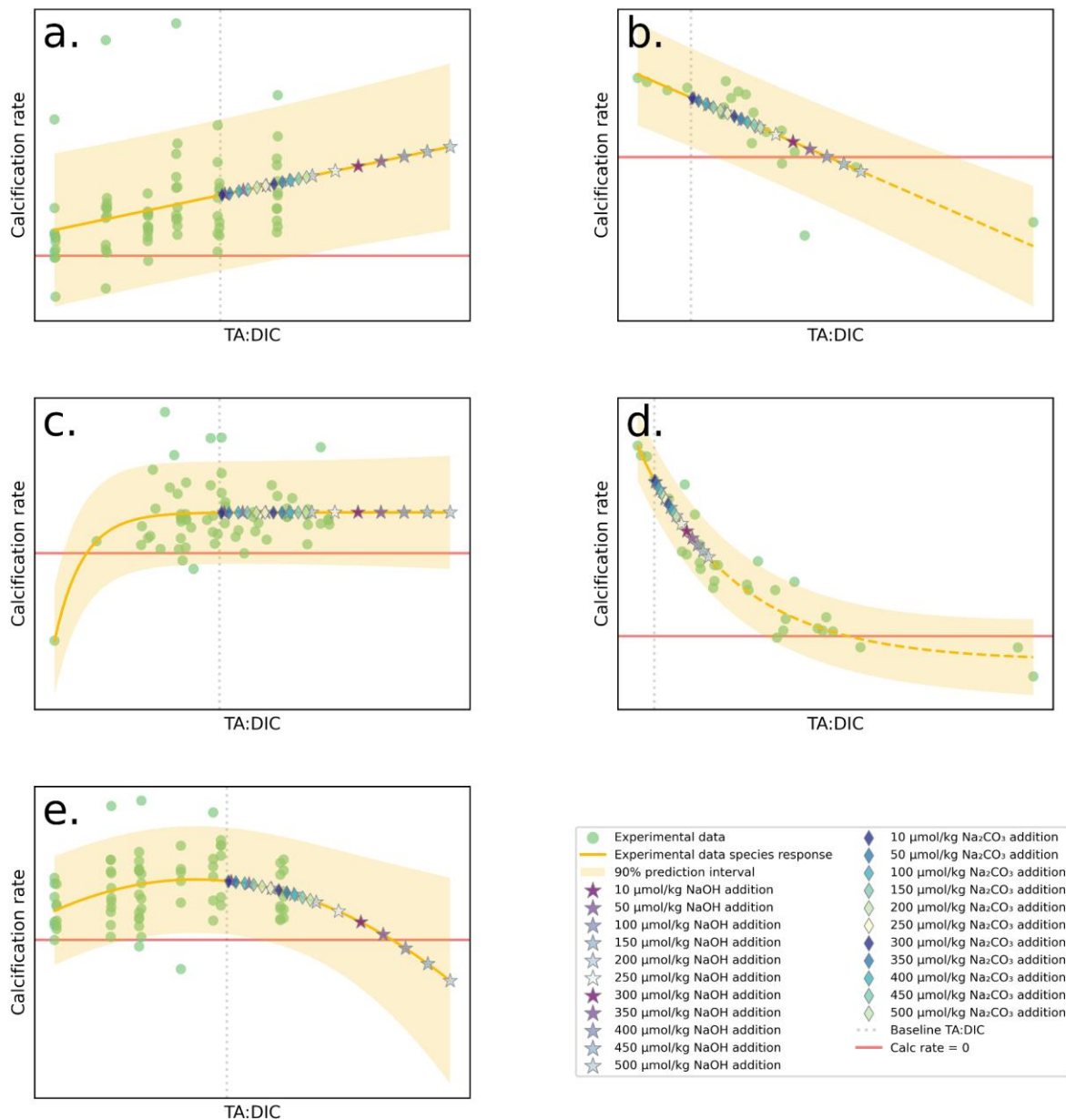
314 **2.6 Evaluation of the biological responses based on alkalinity addition**

315 The individual species with significant correlations were grouped visually based on their best-fit
316 regression models and are classified into positive, negative, and neutral as the following:

317 1) *Positive responders*: species with predicted *linear positive* and *threshold positive* calcification
318 rate response with increased TA addition.

319 2) *Negative responders*: species with predicted *linear negative*, *parabolic* and *threshold negative*
320 *response* in calcification rate upon (a certain amount of) TA addition. For the parabolic responders,
321 a concentration of NaOH was determined that indicates the threshold in TA:DIC beyond which
322 the response becomes negative (see inflection points in Supplemental Table 2).

323 3) *Neutral responders*: species with *no significant correlation* ($p < 0.05$) in calcification rate upon
324 TA addition.



325

326 **Figure 3:** Conceptual diagrams for five types of responses; a) linear positive; b) linear negative;
 327 c) threshold positive; d) threshold negative and e) parabolic response, plotted with experimental
 328 data from OA studies (green dots), predicted values at various additions of alkalinity (stars and
 329 diamonds), the regression line and prediction error margins fitted for a given species. The red
 330 horizontal line indicates zero net dissolution (calcification rate is equal to 0; dissolution rate =
 331 calcification rate). The grey vertical line indicates the baseline from which alkalinity is added.
 332 NaOH and Na_2CO_3 addition is shown up to 500 $\mu\text{mol/kg}$.

333 **2.7 Determining threshold values indicative of negative biological response to OAE**

334 The metrics to evaluate the sensitivity of calcification rate of the negative responders in this study
335 were based on the amount of NaOH or Na₂CO₃ addition required to reduce the current calcification
336 rate by a half. The greater the TA:DIC ratio value was required to trigger half calcification rate
337 reduction, the less sensitive species was to NaOH addition. We refer to this TA:DIC ratio as the
338 biological threshold, which we also report along with corresponding pH and Ω_{ar} and the associated
339 uncertainty. TA:DIC thresholds were converted to their respective pH and Ω_{ar} , which are affected
340 by temperature and salinity. To calculate threshold pH and Ω_{ar} we used the average temperature
341 and salinity per species rate group, as done for calculating the baseline.

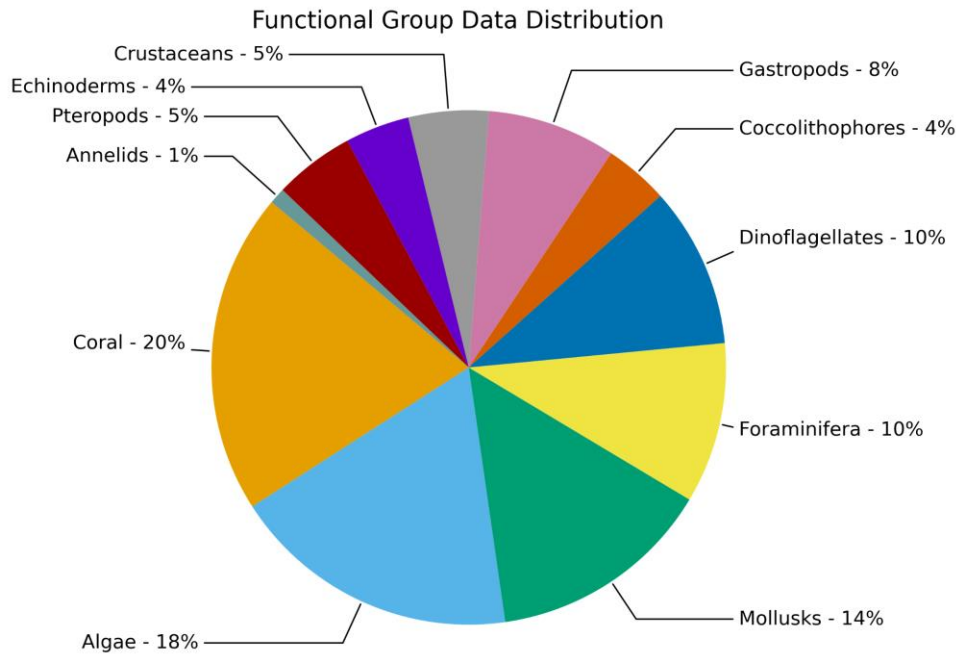
342 **2.8 Extraction of the carbonate chemistry data from the GLODAP dataset**

343 We extracted total alkalinity, dissolved inorganic carbon, Ω_{ar} , and pH_T from the Global Ocean
344 Data Analysis Project GLODAPv2.2023 dataset (<https://glodap.info>). We used the regression
345 application in MATLAB with a second-order polynomial equation to predict Ω_{ar} from the TA:DIC.
346 The regression analysis was performed using data from various depth intervals (0–10m, 0–30m,
347 0–50m, 0–100m, 0–200m) regionally and globally. The regional analysis divided the global oceans
348 into the following groupings: Arctic (north of 65°N), Southern (south of 40°S), North Pacific
349 (north of 40°N), Central Pacific (40°S to 40°N), North Atlantic (North of 40°N), Central Atlantic
350 (40°S to 40°N), and Indian Ocean (north of 40°S).

351 **3. Results**

352 **3.1 Data collection for the calcification rate responses of different biological groups**

353 We examined 68 datasets, which covered 84 different species that were divided into 11 different
354 groups (Fig. 4). These functional groups were corals (20% of datasets), calcifying algae (18%),
355 mollusks (14%), foraminifera (10%), dinoflagellates (10%), coccolithophores (4%), gastropods
356 (8%), crustaceans (5%), echinoderms (4%), pteropods (5%), and annelids (1%). In the mollusks
357 group, we have separated out the gastropod and pteropod because of a higher number of studies
358 that explicitly cover these two groups. The group of gastropods refers to all gastropods that are not
359 pteropods. If all three groups were combined (mollusks, gastropods, pteropods), this group would
360 be the largest.



362

363 **Figure 4:** Percent of studies for multiple groups ($N=11$) with available data for the calcification
 364 rate responses as part of data compilation of 68 studies covering 84 species).

365 3.2 Species-specific responses to NaOH/Na₂CO₃ addition

366 Calcification rate responses of species from different groups were correlated to TA:DIC and
 367 summarized to obtain calcification rate response. The calcification rate responses encompassed
 368 linear (positive and negative), threshold (positive and negative), parabolic, and neutral responses,
 369 with the slope and the intercept of the response determining the type and the magnitude of the
 370 response. We present fitted responses of calcification rate per TA:DIC ratio for each examined
 371 species (Table 1; Supplemental Fig. 4). When possible, we fit a regression to multiple datasets of
 372 the same species that used the same calcification units. We also present the response with the
 373 additions of NaOH and Na₂CO₃ for each species per examined study and corresponding rate unit
 374 and their biological TA:DIC thresholds (Table 2; Supplemental Table 4).

375 **Table 1:** The summary of all the OA studies from which the chemical and biological data was
 376 collected, including the name of the species and group and the accompanying calcification rate
 377 unit. The response for each species rate group was determined by the regression with the lowest

378 *p*-value, where the *p*-value was smaller than 0.05. We also include the *p*-value, goodness of fit (R^2)
379 and Root Mean Square Error (RMSE). Non-significant responses are categorized as having a
380 'neutral' response. The type of response (linear positive or negative, threshold positive or
381 negative, parabolic, and neutral) is indicated, as well as if this response is positive, negative or
382 neutral.

Studies	n	Group	Species	Rate unit	Response	Pos/Neg/ Neut	p-value	R ²	RMSE
Vasquez-Elizondo et al. (2016)	4	Algae	<i>Amphiroa tribulus</i>	mmol/m ² /hr	neutral	Neutral			
Sinutok et al. (2011)	16	Algae	<i>Halimeda cylindracea</i>	mmol/hr	neutral	Neutral			
Comeau et al. (2013)	71	Algae	<i>Halimeda macroloba</i>	mmol/g/hr	parabolic	Negative	0.0127	0.1200	0.0028
Meyer et al. (2015)	24	Algae	<i>Halimeda macroloba</i>	mmol/m ² /hr	neutral	Neutral			
Sinutok et al. (2011)	16	Algae	<i>Halimeda macroloba</i>	mmol/hr	parabolic	Negative	0.0108	0.5000	0.0001
Comeau et al. (2013)	62	Algae	<i>Halimeda minima</i>	mmol/g/hr	neutral	Neutral			
Meyer et al. (2015)	24	Algae	<i>Halimeda opuntia</i>	mmol/m ² /hr	linear +	Positive	0.0080	0.2800	0.0222
Comeau et al. (2013)	72	Algae	<i>Hydrolithon reinboldii</i>	mmol/g/hr	linear +	Positive	0.0053	0.1100	0.0026
Cornwall et al. (2018)	23	Algae	<i>Hydrolithon reinboldii</i>	mmol/m ² /hr	neutral	Neutral			
Comeau et al. (2013)	72	Algae	<i>Lithophyllum flavescens</i>	mmol/g/hr	neutral	Neutral			
Johnson et al. (2021)	420	Algae	<i>Lithophyllum sp.</i>	mmol/g/hr	linear +	Positive	0.0000	0.1000	0.1136
Vasquez-Elizondo et al. (2016)	4	Algae	<i>Lithothamnion sp.</i>	mmol/m ² /hr	neutral	Neutral			
Monserrat et al. (2022)	62	Algae	<i>Neogoniolithon brassica-florida</i>	mmol/m ² /hr	neutral	Neutral			
Ries et al. (2009)	42	Algae	<i>Neogoniolithon sp.</i>	mmol/g/hr	parabolic	Negative	0.0000	0.4100	0.0003
Vasquez-Elizondo et al. (2016), Comeau et al. (2018)	26	Algae	<i>Neogoniolithon sp.</i>	mmol/m ² /hr	neutral	Neutral			
Briggs-Carpenter et al. (2019)	425	Algae	<i>Porolithon onkodes</i>	mmol/m ² /hr	linear +	Positive	0.0010	0.0300	0.8093
Comeau et al. (2018, 2019)	64	Algae	<i>Sporolithon durum</i>	mmol/m ² /hr	parabolic	Negative	0.0012	0.2000	0.1704
Ries et al. (2009)	41	Annelid	<i>Hydroides crucigera</i>	mmol/g/hr	neutral	Neutral			
Fiorini et al. (2011), Langer et al. (2006, 2011)	14	Cocco.	<i>Calcidiscus leptoporus</i>	mmol/#/hr	neutral	Neutral			
*	233	Cocco.	<i>Emiliania huxleyi</i>	mmol/#/hr	parabolic	Negative	0.0000	0.1600	0.0000
Casareto et al. (2009)	14	Cocco.	<i>Pleurochrysis carterae</i>	mmol/m ³ /hr	neutral	Neutral			
White et al. (2018)	118	Cocco.	<i>Pleurochrysis carterae</i>	mmol/#	neutral	Neutral			
Meyer et al. (2016)	24	Coral	<i>Acropora millepora</i>	mmol/m ² /hr	neutral	Neutral			
Camp et al. (2017), Comeau et al. (2013)	74	Coral	<i>Acropora pulchra</i>	mmol/m ² /hr	parabolic	Negative	0.0000	0.2900	1.3257
Agostini et al. (2021)	18	Coral	<i>Acropora solitaryensis</i>	mmol/m ² /hr	neutral	Neutral			
Comeau et al. (2018), Comeau et al. (2019)	81	Coral	<i>Acropora yongei</i>	mmol/m ² /hr	linear +	Positive	0.0000	0.2900	1.9447
Bove et al. (2020)	27	Coral	<i>Duncanopsammia axifuga</i>	mmol/m ² /hr	linear +	Positive	0.0016	0.3300	5.0785
Cornwall et al. (2018)	44	Coral	<i>Goniopora sp.</i>	mmol/m ² /hr	neutral	Neutral			
Maier et al. (2009)	237	Coral	<i>Lophelia pertusa</i>	mmol/g/hr	linear +	Positive	0.0030	0.0400	0.0002
Bove et al. (2020)	65	Coral	<i>Montastraea cavernosa</i>	mmol/m ² /hr	linear +	Positive	0.0154	0.0900	0.5047
Ries et al. (2009)	54	Coral	<i>Oculina arbuscula</i>	mmol/g/hr	parabolic	Negative	0.0000	0.8600	0.0001
Comeau et al. (2013)	72	Coral	<i>Pavona cactus</i>	mmol/m ² /hr	parabolic	Negative	0.0002	0.2200	0.9093
Comeau et al. (2019)	49	Coral	<i>Plesiastrea versipora</i>	mmol/m ² /hr	linear +	Positive	0.0069	0.1500	0.6003
Brown et al. (2022)	4	Coral	<i>Pocillopora damicornis</i>	mmol/g/hr	neutral	Neutral			
Comeau et al. (2013, 2018), Putnam-Gates et al. (2015)	117	Coral	<i>Pocillopora damicornis</i>	mmol/m ² /hr	neutral	Neutral			

Studies	n	Group	Species	Rate unit	Response	Pos/Neg/ Neut	p-value	R ²	RMSE
Evensen-Edmunds et al. (2016)	60	Coral	<i>Pocillopora verrucosa</i>	mmol/m ² /hr	linear +	Positive	0.0132	0.1000	0.8297
Agostini et al. (2021)	18	Coral	<i>Porites heronensis</i>	mmol/m ² /hr	neutral	Neutral			
Comeau et al. (2013)	72	Coral	<i>Porites rus</i>	mmol/m ² /hr	linear +	Positive	0.0020	0.1300	2.0281
Okazaki et al. (2013)	75	Coral	<i>Siderastrea radians</i>	mmol/m ² /hr	linear +	Positive	0.0004	0.1600	2.7886
Okazaki et al. (2013)	64	Coral	<i>Solenastrea hyades</i>	mmol/m ² /hr	threshold +	Positive	0.0004	0.2300	2.0385
Krueger et al. (2017)	36	Coral	<i>Stylophora pistillata</i>	mmol/m ² /hr	neutral	Neutral			
Pansch et al. (2014)	36	Crust.	<i>Amphibalanus improvisus</i>	mmol/g/hr	linear +	Positive	0.0000	0.4300	0.0004
Ries et al. (2009)	36	Crust.	<i>Callinectes sapidus</i>	mmol/g/hr	linear -	Negative	0.0000	0.4000	0.0082
Ries et al. (2009)	18	Crust.	<i>Homarus americanus</i>	mmol/g/hr	linear -	Negative	0.0014	0.4800	0.0079
Ries et al. (2009)	12	Crust.	<i>Penaeus plebejus</i>	mmol/g/hr	linear -	Negative	0.0124	0.4800	0.0006
Findlay et al. (2010)	6	Crust.	<i>Semibalanus balanoides</i>	mmol/g/hr	neutral	Neutral			
Tatters et al. (2013)	45	Dino.	<i>Alexandrium sp.</i>	l/hr	neutral	Neutral			
Hansen et al. (2007)	19	Dino.	<i>Ceratium lineatum</i>	#/hr	linear -	Negative	0.0000	0.6700	0.0043
Tatters et al. (2013)	45	Dino.	<i>Gonyaulax sp.</i>	l/hr	neutral	Neutral			
Hansen et al. (2007)	31	Dino.	<i>Heterocapsa triquetra</i>	#/hr	threshold -	Negative	0.0000	0.9100	0.0027
Wang et al. (2019)	4	Dino.	<i>Karenia mikimotoi</i>	l/hr	neutral	Neutral			
Tatters et al. (2013)	45	Dino.	<i>Lingulodinium polyedrum</i>	l/hr	neutral	Neutral			
Tatters et al. (2013)	45	Dino.	<i>Prorocentrum micans</i>	l/hr	neutral	Neutral			
Hansen et al. (2007)	21	Dino.	<i>Prorocentrum minimum</i>	#/hr	threshold -	Negative	0.0000	0.8800	0.0019
Brading et al. (2011)	175	Dino.	<i>Symbiodinium sp.</i>	#/hr	linear -	Negative	0.0010	0.0600	0.0066
Van de Waal et al. (2013)	12	Dino.	<i>Thracosphaera heimii</i>	mmol/hr	parabolic	Negative	0.0002	0.8500	0.0000
Ries et al. (2009)	17	Echino.	<i>Arbacia punctulata</i>	mmol/g/hr	parabolic	Negative	0.0000	0.8900	0.0003
Courtney et al. (2013)	4	Echino.	<i>Echinometra viridis</i>	%/hr	linear +	Positive	0.0244	0.9500	2.3854
Courtney et al. (2015)	28	Echino.	<i>Echinometra viridis</i>	%	linear +	Positive	0.0009	0.3500	13.0388
Ries et al. (2009)	18	Echino.	<i>Eucidaris tribuloides</i>	mmol/g/hr	threshold +	Positive	0.0000	0.8400	0.0004
Keul et al. (2013)	205	Foram.	<i>Ammonia sp.</i>	mmol/#/hr	linear -	Negative	0.0277	0.0200	0.0000
Prazeres et al. (2015)	32	Foram.	<i>Amphistegina lessonii</i>	%/hr	parabolic	Negative	0.0008	0.3900	0.0010
Kisakurek et al. (2011)	16	Foram.	<i>Globigerinella siphonifera</i>	mmol/hr	neutral	Neutral			
Kisakurek et al. (2011)	14	Foram.	<i>Globigerinoides ruber</i>	mmol/#/hr	neutral	Neutral			
Reymond et al. (2013)	179	Foram.	<i>Marginopora rossi</i>	%/hr	linear +	Positive	0.0000	0.1900	0.0090
Uthicke-Fabricsius et al. (2012)	47	Foram.	<i>Marginopora vertebralis</i>	mmol/g/hr	threshold +	Positive	0.0000	0.4000	0.0004
Sinutok et al. (2011)	16	Foram.	<i>Marginopora vertebralis</i>	mmol/hr	neutral	Neutral			
Prazeres et al. (2015)	32	Foram.	<i>Marginopora vertebralis</i>	%/hr	linear -	Negative	0.0006	0.3300	0.0005
Manno et al. (2012)	192	Foram.	<i>Neogloboquadrina pachyderma</i>	mmol/#/hr	linear +	Positive	0.0000	0.7100	0.0000
Oron et al. (2020)	96	Foram.	<i>Operculina ammonoides</i>	mmol/g/hr	linear -	Negative	0.0031	0.0900	0.0017
Manriquez et al. (2016)	74	Gastropod	<i>Concholepas concholepas</i>	mmol/g/hr	linear +	Positive	0.0000	0.2400	0.0009
Noisette et al. (2016), Ries et al. (2009)	173	Gastropod	<i>Crepidula fornicata</i>	mmol/g/hr	parabolic	Negative	0.0000	0.2100	0.0028
Garilli et al. (2015)	68	Gastropod	<i>Cyclope neritea</i>	mmol/g/hr	linear -	Negative	0.0020	0.1400	0.0037
Ries et al. (2009)	42	Gastropod	<i>Littorina littorea</i>	mmol/g/hr	linear +	Positive	0.0001	0.3400	0.0002
Bibby et al. (2007)	4	Gastropod	<i>Littorina littorea</i>	µm (shell thickness)	neutral	Neutral			
Garilli et al. (2015)	315	Gastropod	<i>Nassarius corniculus</i>	mmol/g/hr	parabolic	Negative	0.0000	0.2500	0.0064
Ries et al. (2009)	21	Gastropod	<i>Strombus alatus</i>	mmol/g/hr	linear +	Positive	0.0000	0.6400	0.0001
Ries et al. (2009)	33	Gastropod	<i>Urosalpinx cinerea</i>	mmol/g/hr	linear +	Positive	0.0000	0.5700	0.0001
Ries et al. (2009)	18	Mollusks	<i>Argopecten irradians</i>	mmol/g/hr	linear +	Positive	0.0097	0.3500	0.0002
Ramajo et al. (2016)	6	Mollusks	<i>Argopecten purpuratus</i>	mmol/g/hr	neutral	Neutral			
Zhang et al. (2011)	5	Mollusks	<i>Azumapecten farreri</i>	mmol/g/hr	linear +	Positive	0.0106	0.9200	0.0001
Ong et al. (2017)	24	Mollusks	<i>Cerastoderma edule</i>	mmol/g/hr	neutral	Neutral			

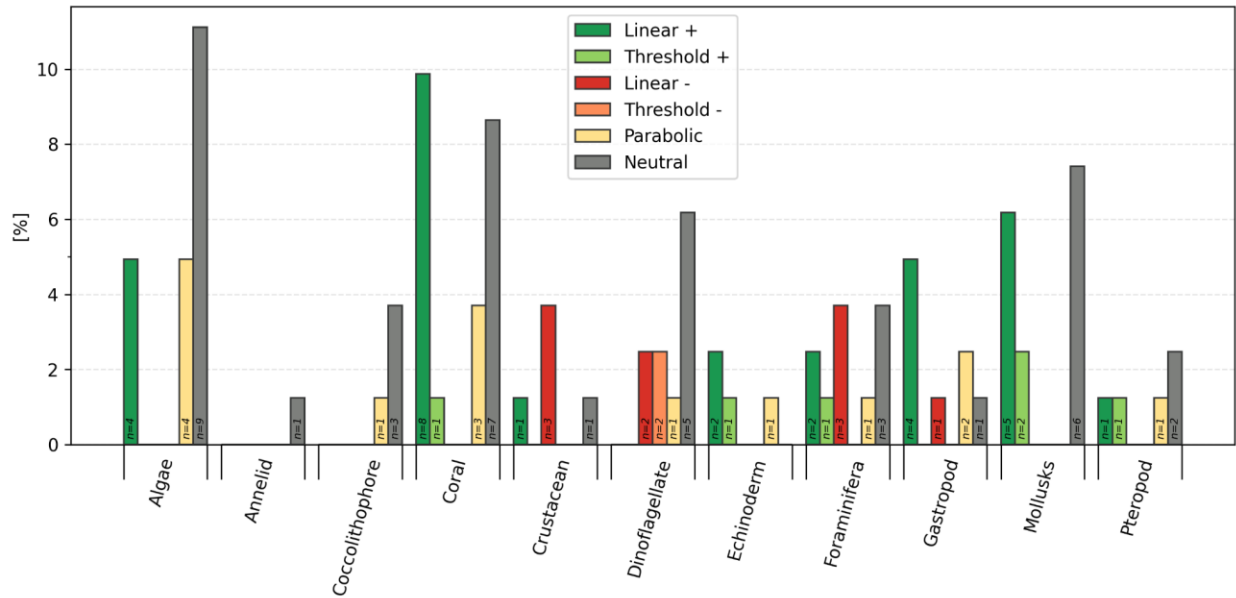
Studies	n	Group	Species	Rate unit	Response	Pos/Neg/ Neut	p-value	R ²	RMSE
Sordo et al. (2021)	27	Mollusks	<i>Chamelea gallina</i>	mmol/g/hr	neutral	Neutral			
Gazeau et al. (2007)	20	Mollusks	<i>Crassostrea gigas</i>	mmol/g/hr	linear +	Positive	0.0001	0.6100	0.0000
Ries et al. (2009), Waldbusser et al. (2011)	28	Mollusks	<i>Crassostrea virginica</i>	mmol/g/hr	threshold +	Positive	0.0000	0.5600	0.0003
Ries et al. (2009)	25	Mollusks	<i>Mercenaria mercenaria</i>	mmol/g/hr	threshold +	Positive	0.0000	0.8300	0.0000
Ries et al. (2009)	14	Mollusks	<i>Mya arenaria</i>	mmol/g/hr	linear +	Positive	0.0001	0.7300	0.0003
Ninokawa et al. (2020)	13	Mollusks	<i>Mytilus californianus</i>	mmol/m ² /hr	neutral	Neutral			
Ries et al. (2009), Gazeau et al. (2007)	86	Mollusks	<i>Mytilus edulis</i>	mmol/g/hr	linear +	Positive	0.0119	0.0700	0.0002
Gazeau et al. (2014)	11	Mollusks	<i>Mytilus galloprovincialis</i>	mmol/g/hr	neutral	Neutral			
Cameron et al. (2019)	30	Mollusks	<i>Pecten maximus</i>	mmol/g/hr	neutral	Neutral			
Comeau et al. (2010b)	5	Pteropod	<i>Cavolinia inflexa</i>	mm (shell length)	neutral	Neutral			
Comeau et al. (2009, 2010a)	12	Pteropod	<i>Limacina helicina</i>	mmol/g/hr	linear +	Positive	0.0000	0.8500	0.0001
Lischka et al. (2011, 2012)	119	Pteropod	<i>Limacina helicina</i>	mm (shell length)	threshold +	Positive	0.0003	0.1300	0.1303
Bednarsek (2021a), Mekkes et al. (2021)	117	Pteropod	<i>Limacina helicina</i>	µm (shell thickness)	parabolic	Negative	0.0000	0.1800	0.0038
Lischka et al. (2012)	28	Pteropod	<i>Limacina retroversa</i>	mm (shell length)	neutral	Neutral			

383 *Barcelos-Ramos et al. (2010), Fiorini et al. (2011), Iglesias-Rodriguez et al. (2008), Richier et al. (2011), Sciandra et al. (2003), Stoll et al. (2012),
384 Gafar et al. (2018), Bach et al. (2011), Sett et al. (2014).

385 Within each of the 11 functional groups, several categories of calcification response occur within
386 each functional group, with the most varied being the group of dinoflagellates and foraminifera,
387 both showing 4 or 5 different categories of calcification responses (Fig. 5). Of the six types of
388 responses of calcification rate vs. TA:DIC, 28% were linear positive (N=27), 9% linear negative
389 (N=9), 6% threshold positive (N=6), 2% threshold negative (N=2), 15% parabolic (N=14) and
390 40% neutral (N=38).

391 Such responses could be further summed up into positive (linear and threshold positive), negative
392 (linear and threshold negative, parabolic) and neutral responses (Fig. 6) when generalized for
393 calcification rate against TA:DIC ratio. A summary of responses includes 34.4% positive (N=33),
394 26.0% negative (N=25), while 39.6% show a neutral response (N=38).

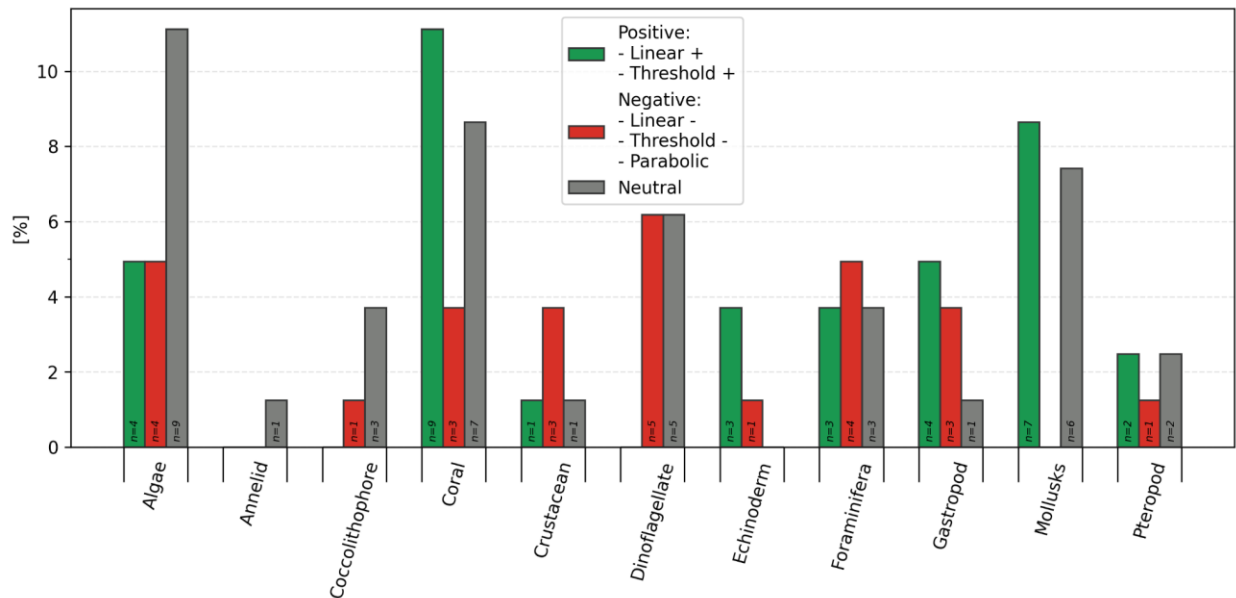
395



396

397 **Figure 5:** Categories of calcification rate responses and percentage (%) response across eleven
 398 groups (calcifying algae, annelids, coccolithophores, corals, crustaceans, dinoflagellate,
 399 echinoderms, foraminifera, gastropods, mollusks, pteropods). The number on the bar indicates
 400 the number of studies of species included.

401



402

403 **Figure 6:** Summary of percentage (%) responses in calcification rates as positive (linear and
 404 threshold positive), negative (linear and threshold negative, parabolic) and neutral across
 405 eleven groups (calcifying algae, annelids, coccolithophores, corals, crustaceans, dinoflagellate,

406 *echinoderms, foraminifera, gastropods, mollusks, pteropods*). The number on the bar indicates
407 the number of studies with species included.

408

409 **3.3 Evaluation of the responses to NaOH/Na₂CO₃ addition**

410 Upon added TA, the calcification rate in positive responders will increase, either in a linear or
411 threshold positive response, where calcification plateaus, with the concentration being dependent
412 on the species-specific rate of response (Fig. 2; Supplemental Fig. 4). The negative responders
413 (linear or threshold negative and parabolic) will be negatively impacted as follows: first, for the
414 linear negative responders, addition of the Na₂CO₃ will linearly decrease calcification rate, but
415 there is no associated threshold to it; second, for the threshold negative responders, calcification
416 rate will decline in an exponential way until reaching a TA:DIC value where the response plateaus;
417 and third, for the parabolic responders, the calcification rate will initially increase until reaching a
418 certain TA:DIC threshold upon which calcification starts declining. The TA:DIC thresholds for
419 negative responders are species-specific (Table 2; Supplemental Table 4).

420 **3.4 Threshold values indicative of negative biological response to OAE**

421 The TA:DIC biological thresholds in Table 2 are determined by the amount of NaOH addition
422 required to reduce calcification rate by a half (see Supplemental Table 4 for Na₂CO₃ thresholds).
423 These thresholds demonstrate the range of carbonate chemistry conditions over which the negative
424 biological effects of OAE deployment might occur and are shown alongside the corresponding
425 pH_T and Ω_{ar} . Uncertainties are higher for the experimental studies where the experimental
426 temperature and salinity ranges were high (see Supplemental Table 5), seeing as we use the average
427 for each species rate group to compute the baseline and thresholds.

428 For the negative responders, TA:DIC thresholds range from 1.13 to 1.74. The majority of species
429 have reached their thresholds by an addition of 500 $\mu\text{mol/kg}$ NaOH, though for 3 species a NaOH
430 addition of more than 500 $\mu\text{mol/kg}$ is required to cross the thresholds in the TA:DIC range of 1.39
431 to 1.74. *Crepidula fornicata* (gastropod), *Neogoniolithon sp.* (algae), *Homarus americanus*
432 (crustacean) and *Oculina arbuscula* (coral) reach their thresholds by 100 $\mu\text{mol/kg}$ addition of
433 NaOH, indicating they are more sensitive to alkalinity addition. Foraminifera, dinoflagellates and

434 coccolithophores generally require higher concentrations of NaOH to reach their thresholds, with
435 the linear negative responder *Ammonia sp.* of the foraminifera group requiring 1400 $\mu\text{mol/kg}$ to
436 reduce calcification rate in half.

437 For some negative responders (*Arbacia punctulata*, *Nassarius corniculus*, *Penaeus plebejus*,
438 *Callinectes sapidus*, *Cyclope neritea*, and *Symbiodinium sp.*), the baseline from which NaOH
439 addition occurs was outside of the range of the experimental data and very close to a calcification
440 rate of 0. These were omitted from Table 2 since our defined threshold does not give an accurate
441 representation of their sensitivity to alkalinity addition. *Limacina helicina* was also omitted since
442 the indicator of calcification (shell thickness) was not an actual rate.

443 **Table 2:** *Studies with negative responders (linear and threshold negative, parabolic) with*
444 *demonstrated TA:DIC thresholds, indicating the amount of NaOH needed to halve the current*
445 *calcification rate (i.e. at the baseline). The value for TA:DIC threshold is used to determine the*
446 *pH_T and Ω_{ar} (at average temperature and average salinity per species). See Supplemental Table 4*
447 *for Na_2CO_3 thresholds.*

Studies	Group	Species	Temp (°C)	Salinity	Rate unit	Thresh old	TA addition	pH _T at threshold	ΔpH _T from baseline	Ω _{ar} at threshold	Exposure time
Noisette et al. (2016), Ries et al. (2009)	Gastropod	<i>Crepidula fornicata</i>	15.31	34.33	mmol/g/hr	1.13	50	8.17	0.07	3.77	6 months 60 days
Ries et al. (2009)	Algae	<i>Neogoniolithon sp.</i>	25.00	31.70	mmol/g/hr	1.17	50	8.16	0.06	4.87	60 days
Ries et al. (2009)	Crustacean	<i>Homarus americanus</i>	25.02	31.96	mmol/g/hr	1.19	100	8.22	0.12	5.49	60 days
Ries et al. (2009)	Coral	<i>Oculina arbuscula</i>	25.01	31.61	mmol/g/hr	1.19	100	8.22	0.12	5.46	60 days
Prazeres et al. (2015)	Foraminifera	<i>Amphistegina lessonii</i>	24.18	33.46	%/hr	1.21	150	8.27	0.17	6.10	30 days
Hansen et al. (2007)	Dinoflagellate	<i>Ceratium lineatum</i>	15.00	30.00	#/hr	1.18	200	8.38	0.28	5.15	14 d acclimation to irradiance; 7 days acclimation to experimental conditions; 14 days exposure to irradiance; 22 days stationary growth phase
Sinutok et al. (2011)	Algae	<i>Halimeda macroloba</i>	27.23	36.27	mmol/g/hr	1.26	200	8.30	0.20	7.38	2 weeks acclimation, 2 weeks incubation
Comeau et al. (2019)	Algae	<i>Sporolithon durum</i>	20.60	35.87	mmol/m ² /hr	1.22	200	8.32	0.22	6.31	27 weeks
Van de Waal et al. (2013)	Dinoflagellate	<i>Thoracosphaera heimii</i>	15.00	34.00	mmol/hr	1.23	300	8.46	0.36	6.56	21 days acclimation, 8 days experiment = total of >10 generations
Oron et al. (2020)	Foraminifera	<i>Operculina ammonoides</i>	25.00	37.00	mmol/g/hr	1.33	400	8.46	0.36	9.44	65 - 120 hours
Prazeres et al. (2015)	Foraminifera	<i>Marginopora vertebralis</i>	24.18	33.46	%/hr	1.33	450	8.53	0.43	9.78	30 days
Camp et al. (2017), Comeau et al. (2013)	Coral	<i>Acropora pulchra</i>	27.30	36.27	mmol/m ² /hr	1.38	500	8.52	0.42	11.05	n/a (natural conditions) 2 weeks acclimation; 2 weeks incubation
Hansen et al. (2007)	Dinoflagellate	<i>Heterocapsa triquetra</i>	15.00	30.00	#/hr	1.30	500	8.66	0.56	8.81	14 d acclimation to irradiance; 7 days acclimation to experimental conditions; 14 days exposure to irradiance; 22 days stationary growth phase
Comeau et al. (2013)	Coral	<i>Pavona cactus</i>	27.23	36.28	mmol/m ² /hr	1.38	500	8.52	0.42	11.03	2 weeks acclimation; 2 weeks incubation
Hansen et al. (2007)	Dinoflagellate	<i>Prorocentrum minimum</i>	15.00	30.00	#/hr	1.39	700	8.81	0.71	11.35	14 d acclimation to irradiance; 7 days acclimation to experimental conditions; 14 days exposure to irradiance; 22 days stationary growth phase
*Coccolithophore		<i>Emiliana huxleyi</i>	17.30	35.12	mmol/#/hr	1.46	850	8.83	0.73	13.65	**
Keul et al. (2013)	Foraminifera	<i>Ammonia sp.</i>	26.00	32.75	mmol/#/hr	1.74	1400	9.11	1.01	22.27	59-96 days of culturing

448 *Barcelos-Ramos et al. (2010), Fiorini et al. (2011), Iglesias-Rodriguez et al. (2008), Richier et al. (2011), Sciandra et al. (2003), Stoll et al. (2012),
449 Gafar et al. (2018), Bach et al. (2011), Sett et al. (2014).

450 **26hrs, Acclimation for 7 generations, experiment/sampling for 2-3 generations, n/a, 8 days, 16 days, Acclimation for 12 generations, Pre-
451 acclimation for 8-12 generations, 9 generations, Acclimated for at ~7 generations (5-15 days)

452 3.5 Regulatory pH_T 9 threshold

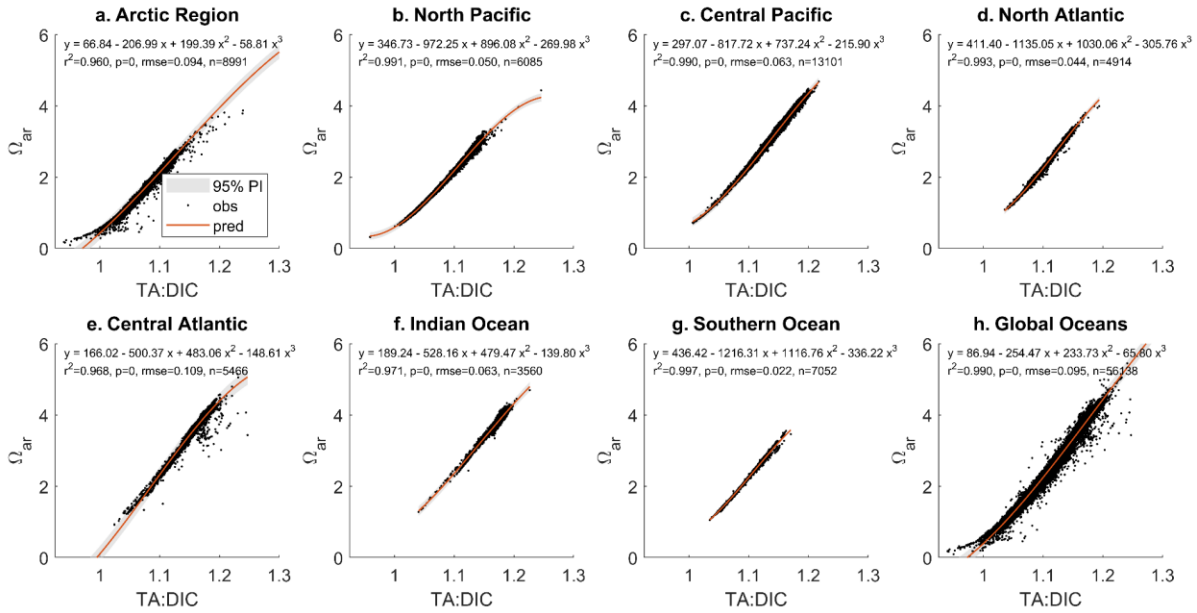
453 We also compute how much NaOH needs to be added before reaching a pH_T threshold of 9, as per
454 the US Environmental Protection Agency's rule for waste water not exceeding a pH_T of 9 when
455 entering the coastal ocean (NPDES manual, 2010). This amount averages at 1200 μmol/kg of
456 NaOH for most of the examined species. For some species (*Amphibalanus improvisus*,
457 *Neogloboquadrina pachyderma*, *Limacina helicina*, *Limacina retroversa*, *Lophelia pertusa*, and

458 *Semibalanus balanoides*), their threshold was reached below 1000 $\mu\text{mol/kg}$, with *Amphibalanus*
459 *improvisus* reaching a threshold at 750 $\mu\text{mol/kg}$.

460 **3.6 Global and regional carbonate chemistry data coverage based on GLODAP datasets**

461 The compilation of chemical observational data (pH, Ω_{ar} , TA, DIC) was done for the GLODAP
462 data across the regional ocean and global scales to determine the range of Ω_{ar} , TA and DIC (as
463 represented by the TA:DIC ratio) and TA:DIC vs Ω_{ar} correlation down to the depths averaged over
464 200 m. This allowed us to apply the thresholds even for the regions for which we do not have
465 sufficient or reliable data or experimental coverage, making the inferences about the OAE impact
466 even in those regions.

467 Here, we focused on showing the results ranging over the 0–50m because this covers most of the
468 biological habitat for examined species and it is where the OAE enhancement would induce the
469 greatest changes. Over the 0–50 m depth, Ω_{ar} ranges from 0.2 to 5 and TA:DIC ranges from 0.1 to
470 1.25 and both parameters are correlated across all the regions, as demonstrated by the fitted second-
471 order polynomial regressions, with R^2 of 0.96 or higher, and all the correlations being significant
472 (Fig. 7), with regional specific relationships not impacting the fit. All the correlation parameters
473 are presented in Supplemental Table 4. Similar fits were found at different depths. The conditions
474 in the higher latitude regions are located at the lower range of Ω_{ar} vs TA:DIC, while the conditions
475 in the low latitudes and temperate regions are at the upper range, with the highest values present
476 in the central Atlantic and Pacific region. Such strong correlation as observed for Ω_{ar} vs TA:DIC
477 does not exist with pH, regardless of the depth interval examined. While the correlations are still
478 significant, they are broadly distributed and represented over a shorter TA:DIC range, with
479 significantly lower goodness of fit (Supplemental Fig. 4), with the correlations being highly
480 regionally dependent due to pH and temperature co-linearity. Because of this, all further biological
481 analyses are only done using the Ω_{ar} vs TA:DIC ratio.



482

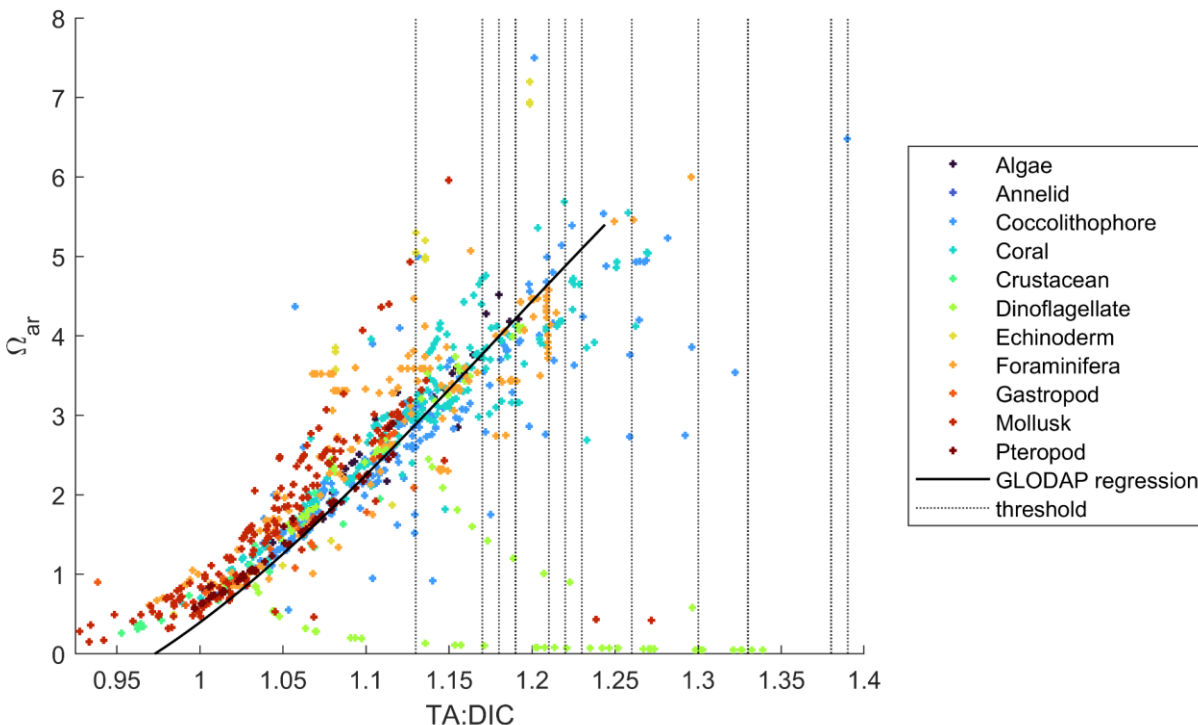
483 **Figure 7:** The range of observed Ω_{ar} , TA and DIC values (as represented by the TA:DIC ratio)
 484 values and the relationship with the best fitted curve between Ω_{ar} vs TA:DIC across regional (a-
 485 g) and global (h) scales based on the observational GLODAP dataset averaged over the 0-50 m
 486 depth range.

487 3.7 TA:DIC vs Ω_{ar} for experimental data and GLODAP

488 We compared the ranges of TA:DIC and Ω_{ar} of biological experimental data with field
 489 biogeochemical data (GLODAP) to examine if similar range of conditions and TA:DIC
 490 correlations are applicable over a broader, global dataset. For this, we plotted Ω_{ar} vs TA:DIC along
 491 with the GLODAP regression line for Ω_{ar} vs TA:DIC (Fig. 8). For each TA and DIC datapoint, the
 492 corresponding salinity and temperature specific values for that data point were used to compute
 493 Ω_{ar} . We show the similarity in the conditions, which gives the validity of our experimentally-
 494 derived thresholds to be extrapolated within the global GLODAP dataset.

495 Figure 8 also shows that various biological groups are clustered around specific TA:DIC ratios,
 496 for example, mollusks, coral and coccolithophores are represented on the lower, mid, and higher
 497 TA:DIC spectra, respectively, while dinoflagellates are randomly scattered off the TA:DIC line.
 498 This indicates that there is a general lack of data distribution in the upper ranges of TA:DIC ratio,
 499 especially for the groups that are lying at the lower and mid end of the TA:DIC ratio spectra.

500 Plotting biological data from the OA datasets against the regional and global TA:DIC gradient
501 derived from GLODAP (Fig. 7), we also observed that experimental data ranges were not always
502 consistent with natural conditions, for example, having a lower Ω_{ar} at a higher TA:DIC ratio.



503
504 **Figure 8:** Ω_{ar} values from experimental biological studies for eleven investigated functional
505 groups (see legend) plotted against TA:DIC, with the latter being computed using experimental
506 TA and DIC. The black line represents the regression line of TA:DIC and Ω_{ar} data from the
507 GLODAP dataset (covering 0-50m depth). See Supplemental Fig. 5 for GLODAP Ω_{ar} vs TA:DIC,
508 from which the black regression line shown here is derived. The vertical dotted lines represent
509 the thresholds shown in Table 2.

510 4. Discussion

511 OAE is a quickly developing strategy that is in the field-testing phase despite extremely limited
512 understanding of the sequestration potential, biological implications and environmental concerns.
513 Hence, gaining insights of potential risks for the biological species and communities is essential
514 and timely. In retrospect, it took decades for the OA research community to get a more accurate
515 and comprehensive understanding leading to predictions of biological responses to OA (Riebesell

516 and Gattuso, 2015). Without a very clear conceptual strategy for the OAE testing, the research
517 community might also need years to decades before OAE-related implications are
518 comprehensively understood. Consequently, there is an essential need to develop an assessment
519 framework of predictive responses and testing strategies that will assist in OAE scaling and risk
520 avoidance. This paper aims at developing such an assessment, where calcification responses
521 against TA:DIC are categorized per species. We propose to use the TA:DIC ratio in the biological
522 studies reporting OAE results, as we believe it simplifies the system and makes it easier to use and
523 translate the carbonate chemistry in the experimental setting. Such a TA:DIC ratio allows to
524 ultimately standardize the biogeochemical and biological data and is useful for easier comparisons
525 among the conducted experiments.

526 **4.1. Identified strengths and limitations of the synthesis approach based on OA studies**

527 Prior to conducting this study, several drawbacks were identified that could potentially limit such
528 a synthesis work: first, an insufficient amount of data at the upper range of carbonate chemistry
529 conditions (high pH, high Ω_{ar}); second, experimental data under conditions with no relevance to
530 natural settings (Fig. 8); and third, an insufficient number of validation studies under high TA
531 conditions to validate the results of this synthesis. To overcome the first two limitations, the
532 decision was made to combine multiple OA datasets for a single species with the aim to achieve a
533 greater range in carbonate chemistry conditions, including higher pH, Ω_{ar} experimental values,
534 which should reduce the uncertainty of the predictions. However, combining raw data on species
535 calcification rate proved to be more challenging because even across the same species the reporting
536 of the calcification rates was highly variable. The use of different measuring approaches of
537 calcification rates while conducting OA studies generated data with divergent units that do not
538 allow for the intercomparison of data and results. As different studies for a single species could
539 not be combined, we chose to increase the number of studies and thus, the number of examined
540 species. Based on the response categories from the OA studies (Ries et al., 2009), our hypothesis
541 was that OAE will elucidate the same categories of responses, i.e. positive, negative and neutral.
542 Within each of the groups examined, multiple categories of predicted calcification response were
543 found. In this way, we demonstrated that it was possible to develop a useful framework for
544 assessing and predicting species-specific OAE responses that can delineate different responders,
545 identify species with greater OAE sensitivity and determine the thresholds where such negative

546 responses could happen.

547 **4.2 Synthesizing biological response under OAE identifies positive and negative responders**

548 The responses were summarized across three emerging groups of responses: positive, negative,
549 and neutral (Fig. 6). We observe species-specific variability at the species level, which is related
550 to various calcification mechanisms across the observed groups. The greatest variability upon
551 NaOH addition within each group in calcification rate was evident in corals, dinoflagellates,
552 foraminifera, gastropods and pteropods, where four to five different categories of responses were
553 found.

554 Positive responders (34%) show an increased calcification rate upon alkalinity addition, observed
555 within all functional groups besides annelids, coccolithophores and dinoflagellates. Corals mostly
556 have positive and neutral responses, suggesting that coral species would not be negatively
557 impacted during OAE field trials. This mostly positive response is validated by increased coral
558 calcification, shown for two coral species of *Acropora* and *Siderastre* in experiments conducted
559 by Palmer et al. (2022).

560 The metrics to evaluate the sensitivity of calcification rate for the negative responders (negative
561 linear and threshold) to alkalinity addition was based on the amount of alkalinity addition required
562 to halve the current calcification rate (Fig. 3; Tables 1, 2). The most negative responses were found
563 in dinoflagellates (6% of all species), algae and foraminifera (both 5% of all species). However,
564 these numbers are affected by the difference in data coverage per functional group. When
565 comparing the ratio of negative to positive and neutral responses, crustaceans and dinoflagellates
566 are expected to be most negatively affected. As such, these groups are one of the priorities for the
567 future OAE experimental work to determine at which TA:DIC negative response happens.
568 Dinoflagellates demonstrate negative response in 5 cases, 5 neutral responses and 0 positive (see
569 Table 1; Supplemental Fig. 4). The reason for negative response to OAE in this group is related to
570 the fact that their growth gets limited at higher pH, with further carbon limitation playing a role at
571 very high pH levels and low DIC concentration (Hansen et al., 2002; 2007). On the other hand,
572 crustaceans only demonstrated positive response in one study (Pansch et al., 2014), while
573 remaining results predict either negative or neutral response. While crustaceans are effective in
574 retaining homeostasis at lower pH, they might be less so at higher pH, which was shown in the

575 OA experiments by Ries et al. (2009) for three crustacean species (*Callinectes sapidus*, *Homarus*
576 *americanus*, *Penaeus plebejus*), confirmed in the OAE study by Cripps et al. (2013) in *Carcinus*
577 *maenas*. While studies are still lacking, physiological acid-base regulation at higher pH is
578 associated with higher costs (Cripps et al., 2013). Crustaceans show a disrupted acid–base balance,
579 evident through the increase in hemolymph pH, K⁺, Na⁺ ions and osmolality, coupled with a
580 decrease in extracellular pCO₂ and HCO₃⁻, indicative of respiratory alkalosis (Truchot,
581 1984;1986). This is often associated with hyperventilation, the aim of which is to flush out the
582 hemolymph CO₂ to increase the affinity of oxygen uptake. However, while this might be a
583 temporary physiological relief it also implies energetic costs, potentially also for calcification.

584 For the neutral responders or groups with no significant correlation between calcification rates and
585 TA:DIC, it is somewhat uncertain to predict if such responses will be retained under OAE. While
586 parabolic responders show a physiologically understandable parabolic type of dose-response,
587 positioning the TA:DIC values where the threshold occurs is also highly species-specific and
588 potentially uncertain, meaning that it might depend on other environmental factors.

589 With respect to the coccolithophores, we note that this was the only group where data compilation
590 on calcification rate across the group was possible because the OA studies were conducted in a
591 more uniform way, using similar approaches, and reporting the result in the same units. When data
592 for *E. huxleyi* across the comparable studies was compiled (Barcelos-Ramos et al., 2010; Fiorini
593 et al., 2011; Iglesias-Rodrigues et al., 2008; Sciandra et al., 2003; Stoll et al., 2012; Richier et al.,
594 2011), a significant parabolic response was obtained (Table 1), although the goodness of fit was
595 fairly low ($R^2=0.16$). Despite lower R^2 , we decided to use the compiled dataset because of the
596 increased statistical power. The parabolic response obtained aligns with Langer et al. (2006) and
597 also with the parabolic type responses found in the synthesis studies by Paul and Bach (2020) and
598 Bach et al. (2015). The threshold indicates the mechanisms of coccolithophore growth that are
599 driven by CO₂, which is shown to decline with alkalinity addition. The threshold based on all
600 studies for *E. huxleyi* combined was positioned at a TA:DIC of 1.46 ($\Omega_{ar} = 13.65$, see Table 2),
601 which would be triggered at 850 $\mu\text{mol/kg}$ of added NaOH and at a pCO₂ of 60 μatm .
602 Comparatively with the phytoplanktonic diatoms, such growth limitation is predicted at a pCO₂
603 amount at 100 μatm (Riebesell et al., 1993). It is important to note that when these studies were
604 analyzed individually, a mixture of different responses was observed. We emphasize the variability

605 within the coccolithophore responses, which are species-specific and inherently related to the
606 strain adaptation to their innate regional settings and dependent on a variety of other factors (Bach
607 et al., 2015; Gafar and Schultz, 2018), including the longevity of the species, the experimental
608 settings used in the study (e.g. nutrient-replete vs nutrient deficient conditions) and the presence
609 or absence of (un)suitable light conditions. Interestingly, for all the coccolithophore species other
610 than *E. huxleyi*, responses were neutral. For validation purposes, the results of our study could not
611 be compared, either because the calcification rates were not studied or the calcification units were
612 not comparable (e.g. Diner et al., 2015).

613 **4.3 Parameters impacting derivation of thresholds and their application**

614 We developed a set of species-specific thresholds in this study, with demonstrated application
615 across the global Ω_{ar} vs TA:DIC conditions (Table 2; Fig. 8). The range of alkalinity additions to
616 result in a threshold of 50% decline in calcification rate varied significantly between the species
617 and the type of response. The TA:DIC thresholds upon TA application ranged between 50 to 1400
618 $\mu\text{mol/kg}$ of NaOH addition, and the pH_T 9 thresholds averaged at 1200 $\mu\text{mol/kg}$ of NaOH for all
619 species. However, there are many parameters that impact threshold derivation and application,
620 which we discuss in greater detail.

621 First, we note that differences in experimental conditions for different species make it difficult to
622 directly compare different species' thresholds among each other. Instead, they are intended to
623 delineate sensitivity to alkalinity addition of individual species at given experimental conditions.
624 In the case that the lab experimental conditions mimic species' natural habitat, this threshold-
625 related sensitivity can be extrapolated to their natural habitats.

626 Second, we emphasize that the threshold application should not only consider the magnitude of
627 NaOH added, but also the duration or exposure time of the experimental study. As such, when
628 applying the thresholds to respective model outputs or observation data, both duration and
629 exposure time should be considered. For all the derived thresholds, we have added duration
630 exposure information to Table 2. Additional parameters that need to be included when applying
631 these thresholds are related to local temperature and salinity. The extracted threshold values are
632 calculated with the temperature and salinity from the experimental conditions, which means that
633 this threshold should not be applied to very different conditions without adjusting for salinity and

634 temperature.

635 Third, we assumed global surface ocean conditions to be standardized at a $p\text{CO}_2$ of 425 ppm and
636 a pH_T of 8.1 as a control point for OAE compound additions. However, we note that in different
637 habitats, pH_T 8.1 may not represent the baseline from where OAE should be considered adding,
638 because the average pH might be different. This means that the amount of TA required to reach a
639 certain threshold could vary and is dependent on the baseline carbonate chemistry at the site of
640 deployment and its variability. This is especially relevant in habitats with a lower baseline pH,
641 where more TA would need to be added for the threshold to be reached, meaning less negative
642 biological implications.

643

644 In addition, physical parameters of importance are related to the dilution effect, mixing, retention
645 capacity, as well as the rate of the equilibration effects of the air-sea CO_2 uptake (Ferderer et al.,
646 2022; He and Tyka, 2023; Schulz et al., 2023; Wang et al., 2023), because they determine relevant
647 exposure duration and the variability of carbonate chemistry parameters across spatial and vertical
648 depths. Therefore, to obtain the most accurate and regionally applicable threshold for the species
649 of interest, it is recommended that the baseline for OAE additions be determined based on local
650 conditions.

651 Lastly, if similar conditions as induced by the OAE field trial are present in the habitats that species
652 inhabit, it is more likely that the species might be pre-adapted to such conditions. However, if
653 species have not been exposed to such conditions, OAE might induce rapid change in conditions
654 and species exposure, which could be more challenging for the species. As such, it is worth
655 considering that OAE deployments could be, at least for the most sensitive species, carried out not
656 as a single high dosage deployment, but rather as a more continuous, lower dosage application.
657 This would eliminate the swings and maxima in conditions, while also allowing more time for
658 species acclimation or migration during the initial injection of the OAE deployment. Ultimately,
659 it is the combination of all these factors that creates baseline exposure conditions that are relevant
660 in the context of biological outcomes (Wang et al., 2023).

661 **4.4 Direction of laboratory OAE experiments should change to incorporate field conditions**

662 The lab OAE experiments that are being conducted right now are done under different conditions

663 than in the field. The former are conducted with the aim of gaining a wide-ranging empirical
664 response, which implies high treatment levels of OAE additions. However, biogeochemical model
665 outputs show that OAE-related concentrations at the injection site are high for a short-time, while
666 the realistic field dosing upon rapid dilution due to mixing is low. Wang et al. (2023) reported that
667 the nearfield maxima in the respective investigation area of the Bering Sea is to increase TA by
668 about 10 $\mu\text{mol/kg}$ in the nearfield and by about 1 $\mu\text{mol/kg}$ of NaOH in the farfield region. As such,
669 we should be more concerned about the threshold of exceedance occurring at the low NaOH
670 dosing, rather than at high NaOH additions, because these are more realistic and point to the most
671 sensitive species. As a result, we explicitly emphasize the importance of including much lower
672 additions of TA in the experimental treatment levels to better support biological understanding and
673 OAE application in the field. In addition, prior to the lab experiments it would be important to
674 identify what type of response is predicted in the experimental species. This is especially pertinent
675 for the groups for which OA experimental data is limited and skewed towards the lowest TA:DIC
676 ratio (Fig. 8; Supplemental Fig. 4).

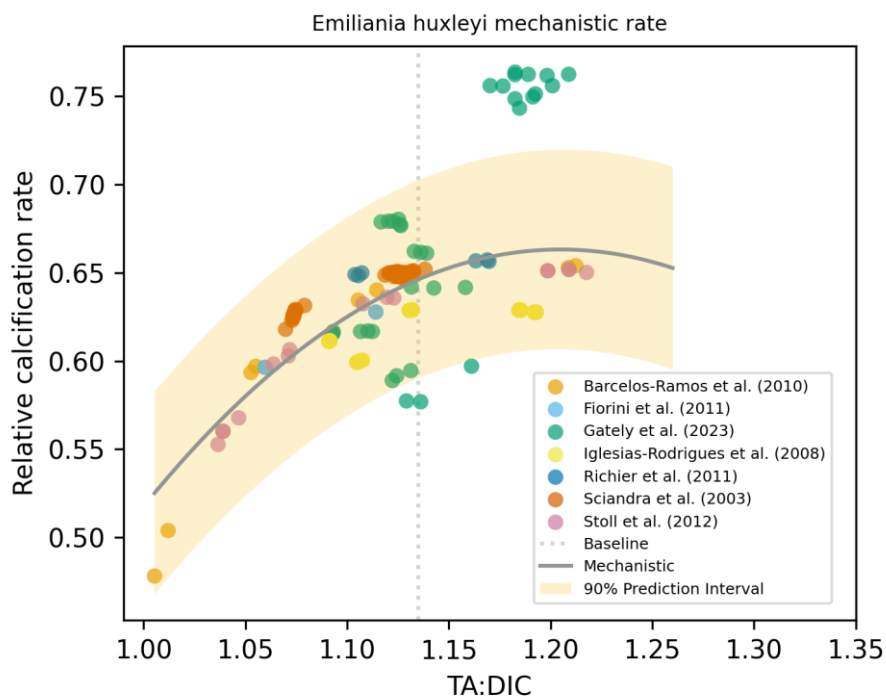
677 What is needed urgently for the safe biological field trial experiments is a set of protocols that are
678 species-, habitat- and local conditions- specific, which would allow for comprehensive and
679 comparative risk analyses and threshold determination. As part of this, we also need to develop
680 regionally specific indicators for biological monitoring. Ideally, such biological and environmental
681 risk monitoring and assessment would be accompanied by the application of the physical mixing
682 models with site-specific biogeochemical processes (Ho et al., 2023; Fennel et al., 2023) that can
683 predict the maximum expected TA increase in the nearfield and farfield regions of the study site,
684 representing a more realistic exposure and better informing further experimental work.

685 **4.5 Validating OAE responses based on the mechanistically-derived calcification**

686 This study establishes the predictions of responses that relied upon empirical experimental studies.
687 A good alternative to validating the predicted responses is to use species-specific mechanistic
688 responses, a more accurate representation of responses compared to empirical studies. Here, we
689 conducted a subset synthesis study for the two species of coccolithophores, using the results from
690 this study and compared it to the literature-derived mechanistic responses where the responses are
691 described with a different set of carbonate chemistry parameters. We wanted to determine to what

692 extent mechanistic relationships can contribute to improved, i.e. more accurate and certain, OAE
693 predictions.

694 For *Emiliana huxleyi*, we used experimental TA and DIC data to calculate the $[\text{HCO}_3^-]$, $[\text{H}^+]$ and
695 $[\text{CO}_2]$ concentrations to be able to use the mechanistic rate equation from Bach et al. (2015). We
696 calculated and plotted the rate derived via mechanistic approach and applied linear, polynomial
697 (second-order) and exponential regressions and chose the best fit based on the lowest p-value,
698 using the same method as for our experimental calcification rate data regressions. Like the
699 mechanistic rate regression based on three carbonate chemistry parameters was a parabolic fit
700 (Bach et al., 2015), we also obtained the same fit using the experimental calcification rate data (see
701 Fig. 9). However, when using the same approach for another coccolithophore species *Calcidiscus*
702 *leptopus* (Bach et al., 2015), our best fit did not align with the proposed mechanistic response;
703 instead, a non-significant relationship was obtained using experimental data (Supplemental Fig.
704 5). Such comparisons reveal species-specific relationships are likely dependent on a lot of
705 parameters, with one equation alone not being operable among different species from different
706 experiments or over varied regional settings.



707
708 **Figure 9:** Mechanistic rate equation and parameters ($a = 9.56e-1$, $b = 7.04e-4$ mol/kg, $c = 2.1e6$
709 kg/mol, $d = 8.27e6$ kg/mol) taken from Bach et al. (2015) and fitted using experimental data for

710 *E. huxleyi* (used data from the studies indicated in legend). Shading represents the 90%
711 prediction interval.

712

713 For most of the species, such mechanistic relationships are not available. Often, substrate-to-
714 inhibitor ratio (SIR) (i.e. the bicarbonate ion to hydrogen ion concentration ratio) has been used to
715 describe a calcification relationship that was based on the single-parameter relationships. To
716 compare if our experimental results can reproduce SIR, we computed and plotted the SIR ratio for
717 the mollusk, coral and coccolithophore groups, and applied a best-fit regression model. We
718 compared these SIR regressions to the respective best-fit regressions based on the empirical data
719 from the experiments. We found large differences between our proposed response and the SIR-
720 proposed mechanisms (Supplemental Fig. 6). For most of the coccolithophore groups, the
721 experimental rate regressions cannot be explained using SIR mechanisms (i.e. the responses are
722 different). Only in the case of *Calcidiscus leptoporus*, the experimental and mechanistic responses
723 remain the same (neutral). Reasons for these discrepancies could potentially be that SIR might
724 insufficiently explain the multitude of biological processes involved in the calcification (e.g. how
725 carbon is provisioned or the ability to regulate calcifying fluid pH), as well as other environmental
726 parameter variations. For mollusks, a third of the mechanistic rate regressions based on the SIR
727 agreed with the experimental calcification rate regressions. The other two-thirds did not agree,
728 especially for the studies with experimental conditions of $\Omega_{ar} > 1$. For corals, the majority of the
729 coral species (N=14) were classified as having a linear positive mechanistic relationship when
730 using SIR relationships. When comparing this to our experimental rate regressions, we only found
731 agreements between the experimental and mechanistic regressions in 6 out of 18 species. Based
732 on these results, the general consensus is that the SIR ratio tends to oversimplify species'
733 calcification rate responses. However, for corals it does seem to validate experimental results.

734

735 Mechanistic models can offer better insights into calcification responses for some species,
736 especially when multiple environmental factors are accounted for, but they are not generally
737 applicable across taxa. Species-specific responses are influenced by unique biological and
738 physiological factors, which can lead to significant deviations between mechanistic and empirical
739 predictions. Therefore, mechanistic approaches will only provide valuable frameworks for species
740 with well-understood calcification processes. By comparing mechanistic studies with experimental

741 data, we hoped to validate the predictive results of our experimental studies. Mechanistic models
742 can offer better insights into calcification responses for some species, especially when multiple
743 environmental factors are accounted for. However, for many species covered in this study the
744 calcification process is not well-understood. This meant the mechanistic relationship did not exist,
745 and when they did exist they were often based on one parameter only. Ninokawa et al. (2024) and
746 Li et al. (2023) emphasized that using only one parameter to describe the calcification process is
747 insufficient and strongly recommended using at least two parameters for more accurate
748 calcification predictions. Our findings agree with Ninokawa et al. (2024), for example, we observe
749 that using SIR relationships to successfully describe calcification was limited to only a few species
750 and that there are no generalizable patterns that could be applicable across multiple groups. This
751 clearly delineates a major gap in the mechanistic understanding of calcification so far, the lack of
752 which significantly limits our ability of ecological and biogeochemical predictions to OAE. As
753 such, more research is urgently needed on broader mechanistic understanding of calcification
754 across different species, and additionally, one parameter calcification processes should be replaced
755 with more accurate and comprehensive methods using two or three parameters.

756 **4.6 Unknowns about ecological and biogeochemical implications call for the precautionary** 757 **approach**

758 The value of calcification as the response proxy is indicative of organismal fitness, which directly
759 relates to OAE effects as harmful or beneficial for the species. From an ecological perspective, a
760 total of 26.0% negative responders demonstrates a potential for negative implications. In addition,
761 we note that this study did not include diatoms in the analyses, which are predicted to be negatively
762 impacted by carbonate-based OAE (Ferderer et al., 2022), leading to possible community-based
763 ecological shifts (Bach et al., 2019). The possibility of the ecological shifts should not be neglected
764 given the variety of the positive responders, understudied effects of OAE in non-calcifiers and
765 their relationship with the calcifiers through the grazing impact, and lastly, unknown and highly
766 unpredictable indirect effects. In addition, the inferences on the neutral responders should also
767 remain cautious.

768 From a biogeochemical perspective, it is reasonable to infer that OAE will introduce changes in
769 calcification rate across species, potentially resulting in changing the carbon export or carbonate

770 counter pump. Species-specific responses in major carbonate producers, i.e. coccolithophores,
771 foraminifera and pteropods show both, negative and positive response, which could have strong
772 effects on biogeochemical fluxes (Riebesell et al., 2017; Bach et al., 2019). Increased calcification
773 could result in thicker and denser shells, contributing to faster sinking and increased carbonate
774 fluxes, while decreased calcification has the opposite effect. This could potentially induce changes
775 on the subsurface total alkalinity at intermediate and deeper depths in the water column, and
776 dissolution at or near the seafloor (Gehlen et al., 2011) or result in a potential feedback of increased
777 CO₂ flux to the atmosphere (Gattuso et al., 2021). The full scope of ecological and biogeochemical
778 shifts remains a high priority topic for future investigations and until these huge uncertainties are
779 resolved, we should exercise a precautionary principle in considering the next steps of OAE field
780 implementations.

781 **4.7 Potential confounding effects**

782 This study only considered the changes in carbonate chemistry due to the addition of NaOH and
783 Na₂CO₃. However, other OAE feedstocks contain compounds that could induce biological toxicity
784 due to the presence of trace metals (Ni, Cu, Ca, Si; Bach et al., 2019), as well as potential negative
785 environmental impacts due to secondary precipitation (Hartmann et al., 2022; Moras et al., 2022).
786 This study also did not focus on the sensitivity across different life stages, even though stage-
787 specific sensitivities to OAE are expected based on previous OA results. Furthermore, we did
788 include data from experimental lab and field studies that involve multiple stressors in their
789 experimental designs. As such, an additional impact of warming, dissolved oxygen, and light
790 intensity on the OAE-induced responses was not determined, although they could elicit different
791 biological pathways than OAE alone or have additional confounding effects.

792 The synthesis of the experimental studies always includes implicit biases that are based on the
793 published experimental studies, the range and species used, regional coverage and heterogeneity.
794 Important consideration is the adaptation of the species used in the experimental studies because
795 their calcification optimum might be pre-determined based on their local habitat conditions. Given
796 that the baseline for the OAE-compound addition was chosen at the global current surface pH
797 value, some of the thresholds might actually be lower than expected.

798 **4.8 Applications within the existing governmental regulations and the guiding principle**

799 Our results, especially related to the use of biological thresholds or NaOH dosing, could have
800 wider applications, most notably with policy-management governmental regulations. For example,
801 we calculated the amount of alkalinity addition required to reach the pH_T threshold of 9, the
802 maximum pH allowed by the US Environmental Protection Agency's for waste water entering the
803 coastal ocean (see NPDES manual, 2010). To reach this threshold, 1200 $\mu\text{mol/kg}$ of NaOH was
804 required on average for all species, with the lowest threshold reached at 750 $\mu\text{mol/kg}$ addition for
805 *Amphibalanus improvisus*. This is a high concentration, and the thresholds for most of the negative
806 responders with identified thresholds (Table 2) will be exceeded far below the regulatory standards
807 of pH_T 9 (Table 2), especially if the exposure occurred over a duration period that matters for
808 calcification and for the organism's physiological status. This case demonstrates discrepancy of
809 the current chemical pH regulation and associated biological effects, where safe biological limits
810 are not considered, and biological harm might not be prevented. Despite the fact that achieving
811 such a high pH through NaOH implementation is unlikely to occur in the field, such regulations
812 currently do not assure safety space for marine biota and they need to be urgently addressed.

813 **5. Conclusions and next steps**

814 Sufficient certainty in predicting biological responses reduces the risks and supports safe operating
815 space for OAE implementation and scaling up. Overall, given that almost 60% of examined species
816 showed non-neutral response (either positive or negative), this calls for careful implementation of
817 OAE until the safe operational temporal and spatial scales are identified and OA mitigation
818 measures are established. The goal of this study is to serve as a baseline for prioritizing
819 experimental and field OAE research and assess environmental risks. Such prioritization identifies
820 those species for which experimental work needs to be conducted first. This would involve species
821 with the greatest OAE-related sensitivity (negative responders), species with the greatest
822 uncertainty in response, as well as the species with very strong predicted positive response that
823 could potentially introduce a shift on the community level. In addition, it would also recognize the
824 species for which the existing knowledge is sufficient and there is less immediate need for the
825 OAE experiments. We hope that all presented tools provide guidance for the practicing and
826 regulatory community considering OAE field application within the safe limits.

827 It is important to emphasize that this study is the first comprehensive synthesis of the effects of

828 OAE. Ongoing updates and additional data would enhance its value, particularly when
829 complemented by further experimental research. Similar datasets on OA exist for various
830 biological parameters, including genetics, physiology, and survival data, as well as for non-
831 calcifying organisms. This availability allows for the exploration of ecological implications and
832 contributes to developing an ecosystem-based predictive risk assessment for OAE.

833 **Data availability**

834 No new data were generated during this study; all data was collected from previously published
835 studies. The compiled data is currently available on request. The Python code used for computing
836 baselines per species, conceptually adding alkalinity in the form of NaOH and Na₂CO₃, predicting
837 calcification rate response, visualizing data and computing thresholds is available in the GitHub
838 repository at https://github.com/hannavdmortel/OAE_calc_response (last access: 1 November
839 2024) and is archived on Zenodo at <https://doi.org/10.5281/zenodo.14024442> (van de Mortel,
840 2024). PyCO₂SYs v1.8.0 (Humphreys et al., 2022) was used to solve for the carbonate system,
841 with software available at <https://doi.org/10.5281/zenodo.3744275> (Humphreys et al., 2023).

842 **Author contributions**

843 NB designed and conceptualized the research and wrote the first draft of the paper. HvdM collected
844 and curated data, conducted formal analyses and provided visualization. GP provided the analyses
845 using GLODAP data, and also provided visualizations and formal analyses. MGR has provided
846 formal statistical analyses and visuals. RAF and AD have provided insights, suggestions, and
847 generated discussion about specific parts of the paper. All have contributed to the writing of this
848 paper.

849

850 **Competing interests**

851 The contact author has declared that none of the authors has any competing interests.

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862 **References**

863 Agostini, S., Harvey, B. P., Milazzo, M., Wada, S., Kon, K., Floc'h, N., Komatsu, K., Kuroyama,
864 M., Hall-Spencer, J. M.: Seawater carbonate chemistry and kelp densities and coral coverages at
865 three study locations and photosynthesis and calcification of corals measured in the laboratory,
866 2021. PANGAEA, <https://doi.org/10.1594/PANGAEA.944056>, Supplement to: Agostini, S. et al.:
867 Simplification, not “tropicalization”, of temperate marine ecosystems under ocean warming and
868 acidification. *Global Change Biology*, 27(19), 4771-4784, <https://doi.org/10.1111/gcb.15749>,
869 2021.

870 Bach, L. T., Riebesell, U., and Schulz, K. G. (2011): Seawater carbonate chemistry, growth rate
871 and PIC and POC production during experiments with *Emiliana huxleyi* (B92/11), 2011 [dataset].
872 PANGAEA, <https://doi.org/10.1594/PANGAEA.771288>, Supplement to: Bach, L. T. et al. (2011):
873 Distinguishing between the effects of ocean acidification and ocean carbonation in the
874 coccolithophore *Emiliana huxleyi*. *Limnology and Oceanography*, 56(6), 2040-2050,
875 <https://doi.org/10.4319/lo.2011.56.6.2040> , 2011.

876 Bach, L. T., and Mackinder, L. C. M.: Experiment: Dissecting the impact of CO₂ and pH on the
877 mechanisms of photosynthesis and calcification in the coccolithophore *Emiliana huxleyi*.
878 PANGAEA, <https://doi.org/10.1594/PANGAEA.830627>, Supplement to: Bach, L. T., Mackinder,
879 L. C. M., Schulz, K. G., Wheeler, G., Schroeder, D. C., Brownlee, C., Riebesell, U.: Dissecting
880 the impact of CO₂ and pH on the mechanisms of photosynthesis and calcification in the
881 coccolithophore *Emiliana huxleyi*, *New Phytol.*, 199(1), 121-134,
882 <https://doi.org/10.1111/nph.12225>, 2013.

883 Bach, L.T., Riebesell, U., Gutowska, M.A., Federwisch, L. and Schulz, K.G.: A unifying concept
884 of coccolithophore sensitivity to changing carbonate chemistry embedded in an ecological
885 framework, *Prog. Oceanogr.*, 135, 125-138, 2015.

886 Bach, L.T., Gill, S.J., Rickaby, R.E., Gore, S. and Renforth, P., 2019. CO₂ removal with enhanced
887 weathering and ocean alkalinity enhancement: potential risks and co-benefits for marine pelagic
888 ecosystems., *Frontiers in Climate*, 1, p.7, 2019.

889 Barcelos e Ramos, J., Müller, M. N., and Riebesell, U.: Seawater carbonate chemistry and
890 processes during experiments with phytoplankton *Emiliania huxleyi* (strain Bergen 2005), 2010.
891 PANGAEA, <https://doi.org/10.1594/PANGAEA.736022>, Supplement to: Barcelos e Ramos, J. et
892 al. : Short-term response of the coccolithophore *Emiliania huxleyi* to an abrupt change in seawater
893 carbon dioxide concentrations. *Biogeosciences*, 7(1), 177-186, [https://doi.org/10.5194/bg-7-177-](https://doi.org/10.5194/bg-7-177-2010)
894 2010, 2010.

895 Bednaršek, N., Feely, R.A., Howes, E.L., Hunt, B.P., Kessouri, F., León, P., Lischka, S., Maas,
896 A.E., McLaughlin, K., Nezhlin, N.P. and Sutula, M.: Systematic review and meta-analysis toward
897 synthesis of thresholds of ocean acidification impacts on calcifying pteropods and interactions with
898 warming, *Frontiers in Marine Science*, 6, 227, 2019.

899 Bednaršek, N., Naish, K. A., Feely, R. A., Hauri, C., Kimoto, K., Hermann, A. J., Michel, C.,
900 Niemi, A., and Pilcher, D.: Integrated Assessment of Ocean Acidification Risks to Pteropods in
901 the Northern High Latitudes: Regional Comparison of Exposure, Sensitivity and Adaptive
902 Capacity, *Frontiers in Marine Science*, 8, 671497, <https://doi.org/10.3389/FMARS.2021.671497>,
903 2021a.

904 Bednaršek, N., Ambrose, R., Calosi, P., Childers, R.K., Feely, R.A., Litvin, S.Y., Long, W.C.,
905 Spicer, J.I., Štrus, J., Taylor, J. and Kessouri, F.: Synthesis of thresholds of ocean acidification
906 impacts on decapods, *Frontiers in Marine Science*, 8, 651102, 2021b.

907 Bednaršek, N., Calosi, P., Feely, R.A., Ambrose, R., Byrne, M., Chan, K.Y.K., Dupont, S., Padilla-
908 Gamiño, J.L., Spicer, J.I., Kessouri, F. and Roethler, M.: Synthesis of thresholds of ocean
909 acidification impacts on echinoderms, *Frontiers in Marine Science*, 8, 602601, 2021c.

910 Bibby, R., Cleall-Harding, P., Rundle, S., Widdicombe, S., & Spicer, J. I.: Seawater carbonate
911 chemistry during experiments with *Littorina littorea*, 2007 [dataset], 2007. PANGAEA,
912 <https://doi.org/10.1594/PANGAEA.716837>, Supplement to: Bibby, R., Cleall-Harding, P.,
913 Rundle, S., Widdicombe, S., & Spicer, J. I.: Ocean acidification disrupts induced defences in the

914 intertidal gastropod *Littorina littorea*. *Biology Letters*, 3(6), 699-701.
915 <https://doi.org/10.1098/rsbl.2007.0457>, 2007.

916 Bove, C. B., Whitehead, R. F., and Szmant, A. M.: Seawater carbonate chemistry and
917 gastrovascular cavity pH, calcification of *Montastraea cavernosa* and *Duncanopsammia axifuga*.
918 PANGAEA, <https://doi.org/10.1594/PANGAEA.927310>, 2020.

919 Brading, P., Warner, M. E., Davey, P., Smith, D. J., Achterberg, E. P., and Suggett, D. J.: Seawater
920 carbonate chemistry and growth rate during experiments with phylotypes of *Symbiodinium*
921 (Dinophyceae), 2011. PANGAEA, <https://doi.org/10.1594/PANGAEA.771293>, Supplement to:
922 Brading, P. et al.: Differential effects of ocean acidification on growth and photosynthesis among
923 phylotypes of *Symbiodinium* (Dinophyceae). *Limnol. Oceanogr.*, 56(3), 927-938,
924 <https://doi.org/10.4319/lo.2011.56.3.0927>, 2011.

925 Briggs, A. A., and Carpenter, R. C.: Seawater carbonate chemistry and photosynthesis and
926 photochemical efficiency of *Porolithon onkodes*. PANGAEA,
927 <https://doi.org/10.1594/PANGAEA.920025>, 2019.

928 Brown, K. T., Mello-Athayde, M. A., Sampayo, E. M., Chai, A., Dove, S., and Barott, K. L.:
929 Seawater carbonate chemistry and endosymbiont density, photosynthesis and net calcification
930 rates of reef-building coral *Pocillopora damicornis*. PANGAEA,
931 <https://doi.org/10.1594/PANGAEA.953058>, 2022.

932 Cameron, L. P., Reymond, C. E., Müller-Lundin, F., Westfield, I. T., Grabowski, J. H., Westphal,
933 H., and Ries, J. B.: Seawater carbonate chemistry and physiology and extrapallial fluid pH,
934 calcification rate, and condition factor of the king scallop *Pecten maximus*. PANGAEA,
935 <https://doi.org/10.1594/PANGAEA.919939>, 2019.

936 Camp, E. F., Nitschke, M. R., Rodolfo-Metalpa, R., Houlbrèque, F., Gardner, S. G., Smith, D. J.,
937 Zampighi, M., and Suggett, D. J.: Seawater carbonate chemistry and calcification rate, net
938 photosynthesis and respiration rate of reef-building corals. PANGAEA,
939 <https://doi.org/10.1594/PANGAEA.880242>, Supplement to: Camp, E.F. et al.: Reef-building
940 corals thrive within hot-acidified and deoxygenated waters, *Sci. Rep.*, 7(1),
941 <https://doi.org/10.1038/s41598-017-02383-y>, 2017.

942 Casareto, B. E., Niraula, M. P., Fujimura, H., and Suzuki, Y.: Seawater carbonate chemistry,
943 primary production, biomass and calcification of plankton and bacteria, 2009. PANGAEA,
944 <https://doi.org/10.1594/PANGAEA.756687>, 2009.

945 Comeau, S., Gorsky, G., Jeffree, R., Teysssié, Jean-Louis, and Gattuso, Jean-Pierre: Seawater
946 carbonate chemistry, shell linear extension and calcification during calcein staining and ⁴⁵Ca
947 experiments with pteropod *Limacina helicina*, 2009. PANGAEA,
948 <https://doi.org/10.1594/PANGAEA.726856>, Supplement to: Comeau, S. et al. : Impact of ocean
949 acidification on a key Arctic pelagic mollusc (*Limacina helicina*), *Biogeosciences*, 6(9), 1877-
950 1882, <https://doi.org/10.5194/bg-6-1877-2009>, 2009.

951 Comeau, S., Jeffree, R., Teysssié, Jean-Louis, and Gattuso, Jean-Pierre: Seawater carbonate
952 chemistry and biological processes during experiments with *Limacina helicina*, 2009. PANGAEA,
953 <https://doi.org/10.1594/PANGAEA.744720>, In: EPOCA Arctic experiment 2009 team (2009):
954 EPOCA 2009 Svalbard benthic experiment. PANGAEA,
955 <https://doi.org/10.1594/PANGAEA.745083>, 2010a.

956 Comeau, S., Gorsky, G., Alliouane, S., and Gattuso, Jean-Pierre: Seawater carbonate chemistry
957 and shell length of Mediterranean pteropod *Cavolinia inflexa* larvae during experiments.
958 Laboratoire d'Océanographie de Villefranche, PANGAEA,
959 <https://doi.org/10.1594/PANGAEA.733905>, Supplement to: Comeau, S. et al. : Larvae of the
960 pteropod *Cavolinia inflexa* exposed to aragonite undersaturation are viable but shell-less, *Marine*
961 *Biol.*, 157(10), 2341-2345, <https://doi.org/10.1007/s00227-010-1493-6>, 2010b.

962 Comeau, S., Edmunds, P. J., Spindel, N. B., and Carpenter, R. C.: The responses of eight coral
963 reef calcifiers to increasing partial pressure of CO₂ do not exhibit a tipping point. PANGAEA,
964 <https://doi.org/10.1594/PANGAEA.833687>, Supplement to: Comeau, S. et al.: The responses of
965 eight coral reef calcifiers to increasing partial pressure of CO₂ do not exhibit a tipping point.
966 *Limnol. Oceanogr.*, 58(1), 388-398, <https://doi.org/10.4319/lo.2013.58.1.0388>, 2013.

967 Comeau, S., Cornwall, C. E., De Carlo, E. H., Krieger, E., and McCulloch, M. T.: Seawater
968 carbonate chemistry and calcification physiology data in coral reef taxa. PANGAEA,
969 <https://doi.org/10.1594/PANGAEA.892655>, Supplement to: Comeau, S. et al.: Similar controls on

970 calcification under ocean acidification across unrelated coral reef taxa, *Glob. Change Biol.*,
971 <https://doi.org/10.1111/gcb.14379>, 2018.

972 Comeau, S., Cornwall, C. E., Pupier, C. A., DeCarlo, Thomas M., Alessi, C., Trehern, R., and
973 McCulloch, M. T.: Seawater carbonate chemistry and calcification rate, calcifying fluid pH,
974 calcifying fluid DIC, photosynthetic rates, metabolic alteration of pH in the DBL of corals and
975 coralline algae. PANGAEA, <https://doi.org/10.1594/PANGAEA.914328>, 2019.

976 Cornwall, C. E., Comeau, S., DeCarlo, T. M., Moore, B., D'Alexis, Q., and McCulloch, M. T.:
977 Seawater carbonate chemistry and resistance of corals and coralline algae to ocean acidification.
978 PANGAEA, <https://doi.org/10.1594/PANGAEA.914886>, 2018.

979 Courtney, T., Westfield, I. T., & Ries, J. B.: Seawater carbonate chemistry and calcification in the
980 tropical urchin *Echinometra viridis* in a laboratory experiment [dataset], 2013. PANGAEA,
981 <https://doi.org/10.1594/PANGAEA.824707>, Supplement to: Courtney, T., Westfield, I. T., & Ries,
982 J. B.: CO₂-induced ocean acidification impairs calcification in the tropical urchin *Echinometra*
983 *viridis*. *Journal of Experimental Marine Biology and Ecology*, 440, 169-175.
984 <https://doi.org/10.1016/j.jembe.2012.11.013>, 2013.

985 Courtney, T., and Ries, J. B.: Impact of atmospheric pCO₂, seawater temperature, and calcification
986 rate on the delta 18O and delta 13C composition of echinoid calcite (*Echinometra viridis*).
987 PANGAEA, <https://doi.org/10.1594/PANGAEA.862558>, Supplement to: Courtney, T., and Ries,
988 J.B. : Impact of atmospheric pCO₂, seawater temperature, and calcification rate on the delta 18O
989 and delta 13C composition of echinoid calcite (*Echinometra viridis*), *Chem. Geol.*, 411, 228-239,
990 <https://doi.org/10.1016/j.chemgeo.2015.06.030>, 2015.

991 Cripps, G., Widdicombe, S., Spicer, J.I. and Findlay, H.S. Biological impacts of enhanced
992 alkalinity in *Carcinus maenas*. *Marine pollution bulletin*, 71(1-2), pp.190-198, 2013

993 Dickson, A. G.: Standard potential of the reaction: , and and the standard acidity constant of the
994 ion HSO₄⁻ in synthetic sea water from 273.15 to 318.15 K. *J. Chem. Thermodyn.*, 22(2), 113–
995 127, [https://doi.org/10.1016/0021-9614\(90\)90074-z](https://doi.org/10.1016/0021-9614(90)90074-z), 1990.

996 Diner, R. E., Benner, I., Passow, U., Iglesias-Rodriguez, M. D., & Robertson, D. L.:

997 Negative effects of ocean acidification on calcification vary within the coccolithophore
998 genus *Calcidiscus*. *Marine Biology*, 162, 1287–1305. [https://doi.org/10.1007/s00227-015-](https://doi.org/10.1007/s00227-015-2669-x)
999 2669-x, 2015.

1000 Eisaman, M. D., Geilert, S., Renforth, P., Bastianini, L., Campbell, J., Dale, A. W., Foteinis, S.,
1001 Grasse, P., Hawrot, O., Löscher, C. R., Rau, G. H., and Rønning, J.: Assessing the technical aspects
1002 of ocean-alkalinity-enhancement approaches. *State Of The Planet*, 2-oae2023, 1–29,
1003 <https://doi.org/10.5194/sp-2-oae2023-3-2023>, 2023.

1004 Evensen, N. R., and Edmunds, P. J.: Interactive effects of ocean acidification and neighboring
1005 corals on the growth of *Pocillopora verrucosa*. *PANGAEA*,
1006 <https://doi.org/10.1594/PANGAEA.867268>, Supplement to: Evensen, N.R., and Edmunds, P.J.:
1007 Interactive effects of ocean acidification and neighboring corals on the growth of *Pocillopora*
1008 *verrucosa*, *Marine Biol.*, 163(7), <https://doi.org/10.1007/s00227-016-2921-z>, 2016.

1009 Feely, R. A., Sabine, C. L., Lee, K., Berelson, W., Kleypas, J., Fabry, V. J., and Millero, F. J. :
1010 Impact of anthropogenic CO₂ on the CaCO₃ system in the oceans. *Science*, 305(5682), 362–366,
1011 <https://doi.org/10.1126/SCIENCE.1097329>, 2004.

1012 Fennel, K., Long, M.C., Algar, C., Carter, B., Keller, D., Laurent, A., Mattern, J.P., Musgrave, R.,
1013 Oschlies, A., Ostiguy, J. and Palter, J.: Modeling considerations for research on Ocean Alkalinity
1014 Enhancement (OAE). *State of the Planet Discussions*, 2023, 1-47, 2023.

1015 Ferderer, A., Chase, Z., Kennedy, F., Schulz, K.G. and Bach, L.T.: Assessing the influence of
1016 ocean alkalinity enhancement on a coastal phytoplankton community. *Biogeosciences*, 19(23),
1017 5375-5399, 2022.

1018 Findlay, H. S., Kendall, M. A., Spicer, J. I., and Widdicombe, S. : Seawater carbonate chemistry
1019 and biological processes during experiments with barnacle *Semibalanus balanoides*, 2010.
1020 *PANGAEA*, <https://doi.org/10.1594/PANGAEA.737438>, Supplement to: Findlay, H.S. et al.:
1021 Relative influences of ocean acidification and temperature on intertidal barnacle post-larvae at the
1022 northern edge of their geographic distribution, *Estuar. Coast. Shelf S.*, 88(4), 675-682,
1023 <https://doi.org/10.1016/j.ecss.2009.11.036>, 2010.

1024 Fiorini, S., Middelburg, J. J., and Gattuso, Jean-Pierre: Seawater carbonate chemistry, nutrients,
1025 particulate carbon and growth rate of *Emiliana huxleyi* (AC472), *Calcidiscus leptoporus* (AC370)
1026 and *Syracosphaera pulchra* (AC418) during experiments, 2011. PANGAEA,
1027 <https://doi.org/10.1594/PANGAEA.773860>, Supplement to: Fiorini, S. et al.: Testing the effects
1028 of elevated pCO₂ on coccolithophores (Prymnesiophyceae): comparison between haploid and
1029 diploid life stages. *J. Phycol.*, 47(6), 1281–1291, [https://doi.org/10.1111/j.1529-](https://doi.org/10.1111/j.1529-8817.2011.01080.x)
1030 [8817.2011.01080.x](https://doi.org/10.1111/j.1529-8817.2011.01080.x), 2011.

1031 Friedlingstein, P., O'sullivan, M., Jones, M.W., Andrew, R.M., Gregor, L., Hauck, J., Le Quéré,
1032 C., Lujikx, I.T., Olsen, A., Peters, G.P. and Peters, W.: Global carbon budget 2022. *Earth System*
1033 *Science Data Discussions*, 2022, 1-15, 2022.

1034 Gafar, N.A. and Schulz, K.G.: A three-dimensional niche comparison of *Emiliana huxleyi* and
1035 *Gephyrocapsa oceanica*: reconciling observations with projections. *Biogeosciences*, 15(11),
1036 3541-3560, 2018.

1037 Garilli, V., Rodolfo-Metalpa, R., Scuderi, D., Brusca, L., Parrinello, D., Rastrick, S. P. S.,
1038 Foggo, A., Twitchett, R. J., Hall-Spencer, J. M., & Milazzo, M.: Physiological advantages
1039 of dwarfing in surviving extinctions in high-CO₂ oceans [dataset], 2015. PANGAEA.
1040 <https://doi.org/10.1594/PANGAEA.847397>, Supplement to: Garilli, V., Rodolfo-Metalpa,
1041 R., Scuderi, D., Brusca, L., Parrinello, D., Rastrick, S. P. S., et al.: Physiological
1042 advantages of dwarfing in surviving extinctions in high-CO₂ oceans. *Nature Climate*
1043 *Change*. <https://doi.org/10.1038/NCLIMATE2616>, 2015.

1044 Gately, J. A., Kim, S. M., Jin, B., Brzezinski, M. A., and Iglesias-Rodriguez, M. D.:
1045 Coccolithophores and diatoms resilient to ocean alkalinity enhancement: A glimpse of hope?
1046 *Science Advances*, 9(24), eadg6066. <https://doi.org/10.1126/SCIADV.ADG6066>, 2023.

1047 Gattuso, J.P., Magnan, A.K., Bopp, L., Cheung, W.W., Duarte, C.M., Hinkel, J., Mcleod, E.,
1048 Micheli, F., Oschlies, A., Williamson, P. and Billé, R.: Ocean solutions to address climate
1049 change and its effects on marine ecosystems, *Frontiers in Marine Science*, 5, 410554, 2018.

1050

1051 Gattuso, J.P., Williamson, P., Duarte, C.M. and Magnan, A.K.: The potential for ocean-based
1052 climate action: negative emissions technologies and beyond. *Frontiers in Climate*, 2, 575716,
1053 2021.

1054 Gazeau, F., Quiblier, C., Jansen, J. M., Gattuso, Jean-Pierre, Middelburg, J. J., and Heip, C. H.R.:
1055 Seawater carbonate chemistry and calcification during incubation experiments with *Mytilus edulis*
1056 and *Grassostrea gigas*, 2006. PANGAEA, <https://doi.org/10.1594/PANGAEA.718130>,
1057 Supplement to: Gazeau, F. et al.: Impact of elevated CO₂ on shellfish calcification. *Geophys. Res.*
1058 *Let.*, 34, <https://doi.org/10.1029/2006GL028554>, 2007.

1059 Gehlen, M., Gruber, N., Gangstø, R., Bopp, L. and Oschlies, A.: Biogeochemical consequences of
1060 ocean acidification and feedback to the earth system. *Ocean acidification*, 1, 230-248, 2011.

1061 González, M.F. and Ilyina, T.: 2016. Impacts of artificial ocean alkalization on the carbon cycle
1062 and climate in Earth system simulations. *Geophysical Research Letters*, 43(12), pp.6493-6502,
1063 2016.

1064 González, M.F., Ilyina, T., Sonntag, S. and Schmidt, H.: Enhanced Rates of Regional Warming
1065 and Ocean Acidification After Termination of Large-Scale Ocean Alkalization, *Geophys. Res.*
1066 *Let.*, 45(14), 7120-7129, 2018.

1067 Hansen, P.J., 2002. Effect of high pH on the growth and survival of marine phytoplankton:
1068 implications for species succession. *Aquatic Microbial Ecology*, 28(3), pp.279-288, 2002.

1069 Hansen, P. J., Lundholm, N., and Rost, Björn.: Seawater carbonate chemistry and growth rate
1070 during experiments with dinoflagellates, 2007. PANGAEA,
1071 <https://doi.org/10.1594/PANGAEA.718182>, In supplement to: Hansen, P.J. et al.: Growth
1072 limitation in marine red-tide dinoflagellates: effects of pH versus inorganic carbon availability.
1073 *Mar. Ecol. Prog. Ser.*, 334, 63-71, <https://doi.org/10.3354/meps334063>, 2007.

1074 Hartmann, J., Suitner, N., Lim, C., Schneider, J., Marín-Samper, L., Arístegui, J., Renforth, P.,
1075 Taucher, J. and Riebesell, U.: Stability of alkalinity in ocean alkalinity enhancement (OAE)
1076 approaches—consequences for durability of CO₂ storage, *Biogeosciences Discussions*, 2022, 1-
1077 29, 2022.

1078 He, J. and Tyka, M.D.: Limits and CO₂ equilibration of near-coast alkalinity enhancement,
1079 Biogeosciences, 20(1), 27-43, 2023.

1080 Ho, D.T., Bopp, L., Palter, J.B., Long, M.C., Boyd, P.W., Neukermans, G. and Bach, L.T.:
1081 Monitoring, reporting, and verification for ocean alkalinity enhancement, State of the Planet, 2, 1-
1082 12, 2023.

1083 Humphreys, M. P., Lewis, E. R., Sharp, J. D., and Pierrot, D.: PyCO2SYS v1.8: marine carbonate
1084 system calculations in Python, Geosci. Model Dev., 15(1), 15–43, [https://doi.org/10.5194/gmd-](https://doi.org/10.5194/gmd-15-15-2022)
1085 15-15-2022, 2022.

1086 Iglesias-Rodriguez, M.D., Halloran, P. R., Rickaby, R.E.M., Hall, I. R., Colmenero-Hidalgo, E.,
1087 Gittins, J. R., Green, D. R. H., Tyrrell, T., Gibbs S.J., von Dassow, P., Rehm, E., Armbrust, E. V.,
1088 and Boessenkool, K. P.: Seawater carbonate chemistry and processes during experiments with
1089 *Emiliania huxleyi*, 2008. PANGAEA, <https://doi.org/10.1594/PANGAEA.718841>, Supplement
1090 to: Iglesias-Rodriguez, D. et al.: Phytoplankton calcification in a high-CO₂ world. Science,
1091 320(5874), 336-340, <https://doi.org/10.1126/science.1154122>, 2008.

1092 Johnson, M. D., Bravo, L., Lucey, N. M., and Altieri, A.H. : Seawater carbonate chemistry and
1093 calcification rate of crustose coralline algae. PANGAEA,
1094 <https://doi.org/10.1594/PANGAEA.939809>, 2021.

1095 Keul, N., Langer, G., de Nooijer, L. J., and Bijma, J.: Seawater carbonate chemistry and benthic
1096 foraminifera *Ammonia sp.* mass, size, and growth rate during experiments, 2013. PANGAEA,
1097 <https://doi.org/10.1594/PANGAEA.821209>, Supplement to: Keul, N. et al.: Effect of ocean
1098 acidification on the benthic foraminifera *Ammonia sp.* is caused by a decrease in carbonate ion
1099 concentration. Biogeosciences, 10(10), 6185-6198, <https://doi.org/10.5194/bg-10-6185-2013>, 2013.

1100 Kheshgi, H.S.: Sequestering atmospheric carbon dioxide by increasing ocean alkalinity, Energy,
1101 20(9), 915-922, 1995.

1102 Kisakürek, B., Eisenhauer, A., Böhm, F., Hathorne, E. C., and Erez, J.: Seawater carbonate
1103 chemistry and biological processes of foraminifera, *Globigerinoides ruber* and *Globigerinella*
1104 siphonifera during experiments, 2011. PANGAEA, <https://doi.org/10.1594/PANGAEA.763297>,

1105 Supplement to: Kisakürek, B. et al.: Controls on calcium isotope fractionation in cultured
1106 planktonic foraminifera, *Globigerinoides ruber* and *Globigerinella siphonifera*. *Geochim.*
1107 *Cosmochim. Ac.*, 75(2), 427-443, <https://doi.org/10.1016/j.gca.2010.10.015>, 2011.

1108 Kroeker, K.J., Kordas, R.L., Crim, R., Hendriks, I.E., Ramajo, L., Singh, G.S., Duarte, C.M. and
1109 Gattuso, J.P. Impacts of ocean acidification on marine organisms: quantifying sensitivities and
1110 interaction with warming. *Global change biology*, 19(6), pp.1884-1896, 2013.

1111 Krueger, T., Horwitz, N., Bodin, J., Giovani, Maria-Evangelia, Escrig, S., Meibom, A., and Fine,
1112 M.: Seawater carbonate chemistry and photosynthesis, respiration and calcification of common
1113 reef-building coral in the Northern Red Sea. PANGAEA,
1114 <https://doi.org/10.1594/PANGAEA.880318>, Supplement to: Krueger, T. et al.: Common reef-
1115 building coral in the Northern Red Sea resistant to elevated temperature and acidification. *Roy.*
1116 *Soc. Open Sci.*, 4(5), 170038, <https://doi.org/10.1098/rsos.170038>, 2017.

1117 Langer, G., Geisen, M., Baumann, Karl-Heinz, Kläs, J., Riebesell, U., Thoms, S., and Young, J.:
1118 Seawater carbonate chemistry, growth rate and processes during experiments with *Coccolithus*
1119 *pelagicus* and *Calcidiscus leptoporus*, 2006. PANGAEA,
1120 <https://doi.org/10.1594/PANGAEA.721107>, Supplement to: Langer, G. et al.: Species-specific
1121 responses of calcifying algae to changing seawater carbonate chemistry. *Geochem. Geophys.*
1122 *Geosys.*, 7, Q09006, <https://doi.org/10.1029/2005GC001227>, 2006.

1123 Langer, G., and Bode, M.: Seawater carbonate chemistry, growth rate and morphology of
1124 *Calcidiscus leptoporus* (RCC1135) during experiments, 2011. PANGAEA,
1125 <https://doi.org/10.1594/PANGAEA.763286>, Supplement to: Langer, G., and Bode, M.: CO₂
1126 mediation of adverse effects of seawater acidification in *Calcidiscus leptoporus*. *Geochem.*
1127 *Geophys. Geosys.*, 12(5), Q05001, <https://doi.org/10.1029/2010GC003393>, 2011.

1128 Leung, J.Y., Zhang, S., and Connell, S.D.: Is ocean acidification really a threat to marine calcifiers?
1129 A systematic review and meta-analysis of 980+ studies spanning two decades. *Small*, 18(35),
1130 2107407, 2022.

1131 Lischka, S., Büdenbender, J., Boxhammer, T., and Riebesell, U.: Seawater carbonate chemistry
1132 and biological processes of *Limacina helicina* during experiments, 2011. PANGAEA,

1133 <https://doi.org/10.1594/PANGAEA.761910>, Supplement to: Lischka, S. et al.: Impact of ocean
1134 acidification and elevated temperatures on early juveniles of the polar shelled pteropod *Limacina*
1135 *helicina*: mortality, shell degradation, and shell growth. *Biogeosciences*, 8(4), 919-932,
1136 <https://doi.org/10.5194/bg-8-919-2011>, 2011.

1137 Lischka, S., and Riebesell, U.: Synergistic effects of ocean acidification and warming on
1138 overwintering pteropods in the Arctic. PANGAEA, <https://doi.org/10.1594/PANGAEA.832422>,
1139 Supplement to: Lischka, S., and Riebesell, U.: Synergistic effects of ocean acidification and
1140 warming on overwintering pteropods in the Arctic. *Glob. Change Biol.*, 18(12), 3517-3528,
1141 <https://doi.org/10.1111/gcb.12020>, 2012.

1142 Lutier, M., Di Poi, C., Gazeau, F., Appolis, A., Luyer, J. L., and Pernet, F.: Revisiting tolerance to
1143 ocean acidification: Insights from a new framework combining physiological and molecular
1144 tipping points of Pacific oyster, *Glob. Change Biol.*, 28(10), 3333–3348.
1145 <https://doi.org/10.1111/gcb.16101>, 2022.

1146 Maier, C., Hegeman, J., Weinbauer, M. G., and Gattuso, Jean-Pierre: Seawater carbonate
1147 chemistry and calcification of *Lophelia pertusa* during experiments, 2009. PANGAEA,
1148 <https://doi.org/10.1594/PANGAEA.767577>, Supplement to: Maier, C. et al.: Calcification of the
1149 cold-water coral *Lophelia pertusa* under ambient and reduced pH. *Biogeosciences*, 6(8), 1671-
1150 1680, <https://doi.org/10.5194/bg-6-1671-2009>, 2009.

1151 Manno, C., Morata, N., and Bellerby, R.G.J.: Seawater carbonate chemistry, survival rate, shell
1152 size, calcification rate of the planktonic foraminifer *Neogloboquadrina pachyderma* (sinistral) in a
1153 laboratory experiment. PANGAEA, <https://doi.org/10.1594/PANGAEA.830908>, Supplement to:
1154 Manno, C. et al. : Effect of ocean acidification and temperature increase on the planktonic
1155 foraminifer *Neogloboquadrina pachyderma* (sinistral). *Polar Biol.*, 35(9), 1311-1319,
1156 <https://doi.org/10.1007/s00300-012-1174-7>, 2012.

1157 Manríquez, P. H., Jara, M. E., Seguel, M.E., Torres, R., Alarcon, E., Lee, M. R., and Dam, H. G.:
1158 Ocean acidification and increased temperature have both positive and negative effects on early
1159 ontogenetic traits of a rocky shore keystone predator species. PANGAEA,
1160 <https://doi.org/10.1594/PANGAEA.869291>, 2016.

1161 Mekkes, L., Renema, W., Alin, S. R., Feely, R. A., Huisman, J., Roessingh, P., &
1162 Peijnenburg, K. T. C. A.: Seawater carbonate chemistry and shell thickness, shell dissolution
1163 of *Limacina helicina* pteropods [dataset]. PANGAEA.
1164 <https://doi.org/10.1594/PANGAEA.930065>, 2021.

1165 Meyer, F. W., Vogel, N., Teichberg, M., Uthicke, S., Wild, C., and Diaz-Pulido, G.: The
1166 physiological response of two green calcifying algae from the great barrier reef towards high
1167 dissolved inorganic and organic carbon (DIC and DOC) availability. PANGAEA,
1168 <https://doi.org/10.1594/PANGAEA.868094>, 2015.

1169 Meyer, F. W., Vogel, N., Diele, K., Kunzmann, A., Uthicke, S., & Wild, C.: Effects of high
1170 dissolved inorganic and organic carbon availability on the physiology of the hard coral
1171 *Acropora millepora* from the Great Barrier Reef [dataset]. PANGAEA.
1172 <https://doi.org/10.1594/PANGAEA.869416>, 2016

1173 Mongin, M., Baird, M.E., Lenton, A., Neill, C. and Akl, J.: Reversing ocean acidification along
1174 the Great Barrier Reef using alkalinity injection, *Environ. Res. Lett.*, 16(6), 064068, 2021.

1175 Monserrat, M., Comeau, S., Verdura, J., Alliouane, S., Spennato, G., Priouzeau, F., Romero, G.,
1176 and Mangialajo, L.: Seawater carbonate chemistry and the recruitment of macroalgal marine
1177 forests. PANGAEA, <https://doi.org/10.1594/PANGAEA.955425>, 2022.

1178 Moras, C.A., Bach, L.T., Cyronak, T., Joannes-Boyau, R. and Schulz, K.G.: Ocean alkalinity
1179 enhancement—avoiding runaway CaCO₃ precipitation during quick and hydrated lime dissolution.
1180 *Biogeosciences*, 19(15), 3537-3557, 2022.

1181 National Academies of Sciences, Engineering, and Medicine: A research strategy for ocean-based
1182 carbon dioxide removal and sequestration, 2021.

1183 Ninokawa, A., Takeshita, Y., Jellison, B. M., Jurgens, L. J., and Gaylord, B.: Seawater carbonate
1184 chemistry and mussel respiration and calcification rates. PANGAEA,
1185 <https://doi.org/10.1594/PANGAEA.915978>, 2020.

1186 Ninokawa, A. T., Saley, A. M., Shalchi, R., and Gaylord, B.: Multiple carbonate system parameters
1187 independently govern shell formation in a marine mussel. *Communications Earth & Environment*,

1188 5(1), <https://doi.org/10.1038/s43247-024-01440-5>, 2024.

1189 Noisette, F., Bordeyne, F., Davoult, D., and Martin, S.: Assessing the physiological responses of
1190 the gastropod *Crepidula fornicata* to predicted ocean acidification and warming. PANGAEA,
1191 <https://doi.org/10.1594/PANGAEA.860508>, Supplement to: Noisette, F. et al.: Assessing the
1192 physiological responses of the gastropod *Crepidula fornicata* to predicted ocean acidification and
1193 warming. *Limnol. Oceanogr.*, 61(2), 430-444, <https://doi.org/10.1002/lno.10225>, 2016.

1194 Okazaki, R., Swart, P. K., and Langdon, C.: Stress-tolerant corals of Florida Bay are vulnerable to
1195 ocean acidification. PANGAEA, <https://doi.org/10.1594/PANGAEA.833005>, Supplement to:
1196 Okazaki, R. et al.: Stress-tolerant corals of Florida Bay are vulnerable to ocean acidification. *Coral*
1197 *Reefs*, 32(3), 671-683, <https://doi.org/10.1007/s00338-013-1015-3>, 2013.

1198 Ong, E. Z., Briffa, M., Moens, T., and Van Colen, C.: Seawater carbonate chemistry and
1199 respiration, clearance and calcification rates of the common cockle *Cerastoderma edule*.
1200 PANGAEA, <https://doi.org/10.1594/PANGAEA.949749>, 2017.

1201 Oron, S., Evans, D., Abramovich, S., Almogi-Labin, A., and Erez, J.: Seawater carbonate
1202 chemistry and calcification, respiration, and photosynthesis of the widespread diatom-bearing LBF
1203 *Operculina ammonoides*. PANGAEA, <https://doi.org/10.1594/PANGAEA.929866>, 2020.

1204 Oschlies, A., Bach, L., Rickaby, R., Satterfield, T., Webb, R.M. and Gattuso, J.P.: Climate targets,
1205 carbon dioxide removal and the potential role of Ocean Alkalinity Enhancement. *State of the*
1206 *Planet Discussions*, 2023, 1-11, 2023.

1207 Palmer, R. M.: Alkalinity enhancement, thermal stress and their impacts on the physiology of three
1208 Caribbean coral species: *Acropora Cervicornis*, *Pseudodiploria strigosa* and *Siderastrea siderea*. In
1209 University of Miami. [https://scholarship.miami.edu/esploro/outputs/graduate/Alkalinity-](https://scholarship.miami.edu/esploro/outputs/graduate/Alkalinity-Enhancement-Thermal-Stress-and-their/991031772314902976)
1210 [Enhancement-Thermal-Stress-and-their/991031772314902976](https://scholarship.miami.edu/esploro/outputs/graduate/Alkalinity-Enhancement-Thermal-Stress-and-their/991031772314902976), 2022.

1211 Pansch, C., Schaub, I., Havenhand, J. N., and Wahl, M.: Habitat traits and food availability
1212 determine the response of marine invertebrates to ocean acidification. PANGAEA,
1213 <https://doi.org/10.1594/PANGAEA.831428>, Supplement to: Pansch, C., Schaub, M., Havenhand,
1214 J. N., and Wahl, M.: Habitat traits and food availability determine the response of marine

1215 invertebrates to ocean acidification. *Glob. Change Biol.*, 20(3), 765-777,
1216 <https://doi.org/10.1111/gcb.12478>, 2014.

1217 Paul, A.J. and Bach, L.T.: Universal response pattern of phytoplankton growth rates to increasing
1218 CO₂. *New Phytologist*, 228(6), 1710-1716, 2020.

1219 Prazeres, M., Uthicke, S., and Pandolfi, J. M.: Ocean acidification induces biochemical and
1220 morphological changes in the calcification process of large benthic foraminifera. *PANGAEA*,
1221 <https://doi.org/10.1594/PANGAEA.848419>, 2015.

1222 Putnam, H. M., and Gates, R. D.: Preconditioning in the reef-building coral *Pocillopora damicornis*
1223 and the potential for trans-generational acclimatization in coral larvae under future climate change
1224 conditions. *PANGAEA*, <https://doi.org/10.1594/PANGAEA.859356>, Supplement to: Putnam,
1225 H.M., and Gates, R.D.: Preconditioning in the reef-building coral *Pocillopora damicornis* and the
1226 potential for trans-generational acclimatization in coral larvae under future climate change
1227 conditions. *J. Exp. Biol.*, 218(15), 2365-2372, <https://doi.org/10.1242/jeb.123018>, 2015.

1228 Ramajo, L., Marbà, N., Prado, L., Peron, S., Lardies, M. A., Rodriguez-Navarro, A., Vargas, C.
1229 A., Lagos, N. A., and Duarte, C. M.: Biomineralization changes with food supply confer juvenile
1230 scallops (*Argopecten purpuratus*) resistance to ocean acidification. *PANGAEA*,
1231 <https://doi.org/10.1594/PANGAEA.860506>, Supplement to: Ramajo, L. et al.: Biomineralization
1232 changes with food supply confer juvenile scallops (*Argopecten purpuratus*) resistance to ocean
1233 acidification. *Glob. Change Biol.*, 22(6), 2025-2037, <https://doi.org/10.1111/gcb.13179>, 2016.

1234 Renforth, P. and Henderson, G.: Assessing ocean alkalinity for carbon sequestration, *Rev.*
1235 *Geophys.*, 55(3), 636-674, 2017.

1236 Reymond, C. E., Lloyd, A., Kline, D. I., Dove, S., and Pandolfi, J. M.: Decline in growth of
1237 foraminifer *Marginopora rossi* under eutrophication and ocean acidification scenarios.
1238 *PANGAEA*, <https://doi.org/10.1594/PANGAEA.833683>, Supplement to: Reymond, C.E., et al.:
1239 Decline in growth of foraminifer *Marginopora rossi* under eutrophication and ocean acidification
1240 scenarios. *Glob. Change Biol.*, 19(1), 291-302, <https://doi.org/10.1111/gcb.12035>, 2013.

1241 Richardson, K., Steffen, W., Lucht, W., Bendtsen, J., Cornell, S.E., Donges, J.F., Drüke, M.,

1242 Fetzer, I., Bala, G., von Bloh, W., and Feulner, G.: Earth beyond six of nine planetary boundaries.
1243 Science Advances, 9(37), p.eadh2458, 2023.

1244 Richier, S., Fiorini, S., Kerros, Marie-Emmanuelle, von Dassow, P., and Gattuso, Jean-Pierre:
1245 Seawater carbonate chemistry, particulate inorganic and organic carbon and growth rate of
1246 *Emiliana huxleyi* (RCC1216) during experiments, 2011. PANGAEA,
1247 <https://doi.org/10.1594/PANGAEA.770439>, In supplement to: Richier, S. et al.: Response of the
1248 calcifying coccolithophore *Emiliana huxleyi* to low pH/high pCO₂: from physiology to molecular
1249 level. Marine Biol., 158(3), 551-560, <https://doi.org/10.1007/s00227-010-1580-8>, 2011.

1250 Riebesell, U., Wolf-Gladrow, D.A. and Smetacek, V.: Carbon dioxide limitation of marine
1251 phytoplankton growth rates. Nature, 361(6409), 249-251, 1993.

1252 Riebesell, U. and Gattuso, J.P.: 2015. Lessons learned from ocean acidification research., Nature
1253 Climate Change, 5(1), pp.12-14, 2015.

1254 Riebesell, U., Bach, L.T., Bellerby, R.G., Monsalve, J.R.B., Boxhammer, T., Czerny, J., Larsen,
1255 A., Ludwig, A. and Schulz, K.G.: Competitive fitness of a predominant pelagic calcifier impaired
1256 by ocean acidification, Nat. Geosci., 10(1), 19-23, 2017.

1257 Ries, J. B., Cohen, A.L., and McCorkle, D.C.: Seawater carbonate chemistry and biological
1258 processes during experiments with calcifying organisms, 2009. PANGAEA,
1259 <https://doi.org/10.1594/PANGAEA.733947>, Supplement to: Ries, J.B., et al.: Marine calcifiers
1260 exhibit mixed responses to CO₂-induced ocean acidification. Geology, 37(12), 1131-1134,
1261 <https://doi.org/10.1130/G30210A.1>, 2009.

1262 Ries, J.B.: 2011. A physicochemical framework for interpreting the biological calcification
1263 response to CO₂-induced ocean acidification. Geochimica et cosmochimica acta, 75(14), pp.4053-
1264 4064, 2011.

1265 Sett, S., Bach, L. T., Schulz, K. G., Koch-Klavsen, S., Lebrato, M., and Riebesell, U.: Temperature
1266 modulates coccolithophorid sensitivity of growth, photosynthesis and calcification to increasing
1267 seawater pCO₂ [dataset]. PANGAEA, <https://doi.org/10.1594/PANGAEA.835214>, Supplement
1268 to: Sett, S et al. (2014): Temperature Modulates Coccolithophorid Sensitivity of Growth,

1269 Photosynthesis and Calcification to Increasing Seawater pCO₂. PLoS ONE, 9(2), e88308,
1270 <https://doi.org/10.1371/journal.pone.0088308>, 2014.

1271 Schulz, K.G., Bach, L.T. and Dickson, A.G.: Seawater carbonate system considerations for ocean
1272 alkalinity enhancement research. State of the Planet Discussions, 2023, 1-24, 2023.

1273 Sciandra, A., Harlay, J., Lefèvre, D., Lemee, R., Rimmelin, P., Denis, M., and Gattuso, Jean-
1274 Pierre: Seawater carbonate chemistry and processes during experiments with *Emiliana huxleyi*
1275 (TW1), 2003. PANGAEA, <https://doi.org/10.1594/PANGAEA.727841>, Supplement to: Sciandra,
1276 A. et al.: Response of coccolithophorid *Emiliana huxleyi* to elevated partial pressure of CO₂ under
1277 nitrogen limitation. Mar. Ecol. Prog. Ser., 261, 111-122, <https://doi.org/10.3354/meps261111>,
1278 2003.

1279 Seabold, S., & Perktold, J.: Statsmodels: Econometric and statistical modeling with Python. In *9th*
1280 *Python in Science Conference* (pp. 57–61). Austin, TX, June 28–July 3, 2010.
1281 <https://doi.org/10.25080/Majora-92bf1922-011>, 2010.

1282 Sinutok, S., Hill, R., Doblin, M. A., Wuhrer, R., and Ralph, P.J. : Seawater carbonate chemistry,
1283 calcification rate, oxygen production, maximum quantum yield, symbiont density, chlorophyll
1284 concentration and crystal width of *Halimeda macroloba*, *Halimeda cylindracea* and *Marginopora*
1285 *vertebralis* during experiments, 2011. PANGAEA, <https://doi.org/10.1594/PANGAEA.774792>,
1286 Supplement to: Sinutok, S. et al. : Warmer more acidic conditions cause decreased productivity
1287 and calcification in subtropical coral reef sediment-dwelling calcifiers. Limnol. Oceanogr., 56(4),
1288 1200-1212, <https://doi.org/10.4319/lo.2011.56.4.1200>, 2011.

1289 Sordo, L., Duarte, C., Joaquim, S., Gaspar, M. B., Matias, D.: Seawater carbonate chemistry and
1290 growth and survival of juveniles of the striped venus clam *Chamelea gallina*. PANGAEA,
1291 <https://doi.org/10.1594/PANGAEA.937477>, 2021.

1292 Stoll, H. M., Cruzado, A., Shimizu, N., and Kanamaru, K. : Seawater carbonate chemistry and
1293 B/Ca, calcification rate of *Emiliana huxleyi* and *Coccolithus braarudii*. PANGAEA,
1294 <https://doi.org/10.1594/PANGAEA.949913>, 2012.

1295 Sulpis, O., Lauvset, S. K., and Hagens, M.: Current estimates of K₁* and K₂* appear inconsistent

1296 with measured CO₂ system parameters in cold oceanic regions. *Ocean Sci.* 16(4), 847–862,
1297 <https://doi.org/10.5194/os-2020-19>, 2020.

1298 Tatters, A. O., Schnetzer, A., Fu, F., Lie, A.Y.A., Caron, D. A., and Hutchins, D. A.: Short- versus
1299 long-term responses to changing CO₂ in a coastal dinoflagellate bloom. PANGAEA,
1300 <https://doi.org/10.1594/PANGAEA.823381>, Supplement to: Tatters, A.O., et al.: Short- versus
1301 long-term responses to changing CO₂ in a coastal dinoflagellate bloom: implications for
1302 interspecific competitive interactions and community structure. *Evolution*, 67(7), 1879-1891,
1303 <https://doi.org/10.1111/evo.12029>, 2013.

1304 Truchot, J.-.: Water carbonate alkalinity as a determinant of hemolymph acid-base balance in the
1305 shore crab, *Carcinus maenas*: a study at two different ambient P CO₂ and P O₂ levels, *J. Comp.*
1306 *Physiol.. B*, 154(6), 601–606. <https://doi.org/10.1007/bf00684414>, 1984.

1307 Truchot, J.: Changes in the Hemolymph Acid-Base State of the Shore Crab, *Carcinus maenas*,
1308 Exposed to Simulated Tidepool Conditions, *Biol. Bull.*, 170(3), 506–518,
1309 <https://doi.org/10.2307/1541858>, 1986.

1310 Uppström, L. R.: The boron/chlorinity ratio of deep-sea water from the Pacific Ocean, *Deep-sea*
1311 *Research And Oceanographic Abstracts*, 21(2), 161–162, [https://doi.org/10.1016/0011-](https://doi.org/10.1016/0011-7471(74)90074-6)
1312 [7471\(74\)90074-6](https://doi.org/10.1016/0011-7471(74)90074-6), 1974.

1313 Uthicke, S., and Fabricius, K.E.: Seawater carbonate chemistry, productivity and calcification of
1314 *Marginopora vertebralis* in a laboratory experiment. PANGAEA,
1315 <https://doi.org/10.1594/PANGAEA.831207>, Supplement to: Uthicke, S., Fabricius, K.E.:
1316 Productivity gains do not compensate for reduced calcification under near-future ocean
1317 acidification in the photosynthetic benthic foraminifer species *Marginopora vertebralis*. *Glob.*
1318 *Change Biol.*, 18. Zenodo. <https://doi.org/10.5281/zenodo.12806137>, 2024.

1319 Van de Waal, D. B., John, U., Ziveri, P., Reichart, Gert-Jan, Hoins, M., Sluijs, A., and Rost, B.:
1320 Seawater carbonate chemistry and growth, calcification of *Thoracosphaera heimii* in a laboratory
1321 experiment. PANGAEA, <https://doi.org/10.1594/PANGAEA.824705>, Supplement to: Van de
1322 Waal, D.B., et al. : Ocean Acidification Reduces Growth and Calcification in a Marine
1323 Dinoflagellate. *pone.0065987*, 2013.

- 1324 Vázquez-Elizondo, R. M., and Enríquez, S.: Coralline algal physiology is more adversely affected
1325 by elevated temperature than reduced pH. PANGAEA,
1326 <https://doi.org/10.1594/PANGAEA.860802>, Supplement to: Vázquez-Elizondo, R.M., and
1327 Enríquez, S.: Coralline algal physiology is more adversely affected by elevated temperature than
1328 reduced pH. *Scientific Reports*, 6, 19030, <https://doi.org/10.1038/srep19030>, 2016.
- 1329 Waldbusser, G. G., Voigt, E. P., Bergschneider, H., Green, M. A., and Newell, R.I. E.: Seawater
1330 carbonate chemistry and calcification rate of eastern oyster *Crassostrea virginica*, 2011.
1331 PANGAEA, <https://doi.org/10.1594/PANGAEA.758181>, Supplement to: Waldbusser, G.G., et al.
1332 : Biocalcification in the Eastern Oyster (*Crassostrea virginica*) in Relation to Long-term Trends in
1333 Chesapeake Bay pH. *Estuar. Coast.*, 34(2), 221-231, <https://doi.org/10.1007/s12237-010-9307-0>,
1334 2011.
- 1335 Wang, X., Feng, X., Zhuang, Y., Lu, J., Wang, Y., Gonçalves, R.J., Li, X., Lou, Y., and Guan, W.:
1336 Seawater carbonate chemistry and physiology and toxicity of dinoflagellate *Karenia mikimotoi*.
1337 PANGAEA, <https://doi.org/10.1594/PANGAEA.923683>, 2019.
- 1338 Wang, H., Pilcher, D.J., Kearney, K.A., Cross, J.N., Shugart, O.M., Eisaman, M.D. and Carter,
1339 B.R. Simulated impact of ocean alkalinity enhancement on atmospheric CO₂ removal in the
1340 Bering Sea. *Earth's Future*, 11(1), p.e2022EF002816, 2023.
- 1341 White, M. M., Drapeau, D.T., Lubelczyk, L. C., Abel, V. C., Bowler, B. C., and Balch, W. M.:
1342 Seawater carbonate chemistry and calcification of an estuarine coccolithophore. PANGAEA,
1343 <https://doi.org/10.1594/PANGAEA.923623>, 2018.
- 1344 Zhang, M., Fang, J., Zhang, J., Li, B., Ren, S., Mao, Y., and Gao, Y.: Seawater carbonate chemistry
1345 and calcification and respiration of *Chlamys farreri*. PANGAEA,
1346 <https://doi.org/10.1594/PANGAEA.949604>, 2011.