

1 **An investigation into atmospheric nitrous acid (HONO)**
2 **processes in South Korea**

3
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18 **Short title:** Atmospheric HONO processes

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24 **Abstract**

25 Nitrous acid (HONO) is a main precursor of hydroxyl radicals (OH), which contribute
26 to the formation of numerous secondary air pollutants in the troposphere. Despite its
27 importance in atmospheric chemistry, HONO chemistry has not been fully incorporated into
28 many chemical transport models (CTMs). Due to the lack of atmospheric HONO processes,
29 CTM simulations often tend to underestimate atmospheric mixing ratios of HONO. This study
30 was undertaken because simulations with current Community Multiscale Air Quality (CMAQ)
31 model have a strong tendency to underestimate the HONO mixing ratio. In search of missing
32 sources of atmospheric HONO, we attempted to sequentially incorporate the following
33 potential HONO sources and processes into the CMAQ modeling framework: (i) gas-phase
34 HONO reactions; (ii) traffic HONO emissions; (iii) soil HONO emissions; (iv) heterogeneous
35 HONO production on the surfaces of aerosols; (v) heterogeneous HONO formation on tree leaf
36 and building surfaces; (vi) photolysis reactions of particulates and deposited HNO₃/nitrates
37 called 'renoxification'. The simulation performances of the modified CMAQ models were then
38 evaluated by comparing the modeled HONO mixing ratios with the HONO mixing ratios
39 observed at the Olympic Park station in Seoul, South Korea. When HONO processes were fully
40 added to the CMAQ model, average daily HONO mixing ratios increased from 0.06 ppb to
41 1.18 ppb. The daytime HONO mixing ratios produced from the CMAQ model run with a full
42 account of atmospheric HONO processes were found to be in better agreement with
43 observations than those from the original CMAQ model (CMAQv5.2.1) runs with improved
44 statistical metrics (e.g., IOA increased from 0.59 to 0.68, while MB decreased dramatically
45 from -0.57 ppb to -0.34 ppb). In addition, we investigated the contributions of individual
46 atmospheric HONO processes to HONO mixing ratios, as well as the impacts of HONO
47 atmospheric processes on the concentrations of other atmospheric species in South Korea. All
48 these issues are also discussed in this manuscript.

49 **Keywords:** Nitrous acid (HONO); Heterogeneous HONO production; CMAQ model; Ozone
50 production rate.

51 **1. Introduction**

52 Hydroxyl radicals (OH) play a key role in atmospheric chemistry. OH radicals oxidize
53 volatile organic compounds (VOCs), sulfur dioxide (SO₂), and nitrogen dioxide (NO₂),
54 contributing to the formation of secondary organic and inorganic aerosols (Pathak et al., 2009).
55 Therefore, accurate determination of the mixing ratio of OH radicals is crucial to understanding
56 atmospheric photochemistry in both polluted and remote areas.

57 Nitrous acid (HNO₂ or HONO) has been recognized as a main precursor of OH radicals
58 via photo-dissociation (R1) (Alicke et al., 2003; Harris et al., 1982; Kleffmann et al., 2005):



60 Several studies have estimated that HONO photolysis reactions contribute 20 – 80% of OH
61 radicals and 30 – 87% of HO_x formation in polluted urban areas (Acker et al., 2006; Alicke et
62 al., 2003; Hendrick et al., 2014; Kim et al., 2014; Kleffmann et al., 2005; Monks et al., 2009;
63 Ren et al., 2003). However, it was also recognized that the HONO chemistry was not yet fully
64 understood.

65 Therefore, many field measurements have been carried out to characterize atmospheric
66 HONO processes (Kim et al., 2014; Lee et al., 2016; Li et al., 2012; Su et al., 2008). These
67 studies showed that the observed HONO mixing ratios were significantly higher than those
68 predicted by atmospheric chemistry-transport model simulations (Lee et al., 2016; Li et al.,
69 2014; Su et al., 2008; VandenBoer et al., 2013). This indicates that there should be missing
70 HONO sources or processes that are not considered in current atmospheric models (CTMs).

71 Recent studies have proposed incorporating several HONO production pathways into
72 chemical transport models to explain the missing HONO processes. Suggested sources include
73 i) traffic HONO emissions (Czader et al., 2015; Kirchstetter et al., 1996; Kurtenbach et al.,
74 2001; Li et al., 2018; Nakashima and Kajii, 2017; Rappenglück et al., 2013; Xu et al., 2015);

75 ii) soil HONO emissions (Meusel et al., 2016; Nagai and Kubota, 1972; Oswald et al., 2013;
76 Weber et al., 2015); iii) HONO emissions from biomass burning (Cheng et al., 2014; Crutzen
77 and Andreae, 1990; Nie et al., 2015); iv) indoor HONO emissions (Gligorovski, 2016; Zhang
78 et al., 2019); and iv) heterogeneous conversion of NO₂ to HONO on the surfaces of aerosols,
79 grounds, and leaves (Han et al., 2017; Reisinger, 2000; Svensson et al., 1987; Stemmler et al.,
80 2016; Wiesen et al., 1995).

81 Among these processes, traffic HONO emissions were reported to be the key factor
82 influencing the HONO mixing ratio in the Beijing–Tianjin–Hebei (BTH) region at night
83 (Zhang et al., 2019). Heterogeneous NO₂ reactions on aerosol surfaces were an important
84 source of HONO during the severe haze period in Beijing (Jia et al., 2020). On the other hand,
85 Zhang et al. (2016) reported that heterogeneous reactions on ground surfaces could be the
86 dominant source of atmospheric HONO, accounting for ~42% of the HONO mixing ratios in
87 Hong Kong suburban areas.

88 These findings may indicate that atmospheric HONO production and a potential cause
89 of discrepancies between modeled and observed HONO mixing ratios may vary temporally
90 and regionally. In addition, no research has been conducted on which sources of HONO control
91 the levels of HONO in Seoul, South Korea. In this context, the aims of this study are three-fold:
92 i) to determine which HONO sources or processes are significant in South Korea; ii) to estimate
93 the budget of the HONO mixing ratios from various HONO sources; and iii) achieving
94 objectives i) and ii) to develop a near-perfect CTM in terms of HONO mixing ratio. To
95 accomplish these goals, we decided to improve the US EPA CMAQ v5.2.1 model by
96 incorporating several HONO production pathways including i) homogeneous HONO reactions;
97 ii) direct HONO emissions from biomass burning, traffic vehicles, and soil; iii) heterogeneous
98 HONO production on the surfaces of atmospheric aerosols, buildings, and tree leaves; and iv)
99 photolysis reactions of particulate and deposited HNO₃/nitrate (renoxification).

100 We then tested the performances of the modified CMAQ models by comparing the
101 modeled HONO mixing ratios with the HONO mixing ratios observed during the Korea-United
102 States Air Quality (KORUS-AQ) campaign. After the comparison analysis, we evaluated the
103 contributions of individual HONO processes to the HONO budget in South Korea and also
104 investigated the effects of the HONO mixing ratios on the levels of other important atmospheric
105 species.

106 **2. Methodology**

107 In this study, we incorporated various HONO sources and reactions into the CMAQ
108 model framework to accurately estimate HONO mixing ratios in the atmosphere. Then, the
109 simulation results of the modified CMAQ models were analyzed, comparing the modeled
110 outputs with observations during the KORUS-AQ campaign. Details of the modifications of
111 CMAQ models, the HONO measurements, and potentially important HONO sources
112 considered in this study are described in the following sections.

113 **2.1. WRF-CMAQ model configuration**

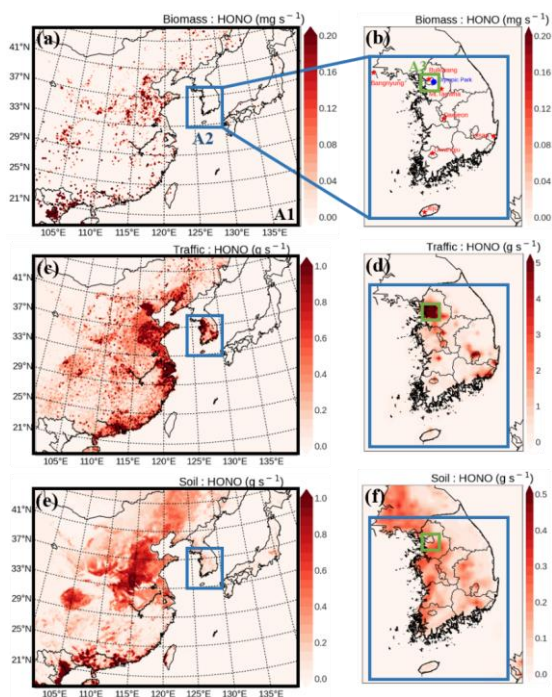
114 Simulation of the Community Multiscale Air Quality (CMAQ) v5.2.1 model (Byun
115 and Schere, 2006) was carried out to estimate the HONO mixing ratios during the period of the
116 KORUS-AQ campaign (9 May – 12 June, 2016). Figure 1 shows the horizontal domain (A1)
117 for the CMAQ model simulation. The spatial domain has 273×204 grid cells with a horizontal
118 resolution of 15 × 15 km² and contains 15 vertical layers with the first layer at ~34 m above
119 the ground.

120 The photochemical mechanism used in the simulation of the CMAQ model was the
121 Statewide Air Pollution Research Center-07 (SAPRC-07 TC) (Carter, 2010; Hutzell et al.,
122 2012). The AERO6 module was used for aerosol calculations (Binkowski and Roselle, 2003).
123 In particular, the heterogeneous reactions considered in this study were embedded into the

124 SAPRC-07 TC via the Korean Flexible Chemistry (KFC) editor. The KFC editor is a chemical
125 mechanism editor in a framework of Graphic User Interface (GUI) developed to quickly
126 implement the modifications of the chemical mechanisms of the CMAQ model. The details of
127 the heterogeneous reactions we considered are discussed in Sects. 2.3.5 – 2.3.6.

128 The Weather Research and Forecasting (WRF) v3.8.1 model (Skamarock et al., 2008)
129 was run to generate meteorological fields that drive the CMAQ model. The physical options
130 used in the WRF run are as follows: i) WRF Single-Moment 6-class Microphysics scheme
131 (Hong and Lim, 2006); ii) Rapid radiative transfer model (RRTMG) for longwave and
132 shortwave radiation (Iacono et al., 2008); iii) the NOAH Land Surface scheme (Chen and
133 Dudhia, 2001); iv) Yonsei University (YSU) Planetary Boundary Layer (PBL) scheme (Hong
134 et al., 2006); v) MM5 surface layer scheme (Jiménez et al., 2012); and vi) the Grell-Freitas
135 Ensemble scheme for cumulus physics (Grell and Freitas, 2014). Initial and boundary
136 conditions for the WRF model runs were obtained from the National Center for Environmental
137 Prediction Final Analysis (NCEP-FNL) every six hours.

138 For anthropogenic emissions, this study used the KORUS v5.0 inventory processed by
139 the Sparse Matrix Operator Kernel Emissions in Asia (SMOKE-Asia; Woo et al., 2012) (Woo
140 et al., 2020). The KORUS v5.0 emissions were developed particularly for CTM runs as part of
141 the KORUS-AQ campaign. Biogenic emissions were generated using the Model of Emissions
142 of Gases and Aerosol from Nature (MEGAN) v2.10 (Guenther et al., 2012). Fire emissions
143 were obtained from the Fire Inventory from NCAR (FINN) v1.5 emission inventory
144 (<https://bai.acom.ucar.edu/Data/fire/>; Wiedinmyer et al., 2011). The various HONO emissions
145 considered in this study are discussed in Sects. 2.3.2. – 2.3.4.



146

147 **Figure 1.** Spatial distributions of HONO emission rates from biomass burning (panels (a) and
 148 (b)), from traffic (panels (c) and (d)), and from soil (panels (e) and (f)) over East Asia (A1),
 149 South Korea (A2), and the Seoul Metropolitan Area (A3). Several super-monitoring stations
 150 are located at Bangnyung-Do, Bulkwang-Dong, Olympic Park, Mt. Taehwa, Daejeon, Gwangju,
 151 Ulsan, and Jeju. The locations of these super-stations are shown in panel (b).

152 2.2 Measurements

153 During the KORUS-AQ campaign period, concentrations of NO₂, O₃, and particulate
 154 matter were measured at several locations such as Olympic Park (37.52N; 127.12E),
 155 Bangnyung (37.96N; 124.64E), Bulkwang (37.61N; 126.93E), Mt. Taewha (37.31N; 127.31E),
 156 Daejeon (36.35N; 127.38E), Gwangju (35.23N; 126.84E), Ulsan (35.53N; 129.31E), and Jeju
 157 (33.32N; 126.40E) (refer to Fig. 1b regarding the locations). In this study, we also used data
 158 observed at approximately 320 stations from the AIR-KOREA network
 159 (<https://www.air.korea.or.kr>), officially managed by the Korean Ministry of Environment in
 160 South Korea.

161 Surface data observed at the Olympic Park station in Seoul were used for direct comparisons
162 between simulated and observed HONO mixing ratios. These HONO mixing ratios were
163 measured using the Monitor for AeRosols and GAses in ambient air (MARGA ADI 2080)
164 (Applikon-ECN, Netherlands) instrument with a time resolution of 1 hour. This measurement
165 is based on the wet-denuder-ion-chromatography (WD/IC) method. In the WD/IC system,
166 HONO molecules absorbed by the solution in the denuder were converted to nitrite (NO_2^-),
167 and then the nitrite concentrations were quantified by ion chromatography (Xu et al., 2019).
168 The detection limit of the MARGA instrument for HONO is ~ 0.02 ppb. At the Olympic Park
169 station, NO_2 and O_3 were also measured using commercially available instruments, EC9841
170 and EC9810, respectively, manufactured by Ecotech. Their detection limits for both species
171 are ~ 0.5 ppb during the daytime. Details on the principles of EC9841 and EC9810 can be
172 found in Keywood et al. (2019). $\text{PM}_{2.5}$ at the Olympic Park station were measured
173 continuously using a Thermo Scientific Continuous Particulate Monitor, FH62C14, based on
174 the beta attenuation method. The detection limit of the instrument is $4\mu\text{g}/\text{m}^3$ in hourly
175 measurements. Further information about instruments is provided in Table S1.

176 Meteorological data on temperature, relative humidity, pressure, wind speeds, and
177 wind directions were also measured by the Automated Synoptic Observing System (ASOS) at
178 the Olympic Park station. In order to test the simulation performances of the WRF-CMAQ
179 model, observed meteorological data was compared with the modeled outputs, which is shown
180 in [Figure S1 and](#) Table S2. In general, WRF model simulations tended to accurately predict
181 meteorological fields.

182 **Table 1.** Comparison of parameterizations of HONO processes between CMAQ v5.2.1 and this study.

	HONO processes	CMAQ v5.2.1	This study	Ref.		
Reaction	(R1) $\text{HONO} + h\nu \xrightarrow{J_{\text{HONO}}} \text{OH} + \text{NO}$	J_{HONO}	J_{HONO}	1		
	(R2) $\text{NPHE} + h\nu \xrightarrow{J_{\text{NPHE}}} \text{HONO} + \text{xPROD2}$	$J_{\text{NPHE}} = 1.50e^{-3} \times J_{\text{NO}_2}$	$J_{\text{NPHE}} = 1.50e^{-3} \times J_{\text{NO}_2}$			
	(R3) $\text{OH} + \text{NO} + \text{M} \xrightarrow{k_3} \text{HONO}$	$k_3 = \left\{ \frac{k_a[\text{M}]}{(1+k_a[\text{M}]/k_b)} \right\} 0.6^{\frac{1}{1+\log_{10}\left(\frac{k_a[\text{M}]}{k_b}\right)^2}}$ <ul style="list-style-type: none"> ▪ $k_a = 7.0 \times 10^{-31} \left(\frac{T}{300}\right)^{-2.6}$, ▪ $k_b = 3.6 \times 10^{-11} \left(\frac{T}{300}\right)^{-0.1}$ 	<ul style="list-style-type: none"> ▪ $k_a = 7.1 \times 10^{-31} \left(\frac{T}{300}\right)^{-2.6}$, ▪ $k_b = 3.6 \times 10^{-11} \left(\frac{T}{300}\right)^{-0.1}$ 	2		
	(R4) $\text{HONO} + \text{OH} \xrightarrow{k_4} \text{H}_2\text{O} + \text{NO}_2$	$k_4 = 2.5 \times 10^{-12} e^{\left(\frac{-260}{T}\right)}$	$k_4 = 3.0 \times 10^{-12} e^{\left(\frac{-250}{T}\right)}$			
	(R5) $2\text{NO}_2 + \text{H}_2\text{O} \xrightarrow{k_{\text{aerosol}}} \text{HONO} + \text{HNO}_3$	$k_{\text{aerosol}} = 1.0 \times 10^{-4} (S/V)$	$k_{\text{aerosol}} = \frac{1}{4} v_{\text{NO}_2} (S/V) \times \gamma_{\text{a,NO}_2}$ Daytime: $\gamma_{\text{a,NO}_2} = 1.3 \times 10^{-4} \times \frac{\text{light intensity}}{900(W \cdot M^{-2})}$ Nighttime: $\gamma_{\text{a,NO}_2} = 8.0 \times 10^{-6}$	3, 4 5		
	(R6) $2\text{NO}_2 + \text{H}_2\text{O} \xrightarrow{k_{\text{ground}}} \text{HONO} + \text{HNO}_3$	$k_{\text{ground}} = 5.0 \times 10^{-5} \times \left(\frac{S_{\text{g,building}}}{V} + \frac{S_{\text{g,leaf}}}{V} \right)$ <ul style="list-style-type: none"> ▪ $\frac{S_{\text{g,leaf}}}{V} = \frac{2 \times \text{LAI}}{H}$ ▪ $\frac{S_{\text{g,building}}}{V} = \text{PURB} \times \frac{0.2 \frac{\text{m}^2}{\text{m}^3}}{100\%}$ 	$k_{\text{ground}} = \frac{1}{8} \times v_{\text{NO}_2} \times \gamma_{\text{g,NO}_2} \times \left(\frac{S_{\text{g,building}}}{V} + \frac{S_{\text{g,leaf}}}{V} \right)$ <ul style="list-style-type: none"> ▪ $\frac{S_{\text{g,leaf}}}{V} = \frac{2 \times \text{LAI}}{H}$ ▪ $\frac{S_{\text{g,building}}}{V} = \text{PURB} \times \frac{0.3 \frac{\text{m}^2}{\text{m}^3}}{100\%}$ Daytime: $\gamma_{\text{g,NO}_2} = 5.8 \times 10^{-6} \times \frac{\text{light intensity}}{900(W \cdot M^{-2})}$ Nighttime: $\gamma_{\text{g,NO}_2} = 5.0 \times 10^{-7}$	6 7 4, 8 9		
		(R7) $p\text{NO}_3 + h\nu \xrightarrow{J_{p\text{NO}_3}} 0.67\text{HONO} + 0.33\text{NO}_2$	-	$J_{p\text{NO}_3} = 118 \times J_{\text{HNO}_3}$	10	
		(R8) $\text{Deposited_HNO}_3/\text{nitrate} + h\nu \xrightarrow{J_{\text{D_HNO}_3/\text{nitrate}}} 0.67\text{HONO} + 0.33\text{NO}_2$	-	$J_{\text{D_HNO}_3/\text{nitrate}} = 48 \times J_{\text{HNO}_3}$		
		Emission	Biomass Burning	-	FINNv1.5	11
			Traffic	-	Gasoline: $\text{HONO}_{\text{traffic}}/\text{NO}_x = 0.8\%$ Diesel: $\text{HONO}_{\text{traffic}}/\text{NO}_x = 2.3\%$	7
Soil	-		$\text{HONO}_{\text{soil}}/\text{NO}_x = f(\text{soil water content})$	12		

1. Burkholder et al. (2015); 2. Burkholder et al. (2020); 3. Xue et al. (2022); 4. Czader et al. (2012); 5. Vandenboer et al. (2013); 6. Sarwar et al. (2008); 7. Zhang et al. (2016); 8. Yu et al. (2021); 9. Yu et al. (2022); 10. Fu et al. (2019); 11. Wiedinmyer et al. (2011); 12. Meusel et al. (2018)

184 **2.3 HONO sources**

185 In this study, we considered several possible missing HONO sources or processes in
 186 the CMAQ model simulations. The possible missing HONO sources include gas-phase HONO
 187 reactions, three HONO emission sources, three heterogeneous HONO reactions, and two
 188 photolytic reactions. The considered possible missing HONO sources are also contrasted to the
 189 current HONO processes embedded in the CMAQ v5.2.1 model in Table 1. The details of each
 190 HONO process are discussed below.

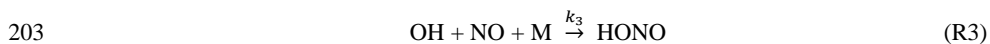
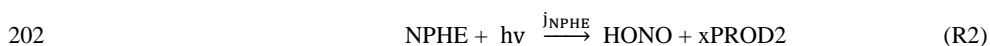
191 **Table 2.** Design for 8 EXP simulations.

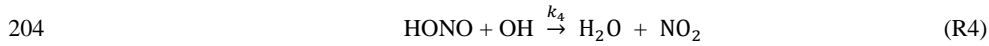
HONO Source	Experiment							
	EXP1	EXP2	EXP3	EXP4	EXP5	EXP6	EXP7	EXP8
GAS ¹⁾	√	√	√	√	√	√	√	√
BioB ²⁾		√	√	√	√	√	√	√
TRAF ³⁾			√	√	√	√	√	√
SOIL ⁴⁾				√	√	√	√	√
HET_A ⁵⁾					√	√	√	√
HET_L ⁶⁾						√	√	√
HET_BD ⁷⁾							√	√
RENO _x ⁸⁾								√

192 ¹⁾ Gas-phase reactions; ²⁾ Biomass Burning Emissions; ³⁾ Traffic Emissions; ⁴⁾ Soil Emissions; ⁵⁾ Heterogeneous reactions on
 193 aerosol surfaces; ⁶⁾ Heterogeneous reactions on the surfaces of leaves; ⁷⁾ Heterogeneous reactions on the surfaces of buildings;
 194 ⁸⁾ Renoxification

195 **2.3.1 Gas Phase reactions (GAS)**

196 We used the SAPRC-07 TC chemical mechanism as base mechanism. A total of 4 gas-
 197 phase HONO-related reactions were considered for HONO formation and dissociation (Carter,
 198 2010;). HONO is produced by i) photolysis of nitrophenol (NPHE) (R2) and ii) reaction of
 199 NO with OH in the presence of the third body (M) (R3). Meanwhile, HONO is removed by
 200 reaction with OH radicals (R4) and photolytic dissociation (R1). All these reactions are shown
 201 below:





205 where, J_{NPHE} (R2) and J_{HONO} (R1) are the photolysis rates constants of NPHE and HONO,
206 respectively, which were adopted from the study of Stockwell et al. (1990). As shown in Table
207 1, J_{NPHE} was calculated, based on J_{NO_2} (i.e., $J_{\text{NPHE}} = 1.50 \times 10^{-3} \times J_{\text{NO}_2}$), which was defined
208 in Bejan et al. (2006). k_3 and k_4 are the reaction rate constants of (R3) and (R4) and were
209 obtained from the National Aeronautics and Space Administration (NASA) Jet Propulsion
210 Laboratory (JPL) Publication 19 (Burkholder et al., 2020). Among these reactions, the reaction
211 rate constants of (R3) and (R4) were updated in our study (refer to Table 1). The effect of these
212 gaseous reactions on HONO mixing ratios was tested in the EXP1 simulation (see GAS in
213 Table 2).

214 **2.3.2 Biomass burning emissions (BioB)**

215 Biomass combustion includes three types of burning events: natural wildfires,
216 agricultural fires, and wood burning (Wiedinmyer et al., 2011). In East Asia, agricultural fires
217 typically occur in early summer and fall (Ryu et al., 2004; Tao et al., 2013; Zhang et al., 2013).
218 The period of the KORUS-AQ campaign coincides with the period of the agricultural residue
219 burning after barley and wheat harvest in East Asia. Biomass burning emissions, including
220 agricultural fire emissions, were obtained from the Fire INventory from NCAR version 1.5
221 (FINN v1.5, Wiedinmyer et al., 2011; Wiedinmyer et al., 2006). This was then considered in
222 the EXP2 simulation (see BioB in Table 2). The spatial distributions of HONO emissions from
223 the biomass burning events in the East Asia domain (A1), South Korea domain (A2), and Seoul
224 Metropolitan Area domain (A3) are presented in Fig. 1a and 1b. However, we found that the
225 HONO emission rates used in the EXP2 simulation were relatively small, compared to the total
226 HONO emission rates presented in Fig. 1 and Table 3.

227

228

229 **Table 3.** HONO emission rates from biomass burning, traffic, and soil. The total HONO
230 emission rates during the period of the KORUS-AQ campaign are shown.

Region	Source			
	Biomass Burning Emission ($\text{g} \cdot \text{s}^{-1}$)	Traffic Emission ($\text{Mg} \cdot \text{s}^{-1}$)	Soil Emission ($\text{Mg} \cdot \text{s}^{-1}$)	Total ($\text{Mg} \cdot \text{s}^{-1}$)
East Asia (A1)	2.46	6.40	5.65	14.51
South Korea (A2)	0.00	0.32	0.06	0.38
Seoul Metropolitan Area (A3)	0.00	0.10	0.01	0.1

231 **2.3.3 Traffic emissions (TRAF)**

232 Traffic emissions are an important HONO source, particularly at night (Zhang et al.,
233 2016). HONO is emitted directly from vehicle exhaust systems. In this study, to estimate the
234 direct HONO emissions from traffic sources, we assumed that the HONO to NO_x emission
235 ratio is 0.8% for gasoline vehicles and 2.3% for diesel vehicles (Sarwar et al., 2008; Zhang et
236 al., 2016). All off-road vehicles were treated as diesel vehicles in the calculations of HONO
237 emissions (Gutzwiller et al., 2002). Table 3 presents the total emission rates for East Asia,
238 South Korea, and the Seoul Metropolitan Area, which are 6.40, 0.32, and 0.1 $\text{Mg} \cdot \text{s}^{-1}$,
239 respectively. Moreover, as shown in Fig. 1c and 1d, HONO emissions from traffic sources are
240 dominant, particularly in metropolitan areas such as Seoul, Beijing, Shanghai, and Hong Kong.
241 The contribution of traffic sources to total HONO emissions was estimated to be dominant in
242 the Seoul Metropolitan Area. In the EXP3 simulation, the impact of the traffic source (see
243 TRAF in Table 2) on the atmospheric HONO mixing ratios was investigated.

244 **2.3.4 Soil emissions (SOIL)**

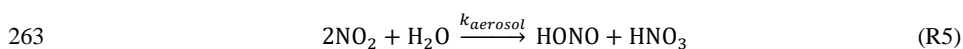
245 Emissions from soil bacterial activity are important sources of HONO. The amount of
246 their emissions depends on the soil type, land category, fertilization, temperature, soil water

247 content (SWC in %), and soil pH (Meusel et al., 2018; Wu et al., 2020). In this study, HONO
248 emissions were estimated based on the ratio of HONO to NO_x emissions from soil (Oswald et
249 al., 2013). SWC was used as a proxy for soil pH due to the technical limitations of direct
250 measurement of the soil pH. The SWCs were set at 0-7.5%, 7.5-15%, 15-20%, 20-30% and 30-
251 40% for HONO-to-NO_x ratios of 1.0, 0.67, 0.75, 0.5 and 0.25, respectively, because the ratio
252 of HONO to NO_x is very sensitive to the water content in the soil. For this estimation, monthly
253 soil NO_x emissions were acquired from the MEGAN v2.10 model.

254 The HONO emission rate from soil was estimated at 0.06 Mg s⁻¹ for South Korea,
255 accounting for ~16% of the total HONO emission rate in South Korea (refer to Table 3). The
256 spatial distributions of emission are presented in Fig. 1e and 1f. The impact of HONO soil
257 emissions (see SOIL in Table 2) was examined in the EXP4 simulation.

258 **2.3.5 Heterogeneous reaction of NO₂ on atmospheric aerosol surfaces (HET_A)**

259 In the EXP5 simulation, we added the heterogeneous reaction of NO₂ on the surface of
260 atmospheric aerosols via reaction (R5) (see HET_A in Table 2), which has been reported to be
261 a possible pathway for HONO formation (Han et al., 2017; Lu et al., 2018; Reisinger, 2000;
262 Svensson et al., 1987; Wiesen et al., 1995).



264 We found a similar diurnal pattern of the concentration ratio of HONO/NO₂ to the
265 HONO mixing ratio at the Olympic Park station. This indicates that the conversion of NO₂ to
266 HONO via reaction (R5) may be a main process for HONO formation (Fig. S2Fig-S4). The
267 HONO/NO₂ ratios at the Olympic Park station in Seoul ranged from 1.9% to 6.8% during the
268 KORUS-AQ campaign, which is also comparable to those observed in Taichung, Taiwan, and
269 Shanghai, China (Hao et al., 2020; Tong et al., 2015).

270 The current AERO6 module in the CMAQv5.2.1 model already considers reaction (R5)
 271 but does not take into account ‘photo-enhancement’. However, several previous studies
 272 suggested that the photo-enhanced reactions should produce more HONO molecules during the
 273 daytime (Colussi et al., 2013; Czader et al., 2012; Fu et al., 2019; Levy et al., 2014; Li et al.,
 274 2010; Sarwar et al., 2008). The potential photo-enhancement of the reaction (R5) was taken
 275 into account by making k_{aerosol} dependent on the magnitude of light intensity:

$$276 \quad k_{\text{aerosol}} = \frac{1}{4} \times v_{\text{NO}_2} \times \frac{S_{\text{aero}}}{V} \times \gamma_{\text{a,NO}_2} \quad (\text{Eq.1})$$

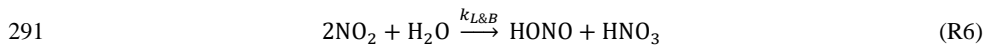
$$277 \quad \gamma_{\text{a,NO}_2} = 8.0 \times 10^{-6} \quad (\text{nighttime})$$

$$278 \quad \gamma_{\text{a,NO}_2} = 1.3 \times 10^{-4} \times \left(\frac{\text{light intensity}}{900} \right) \quad (\text{daytime})$$

279 where, v_{NO_2} , $\frac{S_{\text{aero}}}{V}$, and $\gamma_{\text{a,NO}_2}$ represent the mean molecular velocity of NO_2 ($\text{m} \cdot \text{s}^{-1}$), the
 280 aerosol surface density ($\text{m}^2 \cdot \text{m}^{-3}$), and the NO_2 uptake coefficient on the surface of
 281 atmospheric aerosols, respectively. The values of $\gamma_{\text{a,NO}_2}$ were finally selected from the
 282 sensitivity tests. It should also be noted in Table 1 that the CMAQ v5.2.1 model simply uses a
 283 fixed reaction constant ($= 10^{-4} \times \frac{S_{\text{aero}}}{V}$) for this heterogeneous reaction.

284 **2.3.6 Heterogeneous reactions of NO_2 on tree leaf and building surfaces (HET_L and** 285 **HET_BD)**

286 The heterogeneous reaction of NO_2 can also take place on the ground surfaces (e.g.,
 287 tree leaves and buildings). Several studies have reported that heterogeneous reactions on the
 288 surfaces of tree leaves and buildings via reaction (R6) can contribute to the HONO mixing
 289 ratios in the atmosphere (An et al., 2013; Hou et al., 2016; Karamchandani et al., 2015; Zhang
 290 et al., 2016). Therefore, we also considered these photo-enhanced heterogeneous NO_2 reactions.



292 In this study, $k_{\text{L\&B}}$ was calculated using equation (2), with a modification of the equation:

293
$$k_{L\&B} = \frac{1}{8} \times v_{NO_2} \times \gamma_{g,NO_2} \times \left(\frac{S_{g,building}}{V} + \frac{S_{g,leaf}}{V} \right) \quad (\text{Eq.2})$$

294
$$\gamma_{g,NO_2} = 5.0 \times 10^{-7} \quad (\text{nighttime})$$

295
$$\gamma_{g,NO_2} = 5.8 \times 10^{-6} \times \left(\frac{\text{light intensity}}{900} \right) \quad (\text{daytime})$$

296 where γ_{g,NO_2} is the NO₂ uptake coefficient on the ground surfaces. These values are also
 297 selected from sensitivity tests. Here, $\frac{S_{g,building}}{V}$ represents the ratios of the building surface
 298 area to the volume, which were calculated from equation (3):

299
$$\frac{S_{g,building}}{V} = \text{PURB} \times \frac{0.3 \frac{m^2}{m^3}}{100\%} \quad (\text{Eq.3})$$

300 where, PURB represents the percentage of building area with a maximum value of 0.3
 301 (Kurtenbach et al. 2001). For vegetation areas, $\frac{S_{g,leaf}}{V}$ (the ratio of the leaf surface to volume)
 302 was estimated based on leaf area index (LAI) information, along with equation (4) proposed
 303 by Sarwar et al. (2008):

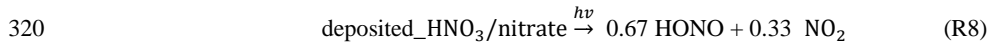
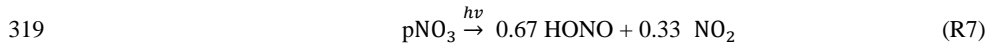
304
$$\frac{S_{g,leaf}}{V} = \frac{2 \times \text{LAI}}{H} \quad (\text{Eq.4})$$

305 where, H represents the height of the first layer of the model simulation (Sarwar et al., 2008;
 306 Yuan et al., 2011; Zhang et al., 2012). The LAI was obtained from improved Moderate
 307 Resolution Spectroradiometer (MODIS) land use data (Yuan et al., 2011).

308 2.3.7. Photolysis reactions (RENO_x)

309 Several measurement studies have reported that the photolytic dissociation of
 310 particulate nitrate (pNO₃) in the atmosphere (R7) may be able to explain the high HONO
 311 mixing ratios observed during the daytime (Romer et al., 2018; Ye et al., 2017). Other studies
 312 suggested that the photolysis reactions of HNO₃ and nitrate deposited on tree canopies and
 313 artificial surfaces (R8) can also be significant sources of daytime HONO, particularly in rural
 314 areas (Ye et al., 2016; Zhou et al., 2011). All these heterogeneous reactions from N(V) to N(III)
 315 or N(IV) are called atmospheric ‘renoxification’. Some studies have also reported that these

316 types of reduction reactions actually take place in the snow (Chen et al., 2019). In order to
317 better estimate the daytime mixing ratios of HONO in the atmosphere, reactions (R7) and (R8)
318 were included in the EXP8 simulation (see RENO_x in Table 2).



321 In the EXP8 simulation, we chose equations for both the photolysis rate constant of particulate
322 NO₃⁻ (denoted by J_{pNO₃}) and the photolysis rate constant of HNO₃/nitrate deposited on
323 surfaces (denoted by J_{D_HNO₃/nitrate}), following the methods proposed by Zhang et al. (2022),
324 and Fu et al. (2019). These equations are presented below:

$$325 \quad J_{\text{pNO}_3} = 118 \times J_{\text{HNO}_3} \quad (\text{Eq.5})$$

$$326 \quad J_{\text{D_HNO}_3/\text{nitrate}} = 48 \times J_{\text{HNO}_3} \quad (\text{Eq.6})$$

327 where, J_{HNO₃} is the reaction rate constant of gaseous HNO₃ photo-dissociation, which is
328 calculated by the photolysis rate preprocessor module (JPROC) in the CMAQ model.

329 **3. Results and Discussions**

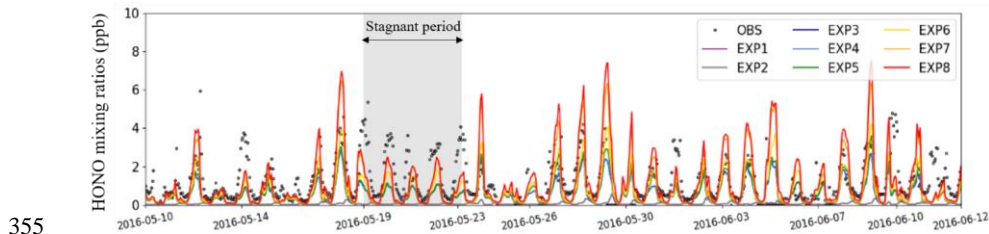
330 In this section, we first evaluated the performances of the modified CMAQ models in
331 terms of HONO mixing ratios by comparing the model outputs with ground-based observations
332 from the Olympic Park station in South Korea. We then carried out sensitivity tests to estimate
333 the contributions of the various atmospheric HONO processes to atmospheric HONO mixing
334 ratios.

335 **3.1 Observed vs Simulated HONO mixing ratios**

336 Figure 2 presents the hourly variations of the HONO mixing ratios at the Olympic Park
337 station. Observations are marked with open black circles, and colored lines represent HONO
338 mixing ratios calculated from the 8 EXP simulations. When HONO sources were added

339 sequentially to the experiments, the HONO mixing ratios averaged over the entire simulation
340 period increased from 0.06 ppb (EXP1 simulation) to 1.18 ppb (EXP8 simulation). The
341 averaged HONO mixing ratios in the EXP8 simulation, which took into account all the HONO
342 processes, were almost comparable to those observed from the ground (1.35 ppb of HONO).

343 The CMAQ-simulated HONO mixing ratios were particularly underestimated from 19
344 May to 23 May, 2016 (refer to gray-shadow period in Fig. 2). This period was characterized
345 by low wind speeds and poor mixing within planetary boundary layer height (PBLH), which
346 can lead to the accumulation of air pollutants (Crawford et al., 2021). On the other hand, the
347 WRF model has a strong tendency to produce higher wind speeds than the actual ones, which
348 may lead to underestimation of air pollutant concentrations (Jo et al., 2017). In particular, the
349 modeled wind speed during the stagnant period is overestimated by 36.3% compared to the
350 observed wind speed, which is significantly higher than the overestimation of 23.4% for the
351 entire KORUS-AQ period. Therefore, the underestimation of the HONO mixing ratios may be
352 caused by the overestimation of the wind speed on a given period. Despite all the discrepancies,
353 the HONO mixing ratios agree relatively well with the observed HONO mixing ratios during
354 the period of the KORUS-AQ campaign.



356 **Figure 2.** Hourly variations of the HONO mixing ratios (unit: ppb) at the Olympic Park station
357 in Seoul. The observations are marked with black circles and the colored lines represent the
358 HONO mixing ratios obtained from the 8 experimental simulations.

359 Figure 3 shows the diurnal variations of averaged HONO mixing ratios estimated from
360 the 8 EXP simulations, together with HONO observations at the Olympic Park station. For the

361 analysis of Fig. 3, daytime and nighttime are defined as 06:00–18:00 and 18:00–06:00 local
362 standard time, respectively. The EXP1 simulation showed slightly elevated HONO mixing
363 ratios during the daytime (purple line in Fig. 3) due to the net production of HONO in the gas
364 phase. The peak mixing ratio of the simulated HONO is ~0.14 ppb, which is significantly lower
365 than the observed mixing ratio. The large differences between EXP1 results and observations
366 suggest that there should be more unaccountable sources of HONO, which should be further
367 taken into account in our model simulations.

368 In the EXP2 simulation, HONO emissions from biomass burning were added. Several
369 studies have reported that direct and indirect sources emitted from biomass burning events
370 could contribute to the primary/secondary HONO formation (Jiang et al., 2023; Wang et al.,
371 2021; Gen et al., 2021). However, the addition of these biomass burning emissions resulted in
372 nearly negligible impact on the HONO mixing ratios, because no major biomass-burning
373 events occurred in South Korea during the period of the KORUS-AQ campaign (refer to Fig.
374 1b). Thus, there are minimal differences between the EXP1 and EXP2 simulations (i.e.,
375 between the purple and grey lines in Fig. 3).

376 EXP3 simulation was then carried out to examine the impact of traffic sources (TRAF)
377 on HONO mixing ratios. The average HONO mixing ratio increases to ~0.55 ppb. As
378 previously discussed in Fig. 1c, HONO emissions from traffic sources can be significant,
379 particularly in the Seoul Metropolitan Area. However, simulated levels of HONO are still much
380 lower than observed levels of HONO.

381 HONO emissions from soil (SOIL) were further included in the EXP4 simulation. As
382 discussed previously, several studies have reported that the consideration of soil emissions can
383 lead to large increases in atmospheric HONO mixing ratios, particularly in East Asia (Fig. 1e).
384 However, it was found that almost no significant changes had occurred in South Korea. This is

385 because the low soil NO_x levels in the South Korea are linked to several factors: (i)
386 geographical feature mainly covered by forest and mountains areas; (ii) use of low nitrogen
387 fertilizers; (iii) the reduced availability of nitrogen in the soil caused by acidic deposition; and
388 (iv) the relatively high soil water content (SWC) over the Korean Peninsula (An et al., 2023;
389 Kim et al., 2008).

390 In the EXP5 simulation, the heterogeneous reactions of NO₂ on the surfaces of
391 atmospheric aerosols (HET_A) were further taken into account. The addition of these reactions
392 was found to have only minor effect on the HONO mixing ratios, because γ_{a,NO_2} used in Eq.
393 (1) is too small to enhance the HONO mixing ratios in reaction (R5). In our study, the
394 heterogeneous reactions on the surface of atmospheric aerosols contribute only ~0.06 ppb. The
395 heterogeneous reactions can be potentially important in more polluted regions where larger
396 aerosol surface areas are available (Zhang et al., 2019).

397 On the contrary, the HONO mixing ratios can be greatly enhanced by NO₂ to HONO
398 conversions on the surfaces of the tree leaves and buildings. These two processes were
399 implemented in the EXP6 and EXP7 simulations (HET_L and HET_BD). In these two cases,
400 there were significant increases in the HONO mixing ratios, particularly during the nighttime
401 (i.e., on average, increases of 0.23 and 0.55 ppb in the HONO mixing ratios were found in the
402 EXP6 and EXP7 simulations, respectively).

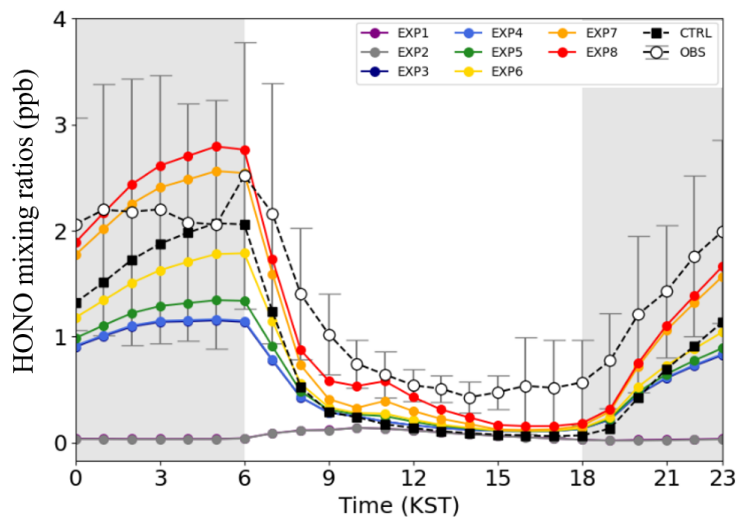
403 Finally, the photolytic renoxification of nitrate was added to the EXP8 simulation. In
404 this EXP8 simulation, the averaged HONO mixing ratios increased by 0.11 ppb. The
405 enhancement in the HONO mixing ratios was particularly large in the early morning (an
406 increase of ~0.23 ppb was found at 6 a.m.). Overall, the EXP8 simulation produced the best
407 HONO mixing ratios (averaged value of 1.18 ppb), compared to observed HONO mixing ratio
408 (1.35 ppb). Also, the estimated HONO mixing ratios were more comparable than those in the

409 CTRL (original CMAQ v5.2.1) model simulation (represented by black squares in Fig. 3).
410 Again, it is noted that our simulations incorporated 'new HONO processes' such as: i) the
411 photo-enhanced HONO production pathway through (R5) and (R6); ii) daytime HONO
412 production from renoxification reactions through (R7) and (R8); and iii) HONO emissions
413 (refer to Table 1).

414 In addition to the graphical comparison in Fig. 3, several statistical metrics were also
415 calculated to evaluate the performances of the 8 EXP and CTRL simulations in Table 4.
416 Significant improvements were found when the HONO processes were sequentially added
417 from the EXP1 to the EXP8 simulations. For example, the index of agreement (IOA) increases
418 from 0.44 to 0.76, and the mean bias (MB) decreases drastically from -1.29 ppb to -0.17 ppb
419 from the EXP1 to the EXP8. In particular, the EXP8 simulation showed the best performance,
420 compared to the CTRL simulation during the daytime. For example, the IOA during the
421 daytime increased from 0.59 to 0.68, while the MB decreases from -0.57 to -0.34, respectively.
422 The root mean square error (RMSE) also decreased from 0.80 to 0.70 during the daytime.

423 Although the EXP8 simulation showed a notable enhancement in HONO production,
424 the HONO mixing ratios were still underestimated during the daytime. Such underestimation
425 of HONO mixing ratios during the daytime could be attributed to stronger HONO photo-
426 dissociation than in real situations. This is possibly due to failure in predicting cloud shades
427 fractions in meteorological modeling and/or due to additional sources that were not considered
428 in this study (e.g., acid displacement for HNO₃ and HCl, [nitrate and Fe\(II\) in iron-organic](#)
429 [complex under irradiation, and renoxification of nitrate in presence of carbonaceous aerosols](#))
430 ([Gen et al., 2021](#); [VandenBoer et al., 2013](#); [Wang et al., 2021](#)). This certainly indicates that
431 additional work is needed to further investigate HONO formation and removal during the
432 daytime.

433



434
 435 **Figure 3.** Diurnal variations of HONO mixing ratios (unit: ppb) at the Olympic Park station
 436 averaged over the period of the KORUS-AQ campaign. Error bars and grey-shaded areas
 437 indicate one standard deviation and nighttime (18 – 06 LST, Local Standard Time),
 438 respectively.

439
 440
 441
 442
 443
 444
 445 **Table 4.** Statistical analysis with modeled and observed HONO mixing ratios at the Olympic
 446 Park station, Seoul, Korea.

Experiment	Observed mean (ppb)	Modeled mean (ppb)	RMSE (ppb)	MB (ppb)	IOA
CTRL	1.35	0.78	1.06	-0.57	0.75
EXP1	1.35	0.06	1.68	-1.29	0.44
EXP2	1.35	0.06	1.68	-1.29	0.44
EXP3	1.35	0.55	1.15	-0.79	0.63
EXP4	1.35	0.56	1.15	-0.79	0.64
EXP5	1.35	0.61	1.12	-0.73	0.66
EXP6	1.35	0.75	1.02	-0.60	0.72
EXP7	1.35	1.07	1.05	-0.28	0.77
EXP8	1.35	1.18	1.12	-0.17	0.76

447
 448 **3.2 Relative contribution of HONO sources**

449 Individual HONO processes affect the HONO mixing ratios in different ways. Figure
450 4 summarizes the relative contribution of HONO processes to the HONO mixing ratios. During
451 the daytime, both GAS and RENO_x contribute significantly to the production of atmospheric
452 HONO molecules. In particular, the contribution of these two processes is the largest between
453 10:00 and 16:00 local time, when sunlight is strong. These two processes account for 29.1%
454 and 29.8% of the daytime HONO production, respectively, but are almost negligible during the
455 nighttime.

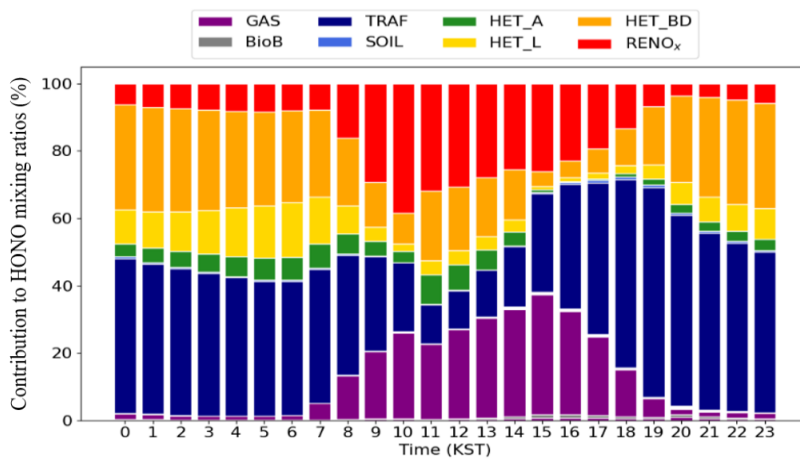
456 During the nighttime, TRAF (denoted by navy color in Fig. 4) contributes the large
457 portion of 47.2% of the total HONO production. However, there is a possibility that TRAF
458 might have been somewhat overestimated during the nighttime since we applied constant
459 diurnal anthropogenic NO_x emissions, including those from traffic source. In turn, HET_BD
460 and HET_L exhibit substantial contributions of 28.5% and 10.6%, respectively during the
461 nighttime. The contributions of other processes such as biomass burning (BioB) and
462 heterogeneous reactions on atmospheric aerosols (HET_A) are minimal. HET_A contributes
463 only 4.3% during the nighttime. ~~Its contribution increases to 4.2% during the daytime.~~ In terms
464 of the average 24-hour contribution, TRAF (41.4%), HET_BD (27.1%), and HET_L (11.1%)
465 are the large sources of atmospheric HONO at the Olympic Park station.

466 Using the same approach, we analyzed the HONO source contributions across South
467 Korea during the period of the KORUS-AQ campaign. As shown in Fig. 5f and 5c, HET_L
468 and TRAF were modeled to have the largest impacts on HONO production, contributing 0.15
469 ppb (41.5%) and 0.08 ppb (18.1%), respectively, across South Korea (also, refer to the
470 incremental ratio in Fig. S4S3).

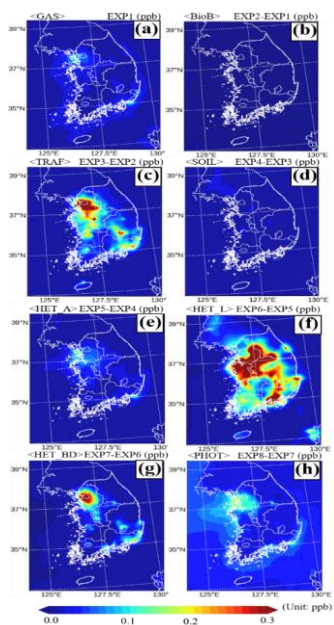
471 Fig. 6 shows the contributions of different sources to the HONO mixing ratios at 8
472 super monitoring stations. As shown in Fig. 6, each station has different characteristics in terms

473 of source contribution. In particular, the contribution of HET_L at the Daejeon is 44.4%. Also,
 474 TRAF in Bulkwang, Olympic Park, Mt.Taehwa, Ulsan, and Gwangju have large contributions
 475 of 41.2%, 41.4%, 29.3%, 29.6%, and 40.5%, respectively. As for TRAF and HET_BD, their
 476 contributions are high only in densely populated cities (refer to Fig. 5c and 5g). On the other
 477 hand, the contributions of BioB, SOIL, HET_A, and RENO_x sources were insignificant, as
 478 shown in Fig. 5b, 5d, 5e, and 5h.

479 Meanwhile, at the Bangnyung and Jeju stations, RENO_x has the largest contribution
 480 of 70.4%, and 33.2%, respectively. This is because the amounts of NO₂ and HONO from direct
 481 emissions (BioB, TRAF, and SOIL) are relatively small. The Bangnyung and Jeju stations are
 482 located on remote and less populated islands.

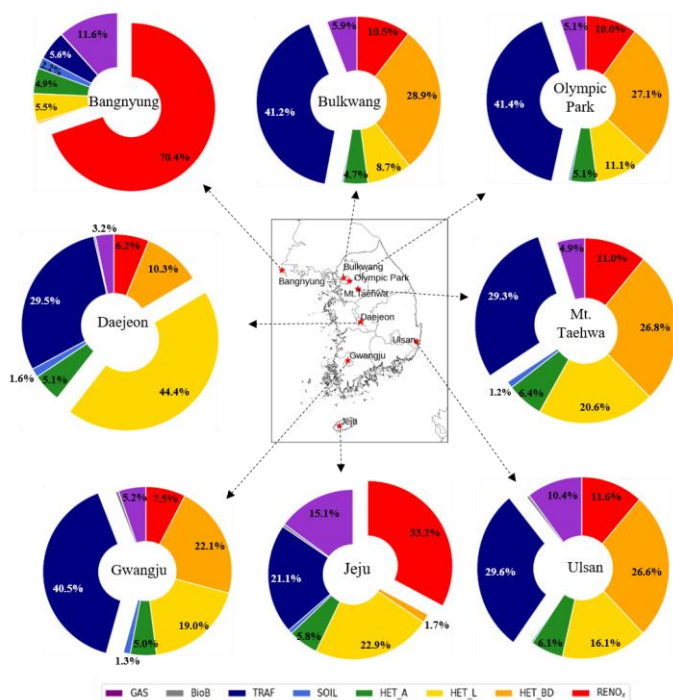


483
 484 **Figure 4.** Diurnal contributions of individual HONO processes to the HONO mixing ratios at
 485 the Olympic Park station during the period of the KORUS-AQ campaign.



486

487 **Figure 5.** Spatial impacts of (a) gas phase reactions; (b) biomass burning emissions; (c) traffic
 488 emissions and (d) soil emissions; (e) heterogeneous reactions on the aerosol surfaces, (f)
 489 heterogeneous reactions on the leaf surfaces, and (g) heterogeneous reactions on the building
 490 surfaces; and (h) renoxification on HONO mixing ratios, based on model simulations during
 491 the period of the KORUS-AQ campaign in South Korea.



492
 493 **Figure 6.** Contributions of individual processes to the average HONO mixing ratios at 8
 494 monitoring stations during the period of the KORUS-AQ campaign.

495 **3.3 Impact of HONO processes on atmospheric species**

496 **3.3.1 Impact on atmospheric species**

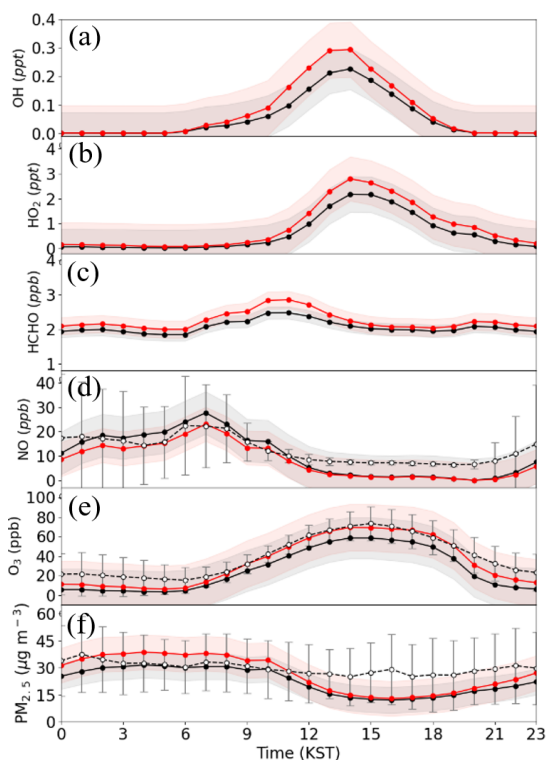
497 We also investigated the effect of HONO processes on atmospheric levels of HO_x
 498 (=OH + HO₂), HCHO, O₃, NO, and PM_{2.5} at the Olympic Park station. Figure 7 presents the
 499 diurnal concentrations of these gaseous and particulate species at the Olympic Park station.
 500 The mixing ratios of OH and HO₂ radicals in the EXP8 simulation increased by 0.02 ppt (35.2%)
 501 and 0.23 ppt (39.2%), respectively, compared to those in the CTRL simulation. This is certainly
 502 due to the enhancement in OH levels due to HONO photo-dissociation, and then HO₂ levels in
 503 the HO_x cycle. As shown in Fig. 7a (and 7b), the OH (and HO₂) mixing ratios increased from
 504 0.21 ppt to 0.29 ppt (1.71 ppt to 2.28 ppt) at 1 p.m. local standard time. Subsequently, the

505 HCHO mixing ratios were also enhanced by 0.18 ppb (8.8%), due to increased VOC oxidation
506 resulting from elevated levels of OH radicals (Fig. 7c). On the contrary, the NO mixing ratios
507 in the EXP8 simulation decreased by 2.13 ppb (20.1%). This may be due to an increase in the
508 mixing ratios of HO₂ and RO₂ radicals (organic peroxy radicals) reacting with NO molecules
509 (Fig. 7d). In other words, the reduced levels of NO indicate active NO to NO₂ conversion via
510 NO+HO₂ and NO+RO₂ reactions. Such active NO to NO₂ conversion increases the mixing
511 ratio of atmospheric ozone because these two reactions are rate-determining reactions for ozone
512 production. This is presented in Fig. 7e. In Fig. 7e, the modeled ozone mixing ratios increased,
513 approaching the observed ozone mixing ratios. This is another good result showing that the
514 incorporation of atmospheric HONO processes may be able to enhance the accuracy of
515 prediction of ozone mixing ratios. More details about ozone production are discussed in Sect.
516 3.3.2.

517 Elevated levels of atmospheric O₃ and HO_x can change the rates of particulate nitrate
518 and sulfate production. In particular, the formation of particulate nitrates and sulfates can also
519 be enhanced by increasing the levels of HNO₃, N₂O₅, and H₂SO₄. In addition, the nitrate
520 concentration can also be enhanced by the HONO reaction (i.e., via NO₂ + H₂O → H⁺ + NO₃⁻
521 + HONO, as accounted for by [R5R6](#)) during the nighttime. In total, the addition of HONO
522 processes increases PM_{2.5} by 4.19 μg m⁻³ (18.6%) at the Olympic Park station. However, PM_{2.5}
523 in the EXP8 simulation was still underestimated by 3.16 μg m⁻³ at the Olympic Park station, as
524 shown in Fig. 7f. There are several potential reasons for this underestimation, such as the
525 underestimation of secondary organic aerosol (SOA) formation (e.g., Murphy et al., 2017).
526 This issue may require further investigation in the future.

527 Figure [S43](#) presents similar results for 320 AIR KOREA monitoring stations in South
528 Korea. The impacts of HONO processes on atmospheric levels of OH, HO₂, O₃, and PM_{2.5} are
529 also presented in Fig. [S54](#). Overall, we found that incorporating HONO chemistry into the

530 modeling system tends to enhance the mixing ratios of HO_x, which in turn increases the mixing
531 ratios of O₃ and PM_{2.5}.



532
533 **Figure 7.** Diurnal variations of the mixing ratios of (a) OH, (b) HO₂, (c) HCHO, (d) NO, (e)
534 O₃, and (f) PM_{2.5} (black lines represent the mixing ratios from the CTRL simulation and the
535 red lines represent those from the EXP8 simulation) and observations (marked with white open
536 cycles) at the Olympic Park station during the period of the KORUS-AQ campaign. Shaded
537 areas represent one standard deviation for each simulation.

538 3.3.2 Impact on net ozone production

539 The ozone mixing ratio is determined by the balance between ozone formation and
540 destruction in the atmosphere. To better understand the impacts of HONO chemistry on ozone
541 production, we quantitatively analyze the rate of net ozone production ($P(O_3)$). The $P(O_3)$ is
542 defined by equation (7):

543
$$P(O_3) = F(O_3) - D(O_3) \quad (\text{Eq. 7})$$

544 where, $F(O_3)$ and $D(O_3)$ represent the rate of ozone formation and destruction, respectively.
 545 $F(O_3)$ and $D(O_3)$ can be calculated from equations (8) and (9), respectively (Mazzuca et al.,
 546 2016; Song et al., 2003).

547
$$F(O_3) = k_{HO_2+NO}[HO_2][NO] + k_{RO_2+NO}[RO_2][NO] \quad (\text{Eq. 8})$$

548
$$D(O_3) = k_{NO_2+OH}[NO_2][OH] + k_{O_3+VOC}[O_3][VOC] + k_{O(1D)+H_2O}[O(^1D)][H_2O]$$

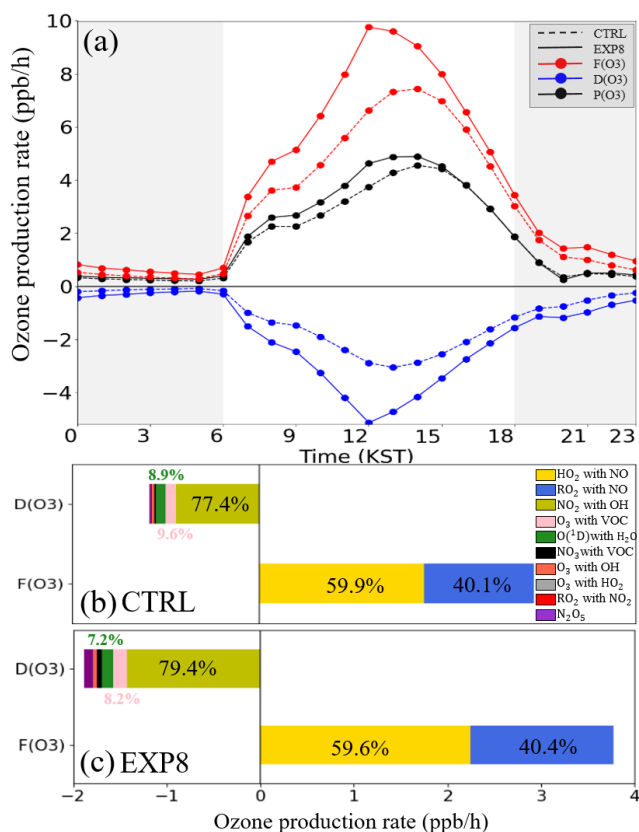
 549
$$+ k_{O_3+OH}[O_3][OH] + k_{O_3+HO_2}[O_3][HO_2] + k_{RO_2+NO_2}[RO_2][NO_2]$$

 550
$$+ 2k_{NO_3+VOC}[NO_3][VOC] + 3k_{het}[N_2O_5] \quad (\text{Eq. 9})$$

551 where k_i represents the reaction rate constants for each reaction i . In particular, k_{het} denotes the
 552 heterogeneous reaction rate constants of N_2O_5 radicals.

553 Figure 8a shows the diurnal variations of $F(O_3)$, $D(O_3)$, and $P(O_3)$ from the CTRL and
 554 EXP8 simulations. Including HONO processes in the EXP8 simulation resulted in an average
 555 $P(O_3)$ that was 10.6% higher than in the CTRL simulation. This is the primary reason for the
 556 ozone enhancement in Fig. 7e.

557 Figures 8b and 8c provide more details about the budget of ozone production. The
 558 main increase in $F(O_3)$ occurred through the reactions of $HO_2 + NO$ and $RO_2 + NO$. On the
 559 other hand, the increase in $D(O_3)$ was mainly controlled by the $NO_2 + OH$ reaction at the
 560 Olympic Park station. The increases in the $HO_2 + NO$ and $RO_2 + NO$ reaction rate exceeded
 561 the increases in the reaction rate of $NO_2 + OH$, leading to the net positive ozone production
 562 (i.e., positive $P(O_3)$) shown in Fig. 8a.



563
 564 **Figure 8.** Diurnal variations of (a) net ozone production rate (P(O₃); black line), ozone
 565 formation rate (F(O₃); red line), and ozone loss rate (D(O₃); blue line). The dashed and solid
 566 lines represent the CTRL and EXP8 simulations, respectively. Cumulative bar chart for D(O₃)
 567 and F(O₃) in case of (b) CTRL and (c) EXP8 simulations at the Olympic Park station during
 568 the period of the KORUS-AQ campaign.

569 **4. Conclusions**

570 In this study, we successfully incorporated the following HONO processes into the
 571 CMAQ modeling framework to enhance the accuracy in the predictions of HONO mixing
 572 ratios: i) gas-phase HONO reactions; ii) HONO emission from biomass burning; iii) HONO
 573 emission from traffic and soil; iv) photo-induced heterogeneous reactions on the surfaces of
 574 atmospheric aerosols, tree leaves, and buildings; and v) photolysis reactions of particulate

575 nitrate and deposited HNO₃/nitrate. The analysis showed that the incorporation of HONO
576 processes into the CMAQ model framework increased the average HONO mixing ratios from
577 0.78 ppb to 1.18 ppb compared to the CTRL simulation. Average mixing ratios of HONO and
578 its diurnal patterns became much more comparable to observations, with large improvements
579 in statistical parameters. Especially during the daytime, IOA increased from 0.59 to 0.68, while
580 the MB decreased from -0.57 ppb to -0.34 ppb, and RMSE dropped from 0.80 ppb to 0.70 ppb,
581 as HONO processes were fully incorporated into the CMAQ model.

582 Several findings also emerged from the sensitivity simulations. First, each HONO process
583 had a different effect on the HONO mixing ratios during the daytime and the nighttime at the
584 Olympic Park station. For example, the GAS (29.1%) and RENO_x processes (29.8%) had major
585 contributions to the mixing ratios of HONO during the daytime, while the TRAF (47.2%) and
586 HET_BD (28.5%) processes had large contributions to the mixing ratios of HONO during the
587 nighttime. During the period of the KORUS-AQ campaign, HONO mixing ratios estimated at
588 the Olympic Park station were enhanced by an average of 41.4% (TRAF), 27.1% (HET_BD),
589 and 11.1% (HET_L).

590 In the experimental simulation including all the HONO processes (i.e., EXP8 simulation),
591 the mixing ratios of OH, HO₂, HCHO, O₃, and PM_{2.5} at the Olympic Park station increased by
592 0.02 ppt (35.2%), 0.23 ppt (39.2%), 0.18 ppb (8.8%), 7.86 ppb (30.8%), and 4.19 μg m⁻³
593 (18.6%), respectively, compared to those from the CTRL simulation. The net ozone production
594 rate was enhanced by 0.19 ppb h⁻¹ (10.6%) with the EXP8 simulation. This increases in P(O₃)
595 were caused mainly by the increased reaction rates of HO₂ + NO.

596 In this study, we improved our understanding of atmospheric HONO processes in
597 South Korea. Nevertheless, we believe that both further field studies and modeling
598 investigations are necessary for many remaining HONO-related issues such as NO₂ uptake

599 coefficient, possible missing HONO sources, and daytime photochemical reaction pathways of
600 HONO. Such studies will also help to further improve the performances of current CTMs.

601 For example, the Airborne and Satellite Investigation of Asian Air Quality (ASIA-AQ)
602 field campaign organized by the National Institute of Environmental Research (NIER) in Korea
603 and the National Aeronautics and Space Administration (NASA) in the U.S. is planned in 2024
604 in South Korea. In this campaign, the HONO mixing ratios are scheduled to be measured in
605 the aircraft and at the ground station. This joint campaign is thus expected to provide a valuable
606 opportunity to expand our knowledge on atmospheric HONO processes and HONO photo-
607 chemistry.

608

609 **Code and data availability**

610 After user registration, the WRF model 3.8.1 (<https://www2.mmm.ucar.edu/wrf/users>, last access: 11
611 April 2024) and CMAQ v5.2.1 (<https://doi.org/10.5281/zenodo.1212601>, last access: 11 April 2024)
612 are available from web page. The observation data we used can be accessed at [https://www-
613 air.larc.nasa.gov/cgi-bin/ArcView/korusaq?GROUND-NIER-OLYMPIC-PARK=1](https://www-air.larc.nasa.gov/cgi-bin/ArcView/korusaq?GROUND-NIER-OLYMPIC-PARK=1) (last access: 11
614 April 2024).

615

616 **Author contribution**

617 **KYK** designed experiments and led manuscript writing and conceptualization. **KMH** and **CHS**
618 supervised this project and contributed to experimental design and manuscript writing. **HJL, RB,**
619 **JHY, GY,** and **BYK** performed research development. **JM** contributed to editing and writing review.
620 **JHW** and **SJC** provided useful datasets.

621

622 **Competing interests**

623 Chul H. Song is a member of the editorial board of *Atmospheric Chemistry and Physics*. The authors
624 declare that they have no conflict of interest.

625

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1182 *Supplementary information for*

1183 **An investigation into atmospheric nitrous acid (HONO)**
1184 **processes in South Korea**

1185

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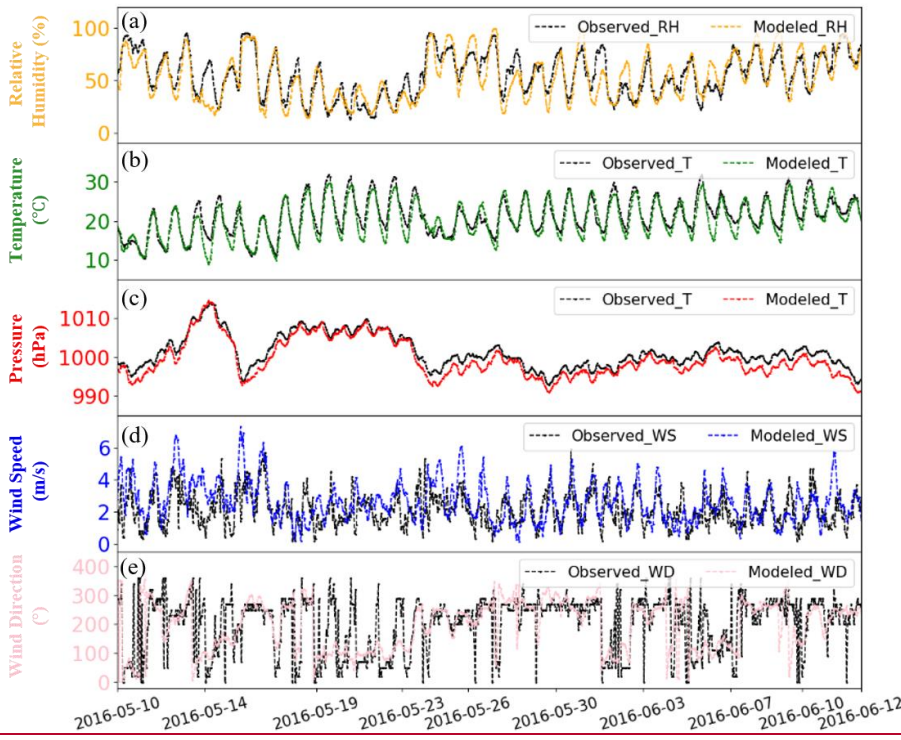
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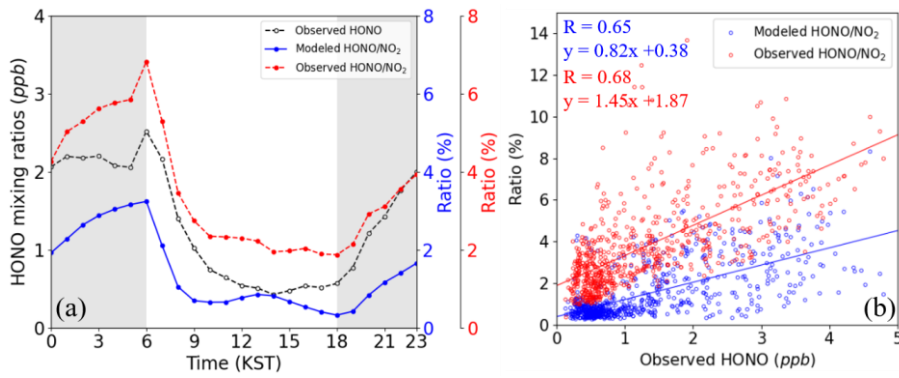
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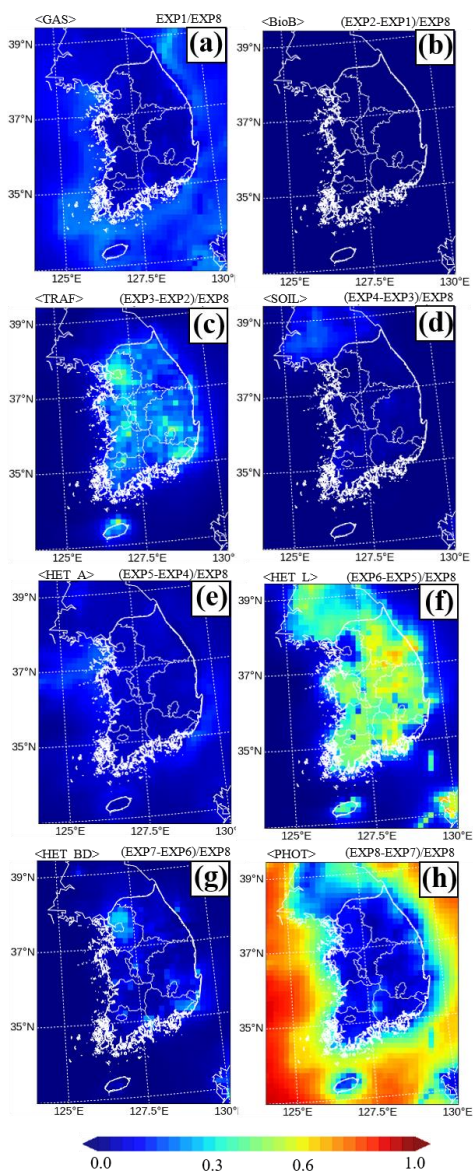
Fig. S1. Temporal variations of (a) relative humidity (RH); (b) temperature (T); (c) pressure; (d) wind speed (WS) at 10m above the surface; and (e) wind direction (WD) at 10m above the surface during period of the KORUS-AQ campaign.

서식 있음: 캡션, 왼쪽



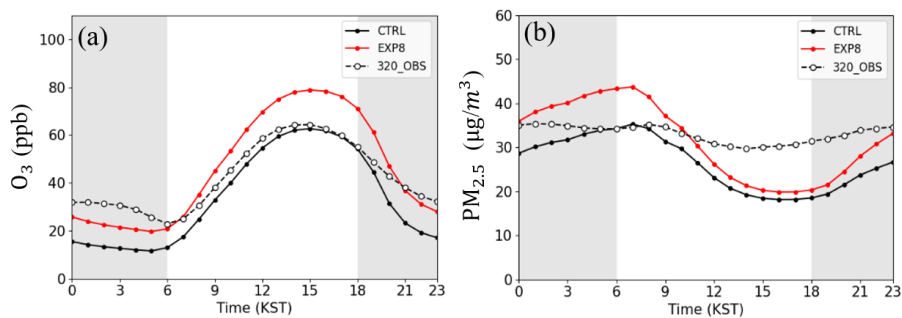
1207

1208 **Fig. S1S2.** (a) Diurnal variations of observed HONO (black open circles, left y-axis), modeled
 1209 HONO to NO₂ ratio (blue line, right y-axis), and observed HONO to NO₂ ratio (red line, right
 1210 y-axis) and (b) their scatter-plots between the observed HONO and the modeled ratio of HONO
 1211 to NO₂ (blue circles), and observed HONO to NO₂ (red circles) at the Olympic Park station
 1212 during the period of the KORUS-AQ campaign.

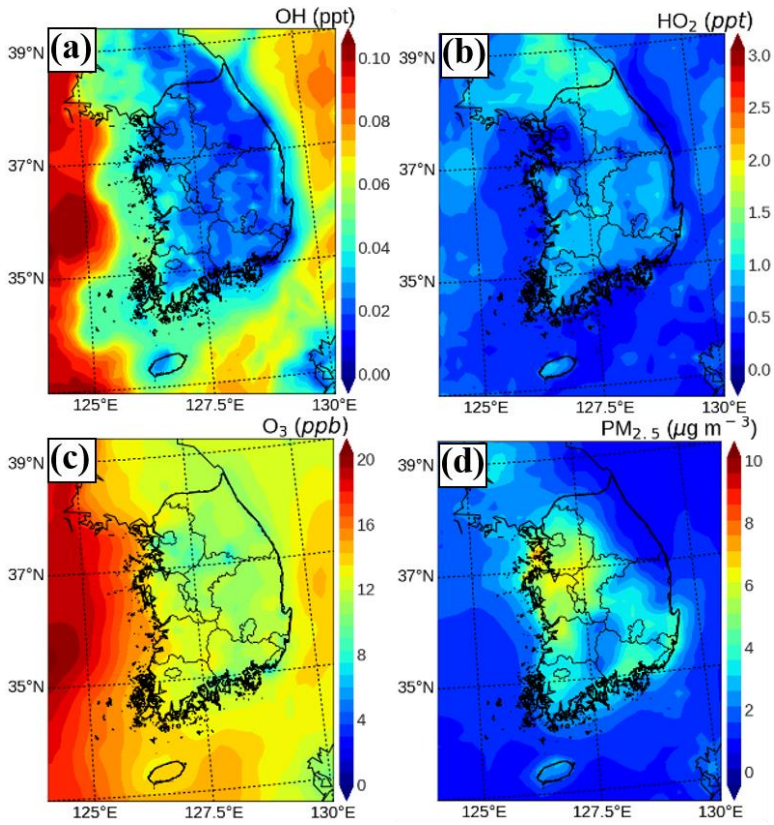


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 1214 **Fig. S2S3.** Incremental ratio of (a) gas phase reactions; (b) biomass burning emissions; (c)
 1215 traffic emissions and (d) soil emissions; (e) heterogeneous reactions on the aerosol surfaces, (f)
 1216 heterogeneous reactions on the leaf surfaces, and (g) heterogeneous reactions on the building
 1217 surfaces; and (h) renoxification on HONO mixing ratios (unit: dimensionless).

1218



1219 **Fig. S4S4.** Comparison of diurnal variations of the mixing ratios of (a) O₃, and (b) PM_{2.5}. Both
1220 are averaged for 320 AIR KOREA monitoring stations during the period of the KORUS-AQ
1221 campaign. The black open circles, black lines, and red lines represent observed values and
1222 values from the CTRL and EXP8 simulations, respectively.



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Fig. S555. Spatial distributions of the differences levels of (a) HONO, (b) OH, (c) HO₂, and (d) PM_{2.5} between the EXP8 and CTRL simulations in South Korea during the period of the KORUS-AQ campaign.

1227 **Table S1.** Detection limits and uncertainties of instruments for observed HONO, NO₂, O₃, and
1228 PM_{2.5} at Olympic Park station, Korea.

Species	Instruments	Detection limit	Uncertainty	Time resolution
HONO	Monitor for Aerosols and Gases in Ambient Air (MARGA, model ADI 2080)	0.02ppbv	±20%	1 hour
NO ₂	Ecotech gas sensor, EC8941	0.5ppbv	±10%	1 hour
O ₃	Ecotech gas sensor, EC9810	0.5ppbv	±5%	1 hour
PM _{2.5}	Thermo Fisher Scientific, FH62C14	4µg/m ³	±10%	1 hour

1229

1230 **Table S2.** Statistical analysis of modeled and observed meteorological parameters at the
 1231 Olympic Park station during the period of the KORUS-AQ campaign.

Parameter	Observed mean	Modeled mean	R	RMSE	MB	IOA
RH (%)	<u>55.81</u>	<u>53.33</u>	<u>0.85</u>	<u>11.95</u>	<u>-2.48</u>	<u>0.92</u>
	55.50	55.60	0.87	11.06	0.09	0.93
T (°C)	<u>21.27</u>	<u>20.28</u>	<u>0.93</u>	<u>1.96</u>	<u>-0.99</u>	<u>0.96</u>
	22.38	21.10	0.94	1.84	-1.28	0.94
Pressure (hPa)	<u>1001.38</u>	<u>999.62</u>	<u>0.98</u>	<u>2.01</u>	<u>-1.77</u>	<u>0.95</u>
	1000.35	999.75	0.97	1.14	-0.60	0.98
WS (m s ⁻¹)	<u>2.14</u>	<u>2.65</u>	<u>0.47</u>	<u>1.30</u>	<u>0.51</u>	<u>0.66</u>
	2.08	2.46	0.45	1.17	0.39	0.65
WD (°)	<u>202.71</u>	<u>196.01</u>	<u>0.53</u>	<u>88.87</u>	<u>-6.70</u>	<u>0.75</u>
	205.40	202.18	0.53	83.49	-3.24	0.74

1232