Droughts of the Early 19th Century (1790-1830) in Northeast Iberian Peninsula: Integration of historical and instrumental data for high-resolution reconstructions of extreme events

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ABSTRACT.

Drought represents a prevalent climate risk in the Mediterranean region. In the context of climate change, an increase in both frequency and intensity is anticipated over the next century. In order to effectively manage future scenarios where global warming overlays natural climate variability, a thorough analysis of the nature of droughts prior to the industrial age is ~~imperative~~crucial. This approach incorporates an extended temporal scale into the study of severe droughts, enabling the identification of low-frequency drought events that occurred before the instrumental period. The objective of this study is to examine the occurrence and magnitude of extreme droughts lasting over a year in the Spanish Mediterranean Basin during the Early 19th Century (1790-1830). To achieve this objective, the research integrates the use of instrumental observations and information derived from historical documentary sources with daily to monthly resolutions (e.g. rogation ceremonies). The findings reveal that drought episodes were more frequent and severe during the Early 19th Century ~~than in the~~ compared to the Late 19th Century~~second half of this century~~. Moreover, drought episodes of similar severity were rare throughout the 20th Century. Only in the current context of climate change, over the last two decades, has a pattern of high drought severity been identified that resembles the severity found during the Early 19th Century (especially between 1812 and 1825). This study underscores the presence of high variability in drought patterns over the last centuries, justifying the need for intensified research on drought episodes with high temporal resolution for extended periods.

KEYWORDS.
Early 19th Century, Documentary Sources, Droughts, Drought Indices, Meteorological records, Spanish Mediterranean Basin.

1. INTRODUCTION

Drought is a climate phenomenon defined as a prolonged absence of precipitation that can last for a few weeks to periods of up to several years (IDMP, 2022). According to the IPCC, drought is an exceptional period of water shortage for existing ecosystems and the human population (due to low rainfall, high temperature and/or wind) (IPCC, ~~2022~~2023). Despite their complexity as a natural phenomenon, droughts should not be confused with aridity, desertification or other related natural risks such as forest fires or heatwaves (~~IDMP, 2022; ~~Van Loon, 2015; IDMP, 2022). Drought, as a prolonged lack of precipitation, can be classified depending on the impacts on the environment and society resulting in distinct types of droughts such as meteorological, hydrological, agricultural and social (Wilhite & Glantz, 1985).

Meteorological drought is defined as a prolonged period with abnormal rainfall deficit for a large region and for a long period of time (Mishra & Singh, 2010; IPCC, ~~2022~~2023). This absence of rain is transmitted to the hydrological system by affecting soil moisture and groundwater input, ultimately reducing surface water levels (Van Loon & Van Lanen, 2012). Thus, hydrological drought is defined as a period with large runoff and water deficits in rivers, lakes and reservoirs (Nalbantis, 2008; IPCC, ~~2022~~2023). It also has effects on groundwater and surface hydrology (~~Mishra & Singh, 2010; ~~Wilhite & Glantz, 1985; Mishra & Singh, 2010). This deficit causes a reduction in water supply to plant roots leading to agricultural and ecological drought (Sivakumar, 2011; Van Loon, 2015; IPCC, 2023). ~~). This type of drought consists of a period of abnormal soil moisture deficit, which results from combined shortage of rainfall and excess evapotranspiration (Sivakumar, 2011a; IPCC, 2022). Consequently, during the growing season, it impinges on crop production, leading to reduced yields and even crop failure (Mishra & Singh, 2010; IPCC, 2022). ~~The different impacts of drought mentioned above, such as the reduction in water levels or crop failures, have a direct effect on human societies (Van Loon, 2015). ~~These effects on society caused by a prolonged drought over time are defined as social drought (Van Loon, 2015). The main impact of drought on society consists on shortages or limitations in the availability of the water resource and the failure of water supply for different uses: the worsening of agricultural production, the decrease in energy or industrial production, problems in the supply of drinking water, or limitations in any recreational or ornamental use of water (Eslamian, et al., 2017; Mishra & Singh, 2010).~~

~~All drought types can be characterised by several items: (1) severity (i.e., expressed through the rainfall values themselves as well as through ceremony levels of *pro pluvia* rogations), (2) duration (i.e., from onset to end), (3) spatial extent (i.e. area of impact) and (4) frequency of occurrence (Nalbantis, 2008). Long droughts can cause serious hydrological imbalance gradually increasing its severity (Wilhite & Glantz, 1985). While magnitude, duration and recurrence are necessary drought features to assess the physical impacts of droughts on a territory, the vulnerability of the society in relation to degree of exposure and strategies to cope with the physical hazard are fundamental for a comprehensive evaluation of climate risk. Beyond contributing to direct water scarcity, droughts affect agriculture, hinder the production of energy, the access to fresh water and may aggravate political tensions connected to water~~

rights (Gorostiza et al., 2021). Moreover, the socio-economic impacts of drought generally persist even when the episode of meteorological or hydrological drought ends and the volume of rainfall returns with regularity (IDMP, 2022).

Despite the importance of droughts and their capacity to seriously affect the economic and productive activities of societies, the level of knowledge on this natural phenomenon contrasts with that of other natural hazards (Van Loon & Van Lanen, 2012). For these reasons it in this study is justified to conduct a more detailed and systematic study of drought events with historic perspective. It will also take into account the analysis of specific episodes of lower frequency and greater severity, which may provide additional information on long term drought behaviour in the Mediterranean region (Olcina, 2001a; Olcina, 2001b). Of particular interest are those that have occurred within the framework of the Mediterranean, where drought is an intrinsic phenomenon of the climate of the region (Olcina, 2001b).

In general terms the Mediterranean climate in the Iberian Peninsula is characterised by a highly irregular rainfall, both inter-annual and intra-annual (Martín-Vide & Olcina, 2001). Another characteristic is the pronounced aridity during the warm season (summer) (WMO, 2023). Additionally, presents important variations in the intra-annual distribution of precipitation depending on the region (Martíin-Vide, 1985). On the eastern of Iberian Peninsula and Balearic Islands, Mediterranean climate type has two main varieties in relation to the seasonal distribution of rainfall:. Autumn and spring maximums in Northern and Central sectors, autumn-winter maximums in Southern sector Between the provinces of Girona and Almería, main rainfalls are in autumn, with a secondary peak in spring. In the southernmost provinces (coast of Granada, Málaga and Cádiz) the most abundant rains are recorded in the autumn and winter months (AEMET, 2011). Autumn and spring rains are mainly linked to cold drop atmospheric situations (Sánchez-Almodóvar et al., 2022). In the spring, along with this cold drop configuration, convective rains (storms) take on a prominent role. In the winter, Atlantic storms with a southern trajectory (Gibraltar strait) are frequent in the south of the Mediterranean strip. In either of these two varieties, summer is always the season with the lowest rainfall contribution (Serrano-Notivoli et al., 2018; Mathbout et al., 2020; Sánchez-Almodóvar et al., 2022). Annual quantities decrease from north to south between Girona and Almería (Serrano-Notivoli et al., 2018). In the southern section of this region, rainfall increases again due to the greater contribution of winter rains (Mathbout et al., 2020). For this reason, spring and autumn are the key seasons that balance the annual water input. Strong droughts occur when the summer and winter lack of precipitations connect due to an extraordinary lack of rainfall on the rainy seasons (autumn and spring). These seasonal aspects determine a high temporal variability of water reserves on the Northern Spanish Mediterranean Basin (Kim & Raible, 2021). In addition, this high vulnerability is magnified by the eventual impacts caused by droughts (González-Hidalgo, et al., 2018). In this regard, together with drought, water management in the Mediterranean region have always been a challenge, but now it is exacerbated within the context of climate change (Hohenthal & Minoia, 2017). Additionally, serious problems have derived from greater water demands in result of population increase and the spread of a lifestyle model based on mass consumption of goods and services (IDMP, 2022). In this respect, the Mediterranean region is a clear example of imbalance between water demand and water availability. As a result, in recent decades it has become one of hotspots areas impacted by climate change. Along with the already detected temperature increase, since the beginning of the 21st Century there is also the added challenge of increased rainfall variability (Barrera-Escoda & Cunillera, 2011). In this context, along with the increased rainfall irregularity, extended dry periods occur with greater frequency and severity (Marcos-García, et al., 2017; Kim & Raible, 2021). Therefore, drought in the Spanish Mediterranean Basin is one of the natural risks with the greatest impact, due to its capacity to cause simultaneous effects on different levels: environmental, economic, social, etc., and also its capacity to last for a long time (Walker et al., 2010).

Because of the impacts of extreme hydrometeorological phenomena in the Mediterranean, such as droughts and floods, observation of their behaviour in the recent past is justified. Previous work on the reconstruction of rainfall on a long-time scale generally shows situations of rainfall shortage (Pauling, et al., 2006; Camuffo, et al., 20123; Smerdon, et al.PAGES, 2017). As well the results obtained from instrumental data series highlight that rain shortage (droughts) are perceived at the seasonal level. The present work aims to analyse the most extreme phenomena detected in the aforementioned paper, but using data that allow us to analyse rainfall deficits at the monthly level. Unfortunately, for the Iberian Peninsula as a whole in the study under review there is only one instrumental data series. For this reason, the instrumental data from the Barcelona station (1786-2022) complement the historical data from the MILLENIUM project ("European climate of the last millennium", Code: IP 017008-2). The combination of instrumental and historical data has been used to study specific periods of anomalous temperature and rainfall conditions. One such widely studied case is the anomaly of the non-summer year of 1816, a consequence of the Tambora volcanic eruption (Trigo, et al., 2009; Luterbacher & Pfister, 2015).

The Spanish Mediterranean basins are currently experiencing a situation of severe rainfall shortage. Due to this serious situation, it is necessary to find references of droughts of equal or greater magnitude in order to understand the characteristics of these phenomena in their most extreme behaviour. Studies carried out on the Iberian Peninsula to study historical droughts using historical data show significant results obtained from the use of rogations as a data (Domíinguez-Castro et al., 2012; Tejedor et al., 2019). These studies make it possible to identify the importance of the 19th century for its study, highlighting specific years such as 1817 or 1824 (Domíinguez-Castro et al., 2012). Despite these results, the data used in these studies were applied to yearly resolutions. The need for knowledge of past droughts adds to the need to expand the detail of existing studies on historical droughts in the study area.

Historical data allow us to observe the behaviour of droughts in much more distant historical periods than those of the instrumental precipitation data series. Therefore, this data would allow us to improve the knowledge of drought natural variability over a long-time scale than the instrumental period. Also, this longer timescale would help to study drought return periods on centennial scale (i.e. lower frequency droughts) and the duration and magnitude of past extreme droughts (note that only a handful are available during the instrumental period). In the case of droughts, it is crucial to know those episodes which occurred in the past and whose severity, extent and duration were exceptional For periods where overlap exists, historical data (rogations ceremonies) can be correlated with early instrumental data. This aspect is novel and important for three main reasons, which motivate focusing this work on such analysis: 1) the existence of long instrumental records is scarce and spatially dispersed in Spain; 2) social changes during the 19th Century make historical records of social impact, based on rogations, demonstrably inconsistent after 1836 (Gil-Guirado et al., 2016; Espín-Sánchez & Gil-Guirado, 2022), which discourages their use as proxies after this date; 3) the available instrumental records and social impact data are contemporaneous.~~(Gil-Guirado, et al., 2016).~~

According to all the reasons exposed above, in the current paper we will discuss the topic of the extreme droughts that affected the Mediterranean Basins of the Iberian Peninsula during the Early 19th Century (1790-1830). The detailed study of drought events during this period is justified by the physical and social reasons that underline their exceptionality. The severity of the different droughts recorded, their cumulative duration and the impact they had on the societies of the Spanish Mediterranean Basins do not have an equal magnitude in the recent collective memory. On the other hand, this period has been studied relatively well, thanks to climate reconstructions for the beginning of the 19th Century based on natural and historical *proxy* data and the first instrumental meteorological data series (Prohom et al., 2016; Brönnimann et al., 2018b~~; Prohom et al., 2016~~).

The novel aspect of the present study consists mainly in the fact that the period chosen (1790-1830) has not been analysed in depth with historical data. Furthermore, it has not been analysed with daily resolution data, as is the case in the present study. This paper focuses on the impacts caused by meteorological droughts because of the nature of the data used. The main sources of information used for the analysis of droughts in the historical period are the historical data of rogations (Spain, with higher density for Catalonia) and the instrumental precipitation data sets of Barcelona (Catalonia, NE Spain). ~~and the historical data of rogations (Spain, with higher density for Catalonia).~~ The case of the rogations differs from that of the instrumental series, since the former focuses on the lack of precipitation while the rogations would allow the analysis of agricultural drought (Brázdil et al., 2018). However, rogations also allow meteorological monitoring of the natural phenomenon, as the ceremonies itself are interrupted when an improvement in rainfall is detected. This is because of the daily level of detail of the rogation system as a source of information (Martín-Vide & Barriendos, 1995). The very etymology of the rogations (*pro pluvia*, to obtain rain as usual) demonstrates the meteorological nature of the ceremony. Their purpose was not directly to obtain a large harvest, but to achieve a good rainfall episode. ~~-~~The historical data at daily resolution used to carry out this study come from the AMARNA database on climate risks (*Arxius Multidisciplinars per a l'Anàlisi del Risc Natural i Antròpic*, from catalan: Multidisciplinary Archives for the Analysis of Natural and Anthropogenic Risk)~~-~~. This is a compilation effort focused on organising climate information from historical proxies in high spatio-temporal resolutions. The total number of records for the period EC 1035-2022 amounts to slightly more than 19,000 cases, organised in more than 5,500 episodes (Tuset et al., 2022). It is originally a database of torrential rainfall and flood events and from the present study, information on pluviometric deficits is also being introduced.

1.1. Research background

The Early 19th Century (1790-1830) occurred during the climate episode named as the Little Ice Age (hereafter, LIA) between the fourteenth and nineteenth centuries (Grove, 1988). This climate oscillation was clearly characterised by lower average temperatures with respect to the previous episode (Medieval Warm Period) and the subsequent episode (Current Global Warming) (Fischer et al., 2007). Another significant aspect of the LIA is the irregular behaviour of rainfall, with a clear increase in the frequency and magnitude of severe hydrometeorological events (Oliva et al., 2018; Barriendos et al., 2019~~,~~; Gorostiza et al., 2021~~ Oliva et al., 2018~~). In the case of the Iberian Peninsula, different oscillations were observed including increases in heavy rains or droughts throughout this period (Barriendos, 1996). One of the most exceptional oscillations is called Maldà Oscillation, which occurred between 1760 and 1800 (Barriendos & Llasat, 2003). The Maldà Oscillation was characterised by simultaneous increases in the frequency of heavy rain events, alternating with droughts. The alternation of extreme rainfall and droughts events had strong social and economic impact on the Iberian Peninsula. Specifically, the sequence of droughts, cold snaps and snowfalls had serious direct consequences on agriculture, while consecutive floods also damaged or destroyed many infrastructures. Furthermore, during the period of the Maldà Oscillation there was an emergence of uncommon epidemic diseases, such as smallpox or yellow fever viruses, occurring at the same time than more common diseases such as epidemic malaria or typhoid (Barriendos & Llasat, 2003; Alberola, 2010; Alberola & Arrioja, 2018).

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Within the LIA, the Early 19th Century was characterised by an abnormally low amount of emitted solar radiation, which generated an overall decrease in the amount of solar radiation arriving to the Earth (Prohom et al., 2016). In addition to this external forcing factor, climate variability at the end of the LIA was also affected by several volcanic eruptions that occurred between 1790 and 1830 (a total of 302 eruptions with Volcanic Explosivity Index (~~herafter~~hereafter, VEI) between 2 and 7) (Fang et al., 2023). Of these 302 eruptions, 247 had an VEI ≥ 2; 35 had a VEI ≥ 3; 16 had a VEI ≥ 4; 2 had a VEI ≥ 5; 1 had a VEI ≥ 6 and 1 had a VEI ≥ 7 (Global Volcanism Program, 2023). Among these volcanic eruptions, stand out a sequence of large explosive volcanic eruptions (Prohom, 2003; Wagner & Zorita, 2005~~; Prohom, 2003~~; Lee & MacKenzie, 2010): *Unknown* (1808), Tambora (1815), Galunggung (1822) and Cosigüina (1835). Some studies indicate that the high intensity volcanic eruptions, occurring between the LIA and the current Global Warming, led to a decrease in temperatures, together with an increase in rainfall irregularity in the study area (Gil-Guirado et al., 2021~~2020~~).

Among the three eruptions of the Early 19th Century, the 1815 Tambora eruption is considered one of the most significant of the past two thousand years in terms of the particles emitted (Raible~~,~~ et al., 2016). Also, it is considered as the cause of the most pronounced climate anomaly of the first third of the 19th Century (Brönnimann et al., 2018b). Due to his outstanding volcanic explosivity (VEI 7), this eruption was the largest and most devastating eruption recorded in the historical age and is considered to be responsible for the "year without a summer" of 1816 reported across Europe and North America (Trigo et al., 2009; Luterbacher & Pfister, 2015). This temperature anomaly ~~afected~~affected severely Central~~ Europe~~, Western, ~~Europe~~ and Northern Europe, with recorded temperatures ~~recorded of~~ between 2 to 3ºC below the average in areas of Spain and Portugal (Pfister & White, 2018). ~~were the greatly affected regions by this "year without a summer".~~ During that summer the number of rainy days almost doubled and cloudy days were more frequent in the whole of Europe and North America. Alterations in the usual general atmospheric circulation pattern and its centres of action were also reported as a result of cooling due to the direct effect of the reflection of incident radiation associated to the presence of volcanic aerosols (Brönnimann et al., 2018b).

~~This study has found a time period in which there is an accumulation of particularly severe drought episodes. This period coincides chronologically with Dalton Solar Minimum and an anomaly in~~

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~~volcanic activity (eruptions of Tambora and other volcanoes mentioned). Obviously, the chronological coincidence does not presuppose any cause-effect relationship between the anomalies in solar and volcanic activity and the pluviometric anomalies under study.~~

1.2. Historical Droughts Studies in Spain

The analysis of historical droughts in Spain dates back to studies by Manuel Rico y Sinobas in the mid-19th Century, in which he analysed the impacts of drought episodes on agriculture. His main objective was to compile records in order to obtain a broad temporal dimension of the phenomenon (Rico y Sinobas, 1851). Subsequently, and until the beginning of the 1990s, only sporadic studies were carried out that were in some way related to events (Bentabol, 1900). One exception is the study by Couchoud (1965), who analysed the region of Murcia in depth (SE Spain) based on a detailed compilation and analytical process. In 1994, two PhD theses on historical climatology that engaged with droughts were defended in Spain (Barriendos, 1994; ~~Sánchez~~ Rodrigo, 1994). They constitute benchmark studies in the research on this topic. From this decade onwards, there has been a proliferation of studies and publications in which drought is taken into consideration (see, among other, ~~Sánchez~~ Rodrigo~~,~~ et al., 1994; Martíin Vide & Barriendos, 1995; ~~Sánchez~~ Rodrigo~~,~~ et al., 1995; Barriendos, 1997; Barriendos & Martíin Vide, 1998; ~~Sánchez~~ Rodrigo, et al., 1998), including manuals on natural risks (Olcina, 2001a). More recently, a new PhD thesis (Gil-Guirado, 2013) once again insisted on the need to study historical droughts in the Spanish Mediterranean Basin based on a quantitative approach.

In addition to PhD theses, there are also recent publications focused on the study of historical droughts using a quantitative approach. An example that actually corresponds to the period analysed in present work is the paper focused on droughts for the Iberian Peninsula (1750-1850) (Domíinguez-Castro et al., 2012). This article approaches the severe episodes of historical droughts by means of rogations at annual resolution. Other studies have continued this line of research in the Iberian Peninsula (Trigo et al., 2009; Fragoso et al., 2018; Tejedor et al., 2019; Bravo-Paredes et al., 2020) and even in more detail for the Ebro basin (Cuadrat et al., 2022). The availability of *pro -pluvia* rogations in the Hispanic Monarchy extended beyond the Iberian Peninsula, as evidenced by works in Mexico and all Central American countries (Garza-Merodio, 2017; Alberola & Arrioja, 2018; Ramírez-Vega, 2021). Rogations are a liturgical mechanism used in other Catholic countries and therefore these studies can be extended to this broader religious sphere (Pfister, 2018; Garnier, 2019~~; Pfister, 2018~~). Finally, the amount of information that is becoming available is already being organised in comprehensive databases such as AMARNA or in international initiatives (Domínguez-Castro et al., 2021).

Parallel to the research based on rogations, the study of historical droughts in the Iberian Peninsula has also been carried out through the analysis of ancient instrumental precipitation data series as well (Prohom et al., 2016). Or with the combination of data on rogations and precipitation series analysed by means of drought indices (Tejedor et al., 2019). These studies allow us to observe severe droughts based on inter-annual variability.

1.3. Objectives

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The main objective of this study is to analyse the patterns of drought episodes that affected the Northeast of the Iberian Peninsula during the Early 19th Century (1790-1830) using instrumental and historical sources. This period that corresponds to the last stages of the Little Ice Age was chosen due to severity of drought occurring in the Mediterranean Basins of the Iberian Peninsula. Additional objectives of this study are: 1) to qualitatively and quantitatively extend the AMARNA database on climate risks ~~(*Arxius Multidisciplinars per a l'Anàlisi del Risc Natural i Antròpic*, from catalan: Multidisciplinary Archives for the Analysis of Natural and Anthropogenic Risk)~~ to incorporate droughts and different social processes linked to environmental impact in addition to hydro-meteorological excesses (Tuset et al., 2022); 2) to compile and describe the variability of extreme hydrometeorological events (heavy rainfall and droughts) in the Spanish Mediterranean Basin during the Early 19th Century. In order to study how the opposite extreme events behave and interact with each other. Also to understand if the behaviour of past hydrometeorological extremes is similar to the modelled behaviour for the future in the study area. In addition, the spatio-temporally coherent periods of climatic anomalies have among their main characteristics the increase in rainfall irregularity in the study area (Gil-Guirado et al., 2016); 3) to characterise the drought episodes, analysed from historical data, considering their duration, extension and severity in high resolution for the period analysed; and 4) to analyse the entire instrumental precipitation data series of Barcelona (1786-2022) for the whole duration of the series in order to characterize periods of drought.

In order to fulfil these objectives, the paper analyses the historical and instrumental data available in the Spanish Mediterranean Basins, using different time and spatial scales. The socio-environmental context during the Early 19th Century is analysed using data compiled from historical documentary sources, namely the records of the *pro pluvia* rogation ceremonies held in the main villages of the affected regions. These data are compared with the analysis of the instrumental precipitation data series of Barcelona (1786-2022) based on different statistical techniques, including the use of three drought indexes: SPI (*Standardized Precipitation Index*) (McKee et al., 1993), SPEI (*Standardized Precipitation Evapotranspiration Index*) (Vicente-Serrano et al., 2010) and Deciles (Gibbs & Maher, 1967).

The article focusses on analysing climate variability during the Early 19th Century period and provide the state of the art on droughts in historical perspective in Spain and Europe as a whole. Subsequently, the results obtained are presented through graphic and cartographic resources.

2. MATERIALS AND METHODS

2.1. Sources of information

The sources of information used to analyse droughts in the Early 19th Century consist mainly of historical data and the Barcelona instrumental precipitation data series ranging from 1786 to the present day. The historical data on droughts in the Spanish Mediterranean Basin during the Early 19th Century was obtained from Documentary sources of public administrations and ecclesiastical institutions complied in the AMARNA database (Barriendos & Barriendos, 2021; Tuset et al., 2022). AMARNA is an archive that compiled climate historical episodes from different documentary sources which are geo-referenced

Con formato: Sangría: Primera línea: 1,25 cm and classified into numerical categories on a daily resolution. The information from AMARNA refers to any type of extreme meteorological event and its social impacts. Events about which there is more information are those relating to water excess (persistent rainfall, pluvial and fluvial floods) and rainfall deficits (droughts). The total number of records for the period EC 1035-2022 amounts to slightly more than 19,000 cases, organised in more than 5,500 episodes (Tuset et al., 2022). Sources of information are mainly administrative and private documentary sources, with direct descriptions of events and their impacts. The institutional documentary sources also provide systematic and continuous records over time throughout the existence of the institution, with resources and conditions that favour the conservation and access to the documents (Martín-Vide & Barriendos, 1995; Brönnimann et al., 2018a). Water deficits are obtained from the records of *pro pluvia* rogation ceremonies (cultural-historical proxy) from municipal and local ecclesiastical sources (Brádzil,Brázdil et al., 2018; Brázdil et al., 2019). Rogations are the main *data proxy* in order to identify and compile information on droughts in the Spanish Mediterranean Basin. The records of these ceremonies are generated and initiated by public authenticators in collegiate administrative bodies (municipal councils, cathedral councils), which guarantees the reliability of the document itself and the veracity of the information contained therein. The rogation records contain reliable and homogeneous information due to their institutional origin and the formal rigidity of the related liturgical procedures (Brádzil,Brázdil et al., 2018). The documentary record of the rogation ceremony informs of the location, the date and duration of the drought conditions. With respect to the severity of the event, the application of a specific methodology based on the type of liturgical acts used enables their classification by categories and their numerical indexing (Martín-Vide & Barriendos, 1995; Barriendos, 1997). As a complement to these administrative sources, AMARNA also uses private personal sources, such as appointment books, memoirs or chronicles.

Rainfall excesses are also found in the same administrative documentary sources as the deficits and their cataloguing and numerical classification procedure is also based on objective indicators. In the 1990s, simple and easy to cross reference classification criteria were proposed for all of the European basins, based on the levels of river overflows and the damage recorded (Barriendos & Martín-Vide, 1998; Brázdil et al., 1999). The first studies that used these information sources in the area of study sought to conduct an overall reconstruction of the climate variability through the generation of weighted annual indices (Barriendos, 1996; Barriendos, 2005). Subsequent studies extended the analysis with annual indices for different locations of the Spanish Monarchy, both on historical floods (Barriendos & Sánchez Rodrigo, 2006) and for droughts (Domínguez-Castro et al., 2008; Domínguez-Castro et al., 2012; Sánchez Rodrigo & Barriendos, 2008,; Domínguez-Castro et al., 2012; Gil-Guirado et al., 2019; Tejedor et al., 2019; Gil-Guirado, et al., 2019).

In addition to the analysis of historical data, the second part of the study consists in the statistical analysis of the instrumental precipitation data series of Barcelona spanning from 1786 to 2022. Unfortunately, the Barcelona set series is the only continuous rainfall series available in the study area for the Early 19th Century. This information is scarce for such a large geographical area, but the Barcelona series is located in the area with the most historical information available for this period. Therefore, the joint analysis of instrumental and historical information is relatively consistent.

The Barcelona rainfall series used in this study comes mostly from the series ~~has been compiled by consulting several documentary sources and different sets of records. The principal source was~~ elaborated by the Meteorological Service of Catalonia, *Servei Meteorològic de Catalunya* (SMC) (Prohom et al., 2016). This series ranges from 1786 to 2014 and was compiled from different institutional observers who generated records during the 18th and 19th centuries in the centre of Barcelona (at around 30 meters above sea level). For the 20th Century the records were generated at the Fabra Observatory, placed outside of the city (at the Tibidabo mountain, at 412 meters above sea level). The analysis of these sources has enabled the homogenisation of the monthly precipitation data ~~sets~~series from 1786 to 2014. For the period 2015-2022 we used rainfall series located in the city of Barcelona, instead of continuing with the series from the Fabra Observatory for the following reasons: it corresponds to an altitude significantly different from that of the flat coastal area of Barcelona; ~~also~~also, the Fabra Observatory is placed far from the city ~~center~~centre; and finally, it began its record measurements in 1913. Taking into consideration the high irregularity of the precipitation in the Mediterranean climate, these ~~diferences~~differences ~~makes~~make advisable to use a landmark closer to the area of the historic centre of the city of Barcelona. These considerations have been the subject of debate for years when defining climate instrument series for Barcelona (Prohom et al., 2016). To complete the SMC series up to the year 2021, instrumental records from a private observatory in the Can Bruixa neighbourhood of Barcelona was used. In order to complete the remaining year, the series was completed with ~~For the year 2022,~~ data ~~was~~ from the official SMC Raval automatic station (University of Barcelona). These two ~~sets~~series are validated by the SMC and their data have been collected in the centre of the city of Barcelona, making their values closer to those collected at the beginning of the series.

2.2. Indexation system of historical climate data

This study is based on the use of information on a daily scale drawn from the historical data obtained from the AMARNA database. This information is organised into cases and episodes. Every episode consists of a group of cases or records of different dates and locations which provide information about the impact and duration of each episode. Cases are the basic units of documentary record in which there is mention of some kind of impact on the water deficit. They may be decisions by the authorities to initiate or continue *pro pluvia* rogations, qualitative records of rainfall within a drought episode, or records of the decisions taken by the authorities to end the rogations once the drought is considered to be over. The cases and episodes are classified into five categories and fifteen sub-categories (Barriendos & Barriendos, 2021) (Table 1). These fifteen thematic subdivisions proposed correspond to the highest degree of detail observable in the documentary and bibliographic sources consulted (Table 1). For the specific case of drought episodes (DR), these come mostly from records of the celebration of *pro-~~ ~~pluvia* rogation ceremonies. These liturgical events, typical of the Catholic Church, are highly institutionalised. The adverse weather situation is detected by the monitoring and reports of the farming guilds. Their evaluations are assessed by the municipal councils which, in view of the severity of the situation, decide to hold the prayers. And finally, the ecclesiastical authorities are responsible for the effective execution of the ceremonies. This procedure has the positive factor of having administrative documentary sources generated by public administrations that guarantee the accuracy of the information, its objectivity and the

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reliability of its contents. It should be noted that the administrative documentation consulted is always validated by public authenticators (secretaries of municipal and ecclesiastical councils who are public notaries). These records provide information on both the duration and severity of drought events. From the documentary recording mechanism of the ceremonies, it is also possible to detail that the rogations present a formal differentiation of their liturgical acts according to the severity of the drought determined by the specialised guild authorities. This liturgical format has remained almost unchanged since the ~~middle ages~~Middle Ages, when continuous municipal and ecclesiastical records were already available. The difference between ceremonies is based on their format, with a total of five levels, always adapted to the cultural singularities of each town (Martín-Vide & Barriendos, 1995; Barriendos, 1997; Tejedor et al., 2019; among other references: Alcoforado et al., 2000; Gil-Guirado et al., 2019; Espín-Sánchez & Gil-Guirado, 2022): 1- Simple prayers inside the churches; 2- Prayers using the exhibition of relics or images inside the churches; 3- Public processions through the public area of the town (planned routes through the main streets of the town); 4- Until 1619, immersion of images or relics in water. From 1619 onwards, due to prohibition of immersions by the Vatican authorities, liturgical acts of similar solemnity were held in public spaces within the town's boundaries; 5- Pilgrimages to sanctuaries of special veneration that required a journey outside the town. By having the dates on which each ceremony is held, we can identify both the beginning and the end of the rogations, along with increases in severity. ~~The type of ceremonies held in order to ask for rain will define the severity of the episode, which are organised between 1 and 5 (Martín-Vide & Barriendos, 1995; Barriendos, 1997).~~ A level 1 rogation marks the beginning of each drought episode and a *Te Deum Laudamus* (gratitude ceremony) marks the end of the episode. Each drought episode will thus have a different duration and its severity will be defined by the ceremonies between the first rogation and the closing of the ceremonies.

CATEGORIES		SUB-CATEGORIES	
Code	**Name**	**Code**	**Name**
ERE	Extraordinary Rainfall Event	FF	Fluvial Flood
		PF	Pluvial Flood
		PR	Persistent Rainfall
		SS	Sea Storm
		DR	Drought
ECE	Extraordinary Convective Event	HE	Hail Event
		ES	Electric Storm
		WS	Wind Storm
ETE	Extraordinary Thermic Event	CW	Cold Wave
		US	Unusual snowfall
		HW	Heat Wave
SIE	Social Impact Event	EE	Epidemic Event
		PE	Plague Event
		FS	Food Shortage
ERR	Technical mistake	ERR	Spurious case

Table 1: Classification system of the AMARNA database (Barriendos & Barriendos, 2021).

The AMARNA database originally only provided data on water excesses recorded in historical periods for the Spanish Mediterranean basins (Tuset et al., 2022). An effort is currently being made to add data on droughts to the AMARNA database. In this regard, the period from 1790 to 1830 has been a test to see how the recently gathered data on droughts and the existing data on excess water fit together. Thus,

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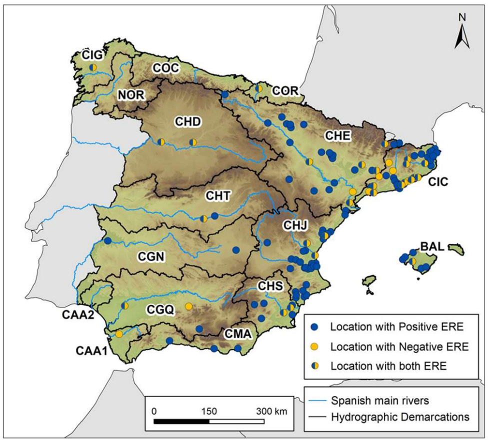
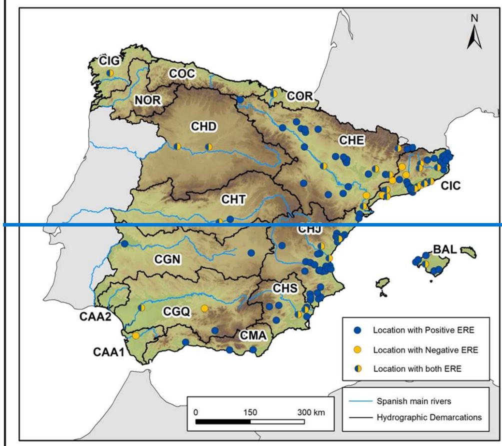
the work proposed in this article supposed progressing from 0 cases and episodes of drought events for the Early 19th Century, to the values with which the study has been carried out (2047 cases, of which 1132 cases correspond to drought events). The AMARNA database is still under development for other historical periods and therefore not yet available for public access.

The georeferencing of all the historical data compiled in the AMARNA database allowed the use of SIG tools for the cartographic representation of this historical information. The distribution of the droughts in the Early 19th Century have been represented both on a municipal level and with the cases grouped by hydrographic basin. These are the Spanish administrative units for managing water resources (Figure 1) (MITECO, 2023). The organization at a municipal level allow the analysis of the time-space distribution of the impacts caused by different drought episodes representative of the period of study. The different efforts to compile data on the AMARNA database on water excesses and droughts have resulted in a very characteristic distribution of data for the case of the Early 19th Century period (Figure 1). Most of the points with information on water excesses collected in AMARNA are located in the Spanish Mediterranean basins. On the other hand, the information on droughts covers points all over Spain, but with a higher density in the territory of Catalonia, between the hydrographic basins of the Ebro River (CHE) and the Catalonia Coastal Rivers (CIC)~~) hydrographic basins~~, and Murcia city (CHS) (Figure 1). This disproportion in the amount of information between the Atlantic and Mediterranean basins is due to the effort focused on the latter, where there is more interest in the study of hydrometeorological phenomena. Therefore, as the title of the paper indicates, the analysis of drought on the Noth-E~~e~~ast Iberian Peninsula uses information from the Atlantic basin of the Iberian Peninsula only as a reinforcement or complement for a better characterisation of the episodes identified.



Basin Code	Basin Name
CIG	Galicia Coastal Rivers
NOR	Northwest (Miño River)
COC	Northern Rivers, Western Side
COR	Northern Rivers, Eastern Side
CHD	Duero River
CHT	Tajo River
CGN	Guadiana River
CGQ	Guadalquivir River
CAA1	Andalucia Atlantic Rivers (Guadalete and Barbate)
CAA2	Andalucia Atlantic Rivers (Tinto, Odiel and Piedras)
CHE	Ebro River
CIC	Catalonia Coastal Rivers
CHJ	Júcar River
CHS	Segura River
BAL	Balearic Islands
CMA	Andalucia Mediterranean Rivers



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CAA2	Andalucia Atlantic Rivers (Tinto, Odiel and Piedras)
CHE	Ebro River
CIC	Catalonia Coastal Rivers
CHJ	Júcar River
CHS	Segura River
BAL	Balearic Islands
CMA	Andalucia Mediterranean Rivers

Figure 1: Spanish hydrographic basins analysed in this study. Locations with historical information for the Early 19th Century. This specifies the locations that have records on positive ERE (FF, PF, PR and SS), negative ERE (DR) or both types of ERE.

2.3. Generation of drought indices

Several drought indices were generated using the Barcelona precipitation data series (1786-2022). In all cases, the indexes were calculated based on monthly values and for groups of 12 months. Due to the irregular distribution of precipitation throughout the year in the North-East of the Iberian Peninsula, the 12-month groupings are the ones that best group and detect drought episodes. Other groupings such as 3 or 6 months can detect a lack of precipitation that is typical of the usual conditions of the seasons and intra-annual variability (Gil-Guirado & Pérez-Morales, 2019). The SPI (*Standardized Precipitation Index*) (McKee et al., 1993) was the first index calculated, which is widely used for classifying droughts (WMO & GWP, 2016). This index enables the analysis of the duration and variability of droughts, as well as of the wet periods and is generated based on the transformation of the temporal precipitation data series in a standardised normal distribution (Lloyd-Hughes & Saunders, 2002; Zargar et al., 2011; Gil-Guirado & Pérez-Morales, 2019~~; Zargar et al., 2011~~). The second index is the SPEI (*Standardized Precipitation Evapotranspiration Index*) (Vicente-Serrano et al., 2010), which is similar to the SPI index, but also uses the average monthly temperature variable (WMO & GWP, 2016). It is a relatively versatile index, simple to apply and enables analyses to be carried out for any climate regime (Stagge et al., 2015). The third index used is the Deciles index (Gibbs & Maher, 1967), which stands out for its applicability and simplicity, due to the facility of the calculations that it requires and the fact that it only requires precipitation data (Steinemann et al.,~~Hayes & Cavalcanti,~~ 2005; Tsakiris~~,~~ et al., 2007). This method is obtained by dividing the distribution of the monthly precipitation data into deciles (WMO & GWP, 2016), which define thresholds for different water deficit conditions (~~Eslamian et al., 2017;~~ Zargar et al., 2011; Eslamian et al., 2017).

The results obtained from analysing the instrumental rainfall series of Barcelona (1786-2022) with the different indices (SPI, SPEI and Deciles) have been statistically analysed. ~~Statistical analyses were conducted with the results of the three indices.~~ Three different statistical tests have been carried out with the monthly rainfall series of Barcelona (Gil-Guirado & Pérez-Morales, 2019): On the one hand, the trends of the series have been calculated using the Mann Kendall test~~;~~. On the other hand, the Sen slope has been obtained~~;~~. Finally, the breakpoints of the series have been analysed using the Pettitt's test. ~~Testing of trends was carried out using the Mann Kendall test and with the Sen slope. Analysis of breakpoints of the monthly series was conducted using the Pettitt Test (Gil-Guirado & Pérez-Morales, 2019).~~ In order to carry out these statistical tests, different scripts in R language have been used, which have been executed in RStudio to obtain the results.

Based on the results obtained from various drought indices, a detailed criterion has been established to classify the different drought episodes identified for the early 19th Century. This criterion relies mainly on the SPI to define each episode, based on the drought thresholds defined by the literature (McKee et al., 1993). Specifically: The start and end of a drought episode are determined by SPI values that cross the threshold of -0.7~~5~~0. This threshold is chosen to capture the transition between drought episodes. Also noticing the transition periods into and out of drought conditions. A drought episode is characterised by having at least ~~six~~five consecutive months in which SPI values are consistently below -1.0, indicating moderate to severe drought conditions. By defining these specific ranges, we ensure a systematic and reproducible approach to identifying and analysing drought episodes.~~Based on the results obtained in the different drought indices, a criterion has been defined to organise the different drought episodes detected for the Early 19th Century. This criterion takes the values obtained with the SPI as a reference to define each episode. Each drought episode must have at least six consecutive months with SPI values lower than "-1". The count of the total number of months of the drought episode starts and ends when the SPI values are below "-0.75".~~

3. RESULTS

3.1. The hydro-meteorological extremes in Spain (1790 - 1830)

This study has found that the a time periodperiod (1790-1830) in which there is an accumulation of particularly severe drought episodes. This period coincides chronologically with the Dalton Solar Minimum and an anomaly in volcanic activity (eruptions of Tambora and other volcanoes mentioned). Obviously, the chronological coincidence does not presuppose any cause-effect relationship between the anomalies in solar and volcanic activity and the pluviometric anomalies under study.

The AMARNA database used in this paper provides a total of 19115 cases spread over 5,551 episodes for the period from 1035 to 2022. For the Early 19th Century (1790-1830), the AMARNA database provides for the whole of the Iberian Peninsula 2047 cases, which are grouped into 708 episodes (Barriendos et al., 2019).

From the 2047 total number of cases 1789 cases correspond to ERE events (Extraordinary Rainfall Event). Within the ERE cases, there is a clear predominance of the subcategory DR (Drought), with 64% of the ERE cases (Table 32).

Subcategories	Number of cases	Percentage
Fluvial Flood (FF)	431	24.09%
Pluvial Flood (PF)	40	2.24%
Persistent Rainfall (PR)	164	9.17%
Sea Storm (SS)	22	1.23%
Drought (DR)	1132	63.28%
Total	1789	

Table 32: Total number of cases of the five groups making up the ERE category (Extreme Rainfall Event). Elaboration from AMARNA database.

The temporal distribution of the ERE episodes throughout the Early 19th Century reveals a predominance of droughts with respect to the other types of ERE, but with a non-homogeneous distribution (Figure 2). For instance, between 1790 and 1800 1805 rainfall was abundant, so floods were more significant than droughts in years such as 1793, 1797 or 1801 (Figure 2). This decade also stands out due to its clear irregularity across different years, which can be related to the final part of an abnormal climate period detected between 1760 and 1800, known as the Maldà Oscillation (Barriendos & Llasat, 2003). The 5-years moving averages show the most pronounced episodes of droughts and water excesses during this period. Figure 2 highlights its temporal distribution: in the first decade, positive extreme peaks were interrupted with the drought of 1798. On the other hand, from the episode of 1807, droughts became predominant, being particularly severe between 1812 and 1825 (Figure 2). The positive EREs cases diminished from 1806 definitively for the rest of the Early 19th Century, while the negative EREs increased from 1812. Between these two well defined periods exists a transition period with low number of heavy rainfalls or droughts.

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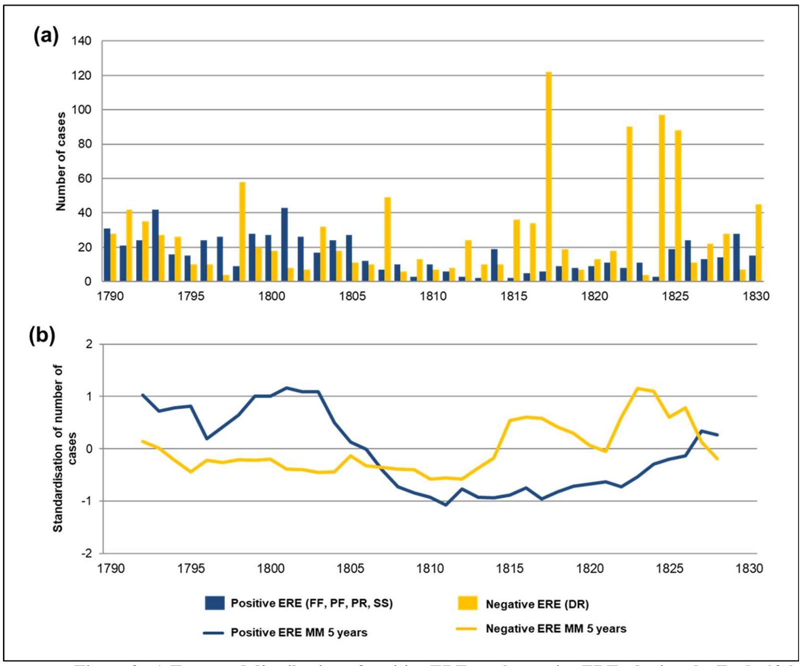
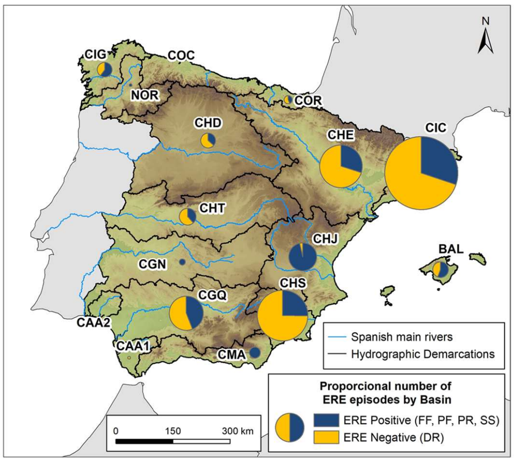


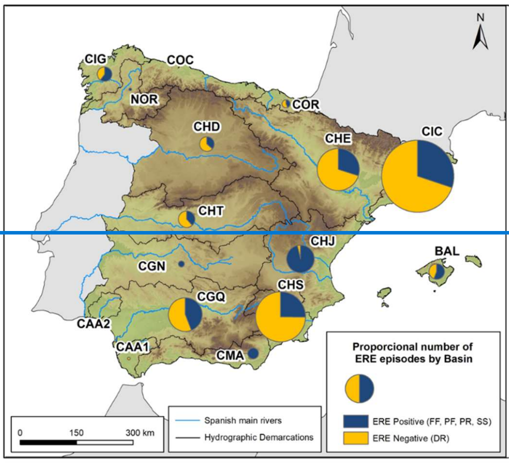
Figure 2: a) Temporal distribution of positive EREs and negative EREs during the Early 19th Century (1790-1830). b) ~~Annual cases of positive EREs and negative EREs during the Early 19th Century (1790-1830).~~ 5-years moving averages of the standardised values for the positive EREs and negative EREs. Elaboration through AMARNA database.

The geographical distribution of ERE cases for this period also provides interesting information. It is highlighted the large number of cases recorded in the Spanish Mediterranean Basin against those recorded in the Atlantic basins for the same period (Figure 3). The Guadalquivir basin (CGQ) is the only Atlantic basin with an important amount of ERE cases. The predominance of drought in the Spanish Mediterranean basins contrasts with the greater impact of the positive ERE episodes in the Atlantic basins. In the Mediterranean area, the Júcar basin (CHJ) stands out as there is a high incidence of positive ERE, unlike the dynamics of the other Mediterranean basins. This bias can be applied to the CHJ, NOR, CGN and CMA basins. For this reason, in the basins that suffer this bias, the majority of the information corresponds only to the episodes of positive ERE.

The towns that account for more than 50 cases of drought were all spatially distributed across the Mediterranean basins, except for Seville, located in the Atlantic watershed (Figure 4). Regarding the drought temporal distribution at the different cities, Murcia case is noteworthy by the regularity of drought episodes compared to the majority of the other cities that exhibit larger temporal variability (Olcina, 2001b). This fact is related to its geographical position in the South-east of Spain. Within this environment, “specific” drought events occur (the so-called “surestinas” south-eastern droughts) related to the lack of precipitation from the Atlantic and absence of Mediterranean rainfall events (Olcina, 2001b).

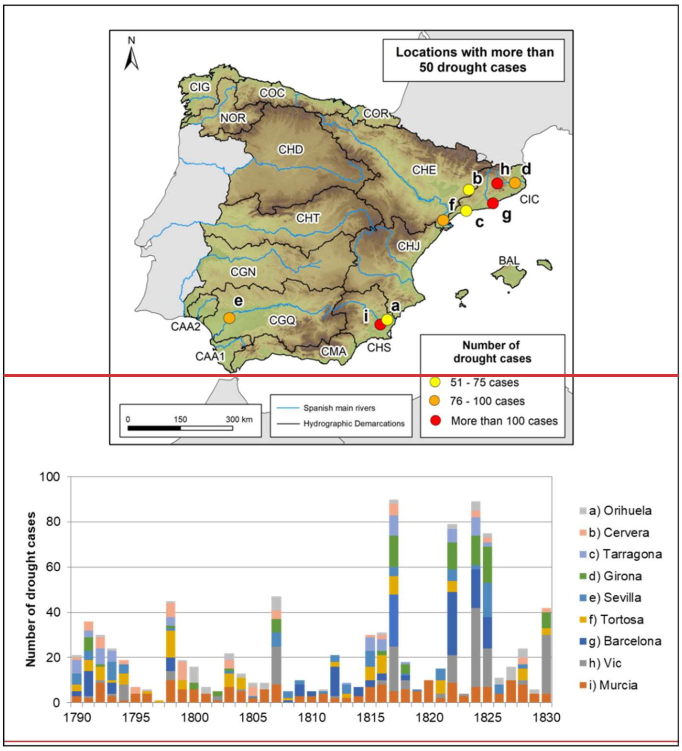


Basin Code	Positive ERE (Nr)	Negative ERE (Nr)
CIG	18	12
NOR	1	0
COC	0	0
COR	4	5
CHD	11	19
CHT	15	24
CGN	5	0
CAA1	0	1
CAA2	0	0
CGQ	71	90
CMA	17	0
CHS	90	265
CHJ	104	4
CHE	75	177
CIC	223	518
BAL	19	15
Total	**653**	**1130**



Basin Code	Positive ERE (Nr)	Negative ERE (Nr)
CIG	18	12
NOR	1	0
COC	0	0
COR	4	5
CHD	11	19
CHT	15	24
CGN	5	0
CAA1	0	1
CAA2	0	0
CGQ	71	90
CMA	17	0
CHS	90	265
CHJ	104	4
CHE	75	177
CIC	223	518
BAL	19	15
Total	**653**	**1130**

Figure 3: Number of positive ERE cases (FF, PF, PR, SS) and negative ERE cases (DR) for the different Spanish river basins during the Early 19th Century (1790-1830). A list of the full names of the basin codes can be found in Figure 1. Elaboration from AMARNA database.



N
Locations with more than 50 drought cases
CIG
COC
NOR
COR
CHD
CHE
CIC
CHT
CHJ
BAL
CGN
CGQ
CAA2
CAA1
CMA
CHS
a
b
c
d
e
f
g
h
i
Number of drought cases
51 - 75 cases
76 - 100 cases
More than 100 cases
0
150
300 km
Spanish main rivers
Hydrographic Demarcations
Number of drought cases
100
80
60
40
20
0
1790
1795
1800
1805
1810
1815
1820
1825
1830
a) Orihuela
b) Cervera
c) Tarragona
d) Girona
e) Sevilla
f) Tortosa
g) Barcelona
h) Vic
i) Murcia

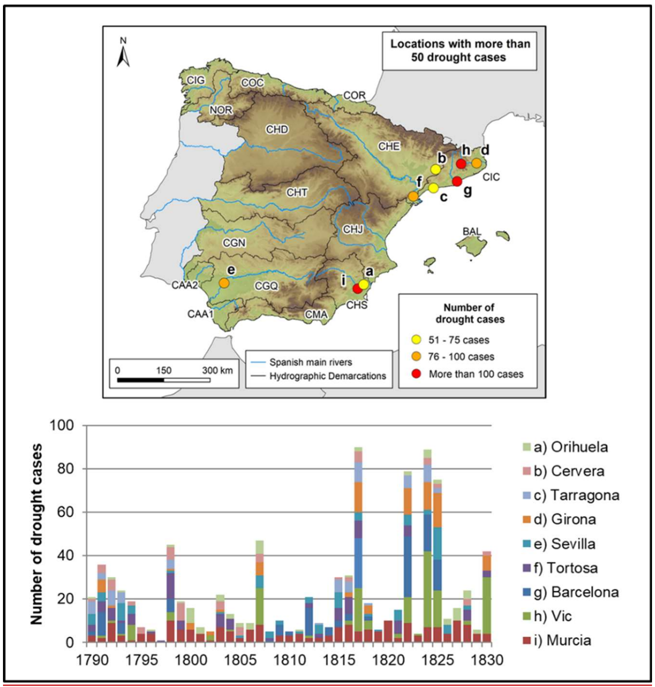


Figure 4: Towns with more than 50 cases of drought during the Early 19th Century. Elaboration from AMARNA database.

3.2. Drought analysis of the Early 19th Century on the Spanish Mediterranean Basin

Table 43 shows historical data from the most severe drought episodes of the Early 19th Century based on all the cases from all the Spanish towns that record rogation ceremonies for each drought episode. In this regard, it will be possible to consider the different nuances that appear in the most representative droughts of the analysed period.

Episode	Year of greatest impact (Nº Cases)	Approximate duration	Total cases
1798 -1799	1798 (58)	25 months	78 cases
1807 -1808	1807 (49)	19 months	55 cases
1812 -1814	1812 (24)	21 months	44 cases
1816 -1818	1817 (122)	37 months	175 cases
1822 -1825	1824 (97)	40 months	279 cases

Table 43: Summary of the severe drought episodes according to historical data for the Early 19th Century. Elaboration from AMARNA database.

The first of these episodes runs from December 1797 to December 1799, with the peak of intensity in March and April 1798. This episode stands out as it occurred several years before the megadrought of 1812-1825 and was possibly an episode still linked to Maldà Oscillation (Barriendos & Llasat, 2003). It affected five hydrographic basins (Catalan basins, Ebro, Segura, Tagus and

Guadalquivir), three of which are Mediterranean (Figure 5). Despite its considerable extension, this episode had a limited duration, with only a few months of rogations. The exception is the municipality of Murcia, where rogations were recorded for 10 of the 25 months that the episode lasted. Furthermore, this episode was noteworthy in this town due to plague outbreaks (Zamora Pastor, 2001).

The second episode of severe drought occurred between January 1807 and July 1808 (Figure 5), with the largest number of cities holding rogations in October 1807. It affected six river basins (Catalan basins, Ebro, Balearic basins, Segura, Duero and Guadalquivir), four of which are Mediterranean. Its main characteristic is that it had a greater impact on towns in the southern sector of the Atlantic and Mediterranean watersheds of the Peninsula, such as Murcia and Seville.

The third episode accumulated less cases of drought but marked the beginning of the megadrought that lasted until 1825, with different regional effects throughout the sequence. It occurred between March 1812 and April 1814 with the peak of greatest severity in April 1812 (Figure 5). Despite the low number of rogations recorded (44), significant effects on crops were documented, causing wheat shortages and widespread famine in the Mediterranean basins. It had a broad impact across the Iberian Peninsula, affecting eight river basins (Catalan basins, Ebro, Balearic basins, Júcar, Segura, Duero, Tagus and Guadalquivir), three of which are in the Atlantic watershed.

The fourth episode runs between December 1815 and November 1818 (Figure 5) and stands out for the impact of the drought during 1817, which was very severe in Catalonia with instrumental records in Barcelona that were unprecedented until then (Moruno, 2021). In this episode, there was an exceptionally dry month (April 1817) in which fourteen of the twenty municipalities recorded *pro pluvia* rogations. This drought affected eight very broadly distributed river basins;: four Mediterranean basins (Catalan basins, Ebro, Balearic basins and Segura) and four Atlantic basins (Galician basins, Duero, Tagus and Guadalquivir). Rogations were made during this drought for many months, particularly in the cities of Murcia and Girona with 12 and 11 months, respectively.

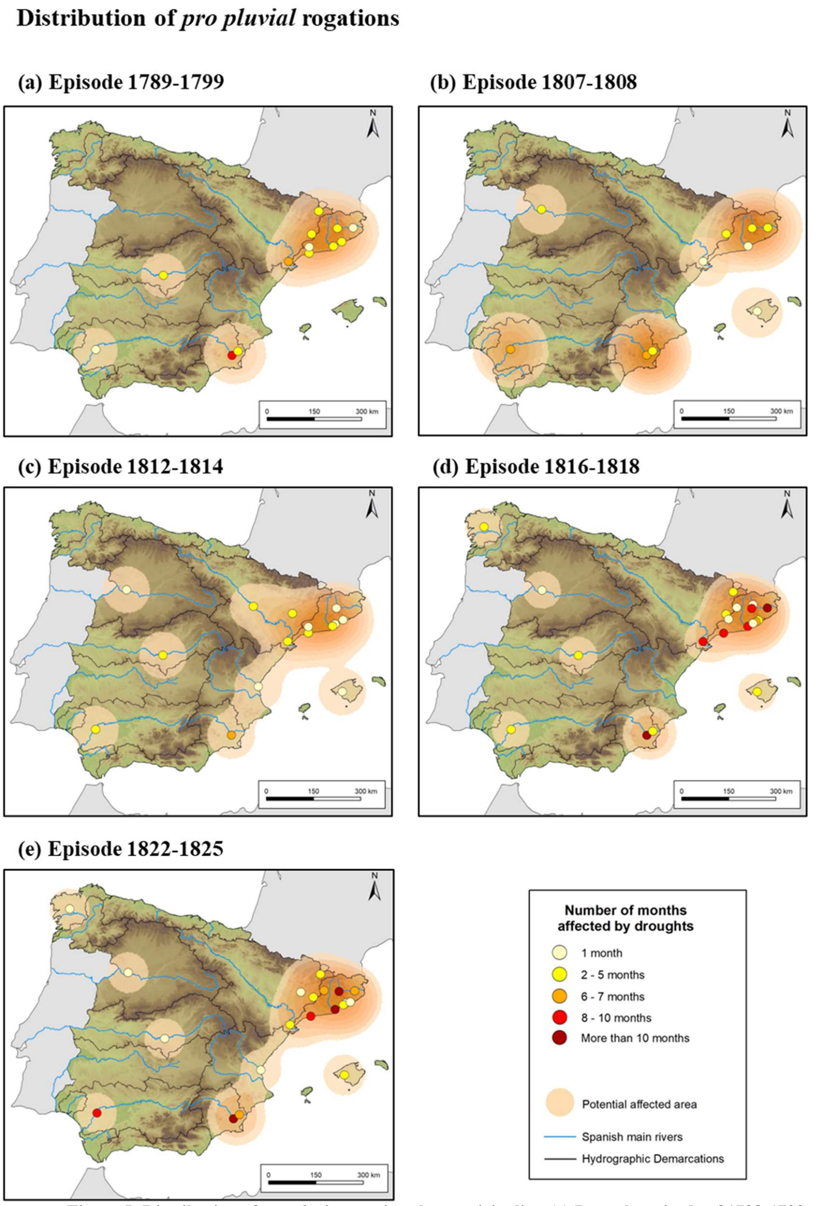


Figure 5: Distribution of *pro pluvia* rogations by municipality. (a) Drought episode of 1798-1799. (b) Drought episode of 1807-1808. (c) Drought episode of 1812-1814. (d) Drought episode of 1816-1818. (e) Drought episode of 1822-1825. Elaboration from AMARNA database using Arc Map GIS Software, applying Kernel Density Tool.

The last episode took place between January 1822 and January 1826 (Figure 5), although the year 1823 recorded a low number of rogations. This drought is noteworthy for being the longest and most persistent of the Early 19th Century (40 months). Three different peaks of severity can be observed: March 1822, April-May 1824 and February 1825. This drought affected eight very broadly distributed river basins: four Mediterranean basins (Catalan basins, Ebro, Balearic basins and Segura) and four Atlantic basins (Galician basins, Duero, Tagus and Guadalquivir). Also significant was the large accumulation of rogations carried out each month in the towns affected. For example, the town of Vic recorded twenty months of rogations, Murcia seventeen months and Barcelona, fifteen. This episode was accompanied by price increases of wheat and the emergence of a locust plague which affected different towns (Azcárate, 1996).

3.3. Analysis of the instrumental precipitation data series of Barcelona (1786-2022)

The analysis of the instrumental precipitation data series of Barcelona (1786-2022) was developed using three different drought indices (SPI, SPEI and Deciles) (Figure 6). The three drought indices reported a significant number of extreme drought events, both in severity and duration, during the Early 19th Century. A dry period between 1812 and 1825 stands out for its significant severity and duration. The three drought records also show values of relative abundant rainfall from the end of the 19th Century until the end of the 20th Century. The beginning of the 21st Century reveals an upturn in the severity and duration of drought episodes with respect to the 20th Century. This dry period that continues to the present day appears to be less intense to those of the Early 19th Century, but may eventually become of similar duration and severity.

The SPI, in comparison with the behaviour of the other two indices, highlights more clearly the peaks of greater severity, both positive and negative (Figure 6). In this regard, 1817 stands out as the driest year in the precipitation data series, with months of maximum severity reaching values close to -4 (-3.91 in the month of August) (Table 65). If we look at the results of this index, it becomes clear that after the Early 19th Century, during the 1830s, the years in drought conditions were prolonged, ending around 1840. From the mid-19th Century, a new phase began with a low presence of prolonged dry periods until the end of the 20th Century. In the 21st Century, severe drought values can be observed again. For example, in 2021, a negative value of the SPI of close to -3 was recorded for the first time since the Early 19th Century. The SPEI shows a different result to the other two indices as it combines rainfall and temperature values. In this respect, it is noteworthy that the most severe year of the series, according to the SPEI was not 1817 but 1822. It is possible that the negative thermal effect of the Tambora eruption (1815) was still significant in 1817, resulting in 1822 having a higher temperature and, consequently, a lower SPEI value. The 1870-1890 drought episode, which does not stand out so much in the other two indices, is also perceived as severe. With regard to the 20th Century, SPEI shows a phase of positive values that lasted twenty years from the 1970s to the 1990s with almost not a single month with negative values. In contrast, for the beginning of the 21st Century there are hardly any years with such positive values (Figure 6). Undoubtedly, the recent thermal warming increases the intensity of negative SPEI values and presents increased problems for water management.

The behaviour of the Deciles index is very similar to that obtained with the SPI index. This index softens the extreme positive and negative behaviours. Thus, the interpretation of rainfall abnormalities ~~do~~does not help, with only the most evident episodes being highlighted.

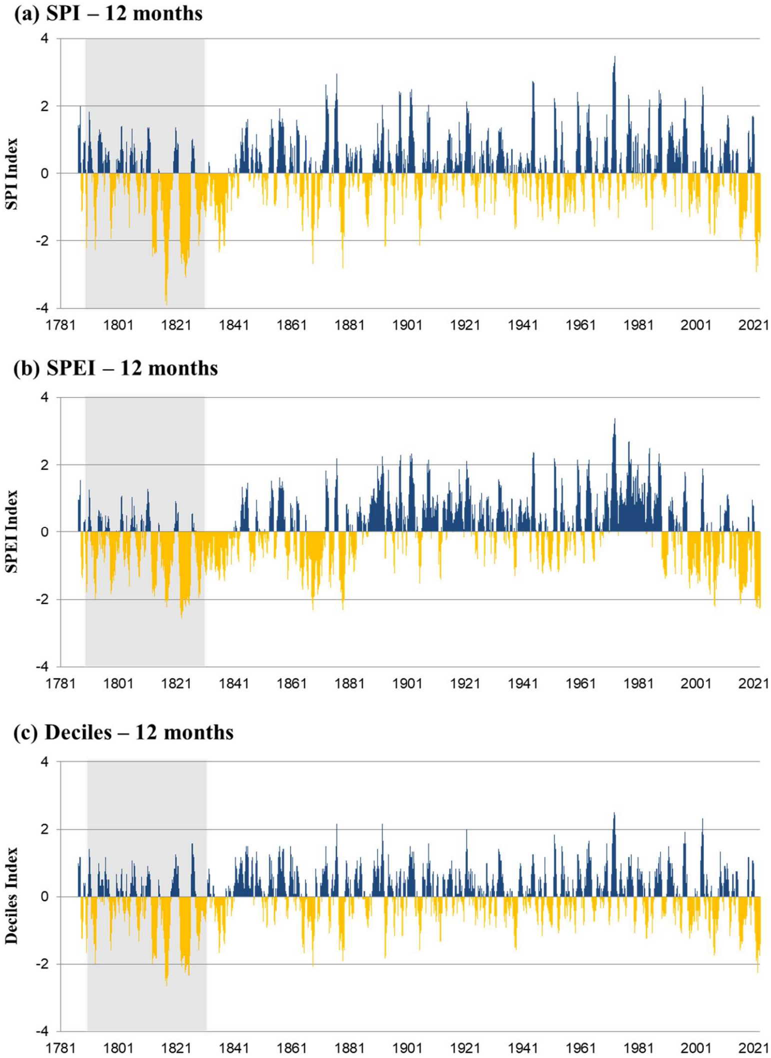


Figure 6: Monthly values of the SPI, SPEI and Deciles indices for the Instrumental precipitation data series of Barcelona (1786-2022). The study period has been shaded in grey. Elaboration with the data obtained from Prohom et al., 2016.

The results obtained with the Pettitt Test are very similar for the SPI and Decile index values, although there are differences with respect to the SPEI index (Table ~~5~~4). The main difference is the position of the first breakpoint which, for the case of the SPI and Deciles, occurred right at the end of the Early 19th Century, in the 1840s. On the other hand, for the SPEI index, this first breakpoint occurred at the end of the nineteenth century, when a strong dry period ended that had lasted from 1860 to 1880 and is much more important in this index than in the other two analysed. With respect to the breakpoint that marks the end of the wet period of the twentieth century, the SPI and Deciles indices coincide with the same period, at the end of 1997. Meanwhile, the SPEI marks it at the end of the 1980s, after the wet phase of the 1970s and 1980s. From this point, the three indices go back to indicating negative averages for their respective series (Table ~~5~~4).

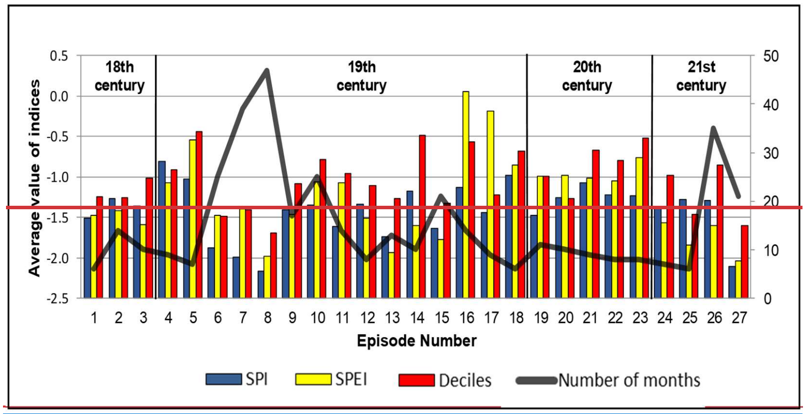
	Pettitt Test Results				
Monthly data series	1rst section's average	1rst breaking point	2nd section's average	2nd breaking point	3rd section's average
SPI	- 0.48 (669 m: 56 yr)	October 1842	0.21 (1862 m: 155 yr)	December 1997	-0.22 (301 m: 25 yr)
SPEI	- 0.43 (1157 m: 96 yr)	June 1883	0.56 (1266 m: 105 yr)	December 1988	-0.52 (409 m: 34 yr)
Deciles	- 0.34 (643 m: 54 yr)	August 1840	0.15 (1886 m: 157 yr)	October 1997	-0.25 (303 m: 25 yr)

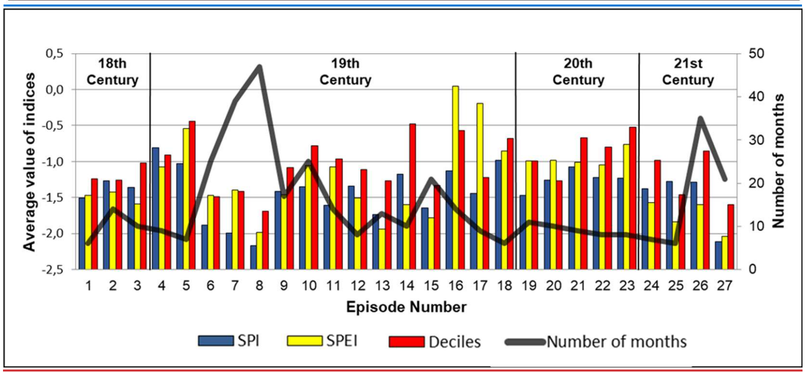
Monthly data series	1rst section's average	1rst Break. Point	2nd section's average	2nd Break. Point	3rd section's average
SPI	- 0.48 (669 m: 56 yr)	OCT 1842	0.21 (1862 m: 155 yr)	DEC 1997	-0.22 (301 m: 25 yr)
SPEI	- 0.43 (1157 m: 96 yr)	JUN 1883	0.56 (1266 m: 105 yr)	DEC 1988	-0.52 (409 m: 34 yr)
Deciles	- 0.34 (643 m: 54 yr)	AUG 1840	0.15 (1886 m: 157 yr)	OCT 1997	-0.25 (303 m: 25 yr)

Table ~~5~~4: Results of the breakpoint analysiss carried out by means ~~according to~~of the Pettitt Test on the three ~~for~~ drought indices used in this study (SPI, SPEI and Deciles). The table describes the date (month and year) on which the first and second breakpoints occur. There are also three more columns indicating the average of the index values between breakpoints, the number of months and the number of years.
Elaboration with the data obtained from Prohom et al., 2016.

Based on the values of the three indices, the drought episodes are summarised for the Barcelona data series (Table ~~6~~5). It reveals a greater number of drought episodes recorded in the 19th Century compared to the ~~21st~~ 20th Century, in which ~~the~~ droughts were not only scarce but also less severe and shorter (Figure 7). This can be confirmed if we consider that in the first twenty years of the 21st Century droughts are being more frequent and severe than during ~~the same number of droughts have been recorded as those occurring throughout the whole of~~ the 20th Century.

The droughts of the Early 19th Century period (Nr. 2 to ~~8~~9) stand out due to their extreme severity, particularly those in the central part of the period, when not only were the droughts severe but also a large number of dry months were concentrated during this time (Table ~~6~~5, Figure 7). For the rest of the drought episodes of the series, we can observe that the majority had shorter duration (Figure 7). According to the average value of indices and the number of dry months only three noteworthy drought episodes are outside of the Early 19th Century: 1877-1879 (Nr. 15), 2015-2018 (Nr. 26) and 2021-2022 (Nr. 27~~-~~).~~Only three noteworthy drought episodes are outside of the Early 19th Century: 1877-1879 (Nr. 14), 2015-2018 (Nr. 23) and 2021-2022 (Nr. 24).~~





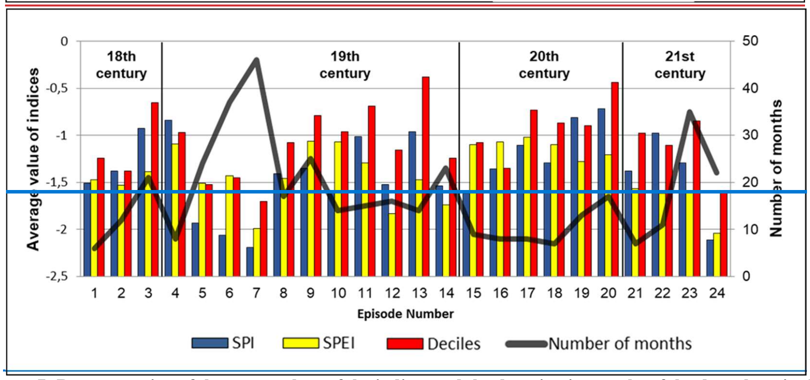


Figure 7: Representation of the mean values of the indices and the duration in months of the drought episodes described in Table 65. Elaboration with the data obtained from Prohom et al., 2016.

Episode Num.	Date		Month Num. *	Averages of index values for each episode			Minimum values of the episodes **			
	Onset	Ending		SPI	SPEI	Dec.	SPI	SPEI	Dec.	Month
1	1789/09	1790/03	6	-1.34	-1.47	-1.24	-2.22	-1.79	-1.67	11/1789
2	1792/05	1793/04	12	-1.38	-1.53	-1.38	-2.29	-2.00	-2.00	01/1793
3	1798/03	1798/12	10	-1.36	-1.59	-1.02	-1.94	-1.76	-1.58	05/1798
4	1807/09	1808/04	8	-0.84	-1.09	-0.97	-1.19	-1.31	-1.33	01/1808
5	1812/05	1814/05	25	-1.88	-1.47	-1.49	-2.46	-1.82	-2.00	10/1812

6	1815/11	1818/11	37	-2.06	-1.43	-1.45	-3.91	-2.24	-2.67	08/1817
7	1822/03	1825/11	45	-2.23	-2.02	-1.75	-3.10	-2.22	-2.17	01/1824
8	1828/01	1829/05	17	-1.41	-1.46	-1.08	-2.14	-1.95	-1.58	10/1828
9	1834/04	1836/04	25	-1.35	-1.06	-0.78	-2.35	-1.44	-1.67	11/1835
10	1836/11	1837/12	14	-1.61	-1.07	-0.96	-2.17	-1.46	-1.42	08/1837
11	1864/04	1864/11	8	-1.34	-1.51	-1.11	-1.71	-1.72	-1.50	09/1864
12	1867/10	1868/10	13	-1.74	-1.94	-1.27	-2.69	-2.32	-2.08	03/1868
13	1869/10	1870/07	10	-1.18	-1.60	-0.48	-1.62	-1.86	-0.75	11/1869
14	1877/05	1879/01	21	-1.64	-1.78	-1.33	-2.80	-2.31	-1.92	08/1978
15	1886/07	1887/08	14	-1.13	0.05	-0.57	-1.60	-0.16	-0.92	03/1887
16	1893/02	1893/10	9	-1.44	-0.19	-1.22	-2.19	-0.79	-1.83	04/1893
17	1904/12	1905/10	11	-1.47	-0.99	-0.99	-2.15	-1.53	-1.58	04/1905
18	1937/11	1938/08	10	-1.26	-0.98	-1.27	-1.65	1.32	-1.50	03/1938
19	1947/05	1948/01	9	-1.07	-1.01	-0.67	-1.40	-1.23	-0.83	10/1947
20	1952/10	1953/05	8	-1.22	-1.05	-0.8	-1.49	-1.20	-1.00	03/1953
21	1965/02	1965/09	8	-1.23	-0.76	-0.52	-1.58	-0.88	-0.75	09/1965
22	2005/03	2005/09	7	-1.38	-1.57	-0.98	-1.79	-1.86	-1.25	07/2005
23	2006/11	2007/04	6	-1.28	-1.84	-1.46	-1.84	-2.15	-1.75	01/2007
24	2015/09	2018/07	35	-1.29	-1.60	-0.85	-2.00	-2.15	-1.50	03/2016
25	2021/04	2022/12	21	-2.11	-2.04	-1.6	-2.92	-2.22	-1.92	09/2021

Episode Num.	Date		Month Num.	Averages of index values for each episode			Minimum values of the episodes **			
	Onset	Ending	*	SPI	SPEI	Dec.	SPI	SPEI	Dec.	Month
1	1789/09	1790/02	6	-1.51	-1.47	-1.24	-2.22	-1.79	-1.67	11/1789
2	1792/05	1793/06	14	-1.27	-1.42	-1.26	-2.29	-2.00	-2.00	01/1793
3	1798/03	1798/12	10	-1.36	-1.59	-1.02	-1.94	-1.76	-1.58	05/1798
4***	1807/09	1808/05	9	-0.81	-1.07	-0.91	-1.19	-1.31	-1.33	01/1808
5	1809/10	1810/04	7	-1.03	-0.54	-0.44	-1.25	-0.74	-0.67	12/1809
6	1812/05	1814/05	25	-1.88	-1.47	-1.49	-2.46	-1.82	-2.00	10/1812
7	1815/09	1818/11	39	-1.99	-1.40	-1.41	-3.91	-2.24	-2.67	08/1817
8	1822/01	1825/11	47	-2.17	-1.98	-1.69	-3.10	-2.22	-2.17	01/1824
9	1828/01	1829/05	17	-1.41	-1.46	-1.08	-2.14	-1.95	-1.58	10/1828
10	1834/04	1836/04	25	-1.35	-1.06	-0.78	-2.35	-1.44	-1.67	11/1835
11	1836/11	1837/12	14	-1.61	-1.07	-0.96	-2.17	-1.46	-1.42	08/1837
12	1864/04	1864/11	8	-1.34	-1.51	-1.11	-1.71	-1.72	-1.50	09/1864
13	1867/10	1868/10	13	-1.74	-1.94	-1.27	-2.69	-2.32	-2.08	03/1868
14	1869/10	1870/07	10	-1.18	-1.60	-0.48	-1.62	-1.86	-0.75	11/1869
15	1877/05	1879/01	21	-1.64	-1.78	-1.33	-2.80	-2.31	-1.92	08/1978
16	1886/07	1887/08	14	-1.13	0.05	-0.57	-1.60	-0.16	-0.92	03/1887
17	1893/02	1893/10	9	-1.44	-0.19	-1.22	-2.19	-0.79	-1.83	04/1893
18***	1896/02	1896/07	6	-0.98	-0.85	-0.68	-1.49	-1.22	-1.17	06/1896
19	1904/12	1905/10	11	-1.47	-0.99	-0.99	-2.15	-1.53	-1.58	04/1905
20	1937/11	1938/08	10	-1.26	-0.98	-1.27	-1.65	1.32	-1.50	03/1938
21	1947/05	1948/01	9	-1.07	-1.01	-0.67	-1.40	-1.23	-0.83	10/1947
22	1952/10	1953/05	8	-1.22	-1.05	-0.8	-1.49	-1.20	-1.00	03/1953
23	1965/02	1965/09	8	-1.23	-0.76	-0.52	-1.58	-0.88	-0.75	09/1965
24	2005/03	2005/09	7	-1.38	-1.57	-0.98	-1.79	-1.86	-1.25	07/2005
25	2006/11	2007/04	6	-1.28	-1.84	-1.46	-1.84	-2.15	-1.75	01/2007
26	2015/09	2018/07	35	-1.29	-1.60	-0.85	-2.00	-2.15	-1.50	03/2016
27	2021/04	2022/12	21	-2.11	-2.04	-1.6	-2.92	-2.22	-1.92	09/2021

Table 65: Drought episodes in the instrumental precipitation data series of Barcelona (1786-2022).

***Number of months determined by the following criteria: Episodes must have at least 6 5 months below "-1" value of SPI Index. The count of months will start and finish with the values below "-0.7570".**

****The month with the lowest value of each episode corresponds to the SPI index. Elaboration with the data obtained from Prohom et al., 2016.**

*****The episodes of 1807-1808 and 1896 have less than 5 monthmonths below -1 but their importance makes it interesting to mention them.**

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4. DISCUSSION

The comparison between the results obtained from the historical data and the instrumental data setseries is part of the main objective of this study. This comparison makes it possible to contrast the reliability of the methods used and to assess the consistency of the results obtained.

The combination of different *proxy* data expands the knowledge on the extreme hydrometeorological events, whether they be excesses or deficits, occurring in the past. In this case, the historical data and the instrumental data setseries of Barcelona have allowed us to analyse one of the driest known periods in the study area (Table 76). The comparison of the standardised values of the historical series with the instrumental indices enables us to observe the synchrony between the historical *proxy* and the instrumental data (Figure 8). The coincidence of the duration of the episodes from the historical data and instrumental seriessets is noteworthy. The only episode for which the durations are different is that of 1807, attributable to the fact that it mainly affected and for longer the southern regions of the Iberian Peninsula. In terms of the severity of the episodes, the coincidence between the two seriessets of data is also noteworthy, with the episodes with most documented cases coinciding with those with a lower SPI index. The only episode that does not follow this pattern is that of 1812, in which the number of negative ERE cases is relatively low. But, on the other hand, according to the SPI, it is the episode with the third lowest mean of the Early 19th Century (1790-1830). The use of elements related to the social vulnerability to drought and extending the length of the data collection in different locations would help to resolve these specific uncertainties and constitute lines of research to be developed in the future.

Episod. Num.	Date according to historical data		Month Num. Hist. Data	ERE DR Num. cases	Date according to instrumental data		Month Num. Inst. Data	SPI episode average
	Onset	Ending			Onset	Ending		
3	1797/12	1799/12	25	78	1798/03	1798/12	10	-1.36
4	1807/01	1808/07	19	55	1807/09	1808/05	9	-0.81
6	1812/03	1814/04	26	44	1812/05	1814/05	25	-1.88
7	1815/12	1818/11	36	175	1815/09	1818/11	39	-1.99
8	1822/01	1826/01	49	279	1822/01	1825/11	47	-2.17

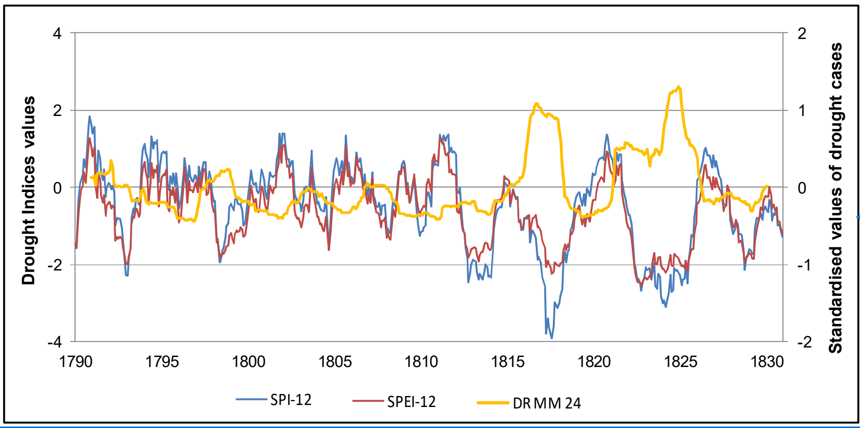
Episod. Num.	Date according to historical data		Month Num.	Date according to instrumental data		Month Num.	Number of cases		SPI episode average
	Onset	Ending		Onset	Ending		ERE Pos.	ERE Neg.	
3	12/1797	12/1799	25	03/1798	11/1799	21	37	78	-0.93
4	01/1807	07/1808	19	09/1807	04/1808	8	17	55	-0.84
5	03/1812	04/1814	21	05/1812	04/1814	24	24	44	-1.93
6	12/1815	11/1818	37	11/1815	11/1818	37	20	175	-2.06
7	01/1822	01/1826	40	01/1822	10/1825	46	41	279	-2.19

Table 76: Characteristics of the five principal drought episodes of the Early 19th Century (1790-1830) according to historical data and instrumental series. Elaboration with the data obtained from Prohom et al., 2016 and the data from the AMARNA database.

Figure 8 shows the coincidence of the droughts according to the historical data (positive values) with the negative oscillations shown by the SPI and SPEI indices. The overlapping of this information highlights the importance of the droughts in the final part of the Early 19th Century, specifically between 1815 and 1825, although the instrumental data indicate that this period could have started in 1812. For this reason, it is desirable to analyse in more detail the three drought episodes in which there is a high degree of alignment between the instrumental data the historical proxy data:

- That of 1798 stands out as it forms part of the rainfall irregularity typical of the Maldà Oscillation (Barriendos & Llasat, 2003). This drought occurred between two phases of intense rainfall. The alternation of floods or heavy rains with droughts is typical in areas with Mediterranean climate. Despite this fact, this is the only drought of the Early 19th Century which precedes and is preceded by flood or heavy rains episodes in the Northeast Iberian Peninsula.
- The drought of 1817 was different as it was the most severe according to the SPI index and was the year during which the most drought cases were recorded in the whole of the Early 19th Century (120). Despite this strong impact, mainly corresponding to the first half of 1817, the episode was not as long as that of 1822-1825 and, for this reason, according to the SPEI index, it was less severe than this latter episode.
- That of 1822-1825 stands out for its duration of around 40 months according to the rogations and 46 months according to the instrumental series. This not only makes it the longest episode of the Early 19th Century but also of the whole of the precipitation series of Barcelona (Table ~~7~~6). Moreover, this episode is the one with the highest severity, both in terms of the accumulation of drought cases (279) and in terms of the SPI average for the episode as a whole. It is also worth mentioning that according to the SPEI index, this episode is the most severe of the entire Barcelona rainfall data series.

Based on the standardised data series of the number of droughts for the Early 19th Century, the correlation coefficient has been calculated with the values of the different drought indices, with which the precipitation sets of Barcelona have been analysed (Table ~~8~~7). To correlate our drought index with the SPI, SPEI and Deciles values, we performed different correlation tests with RStudio software (Posit team, 2024), taking into account the normality or non-normality of the data. In this regard, the results of the Shapiro-Wilk test show that the SPI and Deciles series do not deviate significantly from normality (p-value> 0.05). However, the SPEI series and our drought index, show significant deviations from normality ($p < 0.05$). Given these results, we opted to apply different correlation methods: Pearson's correlation for normally distributed data, Spearman's and Kendall's correlations for data that did not meet the assumption of normality. ~~We can observe that the values of the three indices generate correlations of over -0.5; where the Deciles index has the greatest correlation value, which is inverted (-0.65). The same is confirmed for the coefficients of determination: it verifies that between 35% and 42% of the variability is explained by the behaviour of the variables used (drought indices vs. number of drought cases).~~ Correlation analyses show moderate to weak negative correlations in all cases. In Pearson's correlation, correlation coefficients range from -0.59 to -0.65, with coefficients of determination (R^2) indicating that between 35% and 42% of the variability in the drought indices can be explained by our index. Spearman and Kendall correlations, which do not assume normality of the data, show lower coefficients, suggesting weaker correlations, with R^2 values ranging between 0.08 and 0.22. However, given the specific nature and context of our index, it can be considered a suitable proxy for drought, especially when used in combination with other indices and methods of analysis.



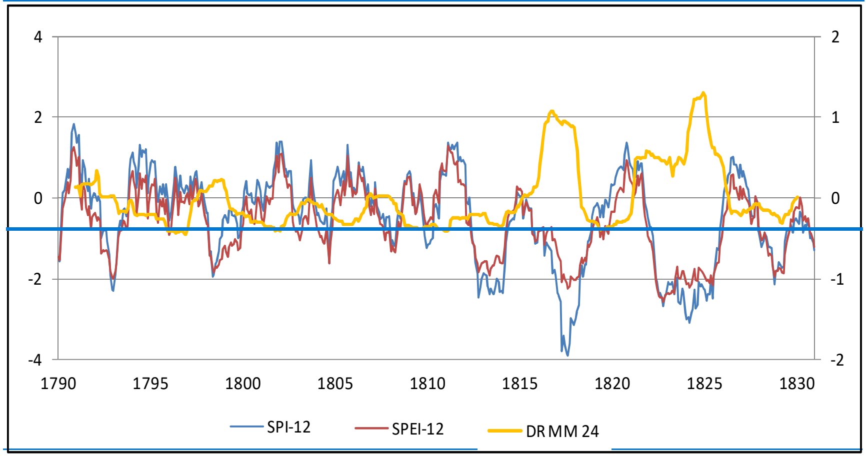


Figure 8: Comparison of the results of the drought indices (SPI and SPEI) and the two-year moving average of the standardised monthly values of the drought cases (DR). Elaboration with the data obtained from Prohom et al., 2016 and the data from the AMARNA database.

~~Index~~	~~Correlation coefficient (R)~~	~~Correlation coefficient (R^2)~~
~~SPI~~	~~-0.62~~	~~0.38~~
~~SPEI~~	~~-0.59~~	~~0.35~~
~~Deciles~~	~~-0.65~~	~~0.42~~

Index	Pearson correlation		Spearman correlation		Kendall correlation	
	Correlation coefficient (R)	Coefficient of determination (R^2)	Correlation coefficient (R)	Coefficient of determination (R^2)	Correlation coefficient (R)	Coefficient of determination (R^2)
SPI	-0.62	0.38	-0.39	0.15	-0.28	0.08
SPEI	-0.59	0.59	-0.44	0.19	-0.31	0.1
Deciles	-0.65	0.42	-0.47	0.22	-0.35	0.12

Table 87: Pearson, Spearman and Kendall correlation and determination coefficients of drought index values and historical data.~~Correlation and determination coefficients.~~

The study of extreme drought episodes in the past is important for understanding the pattern of low frequency episodes and for addressing the droughts occurring in the context of climate change, which

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have erratic behaviour according to the most recent models (IPCC, ~~2021~~2023). Furthermore, the knowledge generated for the study over a long period of time also enables us to better understand the vulnerability of society in different historical contexts and the way in which it has adapted over time to droughts.

Different studies carried out on droughts for the whole of the Mediterranean region for long time periods (Marcos-Garcia et al., 2017; Xoplaki et al., 2018; Kim & Raible, 2021~~; Xoplaki, et al., 2018; Marcos-García, et al., 2017~~) reveal that it is one of the most vulnerable regions to this natural risk within the context of global warming. Taking into account the results of this ~~paper~~study, ~~the~~ the importance of droughts in the Mediterranean region are underlined. ~~are underlined,~~ Given its importance in the current context, it is necessary to analy~~z~~se droughts with ~~have~~ the support of different drought indices, as well as other climatic indicators to determine their severity (Kim & Raible, 2021). The availability of ~~older~~ early instrumental data series is highly important to find a wider range of drought severities and typologies than those found only by analysing the 20th Century. This relationship is evidenced in the research carried out by Erfurt et al., (2020) that combines historical instrumental data with dendrochronological records to analyse the period of the beginning of the 19th Century in south-east Germany. With respect to the use of dendrochronological data to analyse the droughts and megadroughts of the past, the Old-World Drought Atlas is also worth mentioning (Cook et al., 2015). This publication includes a severe drought that occurred at the beginning of the 19th Century, between the Little Ice Age and the Modern Climate period.

Other authors, particularly in the study of the Iberian Peninsula, have used historical data for classifying droughts in the period at the beginning of the nineteenth century (Domínguez-Castro et al., 2012; Gil-Guirado et al., 2019; Gil-Guirado & Pérez-Morales, 2019). It is worth highlighting the article by Domínguez-Castro et al., 2012, in which the historical data is combined with instrumental data to characterise the droughts of the period analysed in Spain. In this case, the same dry periods of great intensity are detected (1817 and 1824) by both the historical and the instrumental data series The authors conclude that the relationship between these droughts and external forcing factors is clear, but more research is also required to confirm it.

Furthermore, the modelling used by (Kim & Raible, 2021) does not show any extraordinary occurrence of droughts for the Mediterranean region as a whole during the Early 19th Century. Neither do these authors relate rainfall patterns with that of those volcanic eruptions emitting more particles into the lower stratosphere, such as Tambora. According to their study, droughts occurring in the Mediterranean are due mainly to the internal dynamics of the climate system and not to external forcing factors (inter-tropical volcanic eruptions and solar radiation variations) (Kim & Raible, 2021). The same conclusion has been obtained for the Eastern Mediterranean region, although for another period than the one studied in this research (Xoplaki~~,~~ et al., 2018). For these reasons, it may be concluded that the relationship between the external forcing factors can lead to different rainfall pattern depending on the region in which specific conditions prevail.

In the case of the Iberian Peninsula, the combination of inter-tropical volcanic eruptions with positive phases of the North Atlantic Oscillation during the first two years after the eruption could result in dry periods for the Iberian Peninsula and in wet phases for Central Europe (Domínguez-Castro et al., 2012). To that, the lack of droughts detected in south-east Germany during the Early 19th Century could reinforce this hypothesis (Erfurt et al., 2020). In this study of droughts for south-east Germany, despite the lack of droughts in the Early 19th Century, there were temporal coincidences with other severe drought episodes, such as those occurring at the end of the 19th Century (between 1857 and 1870) and at the beginning of the 21st Century (2003 to 2018) (Erfurt et al., 2020). This period coincides with two of the most severe episodes of this century according to the records of the instrumental precipitation data series of Barcelona: the drought of 2007-2008 and that of 2015-2018 (see Table ~~6~~5).

5. CONCLUSIONS

The results obtained with this broad time-scale research contribute to a better understanding of drought episodes occurring in the early 21st Century in the study area. Data collection and extension of databases, allows a substantial improvement of knowledge about drought patterns in the study area.

This study has found a time period in which there is an accumulation of particularly severe drought episodes (1812-1825). This period coincides chronologically with Dalton Solar Minimum and an anomaly in volcanic activity (eruptions of Tambora and other volcanoes mentioned). Obviously, the chronological coincidence does not presuppose any cause-effect relationship between the anomalies in solar and volcanic activity and the pluviometric anomalies under study.

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One of the main results achieved in this research is the high negative correlation between the drought historical data and the instrumental precipitation data sets of Barcelona. This correlation validates the historical information for the study of climate droughts in historical perspective. Despite their different origins and methodologies, these two data sources have shown that they can provide information that is comparable, enabling the reinforcement of the importance of the episode recorded, either floods or droughts.

The combined use of instrumental and historical sources shows changes in rainfall variability in specific periods, alternating between periods of heavy rainfall and drought. Accordingly, in the Spanish Mediterranean basins during the Early 19th Century, between 1810 and 1830, the alternation between periods of heavy rainfall and drought is revealed. In contrast, rainfall patterns during the preceding climatic phase of the Maldá Oscillation (1760-1800) were directly opposite to those observed during the Early 19th Century. Additionally, the analysis of instrumental data shows the similar pattern of severe droughts between the end of the LIA and the current context of Global Warming. On the other hand, the 20th Century does not show such a pattern of severe droughts.

The integration of historical documentary sources with instrumental records for identifying severe droughts has yielded promising outcomes. This methodology, leveraging documentary evidence, has been proven viable for periods or regions lacking instrumental data. Building on the success of merging these two climatic information sources, a prospective research direction for the Early 19th Century and other significant climatic epochs involves amalgamating historical data with evidence from other climatic proxies, particularly dendrochronology, alongside instrumental pressure series. Such an approach would enhance our comprehension of the atmospheric processes at a synoptic scale, elucidating the mechanisms behind the most severe drought episodes.

6. REFERENCES

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DATA AVAILABILITY

The historical drought data used in this paper for the period 1790-1830 are presented in a database file.

AUTHOR CONTRIBUTION

Josep Barriendos: Data processing and analysis. Interpretation of the results. Preparation of graphic and cartographic material.

María Hernández Hernández: General revision of the texts and advice on the preparation of the materials.

Salvador Gil-Guirado: Methodological approach and advice on the conceptual criteria for defining drought.

Jorge Olcina Cantos: General review and advice on the conceptual criteria for defining drought.

Mariano Barriendos: Elaboration and organisation of information from historical sources.

COMPETING INTERESTS

The contact author has declared that none of the authors has any competing interests.

ACKNOWLEDGEMENTS

This article has been possible thanks to the support and finance from:

- "Ayudas para estudios de Máster oficial e iniciación a la investigación (AII)", University of Alicante.
- "Ajuts Joan Oró per a la contractació de personal investigador predoctoral en formació (FI)", year 2023. Government of Catalonia.
- "Grupo de investigación de Agua y Territorio", University of Alicante.
- "Grupo de Clima y Ordenación del Territorio", University of Alicante.
- "Departamento de Análisis Geográfico Regional y Geografía Física", University of Alicante.