Coupling the regional climate model ICON-CLM v2.6.6 into the Earth system model GCOAST-AHOI v2.0 using OASIS3-MCT v4.0

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Abstract

Interactions and feedback between compartments of the Earth system can have a significant impact on local and regional climate and its changes due to global warming. These effects can be better represented by regional Earth system models (RESMs) than by traditional stand-alone atmosphere and ocean models. Here, we present the RESM GCOAST-AHOI version 2.0, which includes a new atmospheric component, the regional climate model ICON-CLM, which is coupled with the ocean model NEMO and the hydrological discharge model HD via the OASIS3-MCT coupler. The GCOAST-AHOI model has been developed and applied for climate simulations over the EURO-CORDEX domain. Two 11-year simulations from 2008-2018 of the uncoupled ICON-CLM and GCOAST-AHOI give similar results for seasonal and annual means of near-surface air temperature, precipitation, mean sea level pressure and wind speed at 10 m height. However, GCOAST-AHOI has a cold SST bias of 1-2 K over the Baltic and the North Seas, most pronounced in winter and spring seasons. A possible reason for the cold SST bias could be the underestimation of the downward shortwave radiation at the surface of ICON-CLM with the current model settings. Despite the cold SST bias, GCOAST-AHOI was able to capture other key variables such as those mentioned above well. Therefore, GCOAST-AHOI can be a useful tool for long-term climate simulations over the EURO-CORDEX domain. Compared with the stand-alone NEMO3.6 forced by ERA5 and ORAS5 boundary forcing, GCOAST-AHOI has positive biases in sea ice fraction and salinity, but negative biases in runoff, which need to be further investigated in the future to improve the coupled simulations. The new OASIS3-MCT coupling interface OMCI implemented in the ICON-CLM model adds a possibility to couple ICON-CLM with an external ocean model and an external hydrological discharge model using OASIS3-MCT instead of the YAC coupler. Using OMCI, it is also possible to set up a RESM with ICON-CLM and other ocean and hydrology models possessing the OASIS3-MCT interface for other regions, such as the Mediterranean Sea.

Keyword: GCOAST, ICON-CLM, OASIS3-MCT coupling interface, climate simulations, EURO-CORDEX, RESM

1   Introduction

GCOAST (Geesthacht Coupled cOAstal model SysTem) is an Earth system framework developed at Helmholtz-Zentrum Hereon, Germany (Staneva et al., 2018). GCOAST is a modular system of

different models each developed for a specific component of the Earth system. Based on a specific scientific question, different models from GCOAST can be selected. These models can be plugged together by various couplers, such as OASIS3-MCT (Valcke et al., 2015), ESMF (Earth System Modeling Framework; Hill et al., 2004), or FABM (Framework for Aquatic Biogeochemical Models; https://fabm.net). The coupling can be done at different levels of coupling granularity and the couplers handle the exchange of information between model combinations, individual models, and processes.

GCOAST systems have been applied for several studies covering the Baltic and-North Sea region and part of the North Atlantic. These studies include atmosphere-river-ocean-sea ice coupling (Ho-Hagemann et al., 2020), atmosphere-wave coupling (Wahle et al., 2017; Wiese et al., 2019, 2020), wave-ocean coupling (Staneva et al., 2016; Schloen et al., 2017; Lewis et al., 2019), hydrosphere-biosphere coupling for the Elbe estuary (Pein et al., 2019), the total organic carbon-macrobenthos coupling model (Zhang et al., 2019), and multi-model couplings developed by Lemmen et al. (2018), which have been applied to assess ecosystem impacts of offshore wind farms (Slavik et al., 2019).

So far, the atmospheric model component of GCOAST has been the non-hydrostatic limited area atmospheric model COSMO-CLM v5.0 (Rockel et al., 2008). The COSMO (cOnsortium for Small-scale mOdeling) model was initially developed by the Deutscher Wetterdienst (DWD, the German Meteorological Service) in the 2000s as a limited-area weather forecast model. Later, it was further developed in the Climate Limited-area Modeling Community (CLM-Community) as the regional climate model COSMO-CLM (hereafter referred to as CCLM). In December 2021, COSMO v6.0 was released which is the last version of the COSMO model. With this release, the development of the COSMO model ended after more than two decades. The successor of the COSMO model is the ICON model.

In 2001, a cooperation between the DWD and the Max-Planck Institute for Meteorology (MPI-M) was initiated, with the aim of developing a new modelling system for weather prediction and climate simulations. As one result of this initiative, the global numerical weather prediction model ICON (Icosahedral Nonhydrostatic) was developed (Zängl et al., 2015). Nowadays, with contributions from the Karlsruhe Institute of Technology (KIT) and the German Climate Computing Center (DKRZ), etc., the ICON Earth system framework can include not only the atmospheric, land, river routing, ocean-sea ice, wave, and biogeochemical compartments but also the Aerosols and Reactive Trace gases (ART) model. ICON can be set up to operate on several high-performance computing systems such as Bull ATOS at DKRZ (Hamburg, Germany), NEC-Aurora Tsubasa at DWD (Offenbach, Germany) or BullSequana at Forschungszentrum Jülich (FZJ, Jülich, Germany). ICON can be used on a wide range

of scales from climate projection, climate prediction and numerical weather prediction down to large-eddy simulations (Dipankar et al., 2015; Heinze et al., 2017).

The atmospheric component of ICON includes two different physics packages: the first one is the Numerical Weather Physics package of DWD (i.e. the ICON-NWP model); and the second one is the ECHAM physics package of MPI-M (i.e. the ICON-A model, Giorgetta et al., 2018). The global atmospheric model ICON-A is coupled with the global ocean model ICON-O (Korn, 2017) and the land and biosphere model JSBACH (Reick et al., 2021) within the ICON Earth System Model (ICON-ESM; Jungclaus et al., 2022). ICON-NWP can be also coupled with ICON-O in the ICON-Seamless Earth system coupling framework which has been newly developed during the recent years. In ICON-Seamless, there are two options for the land surface schemes, TERRA and JSBACH, which are coupled via subroutine to the atmospheric component. A new land surface model (ICON-Land) is being developed based on JSBACH and some features of TERRA. Another component of ICON-Seamless is the wave model ICON-Wave. The hydrological discharge model HD can now be used as an external model instead of being coupled as a subroutine of JSBACH.

The components of ICON are coupled together using YAC (Yet Another Coupler; Hanke et al., 2016). However, coupling between ICON components, or the coupling of multiple ICON components to an external model without a YAC coupling interface is not supported due to how the initial communicator splitting is implemented.

ICON can also be used in a configuration with regional grid refinement (2-way nesting) or in limited area mode. ICON-LAM is the Limited-Area Mode of ICON-NWP. Starting in 2017, DWD and the CLM-Community decided to develop the climate limited area mode (ICON-CLM, Pham et al., 2021) based on ICON-LAM. Within ICON-Seamless, a limited area mode of the ocean model (ICON-O-LAM) is being developed and can be coupled with ICON-LAM via YAC.

As mentioned above, coupling a component or multiple components of ICON to an external model that has no YAC coupling interface is not supported and potentially impossible. For the coupling of ICON-CLM as an atmosphere component into GCOAST-AHOI, which includes HD and the ocean model NEMO (Nucleus for European Modelling of the Ocean, Madec et al. 2017), representing the ocean and sea ice components, basically, there were two feasible options: either to implement a YAC interface in NEMO and HD, or to implement an OASIS interface in ICON-CLM. For the option one, the YAC coupling interface was added into the HD source code by M. Hanke (DKRZ) (see Hagemann et al. 2023), but YAC has not yet been available in the NEMO source code. To our current knowledge, there is no RESM with NEMO using the YAC coupler. The NEMO model is already linked to the OASIS coupler, which has been used to couple NEMO with many other model components.

70   Implementing the YAC interface in NEMO would require a larger effort, as the NEMO source code is
71   much more complicated than the HD code. In addition, although the NEMO source code is freely
72   available, we are ordinary users in the NEMO community, not members of the model development
73   team. Therefore, implementing and especially maintaining the YAC interfaces in NEMO is a big
74   challenge. For the second option, HD and NEMO already have the OASIS3-MCT coupling interface
75   (OMCI), so all we had to do was to implement OMCI in ICON. Here, we also have the advantage that
76   we belong to the ICON development team of the CLM-Community, so that we can get great and
77   quick technical support from the development team when coding with ICON. Therefore, in 2021,
78   we started to port the OASIS coupling interfaces from CCLM to ICON-CLM for the coupling to NEMO
79   and HD.

80   Some other groups were using a similar method while coupling ICON into their available coupled
81   system model which does not include the YAC coupler. For example, Bauer et al. (2021) have
82   implemented the ESMF interfaces in an earlier version of ICON-NWP as well as into the ocean model
83   GETM to build up the regional ocean-atmosphere coupling over the Baltic Sea. However, they did
84   not consider sea ice in the coupling. There is an ongoing work at FZJ to couple ICON-CLM with the
85   Community Land Model (CLM) via the OASIS3-MCT coupler (manuscript in preparation) as has been
86   done for the CCLM via the OASIS interface (see Shrestha et al., 2014; Will et al., 2017).

87   The aim of this article is to give a detailed description of the OASIS3-MCT coupling interface
88   (hereafter referred to as OMCI) in ICON-CLM (ICON release version 2.6.6), how to implement OMCI
89   with as little modification of the ICON source code as possible, how to compile it on the high-
90   performance computing system Levante at DKRZ, and how to run the coupled system model
91   GCOAST-AHOI with ICON-CLM for climate simulations over the EURO-CORDEX domain. This
92   information is useful to other groups planning to couple ICON-CLM with NEMO or any other ocean
93   model that already has an OASIS3-MCT interface available. The Earth System Modelling (ESM)
94   Community agrees that ICON and IFS (coupled to FESOM and NEMO) will play a central role within
95   the Helmholtz Association of German Research Centres (HGF). This new OMCI opens more
96   opportunities to use ICON-CLM in ESM applications as well as in other modelling communities. The
97   OMCI can also be applied to couple with a land surface model with minor necessary adaptations.

98   We briefly introduce the coupled system model GCOAST-AHOI in Section 2 and describe the
99   details of OMCI in ICON-CLM in Section 3. Experiment setups are presented in Section 4, followed
100   by an analysis of the model simulations in Section 5. Finally, conclusions and a discussion are given
101   in Section 6.

2 The coupled system model GCOAST-AHOI

GCOAST-AHOI is a subset of GCOAST that includes model components for **A**-Atmosphere and Land, **H**-Hydrological discharge, **O**-Ocean, and **I**-Sea Ice. GCOAST-AHOI version 1.0 (Ho-Hagemann et al., 2020) contains the atmospheric model CCLM v5.0, the ocean model NEMO v3.6 (including the sea ice model LIM3) and the hydrological discharge model HD v4.0 (Hagemann and Dümenil, 1998; Hagemann et al., 2020), coupled via OASIS3-MCT v2.0. A detailed description of CCLM, NEMO and HD as components of GCOAST-AHOI can be found in Ho-Hagemann et al. (2020).

In the GCOAST-AHOI version 2.0, ICON-CLM replaces CCLM as atmospheric model, which is coupled to NEMO v3.6 and HD v5.1 via OASIS3-MCT v4.0. By coupling the atmosphere-ocean-river runoff models in GCOAST-AHOI, we aim to close the water balance in the RESM. Figure 1 illustrates the three models exchanging radiation, wind, pressure, temperature, humidity, water, and sea ice related variables at their interfaces via the OASIS coupler.

The OMCI in NEMO v3.6 has been modified compared to the original one in the officially released version at http://forge.ipsl.jussieu.fr/nemo/wiki/Users/release-3.6 to be able to receive state variables from the atmospheric model (Ho-Hagemann, 2024). Supplement S1 contains a flowchart of the OMCI for NEMO v3.6. This flowchart differs slightly from Figure 9 in Will et al. (2017), who used the older version NEMO v3.3. The OMCI in HD can be found in the source code publication of Hagemann and Ho-Hagemann (2021) and Hagemann et al. (2023). Supplement S2 shows the OMCI of HD. In this article, we describe in detail the OMCI in ICON-CLM.

In Section 3, we demonstrate the construction of the OMCI in ICON-CLM and the optional coupling methods between ICON-CLM and NEMO.

3 The coupling OASIS3-MCT interface in ICON

3.1 Interface structure

Figure 2 shows a flowchart of ICON with the OMCI implemented for coupling with NEMO and HD. 10 levels of ICON's source code are described: the first level is the main program ICON, the second level starts with the *start_mpi*, then *atmo_model* and ends with *stop_mpi*, etc.

Levels 2 to 6, 8 and 9 comprise subroutines of ICON (marked in red) that are modified by the coupling. On levels 3 to 7 and 10, new subroutines (orange boxes B1-B7) have been added with the OMCI. They are organized in three modules (cpl_oas_vardef.f90, cpl_oas_mpi.f90 and cpl_oas_interface.f90) containing about 3000 lines of Fortran code (including the current debug lines). The files have been added to the icon/externals/oasis3-mct directory and linked to the src/atm_phy_nwp directory of the ICON source tree. The detailed description of the interface

134   structure can be found in Supplementary S4.

135   Supplement S5 contains a guide for compiling ICON with this OMCI on Levante at DKRZ. The

136   preparation of OASIS input files for GCOAST-AHOI is described in Supplement S6, which is

137   accompanied by an example of the namcouple file in Supplement S7 and the namelist_cpl_atm_oce

138   in Supplement S8. The command to run GCOAST-AHOI on Levante is provided in Supplement S9.

139   The complete package to conduct experiments for this study is included in the Starter Package for

140   ICON-CLM Experiments (SPICE; Rockel and Geyer, 2022), which is a workflow engine to easily

141   perform long-term simulations. This tool has been further developed from the ICON-CLM_SP starter

142   package (Pham et al. 2021). Some additional parts for coupling with NEMO and HD have been added

143   to the original package.

3.2   Coupling methods

145   In the officially released version of NEMO v3.6, several fluxes and variables, including shortwave

146   (SW) and longwave (LW) radiation fluxes, latent (LH) and sensible heat (SH) fluxes, rain, snow,

147   evaporation, ice sublimation, mean sea level pressure (MSLP) and surface momentum, can be sent

148   from an atmospheric model to NEMO via the OASIS3-MCT coupler. To be able to receive state

149   variables from the atmospheric model, the OMCI in NEMO v3.6 has been modified to allow air

150   temperature and air specific humidity at 2 m height (T_2M and QV_2M respectively) to be sent from

151   the atmospheric model to NEMO. This allows NEMO to use these variables to calculate the LH and

152   SH, as in the case of the stand-alone NEMO using the "CORE bulk formulae" (Large and Yeager,

153   2004). Thus, we have three options for the coupling method between ICON and NEMO:

154   a) CPL_flx: **flux coupling**, which is the default option in the NEMO source code (described above)

155   b) CPL_var: **state variable coupling**, the new method, where SW and LW, T_2M, QV_2M, wind

156   speed at 10 m height (UV_10M), rain, snow, MSLP and surface momentum are sent from

157   ICON-CLM to NEMO. NEMO calculates LH and SH using the "CORE bulk formulae" which is

158   based on the Monin Obukhov similarity theory.

159   c) CPL_mix: **mixture coupling**, the new method, like CPL_var, but ICON-CLM also sends LH and

160   SH to NEMO. NEMO then averages them with the LH and SH calculated using the "CORE bulk

161   formulae".

162   With the modification of OMCI in NEMO v3.6, it is now easy to select the coupling method via

163   the namelist settings. Section 5 considers the simulations using the coupling method 3 (CPL_mix),

164   which was also used in Ho-Hagemann et al. (2020). An extra experiment using the coupling method

165   1 (CPL_flx) is also conducted and analyzed in Section 5.

166   In turn, NEMO sends the sea surface temperature, sea ice fraction and sea ice albedo to ICON-

167   CLM. Figure 3 illustrates how the surface temperature is updated in ICON over the ocean (left side)

168   and over land (right side) in the presence of sea ice and snow. ICON utilizes a tile approach to

169   compute surface fluxes of momentum and scalars. For the "sea-water type" grid boxes, the grid box

170   mean fluxes are computed as a weighted average of the fluxes over ice and over open water, using

171   the fractional ice cover *fice* and the fractional open water cover (*1 – fice*) as the respective weights.

172   Sea ice in each ICON grid box is considered only if *fice* exceeds its minimum value of 0.015.

173   Otherwise, the grid box is treated as ice-free. In ICON, two types of surface temperature are

174   considered: the ground temperature t_g and the surface temperature t_s. If a grid box is covered

175   by sea ice or snow, t_g is the mixed temperature of the free sea ice/free snow surface temperature

176   and the temperature on top of the sea ice/snow. Under the sea ice, t_s is calculated as a mixture of

177   the free sea ice temperature and the salt water freezing temperature of 271.45 K. If there is no sea

178   ice or snow in the grid box, t_g is equal to t_s. In principle, NEMO can send the mixed sea ice and

179   water temperature to ICON to update t_g over the ocean points, as in CCLM in Ho-Hagemann et al.

180   (2020). Or it can send the open water temperature, the sea ice surface temperature and the sea ice

181   fraction so that ICON can calculate t_g as the mixture. However, in the uncoupled mode of the

182   current ICON-CLM version, the sea surface temperature (SST) forcing is read in as the variable

183   t_seasfc (or t_s_w in Fig. 3) and passes through the subroutines *nwp_surface_init* and

184   *process_sst_and_seaice* to calculate t_g. To be consistent with the ICON-CLM updates, we pass the

185   SST (to update the t_seasfc), the sea ice fraction (to update fr_seaice), and the sea ice albedo

186   (alb_si_ext) from NEMO to ICON. ICON will then calculate t_g, t_s, alb_si, etc. using its sea ice

187   scheme. In the future, we may modify this coupling method by using the sea ice temperature from

188   NEMO.

189   **4   Experimental design**

190   In this study, four main experiments are conducted for the period of 2008-2018 (Table 2). The

191   uncoupled ICON (**ICON266**), the coupled GCOAST-AHOI (**ICPL266**), the stand-alone NEMO v3.6

192   (**NEMO3.6**), and the stand-alone HD v5.1 forced by ICON266 runoffs (**HDICON266**). Two additional

193   experiments are conducted: **ICPL266_noNewa** as a sensitivity test for the location of the Newa River

194   mouth on NEMO grid, and **ICPL266_flx** to test the coupling method CPL_flx.

195   Each experiment starts on 01 January 2008 and ends on 01 January 2019, restarting each month.

196   The integration domains of ICON, NEMO and HD are displayed in Fig. 4. The namelist setup of

197   physical parametrization for ICON-CLM is similar to that of the NUKLEUS project (B. Geyer, personal

198   communication). The resolution of ICON is R13B5, with an approximate mesh size of 12 km, using

199   60 vertical levels. The model top height is at 23.5 km. The following physical schemes are used in

200   the current namelist setting of ICON: Radiation scheme ecRad (Hogan and Bozzo, 2018; Rieger et

201   al., 2019); Mass-flux shallow and deep convection scheme (Tiedtke, 1989; Bechtold et al., 2008);

202   Microphysics single-moment scheme (Doms et al., 2004); Planetary boundary layer scheme

203   prognostic TKE (Raschendorfer, 2001; Raupach and Shaw, 1982); Land-surface scheme tiled TERRA

204   (Schrodin and Heise, 2001; Schulz et al., 2016; Schulz and Vogel, 2020). The initial and lateral

205   boundary forcing of ICON is obtained from the ERA5 reanalysis data (Hersbach et al., 2020). The

206   Tegen aerosol climatology (Tegen, 1997), i.e. a monthly aerosol optical depth of sulphate droplets,

207   total dust, organic carbon, black carbon, and sea salt, is used in this study. The initial and daily lateral

208   boundary forcing of NEMO is taken from the ORAS5 reanalysis data (Copernicus Climate Change

209   Service, 2021). The spatial resolution of NEMO is ~3.7 km with 50 vertical levels. In the stand-alone

210   mode, NEMO3.6 is driven by the atmospheric ERA5 data, the ocean ORAS5 data and climatological

211   river runoff data. HD has the resolution of 1/12 degrees, ca. 8 km. More information on the model

212   configuration can be found in Table 1 and in Ho-Hagemann et al. (2020).

213   To estimate the computational performance of the coupled model, we used LUCIA (Maisonnave

214   and Caubel, 2014), which is part of OASIS3-MCT. In Supplement S10 and Fig. S1, we describe how to

215   use LUCIA for GCOAST-AHOI to optimize the computational performance.

5   Evaluation of model simulations

216

217   The first two years 2008-2009 are excluded as spin-up time, and the output data of the two

218   simulations ICON266 and ICPL266 for nine years (2010-2018) are compared with the observational

219   and the ERA5 reanalysis data to assess the model performance. For sea surface temperature (SST),

220   we use the Operational Sea Surface Temperature and Ice Analysis (OSTIA) data (Good et al. 2020)

221   to evaluate the simulated SST of ICPL266. For air temperature at 2 m height (T_2M) and precipitation

222   (TOT_PREC), the daily E-OBS data (Haylock et al. 2008; Van den Besselaar et al. 2011) version 27.0

223   on the grid of 0.11 degree are used. The ERA5 reanalysis data are interpolated onto the E-OBS grid

224   and used as a reference for comparison with the simulated shortwave and longwave surface

225   radiation, turbulent fluxes, mean sea level pressure (PMSL), wind speed at 10 m height (SP_10M),

226   and T_2M. The Surface Radiation Data Set - Heliosat (SARAH) - Edition 2 (Pfeifroth et al. 2017) is

227   used to evaluate the shortwave downward radiation of the simulations.

228   Seasonal means of winter (DJF), spring (MAM), summer (JJA), autumn (SON) and annual means

229   (ANN) of several variables are analyzed in this section.

5.1 Sea surface temperature and 2m air temperature

Over the ocean, the sea surface temperature (SST) of ICON266 is the ERA5 forcing data, which is based on observations, so it's very close to the OSTIA data (not shown). Thus, the SST difference between the coupled and the stand-alone ICON-CLM simulations (Fig. S2) can be interpreted as a bias towards a measurement-based product. In the coupled model, the SST is provided by NEMO over the GCOAST domain. In general, ICPL266 has a cold SST bias of about 1-3 °C compared to OSTIA over the GCOAST domain, except around the British coast in summer (JJA, Fig. 5a). The SST bias of ICPL266_flx (Fig. 5b) is similar to ICPL266 (Fig. 5a). Ho-Hagemann et al. (2020) noted that using CPL_flx when coupling COSMO-CLM with NEMO leads to larger biases in the SST than using CPL_mix. This is not the case here when coupling ICON-CLM. A possible reason for this is that due to the tile approach (cf. Sect. 3.2), the fluxes from ICON-CLM to NEMO are sent separately over water and sea ice, while COSMO-CLM v5.0 doesn't have the tile approach, therefore, the fluxes in each ocean grid box sent from the atmosphere to the ocean are the mixed fluxes of water and sea ice.

The annual mean SST bias of the stand-alone NEMO3.6 is less than 0.5 °C over the Baltic and North Seas, and of about -1 to -2 °C over the North Atlantic compared to the OSTIA data (Fig. 5c). In summer, a positive SST bias of about 1-2 °C is present over the Baltic and North Seas. In this case, the reduction of SST by the coupling reduces the warm bias of the stand-alone NEMO3.6.

The cold SST bias of ICPL266 over the GCOAST domain intensifies the cold T_2M bias (Fig. 6b, Fig. S2b), especially in winter (DJF) and spring (MAM). In summer, ICPL266 reduces the warm T_2M bias of ICON266 (Fig. 6a). In general, the annual (ANN) T_2M bias of ICPL266 is slightly more negative by about 0.5 °C than that of ICON266. ICPL266_flx reproduces a similar T_2M (Fig. 6c) to that of ICPL266. Comparison with the E-OBS data (Fig. S3) shows similar results to Fig. 6, except over Northern Africa and Turkey, where the quality of the E-OBS data is affected by the lack of observations in that region (cf. Fig. 1 in Hagemann and Stacke, 2022).

5.2 Shortwave radiation, longwave radiation, and turbulent fluxes

A possible reason for the SST cold bias of ICPL266 may be that the shortwave and longwave radiation from ICON-CLM sent to NEMO is too low. Figure S4 shows the relative bias (%) of the shortwave downward radiation (SWDN) of ICON266 and ICPL266 compared to the ERA5 data, as well as the relative difference (%) between SARAH2 and ERA5. Figure 7 shows a zoomed section of Fig. S4 over the GCOAST ocean domain (note the adapted color scale). In general, both ICON266 and ICPL266 have a positive SWDN bias of less than 10 % over land compared to ERA5, except for the larger bias of 15-20 % over northern Europe in winter and eastern Europe in autumn (Fig. S4). Over the North Sea, ICON266 and ICPL266 have a small negative bias of about 5-10 % compared to ERA5 (Fig. 7).

263 The area of negative SWDN bias in the North Sea is slightly larger in ICPL266 than in ICON266.

264 Comparing the ERA5 data and the SARAH2 data, SWDN over southern Europe is similar between the

265 two datasets, with SARAH2 being slightly larger over land (Fig. S4). In general, the SWDN of ICON266

266 and ICPL266 over the North Sea is rather close to the SARAH2 data, but slightly overestimated over

267 the Baltic Sea.

268 Figure S5 in the Supplementary Appendix shows a similar plot to Fig. S4, but for the longwave

269 downward radiation (LWDN) and without the SARAH2 data, as it is not available. The modeled

270 LWDN has a negative bias of about 2-4 % annually and a larger bias in winter of about 6-8 %, most

271 pronounced over land. Over the ocean, ICON266 reproduces well the LWDN of the ERA5 data, and

272 ICPL266 has a small negative bias of 2-4 %. Overall, the negative SWDN bias over the North Sea and

273 the negative LWDN bias of ICON-CLM give an indication why the ICPL266 shows an increased cold

274 SST bias.

275 The ERA5 reanalysis data is used as the atmospheric forcing for the uncoupled NEMO3.6, and the

276 namelist settings of the NEMO model used in this study were tuned for an SST close to OSTIA (Fig.

277 5c). If the same namelist settings of NEMO3.6 are used for ICPL266, to reduce the cold SST bias over

278 the North Sea in the coupled simulations, a bias correction for SWDN and LWDN should be done.

279 Figures S6a, b and Figures S7 a, b show the seasonal SWDN and LWDN of ICPL266 and NEMO3.6

280 averaged over the North Sea and Baltic Sea for the period of 2010-2018. Note that we don't show

281 the result of ICPL266_flx in Figs. S6 and S7 because there is no output of LWDN in ICPL266_flx due

282 to the setup of the CPL_flx coupling method. Over the North Sea, the SWDN of ICPL266 is smaller

283 than the ERA5 used for NEMO3.6 in spring and summer (Fig. S6a), which mainly leads to the cold

284 SST bias of ICPL266 (Fig. 5a). Therefore, we plan to increase the SWDN of ICON by about 10 % before

285 sending it to NEMO. However, the cold SST bias over the Baltic Sea does not seem to be directly

286 related to the SWDN as there is no clear SWDN difference between ICPL266 and NEMO3.6 in

287 summer or in any other season (Fig. S6a). The LWDN of ICPL266 is similar to NEMO3.6 in summer

288 but slightly smaller in the other three seasons, both over the North Sea and the Baltic Sea. Increasing

289 the LWDN in ICON-CLM by about 5-10 W/m2 before sending it to NEMO should be tested to reduce

290 the SST bias. Note that the seasonal cycle of LWDN is more pronounced over the North Sea than

291 over the Baltic Sea.

292 We also compare the turbulent heat flux (i.e. the sum of SH and LH) of NEMO3.6 and the flux of

293 NEMO in ICPL266, averaged over the North Sea and the Baltic Sea (Fig. S6c and Fig. S7c), and the

294 net downward heat flux, which is the sum of SWDN, LWDN, SH and LH (Fig. S6d and Fig. S7d). We

295 only consider the turbulent flux because NEMO doesn't write out SH and LH separately, but only the

296 SWDN, LWDN and the net downward heat flux. Note that the turbulent flux from NEMO3.6 is

297 calculated using the "CORE bulk formulae", and due to the CPLmix coupling method used, the

298 turbulent flux in NEMO from ICPL266 is the average of the flux from ICON-CLM and the one

299 calculated inside NEMO (see section 3.2). The results are similar for both seas. The turbulent flux

300 and the net downward flux of the two experiments are quite similar with the largest differences in

301 winter (DJF) and summer (JJA).

302 Using ERA5 as a reference, the SH and LH biases of ICON266 and ICPL266 are shown in Fig. S8

303 and Fig. S9. Over land, the bias of ICON266 is very similar to that of ICPL266. However, over the

304 ocean, the bias of ICPL266 is generally more positive (i.e. the fluxes are less negative) than that of

305 ICON266 with the largest bias over the North Atlantic. Smaller heat fluxes are consistent with lower

306 SSTs in ICPL266, as lower SSTs lead to larger stability and less vertical mixing. Over the North and

307 Baltic Seas, SH and LH of ICPL266 are quite close to ERA5. Despite of the SST forcing from ERA5,

308 ICON266 has a negative SH bias of about -5 to -15 W/m^2 over the North and Baltic Seas, especially

309 in winter. This suggests a future analysis of the difference in air temperature at the lowest level of

310 ICON-CLM and ERA5.

311 Besides the energy flux biases causing the cold SST bias, a spin-up time of 2 years may be too

312 short for NEMO to reach the stable state, leading to the cold SST bias. In addition, NEMO's namelist

313 settings should also be optimized for the coupled simulations.

314 Currently, in the COPAT2 (Coordinated Parameter Testing, phase 2) initiative of the CLM-

315 Community, several parameters of ICON-CLM are being tested in a similar way as done for the

316 COSMO-CLM model (Russo et al., 2024) to find out the recommended settings. For example, the use

317 of the transient aerosol MAC2-SP (Kinne, 2019; Stevens et al., 2017) and a careful adjustment of

318 various namelist settings related to cloud cover, the soil and vegetation scheme and the turbulent

319 transfer will further reduce the T_2M cold bias and improve the shortwave downward radiation.

5.3 *Precipitation, mean sea level pressure, and wind speed*

321 The precipitation biases (Fig. 9) of the three simulations ICON266, ICPL266 and ICPL266_flx

322 compared to the E-OBS data are very similar in general, with a wet bias in winter and spring and a

323 dry bias in summer (JJA) and autumn (SON). Fig. S10 and Fig. S11 in the Supplementary show the

324 biases of PMSL and SP_10M of ICON266 and ICPL266 compared to ERA5. The PMSL and SP_10M

325 figures of ICPL266_flx are not shown because they are very similar to those of ICPL266.

326 ICPL266 tends to overestimate the PMSL throughout the year except in the summer, while

327 ICON266 has only a pronounced positive bias in winter (DJF) and negative bias in summer (JJA). The

328 wind speed of the two simulations is very similar over land (Fig. S11). ICPL266 reduces the wind

329   speed over the GCOAST ocean domain by up to 1.5 m s^{-1} compared to ICON266 (Figs. 8, S11).

330   Therefore, while ICON266 has a positive bias of about 0.5 m s^{-1} over the North Sea and the Baltic

331   Sea in winter (Fig. S11a), ICPL266 is very close to ERA5 (Fig. S11b). In general, ICPL266 produces a

332   cooler SST and weaker wind speed than ICON266, consistent with the smaller SH mentioned above

333   in Section 5.2. This positive feedback is known as the thermal feedback (TFB) mechanism in the

334   atmosphere-ocean surface coupling process (Zhang and Perrie, 2001; Renault et al., 2023).

335   Figure 10 shows monthly climatology of different variables (T_S, T_2M, TOT_PREC, PMSL) over

336   the GCOAST domain or the whole EURO-CORDEX domain, considering only ocean or land points.

337   ICPL266 has a cold T_S bias of about 1-2 $^{\circ}$C over the ocean (Fig. 10a), which also causes the T_2M

338   bias of 0.5-1 $^{\circ}$C over the ocean (Fig. 10c). In winter, ICPL266 is slightly colder over land than ICON266

339   and E-OBS (Fig. 10d). In summer, both simulations are very close to E-OBS. The simulated

340   precipitation of ICON266 tends to be overestimated compared to E-OBS with a maximum in May

341   and June, and slightly underestimated in August and September (Fig. 10b). The coupled run shows

342   1-3 mm/month less precipitation than the atmosphere-only experiment. In previous studies by Ho-

343   Hagemann et al. (2015, 2017), the stand-alone atmospheric model COSMO-CLM has a dry bias in

344   summer and the coupled run reduces the dry bias due to the improvement of the moisture

345   convergence and transport from ocean to land. This situation is not found in the current study,

346   which needs to be thoroughly analyzed in the future.

347   For the PMSL, the whole EURO-CORDEX domain is considered, but separately for ocean points

348   (Fig. 10e) and land points (Fig. 10f). In both cases, ICPL266 has a larger PMSL than ICON266. The

349   higher surface pressure in ICPL266 may be caused by the cooler air near the surface (due to the

350   negative T_2M bias) which leads to a higher density of the air mass and therefore a higher pressure.

351   Over the ocean, the PMSL of ERA5 is better reproduced by ICPL266 than by ICON266. Over land,

352   ICPL266 increases the PMSL positive bias in winter compared to ICON266. ICPL266_flx and ICPL266

353   have similar results (not shown) indicating that the coupling methods used in GCOAST-AHOI v2.0

354   don't affect the simulated climate variables strongly in this study.

355   *5.4   Sea ice*

356   The sea ice fraction bias of ICPL266 is about 0.2-0.3 over the Bothnian Bay and Sea in winter and

357   spring (Fig. 11a), while ICPL266_flx has a larger bias of about 0.3-0.5 (Fig. 11b) and the ERA5-forced

358   NEMO3.6 has a relatively small positive ice fraction bias there (Fig. 11c). The monthly mean sea ice

359   fraction averaged over the Bothnian Bay and Sea from ICPL266 and NEMO3.6 compared to the

360   OSTIA data is shown in Fig. 11d, where the sea ice temporal variation is quite well captured by the

361   two models, with a high peak in spring 2010 and a low peak in spring 2015. However, all three

362 simulations overestimate the sea ice fraction of OSTIA, with the two coupled simulations showing a

363 larger positive bias also in the time series. While ICPL266 has a winter SWDN about 8 % larger than

364 ERA5 (Fig. 7b), the incoming shortwave radiation is relatively small over the high latitudes in winter.

365 Therefore, we don't expect the positive SWDN bias to be the main reason for the overestimation of

366 sea ice. However, the LWDN of ICPL266 is about 10 W/m^2 lower than ERA5 (the forcing for

367 NEMO3.6) in winter over the North and Baltic Seas (Figs. S6b, S7b), and T_2M is about 3 ^{o}C lower

368 than ERA5 over the Scandinavian region surrounding the Bothnian Bay and Sea (Fig. 6b). The cold

369 T_2M bias and negative LWDN bias of ICPL266 may explain its positive sea ice fraction bias. The cold

370 air temperature above the sea ice surface often produces more sea ice in winter and spring,

371 especially over an area with water of low salinity, such as the Bothnian Bay and Sea. Figure 6c shows

372 the larger T_2M cold bias of ICPL266_flx over the Baltic Sea in spring about 1°C more than ICPL266

373 (Fig. 6b) which is consistent with the larger sea ice fraction bias of ICPL266_flx (Fig. 11b) compared

374 with that of ICPL266 (Fig. 11a).

375   Another factor that could contribute to an increase in sea ice cover in spring would be an increase

376 in river runoff, which would result in less salty sea water and therefore more sea ice. These two

377 variables are analyzed in the next section.

5.5   Salinity and river runoff

379 As mentioned in section 4, NEMO3.6 uses a climatological dataset for river runoff. Therefore, a

380 rough verification of the river runoff produced by the HD model in ICPL266 can be made by

381 comparing against this climatological river runoff. Differences in sea surface salinity and river runoff

382 between ICPL266 and NEMO3.6 are shown in Figure 12. The salinity simulated by ICPL266 is about

383 0.3-1 PSU higher than that of NEMO3.6 along the British Isles and the North Sea coast, and about

384 0.9-1.8 PSU higher in the Baltic Sea. Two areas with the largest salinity differences of more than 2

385 PSU are found south of the Kattegat and in the Gulf of Finland (Fig. 12a). The river runoff differences

386 (Fig. 12b) are largest near the Ems and Newa estuaries with more than 0.1 and 0.2 kg/m2/s,

387 respectively. The small river runoff difference between the two models at Kattegat cannot be used

388 to directly explain the increase in salinity there. The river runoff differences near the Ems estuary

389 have opposite signs (blue point overlaid by a red one in Fig. 12b), but very similar values. The reason

390 for this may be the discrepancy in the locations of river mouths between the NEMO3.6 setup, where

391 the river runoff is taken from a climatology, and those in ICPL266, which are defined based on the

392 river mouths in the HD model and the NEMO land-sea mask. In the latter case, the river mouths of

393 the HD model are interpolated onto the NEMO grid by searching for the closest ocean point of

394 NEMO. For example, the Ems River mouth in ICPL266 may not be in the same position as in the

395 climatology data. This discrepancy would lead to a difference in salinity near the coast (see Fig. 12a).

396 The extent of the effect on salinity in the deeper layers of the ocean in a longer-term simulation

397 needs to be analyzed in the future.

398   The large difference in river runoff near the Newa estuary is also caused by a mismatch in the

399 locations of the river mouths. In this case, the mouth of the Newa River in the climatology data

400 (60.1333°N, 29.888°E) is located slightly northwest of its "real" location (Wikipedia: 59.9453°N,

401 30.1708°E). The interpolation program used to define the mouths of the HD rivers on the NEMO

402 grid, by searching for the closest ocean point to the mouth of the HD rivers, found the mouth of the

403 Newa River at (59.95835°N, 30.20825°E), which is very close to the "real" position and at the

404 furthest grid point to the east in the Gulf of Finland. However, in the NEMO model, this eastern

405 boundary point in the Gulf of Finland is masked as a buffer zone. Therefore, the discharge from HD

406 to NEMO at this point in ICPL266_noNewa was ignored in the NEMO calculations, resulting in a lack

407 of freshwater inflow to the Gulf of Finland in ICPL266 (Fig. 12b2) and consequently an increase in

408 salinity (Fig. 12a2). The ICPL266 simulation with the Newa River mouth located on the NEMO buffer

409 zone is referred to ICPL266_noNewa in Figure 12. To overcome the location deficiency, that the

410 Newa River mouth was shifted one grid point to the west on the NEMO grid to allow the large

411 amount of river runoff to enter the Gulf of Finland in the coupled model. Therefore, the salinity

412 difference of ICPL266 compared to NEMO3.6 is reduced (Fig. 12a1) and the river runoff difference

413 shows the shift of the river mouth instead of the missing one (Fig. 12b1). The shift of the Newa River

414 mouth has little effect on the simulated atmospheric variables, but improves the simulated salinity

415 in the Baltic Sea, which is important for ecosystem modelling when a marine biogeochemical or

416 ecosystem model such as ECOSMO (Daewel and Schrum, 2013) is coupled with GCOAST-AHOI in the

417 future.

418   Other river mouths in the Baltic Sea have river runoff differences of less than 1.4 kg/m^2/s when

419 comparing ICPL266 to NEMO3.6. In general, ICPL266 tends to simulate less river runoff than the

420 climatology, leading to increased salinity there. The sources of the river runoff used for NEMO in

421 ICPL266 are the surface and sub-surface runoffs from the land component in ICON-CLM that are

422 transported to the ocean by HD. We applied the HD model to calculate the discharge using the

423 ICPL266 and ICON266 surface and subsurface runoff (Table 3) to evaluate it against the discharge

424 observation. The annual discharge difference of ICPL266 and HDICON266 in the Baltic Sea is about

425 -11 %. However, HDICON266 with a discharge of 12449 m^3/s is about -20 % biased towards the

426 HELCOM (Helsinki Commission; Svendsen and Gustafsson, 2022) value of 15676 m^3/s. Note that for

427 Baltic Sea ocean models, the mean long-term bias of river runoff must be less than 7 % (Hagemann

428 and Stacke, 2022). In the North Sea, ICPL266 discharge is about -4 % compared to HDICON266, which

429 has an annual value of 6366 m^3/s. However, both models have a dry discharge bias compared to the

430 OSPAR data (Farkas and Skarbøvik, 2021), which is 9190 m^3/s.

431   The main driver of the runoff is precipitation. Figure 9 shows that over Scandinavia, ICON266 has

432 a wet bias of about 10-30 mm/month in spring and summer compared to the E-OBS data. Thus, even

433 with the wet precipitation bias, ICON266 has a dry discharge bias. At the same time, ICPL266

434 precipitation is lower than for ICON266 in spring and summer, and therefore closer to the E-OBS

435 data (Fig. 9b). This difference in precipitation between ICPL266 and ICON266 explains the -11 %

436 difference in discharge, which increases the dry discharge bias. The reduction of the precipitation

437 bias in ICPL266 leading to a larger discharge dry bias implies that a better simulation of precipitation

438 compared to observations does not necessarily lead to a better runoff. We note that the runoff from

439 the atmosphere-only ICON-CLM has a general dry bias, which can be attributed to the respective

440 parametrizations in the TERRA land surface scheme used in ICON-CLM (S. Hagemann, pers.

441 communication, 2024). This dry discharge bias can be improved either by using the JSBACH land-

442 surface model in ICON-Seamless, or by applying a discharge bias correction developed by Hagemann

443 et al. (2024).

444   In Section 5.4, it was speculated that an increase in river runoff would lead to less saline sea

445 water and therefore more sea ice over the Baltic Sea. In our study, ICPL266 simulates too little river

446 runoff, leading to increased salinity in the Baltic Sea, which would mean less sea ice. However, the

447 sea ice fraction is increased compared to the ERA5-forced NEMO3.6. Thus, the main factor causing

448 the bias in the sea ice fraction seems to be the cold bias in the air temperature over sea ice in the

449 Baltic Sea in winter and spring (Fig. 6b and c).

450 ## Conclusion and Outlook

451   In the present study, we introduce the regional Earth system model (RESM) GCOAST-AHOI

452 version 2.0, in which a new atmospheric component - the regional climate model ICON-CLM version

453 2.6.6 - is coupled with the ocean model NEMO version 3.6 and the hydrological discharge model HD

454 version 5.1 via the OASIS3-MCT coupler version 4.0.

455   GCOAST-AHOI v2.0 is developed and applied for climate simulations over the EURO-CORDEX

456 domain. Several 11-year simulations from 2008-2018 of the uncoupled ICON-CLM (ICON266) and

457 GCOAST-AHOI (ICPL266, ICPL266_flx, ICPL266_noNewa) yield similar results for seasonal and annual

458 means of near-surface air temperature and precipitation, as well as mean sea level pressure and

459 wind speed at 10 m height. However, GCOAST-AHOI has a cold SST bias of 1-2 °C over the Baltic and

460 the North Seas, most pronounced in winter and spring seasons. The coupling methods CPL_mix and

461 CPL_flx give similar biases of SST and other climate variables like T_2M, precipitation, PMSL, etc.

462 A possible reason for the cold SST bias of GCOAST-AHOI could be the underestimation of the

463 downward shortwave radiation at the surface of ICON-CLM with the current model settings. A

464 deeper analysis of the bias will be done in the next study, especially after re-running the simulations

465 with the optimal settings of ICON-CLM, which will be found within the COPAT2 initiative of the CLM-

466 Community. For example, the performance of ICON-CLM will be tuned by using the transient

467 MACv2-SP aerosol data (Kinne, 2019) and modified namelist parameters related to cloud cover to

468 improve the shortwave downward radiation and reduce the cold bias.

469 Despite the cold SST bias, GCOAST-AHOI was able to capture the distribution of temperature,

470 precipitation, mean sea level pressure and wind speed well, similar to the uncoupled ICON-CLM

471 model. However, GCOAST-AHOI provides larger biases in sea ice fraction and salinity over the Baltic

472 Sea compared to the stand-alone ocean simulation (NEMO3.6) forced by ERA5 and ORAS5. The sea

473 ice fraction bias is related to the cold T_2M bias in ICPL266 and ICPL266_flx. Using the flux coupling

474 method CPL_flx instead of CPL_mix doesn't strongly affect the bias of SST and climate variables but

475 causes a larger sea ice fraction positive bias over the Baltic Sea. In the future study, a new simulation

476 of ICPL266 with the CPL_var coupling method will be conducted and compared with the current

477 ICPL266 and ICPL266_flx experiments to investigate the impact of the coupling methods on the sea

478 ice simulation.

479 The salinity bias is attributed to the dry runoff bias of ICPL266 compared to the climatology, with

480 the largest bias values are found near the Ems and Newa estuaries. The dry runoff bias near the Ems

481 and Newa River mouths is due to a mismatch of the river mouth locations between the climatology

482 and ICPL266. An adjustment of the Newa River mouth location must be made to allow the Newa

483 River runoff to flow into the Gulf of Finland. The effect of the river runoff bias on salinity in the

484 deeper layers of the ocean should be analyzed in the future.

485 In addition, added value of the coupled model compared to the stand-alone model is usually

486 found in the case of extreme events (Ho-Hagemann et al., 2015, 2017, 2020; Wiese et al., 2019,

487 2020). Therefore, we will analyze the model simulations with a focus on extreme events in the next

488 study.

489 Our present study shows that the RESM GCOAST-AHOI can be a useful tool for conducting long-

490 term regional climate simulations. The new OASIS3-MCT coupling interface OMCI implemented in

491 the ICON-CLM model adds a possibility to couple ICON-CLM with an external ocean model and an

492 external hydrological discharge model, not only with NEMO and HD, using OASIS3-MCT instead of

493 YAC. Given that the standalone model components for each the atmosphere and the ocean are

494 available for a specific geographical domain, it is also quite easy to apply GCOAST-AHOI to other

495 regions. Besides preparing the lateral boundary conditions for NEMO over the new domain, and the

496 OASIS input files (as described in Supplementary S6 and S7), it is necessary to prepare several new

497 parameter files so that OASIS3-MCT can exchange the discharge from HD to NEMO without

498 interpolation. On the one hand, these are files for the general setup of the HD model. The creation

499 of these files is described in Sect. 3 of the HD model readme mark down file included in the HD

500 model package (Hagemann et al. 2023). On the other hand, this includes the HD model coupling file,

501 which is used for coupling via OASIS. Instructions for its generation are provided in Section 2.1 of a

502 markdown file dedicated to the HD model coupling exercises (Hagemann et al. 2023).

503   ICON-CLM with OMCI is also used to couple ICON-CLM with NEMO v4.2 over the GCOAST domain

504 (manuscript in preparation) and with NEMO-MED v3.6 over the Mediterranean Sea region in the

505 CLM-Community. OMCI for the older ICON version 2.6.4 can be found in Ho-Hagemann (2022).

506   Recently, the ICON Consortium has developed and released the Community Interface (ComIn)

507 for the ICON model to allow ICON to be coupled with external model components. The main

508 challenge for the external model component coupling is the initial splitting of MPI_COMM_WORLD,

509 which is done in ICON by a grouping of the mpi communicators (MPI-handshake) (M. Hanke, pers.

510 communication, 2024). There are about 40 ComIn entry points in the new release version of ICON.

511 Using the ComIn entry points does not require any additional patching of the ICON source code. A

512 coupling interface to an external model such as OMCI must be moved into a ComIn-plugin to

513 connect to the entry points in the ICON source code. In addition, the communicator splitting using

514 the MPI-handshake algorithm must be implemented in the NEMO and HD source code.

515   Currently, also a limited area mode of the ocean model (ICON-O-LAM) is being developed within

516 the ICON consortium. This can be coupled with ICON-CLM via the YAC coupler in the ICON-Seamless

517 framework. When that RESM will be available in the future and will be applied for the EURO-CORDEX

518 domain, its simulation can be compared with the simulations of GCOAST-AHOI as a good reference.

519 Investigating difference in simulations of the two RESMs could be helpful to understand better the

520 coupling interactions and feedback between model components of the climate system.

521

522 **Supplementary Materials**: Supplementary material is available online together with the submitted manuscript.

523 **Author Contributions:** H.T.M. H.-H. developed the OMCI in ICON-CLM and HD, modified the OMCI in NEMO, designed

524   the experiments and carried them out, analyzed the results; H.T.M. H.-H. prepared the manuscript with contributions

525   from all co-authors; V.M. contributed to analyze the simulations; S.P contributed to develop the OMCI in ICON-CLM; I.F.

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550   **Code and data availability:** ICON is available to the community under a permissive open source licence (BSD-3C). One

551   can download the newest released version at https://gitlab.dkrz.de/icon/icon-model. The source code of ICON v2.6.6

552   including the OMCI is published on Zenodo (https://doi.org/10.5281/zenodo.11057794). Note that ICON v2.6.6 was

553   released before the open source release of ICON. As this version still comprises 3rd party modules with a more restrictive

554   license, we had to change the file access from public to available upon request.

555   The NEMO source code is freely available and distributed under CeCILL license (GNU GPL compatible). To download

556   the NEMO reference version (for now revision 3.6):

557   *svn co http://forge.ipsl.jussieu.fr/nemo/svn/NEMO/releases/release-3.6/NEMOGCM*

558   The modified NEMO v3.6 source code for different coupling methods are published on Zenodo

559   (https://doi.org/10.5281/zenodo.11057794).

560   The HD source code is available at https://doi.org/10.5281/zenodo.4893099.

561   Source code of OASIS3-MCT v4.0 with small modifications in lib/psmile/src/GPTLget_memusage.c and

562   lib/mct/mct/m_AttrVectComms.F90 is published on Zenodo (https://doi.org/10.5281/zenodo.11057794).

563   Input data, run-scripts, evaluation scripts are published on Zenodo (https://doi.org/10.5281/zenodo.11057794).

564   Because of its huge volume, forcing data used for this study is available from the authors upon request.

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568

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Table 1: Model configuration

Configuration	ICON	NEMO	HD	Coupler OASIS3-MCT
Version	v2.6.6	v3.6	v5.1	v4.0
Domain	EURO-CORDEX	North Sea, Baltic Sea, North Atlantic	Europe	-
Resolution	~ 12 km	~ 3.7 km	~ 8 km	-
Grid points	231660	902 x 777	960 x 540	
Time step	100 s	90 s	3600 s	3600 s
Forcing	ERA5	ORAS5, OTIS	-	-

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Table 2: Model experiments

Experiment	Description
ICON266	Uncoupled ICON-CLM v2.6.6 forced by ERA5
ICPL266	Coupled GCOAST-AHOI forced by ERA5 and ORAS5
ICPL266_noNewa	Coupled GCOAST-AHOI forced by ERA5 and ORAS5, Newa River mouth is located on the buffer zone of NEMO grid.
ICPL266_flx	Coupled GCOAST-AHOI forced by ERA5 and ORAS5, coupling method CPL_flx
NEMO3.6	Stand-alone NEMO v3.6 forced by ERA5 and ORAS5
HDICON266	Stand-alone HD v5.1 forced by surface runoff and sub-surface runoff of ICON266

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Table 3: Seasonal discharge (m^3/s) of ICPL266 and HDICON266 summed over the North Sea, the Baltic Sea during the period of 2010-2018. Diff. (%) is the difference of ICPL266 to HDICON266.

Areas	North Sea			Baltic Sea		
Seasons	ICPL266	HDICON266	Diff. (%)	ICPL266	HDICON266	Diff. (%)
DJF	7356	7687	-3.31	9148	10864	-15.80
MAM	6896	7438	-5.42	18788	19884	-5.51
JJA	5103	5482	-3.79	9837	10990	-10.49
SON	4352	4857	-5.05	6755	8056	-16.15
ANN	5927	6366	-4.39	11132	12449	-10.58

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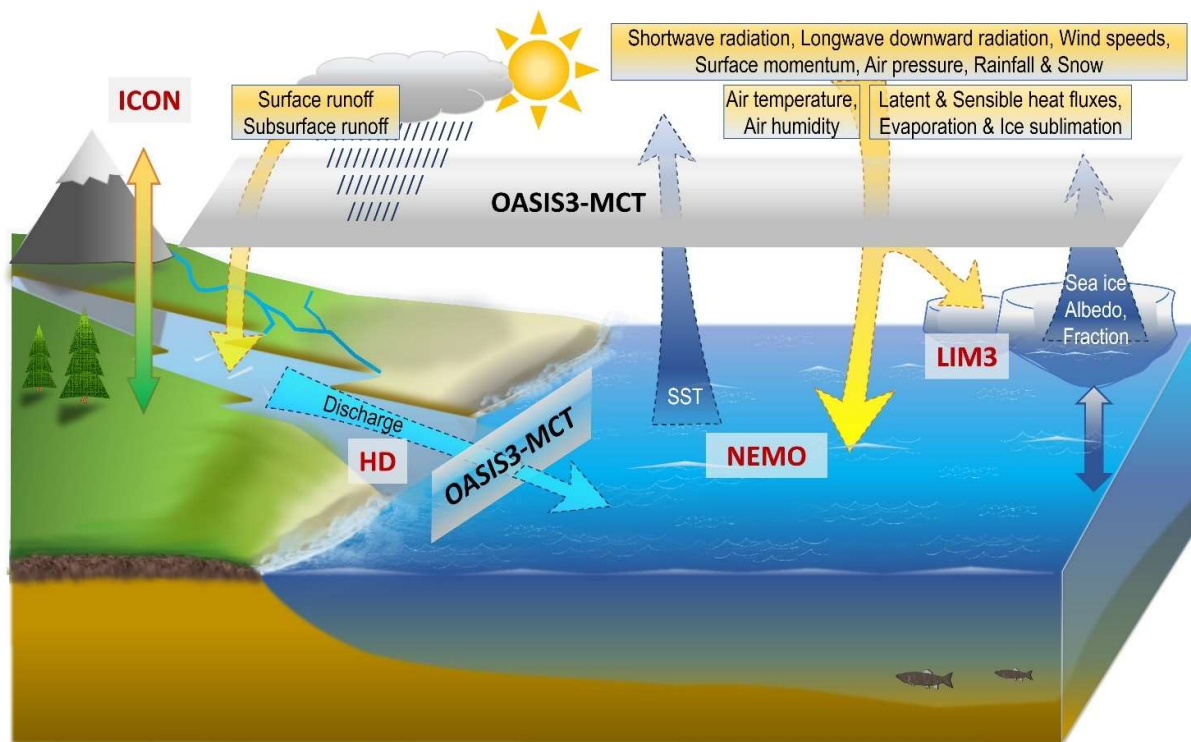
766

Figure 1: Model components of GCOAST-AHOI and variables exchanged via the OASIS3-MCT coupler. Two solid arrows display the communication between atmosphere-land (yellow-green arrow) and ocean-sea ice (gray-blue arrow), which is done via subroutines inside ICON-CLM and NEMO, respectively. Dotted arrows show the transfer between components via the OASIS interface. Yellow arrows present atmospheric transfer to ocean-sea ice and river runoff. The cyan arrow shows the discharge from the river to the ocean. Blue arrows demonstrate the transfer of sea surface temperature (SST) from the ocean as well as the sea ice albedo and sea ice fraction to the atmosphere.

767
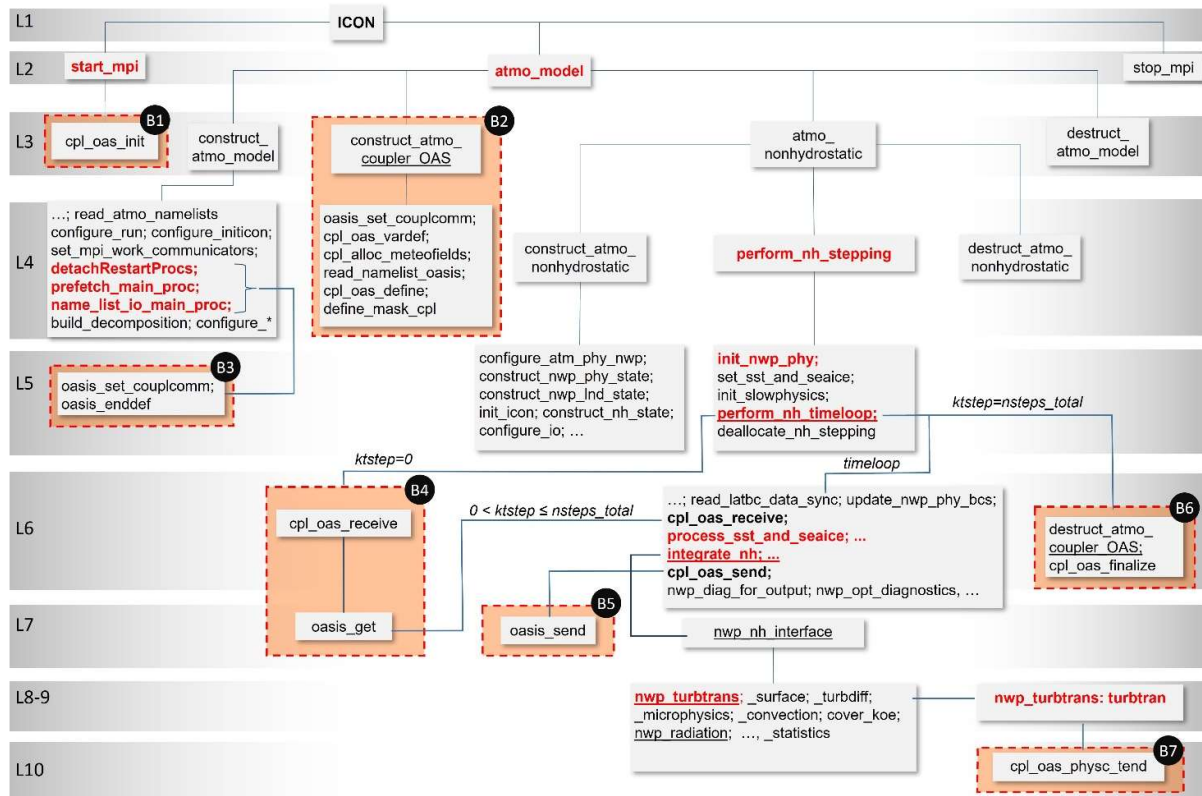
Figure 2: Flowchart of ICON-NWP/ICON-CLM with the OASIS3-MCT coupling interface OMCI. The running sequence is from top to bottom, and from left to right. "L1" indicates the Level 1 – main program ICON, etc. At the levels 2 to 6, 8 and 9, subroutines (in red text) of ICON are modified by the coupling. At the levels 3 to 7 and 10, subroutines added for OMCI are shown in orange boxes (B1-B7).

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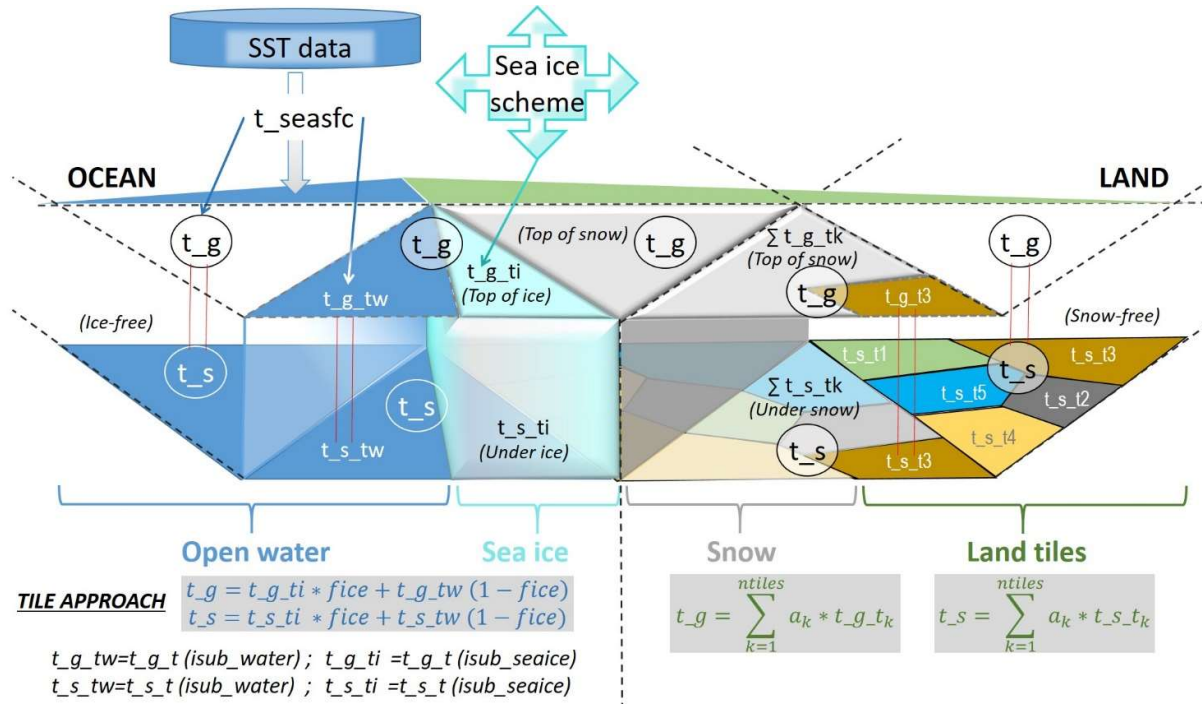
Figure 3: Surface temperature exchange between atmosphere and ocean/land in ICON and GCOAST-AHOI.

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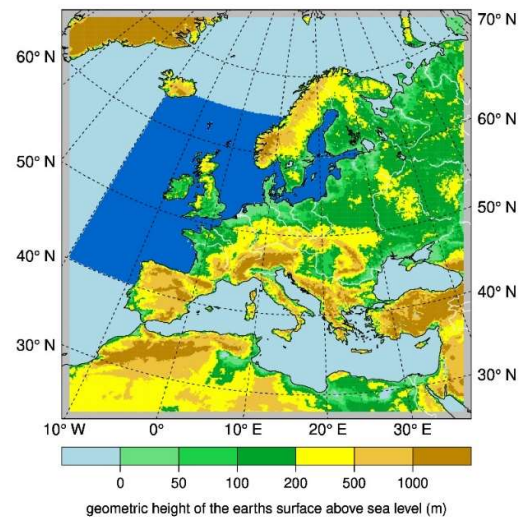
Figure 4: Integration domains of ICON and HD (EURO-CORDEX) and of NEMO-LIM3 (dark blue).

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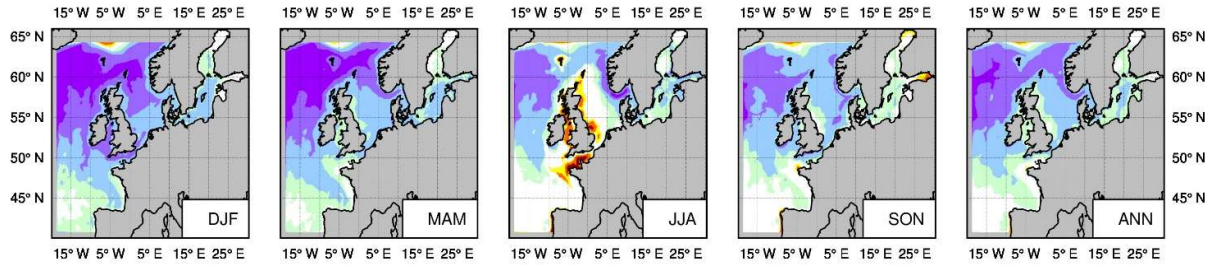
771   a) ICPL266

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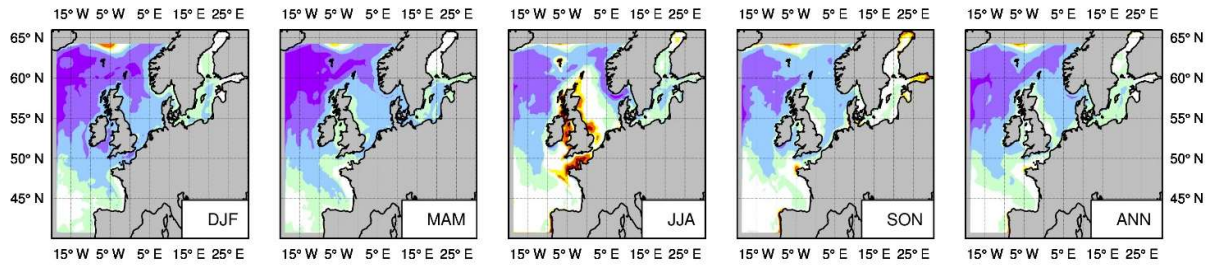
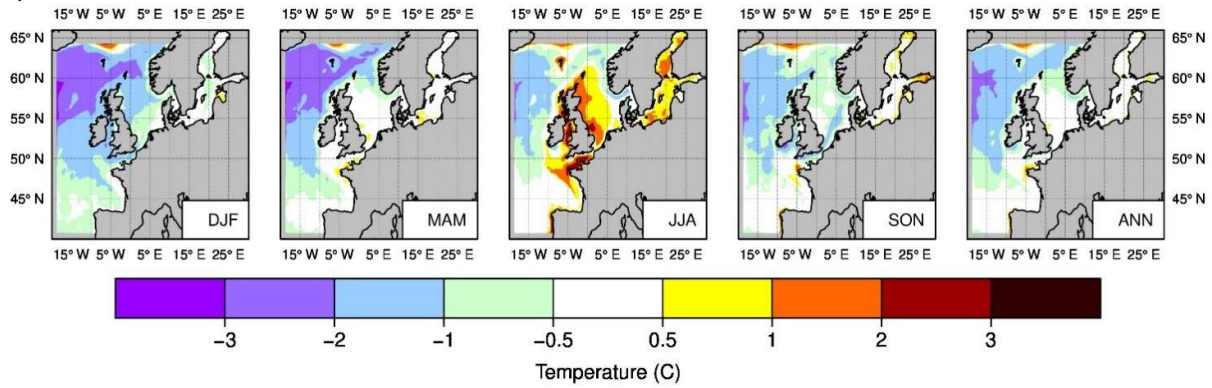
773   b) ICPL266_flx

774
775   c) NEMO3.6

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Figure 5: Seasonal (DJF, MAM, JJA, SON) and annual (ANN) mean of sea surface temperature (K) bias of a) ICPL266, b) ICPL266_flx and c) NEMO3.6 with respect to the OSTIA data for the period of 2010-2018 over the GCOAST domain.

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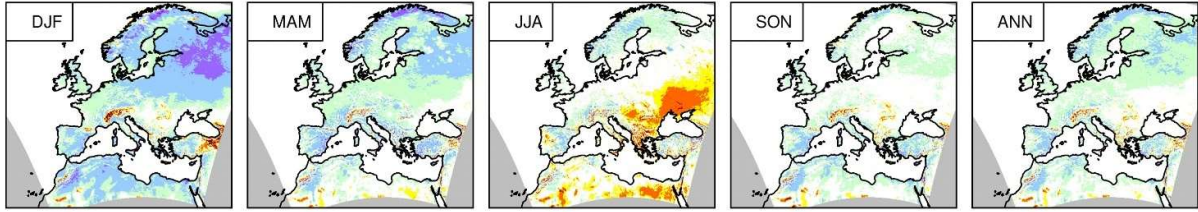
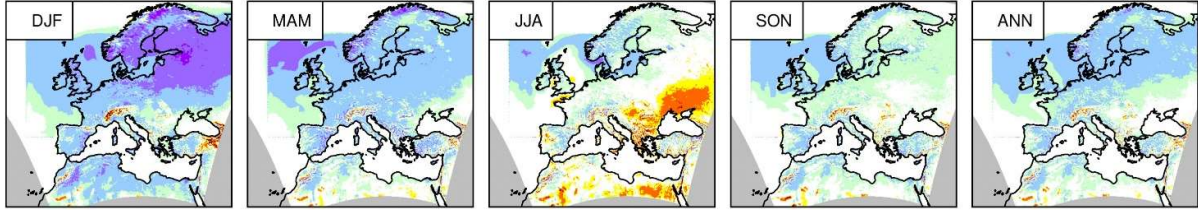
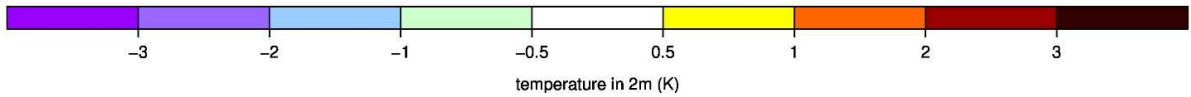
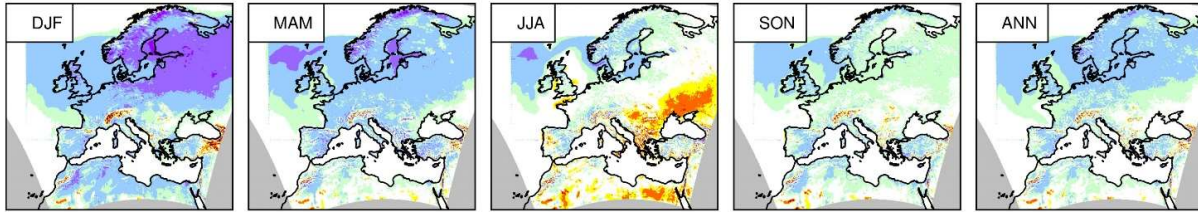
779 a) ICON266

780
781 b) ICPL266

782
783 c) ICPL266_flx

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Figure 6: Seasonal (DJF, MAM, JJA, SON) and annual (ANN) mean of 2m air temperature (K) difference of a) ICON266, b) ICPL266 and c) ICPL266_flx to the ERA5 reanalysis data for the period of 2010-2018.

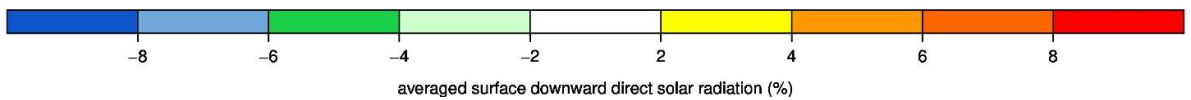
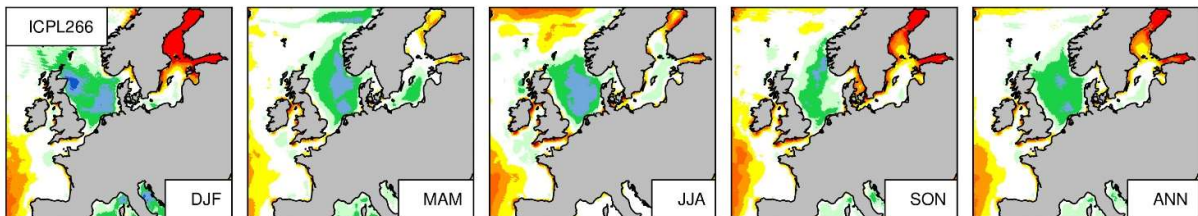
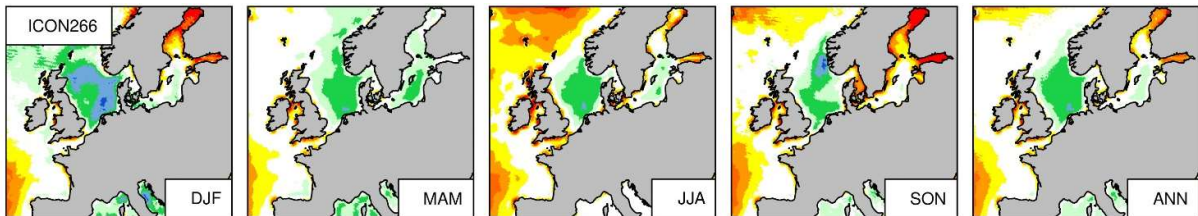
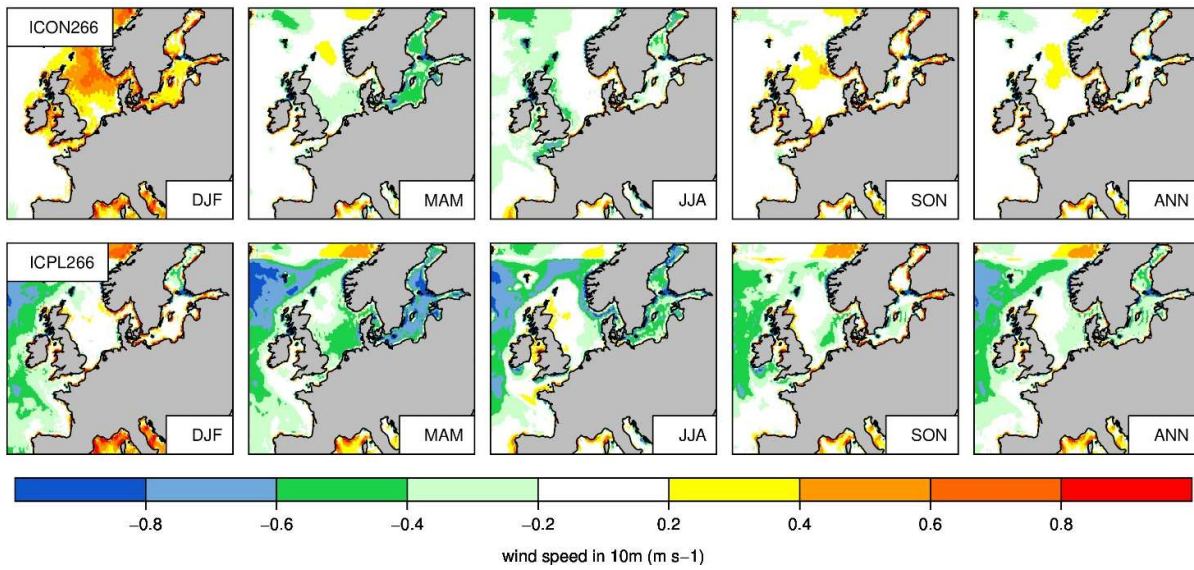



Figure 7: Seasonal (DJF, MAM, JJA, SON) and annual (ANN) mean of shortwave downward radiation bias (%) of ICON266 (top) and ICPL266 (bottom) compared to the ERA5 data for the period of 2010-2018 over the GCOAST domain.

785
786 *Figure 8: Seasonal (DJF, MAM, JJA, SON) and annual (ANN) mean 10-M wind speed bias (m s $^{-1}$) of*
787 *ICON266 (top) and ICPL266 (bottom) compared to the ERA5 data for the period of 2010-2018 over*
788 *the GCOAST domain.*
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790   a) ICON266

791
792   b) ICPL266

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794   c) ICPL266_flx

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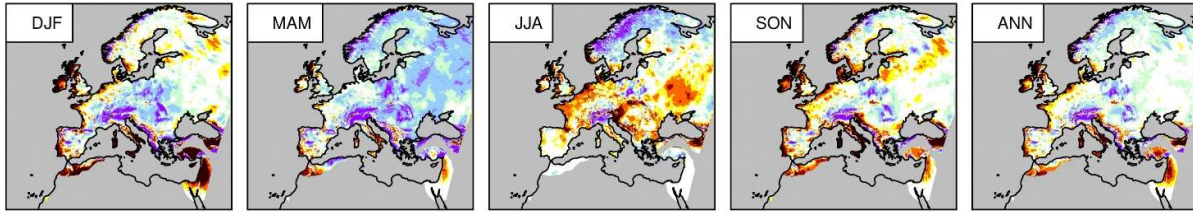
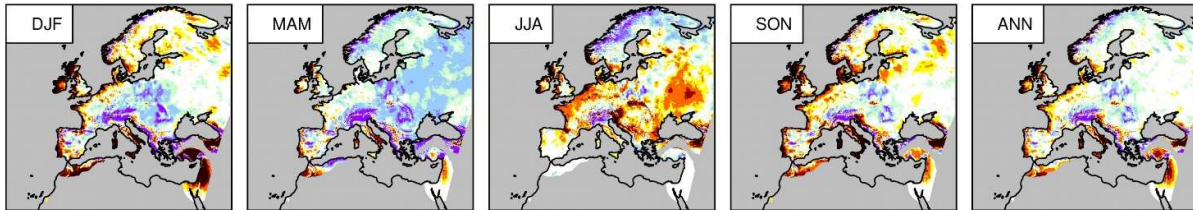
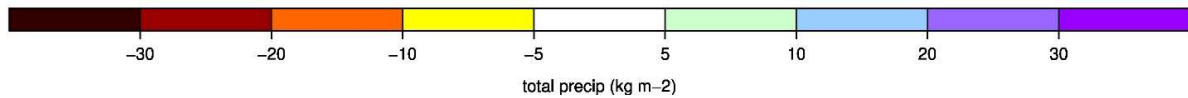
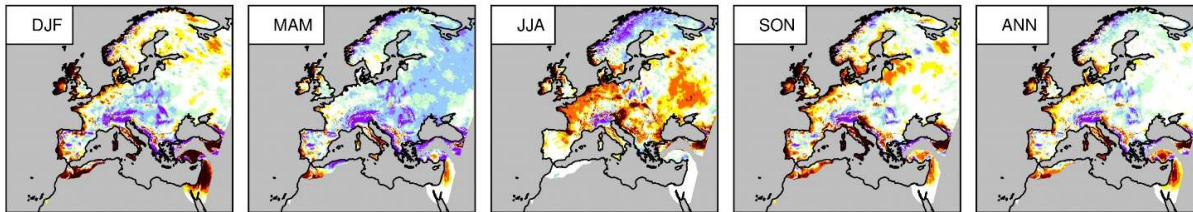
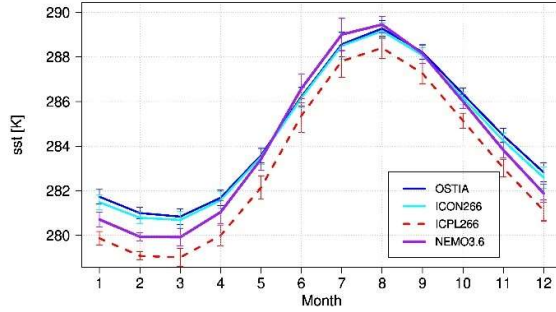


Figure 9: Seasonal (DJF, MAM, JJA, SON) and annual (ANN) difference of precipitation (mm month⁻¹) for a) ICON266, b) ICPL266 and c) ICPL266_flx compared to the E-OBS data for the period of 2010-2018.

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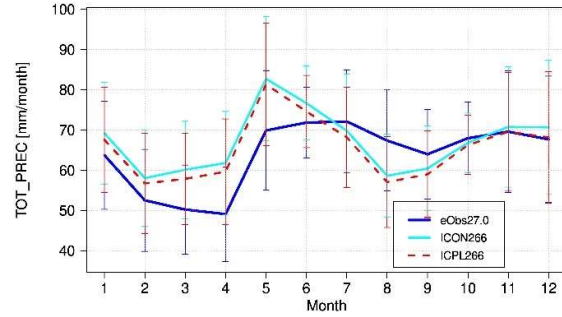
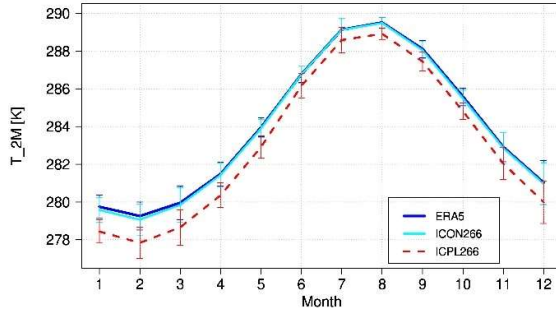
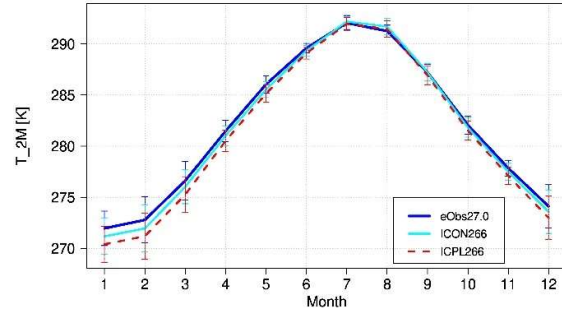
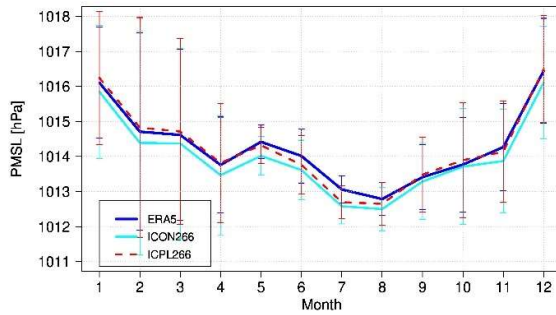
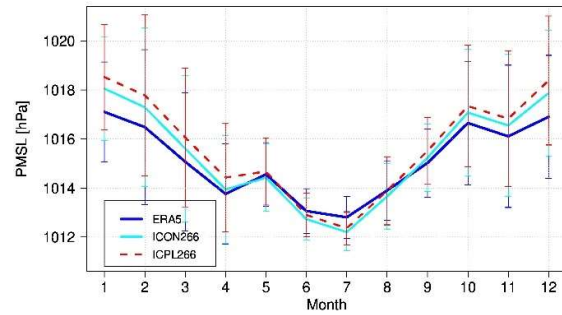
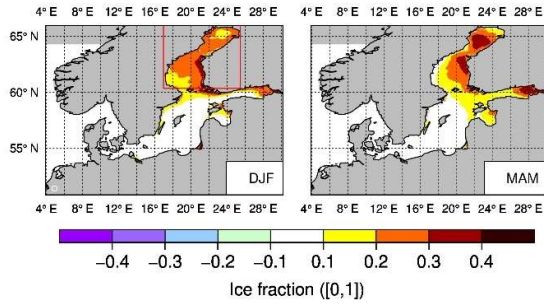
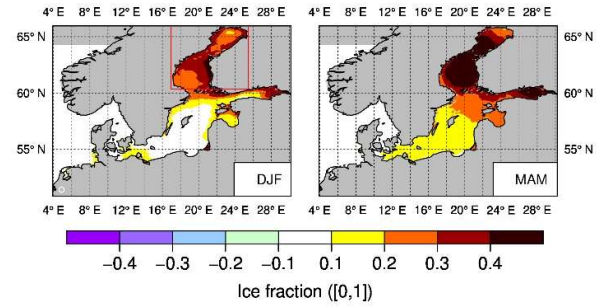
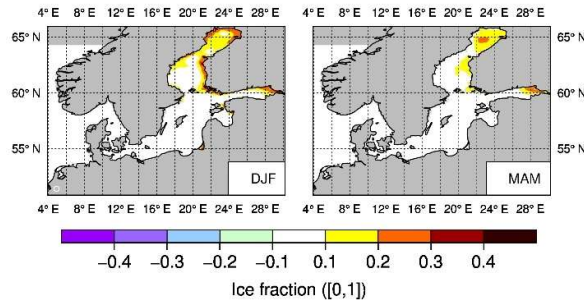


Figure 10: Climatological monthly mean of T_S (K), T_2M (K), PMSL (hPa) and precipitation (mm/month) of ICON266 (cyan solid line) and ICPL266 (red dashed line) compared to the OSTIA, ERA5 and E-OBS data (blue solid line) for the period of 2010-2018. a) includes also the SST of NEMO 3.6 (purple solid line) averaged over the GCOAST domain. Values are averaged over the GCOAST domain (avg. GCOAST) or the whole EURO-CORDEX domain (avg. all), over the ocean or land points only. The vertical bars show the standard deviation of the area mean data.

a) Sea ice fraction bias, ICPL266   b) Sea ice fraction bias, ICPL266_flx

c) Sea ice fraction bias, NEMO3.6

d) Sea ice fraction, Bothnian Bay & Sea

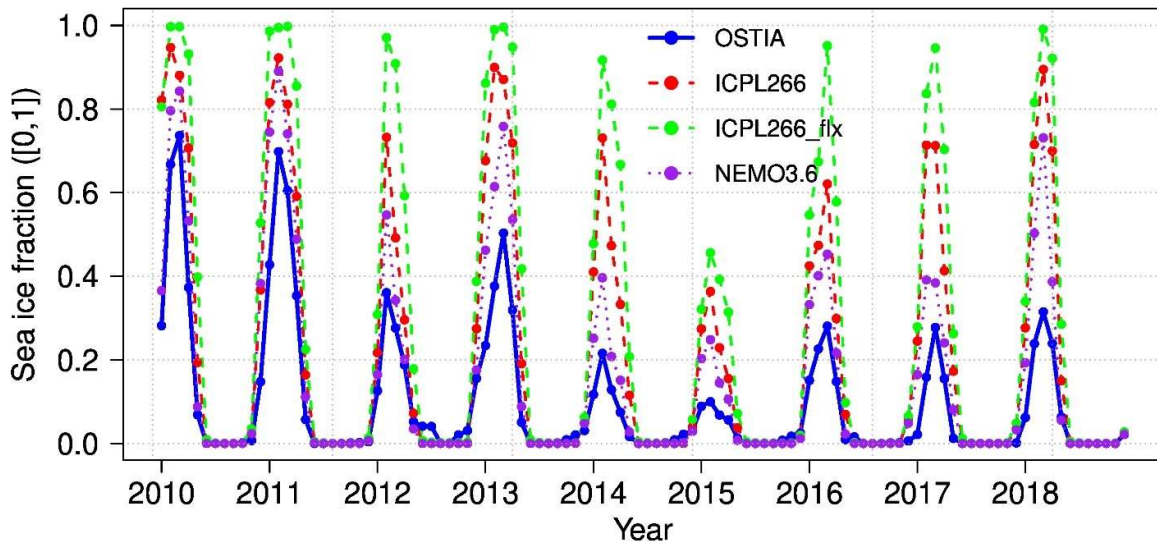



Figure 11: Sea ice fraction bias of a) ICPL266, b) ICPL266_flx, c) NEMO3.6 compared to the OSTIA data in winter (DJF) and spring (MAM) averaged over the period of 2010-2018; d) Monthly mean of sea ice fraction averaged over the Bothnian Bay & Sea (red box in Fig. 11a) of OSTIA, ICPL266, ICPL266_flx and NEMO3.6 during 2010-2018.

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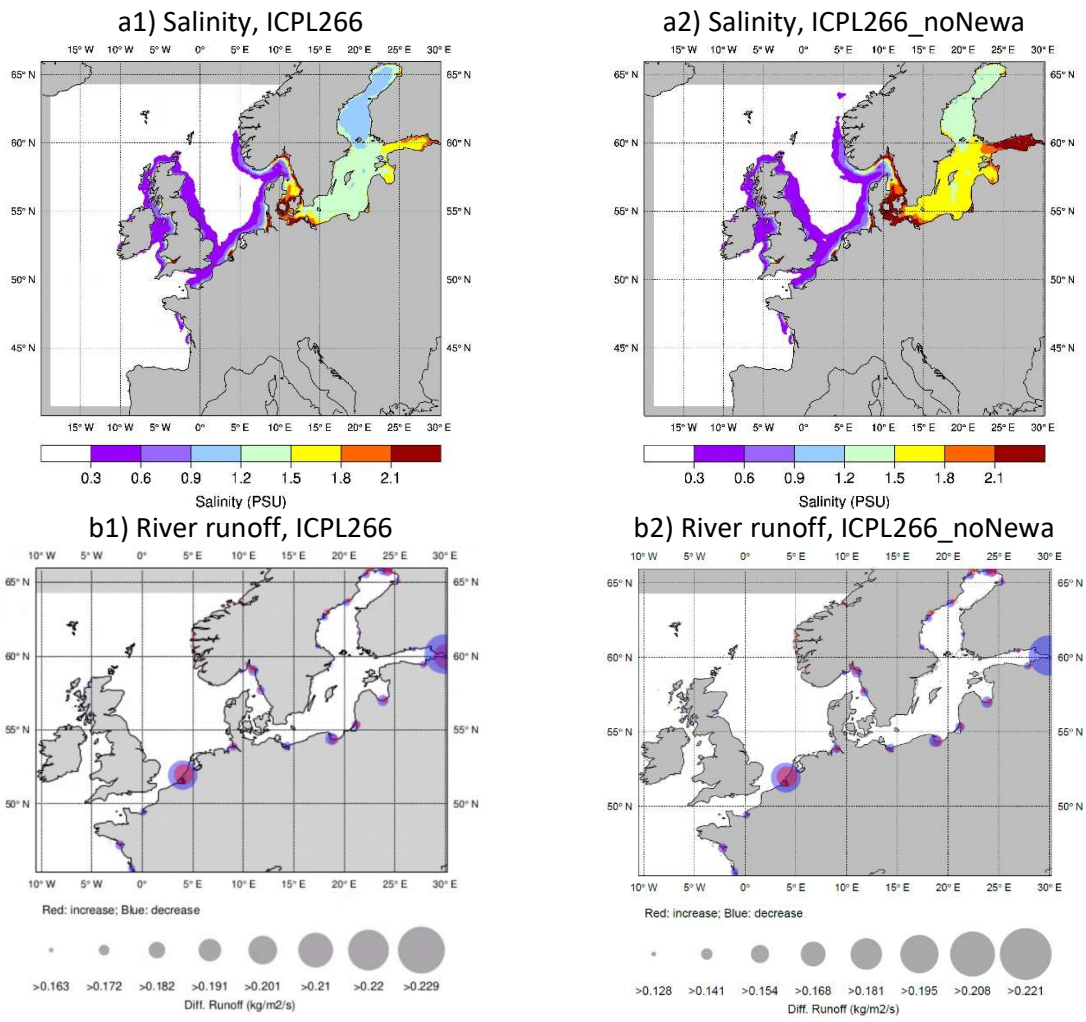
Figure 12: a1) Salinity difference (PSU) and b1) river runoff difference (kg/m2/s) between ICPL266 and NEMO3.6 averaged over the period of 2010-2018. a2) and b2) salinity and river runoff difference of ICPL266_noNewa compared to NEMO3.6. In Fig. 12 b1 and b2, the size of the grey circles indicates the magnitude of the positive (red color) and negative (blue) differences.

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