A quantitative module of avalanche hazard—comparing forecaster assessments of storm and persistent slab avalanche problems with information derived from distributed snowpack simulations

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Abstract. Avalanche forecasting is a human judgment process with the goal of describing the nature and severity of avalanche hazard based on the concept of distinct avalanche problems. Snowpack simulations can help improve forecast consistency and quality by extending qualitative frameworks of avalanche hazard with quantitative links between weather, snowpack, and hazard characteristics. Building on existing research on modeling avalanche problem information, we present the first

5 spatial modeling framework for extracting the characteristics of storm and persistent slab avalanche problems from distributed snowpack simulations. Grouping of simulated layers based on regional burial dates allows us to track them across space and time and calculate insightful spatial distributions of avalanche problem characteristics.

We applied our approach to ten winter seasons in Glacier National Park, Canada, and compared the numerical predictions to human hazard assessments. Despite good agreement in the seasonal summary statistics, the comparison of the daily assessments

- 10 of avalanche problems revealed considerable differences between the two data sources. The best agreements were found in the presence and absence of storm slab avalanche problems and the likelihood and expected size assessments of persistent slab avalanche problems. Even though we are unable to conclusively determine whether the human or model data set represents reality more accurately when they disagree, our analysis indicates that the current model predictions can add value to the forecasting process by offering an independent perspective. For example, the numerical predictions can provide a valuable
- 15 tool for assisting avalanche forecasters in the difficult decision to remove persistent slab avalanche problems. The value of the spatial approach is further highlighted by the observation that avalanche danger ratings were better explained by a combination of various percentiles of simulated instability and failure depth than by simple averages or proportions. Our study contributes to a growing body of research that aims to enhance the operational value of snowpack simulations and provides insight into how snowpack simulations can help address some of the operational challenges of human avalanche hazard assessments.

20 1 Introduction

Avalanche forecasting is a human judgment process where a wide range of observations is synthesized into an overall picture of the nature and severity of avalanche hazard (LaChapelle, 1980; McClung, 2002a, b). The North American Conceptual Model

of Avalanche Hazard (CMAH, Statham et al., 2018a) and similar standards in Europe (EAWS, 2023b) set the foundation for a common language and qualitative framework for assessing avalanche hazard based on the concept of avalanche problems.

- 25 While multiple problems can be present at any given time or location, each avalanche problem is characterized by a set of descriptors: (1) the avalanche problem type, which represents an overarching classification that sets expectations for typical patterns, (2) the location of the problem in the terrain, (3) the likelihood of avalanches of the identified problem type, and (4) their expected destructive size. Avalanche forecasters then typically synthesize the avalanche problem information into an overall assessment of the severity of avalanche hazard using an ordinal five-level danger scale (Statham et al., 2010; EAWS,
- 30 2023a).

Substantial research has recently leveraged data-driven approaches to design decision support tools for avalanche forecasters ranging from predictions of avalanche danger ratings (Pérez-Guillén et al., 2022) to snow instability (Mayer et al., 2022) and avalanche activity (Hendrick et al., 2023; Viallon-Galinier et al., 2023; ?)(Hendrick et al., 2023; Viallon-Galinier et al., 2023; Mayer et al. . One of the key requirements for employing machine learning methods is the availability of large data sets that include the full

35 range of possible events and, ideally, measurable target variables (Guikema, 2020).

Operational experience and recent research has shown that there are considerable differences in how the avalanche danger rating, the CMAH, and the concept of avalanche problems are applied by avalanche forecasters (Lazar et al., 2016; Statham et al., 2018b; Techel et al., 2018; Clark, 2019; Horton et al., 2020c; Hordowick, 2022). Since these inconsistencies can lead to serious miscommunications among forecasters themselves and with the recreational backcountry community, there is a need

- 40 for improving the consistency and quality of the operational use of these cornerstones of avalanche hazard assessments. While the use of predictive models is a possible approach for addressing these challenges, training such models on the existing data sets runs the risk of perpetuating biases and inconsistencies that are contained in the human assessments. Horton et al. (2020c) concluded that a more prescriptive approach might be needed to numerically predict avalanche problem characteristics in an objective way.
- 45 Snowpack simulations that numerically link weather, snowpack, and hazard have great potential to present avalanche forecasters with an independent and reproducible perspective on the possible characteristics of the expected avalanche problems. Extensive research in snowpack modeling for avalanche forecasting dates back over two decades and has led to a variety of operational modeling chains (Morin et al., 2020). While data overload issues and validity concerns have traditionally been the primary hurdles preventing the operational use of snowpack models in Canada (Morin et al., 2020; Herla et al., 2021), sev-
- 50 eral recent studies have focused on making the simulated data more accessible and operationally more relevant by designing visualization tools that better support human sensemaking (Horton et al., 2020b; Nowak et al., 2020; Nowak and Bartram, 2022) and developing algorithms that process snowpack simulations numerically to display relevant summaries in familiar ways (Herla et al., 2021, 2022). A large body of research provides insights into the validation of snowpack simulations from a variety of different angles (Schirmer et al., 2010; Bellaire and Jamieson, 2013; Schmucki et al., 2014; Magnusson et al., 2015;
- 55 Vernay et al., 2015; Quéno et al., 2016; Bellaire et al., 2017; Revuelto et al., 2018; Calonne et al., 2020; Menard et al., 2021; Viallon-Galinier et al., 2020; Morin et al., 2020). Recently, Horton and Haegeli (2022) and ?-Herla et al. (2024b) validated the

simulations on a large scale for their capabilities of representing both new snow amounts and critical avalanche layers, two of the most important aspects for the practitioner community.

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While these studies help forecasters better understand and integrate the simulated snowpack information into their workflows, they do not address the existing challenges in the human analysis process that synthesizes the information into a comprehensive hazard assessment. To address this issue, Reuter et al. (2021) recently established a prescriptive approach for modeling avalanche problem types from simulated snowpack information based on the current understanding of snow instability. In addition, **?** Mayer et al. (2023) developed data-driven models for predicting the probability and size of dry-snow avalanches in the vicinity of weather stations used for snow stratigraphy simulations based on verified data sets of natural avalanche activity and stability tests related to human triggered avalanches. Both of these studies clearly demonstrate the potential of snowpack models for providing avalanche problem information.

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The present study expands on these ideas with two main contributions. First, we present a spatial approach to extracting the characteristics of storm and persistent slab avalanche problems from distributed snowpack simulations that traces individual snowpack layers across space and time and allows the calculation of insightful spatial distributions of avalanche problem

- 70 characteristics. We tailor the output of our numerical predictions to the needs of the North American avalanche community by mirroring concepts included in the CMAH and make the output tangible and relevant by summarizing the simulated information in the familiar format of hazard charts. Second, we examine the agreement between simulations and human assessments for persistent and storm slab avalanche problem situations. We start out with seasonal patterns to compare our results to Reuter et al. (2021) and ?Mayer et al. (2023), but primarily focus on the comparison of daily assessments to simultaneously explore
- 75 the capabilities of the model chain and gain further insight into the strengths and weaknesses of human avalanche hazard assessments. This paper contributes to a growing body of research that aims to enhance the operational value of snowpack simulations and provides insight into how snowpack simulations can help address some of the operational challenges of applying avalanche problems.

2 Data

80 The data sets used in this study consist of snowpack simulations and operational avalanche hazard assessments from avalanche forecasters in western Canada over ten winter seasons (2013–2022) similar to ?Herla et al. (2024b).

The study focuses on the public avalanche forecast region of Glacier National Park that is located in the Columbia Mountains of British Columbia, Canada. Glacier National Park experiences a transitional snow climate with substantial amounts of new snow interspersed with frequent periods of critical layer formation (Haegeli and McClung, 2007; Shandro and Haegeli, 2018).

85 Numerous snowpack modeling studies have been carried out at Glacier National Park (e.g. Bellaire and Jamieson, 2013; Horton et al., 2020c), which is known for high-quality avalanche hazard assessments and observations.

For our simulations, we feed the Canadian numerical weather prediction model HRDPS (Milbrandt et al., 2016, 2.5 km resolution) into the detailed snow cover model SNOWPACK (Bartelt et al., 2002; Lehning et al., 2002b, a) to simulate the snow stratigraphy at 100 grid point locations within the boundaries of Glacier National Park. All simulated snow profiles were



Figure 1. The hazard chart as part of the CMAH allows to quickly understand the severity of distinct avalanche problems. Taken from Statham et al. (2018a).

- 90 valid between 4–5 PM local time representing flat field conditions. For a detailed description of the snowpack simulations used for this study, the interested reader is referred to ?Herla et al. (2024b). Informal conversations with forecasters suggest that the hazard assessments are most supported by observations for the treeline elevation band. Since previous research found most agreement between simulations and assessments also for the treeline elevation band (?)(Herla et al., 2024b), we limit the data set for the present study to grid points between 1800–2100 m asl.
- 95 Avalanche hazard assessments used in this study were issued by public avalanche forecasters every day of the winter season. The assessments represent forecasters' best knowledge of the current conditions (i.e., nowcasts) and were issued in the afternoon for the treeline elevation band in Glacier National Park. Applying the CMAH (Statham et al., 2018a), forecasters partition the avalanche hazard into different avalanche problems and characterize each problem by its type, the likelihood of avalanches, and destructive avalanche size resulting from each avalanche problem. Forecasters express the likelihood of avalanches on a
- 5-level ordinal scale ranging from Unlikelyover, Possible, Likely, Very likely, to Almost certain with half steps. The expected destructive size of avalanches is also expressed on a 5-level ordinal scale ranging from Size 1 to Size 5 with half-sizes (Canadian Avalanche Association, 2016; Statham et al., 2018a). It is common practice in Canada to visualize the assessments of different avalanche problems in a hazard chart that allows for quickly understanding the conditions within a specific location (Fig. 1, taken from Statham et al., 2018a). In addition to the avalanche problem information, the hazard assessments contain danger
- 105 ratings that summarize the hazard from all avalanche problems using the five-level ordinal North American Public Avalanche Danger Scale (Statham et al., 2010), which ranges from *Lowover_, Moderate, Considerable, High*, to *Extreme*.



Figure 2. A flowchart illustrating the methodological steps of this study.

3 Methodology

Our entire analysis is conducted in the R language and environment for statistical computing (R Core Team, 2023) using the open-source software package sarp.snowprofile (Horton et al., 2020a) and consists of three distinct steps. First, individual layers from distributed snowpack simulations are grouped based on date considerations in order to track individual layers across time and maintain knowledge about regional layers across space (Sect. 3.1). Second, we extract avalanche problem information from the snowpack simulations (Sect. 3.2). And third, we compare the simulated information to the human assessment data (Sect. 3.3). Figure 2 illustrates how these individual steps are related.

3.1 Grouping of layers from distributed simulations based on date considerations

- 115 Since persistent weak layers and crusts can cause multiple avalanche cycles, avalanche forecasters typically establish a mental model of where these layers exist and then track the evolution of these layers over time. To facilitate both tracking and communication of these layers, avalanche forecasters in Canada name these layers with date tags and their grain type(s) (e.g., "Jan 17th surface hoar layer"). Reported date tags mostly represent the beginning of snowfall periods that bury layers that were exposed to processes happening at the snow surface before the snowfall and therefore likely contain weak grain types. Sometimes the
- 120 date tags can also represent rain events that form a crust at the snow surface¹.

Since the snowpack builds up chronologically over the winter season, the concept of date tags represents a means to reference specific layers within the snow stratigraphy, similarly to providing the vertical coordinate of a layer (e.g., its depth). However,

¹Please note that other processes leading to crust formation, such as insolation, are not captured by the date tags as presented in this study. In this case, the present approach groups crust layers with other layers based on their burial date



Figure 3. The evolution of a snow stratigraphy over the course of a winter season. The succession of storms and dry periods is illustrated by date tags (dashed vertical lines) that can be used to efficiently track layers across space and over time. Adjacent layers of the same color have been assigned to the same date tags.

the referencing of layers based on dates is more robust, because the vertical coordinate of a specific layer will vary substantially between different locations (Schweizer et al., 2007; Herla et al., 2021) and over time (due to snowpack settlement).

- Perla et al. (2024b) recently applied the concept of layer date tags to group unstable layers from distributed snowpack simulations to determine the spatial prevalence of instability from different layers. In the present study, we apply the concept to our entire data set of layers from one forecast region and elevation band independently of their stability, grain type, or other properties. First, we create a list of all possible date tags for the season based on a) layers that were explicitly tracked by forecasters, and b) simulated precipitation patterns across the region. Analogously to forecaster practice, each date tags for the same storm and were exposed to the same subsequent dry period with one date tag based on its formation date. Figure 3
- illustrates the concept by showing all date tags for the 2019 season and coloring all layers from one profile location according to their date tags. The interested reader is referred to **?** Herla et al. (2024b) for more detailed descriptions of exact rules and thresholds.
- 135 The present approach is well suited to group the layers of large-scale simulations across space and time in a computationally efficient way. Notably, the approach allows us to take into account stable and unstable layers alike, while traditionally, weak layer detection routines only target unstable layers. We exploit that detail in the next step when extracting avalanche problem information from the individual groups of layers.

3.2 Extracting avalanche problem information from snowpack simulations

- 140 To extract the avalanche problem characteristics (i.e., type, location, likelihood of avalanches, and their expected size) from the simulations, we start by selecting all regionally grouped layers from a given location at a given day (Sect. 3.1, Fig. 2). Since our simulations represent flat field conditions and the study domain is limited to one forecast region and elevation band, the location characteristic remains constant throughout this exercise. In a larger-scale application, though, the location would span different regions, elevation bands, and even aspects. In the following paragraphs we describe how we model avalanche problem
- 145 characteristics for each layer at each model grid point separately before aggregating these individual evaluations by date tags, problem types, and finally by all grid points within the relevant location.

Our study focuses on storm and persistent slab avalanche problems (including deep persistent slab avalanche problems), because these problem types can be derived solely from simulated snow profiles. Problem types that cause different avalanches (i.e., wet slab, loose dry, loose wet, cornice, and glide avalanche problems) or require additional weather data (i.e., wind slab

- 150 avalanche problem) are not included in this study. To differentiate between storm and persistent problems, we take into account that forecasters typically issue a storm snow problem for the first few days of the storm, even if a persistent weak layer was buried by the storm and represents the main weakness (Klassen, 2014; Hordowick, 2022). Although our strategy mainly relies on grain type—a persistent layer contributes to a persistent problem, and a new snow layer contributes to a storm problem we assign all persistent layers that have been buried for less than five days to a storm problem instead of a persistent problem.
- According to the CMAH, the likelihood of avalanches emerges from a combination of the spatial distribution² of the avalanche problem within the location bin and the associated sensitivity to triggering avalanches (Statham et al., 2018a). For one isolated model grid point, the likelihood of avalanches simplifies to solely the sensitivity of triggering. To assess the sensitivity of a single layer we use the random forest classifier $p_{unstable}$ developed by Mayer et al. (2022) to characterize dry snow instability for artificial triggering. This model was trained with a high-quality data set of observed snow profiles recorded
- around Davos, Switzerland. Based on the observed instability of the weakest layer in each profile, the model learned to predict the probability of layer instability ($0 \le p_{unstable} \le 1$) from six simulated predictor variables. These predictor variables include both weak layer and slab characteristics, namely the viscous deformation rate, the critical cut length, the sphericity and grain size of the weak layer, the skier penetration depth, and the cohesion of the slab. As suggested by Mayer et al. (2022), we considered layers with $p_{unstable} \ge 0.77$ as critical avalanche layers with poor stability.
- Avalanche size is a function of the lateral and longitudinal extent of the initial slab, release depth, and entrainment along the path (McClung, 2009). While recent research examined the influence of snow mechanical properties like tensile strength and crack propagation speeds on snow instability and avalanche size (e.g. Reuter and Schweizer, 2018; Trottet et al., 2022), there are currently no parametrizations available that can derive avalanche size from large-scale snowpack simulations. We therefore follow the footsteps of McClung (2009) and ? Mayer et al. (2023) and simply use layer depth to characterize the destructive
- 170 potential of simulated avalanche problems. Since the ordinal scale for rating avalanche size is non-linear (Campbell et al.,

²Please note that the term 'spatial distribution' is used slightly differently in North America and Europe. European forecasters primarily use the term 'frequency distribution' and only refer to spatial distribution when they know the exact locations in the terrain (e.g., on a map view). Since this paper builds on the CMAH, we use the term spatial distribution throughout this manuscript.

2016; Statham et al., 2018a) and also the relationship between failure depth and avalanche size has been reported as non-linear (MeClung, 2009; ?)(McClung, 2009; Mayer et al., 2023), we display failure depth on a non-linear axis in all figures of this study by using thresholds suggested by McClung (2009).

After computing the characteristics characteristic avalanche problem type, likelihood of associated avalanches, and depth

- 175 (approximated by $p_{upstable}$), and avalanche size (approximated by anticipated failure depth) for each layer at each grid point with the approaches described above, we then aggregate the information at each grid point by date tag. Hence, the likelihood of avalanches from each group of layers with a specific date tag is represented modeled by the layer with the lowest highest $p_{unstable}$ value, and the release depth corresponds to either the depth of the identified layer (i.e., stable case) or the deepest unstable layer (i.e., if one or more layers are classified as unstable, with $p_{unstable} \leq 0.77 \geq 0.77$). At this point, we know the
- 180 <u>numerical counterparts for all</u> avalanche problem characteristics for each grid point and each date tag, and we can then apply the same strategy to aggregate this information into a single assessment per avalanche problem type at each grid point. Again, the likelihood of avalanches from each problem type is represented modeled by the layer with the lowest highest punstable value, and the release depth corresponds to either the depth of the identified layer or the deepest unstable layer (i.e., punstable $\leq 0.77 \geq 0.77$). The derived avalanche problem information from all grid points in the model domain can now be visualized in
- 185 a similar way similarly to the hazard charts known from the CMAH. The information either characterizes the contribution from each problem type (Fig. 4a) or the contribution from a subset of layers with a specific date tag (Fig. 4b). Data points that are located close to the upper right corner correspond to deeply buried layers that are expected to be triggered easily. Since every data point corresponds to one grid point, the spatial distribution number of potential triggering spots (i.e., spatial distribution in the CMAH) can be gauged from the distribution of the point cloud on the chart. A detailed discussion of the numerical hazard
- 190 chart and its feature of selecting specific date tags is presented in Sect. 5.2.

To aggregate avalanche problem characteristics over a spatial domain, we compute averages medians³ as well as various other percentiles⁴ of likelihood of avalanches $p_{unstable}$ and depth across all individual model grid points (i.e., 10th, 25th, 50th, 75th, and 90th percentiles) taking advantage of our regionally grouped layers that preserve knowledge of stable layers in a meaningful way. While all data points contribute to the computation of the expected likelihood of avalanches $p_{unstable}$ (50th)

- 195 percentile), only data points with poor stability (i.e., $p_{unstable} \ge 0.77$) are considered for the computation of the expected failure depth (50th percentile) and the other percentiles⁵. Taking into account knowledge about both stable and unstable layers allows us to present the spatial distribution of instability in a more comprehensive way and expands on previous approaches, which focused on unstable layers only and were therefore limited to use the proportion of unstable grid points to summarize spatial information (like e.g. ??)(like e.g. Herla et al., 2024b; Mayer et al., 2023).
- To sum up, the present study uses concepts from Reuter et al. (2021), Mayer et al. (2022), and ?-Herla et al. (2024b) to extract avalanche problem characteristics from the simulations. Reuter et al. (2021) demonstrated a prescriptive approach to modeling avalanche problem types primarily informed by physical science principles. Prompted by the findings and conclusions of

³Diamond shapes in Fig. 4a, b, and black lines in Fig. 5c-f.

⁴Grey shading in Fig. 5c–f.

⁵Note that an expected depth is therefore only computed when there is at least one unstable grid point.



Figure 4. Numerical hazard charts derived from snowpack simulations that are similar to the ones produced by forecasters (Fig. 1). As described in the text, the data can be subset by avalanche problem types and date tags, such that each data point corresponds to one grid point location showing the weakest instability ($p_{unstable}$) and the relevant failure depth for the given subset. (a) Showing storm snow and persistent weak layer problems, (b) showing the subset for a specific date tag (Oct 25, the green layer close to the ground in Fig. 3).

Horton et al. (2020c) about existing inconsistencies in the Canadian hazard assessments, we followed the footsteps of Reuter et al. (2021) and designed a prescriptive model-driven approach instead of a data-driven one. Instead of using the processbased stability indices employed in Reuter et al. (2021), though, we applied the random forest classifier p_{unstable} by Mayer et al. (2022), which resulted from a data-driven study using a high-quality data set from Switzerland. We decided to use p_{unstable} instead of the even more recent models proposed by ?-Mayer et al. (2023) that predict the likelihood-probability of natural dry snow avalanche activity and their expected size by modifying p_{unstable}, to build on ?-Herla et al. (2024b) who found encouraging results when applying p_{unstable} to a Canadian data set.

210 3.3 Comparing human assessments to simulations

After extracting and aggregating avalanche problem information from the simulations, this information can be compared to the human avalanche hazard assessment data set. We focused our analysis on the following hazard characteristics: (i) the presence or absence of the problem in either data set with an additional focus on times when the problem was added (i.e., turned on) and removed (i.e., turned off) and removed by forecasters. (ii) trends and (iii) absolute magnitudes of the likelihood of avalanches and their expected size for each avalanche problem type. We used the expected p_{unstable} and the expected failure depth to approximate those characteristics from the simulations (Sect. 3.2). Furthermore, we computed the trends as the strongest trend within a moving five-day window, where the trend could span a single or multiple days. And finally, (iv) we contrasted the expected p_{unstable} and the expected failure depth of both avalanche problem types against the reported danger rating. Since the danger rating synthesizes hazard characteristics from all avalanche problems at a given day, we reduced the data set for this

220 comparison to days when only storm or persistent slab avalanche problems were present but no others. Figure 5 illustrates all



Figure 5. Season summary 2018/19 of human hazard assessments and modeled hazard characteristics for Glacier National Park. (a) Reported danger rating and (b) avalanche problems at treeline elevation, (c), (d) Reported likelihood of avalanches from persistent (red) and storm slab avalanche problems (yellow), respectively. In addition, modeled distribution of $p_{unstable}$ with the envelope of the 10–90th percentiles (light gray shading), the interquartile range (dark gray shading), and the expected value (black line). (e), (f) Analogous to the previous two panels, but showing the reported size of avalanches and the modeled failure depth. (g) modeled interquartile range and median air temperature in units of °C. (h) modeled median height of new snow within 24 hours in units of cm (HN24).

hazard characteristics from the two data sets for the 2019 season: the assigned danger rating and days with reported avalanche problems are shown in (a) and (b), respectively; the assessed likelihood of avalanches from persistent and storm slab avalanche problems is contrasted against the distribution of simulated $p_{unstable}$ in (c) and (d), respectively; and the reported expected size of avalanches is contrasted against the distribution of the simulated depth in (e) and (f).

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To build on previous studies from Reuter et al. (2021) and <u>?Mayer et al. (2023)</u>, we first examined seasonal patterns between the two data sets before analyzing their daily agreement in more detail. For this, we computed the seasonal frequency of storm and persistent slab avalanche problems with the approach presented by Reuter et al. (2021). Next, we explored the multi-seasonal distribution of the expected $p_{unstable}$ and expected failure layer depth on the hazard chart stratified by different avalanche danger ratings in a similar way as $\frac{2}{2}$ Mayer et al. (2023). Finally, we explored the daily agreement between simulated

- 230 and human assessments by computing multi-seasonal distributions of the simulated expected hazard characteristics grouped by forecaster assessments. To examine daily agreement in more detail and make use of the rich information provided by the simulations, we also employed conditional inference trees (CTree) (Hothorn et al., 2006), a type of classification tree that uses a statistical criterion for finding splits. CTrees recursively partition the distribution of a response variable based on the statistically most significant splits along a set of explanatory variables. While the top node of a CTree represents the most
- significant split that divides the entire sample, the resulting subsamples are recursively split into smaller subsamples until the algorithm cannot find any significant splits in the response variable anymore. The resulting terminal nodes describe subsets of the data set with distinct distributions of the response variable that can be linked to specific combinations and thresholds of the explanatory variables. For the present analysis, we used the ctree function in the R package 'partykit' (Hothorn and Zeileis, 2015). We fitted CTrees for both assessed and modeled hazard characteristics as response variables that used explanatory variables from the other data set. The following lists all variables included in the CTree analysis, first the reported variables
- from the human assessments:

- danger rating

- problem type (i.e., storm and persistent)
- problem status (i.e, present, absent)
- 245 expected likelihood of associated avalanches
 - expected size of associated avalanches
 - trend of problem status (i.e., problem got added, removed, or status remained constant)
 - trends of expected likelihood and size

and second the variables extracted from the simulations:

250 – problem type (i.e., storm and persistent)

- expected punstable of problem type and different other percentiles (i.e., 10th, 25th, 50th, 75th, and 90th)
- proportion of unstable grid points with $p_{unstable} \ge 0.77$
- expected depth of problem type and different percentiles (i.e., 10th, 25th, 50th, 75th, and 90th)
- trends of expected punstable and depth



Figure 6. (a), (c) The number of days per season with assessed and modeled persistent weak layer and storm snow problems in GNP. Solid colors refer to human assessments, hatched colors to model results. (b), (d) The relative frequency of assessed and modeled problems. A threshold of the average $p_{unstable} \ge 0.77$ was used to label a problem as modeled.

255 4 Results

4.1 Seasonal patterns

Our ten-year data set contained 1289 days of forecaster assessments that assigned a total of 780 persistent slab avalanche problem days and 572 storm slab avalanche problem days. Using a threshold for the expected $p_{unstable} \ge 0.77$ to classify a problem as modeled, the simulations identified considerably fewer days with avalanche problems, namely 558 and 328

- 260 persistent and storm slab avalanche problem days, respectively. While this offset is evident in all seasons to some degree, the agreement varies between individual seasons (Fig. 6a, b). Overall, the relative frequency between the two avalanche problems is similar between both data sources (Fig. 6c, d). These results are in line with Reuter et al. (2021), whose modeling approach also suggested fewer problems than actually assessed but also found good agreement in the relative frequency of avalanche problems (Fig. 9d–f in Reuter et al., 2021).
- Stratifying the predictions of the numerical hazard chart by the assessed danger rating of the day reveals a steady increase of the median expected values of both failure depth and p_{unstable} (Fig. 7a–d). While the contour maps and their marginal distributions for different danger ratings overlap, the maps for *Low* and *High* occupy distinct areas on the chart. The patterns for *Low* and *Moderate* as well as *Considerable* and *High* show most similarity, particularly since the contour maps are focused on a smaller area for higher danger ratings. The pattern for *Moderate* covers most space on the chart and is therefore most
- 270 strongly characterized by variability.

In addition to the expected $p_{unstable}$ (i.e., 50th percentile), we explored other variants that describe the distribution of $p_{unstable}$, such as the 90th percentile of $p_{unstable}$ or the proportion of unstable grid points (with $p_{unstable} \ge 0.77$) (Fig. A1). While the general patterns remain similar, this comparison illustrates that not one single descriptor best discriminates between the different danger ratings but that a combination might be required (also see CTree analysis further below). Notably, the 90th percentile

275 of *p*_{unstable} best discriminates the danger rating *Low* from the others. While it is located close to the threshold of 0.77 for *Low*, it is substantially higher for all other danger ratings.

Comparing the multi-seasonal patterns of the numerical hazard chart against the human assessments reveals similarities and differences. Most importantly, the medians of the assessed likelihood and size of avalanches increase in a similar manner as in the numerical predictions of Fig. 7, although they are arranged slightly differently due to the categorical nature of the human

- 280 assessments. Furthermore, the contour maps and their marginal distributions show a similar degree of overlap for different danger ratings as the simulated counterparts. Again, the extreme cases of the danger rating occupy distinct areas on the chart. However, the patterns of the human assessments for *Moderate* and *Considerable* are most similar, and the variability is larger for *Considerable* and *High*. Lastly, the visual patterns of the contour maps appear distinctly different between the two data sources. While the numerical contours are circular, the human assessment contours suggest a decreasing trend of the likelihood
- 285 of avalanches for increasing size. The numerical predictions indeed show similar patterns for single days, particularly for layers assigned to different date tags (Fig. 4), but these patterns smooth out when aggregating avalanche problem types and computing multi-seasonal summaries.

4.2 Daily agreement

While examining the ten season summaries (like Fig. 5) qualitatively, we initially found more disagreement than agreement.

For many individual hazard cycles (i.e., a consecutive period of elevated hazard caused by the development and disappearance of an avalanche problem or the cycle of its characteristics), the majority of hazard characteristics showed substantial differences between the human and simulated assessment data sets. For most cycles, only one or two characteristics, such as either absolute magnitudes or trends of different variables of interest, would agree between the two data sets while the other characteristics showed divergent patterns. Hazard cycles with higher levels of agreement in the majority of characteristics were rare. However, a more detailed analysis of the time series that took operational considerations into account revealed more valuable insight.

This paragraph highlights several examples from the 2019 season (Fig. 5). The operational forecasting program started on Dec 01 and instantly reported a persistent weak layer problem. At first sight, this assessment is at odds with the distribution of $p_{unstable}$ which remained at its seasonal minimum for about one week. However, the instability was modeled to be high for the entire week before the forecasting program started and relaxed during the first two days of operations. The human assessments

- 300 during that period most likely took a conservative approach and included the problem due to limited data availability at the beginning of the season, but acknowledged the dormant character of the problem at the same time by publishing danger rating *Low*. Upon loading the persistent layer with several storm cycles between Dec 10 and Jan 05, both data sources agree on the presence of both storm and persistent problems, show comparable trends in hazard characteristics, and suggest a correlation between modeled instability and reported danger rating. A short storm cycle starting on Jan 10 led to storm and persistent
- 305 problems in the human assessment data set. However, since no new snow was modeled during that time period, the modeled hazard characteristics deviated from the assessments for several days. Another brief two-day storm problem starting on Jan 19 was captured by the instability predictions, and the resulting persistent problem was anticipated by the simulations two days earlier than in the assessments. Despite these two problems, the danger rating remained mainly at *Moderate*, until the simulated



Figure 7. (a), (c), (g) Contour maps of the numerical hazard chart and (b), (d), (f), (h) the human assessment hazard chart stratified by different danger ratings from human assessments. The colored square labels highlight the median values of the contour maps and are identical within the left and right columns of the panels, respectively. The violins show the marginal distributions of the contours. While the expected $p_{unstable}$ and the expected depth represent continuous numbers, the assessed likelihood and size are expressed on ordinal 5-level scales with half-steps, which causes the contours to appear binned around the grid intersections.

depth of the weakness increased strongly substantially on Feb 01 when the danger rating also increased to High. After this

- short-lived peak of instability, the danger rating, the reported likelihood of avalanches from persistent problems, and $p_{unstable}$ decreased in concert. After the initial two days of decreasing hazard, the distribution of $p_{unstable}$ started to span a wider range suggesting more variable conditions for triggering. The persistent problem was removed by the forecasters on Feb 15, a week after the modeled interquartile range of $p_{unstable}$ values had decreased below the threshold of 0.77. In the subsequent weeks, several short and moderate mid-level peaks of modeled instability around the threshold of 0.77 were not reflected in the human
- 315 assessments. Each of the peaks was caused by little snowfall amounts below daily averages of 10 cm. A final hazard cycle of the season between Mar 16 and 23 was entirely missed by the simulations. Forecasters issued loose wet avalanche problems and temperatures rose above the freezing level. Although human assessments reported persistent problems, modeled p_{unstable} values remained very low, highlighting that the random forecast classifier was trained for dry snow conditions.

Our qualitative analysis of the seasonal summaries for the other winters revealed the following findings. The modeled instability predictions of persistent problems appear more sensitive to recurrent snow loading than forecaster assessments of likelihood of avalanches. Particularly subtle day-to-day variations seem to agree better with the reported danger rating than the likelihood of persistent avalanches. Interestingly, the 2017 season contained a case when a heavy prolonged snowfall that lasted for longer than two weeks led the modeled instability of persistent layers to decrease considerably, while the instability in storm snow remained high. Not surprisingly, the forecaster assessments listed both problems with peak likelihoods

- 325 of triggering avalanches and the danger rating fluctuated between *Considerable* and *High*. Another notable situation occurred in 2018 when a persistent weak layer problem was dominating the bulletin for nine weeks and simultaneously kept the modeled instability in persistent layers well above the threshold. We also found several instances when an increase in the range of the distribution of p_{unstable} coincided with a decrease of the reported likelihood or danger rating. Although more nuanced, layer-specific information, such as average snow profiles (Herla et al., 2022; ?) (Herla et al., 2022, 2024b) or date tag subsets
- 330 (Fig. 4b), was often helpful to better understand times when persistent problems were added. We also found that the distribution of p_{unstable} added value to the process of understanding the different phases of individual hazard cycles for both storm and persistent problems. Visualizations of the season summaries (like Fig. 5) that are not printed in this manuscript are provided in the code repository (Herla et al., 2024a).

Our quantitative analysis of the multi-seasonal distributions of modeled hazard characteristics using CTrees and direct com-335 parisons supports our qualitative findings. There are distinct differences between storm and persistent slab avalanche problems. The expected p_{unstable} discerns better between periods of reported presence and absence of storm problems than persistent problems (Fig. 8a, b). For both problems, the distributions of the expected p_{unstable} are shifted to significantly larger values when the problem is present (Wilcoxon rank sum test: P < 0.001), but there is considerably more overlap for persistent problems. The addition and removal of storm snow problems is accompanied by mostly increasing and decreasing trends of the 340 expected p_{unstable} (Wilcoxon rank sum test: P < 0.001, Fig. 8c), respectively. In contrast, the distributions of the absolute expected p_{unstable} and the trend of the expected p_{unstable} do not show any differences between days when a persistent problem **as** was added or removed (Wilcoxon rank sum test: P = 0.5, Fig. 8d). Comparing the distributions of the trend of the expected p_{unstable} for different reported trends of the likelihood of avalanches show no significant patterns for either avalanche problem type (Wilcoxon rank sum test: P > 0.17, Fig. 8e, f). Lastly, the absolute values of the expected punstable and the expected failure

- 345 depth show meaningful patterns given their reported counterparts (Fig. 8g–j), although there is substantial overlap within the grouped distributions. Here, the trends are more apparent for persistent problems and expected depth with each modeled median consistently increasing for each increase in the ordinal assessment variable (notched box plots of Fig. 8h, j; for notches see Chambers et al., 2018, p. 61) (Kendall's tau: 0.16, 0.17 for Fig. 8g, i and 0.17, 0.32 for Fig. 8h, j, respectively).
- Our CTree analyses of the hazard characteristics that did not show strong trends in their multi-seasonal distributions mainly uncovered inconsistencies between both data sources and did not reveal any additional insightful findings. We therefore limit our description on the CTrees that use the expected p_{unstable} and the reported danger rating as response variables. To focus readers' attention and highlight patterns more clearly, we limit the visual representation of the CTrees to few levels. Deeper splits that showed insightful relations are mentioned in the text only.
- The CTree for the expected p_{unstable} (Fig. 9) highlights a strong interaction between persistent and storm slab avalanche problems (Node 1). The distributions of the expected p_{unstable} are shifted towards larger values for increasing reported likelihood of avalanches if only persistent problems are reported in the assessments (left branch with Nodes 2–6). However, if both problems are reported, the expected p_{unstable} of persistent layers (Node 8) is yet significantly larger, independently of the reported likelihood. Storm snow problems are generally associated with smaller values of the expected p_{unstable} and are influenced by the reported likelihood more strongly than by the interaction with persistent problems (Nodes 9–13).
- The first CTree for the reported danger rating explores its relation to simulated hazard characteristics (Fig. 10a). The most significant splits of the CTree are driven by characteristics that pertain to the likelihood of avalanches (i.e., percentiles of $p_{unstable}$ and the proportion of unstable grid points, Nodes 1, 2, 5), followed by characteristics that pertain to avalanche size (i.e., percentiles of layer depths, Nodes 6, 9). While the 90th percentile of $p_{unstable}$ discerns low hazard situations (Nodes 1–4), it is the proportion of unstable grid points (with expected $p_{unstable} \ge 0.77$) that initiates the splits for higher hazard situations
- 365 (Node 5). This nicely illustrates that hazard is driven by the weakest instabilities combined with the spatial distribution of instability, but not by stability, which would be expressed by the more stable part of the distribution (i.e., lower percentiles of p_{unstable}). At this point, the depth of the layers becomes important. The 90th percentile of expected failure depth explains the hazard when less than 55 % of grid points are unstable. For depths less than 30 cm the hazard is significantly lower than for situations when avalanches could potentially release deeper (Nodes 7–8). However, if the majority of grid points is are unstable,
- 370 the 10th percentile of expected failure depth identifies the highest hazard situations (Node 9). If avalanches are expected to release at least 35 cm deep, the hazard is almost exclusively rated as *Considerable* or *High* (Node 11). Deeper splits show that the hazard rating generally tends to be higher for storm snow problems than for persistent problems (not shown, but discussed in detail in Sect. 5.1).

We recomputed the CTree for the reported danger rating by using explanatory variables from the human assessments (instead of the simulations) to illustrate similarities and differences between the two data sources (Fig. 10b). In contrast to the previous CTree with the simulated predictors, the most significant split is driven by the presence and absence of avalanche problems (Nodes 1–2). As expected, the hazard is mostly *Low* when no problem is present (Node 3). A sole persistent slab avalanche problem increases the hazard to mostly *Moderate*, while the reported likelihood of avalanches affects the hazard level (Nodes 5–



Figure 8. Multi-seasonal distributions of numerical hazard characteristics grouped by categorical assessments of human forecasters for (a), (c), (e), (g), (i) storm snow problems and (b), (d), (f), (h), (j) persistent weak layer problems.



Figure 9. CTree for the expected punstable examining the interaction effects of explanatory variables extracted from human assessments.

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6). In situations with only a storm snow problem or both problems being present, the likelihood of avalanches being at 'Almost certain' discerns most days with *High* hazard (Nodes 7, 11). Only at a third level, the expected size of avalanches shifts the hazard between *Moderate* and *Considerable* (Nodes 8–10). Despite the substantial importance of presence and absence of avalanche problems in this CTree, both CTree analyses for the reported danger rating suggest that the reported likelihood of avalanches is slightly more influential than their expected size. Furthermore, both CTrees suggest meaningful combinations and thresholds of explanatory variables, while the CTree using human assessment predictors is characterized by slightly less variability than the CTree using numerical hazard characteristics (mean relative frequency of the modes of the terminal nodes 385 equals 0.66 versus 0.51, respectively).

5 Discussion

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The following discussion is structured around the two overarching research objectives. We first discuss the insights and implications from our comparison between simulations and human assessments before reviewing the benefits of the proposed spatial modeling framework and reflecting on the limitations. All sections speak to snowpack modelers and avalanche forecasters alike. In the last section, we outline the additional research steps we intend to complete before submitting the manuscript for publication in a peer-reviewed journal.



Figure 10. CTrees for the reported danger rating examining the explanatory performance of (a) modeled hazard characteristics and (b) human hazard assessments.

5.1 Insights from the comparison between simulations and human assessments

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Reuter et al. (2021) uses different process-based indices for natural versus artificial triggering and dry versus wet slab avalanches,

- 395 whereas ?-<u>Mayer et al. (2023)</u> and Hendrick et al. (2023) derived data-driven models for natural dry slab avalanche and natural wet slab avalanche activity, respectively. Our study only uses one stability index, p_{unstable} as developed by Mayer et al. (2022), to characterize dry snow instability in the context of artificial triggering. Although this decision limits the scope of the present study conceptually, comparisons with the patterns found by Reuter et al. (2021) and ?-<u>Mayer et al. (2023)</u> show very encouraging similarities. Reuter et al. (2021) focused their model validation on a well documented case of critical snow instability
- 400 over the course of ten days as well as on seasonal comparisons against avalanche observations and hazard assessments in Switzerland and Canada spanning multiple years. Although the approaches of characterizing snow instability differ between Reuter et al. (2021) and the present study, seasonal comparisons of modeled storm and persistent slab avalanche problem days against hazard assessments from Glacier National Park, Canada, show very similar patterns (Sect. 4.1, Fig. 6 of this study, and Fig. 9d–f in Reuter et al., 2021). Both modeling approaches suggest fewer avalanche problem days than were actually issued,
- but agree with the human assessments on the relative frequency of the different problem types. This is in line with direct comparisons between $p_{unstable}$ and a combination of process-based instability indices conducted by ? Herla et al. (2024b) who found similar performances in the characterization of weak layers.

? <u>Mayer et al. (2023)</u> nicely illustrate how their data-driven models for predicting the <u>likelihood probability</u> of natural dry slab avalanche activity and their expected size can be used to simulate avalanche problem characteristics. While they did not distinguish between different avalanche problem types, they used the hazard chart of the CMAH (Statham et al., 2018a)

- to demonstrate that their model predictions are in line with multi-seasonal patterns of verified danger ratings from 21 years (Fig. 10 in ?)(Fig. 10 in Mayer et al., 2023). A comparison with Fig. 7a, c, e, g of our study shows very similar patterns. In both studies, the contours of the joint distribution of likelihood and size (or expected punstable and layer depth) migrate nicely from the lower left to the upper right corner with increasing danger ratings, and the medians show clear and comparable trends.
- 415 While the contours show a considerable amount of spread for all danger rating levels, the Canadian data set tends to show more overlap than the Swiss one, which is likely caused by the fact that Canadian danger ratings are of lower quality since they represent operational nowcasts compared to the verified reanalysis ratings in the Swiss data set (?)(Mayer et al., 2023). Interestingly though, the Swiss data set shows the biggest spread for *Considerable*, whereas the Canadian data set has it for *Moderate*. Despite the agreement on general patterns, a detailed comparison of the absolute values along the likelihood
- 420 axis of the contour plots reveals some differences. Since **?**-Mayer et al. (2023) derived the likelihood characteristic from the probability of more than 50 % of grid points being unstable, their numerical predictions span the entire range within [0, 1]. Our likelihood estimates, on the other side are limited to a much smaller effective range since we use the p_{unstable} output of the random forecast classifier. However, this discrepancy can easily be accounted for by shifting the base line of the hazard chart. Lastly, our median predictions of expected failure depth are slightly larger than reported by **?**-Mayer et al. (2023) for danger
- 425 ratings *Low* to *Considerable* and smaller for *High*, which we attribute to differences in model setup (weather model driven versus weather station driven) and the underlying hazard assessment approaches.

Although the seasonal and multi-seasonal comparisons of simulated and reported avalanche problem characteristics presented in this and other studies (Reuter et al., 2021; ?) (Reuter et al., 2021; Mayer et al., 2023) show encouraging agreement, our in-depth comparison of their temporal evolution shows very variable agreement. There are only very few days when most

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hazard characteristics from both data sources agree, but most hazard cycles show some degree of agreement in select characteristics, although these characteristics vary between different cycles. Taking operational considerations into account helped explain the observed differences at times. This general finding is in line with the results of ?-Herla et al. (2024b) who validated snowpack simulations for their capabilities to capture critical layers of operational concern. While they found reasonable patterns overall, the agreement in their seasonal validation was substantially higher than in the validation using daily observations.
 Our analysis also showed that the simulated characteristics of storm problems seem better suited to determine their presence

and absence, while the numerical characteristics of persistent problems could better inform the likelihood and size of persistent

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slab avalanches.

While the human avalanche hazard assessments used in this research aim to accurately represent the current conditions based on the available information and snow science considerations, they are ultimately subjective judgments for the purpose of informing the public about the existing hazard (McClung, 2002a, b; Statham et al., 2018a). As such, they are influenced by several operational considerations, which is different from the simulations that focus purely on physical conditions. For example, forecasters may continue to issue a persistent problem to highlight its lingering character even if they do not expect associated avalanches on the specific day. At other times, forecasters may remove the problem on these low hazard days even

reawakening of the problem (Hordowick, 2022). For these reasons, it is unrealistic to expect that the numerical predictions align closely with all aspects of the human assessments (Horton et al., 2020c).

Given these differences, our comparison poses the question of which data source represents reality better when they disagree. While model simulations have started to outperform observation networks in related geophysical fields (Lundquist et al., 2019), completely mis-assessed avalanche forecasts have been rare for a long time (LaChapelle, 1980). Nonetheless, inherent

if the problem is still present to prevent message fatigue and have a chance to more strongly emphasize the timing of the

- 450 uncertainty in the data sets available to human forecasters and inconsistencies found in the assessments support the notion that high quality forecasts are rare, too (LaChapelle, 1980; Schweizer, 2008). We therefore argue that the model predictions can also add value to situations where the present results show rather variable or even poor agreement, at the very least as independent pieces of information that provoke critical reflection on human assessments.
- One of the situations where simulated avalanche problem characteristics could be particularly helpful for forecasters is the removal of persistent slab avalanche problems. Both Horton et al. (2020c) and Hordowick (2022) found that forecasters struggle with the decision to remove persistent slab avalanche problems. We also found evidence of this issue in our analysis where the timing of forecasters' removal of reported persistent problems was often much later and sometimes appeared somewhat arbitrary when compared mostly uncorrelated to the simulated avalanche problem characteristics. Hence, in these difficult to assess situations, the simulations might provide valuable information about the instability of the relevant weak layers. Another advantage of the numerical predictions is that they depict the evolution of instability and depth more continuously and at a

finer resolution than the coarser, ordinal human assessments. This presents forecasters with a more subtle perspective on the evolution of the hazard characteristics.

Our CTree analysis revealed a strong degree of interaction between storm and persistent slab avalanche problems, which is consistent with the findings of Horton et al. (2020c). Our results show that when there is no reported storm snow avalanche problem, the reported likelihood of persistent slab avalanches increases with increasing p_{unstable} (i.e., decreasing stability) as expected. When a storm snow problem was reported at the same day, however, p_{unstable} was usually highest and did not correlate with the reported likelihood of the persistent problem anymore. This is supported by our qualitative analysis that found modeled instabilities of persistent layers to be more sensitive to new snow loading than the human assessments would suggest. However, the model perceives concurrent instabilities in the storm snow to be more stable (i.e., lower values of

- 470 $p_{unstable}$), which is different from forecasters who tend to assign higher likelihoods of avalanches and higher danger ratings to storm slab avalanche problems (Clark, 2019). This suggests that $p_{unstable}$ is either better suited for persistent layers (which it was developed for), or forecasters generally overestimate the likelihood of storm slab avalanches relative to persistent slab avalanches. Despite these nuances, overall, $p_{unstable}$ seems to characterize the transition from rather stable to rather unstable conditions well. The patterns from the present study are in line with the previously reported threshold of $p_{unstable} \ge 0.77$
- 475 (Mayer et al., 2022) and the logistic regression in $\frac{2}{1000}$, Fig. 5b-Mayer et al. (2023, Fig. 5b) that highlights a rapidly increasing likelihood-probability of natural dry slab avalanches for p_{unstable} $\geq 0.75-0.85$.

Our qualitative and quantitative analyses both suggest that the reported danger rating aligns better with the simulated avalanche problem characteristics than the reported likelihood of avalanches. Comparisons between modeled instability and reported likelihood of avalanches showed a considerable amount of unexplained variability. Nonetheless, the modeled insta-

bilities and layer depths were able to explain the reported danger rating almost as well as the reported hazard characteristics. Although *p*unstable was only trained with snowpack characteristics, the local danger rating was used to filter only for stable and unstable profiles at *Low* and *High* hazard days, respectively, which may create a subtle correlation. Interestingly, the reported characteristic with most explanatory power was the presence or absence of avalanche problems, and not the likelihood or size of avalanches. We hypothesize that this difference observation is at least partially due to the nature of the characteristic that represents a more high-level assessment and, as a binary variable (presence/absence), is easier to assess than a 5-level ordinal scale. This was originally highlighted by Atkins (2004), the conceptual creator of avalanche problem types, who argued that

traditional stability assessments are subject to substantial uncertainties and cannot express all important aspects of the severity of avalanche hazard, even if combined with expected avalanche size.

A notable finding for snowpack modelers is that our CTree analysis suggests that it is better to use detailed distribution information about instability and layer depth to describe avalanche hazard over a spatial domain than just using their average values. More specifically, the combination of the 90th percentile of $p_{unstable}$ and the proportion of unstable grid points (with $p_{unstable} \ge 0.77$) together with the full envelope of the depth distribution (10th to 90th percentiles) was determined by the CTree to explain the hazard best. For instability, this finding can be re-interpreted as "the weakest instabilities paired with the distribution of instability drive hazard". How stable individual grid points were (i.e., lower percentiles) did not emerge as important explanatory variable.

5.2 Benefits of the proposed spatial modeling approach

The modeling approach presented in this paper expands on the methods presented by Reuter et al. (2021) and ?-Mayer et al. (2023) by extracting information from distributed simulations in a way that preserves knowledge about layers across space and time. The approach adopts concepts from the practitioner community to make the output of numerical predictions of avalanche prob-

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lems from large-scale simulations more organized, transparent, and informative for forecasters. By splitting the overall hazard into contributions from different regional layers, the model predictions cater to the existing sensemaking process of forecasters, which will allow them to integrate the simulated information into their mental model more easily.

This approach also makes it easier for forecasters to identify times when the modeled predictions deviate from reality, like when a specific hazard-driving weak layer is is missed in the simulations or a non-existing layer is modeled. In these situations,

- 505 the proposed modeling approach allows forecasters to keep using the simulations as information source for all other regional layers since they are each assessed separately. This is not supported by other approaches that do not group layers based on date considerations and would therefore require a sophisticated data assimilation/model correction scheme to keep using the avalanche problem predictions in an informative way until the influence of the misrepresented layer has vanished.
- There are several additional advantages to the proposed approach. Since we are grouping layers by date tags, we are not limited to extracting unstable layers. Instead, we can, for example, extract the most unstable layer at each grid point that belongs to a specific date tag. This allows us to compute the entire distribution of instability associated with each date tag across all grid points, while other approaches are limited to using the proportion of unstable grid points (??)(Herla et al., 2024b; Mayer et al., 2023) . Our approach also allows the tracking of layer characteristics during the transitions from instability to stability and can continue to provide information about the distribution of the depth of regional layers once they have become dormant. While the
- 515 approach of Reuter et al. (2021) also supports this for individual layers at single locations, the grouping by date tags enables the continuous tracking for spatial simulations. Furthermore, the date tag approach is computationally more efficient and can be resourcefully applied to large-scale simulations. Finally, our approach avoids the known limitation of Reuter et al. (2021) that may miss the faceting of already buried layers.

An additional feature that is not yet implemented into our avalanche problem assessment is the integration of step-down 520 potential. If an individual layer has become dormant, it does not contribute to the expected failure depth. However, forecasters might be interested in knowing whether the release of a shallower avalanche could trigger the layer. Our approach can help support this feature by identifying times and grid points when and where the triggering of a shallower layer is likely.

5.3 Limitations

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While our approach of extracting avalanche problem information from spatial snowpack simulations offers a promising method, and the comparison against human hazard assessments revealed useful insights, our contribution has to be interpreted in light of several limitations. We have already alluded to two of these limitations earlier. First, the potential inconsistencies and biases contained in the human hazard assessments that prevent us from using it as a reliable ground truth during times of disagreement. Hence, our analysis should not be viewed as a complete validation of the simulations. Second, our choice of only focusing on modeled snowpack data and only one stability index limited the analysis of this paper to only storm and persistent slab

- 530 avalanche problems. While the present analysis can be seen as a proof of concept, further effort is needed to integrate other data sources (e.g., wind fields), other stability indices⁶, and eventually other problem types into our analysis framework and to compare the predictions to assessments from other forecast regions and snow climates. Lastly, we want to acknowledge the limitations caused by uncertainties in the simulations. The quality of the numerical avalanche hazard characteristics depends heavily on the correct representation of the slab and the formation of the weak layer. Raleigh et al. (2015) and Richter et al.
- 535 (2020) report that precipitation is the primary source of error for snowpack structure and stability predictions, which was further confirmed by Horton and Haegeli (2022), who examined differences in observed and modeled snow amounts. Combined with our present findings related to the sensitivity of the stability predictions in new snow loading, we can confidently say that the precipitation input is likely the main source of error in the present analysis. The second source of error, the correct formation of the weak layer, is governed by the interplay of several forcing variables combined with the correct timing. Errors in weak
- 540 layer formation only need to happen during a short time window to negatively affect the stability predictions for several weeks thereafter. The interested reader is referred to ?-Herla et al. (2024b) for a more in-depth discussion about model performances in capturing critical layers of concern.

6 Conclusions

We presented a spatial approach to extract the characteristics of storm and persistent slab avalanche problems from distributed snowpack simulations by grouping individual layers based on their regional burial dates. Our approach allows for computationally efficient tracking of instabilities across space and time to compute spatial distributions of hazard characteristics that are consistent with existing avalanche forecasting practices. We applied the approach to ten winter seasons in Glacier National Park, Canada, and compared the numerical predictions to human hazard assessments to quantify seasonal and daily agreement.

Although the seasonal summaries of the numerically predicted avalanche problems showed strong similarities with human
bazard assessments and agreed with the results of existing research (Reuter et al., 2021; ?)(Reuter et al., 2021; Mayer et al., 2023), our comparisons of the daily characteristics of the avalanche problems revealed considerable discrepancies. The best agreements were found in the presence and absence of storm slab avalanche problems and in the likelihood and expected size assessments of persistent slab avalanche problems. However, our qualitative examination also suggested the numerical predictions might have a better handle on the removal of persistent slab avalanche problems, a known operational challenge (Hordowick, 2022). Our analyses also revealed that avalanche hazard was better explained by the the combination of various percentiles of simulated instability and failure depth than by simple averages or proportions, which highlights the value of having access to the full distribution information. Lastly, the comparison of the two data sources with respect to multiple hazard characteristics

led us to build more confidence in the reported danger rating than the reported likelihood of avalanches.

⁶For example, the skier stability index SK38 (Monti et al., 2016) and the critical cut length r_c (Richter et al., 2019) (both for artificial triggering of dry slab avalanches), the expected time to failure (Conway and Wilbour, 1999) (for natural dry slab avalanche activity), the liquid water content index (Mitterer et al., 2013) and the random forest classifier by Hendrick et al. (2023) (for wet slab avalanche activity).

While differences between human assessments and simulated data sets are expected, an important caveat of our study is that

- 560 it is unclear which of the two data sets represents the truth better. Interestingly, our analyses showed that both data sets have their own strengths and weaknesses and can contribute to a better understanding of the conditions. However, it is beyond the present comparison to explain in detail *why* the two data sources disagree. To answer this question and properly validate the numerical predictions (particularly the temporal integrity of existing stability indices and their underlying parametrizations), we need scientific-grade data sets of complete avalanche hazard assessments, which is currently not available in Canada. Such a data set
- 565 could also be used to develop predictive data-driven models. Furthermore, future research should leverage existing datasets on avalanche activity, such as remote detections and manual observations, to better identify times when model simulations align with actual conditions.

To further strengthen avalanche forecasters' familiarity with the strengths and weaknesses of large-scale snowpack simulations, we encourage the use of dashboards that facilitate real-time comparisons between human assessment and model data sets.

- 570 Understanding their current capabilities requires careful study of context and the consideration of operational practices that differ from the purely physical computations of the simulations. Since assessing some hazard characteristics is easier than others, there is potential for gauging the current value of the simulations and integrating them into the reasoning process accordingly. Even at times when forecasters disagree with the numerical predictions, they can be a valuable independent information source that provokes critical reflection.
- 575 *Code and data availability.* The data and code to reproduce the analysis in this paper are available from a DOI repository at https://www. doi.org/10.17605/OSF.IO/94826 (Herla et al., 2024a).

Appendix A: Contour maps based on different variables describing the distribution of punstable

Author contributions. All authors conceptualized the research; FH ran the snowpack simulations and implemented the methods and analysis; all authors contributed to writing the paper; PH acquired the funding.

580 Competing interests. FH, SH, and PM declare they have no competing interests; PH is a member of the editorial board of the journal.



Figure A1. Contour maps of the numerical hazard chart, using the expected $p_{unstable}$ (i.e., 50th percentile) as in Fig. 7a, c, e, g (left column), the 90th percentile of $p_{unstable}$ (middle column), and the proportion of unstable grid points, with $p_{unstable} \ge 0.77$ (right column). The charts are stratified by their reported danger ratings from human assessments (different rows). The colored square labels highlight the median values of the contour maps and are identical within each column. The violins show the marginal distributions of the contours.

References

585

- Atkins, R.: An avalanche characterization checklist for backcountry travel decisions, in: Proceedings of the 2004 International Snow Science Workshop, Jackson Hole, WY, USA, pp. 462–468, https://arc.lib.montana.edu/snow-science/item/1118, 2004.
- Bartelt, P., Lehning, M., Bartelt, P., Brown, B., Fierz, C., and Satyawali, P.: A physical SNOWPACK model for the Swiss avalanche warning: Part I: Numerical model, Cold Reg. Sci. Technol., 35, 123–145, https://doi.org/10.1016/s0165-232x(02)00074-5, 2002.
- Bellaire, S. and Jamieson, J. B.: Forecasting the formation of critical snow layers using a coupled snow cover and weather model, Cold Reg. Sci. Technol., 94, 37–44, https://doi.org/10.1016/j.coldregions.2013.06.007, 2013.
 - Bellaire, S., van Herwijnen, A., Mitterer, C., and Schweizer, J.: On forecasting wet-snow avalanche activity using simulated snow cover data, Cold Reg. Sci. Technol., 144, 28–38, https://doi.org/10.1016/j.coldregions.2017.09.013, 2017.
- 590 Calonne, N., Richter, B., Löwe, H., Cetti, C., Ter Schure, J., Van Herwijnen, A., Fierz, C., Jaggi, M., and Schneebeli, M.: The RHOSSA campaign: Multi-resolution monitoring of the seasonal evolution of the structure and mechanical stability of an alpine snowpack, The Cryosphere, 14, 1829–1848, https://doi.org/10.5194/tc-14-1829-2020, 2020.
 - Campbell, C., Conger, S., Gould, B., Haegeli, P., Jamieson, J. B., and Statham, G.: Technical Aspects of Snow Avalanche Risk Management– Resources and Guidelines for Avalanche Practitioners in Canada, Revelstoke, BC, Canada, 2016.
- 595 Canadian Avalanche Association: Observation Guidelines and Recording Standards for Weather, Snowpack, and Avalanches, Tech. rep., Revelstoke, BC, Canada, 2016.
 - Chambers, J. M., Cleveland, W. S., Kleiner, B., and Tukey, P. A.: Graphical methods for data analysis, CRC Press, https://doi.org/10.1201/9781351072304, 2018.

Clark, T.: Exploring the link between the Conceptual Model of Avalanche Hazard and the North American Public Avalanche Danger Scale,

600 in: MRM Thesis in Resource and Environmental Management, Simon Fraser University, 2019.

Conway, H. and Wilbour, C.: Evolution of snow slope stability during storms1, Cold Reg. Sci. Technol., 30, 67–77, https://doi.org/10.1016/S0165-232X(99)00009-9, 1999.

EAWS: Standards – Avalanche Danger Scale, https://www.avalanches.org/standards/avalanche-danger-scale/, 2023a.

EAWS: Standards - Avalanche Problems, https://www.avalanches.org/standards/avalanche-problems/, 2023b.

- 605 Guikema, S.: Artificial Intelligence for Natural Hazards Risk Analysis: Potential, Challenges, and Research Needs, Risk Anal., 40, 1117– 1123, https://doi.org/10.1111/risa.13476, 2020.
 - Haegeli, P. and McClung, D. M.: Expanding the snow-climate classification with avalanche-relevant information: Initial description of avalanche winter regimes for southwestern Canada, J. Glaciol., 53, 266–276, https://doi.org/10.3189/172756507782202801, 2007.

Hendrick, M., Techel, F., Volpi, M., Olevski, T., Pérez-Guillén, C., Herwijnen, A. V., and Schweizer, J.: Automated prediction of wet-snow avalanche activity in the Swiss Alps, J. Glaciol., 50, 1–14, https://doi.org/10.1017/jog.2023.24, 2023.

Herla, F., Horton, S., Mair, P., and Haegeli, P.: Snow profile alignment and similarity assessment for aggregating, clustering, and evaluating of snowpack model output for avalanche forecasting, Geosci. Model Dev., 14, 239–258, https://doi.org/10.5194/gmd-14-239-2021, 2021.

- Herla, F., Haegeli, P., and Mair, P.: A data exploration tool for averaging and accessing large data sets of snow stratigraphy profiles useful for avalanche forecasting, The Cryosphere, 16, 3149–3162, https://doi.org/10.5194/tc-16-3149-2022, 2022.
- 615 Herla, F., Haegeli, P., Horton, S., and Mair, P.: A quantitative module of avalanche hazard—Data and Code, https://doi.org/10.17605/OSF.IO/W7PJY, 2024a.

- Herla, F., Haegeli, P., Horton, S., and Mair, P.: A large-scale validation of snowpack simulations in support of avalanche forecasting focusing on critical layers, Nat. Hazard Earth Sys., 24, 2727–2756, https://doi.org/10.5194/nhess-24-2727-2024, 2024b.
- Hordowick, H.: Understanding avalanche problem assessments: A concept mapping study with public avalanche forecasters, in: MRM Thesis
 in Resource and Environmental Management, Simon Fraser University, http://www.avalancheresearch.ca/pubs/2022_hordowick_mrm/, 2022.
 - Horton, S. and Haegeli, P.: Using snow depth observations to provide insight into the quality of snowpack simulations for regional-scale avalanche forecasting, The Cryosphere, 16, 3393–3411, https://doi.org/10.5194/tc-16-3393-2022, 2022.
 - Horton, S., Herla, F., and Haegeli, P.: An R package for snow profile analysis and visualization, in: Proceedings of the 2020 Virtual Snow
- Science Workshop VSSW, Fernie, BC, Canada, 2020a.
 Horton, S., Nowak, S., and Haegeli, P.: Enhancing the operational value of snowpack models with visualization design principles, Nat. Hazard Earth Sys., 20, 1557–1572, https://doi.org/10.5194/nhess-20-1557-2020, 2020b.
 - Horton, S., Towell, M., and Haegeli, P.: Examining the operational use of avalanche problems with decision trees and model-generated weather and snowpack variables, Nat. Hazard Earth Sys., 20, 3551–3576, https://doi.org/10.5194/nhess-20-3551-2020, 2020c.
- 630 Hothorn, T. and Zeileis, A.: Partykit: A modular toolkit for recursive partytioning in R, J. Mach. Learn. Res., 16, 3905–3909, http://cran. r-project.org/package=, 2015.
 - Hothorn, T., Hornik, K., and Zeileis, A.: Unbiased recursive partitioning: A conditional inference framework, J. Comput. Graph. Stat., 15, 651–674, https://doi.org/10.1198/106186006X133933, 2006.
 - Klassen, K.: What's the problem? A primer on defining avalanche character., The Avalanche Journal, 105, 10–12, 2014.
- 635 LaChapelle, E. R.: The fundamental processes in conventional avalanche forecasting., J. Glaciol., 26, 75-84, https://doi.org/10.3189/s0022143000010601, 1980.
 - Lazar, B., Trautman, S., Cooperstein, M., Greene, E., and Birkeland, K. W.: NORTH AMERICAN AVALANCHE DANGER SCALE: DO BACKCOUNTRY FORECASTERS APPLY IT CONSISTENTLY?, International Snow Science Workshop, pp. 457–465, https://arc.lib. montana.edu/snow-science/item/2307, 2016.
- 640 Lehning, M., Bartelt, P., Brown, B., and Fierz, C.: A physical SNOWPACK model for the Swiss avalanche warning Part III: Meteorological forcing, thin layer formation and evaluation, Cold Reg. Sci. Technol., 35, 169–184, https://doi.org/10.1016/S0165-232X(02)00072-1, 2002a.
 - Lehning, M., Bartelt, P., Brown, B., Fierz, C., and Satyawali, P.: A physical SNOWPACK model for the Swiss avalanche warning Part II. Snow microstructure, Cold Reg. Sci. Technol., 35, 147–167, https://doi.org/10.1016/S0165-232X(02)00073-3, 2002b.
- 645 Lundquist, J., Hughes, M., Gutmann, E., and Kapnick, S.: Our skill in modeling mountain rain and snow is bypassing the skill of our observational networks, Bull. Am. Meteorol. Soc., 100, 2473–2490, https://doi.org/10.1175/BAMS-D-19-0001.1, 2019.
 - Magnusson, J., Wever, N., Essery, R., Helbig, N., Winstral, A., and Jonas, T.: Evaluating snow models with varying process representations for hydrological applications, Water Resources Res., 51, 2707–2723, https://doi.org/10.1002/2014WR016498, 2015.

Mayer, S., van Herwijnen, A., Techel, F., and Schweizer, J.: A random forest model to assess snow instability from simulated snow stratig-

650 raphy, The Cryosphere, 16, 4593–4615, https://doi.org/10.5194/tc-16-4593-2022, 2022.

- Mayer, S., Techel, F., Schweizer, J., and Van Herwijnen, A.: Prediction of natural dry-snow avalanche activity using physics-based snowpack simulations, Nat. Hazard Earth Sys., 23, 3445–3465, https://doi.org/10.5194/nhess-23-3445-2023, 2023.
- McClung, D. M.: The Elements of Applied Avalanche Forecasting, Part I: The Human Issues, Nat. Hazards, 26, 111–129, https://doi.org/10.1023/a:1015665432221, 2002a.

- 655 McClung, D. M.: The Elements of Applied Avalanche Forecasting, Part II: The Physical Issues and the Rules of Applied Avalanche Forecasting, Nat. Hazards, 26, 131–146, https://doi.org/10.1023/a:1015604600361, 2002b.
 - McClung, D. M.: Dimensions of dry snow slab avalanches from field measurements, J. Geophys. Res. Earth Surf., 114, https://doi.org/10.1029/2007JF000941, 2009.

Menard, C. B., Essery, R., Krinner, G., Arduini, G., Bartlett, P., Boone, A., Brutel-Vuilmet, C., Burke, E., Cuntz, M., Dai, Y., Decharme, B.,

- 660 Dutra, E., Fang, X., Fierz, C., Gusev, Y., Hagemann, S., Haverd, V., Kim, H., Lafaysse, M., Marke, T., Nasonova, O., Nitta, T., Niwano, M., Pomeroy, J., Schädler, G., Semenov, V. A., Smirnova, T., Strasser, U., Swenson, S., Turkov, D., Wever, N., and Yuan, H.: Scientific and human errors in a snow model intercomparison, Bull. Am. Meteorol. Soc., 102, E61–E79, https://doi.org/10.1175/BAMS-D-19-0329.1, 2021.
- Milbrandt, J. A., Bélair, S., Faucher, M., Vallée, M., Carrera, M. L., and Glazer, A.: The pan-canadian high resolution (2.5 km) deterministic
 prediction system, Weather Forecast., 31, 1791–1816, https://doi.org/10.1175/WAF-D-16-0035.1, 2016.
- Mitterer, C., Techel, F., Fierz, C., and Schweizer, J.: An operational supporting tool for assessing wet-snow avalanche danger, in: Proceedings of the International Snow Science Workshop Grenoble–Chamonix Mont-Blanc, pp. 334—-338, https://arc.lib.montana.edu/snow-science/item/1860, 2013.
 - Monti, F., Gaume, J., van Herwijnen, A., and Schweizer, J.: Snow instability evaluation: calculating the skier-induced stress in a multi-layered snowpack, Nat. Hazard Earth Sys., 16, 775–788, https://doi.org/10.5194/nhess-16-775-2016, 2016.
- Morin, S., Fierz, C., Horton, S., Bavay, M., Dumont, M., Hagenmuller, P., Lafaysse, M., Mitterer, C., Monti, F., Olefs, M., Snook, J. S., Techel, F., Van Herwijnen, A., and Vionnet, V.: Application of physical snowpack models in support of operational avalanche hazard forecasting : A status report on current implementations and prospects for the future, Cold Reg. Sci. Technol., 170, 1098–1107, https://doi.org/10.1016/J.COLDREGIONS.2019.102910, 2020.

670

- 675 Nowak, S. and Bartram, L.: I'm Not Sure: Designing for Ambiguity in Visual Analytics, in: Proceedings Graphics Interface, vol. 2022-May, 2022.
 - Nowak, S., Bartram, L., and Haegeli, P.: Designing for Ambiguity: Visual Analytics in Avalanche Forecasting, in: Proceedings 2020 IEEE Visualization Conference, pp. 81–85, Institute of Electrical and Electronics Engineers Inc., https://doi.org/10.1109/VIS47514.2020.00023, 2020.
- 680 Pérez-Guillén, C., Techel, F., Hendrick, M., Volpi, M., Van Herwijnen, A., Olevski, T., Obozinski, G., Pérez-Cruz, F., and Schweizer, J.: Data-driven automated predictions of the avalanche danger level for dry-snow conditions in Switzerland, Nat. Hazard Earth Sys., 22, 2031–2056, https://doi.org/10.5194/nhess-22-2031-2022, 2022.
 - Quéno, L., Vionnet, V., Dombrowski-Etchevers, I., Lafaysse, M., Dumont, M., and Karbou, F.: Snowpack modelling in the Pyrenees driven by kilometric-resolution meteorological forecasts, The Cryosphere, 10, 1571–1589, https://doi.org/10.5194/tc-10-1571-2016, 2016.
- R Core Team: R: A Language and Environment for Statistical Computing, 2023.
 Raleigh, M. S., Lundquist, J., and Clark, M. P.: Exploring the impact of forcing error characteristics on physically based snow simulations within a global sensitivity analysis framework, Hydrol. Earth Syst. Sci., 19, 3153–3179, https://doi.org/10.5194/hess-19-3153-2015, 2015.
 Reuter, B. and Schweizer, J.: Describing Snow Instability by Failure Initiation, Crack Propagation, and Slab Tensile Support, Geophys. Res.

Lett., 45, 7019-7027, https://doi.org/10.1029/2018GL078069, 2018.

690 Reuter, B., Viallon-Galinier, L., Horton, S., van Herwijnen, A., Mayer, S., Hagenmuller, P., and Morin, S.: Characterizing snow instability with avalanche problem types derived from snow cover simulations, Cold Reg. Sci. Technol., 194, 103462, https://doi.org/10.1016/j.coldregions.2021.103462, 2021.

- Revuelto, J., Lecourt, G., Lafaysse, M., Zin, I., Charrois, L., Vionnet, V., Dumont, M., Rabatel, A., Six, D., Condom, T., Morin, S., Viani, A., and Sirguey, P.: Multi-criteria evaluation of snowpack simulations in complex alpine terrain using satellite and in situ observations, Remote Sensing, 10, 1171, https://doi.org/10.3390/rs10081171, 2018.
- Richter, B., Schweizer, J., Rotach, M. W., and Van Herwijnen, A.: Validating modeled critical crack length for crack propagation in the snow cover model SNOWPACK, The Cryosphere, 13, 3353–3366, https://doi.org/10.5194/tc-13-3353-2019, 2019.

695

- Richter, B., Van Herwijnen, A., Rotach, M. W., and Schweizer, J.: Sensitivity of modeled snow stability data to meteorological input uncertainty, Nat. Hazard Earth Sys., 20, 2873–2888, https://doi.org/10.5194/nhess-20-2873-2020, 2020.
- 700 Schirmer, M., Schweizer, J., and Lehning, M.: Statistical evaluation of local to regional snowpack stability using simulated snow-cover data, Cold Reg. Sci. Technol., 64, 110–118, https://doi.org/10.1016/j.coldregions.2010.04.012, 2010.
 - Schmucki, E., Marty, C., Fierz, C., and Lehning, M.: Evaluation of modelled snow depth and snow water equivalent at three contrasting sites in Switzerland using SNOWPACK simulations driven by different meteorological data input, Cold Reg. Sci. Technol., 99, 27–37, https://doi.org/10.1016/j.coldregions.2013.12.004, 2014.
- 705 Schweizer, J.: On the predictability of snow avalanches, in: Proceedings of the 2008 International Snow Science Workshop, Whistler, BC, p. 688, 2008.
 - Schweizer, J., Kronholm, K., Jamieson, J. B., and Birkeland, K. W.: Review of spatial variability of snowpack properties and its importance for avalanche formation, Cold Reg. Sci. Technol., 51, 253–272, https://doi.org/10.1016/j.coldregions.2007.04.009, 2007.
- Shandro, B. and Haegeli, P.: Characterizing the nature and variability of avalanche hazard in western Canada, Nat. Hazard Earth Sys., 18, 1141–1158, https://doi.org/10.5194/nhess-18-1141-2018, 2018.
 - Statham, G., Haegeli, P., Birkeland, K. W., Greene, E., Israelson, C., Tremper, B., Stethem, C., Mcmahon, B., White, B., and Kelly, J.: The North American Public Avalanche Danger Scale, Tech. rep., https://arc.lib.montana.edu/snow-science/item/353, 2010.
 - Statham, G., Haegeli, P., Greene, E., Birkeland, K. W., Israelson, C., Tremper, B., Stethem, C., McMahon, B., White, B., and Kelly, J.: A conceptual model of avalanche hazard, Nat. Hazards, 90, 663–691, https://doi.org/10.1007/s11069-017-3070-5, 2018a.
- 715 Statham, G., Holeczi, S., and Shandro, B.: Consistency and Accuracy of Public Avalanche Forecasts in Western Canada, in: Proceedings of the 2018 International Snow Science Workshop, Innsbruck, Austria, https://arc.lib.montana.edu/snow-science/item/2806, 2018b.
 - Techel, F., Mitterer, C., Ceaglio, E., Coléou, C., Morin, S., Rastelli, F., and Purves, R. S.: Spatial consistency and bias in avalanche forecasts -a case study in the European Alps, Nat. Hazard Earth Sys., 18, 2697–2716, https://doi.org/10.5194/nhess-18-2697-2018, 2018.
- Trottet, B., Simenhois, R., Bobillier, G., Bergfeld, B., van Herwijnen, A., Jiang, C., and Gaume, J.: Transition from sub-Rayleigh anticrack
 to supershear crack propagation in snow avalanches, Nat. Phys., 18, 1094–1098, https://doi.org/10.1038/s41567-022-01662-4, 2022.
 - Vernay, M., Lafaysse, M., Mérindol, L., Giraud, G., and Morin, S.: Ensemble forecasting of snowpack conditions and avalanche hazard, Cold Reg. Sci. Technol., 120, 251–262, https://doi.org/10.1016/j.coldregions.2015.04.010, 2015.
 - Viallon-Galinier, L., Hagenmuller, P., and Lafaysse, M.: Forcing and evaluating detailed snow cover models with stratigraphy observations, Cold Reg. Sci. Technol., 180, 103 163, https://doi.org/10.1016/j.coldregions.2020.103163, 2020.
- 725 Viallon-Galinier, L., Hagenmuller, P., and Eckert, N.: Combining modelled snowpack stability with machine learning to predict avalanche activity, The Cryosphere, 17, 2245–2260, https://doi.org/10.5194/tc-17-2245-2023, 2023.