

Coupling the regional climate model ICON-CLM v2.6.6 into the Earth system model GCOAST-AHOI v2.0 using OASIS3-MCT v4.0

Reply to Reviewer#1 comments:

Thank you very much for the comments! Our answers are given below in blue.

>> General comments:

>> 1) The analysis is too simple because you show only SST from the ocean and no results from the land model. As the freshwater input from land is very important for coastal seas like the Baltic Sea and North Sea, you should evaluate the results of the simulated river discharge from the HD model into the coastal seas. I understood, that the river discharge in the presented coupled simulation is from the HD model and not from observations. If you do not reproduce the integrated freshwater input (precipitation minus evaporation over land and over the ocean), the added value of the regional coupled model for the coastal ocean is rather limited. If the results are not sufficiently good, you may refer to further model development. However, please show these results because they would be interesting for the model performance in the ocean, e.g. for salinity.

We add a new table (Table 3) of discharge and a new figure 12 of salinity and river runoff together with a new sub-section 5.5 (L.378-449, text is highlighted in red in the revised manuscript) to analyse the simulated salinity and river runoff.

L.378-449:

“5.5 Salinity and river runoff

As mentioned in section 4, NEMO3.6 uses a climatological dataset for river runoff. Therefore, a rough verification of the river runoff produced by the HD model in ICPL266 can be made by comparing against this climatological river runoff. Differences in sea surface salinity and river runoff between ICPL266 and NEMO3.6 are shown in Figure 12. The salinity simulated by ICPL266 is about 0.3-1 PSU higher than that of NEMO3.6 along the British Isles and the North Sea coast, and about 0.9-1.8 PSU higher in the Baltic Sea. Two areas with the largest salinity differences of more than 2 PSU are found south of the Kattegat and in the Gulf of Finland (Fig. 12a). The river runoff differences (Fig. 12b) are largest near the Ems and Newa estuaries with more than 0.1 and 0.2 kg/m²/s, respectively. The small river runoff difference between the two models at Kattegat cannot be used to directly explain the increase in salinity there. The river runoff differences near the Ems estuary have opposite signs (blue point overlaid by a red one in Fig. 12b), but very similar values. The reason for this may be the discrepancy in the locations of river mouths between the NEMO3.6 setup, where the river runoff is taken from a climatology, and those in ICPL266, which are defined based on the river mouths in the HD model and the NEMO land-sea mask. In the latter case, the river mouths of the HD model are interpolated onto the NEMO grid by searching for the closest ocean point of NEMO. For example, the Ems River mouth in ICPL266 may not be in the same position as in the climatology data. This discrepancy would lead to a difference in salinity near the coast (see Fig. 12a). The extent of the effect on salinity in the deeper layers of the ocean in a longer-term simulation needs to be analyzed in the future.

The large difference in river runoff near the Newa estuary is also caused by a mismatch in the locations of the river mouths. In this case, the mouth of the Newa River in

the climatology data (60.1333°N, 29.888°E) is located slightly northwest of its “real” location (Wikipedia: 59.9453°N, 30.1708°E). The interpolation program used to define the mouths of the HD rivers on the NEMO grid, by searching for the closest ocean point to the mouth of the HD rivers, found the mouth of the Newa River at (59.95835°N, 30.20825°E), which is very close to the “real” position and at the furthest grid point to the east in the Gulf of Finland. However, in the NEMO model, this eastern boundary point in the Gulf of Finland is masked as a buffer zone. Therefore, the discharge from HD to NEMO at this point in ICPL266_noNewa was ignored in the NEMO calculations, resulting in a lack of freshwater inflow to the Gulf of Finland in ICPL266 (Fig. 12b2) and consequently an increase in salinity (Fig. 12a2). The ICPL266 simulation with the Newa River mouth located on the NEMO buffer zone is referred to ICPL266_noNewa in Figure 12. To overcome the location deficiency, that the Newa River mouth was shifted one grid point to the west on the NEMO grid to allow the large amount of river runoff to enter the Gulf of Finland in the coupled model. Therefore, the salinity difference of ICPL266 compared to NEMO3.6 is reduced (Fig. 12a1) and the river runoff difference shows the shift of the river mouth instead of the missing one (Fig. 12b1). The shift of the Newa River mouth has little effect on the simulated atmospheric variables, but improves the simulated salinity in the Baltic Sea, which is important for ecosystem modelling when a marine biogeochemical or ecosystem model such as ECOSMO (Daewel and Schrum, 2013) is coupled with GCOAST-AHOI in the future.

Other river mouths in the Baltic Sea have river runoff differences of less than 1.4 kg/m²/s when comparing ICPL266 to NEMO3.6. In general, ICPL266 tends to simulate less river runoff than the climatology, leading to increased salinity there. The sources of the river runoff used for NEMO in ICPL266 are the surface and sub-surface runoffs from the land component in ICON-CLM that are transported to the ocean by HD. We applied the HD model to calculate the discharge using the ICPL266 and ICON266 surface and subsurface runoff (Table 3) to evaluate it against the discharge observation. The annual discharge difference of ICPL266 and HDICON266 in the Baltic Sea is about -11 %. However, HDICON266 with a discharge of 12449 m³/s is about -20 % biased towards the HELCOM (Helsinki Commission; Svendsen and Gustafsson, 2022) value of 15676 m³/s. Note that for Baltic Sea ocean models, the mean long-term bias of river runoff must be less than 7 % (Hagemann and Stacke, 2022). In the North Sea, ICPL266 discharge is about -4 % compared to HDICON266, which has an annual value of 6366 m³/s. However, both models have a dry discharge bias compared to the OSPAR data (Farkas and Skarbøvik, 2021), which is 9190 m³/s.

The main driver of the runoff is precipitation. Figure 9 shows that over Scandinavia, ICON266 has a wet bias of about 10-30 mm/month in spring and summer compared to the E-OBS data. Thus, even with the wet precipitation bias, ICON266 has a dry discharge bias. At the same time, ICPL266 precipitation is lower than for ICON266 in spring and summer, and therefore closer to the E-OBS data (Fig. 9b). This difference in precipitation between ICPL266 and ICON266 explains the -11 % difference in discharge, which increases the dry discharge bias. The reduction of the precipitation bias in ICPL266 leading to a larger discharge dry bias implies that a better simulation of precipitation compared to observations does not necessarily lead to a better runoff. We note that the runoff from the atmosphere-only ICON-CLM has a general dry bias, which can be attributed to the respective parametrizations in the TERRA land surface scheme used in ICON-CLM (S.

Hagemann, pers. communication, 2024). This dry discharge bias can be improved either by using the JSBACH land-surface model in ICON-Seamless, or by applying a discharge bias correction developed by Hagemann et al. (2024).

In Section 5.4, it was speculated that an increase in river runoff would lead to less saline sea water and therefore more sea ice over the Baltic Sea. In our study, ICPL266 simulates too little river runoff, leading to increased salinity in the Baltic Sea, which would mean less sea ice. However, the sea ice fraction is increased compared to the ERA5-forced NEMO3.6. Thus, the main factor causing the bias in the sea ice fraction seems to be the cold bias in the air temperature over sea ice in the Baltic Sea in winter and spring (Fig. 6b and c).”

>> 2) I would like to see also an analysis of sea ice in the Baltic Sea, because the sea ice cover is very sensitive to biases in the surface heat and momentum fluxes over the coupling domain.

We add a new figure 11 and a new sub-section 5.4 (L.355-377, text is highlighted in red in the revised manuscript) to analyse the simulated sea ice.

L.355-377:

“5.4 Sea ice

The sea ice fraction bias of ICPL266 is about 0.2-0.3 over the Bothnian Bay and Sea in winter and spring (Fig. 11a), while ICPL266_flx has a larger bias of about 0.3-0.5 (Fig. 11b) and the ERA5-forced NEMO3.6 has a relatively small positive ice fraction bias there (Fig. 11c). The monthly mean sea ice fraction averaged over the Bothnian Bay and Sea from ICPL266 and NEMO3.6 compared to the OSTIA data is shown in Fig. 11d, where the sea ice temporal variation is quite well captured by the two models, with a high peak in spring 2010 and a low peak in spring 2015. However, all three simulations overestimate the sea ice fraction of OSTIA, with the two coupled simulations showing a larger positive bias also in the time series. While ICPL266 has a winter SWDN about 8 % larger than ERA5 (Fig. 7b), the incoming shortwave radiation is relatively small over the high latitudes in winter. Therefore, we don't expect the positive SWDN bias to be the main reason for the overestimation of sea ice. However, the LWDN of ICPL266 is about 10 W/m² lower than ERA5 (the forcing for NEMO3.6) in winter over the North and Baltic Seas (Figs. S6b, S7b), and T_{2M} is about 3 °C lower than ERA5 over the Scandinavian region surrounding the Bothnian Bay and Sea (Fig. 6b). The cold T_{2M} bias and negative LWDN bias of ICPL266 may explain its positive sea ice fraction bias. The cold air temperature above the sea ice surface often produces more sea ice in winter and spring, especially over an area with water of low salinity, such as the Bothnian Bay and Sea. Figure 6c shows the larger T_{2M} cold bias of ICPL266_flx over the Baltic Sea in spring about 1°C more than ICPL266 (Fig. 6b) which is consistent with the larger sea ice fraction bias of ICPL266_flx (Fig. 11b) compared with that of ICPL266 (Fig. 11a).

Another factor that could contribute to an increase in sea ice cover in spring would be an increase in river runoff, which would result in less salty sea water and therefore more sea ice. These two variables are analyzed in the next section.”

>> 3) Furthermore, the integrated heat content changes could be evaluated by temperature profiles. Please use temperature and salinity data from the national monitoring programs for the North Sea and Baltic Sea to evaluate the ocean model performance and surface fluxes in the coupled model. We add four new figures S6, S7, S8, S9 about the surface fluxes in the Supplementary and an analysis in section 5.2 (L.279-310, text is highlighted in red in the revised manuscript).

L.279-310:

“.. Figures S6a, b and Figures S7 a, b show the seasonal SWDN and LWDN of ICPL266 and NEMO3.6 averaged over the North Sea and Baltic Sea for the period of 2010-2018. Note that we don't show the result of ICPL266_flux in Figs. S6 and S7 because there is no output of LWDN in ICPL266_flux due to the setup of the CPL_flux coupling method. Over the North Sea, the SWDN of ICPL266 is smaller than the ERA5 used for NEMO3.6 in spring and summer (Fig. S6a), which mainly leads to the cold SST bias of ICPL266 (Fig. 5a). Therefore, we plan to increase the SWDN of ICON by about 10 % before sending it to NEMO. However, the cold SST bias over the Baltic Sea does not seem to be directly related to the SWDN as there is no clear SWDN difference between ICPL266 and NEMO3.6 in summer or in any other season (Fig. S6a). The LWDN of ICPL266 is similar to NEMO3.6 in summer but slightly smaller in the other three seasons, both over the North Sea and the Baltic Sea. Increasing the LWDN in ICON-CLM by about 5-10 W/m² before sending it to NEMO should be tested to reduce the SST bias. Note that the seasonal cycle of LWDN is more pronounced over the North Sea than over the Baltic Sea.

We also compare the turbulent heat flux (i.e. the sum of SH and LH) of NEMO3.6 and the flux of NEMO in ICPL266, averaged over the North Sea and the Baltic Sea (Fig. S6c and Fig. S7c), and the net downward heat flux, which is the sum of SWDN, LWDN, SH and LH (Fig. S6d and Fig. S7d). We only consider the turbulent flux because NEMO doesn't write out SH and LH separately, but only the SWDN, LWDN and the net downward heat flux. Note that the turbulent flux from NEMO3.6 is calculated using the “CORE bulk formulae”, and due to the CPLmix coupling method used, the turbulent flux in NEMO from ICPL266 is the average of the flux from ICON-CLM and the one calculated inside NEMO (see section 3.2). The results are similar for both seas. The turbulent flux and the net downward flux of the two experiments are quite similar with the largest differences in winter (DJF) and summer (JJA).

Using ERA5 as a reference, the SH and LH biases of ICON266 and ICPL266 are shown in Fig. S8 and Fig. S9. Over land, the bias of ICON266 is very similar to that of ICPL266. However, over the ocean, the bias of ICPL266 is generally more positive (i.e. the fluxes are less negative) than that of ICON266 with the largest bias over the North Atlantic. Smaller heat fluxes are consistent with lower SSTs in ICPL266, as lower SSTs lead to larger stability and less vertical mixing. Over the North and Baltic Seas, SH and LH of ICPL266 are quite close to ERA5. Despite of the SST forcing from ERA5, ICON266 has a negative SH bias of about -5 to -15 W/m² over the North and Baltic Seas, especially in winter. This suggests a future analysis of the difference in air temperature at the lowest level of ICON-CLM and ERA5.”

The temperature profile analysis will be done in a future study when a new simulation with the optimal settings of ICON-CLM will be conducted for a longer time period.

>> 4) It would be good to present a list of scientific and technical arguments for your modeling strategy using OASIS3-MCT instead of YAC. For me the advantages of OASIS3-MCT compared to YAC are not completely clear.

There is no scientific argument for preferring one external coupler over the other in ICON-CLM. Both couplers, OASIS3-MCT and YAC, provide the tools for coupling climate component models of the Earth system.

However, there are several technical and practical arguments for using OASIS3-MCT. Hence, we rewrite the paragraphs:

L.61-74:

“For the coupling of ICON-CLM as an atmosphere component into GCOAST-AHOI, which includes HD and the ocean model NEMO (Nucleus for European Modelling of the Ocean, Madec et al. 2017), representing the ocean and sea ice components, basically, there were two feasible options: either to implement a YAC interface in NEMO and HD, or to implement an OASIS interface in ICON-CLM. For the option one, the YAC coupling interface was added into the HD source code by M. Hanke (DKRZ) (see Hagemann et al. 2023), but YAC has not yet been available in the NEMO source code. To our current knowledge, there is no RESM with NEMO using the YAC coupler. The NEMO model is already linked to the OASIS coupler, which has been used to couple NEMO with many other model components. Implementing the YAC interface in NEMO would require a larger effort, as the NEMO source code is much more complicated than the HD code. In addition, although the NEMO source code is freely available, we are ordinary users in the NEMO community, not members of the model development team. Therefore, implementing and especially maintaining the YAC interfaces in NEMO is a big challenge.”

L.490-493:

“The new OASIS3-MCT coupling interface OMCI implemented in the ICON-CLM model adds a possibility to couple ICON-CLM with an external ocean model and an external hydrological discharge model, not only with NEMO and HD, using OASIS3-MCT instead of YAC.”

And modify the abstract:

“The new OASIS3-MCT coupling interface OMCI implemented in the ICON-CLM model adds a possibility to couple ICON-CLM with an external ocean model and an external hydrological discharge model using OASIS3-MCT instead of the YAC coupler. Using OMCI, it is also possible to set up a RESM with ICON-CLM and other ocean and hydrology models possessing the OASIS3-MCT interface for other regions, such as the Mediterranean Sea.”

>> Minor comments:

>> Line 303 Northern Africa

Corrected. Thank you!

>> Figure 11: “Climatological monthly mean ...” instead of “Annual variability ...”. Please add error bars or shaded areas to your curves denoting the interannual variability. This information is important to judge about the model-data differences.

Done. Thank you!