Response to Reviewers #1, #2 and #3

We like to thank the reviewers for providing helpful comments to improve the manuscript.

We made substantial improvements according to your suggestions. All changes are highlighted in the diff-manuscript below. Added text is wavy-underlined and blue, discarded text is struck out and red. There are also minor changes in some figures that are not highlighted in the diff-manuscript below. Additionally, we slightly changed the algorithms and improved the performance. Therefore, some numbers changed in the manuscript.

The reviewer comments are listed below in black. The author's response is written in blue.

Anonymous Referee #1

General comments:

The manuscript analyses the performance of temperature profile observations by microwave radiometers during rain. This is a relevant topic as more and more of these instruments are operated continuously, and rainy profiles were usually discarded in the past.

I would, however, like to see a more general recommendation on how and under which circumstances these observations can be used. In the discussion, it needs to be clearly stated which channel/angle combination should be used during operation, and which uncertainties are considered to be acceptable, depending on height above ground and rain rate.

We added an additional figure (4) in section 3.2 to highlight which elevation angles and frequencies can be used. Additionally, we added a clear recommendation in the conclusions which retrieval to use.

Concerning the radiative transfer calculations, I would strongly recommend using the updated Rosenkranz gas absorption model from 2022/2023, as this version significantly improves thec radiative transfer results in the lower V-band.

Yes, we agree. We recalculated the radiative transfer with the new Rosenkranz 2022 gas absorption model.

Specific comments:

Lines 29ff: Please provide a more precise statement about which "high frequencies" become opaque? The center V-band frequencies (around 58 GHz) are already opaque without rain.

We have specified the statement as suggested.

Lines 46ff: This motivation can be formulated better, such as: "The method presented here can be applied to standard measurement modes and does not require any changes in measurement setup".

Done as suggested.

Lines 63ff: HATPRO doesn't measure voltages, the detectors convert the antenna signal into voltages. A calibration is necessary to convert the voltages into brightness temperatures (but not "the voltages are then calibrated to brightness temperatures")

Changed as suggested to: "For both absorption bands, HATPRO has its own antenna, which measured signal is converted into voltages of the individual frequencies."

Line 69: Add: "during rain"

Done as suggested.

Section 2.2.: Why did you base your retrieval training on old RS80 radiosondes? These sondes are known to have a dry bias, and therefore might also bias your retrieval. I would strongly recommend using state-of-the-art sondes (Vaisala RS41) which have been used in Lindenberg for quite some time already.

Yes, we agree. Our analysis of the radiative transfer calculations are now based on the stateof-the-art sondes RS41. We adapted the script to the new sounding database.

Section 2.4: It seems you did not use all hourly data from 2000 to 2019 for the retrieval development. That would roughly make 20*365*24=175200 time steps, while you used 58195 profiles. How did you select the data? Are the data evenly distributed over the years and seasons?

Yes, you are right. Originally, we used three hourly data, but we changed now to one hour time resolution for our retrieval development. The number of ERA5 profiles is now 173.088. Data range is now from 2004 to 2023.

Section 3, lines 117ff: Please explain better how spectral retrievals work: Individual channels are highly dependent on each other and can thus be used to retrieve the whole spectrum. Did you apply the spectral consistency check only for zenith observations or for all elevation angles? Please comment on that!

We improved the explanation according to your suggestions. Here in this section, the spectral consistency check is only applied to the zenith observations but can also be applied to other elevation angles. In this case *tbx* retrievals for the specific elevation angle need to be created.

Fig. 3: The y-axis range until +/-60 K makes it difficult to see differences between the channels, especially around 0. I would recommend limiting the range to $+/20$ K. By the way, I don't think that 10 K bias for some channels in the no-rain case (3a) is satisfactory, whereas in 3b for low elevation angles, the agreement is much better. Do you have an explanation for this? For a bias in the input model data, I would expect that the difference changes with elevation angle.

We completely revised this figure according to your suggestions. We added a second row to see both, absolute values (a, b, c) and their difference (d, e, f). We agree that a 10 K bias for Fig. 3.a) is not satisfactory. However we could not find a specific reason for that. That's why we decided to show the non-cloudy case of another day (27 Jul, 2023), where the bias is less, but unfortunately still present.

Fig.3, continued: Furthermore, I couldn't find a clear explanation of the results in 3c, concerning the different fractions of rain used. Can you discuss that a bit more? Don't you consider the solid lines for 51.26 and 52.28 GHz (yellow, blue) for low elevation angles as significant deviations? Concerning the figure style: I cannot really distinguish the different lines in 3c neither.

During our deeper analysis we found out that our LNM disdrometer observation showed unreasonable values during the rain showers. Especially the drop size distribution did not match to the heavy showers on that day in contrast to disdrometer observations at the TROPOS site 5 km away. Therefore, we decided to remove the LNM disdrometer based size distributions and used only rain rates from the HAPRO weather station to assume size distributions as input fir the PAMTRA simulations.

All in all, the illustration is now easier to interpret and recognise.

Page 8, lines 191ff: This paragraph is a bit confusing, try to make your statements more clearly.

We rephrased the paragraph a bit to make it more clear.

Figure 4: Please provide a somewhat larger plot! (same for Fig. 7)

Our intention was here in the preprint to have the same figure size as later in the two-column final version. We believe that the illustration is easy to read.

Section 4.3: Did you classify the rainy cases using rain rates from the model or from observations?

We used the rain rates from the HATPRO weather station. We added that information.

Figure 8: I would strongly recommend to add a comparative figure (putting the red lines from a,c,e,g and b,d,f,h respectively into one figure). Like this, the additional uncertainties depending on the rain rate can be seen much more clearly. What about the performance of 7ν9φ ? I would be interested to see this combination in comparison here as well.

We changed the figure according to your suggestions.

Anonymous Referee #2

I have reviewed the paper titled "Determination of low-level temperature profiles from microwave radiometer observations during rain" by Foth, Lochmann, Saaverdra Garfias and Kalesse-Los. I found the science and its presentation to meet the standards. I hence recommend that this manuscript be accepted with corrections.

Main questions/ concerns:

1. L.65-69: discussion on the wet-radome mitigations on the HATPRO radiometer: What is the HATPRO radome made of? Is it still the blue foam? if so, if it gets wet, it will take time to dry out and, as with sponges, the water will not stop at the top and will fill all the foam by capillarity. This could be problematic for measurements below zenith…

The radome is made of blue foam with an hydrophobic coating. Ageing takes the coating off. We added this information in the manuscript.

2. L.166, L.278 and L.281: On L. 166: "[…] no rain, moderate rain (2.7 mm h−1) and heavy rain(3.7 – 11 mm h–1)". Is moderate defined as 2.7 to 3.6mm/h so that heavy will be defined as 3.7 to 11 mm/h and therefore light will be <2.6 mm/h? your definition is not clear to me. In North America, we have definition for light, moderate and heavy rains: light (< 2.5 mm/h), moderate (2.6 to 7.5 mm/h) and heavy ($>$ 7.6 mm/h); there might be relevant standards in Europe to follow.

Thanks for the hint. We'll apply the definitions from the German weather service:

Light: rain rate \leq 2.5 mm/h moderate: rain rate > 2.5 mm/h up to < 10.0 mm/h heavy: rain rate ≥ 10.0 mm/h very heavy: rain rate ≥ 50.0/h

Although I agree that the new 4n9j scanning strategy can be used up to 2mm/h (L.278) without too much bias, the following lines $(L. 280 - 282)$ in the conclusion:

"In summary, the HATPRO 4ν9φ retrieval method demonstrated in this study achieves unprecedented accuracy of low-level temperature profiling up to 2 km in rain. It was shown that even in heavier rain measurements at elevation angles below 40◦ can be used to derive temperature profiles up to 1.5 km. "

could be interpreted very differently should a casual reader only read the concluding remarks… As such, I would strongly advise that the sentence be re-written much clearer with the limitations of 2mm/h. It is also good practice, even for short papers, to summarize the findings in the conclusion for the casual reader.

Done as suggested.

3. L.250 & Figure 6 : Although I agree that the 4n9j outperforms the other retrievals, there is still some clear influence of the rain in the temperature measurements as seen in Figure 6.h: the retrieved temperature between rain events is smooth like the ECMWF, but during the rain events (3 UTC, 9-13 UTC and 20UTC) the retrieved temperature still shows quite some variability compared to the temperature profiles and still leads to $+/-3K$ temperature difference. This could be because of the wet radome, as the wet radome emissions are likely angle dependent.

Yes, we agree. This might be caused by rain rates above 2.5 mm h^{-1} . We added the discussion in Sec. 4.2.

Minor fixes:

- L. 35 – 37: "Xu et al. (2014) retrieved thermodynamic profiles such as temperature and humidity as well as liquid water profiles by using off-zenith MWR observations at 15◦ elevation to reduce the impact of rain on the measurements. As retrieval technique Xu et al. (2014) used a neural network approach. " would be clearer if written:

Xu et al. (2014) retrieved thermodynamic profiles such as temperature and humidity as well as liquid water profiles by using off-zenith MWR observations at 15◦ elevation to reduce the impact of rain on the measurements [using] a neural network approach.

Done as suggested.

- L. 82: "Their accuracy in contrast to other types of radiosondes is described by Turner et al. (2003). " It is unclear to me where this phrase is going. How does the accuracy of the RS80 compare to other types namely the RS41 which is used by the Lindenberg site (MOL-RAO).

We rephrased this section, since we no longer use the RS80.

- L. 91-95: there is not a single reference to ERA5 papers. At the very least the following paper should be referenced:Hersbach, H. et al. 2020: The ERA5 global reanalysis. Q. J. of the R. Meteorol. Soc., 146, 1999–2049.https://doi.org/10.1002/qj.3803.

Thanks. We added the reference.

- L. 118 – 120: "The second panel (b) shows the results of the spectral consistency checks which is retrieved by the so-called tbx retrievals, which work as follows. There are 14 HATPRO frequencies and only 13 of them are used to estimate the expected value for the $14th$ frequency and then the difference between the estimated and the measured brightness temperature is determined. " would be clearer if written:

The second panel (b) shows the results of the spectral consistency check which is retrieved by the so-called tbx retrieval [. During spectral consistency check (tbx retrievals), 13 of the 14 HATPRO frequencies are used to estimate the value of the unused frequency which is then compared to the measured brightness temperature and the discrepancy is noted.]

Done as suggested.

- Figure caption on Figure 2. "Time series of Moon or Sun and rain quality flag (a), spectral consistency quality flag (b), air temperature and rainfall rate from HATPRO's weather station (c), and height-time series of temperature profiles based HATPRO's firmware radiometer retrieval algorithms in Lindenberg (Germany) on Aug 26, 2020. tb in the colorbar (b) means brightness temperature. "

Should be:

Time series of Moon or Sun and rain quality flag (a), spectral consistency quality flag (b), air temperature and rainfall rate from HATPRO's weather station (c), and height-time series of

temperature profiles based HATPRO's firmware radiometer retrieval algorithms in Lindenberg (Germany) on Aug 26, 2020. tb in the colorbar ([d]) means brightness temperature.

Done as suggested.

Anonymous Referee #3

The paper is about determining temperature profiles with the help of elevation scanning MWRs during rain and how accurate these profiles are. Usually temperature profile retrievals during rain are not possible due to increased opaqueness of the troposphere within the V-band during rain and due to water accumulation on top of the radome. This paper introduces a method on how to retrieve accurate enough temperature profiles in the lower troposphere in spite of these conditions. Key aspects in doing so are only using off-zenith observations and only utilizing the four optical thickest V-band frequencies.

General comments

In my opinion, the necessary quantification of how well the proposed temperature retrieval performs during rainy conditions is missing. Most information is there within the figures but is not stated explicitly in the text.

Overall, the paper is written well and is easy to understand but sometimes details are missing. I will provide more detailed comments and suggestions on what to change below.

We added missing quantification, mentioned in the specific comments below, in the appropriate sections.

General question: What about snowfall? Temperature profiles are usually also not retrieved during snowfall, right? Maybe state in the introduction why you dismiss snowfall and only look into liquid precipitation.

Snow and ice do not emit in the HATPRO frequencies. Cirrus clouds or snow do not disturb the observations. We added the information in the introduction.

Specific comments

Abstract: Quantification missing. What's the accuracy of the new retrieval in different rainfall scenarios?

We added the missing quantification.

32: retrieved Ts from NN approach and 1DVar technique? Here the flow of text seems to suggest that the 1DVar is a form of retrieval (NN and MLR), but it isn't, is it? I think it important to tell a little more what the 1DVar is/does or what it means.

A 1DVAR is a one-dimensional variational approach (also known as optimal estimation technique) and can be seen as an assimilation of an observation in an atmospheric state as first guess. One can retrieve atmospheric profiles by this method. Detailed information are provided by the given reference (Ware et al., 2013).

41: reduced by how much? Your method reduces the error during rain EVEN FURTHER? Should make that clear.

At this point, there is no comparison to our method. Araki et al. (2015) used other rain rate bins and compared their results to soundings, whereas we used statistics based on hourly data of ECMWF profiles. We rephrased the sentences a bit to make it clearer.

Maybe cite Böck et al., 2024 in the introduction section. They look into external measurement uncertainties of scanning HATPROs and what these mean for retrieved temperature profiles.

Done as suggested.

45: "almost saturated" with what? is this quantifiable or is there a source? Or is this something you found out in this study?

By this we mean that the signal in these channels is almost saturated in the sense that there are no more extreme jumps due to liquid water, as is the case in the other channels, for example. The difference in transmissivity is one if the basic principles of MWR temperature profiling.

61: better write: "in the order of seconds". I've seen for exact 5min measurements, that there are only \sim 250 data points and not 300, as expected. So not really a 1s resolution, rather \sim 1.2s.

Done as suggested.

62: K-band, not Ka-band! Please change this in the whole manuscript. 101-102: There are newer Rosenkranz models. Why did you use an older one? Explain.

Done as suggested.

Would using a newer gas absorption model make a difference? For showing what you want to show, the old model is sufficient I guess.

We applied the newer Rosenkranz 2022 absorption model according to suggestions of referee #1.

118: tbx/SPC Retrievals: Are there more details needed for how 13 frequencies predict the 14th?

No, the tbx retrieval here are only based on measured brightness temperature.

150: Maybe explain shortly why only the upper 4 frequencies for elevation scans are used and the lower 3 frequencies for zenith (à optical thickness)

The explanation is given in the following section (3.2) where we describe the selection of frequencies and elevation angles in detail.

166: no range for moderate rain? Why exactly 2.7mm/h?

This refers to Fig. 3 which we already modified. For details see our comments to referee #1. The number refers to the specific rain events in Fig. 3.

177-190: I'd wish for a little bit more quantification here; by how much do TBs differ? (It can be seen in the Figure, but it is not written anywhere).

We added the quantification.

And what is the threshold for significant difference (when does the pink shaded area start and why?)

We added an explanation for the significant difference.

196: "by the less and more transparent..."? get rid of the word less or rephrase.

Done as suggested.

Figure3: y-axis title: change it to Delta TB or something similar, to make clear you're talking about a difference of brightness temperatures here. Maybe just call the shade of color pink instead of rose.

We changed the label of the y-axis.

200: "degreeS of freedom". Please change in the whole manuscript. Also "gives the information content..." sounds strange. I would rephrase.

Done as suggested.

201: You always write "degrees of freedom of signal". Do you always need the word signal or can you omit it?

We introduced the abbreviation DFS, which makes it easier to read.

213-222: Can you quantify the differences a little more in the text? In general, you often describe Figures only qualitatively.

We added more quantification.

229-239 and Figure5: When talking about bias in this context, wouldn't it be better if you evaluate its variance/accuracy as RSME instead of standard deviation? (same for Fig.8). Or is this bias a mean we're talking about and the spread of this mean is then the SD?

Bias is the mean of the absolute differences. We replaced the SD by the RMSE as suggested.

Again: I think it would be better to also quantify your findings in the text.

Yes, we agree and we added more quantifications

240-250: Again: quantify also in the text.

Yes, we agree and we added more quantifications

251-260: Here you do quantify the differences in the text. You should do that everywhere.

Yes, we agree.

Figure8: Again: In this context I'm not sure if you should rather talk about RSME instead of standard deviation. You should check that.

The values won't change much, as the only difference is that you divide by n and not n−1.

It is now the RMSE.

Why is the bias of the 4vz10phi that much worse above 1km in the no rain scenario?

One has to keep in mind that we do not compare here to the truth values (reality), but only to ECMWF model data, which might also be biased. Therefore, the focus here is more on the relative difference between the retrievals and not so much on the absolute values.

276-279: Quantify: How much better does the new 4v9phi retrieval perform and/or with what accuracy during rainfall up to 2mm/h? You only quantify the case in the text with a rainrate of below 0.5mm/h.

We added more quantifications in the manuscript.

280-end: Quantification is missing in the conclusion.

How much Kelvin exactly is the new retrieval method better compared to the standard one?

See comment after next.

E.g. for slight rain below 2km: How much Kelvin is this different to non-rain conditions?

See next comment.

And how much is it different below 1.5km for heavy rain?

A comparison with the standard retrieval during rain is meaningless, as the default retrieval is not intended to be used during rain. The default retrieval is expected to give unreliable results during rain for the reasons mentioned in Sect. 1 and would never be used for that purpose. Temperature profiling during rain is the novelty of our presented approach.

We added more quantification,

In general: What's the general temperature profile accuracy of elevation scanning state-of-the-art HATPROs for no rain scenarios and how does it compare to the new findings?

We added more quantification.

I think this is important for the reader, so they can better classify/categorize the results of this paper.

We agree.

Determination of low-level temperature profiles from microwave radiometer observations during rain

Andreas Foth¹, Moritz Lochmann¹, Pablo Saavedra Garfias¹, and Heike Kalesse-Los¹ ¹Leipzig Institute for Meteorology, Leipzig University, Leipzig, Germany Correspondence: Andreas Foth (andreas.foth@uni-leipzig.de)

Abstract. Usually, microwave radiometer observations have to be discarded during rain. The instrument gets radomes of the receiver antenna get wet which hampers accurate measurements since the retrieval algorithms to derive atmospheric quantities are not trained for rain events. The reason for the latter is, that the rain drops dominate the microwave signal compared to the weaker signal from atmospheric gases. To account for this, radiative transfer simulations need to include the electromagnetic

- 5 properties of rain, which usually requires more complicated and expensive simulations. In this work, the performance of newly developed microwave radiometer retrievals that are not based on rain simulations is evaluated to assess how they work during rain events. It is shown that it is possible to retrieve low-level temperature profiles during rain by omitting certain frequencies and zenith observations. Retrievals with various combinations of elevation angles and frequencies are evaluated. It is presented that, retrievals based on scanning mode observations with angles below $30^\circ 30^\circ$ without zenith observation and only the lesser
- 10 transparent upper four HATPRO microwave radiometer frequencies of the V-band (54.94, 56.66, 57.3, 58 GHz) provides the best results. An analysis of the calculated degrees of freedom of the signal shows that the retrieval of temperature profiles up to 3 km for no rain, $2\frac{1}{2}$ km for light to moderate rain and $1.5\frac{1}{2}$ km for very heavy rain is driven by the HATPRO observation and not by climatology. Finally, the performance of the temperature profile retrieval is explained using a case study in Lindenberg, Germany, and evaluated with temperature profiles from European Center for Medium-range Weather Forecasts (ECMWF)
- 15 model for different rainfall intensities. The results show that the higher the rainfall rate, the larger the deviation of the retrieved microwave radiometer temperature profile retrieval result from the reference from the ECMWF model output. The proposed retrievals for temperature profiles up to at least 1.5 km for rain rates below 0.5 and below 2.5 mm h⁻¹ have uncertainties of less than 1 and 2 K, respectively, compared to ECMWF model output profiles.

1 Introduction

- 20 The continuous development and improvement of weather and climate models poses a great challenge to atmospheric remote sensing. For the evaluation of the models, increasingly better-resolved measurements and retrieval methods are needed, e.g. regarding air temperature profiles. Conventional remote sensing observational approaches mainly fail as they are incapable to provide continuous observations of temperature profiles under all weather conditions and especially during rain. Snow and ice clouds do not emit in the considered spectrum, hence they are not taken into account here. Ground-based Raman
- 25 lidars can usually measure temperature and humidity profiles only below clouds and certainly not during rain (Wandinger,

2005). Radiosondes can provide these atmospheric profiles with high vertical resolution, but they are only routinely available at selected locations and at maximum every 6 hours. Additionally, radiosondes show a significant sonde-to-sonde variability (Nash et al., 2005) as well as a dry bias (Turner et al., 2003).

Multifrequency microwave radiometers (MWR) can provide temporally highly resolved profiles of temperature and humid-30 ity, as well as integrated water vapor and liquid water path (Solheim et al., 1998; Güldner and Spänkuch, 1999; Westwater et al., 2005; Rose et al., 2005). Measurements at different elevation angles increase the accuracy of the derived temperature

- profiles in the atmospheric boundary layer (Crewell and Löhnert, 2007). The measurement uncertainties are described by Böck et al. (2024). Valid retrievals are, however, generally only possible during non-raining conditions (Ware et al., 2004). During rain the atmosphere becomes opaque mainly at high frequencies in the microwave region of the V-band (54.94, 56.66, 57.3,
- 35 58 GHz) and no information can be retrieved from higher altitudes. Additionally, the instrument gets wet and the received signal is dominated by the liquid water accumulated on the instrument. In a previous study, previous studies Cimini et al. (2011) and Ware et al. (2013) compared retrieved profiles of temperature and absolute humidity from a neural network approach (scanning and zenith) and a one-dimensional variational (1DVAR) technique under 15° elevation angle with soundings during all weather conditions. For atmospheric profiling from the surface to 10 km, Cimini et al. (2011) obtained retrieval errors within 1.5 K for
- 40 temperature and 0.5 g m^{-3} for absolute humidity. Xu et al. (2014) retrieved thermodynamic profiles such as temperature and humidity as well as liquid water profiles by using off-zenith MWR observations at $\frac{15°}{15}$ ^o elevation to reduce the impact of rain on the measurements . As retrieval technique Xu et al. (2014) used using a neural network approach. The temperature bias and root mean square error against radiosondes in precipitation were reduced from 3.6 and 4.2 K to 1.3 and 3.1 K, respectively, compared to the zenith MWR observations. Later, Araki et al. (2015) compared the method from Xu et al. (2014)
- 45 with a one-dimensional variational (1DVAR) technique using zenith and off-zenith observation during raining and non-raining conditions. Their results were evaluated with co-located radiosondes . It was shown and they showed that the error in retrieved temperature and water vapor profiles in the low-level troposphere can be reduced by the 1DVAR technique even during rainfall with rain rates less than 1mm h^{-1} , mm h^{-1} by using off-zenith observations. In the presented study, the impact of rain is reduced by using elevation scans only of off-zenith measurements, i.e., at lower elevation angles, because liquid water usually
- 50 accumulates at the top of the MWR. Furthermore, the influence of rain can be reduced by using only the higher frequencies of the oxygen absorption complex (V-band) which in which the signals are almost saturated and will thus not be influenced so strongly by liquid water. The idea of the method presented here is that you can use the common measurement mode that you use for non-rainy situations anywaycan be applied to standard measurement modes and does not require any changes in measurement setup. We show that there is no need to constantly change the measurement mode according to the weather 55 conditions.
	- The structure of the manuscript is as follows: used instruments such as MWR and radiosondes, European Center for Mediumrange Weather Forecasts (ECMWF) model, ERA5 model and radiative transfer models are introduced in the Sec. 2 followed by a description of the retrieval methodology $\lim_{n \to \infty}$ Sec. 3. The retrieval performance based on simulations and observations as well as a comparison of the observations with the ECMWF model output are evaluated in Sec. 4.

Almost all remote sensing data presented in this work were gathered at the Meteorological Observatory Lindenberg - Richard-Assmann-Observatory (MOL-RAO, 52.208°N, 14.118°E) in Lindenberg, Germany, during an instrument intercomparison campaign from July 16, 2020 until October 10, 2020. In addition to that, MWR data presented in Sec. 3 was gathered at the Leipzig Institute for Meteorology, Leipzig University (51.333°N, 12.389°E). The used instruments and models are explained in the 65 following subsections.

2.1 Microwave radiometer HATPRO

The humidity and temperature profiler (HATPRO, generation 5) is a fully automatic microwave radiometer (MWR) from the manufacturer Radiometer Physics GmbH (Rose et al., 2005). It is a passive instrument and measures atmospheric emission at 14 frequencies along the microwave spectrum with a high temporal resolution of 1 sin the order of seconds. Seven frequencies 70 are situated along the upper wing of the water vapor absorption band at 22 GHz (Ka-band) and seven at the lower wing of the oxygen absorption complex at 58 GHz (V-band). For both absorption bands, HATPRO has its own antenna, which measures voltages of measured signal is converted into voltages at the individual frequencies. The voltages are then calibrated to brightness temperatures by automated calibrations (Kazama et al., 1999; Maschwitz et al., 2013; Küchler et al., 2016). The antennae are situated below a radome sheet, which is transparent in the microwave region. It is made of foam with

75 a hydrophobic coating. HATPRO utilizes a rain mitigation system which blows a constant strong air stream over the radome. Nevertheless, during heavy or prolonged rainfall, liquid water might still accumulate on the radome's top, especially if the radome has aged, as is the case during long-term use in the field. This usually An aged radome with a weathered coating absorbs moisture like a sponge. This prevents the accurate determination of atmospheric variables during rain.

In order to estimate column-integrated variables such as the integrated water vapor and liquid water path, as well as vertical

- 80 profiles of temperature and humidity, so-called retrievals must be created (Löhnert and Crewell, 2003). Retrievals are based on artificial neural networks or multi-linear regression models which are trained on relations between measured brightness temperatures and the wanted quantity from radiosondes or numerical weather prediction model output. Observations under different elevation angles enhance the accuracy of the retrieved temperature profile within the atmospheric boundary layer (Crewell and Löhnert, 2007). A sketch showing the HATPROs-HATPRO measurements at default elevation angles color-coded
- 85 by zenith and off-zenith is illustrated in Fig. 1. Those angles were intentionally selected to represent 1, 2, 3, 4, 5, 7, 9, 11, 12, and 14 air masses.

2.2 Radiosondes

Radiosondes provide highly resolved vertical information of atmospheric temperature, humidity and pressure. Here we used a large data set of 9555 Vaisala RS80 soundings from June 1996 to November 2003. 10172 Vaisala RS41 soundings from

90 Lanuary 2015 to April 2024. This serves as input into radiative transfer calculations to create the synthetic brightness temperatures used for the retrieval algorithm to estimate temperature profiles (see Sec. 2.5). Their accuracy in contrast to other

Figure 1. HATPRO's default set of elevation angles. Green and red arrows show off-zenith and zenith elevation angles, respectively.

types of radiosondes is described by Turner et al. (2003). For the comparisons of temperature profiles in Sec. 4.2, Vaisala RS41 radiosondes are used(Sun et al., 2019; Jensen et al., 2016), too (Jensen et al., 2016; Sun et al., 2019). In the presented work, all radiosondes were launched at MOL-RAO.

95 2.3 European Center for Medium-range Weather Forecasts model

In this study, temperature profiles from ECMWF Integrated Forecast System (IFS) are used to evaluate the retrieved temperature profiles from the MWR observations. This is done because the ECMWF-IFSmodel-ECMWF-IFS model data is available in a higher temporal resolution (hourly) than that of the radiosondes. The model data used here are stored in the Cloudnet categorization product (Illingworth et al., 2007) which is freely available at<https://cloudnet.fmi.fi/search/data?site=lindenberg> 100 (last access, 27 Mar 3 Sep , 2024).

2.4 ERA5

ERA5 (ECMWF Reanalysis v5) is the fifth generation of ECMWF's atmospheric reanalysis of global climate (Hersbach et al., 2020) . ERA5 is produced by the Copernicus Climate Change Service (C3S) at ECMWF and covers data from 1940 to present. Here, hourly profiles of temperature, humidityand pressure and cloud liquid with a vertical resolution of 137 pressure

105 levels from the surface up to a height of 80 km are extracted from the global data-set for the MOL-RAO site. 58195 173088 profiles from the ERA5 data set from 2000 to 2019 2004 to 2023 are used as input for the radiative transfer calculation for the temperature retrieval creation.

2.5 Non-scattering microwave radiative transfer model

Based on Simmer (1994), the non-scattering microwave radiative transfer is applied to calculate the brightness temperatures 110 of each profile from 955510172 radio-soundings and 58195173088 ERA5 profiles. This results in a data-set of 183260

profiles with corresponding calculated brightness temperatures which serve as base for the retrieval generation. It uses the gas absorption by Rosenkranz (1998) 2022 Rosenkranz gas absorption (Larosa et al., 2024) and liquid water absorption by Liebe (Liebe et al., 1993). The Rosenkranz gas absorption model is corrected for the water vapor continuum absorption according to Turner et al. (2009). Uncertainty of atmospheric microwave absorption models and their impact on ground-based radiometer

115 simulations and retrievals are extensively described in Cimini et al. (2018). The model code is written in the interactive data language (idl) and was ,e.g. also applied in Löhnert and Crewell (2003); Löhnert et al. (2007); Foth and Pospichal (2017).

2.6 Passive and Active Microwave TRAnsfer PAMTRA

The Passive and Active Microwave TRAnsfer tool (PAMTRA) solves the radiative transfer for passive and active microwave radiation in all-sky conditions, i.e. cloudless, cloudy, and precipitating atmospheres (Mech et al., 2020). In this study, PAMTRA

120 is used to simulate the brightness temperatures at the HATPRO frequencies during rain to investigate the impact of rain in the atmosphere and to assess the effect of liquid water accumulation on the radome (see Sec. 3.2).

3 Methodology

In this section, the problem of retrieving temperature profiles during rain is first shown using an example. Then the theoretical basics of how to create a temperature retrieval are explained. Finally, the procedure to select the most relevant frequencies

- 125 and elevation angles is explained and the results of the information content analysis are shown. Figure 2 (d) illustrates the illustrates a time series of a HATPRO measurement in non-rainy and rainy conditions. The problem of state-of-the-art temperature retrievals during rain, indicated by unrealistic spikes $-\text{is shown in Fig. 2(d)}$. The rain and sun quality flag (a) denotes if rain was detected by HATPRO's weather station or if the Sun or the Moon is directly in the receiver's field of view. Both would affect the quality of the retrieval. The second panel (b) shows the results of the spectral consistency check which is retrieved by
- the so-called *tbx* retrievals, which work as follows. There are Since the signal in the individual channels are highly dependent on each other, they can be used to retrieve the entire spectrum. During spectral consistency check *(tbx* retrievals), 13 of the 14 HATPRO frequencies and only 13 of them are used to estimate the expected value for the 14th frequency and then the difference between the estimated and the value of the unused frequency which is then compared to the measured brightness temperature is determinedand the discrepancy is noted. This procedure is repeated for all 14 frequencies. If the brightness temperature
- 135 difference at a given frequency exceeds the limits of 1 K for $\frac{\text{Ka-band-K-band}}{\text{K-band}}$ and 2 K for V-band, the time steps are flagged with *spectral consistency failed*. This is done here only for zenith observations and it usually happens when nonphysical or unrealistic spectra are measured due to rain or other obstacles in the field of view. During rainy periods none of the frequencies has passed the consistency check, therefore none of the frequencies are reliable to be used. Thus, the state-of-the-art retrieval will not be trustworthy.
- 140 Figure 2 (c) shows the temperature variation and rainfall rate from the HATPRO weather station during the α example day. There are obviously no major no physical temperature gradients during rain events that might explain the height-time series of temperature (d). The presented shown temperature profiles are retrieved by the RPG firmware retrieval for Lindenberg which

Figure 2. Time series of Moon or Sun and rain quality flag (a), spectral consistency quality flag (b), air temperature and rainfall rate from HATPRO's weather station (c), and height-time series of temperature profiles based on HATPRO's firmware radiometer retrieval algorithms in Lindenberg (Germany) on Aug 26, 2020. tb in the colorbar of panel (b) means brightness temperature.

is based on a neural network approach using all 7 V-band frequencies and all 10 elevation angles. This frequency and elevation angle setup corresponds to the state of the art in determining temperature profiles under rain-free conditions.

145 All MWR retrievals, including *tbx* retrievals and for temperature profiles temperature profile retrievals, need to be created for each specific geographic region, as typical atmospheric profiles of temperature and humidity vary across the globe. Walbröl et al. (2022) e.g. created MWR retrievals for low-humidity conditions in the Arctic and Schnitt et al. (2024) for the tropical Atlantic.

3.1 Temperature profile retrieval method

150 The retrieval essentially consists of a series of coefficients that can be based on an artificial neural network or a multi-linear regression model that relates modeled brightness temperatures and temperature profiles (Löhnert and Maier, 2012). In this work we use a regression model. The temperature profiles are based on 955510172 radiosondes and 58195173088 ERA5

output profiles corresponding to the location of the MOL-RAO site in Lindenberg. We decided to use these two different data sources to get a data-set which contains profiles with high vertical resolution (radiosonde) and a large amount of profiles with

- 155 modeled liquid water information (ERA5). From this data-set, temperature, humidity and pressure profiles are extracted. The cloud liquid water content is directly extracted from the ERA5 data. For the radiosonde data a cloud is synthetically determined where 95 % relative humidity is reached (Decker et al., 1978). The modified adiabatic liquid water content is then determined for the altitude range of the cloud according to Karstens et al. (1994). This information is used as input to the non-scattering microwave radiative transfer model (see Sec.2.5). For each input profile the brightness temperatures which would be measured
- 160 by a microwave radiometer under the given input conditions, frequencies and elevation angles are simulated. In total 54200 $146,608$ profiles (80% randomly chosen profiles) were used for the training and 13550 - $36,652$ (20%) to test the regression model to predict the temperature profiles based on simulated brightness temperatures. In this study, different retrieval settings (varying number of frequencies and angles) were generated to contrast the RPG firmware method based on seven frequencies in the V-band (oxygen complex) and ten elevation angles including the zenith direction (90[°]°). Specifically, here a set of 4
- 165 new retrieval setups are proposed that are only based on the upper four HATPRO frequencies in the V-band which exclude the zenith observation (nine angles). The different retrieval setups are listed in Tab. 1. The $4\nu z10\varphi$ retrieval is the most commonly used retrieval for low-level temperature profiling during non-rainy conditions. It uses 10 elevation angles (including the zenith angle) and the upper four frequencies of the V-band. Additionally, the lower three frequencies of the V-band are used at the zenith angle.
- 170 The question how the frequencies and elevation angles for the new temperature retrievals are selected is discussed in the following subsection. The performance during non-raining (cloudy and cloudless) conditions is treated in Sec. 4.1 and is illustrated in Fig. 6.

3.2 Selection of frequencies and elevation angles

To select frequencies and elevations angles for a new temperature retrieval that is less compromised by rain, it is necessary to 175 check which frequencies are less affected by rain accumulated on the radome and by rain in the atmosphere. This was done by a special MWR measurement strategy during a rain event described below. It is worth noting again, that during rain, the atmosphere becomes more opaque with increasing frequency in the V-band.

On July 27, 2023, and on August 1, 2023, on the roof measurement platform of the Institute for Meteorology of Leipzig University, a special measurement was special measurements were performed with the microwave radiometer HATPRO. It

- 180 rained almost There was continuous rain from 9:00 to 15:00 UTC with rain rates, observed by HATPRO weather station, generally below 2 mm h^{-1} on July 27 followed by showers with low intensities. On August 1, it rained continuously from midnight to 8:30 UTC . Rain rates were generally below 2 mm h⁻¹ but occasionally reached 7 mm h⁻¹with rain rates generally below 2mm h^{-1} but occasionally reaching 7mm h^{-1} . Afterwards, there were repeated rain showers and cloudless periods into the night. A scan pattern from θ° until the end of the day. On July 27 at 7:01 UTC as well as on August 1 at
- 185 $\frac{7:41 \text{ UTC}}{2}$ and $\frac{14:14 \text{ UTC}}{2}$ scan patterns from $\frac{0}{2}$ (horizontal) to $\frac{90°-90°}{2}$ (zenith) with $\frac{5°-5}{2}$ elevation angle steps was carried outcontinuouslywere carried out. In addition, PAMTRA simulations of brightness temperatures at all specified elevation an-

Table 1. Retrieval specification. Zenith mode frequencies indicate the frequencies (ν) that are observing only in zenith direction whereas scanning mode frequencies mark those measuring in the directions given by the elevation angle (φ) in the last column. Retrieval name nomenclature: $X\nu[z]Y\varphi$. X: number of frequencies with elevation scanning; Y: number of elevation angles. The index z indicates that, additionally, three zenith observations for 51.26-53.86 GHz have been included in retrieval development (first row). Nomenclature according to Crewell and Löhnert (2007).

	zenith mode frequencies (GHz)	scanning mode frequencies (GHz)	elevation angles $(°)$
$4\nu z10\varphi$	51.26, 52.28, 53.86	54.94, 56.66, 57.3, 58	90, 30, 19.2, 14.4, 11.4
			8.4, 6.6, 5.4, 4.8, 4.2
$4\nu 10\varphi$		54.94, 56.66, 57.3, 58	90, 30, 19.2, 14.4, 11.4
			8.4, 6.6, 5.4, 4.8, 4.2
$7\nu9\n\varphi$		51.26, 52.28, 53.86	30, 19.2, 14.4, 11.4
		54.94, 56.66, 57.3, 58	8.4, 6.6, 5.4, 4.8, 4.2
$4\nu9\varphi$		54.94, 56.66, 57.3, 58	30, 19.2, 14.4, 11.4
			8.4, 6.6, 5.4, 4.8, 4.2

gles were carried out for the three different situations on that day these days: no rain with a thin ice cloud (July 27), moderate rain (2.7mm h^{-1}) and 5.5mm h^{-1} , Aug 1) and very heavy rain $(3.7 - 11 \text{mm h}^{-1}$, (1mm h^{-1}) , (Aug 1) . The ECMWF model output profiles of temperature, pressure and relative humidity in Leipzig from the same day were taken as input 190 for the simulations. Rain drop size distributions for the stratiform rain event early in the day and a for the heavy rain shower around 14:14 UTC were measured estimated by a disdrometer (Type: precipitation laser monitor, LNM from Thies; Fehlmann et al. (2020)) . The spatial variability of rain drop number concentration in the convective afternoon shower needs to be taken into consideration since during elevation scans low and high angles point towards different atmospheric volumes. For that purpose, PAMTRA simulations were run with a) the original rain drop number concentration observed by the LNM, 195 b) a rain drop number concentration which is only 33 modified gamma distribution $(\mu = 2, 2, 1)$ with rain water contents of 0.23 % of a), and c) a rain drop number concentration which is only 66g kg^{-1} and 1.6 % of a) and decreases by half with altitude (from surface up to g kg⁻¹ and number concentrations of 400 m⁻³ and 30 m⁻³, respectively, with uniform rain drop size distributions between cloud base of 2.5 km). This variation in the LNM input allows to account for the heterogeneity

200 simulation. This is important since PAMTRA assumes horizontally homogenious conditions. and the surface. The rain drop size distribution were chosen in a way such that the simulated rain rates match the observations.

The brightness temperatures difference between from HATPRO observations and PAMTRA simulations as from PAMTRA simulations (a,b,c) as well as their difference (d,e,f) as a function of the elevation angle are illustrated in Fig. 3 (a, b, c) for the seven frequencies in the V-band and for three weather conditions (no rain, moderate rain, very heavy rain). It can be seen that

of rain drop number concentration during convective rain and hence helps to assess differences between observation and

205 the simulation and observation fit well for the profile with no rain at $14:337:01$ UTC . The differences originate in the simulation

which on July 27. The differences of around 6 K on average for the lower frequencies and higher elevation angles might be caused by the ECMWF model input which slightly differs from the atmospheric state that was observed by MWR. Additionally, for atmospheric boundary layer scan homogeneous conditions are assumed. If this is not the case, different air masses might be observed by the more transparent channels at 51.26, 52.28, and 53.86 GHz. For the profile at 7:41 UTC on August 1 with

- 210 rain rates $\frac{\text{around } 2.7 \text{ mm h}^{-1}}{2.5 \times 5.5 \text{ mm h}^{-1}}$ (observed) and 5.3 mm h^{-1} (simulated) the brightness temperatures from 51.26, 52.28, and 53.86 GHz -53.86 GHz differ from the simulation above 70° elevation angle 45° elevation angle by up to $26, 18$, and 6 K, respectively. This might be caused by the accumulation of liquid water from rain on the top of the MWR radome. For the heavy rain shower at 14:14 UTC on August ℓ with rain rates between 3.7mm h⁻¹ and 11mm h⁻¹ of 61. ℓ mm h⁻¹ (observed) and 61.7 mm h^{-1} (simulated) the simulated and the observed brightness temperatures at the same three frequencies
- 215 differ by up to $\frac{5036}{28}$ and 10 Kabove 40° , respectively, above 40° elevation angle. 54.94, 56.66, 57.3, and $\frac{58 \text{GHz}}{28}$ GHz as well as all angles below $45^\circ 45^\circ$ are apparently unaffected by the impact of rain and show no significant difference between simulated and observed brightness temperatures. The range of brightness temperature difference at the lower elevation angles (below 45°) is roughly around -5 to 5K. When the brightness temperature difference exceeds this range, this is defined here as significant deviation. That means that all elevation angles below $40°-40°$ and the upper four HATPRO frequencies from the
- 220 V-band can be used to retrieve temperature profiles during rain. It is important to note that most state-of-the-art temperature retrievals from atmospheric boundary layer scans (e.g. HATPRO's firmware) uses the set of elevation angles shown in Fig. 1, thus the majority of elevation angles used by the retrievals are below $40°-40°$ except for the zenith observation.

The spectral consistency check applied to all elevation angles using the corresponding thx retrievals for these angles shows similar results. Figure 4 illustrates the 95th quantile of the brightness temperature difference (observed – retrieved) for different

- 225 elevation angles for all elevation scans that were performed during rain in the observation period. The 95th quantile is used here to exclude outliers and has more significance than median or mean. For 95% of the zenith observations the difference is larger than 2K for all frequencies except for 58 GHz and even for small rain rates. The 58 GHz channel at zenith observation (a) shows small differences since this channel is almost saturated which means that even rain does not increase the observed signal significantly. A typical threshold used for the maximum allowed difference would be 2K as used in the open source
- 230 processing software MWRpy (Marke et al., 2024). Values higher than 2K indicate inconsistency in the spectrum probably caused by rain. For the 30° elevation angle the differences are larger than $3K$ for rain rates above 2.5 mm h⁻¹ and for the first three frequencies of the V-band. Lower elevation angles (below 19.2°) show smaller differences in the brightness temperature and mostly below 2K for all rain rates and frequencies, except the 52.28 GHz channel at 14.4°. This implies that the upper four frequencies of the V-band can be used for temperature retrievals at elevation angles below 30 $^{\circ}$ for rain rates up to 2.5 mm h⁻¹. 235 Disturbances of the observaions by a wet radome would result in larger differences as can be seen at the zenith angle (90°).

One might expect that the addition of the adding the lower HATPRO frequencies of the V-band (i.e. using all seven frequencies in the retrieval) would be more suitableappropriate, as the atmosphere is more transparent at these frequencies. This might be the case if the algorithm pursues to additionally retrieve rain parameters, however for this work we are interested in retrieving only temperature profiles and our analysis However, our analyses have shown that for that purpose the lower V-band

240 frequencies are not optimal and instead increase uncertainties. Horizontally homogeneous conditions are assumed for bound-

Figure 3. Difference between observed Observed and simulated brightness temperatures (a,b,c) as well as their difference (d,e,f) for different frequencies (colors) versus elevation angle for no rain (a, d), moderate rain (b,e), and very heavy rain events (c,f). Different line styles in (c) label different LNM rain inputs into the PAMTRA simulation. Rose rectangle marks the area where the observations significantly differ from the simulation probably caused by wet radome. Note that the y-axis in (c) differs from (a) and (b).

ary layer scans. At low elevation angles, however, different air masses are observed by the less and more transparent channels leading to uncertainties in the retrieved profiles.

3.3 Information content analysis

- To investigate how much information originates from the observations and not from the climatology, an optimal estimation 245 technique has been applied to the case studies (Rodgers, 2000; Maahn et al., 2020). It calculates the degree degrees of freedom of a signal and gives (DFS) and specifies the information content that comes from the measurement itself. The cumulated degree of freedom of signal of DFS of all four retrievals are illustrated in Fig. 5 for the three weather conditions (a, b, c) mentioned in Sec. 3.2. The curves of all retrievals have a similar shape and differ only slightly in the upper layers. With $\phi_{\text{Q, QQ}}$ differ significantly below heights of 3 km. However, with increasing altitude, the difference of cumulative degrees of freedom
- 250 of signal are larger. Once the differences of cumulative DFS increase. Once a DFS curve reaches a vertical line no more information is added by the measurements. The retrievals with fewer frequencies and angles $(4\nu 10\varphi, 4\nu 9\varphi)$ display lower

Figure 4. 95th quantile of brightness temperature difference (observed – retrieved) per frequency (y-axis) and rain rate (x-axis) for different elevation angles (a-j). The grey dashed boxes mark the area of rain impact determined by differences of more than 2 K.

Figure 5. The cumulated degrees of freedom of signal and temperature profiles for no rain (a, d) conditions on Jul 27, 2023, moderate rain (b, e) and heavy rain (c, f) conditions on Aug 1, 2023.

values of the cumulated degrees of freedom of signal DFS under all three weather conditions. This means that there is less information from altitudes above roughly 1.5 km from the measurement and the profile is more driven by the climatology. The more rain there is in the atmosphere, the lower the information content of the measurement, as can be seen in the maximum

- 255 value of the cumulated degree of freedom of signal DES in 3 km which reaches values between 3 and 4 for no rain and between 2.8 and 3.4 during moderate rain and between 2.627 and 3.4 during very heavy rain. Summarizing, the retrieved temperature profile is driven by the measurement at least up to 3 km for no rain, about $2\frac{1}{2}\frac{5}{2}$ km for rain and about $\frac{1}{2}\frac{5}{2}$ km for heavy rain proven by the determined degree of freedom of signal-DES indicated by the point at which the line with lowest information content (red) becomes vertical.
- 260 The retrieved temperature profiles from the four retrievals, as well as the ECMWF temperature output profile for the same three conditions (no rain, rain, moderate rain, very heavy rain) are illustrated in Fig. 5(d, e, f). As expected for non-rainy conditions (d) and shown in section 4.1, all four retrievals show similar deviations from the reference ECMWF profile in the lowest 1.5 km. Above 1.5 km the $4\nu 9\varphi$ differs from 4ν z 4φ 10φ , $4\nu 10\varphi$ and $4\nu 10\varphi$ show the smallest difference to ECMWF output which serves as reference here. But also the $4\nu9\varphi$ performs similarly. Only the $7\nu9\varphi$ underestimates the ECMWF
- 265 temperature in higher altitudes and which may caused by the fact that frequencies with different transparencies observe different air masses at lower elevation angles, as explained above 4ν 9 φ as well as from ECMWF output. For the moderate rain case (e), all retrievals perform similar although the zenith observation is affected by rain as shown in Sec. 3.2, which might indicate that at the observed rain rates of ≤ 2.7 mm h⁻¹ temperature retrieval profiles are less affected by rain in the atmospherebelow about 1 km. Retrievals which use zenith observations (4vz10y and 4v10y) perform worse than the others (7v9y and 4v9y).
- 270 The 7v9x retrieval performs best and shows smallest differences to the ECMWF profile with a difference of 1K below 2 km. For the very heavy rain event (f), the 7ν 9φ and 4ν 9φ retrieval retrievals shows the best performance indicated by the smallest difference to the reference ECMWF model output. As expected 4ν z 10φ and $4\nu10\varphi$ have largest deviations (more than 12 K) in 2 km) from ECMWF model output since they are intentionally made for non-rainy conditions. It is likely that the ECMWF temperature profile does not represent the truth, especially during rain showers. For this reason, no quantitative statement is
- made here and more attention is paid to the intercomparison between the individual retrievals. 275

4 Results

This section first shows the performance of the newly created temperature profile retrievals based on simulations with the test data-set under non-rainy conditions. This is only to show that the new different retrievals produce meaningful results. In section 4.2, the retrieval performance is evaluated on the basis of observations using the ease study MOL-RAO case study of Aug 26, 280 2020 , introduced in Sec. 3. Finally, the retrieved temperature profiles are compared to ECMWF output on a larger data-set.

4.1 Retrieval performance based on simulations during non-raining conditions

The performance of the new approaches $(7\nu 10\varphi, 4\nu 10\varphi, 4\nu 9\varphi)$ in comparison to the common retrieval $(4\nu 210\varphi)$ under nonraining idealized conditions is shown in Fig. 6. This is the result of the test data from the atmospheric profiles from radiosonde

Figure 6. Temperature retrieval performance in terms of bias (a), standard deviation root mean square error (RMSE, b) and coefficient of determination (c) based on synthetic data (trained with radio-soundings and ERA5) during cloudy and cloudless conditions.

- and ERA5 (36 552 profiles). Bias (a), standard deviation (root mean square error (RMSE, b), and coefficient of determination 285 (\mathbb{R}^2_{∞} c) between true values and the prediction of the regression model indicate how much uncertainty is added by omitting frequencies and elevation angles during cloudy and cloudless conditions using profiles from the test data-set. All four sets of retrievals show similar behavior in bias (a), namely just small systematic deviations from zeroat around 2 km. For all four retrievals, standard deviation RMSE (b) increases with altitude while R^2 decreases with altitude, both indicating an increase in uncertainty with height. BiasRMSE and R^2 diverge above 1 km with 4ν 9 φ being worse whereas 4ν z10 φ , standard deviation 290 $\frac{7\nu 10\varphi}{2}$ and $\frac{4\nu 10\varphi}{2}$ almost overlap. Bias, RMSE and \mathbb{R}^2 values are in accordance with Crewell and Löhnert (2007). Highest
- uncertainties are evident for the $4\nu9\varphi$ retrieval. This is an expected behavior since information can be lost by omitting frequencies and zenith observations as shown in Fig. 5, whereas 5. In conclusion, the $4\nu9\varphi$ retrieval does not perform as well as the other retrievals which is expected as it was optimized for rainy conditions.

4.2 Case study based on observations

- 295 If the four versions of the temperature retrieval (The four temperature profile retrievals introduced in Tab. 1) from the previous section are were applied to the MOL-RAO example on $_{\odot}$ Aug 26, 2020, from the problem description (Sec. 2), one can see the improvement by selecting only lower elevation angles and the higher frequencies. 2020. Results are displayed in Figure 7 shows where the height-time plots of the ECMWF model temperature (a), the four temperature retrievals (b, d, f, h), the ECMWF model temperature (a), as well as the difference of the retrieved temperatures to the ECMWF model temperature
- 300 (c, e, g, i) are shown. As introduced above (Sec. 2.1) there are three rain events on that day, early morning around 03 UTC, between 09 and 13 UTC and around 20 UTC (see Fig. 2 (a). One can see that during During all rain events with rain rates between 0 and at maximum 10mm h^{-1} 10mm h^{-1} the spectral consistency check failed (Fig. 2b). The presence of the rain in the lower atmosphere or even-accumulated liquid water on the radome compromises the retrieval output indicated by the unrealistic spikes in the temperature profiles ($Fig. 7b$, d, f) and by a high temperature difference (c, e, g). It is obvious that
- 305 neither). Neither the $4\nu 10\varphi$ nor the $4\nu 10\varphi$ nor the $7\nu 9\varphi$ work can be applied during rain conditions, as can be seen by very large positive temperature differences of more then 10 K above 1 km and values below -3 K below 1 km during a the rain events. However, the 7ν 9φ as well as the 4ν 9φ retrieval can tackle the rain limitation and $\frac{1}{18}$ are able to produce reasonable results in

Figure 7. Height-time series of temperature profiles from ECMWF model (a) and temperature profiles based on different retrieval algorithms (b, d, f, h) and associated temperature difference to ECMWF model (c, e, g, i) in Lindenberg (Germany) on Aug 26, 2020. The radiosonde launch times are indicated by white dashed lines.

comparison to the ECMWF model temperature output (f_{z,}g_z, h, i)with the lowest temperature differences during rainy periods.

Their deviations are mostly below 3 K during rain and mostly between -1 and 1 K for the rest of the day. Nevertheless, during

- 310 the rain events there is some variability in the $7\nu 9\varphi$ and $4\nu 9\varphi$ retrievals in contrast to the ECMWF profile. This is probably caused by a wet radome as the rain rates are larger than 2.5 mm h⁻¹ (see Fig. 2) which is the threshold derived in Fig. 4. Figure 8 illustrates acomparison between the retrieved temperature profiles and three To further estimate the performance of the four temperature profile retrievals, they are compared to **radiosonde launches at MOL-RAO on Aug 26, 2020, 04:45 (a)**, 10:45 (b) and 22:25 UTC (c) in Figure 8. During the launch at 04:45 UTC in non-raining conditions there are no significant
- 315 differences between the retrievals, the sounding and the ECMWF temperature profiles (a). The differences are much higher during the rain event at 10:45 UTC (b) with rain rates around 1.5 mm h^{-1} . Sounding and ECMWF model temperature profile are in good agreement and only the $7\nu 9\varphi$ and the $4\nu 9\varphi$ retrievals fit the sounding as reference within less than 2 K near the s urface and 4 below 1 K and km. Above 1 km the 7ν 9 φ retrieval performs best, since it almost overlaps with the sounding. The 4ν 9 φ retrieval deviates around 3 K , respectively, at 2 km. In contrast, the temperature retrievals from 4ν z10 φ and 7ν 10 φ , are
- 320 eompletely off by over 10 K above 1 km. The temperature profile comparison during the short and light rain event shower with rain rates below 0.5 mm h⁻¹ at around 22:45 UTC in Fig. 8 (c) shows a similar result, the $7\nu9\varphi$ and the $4\nu9\varphi$ retrieval even fit to the reference sounding within the expected sounding uncertainty.

Up to this point, the performance of the retrieval has been evaluated only on the basis of case studies. In the next section it will be evaluated against ECMWF model output using a larger data set.

325 4.3 ECMWF model comparison

In this section the performance of the proposed- $4\nu 9\varphi$, $7\nu 9\varphi$ and the state-of-the-art $4\nu z10\varphi$ temperature retrieval against ECMWF model temperature profiles is investigated. Therefore, all three months of HATPRO observation at MOL-RAO from July to October 2020 are taken into account. Hourly ECMWF model temperatures are interpolated to the measurement grid of approximately 20 minutes per temperature profile, since there is a routine elevation scan every 20 minutes. Figure 9 shows 330 the retrieval performance in terms of bias (left panels) and standard deviation (root mean square errors (RMSE, right panels)

Figure 8. Panels a, b, and c show a comparison of three retrieved temperature profiles obtained from the four different retrievals with radiosoundings launched at MOL-RAO on Aug 26, 2020, at 4:45 (a), 10:45 (b), and 22:45 UTC (c) and ECMWF model output from 5, 11, and 23 UTC.

between ECMWF output and retrievals for non-raining cases (a, b), and raining cases with rain rates smaller than 0.5 mm h⁻¹ 0.5 mm h^{-1} (c, d), rain rates between 0.5 and 2 mm h^{-1} 2.5 mm h^{-1} (e, f) and rain rates larger that 2 mm h^{-1} 2.5 mm h^{-1} (g, h). The rain rates used here, are from the HATPRO weather station. During non-raining conditions (3732 3671 sample profiles) both all retrievals agree well with the ECMWF output (Fig. 9 a,b b, c, dashed) as could be expected from Fig. 6. But

- 335 for small rain rates 64 -below 0.5 mm h⁻¹ (57 sample profiles) the proposed $4\nu 9\varphi$ agrees much better with a bias less than of around 1 K (Fig. 9 c₂ dash-dot) and a standard deviation RMSE ranging between 0.5 and 2 K (d). The state-of-the-art retrieval $(4\nu z10\varphi)$ leads to very high deviations from the ECMWF temperature profiles with biases and standard deviations RMSE's around 5 to 7 K and 5 to 10 K, respectively, apart from altitudes above below 0.5 km. The bias of the 7ν 9 φ and the 4ν 9 φ retrievals increase with height and reach a maximum values of around 4 K in 3 km for rain rates between 0.5 and 2.5 mm h⁻¹.
- **340** The corresponding RMSE's are around 1.5 K within the lowest 1 km and increase up to around 5 K at 3 km. For rain rates above 2.5 mm h⁻¹ the biases and RMSE are largest for each retrieval. The higher the rain rate, the worse the performance of the MWR temperature profile retrievals. Although the 7ν 9 φ and 4ν 9 φ is are significantly better than the common 4ν z 10φ retrieval, it deviates they deviate from the ECMWF outputby 6 K (bias) at 3 km altitude in heavy rain. Of course the ECMWF model output is not the truth, especially $\frac{1}{m}$ during rain, but serves as a reference for comparing the two three retrievals. It should
- 345 be noted that the new 7ν 9 φ and 4ν 9 φ retrieval performs retrievals perform better since the common 4ν z 10φ retrieval setup was intentionally not developed for working under raining conditions. Summarizing, that the new proposed retrieval based on MWR observation under lower elevation angles and only the higher V-band frequencies allows to resolve temperature profiles during rain with rain rates up to $2m + h^{-1}$, \approx mm h⁻¹ which was not possible before with the state-of-the-art retrievals.

5 Conclusions and Outlook

350 In summary, the HATPRO $4\nu9\varphi$ retrieval method demonstrated in this study achieves unprecedented accuracy of low-level temperature profiling up to with a bias of less than 1.5 K and an RMSE below 2 K up to 3 km in rain -with rain rates below 0.5 mm h⁻¹ compared to ECMWF temperature profiles. For rain rates between 0.5 and 2.5 mm h⁻¹ the bias increases up to 2K and RMSE up to 3K in 1.5km. An intercomparison of the different retrievals during non-raining conditions showed a good agreement in bias and RMSE values, respectively. As shown based on ERA5 and radiosonde data, the proposed 4v90 retrieval 355 performs very similar to the state-of-the-art $4\nu z10\varphi$ retrieval up to 1.5 km, during non-rainy conditions. Above these heights, the RMSE increases up to 1.2 K instead of 0.8 K in 3 km as the 4vz10 g. 7v10 g and 4v10 g retrievals which almost overlap. The bias is very similar to the stat-of-the-art retrieval around zero from surface up to 3 km. It was shown that even in heavier $\frac{\text{rain-}\text{very heavy rain}}{61 \text{mm h}^{-1}}$ measurements at elevation angles below 40 \degree can be used to derive temperature profiles up to 1.5 km . The temperature using the 4ν 9 φ . The 7 ν 9 φ partially performs better than the suggested 4ν 9 φ , but in general the 360 4v9q is proposed to be used in most cases. The lower frequencies of the V-band used in the 7v10q are more transparent and hence observe different air masses in the lower elevation angles which might lead to large uncertainties especially in the case of spatially variable precipitation. The recommendation is to use the 4ν 9 φ retrieval for rain rates below 2.5 mm h⁻¹ to retrieve temperature profiles up to 1.5 km with uncertainties less than 2 K.

Figure 9. Bias (left panel) and standard deviation root mean square error (right panel) between retrieved and ECMWF temperature profiles for rainfree cases (a, b), rain free and raining rain cases with different rain rates smaller than 0.5 mm h⁻¹ (e, dlines), rain rates between 0.5 and 2 mm h⁻¹ for different retrievals (erows, f∞olors)and rain rates larger that 2 mm h⁻¹ (g, h). Red lines (dash dot dot) mark the 4ν9φ r etrieval and blue dashed lines mark the common 4ν z 10φ retreival. N denotes the number of time steps taken into account at German Weather Service Observatory in Lindenberg (MOL-RAO) between 16 Jul, 2020 and 8 Oct, 2020. Bias is defined as retrieved minus ECMWF output as reference.The colored area marks values between the 25 and 75 percentile. Note that in (a) the x-axis is different from that in the other panels.

The temperature retrievals can be easily applied with an existing open source software (mwp). In addition, the 365 published software package can be used to create custom retrievals for arbitrary locations (Foth, 2023) user-defined locations (Foth, 2024b). This represents a significant improvement towards the reliability of using MWR for weather nowcasting or forecast. Especially-Improved low-level temperature profile retrievals are of great values for the following applications: investigations of evaporative cooling during precipitation evaporation are often very inaccurate due to incorrect assumptions of temperature which can compromise can be improved by more accurate temperature profile retrievals which can in turn improve 370 the reliability of the evaluation of model parameterizations. Furthermore, the proposed method can be applied retrospectively to correct temperature profiles from long-term observations as long as the MWR scanning brightness temperature data is avail-

- able for the post-processing. In the future , In this way improved climatologies of MWR-based temperature profiles can be derived.
- Several future modifications to even increase the performance of the presented retrieval are envisioned: an optimal es-375 timation method which is also a variational technique should could be used in further investigations. Only In contrast to Cimini et al. (2011), only HATPRO frequencies that pass the consistency check for all elevation angles should be used at each time step independent of the rain situation. Thus, a continuous time series of temperature profiles can be created, which provides physical uncertainties for each time and height range. This might also improve profiles of absolute humidity which is also of interest for the evaporation studies. Additionally, long-term HATPRO observations will enable a quantification of the 380 maximum rain rate at which the new $4\nu 9\varphi$ retrieval can be applied.

Code and data availability. The HATPRO raw data is processed with MWRpy version 0.8.2 [\(https://github.com/actris-cloudnet/mwrpy\)](https://github.com/actris-cloudnet/mwrpy). Also some MWRpy subroutines for plotting are used in this study. The optimal estimation software package pyOptimalEstimation version 1.2 is available under<https://github.com/maahn/pyOptimalEstimation> and described in detail in Maahn et al. (2020). The Passive and Active Microwave TRAnsfer model PAMTRA is also available on github.com [\(https://github.com/igmk/pamtra\)](https://github.com/igmk/pamtra) and is already published in Mech

385 et al. (2019, 2020). ERA5 data is available under<https://cds.climate.copernicus.eu/> (Hersbach et al., 2019). The HATPRO data from the general scans in Leipzig is available at zenodo (Foth, 2024a). The Lindenberg HATPRO and model data used in this study are generated by the Aerosol, Clouds and Trace Gases Research Infrastructure (ACTRIS) and are available from the ACTRIS Data Centre using the following links: [https://doi.org/10.60656/ca8017ee6ef94027,](https://doi.org/10.60656/ca8017ee6ef94027) [https://doi.org/10.60656/E938967BC0524DEE.](https://doi.org/10.60656/E938967BC0524DEE) The retrievals are made with the pyMakeRetrieval routines version 1.2.0 (Foth, 2024b) and are available on github [\(https://github.com/remsens-lim/pyMakeRetrieval\)](https://github.com/remsens-lim/pyMakeRetrieval).

390 *Author contributions.* AF prepared the manuscript in close cooperation with ML, PSG and HKL. AF performed the investigations and data analyses. ML and AF realized the experimental setup in Lindenberg and Leipzig, respectively, and were responsible for the high quality of the HATPRO measurements. The conceptualization was initialized by AF, ML and PSG. All authors have contributed to the scientific discussions.

Competing interests. The authors declare that they have no conflict of interest.

395 *Acknowledgements.* The authors thank the LIM-team and the MOL-RAO team for supporting the HATPRO observations in Lindenberg. The authors also acknowledge the ACTRIS-Cloudnet team and all associated developers for the well documented code around remote sensing especially the HATPRO processing within $mwpyMWRpy$. This research has been supported by the German Science Foundation (DFG) (grant nos. FO 1285/2-11). PSG was funded by the Deutsche Forschungsgemeinschaft (DFG), Transregio-project TR-172 Arctic Amplification, German Research Foundation) – project no. 268020496 – TRR 172, within the Transregional Collaborative Research Center 400 "ArctiC Amplification: Climate Relevant Atmospheric and SurfaCe Processes, and Feedback Mechanisms (AC)³ (grant 268020496)3". sub-project B07 (grant 437153667) and E05, and grant number FO 1285/2-11. This research has been supported by the Federal State of Saxony and the European Social Fund (ESF) in the framework of the program "Projects in the fields of higher education and research" (grant no. 232101734 and 100339509). This work was funded by the Saxon State Ministry for Science, Culture and Tourism (SMWK) [3-7304/44/4-2023/8846].

405 References

- Araki, K., Murakami, M., Ishimoto, H., and Tajiri, T.: Ground-Based Microwave Radiometer Variational Analysis during No-Rain and Rain Conditions, Sola, 11, 108–112, https://doi.org[/10.2151/sola.2015-026,](https://doi.org/10.2151/sola.2015-026) 2015.
- Böck, T., Pospichal, B., and Löhnert, U.: Measurement Uncertainties of Scanning Microwave Radiometers and Their Influence on Temperature Profiling, Atmospheric Meas. Tech., 17, 219–233, https://doi.org[/10.5194/amt-17-219-2024,](https://doi.org/10.5194/amt-17-219-2024) 2024.
- 410 Cimini, D., Campos, E., Ware, R., Albers, S., Giuliani, G., Oreamuno, J., Joe, P., Koch, S. E., Cober, S., and Westwater, E.: Thermodynamic Atmospheric Profiling During the 2010 Winter Olympics Using Ground-Based Microwave Radiometry, IEEE Trans. Geosci. Remote Sens., 49, 4959–4969, https://doi.org[/10.1109/TGRS.2011.2154337,](https://doi.org/10.1109/TGRS.2011.2154337) 2011.
	- Cimini, D., Rosenkranz, P. W., Tretyakov, M. Y., Koshelev, M. A., and Romano, F.: Uncertainty of Atmospheric Microwave Absorption Model: Impact on Ground-Based Radiometer Simulations and Retrievals, Atmospheric Chem. Phys., 18, 15 231–15 259,

415 https://doi.org[/10.5194/acp-18-15231-2018,](https://doi.org/10.5194/acp-18-15231-2018) 2018.

- Crewell, S. and Löhnert, U.: Accuracy of Boundary Layer Temperature Profiles Retrieved With Multifrequency Multiangle Microwave Radiometry, IEEE Trans. Geosci. Remote Sensing, 45, 2195–2201, https://doi.org[/10.1109/TGRS.2006.888434,](https://doi.org/10.1109/TGRS.2006.888434) 2007.
- Decker, M. T., Westwater, E. R., and Guiraud, F. O.: Experimental Evaluation of Ground-Based Microwave Radiometric Sensing of Atmospheric Temperature and Water Vapor Profiles, J. Appl. Meteorol. Climatol., 17, 1788–1795, https://doi.org[/10.1175/1520-](https://doi.org/10.1175/1520-0450(1978)017%3C1788:EEOGBM%3E2.0.CO;2) 420 [0450\(1978\)017<1788:EEOGBM>2.0.CO;2,](https://doi.org/10.1175/1520-0450(1978)017%3C1788:EEOGBM%3E2.0.CO;2) 1978.
	- Fehlmann, M., Rohrer, M., von Lerber, A., and Stoffel, M.: Automated Precipitation Monitoring with the Thies Disdrometer: Biases and Ways for Improvement, Atmospheric Meas. Tech., 13, 4683–4698, https://doi.org[/10.5194/amt-13-4683-2020,](https://doi.org/10.5194/amt-13-4683-2020) 2020. Foth, A.: pyMakeRetrieval v1.1.1, Zenodo, https://doi.org[/10.5281/ZENODO.10014291,](https://doi.org/10.5281/ZENODO.10014291) 2023.

Foth, A.: Brightness Temperature Data and Weather Station Data from General Scans of the Microwave Radiometer HATPRO, 425 https://doi.org[/10.5281/zenodo.13692454,](https://doi.org/10.5281/zenodo.13692454) 2024a.

Foth, A.: pyMakeRetrieval, Zenodo, https://doi.org[/10.5281/zenodo.13692444,](https://doi.org/10.5281/zenodo.13692444) 2024b.

Foth, A. and Pospichal, B.: Optimal Estimation of Water Vapour Profiles Using a Combination of Raman Lidar and Microwave Radiometer, Atmos. Meas. Tech., 10, 3325–3344, https://doi.org[/10.5194/amt-10-3325-2017,](https://doi.org/10.5194/amt-10-3325-2017) 2017.

Güldner, J. and Spänkuch, D.: Results of Year-Round Remotely Sensed Integrated Water Vapor by Ground-Based Microwave Radiometry, J 430 Appl Meteor Clim., 38, 981–988, https://doi.org[/10.1175/1520-0450\(1999\)038<0981:ROYRRS>2.0.CO;2,](https://doi.org/10.1175/1520-0450(1999)038%3C0981:ROYRRS%3E2.0.CO;2) 1999.

Hersbach, H., Bell, B., Berrisford, P., Biavati, G., Horányi, A., Muñoz Sabater, J., Nicolas, J., Peubey, C., Radu, R., Rozum, I., Schepers, D., Simmons, A., Soci, C., Dee, D., and Thépaut, J.-N.: ERA5 Hourly Data on Pressure Levels from 1940 to Present, https://doi.org[/10.24381/cds.6860a573,](https://doi.org/10.24381/cds.6860a573) 2019.

Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas, J., Peubey, C., Radu, R., Schepers, D., Simmons,

435 A., Soci, C., Abdalla, S., Abellan, X., Balsamo, G., Bechtold, P., Biavati, G., Bidlot, J., Bonavita, M., Chiara, G., Dahlgren, P., Dee, D., Diamantakis, M., Dragani, R., Flemming, J., Forbes, R., Fuentes, M., Geer, A., Haimberger, L., Healy, S., Hogan, R. J., Hólm, E., Janisková, M., Keeley, S., Laloyaux, P., Lopez, P., Lupu, C., Radnoti, G., Rosnay, P., Rozum, I., Vamborg, F., Villaume, S., and Thépaut, J.-N.: The ERA5 Global Reanalysis, Quart. J. Roy. Meteor. Soc., 146, 1999–2049, https://doi.org[/10.1002/qj.3803,](https://doi.org/10.1002/qj.3803) 2020.

Illingworth, A. J., Hogan, R. J., O'Connor, E., Bouniol, D., Brooks, M. E., Delanoé, J., Donovan, D. P., Eastment, J. D., Gaussiat, N., Goddard, 440 J. W. F., Haeffelin, M., Baltink, H. K., Krasnov, O. A., Pelon, J., Piriou, J.-M., Protat, A., Russchenberg, H. W. J., Seifert, A., Tompkins,

A. M., van Zadelhoff, G.-J., Vinit, F., Willén, U., Wilson, D. R., and Wrench, C. L.: Cloudnet: Continuous Evaluation of Cloud Profiles in

Seven Operational Models Using Ground-Based Observations, B. Am. Meteorol. Soc., 88, 883–898, https://doi.org[/10.1175/BAMS-88-](https://doi.org/10.1175/BAMS-88-6-883) [6-883,](https://doi.org/10.1175/BAMS-88-6-883) 2007.

Jensen, M. P., Holdridge, D. J., Survo, P., Lehtinen, R., Baxter, S., Toto, T., and Johnson, K. L.: Comparison of Vaisala Radiosondes RS41

- 445 and RS92 at the ARM Southern Great Plains Site, Atmospheric Meas. Tech., 9, 3115–3129, https://doi.org[/10.5194/amt-9-3115-2016,](https://doi.org/10.5194/amt-9-3115-2016) 2016.
	- Karstens, U., Simmer, C., and Ruprecht, E.: Remote Sensing of Cloud Liquid Water, Meteorl. Atmos. Phys., 54, 157–171, https://doi.org[/10.1007/BF01030057,](https://doi.org/10.1007/BF01030057) 1994.

Kazama, S., Rose, T., Zimmermann, R., and Zimmermann, R.: A Precision Autocalibrating 7 Channel Radiometer for Environmental Re-

- 450 search Applications, J. Remote Sens. Soc. Jpn., 19, 265–273, https://doi.org[/10.11440/rssj1981.19.265,](https://doi.org/10.11440/rssj1981.19.265) 1999.
	- Küchler, N., Turner, D. D., Löhnert, U., and Crewell, S.: Calibrating Ground-based Microwave Radiometers: Uncertainty and Drifts, Radio Sci., 51, 311–327, https://doi.org[/10.1002/2015RS005826,](https://doi.org/10.1002/2015RS005826) 2016.
- Larosa, S., Cimini, D., Gallucci, D., Nilo, S. T., and Romano, F.: PyRTlib: An Educational Python-based Library for Non-Scattering Atmospheric Microwave Radiative Transfer Computations, Geosci. Model Dev., 17, 2053–2076, https://doi.org[/10.5194/gmd-17-2053-2024,](https://doi.org/10.5194/gmd-17-2053-2024) 455 2024.
	- Liebe, H. J., Hufford, G. A., and Cotton, M. G.: Propagation Modeling of Moist Air and Suspended Water/Ice Particles at Frequencies below 1000 GHz, AGARD, 1993.
		- Löhnert, U. and Crewell, S.: Accuracy of Cloud Liquid Water Path from Ground-Based Microwave Radiometry 1. Dependency on Cloud Model Statistics, Radio Sci., 38, 8041, https://doi.org[/10.1029/2002RS002654,](https://doi.org/10.1029/2002RS002654) 2003.
- 460 Löhnert, U. and Maier, O.: Operational Profiling of Temperature Using Ground-Based Microwave Radiometry at Payerne: Prospects and Challenges, Atmos. Meas. Tech., 5, 1121–1134, https://doi.org[/10.5194/amt-5-1121-2012,](https://doi.org/10.5194/amt-5-1121-2012) 2012.
	- Löhnert, U., van Meijgaard, E., Baltink, H. K., Groß, S., and Boers, R.: Accuracy Assessment of an Integrated Profiling Technique for Operationally Deriving Profiles of Temperature, Humidity, and Cloud Liquid Water, J. Geophys. Res., 112, D04 205, https://doi.org[/10.1029/2006JD007379,](https://doi.org/10.1029/2006JD007379) 2007.
- 465 Maahn, M., Turner, D. D., Löhnert, U., Posselt, D. J., Ebell, K., Mace, G. G., and Comstock, J. M.: Optimal Estimation Retrievals and Their Uncertainties: What Every Atmospheric Scientist Should Know, Bull. Amer. Meteor. Soc., https://doi.org[/10.1175/BAMS-D-19-0027.1,](https://doi.org/10.1175/BAMS-D-19-0027.1) 2020.
	- Marke, T., Löhnert, U., Tukiainen, S., Siipola, T., and Pospichal, B.: MWRpy: A Python Package for Processing Microwave Radiometer Data, J. Open Source Softw., 9, 6733, https://doi.org[/10.21105/joss.06733,](https://doi.org/10.21105/joss.06733) 2024.
- 470 Maschwitz, G., Löhnert, U., Crewell, S., Rose, T., and Turner, D. D.: Investigation of Ground-Based Microwave Radiometer Calibration Techniques at 530 hPa, Atmos. Meas. Tech., 6, 2641–2658, https://doi.org[/10.5194/amt-6-2641-2013,](https://doi.org/10.5194/amt-6-2641-2013) 2013.
	- Mech, M., Maahn, M., Ori, D., and Orlandi, E.: PAMTRA: Passive and Active Microwave TRAnsfer Tool v1.0, Zenodo, https://doi.org[/10.5281/zenodo.3582992,](https://doi.org/10.5281/zenodo.3582992) 2019.

Mech, M., Maahn, M., Kneifel, S., Ori, D., Orlandi, E., Kollias, P., Schemann, V., and Crewell, S.: PAMTRA 1.0: The Passive and Active

- 475 Microwave Radiative TRAnsfer Tool for Simulating Radiometer and Radar Measurements of the Cloudy Atmosphere, Geosci. Model Dev., 13, 4229–4251, https://doi.org[/10.5194/gmd-13-4229-2020,](https://doi.org/10.5194/gmd-13-4229-2020) 2020.
	- Nash, J., Smout, R., Oakley, T., Pathack, B., and Kurnosenko, S.: WMO Intercomparison of High Quality Radiosonde Systems: Final Report, WMO Rep., p. 118 pp, 2005.
	- Rodgers, C. D.: Inverse Methods for Atmospheric Sounding Theory and Practice, vol. 2, World Scientific Publishing, 2000.
- 480 Rose, T., Crewell, S., Löhnert, U., and Simmer, C.: A Network Suitable Microwave Radiometer for Operational Monitoring of the Cloudy Atmosphere, Atmos. Res., 75, 183–200, https://doi.org[/10.1016/j.atmosres.2004.12.005,](https://doi.org/10.1016/j.atmosres.2004.12.005) 2005.
	- Rosenkranz, P. W.: Water Vapor Microwave Continuum Absorption: A Comparison of Measurements and Models, Radio Sci., 33, 919–928, https://doi.org[/10.1029/98RS01182,](https://doi.org/10.1029/98RS01182) 1998.
	- Schnitt, S., Foth, A., Kalesse-Los, H., Mech, M., Acquistapace, C., Jansen, F., Löhnert, U., Pospichal, B., Röttenbacher, J., Crewell, S.,
- 485 and Stevens, B.: Ground- and Ship-Based Microwave Radiometer Measurements during EUREC⁴A, Earth Syst. Sci. Data, 16, 681–700, https://doi.org[/10.5194/essd-16-681-2024,](https://doi.org/10.5194/essd-16-681-2024) 2024.
	- Simmer, C.: Satellitenfernerkundung Hydrologischer Parameter Der Atmosphäre Mit Mikrowellen, Kovač, 1994.
	- Solheim, F., Godwin, J. R., Westwater, E. R., Han, Y., Keihm, S. J., Marsh, K., and Ware, R.: Radiometric Profiling of Temperature, Water Vapor and Cloud Liquid Water Using Various Inversion Methods, Radio Sci., 33, 393–404, https://doi.org[/10.1029/97RS03656,](https://doi.org/10.1029/97RS03656) 1998.
- 490 Sun, B., Reale, T., Schroeder, S., Pettey, M., and Smith, R.: On the Accuracy of Vaisala RS41 versus RS92 Upper-Air Temperature Observations, J. Atmospheric Ocean. Technol., 36, 635–653, https://doi.org[/10.1175/JTECH-D-18-0081.1,](https://doi.org/10.1175/JTECH-D-18-0081.1) 2019.
	- Turner, D., Cadeddu, M., Löhnert, U., Crewell, S., and Vogelmann, A.: Modifications to the Water Vapor Continuum in the Microwave Suggested by Ground-Based 150-GHz Observations, IEEE Trans. Geosci. Remote Sensing, 47, 3326–3337, https://doi.org[/10.1109/TGRS.2009.2022262,](https://doi.org/10.1109/TGRS.2009.2022262) 2009.
- 495 Turner, D. D., Lesht, B. M., Clough, S. A., Liljegren, J. C., Revercomb, H. E., and Tobin, D. C.: Dry Bias and Variability in Vaisala RS80-H Radiosondes: The ARM Experience, J. ATMOSPHERIC Ocean. Technol., 20, 16, 2003.
	- Walbröl, A., Crewell, S., Engelmann, R., Orlandi, E., Griesche, H., Radenz, M., Hofer, J., Althausen, D., Maturilli, M., and Ebell, K.: Atmospheric Temperature, Water Vapour and Liquid Water Path from Two Microwave Radiometers during MOSAiC, Sci Data, 9, 534, https://doi.org[/10.1038/s41597-022-01504-1,](https://doi.org/10.1038/s41597-022-01504-1) 2022.
- 500 Wandinger, U.: Raman Lidar, in: Lidar Range-Resolved Optical Remote Sensing of the Atmosphere, edited by Weitkamp, C., vol. 102 of *Springer Series in Optical Sciences*, pp. 241–271, Springer Berlin/Heidelberg, 2005.
	- Ware, R., Cimini, D., Herzegh, P., Marzano, F., Vivekanandan, J., and Westwater, E.: Ground-Based Microwave Radiometer Measurements during Preicipitation, in: 8th Specialst Meeting on Microwave Radiometry, p. 3, Rome, Italy, 2004.
- Ware, R., Cimini, D., Campos, E., Giuliani, G., Albers, S., Nelson, M., Koch, S. E., Joe, P., and Cober, S.: Thermodynamic and Liquid 505 Profiling during the 2010 Winter Olympics, Atmospheric Research, 132–133, 278–290, https://doi.org[/10.1016/j.atmosres.2013.05.019,](https://doi.org/10.1016/j.atmosres.2013.05.019)
	- Westwater, E. R., Crewell, S., Mätzler, C., and Cimini, D.: Principles of Surface-Based Microwave and Millimeter Wave Radiometric Remote Sensing of the Troposphere, Quad. Soc. Ital. Elettromagnetismo, 1, 50–90, 2005.
- Xu, G., Ware, R. S., Zhang, W., Feng, G., Liao, K., and Liu, Y.: Effect of Off-Zenith Observations on Reducing the Im-510 pact of Precipitation on Ground-Based Microwave Radiometer Measurement Accuracy, Atmospheric Research, 140–141, 85–94, https://doi.org[/10.1016/j.atmosres.2014.01.021,](https://doi.org/10.1016/j.atmosres.2014.01.021) 2014.

2013.