Supplement of

# Importance of ice elasticity in simulating tide-induced grounding line variations along prograde bed slopes 

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## S.1. Reference measurements

Grounding zone measurements were performed using pairs of DInSAR interferograms at high and low tide; bedrock slope and ice thickness were calculated using Bed Machine Antarctica, and ice flow speed was determined using MEaSUREs (version 2). The measurements, performed along $\sim 20 \mathrm{~km}$-long flow lines (see Figure 1), are provided in Table S1.

Table S1. Grounding zone (GZ), ice thicknesses, bed slope, and ice flow speed measurements over Totten (TOT), Moscow University (MU), and Rennick (REN) glaciers.

| Glacier | Flow line | $\begin{gathered} \mathrm{GZ}, \\ \mathrm{~km} \end{gathered}$ | Bed slope $\boldsymbol{\alpha}$, \% | Ice thickness $\boldsymbol{H}$, km | Ice flow speed $\boldsymbol{v}$, m/yr | Profile type |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| MU | 0 | 0.6 | 5.0 | 1.8 | 322 | Main |
| MU | 1 | 1.3 | 2.8 | 1.8 | 319 | Main |
| MU | 2 | 1.6 | 2.4 | 1.8 | 316 | Main |
| MU | 3 | 1.7 | 2.0 | 1.7 | 301 | Main |
| MU | 4 | 1.6 | 2.8 | 1.7 | 291 | Main |
| MU | 5 | 1.6 | 2.9 | 1.7 | 291 | Main |
| MU | 6 | 1.4 | 2.5 | 1.7 | 291 | Main |
| MU | 7 | 1.6 | 2.2 | 1.7 | 301 | Main |
| MU | 8 | 1.6 | 2.1 | 1.7 | 310 | Main |
| MU | 9 | 1.4 | 2.7 | 1.7 | 317 | Main |
| MU | 10 | 0.8 | 5.0 | 1.8 | 325 | Main |
| MU | 11 | 0.9 | 4.9 | 1.8 | 327 | Main |
| MU | 12 | 1.0 | 4.9 | 1.8 | 334 | Main |
| MU | 13 | 1.3 | 3.3 | 1.9 | 339 | Main |
| MU | 14 | 1.7 | 1.7 | 1.9 | 348 | Main |
| MU | 15 | 1.8 | 1.6 | 1.9 | 361 | Main |
| MU | 16 | 2.2 | 1.0 | 2.0 | 369 | Main |
| MU | 17 | 1.9 | 1.7 | 2.1 | 369 | Main |
| MU | 18 | 1.7 | 1.5 | 2.2 | 372 | Main |
| MU | 19 | 1.6 | 2.3 | 2.2 | 378 | Main |
| MU | 20 | 1.8 | 1.4 | 2.3 | 373 | Main |
| MU | 21 | 1.9 | 1.5 | 2.4 | 373 | Main |
| MU | 22 | 2.0 | 1.5 | 2.4 | 354 | Main |
| MU | 23 | 1.9 | 1.7 | 2.4 | 347 | Main |
| MU | 24 | 1.4 | 3.1 | 2.4 | 342 | Main |
| MU | 25 | 0.7 | 4.5 | 2.4 | 327 | Main |
| MU | 26 | 0.4 | 4.9 | 2.4 | 319 | Main |
| MU | 27 | 0.4 | 5.4 | 2.4 | 307 | Extra |
| MU | 28 | 0.5 | 5.4 | 2.3 | 289 | Extra |
| MU | 29 | 1.1 | 3.9 | 2.3 | 294 | Main |
| MU | 30 | 5.0 | 0.01 | 2.4 | 327 | Main |
| MU | 31 | 5.2 | 0.02 | 2.3 | 314 | Main |
| MU | 32 | 3.9 | 0.4 | 2.3 | 292 | Main |
| MU | 33 | 3.3 | 0.3 | 2.2 | 261 | Main |
| MU | 34 | 3.3 | 0.3 | 2.2 | 243 | Main |
| MU | 35 | 3.4 | 0.3 | 2.2 | 228 | Main |
| MU | 36 | 3.6 | 0.4 | 2.2 | 207 | Main |
| MU | 37 | 3.7 | 0.3 | 2.2 | 185 | Main |
| MU | 38 | 3.8 | 0.3 | 2.1 | 168 | Main |
| MU | 39 | 9.0 | 0.1 | 2.1 | 165 | Main |
| MU | 40 | 10.2 | 0.1 | 2.0 | 158 | Main |


| MU | 41 | 11.5 | 0.1 | 1.9 | 151 | Main |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| MU | 42 | 11.8 | 0.1 | 1.8 | 155 | Main |
| MU | 43 | 12.5 | 0.01 | 1.7 | 149 | Main |
| REN | 0 | 0.8 | 5.2 | 0.7 | 113 | Extra |
| REN | 1 | 1.4 | 2.4 | 0.8 | 135 | Extra |
| REN | 2 | 1.9 | 1.5 | 0.9 | 150 | Extra |
| REN | 3 | 2.1 | 1.2 | 0.9 | 158 | Extra |
| REN | 4 | 2.2 | 1.2 | 1.0 | 171 | Main |
| REN | 5 | 2.4 | 1.0 | 1.0 | 176 | Main |
| REN | 6 | 2.4 | 0.8 | 1.0 | 183 | Main |
| REN | 7 | 2.9 | 0.6 | 1.0 | 186 | Main |
| REN | 8 | 2.8 | 0.7 | 1.1 | 187 | Main |
| REN | 9 | 3.1 | 0.6 | 1.1 | 188 | Main |
| REN | 10 | 2.8 | 0.7 | 1.1 | 187 | Main |
| REN | 11 | 3.3 | 0.3 | 1.1 | 187 | Main |
| REN | 12 | 3.2 | 0.5 | 1.2 | 185 | Main |
| REN | 13 | 2.8 | 0.5 | 1.1 | 183 | Main |
| REN | 14 | 2.3 | 1.0 | 1.2 | 178 | Main |
| REN | 15 | 1.8 | 1.5 | 1.1 | 174 | Main |
| REN | 16 | 1.9 | 1.1 | 1.1 | 169 | Main |
| REN | 17 | 2.2 | 1.0 | 1.0 | 162 | Main |
| REN | 18 | 2.2 | 1.1 | 1.0 | 152 | Main |
| TOT | 0 | 0.9 | 4.6 | 2.3 | 696 | Main |
| TOT | 1 | 1.0 | 4.1 | 2.3 | 674 | Main |
| TOT | 2 | 1.8 | 2.1 | 2.4 | 722 | Main |
| TOT | 3 | 3.5 | 0.4 | 2.4 | 754 | Main |
| TOT | 4 | 4.6 | 0.2 | 2.3 | 758 | Main |
| TOT | 5 | 6.3 | 0.06 | 2.3 | 740 | Main |
| TOT | 6 | 7.0 | 0.03 | 2.3 | 731 | Main |
| TOT | 7 | 7.5 | 0.06 | 2.2 | 709 | Main |
| TOT | 8 | 7.6 | 0.07 | 2.2 | 690 | Main |
| TOT | 9 | 6.0 | 0.03 | 2.1 | 680 | Main |
| TOT | 10 | 5.5 | 0.02 | 2.1 | 683 | Main |
| TOT | 11 | 5.6 | 0.1 | 2.0 | 735 | Main |
| TOT | 12 | 3.0 | 0.5 | 1.9 | 747 | Main |
| TOT | 13 | 2.2 | 1.0 | 1.9 | 683 | Main |
| TOT | 14 | 1.8 | 1.5 | 1.9 | 636 | Main |
| TOT | 15 | 1.7 | 1.6 | 1.8 | 547 | Main |
| TOT | 16 | 1.3 | 2.8 | 1.8 | 569 | Main |

## S.2. Mesh sensitivity analysis

Models' sensitivity to mesh size was analyzed using 200 grounding zone width values (Figure S1), obtained for different mesh sizes of the lower domain surface (from 10 m to 250 m with 10 m step) and constant mesh size of 250 m at the upper domain surface. For both models, the tests were performed for four sets of input parameters: 1) bed slope of $0.5 \%$, glacier thickness of 1 km , and ice inflow speed of $100 \mathrm{~m} /$ year (green dots in Figure S1); 2) bed slope of $5 \%$, glacier thickness of 1 km , and ice inflow speed of $100 \mathrm{~m} /$ year (red dots in Figure S1); 3) bed slope of $5 \%$, glacier thickness of 2.5 km , and ice inflow speed of $100 \mathrm{~m} /$ year (blue dots in Figure S 1 ); 4) bed slope of $5 \%$, glacier thickness of 1 km , and ice inflow speed of $800 \mathrm{~m} /$ year (black dots in Figure S 1 ).


Figure S1. Models mesh sensitivity check. Plots (a) and (c) correspond to the viscous model, and plots marked (b) and (d) represent the viscoelastic model.

## S.3. DInSAR data-inferred parameters as model inputs

Fixing 50 m and 250 m as mesh sizes at the lower and upper domain boundaries, respectively, and keeping the constant glacier domain length of 20 km , each model was executed 192 times using all possible combinations of input parameters, listed in Figure S2.

## Variables



Speed (m/yr)


Figure S2. Schematic representation of the initial parameters. All possible combinations of these variables were examined in the paper.

## S.4. Grounding zone evolution with glacier thickness

The dependence of grounding zone width $(G Z)$ on glacier thickness $(H)$ for each inflow speed and bed slope, approximated with a linear function $G Z=a \cdot H+b$, provides unique coefficients $a$ and $b$ for each model formulation, bed slope, and ice inflow speed. These $a$ and $b$ values are listed in Table S2.

Table S2. Equations of the approximating lines of all considered dependences of the grounding zone magnitude from the glacier thickness.

| Slope, \% | Inflow speed, m/year | Approximating line equation, viscous model | $\mathrm{R}^{2}$ value, viscous model | Approximating line equation, viscoelastic model | $\mathrm{R}^{2}$ value, viscoelastic model |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 5.0 | 100 | $0.119 \cdot x+1.209$ | 0.998 | $0.081 \cdot x+0.799$ | 1.000 |
|  | 350 | $0.071 \cdot x+1.265 ;$ | 0.975 | $0.080 \cdot x+0.779$ | 1.000 |
|  | 600 | $0.033 \cdot x+1.297$ | 0.914 | $0.072 \cdot x+0.779$ | 0.987 |
|  | 800 | $0.027 \cdot x+1.286$ | 0.985 | $0.060 \cdot x+0.776$ | 0.987 |
| 4.5 | 100 | $0.119 \cdot x+1.222$ | 0.999 | $0.088 \cdot x+0.904$ | 0.967 |
|  | 350 | $0.104 \cdot x+1.216$ | 0.999 | $0.076 \cdot x+0.896$ | 0.969 |
|  | 600 | $0.083 \cdot x+1.256$ | 1.000 | $0.076 \cdot x+0.868$ | 0.901 |
|  | 800 | $0.061 \cdot x+1.303$ | 0.990 | $0.064 \cdot x+0.880$ | 0.874 |
| 4.0 | 100 | $0.146 \cdot x+1.316$ | 0.975 | $0.120 \cdot x+0.901$ | 1.000 |
|  | 350 | $0.126 \cdot x+1.318$ | 0.987 | $0.128 \cdot x+0.860$ | 0.996 |
|  | 600 | $0.118 \cdot x+1.294$ | 0.976 | $0.112 \cdot x+0.868$ | 0.942 |
|  | 800 | $0.111 \cdot x+1.269$ | 0.985 | $0.076 \cdot x+0.908$ | 0.901 |
| 3.5 | 100 | $0.225 \cdot x+1.296$ | 0.989 | $0.124 \cdot x+1.001$ | 0.989 |
|  | 350 | $0.167 \cdot x+1.369$ | 0.982 | $0.132 \cdot x+1.001$ | 0.983 |
|  | 600 | $0.090 \cdot x+1.501$ | 0.939 | $0.116 \cdot x+0.977$ | 0.987 |
|  | 800 | $0.081 \cdot x+1.527$ | 0.918 | $0.108 \cdot \mathrm{x}+0.949$ | 0.975 |
| 3.0 | 100 | $0.239 \cdot x+1.415$ | 0.985 | $0.152 \cdot x+1.053$ | 0.997 |
|  | 350 | $0.205 \cdot x+1.440$ | 0.976 | $0.140 \cdot x+1.081$ | 0.972 |
|  | 600 | $0.143 \cdot x+1.545$ | 0.965 | $0.144 \cdot x+1.041$ | 0.973 |
|  | 800 | $0.099 \cdot x+1.624$ | 0.975 | $0.120 \cdot x+1.073$ | 0.987 |
| 2.5 | 100 | $0.227 \cdot x+1.693$ | 0.969 | $0.228 \cdot x+1.049$ | 0.994 |
|  | 350 | $0.183 \cdot x+1.731$ | 0.944 | $0.188 \cdot x+1.129$ | 0.991 |
|  | 600 | $0.194 \cdot x+1.670$ | 0.958 | $0.184 \cdot x+1.121$ | 0.983 |
|  | 800 | $0.176 \cdot x+1.685$ | 0.959 | $0.150 \cdot x+1.191$ | 0.977 |
| 2.0 | 100 | $0.246 \cdot x+1.925$ | 0.982 | $0.276 \cdot x+1.113$ | 0.994 |
|  | 350 | $0.221 \cdot x+1.929$ | 0.975 | $0.248 \cdot x+1.276$ | 0.988 |
|  | 600 | $0.265 \cdot x+1.839$ | 0.979 | $0.244 \cdot x+1.249$ | 0.982 |
|  | 800 | $0.216 \cdot x+1.900$ | 0.968 | $0.218 \cdot x+1.235$ | 0.974 |
| 1.5 | 100 | $0.370 \cdot x+2.010$ | 0.970 | $0.293 \cdot x+1.372$ | 0.989 |
|  | 350 | $0.318 \cdot \mathrm{x}+2.113$ | 0.953 | $0.309 \cdot x+1.424$ | 0.985 |
|  | 600 | $0.350 \cdot x+2.043$ | 0.975 | $0.249 \cdot x+1.568$ | 0.974 |
|  | 800 | $0.293 \cdot x+2.165$ | 0.966 | $0.249 \cdot x+1.532$ | 0.993 |
| 1.0 | 100 | $0.421 \cdot x+2.628$ | 0.966 | $0.492 \cdot \mathrm{x}+1.381$ | 0.999 |
|  | 350 | $0.333 \cdot x+2.865$ | 0.986 | $0.456 \cdot x+1.618$ | 0.997 |
|  | 600 | $0.429 \cdot x+2.565$ | 0.987 | $0.380 \cdot x+1.838$ | 0.991 |
|  | 800 | $0.391 \cdot x+2.762$ | 0.994 | $0.384 \cdot x+1.778$ | 0.988 |
| 0.5 | 100 | $1.002 \cdot x+2.789$ | 0.911 | $0.633 \cdot x+1.705$ | 0.995 |


|  | 350 | $0.930 \cdot \mathrm{x}+3.055$ | 0.951 | $0.709 \cdot \mathrm{x}+2.021$ | 0.989 |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  | 600 | $0.906 \cdot \mathrm{x}+3.271$ | 0.912 | $0.725 \cdot \mathrm{x}+2.181$ | 0.984 |
|  | 800 | $0.827 \cdot \mathrm{x}+3.549$ | 0.902 | $0.685 \cdot \mathrm{x}+2.301$ | 0.975 |
|  | 100 | $2.698 \cdot \mathrm{x}+2.020$ | 1.000 | $1.021 \cdot \mathrm{x}+1.717$ | 0.999 |
|  | 350 | $2.731 \cdot \mathrm{x}+3.276$ | 0.997 | $1.465 \cdot \mathrm{x}+1.950$ | 1.000 |
|  | 600 | $2.673 \cdot \mathrm{x}+4.187$ | 0.991 | $1.565 \cdot \mathrm{x}+2.691$ | 0.999 |
|  | 800 | $2.583 \cdot \mathrm{x}+5.715$ | 0.969 | $1.601 \cdot \mathrm{x}+2.827$ | 0.999 |
|  | 100 | $3.41 \cdot \mathrm{x}+1.707$ | 1.000 | $1.117 \cdot \mathrm{x}+1.838$ | 0.993 |
|  | 350 | $3.891 \cdot \mathrm{x}+2.764$ | 0.997 | $1.690 \cdot \mathrm{x}+2.021$ | 0.999 |
|  | 600 | $3.983 \cdot \mathrm{x}+3.644$ | 0.987 | $1.982 \cdot \mathrm{x}+2.518$ | 0.999 |
|  | 800 | $4.510 \cdot \mathrm{x}+4.792$ | 0.978 | $2.026 \cdot \mathrm{x}+2.418$ | 0.998 |

## S.5. Modifications in modeled grounding zones resulting from input parameters changes

Differences in grounding zone widths for the thickest and thinnest modeled glaciers for every inflow speed and every bedrock slope for both models are provided in Table S3.

Table S3. Grounding zone width difference (in meters) for a 2500 m -thick glacier and a 1000 m -thick glacier.

| $\Delta G Z=G Z_{H=2.5 \mathrm{~km}}-G Z_{H=1.0 \mathrm{~km}}, m$ (viscous model) |  |  |  |  |  |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Speed, m/yr | Bed slope, \% |  |  |  |  |  |  |  |  |  |  |  |
|  | 5.0 | 4.5 | 4.0 | 3.5 | 3.0 | 2.5 | 2.0 | 1.5 | 1.0 | 0.5 | 0.1 | 0.05 |
| 100 | 175 | 175 | 200 | 360 | 390 | 375 | 385 | 635 | 740 | 1866 | 4091 | 5066 |
| 350 | 120 | 160 | 170 | 265 | 350 | 320 | 370 | 550 | 545 | 1651 | 4266 | 6127 |
| 600 | 60 | 125 | 155 | 145 | 240 | 315 | 430 | 575 | 715 | 1696 | 4281 | 6627 |
| 800 | 35 | 85 | 150 | 145 | 135 | 270 | 370 | 510 | 615 | 1565 | 4412 | 6701 |
| Mean | 98 | 136 | 169 | 229 | 279 | 320 | 389 | 568 | 654 | 1695 | 4263 | 6130 |
| $\Delta G Z=G Z_{H=2.5 \mathrm{~km}}-G Z_{H=1.0 \mathrm{~km}}, m$ (viscoelastic model) |  |  |  |  |  |  |  |  |  |  |  |  |
| Speed, m/yr | Bed slope, \% |  |  |  |  |  |  |  |  |  |  |  |
|  | 5.0 | 4.5 | 4.0 | 3.5 | 3.0 | 2.5 | 2.0 | 1.5 | 1.0 | 0.5 | 0.1 | 0.05 |
| 100 | 121 | 141 | 180 | 200 | 240 | 360 | 440 | 481 | 761 | 961 | 1502 | 1521 |
| 350 | 120 | 121 | 201 | 220 | 240 | 300 | 400 | 501 | 720 | 1141 | 2222 | 2543 |
| 600 | 100 | 141 | 201 | 180 | 240 | 300 | 400 | 421 | 600 | 1202 | 2362 | 2963 |
| 800 | 80 | 121 | 141 | 180 | 200 | 260 | 370 | 401 | 620 | 1142 | 2482 | 3143 |
| Mean | 105 | 131 | 181 | 195 | 230 | 305 | 403 | 451 | 675 | 1112 | 2142 | 2543 |

## S.6. Data-derived characteristics of the glaciers

The relative distributions of data-derived bed slopes, ice thicknesses, and ice floe speeds for the glaciers of interest are shown in Figure S3. The empty dots in the boxplots correspond to measurements determined by the IQR (Interquartile Range) method outliers. The IQR-based method identifies a data point as an outlier if it lays outside the $\left[Q_{1}-1.5 \cdot I Q R ; Q_{3}+1.5 \cdot I Q R\right]$ range, where $Q_{1}$ and $Q_{3}$ are the $25^{\text {th }}$ and the $75^{\text {th }}$ percentile of a considered dataset.


Figure S3. Distributions if the ice-bed characteristics, measured along the selected profiles. Subplots (a) - (c) correspond to Moscow University (MU) glacier; Subplots (d) - (f) correspond to Totten (TOT) glacier; Subplots (g) - (i) correspond to Rennick (REN) glacier; Subplots (m) - (o) show the box plots for all three considered glaciers. Left column of subplots shows the glacier distribution based on bed slope, middle column - distribution based on glacier thickness, and right column - distribution based on ice flow speed.

