

Coupling the regional climate model ICON-CLM v2.6.6 into the Earth system model GCOAST-AHOI v2.0 using OASIS3-MCT v4.0

Reply to Reviewer#2 comments:

Thank you very much for the comments! Our answers are given below in blue.

>> General comments:

>> From line 190 to 198, the authors describe the three coupling techniques and, in the end, they chose the mixing strategy, where NEMO averages the LH and SH sent by ICON (fluxes calculated based on tile approach) and LH and SH calculated using bulk formulae. As the authors cite, this approach was also employed by Ho-Hagemann et al. (2020). In Ho-Hagemann et al. (2020), about the coupling of CCLM with NEMO, it was said that the default coupling strategy based on flux exchange from the atmosphere lead to large biases on SST, and that's why they preferred to use the average of fluxes from both models based on different turbulence parameterizations. Even if could be true that the analysis of the best coupling strategy could go beyond the scope of this paper and could be matter of further studies, it could be useful to see an evaluation analysis of the turbulent fluxes component at surface against observations or reanalysis.

We add four new figures S6, S7, S8, S9 about the surface fluxes in the Supplementary and an analysis in section 5.2 (L.279-310, text is highlighted in red in the revised manuscript).

L.279-310:

“.. Figures S6a, b and Figures S7 a, b show the seasonal SWDN and LWDN of ICPL266 and NEMO3.6 averaged over the North Sea and Baltic Sea for the period of 2010-2018. Note that we don't show the result of ICPL266_flux in Figs. S6 and S7 because there is no output of LWDN in ICPL266_flux due to the setup of the CPL_flux coupling method. Over the North Sea, the SWDN of ICPL266 is smaller than the ERA5 used for NEMO3.6 in spring and summer (Fig. S6a), which mainly leads to the cold SST bias of ICPL266 (Fig. 5a). Therefore, we plan to increase the SWDN of ICON by about 10 % before sending it to NEMO. However, the cold SST bias over the Baltic Sea does not seem to be directly related to the SWDN as there is no clear SWDN difference between ICPL266 and NEMO3.6 in summer or in any other season (Fig. S6a). The LWDN of ICPL266 is similar to NEMO3.6 in summer but slightly smaller in the other three seasons, both over the North Sea and the Baltic Sea. Increasing the LWDN in ICON-CLM by about 5-10 W/m² before sending it to NEMO should be tested to reduce the SST bias. Note that the seasonal cycle of LWDN is more pronounced over the North Sea than over the Baltic Sea.

We also compare the turbulent heat flux (i.e. the sum of SH and LH) of NEMO3.6 and the flux of NEMO in ICPL266, averaged over the North Sea and the Baltic Sea (Fig. S6c and Fig. S7c), and the net downward heat flux, which is the sum of SWDN, LWDN, SH and LH (Fig. S6d and Fig. S7d). We only consider the turbulent flux because NEMO doesn't write out SH and LH separately, but only the SWDN, LWDN and the net downward heat flux. Note that the turbulent flux from NEMO3.6 is calculated using the “CORE bulk formulae”, and due to the CPLmix coupling method used, the turbulent flux in NEMO from ICPL266 is the average of the flux from ICON-CLM and the one calculated inside NEMO (see section 3.2). The results are similar for both seas. The turbulent flux

and the net downward flux of the two experiments are quite similar with the largest differences in winter (DJF) and summer (JJA).

Using ERA5 as a reference, the SH and LH biases of ICON266 and ICPL266 are shown in Fig. S8 and Fig. S9. Over land, the bias of ICON266 is very similar to that of ICPL266. However, over the ocean, the bias of ICPL266 is generally more positive (i.e. the fluxes are less negative) than that of ICON266 with the largest bias over the North Atlantic. Smaller heat fluxes are consistent with lower SSTs in ICPL266, as lower SSTs lead to larger stability and less vertical mixing. Over the North and Baltic Seas, SH and LH of ICPL266 are quite close to ERA5. Despite of the SST forcing from ERA5, ICON266 has a negative SH bias of about -5 to -15 W/m² over the North and Baltic Seas, especially in winter. This suggests a future analysis of the difference in air temperature at the lowest level of ICON-CLM and ERA5.”

A new experiment ICPL266_flx using the coupling method CPL_flx is conducted and analyzed in section 5.

We also add some sentences about the different results of coupling COSMO-CLM and ICON-CLM to NEMO:

L237-242:

“Ho-Hagemann et al. (2020) noted that using CPL_flx when coupling COSMO-CLM with NEMO leads to larger biases in the SST than using CPL_mix. This is not the case here when coupling ICON-CLM. A possible reason for this is that due to the tile approach (cf. Sect. 3.2), the fluxes from ICON-CLM to NEMO are sent separately over water and sea ice, while COSMO-CLM v5.0 doesn’t have the tile approach, therefore, the fluxes in each ocean grid box sent from the atmosphere to the ocean are the mixed fluxes of water and sea ice.”

>> The authors should also include an evaluation analysis for the hydrological discharge component, this is extremely important to close the balance of the freshwater. If the results are unsatisfactory, they should be highlighted as current model limitation and, maybe, as further model improvement in future works.

We add a new table (Table 3) of discharge and a new figure 12 of salinity and river runoff together with a new sub-section 5.5 (L.378-449, text is highlighted in red in the revised manuscript) to analyse the simulated salinity and river runoff.

L.378-449:

“5.5 Salinity and river runoff

As mentioned in section 4, NEMO3.6 uses a climatological dataset for river runoff. Therefore, a rough verification of the river runoff produced by the HD model in ICPL266 can be made by comparing against this climatological river runoff. Differences in sea surface salinity and river runoff between ICPL266 and NEMO3.6 are shown in Figure 12. The salinity simulated by ICPL266 is about 0.3-1 PSU higher than that of NEMO3.6 along the British Isles and the North Sea coast, and about 0.9-1.8 PSU higher in the Baltic Sea.

Two areas with the largest salinity differences of more than 2 PSU are found south of the Kattegat and in the Gulf of Finland (Fig. 12a). The river runoff differences (Fig. 12b) are largest near the Ems and Newa estuaries with more than 0.1 and 0.2 kg/m²/s, respectively. The small river runoff difference between the two models at Kattegat cannot be used to directly explain the increase in salinity there. The river runoff differences near the Ems estuary have opposite signs (blue point overlaid by a red one in Fig. 12b), but very similar values. The reason for this may be the discrepancy in the locations of river mouths between the NEMO3.6 setup, where the river runoff is taken from a climatology, and those in ICPL266, which are defined based on the river mouths in the HD model and the NEMO land-sea mask. In the latter case, the river mouths of the HD model are interpolated onto the NEMO grid by searching for the closest ocean point of NEMO. For example, the Ems River mouth in ICPL266 may not be in the same position as in the climatology data. This discrepancy would lead to a difference in salinity near the coast (see Fig. 12a). The extent of the effect on salinity in the deeper layers of the ocean in a longer-term simulation needs to be analyzed in the future.

The large difference in river runoff near the Newa estuary is also caused by a mismatch in the locations of the river mouths. In this case, the mouth of the Newa River in the climatology data (60.1333°N, 29.888°E) is located slightly northwest of its “real” location (Wikipedia: 59.9453°N, 30.1708°E). The interpolation program used to define the mouths of the HD rivers on the NEMO grid, by searching for the closest ocean point to the mouth of the HD rivers, found the mouth of the Newa River at (59.95835°N, 30.20825°E), which is very close to the “real” position and at the furthest grid point to the east in the Gulf of Finland. However, in the NEMO model, this eastern boundary point in the Gulf of Finland is masked as a buffer zone. Therefore, the discharge from HD to NEMO at this point in ICPL266_noNewa was ignored in the NEMO calculations, resulting in a lack of freshwater inflow to the Gulf of Finland in ICPL266 (Fig. 12b2) and consequently an increase in salinity (Fig. 12a2). The ICPL266 simulation with the Newa River mouth located on the NEMO buffer zone is referred to ICPL266_noNewa in Figure 12. To overcome the location deficiency, that the Newa River mouth was shifted one grid point to the west on the NEMO grid to allow the large amount of river runoff to enter the Gulf of Finland in the coupled model. Therefore, the salinity difference of ICPL266 compared to NEMO3.6 is reduced (Fig. 12a1) and the river runoff difference shows the shift of the river mouth instead of the missing one (Fig. 12b1). The shift of the Newa River mouth has little effect on the simulated atmospheric variables, but improves the simulated salinity in the Baltic Sea, which is important for ecosystem modelling when a marine biogeochemical or ecosystem model such as ECOSMO (Daewel and Schrum, 2013) is coupled with GCOAST-AHOI in the future.

Other river mouths in the Baltic Sea have river runoff differences of less than 1.4 kg/m²/s when comparing ICPL266 to NEMO3.6. In general, ICPL266 tends to simulate less river runoff than the climatology, leading to increased salinity there. The sources of the river runoff used for NEMO in ICPL266 are the surface and sub-surface runoffs from the land component in ICON-CLM that are transported to the ocean by HD. We applied the HD model to calculate the discharge using the ICPL266 and ICON266 surface and subsurface runoff (Table 3) to evaluate it against the discharge observation. The annual discharge difference of ICPL266 and HDICON266 in the Baltic Sea is about -11 %. However, HDICON266 with a discharge of 12449 m³/s is about -20 % biased towards the

HELCOM (Helsinki Commission; Svendsen and Gustafsson, 2022) value of 15676 m³/s. Note that for Baltic Sea ocean models, the mean long-term bias of river runoff must be less than 7 % (Hagemann and Stacke, 2022). In the North Sea, ICPL266 discharge is about -4 % compared to HDICON266, which has an annual value of 6366 m³/s. However, both models have a dry discharge bias compared to the OSPAR data (Farkas and Skarbøvik, 2021), which is 9190 m³/s.

The main driver of the runoff is precipitation. Figure 9 shows that over Scandinavia, ICON266 has a wet bias of about 10-30 mm/month in spring and summer compared to the E-OBS data. Thus, even with the wet precipitation bias, ICON266 has a dry discharge bias. At the same time, ICPL266 precipitation is lower than for ICON266 in spring and summer, and therefore closer to the E-OBS data (Fig. 9b). This difference in precipitation between ICPL266 and ICON266 explains the -11 % difference in discharge, which increases the dry discharge bias. The reduction of the precipitation bias in ICPL266 leading to a larger discharge dry bias implies that a better simulation of precipitation compared to observations does not necessarily lead to a better runoff. We note that the runoff from the atmosphere-only ICON-CLM has a general dry bias, which can be attributed to the respective parametrizations in the TERRA land surface scheme used in ICON-CLM (S. Hagemann, pers. communication, 2024). This dry discharge bias can be improved either by using the JSBACH land-surface model in ICON-Seamless, or by applying a discharge bias correction developed by Hagemann et al. (2024).

In Section 5.4, it was speculated that an increase in river runoff would lead to less saline sea water and therefore more sea ice over the Baltic Sea. In our study, ICPL266 simulates too little river runoff, leading to increased salinity in the Baltic Sea, which would mean less sea ice. However, the sea ice fraction is increased compared to the ERA5-forced NEMO3.6. Thus, the main factor causing the bias in the sea ice fraction seems to be the cold bias in the air temperature over sea ice in the Baltic Sea in winter and spring (Fig. 6b and c).”

>> Minor comments:

>> I found the interface structure part (lines 104-180) and the profiling part of the model with the LUCIA tool (lines 246-279) too technical and verbose. These parts could be summarized in the main core of the paper, while all the technicalities, which are extremely useful for the developers and users of both community models and for the reproducibility of the study, could go in the supplement material.

Thank you for this suggestion. We agree with you and move those parts to the Supplementary. The Table 2 is moved to be the Table S3 in the Supplementary. We include a new Table 2 in the main text to list the model experiments.

>> Some of the figure results regarding the NEMO standalone biases are just described and bear the wording “not shown”, maybe also these figures could go in the supplement material? Even if these figures could not give any added value for the results, they could help to visualize the NEMO standalone biases.

We add some figures of stand-alone NEMO in Fig. 5c, Fig. 11, Fig. 12, Fig. S6 and Fig. S7 and the analysis in the section 5.

>> From line 46 to 51, the authors describe why they chose to implement OASIS in ICON and not YAC in NEMO. Could they elaborate on why the implementation of YAC in NEMO is a major challenge than the implementation of OASIS in ICON? I think that, besides the different coupler architectures and implementation issues, the main driver of this decision has been lack of a regional version of the ICON-O component, currently under development as explained in the results.

The reviewer is right. We rewrite the following paragraphs to provide the technical arguments for using OASIS3-MCT:

L.61-74:

“For the coupling of ICON-CLM as an atmosphere component into GCOAST-AHOI, which includes HD and the ocean model NEMO (Nucleus for European Modelling of the Ocean, Madec et al. 2017), representing the ocean and sea ice components, basically, there were two feasible options: either to implement a YAC interface in NEMO and HD, or to implement an OASIS interface in ICON-CLM. For the option one, the YAC coupling interface was added into the HD source code by M. Hanke (DKRZ) (see Hagemann et al. 2023), but YAC has not yet been available in the NEMO source code. To our current knowledge, there is no RESM with NEMO using the YAC coupler. The NEMO model is already linked to the OASIS coupler, which has been used to couple NEMO with many other model components. Implementing the YAC interface in NEMO would require a larger effort, as the NEMO source code is much more complicated than the HD code. In addition, although the NEMO source code is freely available, we are ordinary users in the NEMO community, not members of the model development team. Therefore, implementing and especially maintaining the YAC interfaces in NEMO is a big challenge.”

L.490-493:

“The new OASIS3-MCT coupling interface OMCI implemented in the ICON-CLM model adds a possibility to couple ICON-CLM with an external ocean model and an external hydrological discharge model, not only with NEMO and HD, using OASIS3-MCT instead of YAC.”

And modify the abstract:

“The new OASIS3-MCT coupling interface OMCI implemented in the ICON-CLM model adds a possibility to couple ICON-CLM with an external ocean model and an external hydrological discharge model using OASIS3-MCT instead of the YAC coupler. Using OMCI, it is also possible to set up a RESM with ICON-CLM and other ocean and hydrology models possessing the OASIS3-MCT interface for other regions, such as the Mediterranean Sea.”